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International challenge to model the long-range transport of radioxenon released from medical isotope production to six Comprehensive Nuclear-Test-Ban Treaty monitoring stations



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ABSTRACT

After performing a first multi-model exercise in 2015 a comprehensive and technically more demanding atmospheric transport modelling challenge was organized in 2016. Release data were provided by the Australian Nuclear Science and Technology Organization radiopharmaceutical facility in Sydney (Australia) for a one month period. Measured samples for the same time frame were gathered from six International Monitoring System stations in the Southern Hemisphere with distances to the source ranging between 680 (Melbourne) and about 17,000 km (Tristan da Cunha). Participants were prompted to work with unit emissions in pre-defined emission intervals (daily, half-daily, 3-hourly and hourly emission segment lengths) and in order to perform a blind test actual emission values were not provided to them. Despite the quite different settings of the two atmospheric transport modelling challenges there is common evidence that for long-range atmospheric transport using temporally highly resolved emissions and highly space-resolved meteorological input fields has no significant advantage compared to using lower resolved ones. As well an uncertainty of up to 20% in the daily stack emission data turns out to be acceptable for the purpose of a study like this. Model performance at individual stations is quite diverse depending largely on successfully capturing boundary layer processes. No single modelmeteorology combination performs best for all stations. Moreover, the stations statistics do not depend on the distance between the source and the individual stations. Finally, it became more evident how future exercises

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Received 24 August 2017; Received in revised form 31 January 2018; Accepted 31 January 2018 Available online 08 March 2018 0265-931X/ © 2018 The Authors. Published by Elsevier Ltd. This is an open access article under the CC BY-NC-ND license (http://creativecommons.org/licenses/BY-NC-ND/4.0/). need to be designed. Set-up parameters like the meteorological driver or the output grid resolution should be pre-scribed in order to enhance diversity as well as comparability among model runs.

1. Introduction

The Comprehensive Nuclear-Test-Ban Treaty (CTBT), an international agreement to ban all nuclear tests, has developed a global network of 321 monitoring stations and 16 laboratories for verification purposes (CTBT, 1996), the International Monitoring System (IMS). It monitors seismic, hydroacoustic, infrasound and radionuclide signatures (CTBTO, 2017).

The radionuclide component comprises measurements of aerosolbound radioactivity at 80 locations. Half of the 80 stations shall have additional equipment to measure ambient air concentrations of four radioactive xenon isotopes (Xe-131 m, Xe-133, Xe-133 m, and Xe-135) produced in nuclear explosions. 31 noble gas stations are already in operation, and 25 of those have been certified by the Comprehensive Nuclear-Test-Ban Treaty Organization (CTBTO).

In 1999, the International Noble Gas Experiment (INGE) was launched to determine the feasibility of building and deploying automated systems to detect the four radioactive xenon (radioxenon) isotopes of interest (Auer et al., 2010; Bowyer et al., 2002). Commercial versions of three of the four radioxenon detection systems developed for the INGE are now deployed in the IMS: 1) The Automatic Radioanalyzer for Isotopic Xenon (ARIX), from the Khlopin Radium Institute, Russia (Dubasov et al., 2005), 2) the Swedish Automatic Unit for Noble Gas Acquisition (SAUNA, nowadays produced by Scienta Sauna Systems AB, Uppsala, Sweden), from Totalförsvarets Forskningsinstitut (FOI), Sweden (Ringbom et al., 2003), and 3) the Système de Prélèvement d' Air Automatique en Ligne avec l' Analyse radioXénons atmosphériques (SPALAX) from Departement Analyse, Surveillance, Environnement du Commissariat à l'Énergie Atomique (CEA/DASE), France (Fontaine et al., 2004).

Discrimination between radioxenon releases originating from a nuclear explosion or from civil facilities is a challenging task for the CTBTO. To our knowledge currently at least nine facilities worldwide are in operation: IRE located in Fleurus/Belgium, Mallinckrodt in Petten/the Netherlands, NIIAR in Dimitrovgrad/Russia, BaTek in Jakarta/Indonesia, NECSA in Pelindaba/South Africa, CENA in Ezeiza/ Argentina, HFETR in Chengdu/China, PINSTECH PAAR-1 in Islamabad/Pakistan and ANSTO in Lucas Heights/Australia (Gueibe et al., 2017; Achim et al., 2016). Atmospheric transport modelling (ATM) combined with isotopic ratio analysis (Kalinowski et al., 2010) can be considered as the most important means for achieving this goal. A large number of studies of the release and transport of radioxenon from nuclear power plants and medical isotope production and other man-made radionuclide emission facilities have been conducted to develop an understanding of background levels (Eslinger et al., 2014; Hoffman et al., 2009; Kalinowski et al., 2008; Saey et al., 2010; Wotawa et al., 2010, 2003; Zaehringer et al., 2009; Achim et al., 2016; Schoeppner, 2017). These studies confirm that fission-based production of molybdenum-99 for medical purposes is the largest routine contributor of radioxenon (which comes as a by-product of the production process) in the atmosphere, and that related releases can be detected at large distances. The Mo-99 daughter Tc-99 m is widely used for medical purposes (Peykov and Cameron (2014), approximately 30-40 million medical procedures per year) and a future growth in demand is expected.

