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1 **Behavior of critical metals in metamorphosed Pb-Zn ore**
2 **deposits: example from the Pyrenean Axial Zone**

3

4 Alexandre Cugerone¹, Bénédicte Cenki-Tok¹², Manuel Muñoz¹, Kalin Kouzmanov³, Emilien
5 Oliot¹, Vincent Motto-Ros⁴, Elisabeth Le Goff⁵

6 ¹*Géosciences Montpellier, Université de Montpellier, CNRS, France;*

7 alex.cugerone@gmail.com; benedicte.cenki-tok@umontpellier.fr ;

8 manuel.munoz@umontpellier.fr ; emilien.oliot@umontpellier.fr.

9 ²*Earthbyte Research Group, School of Geosciences, University of Sydney, Sydney, NSW 2006,
10 Australia.*

11 ³*Department of Earth sciences, University of Geneva, Switzerland;*

12 Kalin.Kouzmanov@unige.ch

13 ⁴*Institut Lumière Matière Université Lyon 1, France; vincent.motto-ros@univ-lyon1.fr*

14 ⁵*BRGM, Bureau de Recherches Géologiques et Minières, Montpellier, France ;*

15 e.legoff@brgm.fr

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17 *Accepted Mineralium Deposita*

18

19 **Abstract**

20 Rare metals (Ge, Ga, In, Cd) are key resources for the development of green
21 technologies and are commonly found as trace elements in base-metal mineral deposits. Many
22 of these deposits are in orogenic belts and the impact of recrystallization on rare metal content
23 and distribution in sphalerite needs to be evaluated. Based on laser ablation inductively coupled
24 plasma mass spectrometry (LA-ICP-MS) analyses, and micro-imaging techniques such as laser-
25 induced breakdown spectroscopy (LIBS) and electron backscattered diffraction (EBSD), we
26 investigate the minor and trace element composition related to sphalerite texture for three types
27 of mineralization from the Pyrenean Axial Zone (PAZ). Vein mineralization (Type 2b) appears
28 significantly enriched in Ge and Ga compared to disseminated and stratabound mineralization
29 (Type 1 and Type 2a, respectively). In vein mineralization, the partial recrystallization induced
30 by deformation led to the remobilization of Ge, Ga, and Cu from the sphalerite crystal lattice
31 into accessory minerals. We propose that the association of intra-granular diffusion, and fluid-
32 rock reaction, were likely responsible for the formation of patchy-oscillatory zoning in
33 sphalerite, and the crystallization of Ge-rich accessory minerals. Chemical and textural
34 heterogeneity is common in sphalerite from various world-class deposits and a full
35 understanding of these heterogeneities is now crucial to assess the rare metal potential, and
36 associated extraction processes of deformed base-metal ores.

37 **Introduction**

38 Sphalerite (ZnS) commonly contains high amounts (up to a few 1000s of ppm) of rare
39 metals such as germanium (Ge), indium (In), gallium (Ga) or cadmium (Cd), which are
40 currently extracted as by-products from zinc concentrates (Frenzel et al. 2014, 2016, 2017;
41 Licht et al. 2015) and considered critical for the economy (European Commission 2017; U.S.

42 Geological Survey 2019). These high concentrations of trace elements are generally related to
43 undeformed sphalerite in the form of coarse crystals originating from open-space crystallization
44 or from colloform banded precipitates with fine-grained textures (Cerny and Schroll 1995;
45 Goffin et al. 2015; Henjes-Kunst et al. 2017; Cugeronne et al. 2018a). Undeformed sphalerite
46 occurs mainly in Mississippi Valley Type (MVT) and polymetallic vein deposits (Viets et al.
47 1992; Saini-Eidukat et al. 2009; Murakami and Ishihara 2013; Belissont et al. 2014; Bonnet
48 2014; Bauer et al. 2018). However, the largest sphalerite deposits are found in orogenic domain
49 (Lawrence 1973; Moore et al. 1986; Kelley and Jennings 2004; Kampunzu et al. 2009;
50 Wilkinson 2013; Gibson et al. 2017). This reflects the fact that the ores originally formed as
51 Sediment-Hosted Massive Sulfide (SHMS), Volcanic-Hosted Massive Sulfide (VHMS) or
52 carbonate-hosted deposits (Kipushi-type, more rarely MVT) are more likely to have undergone
53 metamorphism and recrystallization during an orogenic event (Bernstein and Cox 1986; Hughes
54 1987; Murphy 2004; Wagner and Monecke 2005; Huston et al. 2006; Melcher et al. 2006;
55 Monteiro et al. 2006; Kamona and Friedrich 2007; Reiser et al. 2011; Ye et al. 2011).

56 In metamorphic environments, remobilization of minor and trace elements related to
57 dynamic recrystallization, annealing, or even brecciation, has been studied for various sulfide
58 minerals. The mobility of Au, As, Ni, Co, Pb, Bi, and Se has been examined in pyrite (Cook et
59 al. 2009a; Large et al. 2009; Reddy and Hough 2013; Velásquez et al. 2014; Dubosq et al. 2018;
60 Kampmann et al. 2018), arsenopyrite-löllingite (Tomkins and Mavrogenes 2001; Wagner et al.
61 2007; Fougerouse et al. 2016) and pyrrhotite (Vukmanovic et al. 2014). The dynamic
62 recrystallization assisted by metamorphic fluids in gold systems releases invisible gold from
63 the sulfide crystal lattice and subsequent crystallization of metallic gold (Fougerouse et al.
64 2016; Dubosq et al. 2018). Lockington et al. (2014) studied sphalerite mineralization from
65 metamorphosed VHMS and SHMS deposits. Metamorphic recrystallization was shown to lead
66 to remobilization of minor and trace elements resulting, in a loss of compositional zoning, and

homogenization of Cu distribution in the crystal lattice of sphalerite, possibly associated to a depletion in Pb, Bi, Ag and an enrichment in Fe, Cd, Mn, In. These data were acquired in various deposits of different metamorphic grade and the role of primary enrichment for these elements cannot be clearly established. In Neves-Corvos (Iberian Pyrite Belt, Carvalho et al. 2018), indium is primarily incorporated in high-temperature sphalerite and chalcopyrite, but was not lost during recrystallization associated to hydrothermal re-working and/or low-grade tectonometamorphic overprints (Relvas et al. 2006; Carvalho et al. 2018). Nonetheless, roquesite (CuInS_2), locally occurs at sulfide grain boundaries. In sphalerite from Neves Corvo, Ge is reported below ~ 100 ppm Ge (Carvalho et al. 2018). In the Barrigão deposit (Iberian Pyrite Belt), Reiser et al. (2011) and Belissont et al. (2019) suggested late-stage remobilization of germanium associated to low-grade metamorphism (260-300 °C) with germanium enrichments in chalcopyrite and accessory phases localized in fractures. Cugeronne et al. (2020) showed evidence for redistribution of Ge in partly recrystallized sphalerite from the Arre vein deposit (Pyrenean Axial Zone). Recrystallized sphalerite is Ge-depleted but correspond to textures where Ge-minerals occur, compared to dark domains in coarser grains which still contain up to ~ 600 ppm Ge in the sphalerite lattice. Occurrences of rare metals such as Ge, Ga, In, and Cd within deformed/metamorphosed sphalerite ore remains poorly documented and their behavior in such systems needs to be evaluated.

Sphalerite is ductile at relatively low-temperature in various experimental deformation studies (< 400 °C; Clark and Kelly, 1973; Siemes and Borges, 1979; Couderc et al. 1985; Cox, 1987). No experimental study on the impact of sphalerite deformation on trace elements behavior is available.

Currently, the production of rare metals from base-metal deposits remains challenging. Germanium is generally extracted, for example, by hydrometallurgy followed by various mass transfer techniques (Ruiz et al. 2018). However, only a few percent of the total rare metal

92 content in sphalerite (at ppm level) is recovered due to a negative effect which prevents the
93 extraction of zinc (around ~3 % of the Ge extracted; Ruiz et al. 2018; U.S. Geological Survey
94 2019). Chemical extraction from highly concentrated phases (with Ge contents > 1 wt.%),
95 previously separated by magnetic or mechanic techniques, could prove to be much more
96 profitable (Cugeron et al. 2019).

97 This study aims to decipher the behavior of rare metals, mainly Ge, Ga, In, Cd, in sphalerite
98 from different Pb-Zn deposits and mineralization styles of the Pyrenean Axial Zone (PAZ). The
99 Pb-Zn deposits of the PAZ contain sphalerite rich in rare metal (ppm level) and Ge-rich
100 accessory minerals, such as brunogeyerite GeFe_2O_4 , briartite $\text{GeCu}_2(\text{Fe}, \text{Zn})\text{S}_4$, carboirite
101 $\text{GeFeAl}_2\text{O}_5(\text{OH})_2$ or argutite GeO_2 (Laforet et al. 1981; Johan et al. 1983; Bernstein 1985; Johan
102 and Oudin 1986; Pouit and Bois 1986; Julliot et al. 1987; Cugeron et al. 2018a). We have
103 investigated the textural and chemical composition of different types of sphalerite
104 mineralization in the Pyrenees using an in-situ quantitative analysis by LA-ICP-MS and micro-
105 imaging techniques such as EBSD and LIBS mapping. We show a detailed genetic model at
106 vein and granular scales for the partitioning of Ge and associated trace elements between
107 sphalerite and accessory minerals. Finally, we present a comparison with other Ge-enriched Pb-
108 Zn deposits worldwide.

109 **Sampling and analytical methods**

110 Oriented ore samples were collected from underground mines or surface outcrops. We
111 have studied polished sections (either 30 or 150 μm -thick) from six Pb-Zn deposits of three
112 mineralization types.

113 EBSD analysis is used to characterize the texture and the grain size of the sphalerite.
114 This technique requires a well-polished surface. In order to eliminate surface defects, a last step

115 of polishing ($0.25\text{ }\mu\text{m}$) was performed using a Vibromet polisher with 150 g of pressure on a
116 polyurethane layer with colloidal silica. During this process, cleaning the frozen colloidal silica
117 was necessary to avoid scratches on sphalerite. EBSD maps were performed with a Camscan
118 Crystal Probe X500FE SEM-EBSD at Geosciences Montpellier (CNRS-University of
119 Montpellier, France). Operating conditions were 20 kV and 5 nA, with a working distance of
120 25 mm under 2 Pa low vacuum. Sample surfaces were positioned horizontally at 20° to the
121 incident electron beam to improve the collection of the backscattered electrons. Element
122 distribution maps and back-scattered electron (BSE) images were simultaneously acquired by
123 measuring X-ray fluorescence with an EDS (energy dispersive spectrometer). The step size for
124 2D imaging was below $5\text{ }\mu\text{m}$. EBSD mapping were preferentially performed on sphalerite,
125 indexed as a cubic crystal symmetry, while Ge-minerals were localized by EDS maps. The
126 EBSD indexation of quartz, calcite, galena and pyrite was also performed. Match units used for
127 sphalerite indexation were derived from published crystallographic data from Villars and
128 Calvert (1991). The lower misorientation limit for grain boundary indexation is fixed at 20°
129 following careful investigation on misorientation boundary maps superposed on orientation
130 maps. Oxford Instruments softwares AZtec and Channel 5 were used to generate maps using
131 the EDS (chemical maps) and EBSD (structural maps).

132 Electron probe micro-analyses were carried out using a Cameca SX100 at the Service
133 Inter-Régional Microsonde-Sud (Montpellier, France). Major, minor and trace elements were
134 measured with a beam current of 100 nA and accelerating voltage of 20 kV. The analytical
135 procedure and detailed settings for each elemental analysis is detailed in Cugeron et al.
136 (2018a). Zinc concentrations of sphalerite crystals were used as an internal calibration for LA-
137 ICP-MS analyses.

138 LA-ICP-MS was used to determine minor and trace elements concentrations in
139 sphalerite. Type 1 sphalerite was analyzed in one sample from Bentaillou ($n=7$), Type 2a

140 sphalerite in three samples from Bentaillou (n=12), Margalida (n=9) and Victoria (n=11).
141 Type 2b sphalerite was measured in three samples from Arre (n=36), Anglas (n=40) and Pale
142 Bidau (n=30).

143 Analyses were carried out using an Excimer CompEx 102 coupled to a ThermoFinnigan
144 Element XR at the OSU-OREME AETE platform (University of Montpellier, France). Laser
145 ablation was performed using a constant 5 Hz pulse rate at 140 mJ with a spot diameter of
146 26 µm. Each analysis comprises 180 s of background measurement and 60 s of sample ablation
147 (signal measurement), followed by a 60 s retention time to ensure a proper cell washout. Data
148 were processed using the Glitter 4.0 software package (Van Achterbergh et al. 2001). The
149 following isotopes were measured: ^{29}Si , ^{34}S , ^{55}Mn , ^{57}Fe , ^{59}Co , ^{61}Ni , ^{63}Cu , ^{64}Zn , ^{69}Ga , ^{74}Ge , ^{75}As ,
150 ^{77}Se , ^{95}Mo , ^{105}Pd , ^{107}Ag , ^{111}Cd , ^{115}In , ^{118}Sn , ^{121}Sb , and ^{208}Pb . MASS-1 reference material
151 (Wilson et al. 2002) was used as an external standard with a corrected 57 ± 1.75 ppm value for
152 Ge (Dr. Stephen Wilson, personal communication). NIST SRM 610 (Pearce et al. 1997) was
153 used as secondary external standard to identify possible instrumental drift. Concentrations for
154 Ni, As, Mo, Se, Pd and Si were below detection limits in all measurements. Several time-
155 integrated LA-ICP-MS spectra are reported in the electronical supplementary material
156 (ESM 1), illustrating the time intervals used for the measurements. One standard deviation is
157 systematically indicated for median values.

158 Laser Induced breakdown spectroscopy (LIBS) imaging technique was used to highlight
159 the distribution of Ge, Cu, Ga, Zn, Si, Fe, Cd, Al, Mg, and Ti, in a highly zoned sphalerite
160 sample from Arre (ARR03 sample) to map metallic elements at a micrometric spatial resolution,
161 and with a sensitivity at the ppm level (Sancey et al. 2014; Cáceres et al. 2017; Fabre et al.
162 2018). During the analysis, the sample surface is scanned, through single laser pulses, in a pixel-
163 by-pixel manner to induce the breakdown of the material. The light radiation emitted by the
164 plasma is then collected by an optical system and analyzed using a spectrometer. The sample

165 was excited using Nd:YAG laser with a pulse energy of about 600 µJ operating at 100 Hz, and
166 a lateral resolution (i.e., distance between two consecutive laser shots) of 13 µm. The surface
167 analyzed is about 6 cm² (3.6 megapixels) and was measured in 6 h. The spectrometer was
168 configured in the spectral range 250-330 nm to detect Ge (265.1 nm), Cu (324.7 nm) and Ga
169 (294.4 nm). Due to an interference with Al, the Ge line at 269.1 nm was used in the
170 aluminosilicate phases.

171 Geological setting

172 The Pb-Zn ore deposits are hosted in the Variscan PAZ (Fig. 1a), exhumed during the
173 Paleogene collision between the Iberia and Eurasian plates (Zwart 1979; Carreras and Druguet
174 2014). The PAZ is composed of Paleozoic metasediments intruded by Ordovician and Late-
175 Variscan granite (Fig. 1b; Kleinsmiede, 1960; Zwart, 1963; Denèle et al. 2014). The Pb-Zn
176 deposits are concentrated in pluri-kilometric structures such as the Bossost or Pierrefitte
177 anticlinoria (Fig. 1b).

178 The Pb-Zn district of the PAZ was exploited in 20th century. It has produced *ca.* 400 kt of
179 Zn and 180 kt of Pb (BRGM 1984; Ovejero Zappino 1991). Three types of Pb-Zn
180 mineralization are described in Cugeron et al. (2018b). Type 1 is minor, disseminated and
181 laminated mineralization, probably associated to sedimentation during Cambro-Ordovician and
182 Devonian times (Fig. 1c). Type 2a, is a structurally controlled mineralization parallel to S₀-S₁
183 (Fig. 1c), synchronous to Variscan D₁/M₁ regional deformation/metamorphism related to
184 crustal thickening (Zwart 1963, 1979; Mezger et al. 2004; Laumonier et al. 2010) with a peak
185 estimated at 580 °C and 0.55 GPa in the Bossost dome (Mezger et al. 2004). Type 2a
186 mineralization pre-dates Variscan M₂ contact metamorphism, culminating at *ca.* 525 °C and
187 0.2 GPa (Mezger and Passchier 2003; Mezger et al. 2004; de Hoÿm de Marien et al. 2019) in
188 the Victoria-Margalida area. In the field, Type 2a mineralization appears in metric pull-apart

189 type structures in Cambrian marble as in Bentaillou (Fig. 1c) or folded by D₂ deformation in
190 Late-Ordovician metasediments in Victoria (Fig. 1c). It can also be associated to fault damaged
191 zones in Late-Ordovician marble in Margalida (Fig. 1b; Cugeron et al. 2018b).

192 Type 2b (Figure 1) is epigenetic vein mineralization probably late, or post-dating, the D₂/M₂
193 deformation event (Cugeron et al. 2018b). Pb-Zn mineralization is structurally controlled and
194 occurs as vertical decimetric-sized veins parallel to S₂ cleavage, which locally intersect Type 2a
195 mineralization. No alteration halo has been observed along the contact with the host-rock.
196 Type 2b mineralization cuts Late-Variscan aplitic intrusions, and is affected by a cleavage
197 parallel to the Late-Variscan S₂ cleavage in the host rock (Cugeron et al. 2018b). We have
198 studied samples from Type 2b Pale Bidau, Arre and Anglas deposits hosted in calc-schists
199 (Fig. 2b). A late D₃ deformation is reported in the literature (Carreras and Druguet 2014;
200 Cochelin et al. 2017) and corresponds to faults like the Bossost mylonitic fault close to Type 2a
201 Margalida deposit.

202 Results

203 Mineralization types and textures

204 The three types of mineralization are mainly composed of sphalerite and galena, with
205 minor pyrite, chalcopyrite, arsenopyrite and gangue of quartz-carbonate (Fig. 2a).

206 Type 1 is composed of micrometric-size disseminated sulfide crystals. This
207 mineralization is crosscut by Type 2a veins (Bentaillou deposit; Cugeron et al. 2018b). It
208 consists of pluri-micrometric euhedral dark brown crystals.

209 Type 2a sphalerite shows undeformed millimeter-size twinned crystals and polygonal
210 shape that is dark-red in polished slabs (Fig. 2b). The texture is different at Margalida (Fig. 2c),
211 which presents micrometer-scale dark-brown sphalerite grains and forming the matrix of a

212 durchbewegung breccia composed of quartz, calcite and schist clasts. Sphalerite is folded in
213 Victoria (Fig. 2d) and cross-cuts metamorphic centimeter-scale gahnite (Fig. 3; Cugerone et al.
214 2018a). Sphalerite exhibits recrystallized textures in F₂ fold-hinges (Fig. 3).

215 The mineral assemblages of Types 2a and 2b veins (Cugerone et al. 2018b) are distinct
216 with the presence of apatite, ilmenite, tourmaline in Type 2a and Ge-minerals, graphite, Mg-
217 Fe-Mn-Zn carbonates in Type 2b. Sphalerite crystals of Type 1, 2a and 2b mineralizations do
218 not contain chalcopyrite inclusions.

219 Type 2b sphalerite is light to dark brown in transmitted-light and with heterogeneous
220 grain-size resulting from a superimposed cleavage (considered to be S₂) parallel to the vein
221 contact (Figs. 2e and 2f). In the Arre, Anglas and Pale Bidau deposits, sphalerite crystals exhibit
222 various colors from dark-brown zones in coarse, millimetric grains, to light-brown zones in
223 coarse millimetric, or fine micrometric, grain fractions (Figs. 2e and 3f). Quartz and carbonate
224 show comb textures with the occurrence of cockade breccia mainly in Arre and Anglas Type 2b
225 veins. EBSD-grain reference orientation deviation (GROD) maps and transmitted-light
226 microphotographs are superposed in the Figure 4. Coarse grains (> 100 µm) and small
227 recrystallized grains (< 100 µm) are colored in shaded of dark and light blue, respectively. Blue
228 colors are more pronounced where internal deformation is more intense. Three types of
229 sphalerite textures appear in Type 2b mineralization: coarse grains with dark-brown domains,
230 coarse grains with light-brown domains and recrystallized light-brown crystals. Sub-vertical
231 fine recrystallized fraction domains appear mostly in white color in the GROD maps (Fig. 4a)
232 which attests to poor internal deformation and commonly mark the cleavage, for example in PB
233 (Figs. 4a and 4b). Coarse crystals (> 100 µm) with color zones from light to dark-brown are
234 reported in the Arre and Anglas samples (Figs. 4c-f) and commonly contain ductile deformation
235 attested by intense dark-blue color mainly close to grain boundaries (Fig. 4e). In the EBSD-
236 GROD and the grain size maps, it is possible to discriminate between recrystallized and light-

237 brown coarse grains while these two fractions show the same light-orange color in transmitted-
238 light (Fig. 4). Sphalerite from the Pale Bidau deposit exhibits blurred color domains with
239 progressive variations from light to dark brownish tints and no intra-granular color zonation is
240 observed, especially in coarse grains.

241 Minerals in which Ge is a major element have been identified by optical microscopy
242 and EDS maps as brunogeierite (GeFe_2O_4), carboirite [$\text{FeA}_{12}\text{GeO}_5(\text{OH})_2$], briartite [$\text{Cu}_2(\text{Fe},$
243 $\text{Zn})\text{GeS}_4$] and argutite (GeO_2). Briartite is the main Ge-mineral in Arre sphalerite but is rare in
244 Anglas and Pale Bidau sphalerite. Carboirite, brunogeierite and more rarely argutite are the
245 common Ge-minerals in the Anglas and Pale Bidau deposits. These minerals are preferentially
246 hosted along grain boundaries of recrystallized sphalerite grains. In Figures 4 and 5, Ge-
247 minerals occur as small crystals (< 100 μm ; see yellow dots in Figure 4) and are preferentially
248 aligned in the S_2 cleavage. Similarly, micrometric chalcopyrite is preferentially aligned in the
249 same recrystallized areas (Figure 4). These minerals are detected in all Type 2b samples
250 (Figs. 4c and 4e). Carboirite crystals are coarser (from 10 to 100 μm) compared to
251 brunogeierite, argutite, and briartite (5 to 20 μm in size; Fig. 5). Figures 5a-c shows three
252 representative examples of Ge-minerals occurrence characteristic of Type 2b mineralization.
253 Two sphalerite grain size fractions (below and above 100 μm) are considered for the GROD
254 maps (Figs. 5d-f). These maps show that Ge-minerals are only hosted between recrystallized
255 sphalerite grains (< 100 μm , Figs. 5d and 5e) in areas with low GROD angle, at the contact or
256 between coarse-grained sphalerite crystals or in twin boundaries (Fig. 5f).

257 **Chemical composition of sphalerite**

258 The distribution of trace elements for Type 1, 2a and 2b mineralizations is plotted in the
259 Figure 6 and 7 (see dataset in ESM 1). The main characteristic of Type 1 and Type 2a sphalerite
260 mineralizations is the small amount of Ge, mostly below 1 ppm (Fig. 6). In Type 1 disseminated

261 grains from Bentaillou, Fe median content is 2.5 ± 0.2 ppm, and the concentrations of Cu and
262 Ga are low with median values of 24 ± 22 ppm and 12 ± 1 ppm, respectively (Fig. 6).
263 Variations in concentration are relatively large in Cu, Ga, Ag, Sn, Sb and Pb compare to Fe,
264 Co, Ag or Cd with only minor disparities (Fig. 6).

265 In Type 2a, Fe median content is higher than in Type 1, with median value of
266 5.4 ± 0.5 ppm. Copper and Ga contents are generally low with values below 70 and 30 ppm
267 and median contents of 11 ± 13 ppm and 1.4 ± 6 ppm, respectively (Fig. 6). Recrystallized
268 grains in Victoria Type 2a mineralization show lower median contents in Mn, Fe, Co, Ga, and
269 Cd but higher median values in Ag, Pb, Sb, Sn, compared to coarse grains in the same sample
270 from Victoria (Fig. 6; analyzed spot locations are reported in Figure 3). Recrystallized grains
271 from Victoria and Margalida have similar Mn, Co, Cu contents (Fig. 6) but Fe and Cd median
272 values are distinct with 2.6 ± 0.1 wt.% and 952 ± 20 ppm for Victoria recrystallized grains and
273 5.0 ± 0.4 wt.% and 1946 ± 42 ppm for Margalida recrystallized grains (Fig. 6). Silver, Sn, Sb,
274 and Pb concentrations are similar between Victoria and Bentaillou coarse grains compared to
275 Mn, Co, Cu, Ga and Cd contents that show large variations (Fig. 6). For Bentaillou coarse grains
276 median values are 282 ± 19 ppm for Mn, 3.8 ± 0.2 wt.% for Fe, 94 ± 7 ppm for Co,
277 1 ± 0.3 ppm for Ga and 2246 ± 35 ppm for Cd. Manganese content is generally high in
278 Victoria and Bentaillou coarse grains (Figs. 6) with median values of 289 ± 100 ppm. Cobalt
279 content is high in Victoria coarse grains with a median value of 172 ± 10 ppm compared to the
280 grains analyzed in other textural position that yield a median value of 89 ± 8 ppm. Cadmium
281 is highly variable according to deposit and texture. Recrystallized Victoria coarse grains show
282 the lowest contents with a median value of 956 ± 20 ppm (Fig. 6). Victoria coarse grains
283 present higher Cd content with 1504 ± 28 ppm. But recrystallized grains from Margalida
284 contain similar Cd values with median contents of 1557 ± 37 ppm. Indium content is generally
285 below detection limit, when detected the median value is 0.8 ± 0.4 ppm. Silver, Sn, Sb and Pb

286 contents are generally higher in recrystallized grains (and in Type 1 disseminated crystals) than
287 in coarse grains (Fig. 6). In addition, Ag contents are relatively low with total median value of
288 7 ± 3 ppm but locally, Ag can reach 43 ppm in Margalida sphalerite (Fig. 6). The same range
289 of concentrations is observed for Sb, with median value of 7 ± 2 ppm with contents up to
290 45 ± 15 ppm. Tin content is low with a median value of 1 ± 0.2 ppm. Lead is below 100 ppm
291 for most of the studied sphalerite types except for Type 1 sphalerite which shows higher content
292 with a median value of 267 ± 37 ppm.

293 In Type 2b, Fe content ranges from 1.3 to 5 wt.% Fe but no systematic difference is
294 observed between deposits or textures (Fig. 7). In Pale Bidau, a progressive zonation in color
295 is observed (Figs. 4a and 4b), and high Fe contents are reported in dark-brown sphalerite zones
296 with values up to 5 wt.% Fe. High Mn median values are reported for Anglas (151 ± 10 ppm),
297 comparatively to Pale Bidau and Arre (24 ± 4 ppm and 18 ± 2 ppm, respectively). Conversely,
298 low median Cd values are present in Anglas (970 ± 38 ppm) and higher median values for Pale
299 Bidau and Arre (1500 ± 75 ppm and 1110 ± 43 ppm, respectively). No correlation between
300 grain-size or color and the contents in Fe, Mn or Cd has been established. Cobalt content is
301 homogeneous between the three textures (Fig. 7) with a median value of 143 ± 11 ppm. Low
302 In content is measured in these sphalerite types reaching a maximum of 6.2 ppm (Fig. 7) and
303 median value below 1 ppm. Indium content is slightly higher light coarse grains from Pale
304 Bidau with a median value of 4.1 ± 0.2 ppm.

