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Reconstruction of sea-surface temperatures in the Canary Islands during Marine Isotope Stage 11

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Abstract

The study of paleoclimates enables us to improve and better constrain climate models in order to forecast future climate variations. Marine Isotope Stage 11 (MIS11), which began around 425,000 yr BP and lasted about 65,000 yr, is a warm isotope stage of paramount importance, because the astronomical configuration was similar to the one characterizing the Holocene. Therefore, this warm isotope stage is the most appropriate analog to present-day climate known to date. This study aims to provide new data on sea-surface temperatures (SSTs) inferred from the carbon and oxygen isotope compositions of skeletal carbonates of marine invertebrates preserved in two marine deposits of the Canary Islands located at Piedra Alta, Lanzarote, and Arucas, Gran Canaria. According to published isotopic fractionation equations the marine deposit from Arucas recorded SSTs of $15.9 \pm 2.2^\circ\text{C}$ on average, while the tsunamite from Piedra Alta recorded SSTs of $21.2 \pm 1.9^\circ\text{C}$ on average. Absolute dating, mollusc assemblages, and calculated marine temperatures suggest that the Arucas marine deposit corresponds to the beginning of MIS11, while the Piedra Alta tsunamite was formed during MIS11c. These results show that low latitudes also experienced sizable SST changes during interglacial stages.

Keywords: Stable isotope; Sea-surface temperature; Tsunami; Mollusc; Canary archipelago; MIS11; Climate change

INTRODUCTION

The fifth assessment report of the Intergovernmental Panel on Climate Change (IPCC, 2014) reported a global increase in average mean annual temperatures of $+0.85^\circ\text{C}$ between 1880 and 2012, each decade being warmer than the previous one. This increase in temperature, associated with the melting of the northern ice cap and the rise of global sea level, constitutes the central concern of global climate change. Consequently, there is a need to predict the evolution of global climate through the use of numerical models ranging from simple energy-balance models to complex Earth system models (Flato et al., 2013). Paleoclimatic data obtained from various proxies coming from sedimentary archives of interglacial stages are necessary to affix boundary conditions to such numerical models.

Paleontological, geochemical, and sedimentological records reveal climate variations between cold and warm stages for the last 3 Ma. For the last million year, since the mid-Pleistocene transition (Medina-Elizalde and Lea, 2005), the cold stages lasted around 100,000 yr. The glacial and interglacial stages of the last million years, which are called Marine Isotope Stages (MIS), arouse much interest (e.g., Augustin et al., 2004; Tzedakis, 2010; Robinson et al., 2017), because they could contribute to our understanding of the climate dynamics of the Holocene (MIS1) and, especially, the evolution of an interglacial stage devoid of any anthropic influence and its consequent impact upon ecosystems and coastal geomorphology. Many studies have been devoted to two interglacial stages, MIS5e (Oppo et al., 2006; Jouzel et al., 2007) and MIS11c, the latter being usually considered as the best analog to Holocene climate change (Droxler et al., 2003; Loutre and Berger, 2003). Indeed, MIS11 and MIS1 share the same configuration of orbital parameters: low eccentricity, low precession, and weak variations in insolation. MIS11c, which is an interglacial stage within the MIS11 warm isotope stage, lasted a long time (around 27,000 to 30,000 yr) compared with other interglacial stages such as MIS9e ($\approx 12,000$ yr) and MIS5e ($\approx 13,000$ yr)

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115 according to Shackleton (1967), Howard (1997), McManus
 116 et al. (2003), Augustin et al. (2004), and Prokopenko et al.
 117 (2010). Candy et al. (2014) performed a global review of
 118 the climate characteristics of the MIS11 warm isotope
 119 stage. In the framework of our study, it is worth noting that
 120 the record of $\delta^2\text{H}$ in ice cores of Dome C in Antarctica (75°
 121 $06'S$, $123^\circ 20'E$) (Augustin et al., 2004; Jouzel et al., 2007)
 122 indicates that the air temperature optimum during MIS11c
 123 was around 3°C higher than today, although this surface tem-
 124 perature change refers to the high latitudes of the Northern
 125 Hemisphere and is higher than the global mean temperature
 126 change. The study of raised beaches located in tectonically
 127 stable regions and of ice cap dynamics during MIS11
 128 revealed that the sea level was 6 to 13 m higher than today
 129 on average (de Vernal and Hillaire-Marcel, 2008; Elderfield
 130 et al., 2012; Roberts et al., 2012). Moreover, the analysis of
 131 terrestrial records corresponding to MIS11 has shown that
 132 this warm isotope stage was globally humid (Rousseau,
 133 2003; Prokopenko et al., 2010; Melles et al., 2012). The con-
 134 figuration of astronomical parameters during MIS11, espe-
 135 cially the weak precession, induced weaker insolation
 136 variability than for other warm isotope stages (Berger and
 137 Loutre, 2002). Consequently, MIS11 is a warm isotope
 138 stage that is tricky to detect and quantify in the geologic
 139 record, as reflected in the relative lack of data for this stage
 140 compared with other marine isotope stages such as MIS5.

141 Here we provide measurements of the stable carbon and
 142 oxygen isotope ratios of skeletal carbonates from marine
 143 invertebrates sampled from two marine sequences from the
 144 Canary Islands that were deposited during MIS11. The
 145 oxygen isotope ratios are then used to calculate sea-surface
 146 temperatures (SSTs) using an oxygen isotope fractionation
 147 equation for the aragonite–water system. During the two peri-
 148 ods of MIS11 studied, SSTs within tropical latitudes in the
 149 eastern Atlantic Ocean bracketed those documented during
 150 the present day, although the magnitude of seasonal variation
 151 was most likely weaker than at present.

154 GEOLOGIC SETTING AND MATERIAL 155 STUDIED

157 The Canary archipelago is composed of seven volcanic
 158 islands located in the eastern Central Atlantic Ocean between
 159 27°N and 30°N , west of Morocco (Fig. 1a). The formation of
 160 the archipelago is linked to the movement of the African plate
 161 over a magmatic hot spot (Carracedo et al., 1998). Lanzarote,
 162 Fuerteventura, and Gran Canaria, the easternmost and oldest
 163 islands, share an arid and warm climate resulting from their
 164 latitudinal position, their orography, and their proximity to
 165 the Sahara Desert. El Hierro and La Palma, the westernmost
 166 and youngest islands, are under the influence of the humid
 167 northeast trade winds. Different methods have been used to
 168 estimate present-day mean annual temperatures (MATs) and
 169 SSTs for low latitudes such as 28°N . Borges et al. (2004)
 170 analyzed time series of the southeastern North Atlantic sea
 171 surface obtained from the Advanced Very High Resolution

Radiometer (AVHRR) and estimated an average SST close
 to 20.5°C . Comparable average SSTs close to 20.3°C were
 published by DeCastro et al. (2014) for the Moroccan subre-
 gion during the period 1982–2012. Such temperature esti-
 mates have been based on the National Oceanic and
 Atmospheric Administration's (NOAA) high-resolution anal-
 yses of daily SST. Gómez-Letona et al. (2017) estimated a
 SST of 20.3°C for the coastal waters of the Canary archipel-
 ago, using daily Reynolds analyses that combine AVHRR
 and in situ data. The Climate Atlas of the Spanish Met service
 (AEMET, 2012) indicates a MAT of 20°C for the Gran Cana-
 ria and Lanzarote islands. These temperatures are relatively
 low compared with those prevailing in the western part of
 the central Atlantic Ocean, such as the Caribbean Islands,
 which have temperatures close to 27°C – 29°C (Winter et al.,
 1998). Such low temperatures around the Canary archipelago
 are explained by the presence of a southward-flowing oceanic
 current, called the Canary Current, originating from the east-
 ern boundary current of the subtropical North Atlantic gyre
 that is located along the northwest African coast. The current
 splits as it passes the Canary archipelago, with branches
 located east and west of Lanzarote. It is associated with a
 strong coastal upwelling regime along the northwest African
 coast, caused by the northeast trade winds. The coastal
 upwelling is permanent at a latitude of 28°N , with its stron-
 gest activity occurring in summer and early fall, which results
 in lower SSTs over the shelf (Navarro-Pérez and Barton,
 2001; Pardo et al., 2011).

