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1	Combined US-ESR dating of fossil teeth from El Harhoura
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4	
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13	
14 15	Abstract The study of the Middle Stone Age (MSA) in North Africa is according to better understand
15 16	The study of the Middle Stone Age (MSA) in North Africa is essential to better understand the human dispersals during the Late Pleistocene. The timing of the transition from the
10 17	MSA to the Later Stone Age (LSA) in the region is still debated due to the scarcity of sites
17 18	and the limited chronological constraints available. Among the few existing MIS 5 to MIS
19	1 archaeological sites in Northwestern Africa, El Harhoura 2 cave (Morocco) has been
20	extensively excavated during the last decade. It has provided MSA and LSA sequences
21	from which human remains have been systematically discovered in association with faunal
22	and lithic assemblages.
23	The combined US-ESR method has been applied to date five teeth from layer 3 and 4A,
24	which correspond to the most recent MSA occupations at El Harhoura 2 cave. The results
25	suggest the disappearance of the MSA at this site to have occurred around ~ 40 ka, at the
26	end of MIS 3, which is consistent with the palaeoenvironmental interpretation derived
27	from the faunal remains. These new results suggest that the last MSA human occupation
28	are ~15 ka more recent than previously showed by OSL dating. Further investigation is
29	required to fully understand the reason of such discrepancy between the two methods.
30	
31	1. Introduction
32	During the last decades, research in Northwestern Africa has greatly improved the
33	chronology of the North Africa Middle Stone Age (Barton <i>et al.</i> , 2009, 2015; Dibble <i>et al.</i> ,

2012; Doerschner *et al.*, 2016; El Hajraoui *et al.*, 2012; Jacobs *et al.*, 2011, 2012; Klasen *et al.*, 2017; Mercier *et al.*, 2007; Richter *et al.*, 2017, 2010; Schwenninger *et al.*, 2010).
These dating results have contributed to highlight the importance of this region in the study of the origin and dispersal of anatomically modern humans within and out of Africa.

In North Africa, the end of the Late Pleistocene is marked by a cultural and subsistence shift from the Middle Stone Age (MSA) to the Later Stone Age (LSA). The nature and timing of this transition from MSA to LSA remains poorly known due to the paucity of reliable dating data and a lack of well-stratified sites. This results in a large age uncertainty for this transition, dated to between 60 and 25 ka.

El Harhoura 2 cave (Fig.1), located in Rabat-Temara region (Morocco), is one the few 43 sites in Northwestern Africa with a complete archeological sequence spanning from the 44 MSA and the Neolithic (El Hajraoui et al., 2012). This coastal area is well known for its 45 rich Middle to Late Pleistocene (MIS 5 to MIS 1) human fossil record. The MSA and LSA 46 human occupations are documented in Dar-es-Soltane 1, Dar-es-Soltane 2, El Harhoura 2, 47 El Mnasra and Contrebandiers caves. This area has also delivered most of MSA humans 48 49 remains (Debénath, 2000; El Hajraoui et al., 2012). However, for some of the caves such as El Mnasra, the stratigraphic interpretation of the deposits is complicated by post-50 depositional perturbations (Lenoble, 2010). In contrast, the undisturbed archaeo-51 stratigraphic record observed at El Harhoura 2 makes the site a good candidate to 52 53 chronologically constrain the disappearance of the MSA in the region.

In the present study, we applied the combined US-ESR dating method on hydroxyapatite from herbivore fossil teeth (Grün *et al.*, 1988) to directly date human occupation at El Harhoura 2. These results will contribute to the refine the end of the MSA of Rabat-Temara sites.

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59 2. Geological, paleoenvironmental and archeological contexts

El Hahroura 2 cave (33°57'08.9''N, 6°55'32.5''W) is located in Rabat-Temara region, 60 (Nespoulet et al., 2008; El Hajraoui et al., 2012). This karstic cave was dug in local 61 calcarenites by the marine erosion of the Plio-Pleistocene coastal ridge, described as a 62 succession of marine and continental deposits (Chahid et al., 2016). Located 300 m from 63 the shoreline and 20.6 m above the present NGM (actual sea level), the site length is 64 estimated to 22 m, the depth around 9 m and the height ~ 8 m. El Harhoura 2 was 65 66 discovered in 1977 and a preliminary excavation was conducted mainly in the entrance area (El Hajraoui et al. 2012). A first test pit of 8 m² was established in 1996, followed by 67 a lateral extension excavated in 2007, which allowed to describe the complete 68 archeological sequence at the entrance of the cave. In 2012, the excavation was further 69 extended to cover a total area of more than 200 m². 70

The sedimentary cave infilling is homogenous and dominated by sandy-clay matrix (Boudad *et al.*, 2017). Four sedimentary units (Fig.2-A) were identified from bottom to top. Unit I is dominated by grey sandy clays. Unit II is defined by a higher proportion of clays and a lower content of silts. Unit III is described as an intermediate unit between clay and carbonates and the Unit IV described by a lower content of sands.

