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1 **The role of tides in ocean–ice-shelf interactions in the southwestern**
2 **Weddell Sea**

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10 **Key Points:**

- 11 • Near-observed Weddell Sea gyre, continental-shelf & cavity water masses in
12 interactive ocean/sea-ice/tide/ice-shelf melt simulations
- 13 • Tides amplify basal melt-refreeze pattern, raising net melt by 50%, associated
14 with enhanced time-varying kinetic energy at the ice base
- 15 • Impacts of additional meltwater production on Weddell continental shelf sea
16 ice and newly formed bottom water properties are suggested

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Abstract

To investigate the role of tides in Weddell Sea ocean–ice-shelf-melt interactions, and resulting consequences for ocean properties and sea-ice interactions, we develop a regional ocean–sea-ice model configuration, with time-varying ocean boundary and atmospheric forcing, including the deep open ocean (at 2.5-4 km horizontal resolution), the southwestern continental shelf (≈ 2.5 km), and the adjacent cavities of eastern Weddell, Larsen and Filchner-Ronne ice shelves (FRIS, 1.5-2.5 km). Simulated circulation, water mass and ice-shelf melt properties compare overall well with available open-ocean and cavity observational knowledge. Tides are shown to enhance the kinetic energy of the time-varying flow in contact with the ice shelves, thereby increasing melt. This dynamically-driven impact of tides on net melting is to almost 90 % compensated by cooling through the meltwater that is produced but not quickly exported from regions of melting in the Weddell Sea cold-cavity regime. The resulting systematic tide-driven enhancement of both produced meltwater and its refreezing on ascending branches of, especially the FRIS, cavity circulation acts to increase net ice-shelf melting (by 50 % in respect to the state without tides, ≈ 50 Gt/yr). In addition, tides also increase the melt-induced FRIS cavity circulation, and the meltwater export by the FRIS outflow. Simulations suggest attendant changes on the open-ocean southwestern continental shelf, characterized by overall freshening and small year-round sea-ice thickening, as well as in the deep southwestern Weddell Sea in the form of a marked freshening of newly-formed bottom waters.

1 Introduction

The polar sector of the Southern Ocean and its interaction with the cryosphere plays a key role in the global climate system. Ocean–ice-shelf interactions and the associated ice-shelf basal melt are an important cause of the Antarctic ice-sheet mass loss, which is currently accelerating, and is of global concern for current and future sea-level rise (e.g. *The IMBIE team*, 2018).

Ocean–sea-ice–ice-shelf interactions form the densest water masses of the world ocean that sink, fill the bottom of the world ocean, and represent an important engine for the global ocean

43 overturning circulation. In spite of their importance for climate, these intricate interactions and
44 associated mechanisms are still poorly understood, and represent one glaring uncertainty in our
45 current knowledge of climate and its changes.

46 An important block to our understanding of the impact of these polar processes at global scale is
47 the scarcity of polar Southern Ocean observations due to its remoteness, the seasonal presence of
48 sea ice, harsh meteorological conditions, and the difficulties of sampling its ice-shelf cavities. In
49 addition, the current generation of global climate models does not include ocean–ice-shelf
50 interactions, which further limits our process understanding. Recent studies that do account for
51 some of these processes predict, however, considerable global climate impacts and important
52 associated climate feedbacks (*Bronse laer et al.*, 2018; *Golledge et al.*, 2019). Adequate regional
53 modelling tools that capture this regional ocean and its interaction with sea ice and ice shelves are
54 thus called for, to improve our understanding of the physical processes at play, to enable
55 understanding of variability and predictions of responses to change, as well as to inform the
56 accurate representation of these processes in the next generations of global climate models.

57 Here, we present a new regional ocean–sea-ice model configuration of the Weddell Sector of the
58 Southern Ocean, including open-ocean gyre, continental shelf, and cavity ocean interactions with
59 melting ice shelves in presence of tides. The Weddell Sea hosts the largest ice-shelf by volume
60 (the Filchner-Ronne Ice Shelf; FRIS), as well as the largest production of Antarctic Bottom
61 Water (*Orsi et al.*, 1999), parts of which are exported through its deep northern passages (e.g.
62 *Heywood*, 2004) shaping global ocean bottom properties (e.g. *Orsi et al.*, 2002; *Jullion et al.*,
63 2014; *Purkey et al.*, 2018). Whereas currently Weddell Sea ice shelves, to the exception of the
64 Larsen ice shelves adjoining the peninsula, are observed to be amongst the most stable (*Rignot*
65 *et al.*, 2019; *Schröder et al.*, 2019), the region features some of the largest uncertainties in
66 long-term predictions of ice-shelf melting and thereby sea-level contribution (e.g. *Nowicki and*
67 *Seroussi*, 2018; *The IMBIE team*, 2018; *Shepherd et al.*, 2018). Recent modelling results indeed
68 predict the possibility of drastic (and possibly irreversible) changes in local ocean–ice-shelf
69 interactions (*Hellmer et al.*, 2012, 2017; *Naughten et al.*, 2018a), with, if to occur, consequences
70 for global ocean water-mass properties (*Golledge et al.*, 2019) and the (East) Antarctic ice sheet

71 (*Timmermann and Goeller, 2017*).

72 One of the outstanding characteristics of the southwestern Weddell Sea is that it, and especially
73 its continental shelves and ice-shelf seas, features some of the Southern Ocean's most intense
74 tides. The maximum tidal ranges is in excess of 2 m nearly everywhere in the region, peaking at
75 more than 6 m in the FRIS cavity and featuring several zones of intense tidal currents reaching 1
76 m/s (e.g. *Padman et al., 2018*), as known from point-wise observations (e.g. *Foldvik et al., 1990*;
77 *King et al., 2011*) and barotropic tide-only modelling (*Robertson et al., 1998; Makinson and*
78 *Nicholls, 1999*). The presence of these intense tidal circulations is now known to play an
79 important role in ice-shelf dynamics and their exchanges with the ocean (e.g. review by *Padman*
80 *et al., 2018*). In particular, previous modelling work has identified that tides play a key role in the
81 region in terms of exchanges between the ice-shelf cavity and the open ocean, as well as the
82 mixing of water masses on the continental shelf (*Makinson and Nicholls, 1999*), and the
83 exchange of water-masses across the continental slope (*Robertson, 2001a; Pereira et al., 2002*;
84 *Stewart et al., 2018*). Nonetheless, simulations of tides in interaction with the Weddell open and
85 cavity ocean circulation are so far limited to idealized configurations without basal melting
86 (*Robertson, 2001a,b*), a small single ice-shelf subdomain (*Mueller et al., 2012*), or idealized
87 forcing. In particular, the two regional ocean-tide modelling studies of the large FRIS cavity and
88 its adjacent shelf, used either a homogeneous cold or warm shelf state (*Mueller et al., 2018*), or
89 thermohaline restoring on the continental shelf with a fixed spatial pattern and an idealized
90 seasonal modulation (*Makinson et al., 2011*), and did not include substantial deep open-ocean
91 regions. Here, we go one step further by introducing a new realistic configuration of the Southern
92 Weddell Sea sector to address the impact of tides on ice-shelf melting, unveil the processes at
93 play, and investigate the consequences in terms of ocean water-mass characteristics resulting
94 from the ocean-cryosphere interaction in the presence of tides.

95 A large part of our current knowledge of ocean–sea-ice–ice-shelf interactions stems today from
96 regional numerical configurations not including tides (*Dinniman et al., 2016; Asay-Davis et al.,*
97 *2017*). In the Weddell Sea, in particular, several ocean circulation model configurations have
98 been developed including FRIS (*Jenkins and Holland, 2002; Jenkins, 2004; Gerdes et al., 1999*;

99 *Beckmann et al.*, 1999; *Timmermann et al.*, 2002). These simulations were usually done at
100 relatively coarse resolution and often hid the effects of the non-resolved tidal (and non-tidal)
101 circulation on ice-shelf melting in an empirical constant heat exchange velocity (e.g.
102 *Timmermann et al.*, 2002). More recently, global configurations focussing resolution around the
103 Weddell Sea or circumpolar configurations resolving Antarctic ice-shelf cavities have also been
104 developed (*Dinniman et al.*, 2015; *Timmermann et al.*, 2012; *Timmermann and Hellmer*, 2013;
105 *Losch*, 2008; *Schodlok et al.*, 2016; *Mathiot et al.*, 2017). The increasing number of numerical
106 models and configurations starts to allow for intercomparison studies (*Naughten et al.*, 2018b),
107 which begin to shed first light on model dependency, and are, overall, instrumental in refining our
108 understanding of the polar Southern Ocean and its role for climate. Our study complements that
109 diversity of realistic model configurations, and introduces how tides influence the complex
110 ocean–sea-ice–ice-shelf coupled system. The southwestern Weddell Sea configuration designed
111 here features relatively high horizontal resolution ranging from 4.5 km to 1.5 km, with
112 interannually-varying ocean boundary and atmospheric surface forcing, acting on an ocean model
113 coupled to a dynamic-thermodynamic sea-ice model. It captures not only FRIS but also other
114 Weddell cavity systems, namely the Larsen ice shelves downstream and the eastern Weddell ice
115 shelves (EWIS) upstream, and their interaction with the adjacent open continental shelves, and,
116 through the slope front system, the deep open-ocean Weddell gyre (Fig. 1).

117 The model configuration and the experiments designed and analyzed in this study are described
118 in detail in section 2. Simulated ocean – sea-ice – ice-shelf melt interactions in the tidal reference
119 experiment are presented in section 3, including comparisons to available observational
120 estimates. The role of tides in shaping ocean – ice-shelf melt interactions is explicitly assessed,
121 by making use of a non-tidal sensitivity experiment, in section 4, which allows to investigate the
122 physical mechanisms of the tidal impact and assesses wider-reaching impacts of the tide-induced
123 changes on the Weddell climate system. Section 5 provides discussion of results, in particular in
124 comparison to previous modelling, before section 6 concludes.

2 Southwestern Weddell model configuration, experimental design and observational datasets

This section presents the new model configuration (section 2.1) and the experiments (section 2.2) designed for the present study, and introduces various observational datasets discussed in the following in comparison to model results (section 2.3).

2.1 Southwestern Weddell Sea model configuration

The coupled ocean–sea-ice model configuration of the southwestern Weddell Sea developed here includes ice-shelf cavities, continental shelves & the deep open ocean, at a horizontal resolution varying from 4.5 km at 60°S to 1.5 km at the southernmost location in the Ronne cavity (Fig. 1). It is forced at the surface with an interannually varying atmosphere, at the boundaries with interannually varying ocean and sea-ice states as well as tides, all interacting with ice-shelf basal melting in resolved ice-shelf cavities (a detailed presentation of the experimental design choices of the present study follows in section 2.2).

The coupled ocean–sea-ice model used is NEMO (Nucleus for European Modelling of the Ocean) 3.6 (*Madec and NEMO-team, 2016*), with the sea-ice model LIM (Louvain-la-Neuve sea-ice model) 3.6, a dynamic-thermodynamic multi-category multi-layer sea-ice model with elastic-viscous-plastic rheology (*Vancoppenolle et al., 2009; Rousset et al., 2015*), which is used here with 5 ice categories with 5 ice layers each. NEMO’s ocean component is a primitive equation, Boussinesq model on an Arakawa-C grid with a nonlinear equation of state (here we use polynomial TEOS-10; *IOC et al., 2010; Roquet et al., 2015*) and a free surface, which we use here in its nonlinear form, and with a split-explicit formulation of the surface pressure gradient. Corresponding water column thickness variations are distributed through depth with a time-dependant z^* -coordinate (*Adcroft and Campin, 2004*), which has been adapted for the presence of ice shelves, impacting horizontal ocean pressure gradients and modelled here as floating in hydrostatic equilibrium, by *Mathiot et al. (2017)*.

150 Ice shelves in NEMO, as implemented by *Mathiot et al. (2017)*, and so far employed in a global
 151 and an Amundsen regional configuration by *Storkey et al. (2018)* and *Jourdain et al. (2017)*, are
 152 dynamically passive, that is ice loss/gain due to basal melt/freeze are assumed to be
 153 instantaneously equilibrated by ice convergence. This assumption of a static ice-shelf geometry is
 154 a good approximation for simulations on time scales of a few years and decades. The ocean
 155 interacts with ice-shelf basal melting, by driving, and in turn responding to the induced, ice-shelf
 156 melt freshwater and heat fluxes, thermodynamically, and also dynamically through the
 157 meltwater-induced oceanic volume divergence. The thermodynamic interaction, made possible
 158 by the introduction of a diagnostic ice-shelf base interface layer into ocean models as pioneered
 159 by *Hellmer and Olbers (1989)*, is parameterized through a three-equation bulk formulation
 160 (*Holland and Jenkins, 1999; Jenkins et al., 2010; Mathiot et al., 2017*). In this, the oceanic
 161 turbulent heat flux Q_{tbl} flowing upward out of the oceanic cavity top boundary layer (tbl) towards
 162 the ice-shelf base (b) depends not only on the ocean thermal driving (given by difference of
 163 temperature T between tbl and ice base, b), but also on the oceanic tbl current speed c_{tbl} , as
 164 follows

$$Q_{tbl} = \rho_w c_{pw} St_T c_{tbl} (T_{tbl} - T_b), \quad (1)$$

165 and ρ_w and c_{pw} are ocean reference density and heat capacity, and St_T the thermal Stanton
 166 number. Compared to earlier parameterizations (e.g. *Hellmer and Olbers, 1989*), this velocity
 167 dependence of turbulent exchanges has been found to more closely match observations and be
 168 more physically consistent (*Dansereau et al., 2014*). As in the reference experiment of *Jourdain*
 169 *et al. (2017)*, here we use a constant thermal Stanton number $St_T = 0.7 \times 10^{-3}$ in Eq. (1) and
 170 also keep the ratio of thermal and saline Stanton numbers fixed at 35 in the tbl equations.
 171 Simulated melt rates, and related circulation patterns, are sensitive to St_T , a sensitivity that has
 172 been assessed by *Jourdain et al. (2017)* and is not further explored in this study. The
 173 parameterization is applied here to a tbl of 30m thickness (or of the thickness of the top ocean
 174 cell where thicker), and other parameter choices, including the freezing-point linearization
 175 adapted to TEOS-10 ($\lambda_{1,2,3}$), are also as in *Jourdain et al. (2017, cf. their Table 1)*. Although a
 176 simplistic representation that cannot fully account for the dynamic nature of the tbl (*Asay-Davis*

177 *et al.*, 2017; *Jenkins*, 2016), this bulk parameterization remains to date the typical formulation of
178 melt interactions in ocean models with resolved cavities (cf. *Dinniman et al.*, 2016; *Asay-Davis*
179 *et al.*, 2017, for reviews). Furthermore, as common in ocean–ice-shelf melt interactive modelling,
180 the formation of frazil ice is not explicitly represented (e.g. *Bombosch and Jenkins*, 1995), nor is
181 the impact of anticipated spatial heterogeneity in icedraft roughness, and especially its contrasts
182 between melt and freeze regions, on basal melting accounted for (e.g. *Gwyther et al.*, 2015).

183 Ocean model physical parameterizations are also mostly as in *Jourdain et al.* (2017;
184 Suppl. Table S1 lists the few key differences, and provides links to detailed information on the
185 complete sets of choices). Thereby, momentum is advected using an energy- and
186 enstrophy-conserving vector-form scheme, subject to along-geopotential bilaplacian viscosity,
187 and removed through nonlinear bottom and top (under-icedraft) friction, with a frictional drag
188 that varies logarithmically with last model layer depth (e.g. *Dansereau et al.*, 2014, cf. Table S1).
189 Tracers are advected via a flux-corrected transport scheme, and subject to Laplacian iso-neutral
190 diffusion. The vertical eddy diffusivity is $2.0 \times 10^{-6} \text{ m}^2/\text{s}$ (modulated by a horizontal shape
191 function), and effective vertical mixing is obtained via a turbulent kinetic energy closure scheme.
192 Enhanced vertical tracer diffusion is furthermore applied in cases of static instability (cf. Table
193 S1).

194 Our regional configuration uses an isotropic $1/12^\circ$ modified Mercator grid (referred to as
195 eORCA12 in the NEMO community, *Mathiot et al.*, 2017) resulting in a horizontal resolution
196 varying from about 4.5 km at the northern boundary of the gyre, 4 to 2 km over the continental
197 shelves, to less than 2 km in the FRIS cavity (Fig. 1). Thereby it is at the limit of resolving the
198 first baroclinic Rossby radius R_d (of the reference simulation, not shown) in the deep gyre
199 interior as well as the deep Filchner Trough (R_d 4-8 km). On the continental slope and deeper
200 parts of the continental shelves resolution is close to R_d (4-2 km), whereas over shallower areas
201 and in large part of the cavities R_d can be an order of magnitude smaller than the grid scale (≤ 1
202 km). While thus not all mesoscale eddies and their effects on cross-shelf (e.g. *Stewart and*
203 *Thompson*, 2015) and potentially cross-icefront (*Årthun et al.*, 2013) transports are resolved in
204 the current configuration, it is substantially more eddying than other previous ocean–ice-shelf

205 model configurations including the Weddell region. Whereas eddy contributions to cross-icefront
206 exchanges have to date not been assessed in realistic modelling studies, a short eddy-resolving
207 simulation (*Stewart et al.*, 2018) suggests that for ocean-shelf heat transport in the Weddell Sea
208 the role played by eddying circulations (with timescales >1 -day) may be smaller than the roles of
209 the mean circulation and tides. Here neither the Gent-McWilliams parameterization nor
210 parameterizations for smaller-scale surface mixed-layer eddy effects (*Fox-Kemper et al.*, 2011)
211 are used, as eddies are partly resolved over large parts of the domain (and also since interplay
212 with ice shelves requires further developments, in which partial domain parameterizations
213 following e.g. *Hallberg*, 2013, may be useful).