Marine sedimentary deposits of Quaternary age crop out
 along the coasts of several islands in the Canary archipelago,
 namely Lanzarote, Fuerteventura, Gran Canaria, and Tenerife
 (Muhs et al., 2014; Meco et al., 2018). This study focuses on
 two marine deposits attributed to MIS11, the first one being
 located on Arucas (Gran Canaria) and the second one being
 located in Piedra Alta (Lanzarote).

The first marine deposit crops out at Arucas on the northern
 side of the island of Gran Canaria ($28^\circ 08'N$, $15^\circ 29'W$; Figs.
 1b and 2a). First described by Benítez (1912), this sedimen-
 tary sequence extends to about 2 km in length from east to
 west and is deposited above a tephritic basaltic flow including
 some pillow lavas, most likely related to the former activity of
 the Cardones volcano. This lava flow was dated by the K-Ar
 method at 421 ± 20 ka (Meco et al., 2002). This K-Ar age
 allowed Meco et al. (2002) to correlate this marine deposit
 with MIS11. Today, the base of the Arucas deposit crops
 out about 30 m above sea level as a result of a westward tilting
 of Gran Canaria, linked to the lithospheric flexure exerted by
 the growth of the younger island of Tenerife (Perez-Torrado
 et al., 2002; Menéndez et al., 2008; Meco et al., 2011). An
 extralimital species such as *Saccostrea cucullata* (Ostreidae)
 occurs in this deposit and suggests that SSTs were at least 4.2°
 C higher than those prevailing at present according to Monte-
 sinos et al. (2014).

The second marine deposit included in this study is
 exposed on the southwest part of the island of Lanzarote
 ($28^\circ 53'N$, $13^\circ 51'W$; Figs. 1c and 2b). First described by Dris-
 coll et al. (1965), it crops out over 2 km along the coast and at

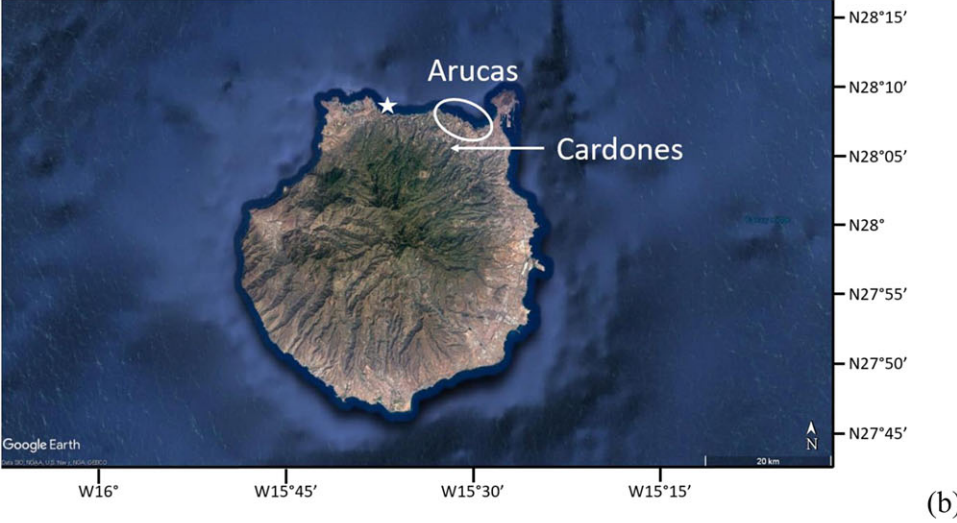
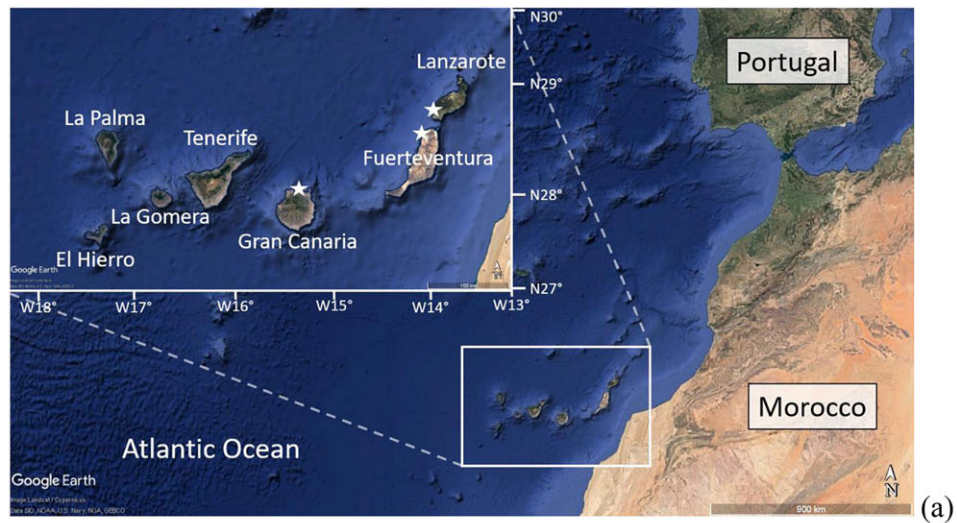


Figure 1. (color online) (a) Location of the Canary archipelago in the Atlantic Ocean; (b) Gran Canaria island, with the location of the Arucas marine deposit circled; and (c) Lanzarote island, with the location of the Piedra Alta tsunamite circled. The locations where coastal waters were sampled are represented by white stars. The datum used for latitude and longitude is WGS84. Pictures were taken from Google Earth.

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Fig. 1 - Côtout online, B/W in print

