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Abstract:

The southern part of the French Alps is studied for years in mapping and understanding of large-scale gravitational deformations. The identification and the knowledge of large-scale slope deformation (deep-seated gravitational slope deformation or DSGSD and deep-seated landslide or DSL) in the previous work of Jomard (2006) and Zerathe (2013) open a new vision of landslide processes with the reinterpretation of their dynamics and the characterization of the time scales involved. We identify DSGSD (10.109 m³) in the Var Valley associated with geological and geomorphological anomalies linked to the Messinian Salinity Crisis (MSC) and the alpine orogenesis. We use field observations, geological information (geological map, boreholes), and topographic analysis performed in a GIS environment in order to describe these anomalies. This old and partly eroded slide mass is associated with three typical DSGSD features: (1) a double-crested ridge, the Sinne Valley, (2) a large formation (2.7 × 10⁸ m³) of slope deposit dated from the Messinian (Carros breccia), and (3) the kilometeric deviation (1 to 2 km) of the Var River. We relate all these anomalies to the MSC (5.97 to 5.46 Ma) and the incision of deep canyons during this period related to this major eustatic variation (≈ 1300 m). The incision of the canyon triggered the collapse process of the DSGSD of the Sinne Valley and so destabilized the entire massif. At present, three DSLs resulting from the DSGSD deformation are still present in the area in a dormant state. Indeed, since the infilling of the Var Canyon during the Pliocene, the activity of the DSGSD has stopped.

Keywords: Deep-seated gravitational slope deformation. Subalpine chains. Messinian Salinity Crisis. Gravitational slope evolution. Eustatic variation.

I- Introduction

For several decades, studies on gravitational phenomena have focused on the geomorphological characteristics and dynamic processes of large landslides to better understand the phenomenon and especially their evolution and their impact on the landscapes. The role played by landslides in mountain ranges is varied, but they are above all a vector of erosion and mass displacement. Our studied gravitational movements are deep-seated gravitational slope deformations (DSGSDs) (greater than 10⁹ m³) (defined by Malgot 1977), and they can involve an entire hill slope. They are complex and require precise three-dimensional or even four-dimensional investigations with time considered. Yet these landslides are still less understood than compared with deep-seated landslides (DSLs) (10⁶ to 10⁷ m³) (Pollet 2004) and we still have to learn how certain DSGSD affect the overall geomorphological evolution of landscapes.

Several authors have linked the stability of these landslides to the long-term evolution of mountain ranges topography (e.g., rock uplift, exhumation rates, and river incision) (Bovis 1982; Crosta and Zanchi 2000; Agliardi et al. 2001; Ballantyne 2002a, b; Korupet al. 2007; Larsen et al. 2010; Larsen and Montgomery 2012; Blöthe et al. 2007; Larsen et al. 2010; Larsen and Montgomery 2012; Blöthe et al. 2015). The localization of DSGSD is strongly controlled by the structural inheritance, the topographic constraints (Radbruch-Hall 1978; Varnes et al. 1990), and the geology (Jomard 2006; El Bedoui et al. 2009; Zerathe and Lebourg 2012). It is also recognized that the seismic activity can control the destabilization or the reactivation of DSGSD (e.g., Beck 1968; Harp and Jibson 1996; McCalpin and Hart 2003; Gutiérrez-Santolalla et al. 2005; Hippolyte et al. 2006; Gutiérrez et al. 2008).

Many DSGSDs in the Alps were triggered in the Holocene period (Tibaldi et al. 2004; Hippolyte et al. 2006; Agliardi et al. 2009; Hippolyte et al. 2009; Pánek et al. 2011; Hippolyte et al. 2012; Jomard et al. 2014) due to valley incision (Bovis and Evans 1996) and flank decompression related to glacier melting (Agliardi et al. 2001; Hippolyte et al. 2006, 2009; Agliardi et al. 2009; Crosta et al. 2013). Some of the largest examples of DSGSDs are in the Alps, like the Flims landslide (8.5 ka) (volume = 8.5–12 10⁹ m³) (Stead and Wolter 2015), or the Sierre landslide (9.6 ka) (volume = 1.5 10⁹ m³) (Pedrazzini et al. 2013). In the central Italian Alps, Frattini et al. (2018), b) highlight by synthetic aperture radar (SAR) interferometry that a majority of the DSGSDs in the region are still active at present. Due to the glaciation/deglaciation cycles, it remains difficult to observe much older DSGSD in the Alps.

The glaciation/deglaciation cycles are not the only events that can trigger DSGSD. Indeed, at present time, Hou et al. (2014) concluded that the primary destabilizing factors of DSGSDs of the Kii Valley in Western Japan are (1) the degradation of rock masses and (2) the incision of the Yoshino River at the bottom of the slope. The incision alone is insufficient to trigger the slide. Furthermore, Tsou et al. (2015) found a relationship between the incision and the development of DSGSD in the mountains crossed by the Dahan River in Northern Taiwan. These observations are also present in the Mediterranean area; Buccolini et al. (2010) and Gargani et al. (2014) highlight the link between marine regressions and slope destabilization resulting from valley incision. This work demonstrates the importance of the valley incision factor in the establishment of DSGSDs.

