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# Seawater $^{87}Sr/^{86}Sr$ ratios along continental margins: patterns and processes in open and restricted shelf domains

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**Highlights:**

- Homogeneous $^{87}Sr/^{86}Sr$ ratios similar to the oceanic value in open shelf contexts
- Noticeable heterogeneity of $^{87}Sr/^{86}Sr$ ratios in lagoons and epeiric seas
- Importance of submarine groundwater discharges which may locally supply 10 to 60% of Sr in lagoons
- Oceanic Sr inputs and $^{87}Sr/^{86}Sr$ homogenization controlled by the coastal morphology

**Abstract**

To better constrain the Sr isotope budget in marginal domains without any fluvial inputs, we analyzed the chemical composition and $^{87}Sr/^{86}Sr$ ratio of waters and shells from four locations: two coastal lagoons, one hemipelagic platform and one open marine shelf. Our results highlight homogeneous $^{87}Sr/^{86}Sr$ ratios typical of oligotrophic oceanic waters (OOW) (i.e., 0.709172 ± 0.000023) in the Pacific Tatakoto atoll and along a Mediterranean shore to offshore transect (~25km off Banyuls-sur-Mer, BSM). This attests that oceanic inputs from oligotrophic areas remain the main Sr source in open shelf areas compared with submarine groundwater discharges (SGD) or particulate dissolution influences. In BSM, only foreshore data are more radiogenic, possibly due to rainwater mixing, local groundwater springs or more efficient particle dissolution in the intertidal zone. In restricted areas, we report variable $^{87}Sr/^{86}Sr$ ratios between the Salses-Leucate (France) and Oualidia (Morocco) lagoons. The first one has homogeneous $^{87}Sr/^{86}Sr$ ratio typical of OOW except close to groundwater discharges. In Oualidia, $^{87}Sr/^{86}Sr$ ratios decrease by $1.2\times10^{-3}$ from OOW values close to the Atlantic inlet to progressively less radiogenic ones upstream within the interior of the lagoon. These differences depend on several factors including the leaky, restricted or choked morphology of lagoons modulating the oceanic Sr inputs, but also SGD fluxes whose $^{87}Sr/^{86}Sr$ ratios and Sr concentrations are highly variable according to the nature of rocks leached in karstic aquifer. In Oualidia, the low $^{87}Sr/^{86}Sr$ ratios correspond to high Sr concentrations (up to 150 $\mu mol \cdot l^{-1}$) issued from the dissolution of Mesozoic evaporites, leading to SGD fluxes accounting for 60% of the local Sr budget. Through data compilation, we show that similar $^{87}Sr/^{86}Sr$ gradients and processes prevail at the whole Mediterranean scale. Finally, we postulate that high coastal water retention times can also account for anomalous coastal $^{87}Sr/^{86}Sr$ ratios and that the combination of water mass restriction, SGD, bioadsorption and early diagenetic processes could decrease seawater Sr concentrations in some marginal areas.

## 1. Introduction

Although the strontium isotope compositions ($^{87}Sr/^{86}Sr$) of river waters and submarine hydrothermal sources strongly vary across the globe (with respective averages of 0.7119 and 0.7037; Palmer and Edmond, 1989; 1992; Bach and Humphris, 1999; Davis et al., 2003; Pearce et al., 2015; Chavagnac et al., 2018), the $^{87}Sr/^{86}Sr$ of oligotrophic oceanic waters (OOW) remains homogeneous worldwide at 0.709172 ± 0.000023 (see El Meknassi et al., 2018 for a recent synthesis). This paradigm is due to the long residence time of Sr in the ocean (i.e., ~2.5 Myr; Hodell et al., 1990) compared to the global ocean mixing duration ranging from 1000 to 2000 years (DeVries and Primeau, 2011). However, the OOW $^{87}Sr/^{86}Sr$ ratios fluctuated markedly through the Phanerozoic, as reported from Sr isotope data of fossil archives (Peterman et al., 1970; Burke et al., 1982; Veizer, 1989; Zaky et al., 2018). These secular variations are ascribed either to global geodynamic events modulating the respective contribution of terrestrial (radiogenic) and hydrothermal (unradiogenic) Sr sources to oceans or to paleogeographic and climatic changes modifying the continental weathering of rocks of different age and lithology (Reeder et al., 1972; Brass, 1976; Goldstein and Jacobsen, 1987; Bataille et al., 2017; Peucker-Ehrenbrink and Fiske, 2019). By assuming a homogeneous OOW $^{87}Sr/^{86}Sr$ ratio at a global scale, these long-term $^{87}Sr/^{86}Sr$ variations are thus widely used for dating sedimentary series through chemostratigraphic calibration of new fossil data (McArthur, 1994; McArthur et al., 2012) or testing paleogeographic and climatic hypotheses (Goddéris et al., 2017).

Despite the popularity of this paleoenvironmental proxy, the question of whether fossil carbonate shells may be used to reconstruct past OOW $^{87}Sr/^{86}Sr$ fluctuations is still a matter of debate as most of these organisms thrived in epeiric and shelf domains potentially disconnected from the global oceanic Sr reservoir. This is partly supported by the apparent heterogeneity of modern shell and (euhaline to brachyhaline) seawater $^{87}Sr/^{86}Sr$ ratios from

worldwide marine shelves, with respective ranges of 0.707636–0.710483 and 0.704638–710648 (precision on measured $^{87}Sr/^{86}Sr$ ratio is better than $10^{-5}$; El Meknassi et al., 2018). Significant offsets between well-dated fossil (or bulk carbonate) $^{87}Sr/^{86}Sr$ ratios and the Phanerozoic isotope curve are also regularly reported, even in sedimentary facies indicating outer shelf contexts and/or normal marine environments (Cochran et al., 2003; Nieto et al., 2008; Sessa et al., 2012; Wierzbowski et al., 2012; Eidvin et al., 2014; Schildgen et al., 2014; Briard et al., 2020). While potential analytical, stratigraphical and diagenetic biases may be discarded (Martin and Scher, 2004: Marcano et al., 2015; Bellefroid et al., 2018; Zaky et al., 2018), these mismatches can indicate that, sporadically, seawater $^{87}Sr/^{86}Sr$ ratios of epeiric and shelf domains are not necessarily representative of the global OOW value. In most case, these variations result from dissolved Sr supplies delivered by river waters, with contrasted influences on seawater $^{87}Sr/^{86}Sr$ ratios according to their Sr concentrations and isotopic ratios linked to the age and lithology of local bedrocks (Palmer and Edmond, 1989, 1992; Bryant et al., 1995; Peucker-Ehrenbrink et al., 2010). This terrestrial Sr contribution is common in restricted environments like epeiric seas, estuaries, fjords, lagoons, and bays where oceanic influences are more restricted and where $^{87}Sr/^{86}Sr$ ratios covary landward with salinity (Andersson et al., 1992; Ingram and Sloan, 1992; Israelson and Buchardt, 1999; Basu et al., 2001; Major et al., 2006; Jones et al., 2014; Beck et al., 2013; Wang and You, 2013; Chakrabarti et al., 2018; Shao et al., 2018). However, this river water influence can remain limited in euhaline marginal contexts as, on average, measurable effects on the seawater $^{87}Sr/^{86}Sr$ ratio are not expected to occur if freshwater inputs do not drop salinity below a threshold value of 12 (Bryant et al., 1995). Only 5 % of world river systems have water $^{87}Sr/^{86}Sr$ ratios sufficiently different from the OOW value to significantly modify the local seawater $^{87}Sr/^{86}Sr$ ratio in euhaline marine environments.

Although poorly documented, seawater and shell $^{87}Sr/^{86}Sr$ ratios ranging from 0.70890 to 0.70921 have been yet reported in deep outer shelf domains and lagoons not directly subject to major river supply (Müller et al., 1990a; 1990b; Peckmann et al. 2001; Major et al. 2006; Huang et al., 2011; El Meknassi et al., 2018). This suggests that submarine groundwater discharges (SGD) or strontium released from particulate dissolution have a much stronger impact on the coastal Sr budget than currently expected (Beck et al., 2013; Jones et al., 2014 ; Trezzi et al., 2017). A direct consequence is that, depending on oceanic mixing and water retention times on shelves (i.e., ranging from a couple of days to decades or centuries; Liu et al., 2019), the Sr concentrations and $^{87}Sr/^{86}Sr$ of marginal areas, could not always be representative of the global oceanic reservoir. In other words, seawater from open shelves and epeiric domains could record Sr isotope anomalies just through cumulative effects of sporadic terrestrial Sr inputs in coastal sub-reservoirs not sufficiently mixed with the global oceanic reservoir.

Here, we test how submarine groundwater discharges (SGD) and water mass restriction can influence the Sr isotope composition of shelf water devoid of river influences by measuring and comparing the $^{87}Sr/^{86}Sr$ ratios of waters and shells from four coastal domains with different levels of connection to the open ocean. These include an open shelf Mediterranean transect, two coastal lagoons influenced by groundwater discharges in the Atlantic and Mediterranean contexts, and a Pacific hemipelagic platform. Through data compilation, our aim is also to better document the worldwide variability of seawater $^{87}Sr/^{86}Sr$ ratios in the coastal ocean and to better constrain the environmental processes acting on seawater $^{87}Sr/^{86}Sr$ patterns in these transition zones.

## 2. Geological Setting

The seawater $^{87}Sr/^{86}Sr$ ratios of three restricted lagoons exhibiting different geomorphological, hydrogeological, hydrodynamical and climatic conditions is investigated in this study (Fig. 1): 1) The Tatakoto atoll in the eastern Tuamotu Archipelago (French Polynesia) in the southern part of the Pacific Ocean; 2) the Salses-Leucate lagoon (France) in the northwestern Mediterranean domain; and 3) the Oualidia lagoon (Morocco) in the North Atlantic domain. These three sites were chosen because they are under the influence of different littoral or reefal barriers, their water resources are issued from various karstic reservoirs, and they have different waves and tide forcing overprinted by different climatic conditions. To get a comparison with an open shelf domain, sampling was also carried out in the Mediterranean Sea along a 25 km-long transect from the rocky coastline of Banyuls-sur-Mer to the submarine Lacaze-Duthiers Canyon (western part of Gulf of Lion).

### *2.1.Oualidia Lagoon (Morocco)*

Formed around 8000 – 6500 BP (Ballouche and Carruesco, 1986), the Oualidia coastal lagoon is located on the western coast of Morocco (Abda Doukkala region) (Fig. 1A) and belongs to a coastal endorheic basin called the Coastal Sahel of Oualidia. The lagoon is 7 km long, 400 m wide, and 4m water deep during the high tide season (Kaid Rassou et al., 2005). It is organized around a main channel of 180 m wide parallel to the coastline and separated from the Atlantic Ocean by a Plio-Quaternary shoal (Ballouche and Carruesco, 1986). The lagoon is connected to the ocean by two inlets to the south (130 and 70 m wide) (El Khalidi et al., 2011), but recent anthropic developments for salt exploitation and eutrophication prevention along the northern part allowed punctual oceanic inflows (Hilmi et al., 2009). Its hydrodynamic balance is influenced by: 1) Atlantic marine inflows linked to important semi-diurnal tidal dynamics (i.e., average amplitude of 0.97 m and main currents >1 $m.s^{-1}$; Hilmi et al., 2005; 2009), 2) submarine groundwater supplies estimated from 0.2 to 1.2 $m^3.s^{-1}$ (Fakir et

148 al., 2019), and 3) evaporation rates of 1.3 to 1.5 m.yr$^{-1}$ characteristic of semi-arid conditions
149 in this region (with precipitation rates of 30 mm.yr$^{-1}$) (Hilmi et al., 2009). Along seasonal
150 climatic and hydrodynamic conditions, the recurrent groundwater discharge to the lagoon is
151 highlighted by a net salinity gradient ranging from 4 to 36 along a north to south transect
152 (Hilmi et al., 2017). These groundwater supplies occur all along the lagoon (Fakir et al., 2019)
153 and are predominantly linked to the extended karstic aquifer of the Coastal Sahel of Oualidia
154 (Fakir et al., 2002; Kaid Rassou et al., 2005; Bouchaou et al., 2017). Its upper part (i.e., the
155 most exploited for freshwater consumption) is composed of a 50 m thick Plio-Quaternary
156 limestone-sandstone overlying a less permeable level of upper Hauterivian red clays. Below,
157 the 30 m thick Dridrate limestone formation represents the lower aquifer which is less
158 exploited but still provide the most abundant freshwater resources (Fakir et al., 2002; Kaid
159 Rassou et al., 2005). Finally, the upper Valanginian marls constitute the substratum of the
160 Dridrate aquifer and covers the Berriasian limestones, the upper Jurassic limestones rich in
161 gypsum levels, and Triassic evaporites (Fakir et al., 2002; Fadili et al., 2015). According to
162 Fakir et al. (2002), the dissolution of evaporites would be responsible for high Sr
163 concentrations in local groundwater, reaching 2 to 2.5 times the OOW value.
164
165 *2.2. Salses-Leucate Lagoon (South of France)*
166 The Salses-Leucate lagoon formed during the Flandrian (i.e., 15 kyr BP; Arnaud and
167 Raimbault, 1969; Clanzig, 1987). It corresponds to a shallow coastal basin separated from the
168 Mediterranean Sea by a sandy barrier interrupted by three narrow and dispatched marine
169 inlets (two of them are artificial ones) (Ladagnous and Le Bec, 1997; Fig. 1B). The lagoon is
170 14 km long and 5 km wide, with an average water depth of 1.7 m reaching 3.7 m in its deepest
171 areas (Stieglitz et al., 2013). The region of Salses-Leucate is characterized by a dry season
172 typical of the Mediterranean climate, with annual rainfalls of 588 mm.yr$^{-1}$ and annual

evaporation rates of 1.5 $m.yr^{-1}$ (Ladagnous and LeBec, 1997). Basically, wind-driven currents control the hydrodynamic balance of the lagoon by provoking lagoon outflows and Mediterranean seawater inflows, both of them overprinted by moderate tidal influences (i.e., tidal range of 0.40 m; Ladagnous and le Bec, 1997). The lagoon hydrology is affected by freshwater inputs from two main karstic systems in its southwestern parts (i.e., the Font Dame and Font Estramar discharges) that are responsible for 98% of freshwater inputs into the lagoon, with fluxes of 8400 and 12500 $m^3.h^{-1}$ respectively (Ladagnous and Le Bec, 1997; Stieglitz et al., 2013). Altogether, these parameters produce spatial and seasonal salinity variations in the lagoon ranging from 20 to 30 (Bejannin et al., 2017). Finally, it is worth noting that the Font Estramar and Font Dame karstic resurgences belong to the well-studied Corbières karstic domain, which is composed of Jurassic and Cretaceous dolomites, limestones and marls covering Triassic evaporites, and constitutes one of the biggest karstic network from Europe (Ladouche and Dörfliger 2004; Aunay et al., 2003).

*2.3.Banyuls-sur-Mer transect (South of France)*

The Banyuls-sur-Mer (BSM) coast is located on the most eastern part of the Pyrenean reliefs (i.e., Massif des Albères) where the substratum lithology is dominated by metamorphic and crystalline continental rocks of Cambrian age (Fig. 1C; Got and Stanley, 1974). Influenced by a Mediterranean climate (with annual rainfalls of 575 $mm.yr^{-1}$), this open rocky coast displays no major freshwater input apart from the Baillaury river (mean flow rate of 0.2 $m^3.s^{-1}$) that can show very important freshwater discharges in rainy season (up to 165 $m^3.s^{-1}$ during floods). BSM is characterized by a low tidal range with an average of 0.40 m. From the BSM beach, a 25 km-long W-E transect was carried out at sea up to the Lacaze-Duthiers (LD) Canyon in order to collect surface and bottom seawater (i.e., up to 450 m water depth in the canyon). The LD canyon is one of the numerous canyons of the Gulf of Lion, with a strong bottom current and high-suspended sediment supplies from seasonal gravity flow

198 events (Heussner et al., 2006). The salinity and temperature of seawater is regularly measured
199 by the Oceanological Observatory of Banyuls-sur-Mer in the framework of the SOMLIT
200 program, both at SOLA (in the Banyuls Bay) and MOLA (offshore above the canyon) buoys.
201 The surface seawater salinity varies between 34 and 38 whereas bottom seawater varies
202 between 37 and 39. The salinity at 450 m depth fluctuates slightly around 38 (Durrieu de
203 Madron et al., 2013).

204 *2.4.The Tatakoto lagoon (French Polynesia, Pacific Ocean)*

205 The Tatakoto lagoon is a semi-enclosed hemipelagic carbonate platform located in the
206 eastern Tuamotu Archipelago in the Pacific Ocean (Fig. 1D). This archipelago spreads over
207 1600 km and is characterized by 77 atolls whose reef-barrier caps volcanic basement dated
208 between 44 and 37.5 Myrs (Rougerie, 1994; Pirazzoli, 1998). The climate is tropical and
209 dominated by a seasonal rainfall regime of 50 to 100 mm monthly from June to October and
210 100 to 200 mm monthly from November to May (Andréfouët et al., 2001). The Tatakoto atoll
211 is about 12 km long and 3 km large, and the lagoon covers an area of ~12 $km^2$ (Gilbert et al.,
212 2006) with a maximal water depth measured at 18 m (C. Brahmi, personal communication).
213 The atoll rim is closed in the northern part and semi-opened in the southern part. Water
214 exchanges between the lagoon and the ocean occur through several shallow inlets located on
215 the southern side and seawater mixing depends on tide, wave and wind regimes (Andréfouët
216 et al., 2012; Van Wynsberge et al., 2017). The seawater exchanges between the atoll inner
217 parts and ocean are weak owing to the low tidal range (with a maximum amplitude of 0.2 to
218 0.35 m). In the Tuamotu islands, water percolation through the reef porosity (i.e., 30 to 60%)
219 is linked to erosion-dissolution and karstification processes (Rougerie, 1994). Note that
220 intermediate oceanic seawaters flow through the porosity network of the barrier by endo-
221 upwelling convection, leading to the formation of interstitial waters representing ~50% of the
222 barrier volume (Rougerie, 1994). Thus, the renewal of the atoll seawater volume requires

several weeks up to several months depending on the morphology, the porosity of the barrier and the seasonal hydroclimatic conditions (Rougerie, 1994; Tartinville et al., 1997). Combined with evaporation rates of 0.5 to 1 $m.yr^{-1}$, all these processes impose lagoonal salinities fluctuating between 30 and 44 (Rougerie et al., 1984).