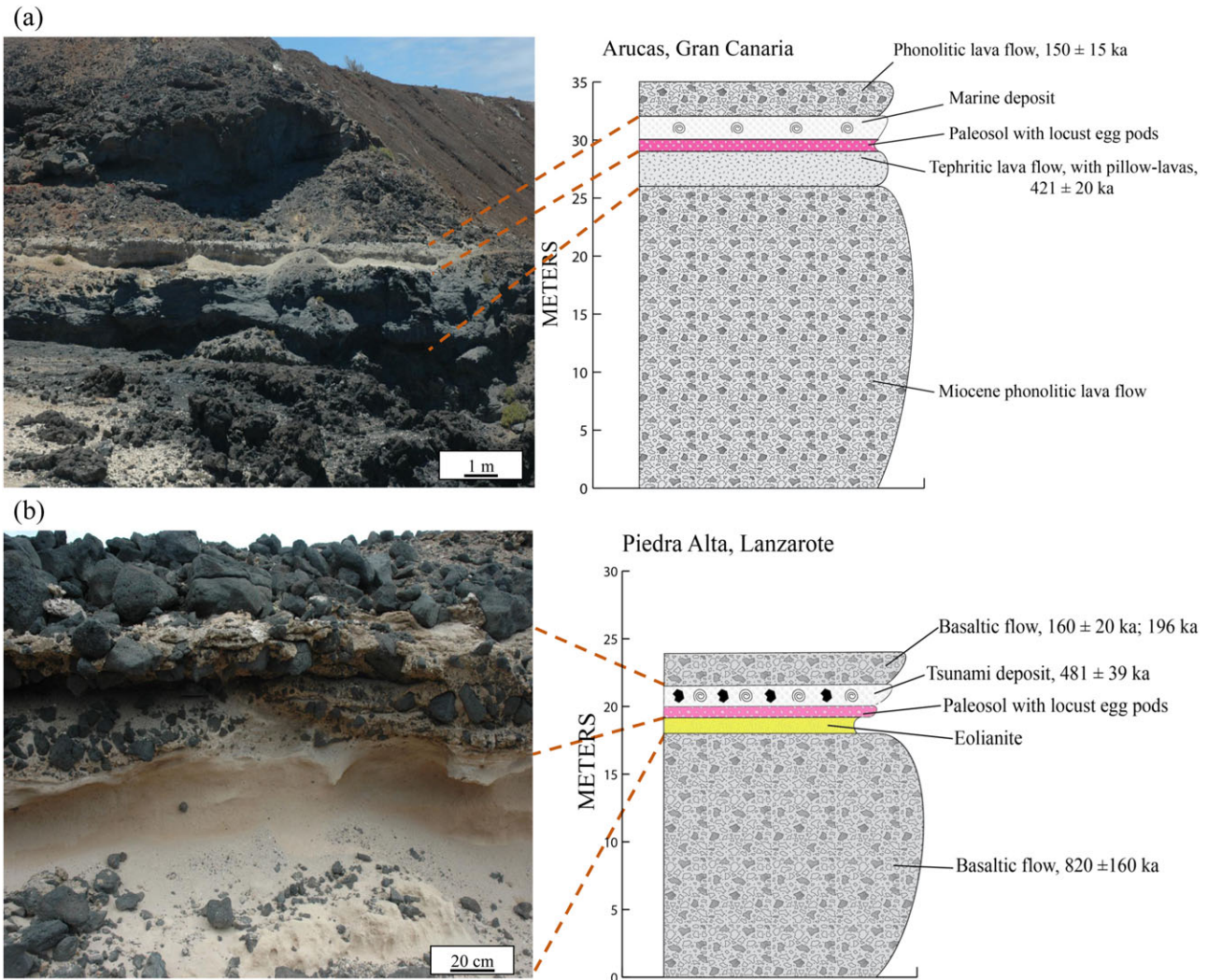


Figure 2. (color online) The Arucas marine deposit (a) and the Piedra Alta tsunamite (b) and their geologic interpretation with stratigraphic logs. Modified after Muhs et al. (2014) and Meco et al. (2011).

least 7 km inland, while the base of the deposit occurs 20 m above the present-day sea level. Zazo et al. (2002) attributed the marine deposit to the Episode IX of their raised marine sequences with a sea level 8 to 10 m higher than today. In terms of chronology, they correlated this marine deposit to MIS11 (Zazo et al., 2002). The marine deposit is interlayered between two basaltic flows that were dated by K-Ar at 820 ± 160 ka for the bottom flow (Meco and Stearns, 1981), while two ages were proposed for the top flow: 160 ± 20 ka by Zazo et al. (2002) and more recently 196 ka by Meco et al. (2011). The marine deposit is poorly sorted, consists of abundant angular or subangular basaltic boulders, and includes large metric basalt boulders as well as finer-grained volcanic clasts. Some of the basaltic clasts are covered by a layer of marine sandstones, while the boulders are cemented with reddish limestones, generally rich in clays (Meco et al., 2008). The sedimentary matrix is very fossiliferous, with abundant remains of marine invertebrate shells, none of which were observed in life position. Meco et al. (2008) noticed that the marine deposit contains a mixture of species with

different ecology, such as the genus *Patella*, which is typical of the intertidal zone; species like *Gadinia garnoti*, which are intertidal and live in the high-tide zone; and fragments of the infralittoral coral *Madracis pharencis* and the circalittoral (at depths greater than 50 m) gastropod *Bursa tenuisculpta*.

These paleontological and sedimentological features allowed Meco et al. (2008) to consider that this deposit was not formed by a slow process of continuous sedimentation but rather by an instantaneous and catastrophic event at the geologic scale, such as a tsunami deposit, like the previously identified Agaete tsunami conglomerate in Gran Canaria and the northwestern coast of Tenerife (Paris et al., 2017, 2018).

We selected 64 specimens of marine invertebrates from the Piedra Alta deposit and 77 specimens from the Arucas deposit for carbon and oxygen isotope measurements of the invertebrate marine fauna contemporaneous with MIS11 (see Table 1). In both cases, the faunal assemblage is composed of several species of molluscs, echinoderms, cnidarians, and bryozoans, whose habitats are infralittoral to intertidal and for which the taxonomy is reported in Table 1.

Table 1. Specimens collected from the Piedra Alta tsunamite and the Arucas marine deposit. The habitats data are extracted from Meco et al. (2008).

Piedra Alta tsunamite					
Taxon	Sample	Phylum	Class	Living environment	Number of specimens
<i>Patella depressa</i>	PA-01	Mollusca	Gastropoda	Intertidal	4
<i>Patella aspera</i>	PA-02	Mollusca	Gastropoda	Intertidal	5
<i>Bolma rugosa</i>	PA-03	Mollusca	Gastropoda	Infralittoral	11
<i>Turritella ligar</i>	PA-04	Mollusca	Gastropoda	Infralittoral	7
<i>Conus vayssieri</i>	PA-05	Mollusca	Gastropoda	Infralittoral	5
<i>Glycymeris glycymeris</i>	PA-06	Mollusca	Bivalvia	Infralittoral	4
<i>Hyotissa hyotis</i>	PA-08	Mollusca	Bivalvia	Infralittoral	9
<i>Saccostrea cucullata</i>	PA-09	Mollusca	Bivalvia	Infralittoral	8
<i>Callista chione</i>	PA-10	Mollusca	Bivalvia	Infralittoral	5
<i>Madracis pharensis</i>	PA-11	Cnidaria	Anthozoa	Infralittoral	2
<i>Smittina cervicornis</i>	PA-13	Bryozoa	Gymnolaemata	Infralittoral	2
<i>Cidaris</i> sp.	PA-14	Echinodermata	Echinoida	Infralittoral	2
Arucas marine deposit					
<i>Saccostrea cucullata</i>	ARU-01	Mollusca	Bivalvia	Infralittoral	3
<i>Stramonita haemastoma</i>	ARU-02	Mollusca	Gastropoda	Infralittoral	5
<i>Cerithium rupestre</i>	ARU-03	Mollusca	Gastropoda	Infralittoral	10
<i>Bolma rugosa</i>	ARU-04	Mollusca	Gastropoda	Infralittoral	1
<i>Tectarius striatus</i>	ARU-05	Mollusca	Gastropoda	Infralittoral	5
<i>Diodora gibberula</i>	ARU-06	Mollusca	Gastropoda	Intertidal	4
<i>Vermetus triquetrus</i>	ARU-07	Mollusca	Gastropoda	Infralittoral	4
<i>Charonia lampas</i>	ARU-08	Mollusca	Gastropoda	Infralittoral	3
<i>Patella aspera</i>	ARU-09	Mollusca	Gastropoda	Intertidal	4
<i>Nucella lapillus</i>	ARU-10	Mollusca	Gastropoda	Intertidal	4
<i>Cardita calyculata</i>	ARU-11	Mollusca	Bivalvia	Intertidal	4
<i>Columbella rustica</i>	ARU-12	Mollusca	Gastropoda	Infralittoral	5
<i>Bursa scrobiculata</i>	ARU-13	Mollusca	Gastropoda	Infralittoral	2
<i>Tritia pfeifferi</i>	ARU-14	Mollusca	Gastropoda	Infralittoral	3
<i>Vexillum zebrinum</i>	ARU-15	Mollusca	Gastropoda	Intertidal	3
<i>Phorcus turbinatus</i>	ARU-16	Mollusca	Gastropoda	Intertidal	4
<i>Phorcus lineatus</i>	ARU-17	Mollusca	Gastropoda	Intertidal	3
<i>Cerithium vulgatum</i>	ARU-18	Mollusca	Gastropoda	Infralittoral	2
<i>Patella vulgata</i>	ARU-19	Mollusca	Gastropoda	Intertidal	4
<i>Paracentrotus lividus</i>	ARU-20	Echinodermata	Echinoida	Infralittoral	4

ANALYTICAL TECHNIQUES

Skeletons of marine invertebrates are mainly composed of either calcite or aragonite. Preservation of the initial mineralogy in a fossil is largely dependent on the temperature-sensitive solubility of these polymorphs in an aqueous phase. Indeed, at 25°C, the solubility products (K_{sp}) are quite high, with values of $10^{-8.48}$ and $10^{-6.40}$ for calcite and aragonite, respectively, according to data from the National Institute of Standards and Technology–Joint Army Navy NASA Air Force (NIST-JANAF) thermodynamic tables (Chase, 1998). Consequently, because aragonite is less stable than calcite, it cannot be excluded that calcitic shells were aragonitic in origin before being diagenetically altered into secondary calcite. It has been known for decades (Land, 1967) that recrystallization operates at ambient temperature in the presence of water during the burial of fossils. Those aqueous fluids may be pore waters, in the case of early diagenesis, or crustal waters of meteoric or metamorphic origin during late diagenesis.

