

Historical Earthquake Scenarios for the Middle Strand of the North Anatolian Fault Deduced from Archeo-Damage Inventory and Building Deformation Modeling

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Seismological Research Letters

Historical earthquake scenarios for the middle strand of the North Anatolian Fault deduced from archeo-damage inventory and building deformation modeling

--Manuscript Draft--

- **Historical earthquake scenarios for the middle strand of the North Anatolian Fault deduced from**
- **archeo-damage inventory and building deformation modeling**
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Abstract

 The city of Iznik (ancient Nicaea), located on the middle strand of the North Anatolian fault zone (MNAF), presents outstanding archeological monuments preserved from the Roman and Ottoman 24 periods ($1st - 15th$ c. AD), bearing deformations which can be linked to past seismic shaking. In order to constrain the date and intensity of these historical earthquakes, a systematic survey of Earthquake Archeological Effects (EAE) is carried out on the city's damaged buildings. Each of the 235 EAE found is given a quality ranking and the corresponding damage is classified according to the EMS98 scale. We show that the walls oriented N-S were preferentially damaged, and that most deformations are perpendicular to the walls' axes. The date of postseismic repairs is constrained with available 30 archeological data and new 14 C dating of mortar charcoals. Three damage episodes are evidenced: (1) 31 between the 6th and late 8th centuries, (2) between the 9th and late 11th centuries AD, (3) after the late 32 14th century AD. The repartition of damage as a function of building vulnerability points toward a global intensity VIII on the EMS98 scale. The 3D modeling of a deformed Roman obelisk shows that only earthquakes rupturing the MNAF can account for this deformation. Their magnitude can be bracketed between Mw 6 and 7. Our archeoseismological study complements the historical seismicity catalogue and confirms paleoseismological data, suggesting several destructive earthquakes along the MNAF since the first century AD. We suggest the fault might still have accumulated enough stress to generate a Mw 7+ rupture.

INTRODUCTION

 Documenting the calendar of past earthquakes in tectonically active regions is an important step to determine their seismic hazard (Guidoboni and Stucchi, 1993; Bilham et al., 2001). In certain contexts, the knowledge of Holocene earthquakes is the only way to estimate the magnitude range expectable in a given area and the recurrence times between big earthquakes. Paleoseismological trenching allows one to identify the last rupturesin the stratigraphy at a given point of the fault (McCalpin, 2009). Historical seismology has been developed to build earthquake catalogues using written descriptions of seismic events (Ambraseys, 2000, 2002). Whereas palaoseismology is primarily concerned with the source of earthquakes, historical seismology interprets the observable consequences in the form of intensities felt at given locations.

 Evidence of coseismic deformation found in the stratigraphy of archeological survey sites and on ancient buildings has been increasingly described and used to complement earthquake catalogues and, in some cases, quantitatively constrain earthquakes thanks to numerical modeling (Korzhenkov and Mazor, 1999; Mazor and Korzhenkov, 2001; Hinzen, 2005). Although the interpretation of archeological evidence is delicate because possible traces of earthquakes are rarely unambiguous, this "archeoseismological" approach allows one to document past earthquakes in regions without written archives, which extends the time scale of the catalogue (Karcz and Kafri, 1978). The archeological record may also preserve smaller events that are poorly preserved in classical trenches

(Rapp, 1986).

 In this study, we apply this archeoseimological approach to the city of Iznik (ancient Nicaea), located 3 km north of the active fault segments of the middle strand of the North Anatolian fault (MNAF, Fig. 1). Since the beginning of the instrumental period, this fault strand has been characterized by a very low seismicity and deformation rate (Özener et al., 2013; Öztürk et al., 2009). However, various datasets documenting the pre-instrumental seismic activity strongly contrast with this modern quiescence. The historical record of ancient earthquakes having affected Iznik (Fig. 2) contains no less than ten destructive earthquakes during the last two millennia (Ambraseys and Finkel, 1991; Ambraseys and Jackson, 1998, 2000; Ambraseys, 2000, 2002, 2009; Guidoboni et al., 1994). The compilation of the paleoseismic trench data along the MNAF suggests the occurrence of at least four Mw 6+ earthquakes during the last two millennia (Barka, 1992, 1993; Dogan, 2010; Ikeda, 1988; Ikeda et al., 1989; Özalp et al., 2013; Yoshioka & Kusçu, 1994). Moreover, the region presents several archeological clues suggesting past earthquake occurrence, such as the submerged basilica recently discovered in Iznik

 Lake (Sahin & Fairchild, 2018), and the numerous traces of deformation and reconstruction on the city's ancient buildings (Benjelloun et al., 2018). Despite this evidence of past earthquakes, more precise data are needed to perform a reliable assessment of seismic hazard for the Iznik region.

 The goal of this study is to use the ancient buildings in Iznik and its surroundings as records of past seismicity. The compilation of many traces of potential earthquake damage provides a complementary dataset to historical seismology and paleoseismology in order to discuss the level of seismic activity of the MNAF during the last 2000 years. The dating of several independent buildings affected by earthquake shaking, and their repairs, allows us to complement the earthquake catalogue. The modeling of archeoseismic damage helps to discuss the epicenter location and strength of the earthquakes responsible for the damage observable today in the city.