Radioxenon levels at IMS noble gas stations resulting from underground nuclear tests can be comparable to background levels (Ringbom et al., 2014; Saey, 2009) and are thus harder to detect in regions under the influence from medical isotope production facilities. A reduction of their radioxenon releases would therefore be useful (Bowyer et al., 2013). Nevertheless, medical isotope production facilities do meet regulatory release limits (Tinker et al., 2010; Hoffmann et al., 2001), thus their operators have little incentive for spending money on reduction measures.

Although atmospheric modelling studies using inert tracers have been conducted since the early 1980s (e.g., Ferber et al., 1986; Gudiksen et al., 1984), detailed source-term data for the simulation of the transport of radioxenon from medical isotope production facilities have not been made available until recently. A 2013 study examined the regional impact of source-term data of different time resolutions on the capability to predict IMS radioxenon detections (Schoeppner et al., 2013). The study utilized emission data down to daily time resolution from the ANSTO facility close to Sydney (Australia) and detections in Australia and New Zealand. It found increasing agreement between simulations and detections from annual down to daily time resolution of the emission data. Little influences from other sources in the Southern Hemisphere were observed. A recent international model comparison (1st ATM Challenge, Eslinger et al. (2016)) used Xe-133 stack emission data from the Institut des Radioéléments (IRE) radiopharmaceutical plant in Fleurus (Belgium) and activity concentration data collected at the IMS noble gas sampler at radionuclide station DEX33 (Schauinsland, Germany). The purpose of that exercise was to ascertain the level of agreement that can be achieved between Atmospheric Transport Models (ATMs) using stack monitoring data and xenon isotopic concentration measurements at IMS stations. One of the conclusions from that exercise was that using stack monitoring data to calculate radionuclide concentrations at a distance of about 400 km can match larger individual simulated sample concentrations (i.e., above 3 mBq m⁻³) to within ± 40% of the measured concentrations if an optimally selected (according to the mean square error) ensemble mean of ATMs is used, and in some cases even lower deviations are achievable. Also, models using source term data in 15 min to 3 h time intervals produced similar agreement with measured concentrations as models using source term data averaged over longer intervals. In addition, even though the releases from IRE dominated the measured concentrations at DEX33, releases from other facilities such as nuclear power plants also influenced the smaller measured concentrations (see also De Meutter et al. (2016)). One of the benefits of that exercise was that it sparked many discussions on which techniques were most suitable, what knowledge and technique gaps exist, and what data fidelity is needed from stack monitors.

This current study also addresses the question of the level of agreement that can be achieved between IMS measurements and those simulated using Xe-133 stack release data and atmospheric transport modelling. Since ATM is a cornerstone of Treaty verification (Becker et al. (2007); Wotawa et al. (2003), including the discrimination between military and civil radionuclide sources) the scenario team of the challenge (made up by ZAMG, CTBTO/IDC and PNNL) sought broad participation of the respective community. The role of ATM in Treaty verification should be underpinned. Having at hand a multitude of simulations an ensemble approach pays off since this is the only way of overcoming individual ATMs' deficiencies and uncertainties and reproducing related measured samples more accurately. Reproducing measured radioxenon samples related to industrial production could be of great benefit to National Data Centers (NDCs, CTBT (1996)) which for verification purposes have to deal every month with a multitude of elevated radioxenon concentrations detected at IMS stations.

The setting of the current challenge is different in several ways compared to the previous one. Concentration data are used from six IMS radionuclide stations rather than just from one, as was used earlier. The distance between the release point in Australia (ANSTO) and the samplers is between 680 and 17,000 km rather than 380 km. The new challenge is set in the Southern Hemisphere which has different weather circulation patterns than the Northern Hemisphere. In addition, only the Koeberg complex in South Africa, the Atucha and Embalse complexes in Argentina and the Angra complex in Brazil have operating nuclear power plants, although releases from other medical isotope production facilities (e.g., Pelindaba in South Africa) can sometimes influence the selected IMS samplers (Eslinger et al., 2014). Overall radioxenon levels are lower in the Southern compared to the Northern Hemisphere (Achim et al., 2016). Just because of these differences between the two challenges it will be interesting to compare the two exercises in terms of qualitative, but also quantitative outcomes in order to deduce common characteristics.

2. Materials and methods

2.1. Participating organizations

Seventeen organizations from ten countries (Australia, Austria, Germany, France, Canada, Sweden, Japan, USA, United Kingdom and Belgium) took part in the 2nd ATM Challenge (Table 1). An overview of the models used is provided in Table 2 together with some key parameters of the model set-ups. Participants were encouraged to submit multiple runs, but were asked to submit at least one run using source term data with daily resolution, as radiopharmaceutical producers would – if at all – probably supply data with this resolution. There were up to six contributions per organization. Participants applied different model set-ups, different meteorological drivers in different spatial and/ or temporal resolutions and different spatial output resolutions. In total 97 simulations were performed, of which 30 were based on daily emission source resolution.

