



**HAL**  
open science

## Seamount effect on circulation and distribution of ocean taxa in the vicinity of La Pérouse, a shallow seamount in the southwestern Indian Ocean

Francis Marsac, Pavanee Annasawmy, Margaux Noyon, Hervé Demarcq, Marc Soria, Njaratiana Rabearisoa, Pascal Bach, Yves Cherel, Jacques Grelet, Evgeny Romanov

### ► To cite this version:

Francis Marsac, Pavanee Annasawmy, Margaux Noyon, Hervé Demarcq, Marc Soria, et al.. Seamount effect on circulation and distribution of ocean taxa in the vicinity of La Pérouse, a shallow seamount in the southwestern Indian Ocean. *Deep Sea Research Part II: Topical Studies in Oceanography*, 2020, 176, pp.104806. 10.1016/j.dsr2.2020.104806 . hal-02910880

**HAL Id: hal-02910880**

**<https://hal.science/hal-02910880>**

Submitted on 22 Aug 2022

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License

1 **Seamount effect on circulation and distribution of ocean taxa in the vicinity of La**  
2 **Pérouse, a shallow seamount in the southwestern Indian Ocean**

3

4 Francis Marsac\*<sup>ab</sup>, Pavanee Annasawmy<sup>ab</sup>, Margaux Noyon<sup>c</sup>, Hervé Demarcq<sup>ab</sup>, Marc Soria<sup>ab</sup>,  
5 Njaratiana Rabearisoa<sup>ab</sup>, Pascal Bach<sup>bd</sup>, Yves Chere<sup>f</sup>, Jacques Grelet<sup>g</sup>, Evgeny Romanov<sup>e</sup>

6

7 <sup>a</sup> *MARBEC, Univ Montpellier, CNRS, Ifremer, IRD, Sète, France*

8 <sup>b</sup> *Institut de Recherche pour le Développement (IRD), Sète, France*

9 <sup>c</sup> *UK-SA NRF/DST Bilateral Research Chair: Ocean Sciences & Marine Food Security, Nelson*  
10 *Mandela University, Port Elizabeth, South Africa*

11 <sup>d</sup> *MARBEC, Univ Montpellier, CNRS, Ifremer, IRD, Victoria, Seychelles*

12 <sup>e</sup> *CAP RUN – CITEB, Le Port, Île de la Réunion, France*

13 <sup>f</sup> *Centre d'Etudes Biologiques de Chizé (CEBC), UMR 7372 CNRS-La Rochelle Université, Villiers-*  
14 *en-Bois, France*

15 <sup>g</sup> *IMAGO, IRD Bretagne, Technopole de Brest Iroise, Plouzané, France*

16

17 \* Corresponding author at: MARBEC, Station Ifremer, Avenue Jean Monnet, 34203 Sète Cedex,  
18 France

19

20 *Email: francis.marsac@ird.fr*

21

22 **Abstract**

23 The La Pérouse seamount (60 m depth) has so far been poorly studied despite it being a short distance  
24 (160 km) from Réunion Island. As part of the MADRidge project, a multidisciplinary cruise was  
25 conducted to evaluate the effect of this shallow seamount on the local hydrology and ecology. Current  
26 measurements, temperature and chlorophyll-*a* profiles, and mesozooplankton and micronekton  
27 samples were collected between the summit and 35 km away. Micronekton data were supplemented  
28 with stomach content of pelagic top predators as well as fisheries statistics from the domestic longline  
29 fleet operating from Réunion. Vertical current profiles revealed distinct patterns between the offshore  
30 and seamount-flanked stations, giving evidence of topographical induced flow instabilities, notably on  
31 its leeward side (west) relative to the east flank. Distinct patterns in temperature and chlorophyll-*a*  
32 vertical profiles suggest the formation of convergent and divergent circulation cells as a result of the

33 irregular and crescent-like summit topography. Spatial differences in zooplankton abundance were  
34 detected with higher biovolumes on the leeward flank. The overall acoustic backscatter for  
35 micronekton over the summit was weaker than offshore, but highly concentrated in the upper layer.  
36 Albacore tuna and swordfish dominate the longline catch west of Réunion, seemingly in association  
37 with a deep (900 m) topographic feature. Yet the largest catch is not directly associated with La  
38 Pérouse which would be too shallow for top predators to aggregate around in the long term. Enhanced  
39 levels of phytoplankton or zooplankton enrichment at La Pérouse were not demonstrated in this study,  
40 nor was there notable diversity of micronekton species. This might explain the relatively limited  
41 importance of this seamount to the tuna fisheries in this region.

42

43 *Keywords:* bathymetry, mesoscale eddies, L-ADCP, mesozooplankton, micronekton, stomach  
44 contents, pelagic longline fisheries.

45

## 46 **1. Introduction**

47

48 The subsea topography of the South West Indian Ocean (SWIO) is rugged, formed of many  
49 banks, ridges (Mascarene Plateau, Mozambique Plateau, Madagascar Ridge, South West Indian Ridge,  
50 among others) and isolated seamounts rising from plateaus or from the deep abyssal plains (Tomczak  
51 and Godfrey, 1994; Demopoulos et al., 2003; Ingole and Koslow, 2005). The summits of these many  
52 seamounts peak at various depths, ranging from >2000 m to just a few metres below the sea surface.  
53 Seamounts are known as biodiversity hotspots and are generally rich in demersal and pelagic  
54 seamount-associated fish of high commercial value (Fock et al., 2002; Clark et al., 2007). Historically,  
55 fish inventories on seamounts worldwide have been developed based on the catches of commercial  
56 trawl surveys, such as the Soviet expeditions carried out in the Indian Ocean in the 1970s (Romanov,  
57 2003). Understanding seamount ecosystems was not a priority during this fishing expansion phase.  
58 However, following rapid depletion of stocks within a few years of exploitation, in particular stocks of  
59 long-lived, slow-maturing species such as the orange roughy (*Hoplostethus atlanticus*) and the  
60 alfonsoinos (*Beryx* spp.) (Morato and Clark 2007), the need to generate integrated ecosystem

61 knowledge was recognized in order to develop management and conservation strategies around these  
62 seamounts (Rogers et al., 2008; Clark et al, 2012).

63 The MADRidge project conducted between 2016 and 2017 (see Roberts et al., 2020) was  
64 designed to investigate the productivity associated with shallow seamounts, in particular the possible  
65 contribution of flow-topographic interactions to enhanced productivity. The project focused on three  
66 seamounts located between 19°S and 34°S — (1) La Pérouse (60 m depth), (2) an unnamed pinnacle  
67 on the Madagascar Ridge (hereafter named MAD-Ridge, 240 m) and (3) the Walters Shoal (18 m).  
68 Multidisciplinary ship surveys were conducted at each seamount. The cruise to La Pérouse was made  
69 in September 2016 (DOI: 10.17600/16004500) and aimed to study physical and biological  
70 processes that potentially drive important ecological interactions around this seamount, which  
71 otherwise is located in an oligotrophic biome — the Indian South Subtropical Gyre (ISSG).

72 La Pérouse is an extinct volcano rising from the abyssal plain (5000 m depth) up to the surface  
73 euphotic layer. Its formation is thought to have occurred 8-10 million years ago, in the same epoch as  
74 Réunion (G. Barruol, pers. comm.), as a result of an intraplaque hotspot (Barruol and Sigloch 2013).  
75 La Pérouse, as an active volcano, was probably an island in the past. However, due to erosion and  
76 seafloor subsidence, it has transitioned to an atoll before becoming a shallow seamount (Deplus et al.,  
77 2016). La Pérouse is well-known by recreational fishers who venture to the seamount from Réunion,  
78 and by professional fishers operating the Réunion-based longline fleet. Apart from fish resources,  
79 significant knowledge gaps exist on the ecosystem and physical environment at and near La Pérouse.  
80 The research cruise in 2016 represents the first integrated approach to study the La Pérouse ecosystem.

81 In this paper, we present and analyse the cruise data that reveal different components of the near  
82 and far environment of La Pérouse — including the circulation, hydrology, plankton, micronekton,  
83 large pelagic fish and megafauna taxa — to assess whether a “seamount effect” underpins an enhanced  
84 ecosystem at this isolated topographic feature on the Mascarene Plain.

85

## 86 **2. Material and methods**

87 The R/V *Antea*, owned by the French Institute of Research for Sustainable Development (IRD),  
88 was used for the survey. It is a 35-m catamaran equipped for physical and biological investigations in

89 the coastal and deep ocean. La Pérouse is located 160 km northwest of Réunion, at latitude 19°43'S  
90 and longitude 54°10'E (Fig. 1a). The R/V *Antea* left Réunion on 16 September 2016 and took 15 h to  
91 reach the study site, which was then investigated for 12 consecutive days. The cruise ended in Réunion  
92 on 29 September 2016.

93

#### 94 2.1. Bathymetric survey

95 As several physical and biological measurements were to be done along the slopes of the  
96 seamount to study the possible disturbances caused by the topographic rise, it was first necessary to  
97 perform a precise bathymetric survey. The GEBCO gridded bathymetric data identifies La Pérouse as  
98 an oval-shaped seamount with regular slopes. In 2012/2013, during the RHUM-RUM geophysical  
99 cruises on board the R/V *Marion-Dufresne* and R/V *Meteor* (DOI: 10.17600/12200070), the  
100 knowledge of seafloor topography in the area was improved by multi-beam surveys along a few  
101 transects, showing a more irregular-shaped seamount compared with the GEBCO dataset. However,  
102 not all of La Pérouse was surveyed. During our cruise in 2016, a more precise bathymetric survey was  
103 conducted using a 12 kHz sounder (SIMRAD EA 500) along latitudinal and longitudinal transects at 1  
104 km intervals over the seamount platform and slopes. This allowed high-resolution mapping of the area  
105 with depth data taken every 3 m along the transects. These data were complemented with other depth  
106 measurements made with the same echosounder along sampling lines in the survey. In all, 216801  
107 depth measurements were collected over a large area stretching from the summit to 15 km to the north  
108 and south, 30 km to the east and 25 km to the west.

109 This new bathymetric dataset was processed using Golden Software Surfer16® to grid the data  
110 across a 20 m mesh to produce a topographic map for the ship survey. The seamount has a crescent-  
111 like shape, suggesting that part of the eastern flanks collapsed or that voluminous landslides may have  
112 occurred during its formation (Fig. 1b). The existence of a flat top may suggest a drowned reef phase  
113 during the seamount's subsidence over time. The platform now lies at 60 m below sea level, and is  
114 about 2.5 km wide and reaches 12 km along its largest extension (Fig. 1c).

115

#### 116 2.2. Current measurements

117 Underway current profiles were measured along the Réunion - La Pérouse transit from 16 to 18  
118 September 2016. These data were collected with a 75 kHz RDI Ocean Surveyor II (S-ADCP) using a  
119 time average of 2 min. Data were processed using the CASCADE software (Le Bot et al., 2011) which  
120 allows flagging and filtering. Measurements were spatially averaged over 5 km intervals for plotting  
121 purposes, and restricted to when the ship was moving at constant speed. No tidal correction was  
122 applied to these data.

123 At 11 hydrographic stations (Fig. 1c), profiles of current measurements were collected by two L-  
124 ADCPs (RDI 300 kHz) mounted on the CTD frame, one facing up, the other down. This particular  
125 configuration maximized the total range of velocity observations, reduced the overall error of the  
126 system and provided redundancy of measurements. The raw data from both instruments were  
127 combined and then processed using the IFM-GEOMAR/LDEO software version 10.16 (Visbeck,  
128 2001). Data processing corrected the motion and oscillation of the CTD frame, data noise and changes  
129 in sound velocity with depth. Ancillary information such as the position (GPS) and the depth (CTD  
130 pressure sensor) were imported into the dataset. A tidal correction was first applied to the dataset  
131 before aggregating measurements to 8-m depth bins. The zonal and meridian components of the  
132 current were transformed to polar coordinates (angle and velocity) and in turn plotted for each station.  
133 A vertical direction change index (VDCI) devised by the authors was defined as the cumulative  
134 changes in direction of the current between each depth level, in the upper 200 and 400 m respectively:

$$135 \quad \text{VDCI} = \sum_{i=1}^n |(D_{i+1} - D_i)|$$

136 where  $i$  is the depth level (8-m intervals) in the 0-200 or 0-400 m layer and  $D$  the angle (0-360°) of the  
137 current vector. VDCI denotes the degree of disruption in the flow which is directly related to the  
138 rotation of the current in the water column. The higher the value, the larger the disruption.

139 Reprocessed daily L-4 Sea Level anomalies and geostrophic currents were extracted from the  
140 CMES (Copernicus Marine Environmental service) database comprising the altimetry data product  
141 SEALEVEL\_GLO\_PHY\_L4\_REP\_OBSERVATIONS\_008\_047 (<http://marine.copernicus.eu>) at a  
142 spatial resolution of 0.25° of latitude and longitude. These data were used to represent the sea level  
143 anomalies and the geostrophic flow during the cruise, and to compute time-series of the daily  
144 geostrophic flow at La Pérouse (a square of 3×3 25 km cells) from January 1993 to June 2018.

145

146 2.3. Hydrographic stations

147 Hydrographic surveys were conducted from 21 to 27 September 2016, using a CTD frame  
148 equipped with a Sea-Bird 911 Plus CTD and 11 Ocean-Test bottles (8 L Niskin type). In all, 11 CTD  
149 casts were done from the surface to a maximum depth of 1500 m (or as close to the bottom as  
150 possible) (Fig. 1c). The CTD measured depth, temperature, conductivity, dissolved oxygen and  
151 fluorescence. Water samples were collected at three depths at all stations for the analytical  
152 determination of salinity and dissolved oxygen. Nutrient samples were collected at six depths, or  
153 more, and pasteurized on board for later analysis using an AXFLOW AA3 continuous flow analyser in  
154 the laboratory (IRD, Brest, France). Four litres of seawater were filtered to collect pigments (filters  
155 preserved at -80°C) at four or five different depths, namely: surface, mixed layer, depth of maximum  
156 fluorescence (given by the CTD Fluorometer (Wetlabs® ECO-FLNTU) and below the thermocline, so  
157 as to cover a wide range of values.

158 Chlorophyll-*a* (Chl *a*), the only pigment used in this study, was later measured in the laboratory  
159 using High Performance Liquid Chromatography (HPLC). The analytical determination of salinity,  
160 dissolved oxygen and Chl-*a* was used to calibrate the CTD measurements. Several proxies were  
161 computed from the calibrated CTD profiles. We *inter alia* used the depth of ‘Temperature at 5 m  
162 minus 1°C’ as a criterion to define the mixed layer depth (MLD) (Sprintall and Cronin, 2009). The  
163 depth at which the sum of nitrate and nitrite concentrations reached 1  $\mu\text{mol kg}^{-1}$  was used to determine  
164 the depth of the nutricline using a linear interpolation between consecutive sampling depths (Dufois et  
165 al. (2014). The deep chlorophyll maximum (DCM) is defined as the depth where the peak chlorophyll  
166 value is found. Finally, the integrated Chl-*a* is the sum of the chlorophyll concentrations of the upper  
167 200 m of the water column.

168 Four stations located away from the seamount were aligned with the cardinal directions — North,  
169 South, West and East (i.e. 1, 2, 8, 10 respectively), and ranged between 10 and 21 km from the  
170 seamount platform (70 m isobaths). These served as control stations and hereafter are referred to as  
171 ‘Offshore’ stations. By contrast, seven stations (hereafter ‘Flank’ stations) were distributed from 1.1 to  
172 3.2 km around the seamount platform (ranging in depth from 800 to 1200 m, except for station 5 at

173 300 m) which enabled investigation into the possible influence of the seamount on the hydrology of  
174 the water column. All stations were then classified accordingly as ‘Offshore’ or ‘Flank’ (see Table 1).

175

#### 176 2.4. Zooplankton sampling

177 Mesozooplankton samples were collected at each hydrological station in the upper 200 m using a  
178 200- $\mu$ m-mesh Bongo net (0.25 m<sup>2</sup> mouth area), equipped with a flowmeter and connected to an  
179 underwater recording unit (attached to the frame). A V-fin depressor hanging below the Bongo frame  
180 allowed the nets to remain correctly orientated during the oblique tows at a vessel speed of 1.5-2  
181 knots. The samples were preserved in buffered formaldehyde (4% final concentration). In the  
182 laboratory, each sample was split until the subsample contained approximately 1000-2000 particles.  
183 This subsample was then scanned using a ZooScan (Hydroptics) following the procedure described in  
184 Gorsky et al. (2010) and Vandromme et al. (2012). An ellipsoid estimated the biovolume of each  
185 organism with a pre-existing training set and the random forest algorithm was used for pre-sorting of  
186 the particles with the online platform Ecotaxa (Picheral et al., 2017). The classification of each  
187 organism was verified manually. Only the biovolume of organisms (in mm<sup>3</sup> m<sup>-3</sup>) is analysed in this  
188 study. Zooplankton were not sampled at station 5 because of its shallowness. All Bongo samples were  
189 collected in daylight.

190

#### 191 2.5. Acoustic survey

192 Acoustic data were collected by both day and night along transects using a 4-frequency Simrad  
193 EK60 echosounder (38, 70, 120 and 200 kHz) at a vessel speed of 8-9 knots. The pulse duration was  
194 set to 0.512 ms. The water column was sampled to a depth of 740 m during data acquisition using the  
195 38 kHz frequency (transmitted power of 1000 W), which is the only frequency used in this paper  
196 because of the larger vertical range relative to the other frequencies. These acoustic data were  
197 processed with the Matecho software (Release 2010b), which is based on IFREMER’s Movies3D  
198 software (Trenkel et al., 2009, Perrot et al., 2018), and considered an open source IRD tool (computed  
199 with MATLAB 7.11.0.184). Background noise, bottom echoes and attenuated signals were removed  
200 using algorithms designed by De Robertis and Higginbottom (2007) and Ryan et al. (2015). The first



201 10 m below the sea surface was removed due to over-amplification of the acoustic signal by surface  
202 bubbles. The along-track sampling unit was 0.1 nautical mile (1 nmi = 1852 m). Echo-integration was  
203 performed at a Sv threshold of  $-80$  dB with 1 m layers intervals. The micronekton acoustic density  
204 was determined by the nautical area scattering coefficient (NASC, or Sa,  $\text{m}^2 \text{nmi}^{-2}$ , related to the  
205 backscattered energy), which defines a proxy of the relative biomass of micronekton, provided certain  
206 assumptions are met (Béahgle et al., 2014). The water column was split into three depth categories:  
207 10–200 m (surface layer), 200–400 m (intermediate layer) and 400–740 m (deep layer), following an  
208 initial analysis of the echograms, which indicated that organisms were organized vertically into these  
209 three main strata. Transects were grouped in distinct spatial domains from the seamount in order to  
210 compare the acoustic densities between domains. Daylight transects were distributed in two domains:  
211 ‘Vicinity’ (3-10 km off the seamount platform) and ‘Summit and flank’ ( $<3$  km), while two night  
212 transects located at 23 and 35 km east and west of the summit, respectively, defined a third domain  
213 named ‘Offshore’ (Fig. 8a later). Here, the ‘Summit and flank’ category, which extends down to 1000  
214 m along the flanks, overlays the ‘Flank’ category used for hydrological stations and Bongo nets, and  
215 the summit platform of La Pérouse.

216

## 217 2.6. Micronekton sampling

218 The first set of micronekton samples was collected during the La Pérouse cruise using a 40 m-  
219 long International Young Gadoid Pelagic Trawl (IYGPT) net, constructed of 8 cm knotless nylon delta  
220 mesh netting at the front tapering to 0.5 cm at the codend (theoretical maximal mouth opening of  $\sim 96$   
221  $\text{m}^2$ ). The trawl was towed at a speed of 3-4 knots for 1 h. Trawl depth was monitored using a Scanmar  
222 depth sensor. Trawl 5 located over the summit was the only one conducted during daylight. The other  
223 nine trawls were performed during night-time at various depths, with maximum depths ranging from  
224 60 to 590 m (Fig. 1c and Table 2). Night trawls were operated mostly after 18:00 local time (twilight).  
225 Midwater trawls were performed to sample the diversity of micronekton organisms and establish links  
226 with the acoustic densities at the different layers. With seven trawls in the ‘Vicinity’ area and one in  
227 each of the ‘Summit’, ‘Flank’ and ‘Offshore’ areas — trawl data were not used to compare areas.