TECTONIC CONTEXT OF THE STUDY: THE MNAF AND IZNIK'S SEISMIC LANDSCAPE

 The study area includes the city of Iznik and its close surroundings, in the southeast of the Marmara Sea (Fig. 1 and 3a). It is located on the middle branch of the North Anatolian fault. The North Anatolian Fault zone (NAFZ) is a major tectonic boundary accommodating the relative motion between the Eurasian and Anatolian plates. It is a 1000-km long, right-lateral strike-slip fault crossing the northern part of Turkey (Sengör, 1979). This fault zone is associated with a significant seismic hazard and with several destructive earthquakes since the beginning of the Common Era (Fig. 2).

89 In its western termination, the NAFZ is divided into three strands (Fig. 1). The northern strand (NNAF) continues along the strike of the main fault zone and crosses the northern part of the Marmara Sea from Izmit Gulf to Gelibolu Peninsula. This strand was especially involved in the last seismic crisis that occurred as a succession of ruptures migrating westwards, from Erzincan in 1939 to Izmit in 1999 (Stein et al., 1997). The middle strand (MNAF) separates from the NNAF in the Akyazi district, borders on the south two successive basins known as the Geyve-Pamukova Basin and Iznik Lake, reaches the Marmara

 Sea through Gemlik Gulf and borders the southern shore of the Marmara Sea (Fig. 3a). With a relative horizontal motion estimated around 5 mm/year with GPS (Ergintav et al., 2014), the MNAF shows a 97 deformation rate five times smaller than the NNAF (~20-25 mm/year) (Reilinger et al., 2006). The southern strand (SNAF) separates from the MNAF near Iznik and continues southwestwards to Ulubat and Manyas Lakes (Fig. 1).

 In addition to the MNAF, smaller structures around Iznik Lake show morphological evidence of recent activity. North of the lake, two oblique fault zones, the NW-SE striking Orhangazi fault (Fig. 3a) and the ENE-WSW striking Boyalica fault, separate the basement rocks of the Armutlu block and the Quaternary alluvia of the lacustrine plain (Öztürk et al., 2009). A small normal fault zone, the Elbeyli fault, was described east of Iznik, forming a 7 km long scarp at the foot of a marble and schist massif, cross-cutting in the south the eastern section of the Iznik aqueduct (Benjelloun et al., 2018).

 The natural environment of Iznik presents cumulative expression of Holocene and historical coseismic effects, such as faceted scarps (Fig. 3b), meter scale recent scarplets evidenced on land (Ikeda et al., 1989) and in the bathymetry of Iznik Lake (Gastineau et al., 2020 PREPRINT), laterally offset geomorphic markers (Ikeda, 1988; Sipahioglu and Matsuda, 1986; Yoshioka and Kusçu, 1994), and transtensional basins (Dogan et al., 2015). This leads geologists to qualify the city's geological environment as a "seismic landscape", as defined by Michetti et al. (2005). The most prominent earthquake source for the Iznik region is the MNAF, which represents a 140 km long dextral fault zone between Akyazi district and Gemlik. A rupture propagating along this whole section would correspond to a Mw 7.6±0.3 earthquake (Wells and Coppersmith, 1994).

ARCHEOSEISMOLOGICAL DATA COLLECTION

 The main principle of an archeoseismological investigation is to look for traces of past earthquake deformation in archeological features, which are commonly called Earthquake Archeological Effects

 (EAE; Rodriguez-Pascua et al., 2011). These EAE can be located on-fault, when an archeological structure is crossed and deformed by an active fault. However, the most numerous archeoseismic observations are located off-fault and concern buildings which have been deformed because of coseismic ground motion. Finally, we include in EAE the evidence for postseismic reconstructions which serve to repair a building damaged during an earthquake and/or to make it more resistant to future earthquakes.

 In Iznik and its surroundings, systematic archeoseismic observations were made on the visible ancient remains (Fig. 4 and 5; for the complete list see Tables S2 and S3 in the electronic supplement). We focused our observations on the critical buildings of the city, such as the aqueduct and the defensive walls, which were especially important for its viability and were likely repaired as soon as damaged. The types of inventoried EAE in Iznik include: wall facing collapses, wall warping, wall tilting, cracks, unusual wall offsets, masonry dislocations, block rotations, block corner expulsions, block shifting, damaged arches, dropped keystones, arch collapses, column tilting, expulsion of building corner, and deformed pavement (Kazmér, 2014; Rodriguez-Pascua et al., 2011). To this dataset we also added suspected traces of post-seismic repairs and reconstructions: presumed "seismoresistant" techniques, buttresses, filling of damaged masonry with new materials, partial reconstruction of a damaged building, filling of windows and crenellations, massive reuse of materials taken from older buildings.

 Each individual observation was given an identification code. This code is composed of a prefix related to the nature or category of building (AQ for the aqueduct, W for the defensive walls, C for the *intra muros* buildings and O for the obelisk) and a number. For suspected post-seismic repairs and constructions, an R is added at the beginning of the ID. The observations were geolocalized and the nature of the associated deformation or reconstruction was reported. Whenever possible, we measured the azimuth of the damaged wall, and the amount of displacement (lateral shift, vertical drop, angle of tilting). For oriented deformations, we also measured the azimuth (i.e. axis of the deformation, between 0 and 180°) and strike (i.e. direction between 0 and 360°) of the deformation.

 In the following sections, when referring to the defensive walls, we use the naming conventions of Schneider and Karnapp (1938) and Foss and Winfield (1986). Each tower is numbered from T1 to T114 starting from Yenisehir Gate in the south and progressing anti-clockwise (Fig. 6b). Each wall section is numbered according to the adjacent towers (e.g. W1/2 for the wall section between T1 and T2).