2.2. Stack emission data

The scenario team of the ATM Challenge received Xe-133 stack emission data from the Australian Nuclear Science and Technology Organization (ANSTO) radiopharmaceutical facility in Lucas Heights, Sydney, Australia (150.98° E and 34.05° S; see Fig. 2). The emission data cover a period of a month, from May, 11th, to June 10th, 2013. This period was chosen due to emission data availability, but also because some stations had outstanding recordings of radioxenon during this period. Emissions at the stack were measured with a sodium-iodide (NaI) system based on the 81.0 keV gamma emission line of Xe-133. The activities were provided for 744 contiguous 1 h release periods, each one being the sum of four 900 s measurements, and are shown together with the daily average of hourly emissions in Fig. 1.

During the period of interest for this ATM Challenge, it is observed that standard release quantities varied by as much as two orders of magnitude. The minimum activity measured by the system during this period is 1.83×10^{9} Bq (May 20th, 2013), the maximum activity is 3.67×10^{13} Bq (exceptional release on May 29th, 2013) and the median activity adds up to 5.46×10^{10} Bq. From one hour to the next, the typical variation of release quantities is by one order of magnitude, but variations by as much as two orders of magnitude are also observed from time to time.

In 2013 there were several process steps which could explain the variability of the releases. One of the main sources of peak emissions was the vacuum buffer tank. Solutions were moved around using vacuum lines and if extra vacuum capacity was required during a production run, any off-gases it may have contained were released. A second source of peaked emissions was the regeneration of the hydrogen convertor. Separately, whilst the facility was equipped with multiple gas storage tanks for trapping the off-gases from target dissolution, there were no large banks of carbon to delay and smooth out any in-cell releases. The decay tanks were released after around seven weeks delay, giving rise to small emissions. Between production runs, maintenance and waste transfer activities were also associated with

Table 1

Participants of the 2nd ATM Challenge. Organizations participating in the 1st Challenge are printed **bold**. *No blind test, involved in drafting the challenge.

Organization Abbreviation	Name(s) of participant(s)	Organization full name	Submission(s)
ARPANSA	Blake Orr	Australian Radiation Protection and Nuclear Safety Agency, Yallambie/Miranda, Australia	ARPANSA
BGR	Ole Ross	Federal Institute for Geosciences and Natural Resources (BGR), Hannover, Germany	BGR
BOKU	Petra Seibert & Anne Philipp	University of Natural Resources and Life Sciences, Institute of Meteorology & University of Vienna, Department of Meteorology and Geophysics; Vienna, Austria	BOKU ₁₋₆
CEA	Sylvia Generoso & Pascal Achim	Commissariat à l'Énergie Atomique, Arpajon, France	CEA ₁₋₂
CTBTO [*]	Jolanta Kusmierczyk-Michulec	Comprehensive Nuclear-Test-Ban Treaty Organization, International Data Center, Vienna, Austria	CTBTO ₁
СТВТО	Michael Schoeppner	Comprehensive Nuclear-Test-Ban Treaty Organization, International Data Center, Vienna, Austria	$CTBTO_2$
ECCC-CMC	Alain Malo	Environment and Climate Change Canada, Meteorological Service of Canada, Canadian Meteorological Center, Environmental Emergency Response Section, RSMC Montreal, Dorval, Québec, Canada	CMC ₁₋₂
FOI	Anders Ringbom	Swedish Defence Research Agency, Stockholm, Sweden	FOI
IRSN	Olivier Saunier, Denis Quèlo, Anne Mathieu	French Institute for Radiation protection and Nuclear Safety, Fontenay-aux-Roses, France	IRSN
JAEA	Yuichi Kijima	Japan Atomic Energy Agency, Tokai, Ibaraki, Japan	JAEA
LLNL	Lee G. Glascoe, Donald D. Lucas, Matthew D. Simpson, Phil Vogt	National Atmospheric Release Advisory Center (NARAC) at the Lawrence Livermore National Laboratory (LLNL), Livermore, California, USA	$LLNL_{1-2}$
Met. Office	Susan J. Leadbetter	Met. Office, Exeter, Devon, UK	METOFFICE
NOAA-ARL	Alice Crawford, Ariel Stein, Tianfeng Chai, Fong Ngan	National Oceanic and Atmospheric Administration Air Resources Laboratory, College Park, Maryland, USA	$NOAA-ARL_{1-4}$
PNNL	Paul W. Eslinger	Pacific Northwest National Laboratory, Richland, Washington, USA	PNNL
PU	Michael Schoeppner	Princeton University, Program on Science and Global Security, Princeton, New Jersey, USA	PU
SCK•CEN RMI	Pieter De Meutter & Andy Delcloo	Belgian Nuclear Research Center, Mol, Belgium & Royal Meteorological Institute of Belgium, Brussels, Belgium	SCKCEN $\mathrm{RMI}_{\mathrm{1-2}}$
UK-NDC	Rich Britton & Ashley Davies	United Kingdom-National Data Center (NDC), Aldermaston, Reading, UK	UK-NDC
ZAMG [*]	Christian Maurer	Zentralanstalt fuer Meteorologie und Geodynamik, Vienna, Austria	ZAMG