305 Gallium content is generally below 100 ppm in the sphalerite samples (Fig. 7) but
306 exhibits relative large dispersion in concentrations between 0.3 and 156 ppm Ga which is not
307 clearly related either to the type of grain or the deposit. Silver, Sn, Sb and Pb contents are
308 commonly below 100 ppm. Some Pb-(Sb-Sn-Ag) micrometric inclusions are recognized in LA-
309 ICP-MS spectra (ESM 2) preferentially in recrystallized domains with Pb values reaching up

310 to 9420 ppm. Except for the presence of micro-inclusions, no clear difference between textures
311 (Fig. 7) and deposits can be defined.

312 Germanium and Cu contents show generally large variations from 102 to 580 ppm for
313 Ge and from 209 to 1265 ppm for Cu. These contents correlate with the type of texture
314 (Figure 4). and are detailed below.

315 In dark domains of coarser grains, Ge and Cu are typically high in the three deposits
316 with median value of 290 ± 19 ppm and 610 ± 61 ppm, respectively (Fig. 7). In Pale Bidau,
317 only few dark domain relicts ($n=4$) are reported and analyzed with median values of
318 152 ± 20 ppm for Ge and 233 ± 45 ppm for Cu. Arre samples show the highest Ge and Cu
319 contents with median contents of 433 ± 21 ppm and 1012 ± 75 ppm, respectively.

320 In light domains of coarse grains, Ge and Cu contents are lower with median values of
321 20 ± 2 and 60 ± 8 ppm, respectively (Fig. 7). Pale Bidau shows the lowest median contents,
322 with 0.7 ± 0.1 ppm for Ge and 29 ± 4.5 ppm for Cu compared to Arre and Anglas which show
323 median contents of respectively 38 ± 1 and 61 ± 3 ppm for Ge, and 102 ± 7 and
324 191 ± 13 ppm for Cu.

325 In recrystallized grains, Ge and Cu are systematically low like in light domains of coarse
326 grains, with median contents of 4.3 ± 1 ppm and 49 ± 10 ppm, respectively (Fig. 7). In Pale
327 Bidau, recrystallized grains may contain similar or larger enrichment in Ge and Cu than in light
328 color domains of coarse grains, with respectively 4.1 ± 0.5 ppm and 102 ± 19 ppm.

329 **Spatial distribution of trace elements in sphalerite**

330 Figure 8 shows Ge, Cu and Ga spatial distribution obtained from LIBS mapping of Arre
331 Type 2b sphalerite. Sphalerite is zoned in color in transmitted-light (Fig. 8a), correlating with
332 the concentration of these trace elements. Germanium and Cu contents are heterogeneous in

333 sphalerite, with patchy zoning, and locally oscillatory zoning (Fig. 8) with numerous Ge- and/or
334 Cu-bearing minerals (Figs 8b and 8c), some of them with Fe correspond to briartite $\text{GeCu}_2(\text{Fe}$,
335 $\text{Zn})\text{S}_4$ (ESM 2). Localized Ge-Cu minerals are preferentially found in sphalerite Ge-Cu-poor
336 domains (ESM 2). Nonetheless, some of them occur in Ge and Cu-rich dark domains and are
337 sub-parallel to S_2 cleavage (Figs. 8b and c).

338 Mass-balance calculations is performed, based on LIBS mapping (see dataset in ESM 1). The
339 bulk Ge concentration hosted in the Ge-rich minerals is 4.72 wt.% Ge, based on 25 wt.% Ge.
340 Light and dark domains present 0.1 and 2.07 wt.% Ge, respectively, considering Ge-poor and
341 Ge-rich sphalerite with 20 and 400 ppm Ge, respectively. Germanium mass fraction is two
342 times higher in Ge-minerals than in sphalerite. The total Ge mass is equivalent to a sphalerite
343 with 690 ppm Ge uniformly distributed in the structure without Ge-minerals.

344 Gallium in sphalerite shows zoned distribution on LIBS maps. Gallium-rich areas are located
345 in dark domains in transmitted-light, correlating with relatively high Cu and Ge contents
346 (Fig. 8d and ESM 2). Gallium-bearing mineral phases have not been detected. Other elements
347 such as Fe and Cd show a homogeneous distribution in sphalerite (ESM 2).

348 Trace elements distribution in quartz spatially associated with sphalerite is reported in Figure 8.
349 Several areas in quartz show enrichment in Ge, up to \sim 75 ppm (Fig. 8b). Copper and Ga have
350 not been detected in quartz (Figs. 8c and 8d). In the host rock, a few Cu-bearing micrometer
351 size grains are identified, and Ge is detected in structures parallel to S_0 - S_1 or S_2 (Fig. 8b) and
352 correlate with the higher Al contents (ESM 2), which correspond to muscovite.

353 **Discussion**

354 **Substitution mechanisms of rare metals in sphalerite**

355 The mechanism of Ge incorporation in sphalerite has been studied by Cook et al.
356 (2009b) who deduced from the apparent lack of correlation with other trace elements that Ge
357 could be present in a divalent form and substitute for Zn in the sphalerite structure ($Zn^{2+} \leftrightarrow$
358 Ge^{2+}). Later, Cook et al. (2015) reported a XANES study performed on Tres Maria Ge-Fe rich
359 acicular sphalerite, demonstrating Ge^{4+} in sphalerite and proposed substitution of Ge^{4+} and a
360 vacancy for Zn^{2+} or Fe^{2+} (i.e., $2Zn^{2+} \leftrightarrow Ge^{4+} + []$). In Central and East Tennessee, Bonnet et
361 al. (2017) showed Ge^{2+} and Ge^{4+} in the same sphalerite crystal, which was interpreted as due to
362 a difference in S_2 and O_2 fugacities. Bauer et al. (2018) demonstrated a coupled substitution
363 between Ge^{4+} and Ag^+ in low-temperature (186 ± 36 °C) Ge-rich sphalerite. Johan (1988)
364 suggested a coupled substitution between tetravalent Ge and combination of monovalent and
365 divalent Cu (i.e., $4Zn^{2+} \leftrightarrow 2Cu^+ + Cu^{2+} + Ge^{4+}$) on low-temperature Ge- and Cu-rich sphalerite
366 from Saint-Salvy. Belissont et al. (2016) μ -XANES study on the same sphalerite confirmed the
367 exclusive occurrence of Ge^{4+} in sphalerite, showing correlation with monovalent Cu^+ ,
368 supporting the hypothesis of a coupled substitution mechanism involving monovalent copper.
369 This substitution $3Zn^{2+} \leftrightarrow 2Cu^+ + Ge^{4+}$ is consistent with the PAZ sphalerite $2Cu - Ge$
370 correlation in Figure 9a. Coarse grains follow this substitution mechanism, however, the
371 recrystallized grains as well as some of the coarse grains with light domains do not plot along
372 the correlation line. For these two Ge-poor domains, such a deviation to the regression line may
373 be explained by the incorporation of Cu coupled to an additional substitution involving Ga
374 (Fig. 9b). Gallium in sphalerite is considered to be trivalent (Cook et al. 2009b; George et al.
375 2016; Wei et al. 2018) and its incorporation occurs following the coupled substitution
376 $2Zn^{2+} \leftrightarrow Cu^+ + Ga^{3+}$.

377 Divalent cations such as Cd²⁺, Mn²⁺, Co²⁺ or Fe²⁺ are incorporated by direct substitution
378 with Zn²⁺ ($X^{2+} \leftrightarrow Zn^{2+}$, Cook et al. 2009b; George et al. 2016). No clear correlation is observed
379 between Cd²⁺ and Fe²⁺ (Fig. 9c), but the Cd-Fe values are mainly varying according to the
380 deposit, indicating poor mobility during vein recrystallization. In Figure 9d, Fe is significantly
381 correlated with Cd+Mn+Co at low values, which attests for their coupled incorporation in
382 sphalerite.

383 **Relationship between sphalerite texture, chemical composition and formation**
384 **temperature**

385 Crystallization temperature of vein assemblages is difficult to assess, but we use relative
386 field chronology to place vein formation in the regional tectono-metamorphic framework. M₁
387 peak is estimated at 580 °C and 0.55 GPa, and M₂ peak at 525 °C and 0.2 GPa (Mezger and
388 Passchier 2003; Mezger et al. 2004). The three sphalerite mineralization types show significant
389 differences in trace element signatures. Types 1 and 2a have low Ge content (<0.5 ppm), as
390 well as Cu, Ga, and Ag, Sb, Sn, Pb but higher Fe, Mn, and Cd contents (Fig. 10). Type 1
391 disseminated mineralization from Bentaillou is considered pre-metamorphic and has been
392 affected by Variscan M₁ regional and M₂ contact metamorphism (Cugeronne et al. 2018b).

393 Type 2a sphalerite from Bentaillou, Victoria and Margalida is syn-M₁ and prior to D₂
394 associated with F₂ folds (Fig. 3) and M₂ metamorphism. During metamorphism, sphalerite can
395 incorporate Fe, Mn, and Cd and in a lesser extent Pb, Bi, Ag, Sn, and Sb (Lockington et al.
396 2014) during static recrystallization. Accordingly, Type 2a sphalerite that was statically
397 (coarse-size fraction) and subsequently dynamically recrystallized (small-size fraction) in
398 Victoria deposit shows respectively high and low Fe, Cd and Mn contents (Figs. 6 and 8c, d).
399 The geothermometer based on Ga-Ge-In-Mn-Fe concentrations in sphalerite lattice (GGIMF,
400 Frenzel et al. 2016) was applied for Type 1 and 2a sphalerite (see dataset in ESM 1) and yields
401 temperatures below 150 °C. Sphalerite underwent static recrystallization at temperatures, likely

402 above 400 °C, and in some areas subsequent dynamic recrystallization, which affect the
403 reliability of the GGIMF geothermometer (Frenzel et al. 2016), such that, this temperature is
404 not that of mineralization or deformation.

405 The chemical composition of Type 2b sphalerite from the Anglas and Arre deposits is
406 heterogeneous. Germanium, Cu and locally Ga are enriched in coarse grains with dark domains
407 (Fig. 8) but depleted in recrystallized grains and coarse crystals with light domains (Cugeronne
408 et al. 2020). At Pale Bidau, no strong chemical zonation is observed in sphalerite, except some
409 color bands that highlight contrasts in Fe-contents (Cugeronne et al. 2018a). However, Ge-poor
410 coarse crystals can be found (Fig. 4b), the latter being interpreted as inherited Ge-poor domains.

411 Type 2b mineralization has low Cd and Fe contents. Variations in Cd content between
412 the Arre-Anglas and Pale Bidau deposits might be linked to differences in the fluid source or in
413 the bulk host rock composition. The Arre and Anglas deposits are hosted by similar Devonian
414 sequences in the same mining district, contrary to Pale Bidau which is hosted by Late-
415 Ordovician calc-schists (Cugeronne et al. 2018b, Fig. 1). Cadmium content in Type 2b
416 mineralization is generally not affected by recrystallization. Nonetheless, Type 2a
417 recrystallized and coarse sphalerite grains from the Victoria deposit present differences in Cd
418 concentrations, as well as in Fe, Mn and Co. Recrystallized grains from Victoria have similar
419 low Cd-Fe concentrations to Type 2b sphalerite from the Anglas deposit. A likely hypothesis
420 is that Type 2b mineralization resulted from remobilization of early Type 2a mineralization in
421 D₂ structures, such as Victoria F₂ fold hinge (Fig. 3). However, this hypothesis is restricted by
422 the lack of Ge-minerals in recrystallized sphalerite and by the relatively low Ge, Cu or Ga
423 contents in the Type 2a mineralization. Late Ge-Cu(-Ga)-rich fluids circulating in S₂ cleavage
424 plane during Type 2b mineralizing event may have been the main Ge-carrier.

425 If Type 2b primary sphalerite had the same chemical composition as coarse grains, the
426 estimated temperature, using GGIMF, is about 250 °C, with large variations of \pm 50 °C due to
427 chemical heterogeneity between dark and light domains. This geothermometer may be applied
428 in dark domains of the coarse grains, if we consider no change in composition after the primary
429 mineralizing stage. However, in light domains, composed of recrystallized grains and some part
430 of the coarse grains, the temperature recorded is inconsistent. Deformation and recrystallization
431 of the sphalerite lattice have contributed to the redistribution of the elements used in GGIMF
432 thermometry. Germanium-rich veins are commonly formed at low temperatures, below 200-
433 250 °C (Munoz et al. 1994; Höll et al. 2007; Belissoent et al. 2014; Frenzel et al. 2016; Bauer et
434 al. 2018). There is no fluid inclusion study yet for the PAZ mineralization to compare with the
435 GGIMF. Johnson et al. (1996) have investigated fluid inclusions in the Cierco deposit where
436 Pb-Zn veins similar to the PAZ Type 2b veins appear (Fig. 1b). At Cierco, some Ge-minerals
437 have been identified in recrystallized sphalerite hosted in Late-Variscan granodiorite. This
438 likely indicates that at least a part of the mineralization, belongs to Type 2b (Johnson et al.
439 1996). Primary or pseudo-secondary fluid inclusion microthermometry in quartz and sphalerite
440 yield temperature between 150 ° and 200 °C with salinities between 3 and 29 wt.% NaCl
441 equivalent. Pressure is assumed to be negligible (Johnson et al. 1996). No clear petrographic
442 relationship or evidences between fluid inclusions and host-minerals is described and their
443 primary or pseudo-secondary origins are questionable in a largely deformed environment.

444 In our samples, pressure-temperature conditions of deformation events, which are syn-
445 or post-mineralization, are difficult to assess. Experimental studies on carboirite brackets its
446 stability field between 340 and 600 °C and 0.75 to 2 MPa whereas brunogaeierite is reported
447 above 510 °C at 2 MPa (Julliot et al. 1987). Quartz crystals appear only slightly deformed,
448 preserving original cockade textures and no alteration halo is observed along the vein contacts.
449 Partial recrystallization of the Pyrenean sphalerite is localized in bands associated to a

450 pronounced cleavage, parallel to the S₂ cleavage of the host-rock (Fig. 4 and 8). Such a partial
451 recrystallization might occur at temperatures below 400 °C and low pressure in comparison
452 with experimental studies on sphalerite that evidence dynamic recrystallization of sphalerite at
453 low-metamorphic grade (Clark and Kelly 1973; Siemes and Borges 1979; Couderc et al. 1985;
454 Cox 1987). Recrystallization of sphalerite at higher-metamorphic grade will result loss of parent
455 grains in favor of recrystallized grains.

456 **Mechanisms of formation of accessory minerals.**

457 Germanium-minerals are exclusively located in Type 2b veins which have been affected
458 by subgrain rotation recrystallization or in close association with weak structures such as twin
459 boundaries (Fig. 8b; Cugeronne et al. 2020). Figure 11a shows a genetic model of their
460 formation. Stage 1 represents the primary Type 2b mineralization hosted in an undeformed vein
461 and Ge is in the sphalerite lattice. Stage 2 is characterized by dynamic recrystallization of
462 sphalerite related to subgrain rotation mechanism (Cugeronne et al. 2020) and the formation of
463 Ge-minerals. Ge-minerals are not included in sphalerite and are commonly located at sphalerite
464 triple junctions (Figs. 5d and f). They rim the grain boundaries of recrystallized sphalerite,
465 which attests of their late formation.

466 Occurrence of intra-granular oscillatory zonations (Plümper et al. 2012) may inform on
467 diffusion mechanisms (Fig. 8 and ESM 2) within sphalerite coarse grains in dark to light
468 domains. Intra-granular diffusion is comprised of different mechanisms such as volume, high
469 diffusivity pathway, or dislocation-impurity pair diffusion (Klinger and Rabkin 1999; Reddy et
470 al. 2007; Plümper et al. 2012; Vukmanovic et al. 2014). In addition, hydrothermal fluids might
471 have percolated and extract Ge, Cu and Ga from the sphalerite lattice along grain boundaries,
472 and twin cleavage planes which constituted nucleation sites as Ge(-Cu) minerals are abundant
473 in such locations (Fig. 5). In hydrothermal fluids, Ge is transported as a Ge(OH)₄ complex, and

474 its solubility increases with temperature (up to 400 °C; Pokrovski and Schott 1998; Pokrovski
475 et al. 2005). In the Arre deposit, direct correlation between Ge and Cu in sphalerite lattice and
476 in accessory minerals (briartite; Figs. 8b and 8c) attests to the similar behavior of these two
477 elements during remobilization. It could be linked to similar ionic radius (Shannon 1976) or
478 specific conditions enhancing the extraction and coupled incorporation in accessory phases. In
479 the Pale Bidau and Anglas deposits (Figs. 4a, b and 4e, f), Ge-minerals such as carboirite and
480 brunogeierite do not contain high Cu contents which may indicate possible differences in Ge
481 and Cu behavior. Gallium is only found in coarse grains (< 100 ppm Ga), mostly correlated to
482 Ge- and Cu-rich areas in the sphalerite lattice (Fig. 8), but no Ga enrichment occurs in Ga-
483 bearing phases which may be due to loss of this metal or to low primary content in sphalerite
484 (Fig. 8). No impact of recrystallization on Cd concentrations is detected in sphalerite (ESM 2).
485 Cadmium is mainly transported as chloride water complexes in hydrothermal fluids
486 ($20 \leq T \leq 450$ °C; Bazarkina et al. 2010) whereas, hydroxyl complexes generally transport Ge
487 ($[GeOH]_4$; Pokrovski and Schott 1998) and more rarely Cu ($[CuOH]$; Hack and Mavrogenes
488 2006) at low-pressure (< 500 MPa). Hydroxyl complexes might be the main rare metal-bearing
489 complexes in hydrothermal fluid during the remobilization stage.

490 Fields of stability of Ge in different minerals according to fO_2 and fS_2 indicate potential
491 fO_2 increase during formation of Ge-minerals during the stage 2 (Fig. 11b). Increasing of fO_2
492 may enhance the release of Ge from sphalerite lattice by circulation of low-temperature fluids
493 which significantly change and diversify the Ge-mineral paragenesis (silicates, oxides,
494 hydroxides, sulfide). It is important to note that hydrothermal fluids responsible for Ge
495 remobilization during deformation do not need to be rich in Ge, if we consider a primary
496 sphalerite with uniform ~ 700 ppm Ge content (Fig. 11a and ESM 1). These arguments allow
497 to demonstrate that a primary Ge-rich sphalerite (and other sulfides) impacted by low-grade

498 deformation below 400 °C with significant increase in fO₂ (Fig. 11b) are potential indicators of
499 Ge-mineral rich deposit.

500 **Comparison with world-class Ge-rich deposits**

501 Type 2b sphalerite vein mineralization from the PAZ present chemical and textural
502 similarities with world-class Pb-Zn(-Ge) districts. In many SHMS, VMS or other Pb-Zn(-Cu)
503 deposits, metamorphism and deformation may have reworked the mineralization (Bodon 1998;
504 Large et al. 2005; Huston et al. 2006; Wilkinson 2013; Vikentyev et al. 2016) as shown by
505 partly dynamically recrystallized sphalerite at the world class George Fisher (Chapman 2004;
506 Murphy 2004) or in the Currawong deposits (Bodon and Valenta 1995). However, too few
507 chemical analyses of rare metals have been performed in these deposits. As an example,
508 Figure 12 reports a compilation of Ge and Cu contents in sphalerite analyzed in Pb-Zn(-Cu)
509 deposits worldwide. These deposits are mostly related to low-temperature hydrothermal fluids
510 (~100-250 °C) with no- or only low-grade deformation. Numerous studies on Pb-Zn(-Cu)
511 deposits deformed or metamorphosed at higher grades (>400 °C) do not report significant rare
512 metals contents in sphalerite or in accessory Ge-minerals and may be related, in some cases, to
513 insufficient textural and chemical investigation. Nonetheless, in all Pb-Zn deposits affected by
514 high grade metamorphism, Ge was lost from sphalerite during high-grade metamorphic
515 processes (>400°C) and recrystallization has potentially concentrated Ge in accessory minerals.
516 This assumption is likely for the granulite facies Broken Hill deposit (Spry et al. 2008; George
517 et al. 2016) where Ge is not reported in sphalerite, while the Ga and In contents are low,
518 suggesting that these rocks should be revisited for accessory minerals rich in rare metals.

519 Examples of Ge-rich sphalerite in the Saint-Salvy deposit, Montagne Noire (France), in
520 the Erzgebirge district (Germany), and in MVT deposits from the Central and East Tennessee
521 indicate that Ge is in bands within a single crystal (Bonnet 2014; Bauer et al. 2018) or in sector

522 zones (Belissont et al. 2014) and may reach up to ~3000 ppm Ge (Fig. 12) with no Ge-minerals
523 reported.

524 In Ge-rich sphalerite, a correlation between Ge and Cu is observed except with the Tres
525 Maria deposit where Ge is mainly correlated to Fe (Fig. 12). Wurtzite, a non-cubic zinc sulfide,
526 is inferred in the zinc ore at Tres Maria (Type I mineralization; Cook et al. 2009b, 2015; Saini-
527 Eidukat et al. 2009) due to an acicular habit, similar to the wurtzite observed in the Kokanee
528 range by Beaudoin (2000).

529 Rare elements (Ge, Ga, In, Cd) are reported in deformed sulfide deposits such as
530 Barrigão, a remobilized Cu-rich VHMS deposit. Germanium-rich chalcopyrite with
531 heterogeneous contents and Sn-Ge-rich phases (Reiser et al. 2011) are probably a result of Ge
532 remobilisation during deformation (Belissont et al. 2019). In China, Pb-Zn deposits, considered
533 to be deformed MVT deposits, contain abundant deformed sphalerite mineralization as in Huize
534 (Han et al. 2006), Niujiaotang (Ye and Tiecheng 1999; Ye et al. 2012), Mengxing (Ye et al.
535 2011), Jinding (Xue et al. 2007) and Lehong (Wei et al. 2019). Germanium and Cu contents
536 from these deposits are plotted in Figure 12 and show that the chemical composition of these
537 deposits is similar to that of the PAZ deposits, but with a wider dispersion along the Ge-Cu
538 arrow. Wei et al. (2019), focusing on the carbonate-hosted Lehong deposit, described dark-
539 brown Ge-rich sphalerite and light-brown Ge-poor sphalerite with Ge content similar to
540 Pyrenean vein mineralization with up to 536 ppm and 118 ppm in dark and light-brown
541 domains, respectively. However, no detail on the texture of these sphalerite grains is provided,
542 and Ge-minerals are not reported. Another example of deformed sphalerite mineralization is
543 reported by Monteiro et al. (2006) in the carbonate-hosted Fagundes Pb-Zn deposit, where large
544 variations in Ge (up to 2390 ppm Ge) were measured. Based on similarities with our
545 observations, this heterogeneity in Ge content may be related to differences in sphalerite texture
546 such as Type 2b sphalerite in the PAZ.

547 Germanium minerals were reported in the Kipushi-type deposits (Emslie 1960;
548 Francotte et al. 1965; Ottemann and Nuber 1972). Belissoont (2016) describes highly variable
549 Ge contents, from 0.2 to 5930 ppm, in sphalerite from the Kipushi mine that presents evidence
550 of dynamic recrystallization (Hughes 1987). At the Black Angel deposit, Greenland (Horn et
551 al. 2018), sphalerite is depleted in Ge and Cu, and the occurrence of briartite is reported. In the
552 Eastern-Alpine base-metal sulfide ores, a LA-ICP-MS study indicates Ge-rich mineralization
553 in carbonate-hosted Pb-Zn deposits (Melcher and Onuk 2019). The analyses reveal that the Ge
554 content in sphalerite is highly variable (from < 1 up to 3700 ppm). Moreover, renierite
555 ($\text{Cu},\text{Zn})_{11}(\text{Ge},\text{As})_2\text{Fe}_4\text{S}_{16}$) is reported as inclusions in bornite from one of these deposits.

556 **Conclusions**

557 Type 2a stratabound mineralization in the PAZ was affected by static recrystallization
558 and local dynamic recrystallization during M₂ metamorphism. In such mineralization,
559 sphalerite is largely depleted in Ge, Cu, and Ga, but generally enriched in Cd. However, Cd
560 content has been found to be variable depending on the type of deposit and the texture, with
561 substantial depletion during dynamic recrystallization.

562 Type 2b mineralization, interpreted as dynamically recrystallized Pb-Zn vein
563 mineralization, shows chemical heterogeneity in sphalerite, particularly in terms of Ge and Cu
564 contents, which are related to deformation at relatively low-temperature (< 400 °C) associated
565 with a possible increase in fO₂. Gallium is depleted in the recrystallized fraction and
566 recrystallization does not have significant impact on Cd concentrations. The interplay between
567 diffusion and dissolution processes in sphalerite may account for large variations in Ge and Cu
568 contents and the precipitation of Ge-minerals.

569 Textural and chemical heterogeneity is frequent in Pb-Zn deposits around the world.
570 Understanding the coupling between microstructures and chemistry has major implications on
571 evaluating rare element potential of deposits. In this study, we show that exploration for rare
572 and critical metals (Ge, Ga, In) must integrate micro-textural and in-situ chemical analyses of
573 sulfides. In cases of dynamic recrystallization of sphalerite, we show in-situ analysis by LA-
574 ICP-MS is not sufficient to provide a good screening of critical metals. In such case, these
575 elements will be accumulated in discrete solid phases that can be found using chemical maps.
576 These small minerals can be often responsible for the high dispersion of bulk and in-situ
577 chemical data. Dynamic recrystallization of sphalerite appears to be widespread in all orogenic
578 fields. The recognition of these textures is essential in a first stage to locate enrichment of rare
579 metals in specific mineral phases. Numerous deformed Pb-Zn(-Cu), such as in SEDEX, MVTs
580 or Kipushi-type deposits should be further explored for their contents of rare metal.

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922 **Figure captions**

923 **Figure 1.** a Location of the Pyrenean Axial Zone in the Variscan belt of Western Europe. b
924 Lithological and structural map of Central Pyrenean Axial Zone with location of Pb-Zn deposits
925 (from Pouit [1985], Bureau de Recherches Géologiques et Minières [<http://infoterre.brgm.fr>],
926 and Instituto Geológico y Minero de España [<http://mapas.igme.es/Servicios/default.aspx>]). c
927 Structural 3D synthetic model with location of the studied Pb-Zn deposits in relation to
928 lithology, structure and metamorphic grade. Chronological interpretation from Cugeron et al.
929 (2018b). Timing of mineralization for the Type 2b vein mineralizations is unknown but
930 deformation imprint is considered as Late-Variscan (D_2).

931 **Figure 2.** Polished samples of Type 2a and Type 2b mineralization (Cal: calcite; Chl: chlorite;
932 Ghn: gahnite; Qz: quartz; Sp: sphalerite). a: Characteristic of the three mineralization types:
933 Type 1, Type 2a and Type 2b. b Sample with millimetric coarse-grained sphalerite (dark-brown
934 reddish), associated to calcite (white) and black-schist (Bentaillou deposit). c Pluri-micrometric
935 small-grained sphalerite (dark-brown reddish). A brecciated texture can be observed with clasts
936 of quartz, calcite and black-schist (Margalida deposit). d Folded (F_2) coarse-grained sphalerite

937 and quartz mineralization. Gahnite crystals are reported in schists and in the mineralization
938 (Victoria deposit). e Light to dark-brown zonation in sphalerite crystals associated to quartz and
939 calc-schist host rock. Set shows coalescence of coarse millimetric and small micrometric
940 crystals appear (Arre deposit). f Light to dark brown sphalerite associated to quartz, calcite and
941 chlorite. Dark brown vertical bands are visible on the sample. Sphalerite grains are locally
942 visible as shown by the black arrow at bottom right (Pale Bidau deposit).

