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# Recent anthropogenic climate change exceeds the rate and magnitude of natural Holocene variability on the Balearic Islands

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#### 37 Highlights

• An 8000-year climatic reconstruction from the Western Mediterranean

• Forty-three years of instrumental climatic data from Palma de Mallorca

40 • A sharp climate contrast between reconstructed and instrumental periods

41 • Present changes are rapid in contrast to the Holocene climate history of Mallorca

42

#### 43 Abstract

Industrial-era warming and aridification have underlined the importance of past climate 44 reconstructions in contextualizing present-day anomalies from a long-term perspective. While 45 46 the issue of climate change is global, studies have long stressed the vulnerability of the Mediterranean basin, especially with regard to its islands with likely acute environmental and 47 socioeconomic impacts. This study uses paleoecological data from Mallorca to quantify climate 48 49 changes during the last eight millennia. We compared and contrasted past ecosystem dynamics with recent instrumental climate data to characterize the emergence of industrial-era warming. 50 We show that anthropogenic warming has transgressed both the rate and magnitude of natural 51 variability. At the scale of the last 8000 years, this study highlights the uniqueness of Mallorca's 52 present climate and emphasizes the rapidity of current changes compared to the climate history 53 of the island. The present-day mean annual temperature appears  $3.3 \pm 1.2$  °C warmer than the 54 reconstructed value for the last eight millennia, with a 33% drop in precipitation between the 55 last 8000 years and the industrial era. This long-term view underscores human impacts on 56 climate in Mallorca, the fourth most populated island of the Mediterranean. It suggests that the 57 ongoing warming and aridification trends will be a significant challenge to Mediterranean 58 islands, at both present and future timescales. 59

60

#### 62 Keywords

Ecosystem dynamics, Pollen-based climate reconstruction, climate change, Industrial-era
 warming, Aridification, island, Holocene, Mallorca, Mediterranean

65

#### 66 **1. Introduction**

The emergence of the industrial economy has triggered climate anomalies since the mid-late 67 19<sup>th</sup> century (Abram et al., 2016; Gleckler et al., 2016). Although the resulting environmental 68 stresses have affected many regions of the world, the International Panel on Climate Change 69 (IPCC, 2001; IPCC, 2014; IPCC 2018) deem the Mediterranean to be "particularly sensitive" 70 71 to the effects of global warming (Fischer and Schär, 2010; Homar et al., 2010; López Mayol et al., 2017; Cramer et al., 2018). For the period 2081-2100, the IPCC's climate models project a 72 2.6-4.8°C rise in average surface temperatures (relative to 1986-2005), accompanied by a 73 pronounced shift in rainfall regimes with anomalies of -4% to -27% in the Mediterranean (from 74 Representative Concentration Pathways 4.5 to 8.5; IPCC, 2014). Such changes will generate 75 exceptional stress on already limited water resources (Guiot and Cramer, 2016), particularly in 76 regions where anthropogenic demands already exceed supply. The projected warming of the 77 Mediterranean is clearly attested on the Balearic Islands, which have witnessed an 78 intensification of temperature anomalies during the last 40 years (~0.6°C per decade) and a 79 clear shortening of the cold season (Cuxart and Guijarro, 2010; Homar et al., 2010; López 80 Mayol et al., 2017). According to the IPCC (2001), a rise in temperature (1.5 to 3.6 °C) and a 81 drop in rainfall amounts (10% to 20%) are forecasted for 2080-2100 on the Balearics, with an 82 increased probability of drought (Kerr, 2005) and extreme weather events (Candela et al., 83 2009). Even if the ocean's thermal inertia mitigates the temperature of the Mediterranean sea 84 surface, delaying the accumulation of heat in coastal zones, notable increases in sea surface 85

temperatures have already been recorded in the Balearic Sea (+1.1 °C during the last 40 years;
Marbà and Duarte, 2010; Jordà et al., 2012; Adloff et al., 2015).

Through the prism of Mallorca, this study aims to answer two main questions. First, to what 88 extent does the industrial-era warming and aridification of Mediterranean islands deviate 89 markedly from long-term Holocene trends? Second, how can an 8000-year palaeoecological 90 archive from the Albufera Natural Park serve as a basis to quantify the current climate 91 92 anomalies (1975-2018 CE)? As climate models project widespread anomalies across the Mediterranean Basin by the end of the 21<sup>st</sup> century (Cook et al., 2015), and because warming 93 and aridification can significantly impact the environment, economy, society (Greve et al., 94 95 2014) and human health (Gale et al., 2010), quantifying the present-day climate trend in relation to the past is critical in improving climate-change projections. It is also important for 96 developing robust hazard assessments and adaptation strategies. 97