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concentration averaging and temporal output resolution (Output resolution), the ATM simulation direction (Simulation direction, forward in time from the source/FWD or backward in time from a receptor/BWD), emission segment resolutions considered (Emission resolution) and the number of particles released per hour (#particles/hour released, k = 1000, M = 1 million). If not indicated via superscripts otherwise particle ranges refer to the particles used for low and high emission Models and set-ups used. Columns indicate the ID of the submission, the name of the atmospheric transport model (ATM), the meteorological model providing the meteorological input (NWP model), the horizontal, vertical (with number of model levels below 2.5 km including the surface level) and temporal resolution of the meteorological input (Meteorological input resolution), the horizontal output resolution, the top height(s) of the ATM output layer(s) above surface used for resolution runs. For the BOKU submissions the difference is only in the horizontal and vertical averaging of dispersion output around the ANSTO facility. Different resolutions at upper model layers for the same meteorological input (e.g., NCEP-2

Ð	ATM	NWP	Meteorolog	Meteorological input resolution		Output resolution	uo		Simulation	Emission	#particles/hour
		Tablott	Δx	Δz (m) (# levels below 2.5 km)	Δ <i>t</i> (h)	۵x	$\nabla \Sigma \left(m ight)$	Δ <i>t</i> (h)		resolution	leteased
ARPANSA	HYSPLIT ver. 0711	NCEP-GFS	1.0°	110-6300 (9)	°	0.5°	50	3	FWD	daily	20.8 k
BOKU	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	10 km	100	1	BWD	all	41.6 to 250 k ^a
BOKU ₂	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	10 km	500	1	BWD	all	41.6 to 250 k ^a
BOKU ₃	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	10 km	1000	1	BWD	all	41.6 to 250 k ^a
BOKU4	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	70 km	500	1	BWD	all	$41.6 \text{ to } 250 \text{ k}^{a}$
BOKUs	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	130 km	1000	1	BWD	all	41.6 to 250 k ^a
BOKU	FLEXPART 9.2beta.r3	ECMWF	0.125°	10-6700 (19)	1	250 km	1000	1	BWD	all	41.6 to 250 k ^a
BGR	HYSPLIT	NCEP-GDAS	0.5°	1-610 (19)	ę	0.5°	100	03	FWD	daily	20.8 k
CEA1	FLEXPART 8.2, variable time	NCEP-GFS	0.5°	110-4700 (9)	9	recep. point	100	1	FWD	all	416.7 k to 1 M
	step					output					
CEA_2	FLEXPART 8.2, fixed time step	NCEP-GFS	0.5°	110-4700 (9)	9	recep. point	100	1	FWD	all	416.7 k to 1 M
			ŝ	40.1500.(15)	ļ	ourpur	001			o the lat the	1 0 000 77 21 17
CIMC1	MILPU	GDPS-analysis	0.22 x0.35°	(c1) 00c1-04	٥	c7.0	100	0.0833 (5 min)	FWD	dauy, nair-dauy, 3- hourly	41./ TO 333.3 K
CMC_2	MLPD	GDPS-forecast	0.22° ×0.35°	40-1500 (15)	ę	0.25°	100	0.0833 (5 min)	FWD	daily, half-daily, 3- hourly	41.7 to 333.3 k
CTBTO1	FLEXPART 9.0.2	ECMWF	0.5°	10-6700 (19)	з	0.5°	150	3	FWD	daily, half-daily, 3-	83.3 to 666.7 k
										hourly	
$CTBTO_2$	FLEXPART-WRF 3.3	WRF (NCEP)	50 to 15 km	no information	1	20 km	100	1	BWD	all	83 k
FOI	HYSPLIT ver. 20150916	NCEP-GDAS	1.0°	110-1500 (9)	ę	1.0°	100	3	FWD	daily, 3-hourly	20.8 and 166.7 k
IRSN	IdX-C3X	ARPEGE	0.5°	40-1000 (7)	e	0.5°	40	1	FWD	all	Eulerian
JAEA	HYSPLIT ver. 4	NCEP-GDAS	0.5°	1-900 (19)	ę	0.5°	$10, 30, 60 \ 100^{b}$	1	FWD	daily, half-daily	16.7 k
LLNL ₁	LODI	NCEP-GFS-ADAPT ^c	0.5°	10-2200 (15)	ŝ	20 km	20-25 (terrain dependent)	12 to 24	FWD	daily	83 k
$LLML_2$	FLEXPART 9.0.2	NCEP-GFS	0.5°	110-4700 (9)		1.0° x0.5°	400 & 500 ^b	1	FWD	daily	167 k
METOFFICE	NAME	Met Office Unified	0.35° v0 22°	10-2000 (19)	ŝ	1.0°	2000	0.25	FWD	all	42 k
	1110501	FONDER FD A Tatania	0.15.0	110 6000 (12)	c	1 0°	0001		CIVILI	11°	0E0 1.
IOAA-AKL ₁		ECIVIWF-EKA-IIILETIII	0./.0	110-0000 (12)	ი o	1.0	1000	CZ.U	TWD		y 007
NUAA-AKL ₂	ΗΥδΡΙΙΙ	NCEP-GDAS	0.1	(6) 0069-011	n	1.0	0001	62.0	FWD	all	X 062
NOAA-ARL ₃	HYSPLIT-GEM	NCEP-GDAS	1.0°	110-5900(9)	იი 1	1.0°	1000	0.25	FWD	all	250 k
NOAA-ARL ₄	HYSPLIT	NCEP-NCAR	2.5°	110-4700(4)	9	1.0°	1000	0.25	FWD	all	250 k
PNNL	HYSPLIT par. ver. 0113	NCEP-GDAS	1.0°	10-890 (12)	e	0.5°	100	1	FWD	all	416.7 to 500 k
DU	FLEXPART 8.2.3	NCEP-GDAS	0.5°	110-6000 (12)	ი	0.5°	100	e	BWD	daily, half-daily, 3- hourly	83 to 166 k ^d
SCKCENRMI1	FLEXPART 9.0.2	ECMWF	1.0°	10-6700 (19)	с	1.0°	100 (starting at 50)	1	FWD	all	100 k
SCKCENRMI ₂	FLEXPART 9.0.2	ECMWF	0.5°	10-6700 (19)	ĉ	0.5°	100 (starting at 50)	1	FWD	all	100 k
UK-NDC	FLEXPART 8.1	ECMWF	1.0°	10-6700 (19)	ĉ	1.0°	150	3	BWD	daily	50 k to 100 k ^e
ZAMG	FLEXPART 8.2.3	ECMWF	0.5°	10-6700 (19)	¢	о 5°	100	-	FWD	all	80.6 k

