

# Chemical Characteristics and Brown Carbon Chromophores of Atmospheric Organic Aerosols Over the Yangtze River Channel: A Cruise Campaign

X. Wang, N. Hayeck, M. Bruggemann, L. Abis, M. Riva, Y. Lu, Biao Wang,

J. Chen, C. George, L. Wang

# ▶ To cite this version:

X. Wang, N. Hayeck, M. Bruggemann, L. Abis, M. Riva, et al.. Chemical Characteristics and Brown Carbon Chromophores of Atmospheric Organic Aerosols Over the Yangtze River Channel: A Cruise Campaign. Journal of Geophysical Research: Atmospheres, 2020, 125 (16), 10.1029/2020JD032497. hal-02943215

# HAL Id: hal-02943215 https://hal.science/hal-02943215

Submitted on 18 Nov 2020

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Chemical characteristics and brown carbon chromophores of
atmospheric organic aerosols over the Yangtze River channel: a
cruise campaign
Xinke Wang <sup>1,2</sup> , Nathalie Hayeck <sup>2\$</sup> , Martin Brüggemann <sup>3</sup> , Letizia Abis <sup>2</sup> , Matthieu Riva <sup>2</sup> , Yiqun
Lu <sup>1</sup> , Buwei Wang <sup>4</sup> , Jianmin Chen <sup>1</sup> , Christian George <sup>2</sup> *, Lin Wang <sup>1,5,6</sup> *
<sup>1</sup> Shanghai Key Laboratory of Atmospheric Particle Pollution and Prevention (LAP <sup>3</sup> ), Department
of Environmental Science & Engineering, Jiangwan Campus, Fudan University, Shanghai 200438,
China
<sup>2</sup> Univ Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626, Villeurbanne,
France
<sup>3</sup> Leibniz Institute for Tropospheric Research (TROPOS), Atmospheric Chemistry Department
(ACD), Permoserstr. 15, 04318 Leipzig, Germany
<sup>4</sup> State Key Laboratory of Operation and Control of Renewable Energy & Storage Systems, China
Electric Power Research Institute. Beijing 100192, China
<sup>5</sup> Collaborative Innovation Center of Climate Change, Nanjing 210023, China
<sup>6</sup> Shanghai Institute of Pollution Control and Ecological Security, Shanghai 200092, China
<sup>\$</sup> Now at the Chemistry Department, Faculty of Arts and Sciences, American University of Beirut,
Beirut, Lebanon
* Corresponding Authors: C.G., email, Christian.George@ircelyon.univ-lyon1.fr; phone,
(33)(0)437423681; fax, (33)(0)472448438. L.W., email, lin_wang@fudan.edu.cn; phone, +86-21-
31243568.
Key Points:
(1) More than one thousand atmospheric particulate organics including seventy-six BrC
chromophores were determined over the Yangtze River channel.
(2) Brown carbon chromophores exhibit high degrees of unsaturation, suggesting being
aromatics, nitro-aromatics, or polycyclic aromatics.
(3) Biomass burning tracers contribute significantly to the total light absorption of
organic aerosols.

#### 29 Abstract.

30 Organic aerosols (OAs) have important influences on the climatic implications and 31 health effects of atmospheric aerosols. Among the complex OA constituents, brown 32 carbon (BrC) accounts for a substantial mass fraction and is of special interest because 33 of its light-absorbing properties. In this study, the chemical composition of atmospheric 34 OAs over the middle-lower Yangtze River (MLYR) channel, as well as the BrC, was investigated during a ship cruise campaign in winter 2015. In total, more than one 35 36 thousand molecular formulas were determined using a combination of ultra-high performance liquid chromatography (UHPLC), a diode array detector (DAD), and 37 38 Orbitrap high-resolution mass spectrometry (HRMS). Large numbers and enhanced signal abundances for known tracers as well as monocyclic and polycyclic aromatics 39 indicate that biomass burning and fossil combustion are important sources of OAs over 40 the MLYR channel. In addition, thirteen chromophores with strong light absorption, 41 42 mostly representing established biomass burning tracers, were unambiguously determined by UHPLC/DAD/HRMS and contributed to 35-37% of the total light 43 absorption of OAs at 290 nm and 58-70% at 350 nm. Sixty-three previously identified 44 biomass burning chromophores were also positively identified in the mass 45 spectrometric analysis here but embedded in the humped signal during the 46 spectroscopic analysis. These BrC chromophores exhibit high degrees of unsaturation, 47 48 suggesting these compounds to be aromatic, nitro-aromatic, and polycyclic aromatic type of species. Our results highlight the significant influence of biomass burning and 49 50 fossil combustion on atmospheric OAs over the MLYR channel in the winter, strongly 51 enhancing light-absorbing properties and decreasing air quality.

# 52 **1. Introduction**

53 Organic aerosols (OAs) contribute to around 20-50% of the submicron particulate mass in continental mid-latitudes and can account for as much as 90% of the particulate 54 55 mass in tropical forested areas, exerting important influences on climatic processes and human health (Andreae & Crutzen, 1997; Kanakidou et al., 2005). OAs can be released 56 57 through direct emissions from natural or anthropogenic sources and/or formed in the atmosphere through chemical reactions of volatile organic compounds (VOCs) and the 58 59 subsequent gas-to-particle conversion of less volatile reaction products (Pöschl, 2005). In addition to their diverse sources, the multiple chemical reactions of OA components 60 61 can significantly increase their complexity (Goldstein & Galbally, 2007). Eventually, 62 these OA components impact the physicochemical properties and the potential toxicity of OAs (Ito et al., 2019; Lei et al., 2014; Tang et al., 2016). 63

64 Light-absorbing OA components, typically referred to as brown carbon (BrC), exhibiting strong light absorption in the near-ultraviolet (UV) and visible wavelength 65 66 regions, thus, affecting climate forcing properties of aerosol particles (Bahadur et al., 2012; Chung et al., 2012). The study of Feng et al. (2013) concluded that the radiative 67 forcing of climate by BrC can contribute up to +0.25 W m<sup>-2</sup>, higher than 25% of that 68 by black carbon (BC). Released from both fossil fuel and biomass burning, BC 69 70 represents soot-like particles with a broad light absorption from the UV up to the 71infrared wavelength range and may be the second most important atmospheric 72 component for global warming in terms of direct forcing (Jacobson, 2001). It also 73 showed that BrC may be the dominant atmospheric light-absorbing material over the 74 regions of biomass burning and biofuel combustion such as South and East Asia, South America, and subtropical Africa (Feng et al., 2013). In addition, global models 75 76 concluded that the absorption of solar radiation by BrC was 27%-70% of that by BC (Lin et al., 2018). Moreover, the BrC absorption was shown to explain almost 30% of 77 the total aerosol absorption at 370 nm (Yang et al., 2009), and even at 550 nm to be as 78 79 high as 20% (Chung et al., 2012).

80 BrC constituents, compared to BC, are less characterized due to their complex composition and properties (Andreae & Gelencsér, 2006). BrC is present in both 81 primary aerosols, emitted from combustion processes, and secondary organic aerosols 82 83 (SOAs). Several studies have idenfied a wide variety of light-absorbing chromophores, including nitro-aromatics, polycyclic aromatic hydrocarbon derivatives, polyphenols, 84 85 and sulfur-containing compounds, by investigating the molecular composition of freshly emitted biomass burning OAs (Budisulistiorini et al., 2017; Lin et al., 2016; 86 Fleming et al., 2018). The exact chemical composition and physicochemical properties 87 of light-absorption BrC chromophores are related to the fuel type that is burned 88 89 (Chakrabarty et al., 2010; Iinuma et al., 2007; Lin et al., 2016). Secondary BrC were shown to be generated from multiphase chemistry, such as acid-catalyzed aldol 90 91 condensation, nitration of aromatic compounds, and the reactions of ammonium ions and amino acids with carbonyls and dicarbonyls (Bones et al., 2010; De Haan et al., 92 93 2009; Kampf et al., 2016; Kwamena & Abbatt, 2008; Shapiro et al., 2009). Furthermore,

numerous lab studies have shown that atmospheric processing such as solar irradiation, 94 95 oxidation, and changes in temperature and relative humidity can alter the composition 96 of BrC and significantly change their optical properties (Lambe et al., 2012; Lee et al., 2013; Lee et al., 2014; Nguyen et al., 2012; Rincón et al., 2010; Zhao et al., 2015). For 97 98 example, Lee et al. (2014) show that SOAs formed from the photo-oxidation of 99 naphthalene under high NOx and from the reaction of limonene with ozone  $(O_3)$  could be potential sources of secondary BrC, but would "photobleach" meaning lose their 100 near-UV absorbance after solar radiation with effective half-lives of ~14 h and <0.5 h, 101 102 respectively. Therefore, based on the local emissions and atmospheric conditions at a specific geographic location, the chemical composition and physicochemical properties 103 104 of atmospheric BrC components can be very different.