98

#### 99 2. Albufera Natural Park

The swamp/wetland of Albufera (elevation 1 m a.s.l.; surface of ~25 km<sup>2</sup>; Fig. 1) is located in 100 the eastern part of a Plio-Quaternary depression, between the northern and southern ridges of 101 the island (dated from the Mesozoic to the Upper Miocene). The swamp is relict of a palaeo-102 103 lagoon formed following the Holocene marine transgression (Fornós et al., 1996; Goy et al., 1997; Giaime et al., 2017). This swamp area is fed by the Sant Miguel and Muro streams and 104 by groundwater inputs. We selected this protected natural reserve to limit any potential 105 anthropogenic influences on the environmental-climate reconstructions. Today, the swamp is 106 mostly populated by Phragmites australis (Cav. Trin. ex Steud.), Cladium mariscus (L. Pohl) 107 and Typha latifolia (L.). Alisma plantago-aquatica (L.) and Sparganium erectum (L.) are also 108 common around the edges, between the water and the marsh. In the water column, Potamogeton 109 pectinatus (L.), Potamogeton crispus (L.), Riccia (L.), Myriophyllum spicatum (L.), 110

Ceratophyllum submersum (L.) and Zannichellia pedunculata (Rchb. in Mössler) occur. 111 Around the park and closer to the dunes, the dominant plants are *Pinus halepensis* (Mill.), 112 Pistacia lentiscus (L.), Rosmarinus officinalis (L.), Phillyrea angustifolia (L.), Cistus 113 salviifolius (L.), Smilax aspera (L.), Asparagus acutifolius (L.), Halimium halimifolium (L.) 114 and Brachypodium retusum (L.). On the seaward flank of the swamp, Juniperus oxycedrus 115 subsp. macrocarpa (Sm. Ball) is the dominant tree species, with a cortege of Cakile maritima 116 (Scop.), Eryngium maritimum (L.), Pancratium maritimum (L.), Matthiola sinuata (L.) and 117 Helichrysum stoechas (L. Moench). The nomenclature derives from the Flora Iberica 118 (http://www.floraiberica.es/). 119

120

#### 121 **3. Materials and methods**

To answer the questions raised in the introduction, we used the instrumental climate dataset from Mallorca's airport weather station and produced a pollen-based climate reconstruction from a core drilled in the Albufera Natural Park (NE Mallorca). We used a <sup>14</sup>C chronology to calculate an age model for the core. We further tested the reliability of our climate reconstruction by estimating the role of anthropogenic impacts in modulating ecosystem dynamics. All data are available in tabulated form in the supplementary information.

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#### 129 *3.1. Present-day climate data*

The instrumental climate dataset was provided by Mallorca's airport weather station (Palma Aeropuerto, 39°33'39"N, 02°44'12"E, 3 m. a.s.l.). The dataset constitutes the Balearic Islands' longest temperature and precipitation records and covers the periods 1975-2018 (temperatures) and 1972-2018 (precipitation). The records from Palma Puerto (39°33'12"N, 02°37'31"E; 1 m a.s.l.; period 1978-2018) and Sa Pobla (39°44'57"N, 03°01'00"E; 27 m a.s.l.; period 2012-2018) were used as independent controls (for extreme values). Previous datasets were already available in the literature (e.g. Cuxart and Guijarro, 2010; Homar et al., 2010; Jansà et al., 2017;
López Mayol et al., 2017).

138

139 *3.2. Core chronology and stratigraphy* 

We drilled an 8-m long core (39°47'44.43"N, 3°05'31.47"E, 1 m a.s.l.; termed POL12) at the 140 heart of the Albufera Natural Park using a hydraulic rotary corer (Fig. 2). The sedimentary 141 deposits mainly consist of clay with inclusions of silt and sand in the upper part. The chronology 142 of the core is based on ten <sup>14</sup>C dates (Beta Analytic and Poznan Laboratory; see Table S1) 143 calibrated to 1 and 2 sigmas ( $\sigma$ ) using Calib Rev 7.0.4 (Fig. 2). The oldest level (790 cm depth) 144 is dated to 7080  $\pm$  40 <sup>14</sup>C year BP, calibrated to 7905 year BP (7980-7830 calibrated  $2\sigma$  BP, 145 7955-7865 calibrated  $1\sigma$  BP). The age of each sample was calculated by interpolating between 146 two adjacent <sup>14</sup>C dates. Our record thus covers ~8000 years of northeastern Mallorcan 147 148 environmental and climate history.

149

#### 150 *3.3. Palaeoecological data*

We prepared a total of 120 samples for pollen analysis using standard procedures for clay 151 deposits (Faegri and Iversen, 1989). Pollen grains were counted under x400 and x1000 152 magnification using an Olympus Bh-2 BHS microscope. Pollen percentages are based on the 153 terrestrial pollen sum, excluding local helophytes (Cyperaceae and Typha), macrophytes and 154 spores of non-vascular cryptogams. Percentages of helophyte-macrophyte taxa were calculated 155 using a sum that includes local freshwater plants and terrestrial vegetation. The mean pollen 156 sum was  $379 \pm 53$  pollen grains, with a minimum of 284 pollen grains. The median value was 157 374 pollen grains, with a 25th percentile of 336 pollen grains and a 75th percentile of 409 pollen 158 grains. The mean number of taxa was  $28 \pm 3$  taxa, with a minimum of 21 taxa. The median 159 value was 27 taxa, with a 25th percentile of 26 taxa, and a 75th percentile of 29 taxa. The 160

