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Redistribution of germanium during dynamic recrystallization of sphalerite

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ABSTRACT

Rare metals are essential to the development of the “green” technologies that are at the core of low-carbon societies. In nature, these metals are frequently present in trace amounts scattered in base metal ore deposits, but the physico-chemical processes that are responsible for their concentration into strategic minerals are still poorly understood. Based on laser-induced breakdown spectroscopy (LIBS), coupled with electron backscattered diffraction (EBSD) analysis, this study shows that plastic deformation and subsequent syntectonic recrystallization of sphalerite (zinc sulfide, ZnS) led to the spatial redistribution of germanium (Ge): from a background level of a few hundreds of parts per million in undeformed primary sphalerite to tens of weight-percent in neocrystallized Ge minerals. During dynamic recrystallization, Ge is likely released from the crystal lattice of parent sphalerite and subsequently concentrated in Ge minerals, leaving behind a Ge-depleted, recrystallized sphalerite matrix. Identifying how rare metals concentrate through deformation and syntectonic recrystallization at the mineral scale is essential to understand the spatial redistribution and localization at the deposit scale. This study highlights the importance of coupling *in situ* chemical mapping analysis with macro- and microstructural characterization when targeting rare metals in deformed ore.

INTRODUCTION

Many rare metals are critical to the sustained development of 21st century technologies; for example, germanium (Ge) is essential to optic-fiber systems used in military and civil telecommunications. The production of rare metals remains challenging, notably because only a few percent of the total rare metal content is recovered from base metal deposits (Ruiz et al., 2018; U.S. Geological Survey, 2019). In this study, we focused on natural processes leading to the spatial redistribution of Ge, a

rare metal for which the background concentration level in the continental crust does not exceed 1.5 ppm (Smith and Huyck, 1999). Redistribution of rare metals during metamorphism and deformation of base metal sulfides may be overlooked by standard mineral exploration approaches. Understanding the redistribution of rare metals that form their own high-grade minerals will lead to a better availability of these elements during ore processing, as well as enhanced extraction and recovery efficiencies. In nature, Ge is associated with sphalerite (zinc sulfide, ZnS) deposits and coal seams at concentrations of a few thousand parts per million (Höll et al., 2007; Frenzel et al., 2014, 2016).

Germanium is generally recovered through pyro- and hydro-metallurgical bulk ore processing (Moskalyk 2004). In some lead, zinc, and copper deposits, small (<100 µm) accessory Ge minerals that can contain up to 70 wt% Ge are locally described, testifying to Ge remobilization and/or concentration at the deposit scale (Johan et al., 1983; Bernstein, 1985; Melcher et al., 2006; Horn et al., 2018). These Ge minerals are found in a large variety of deposits, such as Kipushi-type (breccia-pipe), orogenic veins, volcanic-hosted massive sulfide (VHMS), or epithermal deposits (Ottemann and Number, 1972; Intiomale and Oosterbosch, 1974; Wagner and Monecke, 2005; Belissant et al., 2019), and in Neoproterozoic to Variscan age host rocks (Johan et al., 1983; Melcher et al., 2006; Horn et al., 2018). In addition, Ge minerals typically occur in hydrothermal systems hosted in low-grade metamorphic rocks (Cugerone et al., 2018a). Nevertheless, the physico-chemical mechanisms that induce the remobilization of Ge in sulfide ores are not sufficiently understood, and, consequently, the economic potential of this rare metal occurrence remains difficult to assess.

In the past two decades, numerous studies have shown that trace-element distribution in deformed sulfides may be closely related to microstructures (e.g., in pyrite, pyrrhotite, löllingite-arsenopyrite; Tomkins and Mavrogenes, 2001; Cook et al., 2009a; Vukmanovic et al., 2014). Here, we propose an innovative approach combining (1) electron backscattered diffraction (EBSD), (2) laser-induced breakdown spectroscopy (LIBS), and (3) *in situ* electron probe microanalyzer (EPMA) measurements to illustrate the relationships between Ge distribution and sphalerite microstructure. In naturally deformed sphalerite from the Arre Pb-Zn deposit in the Pyrenean Axial Zone (PAZ), Pyrenees, brunogeierite (GeFe_2O_4) and carboirite [$\text{GeFeAl}_2\text{O}_5(\text{OH})_2$] have already been described (Laforet et al., 1981; Oudin, 1982; Johan et al., 1983; Johan and Oudin, 1986). Therefore, this ore deposit is a first-order target to decipher the links between sphalerite deformation and Ge remobilization.

