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1	Video monitoring of in-channel wood: from flux
2	characterization and prediction to recommendations
3	to equip stations

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8 Abstract

9 Wood flux (piece number per time interval) is a key parameter for understanding wood budgeting, determining the controlling factors, and 10 11 managing flood risk in a river basin. Quantitative wood flux data is critically needed to improve the understanding of wood dynamics and estimate wood 12 13 discharge in rivers. In this study, the streamside videography technique was 14 applied to detect wood passage and measure instantaneous rates of wood 15 transport. The goal was to better understand how wood flux responds to flood 16 and wind events and then predict wood flux. In total, one exceptional wind and 17 7 flood events were monitored on the Ain River, France, and around than 24000 18 wood pieces were detected visually. It is confirmed that, in general, there is a 19 threshold of wood motion in the river equal to 60% of bankfull discharge. 20 However, in a flood following a windy day, no obvious threshold for wood motion

21 was observed, confirms that wind is important for the preparation of wood for 22 transport between floods. In two multi-peaks floods, around two-thirds of the 23 total amount of wood was delivered on the first peak, which confirms the importance of the time between floods for predicting wood fluxes. Moreover, we 24 25 found an empirical relation between wood frequency and wood discharge, 26 which is used to estimate the total wood amount produced by each of the floods. 27 The data set is then used to develop a random forest regression model to 28 predict wood frequency as a function of three input variables that are derived 29 from the flow hydrograph. The model calculates the total wood volume either 30 during day or night based on the video monitoring technique for the first time, 31 which expands its utility for wood budgeting in a watershed. A one-to-one link 32 is then established between the fraction of detected pieces of wood and the dimensionless parameter "passing time × framerate", which provides a 33 34 general guideline for the design of monitoring stations.

Keywords: Fluvial dynamics; Large wood in river; Random forest model; Wind
condition; Multi-peaks discharge; Streamside video monitoring.

37 1. Introduction

38 Floating wood in rivers, known as driftwood, is a significant component of 39 catchments, notably in forested temperate regions (Wohl, 2013; Ruiz-

40 Villanueva et al., 2016a). It is delivered to the rivers by a set of processes 41 (landslides, debris flows, blowdown, bank erosion and so on) which vary from 42 upstream to downstream (Nakamura and Swanson, 1993; Montgomery et al., 43 1996; Abbe and Montgomery, 2003; Gurnell and Petts, 2006). Among different 44 recruitment processes, bank erosion probably delivers most of the large organic 45 material into larger lowland rivers (Keller and Swanson, 1979). These large 46 pieces of wood (i.e. greater than 1 m length and 10 cm diameter) induce 47 variations in hydraulic and sediment dynamics, and contribute to flow resistance 48 and obstructions within the channel (Young, 1991; Gippel, 1995; Shields and 49 Gippel, 1995; Wilcox and Wohl, 2006; Comiti et al., 2008). Especially during a flood, the transport and deposition of large wood pieces represent a potential 50 51 increase in the destructive power of floods, which increases the potential risks to human populations and infrastructures (Lassettre and Kondolf, 2012; De 52 53 Cicco et al., 2018; Mazzorana et al., 2018). For instance, a flow obstruction due 54 to wood accumulation can lead to upstream bed aggradation, channel avulsion, 55 and local scouring processes, which can in turn cause embankment or bridge 56 collapse and floodplain inundation (Diehl, 1997; Lyn et al., 2003; Fischer, 2006; Waldner et al., 2007; Mao et al., 2008; Mazzorana et al., 2009; Comiti et al., 57 58 2012; Ruiz-Villanueva et al., 2014a). Therefore, quantifying wood inputs, 59 transport, deposition, and budgeting in general is crucial for understanding and

60 managing wood risk in rivers.

Understanding the variability and the process-scale dynamics which control 61 62 wood delivery and transport rate is also a critical challenge (Martin and Benda, 63 2001; Benda et al., 2003; Marcus et al., 2011; Schenk et al., 2014; Boivin et al., 64 2015). Wood budgeting can be explored at different time scales. The wood recruitment sites are often observed close to the preferential sites of deposition 65 66 (Schenk et al., 2014; Ravazzolo et al., 2015), but not systematically, as shown along the Isère River, France (Piégay et al., 2017). Some pieces of wood can 67 68 be transported over very long distances during a single flood (Gurnell et al., 69 2002; Gurnell, 2012; Comiti et al., 2016; Kramer and Wohl, 2017). Moreover, 70 the amount of wood can be documented at multi-annual and annual time intervals over long time periods by historical data (Seo et al., 2008; Seo and 71 72 Nakamura, 2009; Ruiz-Villanueva et al., 2014b). Based on this long time scale, 73 however, it is not possible to record continuous series and study wood transport 74 processes during shorter but critical hydrological events such as floods, 75 exceptional wind events, and landslides, which are known to drive wood fluxes 76 in rivers (Lassettre and Kondolf, 2012; Ruiz Villanueva et al., 2014a). 77 To generate wood input series in shorter time scales, Moulin and Piégay,

78 (2004) used weekly time steps to measure the wood stored in a reservoir. The

79 results quantified the timing and magnitude of Large Wood (LW) export during 4 80 flood events in the reservoir and allowed the recruitment and transport processes of LW at the watershed scale to be better understood. Benacchio et 81 82 al. (2017) monitored wood delivery and calculated wood weight in a reservoir by an automated image processing technique using much finer time intervals 83 84 (10 min). In addition to the reservoir-based monitoring, Kramer and Wohl, 85 (2014) showed that in high-discharge, low-velocity rivers, the deployment of 86 monitoring cameras with coarse frame rates (\geq 1 min) enables monitoring of 87 LW transport at large spatial and long temporal scales. However, in smaller and 88 steeper rivers the velocity of wood pieces is higher or the field of view is too 89 small such that low frame rate photography cannot provide accurate estimates 90 of wood delivery.

91 Video monitoring of the water surface can be used to continuously monitor 92 wood flux at a high temporal resolution. Lyn et al. (2003) were the first to apply 93 this technique, using two stream-side video cameras to observe and detect 94 wood accumulation on bridge pier in the Eel River, Unites States. Due to data 95 storage issues, Lyn et al. (2003) downgraded the frame rate to 0.1 fps (frame 96 per second) and applied image compression to the recorded frames through 97 the monitoring period. Such issues were overcome by MacVicar et al. (2009), 98 and MacVicar and Piégay (2012) who established a monitoring station at the 99 Ain River, France, but transferred the full resolution images recorded at 5 fps to 5

100	a remote server for analysis. The high quality and frequency of the data, which
101	is likely necessary in high gradient rivers, allowed them to compare LW
102	dynamics with flood hydrograph and develop a quantitative relation between
103	wood and water discharges. Other studies have implemented similar
104	approaches (Boivin et al., 2015; Kramer et al., 2017; Senter et al., 2017; Ruiz-
105	Villanueva et al., 2018; Ghaffarian et al., 2020a) but overall the technique
106	remains undersubscribed and models of the wood flux as a function of the flow
107	hydrograph remain poorly parameterized.

