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Textural signatures of sediment supply in gravel-bed rivers: revisiting the armour ratio

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Abstract:

The surface of the streambed in gravel-bed rivers is commonly coarser than the underlying bed material. This surface coarsening, or ‘armouring’, is usually described by means of the ratio between surface and subsurface grain-size metrics (the ‘armour ratio’). Such surface coarsening is typical of river reaches that are degrading due to a deficit in sediment supply (e.g. gravel-bed reaches below dams or lakes), but non-degrading gravel-bed streams may also exhibit some degree of armouring in relation to specific hydrological patterns. For instance, selective transport during the recession limbs of long lasting floods may coarsen the bed more significantly than flash floods. Consequently, regional differences in bed coarsening should exist, reflecting in turn the variability in sediment and water regimes. In this paper, we explore the trends linking armour ratios to sediment supply, taking into account the differences in hydrological context. We based our analysis on a large data set of bedload and grain size measurements from 49 natural gravel-bed streams and four flume experiments compiled from the scientific literature. Our main outcome documents how the balances between sediment yields and transport capacities have a quantifiable reflection on the armour ratios measured in the field: we report strong and statistically significant correlations between bedload fluxes and surface grain-size, and an asymptotic rise in armour ratios with the decline of sediment supply. Hydrological controls are also observed, but they are small compared to the signal related to sediment supply. Additionally, the trends observed in the field data are comparable to those previously documented in flume experiments with varying sediment feed. In this regard, different kinds of bedforms and particle arrangements have been commonly described with progressive reductions in sediment inputs and the subsequent coarsening of the streambed. Hence, armour ratios serve as a proxy for the general organization of the streambed of gravel-bed streams, and our results quantify this streambed adjustment to the dominant sediment supply regime.

Keywords: Gravel-bed rivers, Armour ratio, Sediment supply, Bedload

1. Introduction

A longstanding idea in fluvial geomorphology is that balances between sediment transport capacities, water discharge and sediment supply influence channel geometry (Parker et al., 2007; Parker, 2008), bed slope (e.g. Lane, 1955; Borland, 1960; Wilcock et al., 2009), streambed texture (Dietrich et al., 1989; Nelson et al., 2009; Venditti et al., 2017) and planform morphology (Church, 2006). Short-term fluctuations in sediment supply (e.g. hillslope processes, bank erosion, fine release from the bed after large floods) lead to local and temporary adjustments in streambed texture (e.g. Church et al., 1998; Clayton and Pitlick, 2008; Turowski et al., 2011) and channel morphology (Hassan and Zimmerman, 2012). In the long term, the balances between sediment and water yields seem to exert a conspicuous control on dominant channel styles (Montgomery and Buffington, 1997; Buffington, 2012) and river metamorphosis (Métivier and Barrier, 2012).

Seminal flume experiments by Dietrich et al. (1989), and early field observations (e.g. Gessler, 1967; Willets et al., 1988), reported how reductions in sediment supply tend to promote active channel narrowing, surface coarsening, bedload fining and transport rate decrease in gravel-bed rivers. Coarsening and fine sediment depletion in the riverbed is driven by a combination of winnowing during low flows (Gomez, 1983, 1993, 1994), infiltration of fine sediment (Marion and Fraccarollo, 1997; Curran and Waters, 2014; Berni et al., 2018) and kinematic sorting (Wilcock, 2001; Bacchi et al., 2014; Ferdowski et al., 2017) during bed load transport. Substantial subsequent work documents the influence of sediment inputs on the spatial and vertical patterns of grain size sorting (Nelson et al., 2009; 2010), the development of particle arrangements (e.g. Church et al., 1998; Venditti et al., 2017) and overall streambed mobility (Richards and Clifford, 1991; Pfeiffer and Finnegan, 2018).

According to all of the above, surface grain-size responds to decreases in bedload through the expansion of coarse fixed patches (Nelson et al., 2009; Yager et al., 2015), resulting in a general coarsening of the streambed (Dietrich et al., 1989). Hence, the degree of the latter may provide an idea of the dominant sediment supply conditions of a given river reach (Dietrich et al., 1989; Sklar et al., 2009; Venditti et al., 2017). For this reason, fluvial geomorphologists have often used

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62 29 the field observation of surface coarsening as a way to characterize streambed mobility and/or to
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64 30 diagnose the magnitude of bed degradation below dams, for example (see Rollet et al. 2013;
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66 31 Vázquez-Tarrio et al. 2019). However, disparities in the patterns of sediment supply variability
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68 32 (e.g. at the annual, seasonal, intra-flood scale) between different hydrological contexts complicate
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70 33 the interpretation of surface coarsening measured in the field and its application to river diagnosis.
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72 34 For instance, rivers experiencing long-lasting and sustained floods tend to exhaust fine sediment
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74 35 from the riverbed, which in turn enhances the degree of coarsening compared to the situation
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76 36 expected in streams with comparable sediment supplies, but submitted to flash floods (Laronne
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78 37 et al., 1994; Hassan et al. 2006). The hydrological regime is then a source of variability in the
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80 38 degree of surface coarsening that overlaps the signal related to the dominant sediment supply
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82 39 conditions. Moreover, sediment supply is difficult to quantify due to inherent complexities in
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84 40 bedload measurement (Pitlick et al., 2012); therefore, the field assessment of streambed response
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86 41 to sediment supply fluctuations is not easy.

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88 42 In this paper, we build on these ideas in order to quantify the relative weight of sediment supply
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90 43 versus hydrological controls in governing surface armouring. To accomplish this, we based our
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92 44 analysis on a large compilation of grain-size measurements and bedload discharge information
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94 45 extracted from the scientific literature for 49 natural gravel-bed rivers and four flume experiments.
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96 46 We structured this analysis in two steps. First, we performed a wide review of the compiled grain-
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98 47 size data, from which metrics adequate for exploring the links between bed texture and sediment
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100 48 supply were identified. Then, we performed a meta-analysis of the compiled data (based on the
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102 49 metrics introduced in the review) in order to highlight the main trends in armour ratios, and their
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104 50 covariation with channel hydraulics and bedload fluxes. Understanding the linkages between
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106 51 streambed texture and sediment transport regime is a key question in fluvial geomorphology, river
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108 52 ecology, civil engineering and river restoration (Pfeiffer et al., 2017). Consequently, our analysis
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110 53 may have interesting implications for near-future research on gravel-bed rivers.
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121 **54 2. Compiled data**
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123 55 The dataset used in the present paper consists of grain-size and bedload measurements collected
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125 56 at 49 river sites (summarized in Table 1). An important amount of the compiled data derives from
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127 57 the extensive campaign of sediment transport measurements carried out on Idaho, Nevada (King
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129 58 et al., 2004), Colorado and Wyoming rivers (Ryan et al., 2002; 2005), which probably represents
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131 59 some of the best available datasets on bedload in gravel-bed rivers to date. These data have
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133 60 previously been presented and analysed in several papers (Ryan et al., 2002; 2005; King et al.,
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135 61 2004; Barry et al., 2004; Mueller et al., 2005; Muskatirovic, 2008; Pitlick et al., 2008). The
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137 62 remaining data come from comparable measurements in other gravel-bed streams (Milhous, 1973;
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139 63 Emmet and Seitz, 1974; Seitz, 1977; Jones and Seitz, 1980; Reid and Frostick, 1986; Lisle, 1986;
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141 64 1989; Williams and Rosgen, 1986; Gomez, 1988; Kuhle, 1992; Lisle and Madej, 1992; Andrews,
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143 65 1994; Reid et al., 1995; Madej and Ozaki, 1996; McLean et al., 1999; Almedej, 2002; Church and
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145 66 Hassan, 2002; Wilcock and Kenworthy, 2002; Erwin et al., 2011; Mueller and Pitlick, 2014).
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147 67 Grain sizes (surface and substrate/subsurface) for each selected river were obtained from the
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149 68 graphical reading of grain-size curves extracted from the corresponding papers. When bed
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151 69 material was sampled at several locations or moments in the same river, we averaged the results
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153 70 to obtain a characteristic grain-size measure for each case study. Stream discharge information is
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155 71 available for 44 of the 49 selected case studies, together with width-averaged data on the main
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157 72 flow characteristics (velocity, active width). Using this information, we computed bed shear stress
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159 73 based on Rickenmann and Recking's (2011) fit to Ferguson's (2007) friction law (for more details
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161 74 see the supplementary files). We also compiled, for each case study, values for the representative
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163 75 channel-forming or dominant discharge (Table 2), which were derived from the information
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165 76 provided in the original papers about the bankfull discharge (in single-thread channels) or the ~1
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167 77 to 2-year return period discharge (in multi-thread channels) (Table 2). Information on bedload
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169 78 discharges was also available (Table 1); in this regard, we acknowledge the great work of data
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171 79 compilation carried out by Recking (2010; 2013), who provided bedload information for these
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173 80 field sites as supporting files (ibid).
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River/Reach	Source	Channel style	Slope	Surface D ₈₄ (mm)	Subsurface D ₈₄ (mm)	GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Q _b (g/s·m)	N	Sediment sampling method
						Surface	Subsurface						
Big Wood River	<i>King et al. (2004)</i>	Plane-bed	0.091	250	101	Pebble count and core sampling	Core sampling	12.8	0.4-1.1	6.0-30.9	0.0-336.4	100	HS 7.62 or 15.2 cm (0.25 mm mesh)
Blackmare Creek	<i>King et al. (2004)</i>	Plane-bed	0.03	220	97	Pebble count and core sampling	Core sampling	4.94-11.89	0.1-0.5	0.3-4.7	0.0-6.8	88	HS 7.62 or 15.2 cm (0.25 mm mesh)
Boise river	<i>King et al. (2004)</i>	Riffle and pool	0.0038	141	86	Pebble count	Core sampling	52.4-61.0	0.6-2.1	33.7-291.7	0.4-633.5	82	HS 7.62 or 15.2 cm (0.25 mm mesh)
Borgne d'Arolla	<i>Gomez (1988)</i>	Step-Pool	0.03	19	12	Contact sampling technique	10 kg-volumetric sampling	0.3-2.2	0.0-0.1	0.0-0.3	56.2-837.0	31	HS (7.6 cm), 0.5 mm mesh
Chultnana River	<i>Williams and Rossen (1989)</i>	Braided	0.0008	52	184	Pebble count (?)	Dredge sampling	98.5-309.0	1.7-3.6	212.0-7104.0	213.0-2590.0	43	HS (7.62 cm)
Clearwater River	<i>Jones and Seiz (1980)</i>	Riffle and pool	0.00037	70	70	Pebble count	Sieve analysis of dug material	125.0-149.0	3.4-46.3	288.0-3511.0	0.1-284.0	78	HS (7.6 and 15 cm)
Dollar Creek	<i>King et al. (2004)</i>	Plane-bed	0.0146	145	83	Pebble count and core sampling	Core sampling	7.0-11.9	0.2-0.5	0.4-6.4	0.0-9.7	85	HS 7.62 or 15.2 cm (0.25 mm mesh)
East Fork San Juan	<i>Ryan et al. (2005)</i>	Braided	0.008	112	52	Pebble count	Barrel method	15.0-17.2	0.3-0.5	2.8-13.8		77	Wadable version of Elwha sampler, 102 × 203 mm
East Saint-Louis Creek	<i>Ryan et al. (2002)</i>	Step-Pool	0.058	142	23	Pebble count	Barrel method	2.8-3.0	0.1-0.4	0.1-1.24	0.0-21.2	109	HS (7.6 cm)
Fool Creek	<i>Ryan et al. (2002)</i>	Plane-bed	0.053	100	59	Pebble count	Barrel method	1.7-2.1	0.1-0.2	0.0-0.5	0.0-14.7	95	HS (7.6 cm)

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River/Reach	Source	Channel style	Slope	Surface D ₈₄ (mm)	GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Q _b (g/s·m)	N	Sediment sampling method
					Surface	Subsurface D ₈₄ (mm)						
Fraser River	<i>McLean et al. (1999); Ferguson and Church (2009)</i>	Riffle and pool	0.00046	70	68	Pebble count	510	-	1085-11445	0.3-486.3	76	Basket sampler (610 × 255 mm) for high flows and half-size VuV sampler (225 × 115 mm) for lower flows
Goodwin Creek	<i>Kuhnle (1992); Almedeij (2002)</i>	Riffle and pool	0.0021	30	30	Pebble count	11.1-14.6	0.4-1.2	1.4-21.6	0.2-2980.0	357	HS (58 cm2 with trapezoidal shape), 0.25mm net mesh
Harris Creek	<i>Church and Hassan (2002)</i>	Riffle and pool	0.013	100	55	Pebble count (grid-by-number)	15	-	4.2-18.4	0.0-4.3	22	Sediment trap
Jacoby	<i>Lisle (1986); Lisle (1989); Almedeij (2002); Wilcock and Kenworthy (2002)</i>	Riffle and pool	0.0063	95	81	Pebble count	17.2	-	0.6-18.51	0.0-413.0	100	HS (4.4 cm)
Johns Creek	<i>King et al. (2004)</i>	Step-Pool	0.0207	558	63	Pebble count	8.2-14.6	0.3-1.2	1.0-34.3	0.0-10.7	46	HS 7.62 or 15.2 cm (0.25 mm mesh)
Little Buckhorn Creek	<i>King et al. (2004)</i>	Step-Pool	0.0509	340	94	Pebble count and core sampling	1.4-4.6	0.1-0.7	0.1-0.7	0.0-18.5	78	HS 7.62 or 15.2 cm (0.25 mm mesh)

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River/Reach	Source	Channel style	Slope	Surface D ₈₄ (mm)	Subsurface D ₈₄ (mm)	GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Q _b (g/s·m)	N	Sediment sampling method
						Surface	Subsurface						
Little Granite Creek	<i>Ryan et al. (2002)</i>	Plane-bed	0.019	220	41	Pebble count	Barrel method	6.5-11.2	-	0.7-11.6	0.0-128.0	123	HS 7.62 cm
Little Slate	<i>King et al. (2004)</i>	Plane-bed	0.0268	380	141	Pebble count and core sampling	Core sampling	6.7-13.4	0.3-1.0	0.5-18.3	0.0-10.3	157	HS 7.62 or 15.2 cm (0.25 mm mesh)
Lochsa Creek	<i>King et al. (2004)</i>	Plane-bed	0.0023	245	123	Pebble count and core sampling	Core sampling	67.1-83.0	1.8-3.1	110.7-758.9	0.0-48.3	72	HS 7.62 or 15.2 cm (0.25 mm mesh)
Lolo Creek	<i>King et al. (2004)</i>	Plane-bed	0.0097	140	68	Pebble count and core sampling	Core sampling	10.7-16.0	0.3-1.5	1.8-23.0	0.0-13.4	89	HS 7.62 or 15.2 cm (0.25 mm mesh)
Middle Fork	<i>King et al. (2004)</i>	Plane-bed	0.0041	288	140	Pebble count and core sampling	Core sampling	42.7-67.1	1.2-2.01	83.5-433.2	0.1-727.7	64	HS 7.62 or 15.2 cm (0.25 mm mesh)
Middle Fork Piedra River	<i>King et al. (2004)</i>	Riffle and pool	0.011	210	43.3	Pebble count and core sampling	Core sampling	11.4-13.8	0.2-0.5	1-11.0	0.0-216.6	86	HS 7.62 or 15.2 cm (0.25 mm mesh)
Nahal Yatir	<i>Reid et al. (1995)</i>	Riffle and pool	0.0088	13	34	Removing clasts from spray-painted stripes and laboratory sieving	Bulk volume sampling	3.5	0.1-0.6	0.3-0.5	200.0-7050.0	74	Sediment trap
North Fork Clearwater	<i>King et al. (2004)</i>	Plane-bed	0.0005	270	104	Pebble count and core sampling	Core sampling	9.1-93.6	1.7-34.1	100.8-974.1	0.0-732.2	72	HS 7.62 or 15.2 cm (0.25 mm mesh)
Oak Creek	<i>Milhous (1973)</i>	Plane-bed	0.0083	80	52	Pebble count and bottom samplers	Bulk volume sampling	3.7	-	0.0-3.4	0.0-111.0	119	Sediment trap, vortex tube

