



**HAL**  
open science

# Thermophysiologicals of Jurassic marine crocodylomorphs inferred from the oxygen isotope composition of their tooth apatite

Nicolas Séon, Romain Amiot, Jérémy Martin, Mark Young, Heather Middleton, François Fourel, Laurent Picot, Xavier Valentin, Christophe Lécuyer

## ► To cite this version:

Nicolas Séon, Romain Amiot, Jérémy Martin, Mark Young, Heather Middleton, et al.. Thermophysiologicals of Jurassic marine crocodylomorphs inferred from the oxygen isotope composition of their tooth apatite. *Philosophical Transactions of the Royal Society B: Biological Sciences*, 2020, 375 (1793), pp.20190139. 10.1098/rstb.2019.0139 . hal-02991781

**HAL Id: hal-02991781**

**<https://hal.science/hal-02991781v1>**

Submitted on 27 Nov 2020

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

# Thermophysiologicals of Jurassic marine crocodylomorphs inferred from the oxygen isotope composition of their tooth apatite

Nicolas Séon<sup>1</sup>, Romain Amiot<sup>1,\*</sup>, Jeremy E. Martin<sup>1</sup>, Mark T. Young<sup>2</sup>, Heather Middleton<sup>3</sup>,  
François Fourel<sup>4</sup>, Laurent Picot<sup>5</sup>, Xavier Valentin<sup>6,7</sup>, Christophe Lécuyer<sup>1</sup>

<sup>1</sup> UMR 5276, Laboratoire de Géologie de Lyon, Terre, Planètes et Environnement, Université  
Claude Bernard Lyon 1/CNRS/École Normale Supérieure de Lyon, 69622 Villeurbanne Cedex,  
France; nicolasseon@orange.fr ; [romain.amiot@univ-lyon1.fr](mailto:romain.amiot@univ-lyon1.fr); [jeremy.martin@ens-lyon.fr](mailto:jeremy.martin@ens-lyon.fr);  
[christophe.lecuyer@univ-lyon1.fr](mailto:christophe.lecuyer@univ-lyon1.fr)

<sup>2</sup> School of GeoSciences, Grant Institute, University of Edinburgh, James Hutton Road, Edinburgh,  
EH9 3FE, UK; [mark.young@ed.ac.uk](mailto:mark.young@ed.ac.uk)

<sup>3</sup> 16 Rodwell Road, Weymouth, Dorset, DT4 8QL, UK; [heathermidd@yahoo.co.uk](mailto:heathermidd@yahoo.co.uk)

<sup>4</sup> Laboratoire d'Ecologie des Hydrosystèmes Naturels et Anthropisés, CNRS UMR 5023, Université  
Claude Bernard Lyon 1, France; [francois.fourel@univ-lyon1.fr](mailto:francois.fourel@univ-lyon1.fr)

<sup>5</sup> Paleospace, Avenue Jean Moulin, 14640 Villers sur mer, France; [l.picot@paleospace-villers.fr](mailto:l.picot@paleospace-villers.fr)

<sup>6</sup> Laboratoire PALEVOPRIM, UMR 7262 CNRS INEE & University of Poitiers, France;  
[xavier.valentin@univ-poitiers.fr](mailto:xavier.valentin@univ-poitiers.fr)

<sup>7</sup> Palaios, Research Association, 86300 Valdivienne, France

\* Corresponding author

**Keywords:** Metriorhynchidae, Teleosauridae, Jurassic, Thermophysiology, Oxygen and carbon  
isotopes, tooth apatite

## Summary

Teleosauridae and Metriorhynchidae were thalattosuchian crocodylomorph clades that secondarily adapted to marine life and coexisted during the Middle to Late Jurassic. While teleosaurid diversity collapsed at the end of the Jurassic, most likely as a result of a global cooling of the oceans and associated marine regressions, metriorhynchid diversity was largely unaffected, although the fossil record of Thalattosuchia is poor in the Cretaceous. In order to investigate the possible differences in thermophysiology between these two thalattosuchian lineages, we analysed stable oxygen isotope compositions (expressed as  $\delta^{18}\text{O}$  values) of tooth apatite from metriorhynchid and teleosaurid specimens. We then compared them to the  $\delta^{18}\text{O}$  values of coexisting endothermic ichthyosaurs and plesiosaurs, as well as ecto-poikilothermic chondrichthyans and osteichthyans. The distribution of  $\delta^{18}\text{O}$  values suggests that both teleosaurids and metriorhynchids had body temperatures intermediate between those of typical ecto-poikilothermic vertebrates and warm-blooded ichthyosaurs and plesiosaurs, metriorhynchids being slightly warmer than teleosaurids. We propose that metriorhynchids were able to raise their body temperature above that of the ambient environment by metabolic heat production, as endotherms do, but could not maintain a constant body temperature compared to fully homeothermic ichthyosaurs and plesiosaurs. Teleosaurids, on the other hand, may have raised their body temperature by mouth-gape basking, as modern crocodylians do, and benefited from the thermal inertia of their large body mass to maintain their body temperature above ambient one. Endothermy in metriorhynchids might have been a byproduct of their ecological adaptations to active pelagic hunting, and it probably allowed them to survive the global cooling of the Late Jurassic, thus explaining the selective extinction affecting Thalattosuchia at the Jurassic-Cretaceous boundary.

## Introduction

Extant crocodylians are archosaurs that rely on environmental sources of heat in order to raise and maintain their body temperature by behavioural thermoregulation in a restricted range bracketed by “critical minimum” and “critical maximum” temperatures (Cowles and Bogert, 1944; Pough and Gans, 1982). Within this critical range, crocodylians tend to keep their body temperature within a narrower activity range from about 25°C to 40°C as determined empirically for a few extant species (see Markwick (1998) for a review). Consequently, the spatial and temporal distribution of crocodylians is limited by the temperatures of their living environments and by their seasonal fluctuations. Due to their rather conservative growth morphology and their restricted latitudinal distribution today, and in the geologic record, extant and fossil representatives of the crown group have been used as climate proxies for more than a century (Berg, 1965; Crichton, 1825; Markwick, 1998; Owen, 1850). Based on observations of extant representatives of crown group Crocodylia Owen, 1842, Markwick (1998) proposed that the occurrence of fossil representatives imply a living Mean Annual Air temperature (MAT)  $\geq 14.2^{\circ}\text{C}$  and a Coldest Month Mean temperature (CMM)  $\geq 5.5^{\circ}\text{C}$ . These minimum limits have been tentatively used to constrain the climatic environment of long-extinct crocodylomorphs (Amiot et al., 2011), as well as the more distantly related clade Choristodera (Tarduno et al., 1998).

The crocodylomorph clade Thalattosuchia became secondarily adapted to marine life during the Mesozoic, and is subdivided into the families Teleosauridae Saint-Hilaire, 1831 and Metriorhynchidae Fitzinger, 1843. While teleosaurids retained a morphology reminiscent of typical semi-aquatic longirostrine crocodylomorphs, in having extensive osteoderm coverage and limbs adapted for terrestrial locomotion (Eudes-Deslongchamps, 1869), metriorhynchids had a more hydrodynamic body plan, with a hypocercal tail, and hydrofoil-like forelimbs (Fraas, 1902; Young et al., 2010). The bone histology of Callovian teleosaurids and metriorhynchids has been used to hypothesize that thalattosuchians were ectothermic (i.e. they rely on environmental sources of heat in order to raise their body temperature) and poikilothermic (i.e. their body temperature vary along

1  
2 [with that of their environment](#)), with teleosaurids being capable of mouth-gape basking on shore,  
3  
4 whereas metriorhynchids thermoregulated differently by staying close to the water surface (Hua and  
5  
6 De Buffrénil, 1996). Hua and de Buffrénil (1996) did note that young metriorhynchids may have  
7  
8 had a faster growth rate than wild extant crocodylians.  
9

10  
11 Martin et al. (2014) analysed the diversity of marine crocodylomorphs through the Mesozoic  
12  
13 and Paleogene, and observed a significant correlation between Sea Surface Temperatures (SST) and  
14  
15 generic diversity within four lineages, including the teleosaurids. Teleosaurids, at least European  
16  
17 representatives, experienced a diversity crash at the end of the Jurassic during a global cooling of  
18  
19 Tethyan waters (Lécuyer et al., 2003b), but according to Fanti et al. (2016) they may have  
20  
21 continued to survive until at least the Hauterivian along the southern coast of the Tethys Sea.  
22  
23 Metriorhynchids however, appear to have experienced an explosive radiation during the Callovian,  
24  
25 surviving through the Late Jurassic until the Aptian (Chiarenza et al., 2015). [When metriorhynchids](#)  
26  
27 [became extinct is currently unclear](#). Young et al. (2010) hypothesized a two-step extinction for  
28  
29 [Metriorhynchidae](#), first a diversity crash at the end Jurassic, then a final extinction during the cold  
30  
31 [icehouse interval of the Valanginian](#) (Pucéat et al., 2003). However, recent re-evaluations of  
32  
33 [Cretaceous fossils and updated phylogenetic studies have disproved both steps of this hypothesis](#),  
34  
35 [with post-Valanginian metriorhynchid specimens known and no fewer than four metriorhynchid](#)  
36  
37 [lineages crossed the Jurassic-Cretaceous boundary](#) (Chiarenza et al., 2015; Ósi et al., 2018; Young  
38  
39 [et al., 2014a](#)). The global thalattosuchian fossil record however is still poor, the known diversity of  
40  
41 this clade is still heavily biased by the European rock record, as well as global marine record  
42  
43 sampling biases for specific time spans (e.g. Aalenian, Oxfordian, Early Cretaceous). Regardless of  
44  
45 the geologic megabiases, there is still a conspicuous diversity mismatch within Thalattosuchia that  
46  
47 raised the question whether metriorhynchids evolved a distinct thermoregulatory strategy, such as  
48  
49 endothermic capabilities that would explain their diversity and survival under cool SSTs (Martin et  
50  
51 al., 2014).  
52  
53  
54  
55  
56  
57  
58  
59  
60

1  
2 In order to investigate the thermophysiology of metriorhynchids, we analysed the oxygen  
3  
4 isotope composition of the enamel phosphate of their teeth ( $\delta^{18}\text{O}_p$ ). Indeed, the  $\delta^{18}\text{O}_p$  value of  
5  
6 vertebrate apatite (the mineral constituting bone, teeth and some fish scales) depends on the  
7  
8 animal's body water  $\delta^{18}\text{O}_{bw}$  value, as well as its body temperature (Kohn, 1996; Longinelli and  
9  
10 Nuti, 1973a; Luz et al., 1984). For air breathing vertebrates, the body water has a  $\delta^{18}\text{O}_{bw}$  value  
11  
12 controlled by oxygen input coming from drinking water, food and inhaled oxygen (through  
13  
14 metabolic water production), as well as oxygen loss as water vapour through transcutaneous  
15  
16 evaporation, sweat, exhaled vapour, and liquid water in urine and feces, some of these losses being  
17  
18 associated with oxygen isotope fractionation (Kohn, 1996; Langlois et al., 2003). A fractionation  
19  
20 equation that relates the apatite phosphate  $\delta^{18}\text{O}_p$  value to that of body water and body temperature  
21  
22 ( $T_b$ ) can be adapted from the phosphate-water temperature scale previously established by  
23  
24 Longinelli and Nuti (1973a) and recently updated by Lécuyer et al. (2013). Such equation has  
25  
26 proven to be valid for a large range of invertebrate and vertebrate bioapatites (Kolodny et al., 1983;  
27  
28 Lécuyer et al., 1996; Longinelli and Nuti, 1973b, 1973a):  
29  
30  
31  
32  
33  
34  
35

$$36 T_b (\text{°C}) = 117.4 - 4.5 (\delta^{18}\text{O}_p - \delta^{18}\text{O}_{bw}) \quad (1)$$