214 Vertically 75 levels yield a resolution varying from 1 m at the surface, 10–150 m beneath ice
215 shelves and 200 m at 6 km depth (*Mathiot et al.*, 2017). The bathymetry is derived from a
216 combination of ETOPO1 in the open Southern Ocean, IBCSO over the Antarctic continental
217 shelves and BEDMAP2 in ice-shelf cavities for bathymetry and icedraft (as described in *Mathiot*
218 *et al.*, 2017, and references therein), before interpolation to the model grid. Bottom and top
219 partial cells are used to better approach bathymetry and icedraft, and the effective model
220 bathymetry and icedraft (Fig. 1) are modified to eliminate vertically unresolved cavity zones.
221 Here the inclusion of tides has been found to furthermore require a minimum initial water column
222 thickness of 10 m (across two partial-cell layers) to achieve numeric stability and accommodate
223 the tide-induced oscillations of water depth under the hydrostatically floating ice shelves.

224 The domain chosen here (pink lines in Fig. 1) covers roughly the southwest quarter of the
225 asymmetric Weddell gyre (e.g. *Gouretski and Danilov*, 1993; *Orsi et al.*, 1993), with an eastern
226 boundary placed just west of 0° E and a main northern boundary that passes near the northwestern
227 deep escape passages through the South Orkney Islands, i.e. near or to the south of the southern
228 boundary of the ACC (e.g. *Orsi et al.*, 1995; *Heywood*, 2004). In the northwestern corner of the
229 domain, a small western segment of the northern open boundary links the tip of the Antarctic
230 peninsula to Joinville Island, and a western boundary traverses Bransfield Strait from Joinville
231 Island to the South Shetland Islands, from where the major segment of the northern boundary
232 extends eastward. Along the open lateral boundaries the model bathymetry is modified to match

233 that of the lateral forcing simulation over a transition zone (as in *Jourdain et al., 2017*), and the
234 ice-shelf cavity passages linking towards Fimbul along the eastern boundary are grounded (east
235 of the solid pink line in Fig. 1), making Jelbart (without its cavity passages towards Fimbul) the
236 easternmost resolved cavity of the current configuration. Following the gyre, the model domain
237 includes three major ice-shelf formations: (1) the eastern Weddell ice shelves EWIS, with a total
238 model surface of $94 \times 10^3 \text{ km}^2$ and a maximum model ice draft of almost 900 m, include Jelbart,
239 Atka, Ekstroem, Quar and Riiser-Larsen ice shelves and the Brunt/Stancomb-Wills ice-shelf
240 system; (2) FRIS with Filchner and Ronne ice shelves covers 425×10^3 model km^2 reaching to a
241 peak depth of ≈ 1700 m; and (3) the Larsen ice shelves along the peninsula, with a model surface
242 of $63 \times 10^3 \text{ km}^2$ and a shallower maximum depth of ≈ 500 m, are Larsen G, F, E, D, C, as well as
243 the remnant of B after its 2002 breakup, the SCAR inlet glacier, and the remnant of A after its
244 1995 breakup, the Seal Nunatak Glacier (Fig. 1).

245 **2.2 Experimental design**

246 In the experiment used as reference here the Weddell Sea regional configuration introduced above
247 is constrained at its lateral open boundaries with tides (as detailed below) as well as
248 interannually-evolving ocean and sea-ice states from a global $1/4^\circ$ coupled ocean–sea-ice
249 simulation with resolved ice-shelf cavities (GO7, *Storkey et al., 2018*, not including tides and
250 with basal melt rates prescribed using the satellite-derived estimate by *Rignot et al., 2013*). The
251 GO7 lateral ocean boundary conditions are imposed as monthly means, using a flow relaxation
252 scheme for temperature, salinity and baroclinic velocities, and the Flather radiation scheme for
253 sea-surface height and barotropic velocities, applied without regional volume control to allow for
254 tides (see *Madec and NEMO-team, 2016*).

255 Here simulations are initialized with a monthly-average temperature & salinity state from GO7
256 (January 1990 namely). This implies a physically consistent regional ocean initial state
257 throughout the domain including its cavities, obtained via a 14-year spinup in the global $1/4^\circ$
258 GO7 experiment prior to the start of our regional simulations (the GO7 simulation is initialized in

259 1976 from a 1995-2014 climatology of the EN4 objective analysis, *Good et al.*, 2013,
260 extrapolated into the opened ice-shelf cavities).

261 For consistency, atmospheric surface-boundary conditions are taken as in the boundary-forcing
262 simulation, from CORE-2 (*Large and Yeager*, 2009) applied via CORE bulk formulae.

263 Additionally, the surface-freshwater forcing of drifting and melting icebergs is prescribed using
264 interannually-varying monthly-average outputs of GO7's Lagrangian iceberg model (*Marsh et al.*,
265 2015), itself fed along the Antarctic coasts by an iceberg calving-flux climatology (*Rignot et al.*,
266 2013). Furthermore, solar radiation is parameterized to penetrate through depth (*Madec and*
267 *NEMO-team*, 2016) depending on prescribed seasonally-varying climatological chlorophyll-A
268 concentrations (derived from SeaWIFS, *O'Reilly et al.*, 1998). Note that, whereas remote effects
269 of sea-surface salinity restoring, which is applied globally in GO7, are imported through the
270 lateral boundaries, within the regional domain no surface restoring is used in this study.

271 Regarding sea ice, GO7 monthly-average interannually-varying sea-ice concentration, thickness
272 and snow thickness are imposed along the open boundaries (via flow relaxation). Sea ice is also
273 initialized, using above properties capped at 99.7 % ice concentration and thicknesses of 4 m for
274 ice and 50 cm for snow (both for numerical stability and consistency with available observational
275 constraints – GO7 has rather thick sea ice near the ice-shelf edges, especially where ice-shelf
276 waters outflow), as well as in addition ice-surface temperature and ice-depth-average salinity
277 (initial ice-depth-average temperature is set to the average of ice-surface temperature and
278 -1.9°C).

279 Tidal elevations and currents along the open-ocean boundaries are imposed using phases and
280 amplitudes of 18 tidal constituents from FES2012 (*Carrère et al.*, 2012; *Lyard et al.*, 2006), a
281 global finite-element solution of the tidal barotropic equations resolving ice-shelf cavities and
282 assimilating multi-satellite altimetry (note this does not involve many data in the ice-covered
283 Southern Ocean). In comparison to the still sparse tidal records available over Antarctic
284 continental margins and ice shelves, tide models, including FES2012, still show rms errors of
285 typically 15-30 % for the major tidal constituents (*Stammer et al.*, 2014). Indeed, comparing tidal

286 elevation timeseries from GPS-based tide gauge records for two sites on FRIS with
287 reconstruction using five contemporary tide models, including FES2012, reveals instantaneous
288 local model errors up to 40 cm (*King et al.*, 2011, see *Stammer et al.* (2014) their Figure 7).
289 Whereas this is not satisfactory for tidal prediction, such errors are superimposed on tidal
290 elevations of up to 2 m underneath FRIS, and the zero-order tidal signal is overall well
291 represented in both timing and amplitude by all the models including FES2012. Here,
292 furthermore, FES2012 outputs are not used directly underneath ice shelves, but only along the
293 open boundaries of the regional domain, where smaller errors are anticipated, as closer to the
294 regions of data assimilation in FES2012. From the boundaries, the ocean model propagates the
295 tidal signals throughout the domain in full interaction with the three-dimensional oceanic
296 circulation. *Jourdain et al.* (2019), for their Amundsen Sea regional configuration, have shown
297 that the resulting simulated tidal signals compare well with those of the input tidal model near the
298 continental shelf break and underneath ice shelves. Nonetheless it is likely that some localized
299 features of the tidal circulation especially in cavities will be missing or misrepresented, if only
300 due to imperfections in bathymetry and icedraft datasets and their discretizations (e.g. *Padman*
301 *et al.*, 2018; *Rosier et al.*, 2018). As in the most complex simulation of *Jourdain et al.* (2019),
302 here we use 18 tidal constituents. These include the 8 major constituents in the Weddell region,
303 namely the four primary, M2, S2, K1, O1, followed by K2, N2, P1, Q1 (*King et al.*, 2011), as
304 well as in addition S1, 2N2, μ 2, ν 2, L2, T2, M4, Mtm, Mf and Mm (e.g. *Schureman*, 1958).

305 In summary, the regional-configuration reference experiment designed here differs from the
306 global simulation, which is imposed at its open-ocean boundaries and provides a 14-year
307 lower-resolution spinup prior to initializing the former, in three key ways: (i) allowing
308 interactions of the ocean with ice-shelf basal melting, (ii) a higher horizontal resolution, thus
309 capturing smaller-scale oceanic-circulation features, and (iii) the explicit inclusion of tides.

310 With this reference experiment design, a simulation of 8 years, 1990–1997, is run with an
311 ocean-model time step of 4 minutes and a sea-ice-model and surface-forcing time step of 12
312 minutes. As discussed below (section 3, Fig. 2), simulated integral ice-shelf melt displays a
313 spinup of $O(1 \text{ year})$, and the analysis will mostly focus on the last 5 years of the 8-year

314 simulation (1993–1997). In addition, one perturbation experiment is realized here to investigate
315 the role of tides in setting circulation, water-mass structure and Weddell Sea ocean–ice-shelf
316 interactions, by repeating exactly the 8-year reference simulation, but without imposing any tidal
317 boundary forcing. We note that all other experiment and parameterization choices, including the
318 dissipation applied, are unchanged between the two experiments.

319 **2.3 Observation-based datasets**

320 To assess model integral ice-shelf melt rates we compare to observation-based estimates provided
321 for individual ice shelves or groups of ice shelves by *Rignot et al. (2013)*, *Depoorter et al. (2013)*
322 and *Moholdt et al. (2014)*. These estimates, representative of the late 2000s, are obtained using
323 surface mass balance modelling and satellite-based estimates of grounding-line and
324 icefront-calving fluxes, from which basal mass budgets are inferred, either assuming steady state
325 for the Weddell ice shelves (*Depoorter et al., 2013*), or furthermore invoking altimetry-derived
326 ice-shelf thickness changes obtained using Eulerian (*Rignot et al., 2013*) or Lagrangian (*Moholdt*
327 *et al., 2014*) methods.

328 For sea ice, are concentration maps combining measurements from various satellite microwave
329 radiometers and imagers (SMMR, SSM/I & SSMIS, converted using the bootstrap algorithm
330 version 2) are obtained as monthly means from nsidc.org/data/nsidc-0079/versions/2 (*Comiso,*
331 *2000, updated 2015*) for the simulation period. Weddell sea-ice area coverage timeseries are
332 obtained by integrating sea-ice area concentration data on their native 25 x 25 km
333 polar-stereographic grid across the regional model domain using a concentration threshold of 15
334 %.

335 To analyze open-ocean water-mass properties in comparison to the observed open-ocean
336 hydrography, we use available individual in-situ temperature & salinity data from
337 conductivity-temperature-depth (CTD) profiles acquired post 1972 within the model regional
338 domain, compiled as a collection of profiles obtained from the NOAA World Ocean Database
339 (www.nodc.noaa.gov/OC5/SELECT/dbsearch/dbsearch.html), the Pangea database

340 (www.pangaea.de), the Coriolis Argo international program GDAC
341 (ftp://ftp.ifremer.fr/ifremer/argo), and the MEOP instrumented seal consortium (www.meop.net/).
342 We only use profiles that have a quality control flag of 1, containing information on their
343 position, date, temperature, and salinity profiles. Profiles are converted to conservative
344 temperature and absolute salinity using the TEOS-10 tools (*IOC et al.*, 2010). Regarding FRIS
345 cavity water masses, under ice-shelf CTD profile data from hot-water drill sites (mapped in
346 Fig. 1) have been obtained as profiles through depth, covering most of the water column, at
347 different times, during December 1990 at Site 1 (S1, *Nicholls et al.*, 1991), January 1992 at Site 2
348 (S2, *Robinson et al.*, 1994), January 1996 at Site 3 (S3, *Nicholls et al.*, 1997), December 1998 at
349 Sites 4 and 5 (S4 and S5, *Nicholls et al.*, 2001), as well as in December 2002 at F2 and F3 and
350 January 2003 at F1 and F4 (*Nicholls et al.*, 2004), and in January 2016 at FSW1, FSW2, FSE1,
351 FSE2, and in December 2016 at FNE1, FNE2 and FNE3 (*Huhn et al.*, 2018).

342 **3 Characterizing circulation and water masses in the reference** 353 **experiment**

354 The circulation and water masses of the reference experiment, i.e. including tides (cf. section 2),
355 are now assessed, and discussed in comparison to key observation-derived metrics. Investigation
356 of the role played by this tidal forcing, and discussion in comparison to previous modelling
357 studies, follows below (sections 4 and 5).

358 **Ice-shelf melting** Reference experiment integral timeseries of ice-shelf melting are displayed
359 in Fig. 2a (in gigatons per year, Gt/yr, with melting <0) for the three ice-shelf formations within
360 the Weddell regional domain (cf. Figs. 1 & 3): EWIS (cyan), FRIS (red) and Larsen (green). At
361 initialization, regional interactive model melt rates increase drastically from the prescribed melt
362 of the global parent simulation (GO7, cf. section 2.2, indicated by triangles at the time-axis origin
363 in Fig. 2a), by about 150 Gt/yr for each EWIS and Larsen, and by more than 250 Gt/yr for FRIS,
364 within the first 5 days of the simulation (note 45-day running means are displayed in Fig. 2a).

365 This is followed by a slower adjustment phase in which regional integral melt rates decrease back
366 from their initial peak and which lasts about one year (Fig. 2a).

367 For each ice-shelf formation, substantial variability in model basal melt is observed on a range of
368 timescales, including clear variability from year to year, between seasons, and on higher eddy and
369 and tidal frequencies. A sample of the latter are indicated in Fig. 2a by standard-deviation
370 envelopes around the 45-day running means (note that overall similar results are obtained varying
371 the running mean length between 30 and 60 days). These standard deviations include variability
372 at all frequencies between 5 and 45 days, which contains the imprint of tidal-driven variability
373 due to aliasing of tides. While we note that the full tidal-driven melt variability is expected to be
374 larger than indicated by the standard deviations in Fig. 2a, we do not aim here at documenting its
375 exact amplitude, which would require much higher frequency outputs and is beyond the scope of
376 the main objective here.

377 For the chosen set of parameters in Eq. (1), model integral ice-shelf melt rates are close to
378 available recent observational estimates (displayed with their respective errorbars on the rhs of
379 Fig. 2a). In particular time-average net model melt rates (shown in Table 1, using the last 5-year
380 average of the 8-year reference simulation, i.e. 1993–1997) fall within the bracket of the *Rignot*
381 *et al.* (2013) observational estimate (obtained for the late 2000s, see circles on rhs of Fig. 2a) for
382 each of the three Weddell ice-shelf formations considered, if slightly on the high end for EWIS.
383 For FRIS, model average net melt rates (159 ± 28 Gt/yr, where the minus sign, indicating net
384 melting, i.e. a negative basal mass balance, in Figures and Tables, is omitted here and as follows
385 in the text) are close to both the *Rignot et al.* (2013, 155 ± 43 Gt/yr) and the *Moholdt et al.* (2014,
386 124 ± 66 Gt/yr) estimates, and higher than the steady-state estimate by *Depoorter et al.* (2013) of
387 50 ± 40 Gt/yr, the latter being much lower than the other recent non-steady state estimates.

388 Integral FRIS model average melt rates are also consistent with a recent estimate based on noble
389 gases and thereby independent of those by glaciological studies (at 177 ± 95 Gt/yr; *Huhn et al.*,
390 2018).

391 Fig. 2b compares timeseries of ice-shelf melt freshwater fluxes for the Weddell domain total

392 (pink) to the freshwater released by icebergs melting throughout the domain (black). Note the
393 latter is imposed here as freshwater flux obtained from the Lagrangian iceberg model of the
394 boundary forcing simulation and thus in the absence of tides and at lower ocean resolution (see
395 section 2.2). In comparison to ice-shelf melt, iceberg melt shows larger seasonal and interannual
396 variability, whereas ice-shelf melt shows larger high-frequency variability (standard deviation
397 envelopes) due to the presence of tides. In the time average (Table 1), modelled Weddell total net
398 basal melting (257 ± 63 Gt/yr) is slightly smaller than domain-integral iceberg melting (275 ± 220
399 Gt/yr).