calculation of pollen concentrations follows the formula developed by Cour (1974), using 1 161  $cm^3$  of sediment. The mean concentration was  $2728 \pm 382$  pollen grains per  $cm^3$  with a median 162 value of 2692 pollen grains per cm<sup>3</sup>, a 25th percentile of 2419 pollen grains per cm<sup>3</sup>, and a 75th 163 percentile of 2948 pollen grains per cm<sup>3</sup>. Atlases (Reille, 1992, 1995, 1998), books (e.g. Faegri 164 and Iversen, 1989; Moore et al., 1991; Punt and Blackmore, 1991), articles (e.g. Renault-165 Miskovsky et al., 1976; Colombo et al., 1983; Gortemaker, 1986), as well as local reference 166 slides were used for pollen identification. A plate with pollen photos is provided in the 167 supplementary information (see Fig. S1). The fire history of the coastal plain was reconstructed 168 by counting charcoal particles (50-200 µm) on the pollen slides. The 50 µm size criterion was 169 170 chosen to avoid confusion between microscopic charcoal fragments and opaque minerals, which are typically  $< 50 \mu m$  (Parshall and Foster, 2002; Pederson et al., 2005). We followed 171 the 200-300 items (sum of charcoal particles) defined by Finsinger and Tinner (2005). A 172 charcoal concentration curve (fragments per cm<sup>-3</sup>) was drawn on a linear age-scale (Marriner et 173 al., 2019). The mean concentration was  $77 \pm 19$  fragments including the standard error, with a 174 maximum value of 1550 fragments and minimum of 0. The pollen data are detailed and 175 presented as a pollen diagram (Fig. 3), with a focus on cereals, olive trees and fire activity (Fig. 176 4). All of these elements can be consulted in the supplementary information. We curated all the 177 178 pollen slides at the Université Toulouse 3.

179

#### 180 *3.4. Statistical analyses*

We applied a homogeneity test to fire activity, cereals and olive trees to detect potential shifts in their long-term dynamics (Fig. 4). The homogeneity test is based on a null hypothesis that a time series is homogenous. Possible alternative hypotheses such as change in distribution, changes in average or presence of trend can be detected and subdivided. Each change is highlighted and its average is indicated by *mu*. Before any modelling, a CABFAC analysis (Fig.

5) was used to test the potential weight of anthropogenic factors in mediating the ecological 186 signal (e.g. Hammer and Harper, 2006; Kaniewski et al., 2018). The CABFAC factor analysis 187 (Q-mode factor analysis) implements the classic method of factor analysis and environmental 188 regression (CABFAC and REGRESS). We regressed selected environmental data on the 189 CABFAC factors using the second-order (parabolic) method, with cross terms. The 190 environmental regression model (RMA) reports the observed values against the values 191 reconstructed from the factors (Kaniewski et al., 2018). The selected environmental data for 192 this study (wet/dry conditions, warm/cool conditions, anthropogenic activities, and fire) cover 193 a wide range of natural and anthropogenic pressures. The consistency of each reconstructed 194 model is indicated by the coefficient of determination R<sup>2</sup>. The R<sup>2</sup> values were ordered from 195 highest to lowest, and used as a hierarchy of pressures on ecosystems. 196

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#### 198 *3.5. Reconstructions of past climate variables*

We did not consider local helophytes, macrophytes and spores of non-vascular cryptogams in the pollen-based climate reconstruction. The cultivated species (primary anthropogenic indicators: Poaceae cerealia) were also excluded from the pollen-based climate reconstruction in addition to the secondary anthropogenic indicators (e.g. *Plantago*, *Rumex*).

203 Pollen assemblages were used to reconstruct the palaeoclimate history of Mallorca. Our statistical approach presents two key caveats: (1) a fossil pollen assemblage may not have a 204 modern analog sensu stricto; and (2) the climate range of a species/genus can shift in time 205 206 (Cheddadi et al., 2016). Our new method is based on the assumption that the co-occurrence of plant taxa in an assemblage may only occur under a common climate range or an overlap of the 207 climatic range of plant taxa comprising a given assemblage. Based on this assumption, one 208 should not expect two plant species to occur in the same assemblage if their climate ranges do 209 not intersect. We assigned the fossil pollen taxa to modern species, and subsequently calculated 210

a weighted median value and its standard deviations of the climate range encompassed by all
assigned modern species. Such an approach has been used to infer past climate variables using
insects (Elias, 1997), plant fossil remains (Pross et al., 2000), mollusks (Moine et al., 2002),
and ostracods (Horne, 2007) or to derive climate probability density functions (Kühl et al.,
2002) from fossil pollen data.

Our method (Cheddadi et al., 2017) also performs a "leave-one-out calculation" to accommodate the temporal variability of pollen taxa in the fossil record. For each fossil sample, a taxon is removed and a weighted median (using the pollen percentages as a weight) is calculated using all of the remaining taxa. For each sample, we calculated the weighted median as many times as there were taxa in each fossil level. The final figure is the median value of all iterations. In addition, the standard deviations correspond to the median value of all the standard deviations.