GEOLOGICAL SETTING AND SAMPLES

The study area is located in the Paleozoic basement of the PAZ, where numerous Pb-Zn deposits have been exploited in the 20th century (Fig. 1A). Following Denèle et al. (2014), we documented three Variscan deformation stages (D_1 to D_3). A poorly expressed S_1 foliation is associated with regional D_1 deformation and medium-pressure and low-temperature M_1 metamorphism. A well-expressed but local S_2 cleavage is subvertical and is related to a late Variscan D_2 deformation stage, characterized by a low-pressure and medium-temperature M_2 metamorphism (Mezger and Passchier, 2003). Finally, a very late Variscan or Pyrenean D_3 deformation occurs as localized shear zones and faults (Mezger et al., 2012; Carreras and Druguet, 2014; Cochelin et al., 2017).

Three types of Pb-Zn occurrences are recognized by Cugerone et al. (2018b): type 1, type 2a, and type 2b. This study focuses on the Arre locality (Fig. 1B), where epigenetic type 2b veins have been reported (Reyx, 1973). Three subvertical sphalerite veins, striking N060°E to N080°E, are embedded in Lower-Devonian host marble and calc-schist. These veins are structurally located in the hinge zone of a kilometer-scale F_2 anticline and are parallel to the local S_2 cleavage (Figs. 1C and 2A).

The sphalerite occurrence was sampled in the westernmost part of the Arre locality (Fig. 1C). These decimetric veins crosscut S_0 – S_1 structures (Fig. 2A) and are essentially composed of brownish sphalerite (~80 vol%; Fig. 2B) with minor chalcopyrite, pyrite, galena, graphite, and muscovite in a gangue of cockade quartz, carbonate, and rare barite (Fig. 2C).

METHODS

EPMA was carried out using a Cameca SX100 at Géosciences Montpellier (Montpellier, France) following the methodology outlined by Cugerone et al. (2018a). Ge and Cu were analyzed using two large lithiumfluoride (LLIF) monochromators. Detection limits calculated by internal Cameca procedures (PeakSight program (<https://www.cameca.com/service/software/peaksight>) in the SX100 software, using 1σ error) were 85 ppm and 123 ppm for Ge and Cu, respectively (Item DR1 in the GSA Data Repository¹).

Simultaneous energy-dispersive spectrometry (EDS) and EBSD mapping were acquired with a Camscan Crystal Probe X500FE scanning electron microscope at the University of Montpellier (France). Operating conditions were 20 kV and 5 nA with a working distance of 25 mm under 2 Pa low vacuum and a step size of 5 μm . Data were

processed with the Oxford Instruments AZtec and Channel 5 software

(<https://nano.oxinst.com/products/ebsd/post-processing-software>).

LIBS multi-elemental imaging was performed at the Institute Lumière Matière, University of Lyon 1 (France); this all-optical analytical technique was used to map the distribution of Ge in sphalerite (Cáceres et al., 2017; Fabre et al., 2018). The sample was analyzed using a Nd:YAG laser with a pulse energy of 600 μJ operating at 100 Hz and a lateral resolution of 13 μm . The spectrometer was configured to detect an intense line of Ge located at 265.1 nm. The resulting detection limit was 10 ppm.

