

ANALYSIS OF THE KUROSHIO EXTENSION DECADAL VARIABILITY : Internal Mechanisms and External Forcings

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Why study the Kuroshio Extension (KE) from a Climatic Point of View?

•Highest eddy kinetic energy level (EKE) in the North Pacific Ocean. EKE changes lead to high temperature anomalies that are capable of enhancing the variability of the midlatitude coupled ocean– atmosphere system.

•The KE is known to undergo relevant variations on the decadal time scale. The mechanisms that generate the mean structure of the KE and its low frequency variability (LFV) are still under debate.

Bimodal Behavior

Introduction

- Synchronization with external forcings (Qiu and Chen, 2005,2010).
- Many authors have suggested that nonlinear intrinsic oceanic mechanisms can play a fundamental role in the phenomenon (Pierini, 2006).
- The KE LFV is interpreted as a case of intrinsic climate variability paced by external forcing through the coherence resonance mechanism (Pierini, 2014).

QUESTION

What is the role of the intrinsic variability and external forcing in driving the KE LFV?

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METHOD

To explore a *large* ensemble of ocean-only eddy-permitting model simulations: OCCIPUT

OCCIPUT*

KE LFV Indices after post-processing (Nonlinear detrending filter and Lowess Low Pass Filter)

Ind. menters - Ensemble-mean +/- Ensemble-\$td Latitude بياريه Reference SSH isoline of the jet 0.7 meters -2400 **LKE UVS Average in the box** (141-153°E, 32-38°N) EKE **Max total geost. velocity in the box KE Velocity** (141-153°E,30-40°) integrated zonally **FREES**

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Gaussian metrics are appropriate since the climatological PDFs for these variables are pretty gaussian

We take advantage of some definitions and notations (Leroux et al., 2018):

 $f_i(t) = (f_i(t)) + f'_i(t)$

We define $f(t, f)$ as every index after nonlinear detrending, where t stands for the time (yr) and i is the index of the ensemble member. We use overbars for the time-mean operator and angled brackets for the ensemble-mean operator.

$$
\begin{cases}\nA_{intrinsic} = \sqrt{\varepsilon^2} \\
A_{forced} = \sigma_{\langle f \rangle} \\
A_{tot} i = \sigma_{fi} \\
R = \frac{A_{intrinsic}}{A_{forced}}\n\end{cases}
$$

 $\sigma^2(f)$: the time variance coerator applied to the crisemble mean-

 $\varepsilon^2(t)$: ensemble variance operator applied to $f'(t)$ $A_{\ell x, \ell y \ell y \ell z}$: Amplitude of the intrinsic variability $A_{f\text{srrad}}$: Amplitude of the forced variability A_{cot}: Amplitude of the total variability in each member R: Amplitude Natio

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OCCICLIM

STRATEGY:

TO CREATE A PSEUDO-ENSEMBLE FROM OCCICLIM

50 members x 36 yrs have been built randomly from the indices already computed, in order to compare the pure intrinsic variability with the total one given by OCCITENS.

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CONCLUSIONS

- OCCIPUT ensemble simulation : disentangle the intrinsic and forced variabilities
- The KE LFV is a case of intrinsic variability paced by an external forcing
- The external forcing acts reducing the entropy of the dynamical system

WORK IN PROGRESS

What is the ability of state of the art models in reproducing the observed KE LFV: the role of horizontal resolution -PRIMAVERA project- Model data from HIGHRESMIP-CMIP6

** Model data are from the H2020 PRIMAVERA Project

Fedele et al., in preparation

EGL

THANKS FOR YOUR ATTENTION

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