 For each EAE, we used a quality factor between 1 and 3, reflecting the probability of seismic origin of the deformation versus other causes (e. g. anthropic damage, natural decay, soil instability). We assign a quality 1 (good) when the seismic origin is the most probable scenario for the EAE and other processes can be discarded. When the EAE can be associated with one non-seismic process in addition to the seismic scenario, we assign a quality 2 (average). When the seismic origin is dubious because many other processes may be invoked to explain the observed deformation, we assign a quality 3 (bad). The cases of damage were also classified on a damage scale ranging from 1 to 5 (1 corresponding to very slight damage and 5 to complete destruction) and the buildings investigated were classified on a vulnerability scale (Grünthal & Levret, 2001).

 On several EAE, construction material (brick, mortar) was sampled in the original structures and the repairs to constrain the date of damage. Carbonate concretion deposits on the aqueduct formed during potentially coseismic massive leakage episodes were also sampled to date the associated damage.

DESCRIPTION OF THE BUILDINGS STUDIED

The Roman obelisk at Elbeyli

165 6 km north of Iznik, near the village of Elbeyli, stands a Roman obelisk built during the 1st century AD (Fig. 6a and 7). The column is made of five triangular prismatic stones lying on a rectangular pedestal (Fig. 7). Small offsets between the triangular stones suggest translation and rotation motions probably during seismic shaking (Fig. 7b,c).

The aqueduct of Iznik

 The remains of two aqueducts are visible in the east of Iznik. One was discovered in 2019 outside of the city walls, north of the graveyard. The other one, much better preserved, enters the city through Lefke gate (Fig. 6). The best-preserved parts of this aqueduct extend along 600 m, starting from the city gate, with an E-W orientation. 1 km east of the city, the aqueduct runs along the southern border of a carbonate massif where its sources are located (Dereköy village, see Fig. 3a for location). The aqueduct presents an underground section, 400 m long. The aqueduct of Iznik has only recently been studied in detail by Benjelloun et al. (2018). Using a stratigraphic approach coupled with architectural 177 analogies, they concluded that the earliest phases visible were likely to belong to Justinian's period (6th 178 century AD), and identified major restoration works around the $11th$ and $13th$ centuries AD.

The walls of Iznik

 The defensive walls of Iznik are made of two belts surrounding the city (Fig. 6b). They were extensively studied by Schneider and Karnapp (1938) and Foss and Winfield (1986) (the chronology of the fortifications is presented in Fig. S1 in the electronic supplement). The inner belt was first built during 183 the 3rd century AD, after Gothic attacks in 256-257. The remains of these early walls can still be seen at several points (Fig. 4b).

 After an Arab attack in 727, the walls were partly rebuilt by Leo III (717-741). These reconstructions mainly affected the eastern and southern parts. Another important reconstruction campaign followed in 858 under Michael III, especially affecting the eastern and southern sides. After a period of peace, 188 the 10th and 11th century saw the walls damaged on several occasions, until the city was taken by the 189 Turks in the end of the 11th century. In 1097, the city was given back to the Byzantine ruler Alexios I Komnenos, who carried out a few reconstructions (Fig. 4a). The last major reconstructions occurred

191 during the Lascaris and Palaeologue periods in the $13th$ century (Fig. 4c), with the addition of the outer walls during the reign of John III Doukas Vatatzes (1222-1254 AD). Since 2018, important restoring works have been carried out on the defensive walls, erasing most of the traces of previous damage and reconstruction phases. Our observations on the walls were done before the beginning of these works.

Intra muros buildings

 We inventoried several deformations and possible "seismoresistant" techniques in the Iznik theater 198 (Fig. 6b), built during the 2nd century AD, and spoiled significantly as a quarry in later periods. 199 Deformations were also described in Aya Sofia, a church built by Justinian in the $6th$ century AD, 200 restored in the 9th, 11th, and 13th centuries, and converted into a mosque by Orhan shortly after the Ottoman conquest of 1331 (Brounoff, 1925; Schneider, 1943). We also found deformations and technical innovations potentially related to earthquakes in several Ottoman buildings: Yesilcamii 203 mosque built in the late 14th century (Fig. 4e), Mahmut Celebi mosque built in 1442, and Murat II 204 hamam built in the first half of the 15th century (Fig. 4d). Other traces of damage were encountered on a few Ottoman tombstones located in Candarli Hayrettin Pasa mausoleum.

DAMAGE INVENTORY

Location of the damage

 We inventoried 235 cases of damage (Fig. 5). The observations are especially dense on the inner walls (almost 4 observations every 100 m in average) and the outer walls seem much less affected (less than 211 1 observation every 100 m on average). The damage observed is in general regularly distributed all along the inner walls. In some areas the density of observations is much lower. This can be the case when the remains have disappeared, such as in west of the city. In other areas, the dense vegetation or the presence of private areas near the walls prevented the survey, such as in the northeastern section between towers 46 and 50, and in the southwestern section between towers 95 and 104.

Azimuth of the damaged walls

 The relative repartition of the different wall orientations was compared to the relative number of EAE found in these different wall azimuth ranges, no matter the azimuth of the deformation (Fig. 8a). The results do not change significantly when data of lower qualities are included. If all the walls were affected in the same way whatever their azimuth, the relative amount of EAE and the orientations should present similar distributions. This is the case for the E-W and NW-SE oriented walls that show 222 a correspondence between the relative amounts of EAE and the lengths of the walls affected. By contrast, we can see that the N-S oriented walls bear much more EAE than what could be expected from their length: they contain a third of the walls' EAE whereas they represent only 12.7% of the total wall length. Conversely, the NE-SW oriented walls are underrepresented in terms of EAE.