## 3. Material and Methods

### *3.1. Water and mollusc sampling*

In order to check the spatial variability of seawater $^{87}Sr/^{86}Sr$ ratios in coastal domains, 38 water samples and 18 carbonate shells of living molluscs (i.e., bivalves and gastropods) were collected at the four studied sites (Fig. 1; Tables 1 and 2). When and where possible, we collected both the surface and bottom seawater together with molluscs for a given position.

In the field, the water samples were collected in 1L LDPE bottles previously cleaned with ultrapure 1N HCl, filtered *in situ* using 60 ml sterile plastic syringes with 0.2 µm Millipore filters (*in-situ* ultrafiltration was not possible in the field due to logistic constraints). All water samples are void of any microparticles usually found as suspended particle matter in the natural environment. *In-situ* filtration during water collection can significantly reduce microparticle-water interaction and therefore improve the characterization of the dissolved chemical pool (Cotte et al., 2015). They were stored at 4°C in a cold room prior to any sample preparation in a clean laboratory for chemical analyses. In the Oualidia and Salses-Leucate coastal lagoons, the water samples were collected (in April and June 2018, respectively) from 0 to 2 m water depth by diving from a chartered semi-rigid boat at various locations from landward parts affected by groundwater discharges to marine inlets and beaches facing the open marine realm (Fig. 1A,B). Bivalve shells from the Oualidia lagoon include *Crassostrea gigas*, *Venerupis decussata*, and *Solen marginatus*, whereas those from Salses-Leucate correspond to *Mytilus galloprovincialis*, *Crassostrea sp.*, and *Glycymeris glycymeris*. Note

248 that it was possible to collect directly the groundwaters from the Font-Dame and Font-
249 Estramar resurgences at Salses-Leucate lagoon. However, the groundwater discharges of the
250 Oualidia lagoon are mostly submerged, sparse and difficult to locate without *in-situ* chemical
251 and salinity monitoring (Fakir et al., 2019).

252 In the Tatakoto lagoon, two water samples and four *Tridacna maxima* shells have been
253 all collected at 2 m depth from two locations: one close to the southern oceanic inlets and the
254 other one in the northern quiescent part of the inner lagoon (Fig. 1D).

255 Finally, a proximal-distal transect between the BSM beach and LD Canyon was
256 carried out in May 2018 using the oceanographic research ship Néréis II (Station research
257 vessel of Banyuls-sur-Mer, TGIR French Oceanographic Fleet; Fig. 1C), with a regular
258 sampling of both surface and bottom seawater (down to 450 m below sea level; with *ex-situ*
259 water filtration) as well as shells (encompassing unidentified bivalves fragments, *Mytilus*
260 *galloprovincialis* and one gastropod *Turritella communis*) using a sediment grab.

261 *3.2. Chemical composition of waters and shells*

262 All the analytical steps have been performed at the Geosciences Environnement
263 Toulouse laboratory (France). First, the collected water samples were prepared in a clean
264 laboratory for chemical analyses including major (Mg, Ca, Na, K) and trace (Sr) element
265 concentrations as well as the anion (Cl, $SO_4$, Br) concentrations. Anions were determined by
266 anionic chromatography calibrated with an IAPSO seawater standard solution that was diluted
267 at different proportion with MQ-$H_2O$ to cover the entire range of anion concentrations. The
268 IAPSO solution is a seawater standard solution provided by OSIL (UK) and certified for its
269 salinity. The major and trace elements (Ca, Mg, Na, K, Sr) concentrations were measured
270 using an Inductively Coupled Plasma optical emission spectrophotometer (ICP-OES), Horiba
271 Jobin Yvon Ultima 2. We used the IAPSO standard solution to calibrate the instruments
272 (Besson et al., 2014).

The shell samples were abraded, cleaned with Suprapur® acetic acid, and rinsed in Milli-Q water to remove external impurities and periostracum. Smaller bivalves and gastropods samples were entirely crushed with an agate mortar to get a powder, whereas only the inner parts of the umbo (recording the entire life span of the organism) were microdrilled for larger bivalve shells. For major element concentrations of molluscs, we dissolved 10 mg of sample using double-distilled 1 M $HNO_3$ on a hot plate at 70°C. After evaporation of the solution to dryness, it was dissolved in 2N $HNO_3$ and diluted 100 times with MQ-$H_2O$. The solution was then analyzed for Mg, Ca, and Sr concentrations using an ICP-OES. The dissolved elements are 1000 times more concentrated than their limit of detection (>10ppb). The ICP-OES instrument was calibrated using a matrix-matched synthetic standard. One of the calibration solutions was measured every 5 samples to check the instrumental drift during the course of the analyses. The measurements achieved a precision of 2.3% for Ca, 2.8% for Mg and 4.3% for Sr, in line with previous analytical precision on Ca-rich sample (Chavagnac et al., 2005).

*3.3. Sr isotopic compositions*

To determine the $^{87}Sr/^{86}Sr$ ratios, 1 ml of filtered water was transferred to a Savillex beaker prior to evaporation on a hot plate at 70 °C and the residue was dissolved in 0.5 ml of 2N $HNO_3$. For the shells, ~10 mg of powder was dissolved in a Savillex beaker with 2N $HNO_3$ on a hot plate at 70°C and centrifuged to exclude any residual fraction. Strontium was isolated from resulting solutions using Eichrom Sr-Spec column, a conventional elution protocol (Pin et al., 2014). The isolated Sr fraction was dissolved with 2µl of 0.5N $H_3PO_4$ solution of which 1µl was loaded on a tungsten filament together with 1µl of Tantalum activator. The Sr isotopic composition was measured using a Thermo Triton Plus thermal ionization mass spectrometer. The $^{87}Sr/^{86}Sr$ ratios were defined as the average of 150 measurements of ion intensities following the static multi-collection mode normalized to

298 $^{86}Sr/^{88}Sr$ = 0.1194. During the measurement period, the standard NBS 987 gave $^{87}Sr/^{86}Sr$
299 ratios of 0.710280 ± 0.000003 (2SD, standard deviations, n = 5). Over the 2018 year, the
300 measured NBS 987 isotopic compositions provided an external reproducibility of 0.710280 ±
301 0.000012 (2SD; n = 26). Regarding international seawater standards (IAPSO and NASS-6),
302 they were processed in the clean room laboratory in the same manner as for our samples for
303 Sr chemical isolation. The external $^{87}Sr/^{86}Sr$ reproducibilities are 0.709179 ± 0.000007 (2SD;
304 n = 7) for IAPSO and 0.709180 ± 0.000014 (2SD; n = 5) for NASS-6, which agree very well
305 with their certified values of 0.709179 ± 0.000005 and 0.709173 ± 0.000018, respectively
306 (Neymark et al., 2014).

307 For the bivalve shells, we analyzed at repeated times the Sr isotopic composition of
308 the JCt-1 international standard (a giant clam *Tridacna gigas*) treated in the clean room as for
309 our shell samples. The measurements provide an external reproducibility of 0.709176 ±
310 0.000015 (2SD; n = 11) for a certified value of 0.709150 ± 0.000050 (Ohno and Hirata, 2007).
311 For comparison purposes between the $^{87}Sr/^{86}Sr$ ratios reported in this study and the literature
312 data (tables 1 and 2), the $^{87}Sr/^{86}Sr$ ratios were corrected by adjusting the measured NBS 987
313 standard reference value to a certified value of 0.710250.

314 *3.4. Geochemical and Sr isotopic database*

315 To better depict potential influences of groundwater in the Oualidia and Salses-
316 Leucate lagoons, we compiled a dataset of published major and trace elements concentrations
317 (Ca, Mg, Na, K, $SO_4$, Cl, Br and Sr) and $^{87}Sr/^{86}Sr$ ratios corresponding to groundwaters
318 representative of Moroccan aquifers ($n$ = 112) and the Corbières aquifer in France ($n$ = 141)
319 (see supplementary data). In order to analyze the geochemical patterns at a worldwide scale,
320 we also compiled a dataset gathering 522 $^{87}Sr/^{86}Sr$ ratios and 571 Sr concentrations
321 representative of OOW and shelf waters (SW) (see supplementary data). As for our samples,

all $^{87}Sr/^{86}Sr$ ratios reported in the literature were corrected by adjusting their respective NBS 987 standard reference value to a certified value of 0.710250.

## 4. Results

### *4.1. $^{87}Sr/^{86}Sr$ of waters and shells*

The $^{87}Sr/^{86}Sr$ ratios of lagoon waters and shells are reported in Tables 1 and 2. First, at the Tatakoto lagoon, waters and shells exhibit $^{87}Sr/^{86}Sr$ ratios at 0.709165 - 0.709171 and 0.709164 – 0.709182, respectively, which are both indistinguishable from the average OOW value of 0.709172 (Figs. 1D and 2A). Similarly, seawater collected at different water depths along the BSM transect presents a slight dispersion ranging from 0.709166 to 0.709176, but still around the OOW value (Figs. 1C). The only sample deviating from this world value is BA9 collected in the Banyuls-sur-Mer beach: it presents a low salinity of 20 and a $^{87}Sr/^{86}Sr$ ratio of 0.709206 slightly more radiogenic than the OOW range (Fig. 2A). Molluscs from the BSM transect present Sr isotope compositions between 0.709164 and 0.709175, all identical to those of surrounding water (Fig. 2A).

In the Salses-Leucate lagoon, the seawater $^{87}Sr/^{86}Sr$ ratios ranging from 0.709155 and 0.709170 are typical of the OOW variability range. Contrastingly, three water samples (LEU 6, 7, 8) from the western part of the lagoon display values ranging from 0.708903 to 0.709132, that is below the OOW variability range. However, these values are trending towards the signatures of the Font Dame and Font Estramar discharges (LEU 10, 11) at 0.708618 and 0.708401, respectively (Fig. 1B). Again, these relatively low isotopic ratios occur along a salinity drop from 37.4 to 11.8 (Fig. 2A). In this context, the least radiogenic Sr isotope composition is measured for the bottom water sample (LEU 8) compared to the surface one (LEU 7). Mollusc shells sampled in this lagoon present $^{87}Sr/^{86}Sr$ ratios varying between 0.709125 and 0.709177, being either similar to or less radiogenic than the OOW value (Table

2; Fig. 2A). Finally, it is worth mentioning that our isotopic results are within the variability of those obtained for local seawater and groundwater sampled in 1998, 2002 and 2012 by Ladouche et al. (2000), Ladouche and Dörfliger (2004) and Petelet-Giraud et al. (2016).

At the Oualidia lagoon, both seawater and mollusc shells display highly variable $^{87}Sr/^{86}Sr$ ratios ranging from 0.708389 to 0.709168 and 0.708939 to 0.709179, respectively (Tables 1 and 2; Figs. 1A and 2A). Downstream, close to Atlantic seawater incursions by the inlets, the $^{87}Sr/^{86}Sr$ ratios of waters stay within the OOW variability range, while they are less and less radiogenic upstream between the OU12 and OU8 locations (Table 1, Figs. 1 and 2). As in Salses-Leucate and Banyuls bay, this isotopic shift is accompanied by a salinity decrease from 37.2 to 9.9 (Fig. 2A). However, the water sample OU6 deviates from this general tendency: although close to the northern dam, it records a $^{87}Sr/^{86}Sr$ ratio typical of OOW values (Fig. 1A). Finally, we note that most bivalve shells record $^{87}Sr/^{86}Sr$ ratios typical of OOW values but different of surrounding water values (Fig. 2A, Table 2). Only the O4 sample (i.e., an endobenthic *Solen marginatus*) fits the general $^{87}Sr/^{86}Sr$ decrease of lagoonal waters.

### *4.2. Geochemical composition of waters*

All concentrations of anions and major and trace elements measured on the collected waters are reported in Table 1. In order to determine the origin of the waters influencing the $^{87}Sr/^{86}Sr$ of local seawater, only the most discriminating elements such as Sr, $SO_4$, and Ca are used here, plotted as function of salinity, and compared to published values of regional groundwater aquifers (Fig. 2).

In the Pacific Tatakoto atoll, seawater has salinity and $^{87}Sr/^{86}Sr$ ratios typical of OOW while they display lower Sr contents (~73.5 $\mu mol \cdot l^{-1}$ against 88 $\mu mol \cdot l^{-1}$ for OOW; Millero et al. 2008). Slight enrichments in $SO_4$ (30.8 $mmol \cdot l^{-1}$) and Ca (10.3 $mmol \cdot l^{-1}$) are recorded but still within the OOW variability range. These geochemical characteristics also concern most

372 water samples from the Mediterranean BSM transect, whatever the water depth, with
373 respective Sr, $SO_4$ and Ca concentrations of 72.2 – 74.5 $\mu mol \cdot l^{-1}$, 30.9 – 32 $mmol \cdot l^{-1}$, and 11.2
374 – 12.2 $mmol \cdot l^{-1}$. The only exception is the beach sample BA9 which low salinity value of 26.4
375 is mirrored by low element concentrations compared to those of OOW (Table 1 and Fig. 2).
376 Euhaline waters from the inner parts of the Oualidia and Salses-Leucate lagoons are
377 also characterized by similar depletions in Sr and slight enrichments in Ca and $SO_4$ when
378 compared to the areas described above. However, the element concentrations evolve
379 differently in Salses Leucate and Oualidia while the salinity and $^{87}Sr/^{86}Sr$ are decreasing
380 landward. In Salses-Leucate, the Sr, $SO_4$ and Ca contents are dropping down to 17.2 $\mu mol \cdot l^{-1}$,
381 3.8 $mmol \cdot l^{-1}$ and 3.5 $mmol \cdot l^{-1}$, respectively, close to the Font Dame and Font Estramar
382 discharges. In agreement with previous analyses of Ladouche et al. (2000) and Ladouche and
383 Dörfliger (2004), these low concentrations are similar to those prevailing at the scale of the
384 regional aquifer and the covariations between salinity, $^{87}Sr/^{86}Sr$ ratio and these elemental
385 concentrations are consistent with freshwater-seawater binary mixing. In Oualidia, the salinity
386 and $^{87}Sr/^{86}Sr$ decrease while Sr and Ca concentrations remain relatively high and stable (i.e.,
387 ~73 $\mu mol \cdot l^{-1}$ and ~11.5 $mmol \cdot l^{-1}$, respectively), except in the salt marshes where values are
388 much higher (i.e., 117 $\mu mol \cdot l^{-1}$ and 20.5 $mmol \cdot l^{-1}$, respectively). Contrastingly the $SO_4$
389 contents decrease with salinity but toward a brackish groundwater end-member with
390 anomalously high $SO_4$ concentrations (~15 $mmol \cdot l^{-1}$). Oualidia geochemical patterns are
391 compatible with the high Sr, $SO_4$, and Ca contents recorded in some Moroccan groundwater
392 aquifers, and especially those related to the dissolution of Mesozoic carbonate and gypsum
393 levels (see Supplementary data).

## 5. Discussion

Studies with high precision $^{87}Sr/^{86}Sr$ ratio have confirmed that modern oligotrophic oceanic waters (OOW) without terrestrial inputs have a quite homogeneous $^{87}Sr/^{86}Sr$ ratio of 0.709172 with a variability range of only ±0.000023 at the global scale (Andersson et al., 1992; Winter et al., 1997; Huang and You, 2007; Huang et al., 2011; Mokadem et al., 2015; Pearce et al., 2015; Trezzi et al., 2017; El Meknassi et al., 2018). This homogeneity is yet regularly challenged for shelf domains as more radiogenic or unradiogenic terrestrial Sr supplies from rivers or SGD can introduce a dispersion to the measured $^{87}Sr/^{86}Sr$ of $6\times10^{-3}$ or more (El Meknassi et al., 2018) (Fig. 3A). Nevertheless, this estimate mainly concerns estuaries, fjords, bays or epeiric seas supplied by major river systems for which data are relatively abundant (e.g., Ingram and Sloan, 1992; Jørgensen and Banoeng-Yakubo, 2001; Négrel et al., 2005; Jørgensen et al., 2008; Wang and You, 2013; Eissa et al., 2016; Casse et al., 2019). In this context, new Sr isotope data from open shelves and coastal lagoons become crucial to get a better view of processes at play in neritic domains, especially those locations deprived of river and delta influences. These areas represent 40 – 50% of the world coasts (Dürr et al., 2011).