108 Overall, the success of a particular monitoring station will be determined by 109 issues of wood size and image resolution (MacVicar and Piégay, 2012; Ghaffarian et al., 2020a). Ghaffarian et al. (2020a) monitored floods on the Isère 110 River (France) and demonstrated the generalizability of technique to other 111 112 rivers along with limits, constraints, methodological some and 113 recommendations. The oblique angle of the camera means that it is particularly 114 important to understand where wood will pass relative to the camera position 115 (Ghaffarian et al., 2020a). Moreover, a problem remains that there are gaps 116 within the data. Such gaps can occur due to the poor visibility in low light or 117 cloudy weather, lost connections where data is transferred to a remote server 118 for storage (Muste et al., 2008; MacVicar et al., 2009; MacVicar and Piégay, 119 2012; Ghaffarian et al. 2020a), or simply to the time required to extract

120 information about floating wood from videos. Despite some efforts at automatic 121 extraction (Ali and Tougne, 2009; Lemaire et al., 2014), the procedure to date 122 remains predominantly manual. Improved modeling of wood fluxes as a 123 function of flow hydrographs or other environmental conditions could be an 124 effective strategy to reduce sampling effort and fill in missing data such that 125 wood fluxes could be integrated over time to support wood budgeting in 126 watersheds.

127 The aim of the current study is to advance the video monitoring technique for wood flux measurement by addressing the following questions: i) is wood 128 129 transported only above a discharge threshold, and if so, is the threshold a function of antecedent conditions? ii) Can wood flux be modelled as a function 130 131 of the flood hydrograph? and iii) Can we accurately estimate wood flux from sampling? The analysis uses the database assembled by MacVicar and Piégay 132 133 (2012) of sampled periods during three floods on the Ain River but significantly 134 adds to this work by performing a complete analysis of the daytime videos from 135 four new flood events and one period with low flow but an exceptional wind 136 condition, which was then followed by a flood event. This much larger database 137 comprises nearly 180 hours of annotated videos or around than 24,000 138 annotated wood pieces, which substantially expands on the 18 hours and 7800 139 wood pieces monitored by MacVicar & Piégay (2012). The windy day event with 7

140	35-year return period allowed us to address the first research question. A
141	random forest (RF) model was used to answer the second question. Flux and
142	wood measurements from the MacVicar and Piégay (2012) database combined
143	with the data of the present study are used to resolve the third question.

144 **2. Study site**

145 The study site is located on the lower Ain River, a sixth-order piedmont river 146 flowing through a forested corridor in France. The channel is typically single 147 thread with occasional islands, and a wandering system with prominent 148 meander scrolls and cutoff channels (Figure 1.a) (MacVicar et al., 2009). The 149 hydrograph shows a strong seasonal pattern, with low flows in the summer and 150 most of floods occurring between October and April. Bed material sizes are 151 gravel-cobble mix with a median size of 2.5 cm. The unvegetated channel width is 65 m in average at the study site, actively shifting so that significant amount 152 153 of wood is delivered by bank erosion. Tree species established in the floodplain 154 are a mix of soft and hardwood species dominated by black poplar (Populus nigra) that can reach up to 75 cm in diameter and 25 m in height (MacVicar and 155 Piégay, 2012). Along the study site, wood influx has been estimated over 156 several decades from the analysis of aerial photographs at 18 to 38 m³/km/yr 157 158 (Lassettre et al., 2008).

159	Floating wood was counted on the river at Pont de Chazey, where a stream
160	gauge is maintained by a regional authority (Figure 1.b, c). Along the river, the
161	characteristic discharge of 1.5-year return period was $Q_{1.5} = 840$ m ³ /s
162	(Ghaffarian et al., 2020a), and at this study site, an estimated bankfull discharge
163	(Q_{bf}) of 530m ³ /s was confirmed by visual observation (MacVicar and Piégay,
164	2012). At this point the flow discharge is calculated based on the water elevation
165	measured at the gauging station. These data are available online from 1959 at
166	(www.hydro.eaufrance.fr). Mean daily wind speed is also available from the
167	Meteorological Station of Lyon-Bron (1949-2020) (see Figure 2).



Figure 1. Study site at Pont de Chazey: a) Location of the Ain River course in France and location of the gauging and meteorological stations, b) camera position and its view angle in yellow, c) overview of the gauging station with the camera installation point, d) view of the River channel from the camera

168 3. Material and Methods

169 3.1. Stream-side video camera

- 170 Wood pieces were monitored at Pont-de-Chazey gauging station using an
- 171 AXIS P221 Day/Night[™] fixed network camera installed in the spring of 2007.
- 172 Figure 1.d shows the camera field of view on the river surface. The video

173 camera can supply high resolution (HDTV720P) surveillance even in extreme 174 low-light, though not at night time. The camera was located on the side of the 175 river closest to the thalweg to provide a maximum resolution where the majority 176 of wood pieces are observed. The camera elevation is 9.84 m above the base 177 flow surface at a sufficiently wide angle to afford a view of the entire river width 178 during most periods. Ethernet connectivity enables the automatic transfer of 179 recorded videos to a central server where located at CNRS UMR 5600 -Environment Ville et Société, Site of École Normale Supérieure, Lyon, France. 180 181 Videos were recorded continuously at a maximum frequency of 3 to 5 fps. Data 182 was recorded with this camera from 2007 to 2011 at a resolution of 640×480 183 pixels and from 2012 to 2016 at 768×576 pixels. The first three floods (events F1 to F3) thus have a lower resolution than the final four floods and windy period 184 (events F4 to F7 and W1). At minimum compression, each video segment 185 186 occupied approximately 94Mb of memory and approximately 15 minutes so that 187 a 4TB hard drive stored approximately one year of video. Flood levels were 188 reviewed every few days and videos of interest were saved for later analysis.

189 **3.2. Monitored events**

In total, 7 flood events were monitored in this study (Table 1). Three flood
events from 2007 to 2008 were collected from MacVicar & Piégay (2012),

192	referred to herein as events F1 to F3 (Figure 2.a, red lines). A video camera
193	has been recorded video at this location more or less continuously from 2007.
194	For the current work, four additional flood events between 2012 to 2014 were
195	selected for study and sampling and are referred to as events F4 to F7 (Figure
196	2.a, blue lines). The floods range from 578 m ³ /s ($\cong Q_{bf}$) to 1020 m ³ /s ($\cong 2Q_{bf}$).
197	Event F7 was selected to assess whether wind has an effect on the wood
198	delivery because it occurred just two days after an exceptional windy day. The
199	windy day occurred on December 24, 2013 and is referred to herein as event
200	W1 (Figure 2.b). The average daily wind speed on this day was 13.6 m/s, which
201	is considered to be a one in 35 year event based on a Gumbel distribution of
202	the over 70 years of record (Yue et al., 1999).

	ts	Peak flows (m3/s)		wind ity(m/s /zed (hr)		tored on*	Number of floating woods	
Flood periods	Even	total	daylight	Daily veloc	Anal) videc	Moni	Rising limb	falling limb
22 to 24-Nov-2007	F1	578	576	6.6	06:15	09%	2800	38
10 to 12-Dec-2007	F2	616	616	6.3	03:45	05%	968	93
10 to 13-Apr-2008	F3	1050	1007	3.8	07:45	08%	3331	584
01 to 07-Jan-2012	F4	808	807	4.9	57:00	34%	3681	1641
15 to 16-Dec-2012	F5	932	821	4.9	17:15	36%	6901	798
01 to 06-Feb-2013	F6	701	701	8.5	56:30	39%	1040	473
24 to 25-Dec-2013	W1	134	134	13.6	08:45	37%	8	-
25 to 27-Dec-2013	F7	600	580	5.6	25:45	36%	1443	43

Table 1 Wood sampling statistics at the Pont de Chazey for different events.