228 To place the trawl composition of micronekton at La Pérouse in a broader context, we used  
229 unpublished information on fauna identified in the stomach contents of top predators, caught by the  
230 commercial longline fleet based in Réunion. This complementary, historical dataset was collected by  
231 several projects over the past 10 years: IOSSS (Ifremer) in 2009, PROSPER (CAP RUN, IRD) in  
232 2010-2011, DCF-DCMAP (IRD, CAP RUN) in 2012, RAF (CAP RUN, IRD) in 2012-2013,  
233 GERMON (Ifremer, IRD, CAP RUN) in 2014. To add to this, stomach contents of three predators  
234 were also sampled around La Pérouse during the 2016 cruise. This entailed a final dataset consisting  
235 of 93 non-empty stomachs belonging to six top predator species: three tuna species (albacore *Thunnus*  
236 *alalunga*; yellowfin *T. albacares* and bigeye *T. obesus*), swordfish (*Xiphias gladius*), dolphinfish  
237 (*Coryphaena hippurus*) and longnose lancetfish (*Alepisaurus ferox*) (Table 3).

238 The spatial distribution of the longline sets allowed us to discriminate two regions based on the  
239 distance from the seamount. A distance of 20-40 km was considered the radius of influence of La  
240 Pérouse on the large pelagic fisheries (Morato et al, 2008, 2010). The stomachs collected within 10 km  
241 of La Pérouse were in the ‘Vicinity’ area, and those collected from 75 to 280 km were assigned to the  
242 “Open-ocean area” (Fig. 1a). However, all sets located within a radius of 90 km from Réunion were  
243 excluded to minimize any island effect in the micronekton composition. There were 12 stomachs in  
244 the ‘Vicinity’ area, and 81 in the ‘Open-ocean’ area (Table 3). The spatial stratification across the  
245 different data components of this study is summarized in Table 4.

246 Prey items were identified to the lowest possible taxon using the most recent guides for each  
247 group. Identification of highly digested prey remains was based on the diagnostic hard parts (e.g.  
248 otoliths for fish, beaks for cephalopods). For each stomach, the number (N) and occurrence (O) of  
249 each species were recorded. The reconstituted weight (RW), using only fresh remains, was estimated  
250 from allometric equations relating dimensions of identifiable hard parts to the weight of the species  
251 using our own relationships or presented in the literature (Clarke, 1986; Smale et al., 1995; Lu and  
252 Ickeringill, 2002; Potier et al., 2011; E.V. Romanov unpublished). For each identified prey item, the  
253 mean proportion (%) for both number (MN) and reconstituted weight (MRW) were calculated by  
254 considering the proportion of each prey species (or taxa/category) found in the individual stomachs.  
255 Finally, the average species contributions to prey composition across all stomachs were calculated by

256 area, as previously defined. We therefore treated individual stomachs as the sampling unit, allowing us  
257 to compute standard deviation for species relative to the stomachs' statistical population (Chipps and  
258 Garvey, 2007). An index of relative importance (IRI; Pinkas et al., 1971) was calculated to combine  
259 all three measures into a single estimate of the relative importance of food types:

$$260 \quad \text{IRI} = \%O \times (\%MN + \%MRW)$$

261 from which percentage values were used for comparison:

$$262 \quad \%IRI = 100 \sum_{i=1}^n IRI_i$$

263

## 264 2.7. Large pelagic fish indicators

265 No fishing operations targeting top predators were carried out during the cruise. Rather, catch and  
266 effort (C/E) fisheries statistics available in the international database of the Indian Ocean Tuna  
267 Commission (IOTC) ([www.iotc.org/data-and-statistics](http://www.iotc.org/data-and-statistics)) were used to investigate fisheries production  
268 in the proximity of La Pérouse. The approach compared catch distribution throughout the entire South  
269 West Indian Ocean. Only statistics from the Réunion domestic longline fleet were selected because it  
270 is the only longline fleet reporting catch and effort on a 1° grid. Asian longline fleets in the area report  
271 on a 5° grid resolution, which is too coarse to investigate seamount effects on the catch. First the  
272 nominal catches were compared with the C/E statistics to verify the amount of spatial reporting. It  
273 was found that 100% of the fished squares are indeed reported in the database. Four species  
274 accounting for 94% of the total catch were examined: three tuna species (yellowfin, bigeye and  
275 albacore) and swordfish. Catch by species (metric tons) and effort (number of hooks deployed) were  
276 summed per 1° grid cell over the period 2009-2017 to produce average maps.

277

## 278 2.8. Marine mammals and seabird observations

279 Visual census of marine mammals at the sea surface was conducted from the top deck of the  
280 research vessel. Species identification, group size estimate, time and position were recorded. Sighting  
281 positions were obtained from a global position system (GPS). The visual effort (duration) was also  
282 recorded. During the 14-day cruise, only 46 h of observations during just five days were undertaken  
283 due to rough sea conditions.

284

### 285 3. Results

286

#### 287 3.1. Altimetry and regional circulation

288 The altimetry field covering the region between Réunion and La Pérouse, for the duration of the  
289 cruise, was characterized by moderate geostrophic gradients not exceeding 0.36 m (Fig. 2; sea level  
290 anomalies ranged from -0.26 m to +0.10 m). During the first part of the cruise (16-25 September), La  
291 Pérouse was located on the edge of a cyclonic eddy which moved towards the northwest at an average  
292 speed of 12 km d<sup>-1</sup> (Fig. 2a). From 16 to 18 September 2016, the S-ADCP measurements between  
293 Réunion and La Pérouse indicated high current velocities (50-90 cm s<sup>-1</sup>) over a distance of some 55  
294 km (20.5°S to 20°S) in the upper 180 m (Fig. 3a), in a northwesterly direction along the sea level  
295 gradient seen at the edge of the cyclonic eddy. During the last few days of the cruise (26-29  
296 September), the eddy field had changed with a southerly, lower velocity, geostrophic flow moving  
297 over La Pérouse (Fig. 2b). This was confirmed by the S-ADCP measurements between 28 and 29  
298 September 2016 (i.e. 30-50 cm s<sup>-1</sup>) (Fig. 3b). Statistics, calculated from the daily CMEMS altimetry  
299 data product from January 1993 to June 2018, show the geostrophic flow at La Pérouse to exhibit daily  
300 variations ranging from 2 to 55 cm s<sup>-1</sup> with a median of 16 cm s<sup>-1</sup>.

301

#### 302 3.2. Current measurements

##### 303 3.2.1 *L-ADCP profiles at offshore stations*

304 The four 'Offshore' stations were located on the North (1), South (2), East (8) and West (10) of  
305 the seamount. Current velocities in the layer 0-1000 m are displayed in Fig. S1 (Supplementary  
306 information). It is seen that stations 1 and 2 exhibited a similar vertical structure in the upper 350 m.  
307 Flow was in a northwesterly direction in the upper 150-200 m (30-54 cm s<sup>-1</sup>) and gradually rotated to  
308 the west, reaching a depth of 300 m. Below 400 m, the current was eastwards at station 1 (9-20 cm s<sup>-1</sup>)  
309 <sup>1</sup>), but in comparison, southwards then eastwards for the same depth range at station 2 (5-11 cm s<sup>-1</sup>).  
310 At station 8, the vertical current structure was northwesterly in the upper 200 m, similar to station 1  
311 with lower velocities (20-26 cm s<sup>-1</sup>). From 300 to 700 m, the flow was eastwards, i.e. in an opposite

312 direction to that in the upper layer (8-22 cm s<sup>-1</sup>). Station 10 was very different from the other three  
313 ‘Offshore’ stations. This was undertaken during the last part of the cruise, and showed an easterly  
314 current in the upper 150 m with the weakest velocities of all ‘Offshore’ stations (8-16 cm s<sup>-1</sup>). The  
315 sudden rotation of the current at station 10 coincided with a sharp pycnocline at 150 m (Fig. S1).

316

### 317 *3.2.2 L-ADCP profiles at flank stations*

318 None of the seven profiles conducted on the slopes of the seamount showed the same vertical  
319 pattern (Fig. S2). Two groups are considered here, namely stations 3, 4, 5, 6 and 7, that were sampled  
320 on 23 September, and stations 23 and 24 sampled on 27 September, which coincide with the change in  
321 the main surface flow in Fig. 2. Currents appear to be layered in the vertical axis. Average velocities  
322 are hereafter reported with 1 standard deviation.

323 At station 3, the current was highly unsteady in the upper 200 m. The flow was northwards above  
324 100 m (average  $23.6 \pm 5.9$  cm s<sup>-1</sup>) and rotated clockwise (cyclonic) from northeast to southwest in the  
325 layer 100-200 m (average 6.2 cm s<sup>-1</sup>). From 200 to 260 m, the flow was northeastwards with a sharp  
326 reversal to the west until 400 m. At station 6, the current was strong and steady to the west to a depth  
327 of 150 m (average  $24.5 \pm 5.5$  cm s<sup>-1</sup>), where a sharp pycnocline (24.4-25) was observed. Below the  
328 pycnocline, in a layer of about 100 m, the current rotated anti-clockwise (anticyclonic). The current  
329 turned gradually from eastwards to southwards in the layer 250-575 m. At station 7, the flow was  
330 northwestwards in the upper 60 m (average  $17 \pm 2.7$  cm s<sup>-1</sup>), then rotated clockwise (cyclonic) to the  
331 east from 100 to 175 m. Below 200 m, the current remained more or less in the same direction  
332 (northwest). Station 5 on the upper slope of the seamount (200 m bottom depth), was characterized by  
333 a high velocity (up to 52 cm s<sup>-1</sup>) westward flow in the upper 50 m. This then rotated and sharply  
334 decreased in velocity near the bottom. Station 4 showed a high velocity (average  $37.1 \pm 9.3$  cm s<sup>-1</sup>)  
335 towards the northwest in the upper 70 m, then rotated slightly to the northeast down to 110 m. A  
336 sudden directional change (to the west) was observed at 130 m, the same depth at which a strong  
337 density gradient was observed (24.6-25.0). Deeper, the current was steady in direction, keeping a  
338 dominant northward component.

339 The flow in the upper 100 m at both stations 23 and 24 was well established towards the south  
340 (average of  $25.1 \pm 3.7$  and  $24.2 \pm 9.2$  cm s<sup>-1</sup>, respectively). Southward flow was observed down to 200  
341 m at station 23, while a westward flow developed within a thin layer (128-144 m) in association with a  
342 rapid isopycnal change (24.2-25.0) at station 24. At station 23, the flow was unstable in direction from  
343 300 to 400 m with a dominant northward component in the deeper layers. Flow at station 24 was  
344 predominantly eastwards from 160 to 250 m, followed by westwards to 500 m.

345

### 346 *3.2.3 Statistical comparison between offshore and flank stations*

347 ‘Offshore’ and ‘Flank’ stations were compared in terms of velocity and directional changes, in  
348 both the 0-200 m and 0-400 m layers. The full paired Wilcoxon test results for these comparisons are  
349 provided in Tables S1 and S2 (Supplementary information). Velocities are examined first. Within the  
350 ‘Offshore’ stations in the upper 200 m, only stations 1 and 2 were statistically similar ( $p > 0.05$ ) with  
351 higher velocities (median of  $35.7$  and  $40.4$  cm s<sup>-1</sup>, respectively) than at the other two ‘Offshore’  
352 stations ( $25.2$  and  $14$  cm s<sup>-1</sup> for stations 8 and 10, respectively) (Fig. 4a). In the upper 400 m, stations 1  
353 and 2 were also statistically different from stations 8 and 10 ( $p > 0.05$ ) with a generally lower current  
354 velocity along the flanks of the seamount (Fig. 4b). In both layers, the ‘Offshore’ stations had a  
355 significantly different vertical velocity pattern from the ‘Flank’ stations. Statistical different patterns  
356 were observed in 75% of the paired tests in the upper 200 m and 83% in the upper 400 m (Table S1).

357 Variability in the flow direction throughout the water column was similar at ‘Offshore’ stations 1,  
358 2 and 8 in both depth ranges ( $p > 0.05$ ), with few directional changes (Table S2). Station 10 (which  
359 was affected by the westward eddy displacement) exhibited more changes in direction relative to the  
360 other ‘Offshore’ stations (except some similarity with station 1 in the 0-400 m layer;  $p > 0.05$ ). The  
361 directional change also differed between ‘Offshore’ and ‘Flank’ stations, as observed in 78% and 63%  
362 of the cases for the layers 0-200 m and 0-400 m, respectively (Table S2).

363 A vertical directional change index (VDCI) was computed to assess the extent of current rotation  
364 with depth (Table 5). For three of the ‘Offshore’ stations (1, 2 and 8) the VDCI exhibited the lowest  
365 values among all stations. Seamount stations had higher VDCIs than ‘Offshore’ stations. In the upper  
366 400 m, stations 3 and 23, both located on the western flank, had the highest VDCI with station 3 being

367 ranked first. In the range 0-200 m, station 3 had the second highest VDCI after station 5 (eastern flank  
368 at a shallow depth, 200 m, along the slope). Stations 7 and 24, located in the southwest corner of La  
369 Pérouse, showed intermediate VDCI values in both depth ranges. Station 23, located on the western  
370 flank, had the lowest VDCI in the 0-200 m depth range and, by contrast, was the second highest in the  
371 0-400 m depth range, denoting a rather stable direction of the flow in the upper 200 m and a highly  
372 disrupted flow beneath, as shown in Figure S2. Stations 6 (southwest flank) and 4 (eastern flank)  
373 exhibited weak VDCI across both depth ranges.

374

### 375 3.3. Hydrology and chlorophyll concentration

376 The 11 hydrological stations (Fig. S3) were pooled into three groups by visual examination of the  
377 temperature profile shape and depth of the mixed layer (Fig. 5a-c). Group 1 comprised ‘Offshore’  
378 stations, where the MLD is found between 100 and 130 m and the thermocline less pronounced than in  
379 group 2 (Fig. 5a). Groups 2 and 3 are for the ‘Flank’ stations. Group 2, consisting of stations 3, 23 and  
380 7, showed a MLD between 80 and 89 m (Fig. 5b). The temperature gradient at the bottom of the MLD  
381 was weak, as temperature decreased in a quasi-continuous linear trend below 50 m. By contrast, Group  
382 3 comprising stations 4, 24 and 6, had a well-defined and deep thermocline at 150 m (MLD ranges  
383 between 91 and 136 m, Fig. 5c). The surface warming seen at stations 23 and 24 was due to calm  
384 weather conditions on 26 and 27 September, not experienced at any other time during the cruise.

385 Relative to temperature, the fluorescence profiles showed a greater within-group heterogeneity  
386 (Fig. 5d-f). In Group 1, the maximum chlorophyll concentration ( $F_{max}$ ) ranged widely between 0.19  
387 and  $0.4 \text{ mg m}^{-3}$  with a DCM fluctuating between 115 and 141 m (Fig. 5d). In Group 2, the  $F_{max}$  was  
388 twice as high at stations 23 and 7 ( $0.4 \text{ mg m}^{-3}$  and above) than at station 3 ( $0.19 \text{ mg m}^{-3}$ ) and the DCM  
389 fluctuated between 65 and 86 m (Fig. 5e). In Group 3, the  $F_{max}$  was relatively low ( $0.25\text{-}0.32 \text{ mg m}^{-3}$ )  
390 and the DCM was deeper than in Group 2 stations, between 65 and 108 m (Fig. 5f). The position of  
391 the DCM was closely associated with the nutricline for ‘Offshore’ and ‘Flank’ stations ( $r_{sp} = 0.92$ ;  $p <$   
392  $0.05$ ) with  $F_{max}$  located a few metres above the nutricline.

393 MODIS surface Chl-*a* concentration, averaged between 23 and 27 September, was observed to be  
394 slightly higher in the proximity of the seamount than in the surrounding waters (Fig. S4). The

395 distribution was patchy around the seamount, with higher concentrations over the northern platform  
396 and slopes than on the southern part. Nonetheless, the range of Chl-*a* concentration (0.02–0.08 mg m<sup>-3</sup>)  
397 was relatively low and characteristic of an oligotrophic environment. A detailed comparison with  
398 other seamounts in the SWIO is given by Demarcq et al. (2020).

399

#### 400 3.4. Zooplankton

401 The 10 net tows were classified into three groups according to total biovolume, namely Offshore  
402 (4 tows), Western flank (4 tows) and Eastern flank (2 tows) (see Fig. 6). A Kruskal-Wallis ANOVA  
403 test (KW) indicated that all three groups were significantly different ( $H = 7.85$ ,  $p < 0.05$ ), with the  
404 highest biovolume recorded along the western flank (median =  $59 \text{ mm}^3 \text{ m}^{-3}$ ), the smallest along the  
405 eastern flank ( $23.4 \text{ mm}^3 \text{ m}^{-3}$ ) and intermediate values at the offshore tows ( $38.9 \text{ mm}^3 \text{ m}^{-3}$ ). The  
406 relatively high biovolume measured at station 7 should be regarded with caution as it was the only tow  
407 with organisms  $>4$  mm. These comprised large Trachymedusa (Noyon et al., 2020). Overall, the  
408 zooplankton biovolume was inversely related to the integrated Chl-*a* concentrations in the upper 200  
409 m (Fig. 7). The range of zooplankton biovolume and integrated Chl-*a* on the flank tows were  
410 distributed over a wide range of values relative to those for the offshore tows (1, 2, 8, 10). The  
411 correlation between the biovolume and the integrated Chl *a* was highly significant for the flank tows  
412 ( $r_{sp} = -0.94$ ,  $p < 0.05$ ).

413

#### 414 3.5. Micronekton acoustic densities along transects

415 The night-time acoustic density for the entire water column was significantly greater than in  
416 daylight for all transects ( $1025$  and  $499 \text{ m}^2 \text{ nmi}^{-2}$ , respectively, Wilcoxon rank sum test,  $W = 1311$ ,  $p <$   
417  $0.05$ ). However, the daylight acoustic density, similarly for the entire water column, was significantly  
418 different across all transects (Fig. 8b) (KW,  $H=185.49$ ,  $df=12$ ,  $p < 0.05$ ). Transects carried out to the  
419 east (East 2 and 3), northeast and in the western extremity (West 3) of the seamount showed higher  
420 daylight acoustic densities ( $589$ - $635 \text{ m}^2 \text{ nmi}^{-2}$ ) than all the other transects ( $336$ - $499 \text{ m}^2 \text{ nmi}^{-2}$ ) (Fig. 8b).  
421 For almost all transects, the deep scattering layer was denser than the surface ( $52.3\%$  vs  $41.8\%$  on  
422 average). The ‘Summit and flank’ area exhibited the lowest overall acoustic densities, although the



423 proportion in the surface layer was the highest by day (59%, similar to the southwest transect which  
424 was located close to the slope).

425 In a similar way to daylight, the night-time acoustic responses of the whole water column also  
426 differed significantly across all transects (Fig. 8c) (KW,  $H = 182.18$ ,  $df = 11$ ,  $p < 0.05$ ). The transects  
427 in the vicinity of the seamount exhibited the highest acoustic densities, whereas the lowest values were  
428 found on the ‘Summit and flank’ area of the seamount. Most of the night-time acoustic signal was  
429 located in the surface layer, which represented 87% of the total acoustic density on average. The  
430 highest proportion (98%) of surface layer relative to the total acoustic response at night was again  
431 recorded at the ‘Summit and flank’ transects.

432 Over the ‘Summit and flank’, the distribution of the daylight and night-time acoustic responses of  
433 the total water column was asymmetrical with a strong positive skewness (Fig. 9). Such a distribution  
434 was significantly different (with a lower median) from that of the ‘Vicinity’ area (KW,  $H = 66.3$ ,  $df =$   
435  $1$ ,  $p < 0.05$ ). At night, significant differences were also found between the ‘Summit and flank’ and the  
436 ‘Vicinity’ and ‘Offshore’ areas (KW,  $H = 98.44$ ,  $df = 2$ ,  $p < 0.05$ ), which were also significantly  
437 different ( $p < 0.05$ ) (Fig. 9).

438 Echograms of the 38 kHz frequency showed aggregations of organisms closely associated with La  
439 Pérouse’s summit in the top 100 m by day (Fig. 10a). A secondary scattering layer was located below  
440 400 m during the day, whereas at night, the surface scattering layer was intensified in the upper layer  
441 of the water column and therefore over La Pérouse summit (Fig. 10b).

442

### 443 3.6. Composition of micronekton around La Pérouse

#### 444 3.6.1 *Trawl sampling*

445 The best represented group in trawl landings, both in number (36%) and mass (54%), were fish.  
446 Overall, 77 taxa were caught (2272 fish) belonging to 19 families (Cherel et al., 2020). The gelatinous  
447 micronekton was the second most abundant group (33% in number and 24% in mass) at La Pérouse  
448 followed by crustaceans (24% in number and 10% in mass) and cephalopods (7% in number and 12%  
449 in mass). Fish were dominated by myctophids (49%), the most diverse family with 28 species. The  
450 second most represented family was the gonostomatids (21%), with three species (see Cherel et al.,

451 2020, for more detail on the taxonomic composition of the trawls). Cephalopods and crustaceans  
452 delivered 10 and 12 species, respectively.