37  
38  
39  
40

41 This relationship is commonly used to reconstitute past SSTs based on fish apatite as most  
42  
43 of them have a body temperature similar to that of their surrounding water, and a body water  $\delta^{18}\text{O}_{bw}$   
44  
45 value equal to ambient one (Dera et al., 2009; Lécuyer et al., 2003b; Picard et al., 1998; Pucéat et  
46  
47 al., 2003). Air-breathing vertebrates have a body water  $^{18}\text{O}$ -enriched relative to environmental  
48  
49 water, the magnitude of which depends on the amount of body water loss through exhaled  $\text{H}_2\text{O}_{\text{vapor}}$   
50  
51 and transcutaneous evaporation. Direct measurements have shown that terrestrial mammals and  
52  
53 birds have body water from about 4‰ to 7‰  $^{18}\text{O}$ -enriched relative to their drinking water  
54  
55 (Lazzerini et al., 2016; Longinelli, 1984; Wolf et al., 2013) whereas the body waters of semi aquatic  
56  
57  
58  
59  
60

1  
2 to aquatic crocodylians, turtles and birds are about 2‰ to 3‰  $^{18}\text{O}$ -enriched relative to their drinking  
3  
4 water (Amiot et al., 2007; Barrick et al., 1999; Lazzerini et al., 2016).  
5

6 In this study we estimate the body temperatures of metriorhynchid and teleosaurid  
7  
8 thalattosuchians recovered from five Jurassic localities in England and France. By comparing  $T_b$  of  
9  
10 thalattosuchians to those of associated ichthyosaurs or plesiosaurs, as well as to their ambient SST,  
11  
12 we show that metriorhynchids may have evolved some degree of endothermic-like  
13  
14 thermophysiology, whereas teleosaurids retained a more typical ecto-poikilothermy.  
15  
16  
17  
18  
19

## 20 **Materials and methods**

### 21 **Sample collection**

22  
23 We analysed eighty-eight fossil teeth of Jurassic fish and marine reptiles for their oxygen  
24  
25 isotope composition of phosphate ( $\delta^{18}\text{O}_p$ ), as well as for their oxygen ( $\delta^{18}\text{O}_c$ ) and carbon ( $\delta^{13}\text{C}_c$ )  
26  
27 isotope composition of apatite carbonate. The first English locality is Smallmouth Sands, Dorset  
28  
29 (**Figure 1**), from the lower part of the Kimmeridge Clay Formation dated as lower Kimmeridgian  
30  
31 (Young et al., 2014b). The second English locality is the Oxford Clay Formation Peterborough clay  
32  
33 pits, a world-famous fossiliferous collection of sites, dated to the middle Callovian, which has  
34  
35 yielded a rich marine fauna along with terrestrial elements (Martill and Hudson, 1991). The first  
36  
37 French locality is “Les Vaches Noires” of Normandy, dated as upper Callovian, where the “Marnes  
38  
39 de Dives” Formation crops out, consisting of marls where a rich fauna composed of marine and  
40  
41 terrestrial elements have been recovered (Lebrun and Courville, 2013). The second French locality  
42  
43 is the excavation of “Les Lourdines” near the eponymous quarry, dated as middle Callovian, which  
44  
45 is composed of white limestone (calcaire des Lourdines) where a marine fauna and some terrestrial  
46  
47 plants have been recovered (Barale et al., 1974). Finally, the locality of Cintheaux dated as  
48  
49 Bathonian consists of the “Pierre de Caen” limestone that has yielded a marine fauna with some  
50  
51 terrestrial elements (Riout, 1963). Studied tooth specimens include those of metriorhynchid and  
52  
53 teleosaurid thalattosuchians, ichthyosaurs, plesiosaurs, as well as of chondrichthyans and  
54  
55  
56  
57  
58  
59  
60

1  
2 osteichthyans (**Supplementary table 1**). When possible, tooth enamel was sampled using a  
3  
4 spherical diamond-tipped drill bit. For smaller teeth, the bulk enamel and dentin were ground using  
5  
6 an agate mortar and pestle.  
7  
8  
9

## 10 11 **Analytical techniques**

### 12 13 *Oxygen isotope analysis of biogenic apatite phosphate*

14  
15 Apatite powders have been treated following the wet chemistry protocol described by  
16  
17 Crowson et al. (1991) and slightly modified by Lécuyer et al. (1993). This protocol consists in the  
18  
19 isolation of phosphate ( $\text{PO}_4^{3-}$ ) from apatite as silver phosphate ( $\text{Ag}_3\text{PO}_4$ ) crystals using acid  
20  
21 dissolution and anion-exchange resin. For each sample, 20-30 mg of enamel powder was dissolved  
22  
23 in 2 mL of 2 M HF. The  $\text{CaF}_2$  residue was separated by centrifugation and the solution was  
24  
25 neutralized by adding 2.2 mL of 2 M KOH. Amberlite™ IRN 78 anion-exchange resin beads were  
26  
27 added to the solution to isolate the  $\text{PO}_4^{3-}$  ions. After 24 hours, the solution was removed, the resin  
28  
29 was rinsed with deionized water and eluted with 27.5 mL of 0.5 M  $\text{NH}_4\text{NO}_3$ . After 4 hours, 0.5 mL  
30  
31 of  $\text{NH}_4\text{OH}$  and 15 mL of an ammoniacal solution of  $\text{AgNO}_3$  were added and the solutions were  
32  
33 placed in a thermostated bath at 70 °C for 7 hours allowing for the precipitation of  $\text{Ag}_3\text{PO}_4$  crystals.  
34  
35 Oxygen isotope compositions were measured using a high temperature elemental analyzer  
36  
37 interfaced in continuous flow mode to an isotopic ratio mass spectrometer (Fourel et al., 2011) at  
38  
39 the Plateforme d'Ecologie Isotopique du Laboratoire d'Ecologie des Hydrosystèmes Naturels et  
40  
41 Anthropisés (LEHNA – UMR5023, Université Claude Bernard Lyon 1). For each sample, 5  
42  
43 aliquots of 300 µg of  $\text{Ag}_3\text{PO}_4$  were mixed with 300 µg of pure graphite powder loaded in silver foil  
44  
45 capsules. Pyrolysis was performed at 1450 °C using a varioPYROcube™ Elemental Analyzer  
46  
47 (Elementar GmbH) interfaced in Continuous Flow mode with an Isoprime™ Isotopic Ratio Mass  
48  
49 Spectrometer (Elementar UK Ltd). Measurements have been calibrated against silver phosphate  
50  
51 precipitated from the NBS120c (natural Miocene phosphorite from Florida), as well as against the  
52  
53 NBS127 (Barium sulfate precipitated using seawater from Monterey Bay, California, USA). The  
54  
55  
56  
57  
58  
59  
60



1 value of NBS120c was fixed at 21.7‰ (V-SMOW) according to Lécuyer et al. (1993), and that of  
2  
3 NBS127 set at the certified value of 9.3‰ (V-SMOW; (Halas and Szaran, 2001; Hut, 1987)) for  
4  
5 correction of instrumental mass fractionation during CO isotopic analysis. Silver phosphate  
6  
7 precipitated from standard NBS120c along with the silver phosphate samples derived from fossil  
8  
9 bioapatites was repeatedly analyzed ( $\delta^{18}\text{O}_p = 21.72 \pm 0.22\text{‰}$ ,  $n = 20$ ) to ensure that no fractionation  
10  
11 occurred during the wet chemistry. Data are reported as  $\delta^{18}\text{O}$  values with respect to V-SMOW (in  
12  
13  
14  
15 ‰  $\delta$  units).

### 21 *Oxygen and carbon isotope analysis of biogenic apatite carbonate*

22  
23 In order to remove potential organic contaminants as well as secondarily precipitated calcite,  
24  
25 about 10 mg of apatite powder was pre-treated following the protocol of Koch et al. (1997).  
26  
27 Powders were washed with a 2% NaOCl solution to remove organic matter, then rinsed five times  
28  
29 with double deionized water and air-dried at 40 °C for 24 hours. 0.1 M acetic acid was then added  
30  
31 and left for 24 hours, after which the powder was again rinsed five times with double deionized  
32  
33 water and then air-dried at 40 °C for 24 hours. The powder/solution ratio was kept constant at 0.04  
34  
35 g mL<sup>-1</sup> for both treatments. Stable isotope compositions of carbonate oxygen and carbon were  
36  
37 carried out at the Plateforme d'Ecologie Isotopique du Laboratoire d'Ecologie des Hydrosystèmes  
38  
39 Naturels et Anthropisés (LEHNA – UMR5023). The measurement were performed using an  
40  
41 isoFLOW system connected on line in continuous flow mode to a precisION mass spectrometer  
42  
43 (Elementar UK Ltd). For each sample, two aliquots of 2 mg of pretreated apatite powder were  
44  
45 loaded in LABCO Exetainer® 3.7 mL soda glass vials, round bottomed with exetainers caps  
46  
47 (LABCO UK Ltd) and were reacted with anhydrous phosphoric acid. The reaction took place at  
48  
49 90°C in a temperature regulated sample tray. The CO<sub>2</sub> gas generated during the acid digestion of the  
50  
51 carbonate sample was then transferred to the mass spectrometer via the centrION interface. A  
52  
53 calibrated CO<sub>2</sub> gas was used as a monitoring gas. Typical reproducibilities are 0.05‰ and 0.07‰  
54  
55 respectively for  $\delta^{13}\text{C}$  and  $\delta^{18}\text{O}$  measurements. For tooth apatite, the acid fractionation factor  $\alpha$   
56  
57  
58  
59  
60

(CO<sub>2</sub>-apatite carbonate) of 1.00773 determined for the NBS120c phosphate rock reference material has been selected (Passey et al., 2007). Calibrated material used is Carrara Marble ( $\delta^{18}\text{O}_{\text{V-PDB}} = -1.84\text{‰}$ ;  $\delta^{13}\text{C}_{\text{V-PDB}} = +2.03\text{‰}$ ; (Fourel et al., 2016) NBS18 ( $\delta^{18}\text{O}_{\text{V-PDB}} = -23.2\text{‰}$ ;  $\delta^{13}\text{C}_{\text{V-PDB}} = -5.01\text{‰}$ ) and NBS120c ( $\delta^{18}\text{O}_{\text{V-PDB}} = -1.13\text{‰}$ ;  $\delta^{13}\text{C}_{\text{V-PDB}} = -6.27\text{‰}$ ; (Passey et al., 2007) Isotopic compositions are quoted in the standard  $\delta$  notation relative to V-SMOW for oxygen and V-PDB for carbon.