203 * Monitored fraction = monitored duration / total duration of an event



Figure 2. Monitored events a) the daily mean discharge series monitored by MacVicar & Piégay (2012) (red lines) and monitored in this work (blue lines) on the discharge series from 2007 to 2014. b) The daily mean wind velocity series from 1977 to 2013.

204 3.3. Monitoring process

205	In total 183 hours of video was analyzed, including 18 hours monitored by
206	MacVicar & Piégay (2012) (Table 1). For this analysis, floating wood was
207	visually detected by an operator and the position of each piece of wood was
208	digitally annotated frame by frame via a graphical user interface.
209	Two methods were applied for event monitoring: (i) 15-minute monitoring
210	intervals for events F1 to F3; and (ii) continuous monitoring for events F4 to F7
211	and W1. In the first approach, applied by MacVicar and Piégay (2012), a 15-
212	minute video segment was monitored for each daytime hour (e.g. from 8:30 to

213	8:45, 9:30 to 9:45, etc. to ~5:30) and, by multiplying the number of detected
214	pieces by four, the wood flux per hour was extrapolated to be compared with
215	other studies. A problem with this sampling strategy was noted by Ghaffarian
216	et al. (2020a), who showed that a 15 min interval may not be sufficient to reliably
217	estimate the hourly wood flux due to short term variability. For this reason, all
218	daytime periods were monitored by an operator for the flood events added as
219	part of the current analysis.

By extracting the detection time for each piece of wood (indicated on top of each frame, see Figure 4.a), wood flux was calculated as the number of wood pieces within a time interval. In the current study, an hour time interval was selected to model the wood fluxes through the flood events (sections 4.1, and 4.2), again for the reasons highlighted by Ghaffarian et al. (2020a). One and 10-minute time intervals were used for analysis in section 4.3 to assess the importance of shorter-term pulses on overall wood fluxes.

227 **3.4. Observer bias**

The analyzed events in this work are based on two different operators (MacVicar and Zhang). During the detection process, the operator bias can play a role in the quantity of wood fluxes. To check this effect, 13 segments of 15minute videos from events F1 to F3 were selected and wood was detected by 232 both operators following the process used by Ghaffarian et al. (2020a). These 233 video segments were selected such that they cover different light conditions 234 (e.g. sunshine or cloudy weather or different day times) to evaluate the operator 235 visions in this range of conditions. Also, the amount of wood pieces varies 236 greatly among the videos (from 0 to more than 300 pieces), which allowed us 237 to assess whether bias was affected by wood frequency. Overall, there was a 238 ~7% difference in wood flux estimates between the two observers, with most discrepancies occurring when many small wood pieces pass through the image 239 240 within a short time interval.

241 **3.5. Modeling wood flux**

242 A random forest (RF) non-linear regression algorithm was applied to model the link between wood flux and flow discharge in this study. It produces multiple 243 decision trees (here, 500), each of which is trained on a randomly selected 244 245 subset of the data (in-bag portion) while the remaining subset is used to test 246 the regression and assess its performance (out-of-bag portion)(Breiman, 2001; 247 Hastie et al., 2009; Belgiu and Drăgut, 2016). The RF error corresponds to the 248 residual sums of squares averaged across all the out-of-bag portions of the 249 regression trees. The importance of a variable in the RF model can be assessed 250 through a score that corresponds to the total decrease in error due to splits on

that particular variable, averaged across all trees (Breiman, 2001).

252 For the current study, the response variable was the wood flux and the 253 predictor variables were all derived from the flow time series. We considered 254 three predictors that could influence the wood flux during flood including: (i) flow discharge Q(t), (ii) the time elapsed since the last time that Q was higher or 255 256 equal to Q(t), known as T_0 , and (iii) the gradient of discharge over a time lag (5 257 min) dQ/dt. The application of these predictors in the model is presented in the 258 results (section 4.2). Due to gaps in sampling (e.g., during night time), periods 259 where the time interval between two consecutive detections exceeded 10 hr 260 were removed from the data. In cases when several pieces of wood were 261 annotated in the same image frame, we assume a time interval of 0.5 s between wood pieces. 262

263 The RF and all related data-wrangling were carried out using the R software 264 (R Core Team, 2019) and the Random Forest package (Liaw and Wiener, 2002). 265 The random forest consisted of a default number of trees set to 500 and the sampling of in-bag/out-of-bag samples was made with replacement. The R 266 267 notebook **RF-related** available gathering all commands is from https://github.com/lvaudor/wood_flux. 268

269 **3.6. From wood flux to wood discharge**

270 In the study by MacVicar and Piégay (2012), wood discharge was 271 calculated as m³/s by estimating the length and diameter of all detected floating 272 wood pieces. This process is time consuming, and a decision was made for the 273 current study that, rather than completing the size measurements, the wood 274 pieces would only be counted for floods F4 through F7. The wood count allowed 275 the calculation of the wood flux as a frequency (pieces/minute). This approach was justified by considering Figure 3, which shows that there was a strong 276 277 correlation between wood flux and wood discharge for the 15 min video 278 segments (see section 3.3) sampled by MacVicar and Piégay (2012) for F1, F2 and F3 ($R^2 = 0.83$). This strong relation gives confidence that wood discharge 279 and the total wood volume can be reliably estimated from the wood flux to allow 280 281 comparison with other studies and models of the wood budget. Extrapolating 282 this relation for other rivers would be an open question that can be the objective 283 of future comparative works.



Figure 3. Wood discharge as a function of wood flux

284 3.7. Sampling strategy

285 Taking advantage of high temporal resolution videography, it is possible to capture all variations of wood flux during a critical event, while low frame rate 286 287 photography can be used to detect only a fraction of wood fluxes in the river. 288 To understand the link between the detected wood fluxes and the frame rate, 289 here the concept of passing time (PT) is introduced as the time that one piece 290 of wood passes through the camera field of view. As the camera has a large 291 oblique view, PT varies a lot from the foreground to background (right side of 292 Figure 4.a). Therefore, to measure PT, the position where most of wood pieces' 293 pass is used. As it is seen in the left side of Figure 4.a, more than 75% of wood 294 pieces pass from 150 to 250 pixels on j direction. The passing time at this region 295 is around $PT \cong 5s$ (right side of Figure 4.a). Theoretically, in one snapshot of 18

the camera corresponds to time t_i , this object can be detectable from $t_i - \frac{PT}{2}$ to



Figure 4. a) wood flux position on video frame b) link between video time laps Δt and the passing time *PT*

By introducing the frame rate (of frame per second *fps*) as one over the time between two consecutive frames ($fps = \frac{1}{\Delta t} = \frac{1}{t_{i+1} - t_i}$), all the objects that pass from $t_i - \frac{PT}{2}$ to $t_i + \frac{PT}{2}$ or from $t_{i+1} - \frac{PT}{2}$ to $t_{i+1} + \frac{PT}{2}$ can be detected by the observer at each camera snapshot (see Figure 4.b). Consequently, if $\Delta t > PT$,

302	there can be some pieces that cannot be detected by the camera (red region in
303	Figure 4.b, top), while if $\Delta t < PT$, we can be sure that no wood piece is missed
304	between each pair of frames (Figure 4.b, bottom). Therefore, the fraction of the
305	detected wood pieces can be defined as the ratio between detected wood
306	pieces in the green region in Figure 4.b and the summation of detected (green
307	region) and missed pieces (red region). To study the link between the frame
308	rate and the fraction of detected wood pieces, all detections in Table 1 were
309	used. Knowing $\Delta t = 0.2s$ (5 <i>fps</i>) and <i>PT</i> \cong 5 <i>s</i> on the Ain river, we know that
310	$\Delta t \gg PT$ means that there is enough overlap between each pair of frames (the
311	condition presented in the bottom of Figure 4.b) and wood can be detected.
312	Note that while PT changes both with discharge conditions and the transvers
313	position of detection, the value $PT \cong 5 s$ is a rough value for estimating the ratio
314	between frame rate and passing time. Moreover, the 'detectability' of wood
315	pieces does not account for wood that is not clearly visible in the frame (for e.g.
316	small pieces far from the camera), which is a separate issue related to image
317	resolution and camera angle/position. Given the detection time for each piece
318	of wood (as recorded on top of each frame - see Figure 4.a), the effect of the
319	frame rate on the number of detectable wood pieces was assessed by artificially
320	changing the frame rate from 0.001 to 5 fps (0.2 $s < \Delta t < 1000s$). Results are
321	presented in section 4.3.

