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Volcanic Record of the Last Geomagnetic Reversal in a Lava Flow Sequence From the Azores

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We present paleomagnetic and geochronological data of twenty-five stacked lava flows from São Miguel Island (Azores) that were emplaced during the last geomagnetic reversal (Matuyama-Brunhes, M-B). The characteristic direction of natural remanent magnetization was isolated through alternating field and/or thermal demagnetization. From bottom to top, the directional changes display a first phase with reverse, transitional and normal directions which is followed by a second phase with transitional and normal directions. Dating by Potassium-Argon (K-Ar) of fresh separated groundmass provide ages of 768 ± 5 ka and 765 ± 5 ka (1σ) for the first and second transitional episodes, respectively. We infer that this sequence recorded the last reversal and reveal the presence of a rebound that occurred soon after the transition. The age of the transition is 5–15 kyr younger than at other volcanic localities, but in agreement with astronomical ages that were recently proposed for the transitional interval. The reversal is characterized by weak field intensity during the transition, a recovery phase and a subsequent decrease during the rebound. The two transitional virtual geomagnetic poles (VGPs) that characterize the transition are found above the South Atlantic and the North Pacific. When compared to other M-B records, no VGP is found above preferred locations.

Keywords: reversal, paleomagnetism, Matuyama Brunhes reversal, Azores Archipelago, magnetite, precursor, paleosecular variation

INTRODUCTION

After about 50 years of paleomagnetic studies, it is clear that the processes governing the geomagnetic reversals can only be constrained from sedimentary records with very high deposition rates and sequences of stacked volcanic lava flows. Although dating techniques used for lava flows provide a direct and precise radiometric measurement, uncertainties often hinder the reconstruction of the exact temporal succession during fast directional changes in the case of sparse volcanic flows. Therefore, only sequences of superimposed lava flows can be used with confidence to reveal pertinent characteristics of the transitional field. In fact, the major difficulty is to find appropriate successions that were erupted during a reversal period. The last geomagnetic reversal

(Matuyama-Brunhes, M-B) has been studied in marine sedimentary cores (e.g., Bassinot et al., 1994; Tauxe et al., 1996; Oda et al., 2000; Channell et al., 2010) and in lava sequences from the Canary Islands, Chile, Tahiti, Maui, and Guadeloupe (e.g., Chauvin et al., 1990; Carlut et al., 2000; Quidelleur et al., 2003; Brown et al., 2004; Coe et al., 2004; Balbas et al., 2018) but detailed records could only be obtained from Tahiti and Maui volcanic sequences.

The mean age of the M-B transition was recently estimated at ca. 773 ka (with a duration of 4.4 ka) from a set of high-resolution paleomagnetic and oxygen isotope data from marine sites in the northern Atlantic Ocean (Channell et al., 2010), and at ca. 772 ka from a combination of ^{10}Be , paleomagnetic and oxygen isotope analyses (Valet et al., 2019). The volcanic records are characterized by more scattered ages. Potassium-argon ages of 786 ± 3 ka were obtained from a set of seven independent lava flows at La Palma, Canary Islands (Quidelleur et al., 2003), while a mean age of 777 ± 4 ka relied on five independent lava flows from the Caribbean (Ricci et al., 2018). The most recent $^{40}\text{Ar}/^{39}\text{Ar}$ age estimates of 783.4 ± 0.6 ka were published from tephra layers within marine sections from the Indian Ocean (Mark et al., 2017) while ages of 786.0 and 786.1 ka were derived from exposed lacustrine sediments in central Italy (Sulmona section, Sagnotti et al., 2016, 2019). In this last study, the duration of the polarity switch was estimated to have lasted only a few decades, but this result has been questioned by Evans and Muxworthy (2018) due to significant remagnetization of the sequence. Finally, recalculated $^{40}\text{Ar}/^{39}\text{Ar}$ ages spanning ~ 795 to ~ 773 ka from Tahiti, La Palma, Chile, Guadeloupe and Maui lava flows were proposed by Singer et al. (2019) and interpreted as evidence of a complex reversal process and attributed to successive geomagnetic events.

The apparent discrepancies between the ages derived from marine records and from terrestrial lava sequences could be caused by multiple factors: (1) non-dipole structures recorded by lava flows but smoothed out in sediments; (2) potential offset between the two signals due to delayed magnetization resulting from locking depth in some sediments; and (3) uncertainties inherent to radiometric dating (Channell et al., 2010; Singer, 2014) and orbital tuning techniques (Valet et al., 2019). It has been proposed (Valet et al., 2012) that geomagnetic reversals could be described by three successive phases: a precursory event, a 180° polarity switch and a rebound. A precursor has been reported in a few sedimentary records (Channell et al., 2010; Valet et al., 2014). The age spreading from transitional lavas may also indicate that some sequences actually provided a record of the precursor rather than the transition itself (Balbas et al., 2018; Singer et al., 2019).

During the last two decades, numerous studies on the Azores archipelago brought new information on the timing of volcanism at both local and regional scales (e.g., Calvert et al., 2006; Costa et al., 2014; Sibrant et al., 2015; Hildenbrand et al., 2018; Marques et al., 2018). Johnson et al. (1998) and Sibrant et al. (2015) identified lava flow sequences in São Miguel island with ages close to the M-B reversal. These results offered a promising potential for a detailed paleomagnetic study with the hope of finding a record of the last reversal.

REGIONAL SETTING

The Azores Archipelago was built on a relatively shallow seafloor that is usually interpreted as resulting from the interaction between a hot or volatile-rich mantle anomaly and the Mid Atlantic Ridge (MAR, e.g., Métrich et al., 2014). On the eastern side of the MAR, recent volcanism (<2 Ma) has developed along the present diffuse Eurasia-Nubia plate boundary, which comprises an ultra-slow oceanic rift (the Terceira Rift) and a series of horsts and grabens constituting the diffuse Azores Triple Junction (Marques et al., 2013). São Miguel is the largest island of the Azores Archipelago and is located on the easternmost part of the Terceira Rift (**Figure 1**). The island consists of six volcanic complexes (from east to west, Nordeste, Furnas, Eastern Waist Zone, Fogo, Western Waist Zone, and Sete Cidades) that result from westward migration of the volcanic construction during the last 900 kyr (**Figure 1**). A 100 m thick succession of 28 lava flows was sampled in the northeastern part of the Nordeste volcanic complex in NE São Miguel (**Figure 1**). A deep creek has incised the volcanic sequence; 21 lava flows were sampled on the northern slope and 7 on the southern slope (for further details see the Methods section).