400 The integral melt rates discussed so far arise from a spatially heterogenous pattern featuring
401 regions of large and opposing basal melting and freezing (Fig. 3a). For the shallower ice shelves
402 of EWIS and Larsen, refreezing is generally small. Here enhanced melting at average melt rates
403 of typically 2-4 m/yr water equivalent (m.w.e/yr in the following) occurs near the ice fronts. For
404 FRIS enhanced basal melting with peaks above 7 m.w.e/yr occurs near the grounding lines of
405 several icestreams, as well as to the south of Berkner Island ($81^\circ\text{S}/55^\circ\text{W}$) and near the eastern
406 Ronne icefront ($77.5^\circ\text{S}/50^\circ\text{W}$). Refreezing is pronounced and occurs to the north of Korff and
407 Henry ice rises and rumples under central Ronne ($79^\circ\text{S}/62^\circ\text{W}$), as well as in association to the
408 outflows on the western sides of both Filchner and Ronne Troughs ($78.5^\circ\text{S}/43^\circ\text{W}$ and
409 $75.5^\circ\text{S}/63^\circ\text{W}$ respectively). This pattern of simulated FRIS basal melting and freezing agrees
410 very well, in both sign and magnitudes, with observational estimates by *Rignot et al.* (2013) &
411 *Moholdt et al.* (2014), mapped for comparison in Fig. 3c&d. Whereas in these estimates
412 refreezing on the western margin of Ronne is only estimated to occur deeper in the cavity, rather
413 than reaching all the way to the icefront as simulated here, this is also where observational
414 errorbars are amongst the largest (e.g. *Moholdt et al.*, 2014, their Fig. 10). Timeseries of basal
415 melting and freezing, integrated separately over the regions where and when they occur, roughly
416 mirror each other in their variability (not shown), suggesting a tight physical link of refreezing
417 responding to melting without much timelag as part of the cavity mass overturning circulation /
418 ice pump. Whereas integral refreezing is small under the small shallow ice shelves (typically \leq
419 6% of the melting), it amounts to more than a third of the melt underneath FRIS in the

420 time-average (Table 1). The modelled FRIS glacial-ice/meltwater mass overturning rate (with
421 freezing of 89 ± 10 Gt/yr and melting of 248 ± 35 Gt/yr, and thus a freeze/melt ratio of 36 %) is
422 thereby very close to its observational estimate by *Moholdt et al.* (2014, of 77 ± 25 Gt/yr basal
423 accumulation vs 201 ± 41 Gt/yr ablation, and thus a ratio of 38 %). Comparison to previous
424 modelling estimates is provided in the discussion.

425 **Sea ice** Weddell domain integral sea-ice concentration (Fig. 4, black) shows a large seasonal
426 cycle, of about 3.5×10^6 km² in both area and extent, throughout the 8-year reference experiment
427 timeseries. This is superimposed on a smaller-amplitude year-to-year variability, predominately
428 in the spring-through-fall ice cover, with little variability on subseasonal timescales. Simulated
429 integral sea-ice area compares overall well to observations (green curves), capturing the expected
430 annual freezing over of the entire model domain, as well as on average the timings of the onsets
431 of seasonal refreezing and melting, however in most years (but 1994 & 1997) model sea-ice
432 coverage shows a low bias at the summer-time minimum. This corresponds to a low bias in
433 model permanent sea ice, where in observations sea ice consolidates and thickens over
434 southwestern margins of the gyre against the Antarctic peninsula – a known bias of CORE-2
435 forced simulations (e.g. *Downes et al.*, 2015). Indeed, CORE-2 features too weak cold-air
436 outbreaks, and instead generally too strong westerlies crossing the Antarctic peninsula, apportioning
437 too warm air from the west, overly melting sea ice and advecting it eastward, with an especially
438 strong signal in summer. Such biases are typically mediated in higher-resolution
439 reanalysis-forced simulations, for example based on ERA-I (e.g. *Dinniman et al.*, 2015).

440 **Water masses** An overview of simulated water masses is obtained by a 5 year (1993-1997)
441 model volume census in conservative temperature (Θ)-absolute salinity (S_A) space for ice-shelf
442 cavities, as well as several regional subdomains in the open ocean.

443 On the southwestern Weddell Sea open-ocean continental shelf outside the FRIS cavity (shaded
444 white on map in Fig. 5a), the model reference simulation features four major water mass poles
445 (Fig. 5b): the relatively warm and salty (modified) warm deep water ((M)WDW), transiting along

446 the MWDW mixing lines into the fresh pole of winter waters (WW) at the surface freezing point.
447 Sea-ice production and atmospheric cooling on the wide continental shelf form the third major
448 pole, high salinity shelf waters (HSSW), characterized by S_A in excess of 34.8 g kg^{-1} reaching
449 peaks above 35.1 g kg^{-1} . Note these are locally formed and not present on the eastern shelves,
450 where salinities remain below 34.8 g kg^{-1} (see Suppl. Fig. S1). Their interaction with FRIS basal
451 melting at depth produces the freshened ice-shelf waters (ISW) with temperatures below the
452 surface freezing point which, leaving the cavity, form the fourth major water mass pole on the
453 open-ocean continental shelf. Furthermore there are lighter summer-formed Antarctic surface
454 waters (AASW), which, while occupying large extents of the Θ - S_A space, represent only
455 comparatively small fractions of the continental shelf ocean volume.

456 The reference-experiment water masses inside the FRIS cavity are color shaded in Fig. 6 (EWIS
457 and Larsen cavity water masses are displayed and discussed in Suppl. Fig. S2 and
458 Suppl. Note S1). Underneath FRIS, waters above the surface freezing point represent only a
459 small fraction of the total cavity ocean volume. Here major water masses are shown to lie below
460 the surface freezing point along the various meltwater mixing lines describing the transformation
461 of HSSW into ISW through ice-shelf melting (also Gade lines; *Gade, 1979*), with Θ as low as
462 -3.2°C representing melting at the deepest icedrafts. Observed summer-time sub-ice-shelf
463 profile Θ - S_A from boreholes S1–S5 and F1–F4 situated in the Ronne cavity and to the south of
464 Berkner Island, as well as from FSW1–2, FSE1–2 and FNE1–3 situated in the Filchner cavity, as
465 mapped in Fig. 1, are displayed as black dots in comparison to the model's FRIS volumetric
466 water mass census in Fig. 6. Observed FRIS cavity Θ - S_A properties are found to coincide with
467 the major pole of FRIS ISW water masses simulated by the reference experiment. Overall the
468 agreement between available borehole observations and model reference simulation is striking.
469 In particular, modelled Θ - S_A reveal the observed presence of three poles of source water-mass
470 types which enter the cavity with temperatures near the surface freezing point but differing
471 salinities as witnesses of different regimes of HSSW production along the ice front. At lower
472 temperatures, ISW collapses onto one pole elongated along the meltwater mixing line. For the
473 lowest observed Θ , borehole profiles coincide with the fresher end of the modelled pole of ISW

474 (in that temperature range). While the sparse observational coverage prevents a complete view of
475 real-world water masses in the FRIS cavity, we note that sub-glacial runoff, not included in the
476 present reference simulation, provides an additional source of freshening in the real world and
477 may imply a somewhat salty bias in the model (e.g. *Huhn et al.*, 2018).

478 Outside the FRIS cavity, observed Θ - S_A from all available open-ocean hydrographic profiles
479 (cf. section 2.3, mapped in Fig. 5a by green dots for profiles within the southwestern
480 continental-shelf region considered in Fig. 5b) are superimposed on the colored model volumetric
481 Θ - S_A contours in Fig. 5b by gray dots. As within the cavity, also on the open-ocean southwestern
482 Weddell shelf, in-situ observed water-mass characteristics are strikingly consistent with those
483 modelled. We note a bias towards slightly too fresh and cold modelled WDW, which is somewhat
484 more pronounced on the continental shelves in front of FRIS, and EWIS, than in the deep open
485 gyre (see Suppl. Fig. S1). For the interest of brevity we do not further describe the water-mass
486 characteristics in all other regions of the model here, but point the interested reader to
487 Suppl. Fig. S1 & Note S1. We note however that regional changes in water-mass characteristics
488 associated with the production, sinking, and entrainment of bottom waters over the southwestern
489 Weddell continental shelf and slope are consistent with observed profiles.

490 Corresponding maps of time-average simulated Θ - S_A at the sea floor are displayed in Fig. 7a,b
491 (note that, to compare with preponderantly summer-time observational campaigns, February
492 conditions are shown here). They reveal, at regional scale, warm and salty deep and bottom
493 waters of the gyre that stand in stark contrast to the cold continental-shelf waters, the latter
494 typically much closer to or even colder than the surface freezing temperature. At the
495 continental-shelf scale, the eastern Weddell Sea continental shelf adjoining EWIS distinguishes
496 itself as being much fresher than the shelves of the southwestern Weddell Sea adjoining FRIS and
497 Larsen ice shelves (consistent with their classification as fresh versus dense shelves respectively,
498 e.g. *Thompson et al.*, 2018, and references therein). The fresh eastern-shelf waters are advected
499 southwestward following the coast onto the shelf east of the Filchner Trough. With admixture of
500 MWDW accessing the shelf west of Brunt ice shelf, they form a warm and fresh shelf water mass
501 reaching all the way to the eastern Filchner icefront in this 5-year February average. Several other

502 intrusions of MWDW crossing the shelf break deep onto the southwestern continental shelf are
503 clearly visible in the bottom temperature field (Fig. 7a), including in the Central Trough west of
504 Berkner bank, reaching all the way to the icefront (Fig. 7e). In the FRIS cavity, bottom waters are
505 below the surface freezing point everywhere, decreasing below -3°C near the deepest grounding
506 lines. The saltiest waters enter the cavity in the Ronne Trough freshening on their various
507 pathways through FRIS through admixture of ice-shelf melt waters, a transformation also clearly
508 visible in a section below the Filchner ice shelf (Fig. 7c,d – along the green line in Fig. 7b). On
509 the continental shelf just north of the icefront (Fig. 7e,f – section along the cyan line in Fig. 7b),
510 waters below the surface freezing point fill the entire Ronne and Filchner Troughs in this 5-year
511 February average, spilling above the depth of the icedraft, and with the coolest outflowing ISW at
512 this time-of-year located on the eastern side of the Ronne and western side of the Filchner
513 Troughs (Fig. 7e). The outflowing potentially-supercooled ISW, mixing with continental-shelf
514 HSSW, fills the entire Filchner Trough reaching the continental shelf break around 74.5°S
515 (Fig. 7a,c). A signature of the overflow of such dense shelf waters (ISW+HSSW) out of the
516 Filchner Trough may be indicated by a marked gradient on the continental slope downstream of
517 the Filchner sill (Fig. 7a), where warm and salty bottom waters transition westward towards
518 much colder waters. In the vertical, ISW are overlaid by warmer waters (MWDW & WW),
519 forming the characteristic V-shape at the continental shelf break (Fig. 7c,d). This V-shaped layer
520 of MWDW and WW is itself capped by a warm and fresh seasonal thermocline of AASW,
521 extending all the way from the open gyre across the FRIS shelf to the icefront (Fig. 7c,d).
522 Interestingly, all the main features present in Fig. 7 and discussed in this paragraph are in very
523 good agreement with past observational findings: MWDW intrusion pathways onto the
524 continental shelf (e.g. *Nicholls et al.*, 2009; *Darelius et al.*, 2014, 2016; *Ryan et al.*, 2017);
525 HSSW formation on the western Ronne continental shelf and its pathways within the cavity (e.g.
526 *Nicholls et al.*, 2009); the V-shaped layer of MWDW and WW at the continental shelf break (e.g.
527 *Thompson et al.*, 2018); ISW on the continental shelf and its outflow from the northwestern
528 Filchner sill (*Foldvik*, 2004; *Daae et al.*, 2018); and the overall water-mass distribution just
529 outside the Filchner-Ronne ice-shelf front (*Mackensen*, 2001). The two main departures of

530 simulated water-mass regional and vertical repartitions are: (i) the peak HSSW produced in front
531 of western Ronne may not be quite salty enough (e.g. in one given year's summer, February
532 1995, in the deepest Ronne Trough, below 500 m, just outside the icefront, simulated absolute
533 salinities exceed 34.925 g/kg, versus those observed exceed 35 g/kg, *Mackensen, 2001*, yet
534 observational sparsity precludes a full assessment), and (ii) the surface stratification in front of
535 FRIS seems too strong with a thus too thin, too warm and fresh surface layer (*Mackensen, 2001*).

536 **Circulation** The time-average model barotropic streamfunction, mapped in Fig. 8a, shows a
537 cyclonic gyre circulation with total transports exceeding 60 Sv (a local recirculation at the gyre
538 core near 5°W, 61°S increases this to a peak of 75 Sv). About 10–15 Sv of this are carried by
539 coastal and slope branches. Another strong internal branch, which transports about 25 Sv roughly
540 following the 4000 m isobath and not extending beyond 71°S, is associated with the WDW gyre
541 recirculation that is not directly involved in water-mass modifications induced by intense cooling
542 and sea-ice formation on the Weddell Sea continental shelves. These simulated circulations are
543 consistent with the overall observational picture (e.g. *Schröder and Fahrbach, 1999; Armitage*
544 *et al., 2018; Reeve et al., 2019*). In particular, the total flow in the southern branch of the gyre
545 entering the domain from the east is, if somewhat stronger, roughly consistent with the estimates
546 at 0°E by *Schröder and Fahrbach (1999, 66 Sv)* and *Klatt et al. (2005, 56±8 Sv)*. The model
547 annual-mean transport associated with the slope front, across $\approx 17^\circ\text{W}$, is comparable to the
548 early-autumn 1995 estimate by *Heywood et al. (1998)* of 14 Sv. Near the northwestern gyre exit,
549 the model boundary transport (between Joinville Island and 47.5°W), made up of several
550 branches including coastal, slope and Weddell fronts, also compares well to the estimate by
551 *Thompson and Heywood (2008)*, totalling to 46±8Sv (with the remainder of the gyre transport
552 occurring further east). From this predominant cyclonic gyre circulation, cyclonic incursions spin
553 off onto the widening southwestern Weddell continental shelf. In the time-average, these are in
554 particular the coastal current incursion onto the continental shelf east of the Filchner Trough, as
555 well as onto the Ronne continental shelf west of Berkner Bank (Central Trough) and less
556 far-reaching just east of Ronne Trough. Besides the major cyclonic circulation, a

557 smaller-amplitude anticyclonic barotropic circulation (peaks at 0.6 Sv) links the FRIS cavity and
558 the continental shelf. It represents the overall inflow into FRIS in the Ronne Trough and outflow
559 through the Filchner icefront with several branches/pathways, a direct outflow at the western
560 edge of the Ronne Trough, a deep flow to the south of Korff, and in part Henry ice rises, rejoining
561 a shallower more direct branch of the anticyclonic cavity circulation on its path around the south
562 of Berkner Island. The latter is suggested here to also be alimented by direct inflow west of
563 Berkner Island originating from the Berkner shelf.

564 The horizontal barotropic circulation is associated with a vertical circulation that can be
565 described by the time-average meridional overturning streamfunction (Fig. 8b). Two
566 counter-rotating cells dominate the overturning in latitude-pressure space: a counter-clockwise
567 rotating cell mostly representative of the open-ocean gyre, and a clockwise circulation mostly
568 representative of the ice-shelf cavity. The two cells overlap on the continental shelf. The
569 open-gyre cell advects about 0.8 Sv of thermocline waters (WDW) from the central gyre
570 southwards until about 72.5°S, from where a part starts to downwell and another part reaches the
571 average shelf break further south. Above, the open-gyre cell advects an additional 0.8 Sv of
572 lighter surface waters southward onto the continental shelf (half reaching midshelf, a quarter all
573 the way to the ice front), where they downwell as dense shelf waters (HSSW), and are eventually
574 exported down the continental slope northward. About 0.2–0.4 Sv of these dense continental
575 shelf waters are not exported directly but first become associated with the cavity circulation cell,
576 which carries them southward into the FRIS cavity, where they upwell (transformed into ISW)
577 and move back northward following the average ice draft. The cavity cell extends northward
578 beyond the ice front up to the continental shelf break, where cavity waters (ISW) outflow and
579 sink at a rate of a few $O(0.1)$ Sv. Note that this meridional-depth average hides wide regional
580 variability of a system, which is far from being zonally symmetric. So, while helpful in providing
581 an overall idea of the vertical circulation, the meridional overturning streamfunction cannot be
582 interpreted in terms of water-mass circulation and production.

583 Here, the peak 0.4 Sv FRIS volume overturning suggested by Fig. 8b, together with the net
584 production of ice-shelf meltwater of 4.9 mSv in the reference experiment (160 Gt/yr, Table 1),

585 suggest a meltwater composition of the FRIS outflow, and a meltwater production efficiency of
586 the FRIS ice and meltwater pump (used here to signify the volume of meltwater produced per
587 volume of overturning transport into the cavity) of around 1 %. This is in agreement with the
588 observational estimate of the glacial meltwater content in outflowing ISW (*Naveira Garabato*
589 *et al.*, 2016; *Huhn et al.*, 2018).

590 **4 Role of tides**

591 In the following we examine the role of tides in setting the properties of the Weddell oceanic
592 circulation, and of the reference experiment discussed above (in section 3), in particular in terms
593 of ice-shelf melting, its mechanisms and attendant changes in large-scale circulation and water
594 mass characteristics. To do so we analyze a sensitivity experiment, different from the reference
595 experiment only by the absence of tidal boundary forcing, with otherwise the exact same
596 initialization, boundary and atmospheric forcing, as well as the same 8-year time span of
597 simulation 1990–1997. Discussion in comparison to previous modelling studies, in particular
598 *Makinson et al.* (2011) and *Mueller et al.* (2018), follows in section 5.