322 **4. Results**

323 **4.1. Estimate of wood fluxes during critical events**

- 324 Overall, the results show 3-stages of (i) rising from a threshold of motion,
- 325 (ii) high but flat at discharges above the bankfull, and then (iii) around one order
- of magnitude lower on the falling limb (Figure 5 and Table 1). In Figure 5 the
- 327 blue scatters from the new events are quite consistent with the events in red
- 328 from MacVicar and Piégay (2012) which validates the sampling technique.



Figure 5. Comparison between wood flux based on sampling (red) and continuous (blue) monitoring and flood hydrograph (black line). The gray boxes show the night time when video monitoring was impossible. Different symbol shapes show different events and are consistent with some of the next figures.

329 During the exceptional windy day (W1 from 8 to 17 hr) almost no wood was

330 detected on the river (Table 1). This means that the wood flux is only observed 331 during flood events. As it is seen in Figure 5 in all cases but F7, there are almost 332 no wood pieces in the river for flow discharge less than ~300 m³/s. In the case 333 of the flood event F7 following W1 (the exceptional wind event), however, the 334 threshold appears to be much lower or non-existent. For this event only, the 335 wood flux increases immediately following the increase in flow discharge, which 336 demonstrates the potential effect of W1, not in terms of transport of floating 337 wood downstream, but in the wood transfer from the river banks to the channel 338 where it can be readily mobilized in the subsequent flood.

339 In Figure 5, events F4 and F6 are characterized by multi-peak hydrographs. 340 Event F4, for example, is characterized by three peaks with similar discharges 341 (Table 2), during which 3098, 1134 and 839 pieces of wood were observed respectively in the first to third peaks. Event F6 is characterized by two slightly 342 343 lower peaks, and 995 and 427 pieces of wood were observed in two peaks, 344 respectively (Table 2). Despite some differences in the timing of the floods with 345 respect to daylight hours, these results do indicate that around two-thirds of the 346 wood are mobilized in the first peak of a multi-peak flood. As the number of 347 peaks increases, it also appears that the amount of transported wood 348 progressively decreases.

Video monitoring of in-channel wood

Tab	ole 2 Wood f	lux in multi	peak floods	F4 and F6	
Flood event	F4.1(Peak1)	F4.2(Peak2)	F4.3(Peak3)	F6.1(Peak1)	F6.2(Peak2)
Q _{max} (m ³ /s)	801	808	786	701	627
Pieces number	3098	1134	839	995	427
Fraction*	61%	23%	16%	71%	29%
Flux on rising limb (num/hr)	268	211	82	97	35

* Fraction = piece number during one peak / total piece number during an event.

4.2. Predicting wood fluxes from the flow hydrograph

351 As described in section 3.5, three predictors derived from the flow 352 hydrograph that were thought to influence the wood flux during the flood were 353 used to develop a RF model. Figure 6 shows the link between (i) flow discharge (O(t)) (Figure 6.a), (ii) the gradient of discharge over 5 min time lag (dO/dt)354 (Figure 6.b), and (iii) the time elapsed since the last time that Q was higher or 355 356 equal to Q(t) (T_0) (Figure 6.c) from one hand, and the wood flux from the other 357 hand. Regarding the first predictor, as is seen in Figure 6.a, Q(t) has a non-358 linear positive relationship with the wood flux. Wood flux starts to respond to 359 Q(t) from a threshold almost equal to 450 m³/s and reaches its maximum value 360 at around 850 m³/s. These values agree with observed values in Figure 5. For the second predictor, a comparison between positive and negative values of 361 dQ/dt (rising and falling limb) in Figure 6.b shows that while there is a strong 362 effect of flow discharge gradient on the rising limb, there is almost no effect of 363 the discharge gradient on the falling limb. Finally, as seen in Figure 6.c even 364 with a strong initial fluctuation, the wood flux increases with increasing inter-365

366 flood time.



Figure 6 Predicted value of wood flux (in blue) as a function of a) flow discharge Q (m3/s), b) discharge gradient dQ/dt (m3/s/1hr) and c) the time elapsed since the last time that Q was higher or equal to Q(t), T_Q (days). Dashed lines indicate the 90% and 10% quantiles of the data.

367	Figure 7 compares the observed and the modelled wood fluxes time series
368	(aggregated by hour) for continuous (blue) and sampled (red) videos. The clear
369	strength of the model is that the modelled flux is continuous and provides
370	information during the night and other gaps in the wood sampling database. In
371	terms of performance, the number of trees in the RF model (500) was sufficient
372	to show a convergence on the minimum error possible from this data set. The
373	final average R ² for the out-of-bag portion across all trees was 49.5%. When
374	carrying out cross-validation for the RF as a whole (with 80% of the data
375	randomly sampledwithout replacement- as the training set and 20% as the
376	test set) the R^2 for training set was also 49.5% on average across all trees for
377	the training set (estimated on the out-of-bag data) and 69.8% on the test set.
378	The most important predictor is T_Q (responsible for 41% of the total increase in

379	node purity –as measured by the residual sum of squares-), followed by Q (32%)
380	and dQ/dt (27%). To assess the efficiency of the model more objectively,
381	Figure 8 compares observed and modelled data on the rising and falling limbs
382	of the hydrograph at each event. Each data point represents the sum of wood
383	flux values over the entire limb of the flood during the daylight. As shown, the
384	model predicts the observations with a precision estimated to about 95%.



Figure 7. Wood fluxes based on continuous (blue) and sampled (red) videos and



modelled wood fluxes (green line) using RF model as a function of time.

Figure 8. Comparison between observed and modelled piece number: filled and empty scatters show data on the rising and falling limbs of the hydrograph, respectively. Data are compared with a 1:1 line. There are 3 points for F4 and 2 points for F6 due to multiple peak floods.

Based on the field observations and the RF modelled wood fluxes, it is 385 possible to check both the wood mobility during the night and the critical 386 387 threshold of motion. The critical threshold of motion is defined by the discharge 388 which initiates the mobility of wood flux on the rising limb of the flood. Moreover, 389 to be able to compare the wood volume in two different approaches (observation and model) the process described in section 3.6 was used. 390 391 The new phenomenon that is observed here is the exceptional windy day W1 with low flow $(Q < 0.18Q_{bf})$ which is followed by a flood $(Q > Q_{bf})$ F7. 392 During this wind event, almost no wood flux was detected at the video 393

394 monitoring station (only 2 m³). In the subsequent flood, however, the threshold
 28

of wood motion was approximately $0.2Q_{bf}$ (95 m³/s), which is significantly less than the threshold at $0.6Q_{bf}$ for the other flood events (Table 3). Note that in some cases the thresholds occurred during the night, in which case the presented values are the modelled results.