Archeostratigraphy (Fig.2-A) has been mainly described from the test-pit of the 76 entrance area and subdivided into eleven levels (El Hajraoui et al., 2012). Level 1 77 (Neolithic), level 2 (LSA), levels 3, 4, 6, 8 (MSA). Other levels (5, 7, 9, 10 and 11) were 78 excavated only on a small surface ($<3m^2$). For a total number of n=687 lithic artefacts 79 founded in the MSA levels, this number is less than n=14 for levels 5, 7, 9, 10 and 11 (El 80 Hajraoui et al., 2012, p.92). In the current state of excavations, these levels are considered 81 as archeologically sterile. Only the levels 3 and 4 are studied in this paper. The MSA 82 industries of these levels are composed by Levallois and micro-Levallois knapping tools, 83 laminar artifacts, side-scrapers and chopping tools. No tanged tools, characteristic of the 84 Aterian, have been recovered, even in extensively excavated areas. Humans burials were 85

found in the level 1 and 2 and post-cranial isolated bones were found on the levels 3 and 9(El Hajraoui et al., 2012).

Taphonomic studies on abundant and well-preserved micro-faunal remains confirm 88 that no major disturbance occurred in these accumulations, and demonstrate the integrity of 89 the archeological levels (Stoetzel et al., 2011). Zooarcheological studies (Campmas et al., 90 2015) show a diversified faunal assemblage (presence of Alcelaphinae, Equidae, Bovinae 91 92 and Rhinocerotidae). Faunal spectrum is dominated by *Gazella* sp. which represent >70% of the number of identified specimens (NISP) in levels 3 and 4A (and almost 57% in level 93 8). There is also a presence of carnivores (between 10% and 16% of the NISP for the levels 94 3 and 4A), which appear to be the main accumulators of large mammal remains in levels 2 95 96 to 5 (Campmas et al., 2015). These observations, combined with results from the lithic analysis (Stoetzel et al., 2014), show that human occupations in these levels were short and 97 non-intensive. Considering palaeoenvironmental reconstructions, both large and small 98 vertebrate assemblages from the whole archeological sequence show a succession of 99 100 relatively humid (levels 3, 4A, 6 and 8) and arid (levels 2, 5 and 7) phases (Stoetzel et al., 101 2011, 2014).

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104 **3. Material and methods**

Five fossil teeth (Fig.2-B-C) of *Equidae* (*Equus sp.*) and *Bovidae* (*Gazella sp.*) were selected for dating analyses from the Institut National des Sciences de l'Archéologie et du Patrimoine (INSAP, Rabat) collections (three from level 3, one near the limit between levels 3 and 4A and one from level 4A). These teeth were collected during excavations conducted between 2004 and 2010. Combined US-ESR dating analyses were performed at the Geochronology laboratory of the Museum National d'Histoire Naturelle (MNHN) in Paris.

Fossil teeth were prepared following the protocol described in Wagner et al. (2010). 112 Dental tissues (enamel, dentine, cement) were mechanically separated. Enamel samples 113 were cleaned using a dental drill then ground and sieved to 100-200 µm. Depending on the 114 amount of enamel powder available, each sample was split into 8, 9 or 10 aliquots (~50 mg 115 each) and irradiated using a 60 Co γ -ray source at LABRA (CEA, Saclay, France) in 2014. 116 ESR measurements were carried out at room temperature using EMX X-band Bruker 117 spectrometer and the following acquisition parameters: 5 scans, 1024 points resolution, 118 1.013 mW microwave power, 9.80 GHz microwave frequency, 100 kHz modulation 119 frequency, 0.1mT modulation amplitude, 20.48 ms conversion time and 5.12 ms time 120 constant. The ESR measurements were repeated 4 times for each aliquot over different 121 days. ESR intensities were extracted using the Bruker WINEPR System software from the 122 asymmetric ESR signal between the T1-B2 (Grün, 1998) and normalized by the weight of 123 each aliquot. Average ESR intensities form the repeated measurements were used for the 124 evaluation of the equivalent dose (D_e) determination. 125