### *5.1. Seawater $^{87}Sr/^{86}Sr$ ratios in oceanic and open shelf domains*

With an average Sr concentration of 87.4 $\mu mol \cdot l^{-1}$ (de Villiers 1999), oceanic inflows are the main dissolved Sr source to shelf domains and hemipelagic carbonate platforms. If these oceanic inflows remain constant through time and sufficiently mixed with coastal waters, locally homogeneous seawater and shell $^{87}Sr/^{86}Sr$ ratios close to OOW value of 0.709172 would be expected, even in semi-enclosed areas. Far from any terrestrial influence, this is clearly illustrated in the isolated Tatakoto atoll where Pacific seawater is the main dissolved

419 Sr source, conveyed to the shelf by endo-upwelling circulation through the reef porosity
420 followed by surface currents (Rougerie, 1994). Rainwater infiltrations responsible for fossil
421 reef karstification and possibly volcanic basement weathering are important in the Tuamotu
422 archipelago (Guy et al., 1992; Waljeski, 2003). However, unradiogenic Sr supplied from local
423 SGD appears too limited to modify the OOW $^{87}Sr/^{86}Sr$ ratio of studied waters and giant clams
424 in this Pacific atoll (Table 1 and 2).

425 In the nearshore domain exemplified by the 25 km-long Mediterranean BSM transect,
426 seawater and shells from surface down to ~450 mbsl display a homogeneous $^{87}Sr/^{86}Sr$ ratio
427 with an average value at 0.709171 ± 0.000008 (2SD; n=12) indistinguishable from the OOW
428 value (El Meknassi et al., 2018). In the absence of noticeable dissolved Sr inputs from rivers
429 or SGD, this implies that the Sr inputs resulting from boundary-exchange processes acting at
430 the sediment-water interface are not sufficient to modify the seawater $^{87}Sr/^{86}Sr$ ratio in these
431 well-mixed open shelf conditions. This contrasts with estuaries or deltas where dissolved Sr
432 released from suspended particulate dissolution may alter the $^{87}Sr/^{86}Sr$ of seawater in a few
433 months (Jones et al., 2014; Jeandel and Oelkers, 2015). Moreover, the shell data from the
434 BSM transect indicate that the Sr isotope composition of local seawater did not significantly
435 change over life spans of 2 to 3 years. However, these results do not preclude caution in
436 paleoenvironmental interpretation of shell $^{87}Sr/^{86}Sr$ ratios from the BSM coast and more
437 generally nearshore contexts. This is because SGD influence may strongly vary in time and
438 space along continental margins (Burnett et al., 2006), and especially in the Gulf of Lion
439 where SGD are very important (Rodellas et al., 2015).

440 In the foreshore zone, additional factors can impact the seawater $^{87}Sr/^{86}Sr$ ratios. This
441 is illustrated by the BA9 water sample collected on the BSM sandy beach, characterized by a
442 more radiogenic value and a much lower salinity (i.e., 26.4) than the other samples. We
443 propose that this more radiogenic Sr value can be related to three factors typical of intertidal

444 zones: 1) mixing with rainwater whose regional $^{87}Sr/^{86}Sr$ ratios range from 0.708993 to
445 0.709225 (Khaska et al., 2013), 2) sparse groundwater supplies whose Sr concentrations and
446 radiogenic $^{87}Sr/^{86}Sr$ ratios mirror those from Quaternary, Pliocene, and Paleozoic aquifers (i.e.,
447 0.708142 to 0.713782; Petelet et al., 1998; Petelet-Giraud et al., 2016), and 3) important
448 dissolved Sr amounts released from suspended sandy particles continuously shaked in the surf
449 zone (Kalnejais et al., 2010; Jeandel, 2016; Fabre et al., 2019) with particles $^{87}Sr/^{86}Sr$ ratios of
450 0.72232 – 0.72901 reflecting their metamorphic substratum (Brems et al., 2013). At broader
451 scale, similar local processes can explain part of isotopic deviations reported in the rare water
452 and shell data from open shelves (e.g., North Sea, Gulf of Guinea, South China Sea;
453 Jørgensen and Banoeng-Yakubo, 2001; Jørgensen et al., 2008; Huang et al., 2011; El
454 Meknassi et al., 2018) (Fig. 3). Moreover, marginal areas are transition zones where the
455 oceanic mixing is not instantaneous and may take several years according to littoral currents,
456 coastal physiography, and shelf width (Liu et al., 2019) (Fig. 3B). Terrestrial dissolved Sr
457 sources, even remote, could slightly impact local seawater $^{87}Sr/^{86}Sr$ ratios through cumulative
458 effects. Thus, we advocate to reject any water and/or shell sample from beach or intertidal
459 zones to infer modern or past OOW $^{87}Sr/^{86}Sr$ ratios.
460
461 *5.2. $^{87}Sr/^{86}Sr$ patterns in restricted coastal domains*
462 Supplementing sparse literature data (e.g., Müller et al. 1990b; Beck et al. 2013; Shao
463 et al. 2018), the seawater and shell results from the Oualidia and Salses-Leucate lagoons show
464 clear evidence that restricted coastal domains, even without river influence, display very
465 distinct $^{87}Sr/^{86}Sr$ ratios depending on the respective contribution of oceanic and SGD influxes
466 (Fig. 1): 1) quite homogeneous OOW $^{87}Sr/^{86}Sr$ ratios except in close vicinity of SGD in
467 Salses-Leucate; or 2) variable $^{87}Sr/^{86}Sr$ ratios ranging from OOW values at oceanic inlets to
468 unradiogenic signals in the upstream parts of the Oualidia lagoon.

469 At broader scale, such Sr isotopic variabilities are also noticeable in epeiric seas. By
470 compiling literature data and excluding the anomalous $^{87}Sr/^{86}Sr$ ratios (i.e., samples from
471 beach or close to SGD) from Salses-Leucate, Banyuls-sur-Mer and Venice (Fig.4A), we
472 observe that seawater $^{87}Sr/^{86}Sr$ ratios are typical of OOW values in western Mediterranean
473 Sea (i.e., Spain, France, Italy), whereas those from eastern domains like the Bannock Basin,
474 Aegean Sea, Marmara Sea and Black Sea decrease down to 0.70912 (this study; Müller et al.,
475 1990a; Clauer et al., 2000; Peckmann et al., 2001; Major et al., 2006; Beck et al., 2013;
476 Petelet-Giraud et al., 2016; Trezzi et al., 2017; Teichert et al., 2018). Highlighted here for the
477 first time, these regional $^{87}Sr/^{86}Sr$ variations are consistent with the isotopic variability of
478 Mediterranean water masses illustrated by the Nd isotopes (Tachikawa et al. 2004). Whether
479 local or supra-regional, spatial Sr isotopic variability is insightful to decipher the respective
480 influence of SGD and oceanic inputs on the Sr isotope budget of semi-enclosed areas.
481
482 *5.2.1. SGD influences*
483 With relatively high Sr concentrations (i.e., 2.9 $\mu mol \cdot l^{-1}$) compared to river waters (i.e.,
484 0.5 to 1.2 $\mu mol \cdot l^{-1}$), SGDs represent an unradiogenic Sr source that can account for 13 - 30%
485 of the global Sr ocean budget (Basu et al., 2001; Krabbenhöft et al., 2010; Beck et al., 2013).
486 Depending on their geochemical and hydrodynamical characteristics related to the age and
487 nature of weathered rocks but also oceanic intrusions, meteoric recharges, and water residence
488 time in the aquifers (e.g., Négrel et al., 2003; Fadili et al., 2016, 2015; Santoni et al., 2016),
489 their impact on the elemental and isotopic Sr budgets are, however, highly variable in coastal
490 domains (e.g., Beck et al., 2013; Trezzi et al., 2017; Shao et al., 2018). This is illustrated by
491 comparing the Salses-Leucate and Oualidia lagoons.
492 In the Salses-Leucate lagoon, the low salinities and the low Sr, $SO_4$ and Ca
493 concentrations of Font Estramar and Font Dame discharges show that these groundwaters are

494 typical of meteoric freshwaters percolating through the local karstic aquifer (Fig. 2B,C,D). To
495 better constrain the origin of the collected seawaters in the lagoon, we applied the isotopic
496 mixing equation (Albarède, 1996) to geochemical data of regional end members presented in
497 Table 3:

498 $$^{87}Sr/^{86}Sr_{(mix)} = \left( \frac{^{87}Sr/^{86}Sr_{(SW)} \times [Sr]_{(SW)} \times \%_{(SW)}}{[Sr]_{(mix)}} \right) + \left( \frac{^{87}Sr/^{86}Sr_{(GW)} \times [Sr]_{(GW)} \times \%_{(GW)}}{[Sr]_{(mix)}} \right)$$

499 ,where the mixed water $^{87}Sr/^{86}Sr_{(mix)}$ ratio varies between the local seawater (SW) and
500 groundwater (GW) end members of given Sr concentrations and $^{87}Sr/^{86}Sr$ ratios, using mixed
501 water Sr concentrations $[Sr]_{(mix)}$ calculated from cumulative 10% increments of GW to SW.
502 The low $^{87}Sr/^{86}Sr$ ratios of Font Estramar and Font Dame discharges correspond to the lowest
503 isotopic values of regional groundwaters and indicate a preferential dissolution of Jurassic and
504 Cretaceous carbonates characterized by low $^{87}Sr/^{86}Sr$ ratios compared to the Pliocene and
505 Quaternary silicates ones (Fig. 4A; Petelet et al., 1998; Ladouche et al. 2000; Aquilina et al.,
506 2002; Ladouche and Dörfliger 2004; Khaska et al., 2013; Petelet-Giraud et al., 2016). Given
507 the low Sr concentrations of Jurassic and Cretaceous GW endmembers (Table 3; Fig. 2B), our
508 results suggest that the SGD contribution to the lagoon Sr budget is broadly less than 15%
509 despite local SGD fluxes exceeding 5.55 $m^3.s^{-1}$ (Fig. 4B; Ladagnous and Le Bec, 1997;
510 Stieglitz et al., 2013).

511 In Oualidia, the dissolved Sr supply of SGD with much lower estimated fluxes of 0.2 –
512 1.6 $m^3.s^{-1}$ (Hilmi et al., 2005; Fakir et al. 2019) markedly modifies the seawater $^{87}Sr/^{86}Sr$
513 ratios within the upstream parts of the lagoon (Figs. 1A). These induced modifications can be
514 related to considerable unradiogenic Sr inputs while maintaining constant dissolved Sr
515 concentrations (Fig. 2A,B). Indeed, groundwaters from the Coastal Sahel of Oualidia exhibit
516 variable but generally high Sr and Ca concentrations reaching 2 to 2.5 times the OOW value
517 as well as high $SO_4$ amounts (Fakir et al., 2002; Hilmi et al., 2005; Fig. 2B,C,D). Even if
518 oceanic water intrusions are not excluded in the coastal aquifer, specific Ca and $SO_4$ excesses

519 indicate that the higher Sr contents would result from a preferential dissolution of Sr rich
520 gypsum and celestite ($SrSO_4$) present in the Jurassic (and Cretaceous) limestones (Fakir et al.,
521 2002; Hilmi et al., 2005; Kaid Rassou et al., 2005; Bouchaou et al., 2017). This hypothesis is
522 in agreement with the unradiogenic $^{87}Sr/^{86}Sr$ ratio of lagoon waters (i.e., down to 0.708389)
523 whose values are close to Jurassic aquifers of Morocco (Bouchaou et al., 2017; Vinson et al.,
524 2013) (Fig. 4A). We estimate that groundwater fluxes dissolving Mesozoic sulphates spatially
525 contribute from ~10 to 60 % to the lagoon Sr budget, the higher Sr concentrations being
526 related to high evaporation rates in the salt marshes (Fig. 4A). This Sr supply markedly varied
527 in a few years, as shown by the isotopic scatter between local seawater and shell data at a
528 same location or shells sampled in 2006 (El Meknassi et al., 2018) and 2018 (this study) (Fig.
529 2A). In contrast, the low $^{87}Sr/^{86}Sr$ ratio recorded by one endobenthic *Solen marginatus*
530 suggests that groundwater influence remained more constant in the interstitial waters at 80 cm
531 sediment depth (i.e., the maximal burial depth of this species) than in the lagoon waters.
532 With Sr concentrations up to 150 $\mu mol \cdot l^{-1}$ leading to local contributions of 60% to the
533 Sr budget of lagoons, our results call to reconsider the importance of SGD in link to the
534 dissolution of evaporitic deposits. In association with geothermal brines, their impact on the
535 global ocean Sr budget still remains difficult to appraise but could be regionally important
536 given the proximity of Messinian or Triassic evaporites such as those of Mediterranean
537 margins (e.g., Stein et al., 2000). Considering their elevated Sr concentrations, it is also
538 important to note that, contrarily to river-dominated contexts requiring considerable
539 freshwater inputs to alter the coastal seawater $^{87}Sr/^{86}Sr$ ratio, SGD may quickly change this
540 isotopic ratio in euhaline conditions. Indeed, most major world river systems produce
541 measurable effects on seawater $^{87}Sr/^{86}Sr$ ratios when salinity drops below an average
542 threshold of 12 (Bryant et al., 1995). Only ~15, 10 and 5 % of the rivers have significant
543 effects at salinities above 20, 25 and 30, respectively. In contrast, significant $^{87}Sr/^{86}Sr$ shifts

544 occur at salinities close to 32–35 in the studied lagoons (Fig. 2A). From a paleoenvironmental
545 point of view, these results call to be very careful in any paleosalinity deduction from coastal
546 fossil $^{87}Sr/^{86}Sr$ data, especially since saline SGD are difficult to detect from biosedimentary
547 facies or sismic data compared with freshwater SGD (Lecher and Mackey, 2018; Goff, 2019).
548
549 *5.3. Influence of water mass restriction*
550 From small lagoons to larger epeiric areas like Mediterranean and Baltic seas, oceanic
551 mixing is essential to buffer regional SGD and river effects and maintain homogeneous
552 seawater $^{87}Sr/^{86}Sr$ patterns typical of OOW. However, water mass restriction imposed by
553 geographic barriers and water column stratification may be highly variable from one domain
554 to another one. In lagoons, the local OOW renewal is more or less perennial and important
555 depending on their leaky, choked or restricted geomorphology (Kjerfve and Magill, 1989;
556 Umgiesser et al., 2014). Basically, restricted and leaky lagoons are parallel to the coast and
557 present, respectively, few to numerous inlets that allow an efficient oceanic water turnover
558 (e.g. every one to a few days) driven by tidal dynamics and/or wind forcing (Kjerfve and
559 Magill, 1989; Umgiesser et al., 2014). These characteristics contribute therefore to
560 homogenize seawater $^{87}Sr/^{86}Sr$ ratio close to the OOW one like in the Salses-Leucate or
561 Venice lagoons (Beck et al. 2013) (Fig. 1). To the extreme, choked lagoons are characterized
562 by only one or two oceanic inlets with a dominant wind forcing implying several days to
563 several months for a total water renewal in the lagoon (Kjerfve and Magill, 1989; Umgiesser
564 et al., 2014). This situation applies to the Oualidia lagoon, with the exception that tides are an
565 important factor in the local hydrodynamics (Hilmi et al. 2017). Based on numerical
566 modelling of tidal currents at Oualidia, the time necessary for full lagoon water renewal was
567 estimated to be one or a few days in close vicinity of oceanic inlet increasing up to 15 to 30
568 days in the upstream channel and salt marshes, respectively (Hilmi et al., 2005, 2009, 2017).

As a result, the duality and rhythmicity between SGD discharges and seawater tidal influxes explain the $^{87}Sr/^{86}Sr$ gradient observed in the Oualidia lagoon from the oceanic inlet towards the interior of the lagoon.

In epeiric seas, similar water mass restriction processes appear relevant to explain the variability of seawater $^{87}Sr/^{86}Sr$ ratios. The Mediterranean basin is very interesting in this respect as SGD with elevated Sr concentrations (5 - 12 $\mu mol \cdot l^{-1}$) and low $^{87}Sr/^{86}Sr$ ratios (i.e., 0.7078 - 0.7080) prevail all along the northern Mediterranean margin (Rodellas et al., 2015; Trezzi et al. 2017). However, only the eastern Mediterranean waters are significantly affected by these unradiogenic terrestrial Sr inputs (Fig. 3A). We propose that this spatial heterogeneity can be due to an efficient buffering of SGD in the western Mediterranean areas due to Atlantic OOW inflows from the strait of Gibraltar. Indeed, the renewal time of intermediate and deep waters is estimated to be 2.5 to 5 times faster in the western Mediterranean basin than it is in the eastern basin owing to the Sicily Strait (Tanhua et al., 2013).

By summarizing, we identify significant deviations from the OOW $^{87}Sr/^{86}Sr$ ratio in various marine contexts where the interplay between oceanic and terrestrial Sr inputs is a key parameter. This relative variability attests that, depending on their isolation, shelf domains may be regarded as dynamic and transitory Sr sub-reservoirs not always representative of the global ocean reservoir for which the modern Sr residence time is estimated around 2.5 Myr (Hodell et al., 1990). This spatial decoupling in the Sr budget was suggested by de Villiers (1999) for the upper ocean by estimating partial Sr residence times of only 5800–700,000 years in the upper 400 m of the water column due to a group of celestite ($SrSO_4$)-secreting radiolarians (i.e., Acantharia). In semi-enclosed coastal domains with transient dynamics like lagoons or large epeiric seas, we expect that the partial residence time (*sensu* Lin and Liu, 2019) could be in the same order of magnitude as water renewal time in marginal basins (i.e.,

one day to half a month for studied lagoons up to 30–130 years, 2–40 years, and 5–625 years in the Mediterranean, Baltic and Black Sea basins, respectively) (Lee et al., 2002; Stieglitz et al., 2017; Tanhua et al., 2013; Omstedt et al., 2014; Hilmi et al., 2017).