943 **Figure 3.** EBSD-grain reference orientation deviation angle component (GROD) maps related
944 to grain size superimposed onto a transmitted-light microphotograph of Type 2a sphalerite-
945 quartz mineralization hosted in schist deformed by F₂ fold hinge from Victoria deposit. The
946 GROD maps allow to visualize intragranular microstructures and represent the misorientation
947 deviation for each pixel compared to the mean misorientation deviation of the entire sphalerite
948 grain (steps of measurement of 4 and 10 µm for the left and right EBSD maps respectively). In
949 the GROD calculation, twin boundaries are considered as grain boundaries to exclude their
950 large misorientation (55-60°)(Ghn: gahnite; Qz: quartz; Sp: sphalerite). LA-ICP-MS spot
951 analyses in Type 2a coarse and recrystallized sphalerite are indicated

952 **Figure 4.** Microphotographs associated to EBSD-grain reference orientation deviation angle
953 component (GROD) maps related to sphalerite grain size. The GROD maps allow to visualize
954 intragranular microstructures. Small grains contain low internal misorientation compared to
955 coarse grains, which confirm the presence of dynamic recrystallization. Grain boundaries are
956 represented in all maps except in (b) and chemical contents acquired with LA-ICP-MS are
957 reported (Gln: galena; Qz: quartz; Sph: sphalerite). a Scanned microphotograph of sphalerite
958 with superposition of four EBSD-grain size maps (Pale Bidau deposit; step of measurements
959 between 1.5 and 4 µm). Germanium and Cu minerals observed with EDS map are reported.
960 Location of (b) is indicated. b Enlargement of (a) showing the location of LA-ICP-MS spot
961 analyses with Ge contents essentially on two EBSD-GROD maps. Occurrence of Ge and Cu

962 minerals are reported with empty yellow and blue circles to highlight relationships between Ge-
963 Cu minerals and sphalerite (Step of measurement of 1.5 and 4 μm for the left and right EBSD
964 maps respectively). c EBSD-GROD map with location of Ge- and Cu-rich phases (Arre
965 deposit). These minerals are preferentially located in the recrystallized sphalerite domains.
966 Location of (d) is represented (step of measurement of 2.5 μm). d Enlargement of (c) in
967 transmitted-light with LA-ICP-MS spot analyses associated to Ge contents. Note that the color
968 zonations correspond in part to variations in Ge content. e EBSD-GROD map with location of
969 Ge- and Cu-rich minerals in recrystallized sphalerite (Anglas deposit). LA-ICP-MS spot
970 analyses are located with their corresponding Ge contents (step of measurement of 3 μm). f
971 Enlargement of (e) in transmitted-light (150 μm of thickness) with color zonation between light
972 and dark brown. LA-ICP-MS spots are indicated with black or white arrows corresponding to
973 (e).

974 **Figure 5.** Images of Ge-minerals hosted in the interstices between recrystallized sphalerite
975 grains (Cal: calcite; Qz: quartz; Sp: sphalerite). a Five brunogeierite crystals from the Arre
976 deposit (in the black circle; BSE image.). b Carboirite crystals with in sphalerite and calcite
977 from the Pale Bidau deposit (SEM image) c Nine Ge-minerals hosted between sphalerite grains
978 from the Anglas deposit. d-e-f EBSD-GROD maps related to grain size for (a), (b). and (c),
979 respectively (step of measurement of 0.5, 3 and 3 μm for the three EBSD maps, respectively).

980 **Figure 6.** LA-ICP-MS data for Type 1 and 2a sphalerite, showing trace element concentration
981 in three different textures and for the Bentaillou, Victoria and Margalida deposits (dataset in
982 ESM 1). a Individual analyses. b Box plots.

983 **Figure 7.** LA-ICP-MS data for Type 2b vein sphalerite, showing trace element concentration
984 in three different textures and for the Arre, Anglas and Pale Bidau deposits (dataset in ESM 1).
985 a Individual analyses. b Box plots.

986 **Figure 8.** LIBS mapping on Arre Type 2b (Abbreviations: Qz: quartz; Sp: sphalerite). a
987 Microphotograph (thick section, transmitted-light) of Arre sphalerite associated to quartz and
988 calc-schist host-rock. b Ge map in sphalerite, quartz and host-rock. Ge is also shown in purple
989 color in the host-rock (calc-schist) because it is measured on a different spectral line due to Al
990 interferences. White arrow indicates specific Ge zones. c Cu map in sphalerite, quartz and host-
991 rock. d Ga map in sphalerite, quartz and host-rock.

992 **Figure 9.** Content of elements for Type 1, 2a and 2b sphalerite. In coarse grains within dark
993 domains, analyses are close to the Ge – 2Cu line. a Ge vs Cu contents. Recrystallized grains
994 and light domains in the coarse grains do not lie on the correlation line. Pale Bidau coarse grains
995 within light domains contain very low Ge contents. b Ge + Ga vs Cu contents. All the points lie
996 closely to the substitution $\text{Ge} + \text{Ga} = 1\text{Cu}$ or 2 Cu depending on the dominant element in the
997 sphalerite lattice (Ge or Ga). c Cd vs Fe contents. Note that each deposit forms a cluster.
998 Recrystallized grain concentrations from Victoria deposit are close to Anglas deposit. d
999 Cd + Mn + Co vs Fe contents. Note the correlated enrichment between Cd + Mn + Co and Fe
1000 contents. Recrystallized grains compositions from Victoria deposit are close to Anglas deposit.

1001 **Figure 10.** Trace element comparison of coarse sphalerite grains between Type 2b and Type 2a
1002 crystals.

1003 **Figure 11.** a Two stage genetic model of the formation of Ge-minerals and sphalerite textures
1004 and chemical composition in Type 2b Pb-Zn mineralization. Stage 1 represents the
1005 mineralization before D_2/M_2 Variscan deformation. Stage 2 is the mineralization after
1006 deformation. b Behavior of Ge in different geologic environments and minerals according to
1007 $f\text{O}_2$ and $f\text{S}_2$ for the two Type 2b stage described in (a). Experimental data were acquired
1008 between 27 to 327 °C (300 to 600 K) and 0.1 MPa (modified from Bernstein, 1985 with values
1009 of Sahlström et al. 2017). Stage 2 is plots in three different fields according to the deposit:
1010 sulfide, sphalerite and oxide fields. The Arre deposit is restricted to the “Ge-sulfides” field due

1011 to abundance of briartite and rare carboirite and brunogeierite. The Anglas and Pale Bidau
1012 deposits contain both carboirite and brunogeierite but briartite is rare. Germanium in sphalerite
1013 lattice occurs quasi exclusively in Arre and Anglas. Ge-hosts in blue represent the Ge-carrier
1014 in the PAZ Pb-Zn deposits.

1015 **Figure 12.** Germanium vs copper contents comparison between Pyrenean sphalerite
1016 (including Argut-dessus deposit; Cugeron et al. 2018a) and others sphalerite in the world. A
1017 lower Ge limit of 10 ppm is set to compare Ge-rich sphalerite. Germanium and copper
1018 measurements generally follow the line Ge – 2Cu when Ge is enriched in the sphalerite
1019 lattice, with the exception of the Tres Maria zinc sulfide. 1-Niujiatang, Huize, Jinding and
1020 Mengxing deposits: 2-Ye et al. (2011); 3-Central and East Tennessee deposits: Bonnet (2014);
1021 4-Tres Maria, Magura, Rosia Montana, Sacarimb deposit: Cook et al. (2009); 5-Erzgebirge
1022 deposits: Bauer et al. (2018); 6-Saint-Salvy deposit: Belissont et al. (2014).

1023 **Supplementary materials captions**

1024 **ESM 1.** Chemical datasets. **Table S1** In-situ LA-ICP-MS dataset of the Type 1, Type 2a and
1025 Type 2b mineralizations. **Table S2** Mass balance calculation in Ge-rich sphalerite from the Arre
1026 deposit. **Table S3** GGIMF geothermometer calculation of the in-situ LA-ICP-MS datas based
1027 on the equation from Frenzel et al. (2016).

1028 **ESM 2.** Additional figure and LIBS maps. **Figure S1** Time-integrated representative LA-ICP-
1029 MS spectra of sphalerite. Ge spectra is represented in red color. i Spectra in coarse grains with
1030 dark domains (see sample description below). Note the smooth profile for Ge (401 ppm) and
1031 most of the element (except for Sb) which contain micrometric minerals (Anglas sphalerite,
1032 cc_37). ii Spectra in coarse grain with light domains (see sample description below). Note the
1033 smooth and lower signal for Ge (76 ppm) and a small Sb mineral at the beginning of the signal
1034 acquisition (~180 sec.) (Anglas sphalerite, cc_25). iii Spectra in recrystallized grains (see

1035 sample description below) with very low Ge content (0.8 ppm) and relatively smooth signal for
1036 the other elements (Anglas sphalerite, cc_34). iv Spectra in recrystallized grain with the
1037 presence of Fe-Ge-Cu-rich phases. The signal is divided in particular rich-phases field (1) and
1038 sphalerite field (2) (Anglas sphalerite, cc_10). v Spectra in recrystallized grain with occurrence
1039 of high Pb content (9420 ppm) in two micro-phases (Pale Bidau sphalerite, ce_18). **Figure S2**
1040 LIBS mapping on Arre Type 2b ore mineralization. The analyzed area corresponds to the
1041 Figure 8. i Zn map. ii Si map. iii Fe map. iv Cd map. v Al map. vi Mg map. vii Ti map. **Figure**
1042 **S3** Transmitted-light microphotograph of the ore mineralization area analyzed with LIBS.
1043 **Figure S4** Germanium high-resolution LIBS map. **Figure S5** Copper high resolution LIBS
1044 map. **Figure S6** Gallium high resolution LIBS map.

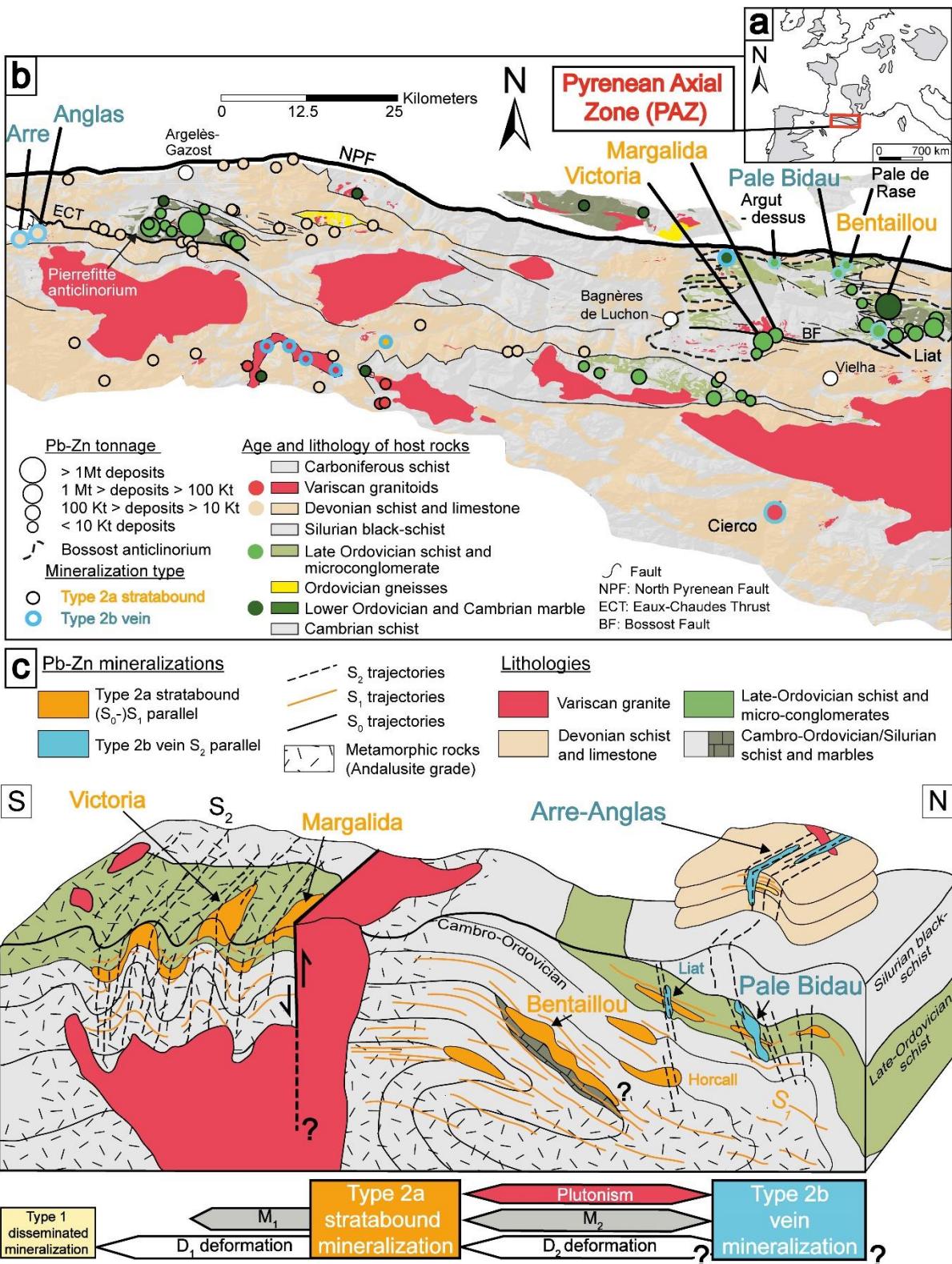


Fig.1

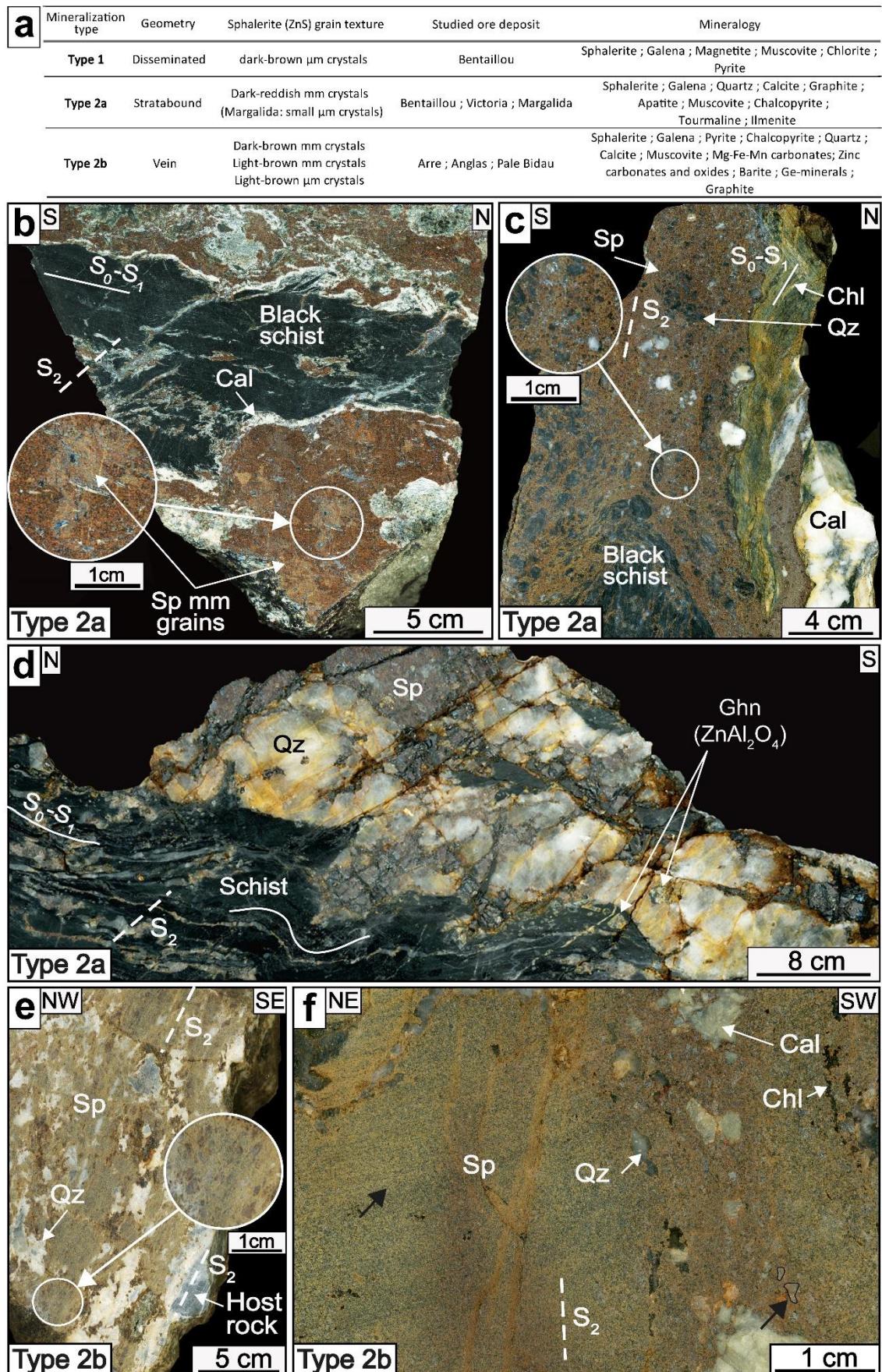


Fig. 2

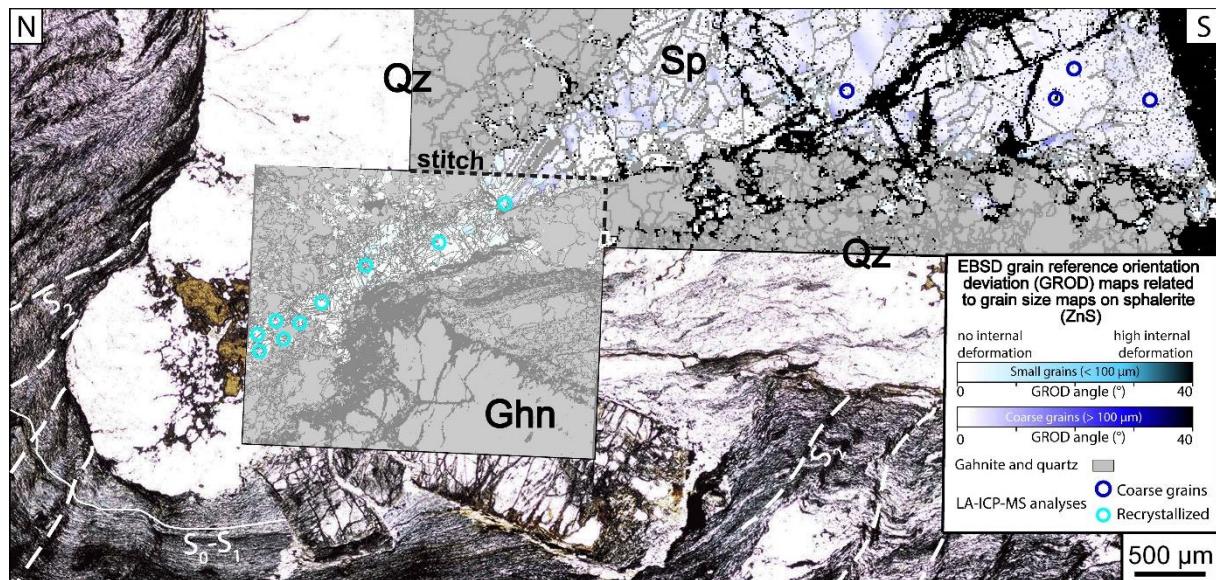


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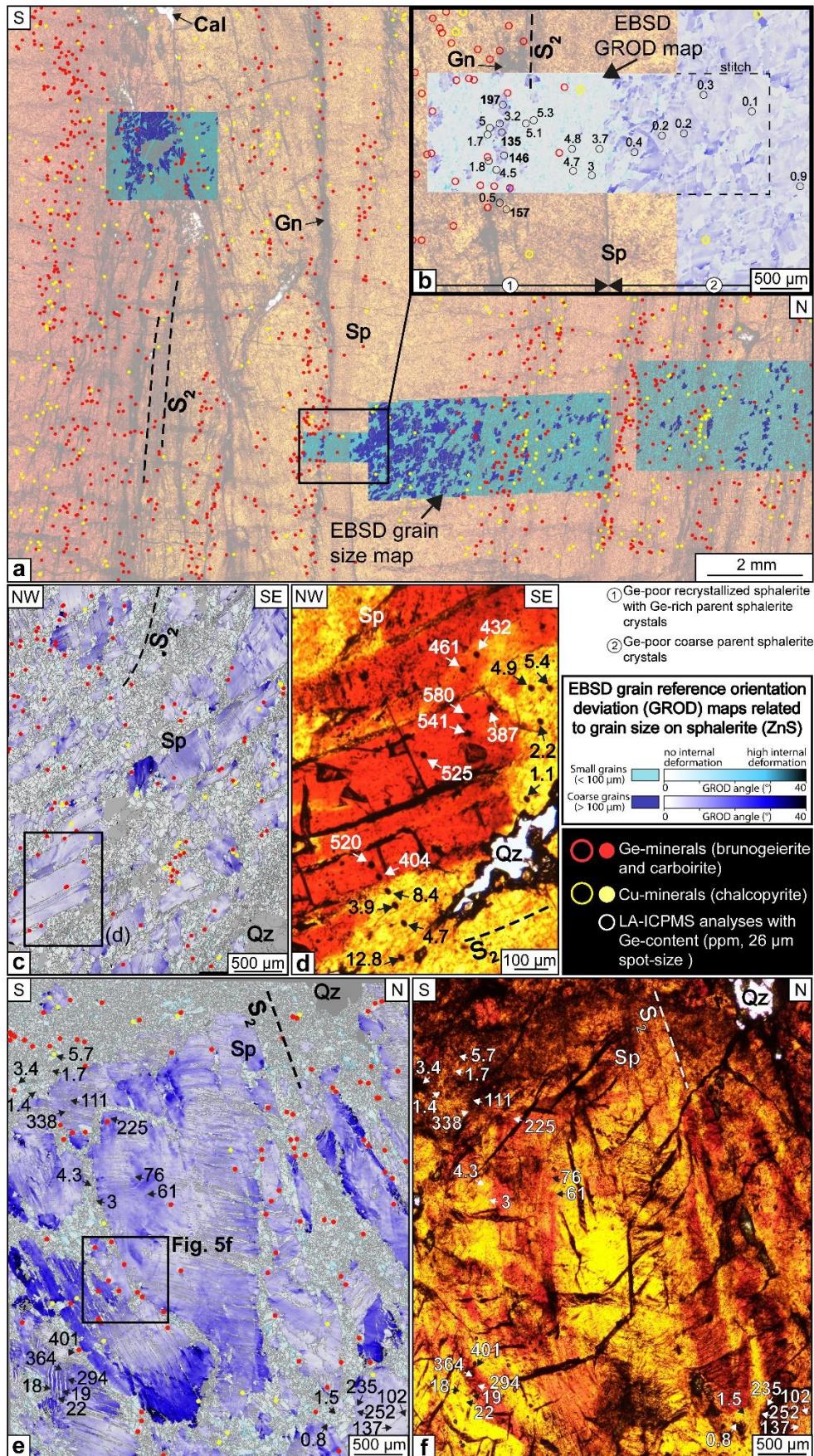


Fig. 4

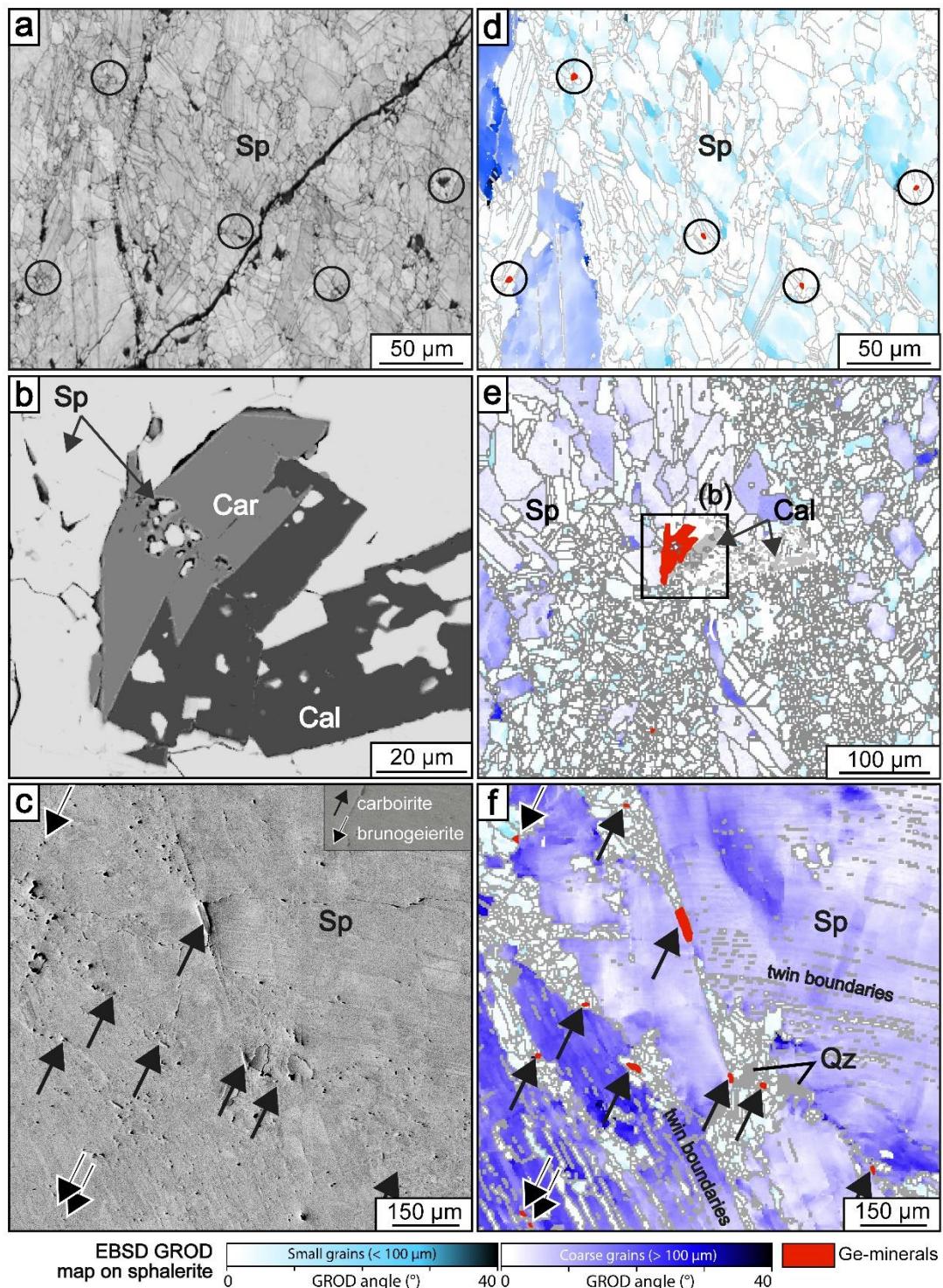


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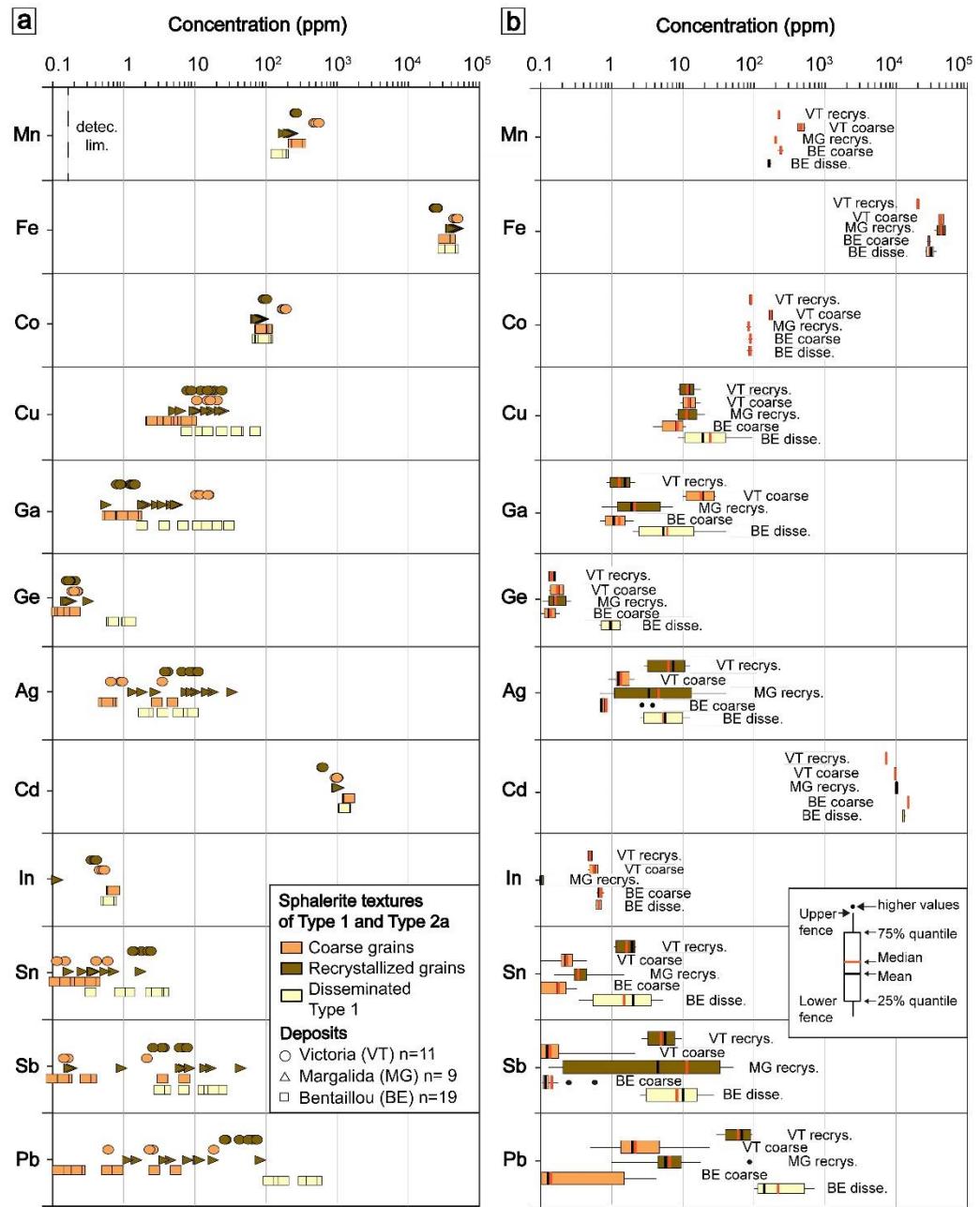