D_e values (Table 1) were obtained by fitting a single saturated exponential (SSE) function through the experimental data points (Yokoyama et al., 1985; Duval and Grün, 2016). Fitting was performed with Origin Pro8 software, with data weighted by the inverse of the squared ESR intensity (Grün and Brumby, 1994). All ESR dose response curves are

130 provided in Fig.1 SM. An example of ESR spectrum is displayed in Fig.2 SM. This 131 spectrum shows a non-horizontal baseline which may potentially impact the D_e (Duval and 132 Martín-Francés, 2017). This was evaluated on one sample (EH2-2730) by performing a 133 baseline correction (using a cubic function) for each aliquot measured.

Alpha-ray spectrometry U-series analyses (Table 2) were carried out at the Musée de 134 l'Homme U-series laboratory. Samples (0.3-1 g) were dissolved in HNO₃ (7 N) and spiked 135 with 233 U, 236 U and 229 Th and charged onto anion exchange resin AG 1×8 (100-200 mesh 136 chloride form) in order to separate and purify U and Th sources (Shao et al., 2011). The 137 columns were rinsed with HCl (0.1 N) to collect uranium and purified using a UTEVA 138 resin column (Horwitz et al., 1992) in HNO₃ (7 N). The was eluted with HCl (8 N). The 139 purified U and Th were then evaporated by mixing TTA/benzene solution and deposed on 140 an iron disk covered by aluminum paper. Alpha counting was run for 8 days. 141

Dose rate provided on Table 1 was determined in situ (Table 3) using a γ -ray 142 portable spectrometer Inspector 1000 Canberra with NaI detector and following the 143 144 threshold technique (Mercier and Falguères, 2007). Additional high resolution γ -145 spectroscopy measurements were performed at the Geochronology laboratory (Paris) on raw sediment samples (Fig.2-C, Table SM.1), using a high resolution low-background Ge 146 detector, to determine the external β -dose rate from specific activity of U, Th, and K. 147 Analyses were performed on about 90 g of homogenized sediment samples sealed in a 148 149 plastic box during at least three weeks to ensure radon equilibrium.

The cosmic dose was evaluated according to Prescott and Hutton (1994). The spatial configuration of the cave in different orientations (North, South, West and East) was taken into account with a 3D model as in Richard et al. (2017).

153 Moisture content was measured in the laboratory by drying at 40°C for ten days. The water content was estimated form the weight difference sediment before and after 154 drying. Sediments presented low water contents, below 5 %, which are quite similar to 155 values presented by Jacobs et al. (2012). However, we suspect these values may 156 157 underestimate the long-term water content, as sediment samples were collected in summer (June 2014) in relatively dry conditions (archeological sections exposed to sunlight). 158 Consequently, we assumed instead a higher value of $10 \pm 5\%$ for age calculation (Grün, 159 1989). 160

161 Combined US-ESR ages were calculated using DATA program (Grün, 2009) and the 162 following parameters: α -efficiency of 0.13 ± 0.02 (Grün and Katzenberger-Apel, 1994), a 163 water content of 5 ± 3 wt% in dentine and cement and null in the enamel. β -attenuation 164 factors are from Brennan et al. (1997) and dose rate conversion factors from Guérin et al., 165 (2011). Age results are given at 1σ confidence level. A cement/enamel/dentine 166 configuration was used for the age calculation of sample EH2-2993, whereas we used a 167 sediment/enamel/dentine configuration for the other samples.

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171 4. **Results and discussion**

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Equivalent doses, doses rates, *p*-values, and US-ESR ages calculated for each sampleare given in Table 1. U-series data and enamel thickness are presented in Table 2.

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177 U-series data

Uranium content in the enamel ranges between 0.170 and 0.060 ppm, i.e. close to the detection limits of α-ray spectrometry (see Falguères et al., 2018). U-content in cement and dentine is between 4 and 8 ppm. 234 U/ 238 U ratios are systematically higher than 1 in dentine and cement tissues. In enamel, this ratio is associated with large relative errors (between 10 and 20%), particularly for samples in which U content is very low. *P-values* for dental tissues are superior than -1 and 230 Th/ 234 U values are lower than unity (between 0.03 and 0.81): they suggest that there is no uranium leaching in the dental tissues.