This additional leave-one-out approach reduces the effects of either large variations in the fossil record or over- or under-representation of some pollen species. The code was written using the software R and exploits the available libraries (Akima and Gebhardt, 2016), RMySQL (Ooms et al., 2016), and the statistical functions which are part of R. The modern plant database is based on georeferenced species from Flora Europaea (Jalas and Suominen, 1973, 1979, 1980) and from GBIF (Heughebaert, 2019). We obtained present climate variables from the WORLDCLIM database (Hijmans et al., 2005).

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231 *3.6. Data availability* 

All data needed to evaluate the conclusions reported in the paper are present in the article and/orthe supplementary information (excel files).

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#### **4. Results**

#### 237 *4.1. The core*

The bottom of the core comprises clay deposits (from 800 cm to 320 cm depth) consistent with a low-energy environment (Fig. 2). We identified a sand lens between 320 and 300 cm depth, which probably results from aeolian transport into the lagoon (Giaime et al., 2017). The upper part (from 300 cm to 170 cm depth) consists of silt associated with freshwater inputs into the area. Some sand particles from the shallowing lagoon were mixed with the silt deposits during this period. The human-modified deposits correspond to the last 170 cm.

244

#### 245 4.2. Pollen analysis

The <sup>14</sup>C dates (Fig. 2) frame the chronology of core POL12. The pollen data allowed us to reconstruct ecosystem dynamics during the past 8000 years at Albufera Natural Park (Fig. 3). We compared and contrasted our data with earlier studies by Burjachs et al. (1994, 2017) in the same area, and with cores drilled on Minorca and Ibiza (Yll et al., 1997; Pérez-Obiol and Sadori, 2007; Pérez-Obiol et al., 2011; Carrión et al., 2012; Burjachs et al., 2017).

The date proposed for the introduction of agricultural practices and human-induced fires in the Albufera Natural Park is ~3600 BP (Fig. 4). According to the 2-sigma calibration, the full range is framed between 3650 and 3550 BP ( $3600 \pm 50$  BP). As stated by Burjachs et al. (2017), cereals are relatively scarce in the area. Data from the core POL12 depict a similar trend, with a maximum value < 10%. In the core POL12, the earliest record of human impacts on the environment thus occurred after the shift from a mesophilous community to a maquis scrub (~5900 BP; Fig. 3).

The abrupt decline of *Juniperus* at ~5900 BP clearly illustrates this phenomenon (Fig. 3). The drop seems to result from environmental pressures because no strong anthropogenic footprints (fire activity or agriculture) were detected (Fig. 4). This decline is associated with greater aridity in the Mediterranean at ~5900 BP (Jalut et al., 2000; Fletcher et al., 2010). As mentioned by
Burjachs et al. (2017), this climate event (along with others at 8.2 and 4.3 ka BP) directly
influenced the composition and evolution of vegetation on the Balearic Islands. According to
Pérez-Obiol and Sadori (2007), the important ecosystems changes ~5000-4000 BP seem to have
been amplified, but not caused, by human action. Most south-western European pollen records
show a general change from wetter to drier climatic conditions (Jalut et al. 2000; Magny et al.
2002). The core POL12 meshes perfectly with previously published data.

A homogeneity test was applied to three curves (fire activities, cereals and olive trees) to detect 268 potential shifts in their long-term dynamics (averages are indicated by mu; Fig. 4). Results 269 270 showed that cereals and fire activity are well correlated, whereas olive trees are mediated by a different dynamic. To further establish whether olive trees were an integral part of early 271 diversified agriculture, we calculated a cross-correlation ( $P_{value} = 0.05$ ). We considered positive 272 correlation coefficients, focusing on the Lag<sub>0</sub> value (with +0.50 as the significant threshold). 273 The test, applied to Poaceae cerealia versus Olea europaea, showed no correlation between the 274 two (Lag<sub>0</sub> = 0.008, P<sub>value</sub> < 0.001). A second test, based on fire activity versus Olea europaea, 275 also demonstrated no correlation (Lag<sub>0</sub> = -0.07, P<sub>value</sub> < 0.001). According to Pérez-Obiol and 276 Sadori (2007), from ~5000 to ~4000 BP, some taxa (Buxus and Corylus) were replaced by 277 278 others such as Olea europaea, a process that reflects an ecological response to global climatic trends. A similar phenomenon was clearly recorded at Albufera (Fig. 3). At Cala'n Porter 279 (Minorca), high occurrences of olive trees have been recorded since ~6500 BP. At Algendar 280 (Minorca) and Prat de Vila (Ibiza), high values are recorded since ~5500 BP (Burjachs et al., 281 2017). It appears that olive-tree dynamics are independent of human influence in the Balearics. 282 Nonetheless, some chronological discrepancies exist between the two studies in the Albufera 283 Natural Park. These differences may result from the absence of dates between the <sup>14</sup>C ages 6270 284  $\pm$  70 and 2440  $\pm$  60 (respectively 7330-6985 and 2625-2350 calibrated 2 $\sigma$  BP) in the sequence 285

published by Burjachs et al. (1994, 2017). Because the quality of an age model is dependent on the number of dated levels, more <sup>14</sup>C dates for the core POL12 have probably reduced the uncertainties for the older part of the sequence (Fig. 2).