Fig. 6

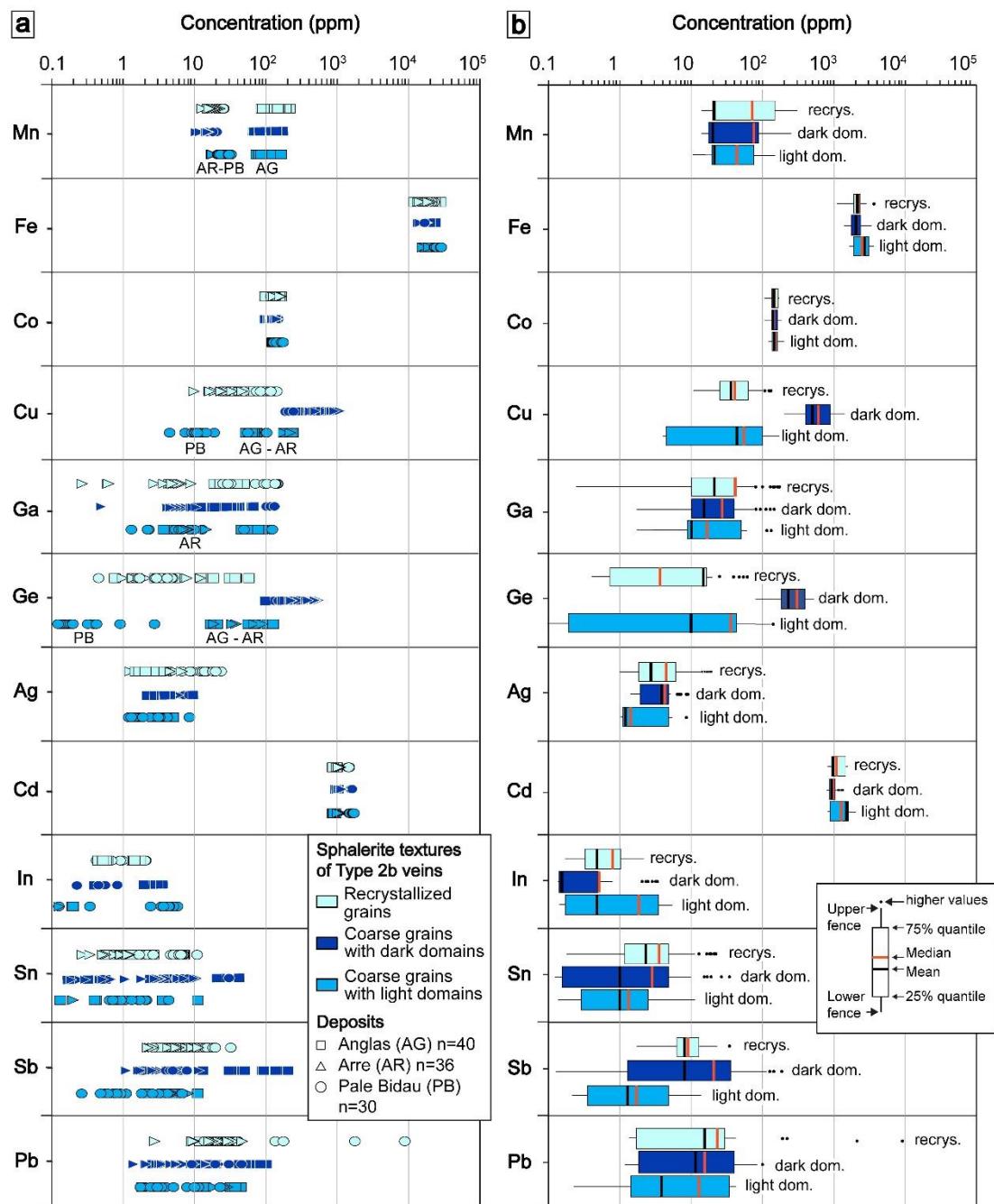


Fig. 7

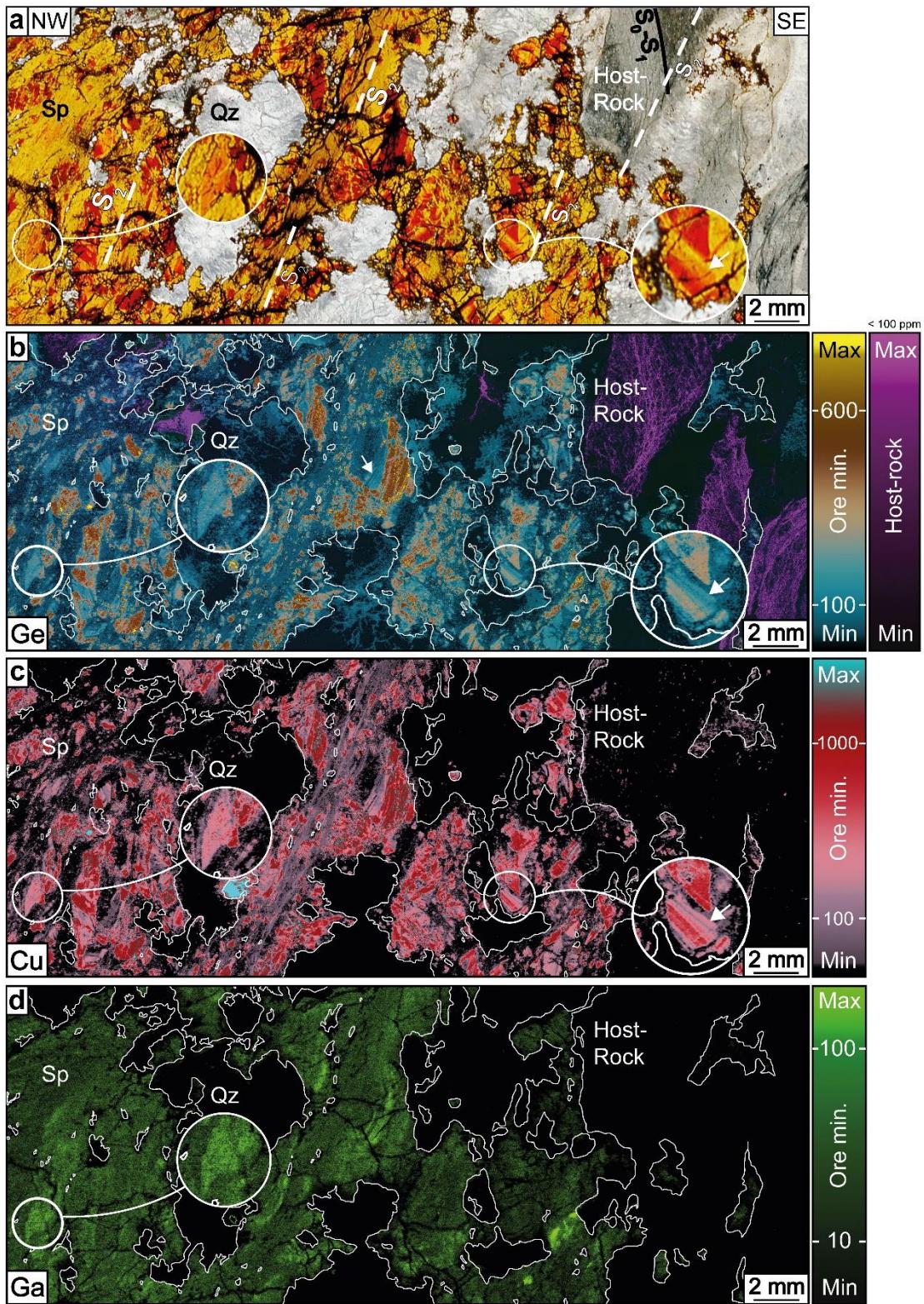


Fig. 8

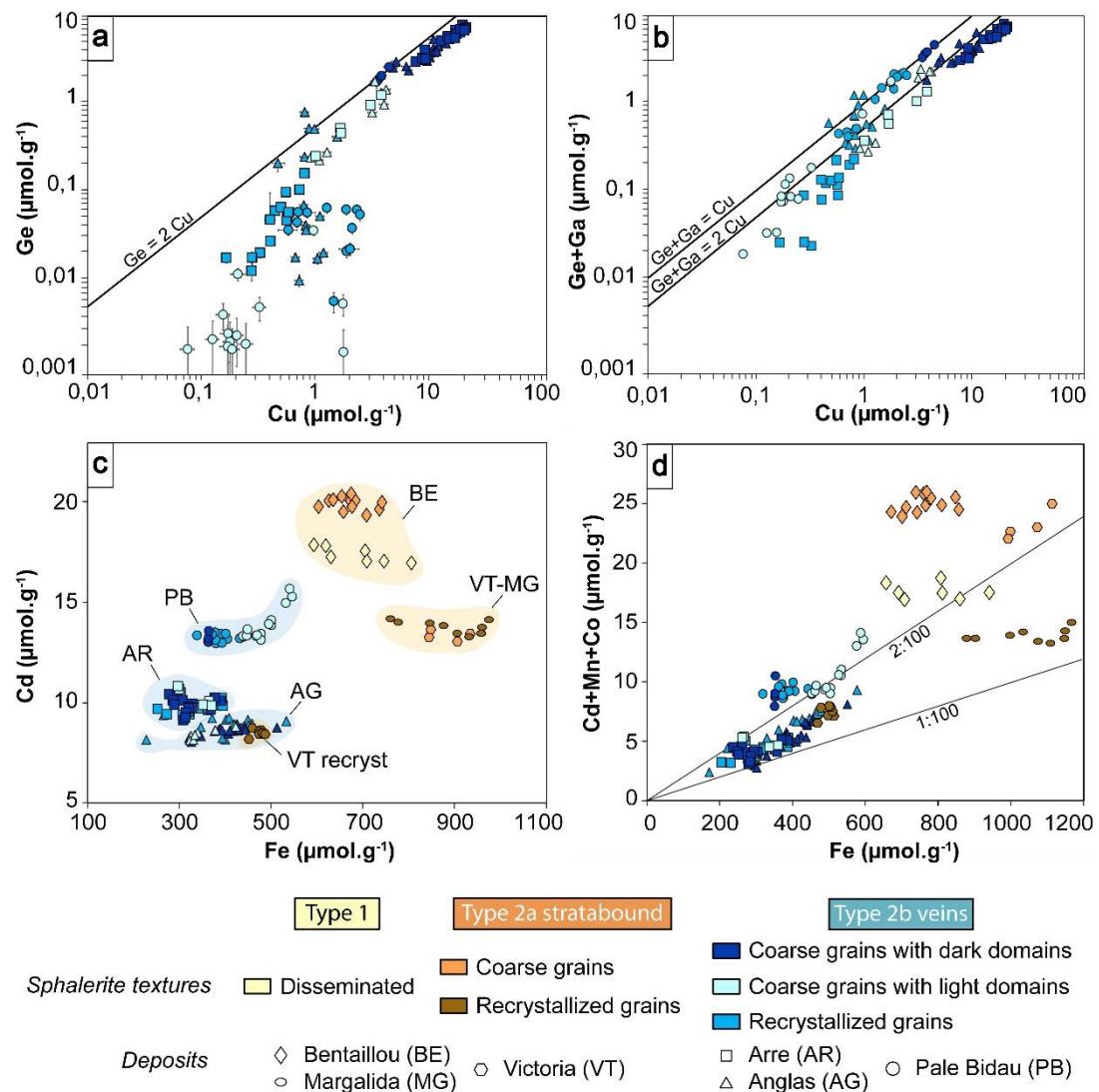


Fig. 9

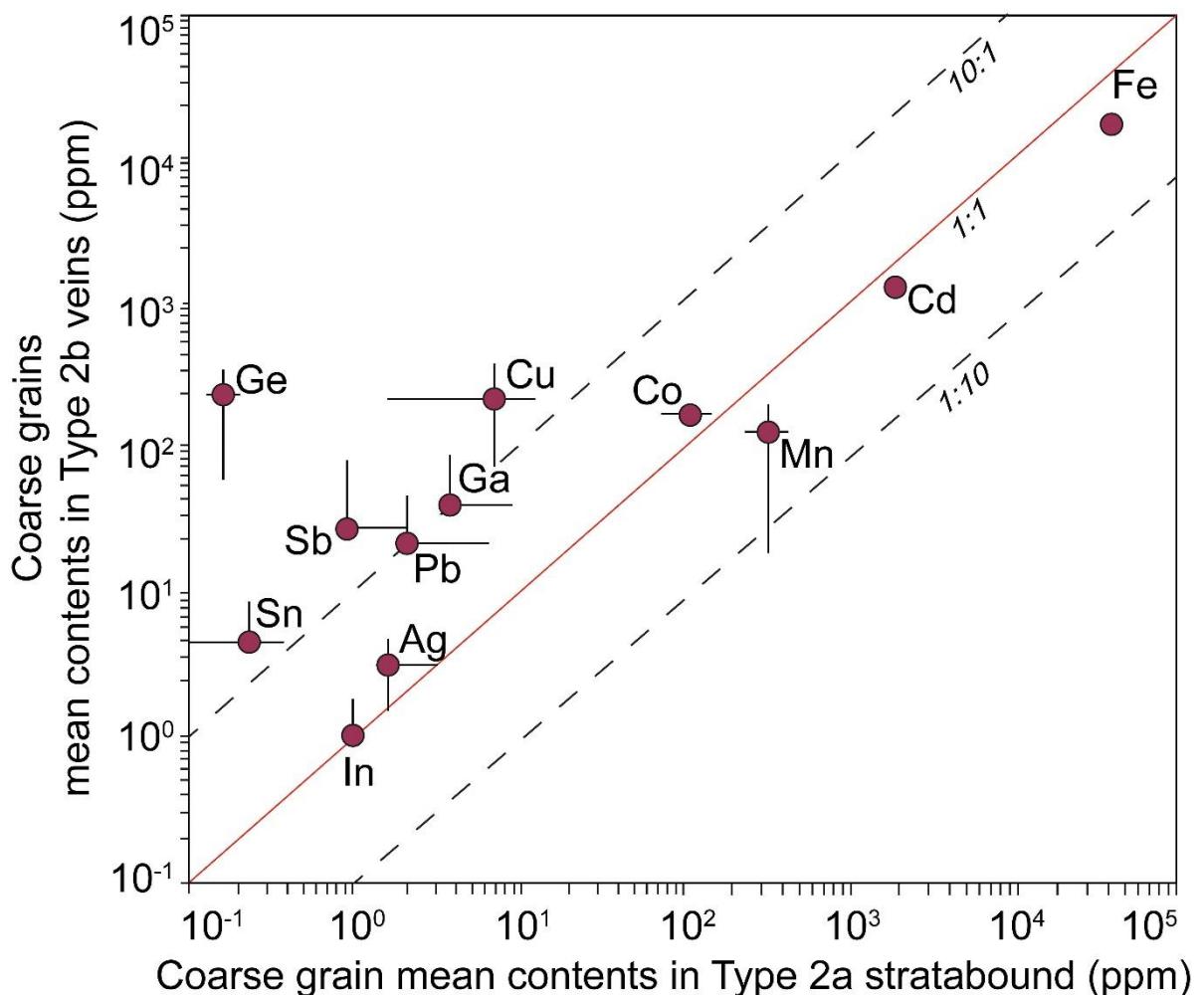


Fig. 10

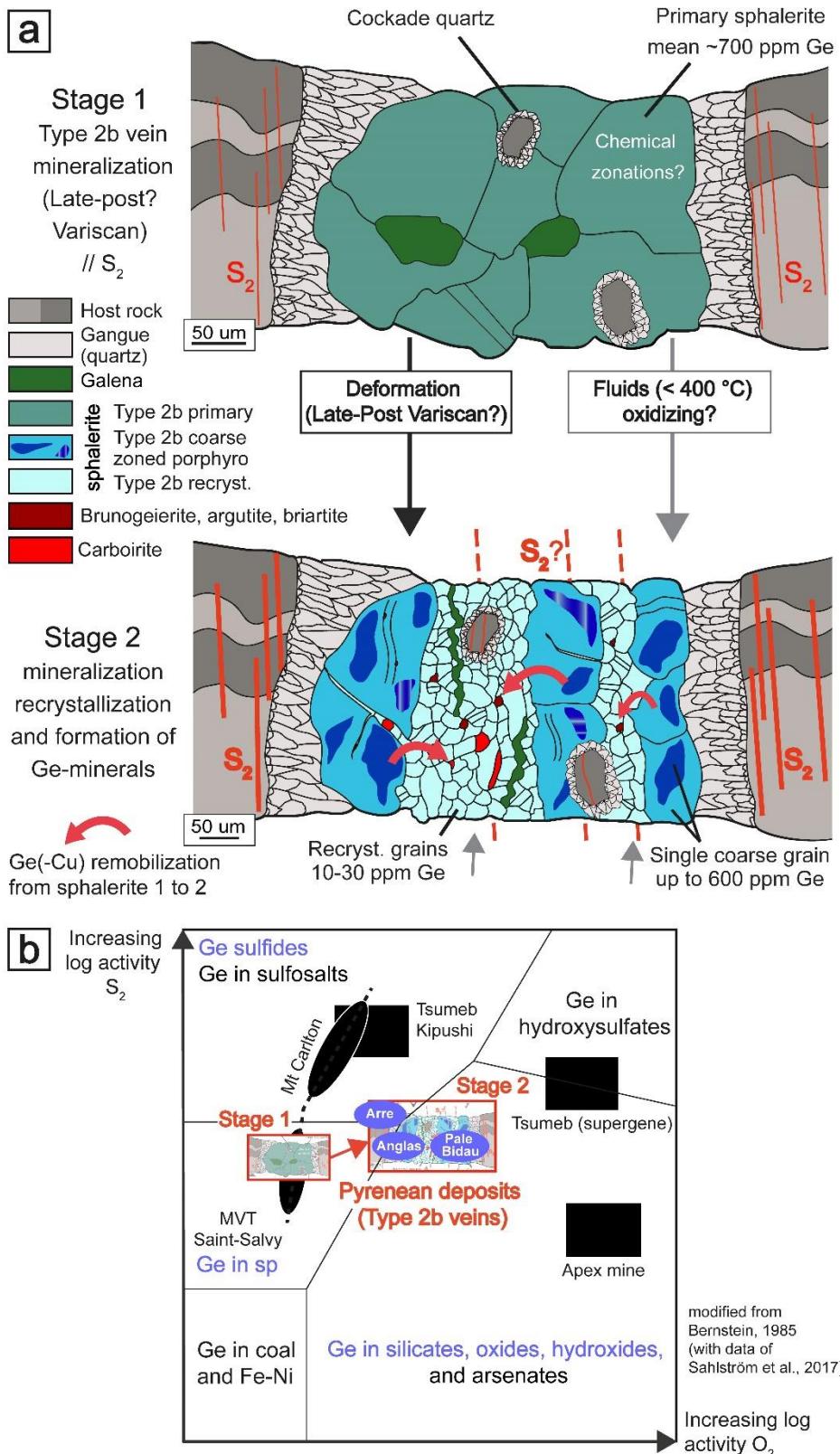


Fig. 11

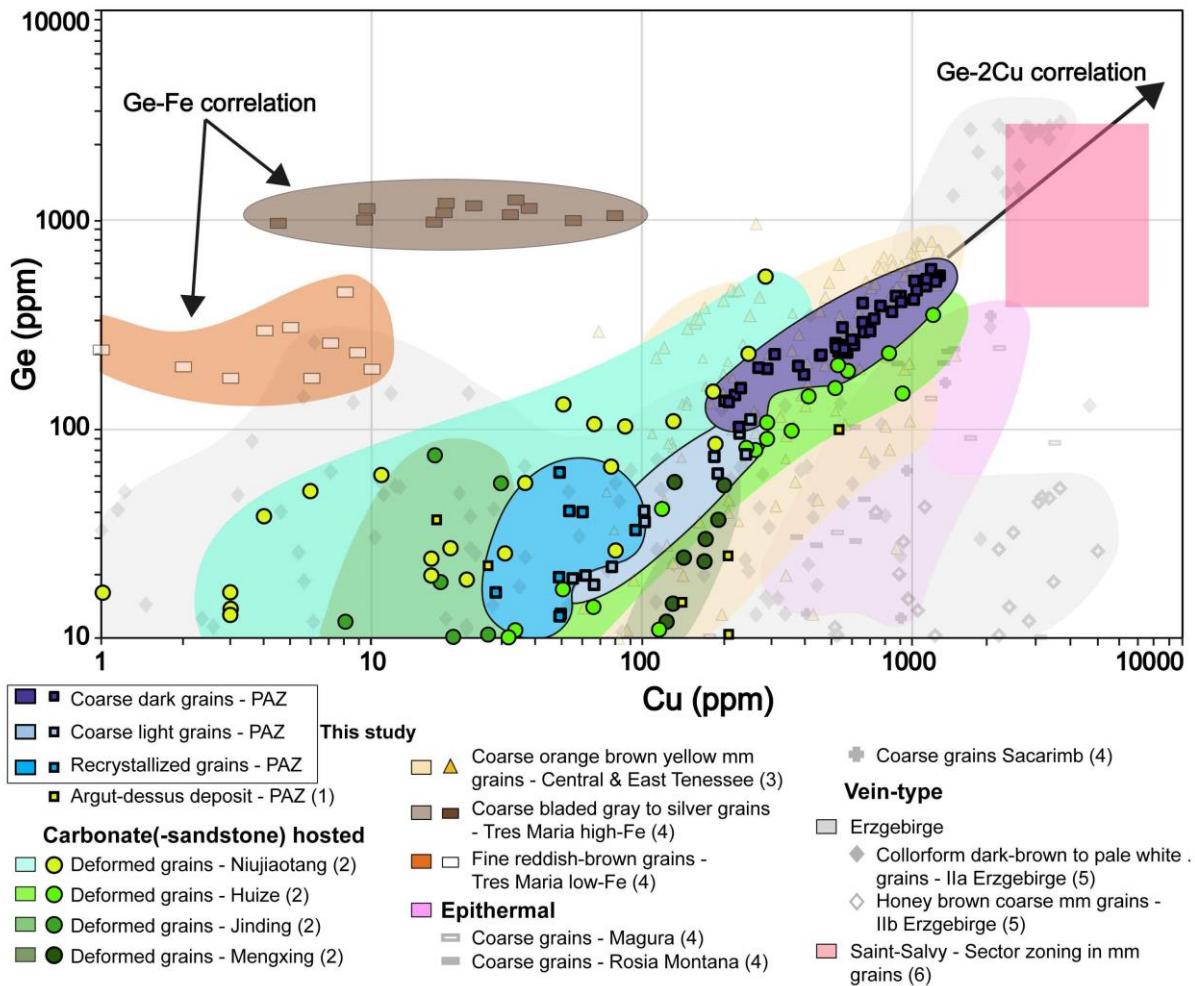


Fig. 12

Electronic supplementary material

ESM1-Table S1

In-situ LA-ICP-MS dataset of the Type 1, Type 2a and Type 2b mineralizations.