Oriented deformations

 105 of our observations present an oriented pattern (e. g. azimuth of collapse or expulsion). Most of the oriented deformations show a strong correlation with the wall azimuth (Fig. 8b): they are either perpendicular to the wall (out-of-plane) or parallel (in-plane). A clear, dominant, N-S azimuth appears, representing 30% of the observations(Fig 8c). This actually mirrors the greater length of walls oriented E-W (Fig. 8a). For almost all the azimuth ranges, deformations exist in both directions, but the deformations oriented NE-SW are characterized by an exclusive NE strike of the damage.

Time constraints on the deformations

 The earliest EAE are found almost all around the inner walls and in the theater (Fig. 5b). The later sections of the walls were also affected by several deformations. The aqueduct shows a significant number of EAE from all periods. EAE from the Ottoman period were found especially intra muros on the mosques and the bathhouse. It is expected that older buildings display a larger amount of damage

238 compared to more recent ones because they are more likely to accumulate damage from several 239 deformation episodes. As mentioned in the previous section, the amount of EAE seems much smaller 240 on the outer walls built in the $13th$ century, which is coherent with the fact that it is one of the latest 241 buildings of the Byzantine period.

242 The repairs were mainly dated relying on the age of the repaired structure and the technique used in the repair. We could find repairs for all the major construction periods of the city, from the 9th century 244 to the Ottoman times (for the list, see Table S3 in the electronic supplement). For the earlier periods, 245 we could only find possible seismoresistant techniques of construction in the theater with the use of 246 polygonal imbricated ashlars (Garduño-Monroy et al., 2012; Hinzen and Montabert, 2017).

247 We were also able to date with 14 C four charcoal samples from mortar filling the dismantled ashlars of tower T57 base and one sample from the new masonry covering the rotated blocks of Istanbul Gate (Fig. 4f, see Fig. 6b for locations). The ages obtained point toward repairs following a deformation episode in the late 11th or early 12th centuries (Table 1). Finally, we tried to date the organic fraction 251 of carbonate concretions sampled on the aqueduct wall, but we obtained a very recent age. The link between the external concretions and historical earthquakes is therefore unlikely.

 Combining the ages of the affected walls and their repairs, we obtain many time ranges of different widths constraining one or several events (Table 2). The narrower, distinct time intervals obtained for 255 possible earthquakes are (1) 527-787, (2) 858-1097, and (3) after the late $14th$ century AD. The first interval corresponds to several good quality EAE but located in Aya Sofya only. The second interval is 257 constrained by one intermediate and one bad quality EAE in the walls, but also matches the ^{14}C ages of the sampled postseismic repairs. It is then probable that the associated damage observed on Istanbul Gate and tower T57 belongs to this episode. The third interval is documented by one good and two intermediate quality EAE in the walls. We can also mention a wider interval, between 730 and 261 1254 AD, including 12 good quality EAE that might belong to the two precedent intervals. Finally, the Ottoman period presents 17 EAE, mostly of good quality, collected on four different buildings. The

 earliest building, Yesilcamii, was built in 1378 AD, but the EAE observed on Celebi camii indicate a deformation occurring after 1442 AD.

Vulnerability analysis and estimation of local intensity

 The investigated buildings were classified into three vulnerability classes A, B and C following the EMS- 98 framework (Grünthal & Levret, 2001). Class A includes the most vulnerable buildings, which use rough rubble stones and a mortar of average quality. Class B includes the buildings using carefully cut stones or terracottas, and a mortar of good quality. Class C includes more resistant buildings using large ashlar. When the building presented visible defaults or weaknesses in its construction, we lowered its vulnerability class by one rank (e.g. A instead of B). For each vulnerability class, we computed the ratio of the total number of damage of a given grade observed over the length of the walls considered (Fig. 9a). The damage repartition is quite similar for classes B and C with a maximum for grade 2 (moderate) and limited damage for higher grades. The buildings of class A appear less affected by individual cases of damage on their total length, but are characterized by a higher severity of the damage, with a maximum for grade 3 (important non-structural damage and moderate structural damage).

 The assessment of building damage following an earthquake makes it possible to derive a local intensity reflecting the severity of ground shaking. For that, the damage distributions obtained for different building vulnerability classes is compared to empirical relationships (damage probability matrices, DPM) linking, for a given intensity, the probability that buildings are affected by damage of different grades (Fig. 9b; Riedel, 2015). In modern postseismic damage surveys, each building is given one grade of damage. In our case, the number of individual buildings that could be surveyed is too low, and long linear buildings like the defensive walls and the aqueduct cannot be considered "individual buildings". Therefore, we chose to assign a damage grade to each EAE and to consider the ratio of EAE per linear meter of wall. It is difficult to directly compare our damage ratios with the DPMs because

287 our counting method should be first calibrated with modern earthquakes, but we can still obtain qualitative estimates of intensity.

 For vulnerability class A, we obtain a flat distribution for quality 1 EAE (Fig. 9a). When lower qualities are included, we observe a larger peak for grade 3 damage and a significant representation of grade 1 damage. The grade 3 damage peak points toward a VII-VIII intensity, but our grade 1 damage is overrepresented compared to the DPM in this range. For vulnerability class B, the obtained distribution is more regular and insensitive to EAE quality, with a maximum for grade 2 followed by a progressive decrease for bigger grades. This distribution is most compatible with an intensity VIII. For vulnerability class C, we obtain a similar distribution to class B, with a clear maximum for grade 2 and much lower 296 ratios for the other grades. This suggests a VIII-IX intensity, with an overrepresentation of grade 4 damage compared to the DPM in this range. As a result, a minimal intensity VIII seems to account best for the three damage distributions (Fig. 9b).