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#### 290 *4.3.* Anthropogenic influence on the signal

Our pollen diagram from the Albufera Natural Park shows that, while human activities may 291 have partially shaped the area since ~3600 BP (attested through an increase in cereal cultivation 292 and fire activity), the CABFAC analyses indicate that the two main factors controlling the 293 ecological signal are humidity ( $R^2 = 0.73$ ,  $P_{value} < 0.001$ ) and temperatures ( $R^2 = 0.65$ ,  $P_{value} < 0.001$ ) 294 0.001). Human impacts ( $R^2 = 0.33$ , Pvalue < 0.001) and fire activities ( $R^2 = 0.31$ , Pvalue < 295 0.001) do not meet the significance threshold of 0.5 (Fig. 5). A further control is provided by 296 the evolution of olive trees versus cereals (Fig. 4). The highest scores of Olea europaea are 297 recorded long before the emergence of human impacts on the environment and also long before 298 the oldest known date of human presence on the island (Burjachs et al., 2017). Olive trees, 299 therefore, are seemingly not a corollary of anthropogenic activities but rather of warmer 300 conditions. The rank of pressures deriving from the CABFAC analyses clearly suggests the 301 dominant role of environmental/climatic agents in mediating the island's palaeoecological 302 dynamics (wet/dry and cool/warm). 303

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#### 305 4.4. Mallorca's mean seasonal temperatures

We first converted past and present-day datasets into mean seasonal temperatures to probe when climate change was most prominent (Fig. 6). According to the instrumental and reconstructed datasets, the sharpest "warming" contrasts between the last eight millennia and the period 1975-2018 occur in winter and autumn, with a respective difference in mean temperature of  $4.8 \pm$  310  $1.8^{\circ}$ C and  $3.9 \pm 1.5^{\circ}$ C. Spring and summer seasons display less pronounced shifts of  $2.1 \pm 1.3^{\circ}$ C and  $2.3 \pm 1.4^{\circ}$ C.

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#### 313 4.5. Mallorca's mean annual temperatures

We then calculated mean annual temperatures (MAT) for the instrumental and reconstructed 314 datasets (Fig. 7A). The present-day MAT is  $16.6 \pm 0.8^{\circ}$ C (Fig. 7A), with a mean high (T<sub>max</sub>) of 315  $22.5 \pm 0.7^{\circ}$ C and a mean low (T<sub>min</sub>) of  $10.6 \pm 0.9^{\circ}$ C. The warming trend is 0.55°C per decade 316 for the period 1975-2018. A previous study based on the whole Balearic archipelago, for the 317 period 1977-2006, showed that T<sub>min</sub> and T<sub>max</sub> are respectively 12.78°C and 21.76°C, with an 318 319 estimated warming trend of 0.5°C per decade (Homar et al., 2010). The two warming estimates are similar, with a minor difference of 0.05°C. In comparison, warming on the island of Tenerife 320 (Canary archipelago, North Atlantic Ocean) is much lower, with an average of  $0.09 \pm 0.04$  °C 321 per decade since 1944 (Martín et al., 2012). Focusing on the last eight millennia, the MAT is 322  $13.3 \pm 1.5$  °C, close to the present-day T<sub>min</sub> of the Balearic Islands (difference of ~0.5 °C; Homar 323 et al., 2010), but with significant temporal variations. The warmest period ( $15.6 \pm 0.4^{\circ}C$ , ~4000-324 3000 BP) is counterbalanced by the coolest phase ( $12.5 \pm 1.1^{\circ}$ C, ~8000-7000 BP), which yields 325 a thermal adjustment that has not been observed since the 20th century. A comparison between 326 327 present-day and past scores shows that only the ~4000-3000 BP values are similar to those recorded during the 1970s (MAT of 15.6  $\pm$  0.2; Fig. 7A). Since the 1980s (MAT of 16.3  $\pm$ 328 0.5°C), according to our pollen-based model, the measured temperatures seem to have no 329 analogue in the Holocene climate history of Mallorca. 330

This finding is supported by the MAT for the last 400 years that clearly shows a major ecological break between the pre-industrial and industrial eras (Fig. 7B). The shift dates back to  $\sim 1900 \pm 30$  (according to our age-model) but changes have been particularly acute since the 1980s (instrumental data). It has been suggested that while the onset of the industrial-era warming can be dated to the 19<sup>th</sup> century (~1830; Abram et al., 2016), the effects were more
acute during the second half of the 20<sup>th</sup> and 21<sup>st</sup> centuries (Brohan et al., 2006; IPCC, 2007;
Abram et al., 2016). In the Balearic Islands, scholars have demonstrated that climate-change
anomalies have mainly increased since the 1980s (Homar et al., 2010; López Mayol et al.,
2017). Our dataset (reconstructed and instrumental data) suggest a transition between the preand post-1980s.