The Yangtze River is the longest river in China and passes by several highly developed economic megacities with major petrochemical complexes and/or steel industries (Li et al., 2018), leading to high emissions of anthropogenic pollutants including particulate matters, VOCs, sulfur dioxide, and nitrogen oxides (Huang et al., 2011). Meanwhile, the Yangtze River also plays an important role in shipping of goods and leisure cruising. Therefore, atmospheric OAs produced/formed over the Yangtze River channel are expected to be complex including abundant BrC components.

Here, we collected filter samples of atmospheric particles during a cruise campaign
carried out along the MLYR channel from November to December 2015 (i.e., Yangtze
River Campaign, YRC), and analyzed the samples to characterize their organic fraction.

In addition, light-absorbing BrC chromophores in the complex OAs were identified via connection of mass spectrometric analysis and spectroscopic analysis. The relative contribution of the strongest BrC chromophores to the total light absorption of OAs was also evaluated. The results obtained help not only to better understand the sources and physicochemical properties of atmospheric aerosols, but also to further characterize the constituents of air pollution at a molecular level over the MLYR channel region.

#### 121 **2. Material and methods**

#### 122 **2.1 Collection of PM<sub>2.5</sub> samples**

Atmospheric aerosol samples were collected onto 90 mm prebaked quartz-fiber 123 filters (Whatman Company, UK) using a middle-flow aerosol sampler (Qingdao 124 Hengyuan Tech Co., Ltd., HY-100) for six hours at a flow rate of 100 L min<sup>-1</sup>, from 125 November 22<sup>nd</sup> to December 5<sup>th</sup>, 2015 along the Yangtze River channel between 126 127 Shanghai and Wuhan (Figure 1). Details about the YRC can be found in a previous 128 study (Li et al., 2018). In order to avoid the possible contaminations from self-emission of the ship, the sampler was placed on the bow of the research vessel. In addition, only 129 130 filter samples during the time period with relative wind directions between -100° and  $100^{\circ}$  of the moving direction of the ship and with a relative wind speed of >1.0 m s<sup>-1</sup> 131 were selected for further analysis. As a result, six samples with sample IDs of 1201M, 132 1201A, 1201N, 1202M, 1202N, and 1203A, respectively, were selected, where the 133 134 numbers represent the sampling date, M denotes "morning", A denotes "afternoon", 135 and N denotes "night". Figure 1 and Table 1 show the route of the vessel during the campaign and the start and end locations of these six samples. Field blank filters were
collected by placing prebaked quartz-fiber filters into a parallel sampler for one hour
without any air flow through the filter substrate. After collection, all filters were stored
at -20°C in a freezer until further analysis to avoid changes in aerosol composition.

140

# 2.2 Aerosol sample analysis

141 Detailed procedures for sample filter extraction can be found in previous studies (Wang et al., 2017; Wang et al., 2016). Briefly, a quarter of each quartz filter was 142 extracted twice with 6 mL of acetonitrile (Optima® LC/MS, Fischer Scientific, USA) 143 144 and agitated for 20 min. Then the combined extracts were filtered through a 0.2 µm polytetrafluoroethylene membrane (13 mm, Pall Corporation, USA), and concentrated 145 146 under a gentle stream of N<sub>2</sub>. The extracts were then reconstituted in 1 mL of a 1:1 vol/vol mixture of water (Optima® LC/MS, Fischer Scientific, USA) and acetonitrile. 147 Lastly, 200 µL of the reconstituted extract was diluted by adding 100 µL of water before 148 149 further analysis due to the high concentrations of organics.

Then the diluted extracts were analyzed by an UHPLC/DAD/HRMS platform, which is the combination of ultra-high performance liquid chromatography (UHPLC, Dionex 3000, Thermo Scientific, USA), a diode array detector (DAD, Dionex UltiMate3000, Thermo Scientific, USA), and Orbitrap high resolution mass spectrometry (HRMS, Q Exactive, Thermo Scientific, Bremen, Germany) using heated electrospray ionization (HESI). Three replicate analyses were performed for each sample extract. The separation was performed using a Waters Acquity HSS T3 column (1.8  $\mu$ m, 100  $\times$  2.1

mm), and the procedures of gradient elution were described in earlier studies (Wang et 157 al., 2017; Wang et al., 2016). In this study, the UV-Vis absorption was measured using 158 159 a DAD over the wavelength range of 280-480 nm with low baseline noise ( $<\pm 10 \,\mu AU$ 160 at 254 nm) and high linearity up to 2.0 AU. For the HRMS analysis, HESI voltages of 161 -2.6 and 3.2 kV were applied in negative and positive ionization mode (ESI- and ESI+) 162 measurements, respectively. The sheath gas flow rate was 42 arbitrary units (a.u.) and the auxiliary gas flow rate was 25 a.u. The temperature of the Orbitrap capillary 163 temperature was set at 350 °C and the temperature of the HESI source at 250 °C. The 164 165 mass resolving power of the Q-Exactive mass spectrometer was 140 000 at m/z 200, and a 2 mM sodium acetate solution was used to provide a series of negative and 166 positive adduct ions in the scanning range of m/z 50–750 Th to perform the daily 167 168 external mass calibration.

Pentafluorobenzylhydroxylamine (PFBHA) derivatization was used to identify organic compounds with carbonyl functional group(s) (Borrás & Tortajada-Genaro, 2012). 200  $\mu$ L of the reconstituted extract were mixed with 800  $\mu$ L of o-(2, 3, 4, 5, 6pentafluorobenzyl) hydroxylamine hydrochloride (Sigma Aldrich,  $\geq$  99.0%) solutions (1 mg/mL). The mixtures were then left in darkness at room temperature for 24 h, and analyzed following the same procedure as earlier described in this session.

## 175 **2.3 Data processing**

Data were acquired using Xcalibur 2.2 (Thermo, USA). The subsequent non-target
screening approach was conducted using an MZmine 2.33 software package. Molecular

178	formulas were assigned to detected signals satisfying the following constraints: $C_{1-50}H_{0-1}$
179	$_{100}O_{0-40}N_{0-5}S_{0-3}$ (C <sub>1-50</sub> H <sub>0-100</sub> O <sub>0-40</sub> N <sub>0-5</sub> S <sub>0-3</sub> F <sub>0-20</sub> for PFBHA derivatives analysis) with a
180	mass tolerance of 2 ppm in ESI- and 3 ppm in ESI+; element count heuristics (i.e., H/C
181	ratio (0.3 $\leq$ H/C $\leq$ 3); NOPS/C ratios (N/C $\leq$ 4, O/C $\leq$ 3, P/C $\leq$ 2, S/C $\leq$ 3); multiple
182	element counts (if each of N, O, P, and S > 1 then N< 10, O < 20, P < 4, S < 3; if each
183	of N, O, and P > 3 then N < 11, O < 22, P < 6; if each of O, P, and S > 1 then O < 14,
184	P < 3, $S < 3$ ; if each of P, S, and $N > 1$ then $P < 3$ , $S < 3$ , $N < 4$ ; if each of N, O, and $S > 1$
185	6 then N < 19, O < 14, S < 8), ring and double bond equivalence (RDBE) restrictions
186	(in the range of 0-25), and isotope pattern matching. It should be noted that the
187	CAMERA algorithm was applied in order to avoid false detection from interferences of
188	multiply charged ions and adducts (Kuhl et al., 2012). Further details on the LC-MS
189	data processing are listed in Tables S1-S2.

190 In addition to RDBE, aromaticity equivalents (Xc) and Kendrick mass defect (KMD) values were calculated, and more details about the calculation can be found in the 191 192 supplemental Text S1 and from our previous study (Wang et al., 2017). Xc has been 193 suggested to help identification and characterization of monocyclic and polycyclic aromatic compounds, with  $Xc \ge 2.50$  and  $Xc \ge 2.71$  as unambiguous minimum criteria 194 195 for the presence of mono- and polycyclic aromatic compounds, respectively (Yassine 196 et al., 2014). KMD is useful to differentiate groups of similar compounds among a large set of molecular formulas. In this study, CH<sub>2</sub> was chosen as a base unit, and thus, 197 molecules with identical KMDs differ only in the number of -CH<sub>2</sub> groups (Hughey et 198 al., 2001). 199

200 The assigned formulas were subdivided into seven groups: compounds containing only carbon, hydrogen, and oxygen atoms in the ESI- and ESI+ (hereafter referred to 201 202 as CHO- and CHO+, respectively); nitrogen-containing organics (hereafter referred to 203 as CHON- and CHON+, respectively); sulfur-containing organics and compounds 204 containing both nitrogen and sulfur in ESI- (hereafter referred to as CHOS- and 205 CHONS-, respectively); and finally, organics without oxygen in ESI+ (hereafter referred to as CHN+). Similarly to our previous study (Wang et al., 2017), CHOS+ and 206 207 CHONS+ compounds were also observed but accounted merely for a few percentages, 208 so they were not further considered here.