### 453 2.6.2 Stomach content analyses

454 Fish was the principal prey item in the stomach contents of pelagic top predators in both the  
455 ‘Vicinity’ and ‘Open-ocean’ areas: 53.7% and 44.6% of IRI respectively (Table 6). Fish prey were  
456 dominated by myctophids in both areas. Unidentified fish (apparently poorly identified groups of  
457 demersal species) were largely represented in stomach contents at La Pérouse vicinity (29.8% of IRI,  
458  $16.0 \pm 31.2\%$  of MRW).

459 Crustaceans were the second most important prey group in the vicinity of La Pérouse and found in  
460 lesser abundance in stomachs in the ‘Open-ocean’. Cephalopods showed an opposite pattern, with  
461 high importance in the ‘Open-ocean’ area both in terms of IRI (40.3%) and MRW (41.5%), and low  
462 importance at in the vicinity of La Pérouse (15.8% of IRI but 41.8% of MRW owing to the occurrence  
463 of few large squid) (Table 6). Among the ‘Open-ocean’ squid, the onychoteuthid whale squid  
464 *Walvisteuthis rancureli* was the second most important prey item (18.8% of IRI,  $6.6 \pm 29.2\%$  of  
465 MRW) followed by purpleback flying squid *Sthenoteuthis oualaniensis* (Ommastrephidae) (9.8% of  
466 IRI and  $4.3 \pm 14.2\%$  of MRW).

467 In the ‘Vicinity’ area, besides the identified fish taxa, prey species comprised of post-larvae reef  
468 lobster *Enoplometopus* spp. (14.6% of IRI,  $15.3 \pm 31.8\%$  of MRW), non-identified fish and fish larvae  
469 (9.3% of IRI,  $4.5 \pm 10.5\%$  of MRW), post-larvae of the stomatopod *Odontodactylus scyllarus* (6.9% of  
470 IRI,  $4.2 \pm 12.3\%$  of MRW), adult hyperiid amphipods (*Platyscelus ovoides*, 5.6% of IRI and  $2.7 \pm$   
471  $6.8\%$  of MRW) followed by whale squid (5.3% of IRI,  $11.0 \pm 26.3\%$  of MRW) and argonauts  
472 (*Argonauta argo*) (4.9% of IRI,  $7.9 \pm 18.6\%$  of MRW). Contrasts between high IRI and low MRW  
473 values (e.g. crustaceans and juvenile fish) are related to high numerical abundance and common  
474 occurrence of prey with low individual weight.

475

### 476 3.7. Distribution of large pelagic catches

477 The Réunion domestic longline fleet mostly operate in the Mascarene basin in the ISSG province,  
478 between Réunion and Madagascar (latitudes 16-27°S). Maps of catch by species and fishing effort

479 (number of hooks) are shown in Fig. S5. The fishery is characterized by two core zones: (1) the east  
480 coast of Madagascar (19-23°S) and (2) an area west of Réunion and south of La Pérouse (19-22°S, 54-  
481 55°E). Some species-specific patterns stand out. Yellowfin tuna catches are of the same magnitude in  
482 both areas. Bigeye tuna are essentially caught east of Madagascar whereas albacore and swordfish are  
483 mostly exploited west of Réunion. To quantify the relative importance of large pelagic catches in the  
484 area close to La Pérouse, we calculated the proportion of catch taken by month in the two 1° squares  
485 adjacent to La Pérouse (19-21°S, 54-55°E) (Fig. S6). On average, the proportion ranged between 10  
486 and 15% during the summer peak fishing season (November-February). The catch pattern follows that  
487 of fishing effort. Among these species, albacore yielded the highest catch in the proximity of La  
488 Pérouse mostly all year round, with exceptions in November and December when it is surpassed by  
489 yellowfin and bigeye tuna. However, the fishing maps in Figure S5 also show the largest catches for  
490 albacore and swordfish from square 21-22°S, 54-55°E located some 100-150 km south of La Pérouse.  
491 Interestingly, another deeper topographic rise, known as ‘Clé à molette’ by the fishers, is found there.  
492 This has a summit around 900 m deep and contributes 15% of the Réunion fleet catch between  
493 November and January.

494

### 495 3.8. Observations of megafauna

496 During the entire cruise, the only marine mammal sighting was a group of two baleen whales.  
497 The sighting was too distant to allow accurate species identification. For seabirds, a total of 40  
498 sightings was recorded. The most frequently encountered seabirds were Barau’s petrels (*Pterodroma*  
499 *barau*) with 10 sightings (12 individuals), and sooty terns (*Onychoprion fuscatus*) with 8 sightings  
500 (29 individuals). Roseate terns (*Sterna dougallii*) and brown noddies (*Anous stolidus*) were the least  
501 commonly sighted — 2 sightings (13 individuals) and 2 sightings (62 individuals), respectively. Other  
502 observations of seabird species included wedge-tailed shearwater (*Puffinus pacificus*), white tailed  
503 tropicbird (*Phaeton lepturus*) and Audubon shearwater (*Puffinus lherminieri*).

504

## 505 4. Discussion

506

507 The research cruise carried out in September 2016 around La Pérouse covered many aspects of the  
508 seamount ecosystem (topography, hydrology, plankton and micronekton distribution) and generated  
509 new datasets on a poorly known topographic rise, only discovered in 1962. The high-resolution map of  
510 La Pérouse produced during this cruise gave evidence of a crescent-shaped summit, possibly as a  
511 result of a huge landslide that occurred on its eastern flank in the distant past. The newly acquired  
512 knowledge of the irregular shape of the summit is potentially key to understanding the circulation  
513 pattern around the seamount.

514

#### 515 4.1. Seamount influence on local circulation

516 The current profiles revealed very different patterns between the ‘Offshore’ (control) and ‘Flank’  
517 stations. ‘Offshore’ stations 1 and 2 (north and south of the study area) were statistically similar with a  
518 current flowing to the west-northwest in the upper 250 m at a speed of about  $36.3 \text{ cm s}^{-1}$ . These  
519 measurements reproduced the regional current pattern depicted by the S-ADCP during the transit from  
520 Réunion to La Pérouse, between 16 and 18 September (Fig. 3a). This was also in agreement with the  
521 altimetry field (Fig. 2a) indicating a sustained westward geostrophic current along a sea level gradient  
522 between a cyclone and an anticyclone located northeast and southwest of La Pérouse, respectively. At  
523 station 8 (western station, sampled on 25 September) the flow velocity was lower than at stations 1  
524 and 2, with an average of  $24.9 \text{ cm s}^{-1}$ . Station 10 (eastern station sampled on 26 September) was  
525 distinct from the other three ‘Offshore’ stations with a low speed and eastward flow (average of  $13.8$   
526  $\text{cm s}^{-1}$ ). This change in current pattern is caused by the evolving altimetry field (Fig. 2b), as confirmed  
527 by S-ADCP measurements on the return transit from La Pérouse to Réunion, on 28 and 29 September  
528 (Fig. 3b). Stations 8 and 10 reflect this displacement of the eddy field. Then, during the last days of the  
529 cruise, two key observations were made, viz.: (i) the altimetry gradient was reduced, resulting in  
530 weaker currents; and (ii) station 10 being in the easternmost position was more influenced by the  
531 ongoing changes than station 8 in the westernmost position. The effect of the changing eddy field also  
532 affected the ‘Flank’ stations. At stations 3-7 (23 September), the flow was mostly northwesterly, but  
533 southerly at stations 23 and 24 (27 September), when the eastern limb of the cyclone was reaching La  
534 Pérouse.

535 The 'Flank' stations showed a larger range of current velocities and directions than the control  
536 stations, giving evidence of flow instabilities caused by the seamount slopes within a distance of some  
537 3 km from the summit platform. Current rotation through the water column was the most obvious flow  
538 disturbance, at times associated with the pycnocline. In general the VDCI was higher at the 'Flank'  
539 stations than at 'Offshore' stations. In the upper 400 m, the highest value was found at stations 3 and  
540 23, both located on the western slopes (leeward side of the summit). In contrast, station 6 had the least  
541 turbulent flow, likely owing to it being the farthest 'Flank' station. These findings clearly indicate that  
542 topographic-induced turbulence in the water column is higher on the leeward side of the seamount,  
543 and diminishes with distance from the summit. Furthermore, changes in the direction of flow below  
544 the seamount summit, i.e. between 100 and 120 m on the eastern flank (stations 4 and 24), show  
545 currents to follow the seamount shape.

546 Deflections of the impinging flow caused by a seamount is a common phenomenon, well-  
547 described by observations and models (Mohn et al., 2009, Lavelle et al., 2010). Taylor columns,  
548 trapped and closed circulations around an obstacle (Taylor, 1923; Hogg, 1973; Huppert, 1975,  
549 Huppert and Bryan, 1976), can only be generated under specific conditions of current speed, water  
550 column stratification and topography. The resulting circulation is anticyclonic which causes trapping  
551 of upwelled water in the water column above the summit, with implications for the local biology  
552 (Genin and Boehlert, 1985; Pitcher and Bulman, 2007). In this regard, vertical flow components  
553 observed around La Pérouse do not show any pattern suggesting the existence of a Taylor column  
554 there, certainly at the time of the cruise. This is not surprising. We recall that the theoretical  
555 calculation to determine whether or not a Taylor column can form at a seamount depends on two  
556 metrics, the Rossby number ( $Ro$ ) and the relative height of the seamount relative to the ocean depth  
557 ( $\alpha$ ) (Chapman and Haidvogel, 1992, White et al., 2007). Both metrics are used to calculate a blocking  
558 parameter ( $Bl$ ) which determines the potential for the generation of a Taylor column over tall  
559 seamounts (when  $\alpha > 0.4$ ,  $Ro < 0.15$  and  $Bl > 2$ , see Chapman and Haidvogel, 1992). In the case of La  
560 Pérouse, the calculation does not support the formation of a Taylor circulation around its summit  
561 (Annasawmy et al., 2020c). Moreover, in reality, the complex crescent-shaped seamount summit,

562 combined with bathymetric protrusions on its flanks, prevent water circulating around the summit with  
563 steady anticyclonic motion, i.e. formation of a Taylor column.

564 It is interesting, that a weak Taylor column was evidenced by direct current measurements at the  
565 shallow Cobb seamount (24 m below the surface, similar to La Pérouse), North Pacific (Freeland,  
566 1994). The difference it seems between the two seamounts concerns the average flow speed which  
567 was  $0.1 \text{ m s}^{-1}$  at Cobb, compared with  $0.3\text{-}0.4 \text{ m s}^{-1}$  at La Pérouse. In other words, La Pérouse appears  
568 to be in a stronger current regime preventing the formation of a Taylor column. Indeed, it should be  
569 remembered that even when theoretically possible, Taylor columns can be shed from seamounts when  
570 strong currents occur (White et al., 2007).

571

#### 572 4.2. Vertical temperature structure and planktonic abundance

573 Two groups of temperature profiles were identified among the ‘Flank’ stations and compared with  
574 ‘Offshore’ stations. The latter, also referred to as ‘control’ stations, represented the typical vertical  
575 temperature structure found in the western ISSG for the season (Longhurst, 1998), characterized by a  
576 deep mixed layer reaching 150 m with moderate thermocline gradient. The nutricline is deep  
577 associated with a distinct DCM. Group 1 of the ‘Flank’ stations characterized by a steady temperature  
578 decrease down the water column, suggested vertical mixing at all depths (stations 3, 23 - west flank-  
579 and 7 - southeast flank). In contrast, Group 2 was typified by a well stratified water column (with a  
580 deep MLD) observed at stations 4, 24 (east flank) and 6 (southwest flank). Overall, the nutricline and  
581 the maximum fluorescence were shallower and more intense (+30% for the fluorescence) in Group 1  
582 relative to Group 2. As previously mentioned, vertical current profiles indicated a cyclonic rotation of  
583 flow at stations 3 and 7, with anticyclonic motion at station 6. This suggests an upwelling-like scenario  
584 at the Group 1 stations with downwelling at Group 2. More generally, this implies alternating  
585 convergence and divergence flow cells sporadically around the seamount, presumably attributable to  
586 the irregular shape of the seamount. These cells however, are likely to be non-permanent and will  
587 evolve according to changes in current velocity and direction. Of course, this could be tested through a  
588 more dense, synoptic array of measurements in future studies.

589 The zooplankton biovolume was higher on the leeward side of the seamount, particularly at  
590 stations 3 and 6. However, those two stations differed in physical properties, i.e. average flow speed in  
591 the upper 200 m, and vertical current patterns (high VDCI at station 3, low at station 6) belonging to  
592 distinct cells (upwelling vs downwelling). Hence, no obvious physical driver is able to account for the  
593 elevated biovolumes. It is possible, though, that the negative relationship between zooplankton  
594 biovolume and integrated Chl *a* (upper 200 m) may suggest a top-down control by zooplankton on the  
595 phytoplanktonic community through enhanced grazing especially at stations 3 and 6 (hence the lowest  
596 Fmax among this group of stations) relative to the other stations around the seamount. At the  
597 ‘Offshore’ stations, no such contrast between zooplankton biovolumes and integrated Chl *a* existed,  
598 suggesting that specific processes (not fully understood through our observations) may have occurred  
599 in association with the seamount.

600 It is intriguing that the MODIS-Aqua data reveal only limited Chl-*a* enhancement at La Pérouse.  
601 Demarcq et al. (2020) demonstrated an increase of 6-8% (relative to the average Chl-*a* in a radius of  
602 90 km from the seamount) in October and during December-March around the slopes of La Pérouse.  
603 Our observations showed the background chlorophyll concentration to be low, denoting overall  
604 oligotrophic conditions (Fig. S4), with the small patches of chlorophyll being transient in nature (Fig.  
605 S7). This does not seem unusual. Many studies carried out on seamounts do not show a consensus on  
606 chlorophyll enhancement associated with seamounts. For example, significant primary production  
607 enrichment was detected at the Cobb seamount (Dower et al., 1992; Comeau et al., 1995) and at  
608 Komahashi N°2 seamount (Odate and Furuya, 1998), among others. However, the surveys at Great  
609 Meteor Seamount (Mouriño et al., 2001) failed to demonstrate persistent increases in depth-integrated  
610 primary production and biomass over the seamount. It therefore seems that seamount effects on  
611 phytoplankton and zooplankton are subject to temporal and spatial variability at seasonal and shorter  
612 time-scales (as shown during our experiment with regards to the evolving altimetry field). However,  
613 studies by Genin and Boehlert (1985), Mohn and White (2007) among others, have shown that  
614 phytoplankton enrichment tends to occur in association with cold domes and Taylor columns which  
615 isolate a water mass from the surrounding areas. As mentioned earlier, elevated flow speeds prevented

616 the formation of any retention feature at La Pérouse, and phytoplankton enhancement might be swept  
617 away from the summit and diluted in the nearby water mass (also see Demarcq et al., 2020).

618

#### 619 4.3. Seamount influence on micronekton

##### 620 4.3.1 *Micronekton distribution investigated using acoustics*

621 The acoustic density, as a proxy of micronekton abundance, highlights clear temporal and spatial  
622 patterns around La Pérouse. The increase of the acoustic responses in the surface layer at night (87%  
623 of acoustic density) reflects the diel vertical migration of several micronekton taxa from below 400 m  
624 to the surface (top 200 m) at dusk and inversely at dawn (Lebourges-Dhaussy et al., 2000; Béhagle et  
625 al., 2014; Annasawmy et al., 2018). The intermediate layer recorded the lowest acoustic densities  
626 across all transects during both day and night because most micronekton taxa are usually vertically  
627 arranged in two well-defined surface (top 200 m) and deep (between 400 and 700 m) layers in the  
628 South West Indian Ocean (Annasawmy et al., 2020b) and only transiently occupy the intermediate  
629 zone during vertical migrations up and down the water column. The average daylight and night-time  
630 acoustic responses of the total water column recorded across the transects conducted at La Pérouse  
631 (from summit to offshore areas) were in the same range of values as recorded south of Mauritius  
632 during the Microton cruise under moderate current speeds and mesoscale variability (Annasawmy et  
633 al., 2018).

634 However, the micronekton acoustic densities were not distributed evenly around La Pérouse. The  
635 ‘Summit and flank’ area of the seamount had a lower median acoustic response compared with the  
636 ‘Vicinity’ and ‘Offshore’ areas. Low total acoustic responses over the summit of seamounts are not  
637 unusual. This was also observed at MAD-Ridge relative to ‘Offshore’ transects (Annasawmy et al.,  
638 2020c). The physical obstruction created by the topography has been hypothesized to explain a  
639 reduced density of deep-dweller animals over the summit and flanks of seamounts (Genin et al., 1988;  
640 Diekmann et al., 2006; De Forest and Drazen, 2009). The lower densities over the summit and slopes  
641 could also be the result of increased predation by seamount-associated predator species. Interviews  
642 carried out with recreational fishers operating at La Pérouse from 2016 to 2019, complemented by the  
643 analysis of high resolution photos provided by collaborative fishers (E.V. Romanov, unpublished),



644 sampling of demersal fish caught by bottom longline fishing on the seamount platform during the  
645 cruise, and results of handline fishing experiments during the South West Indian Ocean Fisheries  
646 Project (SWIOFP, World Bank) in 2008, all indicate a large diversity of benthic and pelagic predators  
647 (30 species from 15 families) on the seamount. A recent inventory by a scuba diving census performed  
648 in October 2019 (Durville et al., 2019) provided a new estimate of at least 120 fish species, including  
649 30 predator species, with the remainder comprising small coral-associated species not sampled by  
650 previous surveys. These fish predators feed primarily upon oceanic micronekton when it gathers over  
651 the summit mostly during the night, and to a lesser extent, by day (Fig. 10a-b). The fish diversity at La  
652 Pérouse denotes a high degree of richness when compared with other seamounts such as the East  
653 Atlantic Seine Seamount, with 16 species (Christiansen et al., 2009) or the Great Meteor seamount, the  
654 largest seamount in the Atlantic, with 35 species (Ehrich, 1977). On the other hand, fish diversity at La  
655 Pérouse is similar to that of the Mozambique Channel seamounts (130 species; Parin et al., 2008) or  
656 the Madagascar Ridge (109 species; Shcherbachev et al., 1989). However, caution must be exercised  
657 when directly comparing seamount diversity, because statistical biases induced by differing sampling  
658 techniques, gear and effort applied can arise.

659 Of note is that the night-time deep-trawl 2 carried out on the western side of La Pérouse (4.2 km  
660 from the summit) sampled a greater proportion of fish (94%), with the seamount-associated fish  
661 species *Diaphus suborbitalis* dominating the catch (Annasawmy et al., 2020b). During our cruise, this  
662 species was only found close to La Pérouse both in the shallow and deep layers (Annasawmy et al.,  
663 2020b; Cherel et al., 2020). Populations of *D. suborbitalis* have also been found associated with  
664 continental and oceanic island slopes, mid-ocean ridges and seamounts (Kawaguchi and Shimizu,  
665 1978; Bekker and Shcherbachev, 1990) such as the Equator seamount in the tropical western Indian  
666 Ocean. *D. suborbitalis* prey on copepods (for small zooplankton organisms), and on other fish species  
667 (e.g. *Cyclothone* sp.) and shrimps for larger individuals (Gorelova and Prut'ko, 1985). At La Pérouse,  
668 *D. suborbitalis* had trophic levels ranging from 3.5 to 3.8 and were found one trophic level higher than  
669 the zooplankton sampled at these stations, confirming the prey-predator relationship (Annasawmy et  
670 al., 2020a).

671

#### 672 4.3.2 Micronekton composition investigated using trawl and stomach sampling

673 Compared with trawl catches (Cherel et al., 2020), biological samplers (i.e. the predators) prey  
674 upon fewer fish species (37 vs. 77). The dominant group in trawl catches was myctophid fish, which  
675 also represented the primary prey group in top predator stomachs in the ‘Open-ocean’ area beyond 75  
676 km from La Pérouse. The trawl samples also revealed high abundance of elongated bristlemouth fish  
677 *Sigmops elongatus* (Gonostomatidae) that were absent from the stomachs of epipelagic top predators.  
678 This species was caught at night in trawl hauls deeper than 100 m (Cherel et al., 2020) but was absent  
679 from the daytime haul. The absence of this species in stomachs can thus be explained by its low  
680 overlap with the daylight habitat of top predators, which occupy the water column higher than the  
681 deep-dwelling *S. elongatus*.