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188

187 *Radioelements content*

Specific activities of ²³⁸U, ²²⁶Ra, ²²²Rn and ²³²Th in the surrounding sediments are presented on Table SM.1. No significant disequilibrium in the U-238 decay chain is observed for EH2-sed 1 and EH2-sed 2 sediments. In contrast, a 17 % loss of Ra and 26 % of Rn-loss observed for EH2-sed 3 sample. The origin of these disequilibria is unknown.

193 194

- 195 Equivalent doses
- 196

197 De values range between 20 and 31 Gy. The ESR spectrum displayed in Fig.SM.2 of 198 the sample EH2-2730 presents a non-horizontal baseline. A sensitivity test performed on sample EH2-2730 shows that baseline correction has, however, a very limited impact on 199 the D_e value (< 3 %). Following Duval and Grün (2016), the influence of the maximum 200 irradiation dose (D_{max}) on the D_e estimate was also evaluated by successively removing the 201 last irradiation points (until 220-240 Gy). Results displayed in Fig.SM.4 show a minor 202 203 impact of D_{max} in the present data set, as the difference between each resulting D_e does not exceed the error range. Finally, repeated measurements (n=4) show relatively limited 204 variability of the individual D_e values derived from each independent measurement (<8% 205 for a given sample; Table SM.3). All these evidence consistently illustrate the robustness 206 207 of the ESR data set collected.

- 208
- 209 *Dose rate*

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211 Results show that 3/4 teeth display homogeneous dose rates, except for EH2-2993 212 sample that exhibits a lower value by about 10%. This difference could be explained by the 213 presence in this tooth of a cement layer, which produced a lower β dose rate compared to 214 that from the surrounding sediment.

Gamma dose rates significantly contribute to the total dose rate by 50-60 % (Table 1,
Fig. SM.3), while cosmic dose rates represent 15-20% of the total dose rate. Consequently,

these two parameters significantly impact the age calculation. The correspondence between 217 sediment and *in situ* measurements with tooth samples is specified in table SM.1. The 218 comparison of our data with in situ gamma dose rates measured by Jacobs et al. (2012) and 219 those measured with TL-dosimeters by Janati-Idrissi et al., (2012) does not a show 220 significant variability (Table 3). For level 3, values overall range from 368.3 ± 18.4 to 221 $310.0 \pm 15.5 \,\mu$ Gy/a. The present work and Jacobs et al. (2012) obtained mean values of 222 223 363.9 ± 18.2 and $346.7 \pm 17.3 \,\mu$ Gy/a, respectively, which representing a variability < 5%. Similar observations can be made for level 4A: values range from 348.0 ± 17.4 to $385.4 \pm$ 224 19.3 µGy/a, with a variability of about 10% (Table 3). As these mean in situ gamma values 225 are very close to the individual in situ measurements, the latter were chosen to calculated 226 227 ages. To evaluate the potential impact of the moisture content on the calculated age, simulations were performed by taking into account different moisture content $(3 \pm 1 \%, 5 \pm$ 228 229 1 %, 8 ± 3 % 10 ± 5%). Some of water content, as $3 \pm 1\%$, 5 ± 1 , $6 \pm 2\%$, 7 ± 2 and $9 \pm 1\%$ 2% have been used by Jacobs et al. (2012). Results are summarized on Fig.SM.5 and show 230 231 a difference of <1%. Consequently, whatever the water content used (between 3 and 10 %), the moisture content has virtually no impact on the calculated ages. 232

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- Combined US-ESR ages

Combined US-ESR range from 39 to 48 ka for the level 3 (with a mean age of 43.7 ± 4.5 ka) and from 46 to 49 ka for the level 4A.

EU-ESR ages, considered as minimum ages (Table SM.2), were also calculated using 239 240 DATA software for comparison. EU-ESR results are very close to US-ESR ages, with differences ranging between 3 and 6 ka (between 7-15% of variability). This is due to the 241 fact that uranium content in dental tissue from El Harhoura 2 samples generated low dose 242 rate from teeth (accounting for 1-15 % of the total dose rate, Fig. 2 SM). Consequently, U-243 244 uptake modelling has a negligible impact on the calculated ages, as often expected for Late Pleistocene tooth samples (Richard, 2015). Combined US-ESR age results appear to be 245 overall stratigraphically consistent. 246

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Comparison with previous results

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Our ages are in agreement with previous combined US-ESR results obtained by Janati-Idrissi et al. (2012) for the level 4A but they are younger than previous OSL dates. Only two of OSL ages overlap with the combined US-ESR ages for the level 3 (Fig.3). Indeed, OSL ages locating the age of the latest MSA human occupations around 57 ± 4 ka as a mean age (Jacobs *et al.*, 2012) and the age of the level 4A at 74 ± 4 ka.