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River/Reach	Source	Channel style	Slope	Surface D ₈₄ (mm)	Subsurface D ₈₄ (mm)	GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Q _b (g/s·m)	N	Sediment sampling method
						Surface	Subsurface						
Pacific creek	<i>Erwin et al. (2011)</i>	Braided	0.0035	45	28	Pebble count	Bulk volume sampling	-	-	-	-	-	-
Rapid River	<i>King et al. (2004)</i>	Plane-bed	0.0108	170	101	Pebble count	Core sampling	11.4-18.6	0.2-0.9	0.9-36.8	0.0-294.3	190	HS 7.62 or 15.2 cm (0.25 mm mesh)
Redwood Creek	<i>Lisle and Madej (1992); Madej and Ozaki (1996)</i>	Riffle and pool	0.0014	18	20	Pebble count	Bulk volume sampling	11.7-70	-	1.8-569	8.1-5067.8	221	HS sampler (7.6 cm)
Saint-Louis Creek	<i>Ryan et al. (2002; 2005)</i>	Plane-bed	0.0110-0.0450	162-543	33-78	Pebble count	Barrel method	5.2-10.3	0.1-0.4	0.4-7.2	0.0-65.7	813	HS (7.6 cm)
Sagehen Creek	<i>Andrews (1994)</i>	Riffle and pool	0.0102	104	96	Pebble count	Bulk volume sampling	2.6	0.3-1.6	1.0-3.1	0.5-34.9	55	HS (15 cm)
Salmon River below Yankee	<i>King et al. (2004)</i>	Plane-bed	0.0034	276	99	Pebble count and core sampling	Core sampling	30.5-38.4	1.2-1.9	38.5-143.6	0.0-98.8	60	HS 7.62 or 15.2 cm (0.25 mm mesh)
Salmon River near Obsidian	<i>King et al. (2004)</i>	Plane-bed	0.0066	128	84	Pebble count and core sampling	Core sampling	12.0-14.3	0.7-0.9	7.5-21.0	0.7-103.4	50	HS 7.62 or 15.2 cm (0.25 mm mesh)
Salmon River near Shoup	<i>King et al. (2004)</i>	Plane-bed	0.0019	174	136	Pebble count and core sampling	Core sampling	46.5-99.5	1.7-2.7	108.4-540.8	1.0-536.1	60	HS 7.62 or 15.2 cm (0.25 mm mesh)
Selway River	<i>King et al. (2004)</i>	Plane-bed	0.0021	265	173	Pebble count and core sampling	Core sampling	82.3-97.8	1.4-2.8	134.8-1067.5	0.0-43.5	72	HS 7.62 or 15.2 cm (0.25 mm mesh)

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River/Reach	Source	Channel style	Slope	Surface		Subsurface		GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Qb (g/s·m)	N	Sediment sampling method
				D ₈₄ (mm)	D ₈₄ (mm)	Surface	Subsurface	Surface	Subsurface						
Silver Creek	<i>Ryan et al. (2005)</i>	Plane-bed	0.0450	73	33	Pebble count	Barrel method	3.8-4.4	0.1-0.3	0.1-1.4	0.0-214.3	57	Wadable version of the Elwha sampler, 102 × 203 mm		
Snake River	<i>Emmett and Seitz (1974); Seitz (1977)</i>	Riffle and pool	0.0009	115	54	Pebble count	Sieve analysis of dug material	155.4-204.2	3.3-6.2	779.0-4559.0	0.0-342.0	63	HS (7.6 and 15 cm)		
Snake River below Jackson Lake	<i>Mueller and Pitlick (2014)</i>	Braided	0.0025	83	58	Pebble count	Bulk volume sampling	-	-	-	-	-	-		
South Fork Payette	<i>King et al. (2004)</i>	Plane-bed	0.004	150	79	Pebble count and core sampling	Core sampling	43.6-51.8	0.4-1.7	20.4-180.9	-	-	HS 7.62 or 15.2 cm (0.25 mm mesh)		
South Fork Red River	<i>King et al. (2004)</i>	Plane-bed	0.0146	150	161	Pebble count and core sampling	Core sampling	5.8-12.2	0.1-0.9	0.2-13.0	0.0-29.0	204	HS 7.62 or 15.2 cm (0.25 mm mesh)		
Squaw Creek USGS	<i>King et al. (2004)</i>	Plane-bed	0.01	72	93	Pebble count and core sampling	Core sampling	3.3-14.1	0.2-0.4	0.1-7.6	0.0-23.0	92	HS 7.62 or 15.2 cm (0.25 mm mesh)		
Sunlight Creek-4	<i>Mueller and Pitlick (2014)</i>	Braided	0.0075	94.6	43	Pebble count	Bulk volume sampling	-	-	-	-	-	-		
Sunlight Creek-11	<i>Mueller and Pitlick (2014)</i>	Riffle and Pool	0.0091	79.2	55	Pebble count	Bulk volume sampling	-	-	-	-	-	-		
Susitna River near Talkeetns	<i>Williams and Rosgen (1986)</i>	Braided	0.0015	96	260	Pebble count, bulk sampling (?)	Bulk sampling (?), dredge sampling	118.0-202.0	1.1-14.0	240.0-1310.0	0.9-156.0	39	HS (7.62 cm)		

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River/Reach	Source	Channel style	Slope	Surface D ₈₄ (mm)	Subsurface D ₈₄ (mm)	GSD measuring method		Flow width (m)	Flow depth (m)	Q (m ³ /s)	Q _b (g/s·m)	N	Sediment sampling method
						Surface	Subsurface						
Susitna River at Sunshine, Alaska	<i>Williams and Rosgen (1986)</i>	Braided	0.0017	88	163	Pebble count, bulk sampling (?)	Bulk sampling (?), dredge sampling	174.0-311.0	2.1-4.4	504.0-2800.0	33.7-1500.0	41	HS 7.62 cm
Talkeetna River near Talkeetna	<i>Williams and Rosgen (1986)</i>	Braided	0.00096	100	184	Bulk sampling (?)	Bulk sampling (?), dredge sampling	-	-	-	-	-	-
Thompson Creek	<i>King et al. (2004)</i>	Plane-bed	0.0153	110	132	Pebble count and core sampling	Core sampling	4.2-6.7	0.2-0.4	0.2-3.5	0.0-38.0	84	HS 7.62 or 15.2 cm (0.25 mm mesh)
Trapper Creek	<i>King et al. (2004)</i>	Step-Pool	0.0414	122	67	Pebble count and core sampling	Core sampling	3.5-6.4	0.1-0.6	0.1-3.8	0.0-31.3	166	HS 7.62 or 15.2 cm (0.25 mm mesh)
Turkey Brook	<i>Reid and Frostick (1986)</i>	Riffle and pool	0.0142	42	35	No information	No information	3	0.1-0.9	0.1-13.8	0.0-50.6	206	Pit traps
Valley Creek	<i>King et al. (2004)</i>	Plane-bed	0.0040	160	78	Pebble count and core sampling	Core sampling	17.7-42.3	0.4-1.3	3.9-40.2	0.0-56.8	192	HS 7.62 or 15.2 cm (0.25 mm mesh)

Table 1. Sources of field data and information compiled for this study. D_{84} : 84-th percentile of the grain-size distribution (GSD). Q : Flow discharges. Q_b : Bedload discharges. *HS*: Helley-Smith sampler. N : number of data.

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River/Reach	Dominant discharge (m ³ /s)	Flow recurrence (years)	River/Reach	Dominant discharge (m ³ /s)	Flow recurrence (years)	River/Reach	Dominant discharge (m ³ /s)	Flow recurrence (years)	River/Reach	Dominant discharge (m ³ /s)	Flow recurrence (years)
Big Wood River	21.7	1.5	Harris creek	19.0	1	Oak Creek	3.4***	No info	Snake River below Jackson Lake	285**	1.5 - 2
Blackmare Creek	4.7*	1.1	Jacoby	19.6	No info	Rapid river	17.7*	1.4	South Fork Payette	86.4*	1.2
Boise river	167.1*	1.7	Johns creek	49.0*	3.4	Pacific creek	60.1	1.5 - 2	South Fork Red River	7.3*	1.5
Borgne d'Arolla	0.2*	<1	Little Buckhorn Creek	0.2	1	Redwood Creek	560*	2 - 5	Squaw Creek USGS	5.1*	1.6
Chulitnana River	1130**	2	Little Granite Creek	5.9*	1.5	Saint-Louis Creek	2.6 - 4.8*	1.5	Sunlight Creek-4	16.5**	1.5 - 2
Clearwater River	2662***	2.2 - 2.3	Little Slate	12.2*	1.4	Sagehen Creek	2	~1	Sunlight Creek-11	14**	1.5 - 2
Dollar Creek	6.4*	1.1	Lochsa Creek	446*	1.5	Salmon River below Yankee	118.1	1.5	Susitna River near Talkeetna	1270**	2
East Fork San Juan	15.7*	1.5	Lolo Creek	11.8*	1.2	Salmon River near Obsidian	12.5	1.5	Susitna River at Sunshine, Alaska	4020**	2
East Saint-Louis Creek	0.9*	1.5	Middle Fork	217	1.5	Salmon River near Shoup	320	1.5	Talkeetna River near Talkeetna	730**	2
Fool Creek	0.3*	1.5	Middle Fork Piedra River	10.1*	1.5	Selway River	651.3*	1.7	Thompson Creek	2.5*	1.6
Fraser River	8760	1	Nahal Yatir	7.9	No info	Silver Creek	1.3*	1.5	Trapper Creek	2.6*	1.9
Goodwin creek	10.6	No info	North Fork Clearwater	453.1*	1.5	Snake River	3426***	2.2 - 2.3	Turkey Brook	19.6	No info
									Valley Creek	24.1*	1.6

Table 2. Information on the dominant channel-forming discharge for the case study selected for the present meta-analysis. (*) Bankfull discharge. (**) Dominant discharge provided by Mueller and Pitlick (2014). (***) Reference discharge provided by Church and Rood (1983)

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507
508 81 We grouped the data following three different criteria. We made an initial classification according
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510 82 to dominant channel morphology, grouping the different case studies as riffle and pool, step-pool
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512 83 and plane-bed channels (after Montgomery and Buffington, 1997). Due to its geomorphological
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514 84 significance, we also defined a separate group for multi-thread rivers, in spite of the fact that each
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516 85 single thread of a braided river commonly shows a riffle and pool or bar-pool morphology. We
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518 86 also classified the different data according to the general sediment supply conditions at the
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520 87 catchment scale (Table 3). Following Recking (2012), we defined three main groups of data: i.
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522 88 low sediment supply, with channels draining highly vegetated watersheds and no clear active
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524 89 sediment sources and/or alluvial material; ii. moderate sediment supply, with rivers located in
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526 90 catchments in which significant bare land areas and/or sparse vegetation, and punctually
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528 91 distributed active sediment sources, are observed; and iii. high sediment supply, with channels
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530 92 well-coupled to landslides/slope deposits, or fed by strong bank erosion and/or bar-edge trimming
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532 93 (e.g. channels with braided morphology). We based this classification on the scarce information
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534 94 (study site description, photographs, etc.) provided by the original studies and our own inspection
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536 95 of the rivers through Google Earth. Finally, we also grouped the compiled data according to
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538 96 dominant flow regime, differentiating between: i. ‘rainfall-dominated’; ii. ‘snowmelt/rain-on-
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540 97 snow-dominated’; iii. ‘snowmelt-dominated’; iv. ‘glacial-fed’ and v. ‘flash-flood dominated’
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542 98 streams (Table 3).

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Flume	Source	S	W (m)	L (m)	h (cm)	D _{50s} (mm)	Shields	Feed rate	r (h)
Tsukuba	<i>Nelson et al. (2009)</i>	0.0035-0.0052	7.5	0.3	10.2-11.3	3.7-4.9	0.049-0.086	1.7-17.4g/min·cm	6-7.5
Berkeley	<i>Nelson et al. (2009)</i>	0.0043-0.0055	28	0.9	21.8-22.8	10.1-11.8	0.045-0.061	0-23.3 g/min·cm	20.7-28.9
UBC	<i>Church et al. (1998)</i>	0.001-0.012	0.5/0.8	6/10	0.5-7.4	1.9-5.1	0.003-0.117	No feed	-
UBC	<i>Hassan and Church (200)</i>	0.006/0.007	0.8	10	0.1-6.7	2.4-4.5	-	1.2-0.644 kg/h	96

100 **Table 4.** Sources of flume data compiled for this study. *S*: Flume slope. *W*: Flume width. *L*:
101 Flume length. *h*: water depth. *D*₅₀: 50-th percentile of sediment GSD. *r*: experiment duration.

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103 Apart from field data, we also benefited from the results of previous flume experiments that
104 explored the role of sediment feed reductions on the surface texture (Church et al., 1998; Hassan

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567 and Church, 2000; Nelson et al., 2009). Table 4 describes the main characteristics of these
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569 106 experiments. Information for these flume investigations was extracted from graphical reading of
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571 107 figures presented in Venditti et al. (2017) (Figures 16.1, 16.2 and 16.5 in that paper).
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575 109 **3. Review of the compiled data**

577 110 ***3.1. Surface coarsening in gravel-bed rivers: introducing the ‘armour ratio’***

579 111 Streambed surface is, in general, coarser than the underlying subsurface grain-size distribution
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581 112 (GSD) in the compiled dataset: the average D_{50} and D_{84} are both coarser on the surface than the
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583 113 subsurface GSD (Figure 1A). The degree of surface coarsening has usually been quantified in
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585 114 fluvial geomorphology through the ‘armour ratio’ (D^*): the ratio between a characteristic grain
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587 115 size (normally, the median size) on the surface GSD and the same characteristic grain size in the
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589 116 subsurface GSD (Bunte and Abt, 2001):

$$591 \quad 117 \quad D_i^* = \frac{D_{is}}{D_{iss}} \quad \text{Eq. 1}$$

593 118 where D_i refers to the i th-percentile of the GSD. A value of D^* equal to 1 means that the surface
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595 119 and subsurface GSD are very similar. When the surface is coarser than the subsurface GSD, then
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597 120 $D^* > 1$; the coarser the surface is compared to the subsurface GSD, the larger the D^* . Average
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599 121 armour ratios are larger in the compiled data if estimated using the median size rather than using
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601 122 the D_{84} (~ 3.1 against ~ 1.8 , respectively) (Figure 2), outlining that differences between both GSD
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603 123 are more important towards finer size fractions. In addition, some different tendencies could be
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605 124 appreciated according to channel morphology (Figure 3).