RESULTS

Sphalerite Microstructures and Locations of Ge Minerals

The vein is dominated by sphalerite associated with quartz (Figs. 2B and 2C). Sphalerite appears intensely zoned with brownish tint variations in transmitted light microscopy (Figs. 2C and 3A). The sphalerite grain size is highly heterogeneous (from $<10 \mu\text{m}$ to $>1 \text{mm}$) and distributed in two distinct populations (Figs. 3B and 4C): The threshold between coarse- and fine-grained size was fixed at 100 μm (Fig. 3D). Combining the EBSD grain size map and optical observations, three grain types could be distinguished: coarse millimeter-size grains with (1) dark or (2) light-brownish domains, and (3) small ($<100 \mu\text{m}$) light-brownish grains. Microstructures in twinned primary sphalerite were mostly internal crystalline strain expressed through deformation bands and lamellae (Figs. 3C and 4A). In a single parent grain, the cumulative internal misorientation could reach a value of 22° (Fig. 3E).

Small sphalerite grains formed aggregates that pervasively mantled and/or crosscut coarse primary sphalerite (Figs. 3C, 3E, and 4A). The misorientation angle

distribution of neighbor pairs revealed a dominance of low values (from 2° to 6°; Fig. 3F), four times higher than the 45° peak of uncorrelated pairs, which followed the predicted random distribution. At thin section scale, fine-grained aggregates outlined the S₂ cleavage (Figs. 3 and 4), and in coarse primary sphalerite, narrow bands composed of small grains were frequently parallel to twins (Figs. 3C and 4A).

The EDS chemical map acquired in this area (Item DR2) indicates that Ge was locally hyperconcentrated in micrometric Ge minerals (reported as yellow dots in Fig. 3B), located mostly in the fine-grained sphalerite aggregates.

Ge Partitioning in Sphalerite

Germanium concentration was mapped in relation to sphalerite texture in a monomineralic layer of fine-grained sphalerite that crosscut an aggregate of large, twinned primary grains (Figs. 4A–4C). The distribution of Ge cations in sphalerite at the thin section scale was highly heterogeneous (Fig. 4B) and mimicked the patchy brownish domains visible in optical microscopy (Item DR4). The most Ge-depleted zones in sphalerite (i.e., blue areas in Fig. 4B) corresponded to both the fine-grained population and random parts of light primary sphalerite (Figs. 4A and 4B). Intermediate Ge concentrations in sphalerite matched the dark part of coarse grains (i.e., brown tints in Fig. 4B). The limit between Ge-rich (dark) and Ge-depleted (light) sphalerite domains was generally progressive and diffuse (Fig. 4B). Zones displaying higher Ge contents (i.e., >15 wt% Ge; yellow areas in Fig. 4B) clearly indicated the presence of Ge minerals (brunogeierite; see inset in Fig. 4C). These were located in twinned, coarse primary grains, and in the fine-grained monomineralic layers that outlined the S₂ cleavage (Figs. 4A–4C; Item DR4).

Ge and Cu contents have been measured in the three sphalerite textures from several vein-type deposits in the PAZ (including the studied Arre deposit; Fig. 4D; Item DR3). The lowest Ge concentrations in sphalerite are below ~150 ppm Ge and are found in light fine-grained aggregates and light domains in coarse primary sphalerite. Dark domains in coarse primary grains display intermediate Ge concentrations from ~150 to ~650 ppm Ge.

INTERPRETATION AND DISCUSSION

Deformation and Dynamic Recrystallization of Sphalerite

The Arre vein displays evidence of strain superimposition, which was verified at the sample scale by a subvertical S_2 cleavage. Strain in sphalerite is primarily accommodated by internal crystalline plasticity forming deformation bands and lamellae (Figs. 3C and 3E). This process indicates active dislocation creep mechanisms in the ZnS (Kollenberg and Siemes, 1983; Hirth and Tullis, 1992). By comparison with experimental deformation in sphalerite (Siemes and Borges, 1979) and deformation mechanisms in pyrite that support a brittle-ductile transition zone as low as 200–260 °C at geological strain rates (Barrie et al., 2011), dynamic recrystallization of sphalerite may occur at lower temperature even if no deformation map is yet available to confirm these conditions.