MATERIALS AND METHODS

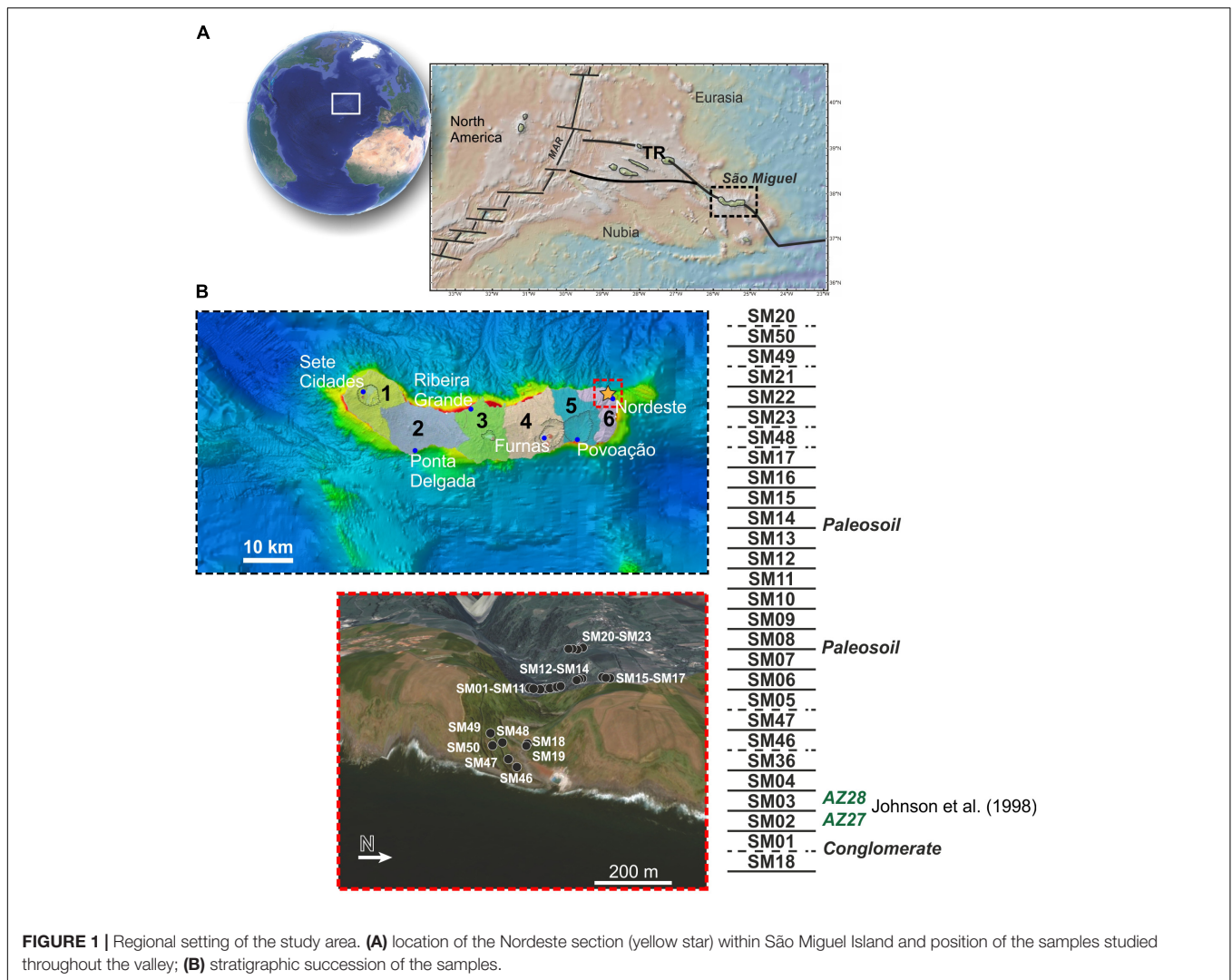
Sampling

The northeastern section of Nordeste comprises three sequences, including a suite of 17 flows in direct stratigraphic relationship (SM01 to SM17), and two smaller sub-sections located down the creek near the shore (SM18 and SM19) and upper to the northwest (flows SM20 to SM23; cf. **Figure 1**). The southern section is characterized by 5 well-exposed lava flows (SM46 to SM50) accessible through the main road going down to the lowermost part of the creek near sea level. The stratigraphic succession of the lava flows from these two parts of the valley was constrained by field observation combined with paleomagnetic and geochronological data (**Figure 1**).

A total of 325 cores (typically 1 inch in diameter) were collected with an average of twelve cores per site (**Table 1**). All samples were magnetically and sun oriented. The mean magnetic declination of -9° (calculated from 261 oriented cores) is indistinguishable from the expected International Geomagnetic Reference Field (-8°). In addition, four fresh lava flows, from bottom to top, were sampled for K-Ar dating. The freshest portion of the lava flows was systematically collected by means of chisels, sledge hammer and rock pick hammer. Samples with obvious traces of weathering were discarded.

Paleomagnetic Treatment

Alternating fields (af) and thermal stepwise demagnetization techniques were carried out in the magnetic shielded room of the Institut de Physique du Globe de Paris (IPGP-Paris). Natural Remanent Magnetization (NRM) was measured using Agico JR-6 spinner magnetometers. Twelve af demagnetization steps were systematically performed at 05, 10, 15, 20, 25, 30,



35, 40, 50, 70, and 90 mT. Thermal demagnetization was carried out at 14 temperature steps up to 580–595°C. The demagnetization diagrams were interpreted using the Paleomac software (Cogne, 2003).

Measurements of low field thermomagnetic susceptibility (κ) were conducted on 1 cm³ sample powders from each flow using an Agico KLY-3 instrument equipped with a CS-3 furnace. Heating-cooling runs were performed from 20 to 650°C to detect mineralogical changes. A first batch of samples was heated in air and a second batch under Argon atmosphere. Hysteresis loops were conducted on one representative sample of each flow using an Alternating Gradient Force Magnetometer (AGFM, Princeton Measurements Corporation) in a 1 T maximum field.

After af demagnetization until 120 mT, two to four samples from each flow were remagnetized by an ARM (Anhysteretic Remanent Magnetization) using an applied magnetic field at 50 μ T. The samples were then af stepwise demagnetized at six steps 10, 20, 30, 40, 50, and 70 mT, the same steps were previously used for NRM demagnetization. Then the demagnetization of the ARM was compared to that of the NRM

for the same AF steps. The steps were chosen so it corresponds to interval representative of the primary magnetization. The NRM/ARM values were calculated using the average from the representative demagnetization steps (examples are given in **Supplementary Figure S1**).

Geochronology

Careful microscopic examination of thin-sections allowed checking the structure, texture and quality of the samples. As most of them are porphyritic basalts, the fresh groundmass was systematically selected. After crushing and sieving at an appropriate size (typically 125 – 250 μ m), extraction of the fresh groundmass in a narrow density interval (typically 2.95 – 3.0) was achieved with heavy liquids and magnetic separation. This procedure is essential to eliminate all phenocrysts (which have mostly formed in a magma chamber and may carry significant unsuitable excess-argon), and so get a meaningful eruption age.

Potassium-Argon analyses were performed at the Geosciences Paris-Sud (GEOPS) Geochronology laboratory using the unspiked Cassinon-Gillot experimental procedure, described

TABLE 1 | Paleomagnetic results.