Article Title: Behavior of critical metals in metamorphosed Pb-Zn ore deposits: example from the Pyrenean Axial Zone

Journal: Mineralium Deposita

Authors: Alexandre Cugerone, Bénédicte Cenki-Tok, Manuel Muñoz, Kalin Kouzmanov, Emilien Oliot, Vincent Motto-Ros, Elisabeth Le Goff

Affiliation of the corresponding author: Géosciences Montpellier, Université de Montpellier, CNRS, France

e-mail address: alex.cugerone@gmail.com

Deposit	Texture	S34	Mn5	Fe57	C95	Cu63	Zn64	Ga69	Ga71	Ge74	Ag107	Cd111	In115	Sn118	Sb121	Pb208	
Content (ppm)																	
ANGL AS	Recrystallized grains	236	19	25	13	94.	604	32	32	33	2.3	10	1.	1.9	12.	23.	
		254	4.	10	7.	5	453	.4	.7	.0	21	60	1.	6	23.	3	
ANGL AS	Recrystallized grains	261	92	12	10	28.	606	28	25	16	1.9	90	1.	2.8	10.	20.	
		026	.3	73	3.	8	147	.3	.2	.7	5	44	1.	0	20.	2	
ANGL AS	Recrystallized grains	230	23	29	17	53.	605	31	30	40	2.2	10	1.	1.1	8.5	22.	
		995	0.	75	2.	8	813	.6	.5	.7	29	73	1.	22.	7	7	
ANGL AS	Recrystallized grains	240	19	25	14	49.	602	31	30	62	1.7	99	1.	1.0	6.1	20.	
		436	5.	55	2.	4	939	.1	.8	.1	9	84	1.	20.	1	1	
ANGL AS	Recrystallized grains	233	19	25	14	50.	605	40	40	13	1.6	99	1.	0.8	6.5	16.	
		791	7.	17	2.	0	185	.3	.0	.1	7	50	1.	16.	6	6	
ANGL AS	Recrystallized grains	233	20	25	14	60.	608	52	52	40	1.6	10	1.	0.8	6.1	13.	
		964	0.	67	5.	1	009	.3	.6	.2	01	51	1.	13.	7	7	
ANGL AS	Recrystallized grains	227	96	14	10	48.	599	32	32	5.	2.3	10	<L	2.6	6.2	23.	
		425	.9	84	4.	6	784	.7	.4	7	50	O	D	2.6	23.	8	
ANGL AS	Recrystallized grains	234	14	22	15	72.	599	38	38	1.	2.8	10	<L	2.4	7.1	29.	
		058	9.	44	2.	0	784	.7	.3	7	23	O	D	2.4	29.	1	
ANGL AS	Recrystallized grains	242	15	23	16	50.	615	29	29	3.	1.9	10	<L	1.3	9.1	21.	
		049	8.	44	0.	4	660	.3	.4	4	03	O	D	1.3	21.	0	
ANGL AS	Recrystallized grains	244	16	24	16	49.	612	35	34	19	1.9	99	<L	0.9	3.7	13.	
		191	6.	26	5.	4	817	.9	.6	.6	3	O	D	0.9	13.	9	
ANGL AS	Recrystallized grains	234	12	19	13	63.	613	41	41	1.	3.5	98	<L	1.4	5.9	22.	
		517	8.	37	9.	7	270	.8	.3	4	2	O	D	1.4	22.	3	
ANGL AS	Recrystallized grains	248	15	22	14	67.	620	24	24	4.	2.8	10	0.	7.7	12.	38.	
		446	1.	79	3.	2	020	.3	.4	3	32	64	0.	5	38.	3	
ANGL AS	Recrystallized grains	241	13	20	13	50.	615	20	20	3.	2.1	10	0.	7.5	7.4	24.	
		187	9.	74	0.	7	250	.3	.4	0	27	61	0.	24.	7	7	
ANGL AS	Recrystallized grains	250	15	22	13	40.	616	25	24	1.	1.8	91	0.	0.6	5.9	12.	
		083	9.	32	4.	9	078	.0	.5	5	8	47	0.	12.	7	7	
ANGL AS	Recrystallized grains	248	15	22	13	44.	614	24	24	0.	2.2	92	0.	0.7	5.2	22.	
		000	5.	01	1.	3	253	.3	.1	8	9	55	0.	22.	7	7	
ARRE	Recrystallized grains	249	18	16	14	33.	627	9.	9.	7.	5.0	11	0.	1.0	7.8	37.	
		315	.6	91	4.	9	187	4	2	9	66	05	0.	37.	7	7	

ARRE	Recrystallized grains	261 254	18 .0	21 79 3. 8 6	16 3. 8 6	10. 3	623 044	0. 6	0. 5	1. 5	1.3	10 92	0. 03	0.3	2.5	2.9
ARRE	Recrystallized grains	249 144	19 .4	20 26 9. 5 4	16 9. 5 4	20. 2	619 786	0. 3	0. 4	1. 7	2.5	10 57	<L O D	0.4	2.4	9.8
ARRE	Recrystallized grains	258 107	23 .8	22 06 2. 7	17 7 7	35. 0	621 670	2. 9	2. 7	4. 2	4.6	10 98	<L O D	1.2	8.2	30. 6
ARRE	Recrystallized grains	249 629	22 .0	21 91 0. 9 2	17 0. 9 2	34. 7	619 784	5. 4	5. 2	3. 8	3.7	11 28	<L O D	1.1	5.2	25. 5
ARRE	Recrystallized grains	275 506	19 .0	18 96 0. 2	16 4	27. 3	623 593	4. 7	4. 4	4. 9	3.8	10 92	<L O D	0.8	5.3	20. 5
ARRE	Recrystallized grains	271 844	20 .3	18 03 3. 7	15 9 9	30. 8	622 920	5. 0	4. 4	5. 4	4.3	10 87	<L O D	0.6	6.4	23. 4
ARRE	Recrystallized grains	260 366	18 .4	18 04 8. 5	15 8 8	24. 9	622 403	4. 0	3. 8	2. 2	3.7	10 60	<L O D	0.6	4.6	19. 1
ARRE	Recrystallized grains	275 685	16 .2	17 20 5. 6	15 5. 6	17. 1	622 974	5. 9	5. 6	1. 1	1.7	10 36	<L O D	0.6	2.8	10. 5
ARRE	Recrystallized grains	256 573	13 .7	14 11 8. 7	12 8. 2	24. 7	623 857	6. 5	6. 5	3. 9	3.2	10 73	<L O D	0.6	5.9	17. 2
ARRE	Recrystallized grains	261 417	20 .6	17 91 4. 1	15 6. 6	44. 6	625 399	6. 9	6. 5	8. 4	5.8	10 80	<L O D	1.3	9.8	37. 2
ARRE	Recrystallized grains	268 070	18 .5	19 09 5. 7	16 0 0	35. 9	624 462	6. 3	6. 1	4. 7	4.8	10 62	<L O D	1.0	10. 0	31. 4
ARRE	Recrystallized grains	263 132	17 .0	15 25 6. 7	13 2 2	49. 5	624 801	5. 2	5. 2	12 .8	7.0	10 46	<L O D	1.3	12. 0	49. 3
ARRE	Recrystallized grains	274 645	17 .4	20 27 0. 3	17. 2	17. 2	622 447	0. 7	0. 3	1. 5	2.0	11 03	<L O D	0.6	2.4	10. 1
PALE BIDAU	Recrystallized grains	230 212	22 .7	21 34 9. 9	12 9. 9	42. 5	625 499	31 .6	32 .0	3. 7	12. 9	15 26	2. 22	6.1	15. 2	38. 4
PALE BIDAU	Recrystallized grains	226 130	23 .8	22 63 9. 7	13 9. 9	43. 5	624 932	26 .8	27 .4	4. 8	15. 5	15 34	1. 97	7.5	16. 5	40. 1
PALE BIDAU	Recrystallized grains	225 249	25 .1	24 14 5. 5	14 5. 5	52. 1	624 932	33 .9	34 .5	4. 7	16. 0	15 27	2. 00	7.2	16. 5	43. 5
PALE BIDAU	Recrystallized grains	232 305	20 .9	21 80 3. 8	13 0	35. 7	624 932	30 .9	31 .7	3. 0	9.8	15 61	2. 15	5.0	11. 0	27. 1
PALE BIDAU	Recrystallized grains	229 688	21 .4	21 23 3. 6	13 6	77. 5	630 610	76 .2	78 .2	5. 3	13. 8	15 00	1. 04	7.0	16. 7	32. 2

PALE	Recrystallized grains	233 949	27 .1	21 95 9	14 2. 7	11 4.3	629 137	10 1. 6	10 4. 3	5. 1	24. 7	15 01	0. 96	11. 5	34. 6	94 20. 7
PALE	Recrystallized grains	231 902	23 .2	20 60 2	13 5. 1	12 8.3	629 729	14 8. 8	15 2. 8	3. 2	14. 8	15 31	0. 66	5.4 5.4	16. 2	18 72. 3
PALE	Recrystallized grains	230 219	25 .6	21 08 3	14 0. 6	14 1.3	628 380	15 6. 4	15 9. 1	5. 0	19. 3	15 69	0. 51	6.4 5	19. 5	14 4.5
PALE	Recrystallized grains	232 679	22 .9	18 93 8	12 6. 1	11 4.8	625 653	15 3. 9	15 7. 6	1. 7	7.1 46	15 49	0. 2.3	6.5 6.5	18 9.3	
PALE	Recrystallized grains	231 699	20 .9	21 20 1	13 8. 6	88. 8	629 729	10 9. 0	11 1. 6	0. 5	4.8 15	15 73	0. 1.8	2.8 2.8	14. 8	
PALE	Recrystallized grains	220 895	26 .4	22 43 4	14 7. 9	15 0.4	624 172	14 6. 5	15 1. 4	4. 5	21. 2	15 46	0. 54	6.9 6.9	20. 9	41. 6
PALE	Recrystallized grains	234 495	26 .4	20 78 1	14 2. 8	12 3.0	621 274	14 2. 9	14 8. 0	1. 8	9.1 26	15 45	0. 2.8	2.8 8.3	22. 2	
ANGL AS	Dark domains-coarse grains	223 644	15 8. 5	22 73 6	11 6. 1	52 0.3	592 870	11 .2	10 .9	25 7. 6	2.8 2.8	94 4	<L O D	0.4 41. 9	12. 1	
	Dark domains-coarse grains	230 354	16 7. 8	23 30 8	11 7. 9	65 5.8	605 197	17 .0	16 .6	32 4. 5	3.9 3.9	96 8	<L O D	0.4 11 0.4	49. 1	
	Dark domains-coarse grains	226 912	15 1. 2	22 88 7	12 3. 2	57 5.1	611 468	40 .9	41 .0	23 2. 0	3.2 3.2	96 4	<L O D	0.2 10 8.6	51. 7	
	Dark domains-coarse grains	232 250	13 5. 8	21 02 6	11 2. 5	66 0.9	607 757	45 .6	45 .7	28 9. 9	3.8 3.8	95 5	<L O D	0.4 12 8.2	44. 7	
	Dark domains-coarse grains	235 823	15 2. 2	23 88 6	12 6. 5	60 9.5	607 757	42 .2	42 .3	25 3. 1	3.6 3.6	96 2	<L O D	0.3 18 3.2	53. 6	
	Dark domains-coarse grains	233 786	18 2. 9	24 62 1	13 9. 8	39 6.7	605 354	37 .9	38 .3	18 1. 9	2.5 2.5	99 5	2. 43	7.4 7.4	35. 0	21. 1
	Dark domains-coarse grains	234 014	17 9. 3	23 99 1	13 7. 3	59 5.5	603 422	23 .9	24 .2	26 8. 1	3.1 3.1	97 7	3. 31	21. 0	96. 7	45. 1

ANGL AS	Dark domains-coarse grains	233 053	17. 5	24 03	13. 7.	37 8.8	603 906	13 .7	13 .4	20 0. 2	2.3	96 7	2. 80	20. 4	32. 6	14. 1
ANGL AS	Dark domains-coarse grains	238 896	18. 3	24 76	13. 9.	30 8.8	603 777	18 .8	18 .9	22 8. 0	2.8	97 3	2. 25	18. 6	14. 4	20. 4
ANGL AS	Dark domains-coarse grains	237 785	21. 9	28 64	15. 6.	29 1.0	606 726	24 .4	24 .7	19 4. 0	2.2	98 2	1. 88	6.6 32. 5	32. 16. 9	
ANGL AS	Dark domains-coarse grains	234 951	13. 2.	21 61	15. 9.	72 0.9	610 619	68 .6	67 .7	33 7. 8	9.5	97 5	0. 38	4.4 14. 2	14. 65. 1	
ANGL AS	Dark domains-coarse grains	244 286	13. 4.	21 53	14. 7.	45 5.2	619 818	12 8. 6	12 8. 3	22 4. 7	4.6	96 9	1. 80	38. 5	2.0 11. 3	
ANGL AS	Dark domains-coarse grains	246 020	14. 7.	21 76	13. 5.	53 7.7	612 171	38 .9	39 .0	25 2. 0	3.3	97 2	0. 37	3.2 99. 4	46. 9	
ANGL AS	Dark domains-coarse grains	243 576	14. 6.	21 97	13. 3.	53 3.5	612 853	19 .4	20 .0	23 5. 1	3.2	96 5	<L O D	2.2 14. 5.0	78. 3	
ANGL AS	Dark domains-coarse grains	249 023	17. 3.	23 77	13. 4.	22 8.4	609 509	35 .7	35 .4	10 2. 3	2.8	10 03	<L O D	0.5 7.4	2.8	
ANGL AS	Dark domains-coarse grains	245 626	81. .9	17 93	13. 1.	69 8.0	618 606	20 .6	20 .5	29 4. 1	3.5	92 5	<L O D	0.2 21. 7.2	10 8.0	
ANGL AS	Dark domains-coarse grains	250 329	83. .9	18 20	13. 2.	84 0.6	621 169	30 .4	30 .0	36 4. 4	3.8	90 5	<L O D	0.1 21. 8.4	89. 7	
ANGL AS	Dark domains-coarse grains	248 434	10. 3.	19 35	13. 4.	65 3.5	619 505	72 .2	72 .4	40 1. 1	3.9	93 6	0. 08	3.3 54. 4	16. 0	
ANGL AS	Dark domains-coarse grains	247 319	15. .9	15 51	13. 3.	55 2.7	627 486	7. 8	7. 8	30 5. 4	3.2	11 55	<L O D	3.7 3.6	16. 1	

ANGL AS	Dark domains- coarse grains	245 895	18 .1	16 52	14 0.	10 19.	624 812	8. 3	7. 9	51 1. 2	3.2	11 51	<L O D	3.1	2.6	9.5
ANGL AS	Dark domains- coarse grains	250 692	17 .4	16 54	13 8.	90 5.6	622 664	5. 6	5. 6	43 2. 8	2.9	11 53	0. 05	3.0	2.1	10. 6
ANGL AS	Dark domains- coarse grains	251 376	15 .3	17 15	13 5.	46 1.3	621 952	4. 9	4. 9	22 7. 1	2.6	11 21	<L O D	0.2	7.0	2.2
ANGL AS	Dark domains- coarse grains	249 074	16 .7	18 30	14 7.	52 2.9	620 573	6. 1	5. 6	24 6. 6	2.6	10 99	<L O D	0.3	8.7	2.3
ANGL AS	Dark domains- coarse grains	261 244	18 .4	20 55	16 4.	55 8.6	626 458	0. 5	0. 1	24 1. 1	2.2	10 85	<L O D	0.2	1.2	1.5
ANGL AS	Dark domains- coarse grains	259 861	19 .6	21 79	17 0.	11 20.	625 362	4. 5	4. 6	48 0. 9	2.5	11 16	<L O D	5.3	2.8	7.2
ANGL AS	Dark domains- coarse grains	248 541	14 .4	15 97	13 9.	76 1.9	620 538	7. 8	7. 7	38 6. 9	2.2	11 21	<L O D	6.4	1.8	6.3
ANGL AS	Dark domains- coarse grains	253 391	13 .1	16 17	14 3.	11 75.	623 367	8. 3	8. 3	57 9. 6	2.5	10 89	<L O D	1.0	5.1	3.2
ANGL AS	Dark domains- coarse grains	257 654	15 .7	18 34	16 1.	12 65.	619 794	6. 9	6. 6	54 0. 8	3.0	10 96	0. 06	1.6	3.1	4.1
ANGL AS	Dark domains- coarse grains	252 882	15 .4	17 87	15 7.	12 25.	617 456	9. 7	9. 6	52 5. 1	2.4	10 71	<L O D	2.9	1.9	3.4
ANGL AS	Dark domains- coarse grains	254 546	16 .2	17 89	15 9.	10 40.	620 317	10 .3	10 .0	46 0. 7	4.6	10 89	<L O D	10. 0	5.6	24. 8
ANGL AS	Dark domains- coarse grains	266 281	15 .1	17 56	15 8.	86 9.2	620 189	11 .3	10 .5	43 1. 7	3.4	10 51	<L O D	8.8	2.6	15. 5

ANGL AS	Dark domains-coarse grains	274 499	15 .5	17 23 6	15 5. 0	11 27. 5	618 876	5. 2	5. 2	51 9. 8	2.8	10 09	<L O D	0.4	4.4	3.8
ANGL AS	Dark domains-coarse grains	270 724	14 .8	17 48 5	15 3. 7	90 8.5	618 790	5. 4	5. 1	40 4. 3	3.3	10 11	<L O D	2.5	2.8	7.6
ANGL AS	Dark domains-coarse grains	291 187	18 .0	20 58 5	16 0. 0	12 22. 7	625 592	6. 0	5. 6	50 7. 3	6.1	10 73	<L O D	0.6	6.0	5.0
ANGL AS	Dark domains-coarse grains	260 770	19 .0	21 08 8	16 4. 9	10 12. 4	625 253	5. 0	4. 3	41 5. 2	4.8	11 15	0. 04	0.5	12. 4	5.9
PALE BIDAU	Dark domains-coarse grains	228 422	20 .8	20 31 1	13 1. 5	22 0.5	629 729	11. 9. 7	12 2. 9	14 5. 9	7.0	14 98	0. 78	7.4	4.3	20. 9
PALE BIDAU	Dark domains-coarse grains	228 410	18 .6	20 30 0	11 0. 3	20 9.3	629 718	10 9. 3	11 1. 7	13 4. 7	4.3	14 87	0. 21	5.0	4.9	9.7
PALE BIDAU	Dark domains-coarse grains	228 433	25 .3	22 32 2	13 2. 6	23 1.6	629 740	12 4. 3	13 4. 0	15 7. 0	7.2	14 09	0. 52	27. 3	5.3	32. 0
PALE BIDAU	Dark domains-coarse grains	228 473	22 .3	50 36 2	17 2. 7	26 9.3	629 780	14 2. 3	17 4. 1	19 7. 1	8.5	18 49	0. 42	3.0	8.3	48. 3
ANGL AS	Light domains-coarse grains	240 197	12 4. 7	21 11 4	15 6. 2	25 1.9	615 298	60 .4	60 .9	11 1. 0	4.0	96 7	0. 21	3.8	4.2	33. 5
ANGL AS	Light domains-coarse grains	246 638	88 .7	18 97 3	14 8. 9	24 2.4	615 787	95 .3	95 .5	75 .8	3.7	95 1	0. 21	11. 8	5.7	3.7
ANGL AS	Light domains-coarse grains	244 049	77 .9	18 12 6	14 2. 2	19 0.5	622 179	82 .9	82 .2	61 .2	1.5	93 3	0. 02	0.9	5.1	1.9
ANGL AS	Light domains-coarse grains	243 835	17 3. 1	23 58 5	13 3. 0	20 1.7	611 658	47 .6	47 .4	13 6. 7	1.7	10 02	0. 02	0.9	11. 8	3.8

ANGL AS	Light domains-coarse grains	251 83 18 13 55. 620 4. 4. 19 3.1 91 0. 0. 0.1 2.2 15. 1 632 .9 01 3. 6 287 5 2 .4 3.1 91 2 02 0.1 2.2 15. 1 6 8
ANGL AS	Light domains-coarse grains	250 90 18 13 77. 623 5. 5. 22 5.3 91 0. 0. 0.4 3.9 47. 6 491 .7 45 1. 8 235 4 3 .1 5.3 91 7 02 0.4 3.9 47. 6 0
ANGL AS	Light domains-coarse grains	249 91 18 13 66. 622 3. 3. 18 4.2 94 0. 0. 0.4 3.9 40. 3 531 .3 62 4. 7 196 9 7 .1 4.2 94 6 02 0.4 3.9 40. 3 4
ARRE	Light domains-coarse grains	241 18 16 15 10 624 15 15 40 4.5 11 0. 0. 1.3 5.6 31. 1 439 .9 77 0. 5 1.6 619 .6 .5 .7 4.5 11 81 03 1.3 5.6 31. 1 2
ARRE	Light domains-coarse grains	251 19 16 13 61. 624 8. 8. 20 3.8 12 0. 0. 1.1 5.1 27. 6 389 .2 60 6. 7 557 8 5 .0 3.8 12 02 0. 1.1 5.1 27. 6 0 8
ARRE	Light domains-coarse grains	247 18 20 15 18 622 6. 6. 73 2.2 11 0. 0. 0.2 8.7 2.4 688 .4 40 8. 4 5.5 585 9 8 .7 2.2 11 14 02 0.2 8.7 2.4 6
ARRE	Light domains-coarse grains	276 19 19 16 10 623 8. 8. 36 3.4 10 0. 0. 1.2 5.0 20. 1 263 .2 67 2. 8 083 4 6 .0 3.4 10 95 07 1.2 5.0 20. 1 8 4
ARRE	Light domains-coarse grains	278 18 20 17 22 625 7. 6. 95 1.4 10 0. 0. 1.6 1.1 2.4 605 .9 69 1. 2 9.7 253 4 0 .5 1.4 10 85 02 1.6 1.1 2.4 2 4
PALE BIDAU	Light domains-coarse grains	227 22 26 14 20. 626 13 13 0. 4.3 15 4. 2.3 4.4 12. 1 075 .7 71 2. 0 392 .6 .8 4 4.3 15 86 40 2.3 4.4 12. 1 0 6
PALE BIDAU	Light domains-coarse grains	225 26 26 14 11. 629 5. 5. 0. 1.5 15 6. 1.1 0.8 3.2 048 .0 05 1. 8 854 7 5 2 1.5 15 85 21 1.1 0.8 3.2 1 7
PALE BIDAU	Light domains-coarse grains	218 24 24 13 7.8 629 2. 2. 0. 1.6 15 6. 1.0 0.8 3.0 609 .1 58 0. 4 646 4 2 2 1.6 15 84 09 1.0 0.8 3.0 5 5
PALE BIDAU	Light domains-coarse grains	223 22 24 13 9.7 629 2. 2. 0. 2.5 15 5. 1.7 2.0 8.3 587 .8 54 2. 0 914 3 2 3 2.5 15 73 34 1.7 2.0 8.3 0 5

PALE BIDAU	Light domains- coarse grains	222 428	23 .1	26 0	14 3	4.7 3	625 508	1. 4	1. 3	0. 1	1.3 1.3	15 58	2. 57	1.0 1.0	0.3 0.3	1.9 1.9
PALE BIDAU	Light domains- coarse grains	224 490	24 .4	24 95	13 6	13. 6	626 989	5. 7	5. 4	1. 0	3.4 3.4	15 83	3. 83	2.2 2.2	3.3 3.3	12. 2
PALE BIDAU	Light domains- coarse grains	227 441	23 .4	25 06	13 7.	12. 7	627 549	10. .5	10. .5	0. 2	1.8 1.8	16 02	4. 27	1.2 1.2	1.2 1.2	4.3 4.3
PALE BIDAU	Light domains- coarse grains	226 121	27 .4	27 95	15 4.	15. 2	626 354	6. 1	6. 0	0. 2	3.3 3.3	16 35	4. 70	1.8 1.8	2.7 2.7	10. 6
PALE BIDAU	Light domains- coarse grains	229 276	29 .7	27 97	15 7.	10. 6	623 396	5. 7	5. 4	0. 2	1.9 1.9	16 58	3. 68	1.4 1.4	1.0 1.0	5.4 5.4
PALE BIDAU	Light domains- coarse grains	233 230	27 .5	27 68	15 4.	10. 7	626 797	6. 5	6. 3	0. 2	1.5 1.5	16 34	4. 63	1.0 1.0	0.9 0.9	3.4 3.4
PALE BIDAU	Light domains- coarse grains	225 045	24 .4	26 24	14 7.	11. 6	623 429	9. 0	9. 2	0. 1	1.4 1.4	15 77	3. 87	0.7 0.7	0.6 0.6	2.6 2.6
PALE BIDAU	Light domains- coarse grains	229 823	34 .3	29 74	17 5.	59. 1	622 995	53 .0	53 .8	3. 0	9.1 9.1	17 56	0. 13	4.7 4.7	7.2 7.2	24. 2
PALE BIDAU	Light domains- coarse grains	227 187	31 .5	45 50	18 4.	10 6.	619 529	12. 5.	12. 7.	0. 5	2.0 2.0	19 87	0. 37	0.7 0.7	0.7 0.7	3.7 3.7
PALE BIDAU	Light domains- coarse grains	230 037	32 .3	30 20	18 3.	10 8.0	617 428	12. 7	13. 8	0. 1	2.0 2.0	18 27	0. 36	0.8 0.8	0.5 0.5	3.2 3.2
VICTO RIA	Recrystallized grains	243 037	25 6.	26 50	92 .3	12. 1	611 699	1. 2	0. 8	<L O	8.5 8	91 9	0. 60	1.8 1.8	3.7 3.7	41. 8
VICTO RIA	Recrystallized grains	249 431	25 2.	25 23	89 .8	7.8 7.8	608 276	1. 2	1. 0	<L O	5.4 3	91 52	0. 52	1.4 1.4	3.5 3.5	27. 0
VICTO RIA	Recrystallized grains	248 918	26 7.	26 75	90 .1	18. 6	606 274	0. 8	0. 6	<L O	13. 8	94 9	0. 55	2.4 2.4	7.1 7.1	70. 9

VICTORIA	Recrystallized grains	245 215	26 2. 5	25 63 0	91 .7 5	16. 8.9	610 786	0. 9	0. 6	<L O D	12. 2	94 5	0. 55	2.3 1.4	6.6 2.6	62. 9
VICTORIA	Recrystallized grains	242 942	26 0. 0	27 06 9	88 .4	8.9	607 974	1. 3	1. 2	<L O D	4.9	94 0	0. 52	1.4 1.4	2.6 2.6	25.
VICTORIA	Recrystallized grains	256 952	26 7. 7	26 93 0	91 .3	15. 1	610 250	0. 9	0. 7	<L O D	11. 1	94 3	0. 59	2.5 2.5	6.2 7.8	56. 73.
VICTORIA	Recrystallized grains	260 557	26 8. 9	26 56 8	98 .1	23. 5	608 419	1. 4	1. 2	<L O D	14. 6	94 8	0. 62	2.5 2.5	7.8 7.8	73. 73.
VICTORIA	Coarse grains	242 840	47 2. 7	47 42 6	16 4. 7	10. 5	577 633	9. 8	9. 7	<L O D	1.2	15 23	0. 76	0.4 0.4	0.2 0.2	2.5
VICTORIA	Coarse grains	246 682	46 8. 9	47 19 1	16 6. 5	20. 1	576 163	11 .3	10 .6	<L O D	4.5	14 81	0. 68	0.6 0.6	2.2 2.2	18. 18.
VICTORIA	Coarse grains	251 850	50 4. 7	50 53 6	17 8. 1	15. 0	576 952	15 .0	14 .7	<L O D	0.9	14 34	0. 74	0.2 0.2	<L O D	0.6
VICTORIA	Coarse grains	247 505	55 2. 0	52 25 1	18 3. 9	16. 4	578 123	14 .5	14 .2	<L O D	1.2	14 95	0. 79	0.1 0.1	0.2 0.2	2.3
MARG ALIDA	Recrystallized grains	227 831	22 9. 8	54 42 .7	89 .4	10. 8	582 370	2. 8	2. 7	<L O D	9.9	15 51	<L O D	0.4	7.0	8.0
MARG ALIDA	Recrystallized grains	230 817	23 0. 3	53 55 2	84 .7	5.1	589 102	1. 8	2. 1	<L O D	1.8	14 92	<L O D	0.1	0.2	1.5
MARG ALIDA	Recrystallized grains	232 839	23 6. 2	53 61 7	83 .7	6.3	589 342	2. 0	1. 8	<L O D	2.3	15 11	<L O D	0.2	0.2	1.1
MARG ALIDA	Recrystallized grains	238 177	20 9. 7	50 56 3	80 .6	25. 1	590 879	4. 4	5. 1	<L O D	23. 2	14 85	<L O D	0.7	17.	17.
MARG ALIDA	Recrystallized grains	241 715	21 6. 2	52 05 0	81 .5	17. 2	584 966	5. 4	5. 4	<L O D	13. 4	14 89	<L O D	0.4	8.6	11.
MARG ALIDA	Recrystallized grains	245 613	22 6. 8	48 86 0	79 .3	17. 9	584 856	5. 1	5. 5	<L O D	11. 5	15 42	<L O D	0.4	6.4	11.
MARG ALIDA	Recrystallized grains	246 778	20 8. 1	47 23 0	77 .4	14. 5	585 846	1. 8	1. 8	<L O D	19. 2	15 20	<L O D	0.6	14.	81.
MARG ALIDA	Recrystallized grains	252 659	17 7. 5	42 42 1	70 .3	22. 1	583 325	0. 5	<L O D	42. 6	15 78	<L O D	1.8	44.	4.5	
MARG ALIDA	Recrystallized grains	255 357	20 6. 9	43 39 7	73 .4	9.9	583 325	3. 4	3. 2	<L O D	3.7	15 68	<L O D	0.3	1.0	3.7
BENT AILLOU	Coarse grains	250 926	28 4. 6	37 70 9	94 .5	4.6	593 659	1. 0	0. 9	<L O D	0.8	21 65	1. 07	0.2	<L O D	<L OD