Raman spectroscopy

We used Raman spectroscopy to characterize the mineralogy of all the skeletal carbonates studied. The two polymorphs of CaCO_3 are readily identified from the low-frequency part of the spectra, and especially from the position and splitting of the symmetric bending mode, which occurs as a single peak at 712 cm^{-1} in calcite and as a doublet at $701\text{--}704\text{ cm}^{-1}$ in aragonite (e.g., Unvros et al., 1991; Gillet et al., 1993). The Raman Spectrometer X'plora, hosted by the Laboratoire de Géologie de Lyon at the University Claude Bernard Lyon 1, was operated using a 1000× objective, an optical network of 1800 lines per mm, a monochromatic laser (532 nm wavelength) filtered by 10% with two acquisitions (60 s per acquisition) performed between 100 and 1600 cm^{-1} .

Sampling strategy

Macroscopic criteria to select the skeletal carbonates for the measurement of stable isotope ratios are unbroken shells,

the absence of mechanical erosion, and dissolution patterns. The selected skeletal carbonates were cleaned by repeated washing with deionized water in an ultrasonic bath to remove any remaining sedimentary particles. For most skeletal carbonates, we performed bulk analysis (hereafter “bulk” protocol), which consists, after mechanically removing the most juvenile parts of the shells (i.e., the umbo area), of powdering the remaining shells in an agate mortar. The juvenile parts are removed because they are biased by a kinetic fractionation linked to the fast growth of the shell at that moment. For two molluscs, *Callista chione* from Piedra Alta (PA-10-04) and *Cerithium vulgatum* from Arucas (ARU-18-02), shells were incrementally sampled with a Dremel micro-drill (hereafter “incremental” protocol) along the main growth axis from the apex of the shell to the aperture to measure the ontogenetic evolution of their carbon and oxygen isotope ratios.

Carbon and oxygen isotope analysis of skeletal carbonates

The carbonate analyses were carried out at the Laboratoire d'Écologie des Hydrosystèmes Naturels et Anthropisés (LEHNA) in Lyon using an isoFLOW system connected online in continuous-flow mode to a precisION mass spectrometer operated by visION software from Elementar UK Ltd. Sample powders were loaded in LABCO Exetainer 3.7 mL round-bottomed, nonevacuated soda-glass vials. For pure carbonate samples, aliquots of 300 to 500 µg for the bulk protocol and aliquots of 15 µg for the incremental protocol were reacted with anhydrous phosphoric acid prepared according to the protocol described by McCrea (1950). Phosphoric acid was automatically introduced in contact with the sample by the use of an acid pump. The reaction took place at 70°C in a temperature-regulated sample tray. The CO₂ gas generated during the acid digestion of the carbonate sample was then transferred to the mass spectrometer via the centrION interface. A calibrated CO₂ gas was used as a monitoring gas. An acid fractionation factor $\alpha(\text{CO}_2\text{-aragonite})$ of 1.0085 was selected according to Kim et al. (2007a). The calibrated material used was Carrara marble ($\delta^{18}\text{O}_{\text{VPDB}} = -1.84\text{‰}$; $\delta^{13}\text{C}_{\text{VPDB}} = +2.03\text{‰}$) (Fourel et al., 2016) and NBS18 ($\delta^{18}\text{O}_{\text{VPDB}} = -23.2\text{‰}$; $\delta^{13}\text{C}_{\text{VPDB}} = -5.01\text{‰}$) (Friedman et al., 1982; Hut, 1987; Stichler, 1995; Coplen et al., 2006). Aliquots of Carrara marble were placed at the beginning and at the end of each analytical batch to correct for potential instrument drift through time. Delta values are expressed with respect to VPDB. External reproducibility (2σ) was lower than $\pm 0.1\text{‰}$ for $\delta^{18}\text{O}$ and $\pm 0.05\text{‰}$ for $\delta^{13}\text{C}$.

Oxygen isotope analysis of water

Coastal marine waters (1–5 m depth) were sampled along the coasts of Lanzarote (Piedra Alta, one sample), Fuerteventura (two samples), and Gran Canaria (Arucas, one sample), as depicted in Figure 1. The $^{18}\text{O}/^{16}\text{O}$ analyses from water

samples were carried out at the LEHNA in Lyon using an isoFLOW system in water equilibration mode (Epstein and Mayeda, 1953; Horita et al., 1989; McCarthy et al., 2005) connected online in continuous-flow mode to a precisION mass spectrometer operated by a visION software from Elementar UK Ltd. Aliquots (200 µL) were loaded in LABCO Exetainer 3.7 mL round-bottomed, nonevacuated soda-glass vials. The sample vial headspace was automatically flushed with helium with the isoFLOW double-needle setup. The equilibration gas was a mixture of 10% CO₂ in helium automatically introduced in contact with the sample through the isoFLOW. The equilibration reaction took place at 40°C in a temperature-regulated sample tray. The equilibrated CO₂ gas generated was then transferred to the mass spectrometer via the centrION interface. A calibrated CO₂ gas was used as a monitoring gas. Calibrated waters were used to anchor the results to the V-SLAP/V-SMOW scale. The calibrated waters used were EdL III ($\delta^{18}\text{O}_{\text{VSMOW}} = -9.34\text{‰}$), EE1 ($\delta^{18}\text{O}_{\text{VSMOW}} = +6.442\text{‰}$), and SSRW ($\delta^{18}\text{O}_{\text{VSMOW}} = -16.0\text{‰}$). During the course of the experiments, the waters used as working standards were calibrated against waters from the Water Isotope Inter-Comparison intercalibration program (Wassenaar et al., 2018). Aliquots of EdL III were placed at the beginning and at the end of each analytical batch to correct for potential instrument drift with time. Delta values are expressed with respect to VSMOW. Typical external precision for $\delta^{18}\text{O}$ analyses from water samples is 0.05‰. Samples were systematically duplicated. When necessary, in the case of oxygen isotope ratios, a correction of -0.27‰ was applied to convert $\delta^{18}\text{O}$ of water from the VSMOW to VPDB scales according to Hut (1987). This correction is necessary to compare measured $\delta^{18}\text{O}$ of CO₂ produced by the reaction of the carbonate with H₃PO₄ and the CO₂ equilibrated with H₂O.

Calculation of SST for the marine deposits

The SSTs were calculated using an oxygen isotope fractionation equation for the aragonite–water system. Two main equations were available for this calculation. The first one (Grossman and Ku, 1986) was determined by analyzing modern molluscs and ambient water in the temperature range $\approx 3^\circ\text{C}$ – 15°C with $\delta^{18}\text{O}_{\text{aragonite}}$ ($\delta^{18}\text{O}_{\text{ar}}$) and $\delta^{18}\text{O}_{\text{seawater}}$ ($\delta^{18}\text{O}_{\text{sw}}$) relative to VPDB:

$$T = 20.6 - 4.34 (\delta^{18}\text{O}_{\text{ar}} - \delta^{18}\text{O}_{\text{sw}}) \quad (\text{Eq. 1})$$

The second oxygen isotope fractionation equation (Kim et al., 2007a, 2007b) was obtained from laboratory experiments with $\delta^{18}\text{O}_{\text{ar}}$ and $\delta^{18}\text{O}_{\text{sw}}$ relative to VSMOW:

$$1000 \ln_{\text{ar-sw}} = 17.88 \left(\frac{10^3}{T} \right) - 31.14 \quad (\text{Eq. 2})$$

T being the temperature in kelvins and α the oxygen fractionation coefficient between aragonite and water:

$$\alpha_{ar-sw} = \frac{\left(\frac{^{18}\text{O}}{^{16}\text{O}}\right)_{ar}}{\left(\frac{^{18}\text{O}}{^{16}\text{O}}\right)_{sw}} \quad (\text{Eq. 3})$$

The errors associated with the linear regression of empirical data ($y = mx + b$), which are not available in the original publications, have been computed as the mean standard error for the y value and are $\pm 1.38^\circ\text{C}$ for the Grossman and Ku (1986) equation and $\pm 1.15^\circ\text{C}$ for the Kim et al. (2007a, 2007b) equation. SSTs calculated from the $\delta^{18}\text{O}$ of modern mollusc shells and $\delta^{18}\text{O}$ of ambient seawater in Crete, Greece, best match those available from oceanic databases available for the Mediterranean Sea when using the Kim et al. (2007a, 2007b) equation (Lécuyer et al., 2018). Grossman and Ku (1986) also acknowledged that their equation may overestimate temperature by about 2°C . Therefore, we chose to use the Kim et al. (2007a, 2007b) equation to infer the SSTs that prevailed in the Canary Islands during MIS11.