607 125 Apart from surface coarsening, some other complementary trends can be identified and seem
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609 126 intimately related to the former (Figures 1C and 1D). For instance, the percentages of fine
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611 127 sediment are, on average, larger in the subsurface ($\sim 15\%$) than on the surface GSD ($\sim 6\%$) (Figure
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613 128 1D). Furthermore, patterns of grain-size sorting also show differences: D_{84}/D_{50} sorting indexes
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615 129 are again larger in the subsurface (~ 3.6 on average) than on the surface GSD (~ 2.1 on average),
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617 130 suggesting that subsurface GSD is more poorly sorted (Figure 1C). All these differences are
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619 131 statistically significant (Welch’s t-test for unequal variances, p-value <0.05).

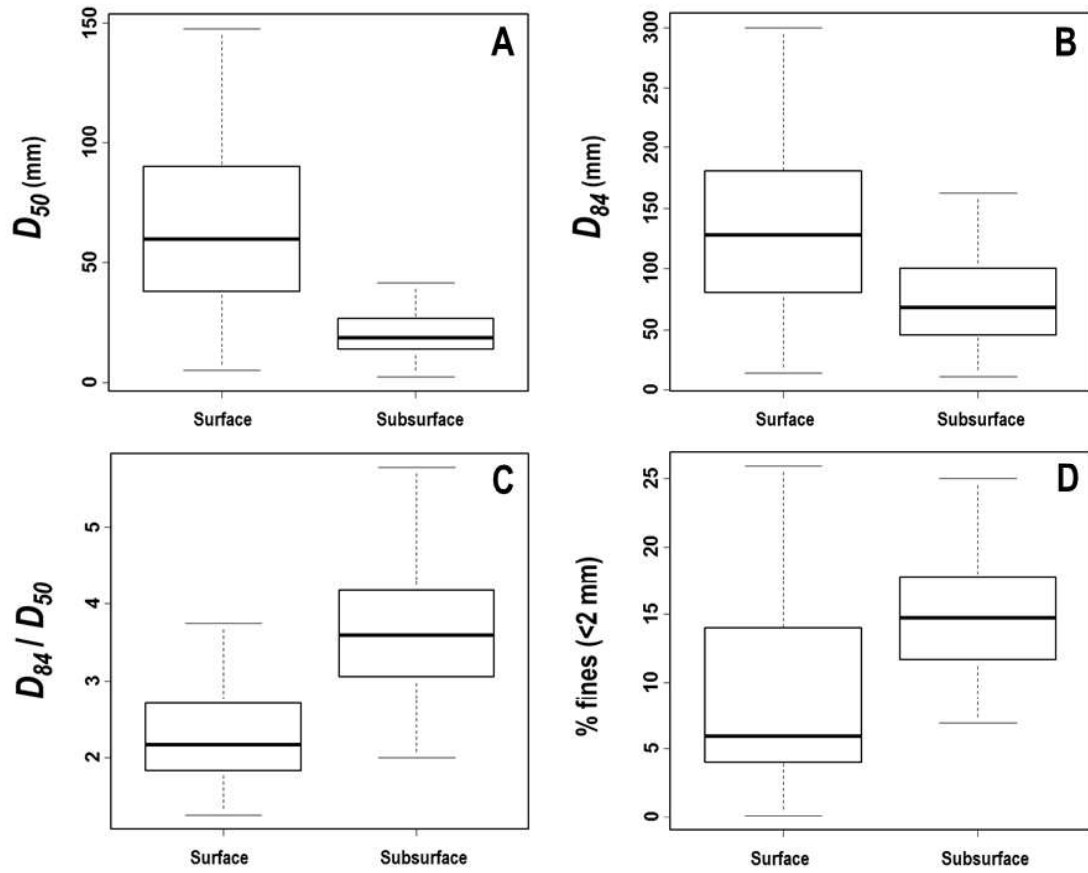


Figure 1. Comparison between surface and subsurface grain-size parameters in the compiled database. A: D_{50} (in mm); B: D_{84} (in mm). C: D_{84}/D_{50} ratio. D: Percentage of fines (< 2mm) present in the sediment.

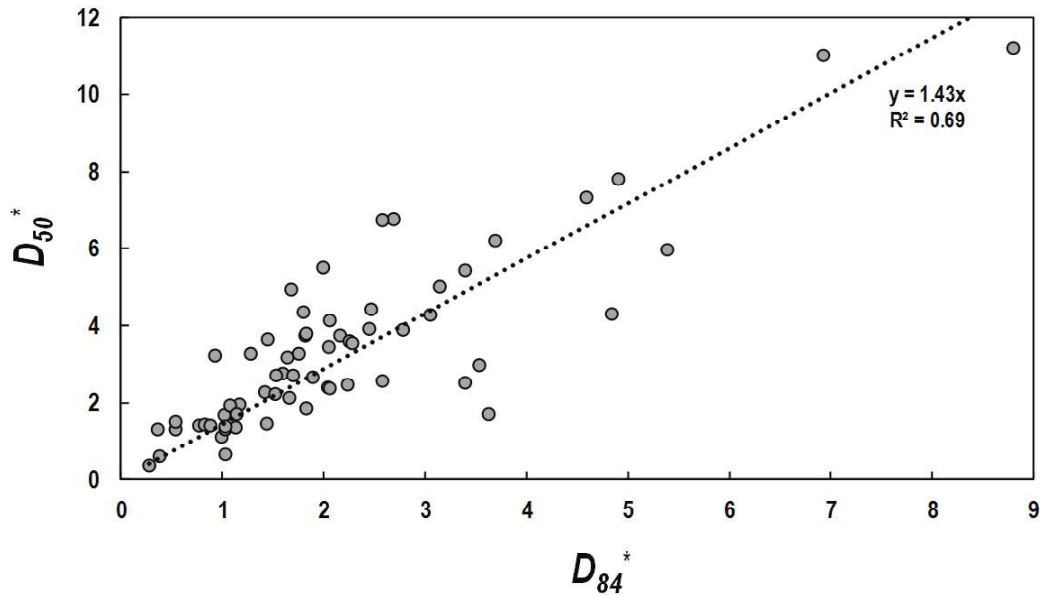


Figure 2. Armour ratios based on D_{50} plotted versus the armour ratios based on D_{84} .

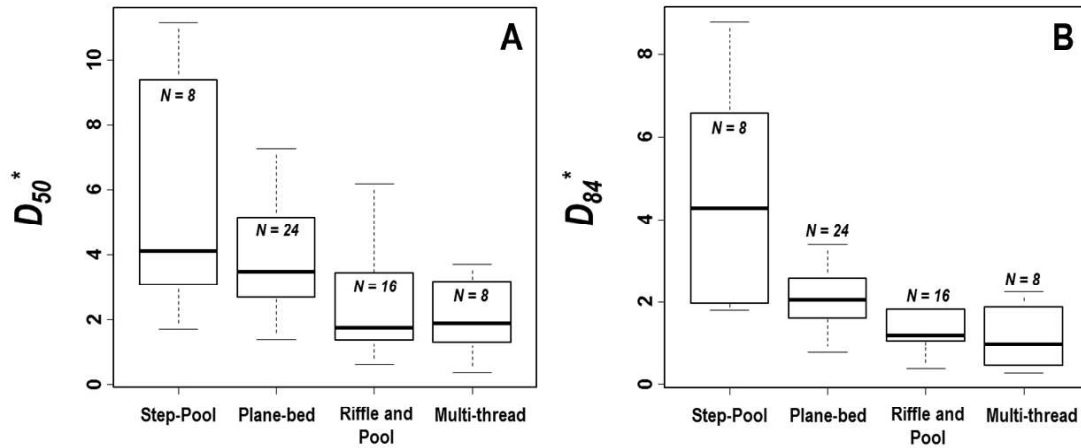


Figure 3. Differences in armour ratios according to channel morphology. A: Armour ratio based on the media size of the bed sediment (D_{50}). B: Armour ratio based on the 84-th percentile (D_{84}) of the grain-size distribution (GSD). N : Number of data.

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133 It could initially be thought that methodological biases may explain a part of these trends.

134 Subsurface GSD is commonly determined by sieving dredged or excavated bulk volume samples

135 (Church et al., 1987). Conversely, a wide diversity of methods have been proposed to approach

136 the surface GSD, such as the classical Wolman (1954) pebble-count method, the photosieving

137 approach (Ibbeken and Schleyer, 1986, Butler et al., 2001, Rubin, 2004, Graham et al., 2005,

138 Buscombe, 2008, Detert and Weitbrecht, 2013) or the more recent protocols based on high-

139 resolution topography (e.g. Heritage and Milan, 2009; Brasington et al., 2012; Vázquez-Tarrió et

140 al., 2017; Woodget et al., 2017). Grain sizes determined from area-by-weight sample methods

141 (like the photosieving approach) are not directly equivalent to grain sizes determined from

142 volume-by-weight sampling procedures (e.g. size sieving of bulk volume samples) (Kellerhals

143 and Bray, 1971; Bunte and Abt, 2001). However, in the study cases compiled here, surface GSD

144 was generally sampled using a grid-by-number pebble count, while subsurface GSD was always

145 obtained from one variant or another (dredging, digging, frozen cores...) of the bulk volume

146 sampling strategy (Table 1). Particle-size distributions determined from volume-by-weight (size-

147 sieving of bulk volume samples) and grid-by-number (pebble count) samples are said to be

148 equivalent (Kellerhals and Bray, 1971; Bunte and Abt, 2001; Rice and Haschenburger, 2004), and

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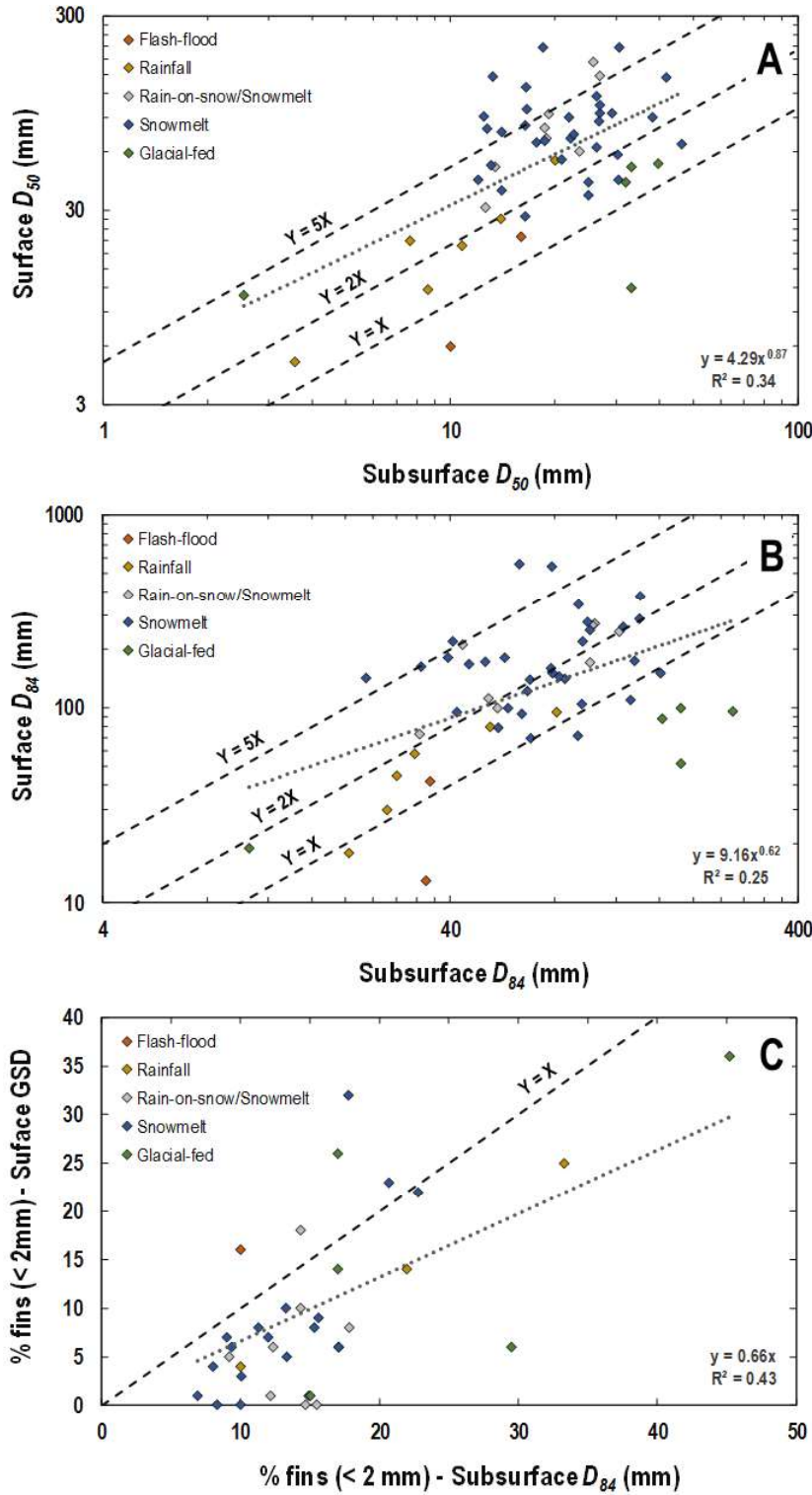
149 so surface coarsening seems to be a real trend in the compiled data, and not an artefact issuing
150 from methodological biases.

151 ***3.2. Sediment supply vs. hydrologic controls on surface coarsening and ‘armour ratios’***

152 The grain size of sediment inputs should exert an obvious control on the grain calibre of the
153 available bed material within a specific river reach, which in turn may condition the sediment size
154 of the streambed’s surface. In this regard, we could consider subsurface GSD as a proxy for the
155 GSD of the average annual sediment inputs (Kuhnle and Willis, 1992; Lisle, 1995; Church and
156 Hassan, 2005; Recking, 2013; Segura and Pitlick, 2015), i.e. the GSD available in the subsurface
157 should condition the GSD of the streambed surface. Within the compiled data, we observe some
158 statistically significant and moderate correlation between surface and subsurface GSD (Figure 4),
159 which highlights this undeniable influence of the sediment supply GSD on the GSD of the
160 streambed’s surface. However, it is interesting to notice how variance in grain-size data is larger
161 for surface than subsurface GSD (see Figures 1A and B). This suggests that grain-size variability
162 introduced by the sediment supplies is diluted by the variability in surface coarsening introduced
163 by some other controls.