F4 Event F1 F2 F3 F5 F6 W1 F7 Modelled wood volume* (m³) 218.69 84.95 680.68 347.08 412.54 52.81 1.88 77.11 Observed wood volume (m³) 88.75 32.41 120.01 118.29 235.05 26.12 0.03 29.36 Threshold (m³/s) 275 300 300 300 350 356 <95 95 Modelled/Observed 0 0 Μ Μ 0 0 Μ Μ

Table 3 Wood volume and threshold of wood motion, modelled (M) or observed (O).

399 * Modeled wood volume includes volume during both day and night time.

400 **4.3. Validation optimal wood flux estimate from sampling**

401 The temporal resolution of video monitoring plays a significant role on the 402 quantity of monitored data. By introducing the passing time *PT* and the frame 403 rate $1/\Delta t$ (as shown in Figure 4, section 3.7), Figure 9 shows the link between 404 the fraction of detected wood fluxes as a function of the dimensionless 405 parameter $PT/\Delta t$. As a note, this figure shows the numerical link between frame rate, passing time, and the fraction of detected objects, while in practice other 406 sources of uncertainty may be important as discussed in section 5. From the 407 408 Figure 9, it is nevertheless clear that frame rates less than the passing time are 409 necessary for a full census monitoring of transported wood.



Figure 9. Effect of frame rate and passing time on the fraction of detected wood pieces.

410	In addition to the fraction of detected wood pieces, the time interval can
411	affect the detection of some short events like wood pulses, defined qualitatively
412	as the delivery of large amount of wood in a short time period (on the order of
413	minutes). Figure 10.a is an example of detected pulses in the event F4 where
414	the wood flux is presented on 1 min intervals. As shown, short term pulses with
415	fluxes much higher than the hourly average are common. To check the quality
416	of detection for such short events, Figure 10.b shows one day detection of wood
417	with one pulse at 10am 3 th Jan 2012. As it is seen, the possibility to detect wood
418	pulse decreases by decreasing frame rate (from red to blue) when considering
419	the flow conditions observed on the Ain River station.



Figure 10. a) Wood fluxes as observed in 1-minute intervals. Beside short fluctuations of wood flux, pulses of wood can be defined qualitatively as the delivery of large amount of wood in a short period of time. The gray boxes show the night time when video monitoring was impossible. b) Effect of the temporal resolution on detecting short time events (a wood pulse).

420 5. Discussions and conclusions

421 5.1. The link between flow hydrographs and wood fluxes

- 422 Our observations confirm that most of the wood pieces are mobilized on the
- 423 rising limb of the hydrograph than the falling limb following MacVicar and Piégay,
- 424 (2012), Kramer and Wohl, (2014) and Ghaffarian et al. (2020a). The peak in
- 425 wood flux is generally reached before the flood peak. These observations

426 demonstrate some hysteresis of water discharge that agrees with MacVicar and 427 Piégay (2012) and Ghaffarian et al. (2020a), who state that the peak discharge 428 and the peak of wood flux do not occur simultaneously and normally wood 429 transport rate decreased before the peak of hydrograph. This result has also 430 been confirmed by the model of Ruiz-Villanueva et al. (2016a). They show that 431 wood flux increases with discharge until it attains an upper threshold or tipping 432 point and then decreases or increases much more slowly.

433 A flood hydrograph can be characterized by several peaks. We observed 434 that the second or the third peaks, even when more intense, carry lower amounts of wood (Table 2). This result agrees with Moulin and Piégay (2004) 435 who indicate that the deposited wood on channel edges from the last event 436 (such as: flood, wind and ice (Boivin et al., 2015)) is transmitted by the first 437 rising of water depth. In addition, Kramer et al. (2017) showed that the 438 439 sequence of peaks, flow discharge and the shape of hydrograph can influence the amount of wood during a flood. As it is seen in the Table 2, more than 60% 440 441 of wood pieces are carried out on the first peak and then, only 30~40% of wood 442 pieces are observed. This decrease in wood flux by increasing the peaks of the 443 flood can be related to the rate of bank erosion and by the initial conditions of 444 the channel in term of wood delivery by external drivers such as wind, ice and 445 tree mortality. The first peak of hydrograph washes most of the wood available 32

within the channel and its edges and prepared over the previous no-flood period, only depositing few wood pieces on channel edges as new recruited material from bank erosion. There is also less green wood which is recruited by a new bank erosion process in the next peaks of hydrograph because the shear stress is not as high as the one observed during the first peak along the eroded bank because channel is now wider and not yet adjusted through vegetation encroachment on the accretion area on the other side of the channel.

453 Moulin and Piégay (2004) show that the wood flux during flood events is not independent from previous floods. For comparison we can look at events 454 F5 and F6, which occur one year and two months after a big flood event 455 respectively. F5 has 5 times the wood flux of F6, which indicates that more 456 457 wood was available for F5, likely from smaller events and external drivers within the inter-flood period that introduced wood pieces that are then flushed by F5. 458 459 Therefore, wood flux can be a combination of fresh material as well as inchannel stored and newly recruited material. These internal mechanisms are 460 461 fairly important as shown by the RF results which showed T_Q is the most 462 important predictor. This agrees with Ruiz-Villanueva et al. (2016a) which 463 shows that a lot of wood material stored upstream of a dam spent some time 464 as deposited wood in the river before being delivered to the reservoir.

465 This is also potentially confirmed by observations done during and after the 33

466	exceptional wind event which played a critical role on wood delivery. In the
467	current study, for example, some of the wood transported in event F7 was likely
468	provided by W1. This result indicates that during a windy period, pieces of wood
469	are recruited into the river, but there is not enough flow velocity and depth for
470	moving these wood pieces further downstream. Later, when water depth and
471	the wetted area of the river increases, the river flow is able to begin to move
472	these wood pieces relatively easily, i.e. at thresholds less than the threshold of
473	wood motion where exceptional wind events did not occur $(0.6Q_{bf})$. Therefore,
474	while the wind is not directly related to the mobility of wood, it can decrease the
475	threshold of motion and prepare wood material to be exported during the next
476	flood playing a significant role to explain the T_Q contribution. This result is the
477	first example in which we were able to detect an effective role of a potential
478	driver within the upper catchment.

A practical recommendation that derives from this improved understanding of wood mobilization is that recording can largely be initiated strictly as a result of flow discharge, for example by setting the camera to record only when Qexceeds $0.6Q_{bf}$, which would minimize the storage needs for videos while capturing by far the largest contributions to the annual wood flux. However, the effect of wind that causes wood transport at lower discharges needs to be more deeply explored using longer time series to explain wood flux differences 34 486 between floods. This factor could be then added in the RF model if applied on487 a longer time series with more flood events.

488 **5.2. Continuous modeling of wood fluxes**

As it is described in section 3.5, a Random Forest model was used to model 489 490 wood pieces during the night, when no wood is visible in relation to three predictors derived from a continuous flow hydrograph. Figure 6 shows that 491 these three predictors and wood flux are correlated. Regarding the first 492 493 predictor Q(t), MacVicar & Piégay, (2012) and Ghaffarian et al. (2020a), both showed that the wood flux is expected to have a non-linear positive relationship 494 495 with flow discharge, which was reflected in Figure 6.a. Also, dQ/dt, as the 496 second predictor, captures the effect of variations in water discharge on wood 497 recruitment during rising (positive values) vs falling (negative values) limb. The direct link between dQ/dt and wood flux on the rising limb in Figure 6.b 498 499 suggests that increasing the water level during the rising limb of the flow 500 hydrograph can be considered as one of the key parameters on wood delivery 501 in rivers.