Correction of SST related to sea-level changes

To correctly estimate the $\delta^{18}\text{O}$ values of coastal seawaters in the Canary archipelago during MIS11, sea-level changes need to be taken into account. There is a long-standing debate about the sea-level range during MIS11. On the one hand, based on the study of raised shorelines in regions such as Bermuda (Hearty et al., 1999; Olson and Hearty, 2009) or the United Kingdom (Bowen, 2003), the high sea-stand levels are estimated to range from 10 to 15 m above modern sea levels and even up to 20 m above modern sea levels according to Roberts et al. (2012). On the other hand, authors of some modeling studies, those devoted to the Red Sea record in particular, argue that sea level during MIS11 was comparable to current sea level (Rohling et al., 2010). For more information about this debate, see Candy et al. (2014).

Muhs et al. (2014) calculated uplift rates for the Canary Islands by using the peak of the last interglacial period (LIG) and the dating of deposits from the LIG and proposed sea-level ranges associated with the different episodes of raised beaches in the Canary archipelago. Accordingly, based on a study of the Arucas deposit, Muhs et al. (2014) proposed a paleo-sea level of +9 m to +24 m relative to the present for Gran Canaria for MIS11. Muhs et al. (2014) also proposed a paleo-sea level of +10 m to +13 m relative to the present for Lanzarote, based on the Piedra Alta deposit, although the authors acknowledge being less confident concerning the latter estimate. Those two sea-level estimates are broadly consistent with each other and are similar to the range of sea-level estimates made for the island of Curaçao (Muhs et al., 2012). We chose to consider the average value of the MIS11 sea-level range for each marine deposit, which is included in the average paleo-sea level range attributed to MIS11 (de Vernal and Hillaire-Marcel, 2008; Elderfield et al., 2012; Roberts et al., 2012) and to integrate the range of uncertainty. Our final estimates correspond to sea-

level ranges of $+16.5 \pm 7.5$ m for the Arucas marine deposit and $+11.5 \pm 1.5$ m for the Piedra Alta tsunamite. If we take an average 3700 m depth for the ocean (Charette and Smith, 2010) coupled to an average isotopic composition of the polar caps of -40‰ , a mass balance calculation indicates that a global sea-level rise of $+16.5 \pm 7.5$ m would correspond to a decrease of seawater $\delta^{18}\text{O}$ by $0.18 \pm 0.08\text{‰}$, while a rise of $+11.5 \pm 1.5$ m would correspond to a decrease of $0.12 \pm 0.02\text{‰}$.

RESULTS

Mineralogical composition of the shells studied

All skeletal carbonates were analyzed by Raman spectroscopy to identify their mineralogy (Table 2). The inorganic fraction (>95%) of shells is entirely composed of either calcite or aragonite, with some exceptions, such as *Patella depressa* (PA-01), *Stramonita haemastoma* (ARU-02), *Nucella lapillus* (ARU-10), and *Patella vulgata* (ARU-19), which contain minute amounts (<5%) of calcite. In the Piedra Alta tsunamite, the shell of *Patella aspera* (PA-02) is made of calcite, while in the Arucas marine deposit, the shell of the same species (ARU-09) is made of aragonite. As aragonite is the original mineralogy of the inner layer of the shell of *Patella aspera* (Parker et al., 2017), calcite is thus interpreted as an alteration product in the case of sample PA-02. Consequently, all the species fully composed of aragonite were selected for stable carbon and oxygen isotope analysis. Moreover, the species *Saccostrea cucullata* (PA-09/ARU-01) was also selected for isotopic analysis, as it is present in both marine deposits in sufficient number to allow the isotopic values to be compared. However, as *Saccostrea cucullata* is composed of calcite, the oxygen isotope ratios were not used to calculate the SST attributed to the marine deposits.

Carbon and oxygen isotope compositions of skeletal carbonates

The isotopic data are available in the Supplementary Material. For the Piedra Alta tsunamite, specimens have $\delta^{18}\text{O}$ values ranging from -1.40‰ to 1.60‰ , with a mean value of 0.46‰ , while $\delta^{13}\text{C}$ values range from -2.71‰ to 2.95‰ , with a mean value of 1.31‰ . For the Arucas marine deposit, the $\delta^{18}\text{O}$ values of skeletal carbonates are between 0.77‰ and 2.83‰ , with a mean value of 1.55‰ , while $\delta^{13}\text{C}$ values are between 0.46‰ and 4.73‰ , with a mean value of 2.42‰ .

Two species, *Callista chione* from Piedra Alta (PA-10-04) and *Cerithium vulgatum* from Arucas (ARU-18-02), were sampled according to the incremental protocol. *Callista chione* has $\delta^{18}\text{O}$ values ranging from 0.13‰ to 0.79‰ , with a mean value of 0.35‰ , while $\delta^{13}\text{C}$ values are between 0.02‰ and 1.90‰ , with a mean value of 0.92‰ . *Cerithium vulgatum* has $\delta^{18}\text{O}$ values ranging from 0.11‰ to 1.52‰ , with a mean value of 1.07‰ , while $\delta^{13}\text{C}$ values are between 0.01‰ and 3.16‰ , with a mean value of 1.98‰ . In both cases, the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values show a pseudo-cyclic

Table 2. Mineralogy of the biogenic carbonates interpreted from Raman spectra. The terms marked with an asterisk (*) specify that the mineralogy of the shell is mainly aragonitic with less than 5% being calcitic.

Sample	Mineralogy of the shell
PA-01	Aragonite*
PA-02	Calcite
PA-03	Aragonite
PA-04	Aragonite
PA-05	Aragonite
PA-06	Aragonite
PA-08	Calcite
PA-09	Calcite
PA-10	Aragonite
PA-11	Aragonite
PA-13	Calcite
PA-14	Calcite
ARU-01	Calcite
ARU-02	Aragonite*
ARU-03	Aragonite
ARU-04	Aragonite
ARU-05	Aragonite
ARU-06	Aragonite
ARU-07	Aragonite
ARU-08	Aragonite
ARU-09	Aragonite
ARU-10	Aragonite*
ARU-11	Aragonite
ARU-12	Aragonite
ARU-13	Aragonite
ARU-14	Aragonite
ARU-15	Aragonite
ARU-16	Aragonite
ARU-17	Aragonite
ARU-18	Aragonite
ARU-19	Aragonite*
ARU-20	Calcite

variation recorded during the ontogenetic evolution of these two specimens studied. However, these isotopic ratios do not show any significant correlation in the case of *Callistichione* ($R^2 = 9 \times 10^{-4}$), while they are significant in the case of *Cerithium vulgatum* ($R^2 = 0.52$).

A few corrections have been applied to the isotopic values in some outliers. Like many corals, *Madracis pharensis* (PA-11) is ^{18}O - and ^{13}C -depleted relative to coexisting molluscs. This isotopic pattern is related to the photosymbiotic relation between the coral and its zooxanthella (Land et al., 1975) and a kinetic isotope effect linked to a discrimination against ^{18}O and ^{13}C during the process of biomineralization, more precisely during the hydration and hydroxylation of CO_2 (McConnaughey, 1989, 2003; Watkins et al., 2014). Therefore, we excluded the isotopic compositions of this coral sample from our data set. All the remaining specimens studied are molluscs, which are known to precipitate their shell carbonate close to carbon and oxygen isotope equilibrium with seawater (Grossman and Ku, 1986; Lécuyer

et al., 2004, 2012). We also subtracted 1.01‰ from each $\delta^{18}\text{O}$ value from *Patella vulgata* (ARU-19) to account for the predictable vital effect of the species (Wang et al., 2012).