Finally, it worth noting that the Sr concentrations measured in the studied marine areas are relatively low (i.e., 60 – 70 $\mu mol \cdot l^{-1}$) compared with those of OOW (i.e., 80 – 90 $\mu mol \cdot l^{-1}$) (Fig. 2B). However, these results are in agreement with the large variability of coastal Sr concentrations (i.e., ±20 $\mu mol \cdot l^{-1}$) reported in worldwide coastal contexts (Fig. 5), sometimes with important annual fluctuations and local heterogeneities in euhaline waters (Ladouche et al., 2000; Brunskill et al., 2003; Elsdon and Gillanders, 2006). This suggests that the slightly non-conservative behaviour of Sr in seawater could be locally enhanced in some coastal and hemipelagic domains. Whether perennial or temporary, the cause of Sr depletions in euhaline waters – already reported the Salses-Leucate lagoon (Ladouche et al. 2000) but also sporadically observed along the Panama coast, in the Venice lagoon (Italy) or in the Gulf of Papua (Papua New Guinea) (Brunskill et al. 2003; Beck et al. 2013) - remains obscure. We hypothesize that, in some water mass restriction contexts, three factors could be involved and combined: 1) binding of Sr at the surface of sedimentary particles depending on grain size, iron and manganese oxides, and organic matter (Takada et al. 2014); 2) fast biological adsorption rates by specific organisms (i.e., see Bowen 1956, de Villiers 1999); or 3) influences of saline SGD (i.e., similar to superficial porewaters) slightly depleted in Sr and with isotopic ratios close to the one of seawater (e.g., Kastner et al. 1990).

## Conclusion

In this study, we analyzed and compared the $^{87}Sr/^{86}Sr$ ratios and the major and trace element concentrations of seawaters and shells from marginal and hemipelagic contexts with different degrees of water mass restriction. Homogeneous $^{87}Sr/^{86}Sr$ ratios typical of OOW (i.e.,

619 0.709172 ± 0.000023) are recorded in the Tatakoto atoll and along the BSM transect. This
620 suggests that, in open shelf context without any river inputs, alternative Sr source like SGD or
621 particulate dissolution are too limited and likely buffered by oceanic Sr inputs. This comforts
622 the use of carbonate fossils from these domains to infer past OOW compositions. However,
623 slight isotopic shifts may be observed in the foreshore area (e.g., BSM beach) that we
624 interpret as resulting from rainwater mixing, local groundwater discharges or particle
625 dissolution in the surf zone. Elevated coastal water retention times could also account for
626 sporadic anomalous $^{87}Sr/^{86}Sr$ ratios in open shelves. In semi-enclosed domains, we report
627 variable Sr isotope patterns between the studied karstic lagoons, with homogeneous seawater
628 $^{87}Sr/^{86}Sr$ ratios typical of OOW (i.e., 0.709155 to 0.709170) in the Salses-Leucate lagoon
629 (except close to groundwater discharges where the values drop to 0.708903) and an important
630 $^{87}Sr/^{86}Sr$ gradient from OOW values close to the Atlantic inlet to less radiogenic values of
631 0.707957 in the upstream parts of the Oualidia lagoon. We relate these differences to two
632 main factors. First, we note that, despite lower fluxes, the Oualidia lagoon is supplied by SGD
633 with unradiogenic $^{87}Sr/^{86}Sr$ ratios and very high Sr concentrations (up to 150 $\mu mol \cdot l^{-1}$) linked
634 to the dissolution of Mesozoic evaporites. These inputs lead to maximal contributions of 60%
635 to the local Sr budget. The second parameter concerns the leaky, restricted or choked
636 morphology of lagoons which controls the oceanic Sr inputs and thus buffers more or less any
637 terrestrial Sr inputs through water mass homogenization. These two parameters appear also
638 relevant to explain the $^{87}Sr/^{86}Sr$ gradient prevailing from west to east in the Mediterranean
639 basin.
640

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647
648 **Figure Captions**
649
650 Fig. 1: Location and $^{87}Sr/^{86}Sr$ ratios of seawater and groundwater samples collected in the four
651 studied sites: Oualidia lagoon (A), Salses-Leucate lagoon (B), Banyuls-sur-Mer (BSM)
652 transect (C), Tatakoto atoll (D). Dashed lines indicate groundwater discharges.
653
654 Fig. 2: $^{87}Sr/^{86}Sr$, Sr, Ca, and $SO_4$ concentrations of waters as a function of salinity. (A) For a
655 same sampling position with a given salinity, the $^{87}Sr/^{86}Sr$ ratios of living mollusc shells are
656 compared to $^{87}Sr/^{86}Sr$ ratios of ambient seawater as well as oysters sampled in 2006 by El
657 Meknassi et al. (2018) (see Table 2 for samples correspondence). Additional literature data
658 from Salses-Leucate (i.e., water sampled in 1998, 2002 and 2012) are represented by squares
659 with white contours (Ladouche et al. 2000; Ladouche and Dörfliger 2004; Petelet-Giraud et al.
660 2016). The isotopic range and the average elemental concentrations of OOW are from El
661 Meknassi et al. (2018) and Millero et al., (2008), respectively. The salinity threshold defining
662 significant influence of river water (RW) inputs on seawater $^{87}Sr/^{86}Sr$ ratios is from Bryant et
663 al. (1995). Seawater and SGD values from the Oualidia and Salses-Leucates lagoons are
664 compared to groundwater data available at the scale of Moroccan aquifers (Fakir et al., 2002;
665 Kaid Rassou et al., 2005; Vinson et al. 2013; Fadili et al., 2015; Bouchaou et al., 2017) and
666 the Corbières aquifer (southern France) (Petelet et al., 1998; Aquilina et al., 2002; Khaska et
667 al., 2013; Petelet-Giraud et al., 2016). The analytical uncertainty bars are smaller than the
668 width of symbols.

669

670 Fig. 3: (A) Worldwide distribution of published and new $^{87}Sr/^{86}Sr$ ratios for euhaline to
671 brackish shelf waters and oligotrophic oceanic waters (OOW). If locally variable, the
672 maximal and minimal $^{87}Sr/^{86}Sr$ ratios are represented by two colours. Box-and-whisker plot
673 corresponds to the OOW $^{87}Sr/^{86}Sr$ variability range published by El Meknassi et al. (2018).
674 Literature data from Müller et al. (1990a, 1990b), Andersson et al. (1992), Ingram and Sloan
675 (1992), Winter et al. (1997), Barker et al. (1998), Israelson and Buchardt (1999), Clauer et al.
676 (2000), Ladouche et al. (2000), Jørgensen and Banoeng-Yakubo (2001), Peckmann et al.
677 (2001), Wang et al. (2001), Négrel et al. (2005), Major et al. (2006), Huang and You (2007),
678 Sharma et al. (2017), Xu and Marcantonio (2017), Jørgensen et al. (2008), Martin and moore
679 (2008), Huang et al. (2011), Patra et al. (2012), Rahaman and Singh (2012), Beck et al. (2013),
680 Uddin et al. (2013), Wang and You (2013), Jones et al. (2014), Mokadem et al. (2015), Pearce
681 et al. (2015), Eissa et al. (2016), Petelet-Giraud et al. (2016), Trezzi et al. (2017), Shao et al.
682 (2018), Teichert et al. (2018), Casse et al. (2019), Danish et al. (2020).

683

684 Fig. 4: $^{87}Sr/^{86}Sr$ vs. 1/Sr diagram for the Salses-Leucate (A) and Oualidia (B) lagoon waters.
685 The data are compared to mixing trends between local seawater and different groundwaters
686 (GW) from Moroccan (A) and Corbières aquifers (B). End members data are presented in
687 Table 3. The mixing curves show the $^{87}Sr/^{86}Sr$ ratios for each 10 % increment. The analytical
688 uncertainty bars are smaller than the width of symbols.

689

690 Fig. 5: Sr concentration of shelf (SW) and oligotrophic oceanic (OOW) waters as a function
691 of salinity. The OOW data are from de Villiers (1999) and the SW data are compiled from
692 various sources (see supplementary data). The red and grey lines represent the linear
693 regression and the 95% confidence interval of Sr concentration values, respectively.

694

695 Table 1: $^{87}Sr/^{86}Sr$ and elemental concentrations of surface and bottom seawater from the
696 studied sites.

697

698 Table 2: $^{87}Sr/^{86}Sr$ and elemental concentrations of studied shells.

699

700 Table 3: Average, minimal and maximal $^{87}Sr/^{86}Sr$ and Sr concentrations of seawater (SW) and
701 groundwater (GW) end members used for mixing model of Fig. 4. GW data refer to waters
702 percolating through different aquifers of the Corbières region (South France) and Morocco.
703 As GW $^{87}Sr/^{86}Sr$ and Sr concentration data are missing or sparse, respectively, for the Coastal
704 Sahel of Oualidia, we used GW data from Moroccan aquifers with host rocks of the same age
705 and lithology (see supplementary data).

706

707 **Bibliography**

708 Albarède, F., 1996. Introduction to geochemical modeling. Cambridge University Press.
709 543pp.
710 Andersson, P.S., Wasserburg, G.J., Ingri, J., 1992. The sources and transport of Sr and Nd
711 isotopes in the Baltic Sea. Earth Planet. Sci. Lett. 113, 459–472.
712 https://doi.org/10.1016/0012-821X(92)90124-E
713 Andréfouët, S., Pagès, J., Tartinville, B., 2001. Water renewal time for classification of atoll
714 lagoons in the Tuamotu Archipelago (French Polynesia). Coral Reefs 20, 399–408.
715 https://doi.org/10.1007/s00338-001-0190-9
716 Andréfouët, S., Ardhuin, F., Queffeulou, P., Le Gendre, R., 2012. Island shadow effects and
717 the wave climate of the Western Tuamotu Archipelago (French Polynesia) inferred from
718 altimetry and numerical model data. Mar. Pollut. Bull. 65, 415–424.

https://doi.org/10.1016/j.marpolbul.2012.05.042

Aquilina, L., Ladouche, B., Doerfliger, N., Seidel, J.L., Bakalowicz, M., Dupuy, C., Le Strat, P., 2002. Origin, evolution and residence time of saline thermal fluids (Balaruc springs, southern France): Implications for fluid transfer across the continental shelf. Chem. Geol. 192, 1–21. https://doi.org/10.1016/S0009-2541(02)00160-2

Arnaud, P., Raimbault, R., 1969. The Salses-Leucate Pond. Its principal physicochemical characteristics and their variations (in 1955-1956 and from 1960-1968). Rev. Trav. Inst. Peches. Marit. 33(4), 335-443.

Aunay, B., Strat, P. le, Duvail, C., 2003. Méthode d'analyse géologique sur la karstification des Corbières orientales et influence des évènements néogènes (tortono-messiniens). Hydrol. Mediterr. semiarid Reg. 278, 124–129.

Ballouche, A., Carruesco, C., 1986. Evolution holocène d'un écosystème lagunaire : la lagune de Oualidia (Maroc atlantique). Rev. géologie Dyn. géographie Phys. 27, 113–118.

Bach, W., & Humphris, S. E., 1999. Relationship between the Sr and O isotope compositions of hydrothermal fluids and the spreading and magma-supply rates at oceanic spreading centers. Geology, 27(12), 1067-1070. https:// doi:10.1130/0091-7613(1999)027<1067:rbtsao>2.3.co;2

Barker, A.P., Newton, R.J., Bottrell, S.H. 1998. Processes affecting groundwater chemistry in a zone of saline intrusion into an urban sandstone aquifer. Applied Geochemistry, 13,6, 735-749. https://doi.org/10.1016/S0883-2927(98)00006-7

Basu, A.R., Jacobsen, S.B., Poreda, R.J., Dowling, C.B., Aggarwal, P.K., 2001. Large groundwater strontium flux to the oceans from the bengal basin and the marine strontium isotope record. Science (5534). 293, 1470–1473. https://doi.org/10.1126/science.1060524

Bataille, C.P., Willis, A., Yang, X., Liu, X., 2017. Continental igneous rock composition: A

major control of past global chemical weathering continental crust to seawater and exerts a direct control on several bio- surfaces have been invoked to explain some more specific features in screened database gath. Sci. Adv. 3, 1–16. https://doi.org/10.1126/sciadv.1602183

Beck, A.J., Charette, M.A., Cochran, J.K., Gonneea, M.E., Peucker-Ehrenbrink, B., 2013. Dissolved strontium in the subterranean estuary – Implications for the marine strontium isotope budget. Geochim. Cosmochim. Acta 117, 33–52. https://doi.org/10.1016/J.GCA.2013.03.021

Bejannin, S., van Beek, P., Stieglitz, T., Souhaut, M., Tamborski, J., 2017. Combining airborne thermal infrared images and radium isotopes to study submarine groundwater discharge along the French Mediterranean coastline. J. Hydrol. Reg. Stud. 13, 72–90. https://doi.org/10.1016/j.ejrh.2017.08.001

Bellefroid, E.J., Planavsky, N.J., Miller, N.R., Brand, U., Wang, C., 2018. Case studies on the utility of sequential carbonate leaching for radiogenic strontium isotope analysis. Chem. Geol. 497, 88–99. https://doi.org/10.1016/j.chemgeo.2018.08.025

Bouchaou, L., Warner, N.R., Tagma, T., Hssaisoune, M., Vengosh, A., 2017. The origin of geothermal waters in Morocco: Multiple isotope tracers for delineating sources of water-rock interactions. Appl. Geochemistry 84, 244–253. https://doi.org/10.1016/j.apgeochem.2017.07.004

Bowen, H.J.M. 1956. Strontium and barium in sea water and marine organisms. J. Mar. Biol. Ass. U.K., 35, 451-460.

Brass, G., 1976. The variations of the marine $^{87}Sr/^{86}Sr$ ratio during Phanerozoic time: interpretation using a flux model. Geochimica et Cosmochimica Acta 40(7), 721-730. https://doi.org/10.1016/0016-7037(76)90025-9

Brems, D., Ganio, M., Latruwe, K., Balcaen, L., Carremans, M., Gimeno, D., Silvestri, A.,

769 Vanhaecke, F., Muchez, P., Degryse, P., 2013. Isotopes on the beach, Part 1: Strontium
770 isotope ratios as a provenance indicator for lime raw materials used in roman glass-
771 making. Archeometry, 55(2), 214-234. doi: 10.1111/j.1475-4754.2012.00702.x
772 Briard, J., Puceat, E., Vennin, E., Daeron, M., Chavagnac, V., Jaillet, R., Merle, D. And De
773 Rafelis, M., 2020, Seawater paleotemperature and paleosalinity evolution in neritic
774 environments of the Mediterranean margin: insights from isotope analysis of bivalve
775 shells. Paleaogeo. Palaeoclim. Palaeoeco, 543, DOI : 10.1016/J.palaeo.2019.109582
776 Brunskill, G.J., Zagorskis, I., Pfitzner, J., 2003. Geochemical mass balance for lithium, boron,
777 and strontium in the Gulf of Papua, new Guinea (Project TROPICS), Geochemica et
778 Cosmochimica Acta, 67 (18), 3365-3383.
779 Bryant, J.D., Jones, D.S., Mueller, A.A., 1995. Influence of freshwater flux on 87Sr/86Sr
780 chronostratigraphy in marginal environments and dating of vertebrate and invertebrate
781 faunas. Journal of Paleontology, 69(1), 1-6.
782 Burke, W.H., Denison, R.E., Hetherington, E.A., Koepnick, R.B., Nelson, H.F., Otto, J.B.,
783 1982. Variation of seawater $^{87}Sr/^{86}Sr$ throughout Phanerozoic time. Geology 10, 516–519.
784 https://doi.org/10.1130/0091-7613(1982)10<516:VOSSTP>2.0.CO;2
785 Burnett, W.C., Aggarwal, P.K., Aureli, A., Bokuniewicz, H., Cable, J.E., Charrette, M.A.,
786 Kontar, E., Krupa, S., Kulkarni, K.M., Loveless, A., Moore, W.S., Oberdorfer, J.A.,
787 Oliveira, J., Ozyurt, N., Povinec, P., Privitera, A.M.G., Rajar, R., Ramessur, R.T.,
788 Scholten, J., Stieglitz, T., Taniguchi, M., Turner, J.V. 2006. Quantifying submarine
789 groundwater discharge in the coastal zone via multiple methods. Science of the Total
790 Environement, 367 (2-3), 498-543, doi: 10.1016/j.scitotenv.2006.05.009
791 Casse, M., Montero-Serrano, J.C., St-Onge, G., Poirier, A., 2019. REE distribution and Nd
792 isotope composition of estuarine waters and bulk sediment leachates tracing lithogenic
793 inputs in eastern Canada. Mar. Chem. 211, 117-130.

https://doi.org/10.1016/j.marchem.2019.03.012

Chakrabarti, R., Mondal, S., Acharya, S.S., Lekha, J.S., Sengupta, D., 2018. Submarine groundwater discharge derived strontium from the Bengal Basin traced in Bay of Bengal water samples. Sci. Rep. 8, 1–10. https://doi.org/10.1038/s41598-018-22299-5

Chavagnac V., German C.R., Milton J. A., Palmer M.R., 2005. Source of REE in sediments cores from the Rainbow vent site (36°14'N, MAR), Chem. Geol., 216 (3-4), 329-352. https://doi.org/10.1016/j.chemgeo.2004.11.015

Chavagnac, V., Leleu, T., Fontaine, F., Cannat, M., Ceuleneer, G., Castillo, A., 2018. Spatial Variations in Vent Chemistry at the Lucky Strike Hydrothermal Field, Mid-Atlantic Ridge (37°N): Updates for Subseafloor Flow Geometry From the Newly Discovered Capelinhos Vent. Geochemistry, Geophys. Geosystems 19, 4444–4458. https://doi.org/10.1029/2018GC007765

Clanzig, S., 1987. Inventaire des invertébrés d'une lagune méditerranéenne des côtes de France, biocénoses et confinement: l'étang de Salses-Leucate (Roussillon). PhD Thesis, Ecole Pratique des Hautes Etudes, Montpellier, France.