502 The third predictor, T_q was introduced to account for wood input processes 503 between floods. The RF modelling showed that this parameter was the most 504 important predictor of wood flux for this data set, which is surprising given the

505	primary focus on water level and rate of increase in previous work. Kramer et
506	al. (2017) do show the strong effect of time between floods on the pulses of
507	wood exported from the Slave River, Canada. Ghaffarian et al. (2020a) also
508	show that time between floods has a logarithmic relation with wood flux, which
509	is confirmed in Figure 6.c. The wood input processes are not modelled explicitly,
510	however, and greater understanding at the process scale may help to develop
511	models that are more readily adapted for different catchments. In conclusion,
512	the good performance of the three predictors (Q, T_Q and dQ/dt) RF model
513	shows that it can be used to predict the wood fluxes on the Ain River given a
514	flow hydrograph. Similar models could be developed in other catchments for
515	comparison and a more general result.

516 **5.3. Selecting an optimized frame rate**

A reduced frame rate may reduce wood detection rates so that considering frame rate and passing time is critical to optimize the wood detection. Because reducing the frame rate is a rational strategy to reduce recording costs, there is always a trade-off between the temporal resolution of video (and computer storage capacity) and the recording and post-processing costs to carefully consider. For example, at a frame rate of twice the passing time, only about 50% of the passing wood pieces are detectable (Figure 9). Figure 11, (solid

524	lines) shows the fraction of detected fluxes as a function of passing time PT
525	based on the model presented in Figure 9 for the frame rates used by three
526	different studies: (i) Kramer and Wohl, (2014) on the Slave River, Canada with
527	0.033 fps, (ii) Ghaffarian et al. (2020a) on the Isère River, France with 1 fps and
528	(iii) MacVicar and Piégay (2012) and this study on the Ain River, France with 5
529	fps. The link between PT and the fraction of detected fluxes is a function of
530	camera frame rate on each river. This function which is presented as solid lines
531	in Figure 11, is compared with the estimated passing time on each river (dashed
532	lines). For all cases an increase of PT results in an exponential increase of the
533	fraction of detected wood pieces, which is governed by Δt (Figure 9). This
534	exponential relation is a strength for the model because the fraction of detected
535	wood pieces is not so sensitive to the PT, so we do not need to select an exact
536	Δt and it can be varied in the same order of magnitude.



Figure 11. Fraction of detected woods based on passing time in different rivers. 37

Video monitoring of in-channel wood

Dashed lines show the estimated passing time on each river, green the Slave, blue, the Isère and red the Ain rivers respectively.

The three studies vary in the fraction of detected wood due to differences 537 in passing time and frame rate. As shown by Ghaffarian et al. (2020a), the 538 539 passing wood in the monitored sections in both the Ain and Isère Rivers was 540 highly concentrated within a relatively narrow band of the wetted width. As a 541 result, PT was relatively consistent, with a mode of ~ 5 and 10 s in the 542 respective channels (Figure 11, red and blue dashed lines respectively). In 543 contrast, wood on the Slave River was laterally dispersed across the channel 544 from 20 to 100 m (Kramer and Wohl, 2014) and the flow velocity on the Slave 545 River was relatively low (~1/10th the velocity in Ain and Isère Rivers). Large 546 variation in transport distance and low flow velocity both result in huge variation 547 of *PT* on this river, roughly from 30s to 120s (Figure 11, green dashed line). As a result, all wood on the Isère and Ain Rivers using the parameters as described 548 549 by Ghaffarian et al. (2020a), MacVicar and Piégay (2012) and this work. In contrast, on the Slave River the frame rate is 0.033 fps ($\Delta t = 30s$) and 30s < 100550 PT < 120s, as shown in Figure 11 (green dashed line), which indicates that not 551 552 all the wood was detectable.

553 To discuss the Slave River, it is necessary to distinguish between time-554 lapse photography and videography techniques. Although time-lapse

555 photography and video monitoring use the same approach (photos are taken per unit time), time-lapse photography is intended to subsample wood flux and 556 557 there is no expectation that all of the wood is recorded. Missing data is expected and planned for. In contrast, video capture is a method to store and 558 559 record the entire sample of wood flux and each piece of wood is expected to be 560 captured in multiple frames. The condition near $PT \cong \Delta T$, however, represents 561 a transition zone between time-lapse photography and video monitoring where 562 errors may occur. For example, at this frame rate a given piece of wood may be seen once or maybe a few times within the series of captured frames. If it 563 564 appears more than once the wood may be double-counted, particularly where visibility is poor, the wood changes in orientation or submergence, or changes 565 566 in the surface reflections and lighting can fool the operator such that they flag the same piece of wood more than once. Higher frame rates will decrease the 567 568 differences between frames and the likelihood of double-counting along with it. Lower frame rates will remove the possibility of double counting and ensure that 569 the monitoring captures only a sub-set of the passing wood, which can then be 570 571 corrected for missing data as is expected for this technique. The Slave River study is thus within the transition zone from time-lapse photography to video 572 573 monitoring where double-counting remains a possibility.

574 As a further practical recommendation, it is important to select an 39

appropriate frame rate for the camera based on a good estimate of velocity 575 conditions in space and in time. Moreover, if the pattern of pulses or the source 576 577 of wood pieces is important, the frame rate should be large enough to 578 continuously detect wood pieces, while if there is a limitation on storage or long-579 term data is needed it is recommended to decrease frame rate and adopt a 580 time-lapse subsampling strategy but considering a continuous acquisition 581 window of at least 30 minutes as shown by Ghaffarian et al. (2020a) to minimize uncertainties in wood frequency estimate due to short term pulses. 582

583 **5.4. Wood pulses**

584 During our observations, it is seen that in some cases the wood flux is 585 mobilized in a sharp pulse, which is typically accompanied by some large pieces of wood that may be recent tree falls or a jam suddenly mobilized. The clarity 586 of these pulses in the video monitoring technique directly relates to the temporal 587 resolution of the camera (Figure 10.b). Moreover, such pulses are fully 588 detectable only if continuous monitoring approach is applied. The difference 589 between continuous monitoring and sampling is visible in Figure 5 where the 590 591 blue scatters show more consistency through each day, which likely is due to 592 the continuous sampling method (samples were the total per hour rather than 593 15 min multiplied by 4 as for the red scatters).

594 It can be hypothesized that wood pulses are the result of either localized or distant wood delivery. Presumably, in such cases of local mobilization, the 595 596 source of wood could be close to the camera and so the wood would be tightly grouped in time. Such pulses are observed at the Ain River location, where 597 598 large wood with visible leaves and root wads are followed by large numbers of 599 smaller pieces of wood. In this case the pulse was attributed to local bank 600 erosion or the sudden mobilization of a wood jam. In contrast, the source of 601 wood could be far upstream from the camera, which dispersed wood in 602 transport tending to agglomerate over longer distances. Such a process of 603 'rafting' or 'clumping' has been observed in the lab and field (Braudrick et al., 1997; Kramer et al., 2017). Therefore, due to the dissipation, the wood pulse 604 605 spreads out during transport in long distances. The pulses at the camera location would therefore be very spread out and come more or less regularly, 606 607 which could mean that the inputs are random or that the distribution has been randomized by dissipation during transport. On the falling limb, despite the bank 608 erosion due to the decrease in the soil pore pressure, the flow might not be 609 610 enough to transport this wood. Also, some wood pieces have already been 611 deposited in the highest possible locations with other wood jams on the 612 upstream floodplain (Ruiz-Villanueva et al., 2016b; Wohl et al., 2018). A careful 613 analysis of wood flux pattern thus provides some potentially key insights about

- 614 the processes that prepare the stock of available wood within a reach and
- 615 should by the subject of further investigation.