682 Despite the difference in species richness between trawls and predator stomachs, the relative  
683 abundance of the main dietary groups at La Pérouse shows similarities between the two sources. Fish  
684 were the most abundant group (53.7% of IRI, 35.9% of trawl collections). Moreover, the second  
685 dominant group in the trawls (gelatinous) was not significantly eaten by top predators. The two  
686 remaining groups, crustaceans and cephalopods, were the second (28.4% of IRI; 23.7% in trawls) and  
687 third (15.8%; 6.9%) most prolific taxonomic groups in the stomachs of the predators, respectively.  
688 More representative samples for elusive prey such as squid are found in the predator stomachs than in  
689 towed gear. This is due to the active avoidance behaviour of squid, leading to their underestimation in  
690 trawls (Tsuchiya et al., 1991). The prey sampled in the stomachs (Table 6) would suggest that  
691 mesopelagic fish (notably the myctophid *D. suborbitalis*, Cherel et al., 2020) and crustaceans  
692 aggregate at La Pérouse.

693

#### 694 4.4. Seamount influence on large pelagic fish and megafauna

695 Seamounts are known by fishers to attract tuna and other large pelagic fish. The importance of  
696 seamounts in purse-seine and longline fisheries has also been documented and analysed extensively in  
697 the Atlantic, Pacific and Indian oceans (Fonteneau, 1991; Holland and Grubbs, 2007; Morato et al.,  
698 2008, 2010; Marsac et al., 2014). Processes leading to prey aggregation are often suggested to explain  
699 these associations (De Forest and Drazen, 2009). The fishing maps shown in Fig. S5 fail to

700 demonstrate that La Pérouse is an outstanding fishing hotspot for large pelagic fish targeted by the  
701 domestic longline fleet to the west of Réunion. Rather high catches are made slightly south of the La  
702 Pérouse seamount. Nonetheless, without necessarily being a true hotspot, the two 1° squares adjacent  
703 to La Pérouse (54-55°E, 19-21°S) still represent some 15% of the regional catch at the peak season. As  
704 mentioned earlier we also observe that the primary fishing area west of Réunion (54-55°E, 21-22°S) is  
705 distinguished by the presence of a deep, elongated shoal at a depth of 900 m. Among the four species  
706 targeted by the Réunion-based longline fleet, two (albacore and swordfish) are more clearly associated  
707 with this feature throughout the year (Fig. S5). Albacore is a species which exhibits a diel pattern,  
708 occupying warm waters in the mixed layer at night and deeper, cooler waters below the MLD (~ 400  
709 to 500 m) during the day (Murray, 1994; Childers et al., 2011, Kiyofuji et al., 2013).

710 Stomach content studies show that albacore are opportunistic, generalist predators mostly preying  
711 on dense aggregations of pelagic fish, cephalopods and sometimes crustaceans (Williams et al., 2015).  
712 In the ISSG province, mesopelagic and meso-bathypelagic species (such as Paralepididae,  
713 Diretmiidae, Microstomatidae and others) constitute up to 23.0% of albacore stomach contents in  
714 terms of relative importance (Romanov et al., 2020). Swordfish are also deep-dwellers by day, diving  
715 as deep as 1200 m (Abecassis et al., 2012), and relying on migrating prey from the deep scatter layer  
716 during its circadian cycle. In the proximity of La Pérouse, the daylight habitat of swordfish lies  
717 between 600 and 800 m depth (Romanov et al., 2016). Mesopelagic cephalopods and fish are the main  
718 prey of swordfish in the ISSG province (Potier et al., 2007). However, swordfish also can reduce their  
719 swimming speed and feed on benthic and benthic-pelagic species at the summits of seamounts,  
720 underwater shoals and off the continental slope (Scott and Tibbo, 1968; Carey and Robison, 1981;  
721 Stillwell and Kohler, 1985). It might be then that these dietary habits and habitat utilization explain  
722 why the albacore and swordfish catch is higher west of Réunion. Moreover, it can be hypothesized that  
723 La Pérouse is perhaps too shallow (60 m) for long-term aggregation of highly mobile (horizontally and  
724 vertically) pelagic fish. Many of the seamounts cited in the literature with notable exploited tuna  
725 aggregations are deeper than La Pérouse. Examples include three seamounts off Senegal, Liberia and  
726 in the Gulf of Guinea in the Eastern Atlantic (250-350 m; Fonteneau, 1991), the Cross Seamount in  
727 the Hawaii chain (330 m; Holland and Grubbs, 2007, Gilman et al. 2012), and Travin Bank on the

728 Coco de Mer ridge, north of the Seychelles, in the equatorial Indian Ocean (187 m; Marsac et al.,  
729 2014), among others.

730 Cherel et al. (2020) report an abundance of seamount-associated myctophids around La Pérouse  
731 (*D. suborbitalis* in particular), along with hatchetfish and gonostomatids. These were identified by the  
732 use of acoustics around the seamount, and may represent a substantial food source for tuna spending  
733 large parts of their diel cycle below 300 m. However, with the exception of the myctophids recorded  
734 in predator stomachs during this study, none of the two other groups contributed significantly to the  
735 diet of albacore tuna and swordfish in the ISSG province (Romanov et al., 2020). Most of the bathy-  
736 and mesopelagic community, vulnerable to towed gears in the proximity of La Pérouse at night, in fact  
737 migrate to deeper layers at dawn, which lie outside the daylight habitat of epipelagic predators.

738 Steep topographies also serve as landmarks for migrating megafauna, such as cetaceans  
739 (Kaschner, 2007), sea turtles (Santos et al., 2007) and seabirds (Monteiro et al., 1996; Pinet et al.,  
740 2012) during their large-scale movements. In the case of humpback whales, the importance of  
741 underwater features was highlighted in the South Pacific (Garrigue et al., 2015), South Atlantic  
742 (Baracho-Neto et al., 2012) and South West Indian Ocean (Dulau et al., 2017). At La Pérouse, satellite  
743 tracking has revealed humpback whales using the seamount as a breeding area during the austral  
744 winter (Dulau et al., 2017). Aerial surveys conducted in the region between Mauritius and Madagascar  
745 showed the La Pérouse area broadly being used by cetacean groups (Mannocci et al., 2014). These  
746 surveys resulted in high encounter rates with marine mammals, including large delphinids, beaked  
747 whales and Globicephalinae. However, despite our cruise taking place during the humpback whale  
748 breeding season, there was no confirmed sighting of humpback whales. The only single marine  
749 mammal sighting was a pair of unidentified baleen whales. It is noteworthy that 2016 was reported as  
750 an anomalous year with very few sightings of baleen whales in the whole SWIO region (V. Dulau,  
751 pers. comm), which would explain the lack of observations in our study area.

752 Previous surveys undertaken in the SWIO demonstrated that Barau's petrels and wedge-tailed  
753 shearwaters occupy the southern part of the tropical Indian Ocean, in particular the Mascarene Plateau  
754 area (Le Corre et al., 2012). Brown terns, tropicbirds and noddies also frequent this area (Mannocci et  
755 al., 2014). The aerial surveys conducted in 2012 showed that the water surrounding La Pérouse is

756 extensively used by several seabird groups (Mannocci et al., 2014). In addition to species for which no  
757 density estimates were available, observations from their study recorded 37 brown terns per 100 km  
758 (most likely sooty terns) and 13 procellariids per 100 km (wedge-tailed and Audubon's shearwaters,  
759 Barau's petrel) around La Pérouse. However, despite a limited observational effort, seabird sightings  
760 during the La Pérouse cruise were congruent with the average estimates produced by previous surveys,  
761 highlighting the fact that La Pérouse is not a hotspot for seabirds, in terms of either diversity or  
762 foraging activity.

763

## 764 **5. Concluding remarks**

765

766 This paper presents new physical and biological data collected in the vicinity of the La Pérouse  
767 seamount in an attempt to assess whether there is a 'seamount effect' on the various trophic levels of  
768 the surrounding ecosystem. The data collected clearly showed that this steep, crescent-shaped  
769 topographic feature, located in the oligotrophic province of the Indian South Subtropical Gyre,  
770 strongly influences the circulation around the summit flanks in the upper 400 m. Consequently, it is  
771 likely that small circulation cells (convergent or divergent) are created along the slopes with  
772 implications for the vertical transport of temperature, nutrients and Chl-*a*. Certainly, satellite  
773 observations showed some Chl-*a* enrichment over the seamount. Although somewhat mismatched  
774 with Chl-*a*, these dynamics would then account for the observed spatial differences in  
775 mesozooplankton biovolume — and notably the high secondary production on the leeward (west)  
776 flank. Acoustic surveys also revealed that micronekton organisms densely aggregate in the upper layer  
777 over the summit, making them available as prey to the diverse benthic fish communities of La  
778 Pérouse, as well as migratory pelagic fish such as albacore and swordfish commonly associated with  
779 seamounts.

780 However in view of these encouraging observations, we could not demonstrate a sustained and  
781 remarkable biological enrichment around La Pérouse. No evidence could be found in the cruise data  
782 that suggests a Taylor column could exist there. Certainly, theoretical calculations, based on the shape  
783 of the seamount, summit depth and impinging circulation, support this finding. Moreover, and

784 contrary to the existing body of literature that frequently relates high productivity with seamounts, our  
785 observations suggest a far more complex and dynamic environment involving spatio-temporal  
786 variability in physical parameters across a range of scales that drive enrichment processes around  
787 seamounts. This highlights experimental design as critical for seamount studies. A full understanding  
788 of physical processes and biological responses requires synoptic and long-term measurements,  
789 repeated hydrological and acoustic transects in different directions from offshore to the seamount  
790 summit, at different seasons, and the development of high-resolution mechanistic and coupled  
791 biophysical models.

792

793

#### 794 **Acknowledgements**

795

796 We thank the officers and crew of the R/V *Antea* for their assistance during the La Pérouse cruise,  
797 which was achieved thanks to the financial support from the Conseil Régional de La Réunion, the  
798 Flotte Océanographique Française (French Oceanographic Fleet) and the Institut de Recherche pour le  
799 Développement (IRD). We are also grateful to the UMS Imago for the logistical support provided in  
800 cruise preparation, to Pierre Rousselot for his assistance in the processing of ADCP data, and to  
801 François Baurand and Sandrine Hillion who performed the chemical analyses (nutrients and pigments)  
802 used in this paper. Finally, we thank the reviewers who improved the quality of the initial manuscript.  
803 Pavanee Annasawmy was funded through a doctoral grant provided by the IRD and the ICMASA  
804 French-South African International Laboratory. Part of this work was supported by NE/P021050/01 -  
805 Sustainable Oceans, Livelihoods and food Security Through Increased Capacity in Ecosystem research  
806 in the Western Indian Ocean (SOLSTICE-WIO) funded by the UKRI Global Challenges Research  
807 Fund (GCRF).

808

809

#### 810 **References**

811

812 Abecassis, M., Dewar, H., Hawn, D., Polovina, F. (2012). Modeling swordfish daytime vertical habitat  
813 in the North Pacific Ocean from pop-up archival tags. *Mar Ecol Prog Ser*, 452, 219-236.

814 Annasawmy, P., Ternon, J.F., Marsac, F., Cherel, Y., Béhagle, N., Roudaut, G., Lebourges-Dhaussy,  
815 A., Demarcq, H., Moloney, C.L., Jaquemet, S., Ménard, F. (2018). Micronekton diel migration,  
816 community composition and trophic position within two biogeochemical provinces of the South  
817 West Indian Ocean: Insight from acoustics and stable isotopes. *Deep-Sea Res Part I*, 138, 85–97.

818 Annasawmy, P., Cherel, Y., Romanov, E., Le Loc'h, F., Ménard, F., Ternon, J.F., Marsac, F. (2020a).  
819 Stable isotope patterns of mesopelagic communities over two shallow seamounts of the south-  
820 western Indian Ocean. *Deep-Sea Res Part II* (this issue).

821 Annasawmy, P., Ternon, J-F., Cotel, P., Demarcq, H., Cherel, Y, Romanov, E., Roudaut, G.,  
822 Lebourges-Dhaussy, A., Ménard, F., Marsac, F. (2020b). Micronekton distribution and  
823 assemblages at two shallow seamounts in the south-western Indian Ocean: Insights from acoustics  
824 and mesopelagic trawl data. *Prog Oceanogr*, 178, 102161.

825 Annasawmy, P., Ternon, J-F., Lebourges-Dhaussy, A., Roudaut, G., Herbette, S., Ménard, F., Cotel,  
826 P., Marsac, F. (2020c). Micronekton distribution as influenced by mesoscale eddies, Madagascar  
827 shelf and shallow seamounts in the south-western Indian Ocean: An acoustic approach. *Deep-Sea  
828 Res Part II* (this issue).

829 Baracho-Neto, C.G., Neto, A.S., Rossi-Santos, M., Wedekin, L.L., Neves, M.C., Lima, F., Faria, D.  
830 (2012). Site fidelity and residence times of humpback whales (*Megaptera novaeangliae*) on the  
831 Brazilian coast. *J Mar Biol Assoc UK*, 92, 1783–1791.

832 Barruol, G., Sigloch, K. (2013). Investigating La Réunion hot spot from crust to core. *EOS T Am  
833 Geophys Un*, 94(23), 205-207.

834 Béhagle, N., Du Buisson, L., Josse, E., Lebourges-Dhaussy, A., Roudaut, G., Ménard, F. (2014).  
835 Mesoscale features and micronekton in the Mozambique Channel: an acoustic approach. *Deep-Sea  
836 Res Part II*, 100, 164–173.

837 Bekker, V.E., Shcherbachev, Y.N. (1990). Bathypelagic species of the families Neoscopelidae and  
838 Myctophidae from the Indian Ocean, with a description of a new species of *Diaphus*. *J Ichthyol*,  
839 30, 122-134.

840 Carey, F.G., Robison, B.H. (1981). Daily patterns in the activities of swordfish, *Xiphias gladius*,  
841 observed by acoustic telemetry. *Fish Bull Wash*, 79(2), 277-292.

842 Chapman, D.C, Haidvogel, D.B (1992). Formation of Taylor caps over a tall isolated seamount in a  
843 stratified ocean. *Geophys Astro Fluid*, 64, 31-65.

844 Cherel, Y., Romanov, E., Annasawmy, P.A., Thibault, D., Ménard, F. (2020). Micronektonic fish  
845 species over three seamounts in the southwestern Indian Ocean. *Deep-Sea Res Part II*,  
846 <https://doi.org/10.1016/j.dsr2.2020.104777>

847 Childers, J., Snyder, S. Kohin, S. (2011). Migration and behavior of juvenile North Pacific albacore  
848 (*Thunnus alalunga*). *Fish Oceanogr*, 20, 157-173.

849 Chipps, S. R., Garvey, J. (2007). Quantitative assessment of food habits and feeding patterns, pp. 473–  
850 514. In: Guy C.S and Brown, M.L. (Eds) *Analysis and Interpretation of Freshwater Fisheries Data*.  
851 American Fisheries Society, Bethesda, MD, 971 p.

852 Christiansen, B., Martin, B., Hirsch, S. (2009). The benthopelagic fish fauna on the summit of Seine  
853 seamount, NE Atlantic: Composition, population structure and diets. *Deep-Sea Res Part II*, 56(25),  
854 2705-2712.

855 Clark, M.R., Schlacher, T.A., Rowden, A.A., Stocks, K.I., Consalvey, M. (2012). Science Priorities  
856 for Seamounts: Research Links to Conservation and Management. *Plos One* 7, e29232.

857 Clark, M.R., Vinnichenko, V.I., Gordon, J.D.M., Beck-Bulat, G.Z., Kukharev, N.N., Kakora, A.F.  
858 (2007). Large-scale distant-water trawl fisheries on seamounts. Chap. 17, pp 363-369. In Pitcher,  
859 T.J., Morato, T., Hart, P.J.B., Clark, M.R., Haggan, N. and Santos, R.S. (Eds) *Seamounts:*  
860 *Ecology, Fisheries and Conservation*. Fish and Aquatic Resources Series, Blackwell, Oxford, UK.

861 Clarke, M.R. (1986) *A handbook for the identification of cephalopod beaks*. Clarendon, Oxford, 286  
862 pp.

863 Comeau, L.A., Vezina, A.F., Bourgeois, M., Juniper, K.S. (1995). Relationship between  
864 phytoplankton production and the physical structure of the water column near Cobb Seamount,  
865 Northeast Pacific. *Deep-Sea Res Part I*, 42, 993–1005.

866 De Forest, L., Drazen, J. (2009). The influence of a Hawaiian seamount on mesopelagic micronekton.  
867 *Deep-Sea Res Part I*, 56(2): 232-250.

868 Demarcq, H., Roberts, M.J., Noyon, M. (2020). Phytoplanktonic enhancements depicted from ocean  
869 colour remote-sensing around seamounts: examples from the south western Indian Ocean. *Deep-*  
870 *Sea Res Part II*, (this issue).

871 Demopoulos, A.W.J., Smith, C.R., Tyler, P.A. (2003). *The Deep Indian Ocean Floor*, pp 219-238.  
872 *Ecosystems of the world*. Elsevier.

873 Deplus, C., Bissessur, D., Dymont, J. (2016). Past and present volcanism around La Réunion: insights  
874 from submarine volcanic structures. RHUM-RUM workshop, La Réunion, 28 Aug-2 Sept 2016.

875 De Robertis, A., Higginbottom, I. (2007). A post-processing technique to estimate the signal to-noise  
876 ratio and remove echosounder background noise. *ICES J Mar Sci*, 64, 1282-1291.



877 Diekmann, R., Nellen, W., Piatkowski, U. (2006). A multivariate analysis of larval fish and paralarval  
878 cephalopod assemblages at Great Meteor Seamount. *Deep-Sea Res Part I*, 53, 1635-1657.

879 Dower, J., Freeland, H., Juniper, K., 1992. A strong biological response to oceanic flow past Cobb  
880 Seamount. *Deep-Sea Res Part I*, 39, 1139–1145.

881 Dufois, F., Hardman-Mountford, N.J., Greenwood, J., Richardson, A.J., Feng, M., Herbette, S.,  
882 Matear, R. (2014). Impact of eddies on surface chlorophyll in the South Indian Ocean. *J Geophys*  
883 *Res-Oceans*, 119, 8061–8077.

884 Dulau, V., Pinet, P., Geyer, Y., Fayon, J., Mongin, P., Cottarel, G., Zerbini, A., Cerchio, S. (2017).  
885 Continuous movement behavior of humpback whales during the breeding season in the southwest  
886 Indian Ocean: on the road again! *Mov Ecol*, 5, 11.

887 Durville, P., Durville, S., Ballesta, L., Holon, F., Mulochau, T., Quod, J-P., Condet, M., Martigne, J-  
888 C., Trentin, F. (2019). Exploration du Mont sous-marin La Pérouse (Banc des 90 miles) au large de  
889 l'île de La Réunion. *Expédition La Pérouse. Rapport de mission. Galaxea/Andromède/Vie Océane*.  
890 25 pp.

891 Ehrlich, S., 1977. Die Fischfauna der Großen Meteorbank. *Meteor Forschungsergeb- nisse, Reihe D*  
892 25, 1–23.

893 Fock, H.O., Matthiessen, B., Zidowitz, H., Westernhaen, H.V., (2002). Diel and habitat-dependent  
894 resource utilization by deep-sea fishes at the Great Meteor seamount: niche overlap and support for  
895 the sound scattering layer interception hypothesis. *Mar Ecol Prog Ser*, 244, 219-233.

896 Fonteneau A (1991) Monts sous-marins et thons dans l'Atlantique tropical est. *Aquat Living Resour*,  
897 4, 13–25.

898 Freeland, H. (1994). Ocean circulation at and near Cobb seamount. *Deep-Sea Res Part I*, 41, 1715-  
899 1732.

900 Garrigue, C., Clapham, P.J., Geyer, Y., Kennedy, A.S., Zerbini, A.N. (2015). Satellite tracking reveals  
901 novel migratory patterns and the importance of seamounts for endangered South Pacific humpback  
902 whales. *R Soc Open Sci*, 2,150489.

903 Genin, A., Boehlert, G.W. (1985). Dynamics of temperature and chlorophyll structures above a  
904 seamount: an oceanic experiment. *J Mar Res*, 43, 907-924.

905 Genin, A., Haury, L., Greenblatt, P. (1988). Interactions of migrating zooplankton with shallow  
906 topography: predation by rockfishes and intensification of patchiness. *Deep-Sea Res Part I*, 35,  
907 151-175.