Still in the Mediterranean area, the Messinian Salinity Crisis (MSC) represents the most relevant valley incision events that ever occur in the region (Clauzon 1979; Clauzon 1982; Lofi et al. 2005; Maillard et al. 2006; Bertoni and Cartwright 2007; Cornée et al. 2008). This eustatic event was characterized by an almost complete desiccation of the Mediterranean Basin after its disconnection from the Atlantic Ocean at 5.97 Ma (Krijgsman et al. 1999). The Messinian Salinity Crisis (MSC) lasted from 5.97 to 5.46 Ma (Gautier et al. 1994; Clauzon et al. 1996a; Krijgsman et al. 1999; Bache et al. 2012; Manzi et al. 2013) but the sea level drops down by about 1500 m compared with present levels

(Gorini et al. 2005) only at 5.6 Ma (Clauzon et al. 1996a). Every river connected to the Mediterranean Sea underwent intense incisions from regressive erosion that penetrated inland in order to adjust to the new base level. Incision rates of up to 10 mm/year have been observed in certain parts of the Rhône (Loget et al. 2006). In France, we observe strong and steep incisions in the principal Mediterranean rivers, the Rhône (Gorini et al. 2005; Lofi et al. 2005; Mocochain et al. 2006), the Durance (Clauzon 1979; Hippolyte et al. 2011), and the Var (Clauzon 1978; Savoye and Piper 1991). The conditions encountered during the MSC are therefore suitable for the establishment of large-scale gravity deformations along the Messinian Canyons. The end of the MSC is marked by the flooding of the Mediterranean Sea which may have been very fast, ranging from a thousand years (Hsü et al. 1973) to only 36 years (Blanc 2002), or may also have occurred in several episodes with different filling speeds more spread over time (Bache et al. 2012). During this flooding, the Messinian Canyons were transformed into deep rias that were filled by sediments prograding in Gilbert-type fan deltas (Clauzon 1990).

In this paper, we present a new case study, the Sinne Valley DSGSD in the lower Var Valley one of the largest mass movements ($\approx 10.109 \text{ m}^3$) but especially one of the oldest in the European Alps (5.6 Ma). The particularity of this Alpine DSGSD is that it has not been triggered by the glaciation/deglaciation cycles but by the creation of the Var Canyon during the MSC. With the help of topographic and geologic analyses, and field observations, we have identified several geological and geomorphological features linked to large-scale gravitational deformation related to alpine orogenesis and canyons creation during Messinian Salinity Crisis. Finally, we describe the complete evolution of an onshore Alpine DSGSD triggered by the MSC.

II- Geological setting

The study area is in the lower Var Valley in the southern Western Alps in the Alpes-Maritimes department (southeast France). The Var River is the largest river of the Alpes-Maritimes; it is 114 km long and its catchment area extends on about 2600 km². The lower Var Valley is defined as the Var sedimentary basin composed of Pliocene conglomerate that filled the Var Messinian Canyon (Fig. 1 and Fig. 2) created during the MSC. This territory includes the southernmost mountain range of the French Alps resulting from two main thrusts: The Castellane arc to the West and the Nice arc to the East (Fig. 1). These two mountain ranges are the outcomes of the exhumation of the Argentera-Mercantour Massif (Fig. 1) during the Miocene and Pliocene (Laurent 1998) and are composed of Mesozoic to Cenozoic folded sedimentary rocks. The regional slope created during this exhumation allows the series to be detached from the crystalline basement via the Keuper level (Upper Triassic), a geological formation that acted as the “soap layer” (Sanchez 2010). From bottom to top, this geological layer is composed of white dolomites interbed with varicolored marls and plastic clays, cagneules, large dolomites, and gypsum layers (Dardeau 1983). The Keuper formation has weak mechanical properties that localize deformation (Young E modulus $\approx 40 \text{ GPa}$ and compressive stresses $\sigma_c \approx 36 \text{ MPa}$, Fabre and Dayre 1982) (Fig. 2). The Castellane and Nice arcs did not evolve synchronously. The southern branch of the Castellane arc is essentially a Miocene structure (Clauzon 1978; Laurent et al. 2000) with a frontal thrust over the Tortonian formation (Miocene) (Irr 1984) (Fig. 2), while the Nice arc shows evidence of Late Neogene thrusting orogeny (Late Pliocene to Early Quaternary) (Clauzon 1978). Still both units underwent the same regional Quaternary uplift.

This study focuses on the Carros and Gattières areas in the Castellane arc (Fig. 2) where we observed geomorphological and geological anomalies.

The main tectonic structures resulting from the thrusting of the Castellane arc are strike-slip fault, oriented N20–40° and dipping 80–90° toward the SE that moved with a sinistral horizontal sense (Fig. 2a). In the Alps, Ambrosi and Crosta (2006) demonstrated a relation between major structural features and DSGSDs. In the Marbrières Range of the Castellane arc (20 km eastward), Zerathe and

Lebourg (2012) observed abnormal vertical shifts and typical gravitational morpho structures on a major N20–40° fault, which suggest the presence of a general gravitational reactivation of the network of inherited faults, with reactivation evolving to scarps or trenches. The lower Var Valley runs through a Gilbert-type fan delta that includes marine bottomset and foreset beds (Clauzon et al. 1990), named “conglomerate of the Var delta” (Dubar 2012) who filled most of the Messinian Var Canyon (Fig. 2). Seismic profiles show that the incision of the Var could be large, since the Messinian Erosional Surface is at 900 m depth near the present coast (Savoye and Piper 1991). To the north, the deep Messinian incision is still attested by 100-m-thick marine Pliocene conglomeratic foreset beds at Clans, more than 30 km upstream from the present coast (Vernet 1982; Clauzon et al. 1996b).

The upper part of the western slope of the Var Valley, the Castellane arc, consists of sub-horizontal Jurassic limestone (dip \approx 5–20°) lying on deformed and thick (> 200 m) Triassic formation (Fig. 2c). The presence of the Keuper evaporate makes the leading edge of the Castellane thrust extremely favorable for large landslides (DSGSD, DSL) (Zerathe 2013). Indeed, the Jurassic limestone is above plastic rocks, vulnerable to the dissolution because of the gypsum. The other outcropping main unit is thick (\approx 200 m) Miocene formation dated from middle Aquitanian to Burdigalian and composed of sandy and micaceous marls (Fig. 2). This formation is representative of a shallow sea depositional environment (Ginsburg et al. 1980).