## Results

Oxygen isotope compositions of apatite phosphate ( $\delta^{18}\text{O}_p$ ), apatite carbonate ( $\delta^{18}\text{O}_c$ ) and carbon isotope compositions of apatite carbonate ( $\delta^{13}\text{C}_c$ ) are reported in [Supplementary table 1](#) along with published  $\delta^{18}\text{O}_p$  values of teeth and bones of marine vertebrates (Anderson et al., 1994; Bernard et al., 2010; Billon-Bruyat et al., 2005). Analysed teeth have  $\delta^{18}\text{O}_p$  values ranging from 18.3‰ to 21.8‰ V-SMOW,  $\delta^{18}\text{O}_c$  values ranging from 24.1‰ to 28.3‰ V-SMOW and  $\delta^{13}\text{C}_c$  values ranging from -11.2‰ to 7.7‰ V-PDB. At the three sites of Smallmouth Sands, Peterborough and Les Vaches Noires, ichthyosaurs and plesiosaurs have slightly lower mean  $\delta^{18}\text{O}_p$  values than those of co-occurring thalattosuchians ([Supplementary table 2](#)).

For each locality, SST has been calculated from the  $\delta^{18}\text{O}_p$  values of fish using equation 1 ([Supplementary table 2](#)) considering that  $T_b \approx T_{\text{sw}}$ , and that  $\delta^{18}\text{O}_{\text{bw}} \approx \delta^{18}\text{O}_{\text{sw}}$  (Kolodny et al., 1983). Because average seawater  $\delta^{18}\text{O}_{\text{sw}}$  value may have varied between -1‰ and 0‰ V-SMOW depending on the amount of seawater stored as polar ice (Shackleton and Kennett, 1975), an average value of -0.5‰ was arbitrarily selected for temperature calculation, keeping in mind that the associated error in temperature calculation is about 2.3°C (based on the slope of 4.5 of equation 1).

The body temperature of studied marine reptiles have also been estimated using equation 1, and assuming a seawater-body water <sup>18</sup>O-enrichment of 2‰, a general enrichment observed among semi-aquatic and aquatic air breathing vertebrates, including extant crocodylomorphs

(Amiot et al., 2007; Barrick et al., 1999; Lazzarini et al., 2016). While ichthyosaurs and plesiosaurs have a body temperature within the 32-40°C range compatible with their known endo-homeothermy (Bernard et al., 2010), teleosaurids have lower  $T_b$  ranging from 27°C to 31°C, and metriorhynchids show intermediate body temperatures ranging from 29°C to 37°C ([Supplementary table 2](#)).

## Discussion

### *Original preservation of the stable isotope compositions*

Before discussing the thermophysiological significance of the oxygen and carbon isotope compositions of vertebrate apatites, pristine preservation of the isotopic record needs to be assessed. Indeed, biotic and abiotic processes leading to the decomposition, burial and fossilization of living organisms may alter the original isotopic composition of bioapatite through processes of secondary precipitation, ion adsorption or dissolution–recrystallization of bioapatite (Blake et al., 1997; Kolodny et al., 1996; Lécuyer et al., 2003a; Trueman et al., 2003; Zazzo et al., 2004a, 2004b). Although no method can definitely demonstrate whether the original isotope compositions have been kept, several ways to assess the preservation state of the isotopic record have been proposed (Fricke et al., 1998; Iacumin et al., 1996; Kolodny et al., 1996; Lécuyer et al., 2003a; Pucéat et al., 2004; Tütken et al., 2008; Zazzo et al., 2004b). In modern skeletal tissues of vertebrates, carbonate and phosphate precipitate close to equilibrium with body water, so the  $\delta^{18}\text{O}_p$  and  $\delta^{18}\text{O}_c$  values are positively correlated. Because isotopic exchange rates between carbonate-water and phosphate-water are significantly different, re-equilibration of both compounds during diagenesis is not expected and altered enamel should show isotopic shifts from the empirical  $\delta^{18}\text{O}_p$ – $\delta^{18}\text{O}_c$  line. Therefore, it is expected that the distribution of pristine  $\delta^{18}\text{O}$  values of fish and reptile tooth enamel should follow a line with a slope close to unity mimicking those established between the  $\delta^{18}\text{O}_c$  and  $\delta^{18}\text{O}_p$  values of modern mammals (Bryant et al., 1996; Chenery et al., 2012; Iacumin et al., 1996; Lécuyer et al., 2010; Zazzo et al., 2004b). Despite the narrow [range](#) in the distribution of oxygen isotope compositions (**Figure 2**), both  $\delta^{18}\text{O}_c$  and  $\delta^{18}\text{O}_p$  values fall within the range observed in

1  
2 extant and fossil marine vertebrates (Kolodny and Luz, 1991; Lécuyer et al., 2003b; Vennemann et  
3 al., 2001). This correlation shows that the oxygen isotope compositions of structural carbonate in  
4 both tooth and bone apatites have preserved to a certain degree its original record.  
5  
6  
7

8  
9 A clue to the primary preservation of stable carbon isotope composition of apatite carbonate  
10 is the systematic and significant difference in  $\delta^{13}\text{C}_c$  values between fish and marine reptiles (**Figure**  
11 **3**). Air breathing reptiles have  $\delta^{13}\text{C}_c$  values mainly reflecting that of their diets with an isotope  
12 fractionation that depends on their digestive physiology (Passey et al., 2005), as expected. In  
13 aquatic environments, the relationship between fish carbonate and diet  $\delta^{13}\text{C}$  values is complicated  
14 as a substantial amount of the carbon may be derived from dissolved inorganic carbon (DIC) of  
15 their ambient water (McConnaughey et al., 1997; Thorrold et al., 1997) with a higher  $\delta^{13}\text{C}$  value  
16 (Santos et al., 2011). Finally, the weight percentage of carbonate in analysed fossil apatites  
17 (**Supplementary table 1**) lies within the expected biological range of modern vertebrate apatites of  
18 2-13% (Brudevold and Soremark, 1967; Rink and Schwarcz, 1995; Vennemann et al., 2001). From  
19 these lines of evidence, we can assume that oxygen isotope compositions of apatite phosphate and  
20 carbonate have kept at least a significant part of their original information, and might be interpreted  
21 in **terms** of seawater temperature and thermophysiology.  
22  
23  
24  
25  
26  
27  
28  
29  
30  
31  
32  
33  
34  
35  
36  
37  
38  
39  
40

### 41 ***Body temperature reconstruction***

42  
43 In a previous study of marine reptile thermophysiology, Bernard et al. (2010) interpreted the  
44 oxygen isotope differences between marine reptiles and coexisting fish from a large range of water  
45 temperatures (estimated from 14°C to 34°C) in terms of  $T_b$  differences. They concluded that the  
46 three lineages Ichthyosauria, Plesiosauria and Mosasauridae were most likely endothermic and  
47 homeothermic marine reptiles, although for mosasaurids this was later debated (Harrell Jr et al.,  
48 2016; Motani, 2010). The method used in Bernard et al. (2010) cannot clearly identify the  
49 thermophysiology of thalattosuchians (**Figure 4**). Indeed, the available sample-set of thalattosuchian  
50 specimens is restricted to five localities with a narrow range of low paleolatitudes from about 29°N  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60

1  
2 to 36°N (**Supplementary table 2**; palaeolatitudes calculated using the online application of van  
3  
4 Hinsbergen et al. (2015)), as well as a narrow range of estimated paleotemperatures from 22±2°C to  
5  
6 27±2°C. Consequently, the sampled sites do not allow the two strategies of thermophysiology to be  
7  
8 clearly distinguished. Most metriorhynchids and teleosaurids show intermediate values between the  
9  
10 expected range for endotherms and ectotherms, with the exception of the metriorhynchids from the  
11  
12 middle Callovian of Les Lourdines which fall within the expected range of endotherms, and one  
13  
14 teleosaurid from the Bathonian of Cintheaux, which has a typical ectotherm signature (**Figure 4**).

15  
16 Using equation 1 and assuming that thalattosuchians have a  $\delta^{18}\text{O}_{\text{bw}}$  value of about 2‰ more  
17  
18 positive than their ambient seawater, both metriorhynchids and teleosaurids have a body  
19  
20 temperature above their environmental one, metriorhynchids being slightly warmer than  
21  
22 teleosaurids in average (**Supplementary table 2**; **Figure 5**). However, thalattosuchian  $T_b$  are  
23  
24 systematically below the calculated ones of co-occurring ichthyosaurs and plesiosaurs for which the  
25  
26 body temperature was calculated from the same equation 1 and using a body water  $\delta^{18}\text{O}_{\text{bw}}$  value  
27  
28 similar to that of thalattosuchians.  
29  
30  
31  
32  
33

34  
35 An endothermic-like thermophysiology interpreted from metriorhynchid  $\delta^{18}\text{O}_p$  values seems  
36  
37 likely according to their known morphology and ecology as active pelagic predators (Andrade et al.,  
38  
39 2010; Fernández and Gasparini, 2000; Hua and De Buffrénil, 1996; Massare, 1987; Young et al.,  
40  
41 2013, 2012). Active predation would require elevated metabolic rates compatible with an  
42  
43 endothermic thermophysiology. This seems plausible considering the large suite of evidence for an  
44  
45 endothermic ancestral condition for Archosauria, and a reversal within the crocodylomorph lineage  
46  
47 to an ectothermic state coming from different fields of biology, including developmental biology  
48  
49 (Seymour et al., 2004), physiology (Farmer and Sanders, 2010), anatomy (Summers, 2005),  
50  
51 palaeohistology (de Ricqlès et al., 2008), and phylogenetic signal extraction (Legendre et al., 2016).  
52  
53 Endothermy within metriorhynchids might have been inherited from their archosaur ancestors, and  
54  
55 “reactivated” along with the acquisition of morphological adaptations to active pelagic predation.  
56  
57 However, body temperature regulation seems to have been limited as observed  $T_b$  vary with varying  
58  
59  
60

1  
2 seawater temperature (**Figure 5**). Limited thermoregulatory capacities would be compatible with  
3  
4 the restricted range of paleolatitudinal occurrences compared to fully endo-homeothermic  
5  
6 ichthyosaurs and plesiosaurs having been found from equatorial to polar seas (Bardet et al., 2014).  
7

8  
9 Interpretation of the oxygen isotope composition of teleosaurid apatite in terms of body  
10  
11 temperature may be strongly biased by their possible semi-aquatic and eurhaline ecology, some  
12  
13 species having been found in estuarine or freshwater environments (Buffetaut, 1982; Martin et al.,  
14  
15 2019, 2016; Wilberg et al., 2019). As most teleosaurids retained a typical semi-aquatic  
16  
17 crocodylomorph morphology (external mandibular fenestrae, extensive osteoderm cover and limbs  
18  
19 adapted to terrestrial locomotion), it is hypothesized that they were ectothermic and poikilothermic  
20  
21 ambush predators spending most of their time motionless, and mouth-gape basking like modern  
22  
23 crocodylians. However, during the late Kimmeridgian-early Tithonian there is evidence for a  
24  
25 subclade of teleosaurids that became more pelagic (Foffa et al., 2019). The oxygen isotope  
26  
27 composition of teleosaurid apatite indicates that they kept a body temperature lower than that of  
28  
29 ichthyosaurs and plesiosaurs, but close to those of metriorhynchids. However, more estuarine living  
30  
31 environments can be characterized by more negative  $\delta^{18}\text{O}$  values as a result of the mixing between  
32  
33 seawater and river waters having negative  $\delta^{18}\text{O}$  values such as in the case of the San Francisco Bay  
34  
35 (Ingram et al., 1996). In this example, waters from the Sacramento river having  $\delta^{18}\text{O}$  values  
36  
37 ranging from about -12‰ to -10‰ mix with seawater of 0‰ and results in estuarine waters that  
38  
39 can have  $\delta^{18}\text{O}$  values of -3‰ to -10‰. Keeping this in mind, then the calculated body temperature  
40  
41 of teleosaurids living in estuarine environments would be lower, and their apatite  $\delta^{18}\text{O}$  values would  
42  
43 fit within the expected range of ectotherms. Moreover, the possible semi-aquatic lifestyle of  
44  
45 teleosaurids could at least partly account for their elevated  $\delta^{18}\text{O}_p$  values as a result of body water  
46  
47 loss through transcutaneous evaporation. In the studied localities where teleosaurid specimens were  
48  
49 found, fossil remains of continental vertebrates and plants have been found, indicating a proximity  
50  
51 to landmasses and a possible estuarine origin of teleosaurids. Gigantothermy and behavioural  
52  
53 thermoregulation may also account for the calculated body temperatures of teleosaurids close to  
54  
55  
56  
57  
58  
59  
60