The textures show coarse sphalerite grains mantled by small grains (<100 μm ; Figs. 3B and 3D), or cut through by S_2 -parallel zones of fine-grained sphalerite. Twins in parent sphalerite often act as inherited structural heterogeneity during deformation (Couderc et al., 1985), inducing the formation of narrow bands of small, aggregated grains (Figs. 3C and 4A). These observations point to syntectonic grain-size reduction,

deriving from dynamic recrystallization. Indeed, small recrystallized grains develop at the expense of large parent grains (Figs. 3C and 4A). With the rising strain and the increase of internal misorientations within parent grains (beyond 20°; Fig. 3E), the high density of low misorientation angles between neighbor pairs (Fig. 3F) suggests that strain-induced subgrain rotation recrystallization was the main deformation mechanism for sphalerite recrystallization (e.g., Guillope and Poirier, 1979).

Mechanisms for Internal Ge Remobilization

Sphalerite integrates tetravalent Ge involving monovalent Cu in a coupled substitution mechanism (Belissont et al., 2016). Germanium generally occurs in relatively low concentrations (up to a few thousands of parts per million), often associated with coarse parent sphalerite grains that display Ge-rich (dark) and Ge-depleted (light) domains (Cook et al., 2009b; Belissont et al., 2014; Bauer et al., 2018). Here, we show that deformation at low- to medium-grade (<400 °C) conditions leads to dynamic sphalerite recrystallization, inducing Ge partitioning between recrystallized Ge-depleted sphalerite and newly crystallized Ge minerals. Although Ge is not homogeneously distributed in parent sphalerite (i.e., rich-dark vs. depleted-light domains), Ge displays higher pervasive concentration in dark coarse grains (up to ~650 ppm) compared to light recrystallized domains (<100 ppm), in which Ge minerals frequently occur (Figs. 3B, 4B, and 4C). We propose that Ge cations were internally remobilized from the parent sphalerite lattice during intracrystalline deformation, most likely via solid-state mechanical transfer (dislocation flow; Marshall and Gilligan, 1993), as a result of increasing internal strain and subsequent recovery. During dynamic recrystallization and S₂ cleavage development, a probable fluid-assisted solution transfer (solution-

precipitation creep) allowed Ge to be remobilized and promoted the neocrystallization of Ge minerals (essentially oxides or chloritoids; Fig. 4C; Johan et al., 1983; Cugerone et al., 2018a). Germanium from an external fluid source is not needed to crystallize the newly formed Ge minerals.

Rare Element Redistribution in Other World-Class Deposits

Germanium minerals and Ge chemical heterogeneity are reported from other deformed sulfide occurrences. For example, Kipushi-type deposits (Melcher et al., 2006; Horn et al., 2018), volcanic-hosted massive sulfide deposits (Reiser et al., 2011), and deformed carbonate-hosted deposits in China (Huize, Niujaotang deposits; Ye et al., 2011) and in Brazil (Monteiro et al., 2006) show recrystallized Ge-rich sphalerite ore. In other chemical systems, formation of the indium-bearing mineral roquesite (CuInS_2) hosted in bornite (Cu_5FeS_4) may be attributed to indium remobilization and diffusion from In-rich sphalerite under metamorphic conditions (Jonsson et al., 2013). Coupling microstructural and *in situ* chemical analysis of deformed sulfides is key to exploring potential rare metal-rich deposits in orogens.

CONCLUSIONS

Multiscale textural and chemical observations show that the Arre vein-type zinc sulfide deposit has undergone deformation that promoted dynamic recrystallization of sphalerite, inducing Ge remobilization and subsequent concentration in Ge minerals. This mechanism, assisted by hydrothermal fluids, led to the development of a fine-grained Ge-depleted sphalerite matrix, in which Ge minerals crystallized. Numerous analogues to the Arre epigenetic sphalerite exist in world-class Pb-Zn deposits, and coupling textural and

microstructural observations with trace-element chemical mapping may be, in the future, a key improvement to help understand the natural occurrence of rare metals.