599 **Tidal signal in Weddell ice-shelf melt** Timeseries of integral ice-shelf melt in the perturbation
600 experiment without tidal boundary forcing are displayed in Fig. 2a & b (by black-edged colored
601 symbols and black-dashed colored lines) in comparison to those of the tide-forced reference
602 experiment discussed above (by plain colored symbols and lines). In the absence of tides, melt
603 rates adjust from the prescribed melt rates of the forcing simulation (triangles) to the interactions
604 with the regional higher-resolution ocean during a quick initial spinup, not discernible from
605 seasonal and year-to-year variability beyond the first 5 days. In contrast, in the simulation with
606 tides the major adjustment occurs over several months during the first year.

607 In the 5-year time-average (over simulated years 4–8, Table 1) the presence of tides enhances net
608 basal melting in the Weddell cavities by about 50 % for both EWIS ($\Delta = 19$ Gt/yr) and FRIS (54
609 Gt/yr), and almost 80 % for Larsen cavities (19 Gt/yr) compared to the state without tides. In the

610 total over the southwestern Weddell region, tides raise ice-shelf meltwater fluxes by almost 100
611 Gt/yr, and thereby to a rate that roughly equals the integral iceberg melt. Note that here icebergs'
612 trajectories and melting are prescribed and do not interact with tides. How tides may impact
613 calving, drift and melting of Antarctic icebergs thereby remains to be assessed in dedicated future
614 modelling configurations.

615 These net tidal signals are the residual of tidal impacts on both melting and freezing. The 5-year
616 time-average maps (Fig. 3a,b) show that tides enhance the intensity of not only melting but also
617 freezing where both occur, and thereby make simulated melt rate patterns closer to observations
618 (Fig. 3c,d). The overall spatial pattern of melting versus freezing regions is similar in presence
619 and in absence of tides. Regions with a change in sign of the basal mass balance from melting to
620 freezing are rare, and include remote parts of Riiser-Larsen and Brunt-Stancomb ice shelves, as
621 well as the refreezing patch north of the central-Ronne ice rises, which besides intensifying,
622 somewhat expands spatially north- and westward in presence of tides. Regarding amplitude, the
623 largest increases occur, for freezing, under central Ronne, and for melting, in the central and
624 eastern Ronne and Filchner Troughs, over the shallow-water regions near the eastern Ronne
625 icefront and in South Channel (south of the central Ronne ice rises, cf. Fig. 1), as well as in
626 several spots in EWIS and Larsen located mostly, but not exclusively, near their icefronts.

627 Integrating the two poles of melting and freezing separately over the cavities (see Table 1) shows
628 that refreezing in presence of tides remains small underneath Larsen, and starts to become
629 noticeable, increasing from 2 % to 6 % of the integral melt, under EWIS. For FRIS, a more than
630 110 Gt/yr increase in melting, representing nearly a doubling (from 136 to 248 Gt/yr), is
631 significantly offset by 60 Gt/yr of additional refreezing, making the latter almost triple (from 31
632 to 89 Gt/yr). Thus, whereas refreezing is overall small under the small shallow ice shelves and
633 remains so with tides, underneath FRIS, where it already amounts to almost a quarter of the melt
634 (23 %) in absence of tides, it enhances to more than a third (36 %) with tides. By augmenting the
635 net FRIS meltwater production from 3.25 to nearly 5 mSv (from 105 to 159 Gt/yr), tides thus
636 significantly enhance the FRIS ice pump and make it closer to what is known from observations
637 (as of the *Moholdt et al.*, 2014, estimate of a FRIS freeze/melt ratio of $77/201 \approx 38\%$). (Note the

638 term ice pump is used here, and in what follows, to signify the melt-driven cavity volume
639 overturning circulation, with more melting driving a more vigorous cavity circulation and, in
640 direct association, enhanced refreezing along its upward branches, e.g. *Lewis and Perkin*, 1986;
641 *Gwyther et al.*, 2015).

642 Net Weddell ice-shelf basal melt rates (Fig. 2a,b) show important year-to-year variations, both in
643 absence and in presence of tides. Whereas any significant differences are not fully
644 characterizable from the present set of 8-year simulations, interannual variations are somewhat
645 dampened in the absence of tides, which is especially visible for the largest simulated anomaly
646 (the warm summer and melting peak of 1993). It is interesting to note here that, while some
647 year-to-year anomalies differ between the three ice shelves considered here, others covary. This
648 points to dynamical processes of different scale at play simultaneously, including both
649 small-scale regional specificity and large-scale driving of the sub-ice shelf circulations and the
650 tidal impact thereon. Importantly, tidal circulations are shown to enhance the integral basal melt
651 water fluxes in all cavities of the Weddell Sea and in all months of the year. In the absence of
652 tides, Weddell basal melt water fluxes are seen to be typically elevated in summer through
653 autumn with a peak in early autumn. The tidal enhancement has itself a marked seasonal
654 distribution, which tends to occur in this season of peak melting (Fig. 2a). Under FRIS, tides
655 induce an additional secondary peak in melting in late winter (Fig. 2a).

656 **Mechanisms of the tidal modulation in ice-shelf melt** The mechanisms through which the
657 presence of tides can impact ice-shelf basal melting are changes in the kinetic energy of ocean
658 currents at the ice-shelf base and changes in the thermohaline properties of the water masses
659 interacting with the ice shelves, the latter possibly arising through a combination of all,
660 tide-driven changes in shelfbreak exchanges, shelf mixing, shelf-cavity exchanges, cavity
661 pathways and cavity mixing and melting (e.g. *Padman et al.*, 2018; *Jourdain et al.*, 2019).

662 The signature of tides in the time-average top boundary layer (tbl) current speed c_{tbl} is displayed
663 in Fig. 9c (note, c_{tbl} is obtained as the time-mean of the instantaneous absolute velocity thus
664 including the speed of the time-mean flow as well as that of high-frequency tidal and other

665 currents; it pertains to the tbl in the cavities and for completeness to the top 30 m flow in the open
666 ocean). It clearly reveals a significant tide-driven increase in kinetic energy at the base of the ice
667 drafts in all cavities of the Weddell Sea. Whereas, in the absence of tides, tbl current speeds c_{tbl}
668 in excess of 10 cm/s occur primarily along the paths of time-mean currents (Fig. 9b), in the
669 presence of tides (Fig. 9a) surface-layer speed is enhanced over the bulk of the Weddell
670 continental shelves (by about 2-10 cm/s), with peak enhancements (≥ 20 cm/s) on the shallower
671 part of the continental shelves just adjoining the shelfbreak, such as to the east and to the west of
672 the Filchner Trough, and along the southwestern corner of the Ronne shelf edge along 72°S. At
673 the ice-shelf bases, tides enhance the tbl speed throughout. Peak enhancements of 20–30 cm/s are
674 reached in several locations of EWIS, and the southern Larsen C. Underneath FRIS they occur
675 within Ronne and Filchner Troughs, in several bottlenecks of the Ronne tributaries, to the north
676 of the Ronne ice rises, and, in particular, in South Channel and along the shallow eastern Ronne
677 ice front. The regions of enhanced tbl tidal speeds simulated here thereby correspond, both on the
678 continental shelves and in the cavities, to the regions of maximum tidal currents (as estimated by
679 barotropic tide models e.g. *Robertson et al.*, 1998; *Padman et al.*, 2018), and are expected to be
680 highly sensitive to cavity geometry (e.g. *Padman et al.*, 2018; *Rosier et al.*, 2018). As a local
681 exception, a decrease of speed surrounding Berkner Island likely reflects a tidal broadening of the
682 Berkner Island boundary current. Importantly, this tidal enhancement of the time-average tbl
683 speed (Δc_{tbl} , Fig. 9c) does overall *not* arise from a residual time-mean circulation change
684 induced by tides (changes in c_{tbl} associated with the 5-year mean tidally-induced circulation
685 alternate in sign and typically do not exceed 5 cm/s in magnitude – not shown). Instead, it
686 predominately reflects the systematic and wide-spread enhanced energy of time-varying tbl
687 currents (zero-mean – not shown) throughout the cavities.

688 Fig. 9f shows overall cooling throughout the tbl of the Weddell cavities. In the absence of tides
689 (Fig. 9e), the tbl of the shallow EWIS and Larsen cavities is typically warmer than the local
690 freezing point (T_b) by 0.1-0.2 K. Also underneath FRIS the time-average reveals a surface tbl
691 above freezing without tides, with peak temperature differences to the local freezing point
692 ($T_{tbl} - T_b$) in excess of +0.2 K, in particular near the grounding lines. In presence of tides

693 (Fig. 9d) regional thermal contrasts reduce yielding a more homogeneous cavity tbl, closer to the
 694 local freezing point throughout. Warm regions, in which tides drive enhanced melting, cool with
 695 tides, and cold regions, with enhanced refreezing, warm with tides. The time-average change in
 696 cavity thermal *driving* thus predominately reflects the *response* of the tbl to the tidal input of
 697 mechanical energy.

698 The respective contributions of these dynamical (Fig. 9a-c) and thermohaline (Fig. 9d-f) tidal
 699 changes in the cavity tbl to the tide-induced change ($\Delta \equiv \text{tide} - \text{notide}$) in ice-shelf basal
 700 melting \mathcal{M} (< 0 melt) can be quantified as (cf. *Jourdain et al.*, 2019)

$$\begin{aligned}
 -\Delta\mathcal{M} &\approx \Delta(Q_{tbl}/L_{fi}) = \Delta(\alpha c_{tbl}(T_{tbl} - T_b)) \\
 &= \alpha\Delta c_{tbl}(T_{tbl} - T_b)|_{no} + \alpha c_{tbl}|_{no}\Delta(T_{tbl} - T_b) + \alpha\Delta c_{tbl}\Delta(T_{tbl} - T_b) \\
 &= -(\Delta\mathcal{M}_{DYN} + \Delta\mathcal{M}_{THERM} + \Delta\mathcal{M}_{COV}), \tag{2}
 \end{aligned}$$

701 where we have used that Q_{tbl} , given by eq. (1), is directly proportional to \mathcal{M} when neglecting
 702 heat diffusion into the ice, we introduced $\alpha \equiv (\rho_w c_{pw} St_T) / L_{fi}$ with L_{fi} the latent heat of
 703 fusion of ice, and $|_{no}$ denotes the cavity state without tides. In the resulting decomposition of
 704 tide-driven basal melt $\Delta\mathcal{M}$, the dynamical contribution $\Delta\mathcal{M}_{DYN}$ reflects a stronger ocean-ice
 705 heat flux due to tide-induced tbl kinetic energy (Δc_{tbl} , Fig. 9c). Thermodynamical
 706 ($\Delta\mathcal{M}_{THERM}$) and covariational ($\Delta\mathcal{M}_{COV}$) components, in turn, reflect tide-induced changes in
 707 tbl temperature ($\Delta(T_{tbl} - T_b)$, Fig. 9f), which either arise in phase with the tidal enhancement of
 708 tbl kinetic energy ($\Delta\mathcal{M}_{COV}$), or not ($\Delta\mathcal{M}_{THERM}$). Here $\Delta\mathcal{M}_{THERM}$ is expected to dominate
 709 where tide-induced warming occurs unrelated to dynamical changes (such as when advected
 710 from the continental shelves) or for a vigorous cavity circulation that quickly evacuates cold
 711 meltwater from regions of melting (such as in warm cavities *Jourdain et al.*, 2019). $\Delta\mathcal{M}_{COV}$, to
 712 the contrary, may dominate where increased tbl kinetic energy produces cold meltwater that is
 713 not quickly exported away (such as in more sluggish cold cavities).

714 Timeseries of the tidal signal in ice-shelf melt $\Delta\mathcal{M}(t)$ for the Weddell domain, that is the
 715 difference between tidal and non-tidal integral melting from Fig. 2b, are displayed in Fig. 9g,
 716 together with the timeseries of the three mechanistic contributions given by Eq. (2). Here the

717 approximation in Eq. (2) is confirmed to be minor, with the sum of the three contributions (gray
 718 dash in Fig. 9g) reproducing near exactly the total change of melt (black in Fig. 9g). The
 719 tide-driven melt ($\Delta\mathcal{M}$, black, < 0 more melt) oscillates, after the initial spin up, seasonally and
 720 from year to year, at ≈ 100 Gt/yr in the Weddell total, and is shown to be solely generated by an
 721 about 8 times larger dynamical enhancement ($\Delta\mathcal{M}_{DYN}$, green), which drives all the seasonal
 722 and interannual variations of the tidal enhancement, and in absence of thermal responses of the
 723 cavity would reflect the net change seen. The tbl however responds thermally to this mechanical
 724 driving, mostly at zero time lag (captured by the covariational term, $\Delta\mathcal{M}_{COV}$, cyan), cooling at
 725 times of enhanced melting and cold ISW production. Together with a smaller non-instantaneous
 726 thermal response ($\Delta\mathcal{M}_{THERM}$, blue) of ≈ 100 Gt/yr, the net thermal response ($\Delta\mathcal{M}_{COV} +$
 727 $\Delta\mathcal{M}_{THERM}$) thereby substantially limits the net effect of tides enhancing Weddell ice-shelf
 728 melt. This decomposition holds for each individual of the three ice-shelf formations, each
 729 showing a very similar repartition into the different mechanistic contributions to the tide-driven
 730 basal melting, as further quantified for the 5-year time average in Table 2: The tidal increase in
 731 kinetic energy at the ice-shelf bases alone enhances melt by around 130 Gt/yr for EWIS and
 732 Larsen, and nearly 500 Gt/yr for FRIS. Thermal adjustments, 82.5-85 % of which occur at
 733 zero-lag (covariational), limit the tide-driven enhancement to the simulated ≈ 20 Gt/yr for the
 734 former and 54 Gt/yr for the latter.

735 **Further implications of tides in the Weddell Sea** Here we briefly focus on integral, barotropic
 736 & overturning, metrics of tide-driven changes in the mean flow, which can reflect both dynamical
 737 (e.g. *Bessières et al.*, 2008; *Padman et al.*, 2018; *Jourdain et al.*, 2019) and melt-induced (e.g.
 738 *Jourdain et al.*, 2019) tidal residual circulations, the latter driven by the tidal imprints on ice-shelf
 739 melting discussed above. Figure 8 maps these tidal signals (in panels e & f respectively) and also
 740 compares their time-average states with (a & b) and without (c & d) tides.

741 Regarding the barotropic circulation (Fig. 8a,c,e), tide-induced changes in the open gyre are
 742 small relative to the total transport (typically less than 5 %), with the strength of the cyclonic gyre
 743 excursions onto the continental shelves, to the east of Filchner and Ronne Troughs as well as onto

744 the Larsen shelf, somewhat reduced in presence of tides. The most significant signal, in terms of
745 relative change, is an enhancement of the predominately anticyclonic barotropic gyre circulation
746 within the FRIS cavity, by around 0.1 Sv of 0.2–0.4 Sv. In association, several local changes in
747 pathways are found, especially in the Ronne local cross ice-front circulation, which becomes
748 more anticyclonic, as well as in remote parts of Ronne and in the Filchner Trough.

749 As to the overturning circulation (Fig. 8b,d,f), tides induce changes in both of its cells, the
750 counterclockwise open-gyre and the clockwise cavity cell. The open-gyre cell weakens at depth,
751 by 0.4 of 2 Sv, or by 20 %. In the upper 500 m, local changes suggest it reaches closer to the
752 surface around 72°S (the latitude of enhanced southwestern shelf break tidal kinetic energy,
753 cf. Fig. 9a,c) and its extension onto the southwestern continental shelf somewhat weakens at
754 mid-shelf, while its furthest extension towards the southern ice fronts instead somewhat
755 intensifies. The most pronounced relative change, again, occurs in the sub-ice-shelf cavities: here
756 the cavity cell, associated to the transformation of HSSW into ISW through ice-shelf melt
757 interactions, shows a marked tide-induced increase almost doubling from peaks at 0.2 to 0.4 Sv
758 near the FRIS ice fronts, and from 0.1 to 0.2 Sv in more remote parts of the FRIS cavity,
759 reflecting a tide-induced enhancement of the FRIS ice pump. It is worth noting here that this
760 enhancement of the mean FRIS circulation is not strong enough to account for the simulated
761 enhancement of tbc flow speed associated with enhanced melting, and the mechanical energy
762 input is instead provided by energized time-variable (tidal) currents. While these integral
763 overturning metrics are useful to highlight such important basic mechanistic understanding, they
764 are unsatisfactory to document water-mass circulations, and so we do not further investigate this
765 aspect here. We note however that by nearly doubling the circulation in the cavity, and therefore
766 substantially enhancing the associated production of ISW, tides may have a direct impact on the
767 characteristics of bottom waters.

768 Here a first insight into such tidal impacts on water-mass characteristics in the Weddell ice-shelf
769 cavities and the adjacent continental shelves is gained from volumetric water-mass differences in
770 the Θ - S_A space (cf. Suppl. Fig. S3, and for a more detailed description Suppl. Note S1). Whereas
771 overall changes, especially in the open ocean, are weak and fragmented, the several more

772 pronounced and systematic tide-induced water-mass changes that stand out are indicative of
773 being a response to the tide-driven enhancement of ice-shelf melting and the associated cooling
774 and freshening. In cavity, shelf and open ocean, these are the following. (I) In the FRIS cavity,
775 the volume of waters above the surface freezing point is reduced and that of sub-freezing point
776 waters increased (also for EWIS), with overall (besides multiple localized poles of changes in the
777 Θ - S_A space) more volume in the fresher varieties of FRIS ISW in presence of tides. (II) The ISW
778 outflowing from FRIS onto the continental shelf show a clear signature of decrease in its saltiest
779 and increase in its freshest varieties. (Other changes on the continental shelf are a reduction of
780 saltiest waters along the dense-shelf water/MWDW mixing line, as well as overall more
781 shelf-produced HSSW, and, for the EWIS open-shelf, more outflowing sub-surface freezing point
782 waters). (III) In the off-shelf deep open ocean, a sole marked water-mass change induced by the
783 presence of tides stands out: a freshening of the densest (deep and bottom) water masses in the
784 western Weddell Sea. This water mass is produced locally through southwestern Weddell shelf
785 and cavity processes (it is not present in the eastern gyre), and the changes in its properties thus
786 reflect the signature of tides on Weddell ocean–ice-shelf interactions and the outflowing shelf
787 water-mass properties.