It should also be noted that although LC separation might help reducing ion 209 suppression effects, only a semi-quantification of organic compounds can be proposed. 210 211 Using ESI, significant uncertainties originate also from different ionization efficiencies of particulate organic components which typically exhibit a wide range of different 212 structures and functional groups (Lin et al., 2012). Due to the large variety of 213 214 compounds and the technical impossibility of providing calibration factors for individual species, we assumed that the determined compounds have a similar signal 215 response and the total peak areas of organics can be compared among different samples. 216

217 **3. Results and discussion** 

## 218 **3.1 Organic aerosols over the MLYR channel**

In this study, 695-1187 and 450-695 molecular formulas were determined in ESIand ESI+, respectively (Table 1). Due to the inherent differences in the ionization

mechanisms between ESI- and ESI+, a minor fraction of the detected compounds 221 222 appears in both modes as shown by the formula lists (see Dataset 1). Data from ESI-223 and ESI+ can provide complementary information to characterize the molecular composition of complex OAs (Lin et al., 2018). However, non-polar compounds, which 224 225 may be part of BrC, are difficult to be ionized by ESI (Kuang et al., 2018). Therefore, contributions from non-polar compounds might be underestimated in this study. The 226 number and percentage of molecular formulas for each subgroup tentatively identified 227 228 in each sample were listed in Table 1. The greatest number of formulas was detected in 229 the 1202M sample, giving 600 more compounds than the 1203A sample, which showed the lowest number of detected compounds. This large difference in the number of 230 231 compounds clearly demonstrates not only the complexity but also the variety of organic 232 species in the aerosol samples. In addition, the number of CHON species including CHON- and CHON+ accounted for significantly higher percentages in the 1202M and 233 1202N samples compared to any other sample, indicating the different sources for these 234235 compound classes.

The pie charts in Figures 2 and S1 show the relative abundances (i.e., based on peak areas) of all subgroups in each sample. It was reported that the relative contributions of CHO-, CHON-, CHON+, CHOS-, and CHN+ species to wintertime OAs in urban Shanghai were close(Wang et al., 2017). However, within the present study, the relative abundances of CHON+ and CHNO- species are much greater, whereas CHOS- species represent a less fraction (Figures 2 and S1). This difference in abundances indicates that nitrogen-containing compounds, such as organonitrates, likely dominated the physicochemical properties of aerosol particles over the MLYR channel region,
whereas sulfur-containing compounds, such as organosulfates, played only a minor role
compared to urban Shanghai areas.

In addition to molecular formulas, the abundances, element ratios, RDBE, the 246 number of carbonyl group, and the values of KMD and Xc for detected species were 247 listed in the Dataset 1. In the present study, large amounts of organics with low 248 saturation degrees were observed in all samples. In both ESI- and ESI+, 62-72% of the 249 determined formulas are characterized by high RDBE values (RDBE  $\geq$  4) and 39-52% 250 of the formulas show Xc values  $\geq 2.50$ , indicating that most of these compounds are 251 aromatics (Wang et al., 2017; Yassine et al., 2014). A further analysis shows that around 252 15-23% of the formulas in ESI+ and 23-28% of the formulas in ESI- were probably 253 polycyclic aromatics ( $Xc \ge 2.71$ ), representing more than half of the number of aromatic 254 molecules. These results suggest that anthropogenic emissions made important 255256 contributions to the atmospheric OAs over the MLYR channel.

Li et al. (2018) classified the air pollution during the cruise campaign into eight distinct episodes based on sampling locations, backward trajectories, and photochemical processes. We note that the sampling duration for the 1201M, 1201A, and 1201N samples corresponds to episode #5 in an overview paper of the cruise campaign (Li et al., 2018), whereas that for the 1202M and 1202N samples overlaps episode #6, and the 1203A sample coincides with episode #7, respectively. Based on the concentrations of carbon monoxide, and trace elements and levoglucosan in

264	particulate matters, episode #5 and episodes #6-7 are cases strongly influenced by coal
265	combustion and biomass burning, respectively (Li et al., 2018). The larger numbers and
266	higher abundances of sulfur-containing compounds in the 1201M and 1201A samples
267	(Figure 2 and Table 1) are probably caused by coal combustion, which is in agreement
268	with Li et al. (2018). In addition to the most intense species detected in ESI- and ESI+
269	(see Figures 2 and S1, and Tables 2 and S3), other organics detected with high
270	abundances in the samples, including C7H7NO, C8H7NO, and C9H7NO were assigned
271	to the benzamide, 4-hydroxy-benzene acetonitrile, and 3-(4-hydroxyphenyl)
272	propionitrile, respectively, which are well known tracers derived from biomass burning
273	(Laskin et al., 2009; Lin et al., 2016; Ma & Hays, 2008). In addition to the cities and
274	factories, there are farmland and rural areas along the Yangtze River. Compound J'
275	(Figure 2, referred to $C_{10}H_{17}NO_7S$ ), with the highest abundance in the 1201M sample,
276	is tentatively assigned to be derived from monoterpene oxidation in the presence of
277	$NO_x$ and $SO_2$ (Brüggemann et al., 2015; Seinfeld et al., 2008; Wang et al., 2018). It
278	should be noted that these tracers may also be formed from chemical processing and
279	cannot be distinguished in the present study, but precursor molecules with aromatic
280	structures likely originated from the same sources. Overall, biomass burning and fossil
281	fuel combustion could be important organic aerosol sources over the MLYR channel
282	(Li et al., 2018).

# 283 **3.2 Molecular characterization of BrC chromophores**

BrC chromophores were distinguished by matching the absorption chromatograms

with the DAD detector to the extracted ion chromatograms. The mass spectrum before 285 and at a major absorption were compared to help identifying the potential 286 287 chromophores, as shown in Figure S2. We note that only chromophores with high ion abundances and strong light absorptions can be unambiguously determined by this 288 289 approach. Indeed, the highly complex composition of the aerosol samples can lead to 290 overlapping peaks in both extracted ion chromatograms and UV-Vis absorption spectra. Note that only the 1201M, 1201A and 1202N samples, which were analyzed with a 291 292 good DAD performance, were chosen to determine the BrC chromophore.

293 UHPLC-DAD absorption chromatograms between 290-480 nm wavelengths were 294 blank-corrected, normalized by the volume of the sample, and subsequently plotted in Figures 3 and S3-S4. Intense absorption at near UV wavelengths (290-350 nm) is 295 296 evident. Much higher light absorption was observed in the 1202N sample than in the 1201M and 1201A samples, likely due to the greater concentrations of BrC 297 298 chromophores in the 1202N sample (Figures 2 and S1). For example, the relative 299 abundances of C<sub>6</sub>H<sub>5</sub>NO<sub>3</sub> in the 1201M and 1201A samples only account for around 17% and 15% of that in the 1202N sample, respectively. The presence of complex mixtures 300 301 of chemicals including isomeric chromophores could have contributed to the humped 302 light absorption in these chromatograms (Budisulistiorini et al., 2017).