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FIGURE CAPTIONS

Figure 1. (A) Map of Variscan massifs in Western Europe, and location of B. (B) Simplified geological map of the Pyrenean Axial Zone (PAZ) and location of main Pb-Zn deposits (yellow stars, from Pouit [1985] and Bureau de Recherches Géologiques et Minières database [<http://infoterre.brgm.fr>] and Instituto Geológico y Minero de España

database [<http://mapas.igme.es/Servicios/default.aspx>]). Study area is highlighted with blue star. (C) Geological and structural map of the Arre deposit area. Pb-Zn ore is hosted in N060°E–N080°E subvertical veins hosted in Lower Devonian marble calc-schist. Location of Figure 2A is indicated.

Figure 2. (A) Outcrop photograph of Arre (Pyrenean Axial Zone, Pyrenees) Pb-Zn vein type. Note S_2 cleavage in the host rock. Location of the studied sample (B) is indicated (modified from Cugerone et al., 2018b). (B) Photograph of ore. Sphalerite (Sph) appears light gray and is associated with quartz aggregate (Qtz) and schistose host rock. Location of the studied section (C) is indicated. (C) Plane-polarized (transmitted) light microphotograph of studied ore (150- μm -thick section). Sphalerite presents color zonations (dark and light brown) and is crosscut by S_2 cleavage. Location of Figure 4 is indicated. Sph—sphalerite, Qtz—quartz, S_2 — S_2 regional cleavage plane.

Figure 3. Sphalerite microstructures and locations of Ge minerals in Arre (Pyrenean Axial Zone, Pyrenees) sample. (A) Plane-polarized (transmitted) light microphotograph of mineralized sample (30 μm thin section). Sph—sphalerite, Qtz—quartz, S_2 — S_2 cleavage plane. (B) Locations of Ge minerals superimposed on sphalerite electron backscattered diffraction (EBSD) grain-size map (misorientation threshold to define grains is fixed at 20°). Sphalerite grains are distinguished into two groups as a function of their grain size. Smaller grains (<100 μm) define aggregated bands parallel to S_2 cleavage. Note that Ge minerals are mainly located in, or close to, these fine-grained bands. (C) Sphalerite EBSD inverse pole figure (IPF) map, plotted with respect to x -axis

of projection sphere, showing internal plastic strain in coarse parent grains that are mantled by recrystallized smaller grains. S_2 cleavage is marked by fine-grained recrystallization bands that crosscut parent sphalerite. Location of cumulative misorientation profile (E) is specified. (D) Histogram distribution and cumulative frequency curve of sphalerite grain size displaying two grain-size populations: coarse ($>100\ \mu\text{m}$, brown) parent grains and small ($<100\ \mu\text{m}$, blue) recrystallized grains. Threshold is considered where cumulative frequency curve becomes subhorizontal. (E) Cumulative misorientation profile relative to start point acquired through single sphalerite parent grain, crosscut by fine-grained band of recrystallized grains. (F) Misorientation angle distribution histograms plotting neighbor-pair (correlated) and random-pair (uncorrelated) misorientations for sphalerite grains in studied area ($>2^\circ$).

Figure 4. Ge concentration in sphalerite and locations of Ge minerals in Arre (Pyrenean Axial Zone, Pyrenees) sample. (A) Electron backscattered diffraction (EBSD) inverse pole figure (IPF) map, plotted with respect to x -axis of projection sphere. Note that S_2 cleavage is marked by fine-grained recrystallization bands that crosscut parent sphalerite. (B) Ge distribution map from laser-induced breakdown spectroscopy (LIBS). Ge minerals are represented by higher Ge concentration (yellow areas). Patchy dark-brown domains correspond to higher Ge concentration (~ 600 ppm Ge) in parent sphalerite. Lighter domains (blue to light brown) indicate intermediate to low Ge concentrations (10–150 ppm Ge) in sphalerite. Recrystallized grains show lower Ge concentration (blue, <100 ppm Ge). Qtz—quartz. (C) Locations of Ge minerals superimposed on sphalerite EBSD grain-size map. Note the presence of Ge minerals in the finest sphalerite fraction