1 those of metriorhynchids. Today, large marine crocodylians, such as ~1000 kg adult individuals of  
2 the saltwater crocodile (*Crocodylus porosus*), are able to raise their body temperature well above  
3 ambient temperatures through mouth-gape basking behaviours and can retain this elevation by the  
4 thermal inertia of their large body size (Grigg et al., 1998; Seebacher et al., 1999). Large  
5 teleosaurids may have used a similar behavioural thermoregulation as extant marine crocodiles (*C.*  
6 *porosus*) and would have raised their body temperatures close to that of Jurassic metriorhynchids  
7 and maintained it within a narrow range. This would explain the apparent tendency of teleosaurid-  
8 fish isotope difference to parallel that of endo-homeothermic ichthyosaurs and plesiosaurs (Figure  
9 4). Most metriorhynchid specimens were of smaller body-size than teleosaurids (Young et al., 2016,  
10 2011), and metriorhynchids would have been unable to mouth-gape bask onshore (Young et al.,  
11 2010).

12 Based on the available isotopic dataset, it seems likely that teleosaurids retained a typical  
13 ecto-poikilothermic thermophysiology in agreement with their morphology and ecology, whereas  
14 metriorhynchids may have been endothermic, being able to raise their body temperature above an  
15 ambient one, and close to that of other warm-blooded marine reptiles. However, metriorhynchids  
16 could not have achieved efficient thermoregulation as suggested from their varying body  
17 temperature along with varying SST.

### 18 **Concluding remarks**

19 The possible difference in thermophysologies between metriorhynchids and teleosaurids  
20 inferred from their stable oxygen isotope composition of apatite can at least partly explain the  
21 peculiar thalattosuchian biodiversity pattern of the Jurassic and Early Cretaceous (Martin et al.,  
22 2014; Young et al., 2014a). Teleosaurid diversity crashed at the end of the Jurassic, a time of global  
23 marine temperature decline, probably as a result of the temperature change or the global regression  
24 affecting their ecological niches (Bardet et al., 2014). Endothermy may have helped  
25 metriorhynchids to cope with global cooling, and due to their pelagic ecology, they might not have

1  
2 been as affected by the marine regressions as teleosaurids. This may explain their success across the  
3  
4 Jurassic-Cretaceous boundary (Chiarenza et al., 2015; Tennant et al., 2017; Young et al., 2014a). At  
5  
6 this point, we cannot only speculate when metriorhynchids became extinction, or why. All we can  
7  
8 mention is that Jurassic metriorhynchids had an imperfect endothermic and poikilothermic  
9  
10 thermophysiology. We hope future studies will investigate whether the Late Jurassic pelagic  
11  
12 subclade of teleosaurids began to develop endothermic capabilities, and test whether Cretaceous  
13  
14 metriorhynchids, the most marine adapted of all thalattosuchians (Hua et al., 2000; Young et al.,  
15  
16 2010), evolved toward enhanced thermoregulatory abilities.  
17  
18  
19  
20  
21

## 22 Acknowledgments

23  
24 The authors would like to thank the Hunterian Museum of Glasgow, Mr and Mrs Penetier  
25  
26 for providing material from England and France and Gilles Cuny for his constructive comments.  
27  
28  
29 MTY is financially supported by a Leverhulme Trust Research Project grant (RPG-2017-167).  
30  
31  
32  
33

## 34 References

- 35  
36 Amiot, R., Lécuyer, C., Escarguel, G., Billon-Bruyat, J.-P., Buffetaut, E., Langlois, C., Martin, S.,  
37 Martineau, F., Mazin, J.-M., 2007. Oxygen isotope fractionation between crocodylian  
38 phosphate and water. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* 243, 412–420.  
39 <https://doi.org/10.1016/j.palaeo.2006.08.013>  
40  
41 Amiot, R., Wang, Xu, Zhou, Z., Wang, Xiaolin, Buffetaut, E., Lécuyer, C., Ding, Z., Fluteau, F.,  
42 Hibino, T., Kusuhashi, N., Mo, J., Suteethorn, V., Wang, Y., Xu, X., Zhang, F., 2011.  
43 Oxygen isotopes of East Asian dinosaurs reveal exceptionally cold Early Cretaceous  
44 Climates. *Proc. Natl. Acad. Sci.* 108, 5179–5183. <https://doi.org/10.1073/pnas.1011369108>  
45  
46 Anderson, T.F., Popp, B.N., Williams, A.C., Ho, L.-Z., Hudson, J.D., 1994. The stable isotopic  
47 records of fossils from the Peterborough Member, Oxford Clay Formation (Jurassic), UK:  
48 palaeoenvironmental implications. *J. Geol. Soc.* 151, 125–138.  
49 <https://doi.org/10.1144/gsjgs.151.1.0125>  
50  
51 Andrade, M.B. de, Young, M.T., Desojo, J.B., Brusatte, S.L., 2010. The evolution of extreme  
52 hypercarnivory in Metriorhynchidae (Mesoeucrocodylia: Thalattosuchia) based on evidence  
53 from microscopic denticle morphology. *J. Vertebr. Paleontol.* 30, 1451–1465.  
54 <https://doi.org/10.1080/02724634.2010.501442>  
55  
56 Barale, G., Cariou, E., Radureau, G., 1974. Etude biostratigraphique et paléobotanique des  
57 gisements de calcaire blanc callovien au Nord de Poitiers. *Geobios* 7, 43–69.  
58 [https://doi.org/10.1016/S0016-6995\(74\)80018-5](https://doi.org/10.1016/S0016-6995(74)80018-5)  
59  
60 Bardet, N., Falconnet, J., Fischer, V., Houssaye, A., Jouve, S., Suberbiola, X.P., Perez-García, A.,  
Rage, J.-C., Vincent, P., 2014. Mesozoic marine reptile palaeobiogeography in response to  
drifting plates. *Gondwana Res.* 26, 869–887. <https://doi.org/10.1016/j.gr.2014.05.005>



- 1  
2 Barrick, R.E., Fischer, A.G., Showers, W.J., 1999. Oxygen isotopes from turtle bone: applications  
3 for terrestrial paleoclimates? *Palaios* 14, 186–191. <https://doi.org/10.2307/3515374>
- 4 Berg, D.E., 1965. Krokodile als Klimazeugen. *Geol. Rundsch.* 54, 328–333.
- 5 Bernard, A., Lécuyer, C., Vincent, P., Amiot, R., Bardet, N., Buffetaut, E., Fourel, F., Martineau,  
6 F., Mazin, J.-M., Prieur, A., 2010. Regulation of body temperature by some Mesozoic  
7 marine reptiles. *Science* 328, 1379–1382. <https://doi.org/10.1126/science.1187443>
- 8 Billon-Bruyat, J.-P., Lécuyer, C., Martineau, F., Mazin, J.-M., 2005. Oxygen isotope compositions  
9 of Late Jurassic vertebrate remains from lithographic limestones of western Europe:  
10 implications for the ecology of fish, turtles, and crocodylians. *Palaeogeogr. Palaeoclimatol.*  
11 *Palaeoecol.* 216, 359–375.
- 12  
13 Blake, R.E., O’Neil, J.R., Garcia, G.A., 1997. Oxygen isotope systematics of biologically mediated  
14 reactions of phosphate: I. Microbial degradation of organophosphorus compounds.  
15 *Geochim. Cosmochim. Acta* 61, 4411–4422.
- 16 Brudevold, F., Soremark, R., 1967. Chemistry of the mineral phase of enamel, in: Mills, A. (Ed.),  
17 *Structural and Chemical Organization of Teeth*, Volume 2. Elsevier, Amsterdam, pp. 247–  
18 277.
- 19  
20 Bryant, D.J., Koch, P.L., Froelich, P.N., Showers, W.J., Genna, B.J., 1996. Oxygen isotope  
21 partitioning between phosphate and carbonate in mammalian apatite. *Geochim. Cosmochim.*  
22 *Acta* 60, 5145–5148.
- 23 Buffetaut, E., 1982. Radiation évolutive, paléoécologie et biogéographie des crocodyliens  
24 mésosuchiens, *Mémoires de la Société géologique de France. Nouvelle série. Société*  
25 *Géologique de France*, Paris.
- 26  
27 Chenery, C.A., Pashley, V., Lamb, A.L., Sloane, H.J., Evans, J.A., 2012. The oxygen isotope  
28 relationship between the phosphate and structural carbonate fractions of human bioapatite.  
29 *Rapid Commun. Mass Spectrom.* 26, 309–319.
- 30  
31 Chiarenza, A.A., Foffa, D., Young, M.T., Insacco, G., Cau, A., Carnevale, G., Catanzariti, R., 2015.  
32 The youngest record of metriorhynchid crocodylomorphs, with implications for the  
33 extinction of *Thalattosuchia*. *Cretac. Res.* 56, 608–616.  
34 <https://doi.org/10.1016/j.cretres.2015.07.001>
- 35  
36 Cowles, R.B., Bogert, C.M., 1944. A preliminary study of the thermal requirements of desert  
37 reptiles. *Bull. Am. Mus. Nat. Hist.* 83, 261–296.
- 38  
39 Crichton, A., 1825. On the Climate of the Antediluvian World and Its Independence of Solar  
40 Influence: And on the Formation of Granite. *Ann. Philos.* 9, 207–217.
- 41  
42 Crowson, R.A., Showers, W.J., Wright, E.K., Hoering, T.C., 1991. Preparation of phosphate  
43 samples for oxygen isotope analysis. *Anal. Chem.* 63, 2397–2400.  
44 <https://doi.org/10.1021/ac00020a038>
- 45  
46 de Ricqlès, A., Padian, K., Knoll, F., Horner, J.R., 2008. On the origin of high growth rates in  
47 archosaurs and their ancient relatives: complementary histological studies on Triassic  
48 archosauriforms and the problem of a “phylogenetic signal” in bone histology, in: *Annales*  
49 *de Paléontologie*. Elsevier, pp. 57–76. <https://doi.org/10.1016/j.annpal.2008.03.002>
- 50  
51 Dera, G., Pucéat, E., Pellenard, P., Neige, P., Delsate, D., Joachimski, M.M., Reisberg, L.,  
52 Martinez, M., 2009. Water mass exchange and variations in seawater temperature in the NW  
53 Tethys during the Early Jurassic: evidence from neodymium and oxygen isotopes of fish  
54 teeth and belemnites. *Earth Planet. Sci. Lett.* 286, 198–207.  
55 <https://doi.org/10.1016/j.epsl.2009.06.027>
- 56  
57 Eudes-Deslongchamps, E., 1869. *Notes Paléontologiques*. Le Blanc Hardel & Savy, Caen and Paris.
- 58  
59 Fant, F., Miyashita, T., Cantelli, L., Mnasri, F., Dridi, J., Contessi, M., Cau, A., 2016. The largest  
60 thalattosuchian (Crocodylomorpha) supports teleosaurid survival across the Jurassic-  
Cretaceous boundary. *Cretac. Res.* 61, 263–274.  
<https://doi.org/10.1016/j.cretres.2015.11.011>
- Farmer, C.G., Sanders, K., 2010. Unidirectional airflow in the lungs of alligators. *Science* 327,  
338–340. <https://doi.org/10.1126/science.1180219>