Following the UHPLC/DAD/HRMS protocol in Figure S2, we identified thirteen BrC chromophores and their formulas were labelled next to the corresponding absorption peaks in Figures 3 and S3-S4. All of them are characterized with high

306	unsaturation degrees (RDBE $\geq 4$ ), indicating the presence of poly-conjugated double
307	bonds. For example, C <sub>6</sub> H <sub>5</sub> NO <sub>3</sub> , presumably nitrophenol, is the most abundant species
308	in the six samples and derived from biomass burning (Mohr et al., 2013). Nitrogen-
309	containing chromophores including C <sub>6</sub> H <sub>5</sub> NO <sub>4</sub> , C <sub>7</sub> H <sub>5</sub> NO <sub>5</sub> , C <sub>7</sub> H <sub>7</sub> NO <sub>3</sub> , C <sub>7</sub> H <sub>7</sub> NO <sub>4</sub> ,
310	$C_8H_7NO_3$ , and $C_7H_6N_2O_5$ also exhibit high degrees of unsaturation and were assigned
311	to nitro-aromatics that are also known to be emitted from biomass burning (Desyaterik
312	et al., 2013; Lin et al., 2016). In addition to the presence of monoaromatics, products
313	formed from the photooxidation of polycyclic aromatic hydrocarbons (PAHs) were
314	observed (Riva et al., 2015). As an example, C <sub>7</sub> H <sub>6</sub> O <sub>2</sub> assigned to benzoic acid was
315	identified in large amount in the different samples. It is worth pointing out that such
316	carbonyl-containing compounds are well-suited candidates to act as photosensitizer,
317	potentially inducing chemical reactions at the surface and in the bulk of aerosol particles
318	(George et al., 2015; McNeill & Canonica, 2016). In total, more than seventy carbonyl-
319	containing formulas were determined based on the PFBHA derivatization, with 6-24%
320	of carbonyl-containing formulas in ESI- and 35-45% in ESI+ belonging to aromatic
321	compounds (i.e., $Xc \ge 2.50$ , see Dataset 1). In addition, polycyclic aromatics including
322	$C_{17}H_{10}O$ (RDBE = 13, Xc = 2.83), $C_{19}H_{10}O$ (RDBE = 15, Xc = 2.86) and $C_{13}H_8O_2$
323	(RDBE = 15, $Xc = 2.75$ ) were contributors to light absorption (Figures 3 and S3-S4),
324	and the latter of which $(C_{13}H_8O_2)$ is an important tracer for biomass burning (Lin et al.,
325	2016). Finally, due to the large presence of $C_{14}H_{11}NO$ (RDBE = 10 and a Xc = 2.78;
326	Figure S1) that is assigned to benzyloxybenzonitrile species, microorganism is another
327	potential source of BrC compounds (Garcia-Alcega et al., 2017) within the Yangtze

328 River channel.

329 In addition to unambiguously determining the thirteen chromophores by the 330 UHPLC/DAD/HRMS analysis, we compared the mass list of organic species in our samples and those in previously observed during biomass burning experiments 331 (Budisulistiorini et al., 2017; Lin et al., 2016). As a result, sixty-three BrC 332 333 chromophores including 31 formulas were observed only in ESI-, 14 formulas only in ESI+, and 18 formulas in both modes. The retention time of all these species 334 335 corresponded to the light absorption region of the chromatogram but did not match an outstanding absorption peak. Their molecular weights are greater than 120 Da, but less 336 337 than 350 Da (Figure S5) and have high RDBE values (ranging from 4 to 12), indicating that such species are likely aromatics. In addition, Figure S6 shows the distribution of 338 339 the number of nitrogen (0-2) and oxygen (1-7) atoms of these seventy-six BrC chromophores. During biomass burning and fuel combustion, OH-initiated oxidation 340 341 may lead to the formation of oxygenated organic compounds, producing light absorbing species (Saleh et al., 2014). Three reduced nitrogen-containing aromatic chromophores 342 with O/N < 3, such as  $C_{14}H_{11}NO$ , were identify in ESI+, whereas more than twenty 343 nitro-aromatic chromophores with  $O/N \ge 3$  were preferably detected in ESI- (Zhang et 344 al., 2013). These nitrogen-containing aromatics were probably produced by 345 heterogeneous reactions of aromatics with NO<sub>2</sub> and nitrate radicals (NO<sub>3</sub>) (Kwamena 346 & Abbatt, 2008) and/or oligomerization of dicarbonyls reacting with ammonium during 347 combustion (Budisulistiorini et al., 2017; Kampf et al., 2016). 348

Figures 4 and S7-S11 show the CH<sub>2</sub>-Kendrick and Van Krevelen (VK) diagrams for 349 the 1201M, 1201A, and 1202N samples with the seventy-six identified BrC 350 351 chromophores being color-coded. Homologous compounds line up a horizontal line in the CH<sub>2</sub>- Kendrick diagrams but a tilted one in the VK diagrams, suggesting that the 352 353 homologues could have variable degrees of saturation and oxidation but possible similar structure and physicochemical properties. This leads to tentative assignments of 354 over 120 species to BrC chromophores. These species were not outstanding either in 355 the mass spectra or in the absorption chromatography, but could have contributed to the 356 357 humped light absorption signal.

358 Many studies show that nitrogen-containing heterocyclic imine and imidazole compounds that are derived from reactions of carbonyls and dicarbonyls with 359 360 ammonium, ammonia or amino acids, can absorb at wavelengths of 300-500 nm (De Haan et al., 2009; Kampf et al., 2016; Nguyen et al., 2013; Nozière et al., 2009; Sareen 361 362 et al., 2009). For example, C<sub>6</sub>H<sub>8</sub>ON<sub>2</sub> could be formed from the reaction of methylglyoxal with ammonium, being an important BrC chromophore (Lin et al., 2015). 363 However, its relative abundances only account for less than 0.6% of the most abundant 364 ion (i.e., C<sub>16</sub>H<sub>33</sub>NO in 1201A sample) in 1201M, 1201A and 1202N samples. In this 365 study, 173-334 CHON+ species and 87-136 CHN+ species were determined according 366 to their exact mass. These CHON+ compounds are characterized with low degrees of 367 368 oxidation with average values of O/N lower than 1.53, indicating the presence of reduced nitrogen in the formula. In addition, 57-59% of CHON+ and 70-78% of CHN+ 369 370 have high unsaturation degrees with RDBE  $\geq$  4, whereas 24-27% of CHON+ and 29371 36% of CHN+ belong to monoaromatics  $(2.71 > Xc \ge 2.50)$  and 6-13% of CHON+ and 372 34-44% of CHN+ belong to polycyclic aromatics ( $Xc \ge 2.71$ ), respectively, suggesting 373 that many of these reduced nitrogen-containing compounds probably contain 374 conjugated systems and possess light absorption abilities.

# 375 **3.3 Light absorption contribution of determined strong chromophores**

Figure 5 shows the light absorption contribution from the 13 strong BrC chromophores unambiguously determined by UHPLC/DAD/HRMS to the total light absorption of OAs at 290-480 nm. The light absorption for each chromophore was summed up based on the corresponding ion chromatogram.

380 Overall, the total light absorption of OAs is the highest for the 1202N sample and the lowest for the 1201A samples. The maximum absorption was observed at 290 nm, and 381 the total absorption decreased as the wavelength increased, showing a similar trend to 382 that in the study of Lin et al. (2016). However, in this study, the light absorptions are 383 384 almost zero at wavelength of 400 nm, and thus, lower than those of the study of Lin et 385 al. (2016). This weakened absorption probably indicates larger concentrations of 386 chromophores in biomass burning aerosols collected in their study compared to the ambient aerosol samples collected over the MLYR channel. At 290 nm, around 35-37% 387 388 of the light absorption of OAs can be attributed to the 13 strong BrC chromophores, whose contribution increased to around 58-70% at 350 nm. In particular, the determined 389 390 nitro-aromatics (i.e., C<sub>6</sub>H<sub>5</sub>NO<sub>3</sub>, C<sub>6</sub>H<sub>5</sub>NO<sub>4</sub>, C<sub>7</sub>H<sub>7</sub>NO<sub>3</sub>, C<sub>7</sub>H<sub>7</sub>NO<sub>4</sub>, C<sub>7</sub>H<sub>5</sub>NO<sub>5</sub>, C<sub>8</sub>H<sub>7</sub>NO<sub>3</sub>, 391 C<sub>8</sub>H<sub>9</sub>NO<sub>3</sub>, and C<sub>7</sub>H<sub>6</sub>N<sub>2</sub>O<sub>5</sub>) alone contribute to about 25-27% at 290 nm and 55-65% at

392 350 nm of the total light absorption of OAs. Moreover, it should be noted that the light 393 absorption of the chromophores from biomass burning (i.e.,  $C_6H_5NO_3$ ,  $C_6H_5NO_4$ , 394  $C_7H_5NO_5$ ,  $C_7H_7NO_3$ ,  $C_7H_7NO_4$ ,  $C_8H_7NO_3$ ,  $C_7H_6N_2O_5$ , and  $C_{13}H_8O_2$ ) account for 27-395 28% at 290 nm and 55-68% at 350 nm, further indicating that biomass burning is 396 potentially the most important sources for light-absorbing species in aerosol particles 397 in the MLYR region.