This last main sedimentary unit is the Carros breccia (late Miocene to early Pliocene in age) a consolidated breccia deposited all along the hill slope and described only in this area (Fig. 2). This breccia lies unconformably on Jurassic and Miocene formations (Fig. 2) (Guglielmi 1993; Ginsburg et al. 1980) and in the upper levels, microbreccias and marl interbeds have been found, revealing Pliocene marine microfauna (Guebhardt 1903).

Finally, a large part of the area is covered with Quaternary scree deposits on steep slopes. Most of these slope deposits are consolidated and covered by vegetation (Ginsburg et al. 1980).

There is no glacial evidence in the southern part of the middle Var Valley. The nearest moraine deposits are located more than 20 km north of the study area (Fig. 2) (Julian 1980; Maisch et al. 1999; Jorda et al. 2000; Darnault et al. 2012). The region presents several DSLs like the La Marbrière Landslide dated from 400 ka triggered by inherited geological anisotropy, and 6 dated from 4.2 ka, triggered by intense climatic events (Zerathe 2013). In terms of regional seismicity, Larroque (2009) shows that the faults of the area can produce earthquakes with magnitude up to 7. This magnitude is enough to trigger DSLs within a 100 km of the epicenter (Keefer 1994; Tibaldi et al. 1995; Jibson et al. 2004, 2006) and can have an impact on DSGSD (Moro et al. 2007).

III- Methods

We based our analysis on the combined use of GIS and field observations. Digital Terrain Model (DTM) and vectorized 1/ 50000 geological map (BD Charm-50 produced by the BRGM available at <http://www.geocatalogue.fr/Detail.do?id=4156> (last accessed on 29/01/2019) analyses are used to perform large-scale geomorphological observations followed by more precise field surveys. The field surveys can confirm, disprove, or detail specific features with the help of GNSS locations, photos, and dip measurements.

The DTM used in this paper is provided by the Conseil departmental des Alpes-Maritimes (available at <https://trouver.datasud.fr/dataset/modeles-numeriques-de-terrain-go-06> (last accessed on 29/01/2020)) with a resolution of 5 m. DTM does not consider the vegetation and can cover the entire department; it allows a spatial 3D representation of the study area through 3D view and elevation profiles and allows many observations that are not possible in the field.

Slope, surface roughness, aspect, hill shades, and curvature are used to describe morphological features like major and minor scarps, ridges and cracks, and landslides boundaries (Morelli et al. 2018). Slope, aspect, hill shades, and curvatures are estimated in ArcGIS 10.3 (Esri, Redlands, CA, USA) using

the spatial Analyst extension, and the surface roughness is estimated at QGIS 2.18 (QGIS development Team 2019) using the Raster Terrain Analysis Plugin. Since slope and surface roughness present the best contrasts in the analysis of DTM data, particularly for escarpments, these two were performed on this study. According to Burrough and McDonnell (1998), the slope is calculated at a given cell by comparing the altitude to its neighbors (3×3 cells), and it represents the maximal variation rate of altitude (from 0° to 90°). According to Riley et al. (1999), the surface roughness is estimated in each cell by calculating the sum change in the altitude of its neighbor (3×3 cells) (flat area has a value close to 0).

Elevation profiles are realized with the QGIS DTM surface plugin Profile tool (developed by Jurgiel et al.). 3D views of the area are realized with the QGIS plugin Qgis2threejs (developed by Akagi), who combine DTM data and images of other data like geology, orthophotos, or surface roughness. Finally, the r.volume model of GRASS plugin (Hinthorne 1988 updated to GRASS 7 by Landa) is used to calculate the volume of the DSGSD resulting from the DTM and the surface area of the DSGSD.

IV- Results

1) The Sinne Valley

The Sinne Valley is a deep valley (≈ 100 m) perched at 700 m high with no permanent watercourses. Its $N20-40^\circ$ orientation and the presence of a strike-slip fault (Fig. 2) point out to a tectonic origin. This family of strike-slip faults is linked to the thrusting of the Castellane arc and ceased their activity in late Miocene.

A set of different gaps or crevasses, 1 to 2 m wide and 1 to 3 m deep, all-oriented $N20-40^\circ$ and tracked on about 100 m was also observed in the Sinne Valley showing NW-SE extension (Fig. 3c).

The Sinne Valley is bordered by two plateaus, one to the east and the other to the west over a distance of ≈ 1.5 km (Fig. 2c, b and Fig. 3b). These two plateaus are composed of the same rocks, the Portlandian and Berriasian lithographic limestone (Fig. 2b, Fig. 3b). We measured bedding planes of the Jurassic formation. On the western plateau, they strike from 205 to 255° and dip from 5° to 20° (14 measurements) (Fig. 3e). They are consistent with those of Baggioni (1971) who described the formation of the summit as sub-horizontal including the plateaus. The bedding planes at the lower part of the Jurassic formation of the Carros/Gattières hill slope strike from 175° to 205° and dip from 25° to 40° (8 measurements) (Fig. 3e). These are consistent with those of Zerathe (2013) (Fig. 3a). Using the DTM and a stochastic approach that traces ten topographic profiles passing through the two plateaus (profiles \pm NW-SE), it appears that the eastern plateau bordering the Sinne Valley has an average elevation of 925 ± 15 m, while the western plateau has an average altitude of 1015 ± 6 m, with a difference of approximately 90 ± 10 m (Fig. 3b). In addition, we perform an analysis on the geological map combine with the DTM and a 30 m stratigraphic offset is observed in the Portlandian- Berriasian (J7), Kimmeridgian (J6), and Sequanian (J5) layers (Fig. 3b). With an approximate 10° dip (Fig. 3b), the offset of the Jurassic series across the valley can be estimated at about 70–80 m, which is close to the 90 m offset between the two plateaus.

We infer that the shape of the Sinne Valley can be assimilated with a “double-crested ridge” (Fig. 3b), a geomorphological structure characteristic of DSGSDs (Agliardi et al. 2012).