Samples	Age [ka]	n/N	Dec.	Inc	k	α 95	VGP long.	VGP lat.	Polarity	Mean (NRM/ARM)
SM19		8/8	206.1	-50.9	289.9	3.3	253.2	-67.7	R	0.68
SM18	789 ± 11	8/10	156.2	-53.9	121.5	5.0	62.1	-70.5	R	1.71
SM01		8/8	172.9	-54.1	753.8	2.2	36.3	-83.3	R	1.05
SM02	776 ± 6*	7/9	168.2	-48.5	283.4	3.6	27.8	-77.1	R	2.16
SM03/04		16/16	181.3	-37.0	155.7	3.0	330.8	-72.8	R	3.45
SM36		8/10	188.9	-34.3	298.1	3.2	310.2	-69.5	R	5.69
SM46		11/12	182.0	-47.8	165.4	3.6	323.9	-80.9	R	1.46
SM47		8/8	187.2	-40.1	356.6	2.9	310.4	-73.8	R	2.33
SM05/06/07		24/28	181.4	-42.1	175.6	2.2	329.6	-76.5	R	2.37
SM08		8/8	206.5	-58.7	266.4	3.4	232.3	-69.3	R	0.85
SM09		8/11	197.3	-59.7	100.8	5.5	228.0	-76.3	R	0.79
SM10		9/12	207.6	-56.3	119.9	4.7	239.0	-68.1	R	0.79
SM11		5/10	164.7	8.4	36.1	12.9	357.0	-45.6	T	0.09
SM12	768 ± 5	5/10	350.3	-37.9	142.1	6.4	165.3	30.2	T	0.11
SM13		7/8	3.2	14.5	52.7	8.4	148.5	59.4	N	0.14
SM14		10/10	35.1	57.0	177.2	3.6	54.4	62.4	N	3.80
SM15	762 ± 11	7/9	20.3	53.7	178.0	4.5	71.2	73.3	N	1.73
SM16		9/11	29.4	57.7	192.5	3.7	54.4	67.0	N	3.55
SM17		9/11	22.0	61.8	110.5	4.9	40.9	72.5	N	1.00
SM48		8/8	358.2	58.8	395.0	2.8	302.3	88.0	N	0.24
SM23	765 ± 5	8/10	331.4	-14.5	82.7	6.1	191.6	37.5	T	0.25
SM22/21		17/18	32.1	20.0	164.8	2.8	100.2	50.2	N	0.86
SM49		8/8	28.1	32.7	257.3	3.5	96.3	58.3	N	0.29
SM50	763 ± 11	8/8	31.8	27.4	255.6	3.5	95.9	53.5	N	0.42
SM20		8/8	15.4	47.1	209.5	3.8	96.9	74.0	N	0.96

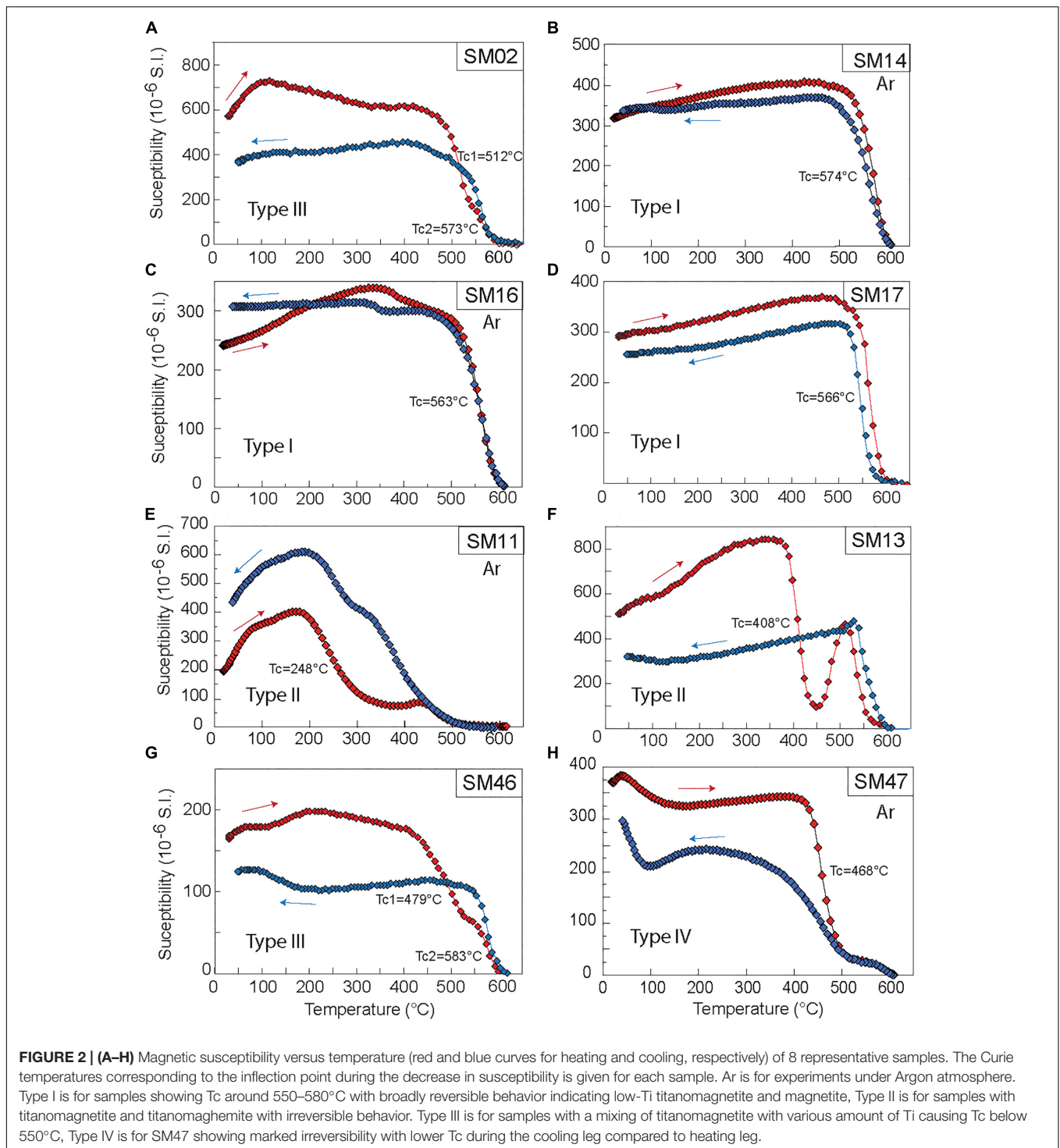
Column headings: sample name; ages $\pm 1\sigma$ uncertainty in ka, all ages are from this study, except SM02 (stars) dated by Johnson et al., 1998; number of samples used/total number of samples measured; Declination; Inclination; Fisher's precision parameter; radius of the 95 per cent confidence cone from Fisher (1953); VGP longitude and latitude, polarity of lava flow, mean value obtained for the NRM/ARM.