#### 398 **4.** Conclusions

Nearly one-third of the Chinese population live along the Yangtze River; the air 399 400 pollution over the MLYR channel region can impact the aquatic ecosystem, the health of local residents, and regional climate. Therefore, it becomes crucial to characterize 401 402 the constituents and sources of air pollutants over the MLYR channel region. To the 403 best of our knowledge, YRC is the first comprehensive cruise observation of air quality along the MLYR channel. The chemical composition of the organic fraction of 404 atmospheric particles over this region was investigated, for the first time, using 405 UHPLC/DAD/HRMS with a special focus on BrC chromophores in the present study, 406 which will significantly improve our understanding of BrC compositions, sources and 407 408 optical properties in this region. So far, only a few studies have investigated atmospheric BrC at such a detailed molecular level. In total, 695-1187 and 450-695 409 molecular formulas were determined in ESI- and ESI+, respectively. In addition, 62-410 411 72% of the organic compounds exhibit high degrees of unsaturation (RDBE  $\geq$  4). Furthermore, 39-52% and 15-23% of these organics have an  $Xc \ge 2.50$  and  $Xc \ge 2.71$ , 412

indicating that they are aromatics and polycyclic aromatics, respectively. Moreover,
high signal abundances and large numbers of known tracers indicate that biomass
burning and fossil fuel combustion are important sources for OAs in the MLYR region.

In total, thirteen strong chromophores were determined by UHPLC/DAD/HRMS 416 analysis. Most of these compounds are probably derived from biomass burning. Apart 417 418 from that, more than sixty chromophores identified in previous biomass burning experiments were also detected in this study, further confirming the abundance of 419 420 biomass burning emissions in the MLYR. All of these chromophores were characterized by a low degree of saturation, and Xc values indicating aromatic, nitro-aromatic, and 421 422 polycyclic aromatic structures. Based on KMD and VK analyses, around one hundred homologues to the identified chromophores were detected, probably having similar 423 structures and light-absorption properties. A large amount of reduced nitrogen-424 containing compounds and potential photosensitizers (i.e., carbonyl-containing 425 426 compounds) were found, probably contributing further to the light absorptions of 427 aerosols in the MLYR.

We found that the determined thirteen strong chromophores contribute to around 35-37% at 290 nm and 58-70% at 350 nm of the overall light absorption of OAs, respectively. The light absorption contributions of chromophores from biomass burning were estimated to 27-28% at 290 nm and 55-68% at 350 nm, highlighting the importance of such anthropogenic emissions for light-absorption properties of aerosol particles over the MLYR channel region. 434

### 435 Acknowledgments

436 This study was financially supported by the National Natural Science Foundation of China (91644213, 21925601 and 41575113), the National Key R&D Program of China 437 (2017YFC0209505), the ANR-RGC Programme (project ANR-16-CE01-0013), the 438 European Union's Horizon 2020 research and innovation program under grant 439 440 agreement No. 690958 (MARSU). The authors declare no competing interests. All data 441 to support the conclusions of this manuscript are included in the main text and supplementary be download 442 materials, data could or at https://nuage.catalyse.cnrs.fr/s/mj7gwfOXpdWnAEL. 443

444

#### 445 **References:**

- Andreae, M O, & Gelencsér, A. (2006). Atmospheric Chemistry and Physics Black carbon or brown
  carbon? The nature of light-absorbing carbonaceous aerosols. *Atmos. Chem. Phys*, 6, 3131–3148.
  Retrieved from www.atmos-chem-phys.net/6/3131/2006/
- 449Andreae, Meinrat O., & Crutzen, P. J. (1997). Atmospheric aerosols: Biogeochemical sources and role450inatmosphericchemistry.Science,276(5315),1052–1058.451https://doi.org/10.1126/science.276.5315.1052
- Bahadur, R., Praveen, P. S., Xu, Y., & Ramanathan, V. (2012). Solar absorption by elemental and brown
  carbon determined from spectral observations. *Proceedings of the National Academy of Sciences*, *109*(43), 17366–17371. https://doi.org/10.1073/pnas.1205910109
- Bones, D. L., Henricksen, D. K., Mang, S. A., Gonsior, M., Bateman, A. P., Nguyen, T. B., et al. (2010).
  Appearance of strong absorbers and fluorophores in limonene-O<inf>3</inf> secondary organic
  aerosol due to NH<inf>4</inf><sup>+</sup> -mediated chemical aging over long time scales. *Journal of Geophysical Research Atmospheres*, *115*(5), 1–14. https://doi.org/10.1029/2009JD012864
- Borrás, E., & Tortajada-Genaro, L. A. (2012). Determination of oxygenated compounds in secondary
  organic aerosol from isoprene and toluene smog chamber experiments. *International Journal of Environmental* Analytical Chemistry, 92(1), 110–124.
  https://doi.org/10.1080/03067319.2011.572164
- Brüggemann, M., Karu, E., Stelzer, T., & Hoffmann, T. (2015). Real-time analysis of ambient organic
  aerosols using aerosol flowing atmospheric-pressure afterglow mass spectrometry (AeroFAPAMS). *Environmental Science and Technology*, 49(9), 5571–5578.
  https://doi.org/10.1021/es506186c
- Budisulistiorini, S. H., Riva, M., Williams, M., Chen, J., Itoh, M., Surratt, J. D., & Kuwata, M. (2017).
  Light-Absorbing Brown Carbon Aerosol Constituents from Combustion of Indonesian Peat and
  Biomass. *Environmental Science and Technology*, 51(8), 4415–4423.
  https://doi.org/10.1021/acs.est.7b00397
- 471 Chakrabarty, R. K., Moosmüller, H., Chen, L. W. A., Lewis, K., Arnott, W. P., Mazzoleni, C., et al. (2010).

- 472 Brown carbon in tar balls from smoldering biomass combustion. *Atmospheric Chemistry and* 473 *Physics*, *10*(13), 6363–6370. https://doi.org/10.5194/acp-10-6363-2010
- Chung, C. E., Ramanathan, V., & Decremer, D. (2012). Observationally constrained estimates of
  carbonaceous aerosol radiative forcing. *Proceedings of the National Academy of Sciences*, 109(29),
  11624–11629. https://doi.org/10.1073/pnas.1203707109
- 477 Desyaterik, Y., Sun, Y., Shen, X., Lee, T., Wang, X., Wang, T., & Collett, J. L. (2013). Speciation of
  478 "brown" carbon in cloud water impacted by agricultural biomass burning in eastern China. *Journal*479 of Geophysical Research Atmospheres, 118(13), 7389–7399. https://doi.org/10.1002/jgrd.50561
- Feng, Y., Ramanathan, V., & Kotamarthi, V. R. (2013). Brown carbon: A significant atmospheric absorber
  of solar radiation. *Atmospheric Chemistry and Physics*, 13(17), 8607–8621.
  https://doi.org/10.5194/acp-13-8607-2013
- 483 Fleming, L. T., Lin, P., Laskin, A., Laskin, J., Weltman, R., Edwards, R. D., et al. (2018). Molecular 484 composition of particulate matter emissions from dung and brushwood burning household 485 cookstoves Haryana, India. Phys, 185194, 2461-2480. in Atmos. Chem. 486 https://doi.org/10.5194/acp-18-2461-2018
- 487 Garcia-Alcega, S., Nasir, Z. A., Ferguson, R., Whitby, C., Dumbrell, A. J., Colbeck, I., et al. (2017). 488 Fingerprinting outdoor air environment using microbial volatile organic compounds (MVOCs) -489 review. **TrAC** Trends Chemistry, 86. 75-83. А in Analytical 490 https://doi.org/10.1016/j.trac.2016.10.010
- George, C., Ammann, M., D'Anna, B., Donaldson, D. J., & Nizkorodov, S. A. (2015). Heterogeneous
  Photochemistry in the Atmosphere. *Chemical Reviews*, 115(10), 4218–4258.
  https://doi.org/10.1021/cr500648z
- Goldstein, A. H., & Galbally, I. E. (2007). Known and Unexplored ORGANIC CONSTITUENTS in the
  Earth's Atmosphere Much remains to be learned about the sources, structure, chemistry, and fate
  of gas-phase and aerosol organic compounds. *Environmental Science & Technology*, (1), 1515–
  1521.
- Gómez-González, Y., Wang, W., Vermeylen, R., Chi, X., Neirynck, J., Janssens, I. A., et al. (2012).
  Chemical characterisation of atmospheric aerosols during a 2007 summer field campaign at
  Brasschaat, Belgium: Sources and source processes of biogenic secondary organic aerosol. *Atmospheric Chemistry and Physics*, 12(1), 125–138. https://doi.org/10.5194/acp-12-125-2012
- De Haan, D. O., Tolbert, M. A., & Jimenez, J. L. (2009). Atmospheric condensed-phase reactions of
  glyoxal with methylamine. *Geophysical Research Letters*, 36(11), 2–6.
  https://doi.org/10.1029/2009GL037441
- Huang, C., Chen, C. H., Li, L., Cheng, Z., Wang, H. L., Huang, H. Y., et al. (2011). Emission inventory
   of anthropogenic air pollutants and VOC species in the Yangtze River Delta region, China.
   *Atmospheric Chemistry and Physics*, 11(9), 4105–4120. https://doi.org/10.5194/acp-11-4105-2011
- Hughey, C. A., Hendrickson, C. L., Rodgers, R. P., Marshall, A. G., & Qian, K. (2001). Kendrick Mass
   Defect Spectrum: A Compact Visual Analysis for Ultrahigh-Resolution Broadband Mass Spectra.
   *Analytical Chemistry*, 73(19), 4676–4681. https://doi.org/10.1021/ac010560w
- 511 Iinuma, Y., Brüggemann, E., Gnauk, T., Müller, K., Andreae, M. O., Helas, G., et al. (2007). Source
  512 characterization of biomass burning particles: The combustion of selected European conifers,
  513 African hardwood, savanna grass, and German and Indonesian peat. *Journal of Geophysical*514 *Research Atmospheres*, *112*(8). https://doi.org/10.1029/2006JD007120
- 515 Iinuma, Yoshiteru, Böge, O., & Herrmann, H. (2010). Methyl-nitrocatechols: Atmospheric tracer