2) The Carros breccia

Along the Var Valley, the Carros breccia is only present in the Carros/Gattières area and corresponds to a thick and highly consolidated slope formation. According to the Roquesteron Geological Map (Ginsburg et al. 1980), the Carros breccia was deposited both at the front of the Castellane thrust and at the edge of the Var Messinian Canyon. Gèze and Nestéroff (1968) described the breccia as synorogenic when the Castellane arc was emplaced. But Clauzon (1978) pointed out that at Le Broc borehole (from 98 to -12 m elevation) (Fig. 2b), the Carros breccia is at least 70 m thick

(from - 12 to 57 m elevation), and is overlain by Pliocene marls (from 57 to 68 m elevation) and Holocene alluvium (from 68 to 98 m elevation). He concluded that the breccia results from the Messinian incision of the Var River, and is not a syn-tectonic breccia because it postdates the emplacement of the Castellane arc. At Le Broc (Fig. 2), we observed northeast prograding subaqueous foreset beds of debris (Fig. 4a) with marine burrows on the limestone blocks (Fig. 4b) which confirms that the breccia continued to deposit after the flooding of the lower Var Valley at the end of the MSC. At present days, the outcropping surface of the Carros breccia is 3.8 km², with an approximate thickness of 70 m (i.e., le Broc borehole) and it represents 2.7×10^8 m³. This volume is an underestimation since the Pliocene conglomerate of the Var delta and Holocene formation may cover it (i.e., le Broc borehole). Anyway, this formation indicates that the western slope of the Var Valley has been locally subject to high degradation during and after the Messinian incision.

3) The DSLs

According to El Bedoui et al. (2009), Hradeck and Pánek (2008), and Jomard et al. (2014), DSGSDs may induce rapid localized failure events like DSLs especially on the slope toe; this interaction can be explained by the long-term deformations of DSGSDs and the resulting steepening of the slopes. Zerathe (2013) already documents a DSL in the area: the Gattières DSL (1 Fig. 2b, Fig. 3a). This DSL has an estimated volume of 1.5×10^6 m³ and a surface of 0.08 km².

On a three-dimensional view of the roughness of the area (Fig. 3e), we can distinguish the presence of three triangular facets on the slope. According to Strak (2012), this geomorphological structure is representative of active normal faults, the southernmost triangular facet corresponding to Gattières DSL (Fig. 3a). This analysis was followed by field campaigns confirming the presence of a fault represented by a scarp that extends for about 2 km of length along the hill slope (Fig. 3e). This escarpment has many karstification indices with many concretions of speleothems covering the whole surface in some places and various evidences of deformation like NW-SE extensional crevasses of different sizes (Fig. 3d). Field campaigns, DTM analysis, and geological map analyses have enabled us to define that the Gattières DSL and two other DSLs are linked to the triangular facets and the so-called Gattières/Carros fault, the Ripert DSL (Fig. 2 and Fig. 3e with a volume estimated at 1.8×10^6 m³) and the Carros DSL (Fig. 2 and Fig. 3e with a volume estimated at 1×10^7 m³). Due to their lengths and shapes, the Gattières/Carros fault and the DSLs are very likely due to the DSGSD. DSGSD activity can create faults along the hill slope to accommodate the deformation, and the fault and the inward dip of the Jurassic formation contribute to fluid circulation to the Triassic gypsum formation whose dissolution leads to the subsidence of the massif and more unstable conditions.

According to Zerathe et al. (2014), who performed cosmic ray exposure dating (CRE) on six DSLs of the region, including the Gattières DSL, their triggers are dated around 4 ka. These datings are correlated with the 4.2 ka climate events, an intense paleohydrological pulse in the Mediterranean area (Zanchetta et al. 2016). These datings correlate the triggering of the DSLs with intense and major climatic events, we will consider that in the absence of such events and of recent deformations observed in the field, the DSLs are considered dormant or suspended.

4) The Var River deviation

The most striking geomorphological structure of the area is the deviation of the Var River toward the east over 8 km north-south between the towns of Gattières and Carros, with a maximum shift of 1 to 2 km toward the east (Figs. 2b, c and 3). Several authors agree that the Var took a much N-S straighter course in the lower Var Valley during the MSC (Savoye and Piper 1991; Clauzon 1975; Dubar 2012). We propose that this deviation is related to the DSGSD of the Sinne Valley. This hypothesis is supported by the long-term evolution of the DSGSD and DSLs, which have continuously

brought large quantities of gravity materials (i.e., Carros breccias, $2.7 \times 10^8 \text{ m}^3$) to the right bank slope of the Var Valley from the Messinian Canyon to its Pliocene infill.

5) Evolution of the Sinne Valley DSGSD

During the late Miocene period, the front of the Castellane arc overlapped the Miocene formations. The Keuper was overthickened because of thrusting and the Jurassic formations resting on the Triassic were tilted northward by about $5\text{--}20^\circ$ and acquired the present inward dip (Fig. 5A). The overlap phase formed a N20–40 strike-slip fault that contributes to the tectonic deformation in the region (Fig. 2, Fig. 5).

During the MSC when the canyon reached a certain depth (at present the canyon is approximately 300 m depth in the study area), slope-undermining phenomena and constraints applied to the right flank of the slope took advantage of the major tectonic fault in the Sinne Valley—a highly altered zone—thus triggering the DSGSD and the related normal faults to accommodate the deformation (Fig. 5B). The Keuper evaporites and clays may have played an important role in the setting up and dynamics of these landslides (DSGSD and DSLs). The counter dip of the Jurassic layers promotes fluid circulation inside the massif, while the normal faults are weakness areas that promoted deep water circulation (Barton et al. 1995) within the Keuper evaporites. The fluid circulation could lead to the dissolution of the gypsum layers and the modification of the mechanical properties of the clay layers of the Keuper Formation, with both phenomena being related to the diminution of stability (Baron et al. 2004; Guerrero et al. 2004; Azañón et al. 2005, 2010; Gutiérrez et al. 2012).