341

#### 342 4.6. Temperature anomalies

We subsequently calculated our reconstructed past climate data as anomalies compared to 343 344 present-day conditions (Fig. 7C), using the period 1900-1980 (mean:  $15 \pm 0.6^{\circ}$ C) as a baseline. Two phases, centered on ~5800-5600 BP and ~4200-3000 BP, were warmer than the first period 345 of industrial-era warming. By contrast, when the second period of instrumental data (1981-346 347 2018; mean:  $16.7 \pm 0.6^{\circ}$ C) is used as a baseline, all reconstructed temperatures from the preindustrial era were lower (i.e. cooler). Only two peaks at  $\sim$ 5665 ± 40 BP (with a range from -348 0.66 to 0.26°C) and ~4065  $\pm$  30 BP (with a range from -0.39 to 0.31°C) are close to the 1981-349 2018 baseline (Fig. 8A). If the mean of the whole instrumental dataset is used as a baseline (Fig. 350 8B), the results illustrate the rapidity of the current changes compared to the climate history of 351 Mallorca. 352

The warming pattern that emerges from our study suggests that the instrumental MAT is  $3.3 \pm 1.2^{\circ}$ C warmer than the reconstructed MAT (for the last eight millennia). Furthermore, the reconstructed MAT is closer to the present-day T<sub>min</sub> (difference of  $2.7 \pm 1.2^{\circ}$ C). With a mean warming rate of  $0.55^{\circ}$ C per decade, the present-day T<sub>min</sub> may become equivalent to the pre-/early industrial-era MAT in about five decades.

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- 359

#### 360 4.7. Mallorca's precipitation

We subsequently reconstructed the dynamics of the rainfall regime during the last 8000 years 361 because variations in humidity were identified as the main forcing agent by the CABFAC 362 analysis (Fig. 5). The MedCLIVAR consortium, and other scholars, have long insisted on the 363 aridification of the Mediterranean area, with a likely increase in heat waves and droughts in the 364 coming decades (Homar et al., 2010; Lionello et al., 2012; Luterbacher et al., 2012; López 365 Mayol et al., 2017). In effect, from 1900 to 2019, droughts and heat waves killed more than 366 62,000 people in the Mediterranean and affected a further > 9 million with total damages 367 exceeding US\$ 29 billion (EM-DAT, 2020). 368

369 We compared the trends highlighted in the reconstructed and instrumental data to assess whether major changes are detectable (Fig. 9). The two average values showed a significant 370 shift, with a 33% drop in precipitation between the last 8000 years and the industrial era. 371 Focusing on the instrumental period, the driest years were recorded in 1983 (258.8 mm), 1995 372 (265.3 mm), 1999 (227.3 mm) and 2000 (255.9 mm). During the past 8,000 years, the driest 373 millennium was ~4000-3000 BP (Fig. 9), with the lowest score recorded at ~3820  $\pm$  30 BP 374  $(579.7 \pm 11.8 \text{ mm})$ . Although important variations were recorded during the Holocene, the 375 highly variable annual precipitation patterns of the industrial era appear to be lower than 376 377 reconstructed values for the last 8000 years.

378

#### 379 *4.8. Holocene climate dynamics*

We finally compared our long-term Holocene reconstruction with other proxy records (Fig. 10).
The temperatures and precipitation data mimic variations in solar activity (Solanski et al., 2005;
Vonmoos et al., 2006; Steinhilber et al., 2012), particularly the warm and dry long interval that
started ~5700 years ago and ended ~2400 years ago (Fig. 10).

#### 385 **5. Discussion**

This paper reports a new climate reconstruction for the Balearics and underlines the fragility and vulnerability of these islands in the face of ongoing industrial-era warming. The long-term temperature and precipitation trend for Mallorca, based on both reconstructed and instrumental data, clearly depicts a marked offset between the present-day situation and the last 8000 years (Figs 6-9).

391

#### 392 *5.1. Caveats*

Nonetheless, some caveats follow. A first limitation to the study results from missing data for 393 394 a 35-yr interval spanning 1940-1975 (Fig. 6) or 1940-1972 (Fig. 9) due to a gap in the instrumental records. A composite record, based on several Balearic weather stations, was not 395 a viable option as no station has a complete time series for this interval. A second limitation 396 relates to the <sup>14</sup>C chronology, which is based on calibrations with varying chronological ranges. 397 Although our reconstructions rely on these chronological data to provide a climate framework, 398 the furnished climate signal remains unchanged. A third limitation relates to the influence of 399 human activities on ecosystem dynamics. As suggested by the CABFAC analyses, human 400 401 impacts are not the main pressures on ecosystem dynamics (for the pre-industrial era) but their 402 influence is evidently recorded and must be accommodated in our reconstruction, because it is impossible to clearly separate natural forcing agents from anthropogenic activities. A final 403 limitation concerns the number of endemic species in Mallorca and the potential non-analogue 404 ecology when the present-day components are compared to the pre-anthropogenic era. Despite 405 these limitations, our reconstructions are relevant because we focus on long-term trends, 406 limiting over-interpretations resulting from a focus on particular events in the past. In effect, 407 several similar climate reconstructions based on pollen (with the same limitations) have already 408 been published and are relevant to ongoing global warming debates (e.g. Feurdean et al., 2008; 409

Guiot and Kaniewski, 2015; Mauri et al., 2015; Guiot and Cramer, 2016; Kaniewski et al.,
2019).