164 In this regard, the compiled data show some statistically significant differences (ANOVA test, p-
165 value < 0.05) in armour ratios according to the amounts of sediment supply at the catchment scale
166 (Figure 5): rivers with high sediment supplies tend to show lower armour ratios than streams with
167 low sediment feeds. Surface coarsening (often named ‘armouring’) has been described in gravel-
168 bed settings for a long time (e.g. Gessler, 1967), and already related to the higher or lower
169 availability of upstream sediment inputs. Armouring has been, thereby, typically reported in
170 degrading beds and river reaches with low or no sediment supply (e.g. gravel-bed reaches below
171 dams or lakes) (Gessler, 1967; Willets et al., 1988; Chin et al., 1994; Gomez, 1994; Vericat et al.,
172 2006). In such cases, armouring is called ‘static’ or ‘pavement’ (Jain, 1990; Yager et al., 2015;
173 Bertin and Freidrich, 2018). Static armours can ‘break up’ during high peaks of flow and/or
174 transport episodes with large sand sediment supplies (Laronne and Carson, 1976; Gomez, 1983;
175 Klaasen, 1988; Chin et al., 1994; Venditti et al., 2005; Vericat et al., 2006; Wang and Liu, 2009;
176 Venditti et al., 2010; Spiller et al., 2012; Curran and Waters, 2014; Orrú et al., 2016; Bertin and

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 803 177 Freidrich, 2018) and re-form during the falling limb of the flood hydrograph, together with a
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 805 178 progressive reduction in sediment mobility and the winnowing of finer and more mobile particles
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 807 179 (Mao, 2012).
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Figure 4. Surface grain-size parameters plotted versus subsurface parameters. A: D_{50} . B: D_{84} . C: Percentage of fine sediment (< 2 mm). We have segregated the data according to hydrological regime.

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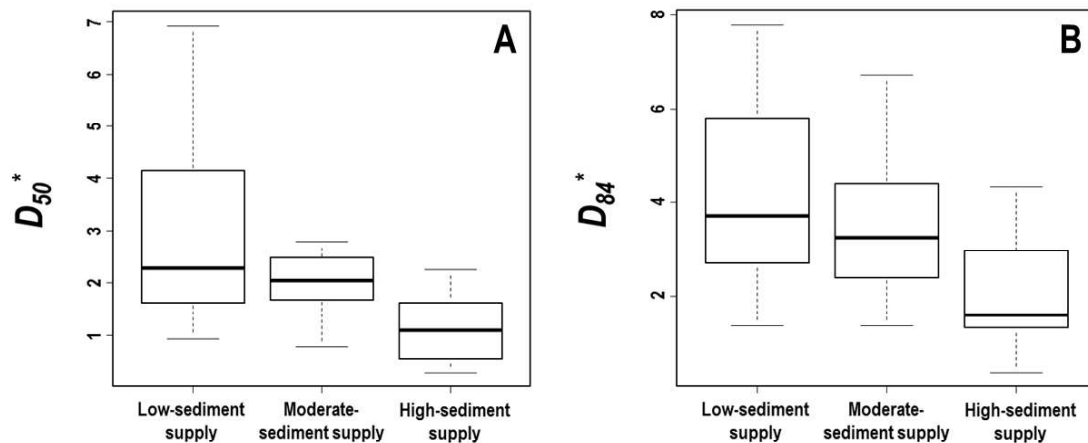


Figure 5. Differences in armour ratios according to sediment-supply conditions at the catchment scale. See main text. A: D_{50}^* . B: D_{84}^* .

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183 However, the compiled data illustrate how a certain degree of surface armouring ('armour ratios'
 184 >1) can also be observed in rivers fed by significant sediment supplies (Figure 5). Therefore, non-
 185 degrading gravel-bed streams with considerable sediment inputs may also exhibit some
 186 armouring. In truth, bedload transport models (Wilcock and DeTemple, 2005), tracer studies
 187 (Haschenburger and Wilcock, 2003), flume experiments (Hassan et al., 2006), and field
 188 observations (Andrews and Erdman, 1986; Clayton and Pitlick, 2008; Haschenburger, 2017)
 189 support the occurrence of such 'mobile' armours and suggest their persistence even during large
 190 floods. The sheltering of small particles into pockets and interstices between coarse grains can
 191 contribute to the development of these coarsened surfaces in well-supplied rivers (Parker et al.,
 192 1982; Parker and Klingeman, 1982; Andrews and Erman, 1986; Andrews and Parker, 1987; Berni
 193 et al., 2018). Surface coarsening represents in these cases an effective mechanism allowing for
 194 the transport of the coarse fractions in the gravel load at the same rate as those which are finer
 195 (Parker and Klingemann, 1982; Parker and Toro-Escobar, 2002; Parker, 2004). As such, it leads
 196 to a progressive equalization between the bedload and the subarmour GSDs (Parker and
 197 Klingemann, 1982; Marion and Fraccarollo, 1997; Mario et al., 2003; Mao et al., 2011; Powell et
 198 al., 2016; Bertin and Freidrich, 2018), a situation referred to as 'dynamic' or 'mobile armours'.
 199 Moreover, in gravel-bed rivers, the amount of entrained bed material and the grain size of the
 200 bedload increase as flow discharge rises (Milhous, 1973; Jones and Seitz, 1980; Kuhnle and

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201 Willis, 1992; Andrews, 1994; Wathen et al., 1995; Powell et al., 2001; Ryan and Emmett, 2002;
202 Wilcock and McArdell, 1993, 1997; Clayton and Pitlick, 2008; Pitlick et al., 2008; Recking,
203 2016). We could, thereby, expect differences in mobility between different grain-size classes with
204 fluctuations in flow discharge. Therefore, selective transport and horizontal winnowing of fines
205 during recession limbs and/or low flows may exhaust the fine sediment and favour the
206 development of a cover of coarse material preventing further sediment transport (Harrison, 1950;
207 Gessler, 1970; Little and Mayer, 1972; Proffitt and Sutherland, 1983; Chin et al., 1994). In this
208 regard, perennial and seasonal streams subjected to long-lasting floods and sustained receding
209 limbs might have more chances of suffering a coarsening of their bed surfaces and thus become
210 more armoured. In the opposite extreme, we could consider streams experiencing flashing-flood
211 hydrology. A flash flood involves a sudden increase of peak flows, bed shear stresses and equal
212 mobility for all the particle sizes represented in the streambed (Laronne and Reid, 1993; Laronne
213 et al. 1994; Reid and Laronne, 1995). In such cases, we should not expect large differences
214 between the surface, subsurface and bedload GSDs (Dietrich et al., 1989; Chin et al., 1994; Powell
215 et al., 2001; Parker, 2008; Venditti et al., 2017; Bertin and Friedrich, 2018). As a result, streams
216 receiving comparable amounts of sediment inputs may exhibit different armour ratios depending
217 on the dominant hydrological regime (Hassan et al., 2006).
218 Unfortunately, the compiled data are mostly from mountain streams dominated by snowmelt flow
219 regimes, probably because planning field-campaigns in order to measure bedload is easier in this
220 kind of rivers compared to rainfall-dominated settings, where it is more difficult to know in
221 advance when a channel-forming flow is going to occur (also to monitor within a fairly deep flow
222 channel with long rising and falling limbs). Thus, it is not easy to explore in depth the influence
223 of hydrological regime with the available data. In any case, snowmelt-dominated rivers, usually
224 submitted to sustained floods during the melting season, tend to show larger armour ratios than
225 rainfall or flash-flood dominated streams (see dotplot shown in Figure 6). Furthermore, the
226 percentage of fine sediment also tends to be lower in these snowmelt-dominated streams (Figure
227 6B), indicating fine sediment depletion during the long and gradually declining limbs of typical
228 snowmelt hydrographs. On the other hand, glacial-fed rivers, which normally have high-sediment

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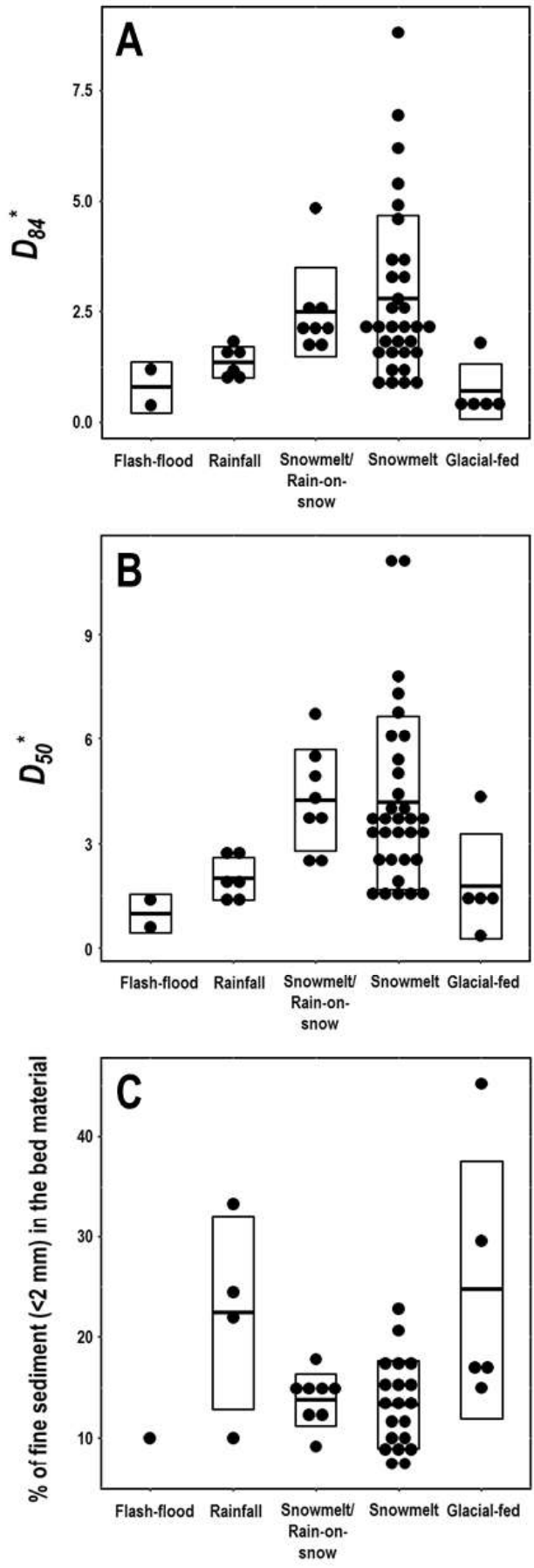


Figure 6. Differences in armour ratios (A and B) and percentages of fine sediment (C) according to the dominant flow regime.

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1039 229 supplies from glacial/periglacial sources (particularly the 'glacial flour' transported from
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1041 230 glaciated upland basins), show relatively low armour ratios (Figure 6B). The influx of fine
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1043 231 sediments represented by the glacial flour may help the coarse grains composing the armour layer
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1045 232 to become more mobile (Cui et al., 2003; Venditti et al. 2010a, b; Yager et al., 2015). In addition,
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1047 233 the flow regime of glacial-fed rivers is often characterised by daily peaks of discharge, rising and
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1049 234 falling successively, and these somewhat shorter, rapidly changing hydrographs are less prone to
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1051 235 promote armouring (Hassan et al., 2006). Hence, in conjunction, both issues may determine the
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1053 236 relatively low armour ratios documented here in the case of glacial-fed rivers.

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1055 237 To summarize, our review of field data shows how sediment supply, together with the shape of
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1057 238 the dominant flood hydrographs, condition surface armouring in gravel-bed rivers. The
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1059 239 hydrological regime controls the rate at which bedload and fine sediment are winnowed and
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1061 240 exhausted during the course of transport events. For instance, we can expect a less armoured
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1063 241 surface if a continuous replenishment of sediment exists during the transport episode, as we can
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1065 242 reasonably expect for glacial fed streams. In the opposite extreme, we can anticipate a better
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1067 243 developed armouring and a streambed surface more depleted in fines in cases where the bedload
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1069 244 is exhausted, which is likely when streams are subjected to long-lasting snowmelt hydrographs.
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1071 245 This is in accordance with the set of flume experiments by Hassan et al. (2006) who investigated
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1073 246 the influence of different hydrographs on surface armouring and observed varying textural
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1075 247 responses to steady vs. gradually varying flows. Nevertheless, fine sediment depletion during
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1077 248 floods may take more or less time depending on the amount of sediment supplied from upstream
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1079 249 sources. Therefore, the rate at which bedload is supplied into the channel may modulate the effects
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1081 250 of flow hydrographs and strongly determine the rate and degree of fine sediment depletion, and
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1083 251 consequently the degree of armour development.

1084 252 ***3.3. Surface armouring and streambed mobility***

1086 253 As stated above, surface coarsening is associated with fine sediment exhaustion from the
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1088 254 streambed and the development of clast arrangements and imbrications (Church et al., 1998;
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1090 255 Venditti et al., 2017). The increase of particle stability linked to surface structuration may involve
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1092 256 a decrease in the frequency of clast mobility and streambed disorganization (Church et al., 1998;
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 1098 Hassan and Church, 2002). To further explore this issue, we introduce a new metric: the ‘transport
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 1100 stage’ ratio (τ^*/τ_c^*). This is defined as the ratio between the peak basal shear stress for a given
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 1102 259 flow discharge and its critical value for incipient sediment motion. Critical stresses for sediment
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 1104 260 entrainment were estimated here based on Recking’s (2013) fit for D_{84} (see supplementary files
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 1106 261 for more information about how we computed transport stage ratios).

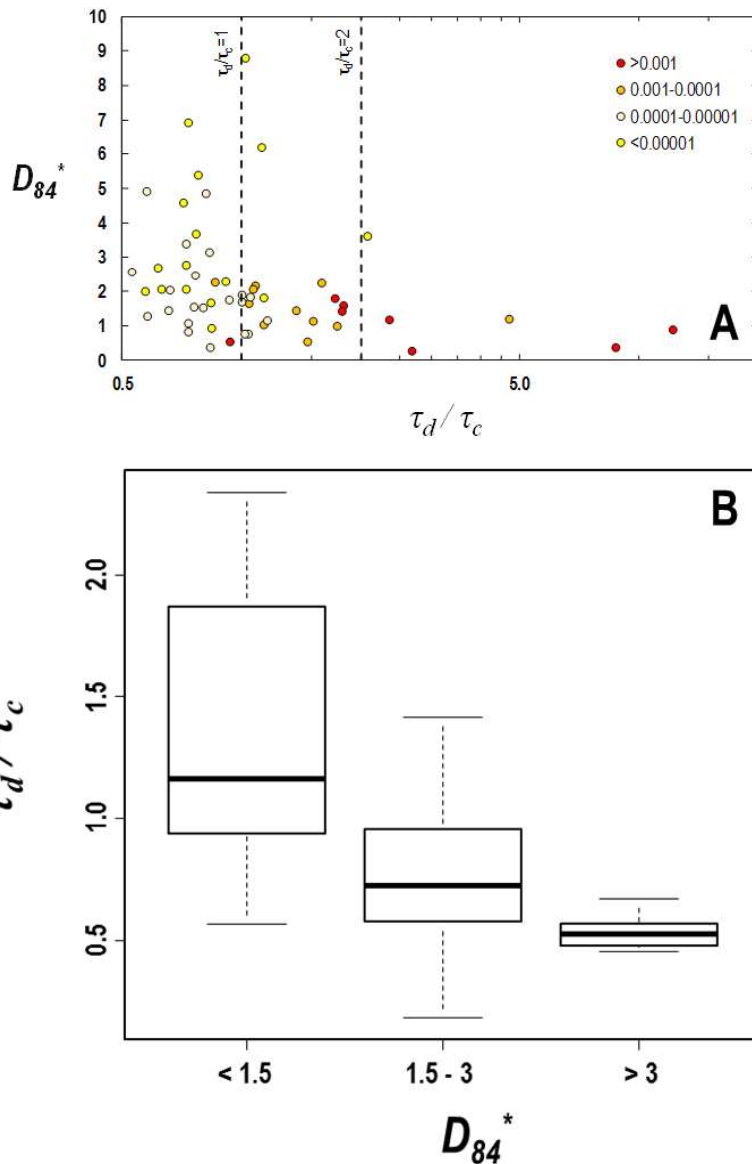


Figure 7. A: Armour ratio (D_{84}^*) plotted versus the transport stage (τ_d/τ_c). B: Transport stages values at the dominant discharge in poorly armoured ($D_{84}^* < 1.5$), normally armoured ($1.5 < D_{84}^* < 3$) and well armoured ($D_{84}^* > 3$) rivers. Critical threshold Shields values (τ_c) were computed using Recking, 2009).