The DSGSD may have proceeded at a very low speed over thousands of years, but the constraint applied to the hill slope was large enough to produce numerous DSL-type landslides, thus allowing the deposition of the Carros breccia during the late Messinian incision (Fig. 5B). Moreover, the crevasses in the Sinne Valley are the result of this activity and that demonstrates that the NW-SE extension was able to develop deep in the Jurassic formations.

These constraints continued during the marine flooding of the Var Valley at the end of the MSC when the Var Canyon was transformed to Ria. At the same time, the exhumation of the Argentera-Mercantour stopped, marking the end of the Castellane arc activity (Fig. 5C). The flooding of the lower Var Valley probably triggered a second phase of activity for the DSGSD and DSLs presented in Fig. 6 by increasing the pore water pressure and thus destabilizing the slopes. This effect may have been greatly accentuated when the water sea level reached the Keuper Formation (Fig. 5C) due to the clay and gypsum rocks being highly vulnerable to the presence of water.

In the Var Valley, the marine conglomeratic foreset beds indicate the elevation of the sea level during the Pliocene. Because the Alps and their foreland have been uplifted since the Pliocene, Pliocene marine deposits are present at various elevations. Despite Quaternary erosion, the Pliocene marine conglomerate at Carros is still present up to 350 m elevation, which is a minimum for the elevation of the Pliocene sea level at this place. At Aspremont, on the eastern flank of the Var Valley, marine conglomeratic foreset beds are present up to 650 m elevation. These elevations are partly due to local tectonic deformation, but they show that Pliocene flooding could have reached the elevation of the thrust detachment in the Keuper formation.

Thus, during the deposition of the Var conglomerate and the Pliocene high stand phase (Fig. 5C) in the lower Var Valley, we assume that the DSGSD and DSLs experienced sustained gravitational activity and brought large numbers of slope breccias to the foreset beds of the Pliocene Gilbert-type delta (Fig. 5C). With the progressive filling of the Messinian Canyon by Gilbert-type delta, the DSGSD had a more stable condition since the thick conglomerate acted as a buttress. We do not know when the DSGSD activity stopped, but we assume that the activity of geomorphological structures with deformations such as crevasses (Fig. 3) has ended before 1.7 Ma.

At present, the DSGSD in the Sinne Valley is still in place, due not only to its size but also to the post-Miocene uplift of about 280 m in the Castellane arc and the quaternary incision of the Var River. An estimation of the DSGSD's volume by the r.volume plugin gives $9.5 \cdot 10^9 \text{ m}^3$ from the Sinne Valley to the Miocene formations (Fig. 5D). The gravitational material resulted from the DSGSD and DSL activity represented by the Carros breccia (estimated at $2.7 \cdot 10^8 \text{ m}^3$ at present days) was an obstacle to the flowing of the Var River during the regional uplift (Fig. 5C) (Clauzon 1996; Schroëtter 1997). This hypothesis of a deviation of the Var River owing to the contribution of gravitational material from the slope is in agreement with the presence of the DSGSD and DSLs. There are no active evidences of deformations observe for the DSGSD and the crevasses on the Sinne Valley were most likely exhumed by erosion. There is also no evidence of active deformations for the three DSLs and the regional dynamics observed by Zerathe et al. (2014) imply major climate events for their triggering so in the current state DSLs are considered dormant.

Since there is no evidence of any deformations link to the DSGSD or the three DSLs, we did not perform any activity measurement in the area. The combination of 5 m resolution DTM analysis and field checking seems to be suitable in particular for the detection of escarpments, which are highlighted by roughness measurements especially over an area of more than 20 km². This technique could easily be applied to the entire front of the Castellane arc to highlight escarpments that could be linked to DSLs. Moreover, due to the unfavorable topography and the land use on the overall slope, it was not possible to set up geophysical profiles such as electrical resistivity tomography (ERT) profiles or seismic reflection profiles. Those techniques would have allowed us to confirm or disproved, and visualized, the presence of the gravity faults highlighted in this paper instead of just relying on the DTM analyses.

IV- Conclusion

In this study, we highlight that the lower Var Valley records gravitational deformations since 5.6 Ma. We provided several arguments in favor of the presence of a DSGSD of about 10 billion m³, the largest and the oldest landslide in the front of the Castellane arc and in the Alpes-Maritimes department.

This DSGSD results from a gravitational process that occurred during the MSC when the Mediterranean Sea level dropped of about 1500 m (compared with present levels) and induced the incision of the Messinian Var Canyon that later was filled by conglomerates during the Pliocene (Fig. 5).

This deformation affected the entire massif and allowed the establishment of DSLs (Carros, Gattières, and Ripert) which actively participated in the creation of the Carros breccia.

The proposed evolution of the DSGSD follows the events that influenced the lower Var Valley over millions of years and explains in a coherent way the various observations and geomorphological characteristics found at present including the Sinne Valley, the Carros breccia, and the Var River deviation.

Landslides and collapse processes are used to be studied in recent times with active signs of deformations but at present days, there is no evidence of DSGSD or DSLs activities in the area. However, in this work, we show that it is possible to observe very old stigmas of collapse protected from erosion by their size with a dynamic driven by tectonic and gravitational events.