- 908 Gilman, E., Chaloupka, M., Read, A., Dalzell, P., Holetschek, J., Curtice, C. (2012). Hawaii longline  
909 tuna fishery temporal trends in standardized catch rates and length distributions and effects on  
910 pelagic and seamount ecosystems. *Aquat Conserv*, 22, 446-488.
- 911 Gorelova, T. A., Prut'ko, V. G. (1985). Feeding of *Diaphus suborbitalis* (Myctophidae, Pisces) in the  
912 equatorial Indian Ocean. *Oceanology-USSR*, 25(4), 523-529.
- 913 Gorsky, G., Ohman, M.D., Picheral, M., Gasparini, S., Stemmann, L., Romagnan, J.B., Cawood, A.,  
914 Pesant, S., García-Comas, C., Prejger, F. (2010). Digital zooplankton image analysis using the  
915 ZooScan integrated system. *J Plankt Res*, 32, 285–303.
- 916 Hogg, N. G. 1973. On the stratified Taylor column. *J Fluid Mech*, 58, 517-537.
- 917 Holland, K. N., Grubbs, R. D. (2007). Fish visitors to seamounts: Tunas and billfish at seamounts.  
918 Chap 10A, pp 189-201. In Pitcher, T.J., Morato, T., Hart, P.J.B., Clark, M.R., Haggan, N. and  
919 Santos, R.S. (Eds) *Seamounts: Ecology, Fisheries and Conservation*. Fish and Aquatic Resources  
920 Series, Blackwell, Oxford, UK.
- 921 Huppert, H. E. (1975). Some remarks on the initiation of internal Taylor columns. *J Fluid Mech*, 67,  
922 397-412.
- 923 Huppert, H. E., Bryan, K. (1976). Topographically generated eddies. *Deep-sea Res. Part I*, 23: 655-  
924 679.
- 925 Ingole, B., Koslow, J.A. (2005). Deep-sea ecosystems of the Indian Ocean. *Indian J Mar Sci*, 34(1),  
926 27-34.
- 927 Kaschner, K. (2007) Air-breathing visitors to seamounts: Marine Mammals. Chapt 12A. pp 230-238.  
928 In Pitcher, T.J., Morato, T., Hart, P.J.B., Clark, M.R., Haggan, N. and Santos, R.S. (Eds)  
929 *Seamounts: Ecology, Fisheries and Conservation*. Fish and Aquatic Resources Series, Blackwell,  
930 Oxford, UK.
- 931 Kawaguchi, K., Shimizu, H. (1978). Taxonomy and distribution of the lanternfishes, genus *Diaphus*  
932 (Pisces, Myctophidae) in the Western Pacific, Eastern Indian oceans and the Southeast Asian seas.  
933 *Bull Ocean Res Inst Univ Tokyo*, 10, 1-145.
- 934 Kiyofuji, H., Okamoto, S., Ijima, H. (2013). Vertical and horizontal changes of North Pacific albacore  
935 derived from archival tag data. ISC Albacore Working Group Intercessional Workshop, 5-12  
936 November 2013, Nat. Res. Inst. Far Seas Fish., Shimizu, Shizuoka, Japan.
- 937 Lavelle, J.W., Thurnherr, A.M., Ledwell, J.R., McGillicuddy Jr., D.J., Mullineaux, L.S. (2010), Deep  
938 ocean circulation and transport where the East Pacific Rise at 9–10°N meets the Lamont seamount  
939 chain, *J Geophys Res-Oceans*, 115, C12073.

940 Lebourges-Dhaussy, A., Marchal., E., Menkès, C., Champalbert, G., Biessy, B. (2000). *Vinciguerria*  
941 *nimbaria* (micronekton), environment and tuna: their relationships in the eastern Tropical Atlantic.  
942 *Oceanol Acta*, 23(4), 515-528.

943 Le Bot, P., Kermabon, C., Lherminier, P. Gaillard, F. (2011). Chaîne automatisée de suivi des  
944 courantomètres acoustiques Doppler embarqués. Cascade V6.1 : logiciel de validation et de  
945 visualisation des mesures ADCP de coque. Rapport OPS/LPO 11-01, Ifremer, France, 93 p.

946 Le Corre, M., Jaeger, A., Pinet, P., Kappes, M.A., Weimerskirch, H., Catry, T., Ramos, J.A., Russell,  
947 J.C., Shah, N., Jaquemet, S. (2012). Tracking seabirds to identify potential Marine Protected Areas  
948 in the tropical western Indian Ocean. *Biol Conserv*, 83–93.

949 Longhurst, A.R. (1998). *Ecological geography of the sea*. Academic Press, 398 pp.

950 Lu, C.C., Ickeringill, R. (2002). Cephalopod beak identification and biomass estimation techniques:  
951 tools for dietary studies of southern Australian finfishes. *Mus Victoria Sci Rep*, 6, 1-65.

952 Mannocci, L., Laran, S., Monestiez, P., Dorémus, G., Van Canneyt, O., Watremez, P., Ridoux, V.  
953 (2014). Predicting top predator habitats in the Southwest Indian Ocean. *Ecography*, 37, 261–278.

954 Marsac, F., Fonteneau, A., Michaud, P. (2014). Le coco de mer, une montagne sous la mer, Chap 7,  
955 pp. 151-163. In : Marsac, F., Fonteneau, A. and Michaud, P., *L’or bleu des Seychelles: histoire de*  
956 *la pêche industrielle au thon dans l’océan Indien*. IRD Editions, 269 pp.

957 Mohn, C., White, M. (2007). Remote sensing and modelling of bio-physical distribution patterns at  
958 Porcupine and Rockall Bank, Northeast Atlantic. *Cont Shelf Res*, 27, 1875-1892.

959 Mohn, C., White, M., Bashmachnikov, I., Jose, F., Pelegri, J.L. (2009). Dynamics at an elongated,  
960 intermediate depth seamount in the North Atlantic (Sedlo Seamount, 40°20’N, 26°40’W). *Deep-*  
961 *Sea Res Part II*, 56, 2582–2592.

962 Monteiro, L.R., Ramos, J.R., Furness, R.W., del Nevo, A.J. (1996). Movements, morphology,  
963 breeding, molt diet and feeding of seabirds in the Azores. *Colon Waterbird*, 19, 82–97.

964 Morato, T., Clark, M.R. (2007). Seamount fishes: ecology and life histories. Chapt 9, pp 170-188. In  
965 Pitcher, T.J., Morato, T., Hart, P.J.B., Clark, M.R., Haggan, N. and Santos, R.S. (Eds) *Seamounts:*  
966 *Ecology, Fisheries and Conservation*. Fish and Aquatic Resources Series, Blackwell, Oxford, UK.

967 Morato, T., Hoyle, S.D., Allain, V., Nicol, S.J. (2010). Tuna longline fishing around West and Central  
968 Pacific seamounts. *Plos One*, 5(12), e14453.

969 Morato, T., Varkey, D. A., Damaso, C., Machete, M., Santos, M., Prieto, A., Santos, R.S., Pitcher, T.  
970 (2008). Evidence of a seamount effect on aggregating visitors. *Mar Ecol Prog Ser*, 357, 23-32.

- 971 Mouriño, B., Fernandez, E., Serret, P., Harbour, D., Sinha, B., Pingree, R. (2001). Variability and  
972 seasonality of physical and biological fields at the Great Meteor Tablemount (subtropical NE  
973 Atlantic). *Oceanol Acta*, 24(2), 167-185.
- 974 Murray, T. (1994). A review of the biology and fisheries for albacore, *Thunnus alalunga*, in the South  
975 Pacific Ocean, pp 188-206. In: Shomura, R.S., Majkowski, J. and Langi, S. (Eds) Interaction of  
976 Pacific tuna fisheries. FAO Fish. Tech. Pap., 336(2), 439 p.
- 977 Noyon, M., Rasoloarijao, Z., Huggett, J., Ternon, J-F., Roberts, M.J. (2020). Comparison of  
978 mesozooplankton communities at three shallow seamounts in the South West Indian Ocean. *Deep-*  
979 *Sea Res Part II*, this issue, <https://doi.org/10.1016/j.dsr2.2020.104759>
- 980 Odate, T., Furuya, K. (1998). Well-developed subsurface chlorophyll maximum near Komahashi No.  
981 2 Seamount in the summer of 1991. *Deep-Sea Res Part I*, 45, 1595–1607.
- 982 Parin, N. V., Timokhin, I. G., Novikov, N. P., Shcherbachev, Y. N. (2008). On the composition of  
983 talassobathyal ichthyofauna and commercial productivity of Mozambique Seamount (the Indian  
984 Ocean). *J Ichthyol*, 48(5), 361–366.
- 985 Perrot Y., Brehmer P., Habasque J., Roudaut G., Béhagle N., Sarre A., Lebourges-Dhaussy A. (2018).  
986 Matecho: an open-source tool for processing fisheries acoustics data. *Acoust Aust*, 46(2), 241-248.
- 987 Picheral, M., Colin, S., Irisson, J-O. (2017). EcoTaxa, a tool for the taxonomic classification of  
988 images. URL <http://ecotaxa.obs-vlfr.fr>
- 989 Pinet, P., Jaquemet, S., Phillips, R.A., Le Corre, M. (2012). Sex-specific foraging strategies  
990 throughout the breeding season in a tropical, sexually monomorphic small petrel. *Anim Behav*, 83,  
991 979-989.
- 992 Pinkas, L., Oliphant, M.S., Iverson, I.L.K. (1971). Food habits of albacore, bluefin tuna, and bonito in  
993 California waters. *Fish Bull US*, 152, 1-105.
- 994 Pitcher, T.J., Bulman, C. (2007). Raiding the larder: a quantitative evaluation framework and trophic  
995 signature for seamount food webs. Chapt 14, pp 282-295. In Pitcher, T.J., Morato, T., Hart, P.J.B.,  
996 Clark, M.R., Haggan, N. and Santos, R.S. (Eds) *Seamounts: Ecology, Fisheries and Conservation*.  
997 *Fish and Aquatic Resources Series*, Blackwell, Oxford, UK.
- 998 Potier, M., Marsac, F., Cherel, Y., Lucas, V., Sabatie, R., Maury, O., Ménard, F. (2007). Forage fauna  
999 in the diet of three large pelagic fish (lancetfish, swordfish and yellowfin tuna) in the western  
1000 equatorial Indian Ocean. *Fish Res*, 83, 60-72.
- 1001 Potier, M., Ménard, F., Benivary, H. D., Sabatié, R. (2011). Length and weight estimates from  
1002 diagnostic hard part structures of fish, crustacea and cephalopods forage species in the western  
1003 Indian Ocean. *Environ Biol Fish*, 92, 413–423.

- 1004 Roberts, M.J., Ternon, J-F., Marsac, F., Noyon, M. (2020). The MADRidge Project : Bio-physical  
1005 coupling around three shallow seamounts in the South Western Indian Ocean. Deep-Sea Res Part II  
1006 (this issue).
- 1007 Rogers, A.D., Clark, M.R., Hall-Spencer, J.M., Gjerde, K.M. (2008). The science behind the  
1008 guidelines: a scientific guide to the FAO draft international guidelines (December 2007) for the  
1009 management of deep-sea fisheries in the High Seas and examples of how the guidelines may be  
1010 practically implemented. IUCN, Switzerland.
- 1011 Romanov, E.V. (2003). Summary and review of Soviet and Ukrainian scientific and commercial  
1012 fishing operations on the deepwater ridges of the southern Indian Ocean. FAO Fisheries Circular.  
1013 No. 991. FAO, Rome, 84 pp.
- 1014 Romanov, E., Sabarros, P.S., Le Foulgoc, L., Bach, P. (2016). A preliminary analysis of swordfish  
1015 (*Xiphias gladius*) habitat and behaviour on migratory track from Reunion Island to equatorial  
1016 waters. IOTC Working Party on Billfish (WPB) Victoria, Seychelles, 06-10 September 2016.  
1017 IOTC–2016–WPB14–16. 20 p.
- 1018 Romanov, E.V., Nikolic, N., Dhurmeea, Z., Bodin, N., Puech, A., Norman, S., Hollanda, S., Bourjea,  
1019 J., West W., Potier M. (2020). Trophic ecology of albacore tuna *Thunnus alalunga* in the western  
1020 tropical Indian Ocean and adjacent waters. Mar Fresh Res, in press.
- 1021 Ryan, T.E., Downie, R.A., Kloser, R.J., Keith, G. (2015). Reducing bias due to noise and attenuation  
1022 in open-ocean echo integration data. ICES J Mar Sci, 72, 2482-2493.
- 1023 Santos, M.A., Bolton, A.B., Martins, H.R., Riewald, B., Bjorndal, K.A. (2007). Air-breathing visitors  
1024 to seamounts: sea turtles. Chapt 12B, pp 239-244. In Pitcher, T.J., Morato, T., Hart, P.J.B., Clark,  
1025 M.R., Haggan, N. and Santos, R.S. (Eds) Seamounts: Ecology, Fisheries and Conservation. Fish  
1026 and Aquatic Resources Series, Blackwell, Oxford, UK.
- 1027 Shcherbachev, A. N. Kotlyar, Abramov, A.A (1989). Ichthyofauna and Fish Stock of Seamounts of  
1028 the Indian Ocean, pp 159-185. In: Biological Resources of the Indian Ocean (Nauka, Moscow).
- 1029 Scott, W.B., Tibbo, S.N. (1968). Food and feeding habits of swordfish, *Xiphias gladius*, in the western  
1030 North Atlantic. J Fish Res Board Can, 25, 903-919.
- 1031 Smale, M.J., Watson, G, Hecht, T. (1995). Otolith Atlas of southern African marine fishes.  
1032 Ichthyological Monographs, 1. J.L.B. Smith Institute of Ichthyology, Grahamstown, 420 pp.
- 1033 Stillwell, C.E., Kohler, N.E. (1985). Food and feeding ecology of the swordfish *Xiphias gladius* in the  
1034 western North Atlantic Ocean with estimates of daily ration. Mar Ecol Prog Ser, 22, 239-247.
- 1035 Sprintall, J., Cronin, M.F. (2009). Upper ocean vertical structure, pp 217-224. Encyclopedia of Ocean  
1036 Sciences (Second Edition), Academic Press, 3900 p.

- 1037 Taylor, G. I. (1923). Experiments on the motion of solid bodies in rotating fluids. Proc R Soc. A, 104,  
1038 213-218.
- 1039 Tomczak, M., Godfrey, J.S. (1994). The Indian Ocean. In: Regional Oceanography: An Introduction.  
1040 Elsevier, UK, 391 p.
- 1041 Trenkel, V.M., Berger, L., Bourguignon, S., Doray, M., Fablet, R., Massé, J., Mazauric, V., Poncet,  
1042 C., Quemener, G., Scalabrin, C., Villalobos, H. (2009). Overview of recent progress in fisheries  
1043 acoustics made by Ifremer with examples from the Bay of Biscay. Aquat Living Resour, 22, 433-  
1044 445.
- 1045 Tsuchiya, K., Piatkowski, U., Okutani, T. (1991). Distribution and re-description of *Abraliopsis*  
1046 *lineata* Goodrich, 1896 (Cephalopoda, Enoploteuthidae) from the Arabian Sea. J Nat Hist, 25(5),  
1047 1121–1133.
- 1048 Vandromme, P., Lars, S., García-Comas, C., Berline, L., Sun, X., Gorsky, G. (2012). Assessing biases  
1049 in computing size spectra of automatically classified zooplankton from imaging systems: A case  
1050 study with the ZooScan integrated system. Meth Oceanogr. 1–2, 3–21.
- 1051 Visbeck M. (2001). Deep velocity profiling using lowered acoustic Doppler current profilers: bottom-  
1052 track and inverse solutions. J Atmos Ocean Tech, 19, 794-807.
- 1053 White, M., Bashmachnikov, I., Arístegui, J., Martins, A. (2007) Physical processes and seamount  
1054 productivity. Chapter 4., pp 65-84. In Pitcher, T.J., Morato, T., Hart, P.J.B., Clark, M.R., Haggan,  
1055 N. and Santos, R.S. (Eds) Seamounts: Ecology, Fisheries and Conservation. Fish and Aquatic  
1056 Resources Series, Blackwell, Oxford, UK.
- 1057 Williams, A.J., Allain, V., Nicol, S.J., Evans, K.J., Hoyle, S.D., Dupoux, C., Vourey, E., Dubosc, J.  
1058 (2015). Vertical behavior and diet of albacore tuna (*Thunnus alalunga*) vary with latitude in the  
1059 South Pacific Ocean. Deep-Sea Res Part II, 113, 154-169.

## List of figures

1060

1061

1062 **Figure 1.** Location of the La Pérouse seamount on a regional map. Red diamond symbols indicate  
1063 stomach samples associated with La Pérouse seamount. Blue circles denote the open-ocean samples.  
1064 (a); 3-D view of the seamount showing the crescent-shaped summit (b); high resolution bathymetric  
1065 map of La Pérouse and immediate surroundings, measured during the Antea-Pérouse cruise, 15-29  
1066 September 2016 (c), showing the location of 11 hydrological stations and zooplankton tows (red  
1067 diamonds), and 10 midwater trawl hauls (blue lines).

1068 **Figure 2.** Sea level anomalies on 19 September (a) and 3 October 2016 (b) with overlaid geostrophic  
1069 currents (red arrows) deduced from satellite altimetry. The velocity scale of the arrows is indicated in  
1070 the legend. Cyclonic eddies are represented with dashed contours. Anticyclonic eddies with solid  
1071 lines. Stations at La Pérouse are represented by blue triangles.

1072 **Figure 3.** Vertical sections for S-ADCP total velocity from Réunion to La Pérouse, 16-18 September  
1073 2016 (a) and from La Pérouse to Réunion, 28-29 September 2016 (b).

1074 **Figure 4.** Box and whisker plots for current velocity determined by the L-ADCP, in the upper 200 m  
1075 (a) and upper 400 m (b) for both the ‘Offshore’ (control) and the “Flank” stations.

1076 **Figure 5.** Grouped vertical temperature (left panels, a-c) and chlorophyll-a (right panels, d-f) profiles  
1077 (1 = Offshore; 2-3 = Flank). The dotted lines indicate the MLD in the left panels and the depth of the  
1078 nutricline in the right panels for each station

1079 **Figure 6.** Spatial distribution of zooplankton biovolume at 10 stations (summit station 5 not sampled).  
1080 Note the relatively high biovolume at station 7 is biased by the presence of a single large-size  
1081 Trachymedusa not recorded in any of the other Bongo net samples.

1082 **Figure 7.** Zooplankton biovolume as a function of the integrated Chl-*a* in the upper 200 m. Numbers  
1083 indicate the stations where Bongo nets tows were deployed. ‘Flank’ stations are denoted by grey  
1084 circles and ‘Offshore’ stations by triangles. The linear fit is shown for the ‘Flank’ stations only.

1085 **Figure 8.** Location of acoustic transects in the ‘Vicinity’ and ‘Offshore’ zones (day= yellow; night=  
1086 cyan) and on the ‘Summit and flank’ of La Pérouse (day= red; night= blue) (a). Micronekton average  
1087 acoustic density (NASC,  $m^2 nmi^{-2}$ ) by transect during the day (b) and at night (c).

1088 **Figure 9.** Box and whisker plot for the overall micronekton acoustic densities, in the day and at night,  
1089 for the aforementioned groups of acoustic transects in Fig. 12a.

1090 **Figure 10.** Typical echograms recorded on the flanks and summit of La Pérouse during daylight (a)  
1091 and at night (b).

1092



1093 Table 1 – Summary of hydrological stations. Note that local time during the cruise is GMT+3. Stations  
 1094 are categorized as offshore (O) or flank (F) based on the distance to the seamount (70 m depth  
 1095 contour).