1143 262 Values of this metric, estimated at the dominant channel-forming discharge, are close to the
 1144
 1145 263 critical thresholds for entrainment of the coarser fraction of bed sediment in well-armoured
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 1147 264 streams (Figure 7A). However, in poorly armoured streams, Shields stresses are well above these
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 1149 265 critical values. These poorly armoured streams also show larger bedload fluxes (see the values
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 1151 266 for transport intensities in Fig. 7A). A turning point can be defined close to an armour ratio of ~2:

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1157 267 there are no channels with large transport intensities ($\phi > 0.0001$) and high transport stages at
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1159 268 dominant discharge ($\tau^*/\tau_c^* > 2$) for armour ratios larger than 2. These observations confirm that
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1161 269 overall streambed mobility would be larger in poorly armoured streams, and document how well-
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1163 270 armoured streambed surfaces are seldom moved. Considering all of the above, the compiled data
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1165 271 support the idea of surface armouring being conditioned by dominant sediment supply conditions
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1167 272 (section 2.2), with ‘static-armours’ characterised by larger armour ratios than ‘mobile-armours’
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1169 273 (Figure 7B).

1171 274 **3.4. Quantifying sediment supply conditions in gravel-bed rivers**

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1173 275 Our review of the literature clearly shows how the GSD of a streambed’s surface, at a given river
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1175 276 reach, adjusts to the volume of sediment inputs and the ability of the channel to export the
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1177 277 available sediment inputs (Figures 6 and 7), an idea already outlined by previous field and flume
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1179 278 studies (Dietrich et al., 1989; Venditti et al., 2017). Introducing metrics allowing the quantitative
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1181 279 description of the sediment supply is then of interest as shown by Dietrich et al. (1989) who
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1183 280 proposed their own metric (q^*) (Dietrich et al., 1989; Montgomery and Buffington, 1997; Venditti
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1185 281 et al., 2017):

$$1186 \quad 282 \quad q^* = \frac{q_{s_s}}{q_{s_{ss}}} \quad \text{Eq. 2}$$

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1189 283 where q_s refers to the bedload transport rate per unit width and the subscripts s and ss to the surface
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1191 284 and subsurface sediment, respectively. The ‘bedload supply’ index as defined in eq. 2 is the ratio
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1193 285 between the sediment transport capacity to mobilize the surface GSD and the capacity to recruit
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1195 286 sediment from a hypothetical surface as fine as the subsurface. When actual bedload rates match
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1197 287 the bedload transport capacity of the channel, q^* should equal 1. Conversely, if sediment inputs
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1199 288 into the channel are lower than transport capacities, then $q^* < 1$. Here we computed the q^* ratio,
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1201 289 based on flow characteristics at the representative channel-forming discharge and on Recking’s
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1203 290 equation (details on q^* calculation are explained in the supplementary information, Subsection
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1205 291 S.3). The obtained values of q^* are in good agreement with our initial classification of compiled
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1207 292 data as high-, moderate- and low-sediment supplied channels (Figure 8A).

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 1216 293 However, the ‘bedload supply’ index q^* has the problem of estimating bedload transport capacity
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 1218 294 of a channel for a hypothetical situation where the streambed’s surface GSD equals its subsurface
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 1220 295 GSD (q_{ss} , denominator in eq. 2). In the field cases compiled here, this required the application of
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 1222 296 a bedload equation for the estimation of q_{ss} , with all the uncertainties linked to choosing the right
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 1224 297 formula. For that reason, we propose a simpler metric to quantify the ratio between the flux of
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 1226 298 sediment entering a given river reach, and the capacity of the channel to transport that sediment
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 1228 299 downstream: the ratio between bedload fluxes (q_s) and stream power (ω) at the representative or
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 1230 300 dominant channel-forming discharge. This metric quantifies the mass of sediment carried by a
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 1232 301 river with a given amount of hydraulic power. Values of this ratio are in good agreement with our
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 1234 302 initial classification of the compiled data as high-, moderate- and low-sediment supply channels
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 1236 303 (Figure 8B), confirming its usefulness for characterizing the dominant sediment supply conditions
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 1238 304 in gravel-bed rivers.

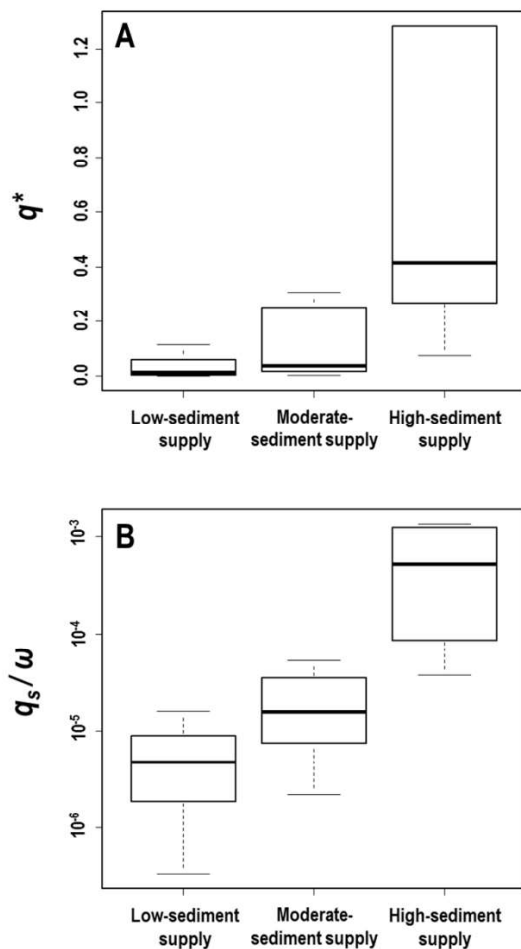


Figure 8. Differences in sediment-supply metrics (q^* and q_s/ω) according to our initial classification of data as low-, moderate- or high-sediment supply streams (see the main text). A: sediment supply metrics proposed by Dietrich et al. (1989). B: our own sediment supply metrics (see main text).

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1274
1275 305 **4. Meta-analysis on the data: linking streambed texture, sediment supply and hydrological**
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1277 306 **regime**

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1279 307 **4.1. Surface armouring driven by sediment supply**
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1281 308 Based on their own flume experiments, Dietrich et al. (1989) already observed that D^* increases
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1283 309 while sediment feed decreases compared to the overall bedload transport capacity of the channel.
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1285 310 Using Meyer-Peter and Muller's (1948) bedload transport equation, these authors described the
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1287 311 links between surface armouring (D_{50s}/D_{50ss}) and sediment supply (q^*) as:

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$$q^* = \left(\frac{\frac{\tau}{\tau_{css}} - D_{50}^*}{\frac{\tau}{\tau_{css}} - 1} \right)^{1.5} \quad \text{Eq. 3}$$

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1293 313 where τ is the basal shear stress and τ_{css} is the critical shear stress for the inception of motion in
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1295 314 the subsurface sediment. At the time when Dietrich et al. (1989) published their work, using
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1297 315 Meyer-Peter and Muller's (1948) bedload equation made sense due to its simplicity and its
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1299 316 common use in sediment transport studies. However, more sophisticated bedload transport
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1301 317 formulas have been proposed since then (e.g. Wilcock and Crowe, 2003; Recking, 2013), which
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1303 318 proved to perform well for 1D computations of bedload in natural gravel-bed streams (e.g. Parker,
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1305 319 2004; Recking, 2010; Camenen et al., 2011; Vázquez-Tarrío and Menéndez-Duarte, 2015; Hinton
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1307 320 et al., 2018). In this paper, based on Recking's (2013) (after Recking et al., 2016) bedload
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1309 321 equation, the following expressions are obtained (see details in the supplementary information):

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$$D_{84}^* = \frac{1}{f} \cdot q^{*-1/5} \quad \text{for partial-mobility conditions} \quad \text{Eq. 4}$$

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1313 323
$$D_{84}^* = q^{*-1} \quad \text{for full-mobility conditions} \quad \text{Eq. 5}$$

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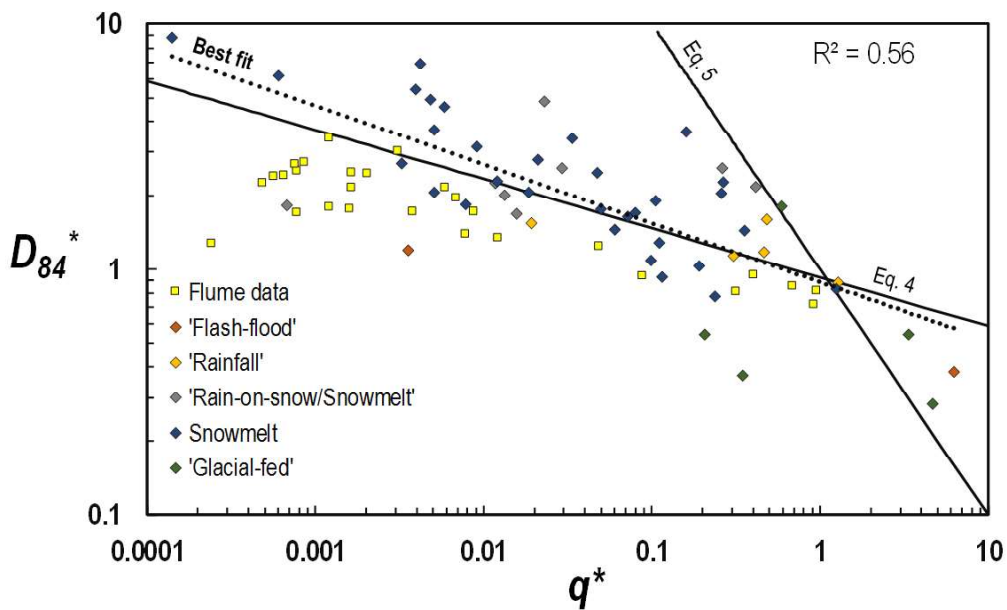
1315 324 where f in eq. 4 is a parameter relating D_{50}^* to D_{84}^* (~ 1.4 , according to Figure 2). Eqs. 4 and 5
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1317 325 provide a simple mathematical description of how D^* should correlate and change with q^* . As
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1319 326 opposed to eq. 3, eqs. 4 and 5 use D_{84}^* rather than D_{50}^* , because Recking's (2013) equation works
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1321 327 with the 84-th percentile of the surface GSD as the reference bed sediment size.

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1323 328 In Figure 9 we have plotted D_{84}^* against q^* for the field data, the latter estimated from flow
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1325 329 characteristics at the dominant channel-forming discharge and Recking's equation (details on q^*
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1327 330 calculation are explained in the supplementary information). We have also incorporated into this
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1332 plot data from flume experiments (extracted from Venditti et al., 2017; see supplementary
 1333 information), which tend to overlap the field data: armour ratios change with bedload supply
 1334 following a similar trend in flume and field data. Additionally, the obtained fit is very close to the
 1335 curve defined by eq. 4 (low-transport stage conditions). This suggests that partial mobility
 1336 conditions may dominate bedload motion at the dominant discharge in well-armoured streams.
 1337 We also plotted armour ratios against q_s/ω ratios (Figure 10A), and the data show a moderate but
 1338 statistically significant correlation between the q_s/ω ratio and the armour ratios. As shown above
 1339 (Figure 4), D_s is partially controlled by D_{ss} . This covariation may be introducing some noise in
 1340 the plot shown in Figure 10A. In order to remove this effect, we propose a new version of the
 1341 armour ratio (D^{**}):

$$D^{**} = \frac{D_s}{D_{ss}^\beta} \quad \text{Eq. 6}$$

1342 where β is the parameter defining how surface grain-size scales to subsurface grain-size.
 1343 According to Figure 4, β is ~ 0.62 and ~ 0.41 for D_{84} and D_{50} , respectively. This ‘corrected’ version
 1344 of the armour ratio shows a stronger correlation to the q_s/ω ratio (Figure 10B). Similar to Figure
 1345 9, Figure 10 shows how streambed surface tends to coarsen with decreasing sediment supplies
 1346 (proxied here by the q_s/ω ratios).



1385 **Figure 9.** Armour ratio plotted against the ‘bedload supply ratio’ q^* (see main text for more
 1386 details).
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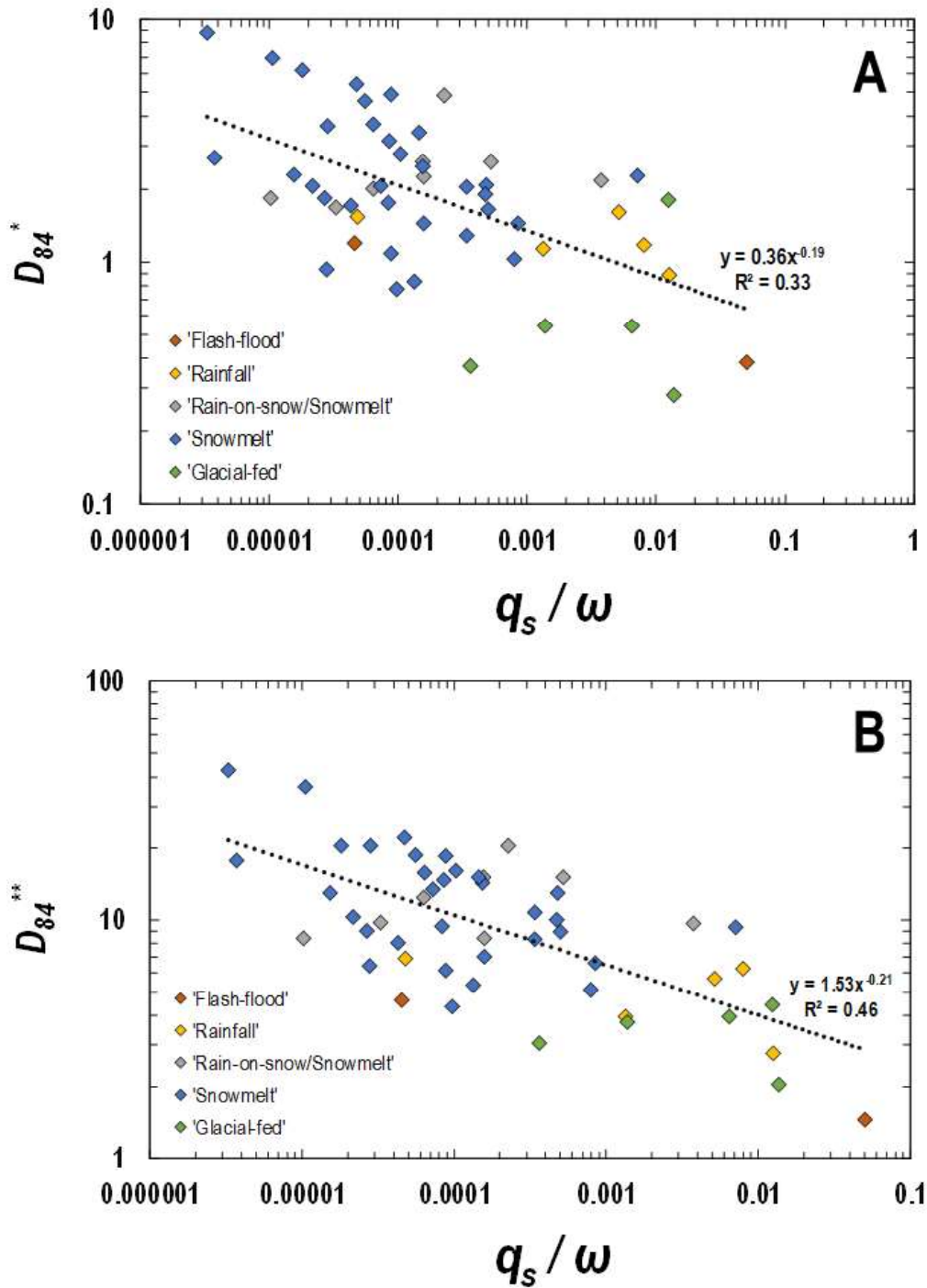


Figure 10. Armour ratio (A) and corrected armour ratio (B) plotted versus the ratio between bedload fluxes and streampower (q_s/ω_d).