Clauer, N., Chaudhuri, S., Toulkeridis, T., Blanc, G., 2000. Fluctuations of Caspian Sea level: Beyond climatic variations? Geology 28, 1015–1018. https:// doi: 10.1130/0091-7613(2000)28<1015:FOCSLB>2.0.CO;2

Cochran, J.K., Landman, N.H., Turekian, K.K., Michard, A., Schrag, D.P., 2003. Paleoceanography of the Late Cretaceous (Maastrichtian) Western Interior Seaway of North America: Evidence from Sr and O isotopes. Palaeogeogr. Palaeoclimatol. Palaeoecol. 191, 45–64. https://doi.org/10.1016/S0031-0182(02)00642-9

Cotte, L., Waeles, M., Pernet-Coudrier, B., Sarradin, P.-M., Cathalot, C., Riso, R.D., 2015. A comparison of in situ vs. ex situ filtration methods on the assessment of dissolved and particulate metals at hydrothermal vents. Deep-Sea Res. 105, 186–194.

819 doi :10.1016/j.dsr.2015.09.005
820 Danish, M., Tripathy G.R., Rahaman, W. 2020. Submarine groundwater discharge to a
821 tropical costal lagoon (Chilika lagoon, India): An estimation using Sr isotopes. Marine
822 Chemistry, 224, 103816, 1-13. https://doi.org/ 10.1016/j.marchem.2020.103816
823 Davis, A.C., Bickle, M.J., Teagle, D.A.H., 2003. Imbalance in the oceanic strontium budget.
824 Earth Planet. Sci. Lett. 211, 173–187. https://doi.org/10.1016/S0012-821X(03)00191-2
825 de Villiers, S., 1999. Seawater strontium and Sr/Ca variability in the Atlantic and Pacific
826 oceans. Earth Planet. Sci. Lett. 171, 623–634.
827 DeVries, T., Primeau, F., 2011. Dynamically and Observationally Constrained Estimates of
828 Water-Mass Distributions and Ages in the Global Ocean. J. Phys. Oceanogr. 41, 2381–
829 2401. https://doi.org/10.1175/jpo-d-10-05011.1
830 Dürr, H.H., Laruelle, G.G., van Kempen, C.M., Slomp, C.P., Meybeck, M., Middelkoop, H.,
831 2011. Worlwide typology of nearshore coastal systems: Defining the estuarine filter of
832 river inputs to the oceans. Estuaries and Coasts 34, 441-458.
833 https://doi.org/10.1007/s12237-011-9381-y
834 Durrieu De Madron, X., Houpert, L., Puig, P., Sanchez-Vidal, A., Testor, P., Bosse, A.,
835 Estournel, C., Somot, S., Bourrin, F., Bouin, M.N., Beauverger, M., Beguery, L., Calafat,
836 A., Canals, M., Cassou, C., Coppola, L., Dausse, D., D'Ortenzio, F., Font, J., Heussner,
837 S., Kunesch, S., Lefevre, D., Le Goff, H., Martín, J., Mortier, L., Palanques, A.,
838 Raimbault, P., 2013. Interaction of dense shelf water cascading and open-sea convection
839 in the northwestern Mediterranean during winter 2012. Geophys. Res. Lett. 40, 1379–
840 1385. https://doi.org/10.1002/grl.50331
841 Eidvin, T., Vinzenz, C., Dybkjær, K., Skovbjerg, E., Piasecki, S., 2014. Discrepancy between
842 Sr isotope and biostratigraphic datings of the upper middle and upper Miocene
843 successions (Eastern North Sea). Palaeogeogr. Palaeoclimatol. Palaeoecol. 411, 267–280.

844 https://doi.org/10.1016/j.palaeo.2014.07.005
845 Eissa, M.A., Thomas, J.M., Pohll, G., Shouakar-Stash, O., Hershey, R.L., Dawoud, M., 2016.
846 Groundwater recharge and salinization in the arid coastal plain aquifer of the Wadi Watir
847 delta, Sinai, Egypt. Appl. Geochemistry 71, 48–62.
848 https://doi.org/10.1016/j.apgeochem.2016.05.017
849 Elsdon, T.S., Gillanders, B.M., 2006. Temporal variability in strontium, calcium, barium, and
850 manganese in estuaries: Implications for reconstructing environmental histories of fish
851 from chemicals in calcified structures. Estuarine, Coastal and Shelf Science, 66, 147-
852 156.
853 El Khalidi, K., Zourarah, B., Aajjane, A., 2011. Evolution récente de la morphologie de delta
854 de flot et son effet sur la dynamique hydro-sédimentaire de la lagune de Oualidia (Côte
855 Atlantique, Maroc): Approche par photographie aérienne. Etudos do Quat. APEQ, Braga,
856 7, 73–78.
857 El Meknassi, S. El, Dera, G., Cardone, T., Rafélis, M. De, Brahmi, C., Chavagnac, V., 2018.
858 Sr isotope ratios of modern carbonate shells : Good and bad news for chemostratigraphy
859 46 (11), 1003–1006. https://doi.org/10.1130/G45380.1
860 Fabre, S., Jeandel, C., Zambardi, T., Roustan, M., Almar, R., 2019. An overlooked silica
861 source of the modern oceans: Are sandy beaches the key? Front. Earth Sci., 7, 231. doi:
862 10.3389/feart.2019.00231
863 Fadili, A., Mehdi, K., Riss, J., Najib, S., Makan, A., Boutayab, K., 2015. Evaluation of
864 groundwater mineralization processes and seawater intrusion extension in the coastal
865 aquifer of Oualidia, Morocco: hydrochemical and geophysical approach. Arab. J. Geosci.
866 8, 8567–8582. https://doi.org/10.1007/s12517-015-1808-5
867 Fadili, A., Najib, S., Mehdi, K., Riss, J., Makan, A., Boutayeb, K., Guessir, H., 2016.
868 Hydrochemical features and mineralization processes in coastal groundwater of Oualidia,

869 Morocco. J. African Earth Sci. 116, 233–247.
870 https://doi.org/10.1016/J.JAFREARSCI.2016.01.014
871 Fakir, Y., El Mernissi, M., Kreuser, T., Berjami, B., 2002. Natural tracer approach to
872 characterize groundwater in the coastal Sahel of Oualidia (Morocco). Environ. Geol. 43,
873 197–202. https://doi.org/10.1007/s00254-002-0644-6
874 Fakir, Y., Claude, C., El Himer, H., 2019. Identifying groundwater discharge to an Atlantic
875 coastal lagoon (Oualidia, Central Morocco) by means of salinity and radium mass
876 balances. Environmental Earth Sciences. 78, 626. https://doi.org/10.1007/s12665-019-
877 8637-x
878 Gilbert, A., Andréfouët, S., Yan, L., Remoissenet, G., 2006. The giant clam *Tridacna maxima*
879 communities of three French Polynesia islands: comparison of their population sizes and
880 structures at early stages of their exploitation. ICES J. Mar. Sci. 63, 1573–1589.
881 https://doi.org/10.1016/j.icesjms.2006.07.001
882 Goddéris, Y., Le Hir, G., Macouin, M., Donnadieu, Y., Hubert-Théou, L., Dera, G., Aretz, M.,
883 Fluteau, F., Li, Z.X., Halverson, G.P., 2017. Paleogeographic forcing of the strontium
884 isotopic cycle in the Neoproterozoic. Gondwana Res. 42, 151–162.
885 https://doi.org/10.1016/J.GR.2016.09.013
886 Goff, J.A., 2019. Modern and fossil pockmarks in the New England mud patch: Implications
887 for submarine groundwater discharge on the middle shelf. Gephysical Research Letters,
888 46 (21), https://doi.org/10.1029/2019GL084881
889 Goldstein, S.J., Jacobsen, S.B., 1987. The Nd and Sr isotopic systematics of river-water
890 dissolved material - Implications for the source of Nd and Sr in seawater. Chemical
891 geology, 66(3-4), 245-272. https://doi.org/10.1016/0168-9622(87)90045-5
892 Got, H., Stanley, D.J., 1974. Sedimentation in two Catalonian canyons, northwestern
893 Mediterranean. Mar. Geol. 16(5), M91-M100. https://doi.org/10.1016/0025-

894 3227(74)90067-X
895 Guy, C., Schott, J., Destrigneville, C., Chiappini, R., 1992. Low-temperature alteration of
896 basalt by interstitial seawater, Mururoa, French Polynesia. Geochim. Cosmochim. Acta
897 56, 4169–4189. https://doi.org/10.1016/0016-7037(92)90259-L
898 Heussner, S., Durrieu de Madron, X., Calafat, A., Canals, M., Carbonne, J., Delsaut, N.,
899 Saragoni, G., 2006. Spatial and temporal variability of download particle fluxes on a
900 continental slope: Lessons from an 8-yr experiment in the Gulf of Lions (NW
901 Mediterranean). Mar. Geol. 234, 63–92.
902 Hilmi, K., Koutitonsky, V.G., Orbi, A., Lakhdar, J.I., Chagdali, M., 2005. Oualidia lagoon,
903 Morocco: an estuary without a river. African J. Aquat. Sci. 30, 1–10.
904 https://doi.org/10.2989/16085910509503828
905 Hilmi, K., Orbi, A., Lakhdar Idrissi, J., 2009. Hydrodynamisme de la lagune de Oualidia
906 ( Maroc ) durant l 'été et l 'automne 2005. Bull. l'Institut Sci. Rabat, Sect. Sci. la Terre,
907 31, 29–34.
908 Hilmi, K., Makaoui, A., Ettahiri, O., Idrissi, M., Larissi, J., Abdellaoui, B., El Ouehabi, Z.,
909 Obri, A., 2017. Fonctionnement Hydrodynamique De La Lagune De Oualidia (Maroc)
910 Avant L'Amenagement De La Souille. Int. J. Adv. Res. 5, 2015–2027.
911 https://doi.org/10.21474/ijar01/4937
912 Hodell, D.A., Mead, G.A., Mueller, P.A., 1990. Variation in the strontium isotopic
913 composition of seawater (8 Ma to present) : Implications for chemical weathering rates
914 and dissolved fluxes to the oceans. Chem. Geol. Isot. Geosci. Sect. 80, 291–307.
915 https://doi.org/10.1016/0168-9622(90)90011-Z
916 Huang, K.F., You, C.F., 2007. Tracing freshwater plume migration in the estuary after a
917 typhoon event using Sr isotopic ratios. Geophys. Res. Lett. 34, 1–5.
918 https://doi.org/10.1029/2006GL028253

Huang, K.-F., You, C.-F., Chung, C.-H., Lin, I.-T., 2011. Nonhomogeneous seawater Sr isotopic composition in the coastal oceans: A novel tool for tracing water masses and submarine groundwater discharge. Geochemistry, Geophys. Geosystems 12, 1–14. https://doi.org/10.1029/2010GC003372

Ingram, B.L., Sloan, D., 1992. Strontium isotopic composition of estuarine sediments as paleosalinity-paleoclimate indicator. Science (5040). 255, 68-72. https://doi: 10.1126/science.255.5040.68

Israelson, C., Buchardt, B., 1999. Strontium and oxygen isotopic composition of East Greenland rivers and surface waters: Implication for palaeoenvironmental interpretation. Palaeogeogr. Palaeoclimatol. Palaeoecol. 153, 93–104. https://doi.org/10.1016/S0031-0182(99)00068-1

Jeandel, C., 2016. Overview of the mechanisms that could explain the "Boundary Exchange" at the land-ocean contact. Philos. Trans. R. Soc. A Math. Phys. Eng. Sci. 374, 1–13. https://doi.org/10.1098/rsta.2015.0287

Jeandel, C., Oelkers, E.H., 2015. The in fl uence of terrigenous particulate material dissolution on ocean chemistry and global element cycles. Chem. Geol. 395, 50–66. https://doi.org/10.1016/j.chemgeo.2014.12.001

Jones, M.T., Gislason, S.R., Burton, K.W., Pearce, C.R., Mavromatis, V., Pogge von Strandmann, P.A.E., Oelkers, E.H., 2014. Quantifying the impact of riverine particulate dissolution in seawater on ocean chemistry. Earth Planet. Sci. Lett. 395, 91–100. https://doi.org/10.1016/j.epsl.2014.03.039

Jørgensen, N.O., Andersen, M.S., Engesgaard, P., 2008. Investigation of a dynamic seawater intrusion event using strontium isotopes ($^{87}Sr/^{86}Sr$). J. Hydrol. 348, 257–269. https://doi.org/10.1016/j.jhydrol.2007.10.001

Jørgensen, N.O., Banoeng-Yakubo, B.K., 2001. Environmental isotopes ($^{18}O$, $^{2}H$, and

$^{87}Sr/^{86}Sr$) as a tool in groundwater investigations in the Keta Basin, Ghana. Hydrogeol. J. 9, 190–201. https://doi.org/10.1007/s100400000122

Kaid Rassou, K., Fakir, Y., Bahir, M., Zouari, K., Marah, M., Monteiro, J.P., 2005. Apport des analyses isotopiques à la compréhension du fonctionnement des aquifères cotiers du bassin hydrologiques de la lagune d'Oualidia (Ocean Atlantique marocain). Comuniceçoes Geol., 92, 129-142.

Kalnejais, L.H., Martin, W.R., Bothner, M.H., 2010. The release of dissolved nutrients and metals from coastal sediments due to resuspension. Mar. Chem. 121, 224–235. https://doi.org/10.1016/j.marchem.2010.05.002

Kastner, M., Elderfield, H., Martin, J.B., Suess, E., Kvenvolden, K.A., Garrison, R.E., 1990. Diagenesis and interstitial-water chemistry at the peruvian continental margin - Major constituents and strontium isotopes. Proceedings of the Ocean Drilling Program, Scientific Results, 112, 413-440.

Khaska, M., Le Gal La Salle, C., Lancelot, J., ASTER, T., Mohamad, A., Verdoux, P., Noret, A., Simler, R., 2013. Origin of groundwater salinity (current seawater vs. saline deep water) in a coastal karst aquifer based on Sr and Cl isotopes. Case study of the La Clape massif (southern France). Appl. Geochemistry 37, 212–227. https://doi.org/10.1016/j.apgeochem.2013.07.006

Kjerfve, B., Magill, K.., 1989. Geographic and hydrodynamic characteristics of shallow coastal lagoons. Mar. Geol. 88, 187–199. https://doi.org/10.1016/0025-3227(89)90097-2

Krabbenhöft, A., Eisenhauer, A., Böhm, F., Vollstaedt, H., Fietzke, J., Liebetrau, V., Augustin, N., Peucker-Ehrenbrink, B., Müller, M.N., Horn, C., Hansen, B.T., Nolte, N., Wallmann, K., 2010. Constraining the marine strontium budget with natural strontium isotope fractionations ($^{87}Sr/^{86}Sr^*$, $\delta^{88}/^{86}Sr$) of carbonates, hydrothermal solutions and river waters. Geochim. Cosmochim. Acta 74, 4097–4109.

https://doi.org/10.1016/j.gca.2010.04.009

Ladagnous, H., Le Bec, C., 1997. La lagune de Salses-Leucate: I- Analyse bibliographique. Direction de l'environnement et de l'aménagement du littoral. R.INT.DEL/97.02/SETE. https://archimer.ifremer.fr/doc/00073/18422/

Ladouche, B., Le Bec, C., Aquilina, L., Bakalowicz, M., Souchu, P., Doerfliger, N., Anus, S. (2000) Recherche de l'origine de la contamination bactériologique de l'étang de Salses-Leucate. Rap. BRGM/RP-50003-FR, 64p

Ladouche, B., Dörfliger, N., 2004. Evaluation des ressources en eau des Corbières. Phase I - Synthèse de la caractérisation des systèmes karstiques des Corbières orientales. Rapport final. Volume 2 - caractérisation géologique et hydrogéologique du système karstique du "synclinal du bas-Agly". BRGM/RP-52919-FR, 196p.

Lecher, A.L., Mackey, K.R.M. 2018. Synthesizing the effects of submarine groundwater discharge on marine biota. Hydrology, 5, 60, doi:10.3390/hydrology5040060

Lee, B.S., Bullister, J.L., Murray, J.W., Sonnerup, R.E., 2002. Anthropogenic chlorofluorocarbons in the Black Sea and the Sea of Marmara. Deep. Res. Part I Oceanogr. Res. Pap. 49, 895–913. https://doi.org/10.1016/S0967-0637(02)00005-5

Li, Y-H., 1982. A brief discussion on the mean oceanic residence time of elements. Geochim. Cosmochim. Acta 46, 2671–2675.

Lin, L., Liu, Z., 2019. Partial residence times: determining residence time compositions in different subregions. Ocean Dynamics, 69, 1023-1036.