BKALL. GBX68 Ior 41.6 K for AUX04, 125 K for AUX09, 250 K for NZX46 and FKX2/

^b Concentrations are averaged over four, respectively two, layers.

 $^{\rm c}$ See supplementary material subsection 1.10. $^{\rm d}$ 83 k for AUX04, AUX09, NZX46; 42 k for FRX27, 166 k for BRX11 and GBX68.

100 k for FRX27, 50 k otherwise.

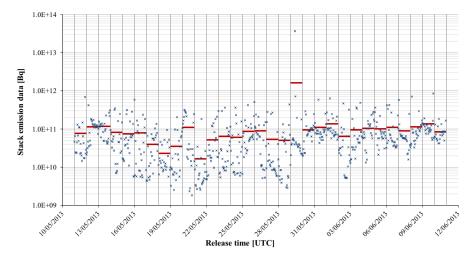


Fig. 1. Hourly (crosses) and daily average (horizontal lines) Xe-133 emissions in Bq from the stack at the ANSTO radiopharmaceutical facility in Sydney, Australia.

small (peaked and smoothed) releases. Since 2013, improvements to the plant and process have helped to reduce the variability of such emissions. ANSTO is currently commissioning its new Mo-99 nuclear medicine facility which includes an extensive abatement system designed to reduce xenon emissions even further.

A crucial aspect of stack emission data is of course its uncertainty. According to ISO Norm 2889, Appendix E, "Evaluating the errors and the uncertainty for the sampling of effluent gases" (ISO, 2010) an overall uncertainty of 10–20% should be achievable if all parameters and system components are suitably qualified and their performance is verified. However, most facilities were operating long before this standard was issued and will not necessarily have upgraded their heating, ventilation, air conditioning and emission monitoring systems to meet the standard, due to the considerable expense involved and the low dose impact of medical isotope production facilities emissions (Tinker et al., 2010; Hoffmann et al., 2001). The latter also implies that the driver from a regulatory viewpoint has historically been low. Thus, the stack emission data measured in 900 s intervals are assumed to have an average uncertainty of at least 20% (Hoffmann, 2017).

The stack monitoring data is subject to uncertainties from the specific measurement devices, their calibration, involved models, software and parameters used to calculate the final result. The systematic bias for the NaI system most importantly includes the resolution of the detector. A region of interest is pre-set at the energy of each target nuclide. The NaI detector has a lower peak resolution than a High Performance Germanium Detector (HPGe) which leads to the overestimation of the true activity present if there are overlapping peaks.

Some of the random and systematic errors can be easily quantified, such as the counting statistics error (random) or the sample extraction plane error (systematic error due to the gas flow not being fully perpendicular to the extraction plane), however the bias that remains is not fully known as some effects are temporal. These include:

- Changes in the flow rates due to declining performance of the extraction fan over time,
- drift in the gain of the NaI detector,
- calibration of the detector and of the flow measurement device,
- and changes in the background or interference present in the plant which may affect the monitoring equipment.

2.3. IMS station data

Six Southern Hemisphere noble gas stations of the CTBTO-IMS network were used in the exercise (Table 3, Fig. 2).

The SPALAX system (Fontaine et al., 2004) at FRX27 collects air

samples of about 80 m³ at ambient temperatures during cycles of 24 h. Each sample is dried, concentrated and purified to produce a final stable xenon volume of about 7.5 ml per air sample. The spectrum acquisition is automatically performed by high resolution γ -spectrometry (HPGe detector). The SAUNA system (Ringbom et al., 2003) at the other stations collects air samples of about 15 m³ at ambient temperatures during cycles of 12 h (by combining two samples of 7.5 m³, each one being collected during 6 h). From each sample, the system extracts a unique stable xenon volume of about 1.5 ml per air sample. The spectrum acquisition is performed by beta-gamma coincidence detection technique (BC404 plastic scintillator combined with a NaI detector). Both detection systems have at least minimum detectable concentrations (MDCs) of about $1mBq m^{-3}$ for Xe-133 as required for a measurement system to be part of the IMS. The reported concentrations for each sample are a decay-corrected average value valid for the sample collection period.