Once we made these modifications to our database, $\delta^{18}\text{O}$ values of mollusc shells range from -0.26‰ to 1.60‰ , with a mean value of 0.55‰ , while the $\delta^{13}\text{C}$ values are between -0.28‰ and 2.95‰ , with a mean value of 1.47‰ for the Piedra Alta tsunamite. Concerning the Arucas marine deposit, $\delta^{18}\text{O}$ values of mollusc shells range from 0.69‰ to 2.83‰ , with a mean value of 1.49‰ , while the $\delta^{13}\text{C}$ values are between 0.46‰ and 4.73‰ , with a mean value of 2.42‰ . We also observed that the various mollusc species collected have clustered $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values, which precludes any kind of identification of their distinct ecology within the littoral environment (Table 1).

Oxygen isotope ratios of seawater samples

The coastal marine waters sampled from Gran Canaria, Fuerteventura, and Lanzarote have $\delta^{18}\text{O}$ values of $0.97 \pm 0.05\text{‰}$, $1.08 \pm 0.07\text{‰}$ (average between 1.09‰ and 1.06‰), and $1.09 \pm 0.04\text{‰}$, respectively. The values are quite homogeneous among those islands, with a mean value of $1.05 \pm 0.07\text{‰}$. The oxygen isotope ratios of coastal waters from Gran Canaria and Lanzarote have been corrected according to a higher sea level during MIS11 (see subsection “Correction of SST related to sea-level changes”) and are $0.79 \pm 0.01\text{‰}$ and $0.97 \pm 0.06\text{‰}$, respectively.

Calculation of the global uncertainty range

There is a need to calculate a global uncertainty range integrating the several uncertainties produced by our calculations: the external reproducibility of the $\delta^{18}\text{O}$ value of the skeletal carbonates ($\pm 0.1\text{‰}$), errors associated with the linear regression of the oxygen isotope fractionation equations ($\pm 1.15^\circ\text{C}$ for the Kim et al. [2007a, 2007b] equation), and uncertainty linked to the correction for sea-level change during MIS11 ($\pm 0.06\text{‰}$ for the Piedra Alta marine deposit and $\pm 0.01\text{‰}$ for the Arucas marine deposit). Consequently, for the SST at the Arucas marine deposit, the total uncertainty range is $\pm 1.4^\circ\text{C}$ for the Kim et al. (2007a, 2007b) equation. For the SST recorded at the Piedra Alta tsunamite, the total uncertainty range is $\pm 1.5^\circ\text{C}$ for the Kim et al. (2007a, 2007b) equation.

SST calculated for the marine deposits

The calculation of SST has been made using the oxygen isotope fractionation equation (Eq. 2) for the aragonite–water system: the data from shells composed of calcite (PA-09 and ARU-01) have been excluded from this calculation. For the 36 specimens analyzed from the Piedra Alta tsunamite (Fig. 3a), we calculated an average SST of $21.2 \pm 1.9^\circ\text{C}$ according to Eq. 2. Similarly, for the 69 specimens analyzed from the Arucas marine deposit (Fig. 3b), we calculated an average SST of $15.9 \pm 2.2^\circ\text{C}$ according to Eq. 2.

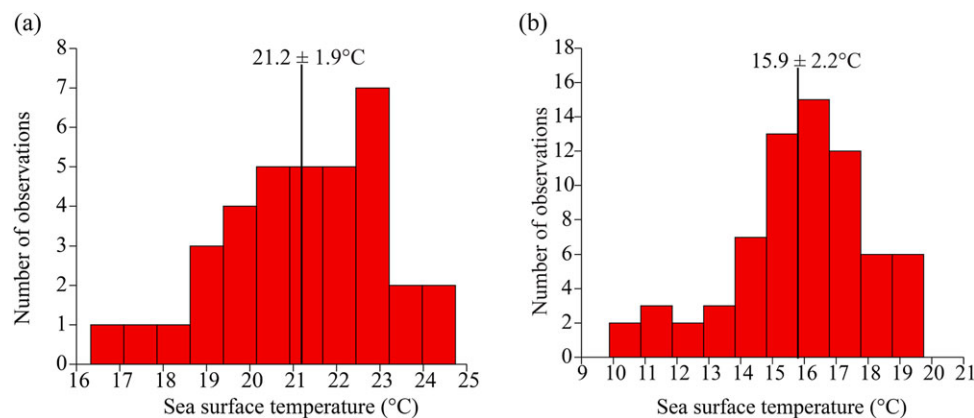


Figure 3. (color online) Distribution of MIS11 SSTs calculated by using the $\delta^{18}\text{O}$ of aragonitic mollusc shells from (a) the Piedra Alta tsunamite with a $\delta^{18}\text{O}_{\text{sw}}$ value of 0.97‰ and (b) the Arucas marine deposit with a $\delta^{18}\text{O}_{\text{sw}}$ value of 0.79‰. Those isotopic temperatures have been calculated according to the aragonite–water oxygen isotope fractionation equation given by Kim et al. (2007a, 2007b).

The ontogenetic profiles, calculated using *Callista chione* from Piedra Alta (PA-10-04) and *Cerithium vulgatum* from Arucas (ARU-18-02), reveal SST sinusoidal-like patterns that can be assigned to seasonal temperature variations (Fig. 4).

DISCUSSION

Analysis of the SST calculated for the marine deposits

We consider that the SST values calculated from a sample of the population of each marine deposit can be considered as representative of the SST of coastal waters during MIS11. In fact, there are no strong differences, in both marine deposits, between the carbon and oxygen isotope ratios recorded by intertidal and infralittoral species (see Table 1 and Supplementary Material). In addition to comparable standard deviations between the two populations of fossil molluscs from MIS11, the SST values follow a normal distribution according to the use of normal probability plots (executed with the software Past3) in which the correlation coefficients, R^2 , are always higher than 0.98 (Filliben, 1975). This means that molluscan assemblages sampled from both the Piedra Alta and Arucas marine deposits correspond to coeval populations, representative of their habitats at a given moment. Such an observation is of paramount importance in the case of the Piedra Alta tsunamite, considering that this catastrophic marine deposit may have included several diachronous populations mixed and intercalated between the lava flows. In the case of the Arucas marine deposit, there may exist time averaging (hundreds or thousands of years) associated with the formation of the marine deposit, but the normal distribution of SSTs indicates that there are no major changes in the SST during this period, and that the values calculated are representative of the SST during MIS11 in the coastal waters of Gran Canaria. The SSTs calculated for the MIS11 are, in the case of Piedra Alta, comparable to the present SST in the

Canary archipelago. In the case of the Arucas marine deposit, the calculated SSTs are lower than present-day recorded SST.

These SST values are in good agreement with those obtained in the eastern Central Atlantic Ocean, such as ODP 958, which ranges from 15°C to 20°C using various methods such as planktonic foraminiferal abundances, alkenone ratios, and oxygen isotope ratios of planktonic and benthic foraminifera (Helmke et al., 2008; Kandiano et al., 2012). They are also in agreement with the general idea that MIS11c was characterized by moderate warmth (Jouzel et al., 2007; Lawrence et al., 2009) compared with other interglacial episodes such as MIS5e (see Candy et al. [2014] for a comparison of SST records between MIS11 and MIS5).

The exoskeleton growth rates for most mollusc species follow a Von Bertalanffy growth mode (Von Bertalanffy, 1957) of the first order (Leveque, 1971; Cloern and Nichols, 1978). Thus, the signal recorded during shell growth is stretched during the juvenile age when the growth is fast, and it contracts when the growth slows down as the organism approaches maturity. Moreover, growth is also dependent on the temperature of ambient seawater. Shell growth may stop under and above some species-dependent threshold seawater temperatures. Consequently, the input signal resulting from the seasonal temperature variations, which is basically sinusoidal, is then distorted in the output oxygen isotope record, with maxima and minima being possibly truncated as the result of a lack of mineralization (Cornu et al., 1993). Such potential bias in the SST record at the seasonal scale implies that the calculated annual amplitude is most likely underestimated in most cases. We note that the seasonal amplitude in our study was between 3°C and 4°C and is weaker than the current seasonal seawater temperature variation, which is estimated to be around 6°C in the Canary Islands (Borges et al., 2004).