Liu, X., Dunne, J.P., Stock, C.A., Harrison, M.J., Adcroft, A., Resplandy, L., 2019. Simulating water residence time in the coastal ocean: A global perspective. Geophysical Research Letters, 46, 13,910–13,919. https://doi.org/10.1029/2019GL085097

Major, C.O., Goldstein, S.L., Ryan, W.B.F., Lericolais, G., Piotrowski, A.M., Hajdas, I., 2006. The co-evolution of Black Sea level and composition through the last deglaciation and its

994 paleoclimatic significance. Quat. Sci. Rev. 25, 2031–2047.
995 https://doi.org/10.1016/j.quascirev.2006.01.032
996 Marcano, M.C., Frank, T.D., Mukasa, S.B., Lohmann, K.C., Taviani, M., 2015. Diagenetic
997 incorporation of Sr into aragonitic bivalve shells: implications for chronostratigraphic
998 and palaeoenvironmental interpretations. Depos. Rec. 1, 38–52.
999 https://doi.org/10.1002/dep2.3
1000 Martin, J.B., Moore, P.J. 2008. Sr concentrations and isotope ratios as tracers of ground-water
1001 circulation in carbonate platforms: Examples from San Salvador and Long Island,
1002 Bahamas. Chemical Geology, 249, 52-65. https://doi.org/
1003 10.1016/j.chemgeo.2007.11.009
1004 Martin, E.E., Scher, H.D., 2004. Preservation of seawater Sr and Nd isotopes in fossil fish
1005 teeth: bad news and good news. Earth Planet. Sci. Lett. 220, 25–39.
1006 https://doi.org/10.1016/S0012-821X(04)00030-5
1007 McArthur, J.M., 1994. Recent trends in strontium isotope stratigraphy. Terra Nov. 6, 331–358.
1008 https://doi.org/10.1111/j.1365-3121.1994.tb00507.x
1009 McArthur, J.M., Howarth, R.J., Shields, G.A., 2012. Strontium Isotope Stratigraphy, in: The
1010 Geologic Time Scale. Elsevier, pp. 127–144. https://doi.org/10.1016/B978-0-444-59425-
1011 9.00007-X
1012 Millero, F.J., Feistel, R., Wright, D.G., McDougall, T.J., 2008. The composition of Standard
1013 Seawater and the definition of the Reference-Composition Salinity Scale. Deep Sea Res.
1014 Part I Oceanogr. Res. Pap. 55, 50–72. https://doi.org/10.1016/J.DSR.2007.10.001
1015 Mokadem, F., Parkinson, I.J., Hathorne, E.C., Anand, P., Allen, J.T., Burton, K.W., 2015.
1016 High-precision radiogenic strontium isotope measurements of the modern and glacial
1017 ocean: Limits on glacial–interglacial variations in continental weathering. Earth Planet.
1018 Sci. Lett. 415, 111–120. https://doi.org/10.1016/J.EPSL.2015.01.036

Müller, D.W., Mueller, P.A., McKenzie, J.A., 1990a. Strontium Isotopic Ratios as Fluid Tracers in Messinian Evaporites of the Tyrrhenian Sea (Western Mediterranean Sea). Proc. Ocean Drill. Program, 107 Sci. Results 107, 603–614. https://doi.org/10.2973/odp.proc.sr.107.194.1990

Müller, D.W., McKenzie, J.A., Mueller, P.A., 1990b. Abu Dhabi sabkha, Persian Gulf, revisited: Application of strontium isotopes to test an early dolomitization model. Geology 18, 618–621. https://doi.org/10.1130/0091-7613(1990)018<0618:ADSPGR>2.3.CO;2

Négrel, P., Casanova, J., Blomqvist, R., Kaija, J., Frape, S., 2003. Strontium isotopic characterization of the Palmottu hydrosystem (Finland): Water-rock interaction and geochemistry of groundwaters. Geofluids 3, 161–175. https://doi.org/10.1046/j.1468-8123.2003.00056.x

Négrel, P., Casanova, J., Blomqvist, R., 2005. $^{87}Sr/^{86}Sr$ of brines from the Fennoscandian Shield: a synthesis of groundwater isotopic data from the Baltic Sea region. Can. J. Earth Sci. 42, 273–285. https://doi.org/10.1139/e04-103

Neymark, L.A., Premo, W.R., Mel'Nikov, N.N., and Emsbo, P., 2014, Precise determination of δ88Sr in rocks, minerals, and waters by double-spike TIMS: A powerful tool in the study of geological, hydrological and biological processes: Journal of Analytical Atomic Spectrometry, v. 29, p. 65–75, https://doi: 10.1039/c3ja50310k.

Nieto, L.M., Ruiz-Ortiz, P.A., Rey, J., Benito, M.I., 2008. Strontium-isotope stratigraphy as a constraint on the age of condensed levels: examples from the Jurassic of the Subbetic Zone (southern Spain). Sedimentology, 55, 1-28. https://doi: 10.1111/j.1365-3091.2007.00891.x

Omstedt, A., Elken, J., Lehmann, A., Leppäranta, M., Meier, H.E.M., Myrberg, K., Rutgersson, A., 2014. Progress in physical oceanography of the Baltic Sea during the

2003-2014 period. Progress in Oceanography, 128, 139-171. https://doi.org/10.1016/j.pocean.2014.08.010

Ohno T., Hirata T., 2007.Simultaneous determination of mass-dependent isotopic fractionation and radiogenic isotope variation of strontium in geochemical samples by multiple collector-ICP-mass spectrometry. Analytical Science, 23, 1275–1280. https://doi.org/10.2116/analsci.23.1275

Palmer, M.R., Edmond, J.M., 1989. The strontium isotope budget of the modern ocean. Earth Planet. Sci. Lett. 92, 11–26. https://doi.org/10.1016/0012-821X(89)90017-4

Palmer, M., Edmond, J., 1992. Controls over the strontium isotope composition of river water. Geochim. Cosmochim. Acta 56, 2099–2111. https://doi.org/10.1016/0016-7037(92)90332-D

Patra, S., Liu, C.Q., Wang, F.S., Li, S.L., Wang, B.L. 2012. Behaviour of major and minor elements in a temperate river estuary to the coastal area. International Journal of Environmental Science and Technology, 9(4), 655-660. https://doi.org/10.1007/s13762-012-0097-8

Pearce, C.R., Parkinson, I.J., Gaillardet, J., Charlier, B.L.A., Mokadem, F., Burton, K.W., 2015. Reassessing the stable ($\delta^{88}/^{86}Sr$) and radiogenic ($^{87}Sr/^{86}Sr$) strontium isotopic composition of marine inputs. Geochim. Cosmochim. Acta 157, 125–146. https://doi.org/10.1016/J.GCA.2015.02.029

Peckmann, J., Reimer, A., Luth, U., Luth, C., Hansen, B.T., Heinicke, C., Hoefs, J., Reitner, J., 2001. Methane-derived carbonates and authigenic pyrite from the northwestern Black Sea. Mar. Geol. 177, 129–150. https://doi.org/10.1016/S0025-3227(01)00128-1

Petelet, E., Luck, J.-M., Ben Othman, D., Négrel, P., Aquilina, L., Brgm, F., 1998. Geochemistry and Water Dynamics of a Medium-Sized Watershed: the Herault, Southern France. Chem. Geol. 150, 63–83. https://doi.org/10.1016/S0009-

2541(98)00053-9

Petelet-Giraud, E., Négrel, P., Aunay, B., Ladouche, B., Bailly-Comte, V., Guerrot, C., Flehoc, C., Pezard, P., Lofi, J., Dörfliger, N., 2016. Coastal groundwater salinization: Focus on the vertical variability in a multi-layered aquifer through a multi-isotope fingerprinting (Roussillon Basin, France). Sci. Total Environ. 566–567, 398–415. https://doi.org/10.1016/j.scitotenv.2016.05.016

Peterman, Z.E., Hedge, C.E., Tourtelot, H.A., 1970. Isotopic composition of strontium in sea water throughout Phanerozoic time. Geochim. Cosmochim. Acta 34, 105–120. https://doi.org/10.1016/0016-7037(70)90154-7

Peucker-Ehrenbrink, B., Miller, M.W., Arsouze, T., Jeandel, C., 2010. Continental bedrock and riverine fluxes of strontium and neodymium isotopes to the oceans. Geochemistry, Geophys. Geosystems 11, 1–22. https://doi.org/10.1029/2009GC002869

Peucker-Ehrenbrink, B., Fiske, G.J., 2019. A continental perspective of the seawater 87Sr/86Sr record: A review. Chem. Geol. 510, 140–165. https://doi.org/10.1016/J.CHEMGEO.2019.01.017

Pin C., Gannounb A., Dupont A., 2014. Rapid, simultaneous separation of Sr, Pb, and Nd by extraction chromatography prior to isotope ratios determination by TIMS and MC-ICP-MS. Journal Analytical Atomic Spectrometry, 29, 1858-1870. https://doi: 10.1039/C4JA00169A

Pirazzoli, P.A., 1998. Sur la vitesse des variations du niveau de la mer, in: Annales de Géographie, 107 (600), 220–232.

Reeder, S.W., Hitchon, B., Levinson, A.A., 1972. Hydrogeochemistry of the surface waters of the Mackenzie River drainage basin, Canada - I. Factor controlling inorganic composition. Geochimica et Cosmochimica Acta, 36(8), 825-865. https://doi.org/10.1016/0016-7037(72)90053-1

1094 Rodellas, V., Garcia-Orellana, J., Masqué, P., Feldman, M., Weinstein, Y., 2015. Submarine
1095 groundwater discharge as a major source of nutrients to the Mediterranean Sea. Proc.
1096 Natl. Acad. Sci. 112, 3926–3930. https://doi.org/10.1073/pnas.1419049112
1097 Rougerie, F., Ricard, M., Mazaury, E., Des, D., Iagonalres, E., 1984. Dynamique et échanges
1098 lagon-océan modèle de circulation interne à travers le socle corallien. Report CEA-R-
1099 5236.
1100 Rougerie, F., 1995. Nature et fonctionnement des atolls des Tuamotu (Polynésie Française).
1101 Oceanol. Acta 18, 61–78.
1102 Santoni, S., Huneau, F., Garel, E., Aquilina, L., Vergnaud-Ayraud, V., Labasque, T., Celle-
1103 Jeanton, H., 2016. Strontium isotopes as tracers of water-rocks interactions, mixing
1104 processes and residence time indicator of groundwater within the granite-carbonate
1105 coastal aquifer of Bonifacio (Corsica, France). Sci. Total Environ. 573, 233–246.
1106 https://doi.org/10.1016/j.scitotenv.2016.08.087
1107 Schildgen, T.F., Cosentino, D., Frijia, G., Castorina, F., Dudas, F.O., Iadanza, A.,
1108 Sampalmieri, G., Cipollari, P., Caruso, A., Bowring, S.A., M. R. Strecker, 2014. Sea
1109 level and climate forcing of the Sr isotope composition of late Miocene Mediterranean
1110 marine basins. Geochemistry, Geophys. Geosystems 15, 4692–4711.
1111 https://doi.org/10.1002/2014GC005332
1112 Sessa, J.A., Ivany, L.C., Schlossnagle, T.H., Samson, S.D., Schellenberg, S.A., 2012. The
1113 fidelity of oxygen and strontium isotope values from shallow shelf settings : Implications
1114 for temperature and age reconstructions. Palaeogeogr. Palaeoclimatol. Palaeoecol. 342–
1115 343, 27–39. https://doi.org/10.1016/j.palaeo.2012.04.021
1116 Shao, Y., Farkaš, J., Holmden, C., Mosley, L., Kell-Duivestein, I., Izzo, C., Reis-Santos, P.,
1117 Tyler, J., Törber, P., Frýda, J., Taylor, H., Haynes, D., Tibby, J., Gillanders, B.M., 2018.
1118 Calcium and strontium isotope systematics in the lagoon-estuarine environments of

South Australia: Implications for water source mixing, carbonate fluxes and fish migration. Geochim. Cosmochim. Acta 239, 90–108. https://doi.org/10.1016/j.gca.2018.07.036

Sharma, M., Balakrishna, K., Hofmann, A.W., Shankar, R. 2007. The transport of Osmium and Strontium isotopes through a tropical estuary. Geochimica et Cosmochimica Acta, 71, 4856-4867. https://doi.org/10.1016/j.gca.2007.08.004

Stein, M., Starinsky, A., Agnon, A., Katz, A., Raab, M., Spiro, B., Zak, I., 2000. The impact of brine-rock interaction during marine evaporite formation on the isotopic Sr record in the oceans. Evidence from Mt. Sedom, Israel. Geochim. Cosmochim. Acta 64(12), 2039–2053. https://doi.org/10.1016/S0016-7037(00)00370-7

Stieglitz, T.C., van Beek, P., Souhaut, M., Cook, P.G., 2013. Karstic groundwater discharge and seawater recirculation through sediments in shallow coastal Mediterranean lagoons, determined from water, salt and radon budgets. Sci. Rep. 8, 73–84. https://doi.org/10.1016/j.marchem.2013.05.005

Tachikawa, K., Roy-Barman, M., Michard, A., Thouron, D, Yeghicheyan, D., Jeandel, C., 2004. Neodymium isotopes in the Mediterranean Sea: Comparison between seawater and sediment signals. Geochemical et Cosmochimica Acta, 68(14), 3095-3106.

Takada, H., Tagami, K., Aono, T., Uchida, S., 2014. Distribution coefficients (Kd) of strontium and significance of oxides and organic matter in controlling its partitioning in coastal regions of Japan. Science of the Total Environment, 490, 979-986.

Tanhua, T., Hainbucher, D., Schroeder, K., Cardin, V., Álvarez, M., Civitarese, G., 2013. The Mediterranean Sea system: A review and an introduction to the special issue. Ocean Sci. 9, 789–803. https://doi.org/10.5194/os-9-789-2013

Tartinville, B., Deleersnijder, E., Rancher, J., 1997. The water residence time in the Mururoa atoll lagoon: Sensitivity analysis of a three-dimensional model. Coral Reefs 16, 193–203.

https://doi.org/10.1007/s003380050074

Teichert, B.M.A., Chevalier, N., Gussone, N., Bayon, G., Ponzevera, E., Ruffine, L., Strauss, H., 2018. Sulfate-dependent anaerobic oxidation of methane at a highly dynamic bubbling site in the Eastern Sea of Marmara (Çinarcik Basin). Deep. Res. Part II Top. Stud. Oceanogr. 153, 79–91. https://doi.org/10.1016/j.dsr2.2017.11.014

Trezzi, G., Garcia-Orellana, J., Rodellas, V., Masqué, P., Garcia-Solsona, E., Andersson, P.S., 2017. Assessing the role of submarine groundwater discharge as a source of Sr to the Mediterranean Sea. Geochim. Cosmochim. Acta 200, 42–54. https://doi.org/10.1016/J.GCA.2016.12.005

Uddin, S., Al Ghadban, A.N., Behbahani, M., 2013. Baseline concentrations of strontium and $^{90}$Sr in seawater from the northern Gulf. Marine Pollution Bulletin, 75 (1-2), 301-304. https://doi.org/10.1016/j.marpolbul.2013.06.042

Umgiesser, G., Ferrarin, C., Cucco, A., De Pascalis, F., Bellafiore, D., Ghezzo, M., Bajo, M., 2014. Comparative hydrodynamics of 10 Mediterranean lagoons by means of numerical modeling. J. Geophys. Res. Oceans. 119 (4), 2212–2226. https://doi.org/10.1002/2013JC009512

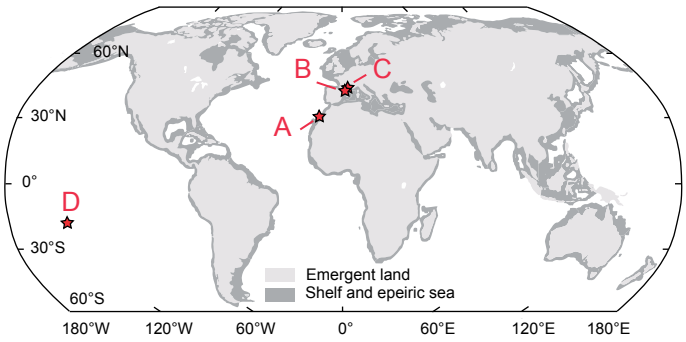
Ünlüata, U., Oguz, T., Latif, M.A., Özsoy, E., 1990. On the physical oceanography of the Turkish straits. In: Pratt, L.J. (ed.) The Physical oceanography of sea straits, NATO/ASI series, pp 25-60. Kluwer Academic Publishers

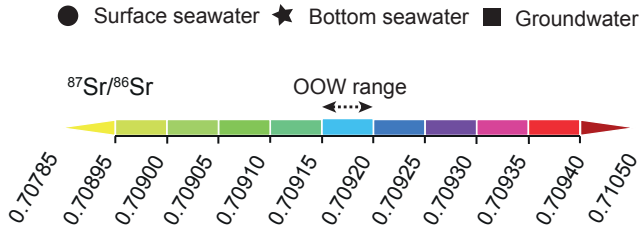
Van Wynsberge, S., Menkes, C., Le Gendre, R., Passfield, T., Andréfouët, S., 2017. Are Sea Surface Temperature satellite measurements reliable proxies of lagoon temperature in the South Pacific? Estuar. Coast. Shelf Sci. 199, 117–124. https://doi.org/10.1016/J.ECSS.2017.09.033

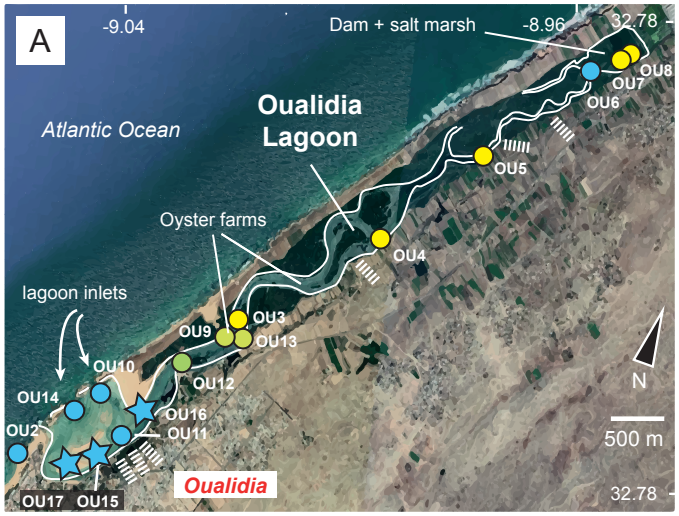
Veizer, J., 1989. Strontium isotopes in seawater through time. Annu. Rev. Earth Planet. Sci. 17, 141–167.