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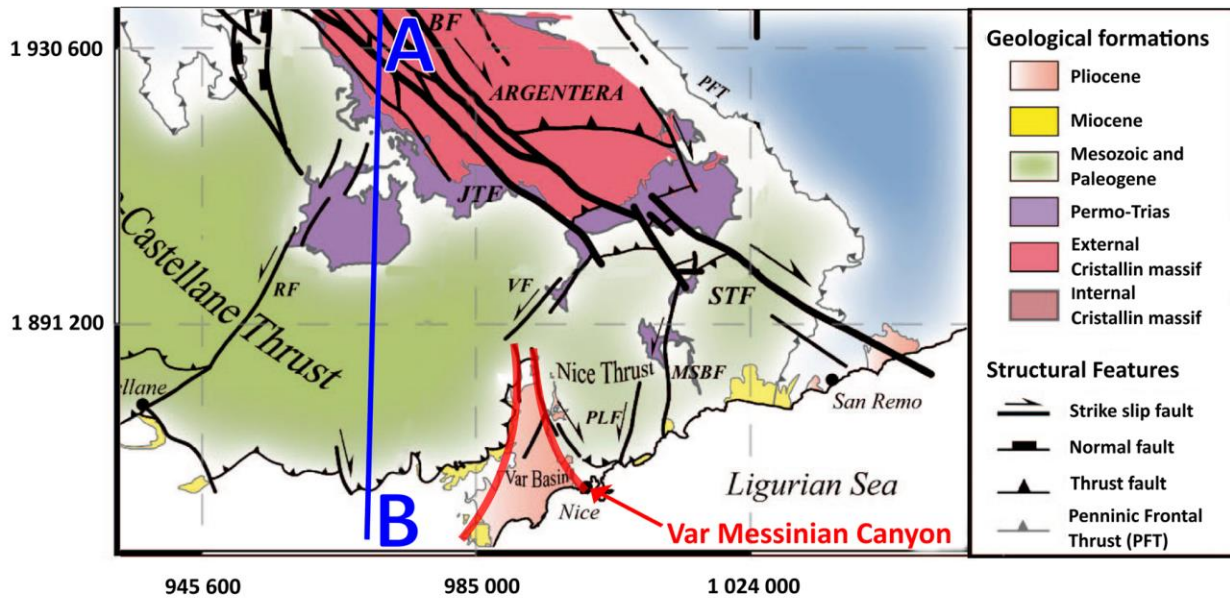
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Fig. 1: (a) Structural map of the Western South Alps (Sanchez 2010) and recent fault systems (in bold, proposed Holocene activity) (Sanchez et al. 2010). (b) Cross-section A–B: Simplified North-South section across the Western South Alps (Sanchez 2010), modified from Laurent (1998).



Description of Figure 1a **BF**: Bersézio fault; **DCFT**: Digne-Castellane Front Thrust; **JTE**: Jausier-Tinée fault; **MSBF**: Monaco-Sospel-Breil fault; **NT**: Nice Thrust; **PFT**: Penninic Frontal Thrust; **PLF**: Peile-Laghet fault; **STF**: Saorge-Taggia ault; **VF**: Vésubie fault. The blue line corresponds to the layout of the exposed section Figure 1b.

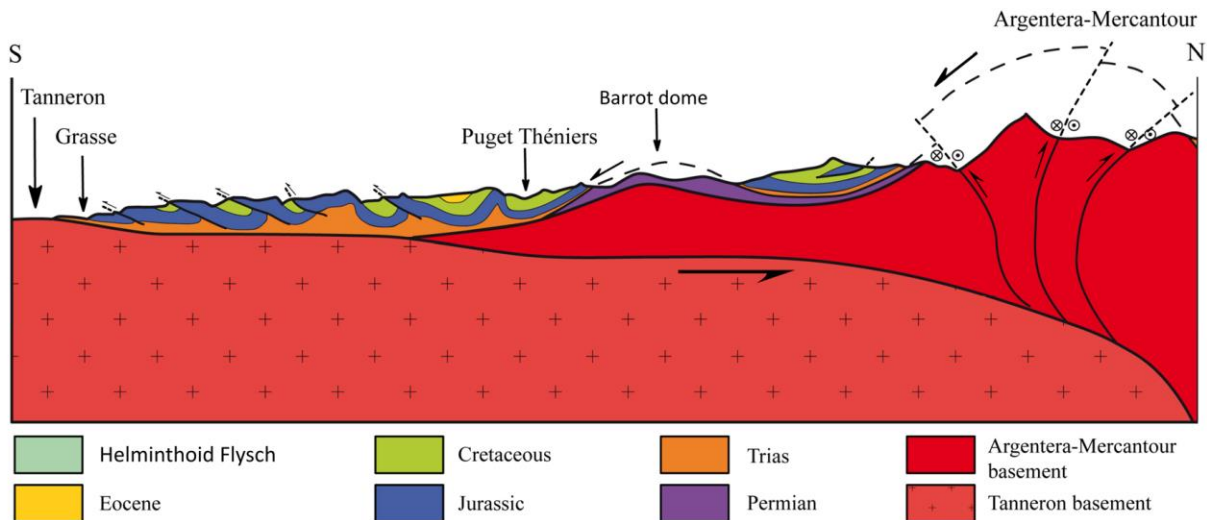


Fig. 2: (a) Geological map of the lower Var Valley at 1/5000000 (modified from Ginsburg et al. 1980). (b) Zoom on the Carros/Gattières area (1) Gattières DSL, (2) Ripert DSL, and (3) Carros DSL. (c) Geological section A–B Fig. 2b.