Saccostrea cucullata (PA-09/ARU-1; Table 1), which is an Ostreidae considered by Montesinos et al. (2014) as an extralimital species, occurs in both marine deposits. Those authors used *Saccostrea cucullata* to propose that the SSTs in the Canary archipelago were at least 4.2°C higher

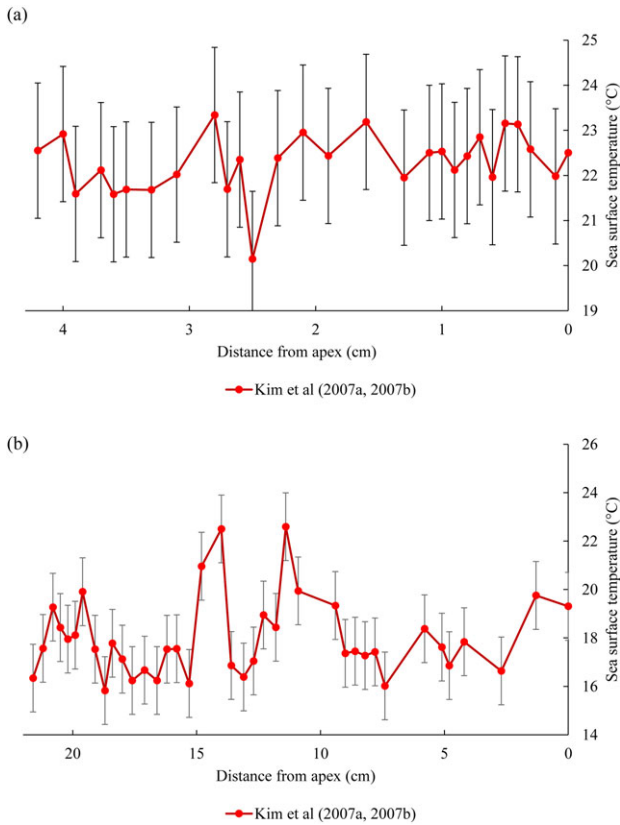


Figure 4. (color online) SST inferred from the incremental $\delta^{18}\text{O}$ of aragonitic mollusc shells from (a) *Callista chione* (PA-10-04) and (b) *Cerithium vulgatum* (ARU-18-02), using the aragonite–water oxygen isotope fractionation equation determined by Kim et al. (2007a, 2007b).

during MIS11 than today. When Eq. 2 is applied to fossil mollusc shells from both marine deposits, we notice that our SST estimates are lower than those proposed by Montesinos et al. (2014). This means that *Saccostrea cucullata* may be more tolerant to SST variations than previously thought and that this species might be able to live in colder environments.

Dating of the marine deposits

SSTs recorded in Arucas are around 5°C lower than those recorded in Piedra Alta and approximately 4°C lower than those measured during the last decades. The Arucas marine deposit is well dated, due to an interlayered basaltic flow dated by K-Ar at ca. 420 ka (Meco et al., 2002, 2011). These dates need to be viewed in the light of the LR04 marine benthic $\delta^{18}\text{O}$ stack record (Lisiecki and Raymo, 2005). The astronomical parameter configuration during MIS11 (low eccentricity, low precession, and low insolation variations) places both the insolation peak and the temperature peak during the middle of the warm isotope stage, that is, around 400 ka BP (Candy et al., 2014). Such relatively low temperatures recorded in Arucas would be in accordance with the deposition of marine sediments during the end of the Termination V between MIS12 and MIS11.

Muhs et al. (2014) dated the marine deposit of Piedra Alta by using the uranium series applied to fragments of the scleratinian coral *Madracis pharensis*, and they obtained an age of 481 ± 39 ka, although a possible bias may be linked to an initial lack of $^{234}\text{U}/^{238}\text{U}$. According to Muhs et al. (2014), the Piedra Alta marine deposit could be correlated to either MIS11 or MIS13. However, the co-occurrence in the Piedra Alta tsunamite of the gastropods *Patella depressa* and *Patella aspera*, which have not been observed in the Canary archipelago before MIS11 and in the Mediterranean basin before the mid-Pleistocene (Meco et al., 2008), suggests that this deposit was more or less contemporaneous with MIS11. The occurrence of a similar paleosol (Fig. 2a and b) that contains innumerable fossilized egg pods from orthopters and terrestrial gastropods in both the Piedra Alta and Arucas marine deposits constitutes an argument to correlate them (Meco et al., 2011). These egg pods are usually linked to humid periods that correspond to the transitions between glacial and interglacial episodes in the Canary Islands (Meco et al., 2011). This paleosol is directly overlain by the tsunamite at Piedra Alta and by the marine deposit at Arucas (Meco et al., 2011). The paleo-sea level of +10 m to +13 m associated with the Piedra Alta tsunamite (Muhs et al., 2014) is also inconsistent with its correlation to MIS13, as most sea-level modeling outputs indicate the maximum sea level during MIS13 to have been about 16 m (Miller et al., 2005, 2011) to 28–30 m (Bintanja and Van de Wal, 2008) below present sea level. Moreover, the presence of species such as *Saccostrea cucullata* and the SST values inferred from mollusc shells from Piedra Alta exclude its correlation with MIS13, which is considered a cold interglaciation, as witnessed by global air temperatures deduced from ice core isotopic records (Augustin et al., 2004; Jouzel et al., 2007) and SSTs calculated using the alkenone record in the Atlantic Ocean (Ho et al., 2012), which are consistently lower than SST values for MIS11 inferred from the study of marine sediments (Past Interglacials Working Group of PAGES, 2016).

Considering the absolute dating and the paleontological, sedimentological, and paleoclimatic data of the two marine deposits, we propose that the Piedra Alta deposit is younger than the Arucas deposit. This hypothesis is indeed in agreement with the absolute dating of the Arucas marine deposit at 421 ± 20 ka (Meco et al., 2002) and the SST values recorded in Piedra Alta mollusc shells, which are around 5°C higher than those calculated in Arucas. Consequently, the Piedra Alta marine deposit could correspond to the MIS11c interglacial stage, while the Arucas marine deposit could correspond to Termination V, between MIS12 and MIS11 (Fig. 5).

$\delta^{13}\text{C}$ of marine species

Carbon in the atmosphere is mostly present as carbon dioxide (CO_2), whereas more than 90% of the carbon in the deep ocean is present as bicarbonate (HCO_3^-), referred to as “dissolved inorganic carbon” (DIC). The reaction for