Station	Date	Local time	Latitude	Longitude	Dist. to seamount (km)	Position attribute
1	21/09/2016	11:09:58	19°32.69' S	54°11.95' E	14.4	O
2	22/09/2016	12:53:29	19°49.58' S	54°05.93' E	9.9	O
3	23/09/2016	06:24:18	19°40.06' S	54°09.09' E	1.5	F
4	23/09/2016	09:13:42	19°42.02' S	54°11.29' E	1.3	F
5	23/09/2016	12:29:08	19°43.67' S	54°09.70' E	1.1	F
6	23/09/2016	13:39:31	19°44.95' S	54°06.30' E	3.2	F
7	23/09/2016	17:52:30	19°46.34' S	54°09.38' E	2.2	F
8	25/09/2016	06:15:47	19°42.55' S	53°56.55' E	21.1	O
10	26/09/2016	06:22:31	19°42.59' S	54°20.02' E	20	O
23	27/09/2016	11:35:11	19°41.90' S	54°06.50' E	2.2	F
24	27/09/2016	18:23:10	19°44.76' S	54°11.06' E	1.1	F

1096

1097 Table 2 – Summary of trawl stations at La Pérouse and shortest distance to the seamount (70 m depth  
 1098 contour). Position attributes: S=Summit, V=Vicinity, F=Flank, O=Offshore

Trawl #	Latitude	Longitude	Max trawl depth (m)	Distance (km)	Position attribute	Start/end local time	Day/Night	Trawling speed (knots)
1	19° 43'S	54° 04'E	590	6.00	V	17:10 18:55	N	3.0
2	19° 44'S	54° 06'E	400	4.20	V	17:20 18:55	N	3.2
3	19° 39'S	53° 52'E	90	28.4	O	18:30 20:00	N	2.7
4	19° 39'S	54° 10'E	110	4.00	V	18:15 19:45	N	3.0
5	19° 44'S	54° 09'E	35	0	S	08:33 09:35	D	3.6
6	19° 42'S	54° 12'E	60	1.15	F	18:44 19:54	N	3.0
7	19° 36'S	54° 08'E	500	7.70	V	17:59 19:50	N	3.0
8	19° 48'S	54° 07'E	430	6.45	V	16:21 18:30	N	3.2
9	19° 49'S	54° 07'E	240	6.76	V	19:05 20:37	N	3.2
10	19° 45'S	54° 15'E	250	7.60	V	19:25 21:10	N	4.0

1099

1100 Table 3 – Summary statistics of stomachs sampled from six different top predator species

Predators	La Pérouse vicinity		Open-ocean	
	3-10 km		75-280 km	
	n	size range (cm)	n	size range (cm)
<i>Thunnus alalunga</i>	5	95-106	31	88-111
<i>Thunnus albacares</i>			3	71-157
<i>Thunnus obesus</i>			5	76-154
<i>Xiphias gladius</i>			9	122-191
<i>Coryphaena hippurus</i>	3	80-98	2	94-100
<i>Alepisaurus ferox</i>	4	77-123	16	56-145
	<b>12</b>		<b>81</b>	

1101

1102 Table 4 – Spatial stratification by distance to the seamount summit used in this study, across the  
 1103 different data sets. The numbers indicate the number of samples, except for acoustic data where the  
 1104 cumulative transect lengths is given (km). The trawl data are not included in this stratification scheme  
 1105 because not used for the spatial analysis, neither fisheries data as a 1° grid was applied for the entire  
 1106 South West Indian Ocean. Acoustic data were not discriminated between summit and flank.

Data sets	Summit	Flank	Vicinity	Offshore	Open-ocean
		0 to 3 km	3 to 10 km	10 to 35 km	75-280 km
Hydrological stations		7		4	
Bongo nets		6		4	
Acoustic surveys	Day	95 km	162 km		
	Night	29 km	110 km	123 km	
Stomachs			12		81

1107

1108

1109 Table 5– Vertical Directional Change Index (VDCI) and deviation to the mean (specific to each group)  
 1110 by station and location ('Offshore' vs 'Flank') for two depth ranges. The highest two VDCI values are  
 1111 indicated in bold.

1112

Location	Station no	Bottom depth (m)	Range 0-200 m		Range 0-400 m	
			VDCI	Deviation to the mean	VDCI	Deviation to the mean
Offshore	1	3968	75	-63	576	+36
	2	2386	74	-63	317	-223
	8	4203	108	-29	512	-28
	10	1654	291	+154	754	+215
Flank	3	1254	<b>630</b>	+208	<b>1255</b>	+412
	23	1260	152	-272	<b>914</b>	+71
	6	1105	240	-182	517	-326
	7	1400	473	+50	814	-29
	24	895	483	+61	835	-8
	5	200	<b>686</b>	+264	-	-
	4	550	294	-129	722	-121

1113

1114 Table 6 – The percentage composition of prey items found in top predator stomachs, expressed as the  
 1115 index of relative importance (IRI) and mean reconstituted weight (MRW) of each large taxonomic  
 1116 group

	La Pérouse vicinity		Open-ocean	
	%IRI	%MRW	%IRI	%MRW
Cephalopods	15.8	41.8	40.3	41.5
Other molluscs	2.1	0.2	3.0	1.4
Crustaceans	28.4	19.7	11.9	4.0
Fish	53.7	38.3	44.6	53.1
Other prey, incl. plastic	0	0	0.1	<0.1

1117

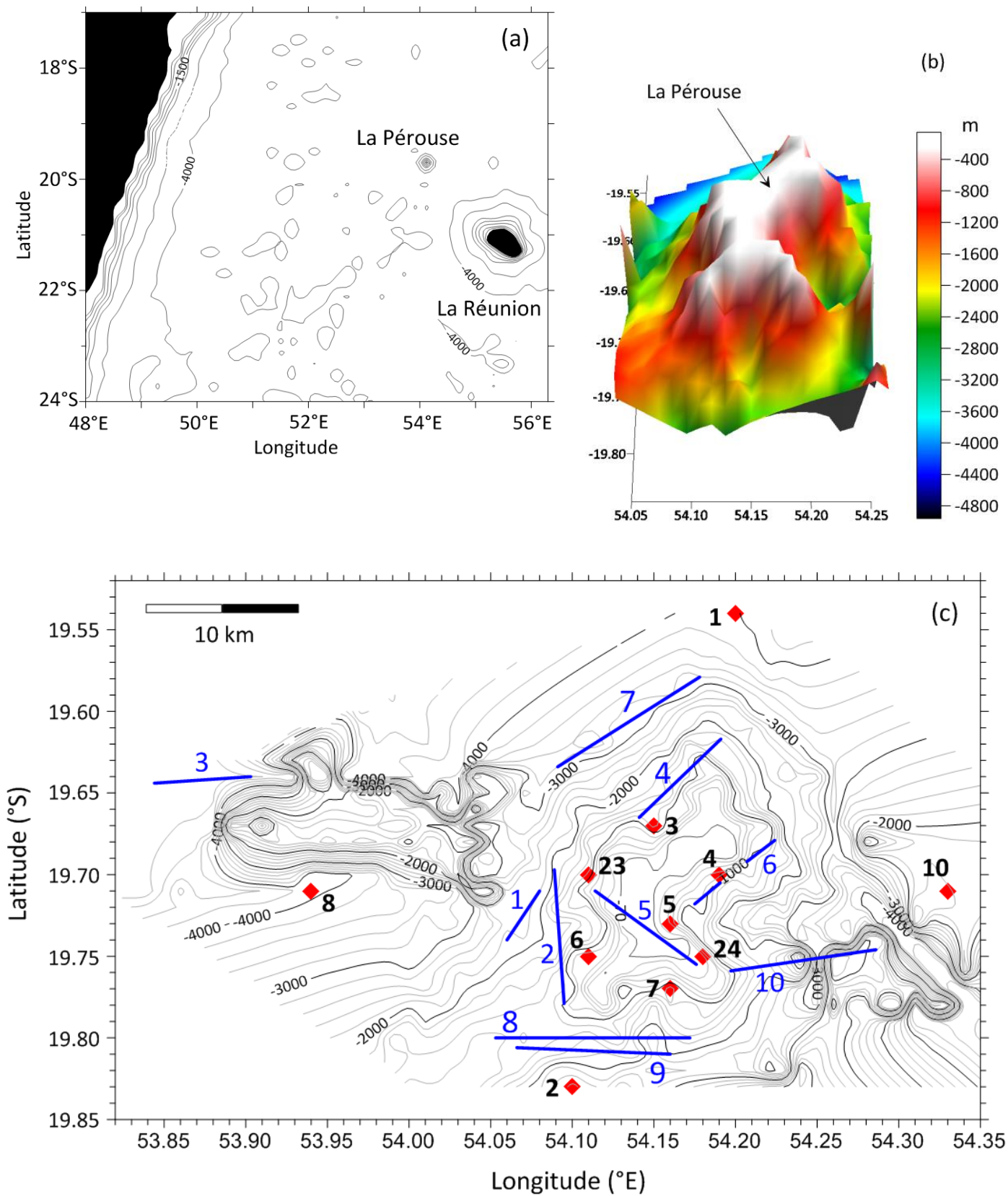


Figure 1

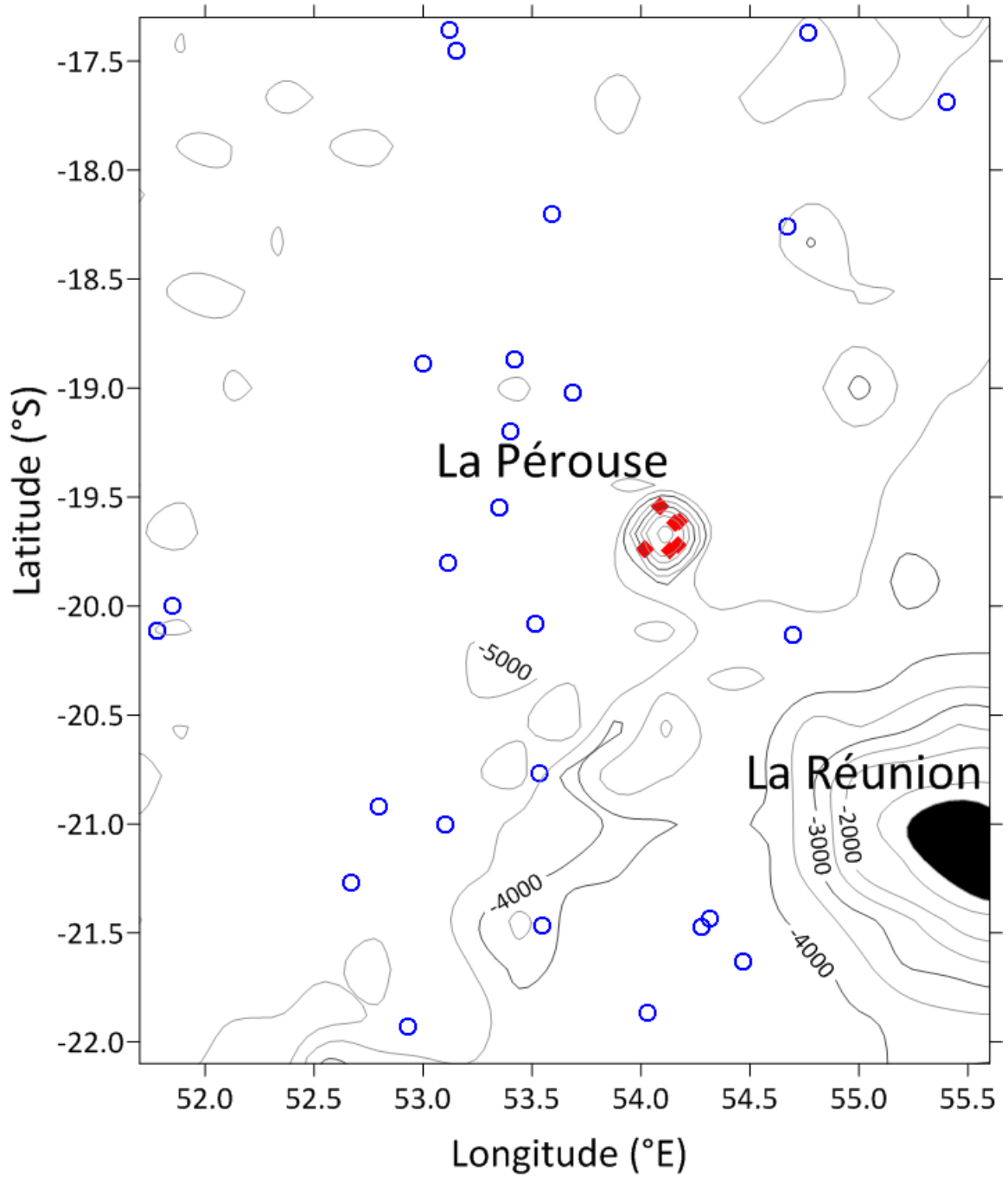


Figure 2

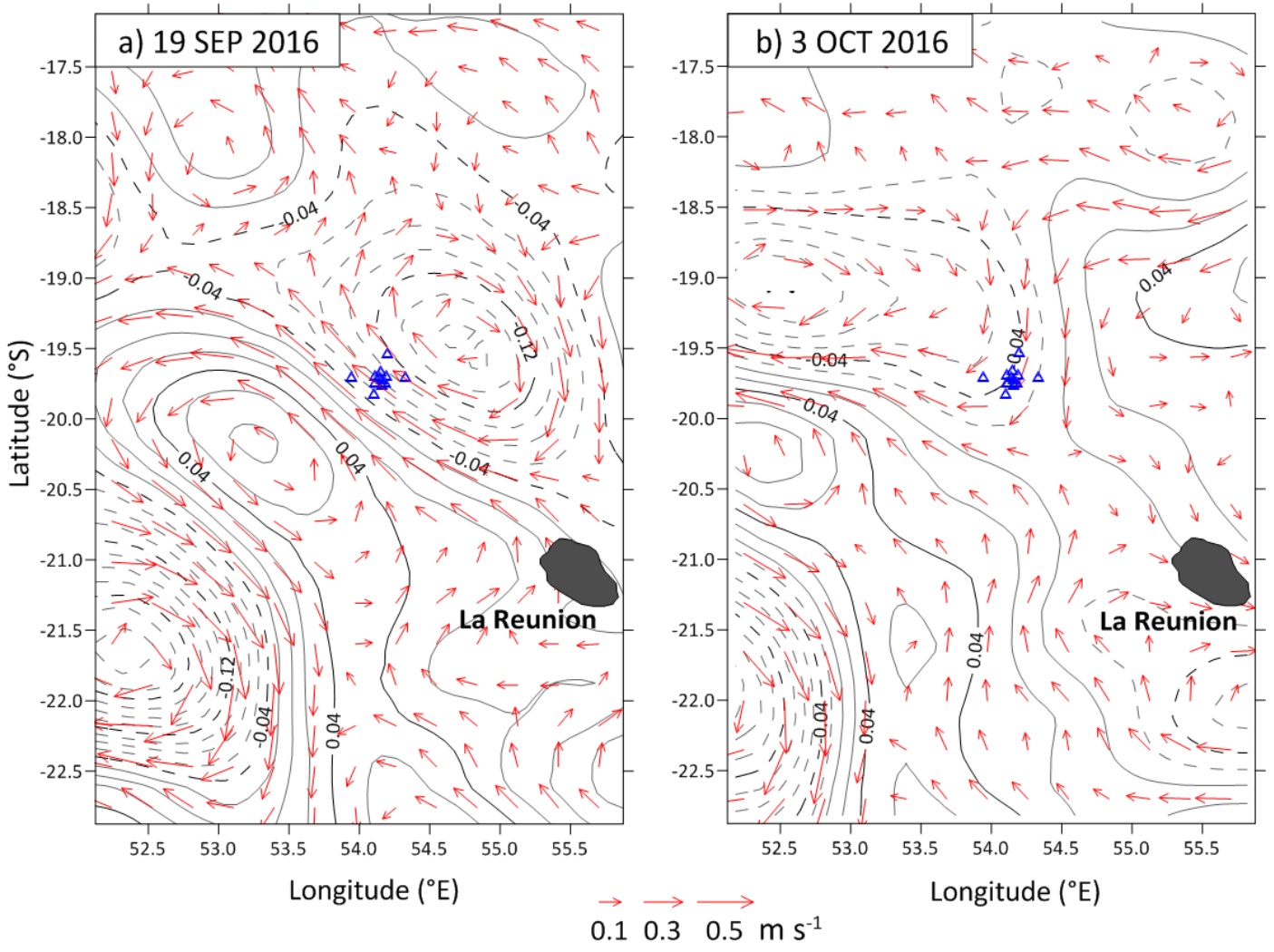


Figure 3

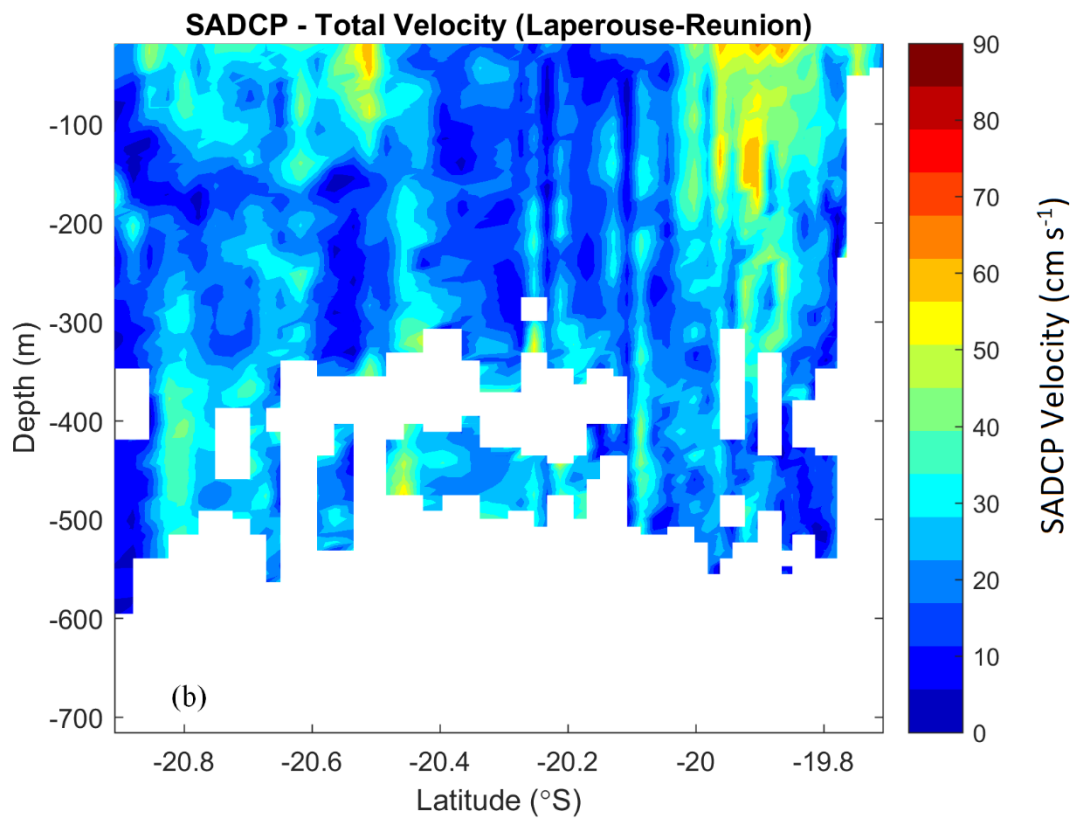
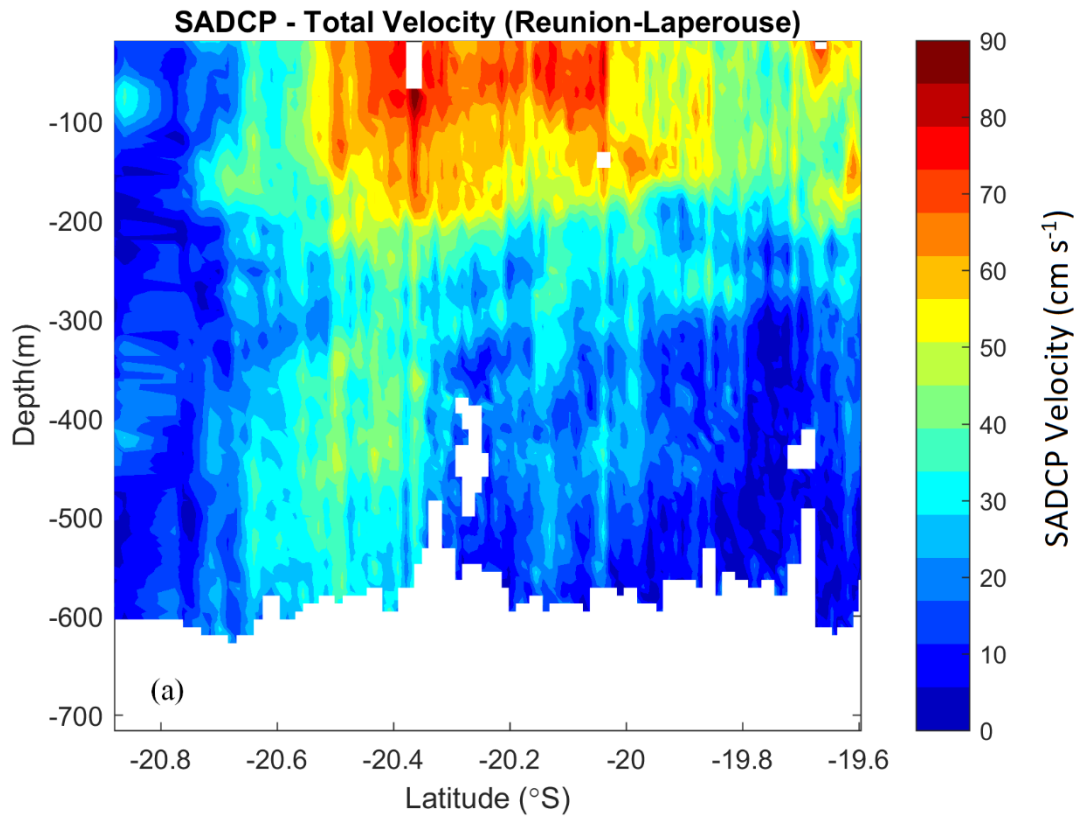