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348 Although the number of plots is not well balanced across the different groups of data, it is worth
 349 pointing out that in Figures 9 and 10 rainfall data tend to plot close to the regression line, while
 350 snowmelt data tend to show a larger scatter and broader range of armour ratios. Snowmelt systems
 351 are typically located in the upper parts of the catchment compared to rainfall streams, which are

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1451
1452 352 usually located downstream, further from the coarser sediment sources. Consequently, this
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1454 353 implies that rainfall streams will be somewhat less prone to develop larger armour ratios even if
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1456 354 their sediment supply is limited. In this regard, large armour ratios may have some dependence
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1458 355 on the upstream/downstream location of the stream and its particular geological context, as
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1460 356 already suggested by Pitlick et al. (2008).

1462 357 ***4.2. Hydrological controls on surface armouring***

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1464 358 Figures 9 and 10 clearly show how streambed surface evolves and becomes progressively
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1466 359 armoured with declining sediment supplies. Nevertheless, there is a certain amount of scatter in
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1468 360 both plots, which highlights that some variability in armour ratios may exist with a given sediment
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1470 361 supply. Indeed, data from snowmelt streams tend to plot in the upper envelope of the point cloud
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1472 362 in Figure 9 and, conversely, data from glacial-fed and flash-flood dominated streams tend to fall
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1474 363 in the lower envelope, while flume data project in the middle of the point cloud, together with
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1476 364 rainfall dominated rivers. Similarly, in Figure 6 snowmelt dominated streams tend to show larger
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1478 365 armour ratios and smaller q_s/ω ratios than rainfall, flash flood and glacial-fed dominated streams.
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1480 366 All these observations highlight the variability in surface armouring and sediment supply
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1482 367 conditions and their interplay with different hydrological regimes.

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1484 368 However, we should also consider that adjustments of the streambed to shifts in sediment supply
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1486 369 conditions are not always synchronous and there is potentially a lag between the two. For instance,
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1488 370 surface coarsening slowly propagates downstream after widespread land cover changes in the
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1490 371 upland basin or downstream from a dam (Rollet et al., 2013), so it can take several years before
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1492 372 we observe riverbed coarsening following a significant upstream perturbation. Furthermore, in
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1494 373 snowmelt systems, some hysteresis in bedload supplies can be observed over the hydrological
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1496 374 year (Moog and Whiting, 1998); for instance, there can be significant bedload transport for a
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1498 375 given discharge in early spring, while bedload rates are reduced for similar flow discharges
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1500 376 occurring at the end of the flood season (Misset et al., in press). We have not controlled for all
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1502 377 these variables, so they probably introduce additional scatter in the relations between surface
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1504 378 armouring and sediment supply reported here.

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1511 379 **4.3. Quantifying the variability in armour ratios**
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1513 380 Based on the analysis performed in Subsections 4.1 and 4.2, we can conclude that surface
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1515 381 coarsening accommodates to the dominant sediment supply conditions, but the rate at which
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1517 382 bedload is emptied during the course of transport episodes may modulate the streambed
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1519 383 adjustment to sediment supply. This may complicate the use of armour ratio for river diagnosis,
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1521 384 insofar as the dominant sediment supply conditions are not the only drivers of surface grain-size
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1523 385 in gravel-bed rivers. The rate at which fine bedload exhausts during floods, the seasonal
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1525 386 variability in sediment availability and the time lag before the riverbed starts to respond after a
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1527 387 significant shift in the upstream bed material supplies also have a non-negligible impact on the
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1529 388 armour ratios documented at a given river reach. All these constraints may be largely dependent
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1531 389 on site-specific controls, such as the cross-sectional extent of armour break-up, the shape and
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1533 390 duration of the flow hydrograph, and the local availability of loose and fine sediment coming
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1535 391 from upstream reaches and bank toes.

1536 392 In this regard, Figures 9 and 10 not only illustrate how surface coarsening scales with sediment
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1538 393 supply, but also show the range of potential variability in armour ratios for a given sediment
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1540 394 supply and observed within the compiled field data. That said, we could report rivers more or less
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1542 395 armoured than the streambed conditions expected for an ‘average behaviour’ with a given bedload
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1544 396 input, reminiscent of the old distinction between ‘static’ versus ‘mobile’ armours (Jain, 1990;
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1546 397 Parker and Sutherland, 1990), so we refer to more (or less) mobile armour layers than the average
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1548 398 mobility expected for a given sediment supply condition (defined by the best fit in Figures 9 and
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1550 399 10). To handle this variability, we performed a quantile regression analysis and plotted the
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1552 400 regression between the different percentiles of the D^* distribution and the sediment supply metrics
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1554 401 (q^* and q_s/ω), obtaining two diagrams that illustrate the likelihood of a given value of D^* (Figure
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1556 402 11A) or D^{**} (Figure 11B) with a given sediment supply. Fluctuations from the 1-st to the 99-th
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1558 403 percentiles are linked to a set of factors, which include differences in the hydrological regime,
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1560 404 seasonal variability and time passed after significant shifts in upstream sediment inputs. Thus,
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1562 405 Figures 11A and 11B can be easily combined with some available qualitative information on a
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1570 406 specific study site in order to have an idea of the most probable values for the armour ratio with
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1572 407 a given sediment supply and hydrological regime.
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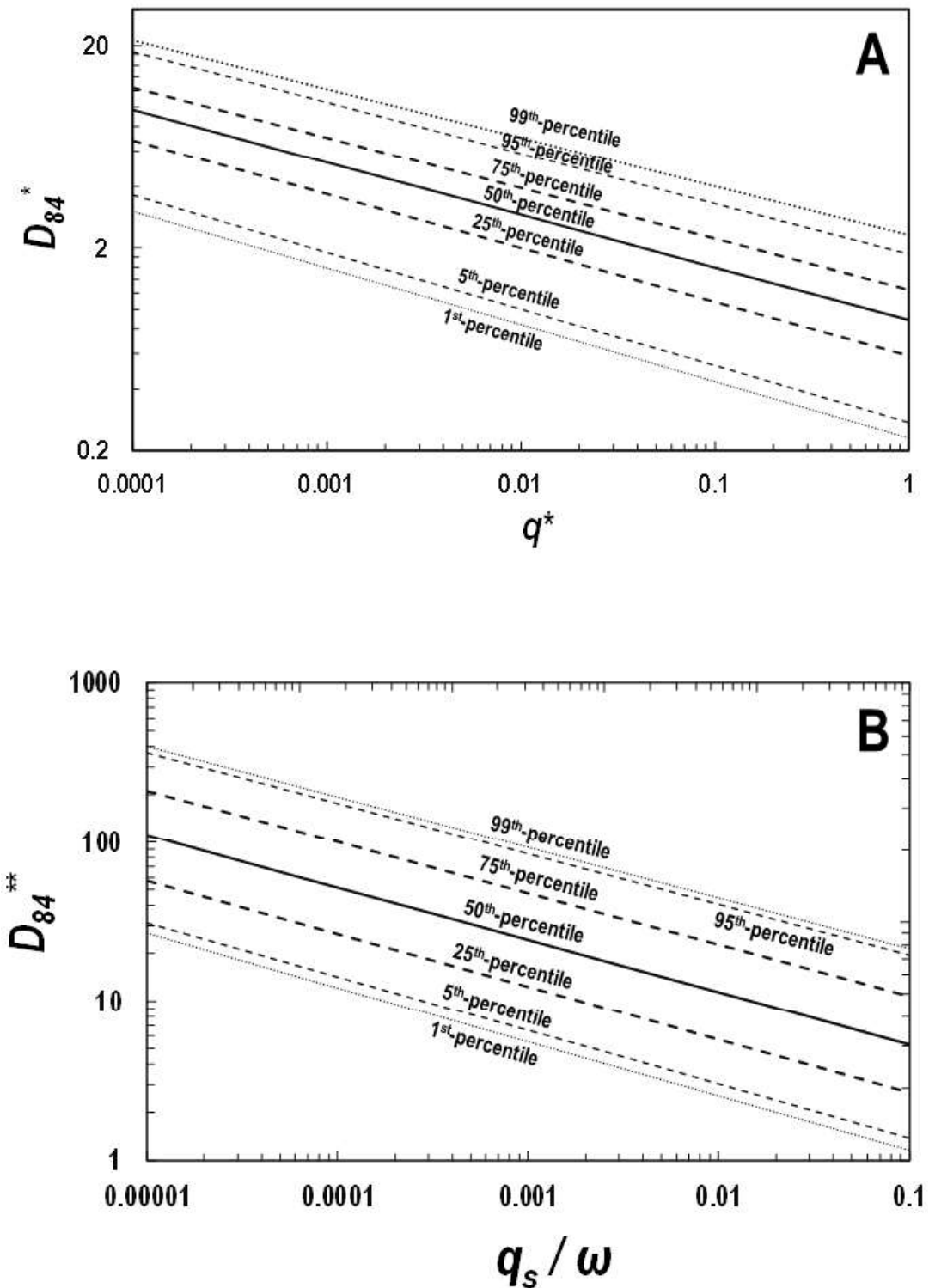


Figure 11. Diagrams issued from quantile regression analysis of the compile data, defining the likelihood of a certain value of the armour ratio according to sediment supply metrics (see main text for more details).

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1629 408 **4.4. Studying links between bedload transport, channel morphology and surface armouring**
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1631 409 Streambed surfaces adjust to available bedload yields (Figures 9 and 10). Reciprocally, surface
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1633 410 coarsening controls the frequency and magnitude of armour breakup and streambed
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1635 411 disorganization during floods (Figure 7). Strongly armoured streambeds are more difficult to
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1637 412 disorganize and more infrequently moved. Therefore, we could expect a lower degree of bed
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1639 413 material recruitment into bedload during regular flows. With this in mind, we expected surface
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1641 414 armouring to have some control over bedload rates. In order to explore this issue, we took the
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1643 415 available information on bedload discharges (Table 1) and performed a multiple regression
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1645 416 analysis linking bedload rates to flow magnitude and armour ratios:

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1647 417
$$\phi = \left(Q/Q_d\right)^a \cdot (D_{84}^*)^b \quad \text{Eq. 7}$$

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1649 418 where ϕ is bedload transport intensity (see supplementary information), and Q/Q_d is the ratio
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1651 419 between the peak discharge for the transport episode and the dominant discharge (hereinafter
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1653 420 called flow ratio). However, the wide diversity of channel morphologies represented amongst the
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1655 421 compiled data should be highlighted, since channel macroforms control flow and sediment
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1657 422 transport patterns, which in turn may potentially affect bedload rates (Ferguson, 2003; Francalanci
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1659 423 et al., 2012; Recking et al., 2016; Vázquez-Tarrío et al., 2018; Vázquez-Tarrío and Batalla, 2019).
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1661 424 For this reason, we introduced a set of three binary indicators into the regression analysis to
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1663 425 incorporate into it the influence of channel morphology:

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1665 426
$$\phi = \left(Q/Q_d\right)^a \cdot (D_{84}^*)^b \cdot e^{c \cdot RP} \cdot e^{d \cdot SP} \cdot e^{d \cdot BR} \quad \text{Eq. 8}$$

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1668 427 where RP, SP and BR are the dummy variables taking a value of 1 in the cases of riffle and pool,
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1670 428 step-pool and braided channels, respectively. In the case of plane-bed channels, the three dummy
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1672 429 variables would be 0.

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1674 430 To test whether eqs. 7 and 8 fit the available bedload information, we used ordinary multiple
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1676 431 regression in linearized form and stepwise procedures, after log transforming all the variables
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1678 432 included in the equation. Both equations explain the variance in the compiled bedload data to a
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1680 433 statistically significant degree, with all the variables included in the regression model being
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1682 434 significant (Table 5). The R^2 values imply that eq. 8 is more robust than eq. 7 ($R^2 = 0.60$ vs. 0.34,

435 respectively), which outlines the existence of some morphological imprint on bedload transport
 436 rates. According to the regression model, transport intensities increase with the 1.7 positive power
 437 of flow ratios. Conversely, bedload rates are negatively correlated to armour ratios, i.e. bedload
 438 rates tend to be considerably weaker in well-armoured channels. With the aim of assessing the
 439 relative importance of each of the independent variables incorporated into the regression model,
 440 we used the method proposed by Lindeman et al. (1980), often recommended for assigning shares
 441 of relative weight of predictors to the R^2 . According to this analysis, channel morphology is the
 442 variable that explains the largest amount of variability in bedload data ($\sim 53\%$ of R^2), followed
 443 by flow magnitude (23% of R^2) and armour ratios (21% of R^2). Differences in dominant channel
 444 morphology have a strong effect on the variability in bedload rates observed between the
 445 compiled data, but differences in surface armouring also have a large impact.