- 516 compounds for biomass burning secondary organic aerosols. *Environmental Science and* 517 *Technology*, 44(22), 8453–8459. https://doi.org/10.1021/es102938a
- Ito, T., Bekki, K., Fujitani, Y., & Hirano, S. (2019). The toxicological analysis of secondary organic
   aerosol in human lung epithelial cells and macrophages, 22747–22755.
- Jacobson, M. Z. (2001). Strong radiative heating due to the mixing state of black carbon in atmospheric
   aerosols. *Nature*, 409(6821), 695–697. https://doi.org/10.1038/35055518
- Jang, M., & Kamens, R. M. (2001). Characterization of secondary aerosol from the photooxidation of
   toluene in the presence of NOx and 1-propene. *Environmental Science and Technology*, *35*(18),
   3626–3639. https://doi.org/10.1021/es010676+
- Kampf, C. J., Filippi, A., Zuth, C., Hoffmann, T., & Opatz, T. (2016). Secondary brown carbon formation:
   Via the dicarbonyl imine pathway: Nitrogen heterocycle formation and synergistic effects. *Physical Chemistry Chemical Physics*, 18(27), 18353–18364. https://doi.org/10.1039/c6cp03029g
- Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., et al. (2005).
   Organic aerosol and global climate modelling: a review. *Atmospheric Chemistry and Physics*, 5(4),
   1053–1123. https://doi.org/10.5194/acp-5-1053-2005
- Kuang, B. Y., Yeung, H. S., Lee, C. C., Griffith, S. M., & Yu, J. Z. (2018). Aromatic formulas in ambient
   PM2.5 samples from Hong Kong determined using FT-ICR ultrahigh-resolution mass spectrometry.
   *Analytical and Bioanalytical Chemistry*, 410(24), 6289–6304. https://doi.org/10.1007/s00216 018-1239-8
- Kuhl, C., Tautenhahn, R., Böttcher, C., Larson, T. R., & Neumann, S. (2012). CAMERA: An integrated
   strategy for compound spectra extraction and annotation of liquid chromatography/mass
   spectrometry data sets. *Analytical Chemistry*, 84(1), 283–289. https://doi.org/10.1021/ac202450g
- Kwamena, N. O. A., & Abbatt, J. P. D. (2008). Heterogeneous nitration reactions of polycyclic aromatic
  hydrocarbons and n-hexane soot by exposure to NO3/NO2/N2O5. *Atmospheric Environment*,
  42(35), 8309–8314. https://doi.org/10.1016/j.atmosenv.2008.07.037
- Lambe, a T., Cappa, C. D., Massoli, P., Onasch, T. B., Forestieri, S. D., Martin, a T., & Cummings,
   M. J. (2012). Relationship between oxidation level and optical properties of secondary organic
   aerosol α -pinene naphthalene guaiacol JP-10, 2467.
- Laskin, A., Smith, J. S., & Laskin, J. (2009). Molecular Characterization of Nitrogen-Containing Organic
   Compounds in Biomass Burning Aerosols Using High-Resolution Mass Spectrometry.
   *Environmental Science & Technology*, 43(10), 3764–3771. https://doi.org/10.1021/es803456n
- Laskin, A., Laskin, J., & Nizkorodov, S. A. (2015). Chemistry of Atmospheric Brown Carbon. *Chemical Reviews*, 115(10), 4335–4382. https://doi.org/10.1021/cr5006167
- Lee, A. K. Y., Zhao, R., Li, R., Liggio, J., Li, S. M., & Abbatt, J. P. D. (2013). Formation of light absorbing
   organo-nitrogen species from evaporation of droplets containing glyoxal and ammonium sulfate.
   *Environmental Science and Technology*, 47(22), 12819–12826. https://doi.org/10.1021/es402687w
- Lee, H. J., Aiona, P. K., Laskin, A., Laskin, J., & Nizkorodov, S. A. (2014). Effect of solar radiation on
  the optical properties and molecular composition of laboratory proxies of atmospheric brown
  carbon. *Environmental Science and Technology*, 48(17), 10217–10226.
  https://doi.org/10.1021/es502515r
- Lei, T., Zuend, A., Wang, W. G., Zhang, Y. H., & Ge, M. F. (2014). Hygroscopicity of organic compounds
  from biomass burning and their influence on the water uptake of mixed organic ammonium sulfate
  aerosols. *Atmospheric Chemistry and Physics*, 14(20), 11165–11183. https://doi.org/10.5194/acp14-11165-2014

- Li, Z., Li, C., Ye, X., Fu, H., Wang, L., Yang, X., et al. (2018). Air quality in the middle and lower reaches
  of the Yangtze River channel: A cruise campaign. *Atmospheric Chemistry and Physics*, 18(19),
  14445–14464. https://doi.org/10.5194/acp-18-14445-2018
- Lin, G., Penner, J. E., Flanner, M. G., Sillman, S., Xu, L., & Zhou, C. (2014). Radiative forcing of organic
   aerosol in the atmosphere and on snow: Effects of SOA and brown carbon. *Journal of Geophysical Research*, *119*(12), 7453–7476. https://doi.org/10.1002/2013JD021186
- Lin, P., Rincon, A. G., Kalberer, M., & Yu, J. Z. (2012). Elemental composition of HULIS in the Pearl
  River Delta Region, China: Results inferred from positive and negative electrospray high
  resolution mass spectrometric data. *Environmental Science and Technology*, 46(14), 7454–7462.
  https://doi.org/10.1021/es300285d
- Lin, P., Laskin, J., Nizkorodov, S. A., & Laskin, A. (2015). Revealing Brown Carbon Chromophores
   Produced in Reactions of Methylglyoxal with Ammonium Sulfate. *Environmental Science and Technology*, 49(24), 14257–14266. https://doi.org/10.1021/acs.est.5b03608
- Lin, P., Aiona, P. K., Li, Y., Shiraiwa, M., Laskin, J., Nizkorodov, S. A., & Laskin, A. (2016). Molecular
   Characterization of Brown Carbon in Biomass Burning Aerosol Particles. *Environmental Science and Technology*, 50(21), 11815–11824. https://doi.org/10.1021/acs.est.6b03024
- Lin, P., Fleming, L. T., Nizkorodov, S. A., Laskin, J., & Laskin, A. (2018). Comprehensive Molecular
   Characterization of Atmospheric Brown Carbon by High Resolution Mass Spectrometry with
   Electrospray and Atmospheric Pressure Photoionization. *Analytical Chemistry*, *90*(21), 12493–
   12502. https://doi.org/10.1021/acs.analchem.8b02177
- Lin, Y. H., Budisulistiorini, S. H., Chu, K., Siejack, R. A., Zhang, H. F., Riva, M., et al. (2014). Lightabsorbing oligomer formation in secondary organic aerosol from reactive uptake of isoprene
  epoxydiols. *Environmental Science and Technology*, 48(20), 12012–12021.
  https://doi.org/10.1021/es503142b
- Ma, Y., & Hays, M. D. (2008). Thermal extraction-two-dimensional gas chromatography-mass
   spectrometry with heart-cutting for nitrogen heterocyclics in biomass burning aerosols. *Journal of Chromatography A*, *1200*(2), 228–234. https://doi.org/10.1016/j.chroma.2008.05.078
- McNeill, K., & Canonica, S. (2016). Triplet state dissolved organic matter in aquatic photochemistry:
   Reaction mechanisms, substrate scope, and photophysical properties. *Environmental Science: Processes and Impacts*, 18(11), 1381–1399. https://doi.org/10.1039/c6em00408c
- Mohr, C., Lopez-Hilfiker, F. D., Zotter, P., Prévoît, A. S. H., Xu, L., Ng, N. L., et al. (2013). Contribution
  of nitrated phenols to wood burning brown carbon light absorption in detling, united kingdom
  during winter time. *Environmental Science and Technology*, 47(12), 6316–6324.
  https://doi.org/10.1021/es400683v
- Nguyen, T. B., Lee, P. B., Updyke, K. M., Bones, D. L., Laskin, J., Laskin, A., & Nizkorodov, S. A.
  (2012). Formation of nitrogen- and sulfur-containing light-absorbing compounds accelerated by
  evaporation of water from secondary organic aerosols. *Journal of Geophysical Research Atmospheres*, *117*(1), 1–14. https://doi.org/10.1029/2011JD016944
- Nguyen, T. B., Laskin, A., Laskin, J., & Nizkorodov, S. A. (2013). Brown carbon formation from
  ketoaldehydes of biogenic monoterpenes. *Faraday Discussions*, 165, 473–494.
  https://doi.org/10.1039/c3fd00036b
- Nozière, B., Dziedzic, P., & Córdova, A. (2009). Products and kinetics of the liquid-phase reaction of
  glyoxal catalyzed by ammonium ions (NH4+). *Journal of Physical Chemistry A*, *113*(1), 231–237.
  https://doi.org/10.1021/jp8078293