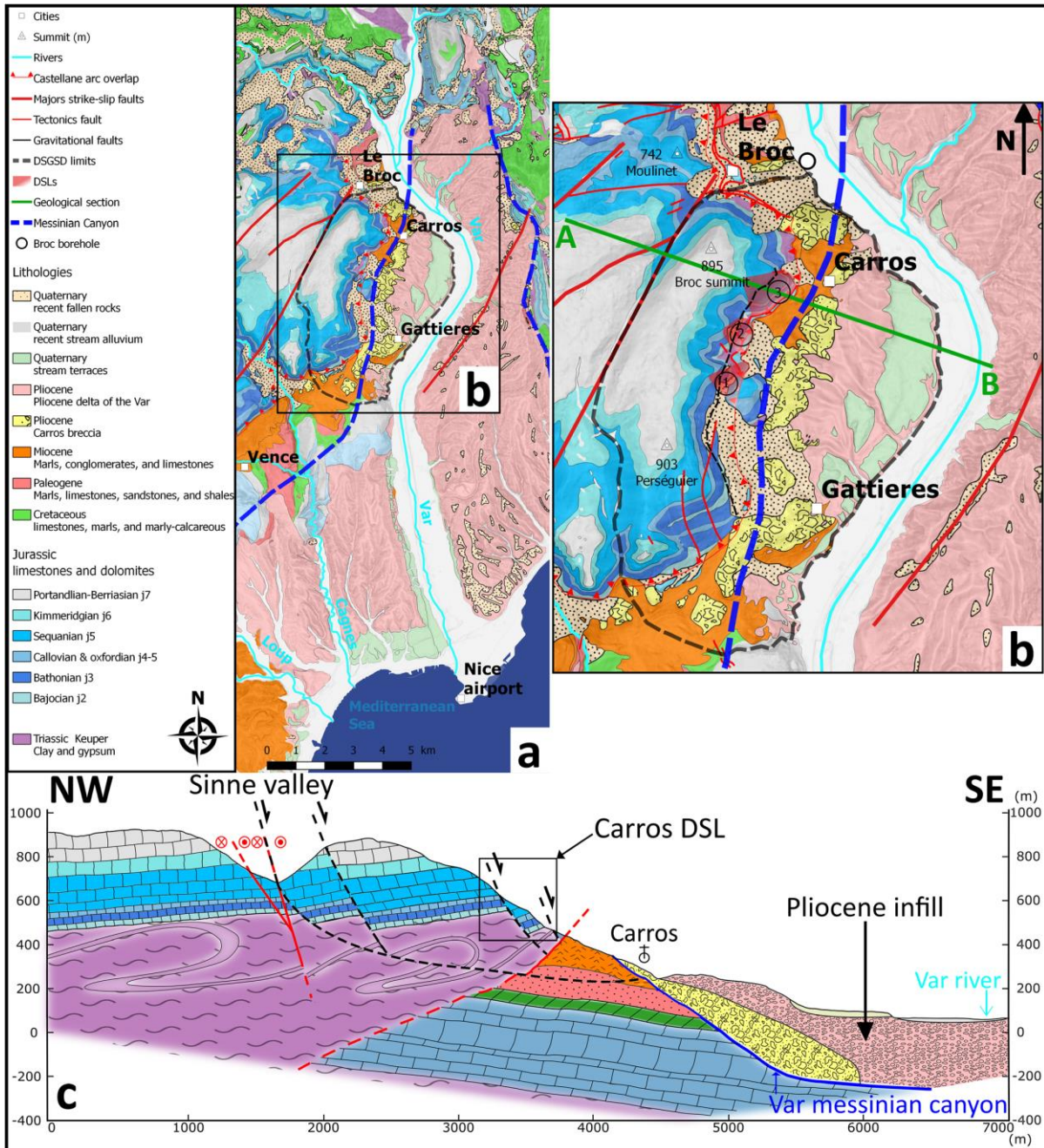


Fig. 3: 3D view of the area. (a) Cross-section M-N (Fig. 3e) Gattières DSL (modified from Zerathe (2013)). (b) Average offset between the two plateaus and the Jurassic limestone formations between the two banks of the Sinne valley without formations dip (up) with formations dip (down). (c) Crevasse toward N20–40° in the limestone formations of the Sinne Valley (up to 2 m wide and 100 m long). (d) Gattières/Carros fault markers, crevasses in NW-SE extension. (e) 3D view of the Gattières/Carros side, geological map (up) roughness map (down)