Regarding these data, the difference observed cannot be explained by the gamma dose rates, which are quite similar for ESR and OSL samples. Considering the equivalent doses, many remarks can be made. First of all, the D_e values determined using SG-SAR protocol by Jacobs et al. (2012) lead to the identification of 2-3 quartz grain populations using Finite Mixture Model (FMM). The presence of different components could be explained

either as mixing grains from underlying/overlying levels even if "mixing mechanism is not 261 clear", or "by the effect of variations in the beta dose rate to individual grains" (Jacobs et 262 al., 2012, p.385-386). It is difficult to understand the meaning of these different population 263 264 grains in term of geology. In fact, the sedimentary processes in karstic coastal caves of this area are not well understood in terms of chronology and origin, due to the impossibility to 265 differentiate quartz from different origins (aeolian, marine, peeling of the cave)(Niftah et 266 267 al., 2005). The final OSL D_e chosen for age calculation correspond to the quartz grain population statistically well represented (for level 3, between 63.7 ± 4.7 to $48.5 \pm 5\%$ of 268 representation). Consequently, the resulting OSL age estimates are about 20% higher than 269 our combined US-ESR ages. Using instead the D_e from the less representative quartz grain 270 271 population (between 5.0 \pm 2.7 to 19.2 \pm 4.2 % of representation) would allow to recalculate OSL ages between ~ 19 ka and 38 ka for level 3. For the level 4A, D_e from the less 272 representative quartz grain population (8.7 \pm 3.7 % of representation) provides an OSL age 273 274 of about 42 ka. This result would be closer to the combined US-ESR ages.

275 The discrepancy of ~ 20% between previous OSL ages (Jacobs et al., 2012) and 276 combined US-ESR ages could also be explained by an underestimation of the D_e values obtained for the fossil teeth, due to the presence of unstable radicals in the hydroxyapatite 277 as observed in Joannes-Boyau and Grün (2011) and Richter et al. (2017) on enamel 278 279 fragment. However, ESR dose estimate were carried out in the present study on enamel 280 powder, which results in randomly spatial distribution of the enamel crystals that renders difficult such a comparison. The impact on these unstable components remains 281 nevertheless unclear and needs to be further investigated. It is indeed sample dependent, 282 and some teeth are simply not affected by this issue (e.g. Dirks et al., 2017). 283

Combined US-ESR ages presented in this work suggest that the last MSA occupation in Rabat-Temara is dated to around 40 ka. Other Northwestern African sites display similar chronologies for late MSA occupations: ~ 29 ka using radiocarbon method in Taforalt (Barton *et al.*, 2015) and 30.9 ± 2.5 ka with OSL (Weisrock *et al.*, 2008) for the end of MSA without tanged tools at Wadi Noun.

In contrary, others sequences in the Rabat-Temara region show different chronologies: for example, the latest MSA in Contrebandiers cave is older than 90 ka by OSL and 70 ka by TL (Dibble *et al.*, 2012). However, erosion surface and the disruption in the deposits of the Contrebandiers cave (Aldeias *et al.*, 2014) limit the discussion. Nevertheless, these combined US-ESR ages of El Harhoura 2 cave match with previous TL ages obtained for the end of MSA in El Harhoura 1 cave placed between 32.15 ± 4.8 ka and 41.16 ± 3.5 ka (Debénath *et al.*, 1986).

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297 5. Conclusion

The numerical ages obtained in the present study correlate the latest MSA human occupations at El Harhoura 2 to the end of the MIS 3 (Fig.3). This attribution is in good agreement with paleoenvironmental data derived from the study of micro fauna (Stoetzel *et al.*, 2011), which position these levels in humid phase and associated to a tree-filled meadow environment.

Combined US-ESR dating of fossil teeth provides a direct chronological constraint on the archaeological levels at El Harhoura 2 and suggest that the latest MSA human

305 occupation may have occurred at least 15 ka after the OSL ages previously estimated by
 306 Jacobs et al. (2012). The origin of the discrepancy between the results from the two
 307 methods should however be further investigated in the future.