788 The tide-induced enhancement of ISW outflow (Figs. 2, 8f, S3) may be furthermore anticipated
789 to impact the continental shelf stratification and accordingly its sea-ice cover. Such tide-driven
790 changes in sea-ice volume in the southwestern Weddell Sea are quantified in Fig. 10. Sea-ice
791 volume on the southwestern Weddell Sea continental shelf is enhanced in all months of the year
792 (Fig. 10b), by about 5 %. Given that tides do not systematically impact area coverage (see Fig. 4
793 for Weddell integrals, mapped concentrations also do not show any marked large-scale changes –
794 not shown), this reflects a tidal impact on the thickness of sea ice (actual thickness, not ice
795 volume per unit area). Thickness increases by on average 5 cm on top of a seasonally-varying
796 continental shelf sea-ice thickness of 1–2 m, and the peak enhancement of 10 cm occurs during
797 the season of thickest ice in summer (mapped in Fig. 10a). Rather than reflecting a local tidal
798 signature on for example shelf mixing, here we suggest that this tidal imprint on continental-shelf
799 sea-ice volume passes via tide–ocean–ice-shelf interactions in Weddell cavities through the tidal

800 enhancement of glacial freshwater production and export enhancing shelf stratification. This
801 mechanism is consistent with previous circumpolar models reporting Weddell sea-ice thickening
802 in response to enhanced meltwater forcing (e.g. *Merino et al.*, 2018). In particular, results by
803 *Hellmer* (2004) of a 100 % meltwater addition, from opening ice-shelf cavities, leading to a 20
804 cm sea-ice thickening (about 10–20 %) are comparable, also in magnitude, to present results of a
805 50 % meltwater enhancement, here tide-driven, leading to a 10 cm thickening (about 5-10%).

806 **5 Discussion**

807 We have presented a first assessment of tide-induced changes in Weddell Sea ocean–ice-shelf
808 interactions in realistic simulations of the southwestern Weddell Sea gyre – continental-shelf –
809 cavity system forced with interannually-varying atmospheric and ocean-boundary conditions.
810 Previous modelling assessments of the tidal impact on FRIS ice-shelf melting have considered
811 solutions (i) to homogeneous and constant FRIS continental-shelf forcing (*Mueller et al.*, 2018)
812 in cold and warm shelf conditions, and (ii) to climatological ocean forcing in combination with
813 spatially-heterogeneous seasonally-modulated FRIS continental-shelf surface thermohaline
814 relaxation (*Makinson et al.*, 2011). Both used ocean-only simulations (no sea ice) across
815 relatively-focussed regional domains (not extending much beyond the continental shelf break and
816 not including other Weddell Sea cavities) at horizontal resolutions coarser than in the present
817 study (2.5 – 4 times).
818 In comparison to *Makinson et al.* (2011)’s results, simulations analyzed here suggest a smaller
819 *relative* net enhancement of FRIS basal melt fluxes in response to tides. In magnitude,
820 tide-induced changes of the FRIS ice pump however closely agree between the two studies, in
821 terms of both net melt and basal melt-to-freeze ratios (*Makinson*’s FRIS net increase of 50 Gt/yr
822 reflecting a ratio of changes in freezing versus melting, $\Delta_{\text{freeze}}/\Delta_{\text{melt}}$, of 53.3 %, compared to
823 this study’s net increase of 54 Gt/yr reflecting a ratio of 51.6 %). In the present study, these
824 tide-induced changes act on a larger non-tidal net melt (105 Gt/yr here versus 41 Gt/yr in
825 *Makinson et al.*, 2011), and this implies that here the net tide-driven change reflects a 50 %

826 increase (compared to the 100 % increase obtained by Makinson), and similarly that here
827 changes in refreezing and melting are ≤ 3 -fold and ≤ 2 -fold (compared to 6-fold and 3-fold in
828 their study). The spatial patterns of non-tidal and tidal FRIS melt rates are also overall similar
829 between the two studies.

830 In contrast, results of this study differ from the homogeneous-shelf condition-forced FRIS
831 ocean-tide modelling results by *Mueller et al. (2018)* in terms of both the non-tidal background
832 and the tide-driven changes, in all the net signal, the ice pump as reflected by the
833 refreezing-to-melt ratio, as well as the spatial distribution of freezing and melting regions. This is
834 to be expected given that the thermohaline-forced west-to-east anticyclonic cavity circulation is
835 absent (in *Mueller et al., 2018*), a circulation suggested by available observations, simulated in
836 this study as well as in *Makinson et al. (2011)*, and likely to represent the dominant contribution
837 to the cavity mean flow especially within Ronne & Filchner Troughs, where tidal residuals are
838 estimated to contribute less than 10% (*Makinson and Nicholls, 1999*). Under the uniform
839 thermohaline shelf forcing of *Mueller et al. (2018)* and the associated east-to-west cyclonic
840 cavity circulation, the absence of Ronne inflows leads to widespread refreezing over the majority
841 of the Ronne sector of FRIS. The latter is a likely reason for the near-compensating tidal
842 increases in melting and refreezing, and the near-zero tidal signal in the net FRIS basal mass
843 balance, obtained by *Mueller et al. (2018)*. In this context, it is interesting to note that several of
844 the tide-induced changes in FRIS circulation patterns obtained here in full tide–ocean–ice-shelf
845 interactions agree with the tide-only results obtained in FRIS barotropic experiments by
846 *Makinson and Nicholls (1999)*, including tide-driven changes in the eastern Ronne cross-icefront
847 transports, an enhancement of the anticyclonic circulation around Berkner Island as well as a
848 reduction in eastward transport through South Channel (Fig. 8e). Similarly, there are also
849 similarities to the constant thermohaline and tide-forced results of *Mueller et al. (2018)*, and they
850 include enhanced melting in the shallow South Channel and western Ronne icefront, mid-Ronne
851 enhanced freezing and a predominately mechanical driving of the integral tide-induced changes.
852 This study provides an assessment of the dynamical and thermodynamical mechanisms at play in
853 ice-shelf cavities' top boundary layer associated with tide-induced melt, for FRIS and other

854 Weddell Sea ice-shelf cavities, and under realistic thermohaline forcing scenarios. For the
855 Amundsen Sea, a previous realistic-simulation assessment (*Jourdain et al., 2019*) finds a
856 mechanistically comparable tidal impact: in both cases, top boundary-layer thermodynamic
857 adjustment diminishes the mechanical tidal driving of increased melt. Interestingly, in the cold
858 Weddell Sea cavities assessed here, in which thermal forcing of ice-shelf melting is weak in the
859 first place, the thermal adjustment response is found to be more pronounced (compensating
860 dynamical tidal forcing to about 85-90 %) than in the warm Amundsen Sea cavities (about 40 %).
861 Results presented here, consistent with *Jourdain et al. (2019)*, suggest that the nature of the tidal
862 signature on melt is primarily local, associated with enhanced kinetic energy at the icedraft, with
863 the thermal top boundary layer signature reflecting the response to this mechanical enhancement
864 of ocean-ice-shelf interactions, rather than being induced advectively by tidal residual
865 circulations. Here it is worth to note that the heat lost from the top boundary layer through
866 increased kinetic energy locally enhancing melt, must be reflected by tbl cooling or, especially in
867 the case of a cold cavity characterized by limited heat available for melting ice, be replenished by
868 heat entering the cavity. A detailed analysis of cavity heat budgets and their responses to tides
869 remains a topic of interest for future study in this regard. Here we simply note that any required
870 enhanced cross-icefront heat transport may reflect local changes in water-mass properties and
871 advective pathways induced by tides independently of their interaction with ice-shelf melting, but
872 it may also arise as a direct response to the tide-induced melting via melt-induced changes in
873 cavity overturning and in the water masses exported, such as those reported here.

874 Critical-latitude effects and internal tide generation through tidal interactions with topography
875 and ocean stratification, as well as its seasonal modulation, are all to some extent included in the
876 presented simulations and known to contribute importantly to shape distributions of tidal energy
877 (e.g. *Robertson, 2001a,b; Makinson et al., 2006; Semper and Darelus, 2017*). While this study
878 has focussed on assessing the integrated tidal impact, including these effects, on ocean-ice-shelf
879 interactions and associated water mass pathways and transformations, further differentiating their
880 individual contributions would provide interesting refinements to the present analysis. Analysis
881 presented here, noting sensitivity to uncertainties, discretization and errors in cavity geometry

882 (e.g. *Rosier et al.*, 2018; *Padman et al.*, 2018, and references therein), also points towards the role
883 of tides in the renewal of FRIS water masses (cf. also Suppl. Note S1 for further discussion of
884 this aspect). In this regard it is worth to mention that a faster cavity flushing timescale would
885 diminish the effective ocean thermal adjustment, and therefore strengthen the tidal enhancements
886 of both ice-shelf melting and freezing where they occur, and thus, given the respective weights of
887 the latter established here, strengthen the tide-driven net melting and meltwater production. This
888 suggests a critical role of both cavity vertical mixing and the top boundary layer heat budget in
889 the ice-shelf melt response to tides (e.g. *Gwyther et al.*, 2015), which points to the latter's likely
890 sensitivity to the parameterizations employed, and in particular to the bulk parameterization of
891 ice-shelf melt rates (and its known limitations, e.g. *Asay-Davis et al.*, 2017). Important
892 outstanding advances thus include improved observational constraints for or elimination of
893 remaining largely-unconstrained parameters in the bulk melt parameterization (such as Stanton
894 numbers and especially boundary layer depth, e.g. *Jourdain et al.*, 2017), and the representations
895 of upward sedimenting frazil-ice platelets freezing throughout the water column (e.g. *Bombosch
896 and Jenkins*, 1995; *Smedsrud and Jenkins*, 2004; *Galton-Fenzi et al.*, 2012) as well as dynamical
897 feedbacks between tides, ocean circulation, ice-shelf melting and water & ice depths as
898 accessible in ocean-ice-sheet coupled simulations (e.g. *Mueller et al.*, 2018; *Reese et al.*, 2018;
899 *Timmermann and Goeller*, 2017).

900 This study provides a carefully-designed high-resolution regional-model configuration capturing
901 relevant high-latitude ocean-cryosphere processes whose simulations given their high
902 computational cost are necessarily rather short. Given the expected long exchange timescale of
903 the large FRIS cavity (based on the simulated circulations in Fig. 8 we obtain a rough scaling
904 estimate of ≈ 13 years; observational studies using oceanographic and isotopic constraints have
905 estimated between 4–10 years, cf. *Nicholls and Østerhus*, 2004; *Huhn et al.*, 2018), here we
906 therefore rely on a spin-up of cavity water masses obtained in the global boundary-forcing
907 simulation GO7, which as described in section 2 uses an overall very similar set up as the
908 regional simulations here except its lower resolution and absence of explicit tides. Departing with
909 thus already physically-consistent states including in cavities, regional-simulation integral

910 ice-shelf melt (Fig. 2), as well as integral temperature and salinity of both open and cavity ocean
911 (not shown), do not show any adjustment processes that can be differentiated from year-to-year
912 variability beyond their first 1-2 years. In contrast, initialization from available ocean reanalyses,
913 which to date typically end in walls either at the entrance of cavities or at fixed latitudes cutting
914 through continental shelves, necessarily implies extrapolation of attendant biased
915 temperature-salinity properties far into cavities and thus demands long regional-simulation
916 spin-ups before physically consistent water mass properties can be established. This, together
917 with the risk that initiating with too warm continental shelf temperatures may trigger irreversibly
918 warm states in the FRIS system (*Hellmer et al., 2017*), has motivated our choice to not constrain
919 the present reference experiment with reanalyses. We hope that regional high-resolution results
920 produced and presented here may provide a useful benchmark, as further model developments
921 will allow to partially parameterize the now known-to-be important high-latitude cryospheric
922 processes, in cavities and shelves and in interaction with tides, in lower-resolution climate-type
923 model simulations thereby allowing for long and regionally-unconstrained integrations. Model
924 intercomparison studies (building on e.g. *Naughten et al., 2018b*) between different types of
925 high-resolution regional, lower-resolution resolved-cavity circumpolar, and global simulations
926 with partially-parameterized processes, will thereby be of immense value to further
927 understanding of relevant processes and systems' variability.

928 **6 Conclusions**

929 The present study assesses for the first time the impact of tides on Weddell Sea ocean–ice-shelf
930 interactions in a coupled ocean–sea-ice model with realistic interannually-varying boundary and
931 atmospheric forcing, fully representing two-way exchanges between the southwestern Weddell
932 Sea open-ocean gyre, its continental shelves and ice-shelf cavities, and at higher spatial
933 resolution than in previous ocean-only simulations with steady (*Mueller et al., 2018*) or idealized
934 climatological forcing (*Makinson et al., 2011*).

935 In comparison to the observed Weddell Sea system, as far as known, overall circulation and water

936 mass features in all its three components, deep, shelf and cavity oceans, are well represented in
937 the simulations assessed here. A key bias is a too warm summer surface layer with a too strongly
938 stratified summer seasonal thermocline and associated low bias in summer sea-ice coverage,
939 especially on the southwestern Weddell shelf. Nonetheless key dense shelf water masses
940 produced by local atmosphere–ocean–tide–sea-ice–ice-shelf interactions are overall well
941 represented. Warm Deep Water intrusions onto the continental shelves are also simulated where
942 observed (with a seasonal and year-to-year variability that has not been fully presented here),
943 although possibly slightly too modified (cooled & also freshened) compared to observations as
944 they reach the remote FRIS continental shelf and its icefronts.

945 Key results are as follows:

- 946 • Net basal meltwater fluxes show substantial variability in all Weddell Sea ice-shelf cavities
947 on all simulated timescales, including tidal, seasonal and year-to-year periods, where the
948 latter are presented here for the first time in simulations of the region including tides and
949 are suggested slightly amplified in presence of tides.
- 950 • Tides drive an increase in time-average net ice-shelf basal melt by 50 % in EWIS (+19
951 Gt/yr melting) and FRIS (+54 Gt/yr) and 80 % in Larsen cavities (+19 Gt/yr). Whereas
952 refreezing is small under the small shallow ice shelves, it amounts to a quarter of melting
953 underneath FRIS, and increases to more than a third with tides, a ratio comparable to
954 observations (*Moholdt et al.*, 2014). This reflects a tide-driven intensification of the
955 existing pattern of basal melting and refreezing (with integral increases in freezing
956 reaching about 50% those in melting) and of the underlying FRIS sub-ice-shelf circulation,
957 in terms of both the primarily anticyclonic main barotropic gyre circulation (although there
958 is local structure, e.g. a reduction through South Channel reflecting a more cyclonic
959 circulation there) and, especially, the cavity volume overturning circulation.
- 960 • Tide-induced changes in FRIS ocean–ice-shelf interactions and ice pump closely match the
961 coarser-resolution isopycnal-model idealized-forcing one-year-average results obtained by
962 the only previous FRIS ocean-tide modelling study including a representation of the

963 (observed) continental-shelf thermal gradients (*Makinson et al.*, 2011). Simulated FRIS
964 basal melt patterns are also overall similar, with the main difference being refreezing
965 outflows reaching the mid-Ronne icefront in the former study, where simulations here
966 instead feature a narrow western-boundary outflow feature in the Ronne Trough and melt
967 conditions all along the Ronne icefront to the east, implying somewhat different circulation
968 regimes in the Ronne cavity.

- 969 • The tide-induced increase in Weddell ice-shelf melting is associated, both in its mean and
970 its time-variability, with a tidal enhancement in the kinetic energy of time-varying ocean
971 currents at the ice base, which mechanically raise the heat exchanged across the
972 background (non-tidal) thermal gradient between the ocean–sub-ice-shelf boundary layer
973 and the ice shelf. Tide-induced changes in this thermal gradient act to significantly damp
974 (rather than enhance) the mechanically-induced tidal enhancement of melting (by 85-90 %,
975 depending on cavity), since they reflect on average a response to the latter, with cooling in
976 regions of enhanced melting and warming in regions of enhanced refreezing – a response
977 which is to more than 80 % instantaneous, the remainder being not zero-lag correlated,
978 inducing a slight lag to this thermal damping. The net tidal enhancement of melting may be
979 sustained by a transition to a cooler cavity state and by tide-driven increases in heat
980 transport into the cavity. The latter, while a detailed assessment of cavity heat budgets’
981 sensitivity to tides is subject to future study, may be achieved, both, independently of melt
982 through local changes of properties and their advective pathways into the cavity, and, in
983 response to tide-induced melting through the melt-induced circulation and water-mass
984 changes described in this study.

- 985 • Changes attendant to the tidal enhancement of FRIS basal meltwater production and cavity
986 circulation include a slight shift towards a fresher ISW exported onto the southwestern
987 Weddell Sea continental shelf, and a freshening and lightening of the densest deep water
988 masses in the open gyre produced in response to the downwelling dense shelf waters, as
989 well as a year-round increase in southwestern shelf sea-ice volume (by 5–10 %). Further
990 refinement of diagnostics is necessary to quantify more reliably the integral tidal

991 contributions to the water-mass circulation in terms of transformation rates and exchanges
992 across shelfbreak and icefronts.