- 1  
2 Fernández, M., Gasparini, Z., 2000. Salt glands in a Tithonian metriorhynchid crocodyliform and  
3 their physiological significance. *Lethaia* 33, 269–276.  
4 <https://doi.org/10.1080/002411600750053835>  
5  
6 Fitzinger, L., 1843. *Systema Reptilium, Fasciculus Primus, Amblyglossae*. Braumüller et Seidel,  
7 Vienna.
- 8 Foffa, D., Johnson, M.M., Young, M.T., Steel, L., Brusatte, S.L., 2019. Revision of the Late  
9 Jurassic deep-water teleosauroid crocodylomorph *Teleosaurus megarhinus* Hulke, 1871 and  
10 evidence of pelagic adaptations in Teleosauroidea. *PeerJ* 7, e6646.  
11 <https://doi.org/10.7717/peerj.6646>  
12  
13 Fourel, F., Martineau, F., Lécuyer, C., Kupka, H.-J., Lange, L., Ojeimi, C., Seed, M., 2011.  $^{18}\text{O}/^{16}\text{O}$   
14 ratio measurements of inorganic and organic materials by elemental analysis–pyrolysis–  
15 isotope ratio mass spectrometry continuous-flow techniques. *Rapid Commun. Mass*  
16 *Spectrom.* 25, 2691–2696. <https://doi.org/10.1002/rcm.5056>  
17  
18 Fourel, F., Martineau, F., Tóth, E.E., Görög, A., Escarguel, G., Lécuyer, C., 2016. Carbon and  
19 oxygen isotope variability among foraminifera and ostracod carbonated shells. *Ann. Univ.*  
20 *Mariae Curie-Sklodowska Sect. AAA–Physica* 70, 133–156.
- 21 Fraas, E., 1902. Die Meer-Krocodilier (Thalattosuchia) des oberen Jura unter specieller  
22 Berücksichtigung von *Dacosaurus* und *Geosaurus*. *Palaeontographica* 49, 1–72.
- 23 Fricke, H.C., Clyde, W.C., O’Neil, J.R., Gingerich, P.D., 1998. Evidence for rapid climate change  
24 in North America during the latest Paleocene thermal maximum: oxygen isotope  
25 compositions of biogenic phosphate from the Bighorn Basin (Wyoming). *Earth Planet. Sci.*  
26 *Lett.* 160, 193–208.
- 27 Grigg, G.C., Seebacher, F., Beard, L.A., Morris, D., 1998. Thermal relations of large crocodiles,  
28 *Crocodylus porosus*, free-ranging in a naturalistic situation. *Proc. R. Soc. Lond. B Biol. Sci.*  
29 265, 1793–1799. <https://doi.org/10.1098/rspb.1998.0504>  
30  
31 Halas, S., Szaran, J., 2001. Improved thermal decomposition of sulfates to  $\text{SO}_2$  and mass  
32 spectrometric determination of  $\delta^{34}\text{S}$  of IAEA SO-5, IAEA SO-6 and NBS-127 sulfate  
33 standards. *Rapid Commun. Mass Spectrom.* 15, 1618–1620. <https://doi.org/10.1002/rcm.416>  
34  
35 Harrell Jr, T.L., Pérez-Huerta, A., Suarez, C.A., 2016. Endothermic mosasaurs? Possible  
36 thermoregulation of Late Cretaceous mosasaurs (Reptilia, Squamata) indicated by stable  
37 oxygen isotopes in fossil bioapatite in comparison with coeval marine fish and pelagic  
38 seabirds. *Palaeontology* 59, 351–363. <https://doi.org/10.1111/pala.12240>  
39  
40 Hua, S., De Buffrénil, V., 1996. Bone histology as a clue in the interpretation of functional  
41 adaptations in the Thalattosuchia (Reptilia, Crocodylia). *J. Vertebr. Paleontol.* 16, 703–717.  
42 <https://doi.org/10.1080/02724634.1996.10011359>  
43  
44 Hua, S., Vignaud, P., Atrops, F., Clément, A., 2000. *Enaliosuchus macrospondylus* Koken, 1883  
45 (Crocodylia, Metriorhynchidae) du Valanginien de Barret-le-Bas (Hautes Alpes, France): un  
46 cas unique de remontée des narines externes parmi les crocodyliens. *Géobios* 33, 467–474.  
47 [https://doi.org/10.1016/S0016-6995\(00\)80080-7](https://doi.org/10.1016/S0016-6995(00)80080-7)  
48  
49 Hut, G., 1987. Consultants’ group meeting on stable isotope reference samples for geochemical and  
50 hydrological investigations. accessible at :  
51 [http://www.iaea.org/inis/collection/NCLCollectionStore/\\_Public/18/075/18075746.pdf](http://www.iaea.org/inis/collection/NCLCollectionStore/_Public/18/075/18075746.pdf) 1–  
52 43.
- 53 Iacumin, P., Bocherens, H., Mariotti, A., Longinelli, A., 1996. Oxygen isotope analyses of co-  
54 existing carbonate and phosphate in biogenic apatite: a way to monitor diagenetic alteration  
55 of bone phosphate? *Earth Planet. Sci. Lett.* 142, 1–6.
- 56 Ingram, B.L., Conrad, M.E., Ingle, J.C., 1996. Stable isotope and salinity systematics in estuarine  
57 waters and carbonates: San Francisco Bay. *Geochim. Cosmochim. Acta* 60, 455–467.  
58 [https://doi.org/10.1016/0016-7037\(95\)00398-3](https://doi.org/10.1016/0016-7037(95)00398-3)  
59  
60 Koch, P.L., Tuross, N., Fogel, M.L., 1997. The effects of sample treatment and diagenesis on the  
isotopic integrity of carbonate in biogenic hydroxylapatite. *J. Archaeol. Sci.* 24, 417–429.

- 1  
2 Kohn, M.J., 1996. Predicting animal  $\delta^{18}\text{O}$ : Accounting for diet and physiological adaptation.  
3 Geochim. Cosmochim. Acta 60, 4811–4829. [https://doi.org/10.1016/S0016-7037\(96\)00240-](https://doi.org/10.1016/S0016-7037(96)00240-2)  
4 2
- 5 Kolodny, Y., Luz, B., 1991. Oxygen isotopes in phosphate of fossil fish — Devonian to recent, in:  
6 Taylor, H.P., O'Neil, J.R., Kaplan, I.R. (Eds.), Stable Isotope Geochemistry: A Tribute to  
7 Samuel Epstein. Geochemical Society, University Park, pp. 105–119.
- 8 Kolodny, Y., Luz, B., Navon, O., 1983. Oxygen isotope variations in phosphate of biogenic  
9 apatites, I. Fish bone apatite-rechecking the rules of the game. Earth Planet. Sci. Lett. 64,  
10 398–404. [https://doi.org/10.1016/0012-821X\(83\)90100-0](https://doi.org/10.1016/0012-821X(83)90100-0)
- 11 Kolodny, Y., Luz, B., Sander, M., Clemens, W.A., 1996. Dinosaur bones: fossils or pseudomorphs?  
12 The pitfalls of physiology reconstruction from apatitic fossils. Palaeogeogr. Palaeoclimatol.  
13 Palaeoecol. 126, 161–171.
- 14 Langlois, C., Simon, L., Lécuyer, C., 2003. Box-modeling of bone and tooth phosphate oxygen  
15 isotope compositions as a function of environmental and physiological parameters. Isotopes  
16 Environ. Health Stud. 39, 259–272. <https://doi.org/10.1080/10256010310001621146>
- 17 Lazzerini, N., Lécuyer, C., Amiot, R., Angst, D., Buffetaut, E., Fourel, F., Daux, V., Betancort, J.F.,  
18 Sánchez Marco, A., Lomoschitz, A., 2016. Oxygen isotope fractionation between bird  
19 eggshell calcite and body water: application to fossil eggs from Lanzarote (Canary Islands).  
20 Sci. Nat. 103, 81. <https://doi.org/10.1007/s00114-016-1404-x>
- 21 Lebrun, P., Courville, P., 2013. Le Jurassique des falaises des Vaches-noires. Fossiles Hors serie 4,  
22 16–33.
- 23 Lécuyer, C., Amiot, R., Touzeau, A., Trotter, J., 2013. Calibration of the phosphate  $\delta^{18}\text{O}$   
24 thermometer with carbonate–water oxygen isotope fractionation equations. Chem. Geol.  
25 347, 217–226. <https://doi.org/10.1016/j.chemgeo.2013.03.008>
- 26 Lécuyer, C., Balter, V., Martineau, F., Fourel, F., Bernard, A., Amiot, R., Gardien, V., Otero, O.,  
27 Legendre, S., Panczer, G., 2010. Oxygen isotope fractionation between apatite-bound  
28 carbonate and water determined from controlled experiments with synthetic apatites  
29 precipitated at 10–37°C. Geochim. Cosmochim. Acta 74, 2072–2081.  
30 <https://doi.org/10.1016/j.gca.2009.12.024>
- 31 Lécuyer, C., Bogey, C., Garcia, J.-P., Grandjean, P., Barrat, J.A., Floquet, M., Bardet, N., Pereda-  
32 Superbiola, X., 2003a. Stable isotope composition and rare earth element content of  
33 vertebrate remains from the Late Cretaceous of northern Spain (Laño): did the  
34 environmental record survive? Palaeogeogr. Palaeoclimatol. Palaeoecol. 193, 457–471.
- 35 Lécuyer, C., Grandjean, P., Emig, C., 1996. Determination of oxygen isotope fractionation between  
36 water and phosphate from living lingulids: Potential application to palaeoenvironmental  
37 studies. Palaeogeogr. Palaeoclimatol. Palaeoecol. 126, 101–108.
- 38 Lécuyer, C., Grandjean, P., O'Neil, J.R., Cappetta, H., Martineau, F., 1993. Thermal excursions in  
39 the ocean at the Cretaceous-Tertiary boundary (northern Morocco):  $\delta^{18}\text{O}$  record of  
40 phosphatic fish debris. Palaeogeogr. Palaeoclimatol. Palaeoecol. 105, 235–243.  
41 [https://doi.org/10.1016/0031-0182\(93\)90085-W](https://doi.org/10.1016/0031-0182(93)90085-W)
- 42 Lécuyer, C., Picard, S., Garcia, J.-P., Sheppard, S.M., Grandjean, P., Dromart, G., 2003b. Thermal  
43 evolution of Tethyan surface waters during the Middle-Late Jurassic: Evidence from  $\delta^{18}\text{O}$   
44 values of marine fish teeth. Paleoceanography 18, 1076–1091.  
45 <https://doi.org/10.1029/2002PA000863>
- 46 Legendre, L.J., Guénard, G., Botha-Brink, J., Cubo, J., 2016. Palaeohistological evidence for  
47 ancestral high metabolic rate in archosaurs. Syst. Biol. 65, 989–996.  
48 <https://doi.org/10.1093/sysbio/syw033>
- 49 Longinelli, A., 1984. Oxygen isotopes in mammal bone phosphate: A new tool for  
50 paleohydrological and paleoclimatological research? Geochim. Cosmochim. Acta 48, 385–  
51 390. [https://doi.org/10.1016/0016-7037\(84\)90259-X](https://doi.org/10.1016/0016-7037(84)90259-X)
- 52 Longinelli, A., Nuti, S., 1973a. Revised phosphate-water isotopic temperature scale. Earth Planet.  
53 Sci. Lett. 19, 373–376. [https://doi.org/10.1016/0012-821X\(73\)90088-5](https://doi.org/10.1016/0012-821X(73)90088-5)