Variable	Coefficient	Standard Error	t	p-value	VIF ¹
Intercept	2.991×10^{-5}	0.077	-135.680	0.000	
Q/Q_d	1.671	0.035	45.156	0.000	1.099
D_{84}^*	-2.371	0.063	-37.520	0.000	1.565
BR	0.809	0.264	3.067	0.002	1.257
RP	4.482	0.077	57.921	0.000	1.155
SP	1.557	0.125	12.453	0.000	1.273

Residual standard error: 2.579 on 5470 degrees of freedom

Multiple $R^2 = 0.60$; Adjusted $R^2 = 0.60$

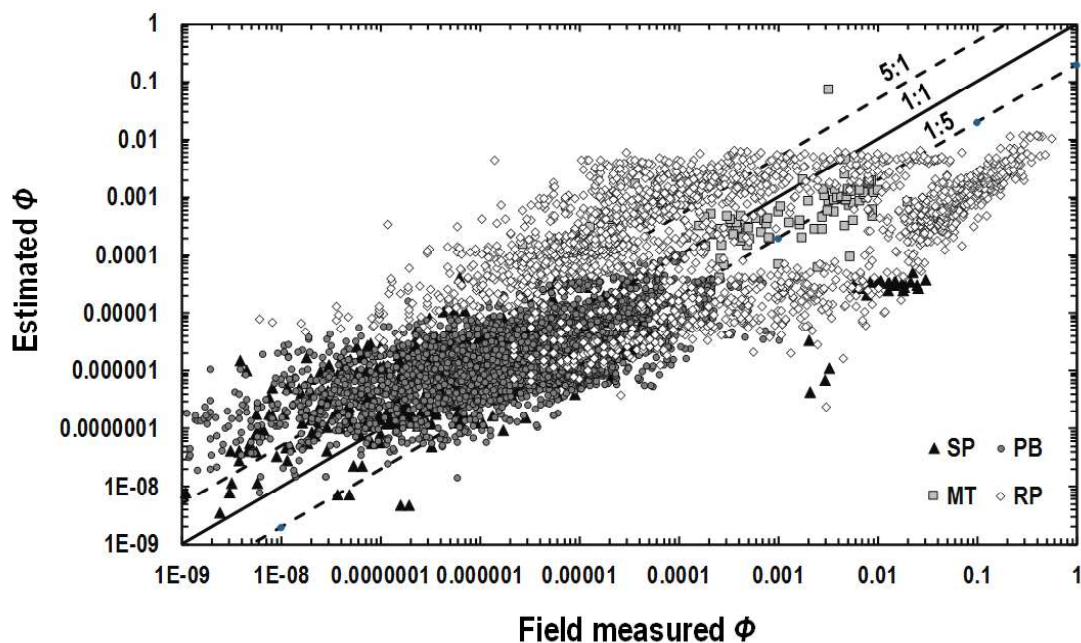
F-statistic: 1611 on 5 and 5470 degrees of freedom. $p\text{-value} = 2.2 \times 10^{-16}$

¹ Variance Inflation Factor

Table 5. Results from the multiple regression model based on eq. 8 (see main text).

449 Different channel morphologies and macro-bedforms result from the adjustment of channel
 450 macro-roughness to different balances between sediment supply and transport capacity
 451 (Montgomery and Buffington, 1997); similarly, armouring is the consequence of the adjustments
 452 in roughness at the grain-scale to the same balances between sediment supply and capacity
 453 (Dietrich et al., 1989; Venditti et al., 2017). Consequently, both dominant channel morphology
 454 and armour ratios constitute a proxy of sediment supply and should be related. Interestingly
 455 enough, we have observed systematic differences in armour ratios between the different channel
 456 morphologies (see Figure 3). In general, multi-thread channels and rivers with riffles and pools

1745
 1746
 1747 457 exhibit a lower degree of surface armouring, while plane-bed and step-pool channels show higher
 1748
 1749 458 armour ratios. Higher armour ratios in step-pool channels could be related to the existence of
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 1751 459 jammed stones and large immobile boulders that may contribute to larger armour ratios. In plane
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 1753 460 bed channels, large protruding grains and immobile stones defining transversal ribs and cells are
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 1755 461 also common (Church et al., 1998), and may help define a surface coarser than in riffle and pool
 1756
 1757 462 and braided channels. Following this, eq. 8 would provide a mathematical expression describing
 1758
 1759 463 how the increase in bedload rates with flow discharge is modulated by the differences in sediment
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 1761 464 supply, which is parameterised through the armour ratio and the dummy variables (i.e. dominant
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 1763 465 macro-bedforms). In Figure 12, we compare the transport intensities estimated using eq. 8 to the
 1764
 1765 466 true values measured in the field. Regression predictions plot close to the 1:1 line, with an average
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 1767 467 0.9 ratio between predictions and truth-values and 43% of points showing a discrepancy lower
 1768
 1769 468 than one order of magnitude. Riffle and pool and braided data show a larger scatter than plane-
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 1771 469 bed and step-pool morphologies; this is probably related to the larger cross-sectional variability
 1772
 1773 470 in shear stress and sediment paths, typical of these channel settings (Recking et al., 2016;
 1774
 1775 471 Vázquez-Tarri6 et al., 2018).



1796 **Figure 12.** Bedload transport intensities estimated using eq. 8 plotted versus transport
 1797 intensities measured in the field. MT: Multi-thread channels; PB: Plane-bed; RP: Riffle and
 1798 pool; SP: step-pool streams.
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1806 472 **5. Discussion**
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1808 473 **5.1. Sediment supply accommodated through armouring in gravel-bed rivers**
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1810 474 Our meta-analysis shows a decline in the armour ratio with increasing bedload supplies
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1812 475 (quantified through the q^* -metrics and q_s/ω_{dd}) in the compiled field data (Figures 7 and 8), which
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1814 476 is similar to the trends observed in previous flume experiments (Dietrich et al., 1989; Venditti et
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1816 477 al., 2017). In this regard, plots shown in Figures 7 and 8 quantify the important imprint of
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1818 478 sediment supply over surface armouring in gravel-bed rivers. It should be noted that q^* was
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1820 479 computed based on subsurface GSD, using the common approach of equalling subsurface GSD
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1822 480 to the average GSD of the bedload. However, bedload capacities would be more adequately
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1824 481 computed based on the actual GSD of the bedload. There may be situations where shifts in
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1826 482 sediment sources and catchment patterns of sediment production could lead to changes in the
1827
1828 483 grain size of the sediment supplies; for instance, mountain watersheds submitted to massive
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1830 484 afforestation (or forest clearance), or river reaches fed by largely managed (e.g. dammed,
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1832 485 dredged) tributaries. In such situations, we would expect changes in the GSD of the bedload
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1834 486 (fining or coarsening) that may derive in a change in bedload transport capacities; this would in
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1836 487 turn influence surface GSD together with changes in the amount of sediment inputs.

1837 488 There is a certain amount of unexplained scatter within the data in Figures 7 and 8, so there should
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1839 489 be other contributing and explanatory variables for data variance in them apart from bedload
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1841 490 supply. We believe that the variability introduced by differences in the shape of the dominant
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1843 491 competent flow hydrograph and flow duration (Hassan et al., 2006; Phillips et al., 2018) are the
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1845 492 most probable candidates. These flow parameters would be largely variable amongst the selected
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1847 493 study reaches and could not be described by a single q^* or q_s/ω ratio based only on peak discharge.
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1849 494 As pointed out above, gravel-bed streams submitted to long receding hydrograph limbs and
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1851 495 experiencing several floods per year are likely to develop a higher degree of armouring than
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1853 496 equivalent streams in terms of sediment supply, but experiencing less frequent flooding and a
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1855 497 short falling limb (Hassan et al., 2006). Therefore, flood hydrograph may exert an important
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1857 498 control on armouring, together with sediment supply. Although the available data do not let us
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1859 499 explore this issue in much depth, in general they are in agreement with this general picture:
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500 snowmelt-dominated data tend to show larger armour than rainfall-dominated rivers in Figure 6,
501 although the possibility of time lags between previous changes in upstream sediment inputs and
502 streambed adjustment, or even some seasonal variability in sediment supply, should not be
503 dismissed, since both may play a role in the variability documented in the compiled data. The
504 influence on armouring of these time lags and seasonal variability in sediment inputs may be more
505 important in those channel morphologies better coupled or connected with upstream sediment
506 sources (e.g. braided channels), where one-shot observations of armour ratios may not detect well
507 the average trends.

508 However, despite the scatter in the data, the overall picture shown by Figures 9 and 10 illustrates
509 how the signal linked to differences in sediment yields is the dominant one, and it is globally well-
510 recorded through the surface grain size and the armour ratio in gravel-bed rivers. Our analysis
511 suggests that sediment supply drives streambed surface GSD in gravel-bed rivers in two different
512 ways: i. the GSD of the sediment supplied and introduced into the channel network exerts an
513 obvious control, since it defines the range of grain sizes that are available in the bed material
514 (Figure 4); ii. streambed surface represents the sediment layer in more intimate contact with fluid
515 forces, so it adjusts to the balances between volumes of sediment supplied into the channel and
516 the streams' capacity to transfer these sediment inputs downstream out from the reach (Figures 9
517 and 10). In principle, stream ability to convey sediment inputs downstream is size-selective (Paola
518 and Seal, 1995), and this partially explains why the surface layer tends to coarsen. This influence
519 of hydraulic conditions could be eventually more important and partially blur the primary
520 influence of the GSD of the bedload.

521 Based on our data and quantile regression analysis, we have created two diagrams illustrating the
522 likelihood of a certain armour ratio for a given sediment supply (Figure 11). We believe these
523 diagrams provide a useful frame to interpret armour ratios measured in the field, and to
524 characterize streambed state and mobility in a quick way. For many applied issues (channel-
525 design, river restoration), an adequate diagnosis of the hydro-morphological status of a given
526 reach is fundamental. However, information of bedload fluxes is lacking in many cases. In such
527 cases, armour ratios measured in the field may provide a quick characterization of the sediment

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528 supply conditions and streambed mobility of the considered river reach, by combining diagrams
529 such as those shown in Figure 9 with some qualitative knowledge about the dominant
530 hydrological regime.

531 ***5.2. Implications for equilibrium channel geometry***

532 We observed how in most of the compiled data shear stresses at the dominant discharge are close
533 to critical thresholds for sediment motion (Figure 10). These observations reflect the streambed's
534 surface adjustment to dominant hydraulic conditions. In this respect, theoretical work (Parker,
535 1978; Parker, 1979) and decades of observations in gravel-bed alluvial channels (Parker, 2004;
536 Mueller and Pitlick, 2005; Parker et al., 2007) supported the hypothesis of threshold or near-
537 threshold channels (Parker et al., 2007; Phillips and Jerolmack, 2016; Métivier et al., 2017).
538 According to this idea, alluvial channel dimensions adjust in such a way that the threshold for the
539 motion of median-size grains occurs close to bankfull flows (e.g. MacKenzie et al., 2018).

540 However, in our data Shields values at the dominant discharge are close to the critical thresholds
541 for the entrainment of the coarser fraction of bed sediment only in well-armoured streams (Figure
542 10); otherwise, Shields stresses are well above these critical values in poorly armoured streams.
543 Pfeiffer et al. (2017) observed that only supply-limited streams meet the threshold channel
544 condition, while in capacity-limited systems they reported Shields stresses at bankfull
545 considerably over the critical levels. Based on these observations, these authors concluded that
546 the common observation of channels adjusting to threshold conditions simply reflects the fact that
547 the channels most commonly surveyed in the field are subject to modest sediment supplies.

548 Nevertheless, we propose an alternative interpretation for the trends shown in Figure 10 and
549 Pfeiffer et al. (2017). According to the data analysed in this manuscript, high-sediment supply
550 systems correspond mainly to multi-thread and riffle-pool streams (Figure 3). These kinds of
551 settings are characterized by bar morphologies, which involve complex 3D flow structures
552 (Francalanci et al., 2012), specific patterns of grain size sorting (Paola, 1989; Lisle et al., 1992)
553 and a great variability in shear stress distributions across their cross-section (Recking, 2009;
554 Recking et al., 2016). Furthermore, gravel dunes and bedload sheet propagation (Venditti et al.,
555 2017) represent an additional source of form roughness in these rivers. Consequently, we expect

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556 larger biases in bedload estimates when averaging shear stresses across the cross section of these
557 streams (Recking, 2013). Therefore, it is possible that 1D averaged shear stresses are not an
558 adequate proxy for the actual variability in shear stresses acting at the local and grain scales in
559 riffle and pool and multi-thread settings. This may explain the observed deviation from near-
560 threshold conditions, which could be to a large degree an artefact resulting from the assumptions
561 made when estimating cross-section averaged shear stresses (Yager et al., 2018).

562 ***5.3. Armour ratio: a proxy for the channel planform configuration***

563 Our review outlined differences in armour ratios according to channel morphology (Figure 3),
564 with multi-thread and riffle and pool rivers exhibiting, in general, a lower degree of surface
565 armouring than plane-bed and step-pool channels. The scale of the roughness and protruding
566 elements in step-pool channels is in general larger, with channel-spanning ribs (steps) composed
567 by an accumulation of jammed cobbles and boulders transverse or oblique to the channel
568 (Zimmermann and Church, 2001; Chin and Wohl, 2005), alternating with pools (Church and
569 Zimmermann, 2007; Lamarre and Roy, 2008). Step-pool streams are normally close to headwater
570 areas and are largely dependent on the accumulation of relatively fine colluvium inputs into pools
571 and around protruding boulders (Turowski et al., 2011; Recking, 2012; Recking et al., 2012; Piton
572 and Recking, 2017). Accordingly, protruding cobbles and boulders can be found on the bed
573 surface of step-pool channels, even with large sediment supplies and thick alluvial covers. This
574 may explain the generally larger armour ratios observed for step-pool streams. Similarly, the data
575 for braided rivers tend to show lower armour ratios (Figure 3), which results from the large
576 sediment inputs typical of braided streams.

577 The trends in armour ratios with channel morphology documented here suggest that planform
578 style, dominant macrobedforms, streambed mobility and grain size adjust together to dominant
579 sediment supply conditions. In this regard, a typical sequence of reach types observed in many
580 mountain basins corresponds to a downstream progression from step pool at headwaters, plane-
581 bed (or forced pool-riffle) to pool-riffle and/or multithreaded channel morphologies at the
582 piedmont valley (e.g. Warburton, 2007). In their seminal paper, Montgomery and Buffington
583 (1997) suggested that this kind of longitudinal sequence describes opposing trends between

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2042 584 sediment supply and transport capacities in the downstream direction. In this regard, Pitlick et al.
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2044 585 (2008) documented a downstream trend to lower armour ratios in rivers from Colorado and Utah.
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2046 586 The trends observed here are in good agreement with this general scheme.
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2048 587 Additionally, gravel-bed rivers exhibit a wide diversity of bed features that are larger than
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2050 588 individual clasts and smaller than reach-scale patterns. Sediment supply plays a critical role in the
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2052 589 development of one kind of bed structuration or another (Buffington and Montgomery, 1999;
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2054 590 Venditti et al., 2017.). Dominant bedforms in gravel bed rivers evolve from gravel dunes (Carling,
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2056 591 1999) and bedload sheets in poorly armoured streams (Whiting et al., 1988; Nelson et al., 2009;
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2058 592 Recking et al., 2009), to a sequence of pebble clusters (Brayshaw, 1984), transverse ribs (Koster,
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2060 593 1978; Allen, 1984), stone lines (Laronne and Carson, 1976) and reticulate stone cells (Church et
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2062 594 al., 1998; Hassan and Church, 2000) with decreasing sediment supplies (Venditti et al., 2017).
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2064 595 Indeed, the stability of a gravel streambed is increased by the presence of particle arrangements
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2066 596 and clusters (e.g., Reid and Frostick, 1984; Church et al., 1998; Hassan and Church, 2000; Piedra
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2068 597 et al., 2012; Ockelford and Haynes, 2013; Heays et al., 2014). Therefore, streambed textures
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2070 598 adjust to sediment supply reductions not only through surface coarsening, but also through a
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2072 599 decrease in streambed mobility and a different bed surface organization. Venditti et al. (2017)
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2074 600 have proposed a phase diagram for bedforms in gravel bed rivers, relating q^* to the armour ratio
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2076 601 (Figures 16.5 and 16.9 in Venditti et al., 2017) that is in some way comparable to our Figure 7. If
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2078 602 armour ratio varies in parallel to channel morphology, dominant macroforms, bed sediment
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2080 603 mobility and streambed structures (as our analysis shows), then the armour ratio should be
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2082 604 considered not only as some kind of ‘textural imprint’ of sediment supply conditions, but also as
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2084 605 a proxy of the streambed’s organization and structuration (Venditti et al., 2017).
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2087 607 **6. Concluding Remarks**

2089 608 A large body of research in fluvial geomorphology has contributed to establishing the general
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2091 609 idea that sediment supply, bedload fluxes, channel morphology, bankfull shear stresses and
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2093 610 surface grain size are intimately related in gravel bed rivers. In this paper, we aimed to quantify
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2095 611 the existing links between sediment supply and surface coarsening. Based on the re-analysis of a
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612 large database of bedload discharge information for gravel-bed streams, we have proposed semi-
613 empirical relations describing how surface grain-size and armour ratios evolve with the balances
614 between bedload yields and channel sediment transport capacities. Armour ratios increase with
615 decreasing sediment inputs, as inferred from the dimensional analysis of bedload equations and
616 as already shown by previous flume experiments.