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998 code is freely available at www.nemo-ocean.eu. Here we use the development branch
999 dev_v3_6_STABLE_r6506_AGRIF_LIM3. Model customizations and parameter choices required
1000 to run the experiments presented in this paper are provided at [https://doi.org/10.5281/](https://doi.org/10.5281/zenodo.3384022)
1001 zenodo.3384022. Observational data used in this study are publicly available as indicated in
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1350 **List of Tables**

1351	1	Southwestern Weddell Sea ice-shelf melt: integral statistics, and role of tides.	58
1352	2	Mechanisms of the tidal impact on Weddell Sea ice-shelf melting.	59

1353 **Figure Captions**

1354	1	Regional model configuration in its geographical context (pink dashes: computational domain, pink solid lines: locations of model open boundaries, outside of which ocean regions are masked in the configuration, but not in the present figure). Ice fronts are contoured in dark gray. Colors show model ocean water column depth h_{water} , which, in ice-shelf cavities, is given by the difference between model bathymetry (light gray contours) and ice draft. Figure axes are parallel to the model grid, longitude and latitudes are dotted, and model isotropic horizontal resolution (km) contoured in purple. For context, observed climatological March & September sea-ice edges are also contoured (white dotted & dashed). Place names indicated are Joinville Island (JI), South Shetland Islands (sSI), Bransfield Strait (BS), South Orkney islands (sOI); Fimbul (Fi), Jelbart (Je), Atka (At), Ekstroem (Ek), Quar (Q), Riiser-Larsen (Ri-L), Brunt/Stancomb-Wills (B/St-W), and Larsen G, F, E, D, C ice shelves, as well as the the SCAR inlet & Seal Nunatak glaciers (remnants of Larsen A & B: rB & rA). Furthermore within the Filchner-Ronne ice-shelf (FRIS) cavity: Berkner Island (BI), South Channel (SCh), Korff & Henry ice rises (Kir & Hir) and Doake rumples (Dr). FRIS sub-ice-shelf mooring locations are indicated by '+'s.	60
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1371 **2 Weddell Sea integral ice-shelf melt timeseries** $\mathcal{M}(t)$ (<0 melt). **a** Integrated by
 1372 ice-shelf formation (colored as labelled): in the reference experiment with tides
 1373 (color spheres give 45-day running means & lines their one-standard-deviation
 1374 envelopes of 5-day-average timeseries), contrasted to the experiment without tides
 1375 (black-edge color spheres & darker colored standard-deviation lines, here often
 1376 coinciding with symbol size). Symbols on the rhs show corresponding available
 1377 long-term mean observational estimates, by *Rignot et al.* (2013, R13, \circ), *De-*
 1378 *poorter et al.* (2013, D13, $*$), *Moholdt et al.* (2014, M14, \square) and *Huhn et al.*
 1379 (2018, H18, \diamond). Triangles at the timeaxis origin give the constant prescribed melt
 1380 of the forcing simulation. **b** As **a**, but for the Weddell domain total, and in compar-
 1381 ison to the interannually-varying melt of drifting icebergs integrated over the
 1382 model domain. 61

1383 **3 Ice-shelf basal melt maps.** Maps of the last 5-year (1993-1997) average basal
 1384 melt \mathcal{M} (<0 melt) in the reference experiment with tides (**a**), contrasted to the
 1385 perturbation experiment without tides (**b**). For comparison, available long-term
 1386 mean observational estimates by *Rignot et al.* (2013, R13) & *Moholdt et al.* (2014,
 1387 M14) are shown in (**c**) & (**d**), the latter available for FRIS only. 62

1388 **4 Sea-ice concentration.** Timeseries of sea-ice area in the regional configuration
 1389 reference experiment (as 5-day averages, black; corresponding no-tide perturba-
 1390 tion experiment sea-ice area timeseries, shown in purple, differ little), in compar-
 1391 ison to monthly-mean microwave observations (green, bootstrap version 2), both
 1392 integrated over the Weddell ocean model domain. (The sea-ice area metric shown
 1393 here integrates the fractional sea-ice covered area of grid points for which area
 1394 concentration exceeds a 15 % threshold. For reference, dashed lines indicate the
 1395 domain's total ocean area in the two grids). 63

- 1396 **5 Open-ocean water masses.** Reference experiment 5-year average (years 4-8,
 1397 i.e. 1993-1997) volumetric Θ - S_A for the south-western Weddell continental shelf
 1398 (light-shaded region in **a**) displayed by colored contours in **b** (as log fraction of
 1399 the region's total ocean volume; gray dotted contours indicate σ_θ). All-time avail-
 1400 able in-situ observations collected in the region (locations mapped by green dots
 1401 in **a**) are superimposed on the regional Θ - S_A diagram in **b** by gray dots. Major
 1402 water-mass types are: modified warm deep water (MWDW), winter water (WW),
 1403 high-salinity shelf water (HSSW), ice-shelf water (ISW) and Antarctic surface wa-
 1404 ter (AASW). (See Supplement for corresponding simulated water-mass diagrams
 1405 in the other open-ocean regions and their comparison to the in-situ observations
 1406 collected there, at the locations mapped by black dots in **a**). 64
- 1407 **6 FRIS cavity water masses.** Reference experiment 5-year average (years 4-8, i.e.
 1408 1993-1997) volumetric Θ - S_A for the FRIS cavity, shown as log fraction of the
 1409 cavity's total ocean volume (color shading). FRIS sub-ice-shelf CTD data from
 1410 mooring boreholes (mapped in Fig. 1, see text for details) are superimposed for
 1411 comparison (black crosses). (See Supplement for corresponding water-mass dia-
 1412 grams of EWIS and Larsen ice-shelf cavities.) 65

- 1413 **7 Illustration of the water mass structure in the reference experiment.** 5-year
 1414 average February (years 4-8, i.e. 1993-1997) conservative temperature Θ & ab-
 1415 solute salinity S_A : **a** & **b** at the sea floor, **c** & **d** along a vertical section crossing
 1416 from the open gyre in the north, across Filchner sill & Trough through the Filch-
 1417 ner cavity towards the Foundation ice stream grounding line (along the green line
 1418 in **b**), and **e** & **f** along a vertical section across the southwestern Weddell Sea
 1419 continental shelf just north of the FRIS ice front (along the cyan line in **b**, which
 1420 near-coincides with the ice front contoured in dark gray just to its south). Contours
 1421 in **a,b** show bathymetry (light gray) and ice fronts (dark gray), in **c-f** Θ isotherms
 1422 of -1.9 & -1.4 °C (solid & dashed black, the latter often barely visible as close to
 1423 the surface), and additionally highlight in **d,f** selected S_A contours (gray), and in
 1424 **e** the depth of the ice draft just to the south (white-gray). 66
- 1425 **8 Integral metrics of circulation.** **a,c,e** Barotropic, and **b,d,f** meridional overturn-
 1426 ing streamfunctions, for the years 4-8 average (1993-1997), both colored and con-
 1427 toured in Sv (dashed contours denote positive streamfunction values, and note the
 1428 reversed color scale in b,d,f vs a,c,e): **a,b** in presence of tides, and **c,d** in absence
 1429 of tides; **e,f** display the tidal signal itself ($\Delta \equiv$ tides – no tides). Additional con-
 1430 tours are in **a,c,e** 400, 1500 & 4000m isobaths (purple), icefronts (gray), as well
 1431 as latitude, longitude (dotted), and in **b,d,f** the weighted-average ice draft & sea
 1432 floor depths (gray line & black dots). 67

1433 **9 Mechanisms of the tidal impact on Weddell Sea ice-shelf melting.** Top panels
 1434 **(a,b,c)** map the sub-ice-shelf ocean boundary layer flow speed c_{tbl} (in m/s; in the
 1435 open ocean top 30 m speed is mapped for reference), panels below **(d,e,f)** the de-
 1436 parture of the sub-ice-shelf boundary layer ocean temperature from the freezing
 1437 point at the ice base (in K): on the lhs **(a,d)** in presence of tides, in the middle
 1438 **(b,e)** without tides, and on the rhs **(c,f)** the tidal signature Δ . All maps show
 1439 year 4–8 averages, and contour 500 & 1000 m isobaths for reference. Panel **g**
 1440 displays timeseries of the tide-driven Weddell ice-shelf melt $\Delta\mathcal{M}(t)$ (black, <0
 1441 melt), and reconstructs these, following Eq. (2), from dynamical (green), ther-
 1442 modynamical (blue) and covariational (cyan) mechanistic contributions, with the
 1443 sum of the three latter (gray dots) visibly coincident with the net change. (As in
 1444 Fig. 2, 45-day running means of 5-day averages and their one standard deviation
 1445 envelopes are shown.) 68

1446 **10 Impact of tides on southwestern Weddell sea-ice distribution,** estimated from
 1447 the year 4-8 5-year average of the reference versus no-tide simulations. **a** Map
 1448 of the tidal signal Δ in February sea-ice volume (divided by grid-cell area). **b**
 1449 5-yr average monthly timeseries of the southwestern Weddell continental shelf
 1450 (east of 25°W, south of 62.5°S, and inside the 1500 m isobath) integral sea-ice
 1451 volume (solid lines, in 10^{12} m³ – as reference the horizontal dotted line shows the
 1452 volume of a 1 m slab covering the whole region) and average sea-ice thickness
 1453 (actual thickness, not volume per area, dotted lines): with tides (dots), without
 1454 tides (circles), and the tidal signal Δ (triangles). 69

Table 1 | Southwestern Weddell Sea ice-shelf melt: integral statistics, and role of tides.^a

		EWIS	FRIS	LARSEN	Weddell
\mathcal{M}_{tide}	net	-55.1 ± 35.6	-159.2 ± 27.8	-42.9 ± 29.7	-257.2 ± 63.0^b
	melt	-58.8 ± 35.4	-247.9 ± 34.8	-43.8 ± 29.5	-350.5 ± 67.5
	freeze	3.6 ± 0.6	88.7 ± 9.9	0.9 ± 0.2	93.3 ± 9.9
\mathcal{M}_{notide}	net	-36.2 ± 24.5	-105.2 ± 15.1	-24.0 ± 12.5	-165.4 ± 32.3
	melt	-36.9 ± 24.4	-136.4 ± 17.0	-24.5 ± 12.4	-197.8 ± 33.4
	freeze	0.7 ± 0.2	31.2 ± 4.6	0.5 ± 0.2	32.4 ± 4.7
$\Delta\mathcal{M}$	net	-18.9 ± 13.5	-54.0 ± 17.0	-18.9 ± 18.3	-91.8 ± 38.6
	melt	-21.9 ± 13.5	-111.5 ± 21.0	-19.3 ± 18.3	152.6 ± 41.3
	freeze	3.0 ± 0.5	57.5 ± 7.0	0.4 ± 0.1	60.8 ± 7.1

Notes: ^a Year 4-8 (1993-1997) 5-year averages of integral ice-shelf melt \mathcal{M} (< 0 for melting), \pm one standard deviation of 45-day running means around the 5-year average, all in Gt/yr : in the reference experiment with tides, in the perturbation experiment without tides, and their difference $\Delta \equiv \text{tide} - \text{notide}$. For each, the first row gives the net integral, and the following two the integral obtained separately for regions of basal melting and basal freezing. ^b In comparison, the net calving iceberg melt freshwater flux integrated over the domain amounts to $-274.8 \pm 219.7 Gt/yr$.

Table 2 | Mechanisms of the tidal impact on Weddell Sea ice-shelf melting.^a

	EWIS	FRIS	LARSEN	Weddell
$\Delta\mathcal{M}$	-18.9 ± 13.5	-54.0 ± 17.0	-18.9 ± 18.3	-91.8 ± 38.6
$\Delta\mathcal{M}_{DYN}$	-132.8 ± 37.0	-494.1 ± 54.3	-132.8 ± 56.8	-759.6 ± 111.0
$\Delta\mathcal{M}_{THERM}$	17.0 ± 8.6	76.9 ± 9.5	17.0 ± 7.0	110.9 ± 16.6
$\Delta\mathcal{M}_{COV}$	96.3 ± 19.8	362.9 ± 40.4	96.4 ± 33.0	555.6 ± 69.7
<i>residual</i>	0.6 ± 0.1	0.3 ± 0.2	0.5 ± 0.2	1.4 ± 0.2

Notes: ^a The net tidal signature ($\Delta \equiv \text{tide} - \text{notide}$) in integral ice-shelf melt \mathcal{M} (< 0 melt) (row repeated from Table 1), and its dynamical $\Delta\mathcal{M}_{DYN}$, thermodynamical $\Delta\mathcal{M}_{THERM}$ and covariational $\Delta\mathcal{M}_{COV}$ contributions, as given by (2), as well as a *residual* (net minus sum of contributions). Otherwise as in Table 1, i.e. based on year 4–8 5-year averages (1993–1997), and in *Gt/yr*.

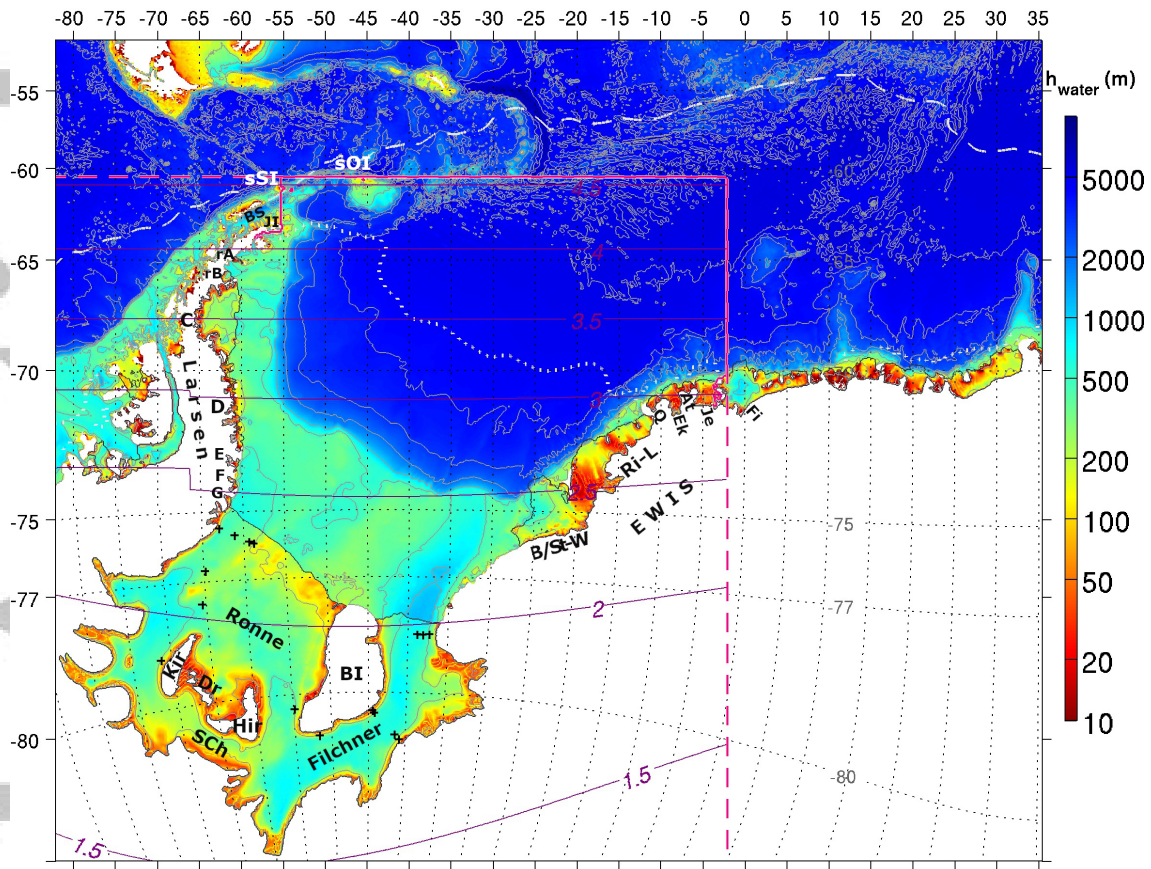


Figure 1 | Regional model configuration in its geographical context (pink dashes: computational domain, pink solid lines: locations of model open boundaries, outside of which ocean regions are masked in the configuration, but not in the present figure). Ice fronts are contoured in dark gray. Colors show model ocean water column depth h_{water} , which, in ice-shelf cavities, is given by the difference between model bathymetry (light gray contours) and ice draft. Figure axes are parallel to the model grid, longitude and latitudes are dotted, and model isotropic horizontal resolution (km) contoured in purple. For context, observed climatological March & September sea-ice edges are also contoured (white dotted & dashed). Place names indicated are Joinville Island (JI), South Shetland Islands (sSI), Bransfield Strait (BS), South Orkney islands (sOI); Fimbul (Fi), Jelbart (Je), Atka (At), Ekstroem (Ek), Quar (Q), Riiser-Larsen (Ri-L), Brunt/Stancomb-Wills (B/St-W), and Larsen G, F, E, D, C ice shelves, as well as the the SCAR inlet & Seal Nunatak glaciers (remnants of Larsen A & B: rB & rA). Furthermore within the Filchner-Ronne ice-shelf (FRIS) cavity: Berkner Island (BI), South Channel (Sch), Korff & Henry ice rises (Kir & Hir) and Doake rumples (Dr). FRIS sub-ice-shelf mooring locations are indicated by '+'s.