in detail in Ricci et al. (2017) and summarized below. The K content was measured by flame absorption spectrometry, and was compared with reference values for the MDO-G (Gillot et al., 1992) and BCR-2 (Rackzec et al., 2001) standards. For each sample, at least two independent measurements of Argon from distinct aliquots of the same preparation have been carried out with a multi-collector 180° sector mass spectrometer following the approach described in Gillot and Cornette (1986). Standard HD-B1, with a recommended age of 24.21 Ma (Hess and Lippolt, 1994), was used for calibration of our mass spectrometer. We note that an age value of 24.18 Ma was proposed more recently for this standard (Schwarz and Trieloff, 2007), which has no significant incidence on age determination in our case, as the relative difference between the two values is of about one per mill, i.e., below 1000 years for the ages obtained in this study. Atmospheric contamination was determined by the comparison of the $^{40}\text{Ar}/^{36}\text{Ar}$ ratios of the sample with an atmospheric dose. We used the decay constants and isotopic ratios of Steiger and Jäger (1977). The uncertainty on each age determination was obtained from the quadratic sum of all three independent sources of uncertainty involved in the calculation: (1) relative uncertainty on K-content determination, (2) relative uncertainty on calibration of the ^{40}Ar signal, and (3) relative uncertainty on the correction of the atmospheric contamination (full description in Hildenbrand et al., 2018).

RESULTS

Rock Magnetic Properties

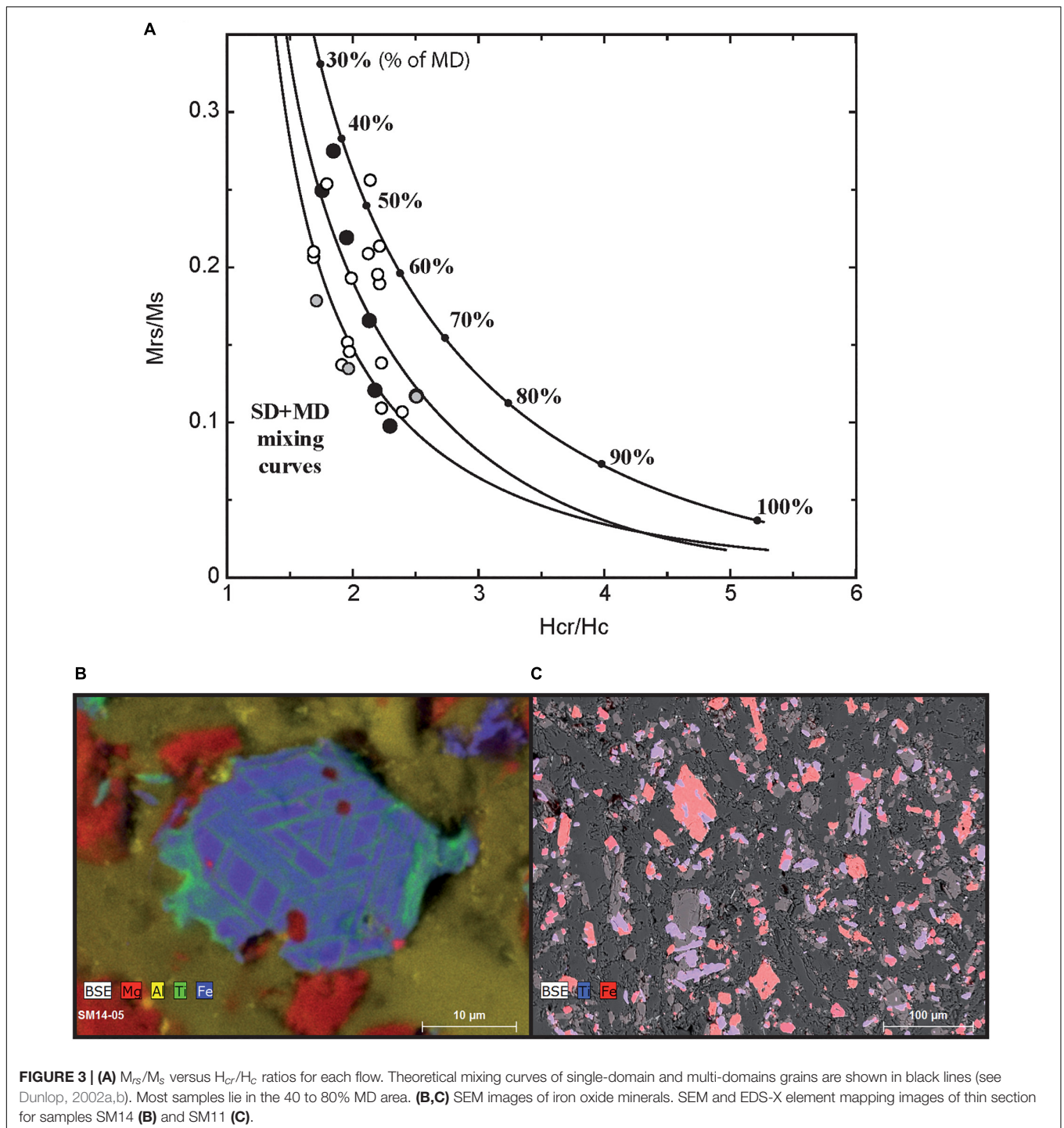
Thermomagnetic experiments were conducted on two samples from each flow, one under air and one under Ar atmosphere. Most samples show a spectrum of Curie temperatures (T_c) in the range 550–580°C as determined using the inflection point on the decreasing branch of the susceptibility curve (e.g., Petrovský and Kapička, 2006), and a broadly reversible behavior in terms of shape, with little difference between the Ar batch compared to air. This is indicative of Ti-poor titanomagnetite phases and/or of pure magnetite (typical examples are shown in **Figure 2**, Type I samples). A more complex mineralogy occurs for SM11 and SM13 (Type II in **Figure 2**) where a decrease in susceptibility around 250°C (SM11) and 410°C (SM13) is followed by a peak around 450°C (SM11) or 500°C (SM13) indicating that unstable cation deficient titanomagnetite exsolved into Ti-poor and Ti-rich phases upon heating in air (Readman and O'Reilly, 1972). The reaction also occurs under Ar atmosphere but is less pronounced (**Figure 2E** compared to **Figure 2F**). SM09, SM10, SM12, SM48, SM02, and SM46 are characterized by a decrease in susceptibility in the 350–550°C range (**Figure 2**, Type III) followed by a second decrease around 570–580°C. For these samples mixtures of titanomagnetite with varying amount of Ti is believed to be responsible for this



behavior. Finally SM47 (Figure 2, Type IV) show marked irreversibility, both in Air and under Ar atmosphere, with the destruction of low Ti titanomagnetite and creation of phases with lower T_c .