Detection limit (ppm)																		
BENT AILLO U	Coarse grains	251 532	28 9.	38 06	94 .2	5.4	591 000	1. 3	1. 2	<L O D	0.8	21 41	1. 02	0.2	0.2	0.2		
BENT AILLO U	Coarse grains	250 142	28 8. 5	36 51 2	95 .9	8.8	593 758	0. 9	0. 8	<L O D	6.3	21 72	1. 07	0.4	3.6	5.4		
BENT AILLO U	Coarse grains	254 588	29 9. 9	37 48 6	95 .7	4.8	591 526	1. 0	1. 0	<L O D	0.8	21 54	1. 08	0.2	0.1	0.2		
BENT AILLO U	Coarse grains	251 468	30 6. 7	41 09 .6	99 4. .6	4.2	593 349	1. 4	1. 3	<L O D	0.7	21 26	1. 10	0.2	<L O D	<L OD		
BENT AILLO U	Coarse grains	247 818	27 7. 6	36 67 4	90 .7	3.9	593 520	1. 3	1. 3	<L O D	0.7	21 17	1. 10	0.2	<L O D	<L OD		
BENT AILLO U	Coarse grains	246 018	28 2. 0	37 67 6	93 .5	7.6	593 028	0. 7	0. 6	<L O D	3.8	21 42	1. 09	0.4	7.3	2.6		
BENT AILLO U	Coarse grains	253 104	30 4. 7	39 49 1	96 .6	3.0	590 901	0. 6	0. 7	<L O D	0.8	21 02	1. 10	0.2	<L O D	<L OD		
BENT AILLO U	Coarse grains	254 077	26 3. 2	41 44 6	87 .0	2.5	589 281	0. 6	0. 6	<L O D	0.8	21 61	1. 07	0.0	0.1	0.1		
BENT AILLO U	Coarse grains	253 059	26 5. 6	35 36 6	84 .2	3.0	592 289	0. 6	0. 4	<L O D	0.9	21 88	1. 05	0.2	0.4	0.6		
BENT AILLO U	Coarse grains	250 675	24 3. 5	34 99 6	80 .3	2.5	589 816	0. 9	0. 8	<L O D	0.8	21 87	1. 03	0.1	0.3	0.8		
BENT AILLO U	Coarse grains	252 247	27 3. 0	33 72 0	80 .9	2.5	590 923	0. 9	0. 7	<L O D	0.8	21 93	1. 07	0.2	0.1	0.2		
BENT AILLO U	Type 1 dissemina ted	209 949	17 1. 1	39 40 4	10 0. 4	12. 0	592 975	28 .8	2. 9	<L O D	4.6	19 04	0. 98	1.2	7.1	50 3.2		
BENT AILLO U	Type 1 dissemina ted	203 360	13 8. 4	34 55 0	72 .8	7.6	584 228	3. 5	0. 7	<L O D	2.9	19 31	0. 91	0.9	3.9	16 9.1		
BENT AILLO U	Type 1 dissemina ted	204 455	16 3. 1	39 55 6	87 .7	69. 8	589 383	10 .5	2. 5	<L O D	10. 2	18 33	0. 88	3.7	23. 8	10 6.2		
BENT AILLO U	Type 1 dissemina ted	206 852	16 2. 1	44 96 4	95 .9	23. 9	588 310	6. 4	1. 5	<L O D	7.6	18 37	0. 94	2.5	13. 2	33 1.6		
BENT AILLO U	Type 1 dissemina ted	213 062	13 8. 4	35 20 2	80 .0	39. 8	591 454	19 .4	4. 0	<L O D	10. 6	18 66	0. 84	3.2	16. 3	15 0.3		
BENT AILLO U	Type 1 dissemina ted	200 802	15 9. 2	33 09 6	81 .5	14. 8	589 383	1. 7	0. 8	<L O D	2.5	19 25	0. 96	0.4	3.2	17 2.7		
BENT AILLO U	Type 1 dissemina ted	207 419	14 0. 9	41 61 1	89 .1	37. 5	589 383	13 .7	3. 5	<L O D	12. 0	18 94	0. 91	3.0	18. 4	43 6.5		

Deposit	Texture	S34	Mn5	Fe57	C059	Cu63	Zn64	Ga69	Ga71	Ge74	Ag107	Cd111	In115	Sn118	Sb121	Pb208
ANGL AS	Recrystallized grains	86.5	0.20	12.90	0.27	0.16	0.73	0.07	0.03	0.15	0.02	0.11	0.01	0.05	0.01	0.00
ANGL AS	Recrystallized grains	152.7.2	3.70	23.3.61	4.63	2.86	12.42	1.31	0.47	2.71	0.2	0.30	0.23	0.03	0.88	0.20
ANGL AS	Recrystallized grains	52.3	0.13	7.95	0.16	0.10	0.42	0.04	0.02	0.09	0.1	0.7	0.01	0.3	0.1	0.03
ANGL AS	Recrystallized grains	79.8	0.19	11.94	0.24	0.15	0.65	0.07	0.03	0.14	0.2	0.1	0.01	0.4	0.1	0.04
ANGL AS	Recrystallized grains	94.4	0.23	14.12	0.29	0.17	0.82	0.08	0.03	0.17	0.2	0.3	0.02	0.5	0.1	0.05
ANGL AS	Recrystallized grains	97.4	0.23	14.47	0.30	0.18	0.83	0.08	0.03	0.17	0.2	0.4	0.02	0.5	0.1	0.05
ANGL AS	Recrystallized grains	182.9	0.44	27.67	0.56	0.33	1.58	0.15	0.06	0.32	0.4	0.5	0.03	0.0	0.3	0.10
ANGL AS	Recrystallized grains	90.5	0.22	13.95	0.29	0.17	0.82	0.08	0.03	0.16	0.2	0.3	0.01	0.5	0.1	0.05
ANGL AS	Recrystallized grains	91.8	0.23	14.20	0.29	0.17	0.99	0.08	0.03	0.16	0.2	0.5	0.02	0.5	0.1	0.05
ANGL AS	Recrystallized grains	78.0	0.19	12.03	0.25	0.15	0.79	0.07	0.03	0.14	0.2	0.2	0.01	0.4	0.1	0.04
ANGL AS	Recrystallized grains	133.0	0.33	20.67	0.43	0.26	1.25	0.12	0.05	0.24	0.3	0	0.02	0.7	0.2	0.07
ANGL AS	Recrystallized grains	94.8	0.24	14.56	0.32	0.18	1.00	0.09	0.03	0.17	0.2	0.6	0.02	0.5	0.1	0.05
ANGL AS	Recrystallized grains	112.4	0.28	17.32	0.38	0.21	1.22	0.10	0.04	0.20	0.2	0.8	0.02	0.6	0.2	0.06
ANGL AS	Recrystallized grains	102.7	0.25	16.06	0.35	0.20	1.05	0.10	0.04	0.19	0.2	0.6	0.02	0.5	0.2	0.05
ANGL AS	Recrystallized grains	99.8	0.25	15.42	0.34	0.19	1.05	0.09	0.03	0.18	0.2	0.6	0.02	0.5	0.2	0.05
ARRE	Recrystallized grains	91.1	0.17	15.43	0.41	0.21	0.47	0.11	0.02	0.17	0.3	0	0.02	0.6	0.1	0.06
ARRE	Recrystallized grains	69.7	0.13	12.17	0.38	0.18	0.61	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0
ARRE	Recrystallized grains	72.9	0.13	12.74	0.39	0.09	0.67	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0
ARRE	Recrystallized grains	73.2	0.13	12.72	0.39	0.09	0.67	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0
ARRE	Recrystallized grains	63.3	0.11	10.73	0.34	0.06	0.59	0.09	0.02	0.12	0.2	0.1	0.01	0.4	0.1	0.05
ARRE	Recrystallized grains	76.4	0.14	13.45	0.43	0.01	0.44	0.12	0.03	0.14	0.3	0.4	0.02	0.5	0.2	0.05
ARRE	Recrystallized grains	83.7	0.15	14.74	0.47	0.03	0.22	0.13	0.03	0.16	0.3	0.5	0.02	0.5	0.2	0.06
ARRE	Recrystallized grains	76.2	0.13	13.32	0.43	0.01	0.55	0.12	0.03	0.14	0.3	0.5	0.02	0.5	0.2	0.06

ARRE	Recrystallized grains	81.4	0. 15. 14. 97. 0. 48. 2. 3. 0. 5. 0. 13. 0. 03. 0. 16. 0. 3. 0. 6. 0. 02. 0. 5. 0. 2. 0. 06
ARRE	Recrystallized grains	86.2	0. 08. 13. 79. 4. 91. 2. 6. 0. 5. 0. 42. 0. 38. 0. 36. 0. 0. 0. 3. 0. 1. 0. 0. 0. 1. 0. 4. 2. 2. 60
ARRE	Recrystallized grains	79.9	0. 14. 13. 87. 0. 45. 2. 0. 0. 12. 0. 03. 15. 3. 0. 5. 0. 0. 0. 1. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
ARRE	Recrystallized grains	72.6	0. 13. 12. 55. 0. 40. 0. 4. 0. 11. 0. 02. 13. 3. 3. 4. 0. 0. 0. 1. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
ARRE	Recrystallized grains	72.4	0. 10. 12. 45. 5. 28. 7. 3. 0. 39. 0. 35. 86. 4. 5. 0. 0. 0. 1. 0. 0. 1. 0. 4. 5. 1. 70
ARRE	Recrystallized grains	81.7	0. 14. 14. 47. 0. 45. 2. 3. 0. 13. 0. 03. 15. 3. 0. 6. 0. 0. 0. 1. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0
PALE	Recrystallized grains	84.4	0. 19. 14. 08. 0. 24. 1. 5. 0. 9. 0. 08. 0. 03. 17. 2. 0. 4. 0. 01. 4. 1. 0. 0. 0. 0. 0. 0. 0. 0. 0
BIDAU	Recrystallized grains	69.7	0. 16. 11. 47. 0. 20. 2. 2. 0. 7. 0. 07. 0. 03. 14. 1. 2. 0. 01. 4. 1. 04
PALE	Recrystallized grains	57.6	0. 13. 9. 4. 0. 17. 0. 6. 0. 05. 0. 02. 12. 1. 0. 0. 01. 3. 1. 03
BIDAU	Recrystallized grains	81.0	0. 19. 13. 28. 0. 23. 4. 4. 0. 08. 0. 03. 16. 2. 0. 4. 0. 01. 4. 1. 04
PALE	Recrystallized grains	72.5	0. 16. 11. 87. 0. 21. 3. 7. 0. 07. 0. 03. 14. 1. 3. 0. 01. 4. 1. 04
BIDAU	Recrystallized grains	63.6	0. 14. 10. 34. 0. 18. 1. 2. 0. 06. 0. 02. 13. 1. 1. 0. 01. 3. 1. 04
PALE	Recrystallized grains	82.1	0. 19. 13. 29. 0. 23. 4. 1. 0. 08. 0. 03. 16. 2. 0. 4. 0. 01. 4. 1. 13
BIDAU	Recrystallized grains	64.8	0. 15. 10. 41. 0. 18. 1. 4. 0. 06. 0. 02. 13. 1. 1. 0. 01. 3. 1. 09
PALE	Recrystallized grains	92.0	0. 21. 14. 77. 0. 26. 6. 2. 0. 09. 0. 04. 18. 2. 0. 6. 0. 01. 5. 1. 10
BIDAU	Recrystallized grains	72.1	0. 17. 11. 53. 0. 21. 3. 4. 0. 07. 0. 03. 15. 2. 0. 3. 0. 01. 4. 1. 07
PALE	Recrystallized grains	51.0	0. 12. 8. 1. 0. 15. 9. 4. 0. 05. 0. 02. 10. 1. 0. 9. 0. 01. 3. 1. 04
BIDAU	Recrystallized grains	62.8	0. 15. 9. 9. 0. 19. 1. 6. 0. 06. 0. 03. 13. 1. 0. 1. 0. 01. 3. 1. 05
ANGL AS	Dark domains-coarse grains	100.3	0. 22. 14. 07. 0. 28. 0. 7. 0. 6. 0. 08. 0. 03. 0. 16. 0. 2. 0. 8. 0. 01. 0. 0. 0. 0. 0. 0. 0. 0. 05
ANGL AS	Dark domains-coarse grains	98.6	0. 22. 13. 78. 0. 28. 0. 7. 0. 6. 0. 08. 0. 03. 0. 17. 0. 2. 0. 0. 0. 01. 0. 0. 0. 5. 1. 0. 0. 05
ANGL AS	Dark domains-coarse grains	76.8	0. 17. 11. 00. 0. 22. 0. 3. 0. 8. 0. 06. 0. 02. 0. 13. 0. 2. 0. 8. 0. 01. 0. 0. 0. 4. 1. 0. 0. 04

ANGL AS	Dark domains-coarse grains	106.0	0.23	15.15	0.30	0.19	0.11	0.72	0.08	0.03	0.18	0.02	0.12	0.01	0.06	0.01	0.06
ANGL AS	Dark domains-coarse grains	71.2	0.16	10.23	0.21	0.12	0.52	0.06	0.02	0.12	0.01	0.99	0.01	0.44	0.01	0.04	
ANGL AS	Dark domains-coarse grains	96.3	0.22	14.14	0.28	0.17	0.74	0.08	0.03	0.16	0.02	0.22	0.01	0.55	0.01	0.05	
ANGL AS	Dark domains-coarse grains	91.9	0.21	13.53	0.27	0.17	0.73	0.08	0.03	0.16	0.02	0.11	0.01	0.55	0.01	0.05	
ANGL AS	Dark domains-coarse grains	90.8	0.21	13.56	0.28	0.17	0.75	0.08	0.03	0.16	0.02	0.22	0.01	0.55	0.01	0.05	
ANGL AS	Dark domains-coarse grains	86.1	0.20	12.95	0.27	0.16	0.77	0.07	0.03	0.15	0.02	0.11	0.01	0.55	0.01	0.04	
ANGL AS	Dark domains-coarse grains	57.4	0.14	8.67	0.18	0.11	0.49	0.05	0.02	0.10	0.01	0.88	0.01	0.33	0.01	0.03	
ANGL AS	Dark domains-coarse grains	93.5	0.23	14.52	0.31	0.18	0.97	0.08	0.03	0.17	0.02	0.55	0.02	0.55	0.01	0.05	
ANGL AS	Dark domains-coarse grains	96.7	0.24	14.83	0.32	0.18	0.95	0.09	0.03	0.17	0.02	0.55	0.02	0.55	0.02	0.05	
ANGL AS	Dark domains-coarse grains	95.2	0.24	14.83	0.32	0.18	1.01	0.09	0.03	0.17	0.02	0.66	0.02	0.55	0.01	0.05	
ANGL AS	Dark domains-coarse grains	98.0	0.25	15.52	0.33	0.19	1.10	0.09	0.03	0.18	0.02	0.66	0.02	0.55	0.02	0.05	
ANGL AS	Dark domains-coarse grains	95.5	0.24	14.78	0.32	0.18	0.99	0.09	0.03	0.17	0.02	0.55	0.02	0.55	0.02	0.05	

ANGL	Dark domains-coarse grains	92.0	0.23	14.17	0.31	0.8	0.97	0.09	0.03	0.16	0.02	0.4	0.10	0.02	0.5	0.01	0.005
ANGL	Dark domains-coarse grains	99.8	0.24	15.27	0.34	0.9	1.14	0.09	0.03	0.18	0.02	0.7	0.10	0.02	0.5	0.02	0.005
ANGL	Dark domains-coarse grains	94.4	0.23	14.67	0.32	0.8	1.06	0.09	0.03	0.17	0.02	0.5	0.10	0.02	0.5	0.02	0.005
ARRE	Dark domains-coarse grains	109.6	0.20	18.55	0.54	0.27	0.75	0.14	0.03	0.20	0.04	0.5	0.10	0.02	0.7	0.02	0.007
ARRE	Dark domains-coarse grains	93.4	0.17	15.93	0.47	0.23	0.70	0.12	0.03	0.18	0.03	0.4	0.10	0.02	0.6	0.01	0.007
ARRE	Dark domains-coarse grains	86.6	0.16	14.81	0.44	0.22	0.67	0.11	0.02	0.16	0.03	0.2	0.10	0.02	0.5	0.01	0.007
ARRE	Dark domains-coarse grains	109.9	0.20	18.80	0.55	0.27	0.88	0.14	0.03	0.20	0.04	0.6	0.10	0.02	0.7	0.02	0.008
ARRE	Dark domains-coarse grains	91.2	0.17	15.54	0.47	0.23	0.75	0.12	0.03	0.17	0.03	0.4	0.10	0.02	0.6	0.02	0.007
ARRE	Dark domains-coarse grains	74.4	0.13	12.83	0.40	0.20	0.67	0.11	0.03	0.14	0.03	0.2	0.10	0.02	0.5	0.01	0.006
ARRE	Dark domains-coarse grains	73.0	0.13	12.77	0.39	0.19	0.69	0.11	0.02	0.13	0.03	0.3	0.10	0.02	0.4	0.01	0.005
ARRE	Dark domains-coarse grains	102.5	0.18	17.80	0.55	0.27	0.97	0.15	0.04	0.19	0.04	0.8	0.10	0.02	0.6	0.02	0.007
ARRE	Dark domains-coarse grains	99.2	0.18	17.25	0.55	0.26	0.95	0.14	0.03	0.18	0.04	0.7	0.10	0.02	0.6	0.02	0.007

		Dark domains- coarse grains		73.9	0.	13	12.	40	0.1	0.7	0.	0.1	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		72.0	0.	13	12.	39	0.1	0.7	0.	0.1	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		76.9	0.	14	13.	42	0.2	0.7	0.	0.7	0.0	0.0	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		79.1	0.	14	13.	44	0.2	0.7	0.	0.4	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		77.4	0.	14	13.	40	0.2	0.7	0.	0.4	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		74.8	0.	13	13.	40	0.2	0.7	0.	0.3	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		78.0	0.	14	13.	45	0.2	0.8	0.	0.1	0.0	0.1	0.	0.0	0.0	0.0
ARRE		Dark domains- coarse grains		84.4	0.	15	14.	49	0.2	0.8	0.	0.9	0.0	0.1	0.	0.0	0.0	0.0
PALE BIDAU		Dark domains- coarse grains		90.2	0.	21	14.	47	0.1	0.9	0.	0.9	0.0	0.1	0.	0.0	0.0	0.0
PALE BIDAU		Dark domains- coarse grains		79.0	0.	20	3.2	7	0.1	0.9	0.	0.9	0.0	0.1	0.	0.0	0.0	0.0
PALE BIDAU		Dark domains- coarse grains		101.	0.	3	25.	20	0.1	0.9	0.	0.9	0.0	0.1	0.	0.0	0.0	0.0
PALE BIDAU		Dark domains- coarse grains		141.	0.	4	65.	19	0.1	0.9	0.	0.9	0.0	0.1	0.	0.0	0.0	0.0

ANGL AS	Light domains-coarse grains	94.2 0. 24 14. 31 0.1 8 0.8 0.9 0.09 0.03 0.17 0.2 0.0 0.4 0.02 0.5 0.1 0.0 0.05
ANGL AS	Light domains-coarse grains	93.2 0. 23 14. 31 0.1 8 0.9 0.3 0.08 0.03 0.17 0.2 0.0 0.5 0.02 0.5 0.1 0.0 0.05
ANGL AS	Light domains-coarse grains	92.3 0. 23 14. 30 0.1 8 0.9 0.4 0.09 0.03 0.17 0.2 0.0 0.4 0.02 0.5 0.1 0.0 0.04
ANGL AS	Light domains-coarse grains	97.1 0. 24 15. 32 0.1 9 0.9 0.8 0.09 0.03 0.18 0.2 0.0 0.4 0.02 0.5 0.2 0.0 0.05
ANGL AS	Light domains-coarse grains	98.8 0. 24 15. 34 0.1 9 1.0 0.8 0.09 0.03 0.18 0.2 0.0 0.7 0.02 0.5 0.2 0.0 0.05
ANGL AS	Light domains-coarse grains	96.4 0. 24 14. 32 0.1 8 1.0 0.7 0.09 0.03 0.18 0.2 0.0 0.6 0.02 0.5 0.2 0.0 0.05
ANGL AS	Light domains-coarse grains	98.2 0. 24 15. 34 0.1 9 1.1 0.5 0.09 0.04 0.17 0.2 0.0 0.6 0.02 0.5 0.2 0.0 0.05
ARRE	Light domains-coarse grains	72.2 0. 13 12. 48 30 0.1 7 0.2 0.8 0.09 0.02 0.14 0.2 0.0 0.7 0.01 0.5 0.1 0.0 0.06
ARRE	Light domains-coarse grains	91.1 0. 17 15. 43 41 0.2 1 0.4 0.7 0.11 0.02 0.17 0.3 0.0 0.1 0.02 0.6 0.1 0.0 0.06
ARRE	Light domains-coarse grains	89.6 0. 16 15. 11 43 0.2 1 0.5 0.11 0.02 0.17 0.3 0.0 0.1 0.02 0.6 0.1 0.0 0.07
ARRE	Light domains-coarse grains	71.9 0. 13 12. 43 37 0.1 8 0.6 0.1 0.10 0.02 0.13 0.2 0.0 0.1 0.02 0.4 0.1 0.0 0.05
ARRE	Light domains-coarse grains	83.9 0. 14 14. 71 48 0.2 3 0.8 0.6 0.13 0.03 0.16 0.3 0.0 0.1 0.02 0.5 0.2 0.0 0.06

PALE BIDAU	Light domains- coarse grains	66.1	0. 15	11. 04	0. 18	0.1 2	0.5 8	0. 06	0. 02	0. 13	0.0 1	0.1 0	0. 01	0. 3	0.0 1	0.0 03
PALE BIDAU	Light domains- coarse grains	67.6	0. 15	11. 06	0. 19	0.1 2	0.6 1	0. 06	0. 03	0. 13	0.0 1	0.1 1	0. 01	0. 4	0.0 1	0.0 04
VICTO RIA	Recrystalli zed grains	83.5	0. 17	15. 73	0. 51	0.2 4	0.7 9	0. 14	0. 03	0. 17	0.0 3	0.1 6	0. 03	0. 7	0.0 2	0.0 08
VICTO RIA	Recrystalli zed grains	96.5	0. 20	17. 78	0. 61	0.2 7	0.9 1	0. 16	0. 04	0. 20	0.0 4	0.0 0	0. 03	0. 8	0.0 2	0.0 09
VICTO RIA	Recrystalli zed grains	74.5	0. 16	13. 88	0. 48	0.2 1	0.7 1	0. 13	0. 03	0. 15	0.0 3	0.1 6	0. 02	0. 6	0.0 2	0.0 06
VICTO RIA	Recrystalli zed grains	81.7	0. 17	15. 31	0. 53	0.2 3	0.8 6	0. 14	0. 03	0. 17	0.0 3	0.1 7	0. 03	0. 7	0.0 2	0.0 07
VICTO RIA	Recrystalli zed grains	80.0	0. 17	14. 80	0. 52	0.2 3	0.8 0	0. 14	0. 03	0. 16	0.0 3	0.1 6	0. 03	0. 6	0.0 2	0.0 07
VICTO RIA	Recrystalli zed grains	80.7	0. 17	15. 19	0. 53	0.2 3	0.8 2	0. 14	0. 03	0. 17	0.0 3	0.1 7	0. 03	0. 6	0.0 2	0.0 05
VICTO RIA	Recrystalli zed grains	77.9	0. 16	14. 84	0. 49	0.2 2	0.7 6	0. 14	0. 03	0. 17	0.0 3	0.1 6	0. 03	0. 6	0.0 2	0.0 07
VICTO RIA	Coarse grains	104. 6	0. 21	19. 68	0. 65	0.2 9	1.1 2	0. 18	0. 04	0. 22	0.0 4	0.2 1	0. 03	0. 8	0.0 3	0.0 10
VICTO RIA	Coarse grains	100. 1	0. 21	18. 85	0. 65	0.2 8	1.0 0	0. 17	0. 04	0. 21	0.0 4	0.0 0	0. 03	0. 8	0.0 3	0.0 08
VICTO RIA	Coarse grains	86.8	0. 18	16. 10	0. 57	0.2 4	0.8 6	0. 15	0. 03	0. 18	0.0 4	0.1 9	0. 03	0. 7	0.0 2	0.0 08
VICTO RIA	Coarse grains	93.0	0. 19	17. 49	0. 62	0.2 6	0.9 4	0. 17	0. 04	0. 20	0.0 4	0.2 1	0. 03	0. 7	0.0 2	0.0 08
MARG ALIDA	Recrystalli zed grains	102. 0	0. 17	15. 21	15. 0	0.2 1	0.9 3	0. 63	0. 39	0. 31	0.4 2	0.2 4	0. 03	0. 2	0.1 0	0.5 00
MARG ALIDA	Recrystalli zed grains	101. 3	0. 14	10. 23	14. 6	0.2 9	0.3 2	0. 41	0. 28	0. 15	0.0 8	0.2 3	0. 02	0. 5	0.0 2	0.4 10
MARG ALIDA	Recrystalli zed grains	95.6	0. 12	14. 23	15. 8	0.2 0	0.5 2	0. 47	0. 25	0. 18	0.1 1	0.2 3	0. 02	0. 6	0.0 3	0.0 20
MARG ALIDA	Recrystalli zed grains	96.5	0. 23	17. 20	15. 6	0.2 8	1.6 4	0.8 0	1. 04	0.7 69	0.0 15	0.7 9	0.2 4	0. 02	0.0 7	1.2 30
MARG ALIDA	Recrystalli zed grains	96.8	0. 28	17. 30	15. 4	0.2 3	1.2 2	0.5 32	0.0 77	0. 18	0.4 8	0.2 5	0. 02	0. 7	0.6 50	
MARG ALIDA	Recrystalli zed grains	98.6	0. 21	17. 60	16. 6	0.2 3	1.3 4	0.8 31	0.0 81	0. 17	0.4 2	0.2 6	0. 02	0. 6	0.5 40	
MARG ALIDA	Recrystalli zed grains	100. 2	0. 23	17. 80	16. 3	0.2 4	1.1 4	0.5 4	0.0 49	0. 29	0.6 18	0.2 9	0. 03	0. 7	1.1 35	
MARG ALIDA	Recrystalli zed grains	104. 3	0. 24	17. 40	15. 6	0.2 4	1.7 4	0.6 4	0. 18	0. 09	1.4 18	0.2 9	0. 03	0. 3	3.8 67	