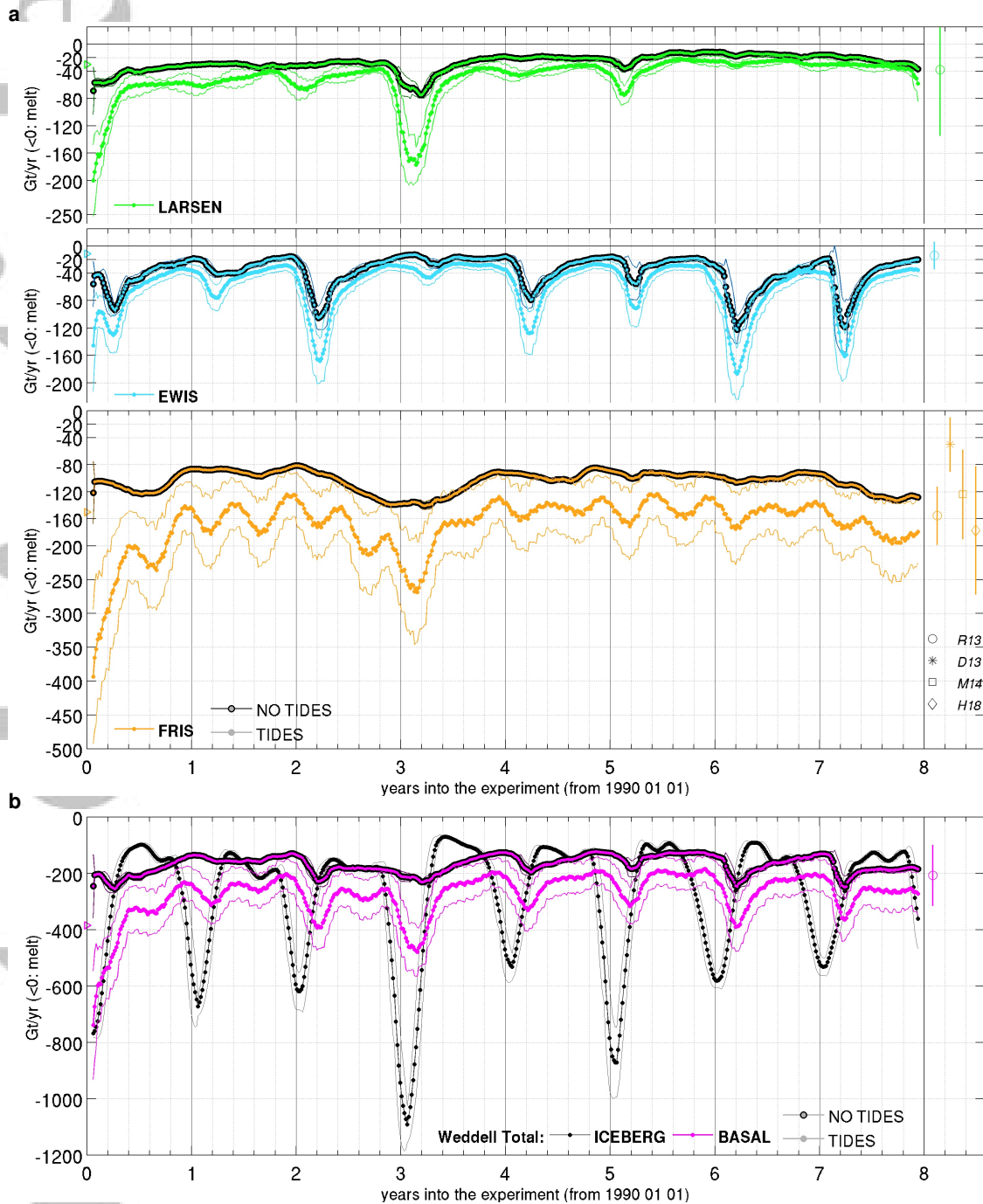


Figure 2 | Weddell Sea integral ice-shelf melt timeseries $\mathcal{M}(t)$ (<0 melt). **a** Integrated by ice-shelf formation (colored as labelled): in the reference experiment with tides (color spheres give 45-day running means & lines their one-standard-deviation envelopes of 5-day-average timeseries), contrasted to the experiment without tides (black-edge color spheres & darker colored standard-deviation lines, here often coinciding with symbol size). Symbols on the rhs show corresponding available long-term mean observational estimates, by *Rignot et al.* (2013, R13, \circ), *Depoorter et al.* (2013, D13, *), *Moholdt et al.* (2014, M14, \square) and *Huhn et al.* (2018, H18, \diamond). Triangles at the timeaxis origin give the constant prescribed melt of the forcing simulation. **b** As a but for the Weddell domain total, and in comparison to the interannually-varying melt of drifting icebergs

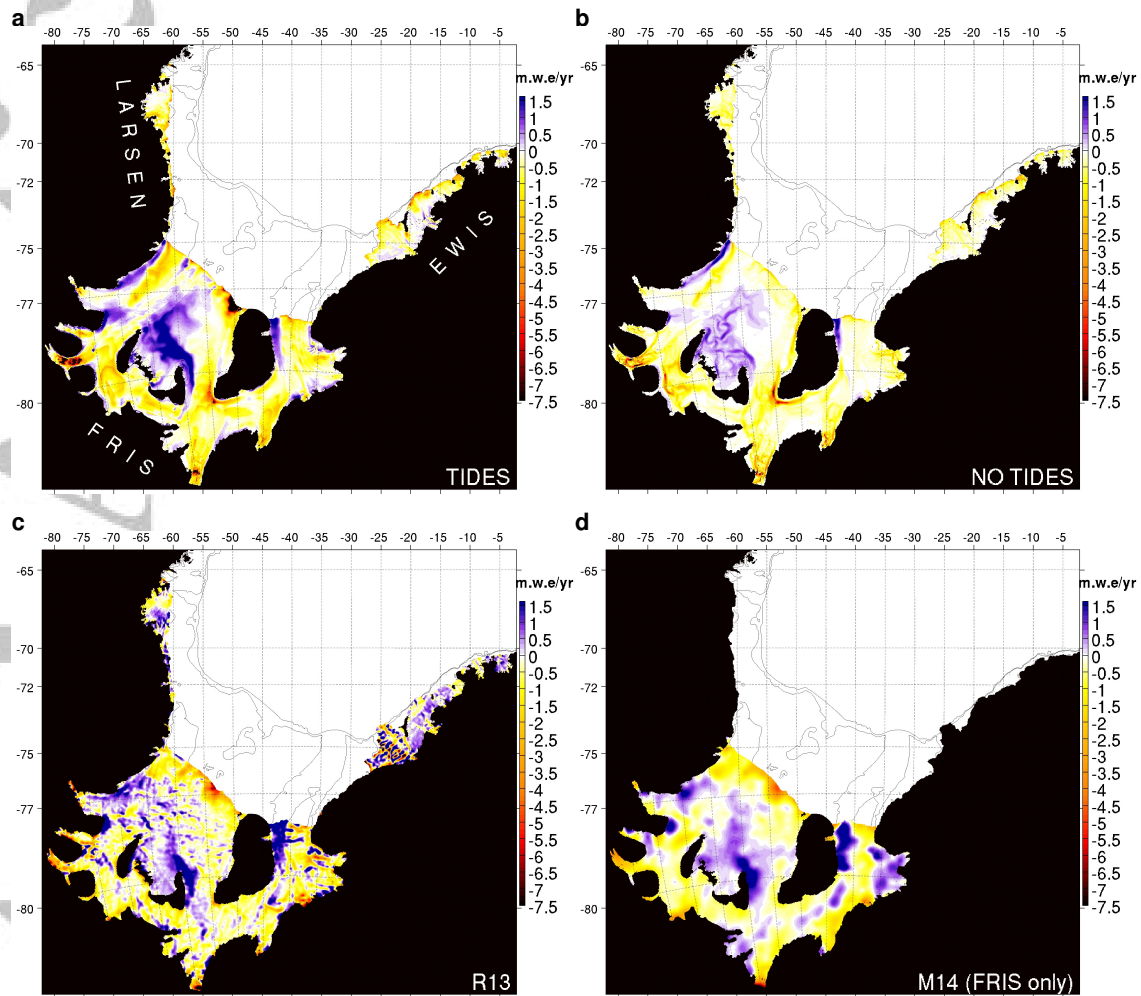


Figure 3 | Ice-shelf basal melt maps. Maps of the last 5-year (1993-1997) average basal melt \mathcal{M} (<0 melt) in the reference experiment with tides (a), contrasted to the perturbation experiment without tides (b). For comparison, available long-term mean observational estimates by *Rignot et al.* (2013, R13) & *Moholdt et al.* (2014, M14) are shown in (c) & (d), the latter available for FRIS only.

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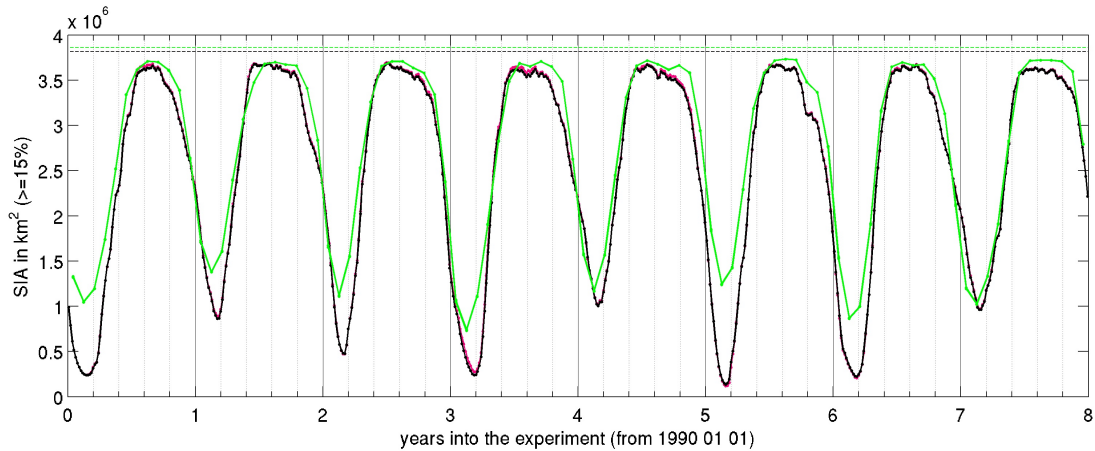


Figure 4 | Sea-ice concentration. Timeseries of sea-ice area in the regional configuration reference experiment (as 5-day averages, black; corresponding no-tide perturbation experiment sea-ice area timeseries, shown in purple, differ little), in comparison to monthly-mean microwave observations (green, bootstrap version 2), both integrated over the Weddell ocean model domain. (The sea-ice area metric shown here integrates the fractional sea-ice covered area of grid points for which area concentration exceeds a 15 % threshold. For reference, dashed lines indicate the domain's total ocean area in the two grids).

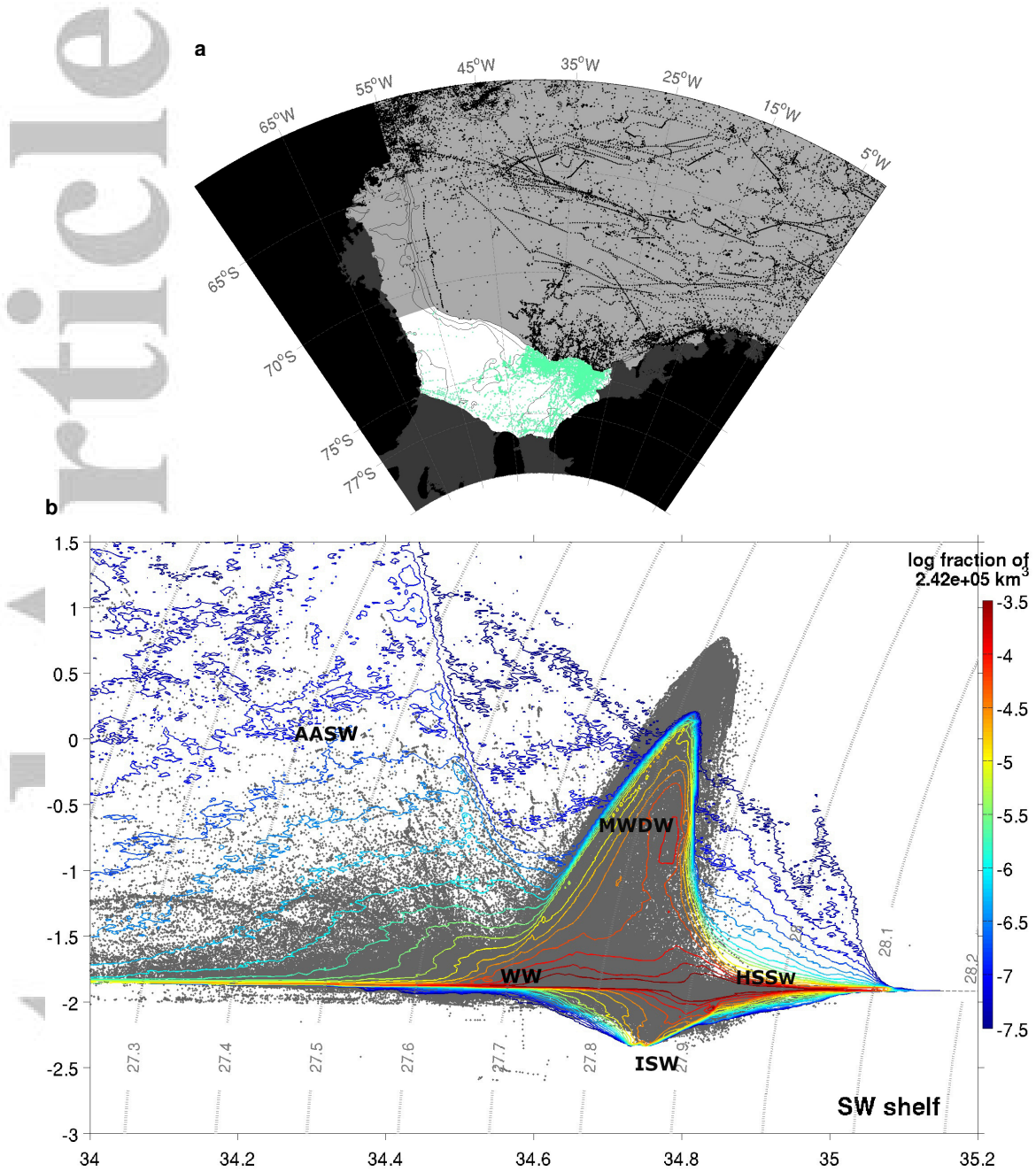


Figure 5 | Open-ocean water masses. Reference experiment 5-year average (years 4-8, i.e. 1993-1997) volumetric Θ - S_A for the south-western Weddell continental shelf (light-shaded region in **a**) displayed by colored contours in **b** (as log fraction of the region's total ocean volume; gray dotted contours indicate σ_0). All-time available in-situ observations collected in the region (locations mapped by green dots in **a**) are superimposed on the regional Θ - S_A diagram in **b** by gray dots.

Major water-mass types are: modified warm deep water (MWDW), winter water (WW), high-salinity shelf water (HSSW), ice-shelf water (ISW) and Antarctic surface water (AASW). (See Supplement for corresponding simulated water-mass diagrams in the other open-ocean regions and their comparison to the in-situ observations collected there, at the locations mapped by black dots in **a**).

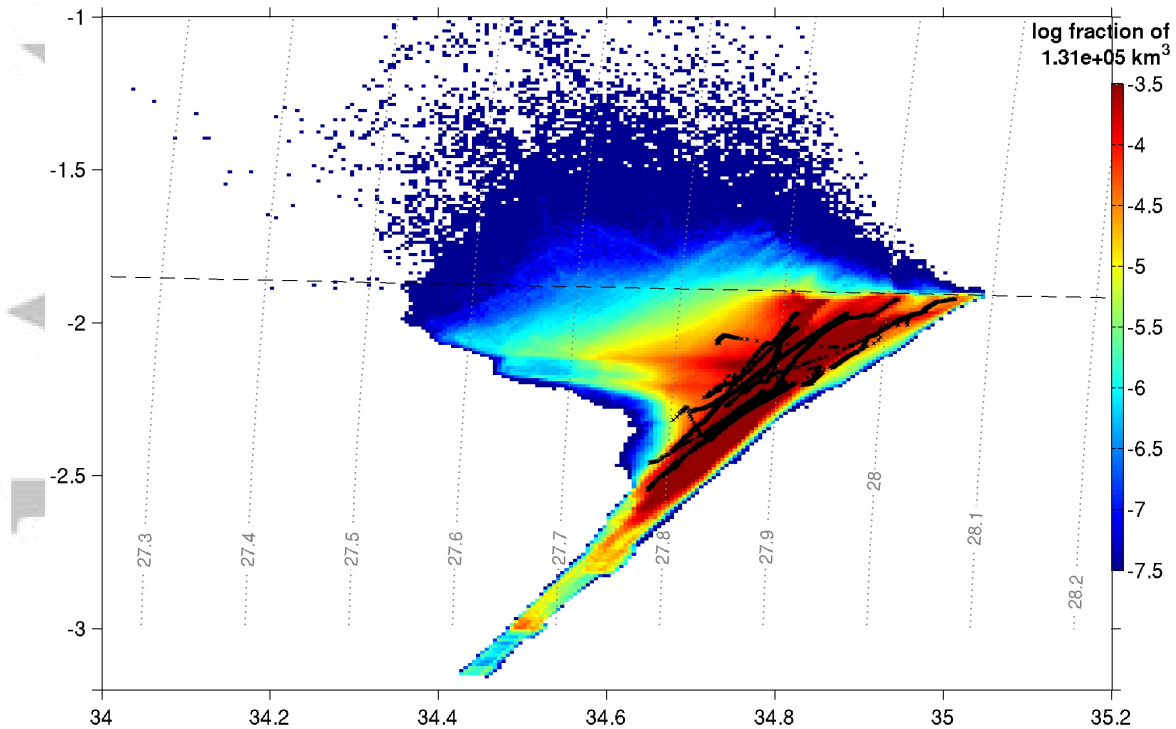


Figure 6 | FRIS cavity water masses. Reference experiment 5-year average (years 4-8, i.e. 1993–1997) volumetric Θ - S_A for the FRIS cavity, shown as log fraction of the cavity’s total ocean volume (color shading). FRIS sub-ice-shelf CTD data from mooring boreholes (mapped in Fig. 1, see text for details) are superimposed for comparison (black crosses). (See Supplement for corresponding water-mass diagrams of EWIS and Larsen ice-shelf cavities.)