Hysteresis loops are rather narrow and rapid saturation of isothermal remanent magnetization below 0.2 T reveal low coercivity ferrimagnetic minerals. Magnetic hysteresis

parameters are calculated after paramagnetic slope corrections. The ratio of saturation remanence to saturation magnetization (M_{rs}/M_s) is plot against the ratio of remanent coercive force to coercive force (H_{cr}/H_c) (as proposed by Day et al., 1977 and further in Dunlop, 2002a,b). Results are indicative of a variable mixture of single-domain (SD) and multi-domains (MD) magnetite particles (Figure 3A).



Scanning electron microscopy (SEM) observations and EDS-X analyses were performed on SM05, SM10, SM11, SM12, SM13, SM14, SM33, and SM34. Images revealed a mixture of dendritic and euhedral (sometimes elongated) Ti-Fe oxide grains of small to moderate size (typically less than 20 μ m) in a silicate matrix. Oxy-exsolution features forming alternating lamellae of magnetite and ilmenite are observed (as in SM14, **Figure 3B**), but in most cases the Ti-Fe phases are well separated between

homogeneous iron oxide phases with low to moderate Ti (Ti/Fe ratio in the range 0.05–0.3 depending on samples) and ilmenite (**Figures 3B,C**).

Paleomagnetic Directions and Relative Paleointensity

The characteristic component of remanent magnetization (ChRM) of 275 samples was determined after both alternating

field (af) and thermal (Th) stepwise demagnetization. Examples of demagnetization diagrams in **Figure 4** show that the ChRM was successfully isolated between 240 and 580°C (or 15–80 mT). In some cases, a secondary component was removed after applying a 15 mT af peak or after heating at 240°C. Most samples lost their magnetization between 540 and 580°C, suggesting low-Ti titanomagnetite and/or magnetite as the main carrier of the ChRM. A median destructive field of about 35 mT is consistent with the presence of pseudo-single domain magnetic grains (Özdemir and Dunlop, 1997).

The mean flow directions were obtained using Fisher statistics (Fisher, 1953) and are summarized in **Table 1**. Data quality (Cromwell et al., 2018) can be assessed from the Fisher precision parameter, which is above 50 for all flows (**Table 1**), except flow SM11 with a value of 36.1. Two lava flows from the bottom of the section (SM02 and SM03) have previously been analyzed by Johnson et al. (1998). Our results are in close agreement with these previous measurements (**Table 1**). The declinations and inclinations of SM03 and SM04, SM05 and SM06, and those from SM21 and SM22 differ by less than 2.5°, and can thus be considered as one single unit that was emplaced over a very short time period (e.g., Carlut et al., 2000). This approach yields twenty-five independent directional units (**Table 1**). Twelve flows display a reverse polarity and ten a normal polarity (**Table 1** and **Figure 5**). The directions of the three lava flows SM11, SM12 and SM23 (with a reversal angle that deviates by more than 30° from full polarity and VGP latitudes lower than 50°) are referred as transitional directions (**Table 1**).

Due to the large number of samples required for the magnetic measurements as well as for dating the remaining collection of samples was not large enough to carry out detailed experiments of absolute paleointensity. Instead, we attempted relative paleointensity experiments (see e.g., Levi and Banerjee, 1976; Tauxe et al., 1995) assuming that the very large drop of field intensity that prevails during a reversal has a much larger consequence on the evolution of magnetization intensity between flows than possible changes induced by mineralogical changes. In order to compensate for changes in magnetic concentration, we used the NRM versus ARM values average over several demagnetization step (see the Paleomagnetic treatment section) our best approximation for relative paleointensity (examples are shown in **Supplementary Figure S1**). The lack of correlation between NRM/ARM and the hysteresis parameters as well as with the coercive field $H_{1/2}$ further indicates the absence of direct control of the slopes by rock magnetic changes (see **Supplementary Figure S2**). Note also that the amplitude of the ARM changes is low compared to the NRM variations. For all these reasons, we are thus confident that NRM/ARM can be interpreted as a first-order in terms of evolution of the field intensity changes that prevailed during this period.

The results of relative paleointensity (**Figure 5D**) show the dominance of very low field during the emplacement of flows SM11, SM12, and SM13 and again at the end of the section for flows SM48, SM23, SM49, and SM50. In between these two periods, a recovery is indicated by several flows (SM14, SM15, and SM16) with higher NRM/ARM values.

K-Ar Dating

For each sample, potassium (K) and argon (Ar) were measured at least twice, and the mean age was obtained when independent measurements were consistent at the 1 σ level. Six independent measurements on the HDB-1 standard were performed during the present study. Their mean age of 24.19 Ma \pm 0.12 Ma (Standard Deviation 1 σ) is in full agreement with the recommended value of 24.21 Ma (Hess and Lippolt, 1994). These results rule out the possibility of significant bias linked to the calibration of the mass spectrometer.

The geochronological results are summarized in **Table 2**. Five new K-Ar ages were obtained between the bottom and top of the section. For samples SM12 and SM23, four and seven independent measurements of K and Ar, respectively, were performed to improve the statistical error on the mean age. Sample SM18 located at the bottom of the section yielded a K-Ar age of 789 \pm 11 ka, which is in full agreement with the overlying flow $^{40}\text{Ar}/^{39}\text{Ar}$ age of 776 \pm 6 ka (sample AZ28, Johnson et al., 1998) recalculated at 789 \pm 6 ka (relative to HD-B1 at 24.21 Ma, using the intercalibration of Kuiper et al. (2008) for Taylor Creek sanidine compared to MMhb-1, and that of Schwarz and Trierloff (2007) for the intercalibration between MMhb-1 and HD-B1 standards). The transitional lava flow SM12 has an age of 768 \pm 5 ka. An age of 762 \pm 11 ka was obtained for SM15, a normal polarity lava flow that lies four flows above SM12. The transitional flow SM23 has a similar age of 765 \pm 5 ka while the normal polarity sample SM50 at the top of the section has an age of 763 \pm 11 ka.

Based on recent compilations of sedimentary records, the Matuyama-Brunhes reversal has been dated around 773 ka (e.g., Valet et al., 2019). It is characterized by rather complex field behavior (Channell, 2017) and a period of low intensity for about 12 kyr (e.g., Channell et al., 2004). The most detailed sedimentary records show the occurrence of a precursor event about 15 kyr before the actual reversal at around 790 ka (Channell et al., 2008). This precursor was likely also recorded in lavas from: - Chile, with mean $^{40}\text{Ar}/^{39}\text{Ar}$ ages for transitional flows at 792 \pm 3 ka (Brown et al., 2004); - Tahiti, Punaruu North with a collection of $^{40}\text{Ar}/^{39}\text{Ar}$ ages from transitional and normal flows older than 785 ka (Singer et al., 2005, 2019; Balbas et al., 2018) - La Palma, with mean K-Ar ages bracketing the reverse to normal transition at 786 \pm 3 ka (Quidelleur et al., 2003), and between 780.3 \pm 10.3 and 803.3 \pm 9.3 ka for a group of 3 transitional flows at Baranco de Los Tilos ($^{40}\text{Ar}/^{39}\text{Ar}$ ages, Singer et al. (2002), see also younger estimation at around 773 \pm 3 ka in Singer et al., 2019) - Guadeloupe, with mean K-Ar ages at 781 \pm 18 ka from a section of 3 lavas that record a succession of transitional-reverse-normal directions (Carlut et al., 2000), and $^{40}\text{Ar}/^{39}\text{Ar}$ ages varying from 779 \pm 6 ka (Singer, 2014) to 784.4 \pm 4.3 ka (Brown et al., 2013).