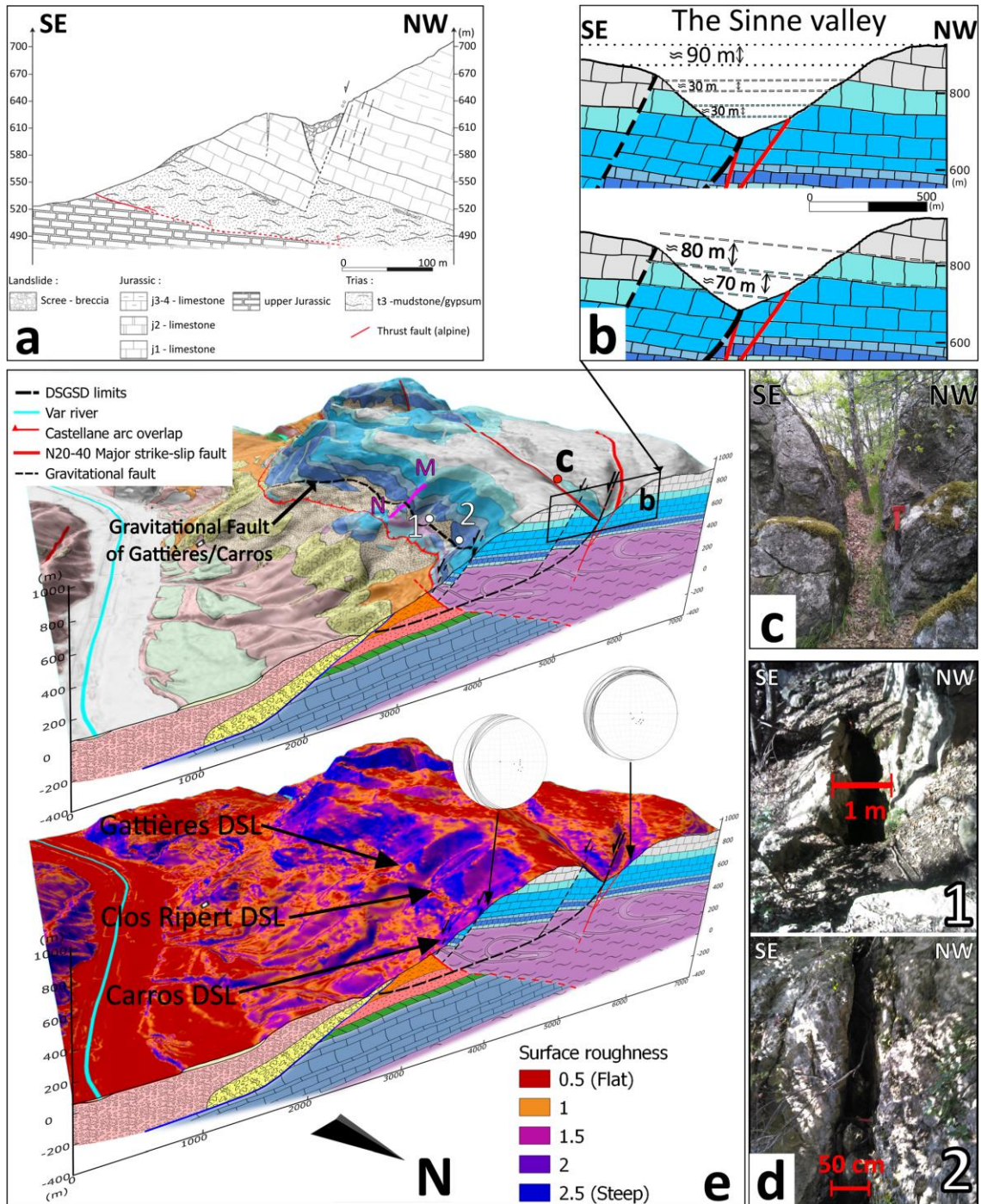


Fig. 4: (a) Foreset beds of Pliocene conglomerate at Le Broc (Carros breccia, Fig. 2). (b) Marine burrows on limestone blocks of the Carros breccia.

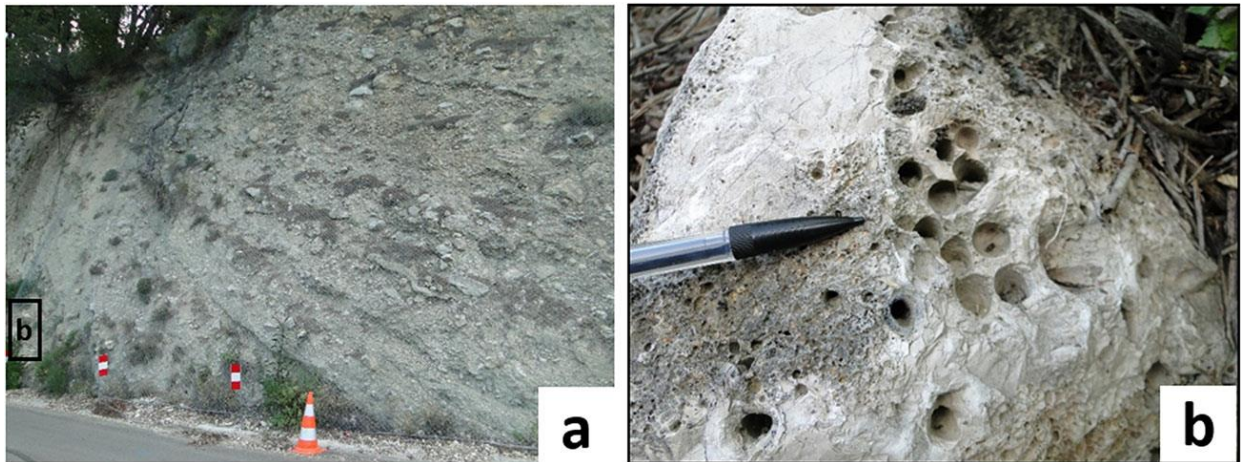


Fig. 6: Schematic interpretation of the processes that triggered the landslides in the case of the Gulf of Lion landslide (A–C) (~ 1.108 m³), the Nile valley landslide (D–F) (~ 2.108 m³), and the Alboran landslide (G–I). The river erosion increased the slope gradient along the side slopes of valleys and the apparent relief H. The flooding of the Mediterranean Sea and incised valleys by seawater increased the pore pressure along the Messinian incised valleys (Gargani et al. 2014).

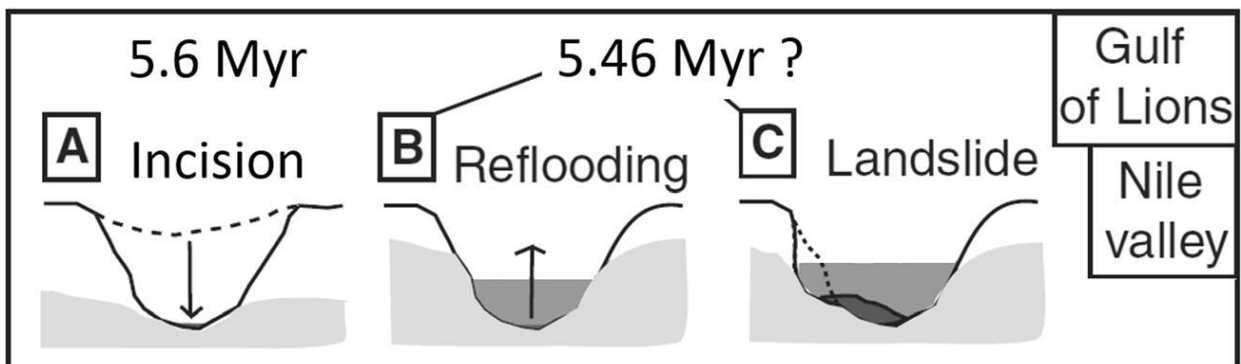


Fig. 5: Interpretations of geological events on the right bank of the Var at Carros from the Messinian (5.5 Ma). (A) Late Miocene. (B) Messinian crisis. (C) Pliocene. (D) Current. Geological layer and section position Fig. 2.