MODELING OF THE DEFORMATION OF THE OBELISK

 The obelisk located north of Iznik was chosen to test numerically whether an earthquake rupturing the MNAF is able to cause the displacement of the blocks, as is observed today. The constituting elements of the obelisk were measured in order to build a 3D model of the obelisk (for the geometry of the model, see Fig. S2 in the electronic supplement). The numerical model was built using a non-linear transient finite-element analysis (Baillet et al., 2005). The marble blocks were assigned a Young modulus of 50 GPa, a Poisson's ratio of 0.2 and a density of 2.7. Two values of Coulomb friction coefficient between each block were tested at 0.5 and 0.7. We selected as input five three-component ground acceleration signals, recorded during four real strike-slip earthquakes (for the list and characteristics of these earthquakes, see Table S5 in the electronic supplement). The range of magnitude of these earthquakes is Mw 5.9-7.3. The signals were selected so that the distance between the recording station and the epicentre was in the same range as the distance between the obelisk and

 the closest MNAF segments (5-20 km) and with a similar site effect category (site class C in Eurocode- 8). The acceleration signals were converted into displacement signals with two integrations. Before each integration, we subtracted the mean and the trend to prevent a drift of the signals, and smoothed the high-frequency peaks with a 4-order bandpass Butterworth between 0.1 and 20 Hz. The signals were eventually resampled to the 0.5e-5 s time step of the numerical modeling approach. They were then imposed at the base surface of the obelisk. In a second set of simulations, we applied the earthquake signals twice in a row to reproduce examples of cumulative deformation (Fig. 10).

 Most of the energy of the seismic signals is focused below 1 Hz for the Düzce earthquake and 4 Hz for the others. The modal analysis of the modeled structure shows two main resonance frequencies at 7.8 and 7.84 Hz in flexional modes. The third mode corresponds to a frequency over 30 Hz, which is not excited by the selected seismic signals.

 The obelisk collapses only for the signal of Mw 7.3 Düzce earthquake (Table 3). For the other signals, the obelisk still shows irreversible relative displacement between blocks. The most significant values of displacement between adjacent blocks are obtained for the Mw 6.1 Azores Islands and Mw 5.9 Kyllini signals. At the end of the single-event simulations, these values are slightly lower than the centimeter offsets observed on the field. The relative rotations obtained between adjacent blocks are less than 1° and hardly reproduce the actual rotation component, clearly visible on the monument. At the end of the double-event simulations, the obelisk remains standing for all events but Düzce, and shows a cumulative deformation, which is twice the deformation of the single-event simulations. The final values obtained for two Azores Islands events better reproduce the field observations.

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DISCUSSION

Local intensity and associated earthquakes

 Our intensity estimation can be impacted by two biases. (1) Contrary to postseismic surveys, we did not study a representative sample of the buildings that used to stand in Iznik in ancient and medieval times. It is probable that the buildings still visible today are among the most resistant, and that the more vulnerable buildings have disappeared and were replaced by modern constructions. The empirical damage distributions suggest that the heaviest structural damage is concentrated on these vulnerable buildings, while the more resistant buildings tend to mainly bear slighter damage (Fig. 9b). Therefore, the successive cycles of repairs and reconstructions will tend to erase the traces of the heaviest damage (grades 4 and 5), while the slighter damage will be preferentially preserved. This selection bias will lead to underestimation of the intensity felt in Iznik during past earthquakes. This means that we constrain a minimal intensity. (2) Our damage dataset results from the cumulative effects of several earthquakes, and our qualitative comparison with the DPMs relies on the assumption that these earthquakes produced similar damage distributions. This seems to be the case as the obtained damage distributions are close to the empirical DPMs (Fig. 9). We can also assume that the final curve will mostly reflect the bigger events, which leave most traces of damage. Our vulnerability analysis may thus only constrain a minimal intensity, around VIII, for the biggest ruptures suffered in Iznik in the last two millennia.

 With an epicentre located on the fault segments 3 km south of the city on the MNAF segments, an intensity VIII in Iznik can be produced by a Mw 6+ earthquake (Erdik et al., 1985). The deformations observed on the obelisk suggest magnitudes over 6, probably between 6 and 7, in the case of events rupturing the same segments south of Iznik, in agreement with the vulnerability analysis. However, the modeling also shows that a stronger earthquake near Iznik would have destroyed the obelisk. We may therefore rule out the occurrence of this kind of Mw 7+ event during the historical period for the fault segments close to Iznik.

Historical seismicity scenarios

 The vulnerability analysis and obelisk modeling argue for a significant seismogenic potential of the MNAF. In this section, we discuss the deformation episodes identified by archeoseismology and examine together independent evidence gathered from the historical seismicity catalogues and paleoseismic trenching studies.

Possible deformations before the 6th century AD

368 Although it was not possible to find repairs anterior to the 8th century AD, a part of the EAE affecting the older buildings of Iznik may have been caused by earlier earthquakes. The historical catalogue includes six events with damage reported in Nicaea (Fig. 2). Three events were attributed to the MNAF, in 29-32 AD, 121 AD and 368 AD. The magnitudes estimated are around 7 and it seems that the 121 AD event was the strongest, possibly of M~7.4. While the first two events are said to have caused partial destruction in Iznik, the 368 AD event is said to have caused complete destruction. This last 374 description is likely exaggerated as several buildings from this period, such as the theater and the 3^{rd} century walls, are still standing today.