In fact, the actual transition would have only been recorded by transitional flows from Punaruu South and Maui as suggested by $^{40}\text{Ar}/^{39}\text{Ar}$ ages at, respectively, 776 \pm 3 ka and 773 \pm 2 ka; Singer et al., 2019). A single transitional flow from Martinique (Caribbean islands) with a K-Ar age of

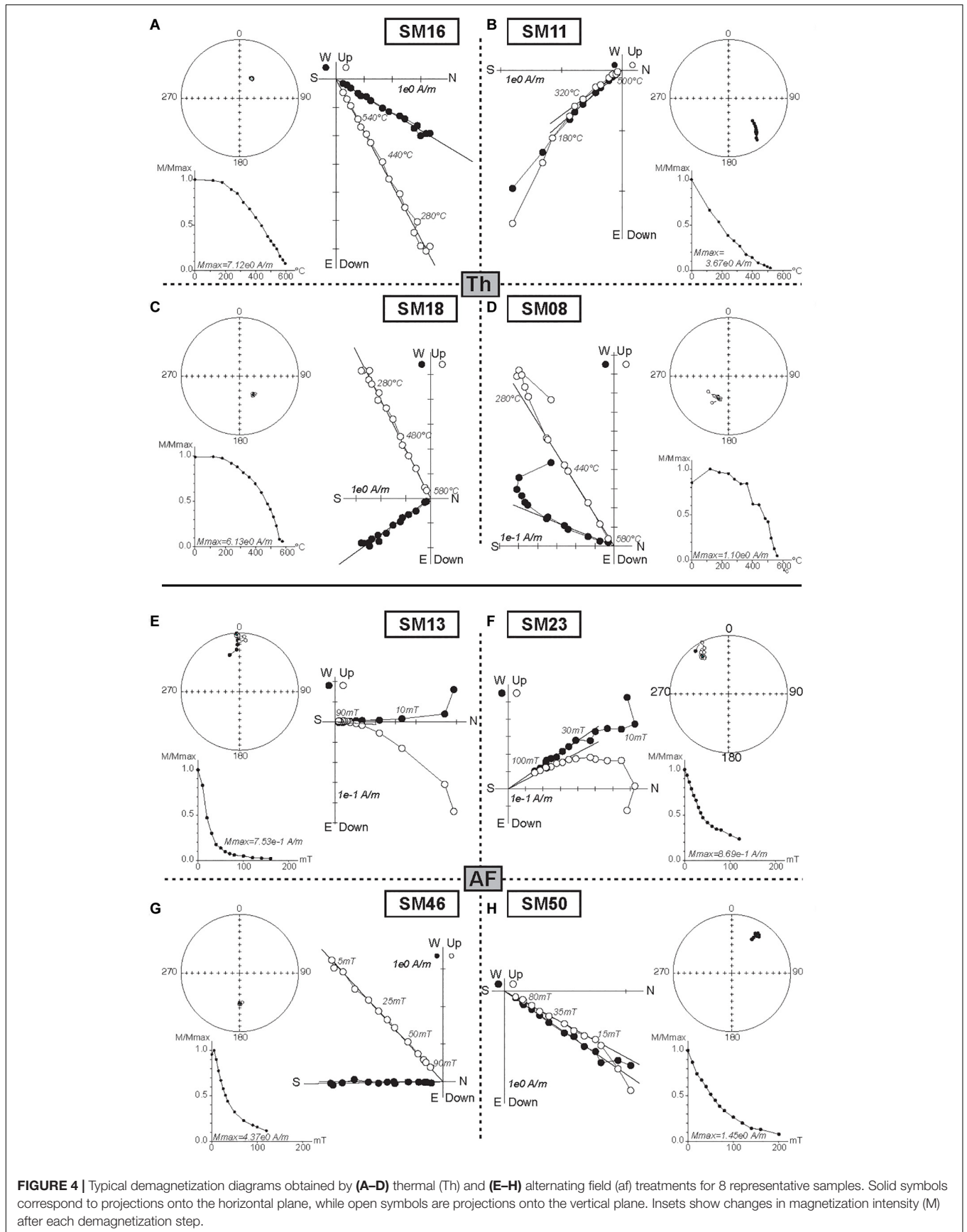


FIGURE 4 | Typical demagnetization diagrams obtained by (A–D) thermal (Th) and (E–H) alternating field (af) treatments for 8 representative samples. Solid symbols correspond to projections onto the horizontal plane, while open symbols are projections onto the vertical plane. Insets show changes in magnetization intensity (M) after each demagnetization step.