Figure 4

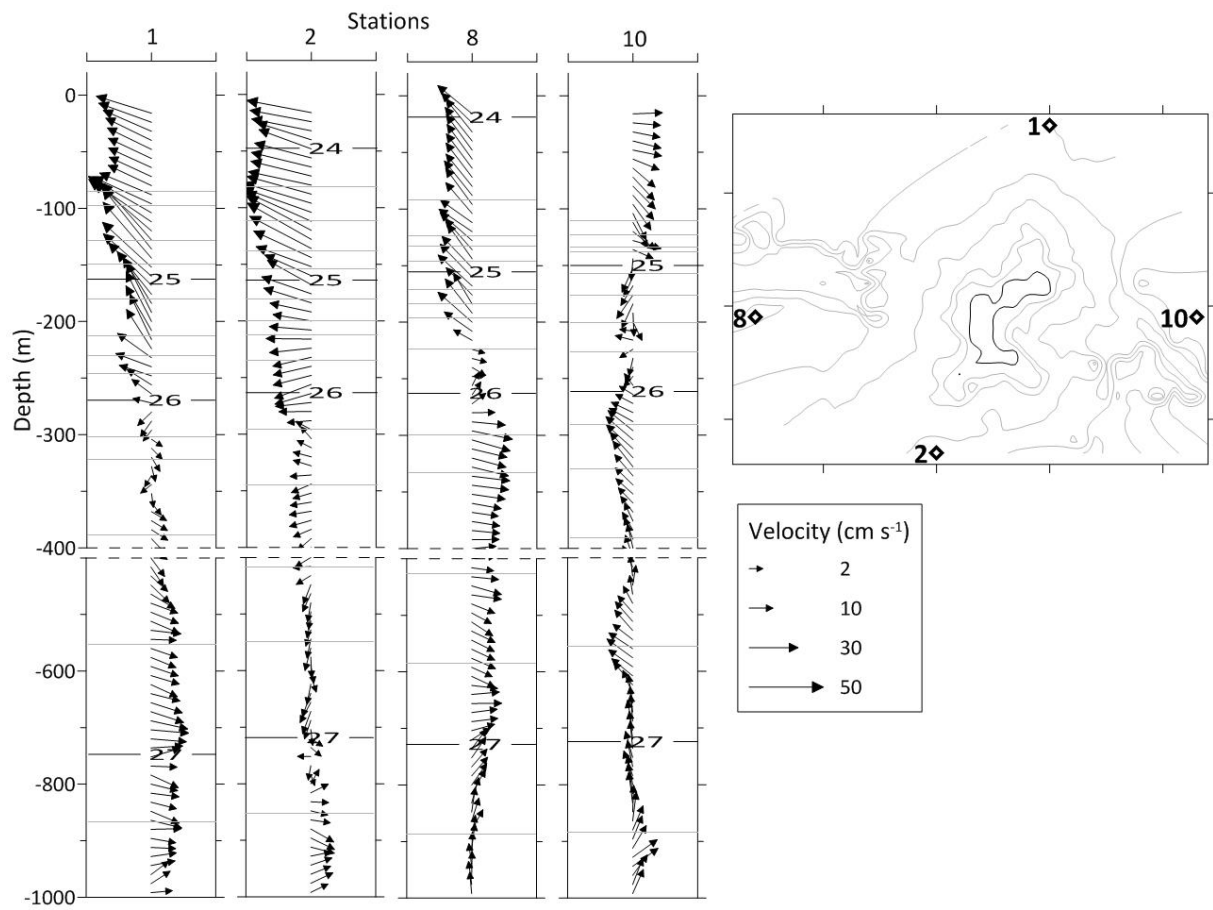


Figure 5

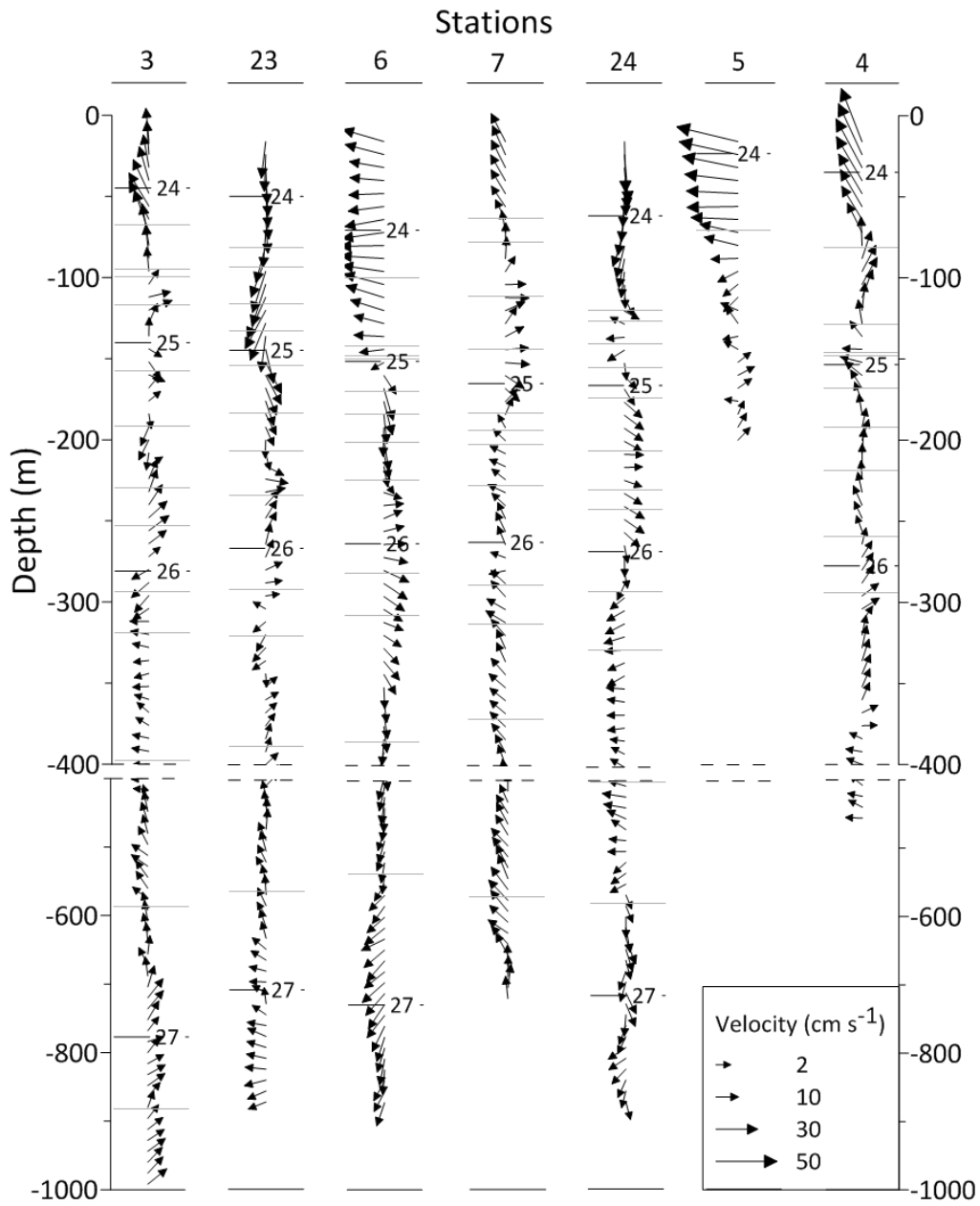
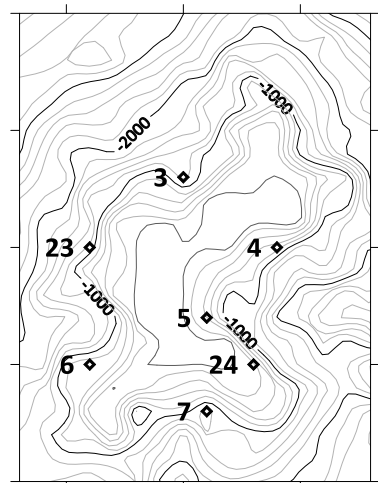


Figure 6





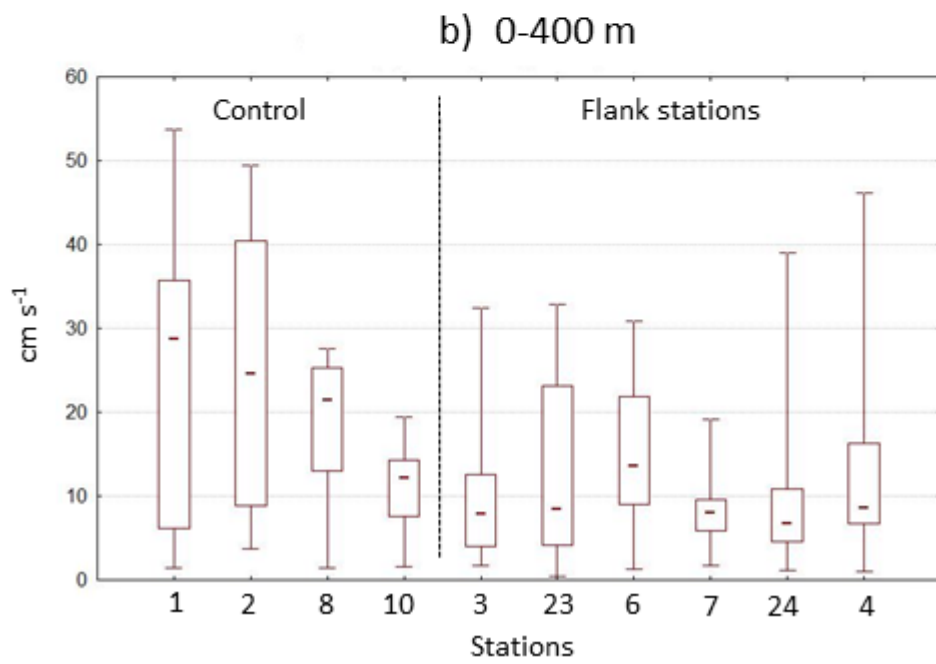
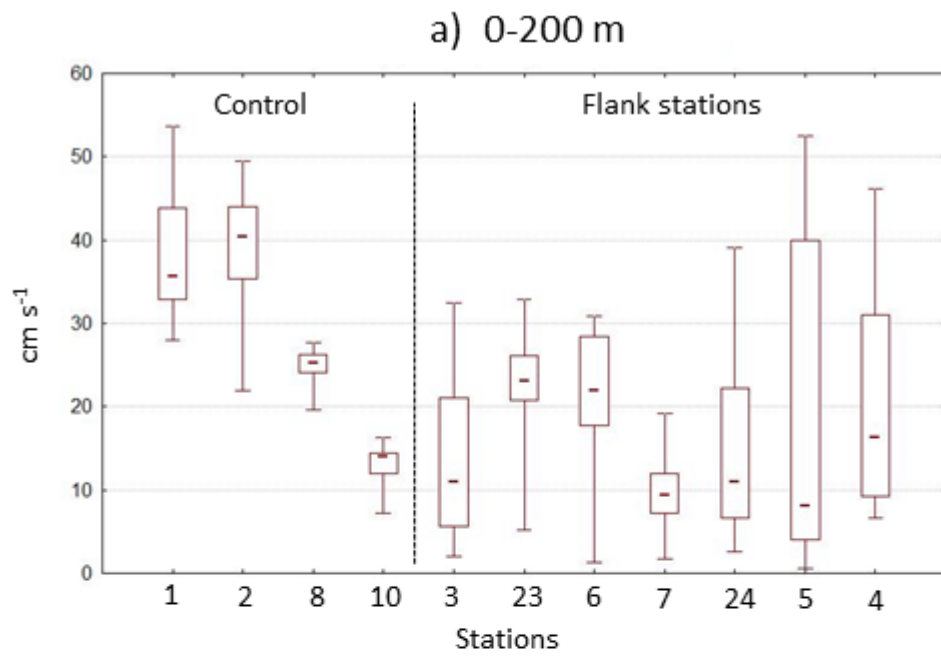