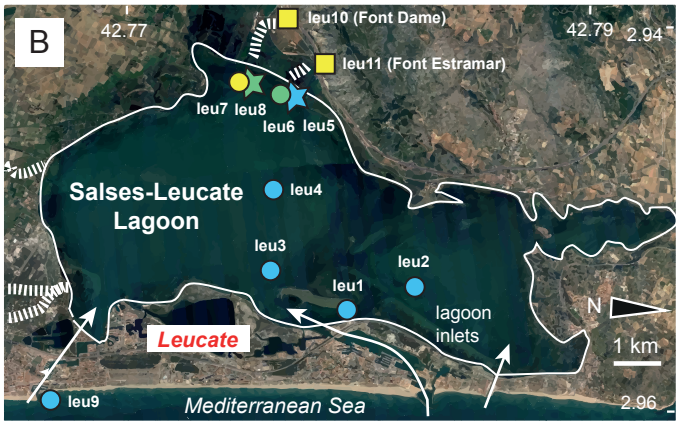
1169 Vinson, D.S., Tagma, T., Bouchaou, L., Dwyer, G.S., Warner, N.R., 2013. Applied
1170 Geochemistry Occurrence and mobilization of radium in fresh to saline coastal
1171 groundwater inferred from geochemical and isotopic tracers. Appl. Geochemistry 38,
1172 161–175. https://doi.org/10.1016/j.apgeochem.2013.09.004
1173 Waljeski, C., 2003. Systematic karst pool erosion within the conglomerate platforms of
1174 Moorea, French Polynesia. Biol. Geomorphol. Trop. Islands 12, 117–128.
1175 Wang, Z., Liu, C., Han, G., Xu, Z. 2001. Strontium isotopic geochemistry of the Changjiang
1176 estuarine waters: Implication for water-sediment interaction. Science in China (Series E),
1177 44, 129-133. https://doi.org/10.1007/BF02916803
1178 Wang, R.M., You, C.F., 2013. Uranium and strontium isotopic evidence for strong submarine
1179 groundwater discharge in an estuary of a mountainous island: A case study in the
1180 Gaoping River Estuary, Southwestern Taiwan. Mar. Chem. 157, 106–116.
1181 https://doi.org/10.1016/j.marchem.2013.09.004
1182 Wierzbowski, H., Anczkiewicz, R., Bazarnik, J., Pawlak, J., 2012. Strontium isotope
1183 variations in Middle Jurassic (Late Bajocian–Callovian) seawater: Implications for
1184 Earth's tectonic activity and marine environments. Chem. Geol. 334, 171–181.
1185 https://doi.org/10.1016/J.CHEMGEO.2012.10.019
1186 Winter, B.L., Johnson, C.M., Clark, D.L., 1997. Strontium, neodymium, and lead isotope
1187 variations of authigenic and silicate sediment components from the Late Cenozoic Arctic
1188 Ocean: Implications for sediment provenance and the source of trace metals in seawater.
1189 Geochim. Cosmochim. Acta 61, 4181–4200. https://doi.org/10.1016/S0016-
1190 7037(97)00215-9
1191 Xu, Y., Marcantonio, F. 2007. Strontium isotope variations in the lower Mississipi River and
1192 its estuarine mixing zone. Marine Chemistry, 105, 118-128.
1193 https://doi.org/10.1016/j.marchem.2007.01.004

1194 Zaky, A.H., Brand, U., Buhl, D., Blamey, N., Bitner, M.A., Logan, A., Gaspard, D., Popov,
1195 A., 2018. Strontium isotope geochemistry of modern and ancient archives: tracer of
1196 secular change in ocean chemistry. Can. J. Earth Sci. 56, 245–264.
1197 https://doi.org/10.1139/cjes-2018-0085
1198


60°N
30°N
0°
30°S
60°S
180°W
120°W
60°W
0°
60°E
120°E
180°E
A
B
C
D
Emergent land
Shelf and epeiric sea


Surface seawater
Bottom seawater
Groundwater
87Sr/86Sr
OOW range
0.70785
0.70895
0.70900
0.70905
0.70910
0.70915
0.70920
0.70925
0.70930
0.70935
0.70940
0.71050


A
-9.04
Dam + salt marsh
-8.96
32.78
Atlantic Ocean
Oualidia Lagoon
OU8
OU7
OU6
OU5
OU4
Oyster farms
lagoon inlets
OU3
OU9
OU13
OU10
OU12
OU14
OU16
OU11
OU2
OU17
OU15
Oualidia
N
500 m
32.78


B
42.77
leu10 (Font Dame)
42.79
2.94
leu11 (Font Estramar)
leu7
leu8
leu6
leu5
Salses-Leucate Lagoon
leu4
leu3
leu2
leu1
lagoon inlets
Leucate
N
1 km
leu9
Mediterranean Sea
2.96


C
3.15
Lacaze-Duthiers Canyon
3.43
42.52
BA4
BA9
BA3
BA6
BA5
BA8
BA7
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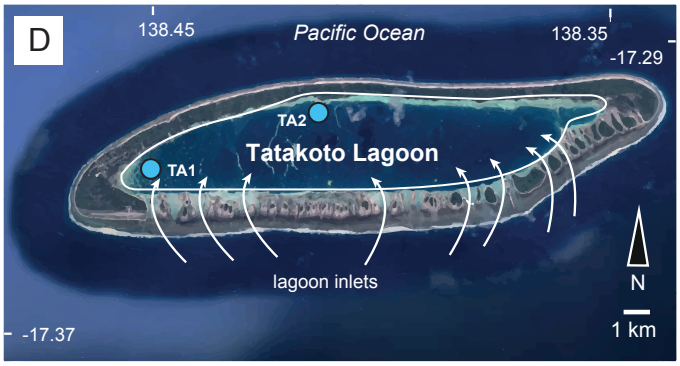
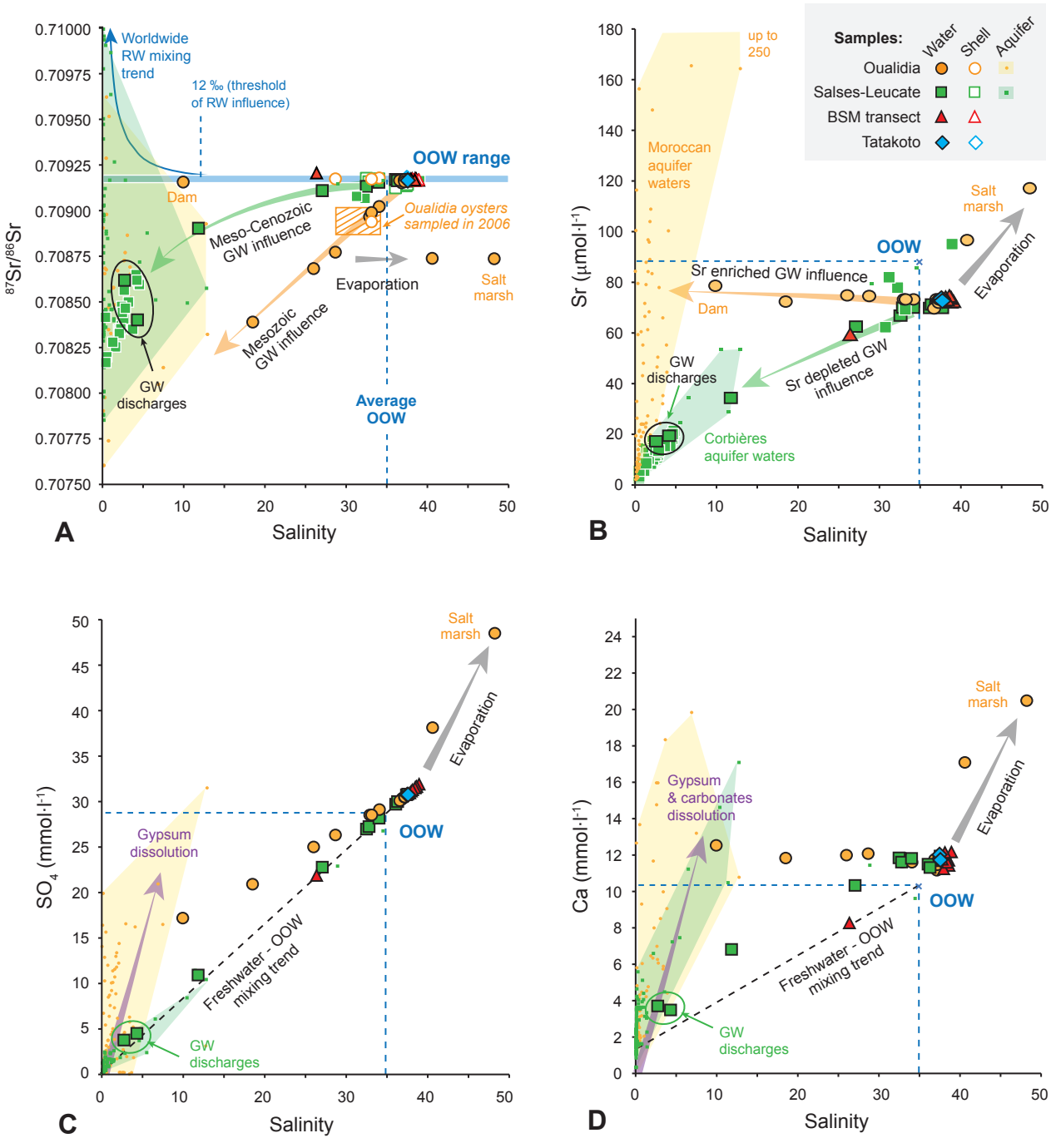
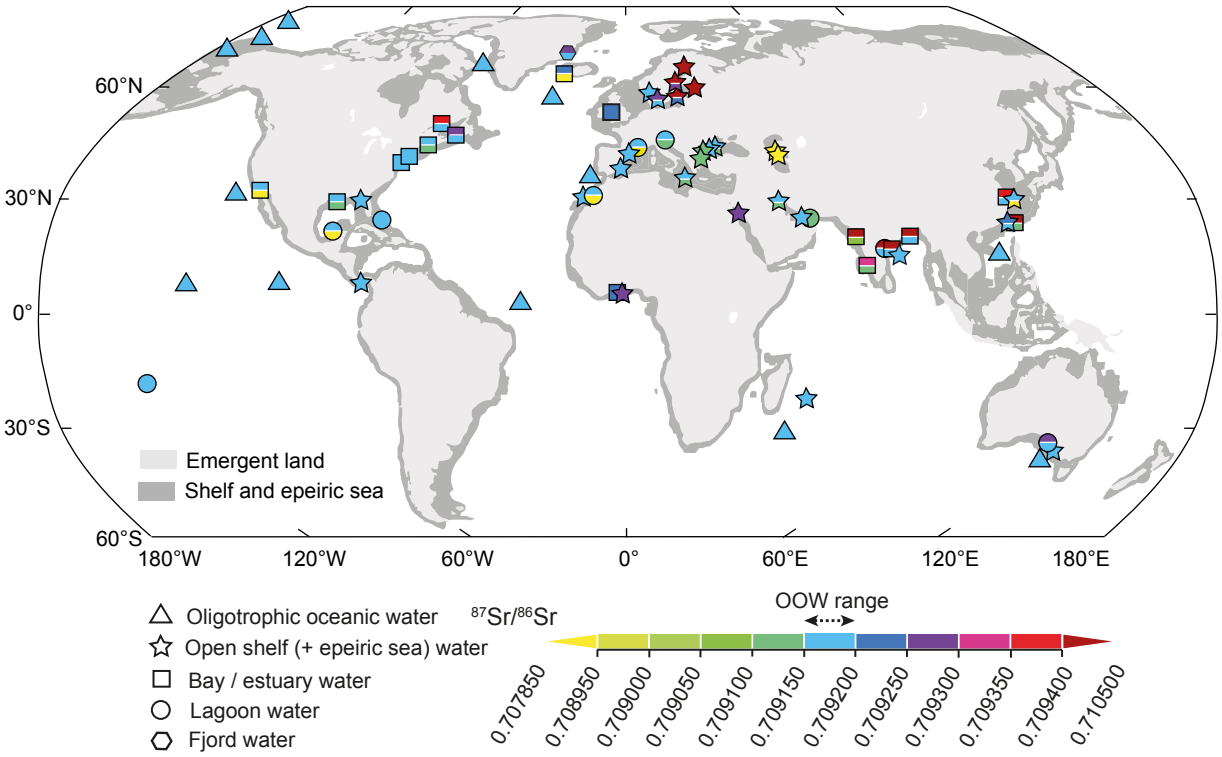
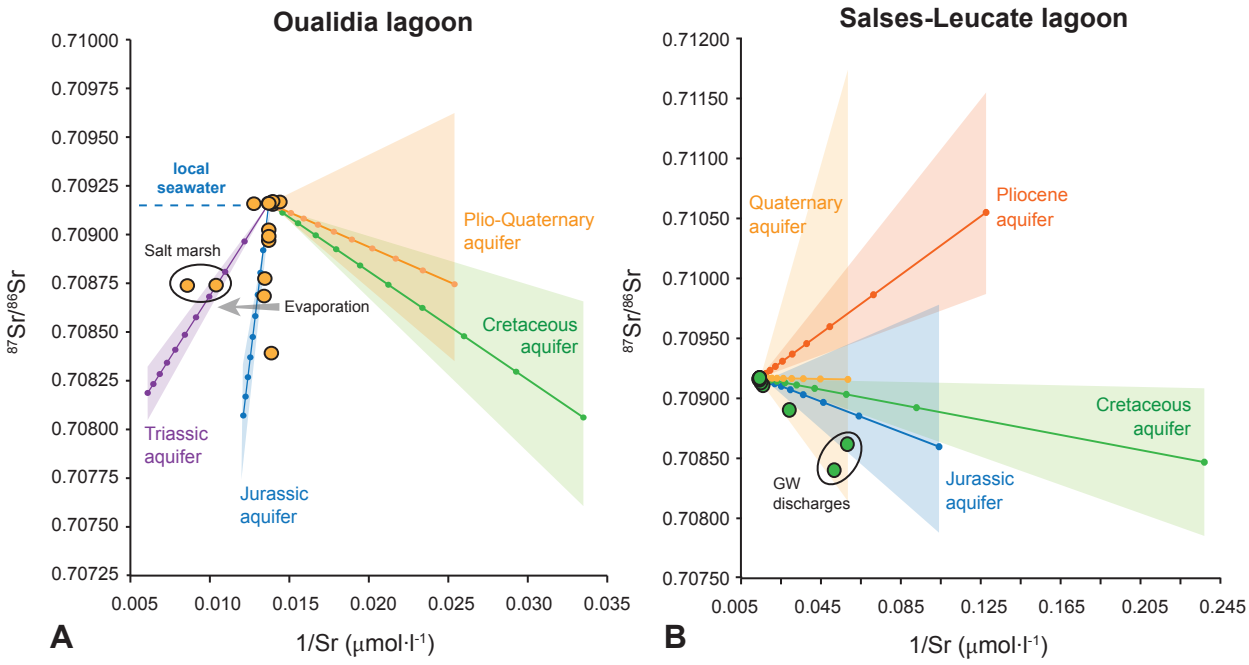
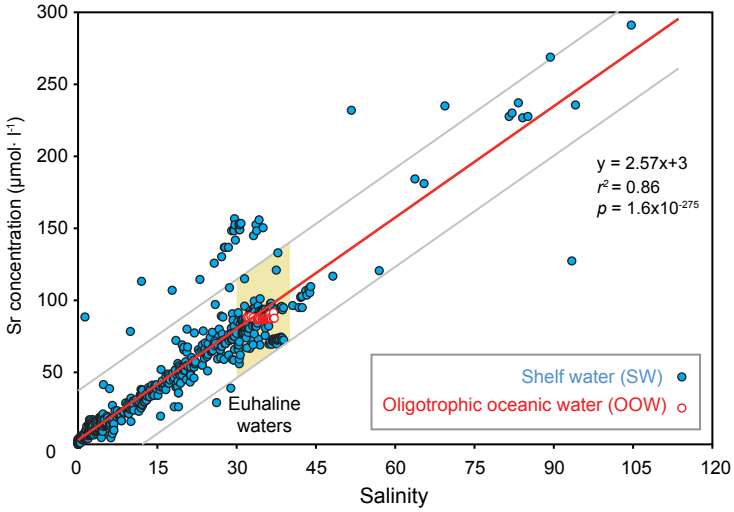
BA2
BA1
Mediterranean Sea
N
5 km
-42.46