The direct dating of the last MSA human occupation at El Harhoura 2 cave provides an important contribution to a better understanding of the timing of Late Pleistocene human dispersals in the coastal environments of Northwestern Africa. Future investigations at Rabat-Temara sites should focus on both the revision of old chronological data and new dating application studies of well-identified and well-contextualized archeological material in order to improve the existing chronological framework of the area.

314 315

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528 529	Supplementary material
530	Table SM.1: ²³⁸ U, ²²⁶ Ra, ²²² -Rn, ²³² Th and ⁴⁰ K content (dpm/g or %) measured and
531	converted using dose rate factors published in Guérin et al. (2011). Sample EH2-sed 3 has
532	been measured but not used on this calculation because of its geographical distance from
533	the teeth samples.
534	
535	Table SM.2: EU-ESR and combined US-ESR ages comparison for tooth samples dated in
536	this work. Errors are displayed at 1 σ .
537	this work. Errors are displayed at 10.
538	Table SM.3: D_e values over the repeated ESR measurements (n=4). The dispersion of the
539	induvial D_e per sample is not significantly important and remains rarely more than ~8 % (<
540	3 Gy). RSD: relative standard deviation. Data are associated with 1 σ .
F 4 4	
541	Figure SM 1: ESD does response out to for complex dated in this work computing by
542 543	Figure SM.1: ESR dose response curve for samples dated in this work computing by Origin Pro 8 software, using SSE function (Yokoyama et al., 1985; Duval and Grün,
543 544	2016). All D_e presented range between 20 and 31 Gy.
	2010). The De presented range between 20 and 51 Gy.
545	
546	Figure SM.2: ESR spectra measured in sample EH2-2730.
547	
548	Figure SM.3: proportion of the different dose rate components. Key: linternal = dose rate α
549	+ β contribution from the enamel, $\beta 1$ = beta contribution from the dentine, $\beta 2$ = beta
550	contribution from cement or sediment.
551	
552	Figure SM.4: Evolution of D_e as a function of D_{max} for all samples fitted with a SSE
553	function. 1 σ errors are displayed. The D _e value used for the US-ESR calculation is shown
554	with a dash line.
555	
556	Figure SM.5: recalculated combined US-ESR ages in taking account different water
557	content of 3 ± 1 %, 5 ± 1 %, 8 ± 3 % and 10 ± 5 %.

Sample	Sampl	à (DLay Gy)	er_	Z (cm)	Tissue	U	-cont	ent (J	opm)	± Dose	²³⁴ U/ ² rate (µ(38U Gy/a)	±	²³⁰ Th/ ²³⁴	⁴ U	±US-E	Initial enamel SR Abi c kn ess ± (µm)	Enamel removed side (µm)	Enam parameter (p-val 2 (µn	side
					Internal $(\alpha + \beta)$	±	β1	±	β2	±	Gamma	±	Cosmic	±	Total	±		Y	е	d	с
EH2-29	3 3	1	9.95	0.8	7	2	16	4	11	2	378	27	104	10	516	27		39 3	1.27 ± 0.33	0.58 ± 0.39	1.27 ± 0.32
EH2-398	36 3	2	7.25	3.8	11	5	50	22	91	8	362	36	104	10	618	43	· ·	44 7	$\textbf{-0.64} \pm 0.1$	$\textbf{-0.43} \pm 0.21$	-
EH2-27.	30 3	2	8.83	0.8	0	0	26	6	108	10	368	39	104	10	606	41		48 4	19.56 ± 2.02	0.72 ± 0.38	-
EH2-102	21 4A	3	0.19	1.7	17	14	30	6	90	9	389	22	93	10	619	31		49 5	$\textbf{-0.87} \pm 0.09$	0 ± 0.20	-
EH2-109	84 4A	2	9.13	1.1	12	5	50	7	103	11	372	27	93	10	630	33		46 3	-0.74 ± 0.15	$\textbf{-0.22} \pm 0.10$	-

Table 1: Equivalent dose (D_e), Dose rate and combined US-ESR age estimation presented at 1 σ confidence level. Key: e= enamel; d=dentine; c=cement. The dose rate components presented are: internal dose rate ($\alpha + \beta$ of the enamel), β 1 (beta contribution from the dentine), β 2 (beta contribution from cement or sediment), gamma and cosmic dose rates.