(<100 μm) or in twin planes. Reflected plane-polarized light, close-up microphotograph shows brunogeierite (GeFe_2O_4) crystal. (D) Ge versus Cu concentrations (expressed in ppm) in sphalerites obtained from electron probe microanalyzer (EPMA) measurements, in dark, light, and recrystallized sphalerite from the Arre deposit (colored symbols) and other Pyrenean Axial Zone (PAZ) vein-type sphalerite (white symbols). Color code corresponds to that used in LIBS map in B.

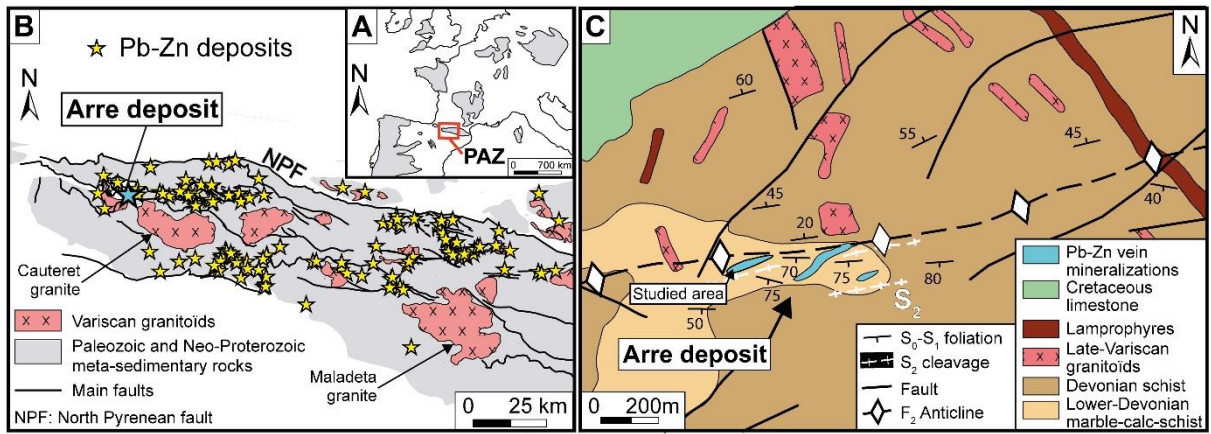


FIG. 1

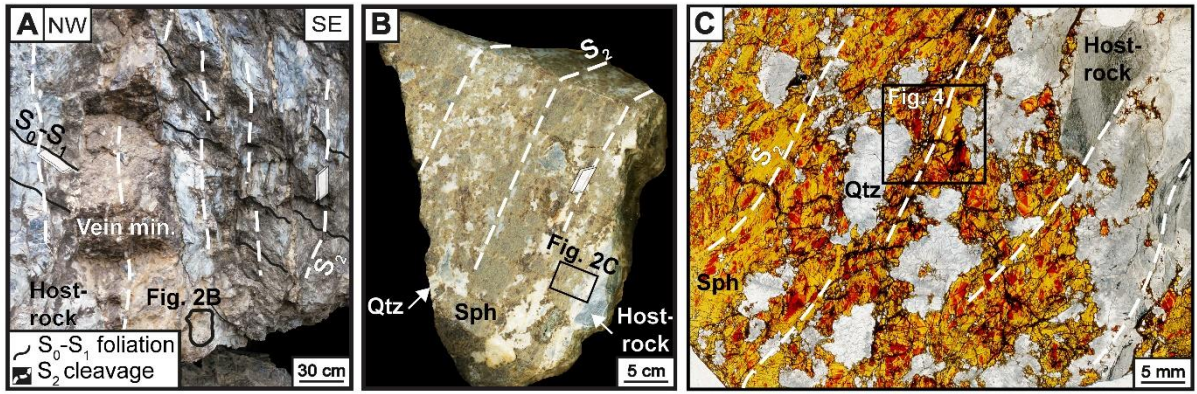


FIG. 2

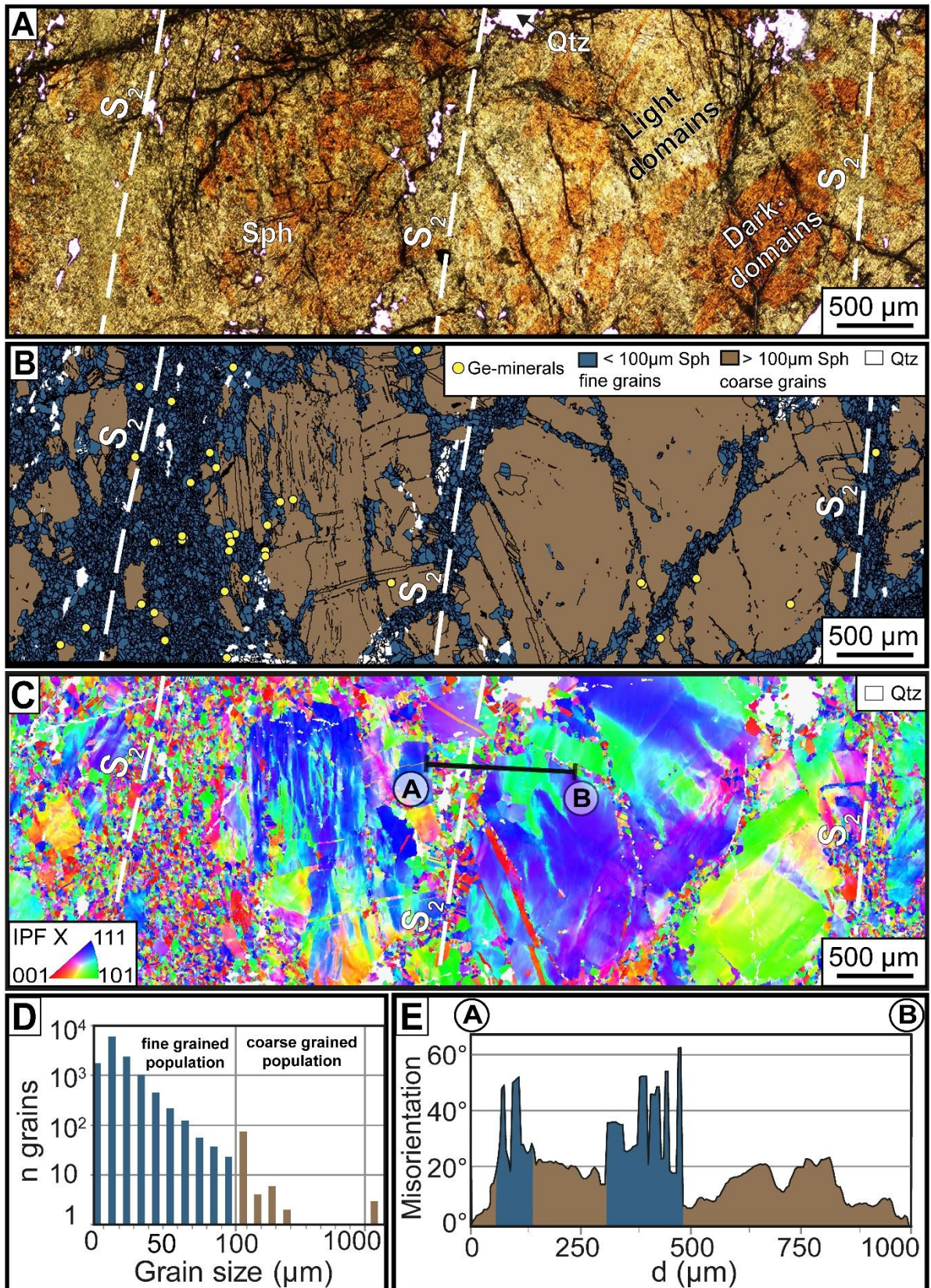


FIG. 3

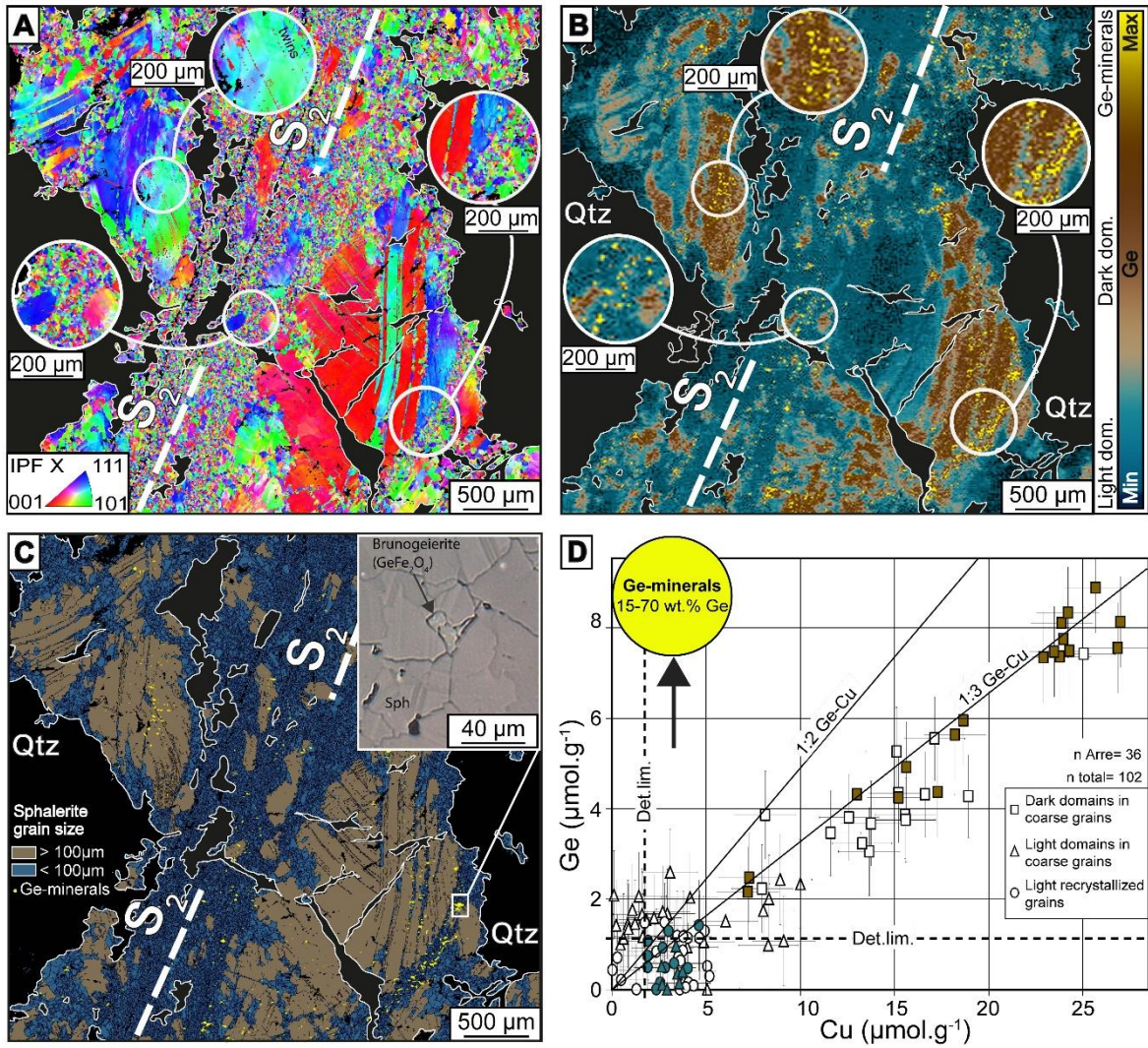


FIG. 4