- 1  
2 Longinelli, A., Nuti, S., 1973b. Oxygen isotope measurements of phosphate from fish teeth and  
3 bones. *Earth Planet. Sci. Lett.* 20, 337–340. [https://doi.org/10.1016/0012-821X\(73\)90007-1](https://doi.org/10.1016/0012-821X(73)90007-1)
- 4 Luz, B., Kolodny, Y., Horowitz, M., 1984. Fractionation of oxygen isotopes between mammalian  
5 bone-phosphate and environmental drinking water. *Geochim. Cosmochim. Acta* 48, 1689–  
6 1693. [https://doi.org/10.1016/0016-7037\(84\)90338-7](https://doi.org/10.1016/0016-7037(84)90338-7)
- 7  
8 Markwick, P.J., 1998. Fossil crocodylians as indicators of Late Cretaceous and Cenozoic climates:  
9 implications for using palaeontological data in reconstructing palaeoclimate. *Palaeogeogr.*  
10 *Palaeoclimatol. Palaeoecol.* 137, 205–271. [https://doi.org/10.1016/S0031-0182\(97\)00108-9](https://doi.org/10.1016/S0031-0182(97)00108-9)
- 11 Martill, D.M., Hudson, J.D. (Eds.), 1991. *Fossils of the Oxford Clay*, Palaeontological Association  
12 Field guide to fossils. The Palaeontological Association, London.
- 13 Martin, J.E., Amiot, R., Lécuyer, C., Benton, M.J., 2014. Sea surface temperature contribute to  
14 marine crocodylian evolution. *Nat. Commun.* 5, 1–7. <https://doi.org/10.1038/ncomms5658>
- 15 Martin, J.E., Deesri, U., Liard, R., Wattanapituksakul, A., Suteethorn, S., Lauprasert, K., Tong, H.,  
16 Buffetaut, E., Suteethorn, V., Suan, G., Telouk, P., Balter, V., 2016. Strontium isotopes and  
17 the long-term residency of thalattosuchians in the freshwater environment. *Paleobiology* 42,  
18 143–156.
- 19  
20 Martin, J.E., Suteethorn, S., Lauprasert, K., Tong, H., Buffetaut, E., Liard, R., Salaviale, C., Deesri,  
21 U., Suteethorn, V., Claude, J., 2019. A new freshwater teleosaurid from the Jurassic of  
22 northeastern Thailand. *J. Vertebr. Paleontol.* e1549059.  
23 <https://doi.org/10.1080/02724634.2018.1549059>
- 24  
25 Massare, J.A., 1987. Tooth morphology and prey preference of Mesozoic marine reptiles. *J.*  
26 *Vertebr. Paleontol.* 7, 121–137. <https://doi.org/10.1080/02724634.1987.10011647>
- 27 McConnaughey, T.A., Burdett, J., Whelan, J.F., Paull, C.K., 1997. Carbon isotopes in biological  
28 carbonates: respiration and photosynthesis. *Geochim. Cosmochim. Acta* 61, 611–622.
- 29 Motani, R., 2010. Warm-blooded “sea dragons”? *Science* 328, 1361–1362.
- 30 Ósi, A., Young, M.T., Galácz, A., Rabi, M., 2018. A new large-bodied thalattosuchian  
31 crocodyliform from the Lower Jurassic (Toarcian) of Hungary, with further evidence of the  
32 mosaic acquisition of marine adaptations in Metriorhynchoidea. *PeerJ* 6, e4668.  
33 <https://doi.org/10.7717/peerj.4668>
- 34  
35 Owen, R., 1850. On the fossil crocodylia of England. *Edinb. New Philos. J.* 49, 248–250.
- 36 Owen, R., 1842. Report on British Fossil Reptiles. Part II. Rep. Br. Assoc. Adv. Sci. Plymouth  
37 Meet. 1841 60–240.
- 38  
39 Passey, B.H., Cerling, T.E., Levin, N.E., 2007. Temperature dependence of oxygen isotope acid  
40 fractionation for modern and fossil tooth enamels. *Rapid Commun. Mass Spectrom.* 21,  
41 2853–2859.
- 42 Picard, S., Garcia, J.-P., Lécuyer, C., Sheppard, S.M.F., Cappetta, H., Emig, C., 1998.  $\delta^{18}\text{O}$  values  
43 of coexisting brachiopods and fish: Temperature differences and estimates of paleo-water  
44 depths. *Geology* 26, 975–978.
- 45  
46 Pough, F.H., Gans, C., 1982. The vocabulary of reptilian thermoregulation, in: Gans, C. (Ed.),  
47 *Biology of the Reptilia* Vol. 12. Physiology, C. Physiological Ecology. Academic Press,  
48 London, pp. 17–23.
- 49 Pross, J., Link, E., Ruf, M., Aigner, T., 2006. Delineating sequence stratigraphic patterns in deeper  
50 ramp carbonates: Quantitative palynofacies data from the Upper Jurassic (Kimmeridgian) of  
51 southwest Germany. *J. Sediment. Res.* 76, 524–538. <https://doi.org/10.2110/jsr.2006.031>
- 52  
53 Pucéat, E., Lécuyer, C., Sheppard, S.M., Dromart, G., Reboulet, S., Grandjean, P., 2003. Thermal  
54 evolution of Cretaceous Tethyan marine waters inferred from oxygen isotope composition  
55 of fish tooth enamels. *Paleoceanography* 18, 1029. <https://doi.org/10.1029/2002PA000823>
- 56  
57 Pucéat, E., Reynard, B., Lécuyer, C., 2004. Can crystallinity be used to determine the degree of  
58 chemical alteration of biogenic apatites? *Chem. Geol.* 205, 83–97.
- 59  
60 Rink, W.J., Schwarcz, H.P., 1995. Tests for diagenesis in tooth enamel: ESR dating signals and  
carbonate contents. *J. Archaeol. Sci.* 22, 251–255.

- 1  
2 Rioult, M., 1963. Le Calcaire de Caen, dépôt de rivage du Bathonien normand. Bull. Société Linn.  
3 Normandie 3, 119–141.
- 4 Saint-Hilaire, G., 1831. Recherches sur de grands sauriens trouvés à l'état fossile aux confins  
5 maritimes de la Basse-Normandie, attribués d'abord au crocodile, puis déterminés sous les  
6 noms de Teleosaurus et Steneosaurus. Mém. Académie Sci. 12, 1–138.
- 7 Santos, G.M., Ferguson, J., Acaylar, K., Johnson, K.R., Griffin, S., Druffel, E., 2011.  $\Delta^{14}\text{C}$  and  $\delta^{13}\text{C}$   
8 of seawater DIC as tracers of coastal upwelling: a 5-year time series from Southern  
9 California. Radiocarbon 53, 669–677.
- 10 Seebacher, F., Grigg, G.C., Beard, L.A., 1999. Crocodiles as dinosaurs: behavioural  
11 thermoregulation in very large ectotherms leads to high and stable body temperatures. J.  
12 Exp. Biol. 202, 77–86.
- 13 Seymour, R.S., Bennett-Stamper, C.L., Johnston, S.D., Carrier, D.R., Grigg, G.C., 2004. Evidence  
14 for endothermic ancestors of crocodiles at the stem of archosaur evolution. Physiol.  
15 Biochem. Zool. 77, 1051–1067.
- 16 Shackleton, N.J., Kennett, J.P., 1975. Paleotemperature History of the Cenozoic and the Initiation of  
17 Antarctic Glaciation: Oxygen and Carbon Isotope Analyses in DSDP Sites 277, 279 and  
18 281. DSDP Initial Rep. - Deep Sea Drill. Proj. 29, 743–756.  
19 <https://doi.org/10.2973/dsdp.proc.29.117.1975>
- 20 Summers, A.P., 2005. Evolution: warm-hearted crocs. Nature 434, 833.  
21 <https://doi.org/10.1038/434833a>
- 22 Tarduno, J.A., Brinkman, D.B., Renne, P.R., Cottrell, R.D., Scher, H., Castillo, P., 1998. Evidence  
23 for extreme climatic warmth from Late Cretaceous Arctic vertebrates. Science 282, 2241–  
24 2244.
- 25 Tennant, J.P., Mannion, P.D., Upchurch, P., Sutton, M.D., Price, G.D., 2017. Biotic and  
26 environmental dynamics through the Late Jurassic–Early Cretaceous transition: evidence  
27 for protracted faunal and ecological turnover. Biol. Rev. 92, 776–814.  
28 <https://doi.org/10.1111/brv.12255>
- 29 Thorrold, S.R., Campana, S.E., Jones, C.M., Swart, P.K., 1997. Factors determining  $\delta^{13}\text{C}$  and  $\delta^{18}\text{O}$   
30 fractionation in aragonitic otoliths of marine fish. Geochim. Cosmochim. Acta 61, 2909–  
31 2919.
- 32 Trueman, C., Chenery, C., Eberth, D.A., Spiro, B., 2003. Diagenetic effects on the oxygen isotope  
33 composition of bones of dinosaurs and other vertebrates recovered from terrestrial and  
34 marine sediments. J. Geol. Soc. 160, 895–901.
- 35 Tütken, T., Vennemann, T.W., Pfretzschner, H.U., 2008. Early diagenesis of bone and tooth apatite  
36 in fluvial and marine settings: Constraints from combined oxygen isotope, nitrogen and REE  
37 analysis. Palaeogeogr. Palaeoclimatol. Palaeoecol. 266, 254–268.
- 38 van Hinsbergen, D.J., de Groot, L.V., van Schaik, S.J., Spakman, W., Bijl, P.K., Sluijs, A.,  
39 Langereis, C.G., Brinkhuis, H., 2015. A paleolatitude calculator for paleoclimate studies.  
40 PloS One 10, e0126946. <https://doi.org/10.1371/journal.pone.0126946>
- 41 Vennemann, T.W., Hegner, E., Cliff, G., Benz, G.W., 2001. Isotopic composition of recent shark  
42 teeth as a proxy for environmental conditions. Geochim. Cosmochim. Acta 65, 1583–1599.
- 43 Wilberg, E.W., Turner, A.H., Brochu, C.A., 2019. Evolutionary structure and timing of major  
44 habitat shifts in Crocodylomorpha. Sci. Rep. 9, 514. <https://doi.org/10.1038/s41598-018-36795-1>
- 45 Wolf, N., Newsome, S.D., Fogel, M.L., Del Rio, C.M., 2013. The relationship between drinking  
46 water and the hydrogen and oxygen stable isotope values of tissues in Japanese Quail  
47 (*Cortunix japonica*). The Auk 130, 323–330. <https://doi.org/10.1525/auk.2013.12075>
- 48 Young, M.T., Bell, M.A., De Andrade, M.B., Brusatte, S.L., 2011. Body size estimation and  
49 evolution in metriorhynchid crocodylomorphs: implications for species diversification and  
50 niche partitioning. Zool. J. Linn. Soc. 163, 1199–1216. <https://doi.org/10.1111/j.1096-3642.2011.00734.x>
- 51  
52  
53  
54  
55  
56  
57  
58  
59  
60