412

#### 413 5.2. Local environmental significance

Our data provide a long-term perspective on current climate change on the Balearic Islands, as 414 evoked by several authors (Cuxart and Guijarro, 2010; Homar et al., 2010; Jansà et al., 2017; 415 López Mayol et al., 2017). The rate and the magnitude of observed changes during the industrial 416 era are significant when compared to the Holocene climate history of Mallorca (Figs 6-9). 417 Moreover, the CABFAC analysis (Fig. 5) underscores that human activity was not the main 418 419 pressure behind observed changes in the northeastern part of the island during the pre-industrial period. This argument was previously raised by Burjachs et al. (1994; 2017) for the same area. 420 Our study also stresses that, during the pre-industrial period, warmer and drier conditions 421 developed from 5700 to 2400 BP (Figs 7 and 10). The same chronological boundary of 5700 422 BP was indicated by Bar-Matthews and Ayalon (2011), with a major shift in aridity in the 423 Levant, considered by the authors as the driest event of the mid-Holocene. Anderson et al. 424 (2011) also stressed this boundary for the Sierra Nevada (Laguna de Río Seco, southern Spain), 425 where drier conditions commenced regionally after 5700 BP, marked by the onset of 426 427 aridification of the flora of southern Spain. Our precipitation reconstruction tightly meshes with this observation (Fig. 9). According to Pérez-Sanz et al. (2013), milder climate conditions 428 characterized northern Spain from 8100 to 5700 BP. Mediterranean climate (warmer and drier) 429 has spread since 5700 BP. Guiot & Kaniewski (2015) previously demonstrated that 430 temperatures increased during the interval 5600-5700 BP. Although the period 5700-2400 BP 431 was warm and dry, it never exceeded the values recorded since the end of the 20th century and 432 those of the 21st century CE. 433

A major consequence of the ongoing warming is falling water availability in island contexts, 434 whose environmental pressure absorption capacities are much lower than for mainland areas 435 (Marriner et al., 2019). As stated by Kent et al. (2002), water shortages are a key environmental 436 problem in Spanish islands, attracting both popular and political attention. Our pollen-based 437 climate reconstruction supports these challenges, suggesting that current precipitation levels are 438 substantially lower than reconstructed values for the Holocene (Fig. 9). For much of the 439 Mediterranean, recent studies have indicated that, for each 1°C of global warming, mean rainfall 440 will probably decrease by about 4% (Cramer et al., 2018; Lionello and Scarascia, 2018). 441

442

#### 443 5.3. Global significance

This study suggests two different climate pacemakers for the Balearic Islands. The climate 444 oscillations playing out during the Holocene seem to result from an ocean-atmosphere coupling 445 with a clear imprint of solar activity as the main driver. A connection between solar irradiance 446 and water balance has previously been suggested for Spain (Morellón et al., 2011). On 447 Mallorca, the warmest and driest interval (2400-5700 BP) during the pre-industrial period 448 correlates with increases in solar activity (Fig. 10) as reconstructed from total solar irradiance 449 ( $\Delta$ TSI; Steinhilber et al., 2012), sunspot numbers (Solanski et al., 2005) and from the solar 450 modulation function  $\Phi$  [MeV] based on cosmogenic radionuclides (<sup>10</sup>Be production rates; 451 Vonmoos et al., 2006). The industrial era shows a different trend with a rate and magnitude of 452 changes that exceed the long-term natural variability (Fig. 7; IPCC, 2014, 2018). Changes in 453 solar activity appear to have been a primary pacemaker of long-term climate variations in 454 Mallorca. By contrast, anthropogenic change has been the primary pacemaker since the 455 industrial era, with a clear acceleration of warming and aridification trends. 456

457

#### 459 5.4. Social and political implications

In the coming decades, warming will have a significant impact on the local economy of the 460 Balearics, which is heavily dependent on tourism (e.g. Kent et al., 2002; Amelung and Viner, 461 2006). Since the 1960s, tourism has reached an apex on Mallorca and Ibiza, with approximately 462 11 million visitors in 2000-2001 and more than 12 million in 2011 (Hof and Blázquez-Salom, 463 2013). As shown by a recent tourism model, climate is an important variable in estimating the 464 economic sustainability of these islands (Álvarez-Díaz and Rosselló-Nadal, 2010). Different 465 climate change scenarios coupled with the Tourism Climatic Index suggest that, during the 21st 466 century, the Balearic Islands, in addition to the wider Mediterranean region, will probably 467 468 become unsustainably hot during the summer months, while, by contrast, northern Europe will be climatically more attractive for tourism (Amelung and Viner, 2006). Socially and 469 economically, the elucidated changes could become detrimental to tourist-based economies 470 throughout the Mediterranean (Amelung and Viner, 2006). 471

Water management on Mallorca is already under pressure from increasing variability in
seasonal and inter-annual rainfall patterns (Candela et al., 2009; Homar et al., 2010; López
Mayol et al., 2017), creating an unbalanced distribution of water/groundwater resources
(domestic supply and irrigation-based agriculture versus natural ecosystem replenishment).
This situation is likely to present serious challenges to the future sustainability of the island and
its economic model based on mass tourism (Candela et al., 2009; Hof and Schmitt, 2011).