- Pöschl, U. (2005). Atmospheric aerosols: Composition, transformation, climate and health effects.
  Angewandte Chemie International Edition, 44(46), 7520–7540.
  https://doi.org/10.1002/anie.200501122
- Rincón, A. G., Guzmán, M. I., Hoffmann, M. R., & Colussi, A. J. (2010). Thermochromism of model
  organic aerosol matter. *Journal of Physical Chemistry Letters*, 1(1), 368–373.
  https://doi.org/10.1021/jz900186e
- Riva, M., Robinson, E. S., Perraudin, E., Donahue, N. M., & Villenave, E. (2015). Photochemical aging
  of secondary organic aerosols generated from the photooxidation of polycyclic aromatic
  hydrocarbons in the gas-phase. *Environmental Science and Technology*, 49(9), 5407–5416.
  https://doi.org/10.1021/acs.est.5b00442
- Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., et al. (2014). Brownness of
  organics in aerosols from biomass burning linked to their black carbon content. *Nature Geoscience*,
  7(9), 647–650. https://doi.org/10.1038/ngeo2220
- Sareen, N., Shapiro, E. L., Schwier, A. N., & McNeill, V. F. (2009). Secondary organic material formed
  by methylglyoxal in aqueous aerosol mimics Part 2: Product identification using AerosolCIMS. Atmospheric Chemistry and Physics Discussions, 9(4), 15567–15594.
  https://doi.org/10.5194/acpd-9-15567-2009
- Seinfeld, J. H., Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeylen, R., Shahgholi, M., et al.
  (2008). Organosulfate formation in biogenic secondary organic aerosol. *Journal of Physical Chemistry A*, *112*(36), 8345–8378. https://doi.org/Doi 10.1021/Jp802310p
- Shapiro, E. L., Szprengiel, J., Sareen, N., Jen, C. N., Giordano, M. R., & McNeill, V. F. (2009). Lightabsorbing secondary organic material formed by glyoxal in aqueous aerosol mimics. *Atmospheric Chemistry and Physics*, 9(7), 2289–2300. https://doi.org/10.5194/acp-9-2289-2009
- Tang, M., Alexander, J. M., Kwon, D., Estillore, A. D., Laskina, O., Young, M. A., et al. (2016). Optical
  and Physicochemical Properties of Brown Carbon Aerosol: Light Scattering, FTIR Extinction
  Spectroscopy, and Hygroscopic Growth. *Journal of Physical Chemistry A*, *120*(24), 4155–4166.
  https://doi.org/10.1021/acs.jpca.6b03425
- Wang, X., Hayeck, N., Brüggemann, M., Yao, L., Chen, H., Zhang, C., et al. (2017). Chemical
  Characteristics of Organic Aerosols in Shanghai: A Study by Ultra-High-Performance Liquid
  Chromatography Coupled with Orbitrap Mass Spectrometry. *Journal of Geophysical Research: Atmospheres*, 1–20. https://doi.org/10.1002/2017JD026930
- Wang, X. K., Rossignol, S., Ma, Y., Yao, L., Wang, M. Y., Chen, J. M., et al. (2016). Molecular
  characterization of atmospheric particulate organosulfates in three megacities at the middle and
  lower reaches of the Yangtze River. *Atmospheric Chemistry and Physics*, 16(4), 2285–2298.
  https://doi.org/10.5194/acp-16-2285-2016
- Wang, Y., Hu, M., Lin, P., Guo, Q., Wu, Z., Li, M., et al. (2017). Molecular Characterization of NitrogenContaining Organic Compounds in Humic-like Substances Emitted from Straw Residue Burning. *Environmental* Science and Technology, 51(11), 5951–5961.
  https://doi.org/10.1021/acs.est.7b00248
- Wang, Y., Hu, M., Guo, S., Wang, Y., Zheng, J., Yang, Y., et al. (2018). The secondary formation of
  organosulfates under interactions between biogenic emissions and anthropogenic pollutants in
  summer in Beijing. *Atmospheric Chemistry and Physics*, 18(14), 10693–10713.
  https://doi.org/10.5194/acp-18-10693-2018
- 447 Yang, M., Howell, S. G., Zhuang, J., & Huebert, B. J. (2009). Attribution of aerosol light absorption to

- black carbon, brown carbon, and dust in China Interpretations of atmospheric measurements
  during EAST-AIRE. *Atmospheric Chemistry and Physics*, 9(6), 2035–2050.
  https://doi.org/10.5194/acp-9-2035-2009
- Yassine, M. M., Harir, M., Dabek-Zlotorzynska, E., & Schmitt-Kopplin, P. (2014). Structural
  characterization of organic aerosol using Fourier transform ion cyclotron resonance mass
  spectrometry: aromaticity equivalent approach. *Rapid Communications in Mass Spectrometry*: *RCM*, 28(22), 2445–2454. https://doi.org/10.1002/rcm.7038
- Zhang, X., Lin, Y. H., Surratt, J. D., & Weber, R. J. (2013). Sources, composition and absorption
  Ångström exponent of light-absorbing organic components in aerosol extracts from the los angeles
  basin. *Environmental Science and Technology*, 47(8), 3685–3693.
  https://doi.org/10.1021/es305047b
- Zhao, R., Lee, A. K. Y., Huang, L., Li, X., Yang, F., & Abbatt, J. P. D. (2015). Photochemical processing
  of aqueous atmospheric brown carbon. *Atmospheric Chemistry and Physics*, *15*(11), 6087–6100.
  https://doi.org/10.5194/acp-15-6087-2015

#### 663 Supporting References:

- Brüggemann, M., Hayeck, N., Bonnineau, C., Pesce, S., Alpert, P. A., Perrier, S., et al. (2017). Interfacial
  photochemistry of biogenic surfactants: A major source of abiotic volatile organic compounds. *Faraday Discussions*, 200, 59–74. https://doi.org/10.1039/c7fd00022g
- 667 Cheng, Y., Fai Ho, K., Jing Wu, W., Hang Ho, S. S., Cheng Lee, S., Huang, Y., et al. (2012). Real-time 668 characterization of particle-bound polycyclic aromatic hydrocarbons at a heavily trafficked 669 roadside site. Aerosol and Air Quality Research, 12(6), 1181-1188. 670 https://doi.org/10.4209/aagr.2011.11.0223
- 671 Fleming, L. T., Lin, P., Laskin, A., Laskin, J., Weltman, R., Edwards, R. D., et al. (2018). Molecular 672 composition of particulate matter emissions from dung and brushwood burning household 673 in 185194, 2461-2480. cookstoves Haryana, India. Atmos. Chem. Phys, 674 https://doi.org/10.5194/acp-18-2461-2018
- 675 Garcia-Alcega, S., Nasir, Z. A., Ferguson, R., Whitby, C., Dumbrell, A. J., Colbeck, I., et al. (2017). 676 Fingerprinting outdoor air environment using microbial volatile organic compounds (MVOCs) -677 Α review. **TrAC** Trends in Analytical Chemistry, 86. 75-83. 678 https://doi.org/10.1016/j.trac.2016.10.010
- 679 Ge, X., Wexler, A. S., & Clegg, S. L. (2011). Atmospheric amines Part I. A review. *Atmospheric* 680 *Environment*, 45(3), 524–546. https://doi.org/10.1016/j.atmosenv.2010.10.012
- Riva, M., Robinson, E. S., Perraudin, E., Donahue, N. M., & Villenave, E. (2015). Photochemical aging
  of secondary organic aerosols generated from the photooxidation of polycyclic aromatic
  hydrocarbons in the gas-phase. *Environmental Science and Technology*, 49(9), 5407–5416.
  https://doi.org/10.1021/acs.est.5b00442
- Teich, M., Van Pinxteren, D., Kecorius, S., Wang, Z., & Herrmann, H. (2016). First Quantification of
  Imidazoles in Ambient Aerosol Particles: Potential Photosensitizers, Brown Carbon Constituents,
  and Hazardous Components. *Environmental Science and Technology*, *50*(3), 1166–1173.
  https://doi.org/10.1021/acs.est.5b05474
- Wang, X., Hayeck, N., Brüggemann, M., Yao, L., Chen, H., Zhang, C., et al. (2017). Chemical
   Characteristics of Organic Aerosols in Shanghai: A Study by Ultra-High-Performance Liquid
   Chromatography Coupled with Orbitrap Mass Spectrometry. *Journal of Geophysical Research:*