2.4. Meteorological data

As can be seen from Table 2 the meteorological input data to drive the atmospheric transport models was quite diverse, but similar to the ones used in the 1st ATM Challenge (Eslinger et al., 2016). Participants ran their models mainly with European Center for Medium-Range Weather Forecasts (ECMWF, Simmons et al. (1989)) and the U.S. National Oceanic and Atmospheric Administration's (NOAA) National Weather Service's National Centers for Environmental Prediction (NCEP, NCEP (2003); Saha et al. (2011)) short-term forecasts, analyses and re-analyses. Four participants employed other NWP data, i.e. the Weather Research and Forecasting (WRF) model (Done et al., 2004; Michalakes et al., 2001; Skamarock et al., 2008), the Action de Recherche Petit Echelle Grand Echelle (ARPEGE) global model (Déqué et al., 1994; Déqué and Piedelievre, 1995), the Canadian Meteorological Center (CMC) Global Deterministic Prediction System (GDPS) model (Buehner et al., 2013, 2015; Charron et al., 2012) and the Met. Office Unified Model-Global (Davies et al., 2005). Horizontal resolution ranges mostly between 0.125° and 1.0° (one submission has a resolution of only 2.5°), temporal resolution between one and 6 h. It should be noted that for some meteorological input (i.e., for ECMWF, NCEP-GDAS and NCEP-GFS) which is listed with different horizontal resolutions in Table 2 the term resolution refers to the extracted resolution rather than to a model resolution, because the underlying model with its specific resolution at which the model is actually run is the same.

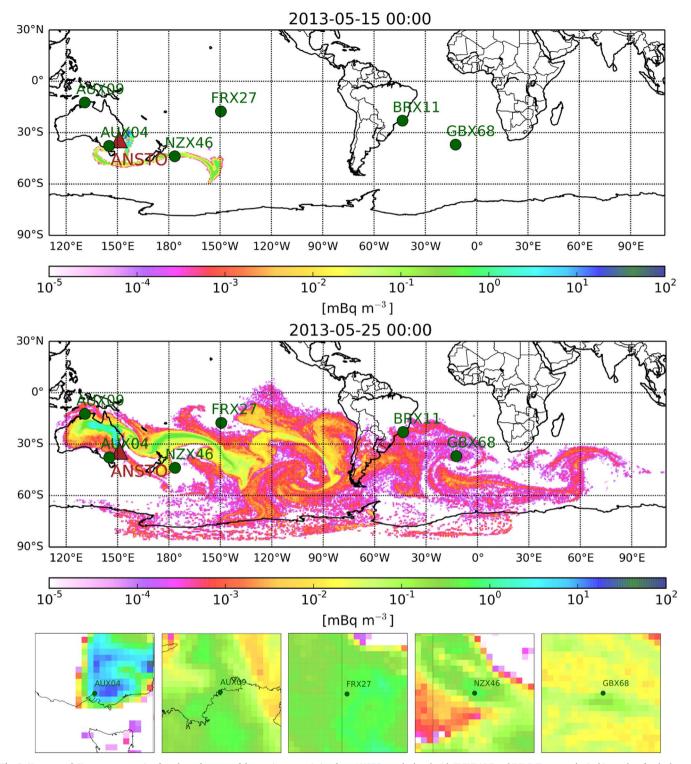


Fig. 2. Upper panel: Xenon concentration four days after start of the continuous emission from ANSTO as calculated with FLEXPART and ECMWF meteorological input data for the lowest 100 m a.g.l. Brown triangle: ANSTO facility; dark green labelled dots: IMS stations selected for the challenge. Middle panel: Same as upper panel, but valid for day 14 after the emission start. Lower panel: Plume dispersion at the IMS stations at the time of the respective maximum simulated concentrations (AUX04: May, 28th, 01 UTC; AUX09: May, 28th, 02 UTC; FRX27: June, 7th, 02 UTC; NZX46: June, 1st, 20 UTC; GBX68: June, 10th, 15 UTC).

2.5. Atmospheric transport models

The models employed for the challenge were similar to those of the 1st ATM Challenge (Eslinger et al., 2016). Five Lagrangian models as well as one Eulerian model and one mixed model (HYSPLIT-GEM) were used. The majority of simulations was accomplished with FLEXPART (16 submissions, Stohl et al. (1998, 2005)) and HYSPLIT (8

submissions, Stein et al. (2015)). Six submissions are based on the models MLPD (D'Amours et al., 2010, 2015), IdX (Tombette et al., 2014), NAME (Jones et al., 2007), HYSPLIT-GEM (Stein et al., 2015) and LODI (Ermak and Nasstrom, 2000; Larson and Nasstrom, 2002). Eight organizations employed FLEXPART, six HYSPLIT, one MLPD, one IdX-C3X, one LODI, one NAME and another one HYSPLIT-GEM. Horizontal output grid resolution ranged between receptor points (i.e.,

Table 3

IMS radionuclide stations considered with noble gas measurement system used.