MARG ALIDA	Recrystalli- zed grains	93.0	0.	27	14.	60	17. 5	0.8	0.6	1.	0.	0.5	0.	0.1	0.2	0.	0.0	0.1	1.4
BENT AILLO U	Coarse grains	72.2	0.	10.	0.	0.1	0.5	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	63.3	0.	9.4	0.	0.1	0.5	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	77.7	0.	11.	0.	0.1	0.6	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	71.6	0.	10.	0.	0.1	0.6	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	60.4	0.	9.2	0.	0.1	0.5	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	85.3	0.	13.	0.	0.1	0.7	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	82.2	0.	12.	0.	0.1	0.7	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	66.2	0.	10.	0.	0.1	0.5	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	69.4	0.	10.	0.	0.1	0.6	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	94.3	0.	15.	0.	0.1	0.8	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	88.0	0.	13.	0.	0.1	0.8	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Coarse grains	102.	0.	16.	0.	0.2	0.9	0.	0.	0.	0.	0.	0.0	0.1	0.	0.0	0.0	0.0	0.0
BENT AILLO U	Type 1 dissemina- ted	494.	1.	11	1.	1.2	4.6	0.	0.	1.	0.1	0.9	0.	0.3	0.0	0.0	0.0	0.0	0.0
BENT AILLO U	Type 1 dissemina- ted	311.	0.	71.	1.	0.7	2.8	0.	0.	0.	0.0	0.6	0.	0.2	0.0	0.0	0.0	0.0	0.0
BENT AILLO U	Type 1 dissemina- ted	304.	0.	68.	1.	0.7	2.8	0.	0.	0.	0.0	0.5	0.	0.2	0.0	0.0	0.0	0.0	0.0
BENT AILLO U	Type 1 dissemina- ted	310.	0.	72.	1.	0.8	2.9	0.	0.	0.	0.0	0.5	0.	0.2	0.0	0.0	0.0	0.0	0.0
BENT AILLO U	Type 1 dissemina- ted	492.	1.	11	2.	1.2	4.6	0.	0.	1.	0.1	0.8	0.	0.3	0.0	0.0	0.0	0.0	0.0

BENT AILLO U	Type 1 disseminated	312. 6	0. 85	72. 51	1. 28	0.8 1	2.9 1	0. 41	0. 15	0. 69	0.0 8	0.0 8	0.5 0.7	0. 1	0.2 1	0.0 6	0.0 18
BENT AILLO U	Type 1 disseminated	541. 6	1. 48	12. 3.8	2. 23	1.3 6	5.0 4	0. 72	0. 26	1. 21	0.1 5	1.0 2	0. 12	0.3 6	0.1 0	0.0 33	
1 sigma (ppm)																	
Deposit	Texture	M S34	n5 5	Fe 57	C 05 9	Cu 63	Zn6 4	G a6 9	Ga 71	Ge 74	Ag 10 7	Cd 11 1	In 11 5	Sn 11 8	Sb 12 1	Pb 20 8	
ANGL AS	Recrystallized grains	356 83.6	12 .7	18. 13. 4	5. 7	8.0	14. 7	1. 6	1. 8	1. 8	0.1 0.1	83. 7	0. 14	0.1 0.1	0.7 0.7	1.7 1.7	
ANGL AS	Recrystallized grains	394 79.7	7. 4	97. 9.1	7. 1	4.2	75. 5	2. 2	2. 0	3. 2	0.4 0.4	75. 7	0. 31	0.9 0.9	0.7 0.7	1.6 1.6	
ANGL AS	Recrystallized grains	348 90.4	15 .0	21. 50. 8	7. 2	4.6	58. 1	1. 5	1. 7	2. 3	0.2 0.2	84. 4	0. 16	0.1 0.1	0.5 0.5	1.7 1.7	
ANGL AS	Recrystallized grains	363 17.4	12 .8	18. 48. 9	6. 0	4.3	190. 67. 5	1. 5	1. 8	3. 4	0.1 0.1	82. 1	0. 17	0.1 0.1	0.3 0.3	1.5 1.5	
ANGL AS	Recrystallized grains	353 10.9	12 .9	18. 18. 1	5. 8	4.2	191. 37. 8	1. 9	2. 2	0. 7	0.1 0.1	81. 7	0. 13	0.1 0.1	0.3 0.3	1.2 1.2	
ANGL AS	Recrystallized grains	353 37.2	13 .1	18. 54. 1	6. 0	5.1	192. 27. 1	2. 5	2. 9	2. 2	0.1 0.1	82. 1	0. 13	0.1 0.1	0.3 0.3	1.0 1.0	
ANGL AS	Recrystallized grains	343 51.9	6. 4	10. 75. 1	4. 4	4.2	189. 67. 6	1. 6	1. 9	0. 5	0.2 0.2	86. 2	0. 05	0.2 0.2	0.3 0.3	1.7 1.7	
ANGL AS	Recrystallized grains	353 51.4	9. 8	16. 21. 1	6. 3	6.1	189. 67. 0	1. 8	2. 1	0. 1	0.2 0.2	83. 9	0. 04	0.2 0.2	0.4 0.4	2.1 2.1	
ANGL AS	Recrystallized grains	365 58.2	10 .4	16. 93. 1	6. 6	4.3	194. 69. 0	1. 4	1. 6	0. 2	0.1 0.1	82. 3	0. 04	0.1 0.1	0.5 0.5	1.5 1.5	
ANGL AS	Recrystallized grains	368 83.7	10 .9	17. 55. 0	6. 9	4.3	193. 79. 7	1. 8	2. 0	1. 2	0.1 0.1	81. 5	0. 05	0.1 0.1	0.2 0.2	1.0 1.0	
ANGL AS	Recrystallized grains	354 21.3	8. 4	14. 00. 2	5. 8	5.4	193. 93. 6	2. 0	2. 3	0. 2	0.2 0.2	80. 6	0. 04	0.1 0.1	0.3 0.3	1.6 1.6	
ANGL AS	Recrystallized grains	375 24.5	9. 9	16. 46. 4	5. 9	5.7	196. 06. 9	1. 2	1. 4	0. 3	0.2 0.2	84. 6	0. 06	0.5 0.5	0.6 0.6	2.8 2.8	
ANGL AS	Recrystallized grains	364 28.8	9. 1	14. 98. 9	5. 4	4.3	194. 56. 2	1. 0	1. 2	0. 3	0.1 0.1	84. 3	0. 06	0.5 0.5	0.4 0.4	1.8 1.8	
ANGL AS	Recrystallized grains	377 71.8	10 .4	16. 12. 5	5. 5	3.5	194. 82. 2	1. 2	1. 4	0. 2	0.1 0.1	75. 3	0. 04	0.1 0.1	0.3 0.3	0.9 0.9	

ANGL AS	Recrystallized grains	374 57.1	10 .2	15 89. 6	5. 4	3.8	194 24. 5	1. 2	1. 3	0. 1	0.1	76. 2	0. 05	0.1	0.3	1.6
ARRE	Recrystallized grains	437 83.2	1. 1	83 3.5	5. 4	3.0	198 33. 6	0. 5	0. 5	0. 5	0.4	18 3.1	0. 01	0.1	0.6	4.9
ARRE	Recrystallized grains	458 80.3	1. 0	10 74. 5	6. 2	1.0	197 02. 7	0. 1	0. 1	0. 2	0.1	17 1.4	0. 02	0.1	0.2	0.4
ARRE	Recrystallized grains	437 53.9	1. 1	10 00. 0	6. 4	1.8	195 99. 9	0. 1	0. 1	0. 2	0.2	16 5.9	0. 02	0.1	0.2	1.3
ARRE	Recrystallized grains	453 27.9	1. 4	10 88. 4	6. 5	3.1	196 59. 4	0. 2	0. 2	0. 3	0.4	17 2.4	0. 02	0.1	0.6	4.0
ARRE	Recrystallized grains	438 39.7	1. 3	10 82. 9	6. 5	3.1	196 00. 0	0. 4	0. 3	0. 3	0.3	17 7.2	0. 02	0.1	0.4	3.3
ARRE	Recrystallized grains	483 83.4	1. 1	93 5.8	6. 1	2.5	197 20. 2	0. 3	0. 3	0. 4	0.4	17 1.4	0. 02	0.1	0.4	2.7
ARRE	Recrystallized grains	477 40.8	1. 2	89 1.5	5. 9	2.8	196 99. 1	0. 3	0. 3	0. 4	0.4	17 0.7	0. 02	0.1	0.5	3.1
ARRE	Recrystallized grains	457 25.1	1. 1	89 1.8	6. 1	2.3	196 82. 8	0. 3	0. 3	0. 3	0.3	16 6.5	0. 02	0.1	0.4	2.5
ARRE	Recrystallized grains	484 15.7	1. 0	85 1.7	6. 0	1.6	197 01. 0	0. 4	0. 4	0. 2	0.2	16 2.7	0. 03	0.1	0.2	1.4
ARRE	Recrystallized grains	256 573. 1	1. 4	81 6.6	5. 2	2.5	197 02. 3	6. 5	6. 5	3. 9	3.2	10 73. 4	0. 04	0.6	5.9	17. 2
ARRE	Recrystallized grains	459 08.8	1. 2	88 3.0	5. 8	4.0	197 77. 2	0. 4	0. 4	0. 5	0.5	16 9.6	0. 02	0.1	0.7	4.9
ARRE	Recrystallized grains	470 78.7	1. 1	94 5.3	6. 3	3.3	197 48. 1	0. 4	0. 4	0. 4	0.5	16 6.8	0. 03	0.1	0.8	4.1
ARRE	Recrystallized grains	263 132. 0	1. 7	85 6.5	5. 2	5.0	624 80. 6	5. 2	5. 2	1. 3	7.0	10 46. 1	0. 05	1.3	12. 0	49. 3
ARRE	Recrystallized grains	482 32.7	1. 0	10 01. 8	6. 5	1.6	196 84. 2	0. 2	0. 1	0. 2	0.2	17 3.2	0. 03	0.1	0.2	1.3
PALE BIDAU	Recrystallized grains	224 74.5	3. 3	21 90. 4	18 .5	7.3	197 80. 1	3. 0	2. 8	0. 5	2.5	26 6.4	0. 12	0.7	2.2	11. 8
PALE BIDAU	Recrystallized grains	224 58.2	3. 5	23 56. 9	20 .3	7.7	197 62. 2	2. 5	2. 5	0. 6	3.0	27 3.3	0. 11	0.9	2.4	12. 6
PALE BIDAU	Recrystallized grains	227 61.5	3. 8	25 51. 6	21 .6	9.4	197 62. 3	3. 3	3. 2	0. 6	3.2	27 7.9	0. 11	0.9	2.5	13. 9

PALE	Recrystallized grains	238	3.	23	20	6.6	197	3.	3.	0.	2.0	29	0.	0.6	1.7	8.9
BIDAU		86.5	2	38. 9	.1	6.6	62. 2	0	0	4		0.2	12			
PALE	Recrystallized grains	240	3.	23	20	4.5	199	7.	7.	0.	2.9	28	0.	0.9	2.6	10. 7
BIDAU		34.1	3	11. 8	.5		41. 8	6	4	7		4.8	06			
PALE	Recrystallized grains	249	4.	24	22	1.8	198	10	10	0.	5.2	29	0.	1.5	5.4	32 13. 3
BIDAU		13.0	3	26. 4	.4		95. 2	.3	.0	7		1.2	06			
PALE	Recrystallized grains	251	3.	23	21	8.1	199	15	14	0.	3.2	30	0.	0.7	2.6	65 2.6
BIDAU		32.1	8	10. 4	.6		13. 9	.3	.9	4		3.4	04			
PALE	Recrystallized grains	253	4.	23	22	8.1	198	16	15	0.	4.3	31	0.	0.8	3.2	51. 5
BIDAU		91.5	2	99. 7	.9		71. 3	.4	.8	7		7.6	03			
PALE	Recrystallized grains	261	3.	21	20	3.3	197	16	15	0.	1.6	31	0.	0.3	1.1	68. 9
BIDAU		16.4	9	87. 7	.9		85. 1	.4	.9	3		9.7	03			
PALE	Recrystallized grains	269	3.	25	23	5.0	199	12	11	0.	1.1	32	0.	0.3	0.5	5.6
BIDAU		29.5	7	21. 4	.8		13. 9	.0	.6	1		6.8	05			
PALE	Recrystallized grains	261	4.	27	25	2.8	197	16	16	0.	5.1	34	0.	1.0	3.7	16. 1
BIDAU		23.7	7	07. 6	.9		38. 5	.4	.1	7		0.5	04			
PALE	Recrystallized grains	282	4.	25	25	24.	196	16	16	0.	2.3	34	0.	0.4	1.5	8.8
BIDAU		13.2	8	43. 4	.5	1	46. 6	.2	.0	3		3.3	03			
ANGL AS	Dark domains-coarse grains	337	10	16	4.	44.	187	0.	0.	13	0.2	77.	0.	0.0	2.1	0.9
		78.4	.3	42. 0	8	1	48. 3	5	6	.8	74	01				
ANGL AS	Dark domains-coarse grains	347	11	16	4.	55.	191	0.	0.	17	0.2	79.	0.	0.0	5.6	3.6
		91.9	.0	83. 3	9	6	38. 1	8	9	.3	73	01				
ANGL AS	Dark domains-coarse grains	342	9.	16	5.	48.	193	2.	2.	12	0.2	79.	0.	0.0	5.5	3.7
		72.0	9	52. 9	1	7	36. 4	0	3	.4	71	01				
ANGL AS	Dark domains-coarse grains	350	8.	15	4.	56.	192	2.	2.	15	0.2	78.	0.	0.0	6.5	3.2
		78.2	9	18. 5	6	0	19. 1	2	5	.5	73	01				
ANGL AS	Dark domains-coarse grains	356	9.	17	5.	51.	192	2.	2.	13	0.2	78.	0.	0.0	9.3	3.9
		18.1	9	25. 2	2	6	19. 1	0	3	.5	79	01				
ANGL AS	Dark domains-coarse grains	353	11	17	5.	33.	191	1.	2.	9.	0.2	81.	0.	0.5	1.8	1.5
		10.3	.9	78. 0	7	6	43. 1	8	1	7		621				

ANGL AS	Dark domains- coarse grains	353 44.7	11 .7	17 32. 6	5. 6	50. 5	190 82. 0	1. 1	1. 3	14 .3	0.2 0.2	80. 1	0. 29	1.3 1.3	4.9 3.3
ANGL AS	Dark domains- coarse grains	351 99.6	11 .7	17 35. 6	5. 6	32. 1	190 97. 3	0. 7	0. 8	10 .7	0.1 0.1	79. 3	0. 24	1.3 1.3	1.7 1.0
ANGL AS	Dark domains- coarse grains	360 81.9	12 .0	17 88. 2	5. 8	26. 2	190 93. 2	0. 9	1. 1	12 .2	0.2 0.2	79. 8	0. 19	1.2 1.2	0.7 1.5
ANGL AS	Dark domains- coarse grains	359 14.7	14 .2	20 68. 8	6. 5	24. 7	191 86. 6	1. 2	1. 4	10 .4	0.1 0.1	80. 6	0. 16	0.4 0.4	1.7 1.2
ANGL AS	Dark domains- coarse grains	354 86.3	8. 6	15 61. 0	6. 5	61. 1	193 09. 6	3. 3	3. 7	18 .0	0.6 0.6	79. 9	0. 04	0.3 0.3	0.7 4.7
ANGL AS	Dark domains- coarse grains	368 96.1	8.	15 54. 9	6. 1	38. 6	196 00. 5	6. 1	7. 1	12 .0	0.3 0.3	79. 4	0. 16	2.4 2.4	0.1 0.8
ANGL AS	Dark domains- coarse grains	371 58.0	9. 6	15 72. 2	5. 6	45. 6	193 58. 7	1. 9	2. 2	13 .5	0.2 0.2	79. 7	0. 04	0.2 0.2	5.1 3.4
ANGL AS	Dark domains- coarse grains	367 88.9	9. 6	15 87. 2	5. 5	45. 2	193 80. 3	0. 9	1. 1	12 .6	0.2 0.2	79. 1	0. 02	0.1 0.1	7.4 5.7
ANGL AS	Dark domains- coarse grains	376 11.6	11 .3	17 16. 9	5. 5	19. 4	192 74. 5	1. 7	2. 0	5. 5	0.2 0.2	82. 2	0. 02	0.0 0.0	0.4 0.2
ANGL AS	Dark domains- coarse grains	370 98.5	5. 4	12 95. 4	5. 4	59. 1	195 62. 2	1. 0	1. 1	15 .7	0.2 0.2	75. 8	0. 01	0.0 0.0	11. 0
ANGL AS	Dark domains- coarse grains	378 08.9	5. 5	13 14. 7	5. 5	71. 2	196 43. 2	1. 5	1. 7	19 .5	0.2 0.2	74. 2	0. 01	0.0 0.0	11. 6.5
ANGL AS	Dark domains- coarse grains	375 22.6	6. 7	13 98. 1	5. 5	55. 4	195 90. 6	3. 4	4. 0	21 .4	0.2 0.2	76. 7	0. 01	0.2 0.2	2.8 1.2

ARRE	Dark domains-coarse grains	434 32.7	0. 9	76 4.7	5. 0	48. 9	198 43. 0	0. 4	0. 4	16 .6	0.3 .3	18 1.4	0. .01	0.3 .3	0.3 .3	2.1
ARRE	Dark domains-coarse grains	431 82.6	1. 0	81 3.8	5. 3	90. 2	197 58. 5	0. 5	0. 4	27 .8	0.3 .3	18 0.7	0. .01	0.3 .3	0.2 .2	1.2
ARRE	Dark domains-coarse grains	440 25.1	1. 0	81 5.2	5. 2	80. 1	196 90. 6	0. 3	0. 3	23 .5	0.3 .3	18 1.0	0. .01	0.3 .3	0.2 .2	1.4
ARRE	Dark domains-coarse grains	441 45.1	0. 9	84 4.9	5. 1	40. 8	196 68. 0	0. 3	0. 3	12 .4	0.2 .2	17 5.9	0. .01	0.0 .0	0.5 .5	0.3
ARRE	Dark domains-coarse grains	437 40.9	1. 0	90 1.6	5. 5	46. 3	196 24. 4	0. 4	0. 3	13 .4	0.2 .2	17 2.6	0. .01	0.0 .0	0.7 .7	0.3
ARRE	Dark domains-coarse grains	458 78.2	1. 1	10 12. 8	6. 2	49. 4	198 10. 6	0. 1	0. 0	13 .1	0.2 .2	17 0.3	0. .01	0.0 .0	0.1 .1	0.2
ARRE	Dark domains-coarse grains	456 35.4	1. 1	10 74. 2	6. 4	99. 1	197 76. 0	0. 3	0. 3	26 .2	0.2 .2	17 5.1	0. .02	0.5 .5	0.2 .2	0.9
ARRE	Dark domains-coarse grains	436 47.2	0. 8	78 6.9	5. 2	67. 4	196 23. 3	0. 4	0. 4	21 .0	0.2 .2	17 6.0	0. .02	0.6 .6	0.1 .1	0.8
ARRE	Dark domains-coarse grains	444 98.8	0. 8	79 7.0	5. 4	10 4.0	197 12. 8	0. 5	0. 4	31 .5	0.2 .2	17 0.9	0. .02	0.1 .1	0.4 .4	0.4
ARRE	Dark domains-coarse grains	452 48.3	0. 9	90 5.3	6. 1	11 2.0	196 00. 1	0. 4	0. 4	29 .4	0.3 .3	17 2.1	0. .02	0.2 .2	0.2 .2	0.5
ARRE	Dark domains-coarse grains	444 10.4	0. 9	88 2.4	6. 0	10 8.4	195 26. 2	0. 6	0. 5	28 .6	0.2 .2	16 8.2	0. .02	0.3 .3	0.2 .2	0.5
ARRE	Dark domains-coarse grains	447 03.0	1. 0	88 4.5	6. 1	92. 0	196 16. 8	0. 6	0. 6	25 .1	0.4 .4	17 1.1	0. .02	0.9 .9	0.4 .4	3.3

		Dark domains-coarse grains	467 63.0	0. 9	86 6.2	6. 0	76. 9	196 12. 4	0. 6	0. 6	23 .5	0.3 5.1	16 0.2	0. 02	0.8 0.2	0.2 2.0
ARRE		Dark domains-coarse grains	482 06.5	0. 9	85 0.9	5. 9	99. 8	195 71. 1	0. 3	0. 3	28 .3	0.3 8.4	15 0.2	0. 02	0.1 0.3	0.5
ARRE		Dark domains-coarse grains	475 43.4	0. 9	86 2.8	5. 8	80. 4	195 68. 3	0. 3	0. 3	22 .0	0.3 8.7	15 0.2	0. 02	0.2 0.2	1.0
ARRE		Dark domains-coarse grains	511 37.6	1. 1	10 16. 6	6. 1	10 8.2	197 83. 6	0. 4	0. 3	27 .6	0.6 8.5	16 0.2	0. 02	0.1 0.5	0.7
ARRE		Dark domains-coarse grains	457 96.2	1. 1	10 42. 2	6. 3	89. 6	197 73. 0	0. 4	0. 3	22 .7	0.5 5.1	17 0.3	0. 03	0.1 0.9	0.8
PALE BIDAU		Dark domains-coarse grains	260 92.6	3. 6	23 82. 2	22 .2	25. 9	199 14. 4	12 .9	12 .6	19 .9	1.6 6.4	31 0.06	1.0 0.06	0.7 7.8	
PALE BIDAU		Dark domains-coarse grains	260 97.7	3. 6	23 81. 2	22 .3	23. 5	199 15. 5	13 .0	12 .7	19 .9	1.6 7.3	31 0.05	1.1 0.05	0.7 7.5	
PALE BIDAU		Dark domains-coarse grains	260 97.6	3. 5	23 87. 3	22 .1	21. 3	199 19. 3	12 .3	12 .4	19 .8	1.7 9.4	31 0.06	1.1 0.06	0.6 7.5	
PALE BIDAU		Dark domains-coarse grains	260 98.4	3. 6	23 82. 8	22 .6	26. 3	199 16. 2	13 .3	12 .5	19 .6	1.4 4.6	32 0.05	1.1 0.05	0.9 7.7	
ANGL AS		Light domains-coarse grains	362 78.5	8. 1	15 24. 9	6. 4	21. 4	194 57. 6	2. 9	3. 4	5. 9	0.2 0.2	79. 3	0. 02	0.3 0.2	2.4
ANGL AS		Light domains-coarse grains	372 51.4	5. 8	13 70. 3	6. 1	20. 5	194 73. 0	4. 5	5. 3	4. 1	0.2 0.2	78. 0	0. 02	0.8 0.3	0.3
ANGL AS		Light domains-coarse grains	368 60.3	5. 1	13 09. 2	5. 9	16. 2	196 75. 2	3. 9	4. 5	3. 3	0.1 0.1	76. 5	0. 01	0.1 0.3	0.1

ANGL AS	Light domains-coarse grains	368 28.1	11 .3	17 03. 3	5. 5	17. 1	193 42. 5	2. 3	2. 6	7. 3	0.1	82. 2	0. .01	0.1	0.6	0.3
ANGL AS	Light domains-coarse grains	380 05.7	5. 5	13 01. 3	5. 5	4.7	196 15. 3	0. 2	0. 3	1. 1	0.2	74. 8	0. .01	0.0	0.1	1.1
ANGL AS	Light domains-coarse grains	378 33.3	5. 9	13 32. 6	5. 4	6.6	197 08. 6	0. 3	0. 3	1. 2	0.3	75. 2	0. .01	0.0	0.2	3.4
ANGL AS	Light domains-coarse grains	376 88.4	6. 0	13 45. 2	5. 5	5.6	196 75. 7	0. 2	0. 2	1. 0	0.3	77. 6	0. .01	0.0	0.2	2.9
ARRE	Light domains-coarse grains	424 00.5	1. 1	82 7.2	5. 7	9.0	197 52. 5	0. 9	0. 8	2. 3	0.4	18 5.4	0. .01	0.1	0.4	4.1
ARRE	Light domains-coarse grains	441 47.5	1. 1	81 7.9	5. 1	5.5	197 50. 4	0. 5	0. 4	1. 1	0.3	18 8.6	0. .01	0.1	0.4	3.6
ARRE	Light domains-coarse grains	434 97.6	1. 0	10 05. 4	6. 0	16. 4	196 88. 1	0. 4	0. 4	4. 0	0.2	17 4.8	0. .01	0.0	0.7	0.3
ARRE	Light domains-coarse grains	485 16.5	1. 1	97 1.4	6. 2	9.1	197 04. 1	0. 5	0. 5	2. 0	0.3	17 1.9	0. .03	0.1	0.4	2.6
ARRE	Light domains-coarse grains	489 28.5	1. 1	10 23. 5	6. 6	20. 4	197 73. 1	0. 5	0. 4	5. 3	0.2	17 0.5	0. .03	0.2	0.1	0.3
PALE BIDAU	Light domains-coarse grains	176 90.4	2. 6	22 86. 5	15 .8	2.6	198 08. 4	1. 0	1. 0	0. 1	0.6	20 7.1	0. .20	0.2	0.5	2.8
PALE BIDAU	Light domains-coarse grains	176 25.5	3. 0	22 35. 2	15 .8	1.5	199 17. 9	0. 4	0. 4	0. 1	0.2	20 8.6	0. .28	0.1	0.1	0.7
PALE BIDAU	Light domains-coarse grains	172 32.1	2. 8	21 15. 9	14 .7	1.0	199 11. 3	0. 2	0. 2	0. 1	0.2	21 0.5	0. .27	0.1	0.1	0.7

PALE BIDAU	Light domains- coarse grains	177 58.8	2. 7	21 22. 2	15 .0	1.3	199 19. 7	0. 2	0. 2	0. 1	0.4	21 1.5	0. 24	0.2	0.2	2.0
PALE BIDAU	Light domains- coarse grains	178 20.7	2. 7	23 22. 9	16 .4	0.7	197 80. 5	0. 1	0. 1	0. 1	0.2	21 1.9	0. 12	0.1	0.0	0.5
PALE BIDAU	Light domains- coarse grains	181 60.4	2. 9	21 86. 5	15 .8	1.8	198 27. 3	0. 5	0. 4	0. 2	0.5	21 8.1	0. 18	0.2	0.4	2.9
PALE BIDAU	Light domains- coarse grains	185 95.4	2. 8	22 13. 8	16 .1	1.8	198 45. 0	0. 8	0. 8	0. 1	0.3	22 4.0	0. 20	0.1	0.1	1.1
PALE BIDAU	Light domains- coarse grains	187 01.4	3. 3	24 91. 8	18 .4	2.1	198 07. 2	0. 5	0. 5	0. 1	0.5	23 2.0	0. 22	0.2	0.3	2.6
PALE BIDAU	Light domains- coarse grains	191 97.1	3. 7	25 18. 8	19 .0	1.5	197 13. 7	0. 5	0. 4	0. 1	0.3	23 9.1	0. 17	0.1	0.1	1.4
PALE BIDAU	Light domains- coarse grains	197 84.7	3. 4	25 19. 3	18 .9	1.6	198 21. 3	0. 5	0. 5	0. 1	0.2	23 9.6	0. 22	0.1	0.1	0.9
PALE BIDAU	Light domains- coarse grains	193 54.1	3. 1	24 16. 2	18 .3	1.7	197 14. 7	0. 8	0. 7	0. 1	0.2	23 5.3	0. 19	0.1	0.1	0.7
PALE BIDAU	Light domains- coarse grains	213 37.8	4. 7	29 22. 5	23 .8	0.4	197 01. 0	0. 5	4. 6	0. 4	1.6	28 8.2	0. 01	0.5	1.0	7.0
PALE BIDAU	Light domains- coarse grains	214 43.9	4. 4	30 40. 2	25 .4	0.5	195 91. 4	0. 8	11 .0	0. 1	0.4	29 9.3	0. 02	0.1	0.1	1.1
PALE BIDAU	Light domains- coarse grains	220 79.6	4. 6	30 53. 7	25 .7	0.2	195 25. 0	0. 6	11 .5	0. 1	0.4	31 2.3	0. 02	0.1	0.1	1.0
VICTO RIA	Recrystalli zed grains	256 98.4	13 .1	16 68. 0	5. 6	1.0	193 44. 1	0. 2	0. 1	0. 2	0.4	84. 4	0. 06	0.1	0.2	3.7
VICTO RIA	Recrystalli zed grains	263 73.6	12 .9	15 87. 4	5. 4	0.7	192 35. 6	0. 1	0. 1	0. 1	0.3	83. 8	0. 05	0.1	0.2	2.4