- 1  
2 Young, M.T., Brusatte, S.L., Beatty, B.L., De Andrade, M.B., Desojo, J.B., 2012. Tooth-on-tooth  
3 interlocking occlusion suggests macrophagy in the Mesozoic marine crocodylomorph  
4 *Dakosaurus*. *Anat. Rec. Adv. Integr. Anat. Evol. Biol.* 295, 1147–1158.  
5 <https://doi.org/10.1002/ar.22491>  
6  
7 Young, M.T., Brusatte, S.L., Ruta, M., de Andrade, M.B., 2010. The evolution of  
8 Metriorhynchoidea (Mesoeucrocodylia, Thalattosuchia): an integrated approach using  
9 geometric morphometrics, analysis of disparity, and biomechanics. *Zool. J. Linn. Soc.* 158,  
10 801–859. <https://doi.org/10.1111/j.1096-3642.2009.00571.x>  
11  
12 Young, M.T., de Andrade, M.B., Brusatte, S.L., Sakamoto, M., Liston, J., 2013. The oldest known  
13 metriorhynchid super-predator: a new genus and species from the Middle Jurassic of  
14 England, with implications for serration and mandibular evolution in predacious clades. *J.*  
15 *Syst. Palaeontol.* 11, 475–513. <https://doi.org/10.1080/14772019.2012.704948>  
16  
17 Young, M.T., de Andrade, M.B., Cornée, J.-J., Steel, L., Foffa, D., 2014a. Re-description of a  
18 putative Early Cretaceous “teleosaurid” from France, with implications for the survival of  
19 metriorhynchids and teleosaurids across the Jurassic-Cretaceous Boundary. *Ann. Paléontol.*  
20 100, 165–174. <https://doi.org/10.1016/j.annpal.2014.01.002>  
21  
22 Young, M.T., Rabi, M., Bell, M.A., Foffa, D., Steel, L., Sachs, S., Peyer, K., 2016. Big-headed  
23 marine crocodyliforms and why we must be cautious when using extant species as body  
24 length proxies for long-extinct relatives. *Palaeontol. Electron.* 19, 1–14.  
25 <https://doi.org/10.26879/648>  
26  
27 Young, M.T., Steel, L., Middleton, H., 2014b. Evidence of the metriorhynchid crocodylomorph  
28 genus *Geosaurus* in the Lower Kimmeridge Clay Formation (Late Jurassic) of England.  
29 *Hist. Biol.* 26, 551–555. <https://doi.org/10.1080/08912963.2013.801468>  
30  
31 Zazzo, A., Lécuyer, C., Mariotti, A., 2004a. Experimentally-controlled carbon and oxygen isotope  
32 exchange between bioapatites and water under inorganic and microbially-mediated  
33 conditions. *Geochim. Cosmochim. Acta* 68, 1–12.  
34  
35 Zazzo, A., Lécuyer, C., Sheppard, S.M.F., Grandjean, P., Mariotti, A., 2004b. Diagenesis and the  
36 reconstruction of paleoenvironments: A method to restore original  $\delta^{18}\text{O}$  values of carbonate  
37 and phosphate from fossil tooth enamel. *Geochim. Cosmochim. Acta* 68, 2245–2258.  
38 <https://doi.org/10.1016/j.gca.2003.11.009>  
39  
40  
41  
42  
43  
44  
45  
46  
47  
48  
49  
50  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60

## Figure Caption

**Figure 1:** Palaeogeography of Northwestern Europe during the Middle-Late Jurassic (modified from Pross et al. (2006). The red stars show the studied localities. 1: Peterboroug; 2: Smallmouth Sands; 3: Les Vaches Noires; 4: Les Lourdines excavation; 5: Cintheaux.

**Figure 2:** Oxygen isotope compositions of tooth phosphate reported against their corresponding oxygen isotope composition of structural carbonate, as well as published values of modern and fossil fish for comparison (Kolodny and Luz, 1991; Lécuyer et al., 2003b; Vennemann et al., 2001). The dashed line with a slope  $a=1$  illustrates the correlation between oxygen isotope composition of phosphate and carbonate.

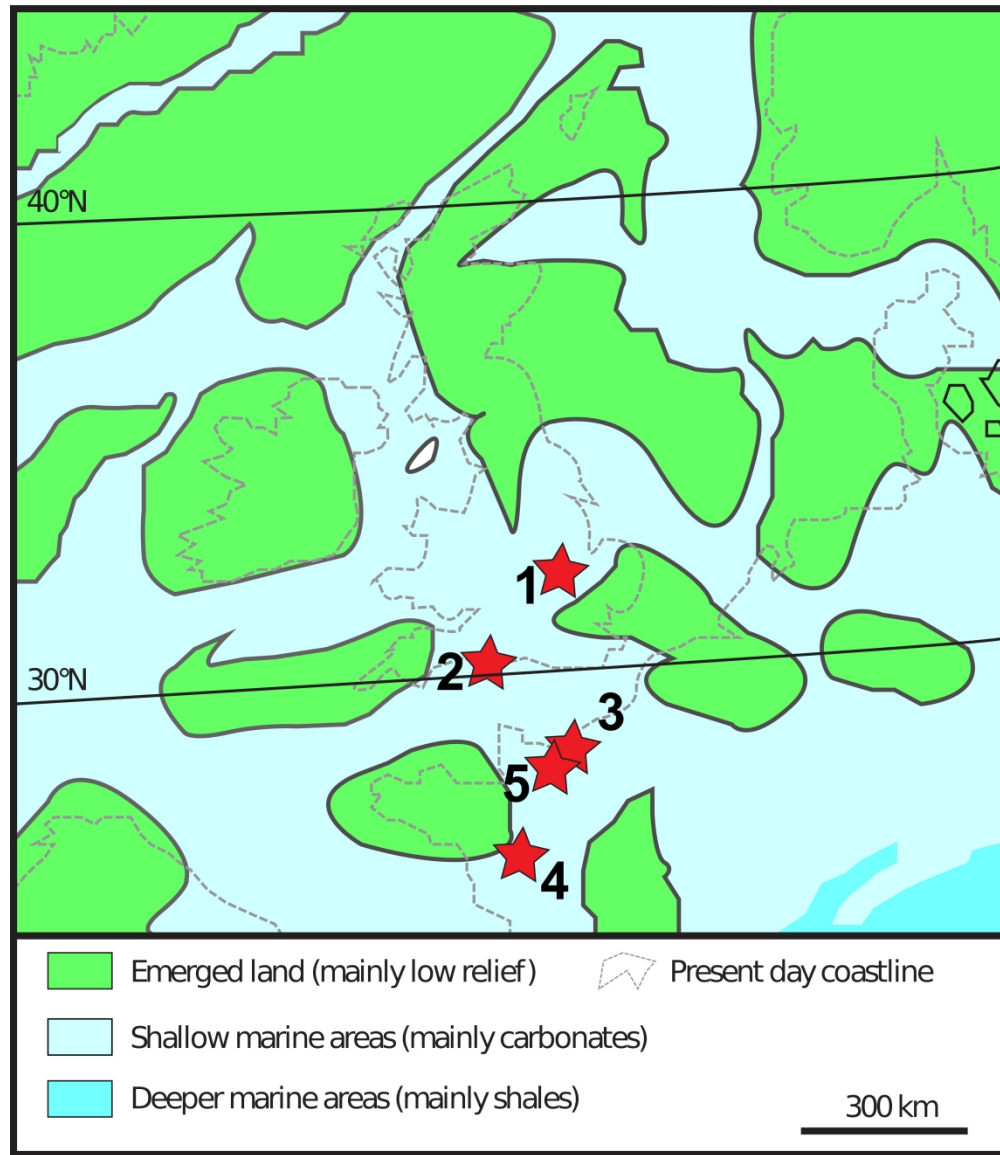
**Figure 3:** Carbon isotope compositions of apatite carbonates ( $\delta^{13}C_c$ ) from fish and marine reptile samples. For each locality, the sizable difference between fish and coexisting marine reptiles  $\delta^{13}C_c$  values is considered as evidence for primary preservation of the stable isotope compositions of studied specimens.

**Figure 4:** Model variation of the differences in the  $\delta^{18}O_p$  values of tooth phosphate between marine reptiles and fish against the variation of the  $\delta^{18}O_p$  values of fish teeth, assuming (1) an ectothermic and poikilothermic reptile [body water  $\delta^{18}O_{bw}$  values 2‰ enriched relative to a seawater value and body temperature (T) equal seawater temperature]; (2) an endothermic reptile with body temperature ranging from 35°C to 39°C and body water 2‰ enriched relative to a seawater value ranging from -1‰ to 0‰ (modified from (Bernard et al., 2010). For comparison, metriorhynchoids and teleosaurids values are reported, along with newly measured (black border) and published (grey border) ichthyosaurs and plesiosaurs values.

1  
2 **Figure 5:** Estimated body temperature of marine reptiles (left axis) and corresponding  $\delta^{18}\text{O}_p$  values  
3  
4 (right axis) are plotted against their environmental sea surface temperature estimated from fish  
5  
6  $\delta^{18}\text{O}_p$  values.  
7  
8  
9  
10  
11  
12  
13  
14  
15  
16  
17  
18  
19  
20  
21  
22  
23  
24  
25  
26  
27  
28  
29  
30  
31  
32  
33  
34  
35  
36  
37  
38  
39  
40  
41  
42  
43  
44  
45  
46  
47  
48  
49  
50  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60

For Review Only





43  
44  
45  
46  
47  
48  
49  
50  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60

Figure 1: Palaeogeography of Northwestern Europe during the Middle-Late Jurassic (modified from Pross et al. (2006)). The red stars show the studied localities. 1: Peterborough; 2: Smallmouth Sands; 3: Les Vaches Noires; 4: Les Lourdines excavation; 5: Cintheaux.