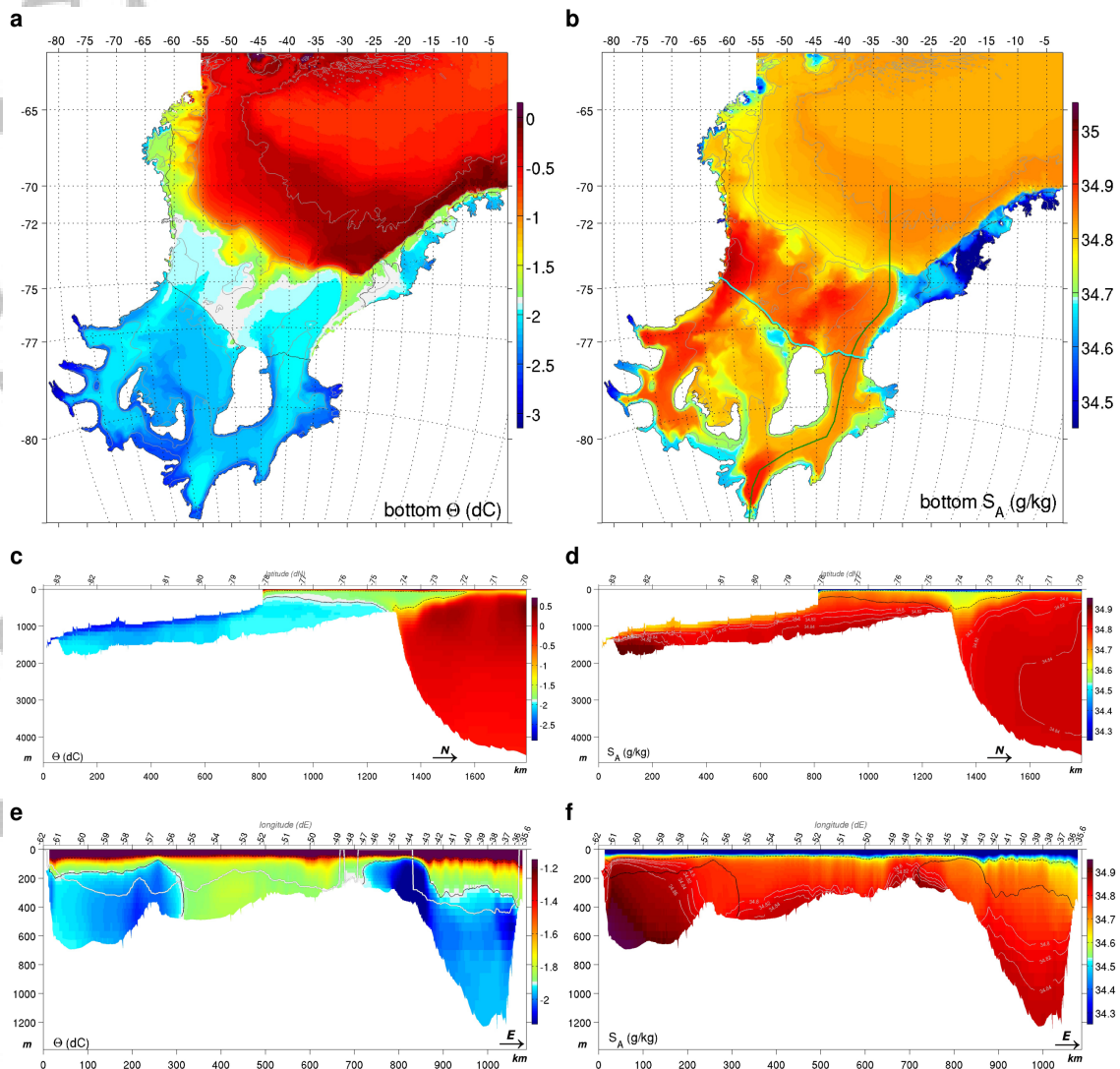


Figure 7 | Illustration of the water mass structure in the reference experiment. 5-year average February (years 4-8, i.e. 1993-1997) conservative temperature Θ & absolute salinity S_A : **a** & **b** at the sea floor, **c** & **d** along a vertical section crossing from the open gyre in the north, across Filchner sill & Trough through the Filchner cavity towards the Foundation ice stream grounding line (along the green line in **b**), and **e** & **f** along a vertical section across the southwestern Weddell Sea continental shelf just north of the FRIS ice front (along the cyan line in **b**, which near-coincides with the ice front contoured in dark gray just to its south). Contours in **a,b** show bathymetry (light gray) and ice fronts (dark gray), in **c-f** Θ isotherms of -1.9 & -1.4 °C (solid & dashed black, the latter often barely visible as close to the surface), and additionally highlight in **d,f** selected S_A contours (gray), and in **e** the depth of the ice draft just to the south (white-gray).

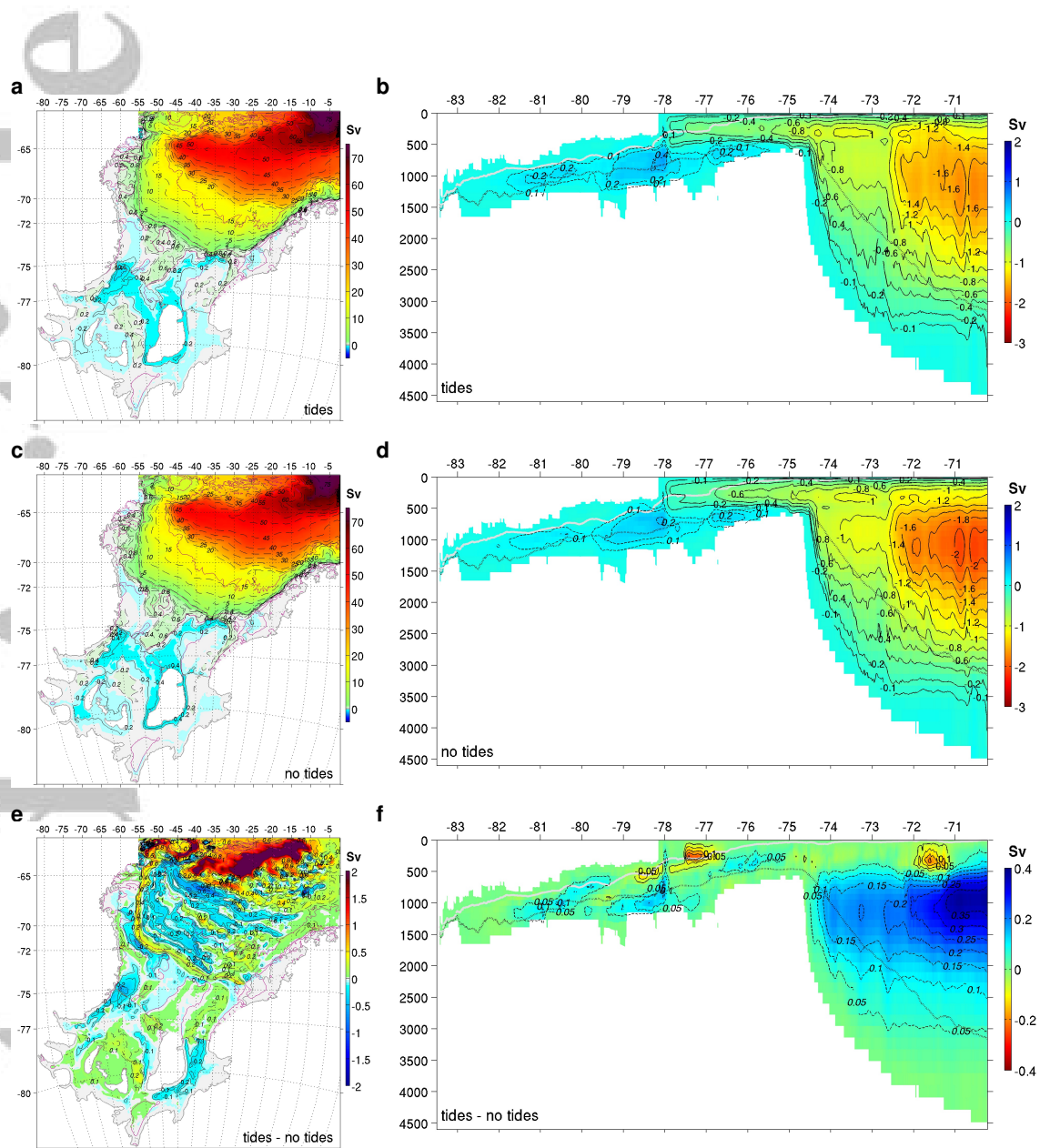


Figure 8 | Integral metrics of circulation. **a,c,e** Barotropic, and **b,d,f** meridional overturning streamfunctions, for the years 4-8 average (1993-1997), both colored and contoured in Sv (dashed contours denote positive streamfunction values, and note the reversed color scale in **b,d,f** vs **a,c,e**): **a,b** in presence of tides, and **c,d** in absence of tides; **e,f** display the tidal signal itself ($\Delta \equiv \text{tides} - \text{no tides}$). Additional contours are in **a,c,e** 400, 1500 & 4000m isobaths (purple), icefronts (gray), as well as latitude, longitude (dotted), and in **b,d,f** the weighted-average ice draft & sea floor depths (gray line & black dots).

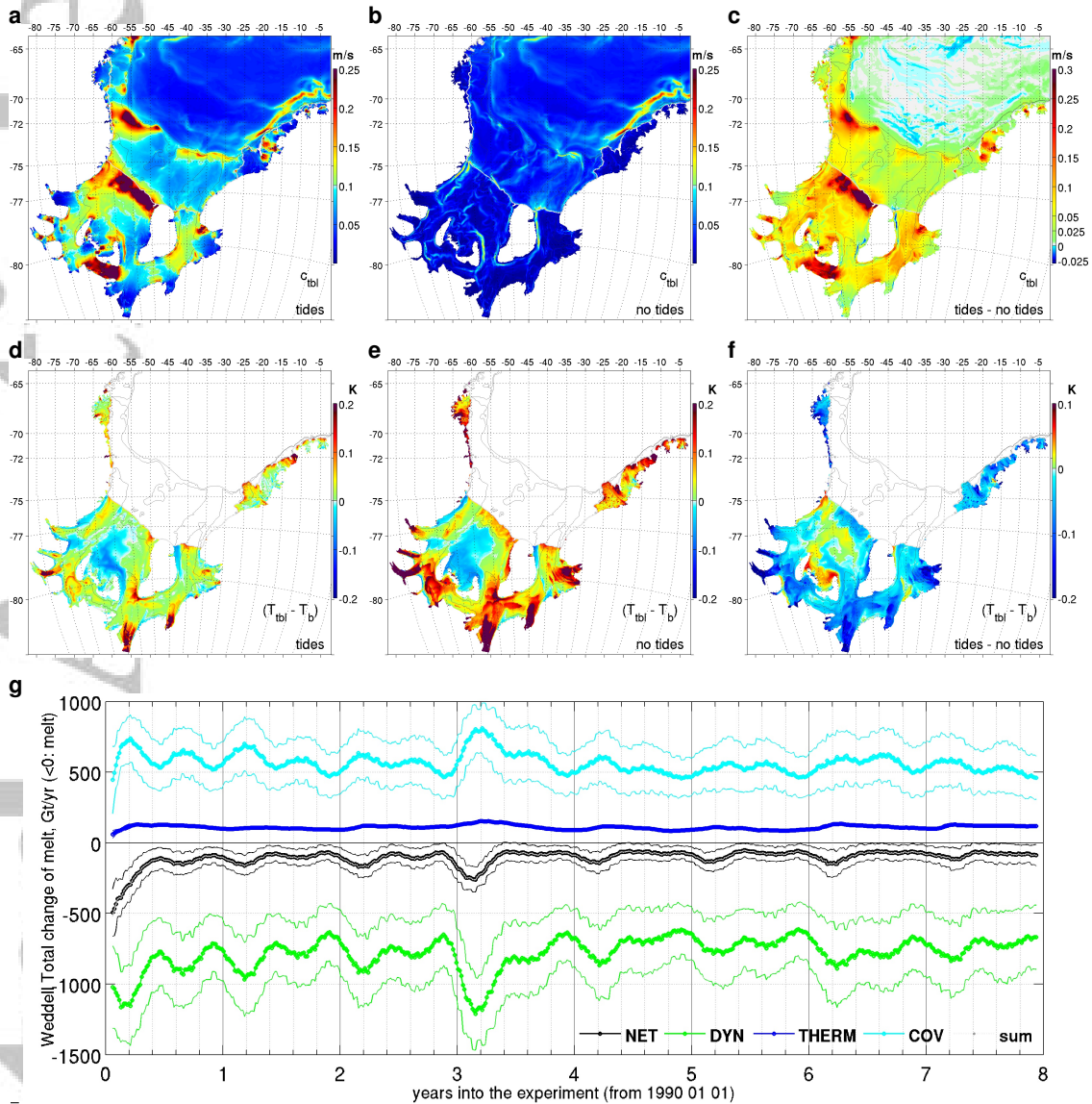


Figure 9 | Mechanisms of the tidal impact on Weddell Sea ice-shelf melting. Top panels (a,b,c) map the sub-ice-shelf ocean boundary layer flow speed c_{tbl} (in m/s; in the open ocean top 30 m speed is mapped for reference), panels below (d,e,f) the departure of the sub-ice-shelf boundary layer ocean temperature from the freezing point at the ice base (in K): on the lhs (a,d) in presence of tides, in the middle (b,e) without tides, and on the rhs (c,f) the tidal signature Δ . All maps show year 4–8 averages, and contour 500 & 1000 m isobaths for reference. Panel g displays timeseries of the tide-driven Weddell ice-shelf melt $\Delta\mathcal{M}(t)$ (black, <0 melt), and reconstructs these, following Eq. (2), from dynamical (green), thermodynamical (blue) and covariational (cyan) mechanistic contributions, with the sum of the three latter (gray dots) visibly coincident with the net change. (As in Fig. 2, 45-day running means of 5-day averages and their one standard deviation envelopes are shown.)

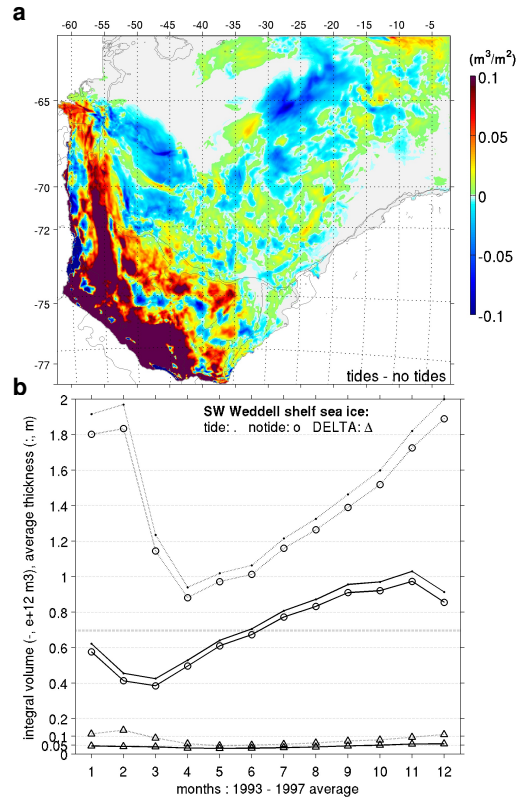


Figure 10 | Impact of tides on southwestern Weddell sea-ice distribution, estimated from the year 4-8 5-year average of the reference versus no-tide simulations. **a** Map of the tidal signal Δ in February sea-ice volume (divided by grid-cell area). **b** 5-yr average monthly timeseries of the southwestern Weddell continental shelf (east of 25°W , south of 62.5°S , and inside the 1500 m isobath) integral sea-ice volume (solid lines, in 10^{12} m^3 – as reference the horizontal dotted line shows the volume of a 1 m slab covering the whole region) and average sea-ice thickness (actual thickness, not volume per area, dotted lines): with tides (dots), without tides (circles), and the tidal signal Δ (triangles).

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