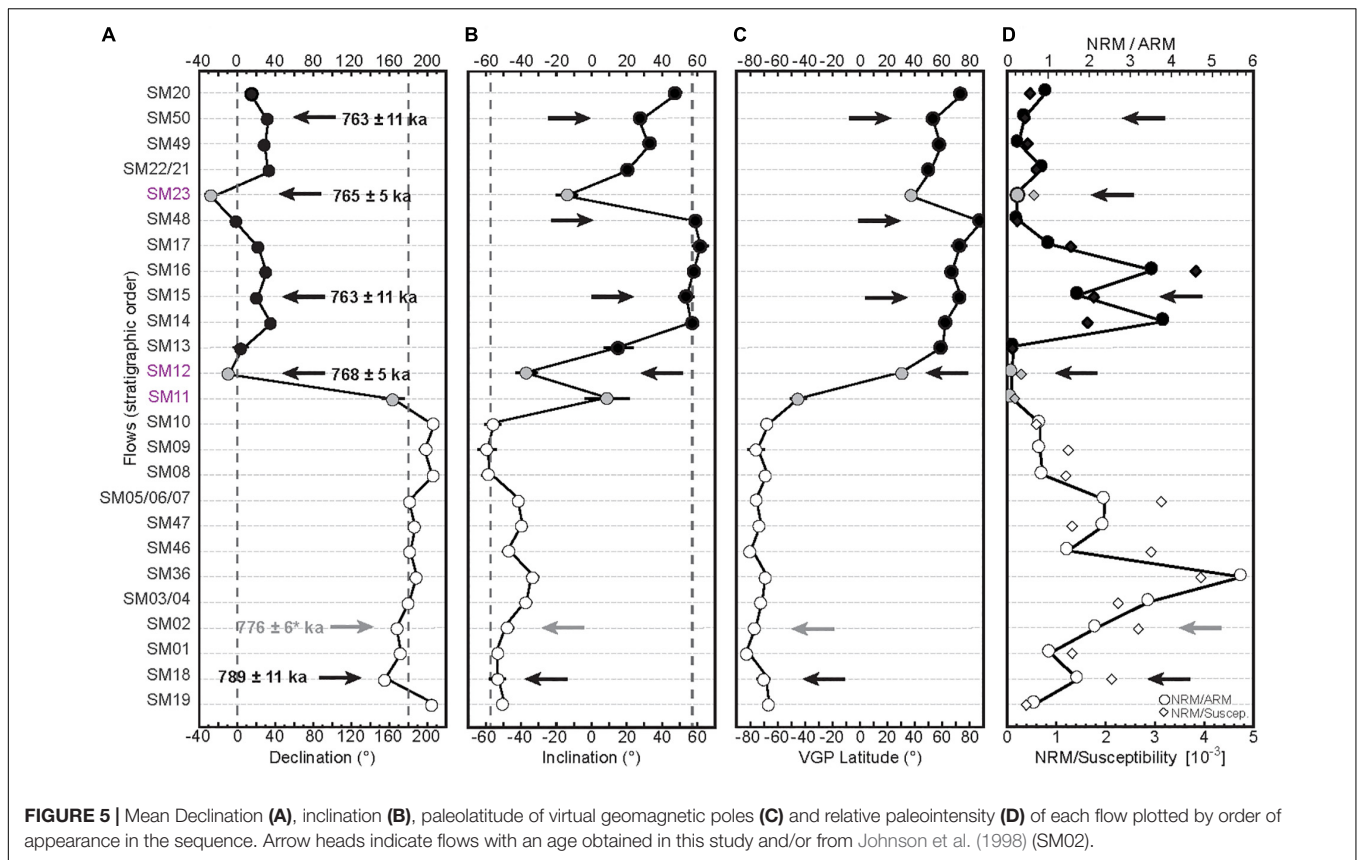


TABLE 2 | Dating results.

Sample	K (%)	$^{40}\text{Ar}^* \times 10^{12}$ (atom/g)	$^{40}\text{Ar}^*$ (%)	Age $\pm 1\sigma$ (ka)	Mean age $\pm 1\sigma$ (ka)
SM18	1.538	1.273	26.1	793 \pm 12	789 \pm 11
		1.262	31.3	786 \pm 11	
SM12	1.871	1.509	33.5	772 \pm 11	768 \pm 5
		1.505	35.7	770 \pm 11	
		1.450	32.8	766 \pm 11	
		1.489	33.7	762 \pm 11	
		1.511	35.8	773 \pm 11	
SM15	1.444	1.450	36.6	766 \pm 11	762 \pm 11
		1.151	26.1	763 \pm 11	
		1.147	29.4	760 \pm 11	
		1.174	31.7	769 \pm 11	
SM23	1.574	1.253	39.9	762 \pm 11	765 \pm 5
		1.260	34.8	766 \pm 11	
		1.642	36.3	769 \pm 11	
		1.259	31.7	766 \pm 11	
		1.258	36.1	765 \pm 11	
		1.265	26.3	769 \pm 11	
SM50	1.675	1.334	19.9	762 \pm 11	763 \pm 11
		1.337	26.6	764 \pm 11	

Column headings: Sample name; Potassium content (K) in percent; radiogenic ^{40}Ar ($^{40}\text{Ar}^*$) content in atoms per gram ($\times 10^{12}$ at.g $^{-1}$) and in percent; Ages $\pm 1\sigma$ uncertainty (ka); and weighted (using the radiogenic content) mean ages $\pm 1\sigma$ uncertainty.

770 \pm 11 ka (Tantý et al., 2015) may be added to this set of M-B transitional lavas. These ages for the actual transition are also consistent with the results obtained from the North Atlantic

Ocean sediments (Channell, 2017) and agree with the U-Pb and astrochronological age of 770.2 \pm 7.0 ka (2 σ) obtained for the Chiba section (Suganuma et al., 2015).

A rebound following the reversal has been predicted on statistical hypothesis (Valet et al., 2012). It has been observed in sediments from the Chiba composite section (Okada et al., 2017). Assuming that the precursor and rebound would both be included in the reversal process, the whole M-B event could have lasted some 30 kyr or more (e.g., Valet et al., 2014; Balbas et al., 2018).

The 25 independent flows from the São Miguel Nordeste section with ages varying from 789 ± 11 to 763 ± 11 ka (1σ) bracket the M-B boundary and provide an opportunity to constrain further the instabilities associated with the reversal. The agreement between the age of the transitionally magnetized flow SM12 (768 ± 5 ka, 1σ) and the recently proposed astrochronological age at 772.4 ± 6.6 ka (1σ , Valet et al., 2019) indicates that these lavas recorded the M-B transition itself.

DISCUSSION

Characteristics of the Transitional Field

The transitional VGPs from the Nordeste section switch from reverse to normal polarity across the South Atlantic Ocean, the North Pacific Ocean, and eastern Russia. It has been suggested that VGPs clusters over Australia, South America, northeast Asia and northeastern Canada (Hoffman, 2000) during the M-B transition would indicate a control of the field by the lower mantle. The expression of non-dipole field components and the fragmentary temporal succession of the lavas hamper a direct comparison between sites (e.g., Mochizuki et al., 2011). Keeping this in mind, we compile below the available transitional VGPs with ages compatible with the transition (i.e., clustering around 770 ka).

The small set that emerges from this work is shown in **Figure 6** and incorporates VGPs from SM11, SM12, and SM23

(this study), MT06 from Martinique (Tanty et al., 2015), flows 42 to 59 from Maui (Coe et al., 2004) and R1T to T1T from Tahiti (Chauvin et al., 1990). We compared the volcanic VGPs with those derived from the high-resolution record of the last reversal in the North Atlantic Ocean (Channell, 2017). However, we restrained the analysis to those VGPs that were derived from measurements of discrete samples at site U1306. Indeed, the U-channel measurements that were systematically performed at the other sites may have artificially generated erroneous transitional directions (Philippe et al., 2018). This selection is evidently far from being exhaustive, but is a conservative first step for comparison with the best resolution sedimentary records. The four groups that can be discerned from the sediment directions are shown in gray in **Figure 6**. They are located over the south and central Atlantic Ocean, eastern Asia and the central Pacific Ocean. The VGPs from SM12, SM23, and MT06 are located in the northern Pacific Ocean, which may indicate that a “stationary feature” was recorded in the lava flows from North Atlantic Ocean sites. The occurrence of such “stationary features” of the geomagnetic field has been suggested by Hoffman and Singer (2004) who observed geographically “preferred” VGP behavior apparent from regionally partitioned paleomagnetic data. Also, Channell et al. (2004) proposed that North Atlantic IODP sites show clusters of VGPs and that polarity transition fields have characteristics similar to the modern non-axial-dipole (NAD) field. In this view large regions could be dominated by mantle-controlled flux and show cluster of VGPs. However, in our North Atlantic set VGPs from SM12, SM23, and MT06 are not mirrored in the sedimentary record from site U1306. At the onset of the reversal, the SM11 (Azores) and 42 (Maui) flows both have a VGP in the South Atlantic Ocean falling within a large VGP cluster in the site U1306 record, but on the other hand the VGP cluster over New Zealand and north America in the Maui data does not correlate with other sites (**Figure 6**).