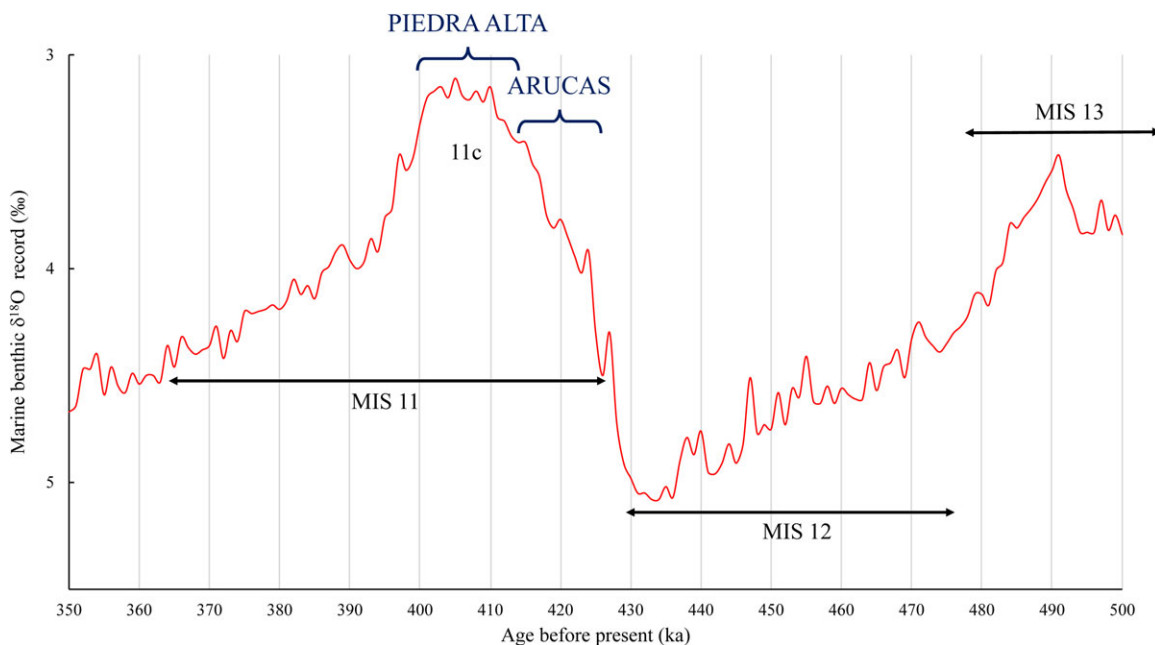
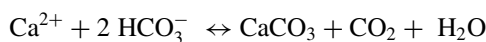


Figure 5. (color online) The LR04 marine benthic $\delta^{18}\text{O}$ record normalized and plotted on the LR04 time scale, with the proposed chronology of the Piedra Alta tsunamite and the Arucas marine deposit with respect to MIS11, MIS12, and MIS13. After Lisiecki and Raymo (2005).

precipitation of carbonates in the seawater, known as the “carbonate pump,” is:



This chemical reaction involves little carbon isotopic fractionation compared with that taking place between carbon dioxide and bicarbonate ions, which constitute the main reservoir of DIC in the oceans. Thus, the $\delta^{13}\text{C}$ of inorganically and biologically precipitated carbonate in the oceans is very close to that of DIC in the oceans (Maslin and Swann, 2006). The species *Saccostrea cucullata* (PA-09/ARU-01) and *Bolma rugosa* (PA-03/ARU-04) are common in the two marine deposits studied (Table 1), even though the number of specimens analyzed is more important for the Piedra Alta tsunamite (11 specimens of *Bolma rugosa* and 8 specimens of *Saccostrea cucullata*) than for the Arucas marine deposit (1 specimen of *Bolma rugosa* and 3 specimens of *Saccostrea cucullata*). As only one specimen of *Bolma rugosa* from the Arucas marine deposit has been analyzed, which prevents a comparison between the two marine deposits, we compared the $\delta^{13}\text{C}$ values of *Saccostrea cucullata* for Piedra Alta and Arucas. This comparison allows us to dismiss the species-dependent carbon isotope fractionation between calcium carbonate and DIC. Consequently, a $\delta^{13}\text{C}$ difference between the shells of *Saccostrea cucullata* from the two marine deposits studied should represent a similar $\delta^{13}\text{C}$ difference in seawater DIC of these marine deposits that could be attributed to a difference in interrelated variables: seawater temperature, the atmospheric pCO_2 , and the productivity of surface waters (Maslin and Swann, 2006). We observed that specimens of Arucas mollusc shells have $\delta^{13}\text{C}$ values about

0.6‰ higher on average than those present in Piedra Alta, which we consider to reflect a similar difference in the $\delta^{13}\text{C}$ values of DIC of the Canary Islands coastal waters. Such an isotopic difference could be the consequence of SST being 6°C lower according to the gradient of carbon isotope fractionation between atmospheric CO_2 and DIC of 0.1‰ per 1°C, although this equilibrium is not achieved anywhere, because surface waters are replaced too quickly (Saltzman and Thomas, 2012). A higher atmospheric pCO_2 prevailing during the Piedra Alta tsunamite deposition event could have lowered the $\delta^{13}\text{C}$ values of DIC subsequent to equilibration with seawater. Another scenario could involve stronger productivity during the Arucas marine deposit than during the MIS11c interglaciation, when the upwelling may have been diminished, resulting in colder waters that allowed more CO_2 to be dissolved.

The two mollusc shells were selected for obtaining isotopic ontogenetic profiles provide complementary information at the seasonal scale between SST and the $\delta^{13}\text{C}$ values of DIC. In the case of the mollusc *Cerithium vulgatum* (ARU-18-02), the $\delta^{13}\text{C}$ values of the shell and the SST estimates inferred from the $\delta^{18}\text{O}$ values of the shell are linearly correlated, with a slope of -0.31 and an R^2 of 0.52. As we already proposed, this correlation may also indicate that there is a temperature-dependent isotopic fractionation between air CO_2 and dissolved bicarbonate in the case of the Arucas marine deposit (Bojar et al., 2004). At the seasonal scale, it means that during summer, when SSTs are the highest (Borges et al., 2004), $\delta^{13}\text{C}$ values of DIC are at their lowest. The variation in $\delta^{13}\text{C}$ values of DIC can be assigned, at the seasonal scale, to a change in the activity of the upwelling of the Canary Current, which brings to the surface ancient and ^{13}C -depleted waters relative to surface

waters. Therefore, the Canary Current may have been stronger during the warm season in the Canary Islands during MIS11, which is compatible with the observations made by Mittelstaedt (1991), Navarro-Pérez and Barton (2001), and Pardo et al. (2011) concerning the activity of the Canary Current.

CONCLUSION

MIS11 is a warm isotope stage considered to be the best reference for understanding the evolution of the MIS1 climate. We have studied two sedimentary marine deposits located in two eastern islands of the Canary archipelago, Lanzarote (Piedra Alta tsunamite) and Gran Canaria (Arucas sedimentary deposit). These marine deposits contain many biogenic skeletal remains of marine origin, which we have studied using Raman spectroscopy to select only the aragonitic skeletal remains for the stable isotope analysis.

Using the bulk oxygen isotope analysis of the aragonitic mollusc shells, SST estimates for MIS11 are $21.2 \pm 1.9^\circ\text{C}$ for the Piedra Alta tsunamite (Fig. 3a) and $15.9 \pm 2.2^\circ\text{C}$ for the Arucas marine deposit (Fig. 3b) according to available oxygen isotope fractionation equations. The normal distribution of SSTs associated with similar standard deviations for both sites means that coeval populations of molluscs were sampled in each sedimentary marine deposit. These SST estimates, combined with absolute chronological dating and sedimentary and paleontological observations, argue in favor of the deposition of the Arucas marine deposit at the end of Termination V and the formation of the Piedra Alta tsunamite during MIS11c, which is the warmest period of the interglacial stage MIS11. The SSTs we calculated are consistent with other data suggesting warm temperatures in the central Southern Ocean during MIS11, although these temperatures are not exceptional compared with SSTs recorded for MIS5e.

The carbon isotope measurements for *Saccostrea cucullata* and *Bolma rugosa*, which are present in both marine deposits, show that species from the Arucas deposit are ^{13}C -enriched relative to those from Piedra Alta. This isotopic difference may be linked to temperature, atmospheric pCO_2 , or marine productivity variations that took place between the formation of these two deposits.

The incremental oxygen isotope measurements for *Callista chione* and *Cerithium vulgatum* (Fig. 4) revealed sinusoidal-like signals that we interpreted as seasonal variations of SST between 3°C and 4°C , an amplitude slightly weaker than the present-day SST documented in the Canary Islands. The carbon isotope measurements performed on *Cerithium vulgatum* (ARU-18-02) are linearly and negatively correlated to the SST variations, which highlights a stronger activity of the Canary upwelling current during the warm season.

The paleoclimatological data characterizing the early and mid-Pleistocene isotope stages are commonly derived from ice cores and marine and lacustrine sediments, with high latitudes sampled in most cases. However, this study shows that, during MIS11c, the global warming also significantly impacted low-latitude marine regions.

SUPPLEMENTARY MATERIAL

The supplementary material for this article can be found at <https://doi.org/10.1017/qua.2019.65>.

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