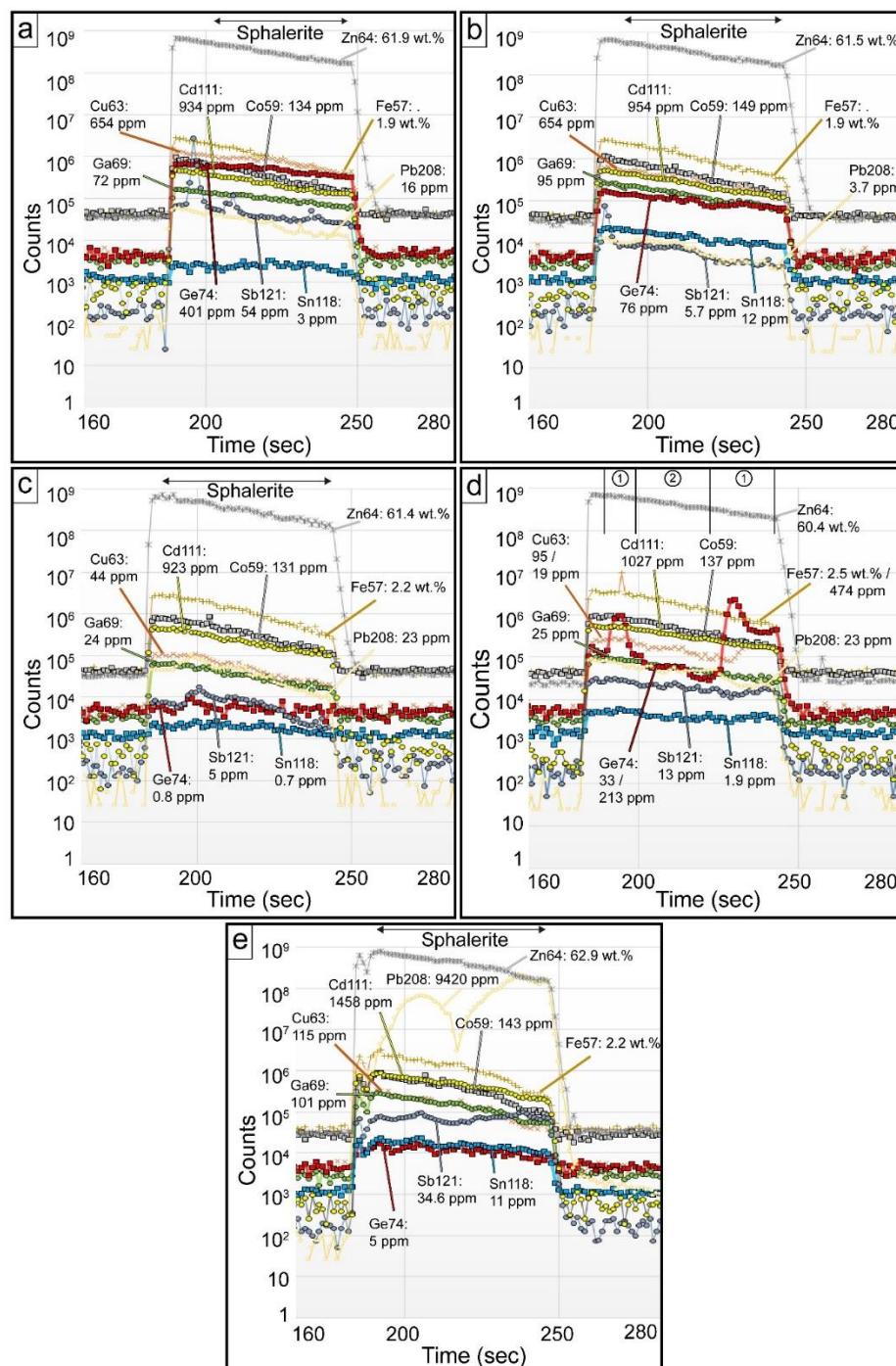
VICTORIA	Recrystallized grains	263 20.0	13 .7	16 83. 3	5. 4	1.5 72. 5	191 193 192 192 192	0. 1	0. 1	0. 2	0.7 0.6 0.3 0.6 0.7	87. 1	0. 06	0.1 0.1	0.3 0.1	6.2 5.5 2.3 5.0 6.4
VICTORIA	Recrystallized grains	259 28.1	13 .4	16 12. 3	5. 5	1.3 15. 1	193 192 192 192 192	0. 1	0. 1	0. 1	0.6 0.3 0.6 0.6 0.7	86. 7	0. 06	0.1 0.1	0.3 0.1	5.5 2.3 5.0 6.4
VICTORIA	Recrystallized grains	256 88.0	13 .3	17 02. 9	5. 3	0.8 26. 2	192 192 192 192 192	0. 2	0. 1	0. 1	0.3 0.3 0.6 0.7 0.7	86. 3	0. 05	0.1 0.1	0.1 0.1	2.3 5.0 6.4
VICTORIA	Recrystallized grains	271 69.6	13 .7	16 94. 7	5. 5	1.2 98. 3	192 192 192 192 192	0. 2	0. 1	0. 2	0.6 0.6 0.6 0.7 0.7	86. 6	0. 06	0.1 0.1	0.3 0.3	5.0 5.0 6.4
VICTORIA	Recrystallized grains	275 50.4	13 .8	16 71. 6	5. 9	1.9 40. 3	192 192 192 192 192	0. 2	0. 1	0. 2	0.7 0.7 0.7 0.7 0.7	87. 0	0. 06	0.1 0.1	0.4 0.4	6.4 6.4
VICTORIA	Coarse grains	256 76.8	24 .2	29 81. 6	9. 9	0.9 66. 6	182 182 182 182 182	0. 6	0. 6	0. 2	0.1 0.1 0.2 0.1 0.1	13 9.7	0. 07	0.1 0.1	0.0 0.0	0.2 1.6 0.1 0.0
VICTORIA	Coarse grains	260 83.1	24 .0	29 67. 0	10 .0	1.6 20. 2	182 182 182 182 182	0. 7	0. 7	0. 2	0.2 0.2 0.1 0.1 0.1	13 5.9	0. 07	0.1 0.1	0.1 1.6	1.6 1.6 0.1 0.0
VICTORIA	Coarse grains	266 30.2	25 .8	31 78. 1	10 .7	1.2 45. 3	182 182 182 182 182	0. 9	0. 9	0. 2	0.1 0.1 0.1 0.1 0.1	13 1.6	0. 07	0.1 0.1	0.0 0.0	0.1 0.1
VICTORIA	Coarse grains	261 70.6	28 .2	32 85. 6	11 .0	1.3 82. 3	182 182 182 182 182	0. 9	0. 9	0. 2	0.1 0.1 0.1 0.1 0.1	13 7.2	0. 08	0.1 0.1	0.0 0.0	0.2 0.2
MARGALIDA	Recrystallized grains	227 83.1	23 .0	54 42. 7	2. 4	0.8 84. 3	182 182 182 182 182	2. 8	2. 7	0. 1	9.9 9.9	13 7.6	0. 02	0.4 0.4	7.0 7.0	8.0
MARGALIDA	Recrystallized grains	230 81.7	23 .0	53 55. 2	2. 7	0.5 87. 0	182 182 182 182 182	1. 8	2. 1	0. 1	1.8 1.8	13 7.5	0. 01	0.1 0.1	0.2 1.5	1.5 1.5
MARGALIDA	Recrystallized grains	232 83.9	23 .6	53 61. 7	2. 7	0.6 84. 4	182 182 182 182 182	2. 0	1. 8	0. 0	2.3 2.3	13 4.2	0. 06	0.2 0.2	0.2 1.1	1.1 1.1
MARGALIDA	Recrystallized grains	238 17.7	21 .0	50 56. 3	2. 6	0.3 81. 2	182 182 182 182 182	4. 4	5. 1	0. 1	23. 23.	12 4.6	0. 05	0.7 0.7	17. 17.	17. 17.
MARGALIDA	Recrystallized grains	241 71.5	21 .6	52 05. 0	1. 5	1.7 94. 3	182 182 182 182 182	2. 4	5. 4	0. 3	13. 13.	11 7.4	0. 03	0.4 0.4	8.6 8.6	11. 11.
MARGALIDA	Recrystallized grains	245 61.3	22 .7	48 86. 0	2. 3	1.8 84. 2	182 182 182 182 182	2. 1	5. 5	0. 3	11. 11.	13 7.1	0. 01	0.4 0.4	6.4 6.4	11. 11.
MARGALIDA	Recrystallized grains	246 77.4	20 .8	47 23. 0	2. 4	1.4 87. 5	182 182 182 182 182	1. 8	1. 8	0. 1	19. 19.	13 9.4	0. 10	0.6 0.6	14. 14.	81. 81.
MARGALIDA	Recrystallized grains	252 65.9	17 .7	42 42. 1	2. 3	2.2 88. 4	182 182 182 182 182	0. 5	0. 5	0. 7	12. 12.	13 5.4	0. 09	1.8 1.8	14. 14.	4.5 4.5
MARGALIDA	Recrystallized grains	255 35.7	20 .7	43 39. 7	2. 4	1.0 82. 3	182 182 182 182 182	3. 4	3. 2	0. 2	3.7 3.7	13 7.3	0. 07	0.3 0.3	1.0 1.0	3.7 3.7

BENT AILLO U	Coarse grains	312 85.9	35 .0	13 6	8. 9	0.7 4	187 73.	0. 2	0. 1	0. 1	0.1 0.1	37 2.6	0. 15	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	321 26.7	36 .8	13 56. 6	9. 1	0.8 1	186 89. 3	0. 3	0. 1	0. 1	0.1 0.1	37 9.5	0. 15	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	328 55.0	38 .1	13 12. 1	9. 6	1.3 4	187 76. 4	0. 2	0. 1	0. 1	0.9 0.9	39 8.0	0. 17	0.1 0.1	0.4 0.4
BENT AILLO U	Coarse grains	344 96.9	41 .3	13 60. 3	9. 9	0.8 0.9	187 05. 9	0. 2	0. 1	0. 1	0.1 0.1	40 9.7	0. 18	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	352 41.3	44 .1	15 07. 9	10 .7	0.7 0.7	187 63. 6	0. 3	0. 1	0. 1	0.1 0.1	42 0.7	0. 19	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	359 88.7	41 .8	13 61. 4	10 .1	0.7 0.7	187 68. 9	0. 3	0. 1	0. 1	0.1 0.1	43 6.5	0. 19	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	370 74.6	44 .5	14 17. 3	10 .8	1.3 1.3	187 53. 3	0. 2	0. 0	0. 1	0.6 0.6	46 0.7	0. 20	0.0 0.0	0.3 0.2
BENT AILLO U	Coarse grains	396 18.5	50 .4	15 07. 5	11 .7	0.6 0.6	186 86. 1	0. 2	0. 1	0. 1	0.1 0.1	47 2.0	0. 21	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	413 32.8	45 .8	16 06. 7	10 .9	0.5 0.5	186 34. 9	0. 2	0. 1	0. 1	0.1 0.1	50 6.6	0. 22	0.0 0.0	0.0 0.0
BENT AILLO U	Coarse grains	427 94.1	48 .5	13 93. 1	11 .0	0.6 0.6	187 30. 0	0. 2	0. 0	0. 1	0.2 0.2	53 5.6	0. 22	0.0 0.0	0.1 0.1
BENT AILLO U	Coarse grains	440 66.1	46 .8	14 03. 0	10 .9	0.5 0.5	186 51. 8	0. 3	0. 1	0. 1	0.2 0.2	55 8.8	0. 23	0.0 0.0	0.0 0.1
BENT AILLO U	Coarse grains	460 84.2	55 .1	13 75. 8	11 .5	0.5 0.5	186 86. 8	0. 3	0. 1	0. 1	0.2 0.2	58 4.1	0. 25	0.0 0.0	0.0 0.0
BENT AILLO U	Type 1 dissemina ted	190 52.3	18 .1	23 56. 5	9. 4	3.1 3.1	187 53. 5	1. 5	0. 3	0. 9	0.9 0.9	18 8.7	0. 16	0.3 0.3	73. 0
BENT AILLO U	Type 1 dissemina ted	184 46.3	14 .6	20 58. 7	6. 7	1.9 1.9	184 75. 7	0. 3	0. 1	0. 5	0.5 0.5	19 1.1	0. 13	0.2 0.2	24. 5
BENT AILLO U	Type 1 dissemina ted	185 49.9	17 .2	23 61. 5	8. 2	16. 7	186 39. 4	0. 7	0. 3	0. 6	1.9 1.9	18 1.6	0. 14	0.3 0.8	15. 4
BENT AILLO U	Type 1 dissemina ted	187 64.8	17 .1	26 79. 7	8. 9	5.8 5.8	186 05. 1	0. 5	0. 2	0. 5	1.4 1.4	18 1.9	0. 14	0.3 0.5	48. 1
BENT AILLO U	Type 1 dissemina ted	193 33.6	14 .6	21 05. 8	7. 6	9.6 9.6	187 05. 2	1. 1	0. 4	0. 9	2.0 2.0	18 4.9	0. 15	0.4 0.4	21. 8
BENT AILLO U	Type 1 dissemina ted	182 16.4	16 .8	19 74. 9	7. 6	3.6 3.6	186 39. 0	0. 4	0. 2	0. 6	0.5 0.5	19 0.6	0. 14	0.2 0.2	25. 1

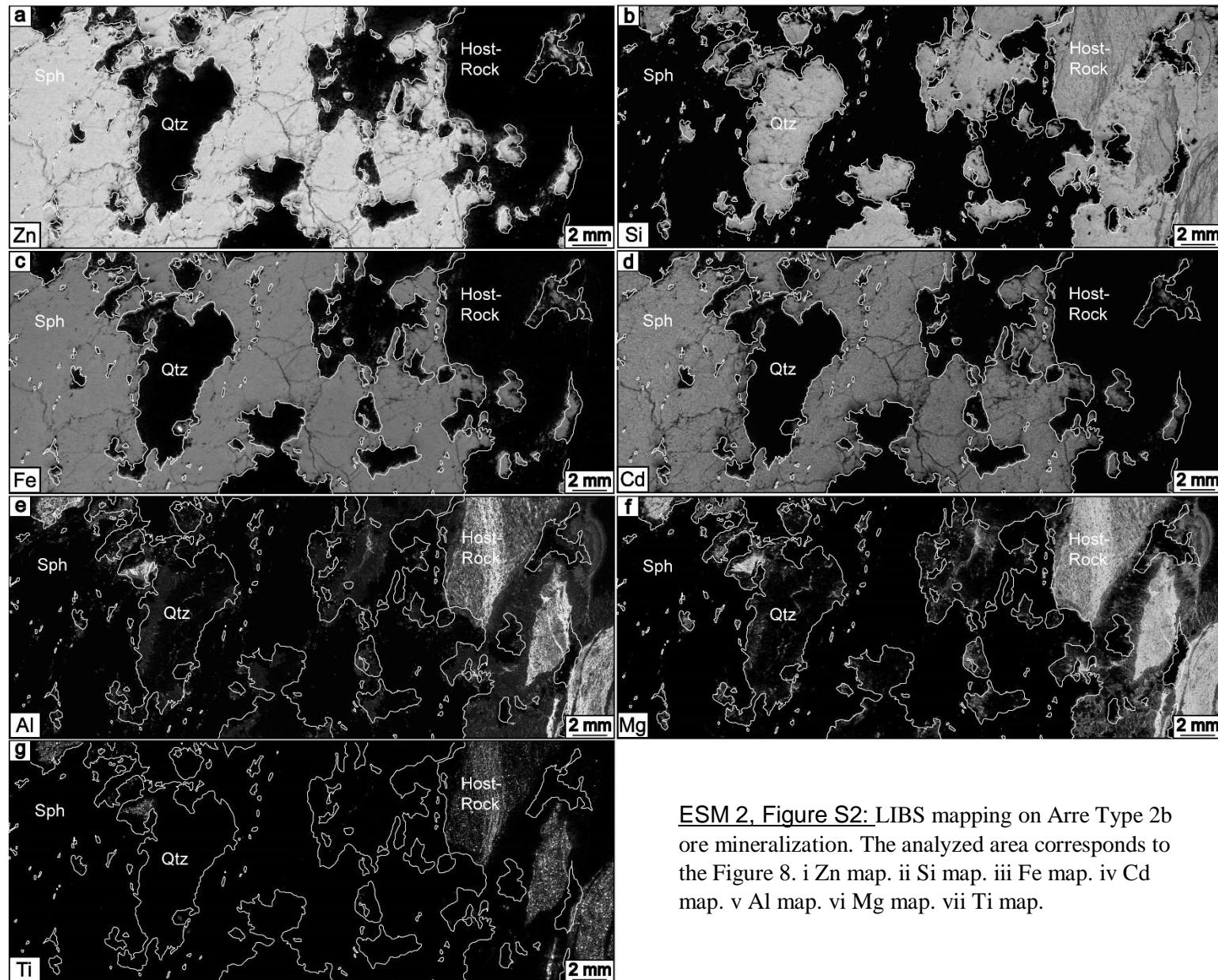
BENT AILLO U	Type 1 dissemina ted	188 40.0	15 .0	25 9	8. 9	9.2 4	186 42.	1. 2	0. 5	1. 4	2.3 8.1	18 8.1	0. 20	0.5 0.5	0.8 3	63. 3
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Electronical supplementary material: Article Title: Behavior of critical metals in metamorphosed Pb-Zn ore deposits: example from the Pyrenean Axial Zone / Journal: Mineralium Deposita / Authors: Alexandre Cugeronne, Bénédicte Cenki-Tok, Manuel Muñoz, Kalin Kouzmanov, Emilien Oliot, Vincent Motto-Ros, Elisabeth Le Goff / Affiliation of the corresponding author: Géosciences Montpellier, Université de Montpellier, CNRS, France / mail: alex.cugeronne@gmail.com

ESM 2, Figure S1: Time-integrated representative LA-ICP-MS spectra of sphalerite. Ge spectra is represented in red color. i Spectra in coarse grains with dark domains. Note the smooth profile for Ge (401 ppm) and most of the element (except for Sb) which contain micrometric minerals (Anglas sphalerite, cc_37). ii Spectra in coarse grain with light domains. Note the smooth and lower signal for Ge (76 ppm) and a small Sb mineral at the beginning of the signal acquisition (~ 180 sec.) (Anglas sphalerite, cc_25). iii Spectra in recrystallized grains with very low Ge content (0.8 ppm) and relatively smooth signal for the other elements (Anglas sphalerite, cc_34). iv Spectra in recrystallized grain with the presence of Fe-Ge-Cu-rich phases. The signal is divided in specific rich-phases field (1) and sphalerite field (2) (Anglas sphalerite, cc_10). v Spectra in recrystallized grain with occurrence of high Pb content (9420 ppm) in two micro-phases (Pale Bidau sphalerite, ce_18).



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ESM 2, Figure S2: LIBS mapping on Arre Type 2b ore mineralization. The analyzed area corresponds to the Figure 8. i Zn map. ii Si map. iii Fe map. iv Cd map. v Al map. vi Mg map. vii Ti map.

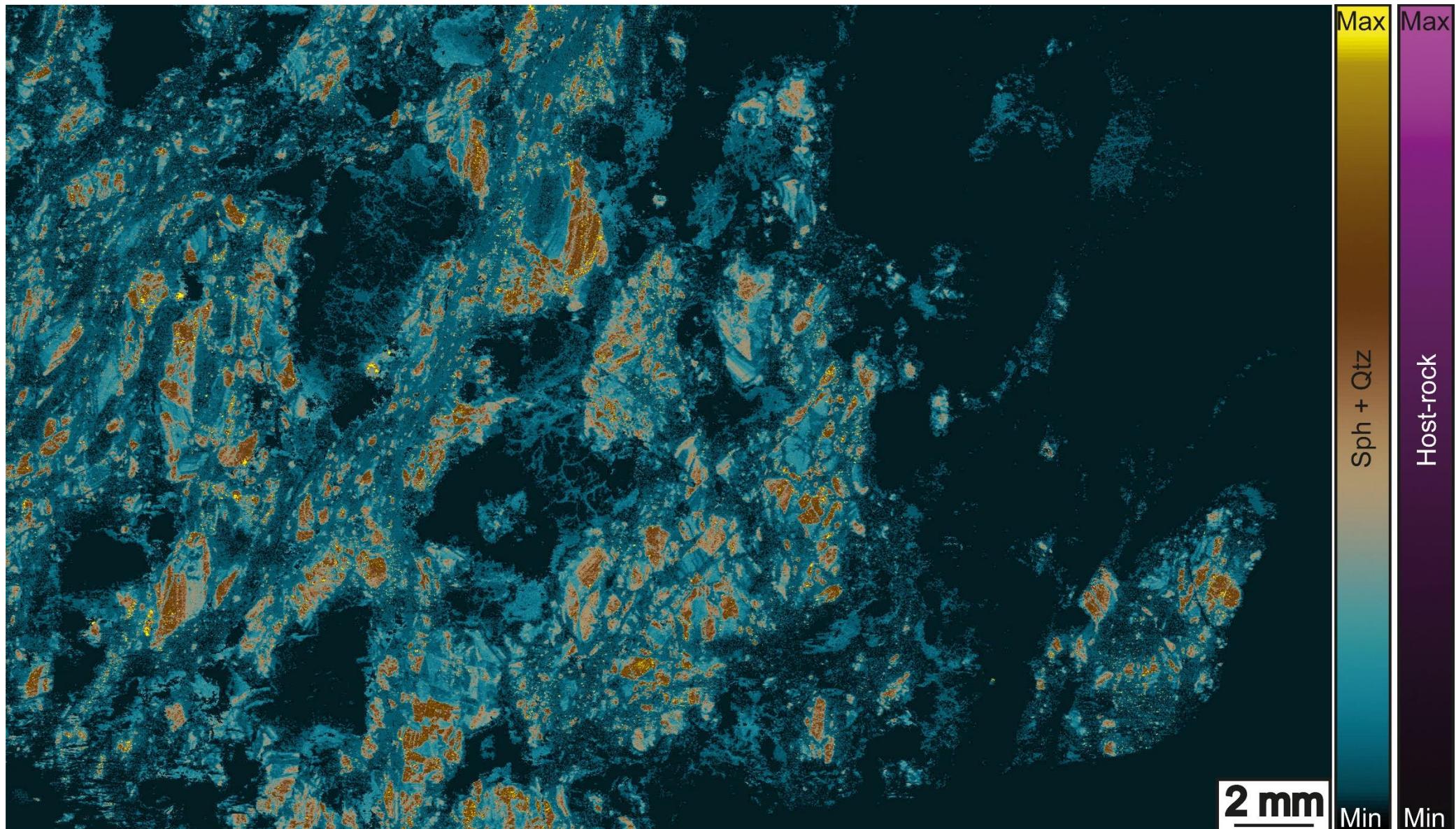
Electronical supplementary material: Article Title: Behavior of critical metals in metamorphosed Pb-Zn ore deposits: example from the Pyrenean Axial Zone / Journal: Mineralium Deposita / Authors: Alexandre Cugeronne, Bénédicte Cenki-Tok, Manuel Muñoz, Kalin Kouzmanov, Emilien Oliot, Vincent Motto-Ros, Elisabeth Le Goff / Affiliation of the corresponding author: Géosciences Montpellier, Université de Montpellier, CNRS, France / mail: alex.cugeronne@gmail.com

ESM 2, Figure S3: Transmitted-light microphotograph of the ore mineralization analyzed with LIBS (Arre sample).



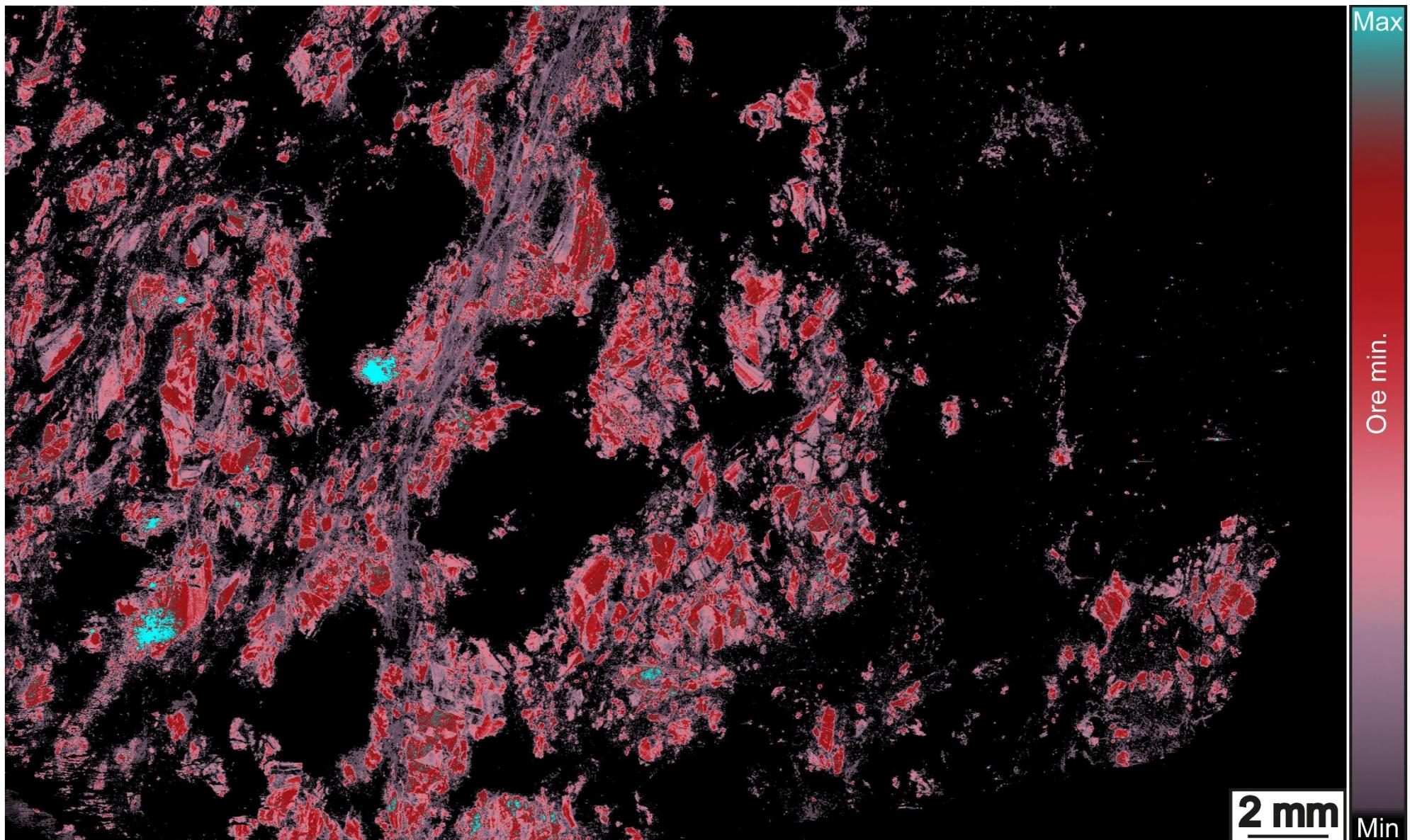
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ESM 2, Figure S4: Germanium high-resolution LIBS map (Arre sample).



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ESM 2, Figure S5: Copper high-resolution LIBS map (Arre sample).



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ESM 2, Figure S6: Gallium high-resolution LIBS map (Arre sample).