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623 Data availability statement

624 Research data not shared.

625 **References**

Abbe TB, Montgomery DR. 2003. Patterns and processes of wood debris
accumulation in the Queets river basin, Washington. *Geomorphology* 51 (1-3):
81–107.

Ali I, Tougne L. 2009. Unsupervised Video Analysis for Counting of Wood
in River during Floods. *In Advances in Visual Computing*, Bebis G et al. (eds).
Springer Berlin Heidelberg: Berlin, Heidelberg; 578–587.

632 Belgiu M, Drăguţ L. 2016. Random forest in remote sensing: A review of 633 applications and future directions. *ISPRS Journal of Photogrammetry and* 634 *Remote Sensing* **114** : 24–31.

635 Benacchio V, Piégay H, Buffin-Bélanger T, Vaudor L. 2017. A new 636 methodology for monitoring wood fluxes in rivers using a ground camera: 637 Potential and limits. *Geomorphology* **279** : 44–58.

Benda L, Miller D, Sias J, Martin D, Bilby R, Veldhuisen C, Dunne T. 2003. *In The Ecology and Management of Wood in World Rivers*, SV Gregory, KL
Boyer, AM Gurnell (eds). Symposium 37. American Fisheries
Society: Bethesda, MD; 49– 73.

Boivin M, Buffin-Bélanger T, Piégay H. 2015. The raft of the Saint-Jean
River, Gaspé (Québec, Canada): A dynamic feature trapping most of the wood
transported from the catchment. *Geomorphology* 231 : 270–280.

645 Braudrick CA, Grant GE, Ishikawa Y, Ikeda H. 1997. Dynamics of wood 646 transport in streams: a flume experiment. *Earth Surface Processes and* 647 *Landforms*: The Journal of the British Geomorphological Group **22** : 669–683.

648 Breiman L. 2001. Random Forests. Machine Learning **45** : 5–32. 649 Doi:10.1023/A:1010933404324

Comiti F, Andreoli A, Mao L, Lenzi MA. 2008. Wood storage in three
mountain streams of the Southern Andes and its hydro-morphological effects. *Earth Surface Processes and Landforms* 33 (2): 244–262.

Comiti F, D'Agostino V, Moser M, Lenzi MA, Bettella F, Dell'Agnese A,
Rigon E, Gius S, Mazzorana B. 2012. Preventing wood-related hazards in
mountain basins: from wood load estimation to designing retention structures. *Proceedings, 12th Congress INTERPRAEVENT 2012,* Grenoble, France; 651–
662.

658 Comiti F, Lucía A, Rickenmann D. 2016. Large wood recruitment and 659 transport during large floods: A review. *Geomorphology* **269** : 23–39.

660 De Cicco PN, Paris E, Ruiz-Villanueva V, Solari L, Stoffel M. 2018. In-661 channel wood-related hazards at bridges: A review. *River Research and* 662 *Applications* **34** : 617–628.

Diehl TH. 1997. Potential Drift Accumulation at Bridges. Publication No.
FHWA - RD - 97 - 028. US Department of Transportation, Federal Highway
Administration, Research and Development, Turner - Fairbank Highway
Research Center: McLean, Virginia, USA.
http://www.tn.water.usgs.gov/pubs/FHWA-RD-97-028/drfront1.htm

Fischer M. 2006. Driftwood During the Flooding in Klosters in 2005, Report,
 HSW Wädenswil, Switzerland (in German).

670 http://www.wsl.ch/fe/walddynamik/projekte/schwemmholzablagerungen/index 671 _EN

Ghaffarian H, Lopez D, Mignot E, Piégay H, Rivière N. 2020b. Dynamics of
floating objects at high particulate Reynolds numbers. *Physical Review Fluids*5 (5): 054307.

675 Ghaffarian H, Piegay H, Lopez D, Rivière N, MacVicar B, Antonio A, Mignot 676 E. 2020a. Video-monitoring of wood discharge: first inter-basin comparison and 677 recommendations to install video cameras. *Earth Surface Processes and* 678 *Landforms*. Available from: 679 https://oplinelibrary.wiley.com/doi/abs/10.1002/opp.4875

679 https://onlinelibrary.wiley.com/doi/abs/10.1002/esp.4875

680 Gippel CJ. 1995. Potential of turbidity monitoring for measuring the 681 transport of suspended solids in streams. *Hydrological Processes* **9** (1): 83–97.

682 Gurnell A. 2012. Wood and river landscapes. *Nature Geoscience* **5** (2): 93– 683 94.

684 Gurnell A, Petts G. 2006. Trees as riparian engineers: the Tagliamento river, 685 Italy. *Earth Surface Processes and Landforms* **31** (12): 1558–1574.

686 Gurnell AM, Piégay H, Swanson FJ, Gregory SV. 2002. Large wood and 687 fluvial processes. *Freshwater Biology* **47** (4): 601–619.

Hastie T, Tibshirani R, Friedman J. 2009. The elements of statistical
learning: data mining, inference, and prediction. Springer Science & Business
Media

Hutengs C, Vohland M. 2016. Downscaling land surface temperatures at
regional scales with random forest regression. *Remote Sensing of Environment* **178**: 127–141.

Keller EA, Swanson FJ. 1979. Effects of large organic material on channel
form and fluvial processes. *Earth Surface Processes* 4 (4): 361–380.

Kramer N, Wohl E. 2014. Estimating fluvial wood discharge using timelapse photography with varying sampling intervals. *Earth Surface Processes and Landforms* **39** (6): 844–852.

Kramer N, Wohl E. 2017. Rules of the road: A qualitative and quantitative
synthesis of large wood transport through drainage networks. *Geomorphology* **279**: 74–97.

Kramer N, Wohl E, Hess-Homeier B, Leisz S. 2017. The pulse of driftwood
export from a very large forested river basin over multiple time scales, Slave
River, Canada. *Water Resources Research* 53 (3): 1928–1947.

Lassettre NS, Kondolf GM. 2012. Large woody debris in urban stream
channels: redefining the problem. *River Research and Applications* 28 (9):
1477–1487.

Lassettre NS, Piégay H, Dufour S, Rollet A-J. 2008. Decadal changes in
distribution and frequency of wood in a free meandering river, the Ain River,
France. *Earth Surface Processes and Landforms* 33 (7): 1098–1112.

Lemaire P, Piegay H, MacVicar B, Mouquet-Noppe C, Tougne L. 2014.
Automatically monitoring driftwood in large rivers: preliminary results. *Presented at the 2014 AGU Fall Meeting*. 19 December, San Francisco, USA.
Available from: https://agu.confex.com/agu/fm14/meetingapp.cgi/Paper/22487

Liaw A, Wiener M. 2002. Classification and Regression by RandomForest.
2:5.

Lyn D, Cooper T, Yi Y-K. 2003. Debris accumulation at bridge crossings:
laboratory and field studies. Purdue University: West Lafayette. Available from:
http://docs.lib.purdue.edu/jtrp/48

MacVicar B, Piégay H. 2012. Implementation and validation of video monitoring for wood budgeting in a wandering piedmont river, the Ain River (France). *Earth Surface Processes and Landforms* **37** (12): 1272–1289.