 We also need to examine the possibility that Nicaea was affected by earthquakes rupturing the NNAF. Historical descriptions do include such occurrences, e. g. in 358 AD, 362 AD and 478 AD with magnitudes estimated between 6.8 and 7.4, but the precise nature and severity of the destruction caused are difficult to determine (Fig. 2). A Mw 7.5 earthquake rupturing the fault segments crossing Izmit Gulf may produce an intensity VIII in Iznik (Erdik et al., 1985). However, during the 1999 Mw 7.6 Izmit earthquake, intensities VI were reported in the Iznik area and the local seismic station recorded a peak horizontal acceleration of 0.12g only (Celebi et al., 2000). Accordingly, the damage in Iznik was slight: the chimneys of some houses fell down or were bent and a very small part of the defensive walls collapsed (pers. comm. Cengiz Celik). Therefore, the earthquakes rupturing the NNAF are unlikely to account for a significant part of our damage dataset.

The 527-787 AD episode

 This period of time is associated with two historical earthquakes, both of them located on the NNAF (Fig. 2). In 554 AD, a Ms~7 earthquake severely affected Constantinople and Nicomedia (Izmit). Partial damage is reported in Iznik, but by one source only. In 740 AD, a major Ms 7.1 event affected the eastern part of the Marmara Sea and it is reported that "a single church survived" in Iznik which was totally destroyed. For the same reasons as exposed above, this testimony cannot be taken literally. But given that no other earthquake is known in the region, the EAE associated to this episode may result from these earthquakes. It should be emphasized that our constraints for this episode are scarce and limited to one building only.

The 858-1097 episode

396 Iznik was affected by war damage during this period of time. In the late $10th$ century, the Byzantine general Bardas Skleros led a wide-scale rebellion in Asia Minor. In 978, he besieged Iznik and some parts of the walls, especially the tower T106 in the southwest were heavily damaged (Holmes, 2005). One century later, the city was temporarily seized by the Turks and was attacked by the Crusaders in 1097 AD. Most fights happened around the southern gate, damaging the walls and destroying tower T106 (Cahen, 1948; Foss, 1996; Norwich, 1993).

 Our evidence for this deformation episode includes good quality EAE and rebuilding on tower T57 and Istanbul Gate in the northern section. The rebuilding closely postdates the only historical earthquake documented in Iznik for this time interval, in 1065 AD. With a magnitude estimated at 6.8, it caused widespread damage to churches, the city walls, and numerous houses. No other locality seems to have been damaged, which suggests a local event rupturing the MNAF. Sediment cores recently sampled in the southeastern part of Iznik Lake corroborate the occurrence of a seismic event rupturing the fault

408 segments south of Iznik during the 11th century AD (Gastineau et al., 2020 PREPRINT). We can consider that our EAE dataset most probably includes the archeological signature of this event.

The Ottoman episode (after 1330 AD)

 The traces of earthquake damage found on several Ottoman buildings indicate that Iznik was affected 412 by at least one significant earthquake since the 14th century AD. Three different scenarios can be proposed.

 (1) This damage may have resulted from earthquakes rupturing the NNAF. Several historical earthquakes with damage reported at Iznik are known (Fig. 2). In 1754 AD, a strong Ms 7 earthquake affected the region of Izmit on the NNAF, destroying a few houses in Iznik, apparently without casualties. In 1878 and 1893 AD, two smaller earthquakes in the same area of the NNAF caused slight damage in Iznik, but with much lower magnitudes, between 5 and 6. Finally, in 1894 AD, a Ms 7.3 earthquake ruptured the Izmit Gulf faults and caused liquefaction and landslides in the epicentral region and widespread destruction between Istanbul, Gemlik and Sapanca. The effects in Iznik are not well known.

 (2) Another source of damage comes from a cluster of earthquakes that affected the SNAF and MNAF 423 west of Iznik in the second half of the 19th century (Fig. 2). Most of the activity extended between Bursa and Gemlik, but it seems that some ruptures were also located closer to Iznik, probably along the faults bordering the southern shore of Iznik Lake. The biggest earthquakes happened in the Bursa region and 426 did not affect the area east of Gemlik significantly. The other events correspond to much smaller earthquakes, with magnitudes under 6. Only slight damage is mentioned in Iznik on March 3, 1893 AD.

 (3) Interestingly, the Ottoman episode matches the paleoseismic evidence of two contemporary 429 ruptures (Fig. 2), sometime between the $12th$ and $18th$ century AD in Gemlik (Özalp et al., 2013), and 430 sometime between the 14th and 18th century AD in the west of Mekece (Ikeda, 1988). It can thus be proposed that the Ottoman damage in Iznik was caused by one or several earthquakes rupturing the

432 MNAF segments between the $14th$ and $18th$ centuries AD. The trench data do not constrain the 433 magnitudes of these events, but the obelisk modeling argues against a Mw 7+ earthquake south of Iznik, which suggests more moderate and/or distant events. The historical catalogues do not present 435 any mention of Iznik between the $12th$ and $18th$ centuries AD, but it can be questioned whether this 436 reflects a true period of seismic quiescence for the Iznik region. It is known that at the end of the 13th 437 centuries, Iznik suffered significant emigration because of the Ottoman conquests in Anatolia, so much so that after the city was conquered, in 1331 AD, it was described as a ruined city only inhabited by few people (Ibn Batuta, 2012). The city eventually lost its political and strategic importance, which might explain why it disappeared from the archives of past seismicity. In comparison, important centers like Bursa or Gemlik continued to suffer damage from repeated earthquakes, such as in 1419 AD. During this period of time, the historical catalogue also includes events which were felt in Istanbul but with too little information to precisely locate the epicenter, e.g. in July 1402 and January 1656 AD (Ambraseys, 2009).