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EH2-2993	3		D	4.72	0.17	1.05	0.04	0.13	0.01			
		-415	С	4.78	0.22	1.14	0.06	0.03	0.03			
			Е	0.17	0.02	0.94	0.14	0.09	0.09	830 ± 83	39 ± 4	57 ± 6
EH2-3986	3	100	D	7.90	0.47	1.05	0.06	0.29	0.06	7		
		-426	Е	0.06	0.01	0.90	0.21	0.25	0.82	720 ± 72	43 ± 4	94 ± 9
EH2-2730	3	-437	D	7.44	0.44	1.06	0.06	0.14	0.02			
		-437	Е	0.07	0.01	1.29	0.26	0.03	0.16	690 ± 69	53 ± 5	40 ± 5
EH2-10221	4A	150	D	6.59	0.29	1.05	0.05	0.75	0.17			
		-458	Е	0.15	0.04	1.05	0.35	0.81	0.32	710 ± 71	92 ± 9	64 ± 6
EH2-10984	4A	-470	D	7.98	0.31	1.16	0.05	0.18	0.02			
		-470	Е	0.09	0.02	0.72	0.24	0.67	0.38	700 ± 70	58 ± 6	54 ± 5

Table 2: U-series data obtained by alpha-spectrometry analyses of El Harhoura 2 tooth samples. Data are presented with 2σ . Key: D= dentine; C= cement; E = enamel; side 1 = dentine and side 2 = sediment or cement. For enamel in low U-content, isotopic ratios (italic) are associated with large relative errors (between 10 and 20%).

Gamma Dose rate (µGy/a)

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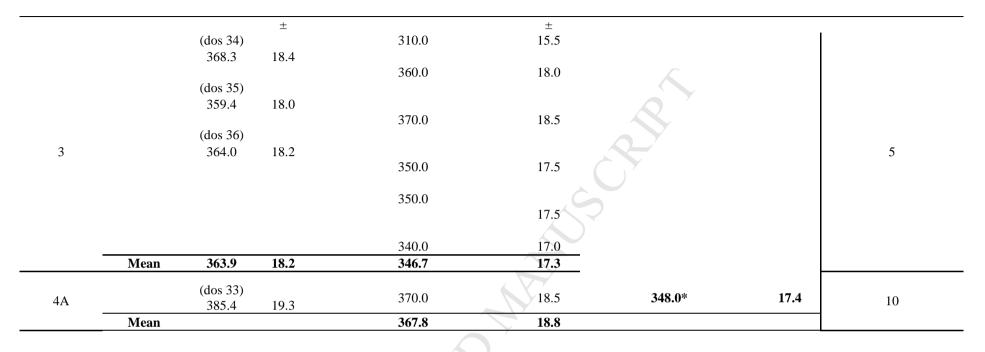
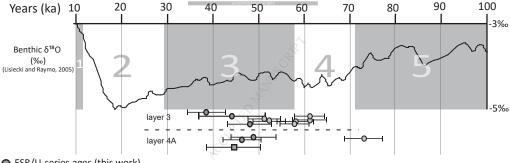
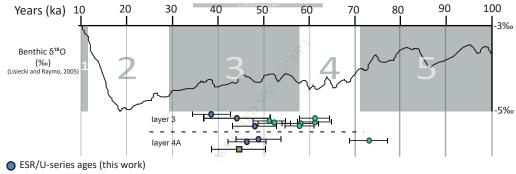


Table 3: *In situ* gamma dosimetry measured with a NaI detector (Jacobs *et al.*, 2012; this work) and TL dosimeters (Janati-Idrissi et al., 2012) in levels 3 and 4A. A mean value was calculated for each level. Error associated to mean value is the mean of error. A 5% error was assumed for each individual measurement. For level 3, 5 % variability was estimated from the two mean values (346.7 ± 17.3 and 363.9 ± 18.2). For level 4A, 10 % variability was estimated from the two values presented (370.0 ± 18.5 and 348.0 ± 17.4); *: initial value given was 438 µGy/a and the cosmic dose rate of 90 µGy/a was subtracted from this value.



- ESR/U-series ages (this work)
- SG-OSL ages (Jacobs et al., 2012)
- ESR/U-series ages (Janati-Idrissi et al., 2012)



- SG-OSL ages (Jacobs et al., 2012)
- ESR/U-series ages (Janati-Idrissi et al., 2012)