MacVicar B, Hauet A, Bergeron N, Tougne L, Ali I. 2012. River Monitoring
with Ground-based Videography. *In Fluvial Remote Sensing for Science and Management*, PE Carbonneau, H Piégay (eds). Wiley: Chichester,
UK; 367–383.

MacVicar BJ, Piégay H, Henderson A, Comiti F, Oberlin C, Pecorari E. 2009.
Quantifying the temporal dynamics of wood in large rivers: field trials of wood
surveying, dating, tracking, and monitoring techniques. *Earth Surface Processes and Landforms* 34 (15): 2031–2046.

Mao L, Burns S, Comiti F, Andreoli A, Urciuolo A, Gaviño-Novillo M,
Iturraspe R, Aristide Lenzi M. 2008. Acumulaciones de detritos leñosos en un
cauce de montaña de Tierra del Fuego: análisis de la movilidad y de los efectos
hidromorfológicos. Bosque (Valdivia) 29 : 197–211.

Marcus WA, Rasmussen J, Fonstad MA. 2011. Response of the Fluvial
Wood System to Fire and Floods in Northern Yellowstone. *Annals of the Association of American Geographers* **101** (1): 21–44.

Martin DJ, Benda LE. 2001. Patterns of Instream Wood Recruitment and
Transport at the Watershed Scale. *Transactions of the American Fisheries*Society 130 (5): 940–958.

Mazzorana B, Ruiz-Villanueva V, Marchi L, Cavalli M, Gems B, Gschnitzer
T, Mao L, Iroumé A, Valdebenito G. 2018. Assessing and mitigating large woodrelated hazards in mountain streams. *Journal of Flood Risk Management* 11
(2): 207–222.

Mazzorana B, Zischg AP, Largiader A, Hübl J. 2009. Hazard index maps
for woody material recruitment and transport in alpine catchments. *Natural Hazards and Earth System Sciences* 9 (1): 197–209.

Montgomery DR, Abbe TB, Buffington JM, Peterson NP, Schmidt KM,
Stock JD. 1996. Distribution of bedrock and alluvial channels in forested
mountain drainage basins. *Nature* 381 : 587–589.

Moulin B, Piegay H. 2004. Characteristics and temporal variability of large
woody debris trapped in a reservoir on the River Rhone(Rhone): implications
for river basin management. *River Research and Applications* 20 (1): 79–97.

Muste M, Fujita I, Hauet A. 2008. Large-scale particle image velocimetry for measurements in riverine environments. *Water resources research* **44** (4): W00D19.

Nakamura F, Swanson FJ. 1993. Effects of coarse woody debris on
morphology and sediment storage of a mountain stream system in western
Oregon. *Earth Surface Processes and Landforms* 18 (1): 43–61.

Piégay H, Moulin B, Hupp CR. 2017. Assessment of transfer patterns and
origins of in-channel wood in large rivers using repeated field surveys and wood
characterisation (the Isère River upstream of Pontcharra, France). *Geomorphology* 279 : 27–43.

Ravazzolo D, Mao L, Picco L, Lenzi MA. 2015. Tracking log displacement
during floods in the Tagliamento River using RFID and GPS tracker devices. *Geomorphology* 228 : 226–233.

767 Ruiz-Villanueva V, Bladé Castellet E, Díez-Herrero A, Bodoque JM, 46 Sánchez-Juny M. 2014a. Two-dimensional modelling of large wood transport
during flash floods. *Earth Surface Processes and Landforms* **39** (4): 438–449.

Ruiz-Villanueva V, Bürkli L, Mazzorana B, Mao L, Ravazzolo D, Iribarren P,
Wohl E, Nakamura F, Stoffel M. 2018. Defining and characterizing wood-laden
flows in rivers using home videos. In *E3S Web of Conference*, vol.40, p.02014.
EDP Sciences.

Ruiz-Villanueva V, Piégay H, Gurnell AM, Marston RA, Stoffel M. 2016a.
Recent advances quantifying the large wood dynamics in river basins: New
methods and remaining challenges: Large Wood Dynamics. *Reviews of Geophysics* 54 (3): 611–652.

Ruiz-Villanueva V, Stoffel M, Piégay H, Gaertner V, Perret F. 2014b. Wood
density assessment to improve understanding of large wood buoyancy in rivers. *River Flow 2014–Schleiss et Al.(Eds)* pp. 2503–2508.

Ruiz-Villanueva V, Wyżga B, Mikuś P, Hajdukiewicz H, Stoffel M. 2016b.
The role of flood hydrograph in the remobilization of large wood in a wide
mountain river. *Journal of Hydrology* 541 : 330–343.

Schenk ER, Moulin B, Hupp CR, Richter JM. 2014. Large wood budget and
transport dynamics on a large river using radio telemetry. *Earth Surface Processes and Landforms* **39** (4): 487–498.

787 Senter A, Pasternack G, Piégay H, Vaughan M. 2017. Wood export
788 prediction at the watershed scale. *Earth Surface Processes and Landforms* 42
789 (14): 2377–2392.

Seo JI, Nakamura F. 2009. Scale-dependent controls upon the fluvial
export of large wood from river catchments. *Earth Surface Processes and Landforms* 34 (6): 786–800.

Seo JI, Nakamura F, Nakano D, Ichiyanagi H, Chun KW. 2008. Factors
controlling the fluvial export of large woody debris, and its contribution to
organic carbon budgets at watershed scales. *Water Resources Research* 44
(4): W04428

Shields FD, Gippel CJ. 1995. Prediction of Effects of Woody Debris
Removal on Flow Resistance. *Journal of Hydraulic Engineering* **121** (4): 341–
354.

800 Team RC. 2019. R: A language and environment for statistical computing 47 801 (version 3.1. 2). Vienna, Austria.

Turowski JM, Badoux A, Bunte K, Rickli C, Federspiel N, Jochner M. 2013.
The mass distribution of coarse particulate organic matter exported from an
alpine headwater stream. *Earth Surface Dynamics Discussions* 1 (1): 1–29.

Waldner P, Rickli C, Köchlin D, Usbeck T, Schmocker L, Sutter F. 2007.
Schwemmholz. Ereignisanalyse Hochwasser 2005–Teil 1: Prozes se, Schäden
und erste Einordnung (in German). Bundesamt für Umwelt BAFU, *Eidgenössische Fors chungsanstalt WSL*. Bezzola GR, Hegg C. UmweltWissen 825 : 181–193.

Wilcox AC, Wohl EE. 2006. Flow resistance dynamics in step-pool stream
channels: 1. Large woody debris and controls on total resistance. *Water Resources Research* 42 (5):W05418.

Wohl E. 2013. Floodplains and wood. *Earth-Science Reviews* **123** : 194– 212.

Wohl E, Cadol D, Pfeiffer A, Jackson K, Laurel D. 2018. Distribution of large
wood within river corridors in relation to flow regime in the semiarid western US. *Water Resources Research* 54 (3): 1890–1904.

Young WJ. 1991. Flume study of the hydraulic effects of large woody debris
in lowland rivers. *Regulated Rivers: Research & Management* 6 (3): 203–211.

Yue S, Ouarda TBMJ, Bobée B, Legendre P, Bruneau P. 1999. The
Gumbel mixed model for flood frequency analysis. *Journal of Hydrology* 226 (1):
88–100.