D
138.45
Pacific Ocean
138.35
-17.29
TA2
Tatakoto Lagoon
TA1
lagoon inlets
N
1 km
-17.37


A
$^{87}Sr/^{86}Sr$
Salinity
Worldwide RW mixing trend
12 ‰ (threshold of RW influence)
OOW range
Dam
Meso-Cenozoic GW influence
Oualidia oysters sampled in 2006
Evaporation
Salt marsh
Mesozoic GW influence
GW discharges
Average OOW
B
Sr (µmol·l$^{-1}$)
Salinity
Samples:
Water
Shell
Aquifer
Oualidia
Salses-Leucate
BSM transect
Tatakoto
up to 250
Moroccan aquifer waters
Salt marsh
OOW
Evaporation
Sr enriched GW influence
Dam
GW discharges
Sr depleted GW influence
Corbières aquifer waters
C
$SO_4$ (mmol·l$^{-1}$)
Salinity
Salt marsh
Evaporation
OOW
Gypsum dissolution
Freshwater - OOW mixing trend
GW discharges
D
Ca (mmol·l$^{-1}$)
Salinity
Salt marsh
Evaporation
Gypsum & carbonates dissolution
OOW
Freshwater - OOW mixing trend
GW discharges


60°N
30°N
0°
30°S
60°S
180°W
120°W
60°W
0°
60°E
120°E
180°E
Emergent land
Shelf and epeiric sea
Oligotrophic oceanic water
Open shelf (+ epeiric sea) water
Bay / estuary water
Lagoon water
Fjord water
87Sr/86Sr
OOW range
0.707850
0.708950
0.709000
0.709050
0.709100
0.709150
0.709200
0.709250
0.709300
0.709350
0.709400
0.710500


Oualidia lagoon
0.71000
0.70975
0.70950
0.70925
0.70900
0.70875
0.70850
0.70825
0.70800
0.70775
0.70750
0.70725
87Sr/86Sr
local seawater
Salt marsh
Evaporation
Plio-Quaternary aquifer
Cretaceous aquifer
Triassic aquifer
Jurassic aquifer
0.005
0.010
0.015
0.020
0.025
0.030
0.035
1/Sr (µmol·l-1)
A
Salses-Leucate lagoon
0.71200
0.71150
0.71100
0.71050
0.71000
0.70950
0.70900
0.70850
0.70800
0.70750
87Sr/86Sr
Quaternary aquifer
Pliocene aquifer
Cretaceous aquifer
GW discharges
Jurassic aquifer
0.005
0.045
0.085
0.125
0.165
0.205
0.245
1/Sr (µmol·l-1)
B


Sr concentration (µmol· l-1)
Salinity
y = 2.57x+3
r2 = 0.86
p = 1.6x10-275
Euhaline waters
Shelf water (SW)
Oligotrophic oceanic water (OOW)
0
50
100
150
200
250
300
0
15
30
45
60
75
90
105
120

| Studied site | Sample | Water type (environment) | Water depth (mbsl) | Salinity | Ca mmol·l$^{-1}$ | Mg mmol·l$^{-1}$ | Na mmol·l$^{-1}$ | K mmol·l$^{-1}$ | $SO_4$ mmol·l$^{-1}$ | Cl mmol·l$^{-1}$ | Li µmol·l$^{-1}$ | Br µmol·l$^{-1}$ | Sr µmol·l$^{-1}$ | $^{87}Sr/^{86}Sr$ | 2 std dev |
|---|---|---|---|---|---|---|---|---|---|---|---|---|---|---|---|
| **Oualidia lagoon, Morocco** | OU2 | Atlantic seawater (beach) | 0 | 37.2 | 11.47 | 59.23 | 567 | 9.97 | 30.51 | 581 | 48.72 | 865 | 73.0 | 0.709158 | 0.000004 |
| | OU3 | Lagoon water | 0 | 28.7 | 12.08 | 47.85 | 453 | 7.83 | 26.33 | 448 | 36.26 | 667 | 74.4 | 0.708772 | 0.000004 |
| | OU4 | Lagoon water | 0 | 26.0 | 11.99 | 42.47 | 399 | 6.96 | 25.01 | 406 | 31.53 | 597 | 74.6 | 0.708682 | 0.000004 |
| | OU5 | Lagoon water | 0 | 18.5 | 11.84 | 30.88 | 276 | 4.85 | 20.92 | 289 | 23.17 | 420 | 72.2 | 0.708389 | 0.000004 |
| | OU6 | Lagoon water (close to the dam) | 0 | 9.9 | 12.53 | 18.48 | 138 | 2.60 | 17.19 | 155 | 15.85 | 223 | 78.4 | 0.709156 | 0.000004 |
| | OU7 | Lagoon water (salt marsh) | 0 | 40.6 | 17.09 | 71.32 | 653 | 11.55 | 38.13 | 634 | 54.53 | 1012 | 96.4 | 0.708738 | 0.000004 |
| | OU8 | Lagoon water (salt marsh) | 0 | 48.2 | 20.49 | 92.47 | 753 | 15.36 | 48.51 | 753 | 80.57 | 1302 | 116.8 | 0.708736 | 0.000004 |
| | OU9 | Lagoon water | 0 | 33.1 | 11.83 | 53.14 | 535 | 9.37 | 28.53 | 517 | 42.62 | 768 | 73.0 | 0.708989 | 0.000004 |
| | OU10 | Lagoon water | 0 | 37.1 | 11.16 | 57.34 | 586 | 10.08 | 30.49 | 580 | 47.44 | 862 | 71.7 | 0.709168 | 0.000004 |
| | OU11 | Lagoon water | 0 | 36.6 | 11.38 | 58.32 | 580 | 10.09 | 30.10 | 571 | 45.66 | 862 | 69.5 | 0.709165 | 0.000004 |
| | OU12 | Lagoon water | 0 | 34.1 | 11.60 | 53.28 | 542 | 9.34 | 29.12 | 532 | 43.26 | 786 | 73.0 | 0.709022 | 0.000004 |
| | OU13 | Lagoon water | 0 | 32.9 | 11.61 | 52.32 | 524 | 9.04 | 28.53 | 514 | 41.77 | 762 | 73.0 | 0.708967 | 0.000004 |
| | OU14 | Lagoon water | 0 | 37.0 | 11.54 | 58.44 | 591 | 10.41 | 30.38 | 578 | 48.10 | 856 | 72.1 | 0.709162 | 0.000004 |
| | OU15 | Lagoon water | 3 | 37.1 | 11.34 | 57.23 | 598 | 10.20 | 30.47 | 579 | 49.57 | 855 | 71.8 | 0.709164 | 0.000004 |
| | OU16 | Lagoon water | 3 | 36.8 | 11.51 | 57.00 | 601 | 10.22 | 30.26 | 575 | 46.99 | 846 | 71.7 | 0.709152 | 0.000004 |
| | OU17 | Lagoon water | 3 | 36.9 | 11.76 | 58.65 | 600 | 10.26 | 30.35 | 576 | 47.06 | 833 | 73.1 | 0.709159 | 0.000004 |
| **Salses-Leucate lagoon, France** | LEU9 | Mediterranean seawater (beach) | 0 | 37.6 | 11.95 | 61.04 | 598 | 10.71 | 30.82 | 586 | 46.71 | 871 | 69.7 | 0.709170 | 0.000004 |
| | LEU1 | Lagoon water (mussel farm) | 0.5 | 36.3 | 11.31 | 57.40 | 577 | 10.24 | 30.00 | 567 | 45.72 | 848 | 71.0 | 0.709164 | 0.000004 |
| | LEU2 | Lagoon water | 0 | 36.1 | 11.50 | 57.25 | 579 | 10.39 | 29.71 | 563 | 46.54 | 849 | 69.7 | 0.709170 | 0.000004 |
| | LEU3 | Lagoon water | 0 | 37.4 | 11.69 | 59.52 | 593 | 10.61 | 30.75 | 584 | 47.81 | 880 | 71.9 | 0.709161 | 0.000004 |
| | LEU4 | Lagoon water | 0 | 34.0 | 11.82 | 56.83 | 576 | 9.95 | 28.17 | 531 | 42.88 | 786 | 69.8 | 0.709155 | 0.000004 |
| | LEU5 | Lagoon water | 1.5 | 32.8 | 11.61 | 54.76 | 553 | 9.65 | 27.25 | 512 | 39.48 | 762 | 69.8 | 0.709169 | 0.000004 |
| | LEU6 | Lagoon water | 0 | 32.5 | 11.85 | 54.75 | 556 | 9.67 | 26.97 | 507 | 41.36 | 746 | 66.7 | 0.709132 | 0.000004 |
| | LEU7 | Lagoon water | 0 | 11.8 | 6.82 | 20.23 | 186 | 3.25 | 10.93 | 185 | 11.42 | 235 | 34.2 | 0.708903 | 0.000004 |
| | LEU8 | Lagoon water | 2.5 | 27.1 | 10.33 | 44.94 | 454 | 7.73 | 22.79 | 422 | 30.55 | 618 | 62.3 | 0.709109 | 0.000004 |
| | LEU11 | Groundwater (Font Estramar dischar | 0 | 4.3 | 3.49 | 6.83 | 39 | 1.07 | 4.52 | 67 | 4.78 | 86 | 19.3 | 0.708401 | 0.000004 |
| | LEU10 | Groundwater (Font Dame discharge) | 0 | 2.7 | 3.71 | 4.37 | 9 | 0.62 | 3.78 | 42 | 2.77 | 51 | 17.2 | 0.708618 | 0.000004 |
| **Banyuls-sur-Mer transect, France** | BA1 | Mediterranean seawater (open shelf) | 22 | 38.1 | 12.15 | 60.73 | 622 | 10.57 | 31.29 | 595 | 45.09 | 894 | 73.1 | 0.709174 | 0.000004 |
| | BA2 | Mediterranean seawater (open shelf) | 0 | 37.8 | 11.76 | 58.50 | 602 | 10.25 | 30.94 | 590 | 48.45 | 884 | 72.7 | 0.709167 | 0.000004 |
| | BA3 | Mediterranean seawater (open shelf) | 40 | 38.4 | 11.60 | 59.69 | 588 | 10.46 | 31.49 | 600 | 47.76 | 913 | 74.5 | 0.709167 | 0.000005 |
| | BA4 | Mediterranean seawater (open shelf) | 0 | 38.0 | 11.24 | 59.91 | 586 | 10.33 | 31.14 | 593 | 45.55 | 893 | 74.0 | 0.709173 | 0.000004 |
| | BA5 | Mediterranean seawater (open shelf) | 140 | 38.6 | 11.43 | 59.52 | 579 | 10.40 | 31.66 | 603 | 50.66 | 900 | 73.9 | 0.709170 | 0.000004 |
| | BA6 | Mediterranean seawater (open shelf) | 0 | 38.5 | 11.75 | 60.35 | 594 | 10.60 | 31.58 | 601 | 49.11 | 914 | 73.6 | 0.709176 | 0.000004 |
| | BA7 | Mediterranean seawater (open shelf) | 510 | 38.9 | 12.17 | 63.26 | 615 | 11.19 | 31.91 | 608 | 49.62 | 904 | 72.4 | 0.709166 | 0.000004 |
| | BA8 | Mediterranean seawater (open shelf) | 0 | 38.6 | 11.71 | 60.22 | 612 | 10.75 | 31.64 | 603 | 47.38 | 911 | 72.2 | 0.709174 | 0.000004 |
| | BA9 | Mediterranean seawater (beach) | 0 | 26.4 | 8.27 | 42.21 | 432 | 7.40 | 21.87 | 412 | 30.34 | 605 | 59.3 | 0.709206 | 0.000004 |
| **Tatakoto lagoon, French Polynesi** | T1 | Lagoon water | 2 | 37.6 | 11.73 | 61.81 | 599 | 10.84 | 30.75 | 587 | 50.84 | 889 | 72.4 | 0.709165 | 0.000004 |
| | T2 | Lagoon water | 2 | 37.5 | 12.02 | 62.77 | 619 | 11.16 | 30.75 | 585 | 50.03 | 887 | 73.5 | 0.709171 | 0.000004 |
| **Global ocean#** | - | Oligotrophic Oceanic water (OOW) | - | 34.9 | 10.28 | 54.10 | 468 | 10.21 | 28.30 | 545 | 24.6 | 873 | 88.0 | 0.709172 | 0.000023 |

| Studied site | Shell sample | Corresponding water sample | Species | Ca ($mg \cdot g^{-1}$) | Mg ($\mu g \cdot g^{-1}$) | Sr ($\mu g \cdot g^{-1}$) | $^{87}Sr/^{86}Sr$ | 2 std dev |
|---|---|---|---|---|---|---|---|---|
| **Oualidia lagoon, Morocco** | O 1 | OU 3 | *Crassostrea gigas* | 370 | 3888 | 1091 | 0.709174 | 0.000004 |
| | O 2 | OU 9 | *Crassostrea gigas* | 371 | 1361 | 575 | 0.709173 | 0.000004 |
| | O 3 | OU 12 | *Venerupis decussata* | 334 | 118 | 1689 | 0.709179 | 0.000004 |
| | O 4 | OU 9 | *Solen marginatus* | 361 | 268 | 2797 | 0.708939 | 0.000004 |
| **Salses-Leucate lagoon, France** | S1 | LEU 9 | *Mytilus galloprovincialis* | 354 | 483 | 746 | 0.709143 | 0.000004 |
| | S2 | LEU 2 | *Glycymeris glycymeris* | *na* | *na* | *na* | 0.709125 | 0.000004 |
| | S3 | LEU 3 | *Crassostrea sp.* | *na* | *na* | *na* | 0.709136 | 0.000004 |
| | S4 | LEU 4 | *Mytilus galloprovincialis* | 364 | 945 | 969 | 0.709174 | 0.000004 |
| | S5 | LEU 5 | *Mytilus galloprovincialis* | 378 | 1291 | 932 | 0.709158 | 0.000004 |
| | S6 | LEU 6 | *Mytilus galloprovincialis* | 349 | 135 | 1511 | 0.709177 | 0.000004 |
| **Banyuls-sur-Mer, France** | Bs 1 | B 1 | Bivalve fragments | 360 | 792 | 1513 | 0.709174 | 0.000004 |
| | Bs 2 | B 3 | *Turritella communis* | 362 | 413 | 1645 | 0.709173 | 0.000004 |
| | Bs 3 | B 5 | Bivalve fragments | *na* | *na* | *na* | 0.709165 | 0.000004 |
| | Bs 4 | B 7 | Bivalve fragments | 381 | 11554 | 2214 | 0.709171 | 0.000004 |
| **Tatakoto lagoon, French Polynesia** | Ts 1 | T1 | *Tridacna maxima* | *na* | *na* | *na* | 0.709178 | 0.000004 |
| | Ts 2 | T1 | *Tridacna maxima* | 365 | 810 | 1587 | 0.709164 | 0.000004 |
| | Ts 3 | T2 | *Tridacna maxima* | *na* | *na* | *na* | 0.709182 | 0.000004 |
| | Ts 4 | T2 | *Tridacna maxima* | 357 | 205 | 1644 | 0.709172 | 0.000004 |

| Studied sites | End-members | Water type | Average Sr ($\mu mol \cdot l^{-1}$) | Minimal Sr ($\mu mol \cdot l^{-1}$) | Maximal Sr ($\mu mol \cdot l^{-1}$) | Average $^{87}Sr/^{86}Sr$ | Minimal $^{87}Sr/^{86}Sr$ | Maximal $^{87}Sr/^{86}Sr$ | References |
|---|---|---|---|---|---|---|---|---|---|
| **Morocco** | Plio-Quaternary aquifer | GW | 39.4 | 1.4 | 148.0 | 0.70874 | 0.70835 | 0.70962 | Fakir et al., 2002; Kaid Rassou et al., 2005; Vinson et al., 2( |
| | Cretaceous aquifer | GW | 29.8 | 3.0 | 116.4 | 0.70760 | 0.70760 | 0.70866 | Fakir et al., 2002; Fadili et al., 2015; Bouchaou et al., 2017 |
| | Jurassic aquifer | GW | 82.7 | 36.3 | 249.9 | 0.70807 | 0.70774 | 0.70833 | Kaid Rassou et al., 2005; Bouchaou et al., 2017 |
| | Triassic aquifer | GW | 164.9 | 164.4 | 165.5 | 0.70819 | 0.70805 | 0.70832 | Bouchaou et al., 2017 |
| | Atlantic water at Oualidia | SW | 73.0 | - | - | 0.70916 | - | - | *this study* |
| **Corbières region, France** | Quaternary aquifer | GW | 17.1 | 1.8 | 85.5 | 0.70916 | 0.70814 | 0.71174 | Petelet et al., 1998; Petelet-Giraud et al., 2016 |
| | Pliocene aquifer | GW | 7.8 | 4.7 | 15.2 | 0.71055 | 0.70987 | 0.71155 | Petelet-Giraud et al., 2016 |
| | Cretaceous aquifer | GW | 4.2 | 1.2 | 9.0 | 0.70847 | 0.70753 | 0.70908 | Khaska et al., 2013 |
| | Jurassic aquifer | GW | 9.6 | 0.3 | 53.0 | 0.70860 | 0.70788 | 0.70978 | Aquilina et al., 2003; Petelet-Giraud et al., 2016 |
| | Mediterranean water at Salses-Leucate | SW | 69.7 | - | - | 0.70917 | - | - | *this study* |

)11