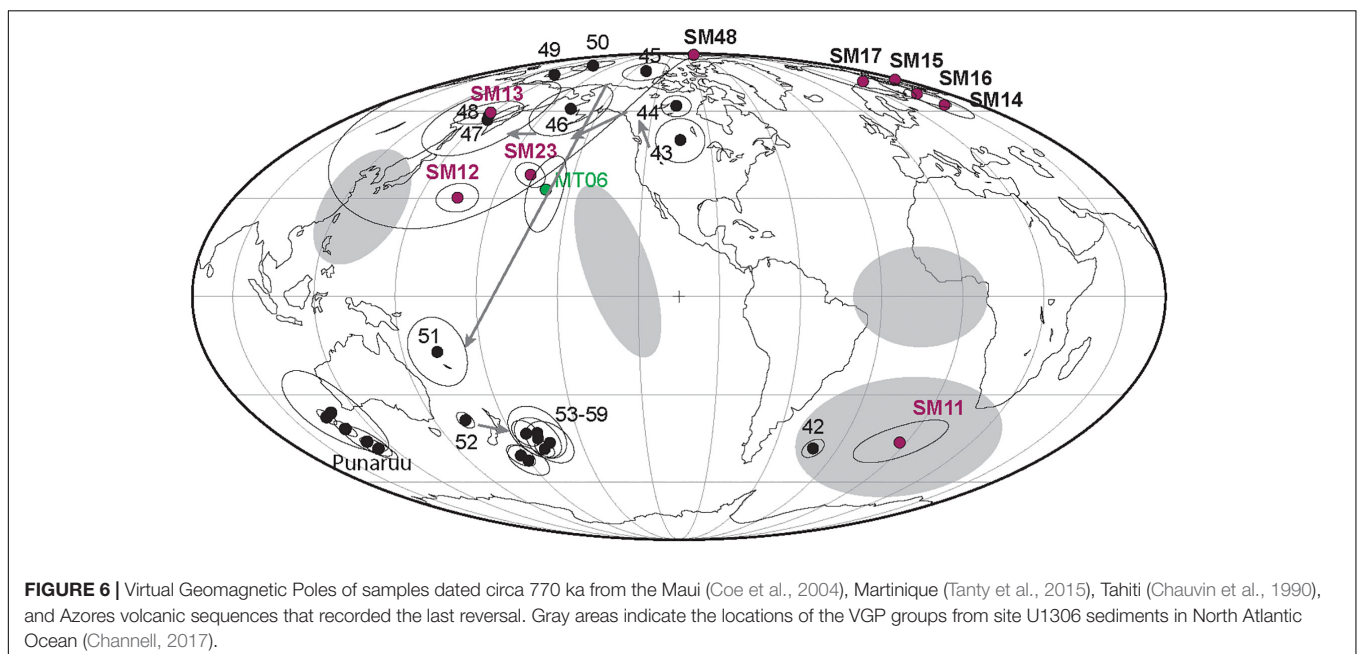


FIGURE 6 | Virtual Geomagnetic Poles of samples dated circa 770 ka from the Maui (Coe et al., 2004), Martinique (Tanty et al., 2015), Tahiti (Chauvin et al., 1990), and Azores volcanic sequences that recorded the last reversal. Gray areas indicate the locations of the VGP groups from site U1306 sediments in North Atlantic Ocean (Channell, 2017).

Summarizing, this small dataset provides at least 7 groups of VGPs without any real statistical significance, which are therefore simply representative of a complex transitional field geometry.

Is the M-B Reversal a Three Phases Process?

The reversal process has been depicted in three phases that incorporate a precursor, the reversal itself, and a rebound (e.g., Valet et al., 2012). To our knowledge, no volcanic record of the rebound has been firmly reported so far. However, a rebound may be present in the Maui record where the VGPs from flows 51–59 depict a loop over the South Pacific Ocean after the reverse-normal polarity transition documented by the VGPs from flows 42 to 44 (Figure 5 and data from Coe et al., 2004). The transitional flows SM11 and SM12 from this study are followed by six normal polarity lava flows with four northern-latitude VGPs and strong relative paleointensity which testify the restoration of a dipolar field after completion of the transition. Overlying these flows, the transitional VGP in the Northern Pacific Ocean and the low field intensity of flow SM23 may be a signature of the rebound that followed the reversal.

The twelve reverse polarity flows in the lower part of the São Miguel section are characterized by a mean inclination of -49.5° ($\alpha_{95} = 7.1^\circ$) that is $\sim 8^\circ$ lower than the Geocentric Axial Dipole (GAD) inclination at the site (57.2°) while low relative paleointensities are recorded by the lowermost three flows (SM19, SM18 and SM01). The ten normal polarity flows following the transition display a mean inclination of 46.8° ($\alpha_{95} = 11.8^\circ$), which is also different from the GAD expected inclination. We thus may have sampled a period of geomagnetic instability that would extend to flows preceding and following the transition, and therefore associated with lower dipolar field and enhanced contribution of non-dipolar components. However, only confirmation by absolute paleointensities will allow to determine if these flows have recorded a peculiar pre- and post-reversal field.

After integrating the São Miguel record with other volcanic and high-resolution sediment records of the last reversal, we thus favor a scenario in which three successive phases would have punctuated the period of the reversing field during the M-B transition. Owing to their age the São Miguel and Maui records could be the only two volcanic sequences that document the actual reverse-normal transition. The ages of the transitional directions from the other volcanic records of the last reversal are in fact more consistent with the precursor that occurred around 780–790 ka. In these two sections the occurrence of transitional

VGPs after the actual reversal might represent a partial signature of a rebound that occurred soon after the transition. These results highlight the complex character of the M-B reversal, and bring additional insights to our understanding of the magnetic field during a reversal.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/**Supplementary Material**, further inquiries can be directed to the corresponding author/s.

AUTHOR CONTRIBUTIONS

JR, JC, FM, and AH designed the project and contributed to the sampling. J-PV managed the ERC project. JR performed the magnetic experiments. JR, JC, and J-PV analyzed the results. JR and AH performed the geochronological analyses. All authors contributed to writing and editing of the manuscript.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2020.00165/full#supplementary-material>

FIGURE S1 | NRM/ARM determined at different demagnetization steps for 8 samples.

FIGURE S2 | $H_{1/2}$ and M_{rs}/M_s versus NRM/ARM.

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Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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