Figure 7

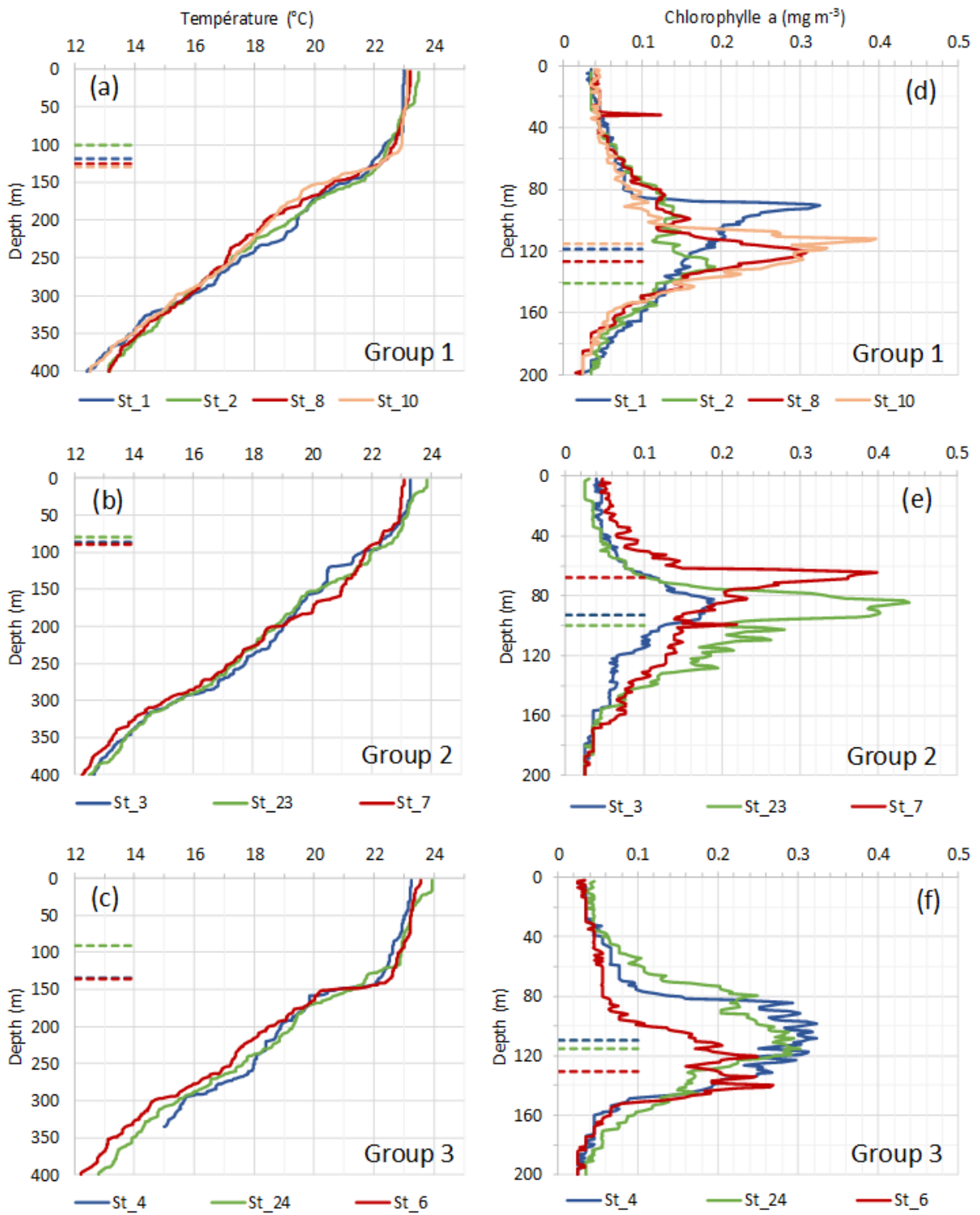


Figure 8

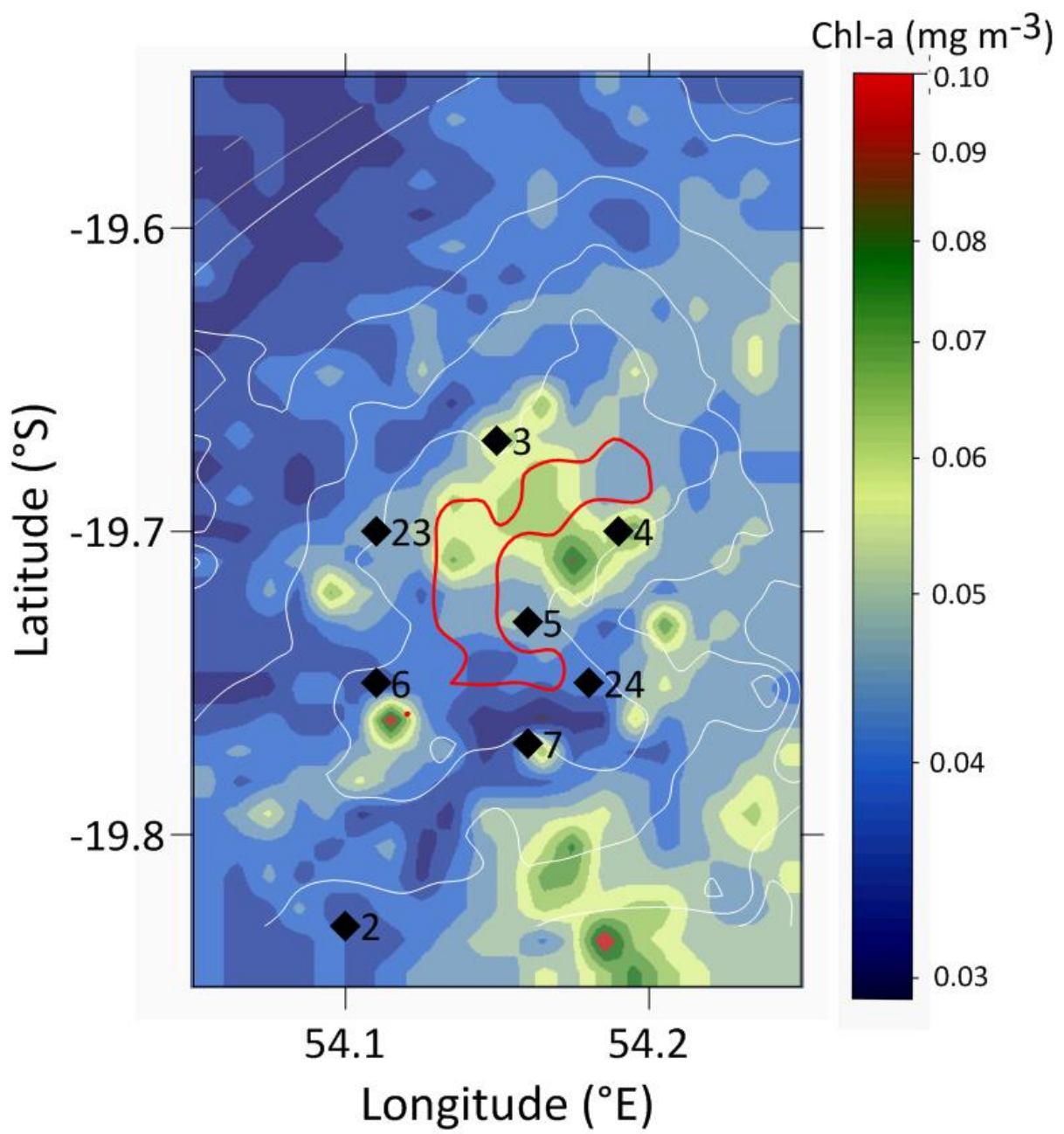


Fig. 9

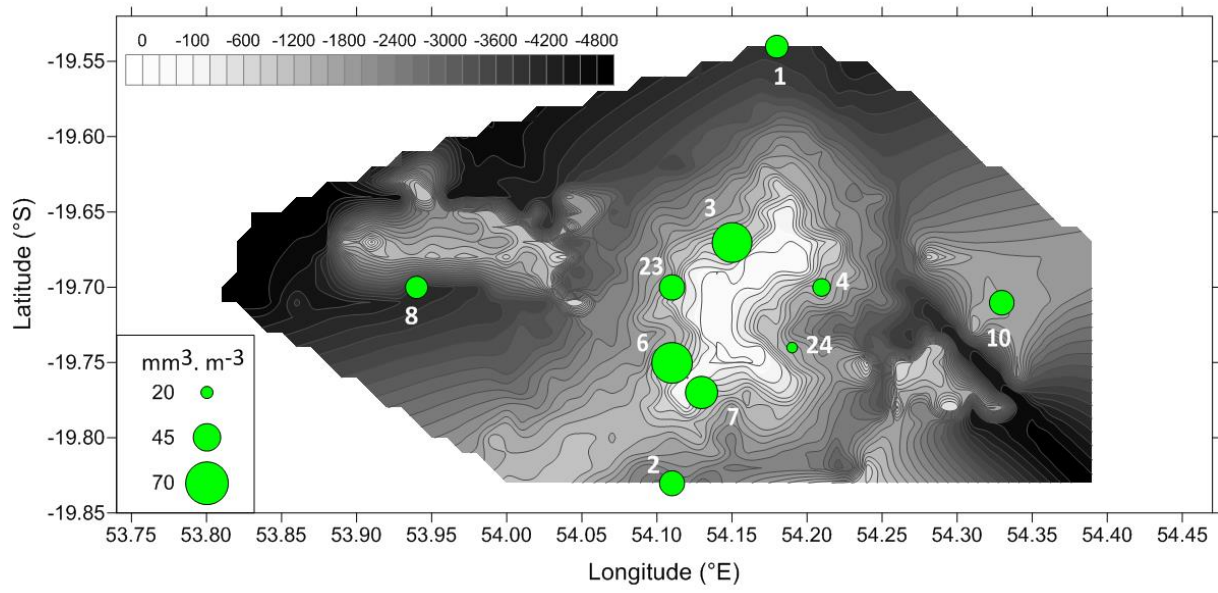


Figure 10

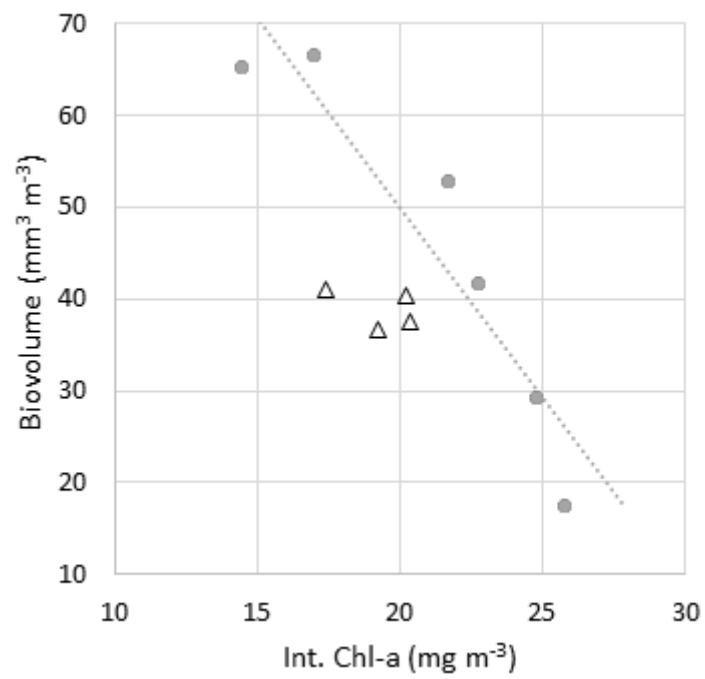


Figure 11

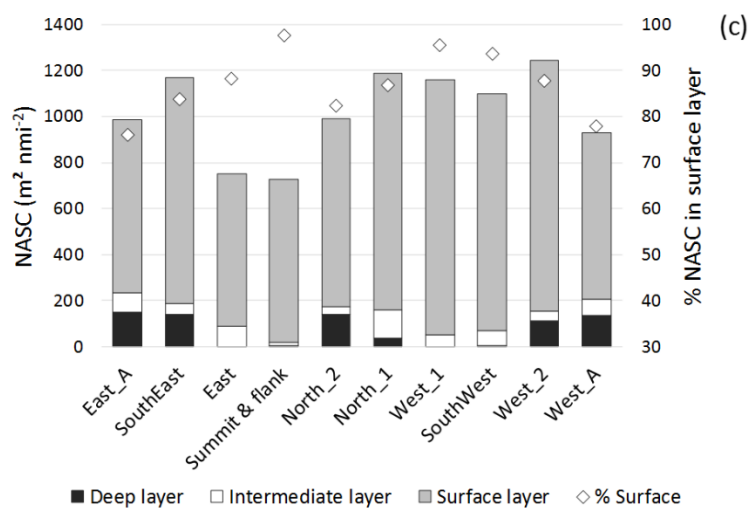
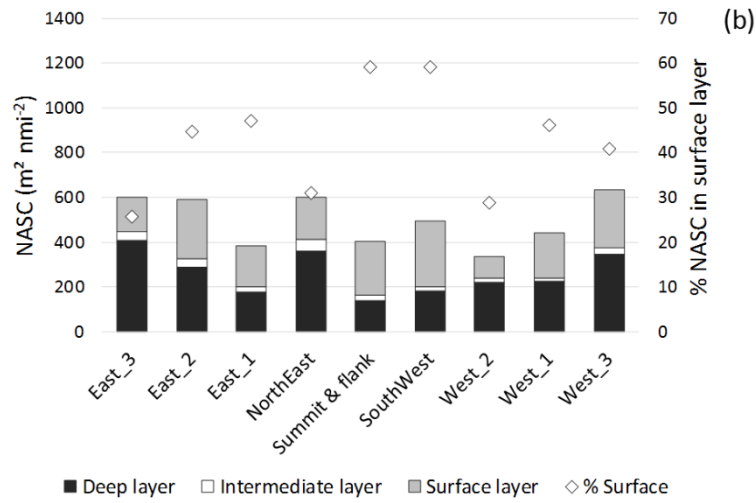
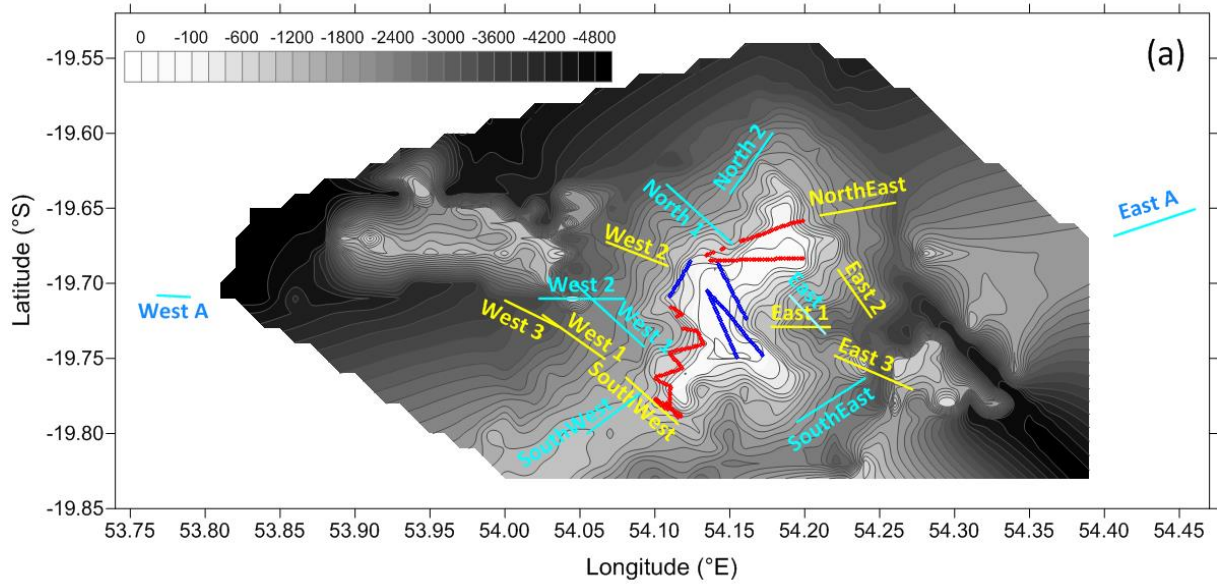


Figure 12

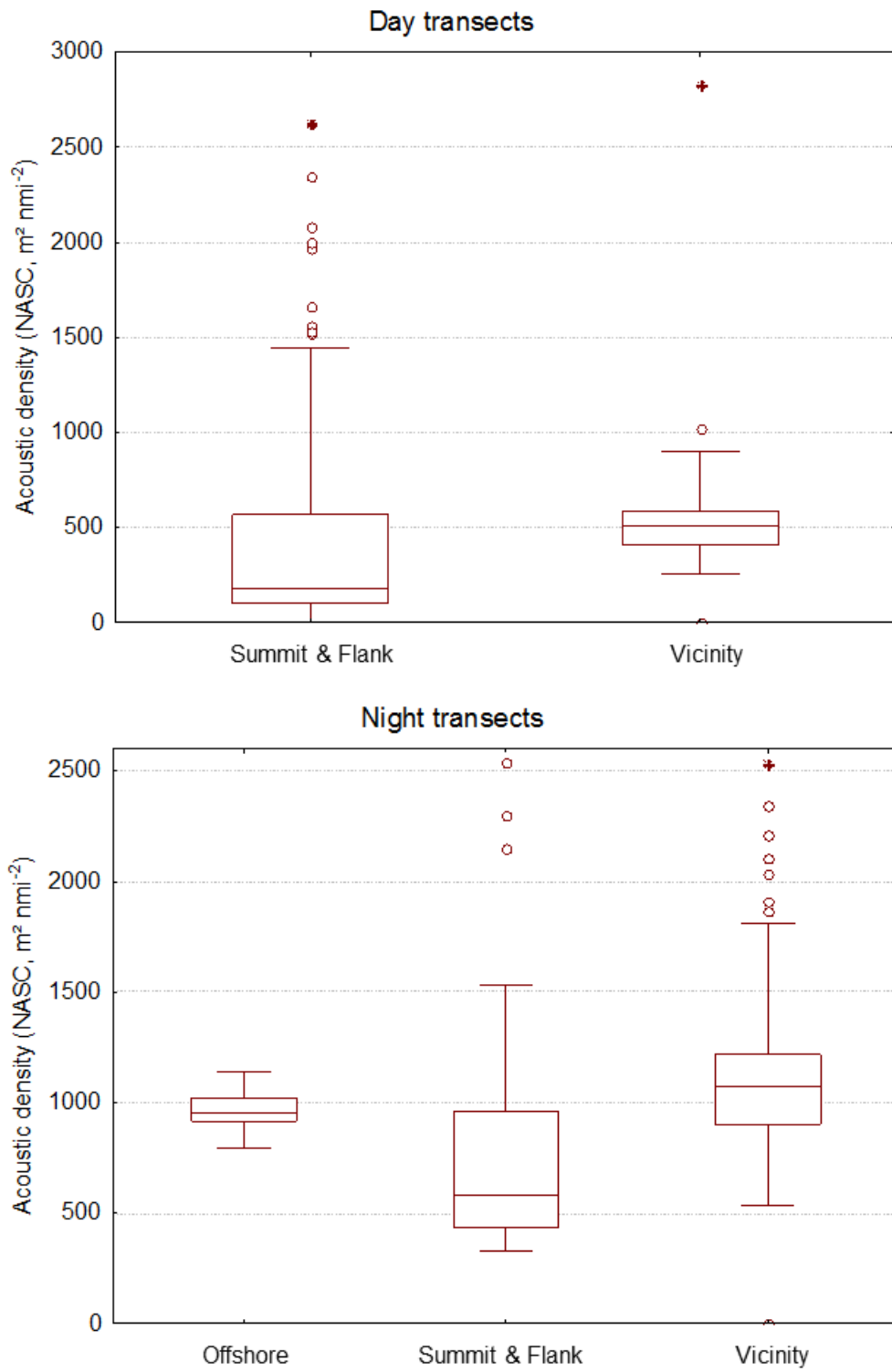


Figure 13

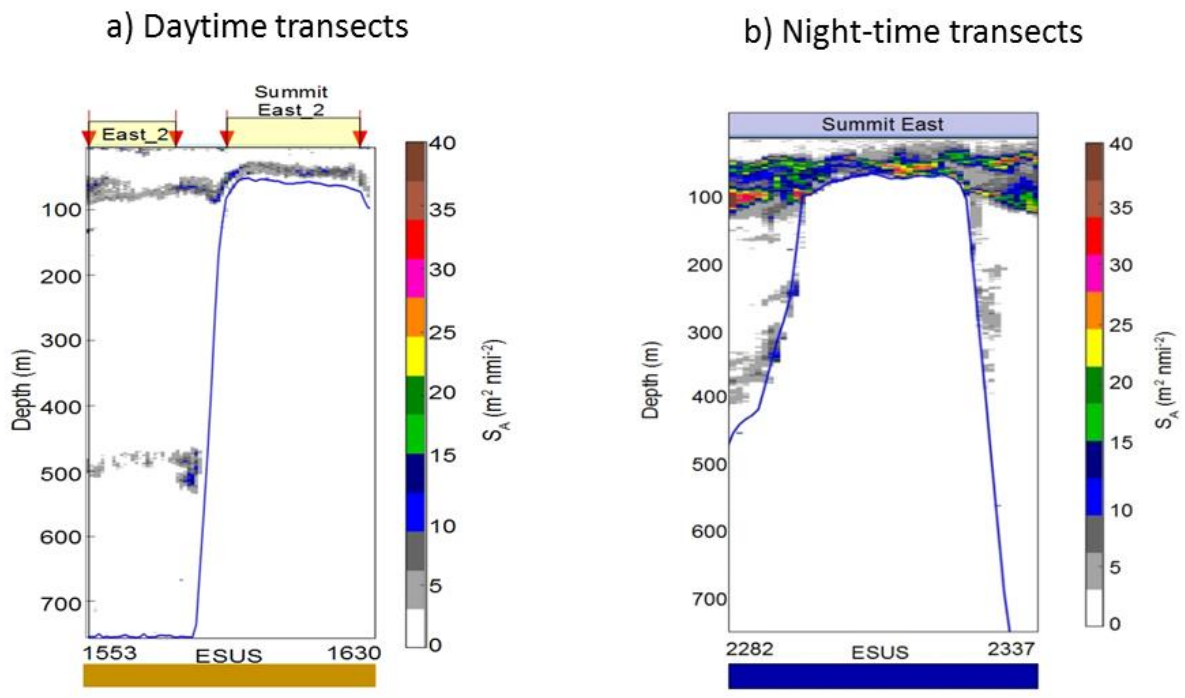


Figure 14

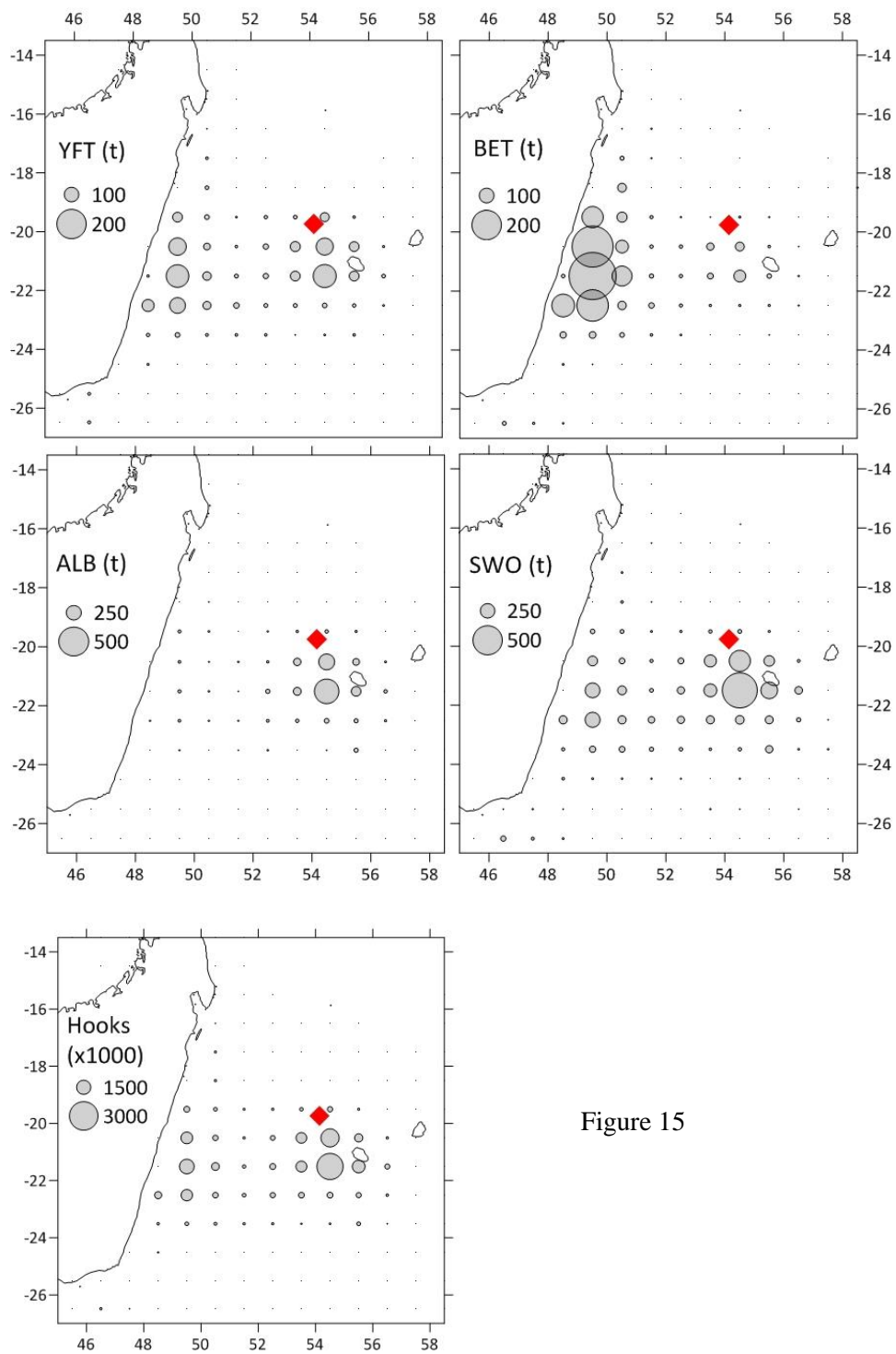


Figure 15



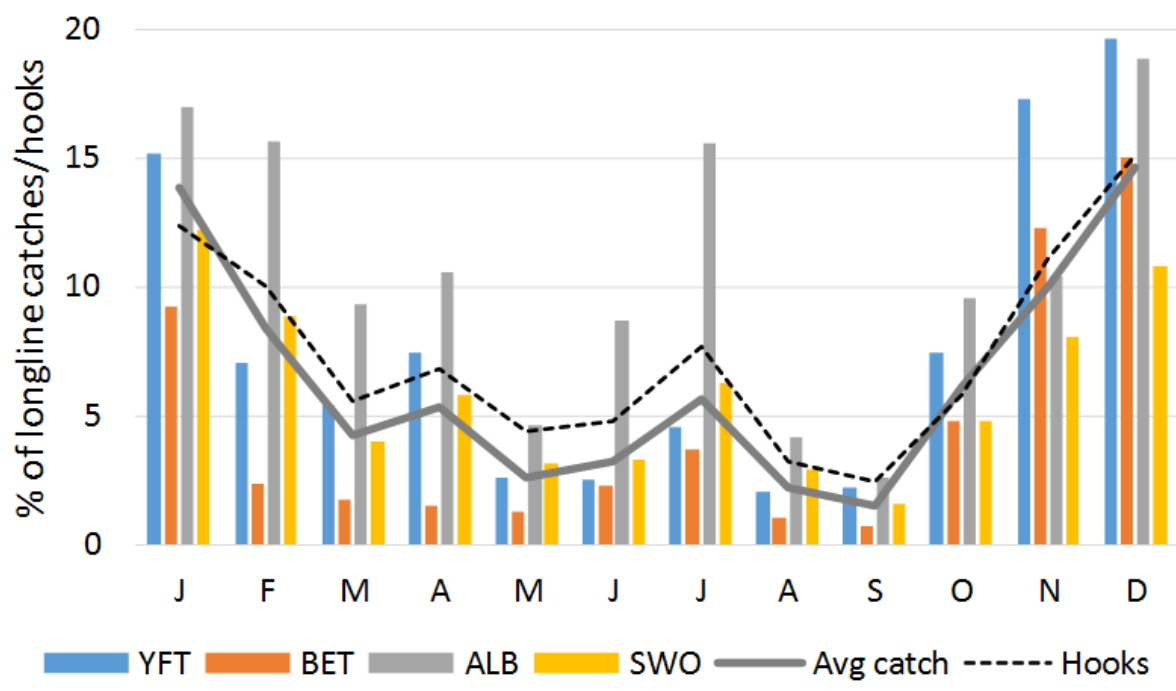


Figure 16