617 Accounting for armouring is important for many reasons, since it influences the local availability
618 of bedload, hydraulic roughness, bed permeability, and physical conditions for aquatic organisms.
619 We believe that the empirical relationships found here between bedload yields and armour ratios
620 have the potential to provide a quantitative frame for exploring the links between surface
621 armouring, hydraulics and sediment availability in specific gravel-bed reaches and are a step
622 towards predicting textural adjustments to changes in sediment supply (for example, downstream
623 dams). Additionally, since armouring can provide important qualitative information about the
624 bedload regime (e.g. partial mobility vs. full mobility, sediment supply) of a specific river reach,
625 our results may potentially provide some interesting clues for palaeohydrological analysis.

626

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Notations

D :	Grain-size (particle diameter)
D_{is} :	i -th percentile of the surface grain-size distribution
D_{iss} :	i -th percentile of the subsurface grain-size distribution
D_i^* :	Armour ratio, or the ratio between the i -th percentiles of the surface and subsurface grain-size distributions
D_i^{**} :	Corrected version of the armour ratio, i.e. armour ratio computed accounting for the inherent covariation between surface and subsurface grain sizes
f :	Ratio between the armour ratio estimated based on the 84-th percentiles (D_{84}) and the armour ratio based on the median size (D50) of the grain size distribution
GSD:	Grain size distribution
ϕ :	Bedload transport intensity (Einstein parameter)
q_s :	Bedload transport rate per unit width
q_{ss} :	Channel's bedload transport capacity (per unit width) to mobilize the subarmour sediment
q_{sss} :	Channel's bedload transport capacity (per unit width) to mobilize the subarmour sediment
q^* :	Bedload 'supply index', or the ratio between q_{ss} and q_{sss}
Q :	Peak discharge
Q_d :	Peak discharge for the dominant discharge
Q/Q_d :	Flow ratio
ω :	Specific streampower
τ :	Section-averaged bed shear stress
τ^* :	Dimensionless (section-averaged) Shields shear stress
τ_c :	Critical (section averaged) bed shear stress for sediment entrainment
τ_{cs} :	Critical (section averaged) bed shear stress for the inception of motion of the surface sediment particles
τ_{css} :	Critical (section averaged) bed shear stress for the inception of motion of the subsurface sediment particles
τ_{ci}^* :	Critical threshold Shields stress for entrainment particle with sizes corresponding to the i -th percentile of the grain-size distribution

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3045 **1087 Supplementary information to Vázquez-Tarrío et al. (*Textural signatures***
3046 **1088 *of sediment supply in gravel-bed rivers: re-visiting the armour ratio*)**

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3049 **1089 S. 1. Computing section-averaged bed shear stresses for the compiled data**

3050 1090 First, we used the available discharge information and Rickenmann and Recking's (2013) fit to
3051 1091 Ferguson's (2007) friction law to estimate average flow velocity (U):

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$$\frac{U}{\sqrt{g \cdot S \cdot D_{84}}} = 1.44 \cdot q_l^{*0.6} \cdot \left[1 + \left(\frac{q_l^*}{43.8} \right)^{0.82} \right]^{-0.24} \quad \text{Eq. S1}$$

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3059 1093 where g is the gravitational acceleration, S is the channel slope, D_{84} the 84th percentile of the
3060 1094 surface GSD and q_l^* the dimensionless specific discharge:

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$$q_l^* = \frac{ql}{\sqrt{g \cdot S \cdot D_{84}^3}} \quad \text{Eq. S2}$$

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3067 1096 where q_l is the flow discharge normalized by channel width. Then, we estimated the water depth
3068 1097 (d) from:

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$$d = \frac{ql}{U} \quad \text{Eq. S3}$$

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3075 1099 For the hydraulic radius (R), we assumed a rectangular cross-section:

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$$R = \frac{d \cdot w}{2 \cdot d + w} \quad \text{Eq. S4}$$

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3081 1101 where w is channel width. Finally, we computed the cross-section averaged basal shear stresses
3082 1102 from the hydraulic radius-slope product:

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$$\tau = \rho \cdot g \cdot R \cdot S \quad \text{Eq. S5}$$

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3089 **1105 S. 2. A simple model for armour ratios based on bedload formulae**

3090 1106 As explained in the main text, Dietrich et al. (1989) developed a metric (q^*) quantifying the
3091 1107 balances between the sediment supplied into a reach and its actual bedload transport capacities
3092 1108 (Dietrich et al., 1989; Montgomery and Buffington, 1997; Venditti et al., 2017):

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$$q^* = \frac{q_{ss}}{q_{ssS}} \quad \text{Eq. S6}$$

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 3104 1110 where q_s refers to the bedload transport rate per unit width and the subscripts s and ss to the surface
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 3106 1111 and subsurface sediment, respectively. Combining eq. 1 with the Meyer-Peter and Muller's
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 3108 1112 (1948) bedload transport equation, Dietrich et al. (1999) suggested:

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$$q^* = \left[\frac{(\tau - \tau_{cs})}{(\tau - \tau_{css})} \right]^{1.5} \quad \text{Eq. S7}$$

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 3114 1114 where τ is the bed shear stresses and τ_c the critical stress for incipient motion. Writing shear
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 3116 1115 stresses in their non-dimensional form (τ^*):

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$$\tau^* = \frac{\tau}{(\rho_s - \rho) \cdot g \cdot D} \quad \text{Eq. S8}$$

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 3123 1117 Dietrich et al. (1989) finally arrived at:

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$$q^* = \left(\frac{\frac{\tau}{\tau_{css}} \frac{D_{50s}}{D_{50ss}}}{\frac{\tau}{\tau_{css}} - 1} \right)^{1.5} \quad \text{Eq. S9}$$

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 3130 1119 This expression quantifies the links between surface armouring (D_{50s}/D_{50ss}) and sediment supply
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 3132 1120 (q^*), based on Meyer-Peter and Müller's bedload equation. The same reasoning could be applied
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 3134 1121 to any other bedload equation. In this paper, we decided to use Recking's (2013b) equation (after
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 3136 1122 Recking et al., 2016), since this equation has been tested and validated against bedload discharge
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 3138 1123 information from natural gravel-bed rivers (e.g. Hinton et al., 2018). Recking (2013b) proposed
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 3140 1124 the following expression for bedload rates:

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$$\phi = A \cdot \frac{\tau^{*\alpha}}{1 + \left(\frac{\tau_m^*}{\tau^*} \right)^\beta} \quad \text{Eq. S10}$$

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 3147 1126 where ϕ is the dimensionless transport rate, estimated using the Einstein parameter:

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$$\phi = \frac{q_s}{\sqrt{g \cdot (s-1) \cdot D_{84}^3}} \quad \text{Eq. S11}$$

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 3153 1128 Coefficients A , α and β are three model parameters, for which values of 14, 2.5 and 4 were
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 3155 1129 proposed, respectively (Recking, 2013b, Recking et al., 2016). τ_m^* is a mobility shear stress that

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 3163 1130 allows to determine low transport conditions (with partial transport and bed clustering) apart from
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 3165 1131 higher transport ones (Recking et al., 2016). Recking (2013b) estimates τ_m^* from:
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$$\tau_m = (5 \cdot S + 0.06) \cdot \left(\frac{D_{84}}{D_{50}}\right)^{4.4\sqrt{S}-1.5} \quad \text{Eq. S12}$$

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 3172 1133 Eq. S10 is such that for low transport conditions ($\tau_{ref}^* \ll \tau_m^*$):
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$$\phi = A \cdot \frac{\tau_m^{\alpha+\beta}}{\tau_m^\beta} \quad \text{Eq. S13}$$

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3178 1135 and for higher transport conditions ($\tau_{ref}^* \ll \tau_m^*$):
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 3181 1136
$$\phi = \tau_m^{\alpha} \quad \text{Eq. S14}$$

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3183 1137 We can now combine equation S6 with the expression for the Einstein parameter (eq. S11):
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$$q^* = \frac{q_{ss} - \phi_s}{q_{ss} - \phi_{ss}} \cdot \left(\frac{D_{84ss}}{D_{84s}}\right)^{3/2} \quad \text{Eq. S15}$$

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3190 1139 Incorporating now eq. S13 into eq. S15, we can deduce:
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$$q^* = \left(\frac{\tau_s^*}{\tau_{ss}^*}\right)^{\alpha+\beta} \cdot \left(\frac{\tau_{ms}^*}{\tau_{mss}^*}\right)^\beta \cdot \left(\frac{D_{84s}}{D_{84ss}}\right)^{3/2} \quad \text{Eq. S16}$$

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3197 1141 for low transport conditions. Taking into account the expression for τ_m^* (eq. S12) and assuming a
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 3199 1142 similar bed slope for the surface and subsurface case:
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$$\frac{\tau_{ms}^*}{\tau_{mss}^*} = \left(\frac{D_{84s}}{D_{84ss}}\right) \cdot \left(\frac{D_{50ss}}{D_{50s}}\right) \quad \text{Eq. S17}$$

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3205 1144 We could now introduce the idea that the armour ratio based on the D_{50} is linearly related to the
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 3207 1145 value based on the D_{84} (see Figure 2 in the main text of the paper):
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$$\left(\frac{D_{50s}}{D_{50ss}}\right) = f \cdot \left(\frac{D_{84s}}{D_{84ss}}\right) \quad \text{Eq. S18}$$

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3213 1147 Incorporating this idea into eq. S17, then:
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$$\frac{\tau_{m_s}^*}{\tau_{m_{ss}}^*} = f \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right) \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{-1} = f$$
 Eq. S19
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3226 1149 Now, we can simplify eq. S16 based on eq. S19:
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$$q^* = f \cdot \left(\frac{\tau_s^*}{\tau_{ss}^*}\right)^{\alpha+\beta} \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{3/2}$$
 Eq. S20
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3232 1151 Combining eq. S20 with that for the dimensionless Shields stress based on the D_{84} (eq. S8), then:
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$$q^* = f \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{3/2-\alpha-\beta} = f \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{-5} \rightarrow \left(\frac{D_{84_s}}{D_{84_{ss}}}\right) = \frac{1}{f} \cdot q^{*-1/5}$$
 Eq. S21
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3239 1153 at low transport conditions.
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3242 1154 At high transport conditions (when, $\tau_{ref}^* \gg \tau_{m_s}^*$), we should combine eq. S14 with eq. S15, and
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3245 1155 we arrive at:
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3247 1156
$$q^* = \left(\frac{\tau_s^*}{\tau_{ss}^*}\right)^\alpha \cdot \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{3/2}$$
 Eq. S22
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3250 1157 Combining this expression with that for the dimensionless Shields stress:
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3253 1158
$$q^* = \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{3/2-\alpha} = \left(\frac{D_{84_s}}{D_{84_{ss}}}\right)^{-1} \rightarrow \left(\frac{D_{84_s}}{D_{84_{ss}}}\right) = q^{*-1}$$
 Eq. S23
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3257 1159 at full mobility conditions.
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3260 1160 **S. 3. How did we compute the ‘bedload supply index’ for the compiled dataset?**

3261 1161 In Section 5 of the paper, we investigated and discussed whether available field measures of
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3264 1162 bedload fluxes may explain some trends in armour ratios (D^*) in the compiled data. More
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3267 1163 specifically, we analysed if the ‘bedload supply index’ (q^*) (at bankfull) could be correlated to
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3270 1164 the armour ratio in the available data and how well it adjusted to eq. S22 and/or S23 (as concluded
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3273 1165 in S. 2). According to Dietrich et al. (1989), q^* can be estimated from eq. S6. Estimating q^* from
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3276 1166 eq. S6 needs two inputs: i. actual bedload rates (in the numerator); and ii. bedload capacities in
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the denominator. In this paper, the numerator was defined from the bedload rates measured at

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1168 (close to) dominant flows for each compiled study case. Bedload rates at dominant discharges
1169 were calculated following these steps: i. for each case study we defined the best power fit relating
1170 bedload to water discharge (based on the available bedload discharge information, see Table 1 in
1171 the paper); ii. we applied the obtained regression equation to the bankfull discharge, and we
1172 computed the bedload discharge at dominant flows. Concerning the denominator q_{ss} in eq. S6, we
1173 approached channel transport capacities (for all the selected case studies) from the bedload rates
1174 estimated using Recking's formula (eqs. S10-S11), but based on subsurface GSD.
1175 The approach followed when defining q^* was obviously different in the case of flume data. As
1176 outlined in the main text of the paper, we also benefited from the results of some flume
1177 experiments that explored the role of sediment feed reductions on surface texture (Church et al.,
1178 1998; Church and Hassan, 2000; Nelson et al., 2009). Data from these flume experiments was
1179 extracted from the graphical reading of figures presented in Venditti et al. (2017). Venditti et al.
1180 (2017) estimated q^* for these flume experiments as the ratio between the bedload transport rate
1181 after the bed has adjusted to a new sediment feed divided by the transport rate for an unarmoured
1182 bed, before any feed reduction.

1183 **S. 4. How did we compute transport-stage ratios for the compiled dataset?**

1184 In Figures 8 and 10, we worked with and analysed transport stage ratios, which are defined as the
1185 ratio between Shields stresses and the critical (or reference) Shields stresses for incipient sediment
1186 motion (Wilcock et al., 2009):

$$1187 \text{Transport stage} = \frac{\tau^*}{\tau_c^*} \quad \text{Eq. S24}$$

1188 Shields stresses are calculated for the 84th percentile of GSD in this paper, since D_{84} provides a
1189 measure of the largest grains in the bed, which probably control hydraulic roughness and
1190 morphological stability (MacKenzie et al., 2018). In eq. S24 (and Figures 8 and 10 in the paper),
1191 τ_c^* is the critical Shields stress below which no sediment transport is assumed to occur. This
1192 parameter is grain-size and slope dependent (Lamb et al., 2008; Pitlick et al., 2008; Recking,
1193 2009). Here, we estimated it based on Recking (2009):

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1194 $\tau_c^* = (1.32 \cdot S + 0.037) \cdot \left(\frac{D_{84}}{D_{50}}\right)^{-0.93}$

Eq. S25