478

#### 479 *5.4. Reducing the impact*

Climate change concerns have led to major political changes on the Balearic Islands. Its parliament, for instance, has recently passed a law on energy transition and climate change that commits local tourism to draw all of its energy needs from renewable sources by 2050. This 483 study reinforces the idea that the Balearics must rapidly adopt such measures to ensure the 484 medium- to long-term viability of its economic model based on mass tourism.

485

#### 486 6. Conclusions

In conclusion, the analyses presented above provide answers to the two research questions 487 posed in this paper. First, through the prism of Mallorca, results have shown that the industrial-488 era warming and aridification of Mediterranean islands deviates markedly from the long-term 489 Holocene trends. The instrumental MAT appears  $3.3 \pm 1.2$  °C warmer than the reconstructed 490 MAT for the last eight millennia, with a 33% drop in precipitation between the last 8000 years 491 492 and the industrial era. Second, this study has also demonstrated the importance of palaeoecological archives in contextualizing current climate anomalies. It has revealed the slow 493 pace of climate dynamics on Mallorca over the past 8000 years, both in rate and magnitude, 494 compared to the speed of changes during the industrial era. This contrast also suggests that, 495 despite human colonization of the Balearic Islands around 4400 years ago, the "Anthropocene" 496 of accelerating human interactions with Earth systems closely intertwines with the industrial 497 era. A human imprint on climate is clearly recorded in the data since 1900 CE, with a sharp 498 acceleration in temperatures and aridity since the 1980s. When temperatures and precipitation 499 during the last 40 years are used as benchmarks (period 1980-2018), the pre-industrial climate 500 of Mallorca appears cooler and wetter than anything that has been recorded since. 501

502

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522

#### 523 Declaration of Competing Interest

524 The authors declare that they have no known competing financial interests or personal 525 relationships that could have appeared to influence the work reported in this paper.

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- 774
- 775 Figures

Figure 1. Geographical location the Albufera Natural Park in Mallorca. The study area is
denoted by a red star. The geomorphological map of the Bay of Alcúdia and the swamp area
(Albufera Natural Park) is adapted from Giaime et al. (2017).

Figure 2. Age-model for the core POL12. The lithology of the core is described in relation to
 depth. The radiocarbon dates are depicted as intercepts and 2-sigma calibrations (95%
 probability).

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Figure 3. Detailed pollen diagram (arboreal pollen – AP; non-arboreal pollen - NAP) of the
core POL12. The arboreal and herbaceous taxa are plotted on a linear age-scale. The vertical
lines indicate the main transitions.

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**Figure 4.** Fire activity, agriculture (Poaceae cerealia) and olive trees (*Olea europaea*) plotted on a linear age-scale. The vertical lines denote the results of the homogeneity test (each average is indicated by the letters mu).

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**Figure 5.** Environmental regression models from the Albufera Natural Park. The selected environmental data are indicated in each graph. The consistency of each reconstructed model is indicated by the R2. The lower graph corresponds to a ranking of the environmental data based on the R2 values.

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**Figure 6.** Long-term trends in Mallorca's mean seasonal temperatures. The chronological scales are expressed in years BP and Before Common Era - Common Era (BCE - CE). The modelled temperatures for each season are displayed, with the standard deviation for each data point. The univariate analyses (mean, 25<sup>th</sup> percentile-Q1, median-Q2, 75<sup>th</sup> percentile-Q3, minimum value, maximum value) are summarized using a boxplot. The period with no data (1940 to 1975 CE) is denoted by a vertical grey bar.

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**Figure 7.** Long-term evolution of mean annual temperatures on Mallorca. (**A**) The modelled and measured values are presented as curves and smoothing splines with the standard deviation for each reconstructed data point. The mean value for each dataset is represented in purple. (**B**) Focus on the period 1600-2018, showing the mean annual temperatures and associated standard errors. The curves are continuous despite the absence of data between 1940 and 1975. (**C**) Temperature anomalies for the pre-industrial era with the mean 1900-1980 used as the baseline.

**Figure 8.** Long-term temperature anomalies on Mallorca. (**A**) Temperature anomalies for the pre-industrial era, with the mean 1981-2018 (16.71°C) used as the baseline. A smoothing spline and a polynomial curve were fitted to the sequence. (**B**) Temperature anomalies for the preindustrial and industrial eras with a mean calculated on the whole dataset (14.12°C). This value was used as the baseline. A smoothing spline and a polynomial curve were added to the sequence. The age is expressed as b2k.

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**Figure 9.** Long-term trends in Mallorca's precipitation. The chronological scales are expressed in years BP and the Common Era (CE). The reconstructed annual precipitation values are displayed with the standard deviation. The univariate analyses (mean, 25th percentile-Q1, median-Q2, 75th percentile-Q3, minimum value, maximum value) are summarized using boxplots. The period with no data (1940 to 1972 CE) is denoted by a vertical grey bar. Two different vertical scales are used (right: reconstructed values; left: instrumental data).

- Figure 10. Temperature and precipitation trends on Mallorca during the pre-industrial period
- compared with the long-term solar activity (Solanski et al., 2005; Steinhilber et al., 2012). The
- 827 yellow shaded part underscores the warmer and driest phase during the last 8,000 years.



