Name	Country	Station code	Latitude	Longitude	Distance to source	Elevation	System
Melbourne	Australia	AUX04	37.73° S	145.10° E	680 km	40 m	SAUNA
Darwin	Australia	AUX09	12.43° S	130.89° E	3150 km	10 m	SAUNA
Rio de Janeiro	Brazil	BRX11	22.99° S	43.42° W	15,500 km ^a	8 m	SAUNA
Papeete	Tahiti/France	FRX27	17.57° S	149.57° W	6140 km	300 m	SPALAX
Chatham Island	New Zealand	NZX46	43.82° S	176.48° W	3000 km	17 m	SAUNA
Tristan da Cunha	United Kingdom	GBX68	37.07° S	12.31° W	17,000 km ^a	56 m	SAUNA

^a Not shortest distance, but rather the distance a plume would travel.

station locations) and around $2.0^\circ\!,$ temporal resolution between 5 min and 24 h.

2.6. Challenge scenario

Apart from gathering emission data from a radiopharmaceutical facility related IMS samples preferably exclusively influenced by the known source are needed. In order to estimate the influence of the main emitters in the Southern Hemisphere (i.e., BaTek in Jakarta (Indonesia), ANSTO in Lucas Heights (Australia), NECSA in Pelindaba (South Africa) and CNEA in Ezeiza (Argentina)) on IMS stations, a new feature of the CTBTO/IDC software WEBGRAPE (CTBTO, 2016) was used. WEBGRAPE (Web connected Graphics Engine) allows users to post-process and visualize source-receptor-sensitivity fields (SRS, Wotawa et al. (2003)), generated by the FLEXPART model and operationally calculated at the IDC. Thus, the six stations listed in Table 3 were selected and can be grouped as follows:

- Melbourne (AUX04), Darwin (AUX09), Chatham Island (NZX46) and Papeete/Tahiti (FRX27): In general, depending on the season, the four MIPFs considered in this work may influence the radio-xenon measurements at these stations (Kuśmierczyk-Michulec et al., 2017). For the relevant time frame, according to the WEBGRAPE analysis, the ANSTO facility was identified as the main emitter. However, contributions from unknown sources as well as model deficiencies in the FLEXPART SRS fields may be expected. There may be samples above the MDC which are not caused by ANSTO and thus even a perfect simulation would exhibit a non-perfect score.
- Tristan da Cunha (GBX68): For this station only slight, short ANSTO influences are visible within several periods which are dominated by CNEA (see Fig. 3). Impaired statistics are even more likely in this case. However, this station is very interesting since it is located 17,000 km away from the source but was nevertheless evidently influenced by the exceptional ANSTO release on May, 29th.
- Rio de Janeiro (BRX11): This station has no reliable ANSTO

influence at all and thus was not included into the statistical evaluation. It was only chosen to assure that none of the submitted runs would produce above MDC values where no measurements related to the ANSTO emissions can be found.

2.7. Blind test

The 2nd ATM Challenge was divided into a Blind Phase and an Open Phase. The paper exclusively deals with results gathered during the first of the two phases. During the Blind Phase participants had no access to the real emission data. These data could only be accessed after the submission of results via signing an agreement with CTBTO. Instead participants were asked to perform their simulations with unit emissions for four different pre-described emission time resolutions, i.e. daily, half-daily, 3-hourly and hourly. To ensure consistency between the submissions participants received templates with emission time intervals. According to the period of interest, May 11th, 2013 to June 10th, 2013 these templates contained 31, 62, 248 and 744 individual time intervals, respectively release sections. The reasoning behind that procedure was to make it practically impossible that simulations are guided by expectations following from inspecting the station measurements. Since many participants are members of National Data Centers (NDCs) access to the observations - in clear contrast to the emissions - could not be precluded. For each release section, SRS values per IMS station sampling intervals had to be calculated. The output format pre-scribed was not common to all the participants and also any (emergency) operational model set-ups could hardly be used. Therefore all runs underwent a sanity check during post-processing. For evaluating the runs, all the SRS values per release section were multiplied by the corresponding actual release value (including the outlier on May, 29th) and consequently these products were summed up over all releases per sample collection time in order to yield the final predicted radioxenon concentration values.

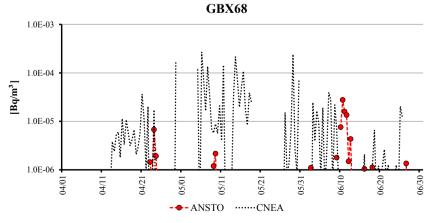


Fig. 3. Sensitivity analysis (using daily average emissions) based on FLEXPART simulations and the CTBTO/IDC software WEBGRAPE for station Tristan da Cunha (GBX68) for the time period April, 1st, to July 1st, 2013.

2.8. Statistical measures

In order to make the results of the 1st and 2nd ATM Challenge as comparable as possible, similar statistical scores (Chang and Hanna, 2004; Draxler, 2006) were evaluated and four of them combined into a rank number:

$$\operatorname{Rank} = R^{2} + \left(1 - \frac{|\operatorname{FB}|}{2}\right) + F5 + ACC$$
(1)

The squared correlation coefficient R^2 is the fraction of measurement variance explained by a linear relationship between measured and predicted values. The fractional bias FB is the bias (mean predicted minus mean measured concentration) normalized by the sum of the two means and multiplied by 2. This score, in the range -2 to 2, is less sensitive to small measured concentrations related to releases other than the ones from the facility under investigation (i.e., ANSTO).