CONCLUSION

 Using an archeoseismological approach, we identified 235 traces of possible earthquake damage in Iznik, and more than 100 of them can be attributed to seismic shaking with high confidence. The different walls of the city, depending on their orientation or their age, show different patterns of damage. The damage was more severe for N-S oriented walls, while NE-SW walls show less deformation. The older constructions are also more affected than more recent walls. Three distinct episodes of earthquake damage were identified, two before the Ottoman conquest of 1331 AD, and one after. The dating of several repairs associated with earthquake damage enabled us to constrain the time of occurrence of two individual events, between 527 and 787 AD, and between 858 and 1097 AD. While the latter event may correspond to the 1065 earthquake described by ancient sources, the two others can be accounted for by different earthquake scenarios. We have estimated that the

 damage observed corresponds to a minimal intensity VIII, most compatible with ruptures along the MNAF segments of magnitudes above 6. The modeling of an obelisk a few kilometres north of Iznik also suggests magnitudes over 6, but is incompatible with a stronger, Mw 7+ rupture along the MNAF segments south of Iznik. This last result implies that the fault segments south of Iznik have not known a big rupture for nearly two millenia, which corresponds to several meters of accumulated stress. This 462 study shows that in favourable contexts, archeoseismology can be used complementarily with historical seismology and paleoseismology to refine past earthquake scenarios and contribute to a better seismic hazard assessment.

Data and Resources

 The bathymetric map of Marmara was made available by the EMODnet Bathymetry project, [\(https://www.emodnet.eu/bathymetry\)](https://www.emodnet.eu/bathymetry), using the EMODnet Bathymetry portal (EMODnet Bathymetry Consortium (2016): EMODnet Digital Bathymetry (DTM), [http://doi.org/10.12770/c7b53704-999d-4721-b1a3-04ec60c87238\)](http://doi.org/10.12770/c7b53704-999d-4721-b1a3-04ec60c87238) funded by the European Commission Directorate General for Maritime Affairs and Fisheries. The large-scale topography of the study area was obtained from the Shuttle Radar Topography Mission 1 Arc-Second Global (DOI: 473 /10.5066/F7PR7TFT). The bathymetry of Iznik Lake was provided by and can be obtained from the Turkish General Directorate of Hydraulic Works (DSI). The raw acceleration signals can be obtained from the Engineering Strong Motion Database [\(https://doi.org/10.13127/ESM.2;](https://doi.org/10.13127/ESM.2) Luzi et al., 2020). The EAE photos, the 3D displacement signals used as input for the obelisk modeling and the complete output file can be obtained on the Zenodo repository at [https://zenodo.org/record/3966773.](https://zenodo.org/record/3966773) The maps of Figures 5 and 6 are from Google Earth satellite imagery. The diagrams of Figures 8 and 9 were 479 prepared using the Matplotlib 3.3.0 package for Python [\(https://matplotlib.org/;](https://matplotlib.org/) Hunter, 2007). Supplemental Material for this article includes the lists of mentioned historical earthquakes and EAE,

 additional information on the walls chronology and obelisk model, and the complete data of wall orientations.

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- 681 **Table 1:** Sample list and radiocarbon dating in Iznik. The dating was done by the LMC14/LSCE. The
- 682 ages were calibrated with the Oxcal 4.2 program (Bronk Ramsey, 2009) using the IntCal13 calibration
- 683 (Reimer et al., 2013).

*due to the very low carbon content, this age result has a low confidence level

	Terminus post quem Terminus ante quem	$Q1*$	$Q1+Q2$ ⁺	Q1+Q2+Q3‡
2nd c.	858	1	1	1
2nd c.	1330	1	1	1
3rd c.	858	5	5	5
3rd c.	1254	6	9	9
3rd c.	1185	$\overline{2}$	$\overline{2}$	$\overline{2}$
3rd c.	1065	0	0	$\mathbf{1}$
3rd c.	1097	11	15	16
3rd c.	1254	7	8	9
3rd c.	1330	Ω	2	$\overline{2}$
527	787	3	3	3
730	1254	12	13	13
858	1097	0	1	$\overline{2}$
858	1222	1	1	1
858	1254	U	1	1
858	1330	0	2	2

Table 2: EAE constrained by post-seismic repairs, separated by quality categories.

*Quality rank 1 (good). †Quality rank 2 (average). ‡Quality rank 3 (bad).

685

List of Figure Captions

Figure 1: Morphology and geometry of the NAFZ in the Marmara region. The active fault segments,

drawn in black, were taken from Emre et al. (2011) and Benjelloun et al. (2018). The Mw>6.5

690 ruptures during the 20th century are shown in colors (Hubert-Ferrari et al., 2000; Pondard et al.,

 2007). The main cities are located with dots. NNAF, MNAF and SNAF: northern, middle and southern branches of the NAF.

 Figure 2: Chronological frieze of Nicaea's history. The colors show the main historical periods of the city: Roman (blue), Seljuk and Ottoman Turk (green) and Lascarid (yellow). The red stars correspond to historical earthquakes having damaged the city according to historical descriptions, and the white stars indicate the historical events for which the reality or amount of seismic damage is uncertain (for the complete list, see Table S1 in the electronic supplement). The two bars under the frieze refer to earthquakes identified in paleoseismological trenches done on the MNAF.