- 692 *Atmospheres*, 1–20. https://doi.org/10.1002/2017JD026930
- Yang, H., Mei, L., Wang, P., Genereux, J., Wang, Y., Yi, B., et al. (2017). Photocatalytic degradation of
  norfloxacin on different TiO <sub>2-x</sub> polymorphs under visible light in water. *RSC Adv.*, 7(72), 45721–
  45732. https://doi.org/10.1039/C7RA09022F
- Yassine, M. M., Harir, M., Dabek-Zlotorzynska, E., & Schmitt-Kopplin, P. (2014). Structural
  characterization of organic aerosol using Fourier transform ion cyclotron resonance mass
  spectrometry: aromaticity equivalent approach. *Rapid Communications in Mass Spectrometry* : *RCM*, 28(22), 2445–2454. https://doi.org/10.1002/rcm.7038
- 700

Comula		Start point		Number of Formulas (Fraction %)								
ID	Time		End point	СНО-	CHON-	CHOS-	CHONS-	All in	All in	CHON+	CHN+	All in
								ESI-	CHO+			ESI+
1201M	1 Dec, 2015, 7:30	(30.10°N,	(29.82°N,	261	238	253	243	995	209	202	90	501
	am -12:00 am	115.32°E)	115.73°E)	(26.2%)	(23.9%)	(25.4%)	(24.4%)	(100%)	(41.7%)	(40.3%)	(18.0%)	(100%)
1201 4	1 Dec, 2015, 12:00	(29.82°N,	(29.90°N,	251	259	241	208	959	190	173	87	450
1201A	am -6:00 pm	115.73°E)	116.48°E)	(26.2%)	(27.0%)	(25.1%)	(21.7%)	(100%)	(42.2%)	(38.4%)	(19.3%)	(100%)
1201N	1 Dec, 2015, 6:00	(29.90°N,	(30.37°N,	295	323	253	242	1113	232	246	133	611
1201N	pm -12:00 pm	116.48°E)	116.89°E)	(26.5%)	(29.0%)	(22.7%)	(21.7%)	(100%)	(38.0%)	(40.3%)	(21.8%)	(100%)
1202M	2 Dec, 2015, 6:00	(30.63°N,	(31.10°N,	291	367	258	271	1187	235	269	139	643
	am -12:00 am	117.24°E)	117.76°E)	(24.5%)	(30.9%)	(21.7%)	(22.8%)	(100%)	(36.6%)	(41.8%)	(21.6%)	(100%)
1202N	2 Dec, 2015, 6:00	(31.42°N,	(32.05°N,	289	353	222	244	1108	225	334	136	695
	pm -12:00 pm	118.35°E)	118.69°E)	(26.1%)	(31.9%)	(20.0%)	(22.0%)	(100%)	(32.4%)	(48.1%)	(19.6%)	(100%)
1203A	3 Dec, 2015, 12:00	(32.22°N,	(32.21°N,	201	209	155	130	695	209	191	116	516
	am -6:00 pm	119.21°E)	119.55°E)	(28.9%)	(30.1%)	(22.3%)	(18.7%)	(100%)	(40.5%)	(37.0%)	(22.5%)	(100%)

Table 1. Summary of the sampling location, sampling time, and the number of assigned formulas in both ESI- and ESI+ mass spectra.

ID	Retention Time (min)	Neutral Mass (Da)	Formula	<b>RDBE</b> <sup>a</sup>	Xc <sup>b</sup>	Potential Identity Potential Source/ Precursor		Reference
A'	6.09; 6.67;7.21	122.0366	$C_7H_6O_2$	5	2.50	Benzoic acid	Naphthalene	(Riva et al., 2015)
B'	8.56	139.0268	C <sub>6</sub> H <sub>5</sub> NO <sub>3</sub>	5	2.50	Nitrophenol	Biomass burning	(Mohr et al., 2013)
C'	9.15; 9.42	153.0424	C7H7NO3	5	2.50	Methyl nitrophenol	Biomass burning	(Mohr et al., 2013)
D'	7.73; 8.57	155.0217	C <sub>6</sub> H <sub>5</sub> NO <sub>4</sub>	5	2.33	Nitrocatechols	Biomass burning and vehicle emissions	(Mohr et al., 2013)
E'	6.02; 7.60; 9.15; 9.53	165.0553	C <sub>8</sub> H <sub>7</sub> NO <sub>3</sub>	6	2.60	Nitroacetophenone	Biomass burning	(Lin et al., 2016)
F'	6.10; 6.67; 6.76	166.0264	$C_8H_6O_4$	6	2.50	Phthalic acid	Naphthalene	(Riva et al., 2015)
G'	7.39; 8.32; 8.58; 8.70; 8.82; 9.47	169.0473	C7H7NO4	5	2.33	Methyl nitrocatechols	Biomass burning and diesel exhaust	(Yoshiteru Iinuma et al., 2010)
H'	7.84; 8.10; 9.01; 9.34; 9.55; 10.15	183.0699	C <sub>8</sub> H <sub>9</sub> NO <sub>4</sub>	5	2.33	Dimethoxynitrobenzene	Biomass burning	(Desyaterik et al., 2013)
I'	9.80; 9.96; 10.42	198.0492	$C_7H_6N_2O_5$	6	2.50	Methyl dinitrophenol	Toluene; biomass burning	(Jang & Kamens, 2001; Desyaterik et al., 2013)
J'	8.33; 8.61; 8.79; 8.91	295.0721	C10H17NO7S	3	0	A nitrooxy organosulfate	α-pinene; β-pinene; α-terpinene; Terpinolene	(Gómez-González et al., 2012)

Table 2. Potential identities and sources for the most abundant species in the ESI- mass spectra.

<sup>714</sup> <sup>a</sup>Ring and double bond number. <sup>b</sup>Aromaticity equivalents.

715	Figure Captions
716	Figure 1. The start and end points for the filter sampling along the Yangtze River
717	channel.
718	
719	Figure 2. Mass spectra of detected CHO-, CHOS-, CHON-, and CHONS- compounds,
720	reconstructed from extracted ion chromatograms (UHPLC-Orbitrap MS analysis, ESI-).
721	Note that the abundance of $C_6H_5NO_3$ in the 1202N sample (the most abundant one in
722	all samples) was set arbitrarily to 100%. The sizes of pie charts are proportional to the
723	total abundance of all subgroups in each sample.
724	
725	Figure 3. UHPLC-DAD absorption chromatograms for the 1202N aerosol sample.
726	Panel (a) shows the blank-subtracted absorption chromatograms at two selected
727	wavelengths (i.e., 300 and 350 nm). Panel (b) shows the blank-subtracted absorption
728	chromatograms at 290-480 nm. The absorption peaks are labeled with the most likely
729	formula of the light-absorbing chromophore. Molecular formulas in red, blue, and black
730	colors represent detection in ESI-, ESI+, and both modes, respectively.
731	
732	Figure 4. (a-b) CH <sub>2</sub> -Kendrick diagrams and (c-d) Van Krevelen diagrams for CHO-
733	and CHON- species in the 1202N aerosol sample, respectively. Purple and orange
734	points denote the light-absorbing CHO- and CHON- compounds, respectively. The size
735	of the points corresponds to the abundance of the corresponding ion signal.
736	
737	Figure 5. Relative contribution of the thirteen BrC chromophores to the total light
738	absorption of OAs for the (a) 1201M, (b) 1201A, and (c) 1202N samples. Note that the
739	light-absorption at 290 nm in the 1202N sample (the highest light absorption signal in
740	all samples) was set arbitrarily to 100%.





Figure 2





Figure 4



Figure 5