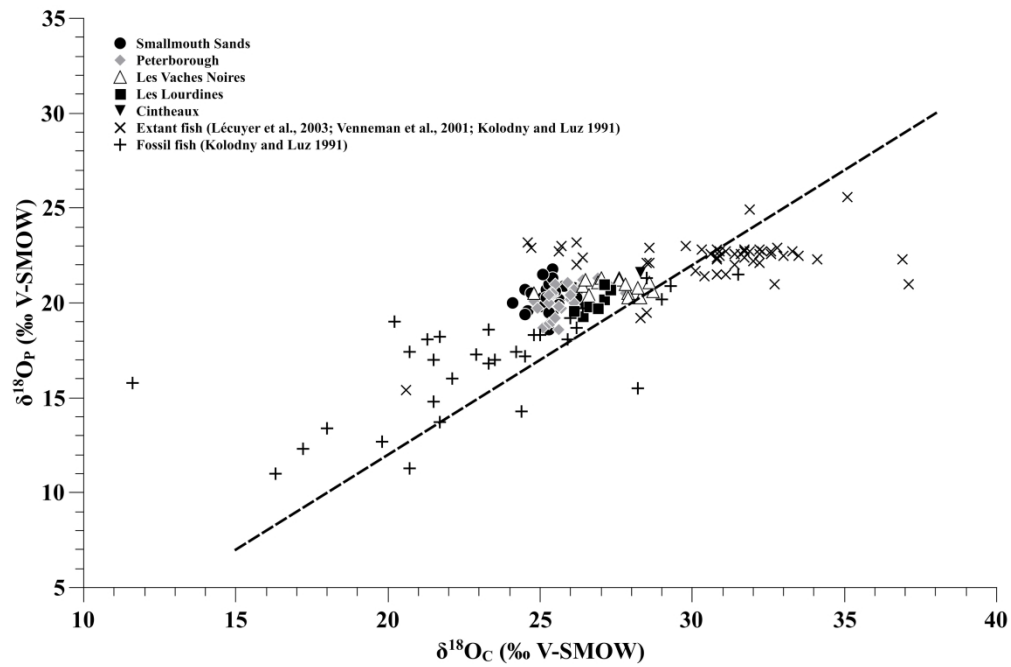


Figure 2: Oxygen isotope compositions of tooth phosphate reported against their corresponding oxygen isotope composition of structural carbonate, as well as published values of modern and fossil fish for comparison (Kolodny and Luz, 1991; Lécuyer et al., 2003b; Vennemann et al., 2001). The dashed line with a slope  $a=1$  illustrates the correlation between oxygen isotope composition of phosphate and carbonate.

297x196mm (300 x 300 DPI)

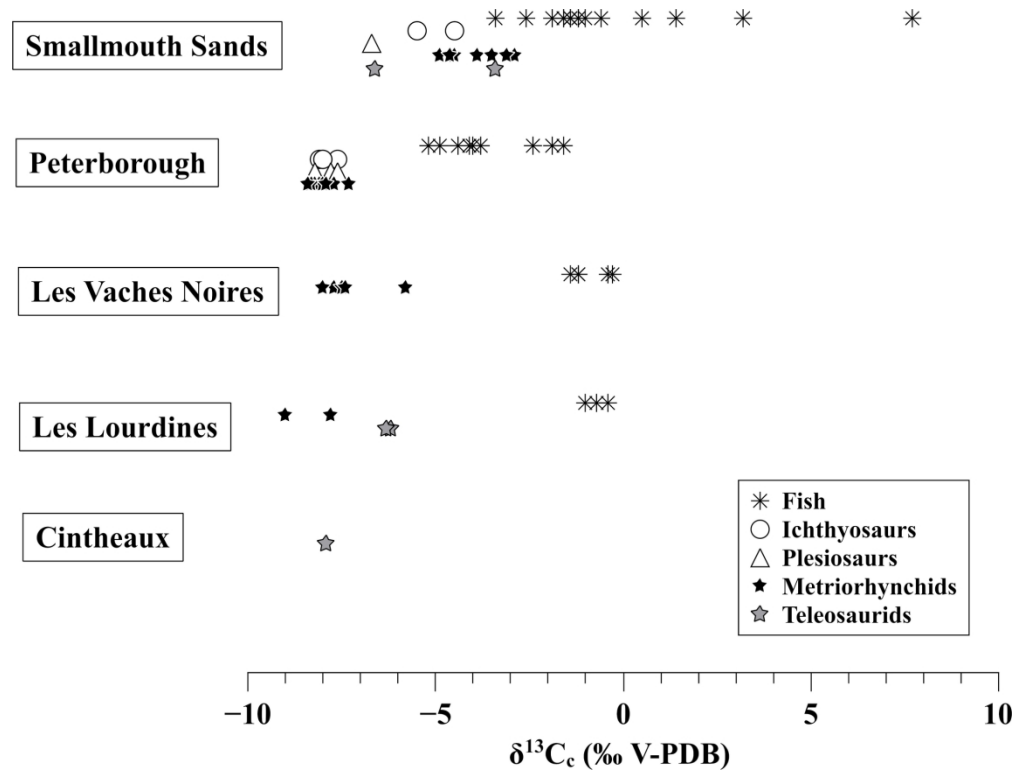


Figure 3: Carbon isotope compositions of apatite carbonates ( $\delta^{13}C_c$ ) from fish and marine reptile samples. For each locality, the sizable difference between fish and coexisting marine reptiles  $\delta^{13}C_c$  values is considered as evidence for primary preservation of the stable isotope compositions of studied specimens.

188x143mm (300 x 300 DPI)

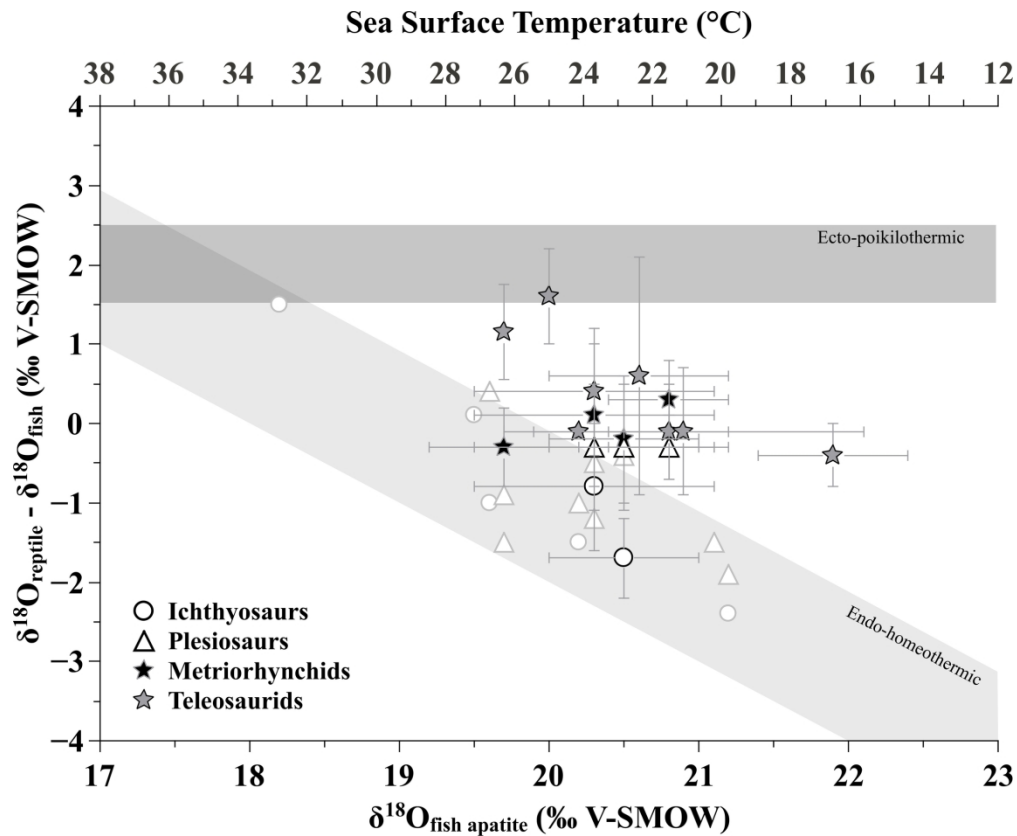


Figure 4: Model variation of the differences in the  $\delta^{18}\text{O}_p$  values of tooth phosphate between marine reptiles and fish against the variation of the  $\delta^{18}\text{O}_p$  values of fish teeth, assuming (1) an ectothermic and poikilothermic reptile [body water  $\delta^{18}\text{O}_{\text{bw}}$  values 2‰ enriched relative to a seawater value and body temperature (T) equal seawater temperature]; (2) an endothermic reptile with body temperature ranging from 35°C to 39°C and body water 2‰ enriched relative to a seawater value ranging from -1‰ to 0‰ (modified from (Bernard et al., 2010)). For comparison, metriorhynchoids and teleosaurids values are reported, along with newly measured (black border) and published (grey border) ichthyosaurs and plesiosaurs values.

192x158mm (300 x 300 DPI)

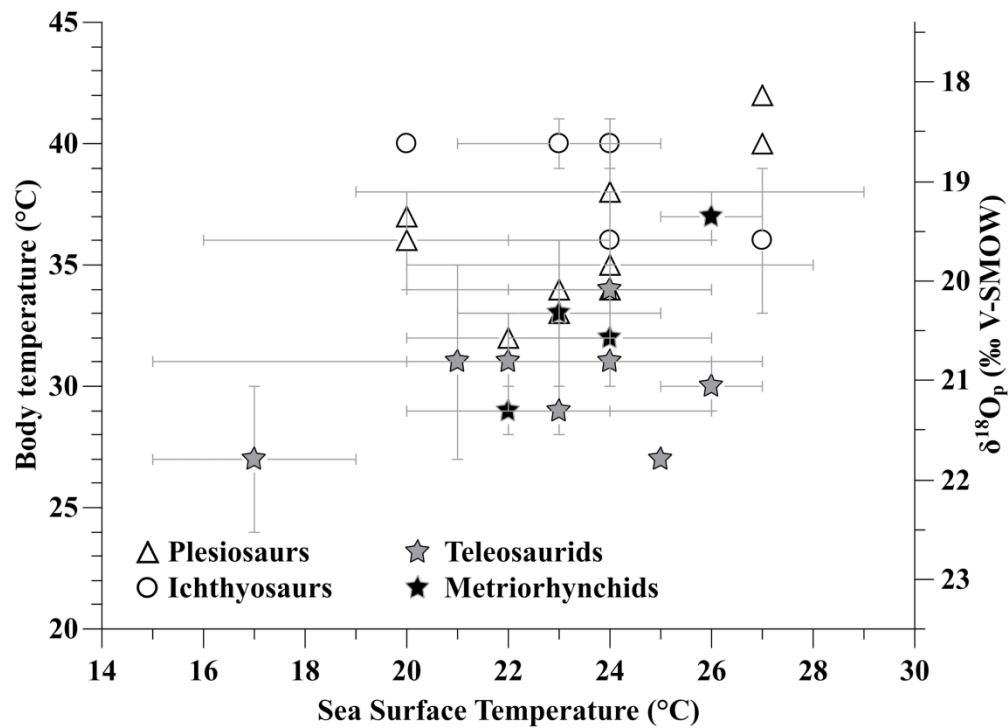


Figure 5: Estimated body temperature of marine reptiles (left axis) and corresponding  $\delta^{18}\text{O}_p$  values (right axis) are plotted against their environmental sea surface temperature estimated from fish  $\delta^{18}\text{O}_p$  values.

213x154mm (300 x 300 DPI)