 Figure 3: Field examples of active faulting in the vicinity of Iznik. (a) Shaded DEM of the MNAF region, with the locations of the field photograph. The MNAF segments are mapped in red, and other active faults in black. The lines in the lake correspond to a 5-m interval bathymetry (see Data and Resources). The fault traces in Iznik Lake are taken from Gastineau et al. (2020 PREPRINT). (b) Triangular facets (black arrows) at the southeast corner of Iznik lake. The vertical faulting component of the MNAF is more visible in the landscape in the vicinity of Iznik Lake.

 Figure 4: Examples of EAE in Iznik. The white arrows indicate the deformation azimuth and strike. The quality category and identifier are indicated on each picture. (a) Tilted base of the defensive wall, the upper part collapsed and was rebuilt in a very distinct masonry. (b) Conjugated cracks in a vaulted

 Figure 5: Inventory and locations of the EAE in Iznik, classified by quality (a) and by age of the wall affected (b).

 Figure 6: (a) Satellite image of Iznik surroundings with the active faults drawn in red and the main aqueduct in yellow. (b) Simplified map of Iznik. The historical city is delimited by a double belt of defensive walls and used to be supplied with water by two aqueducts in the east. The other buildings investigated and mentioned in the text are drawn in grey. The numbers refer to the nomenclature of 727 the towers used in the text, following Foss and Winfield (1986).

 Figure 7: (a) General view of the obelisk near Elbeyli village. See Fig. 6a for location. (b, c) Detailed views of rotated and offset blocks. The red arrows point to the parts of the pictures where the deformation is most visible.

 Figure 8: (a) Comparison between the relative distribution of wall length (bars) and quality 1 EAE (line) for different wall azimuths. The percentages in length and EAE were computed on the total wall length and total number of EAE of quality 1, respectively. For the other qualities, see Table S4 in the electronic supplement. (b) Azimuths of the damage plotted against the azimuths of the wall affected. Note the general correlations between the two values. This correlation is linear for the in-plane deformations (i.e. parallel to the wall) and affine with a ±90 shift for the out-of-plane deformations (i.e. perpendicular to the wall). (c) Rose diagrams of the azimuths of damage observed in Iznik (118 EAE). Note the prominent N-S azimuths.

 Figure 9: (a) Distribution of the damage of various grades for the three vulnerability classes of buildings in Iznik. (b) Quantitative translations of the EMS98 damage classification according to vulnerability class and macroseismic intensity, adapted from Riedel (2015). The vulnerability class A to C correspond to the buildings studied in Iznik that use terra cotta, rubble and ashlar. The 746 probability distributions of the different damage grades are shown with different colors. For each vulnerability class, we try to find the intensity which best replicates our observed damage distribution (reported on the graphs as vertical bars). For example, for our vulnerability B buildings, we observe a majority of D2 damage. This is compatible with an intensity VII or VIII. Then, the fact that we have much more D3 than D1 damage rather suggests an intensity VIII.

 Figure 10: Modeled N displacement (a), E displacement (b), and rotational displacement (c) of the obelisk for the two-event simulation of the Mw 6.1 Azores Islands earthquake. The blocks are numbered from base to top.

Historical earthquake scenarios for the middle strand of the North Anatolian Fault deduced from

archeo-damage inventory and building deformation modeling

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Supplemental Material

Table S1: Catalogue of earthquakes with effects in Iznik and neighbouring areas. Ms=surface wave magnitude. For the magnitude and epicentre location, we provide the different ranges found among the references.

Table S2: Description of the damage EAE observed in Iznik area.

Table S3: Description of the repair EAE observed in Iznik area.

Table S4: Length of walls and EAE of different qualities for each wall orientation range. The percentages in length and EAE were computed on the total wall length and total number of EAE in each quality class respectively.

Table S5: Characteristics of the earthquake used for the simulations on the obelisk.

Figure S1: Summary of the construction history of Iznik walls, after Schneider and Karnapp (1938) and Foss and Winfield (1986). After the construction of the walls in the 3rd century AD, the successive interventions and restoration works are drawn in color. The state of the walls preceding each intervention is drawn in grey.

Figure S2: (a) Drawing of Iznik obelisk with the dimensions and arrangement of the different blocks as measured on the field. The western face is given as an example on the left. (b) 3D mesh of the obelisk used for numerical modeling. (c, d) First and second main vibration modes for eigen frequencies of 7.8 and 7.84 Hz respectively. The initial shape of the obelisk is drawn in black, and the distribution of motion in color.

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References

1=Ambraseys (2002) 2=Ambraseys (2009) 3=Guidoboni et al. (1994) 4=Ambraseys and Finkel (1991) 5=Ambraseys (2000) 6=Ambraseys and Jackson (1998)

40.7 27.5-29.6 very well constrained North

contemporaneous and the contemporaneous contemporaneous and the contemporaneous contemporaneous and $1, 2$

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Damage EAE dataset

- O1 40.487589 29.701961 Obelisk
- B1 40.42575 29.7095 Lake basilica

Vertical crack + slight shift 90 Shifted and rotated blocks Destruction

0

left-lateral 1387

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Reconstruction EAE dataset

Table S4 [Click here to access/download;Supplemental Material \(Main](https://www.editorialmanager.com/srl/download.aspx?id=228756&guid=2a849014-eec4-492f-8e3b-95407116097d&scheme=1) Page, Tables, and Figures);TableS4.xlsx

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Michael III reconstructions (c. 858

Palaeologue reconstructions (1254-1331)

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