The fraction within a factor of 5 (F5) is the fraction of predicted values that is at most a factor larger (5) or smaller (0.2) than the measured values. The latter threshold is applied since this statistic can heavily be influenced by modelled values around zero and measured values, not connected to the known source, at or just above the MDC. It was relevant to define a modified criterion for normalization, because many quotients are not defined when replacing measured samples below the MDC with zero values (see paragraph below). Thus, in the

modified definition of F5 the numerator contains the sum of modelmeasurement pairs satisfying the ratio threshold (0.2–5) and the denominator contains only the sum of pairs where at least either the simulated sample or the measured sample (or both of them) is/(are) greater than or equal to the MDC, regardless of whether a quotient is defined or not (i.e., if the measured sample falls below the MDC and is set to zero). In that way, time series with modelled values below the MDC compared to those ones above the MDC were promoted in case the measurements fell below the MDC and were set to zero.

The accuracy ACC (Swets, 1988) represents the ratio between the sum of true positives and negatives on one side and true and false positives and negatives on the other side. The criterion adopted for true positives and negatives is whether the predicted and measured values lie simultaneously at/above or below the MDC. As it is well known that predicted samples may strongly deviate in magnitude and even in phase from the observed values, the ACC measure accounts only for the fact that if there was something relevant (or likewise nothing relevant) observed, the model manages to reproduce this important, basic information.

The maximum time-lagged correlation TLRmax in comparsion to *R* indicates whether the model predictions exhibit a phase shift of amount \hat{t} in relation to the measured samples. \hat{t} represents the time shift Δt in samples yielding the maximum time-lagged correlation TLR (Δt).

The Kolmogorov-Smirnov parameter KS, unlike for the 1st ATM

Table 4

Average statistics per submission-ID over all time resolutions and stations AUX04, AUX09, FRX27, NZX46 and GBX68 ordered by rank. *: GBX68 not provided. *: GBX68 not provided. *: GBX68 not considered. *: Undefined statistical scores for GBX68. Bold numbers for individual metrics (rank excluded) mark the best value for the specific metric among all submissions and organizations.

Submission-ID	R	FB	F5 [%]	RMSE	NMSE	ACC [%]	NAAD [%]	TLRmax (\hat{t})	Rank
ARPANSA	0.73	0.02	55	0.23	14	88	125	0.73 (0)	2.66
NOAA-ARL ₃	0.60	-0.14	49	0.17	13	87	90	0.70 (12)	2.51
SCKCENRMI ₁	0.68	-0.04	49	0.21	16	84	122	0.77 (3)	2.48
SCKCENRMI ₁₋₂	0.66	-0.11	44	0.28	21	84	146	0.79 (5)	2.41
SCKCENRMI ₂	0.64	-0.18	39	0.36	26	84	170	0.80 (7)	2.33
METOFFICE	0.56	-0.15	36	0.20	15	82	129	0.70 (11)	2.27
PNNL	0.57	0.10	35	0.23	25	82	134	0.78 (6)	2.23
BOKU ₆	0.51	-0.01	30	0.29	25	85	146	0.67 (3)	2.16
CEA ₂	0.53	0.47	38	0.52	34	83	206	0.73 (4)	2.15
CEA_{1-2}	0.55	0.67	39	0.81	38	82	293	0.72 (8)	2.14
BOKU ₅	0.48	0.02	29	0.34	29	84	161	0.67 (1)	2.12
ZAMG	0.53	-0.36	33	0.24	23	83	114	0.76 (5)	2.12
NOAA-ARL ₂	0.56	0.00	28	0.25	21	83	140	0.72 (6)	2.11
CEA ₁	0.58	0.92	40	1.11	48	82	404	0.73 (9)	2.11
BOKU ₁	0.48	0.27	30	0.55	38	84	245	0.70(1)	2.10
PU	0.55	0.23	33	0.47	37	82	245	0.74 (6)	2.09
BOKU ₄	0.47	0.12	30	0.42	33	85	188	0.68 (-3)	2.08
LLNL ₂	0.42	0.23	31	0.26	18	82	164	0.70 (6)	2.08
JAEA	0.49	0.28	41	0.43	45	81	365	0.67 (6)	2.06
FOI	0.56	0.08	29	0.35	54	85	162	0.70 (6)	2.06
LLNL ₁	0.58	-0.50	23	0.18	52	84	114	0.71 (6)	2.06
NOAA-ARL ₁₋₄	0.47	-0.23	28	0.23	28	83	138	0.71 (2)	2.03
CTBTO ₁	0.50	1.00	39	1.13	41	81	389	0.70 (9)	2.03^{*}
BGR	0.56	0.09	35	0.32	115	81	261	0.73 (6)	2.03
$LLNL_{1-2}$	0.48	-0.17	27	0.22	35	83	140	0.68 (6)	2.02
BOKU ₁₋₆	0.48	0.51	29	1.19	85	82	511	0.69 (-1)	1.98
CMC_1	0.42	-0.51	25	0.25	210	81	141	0.66 (13)	1.97
CTBTO ₁₋₂	0.48	0.60	25	0.74	44	87	305	0.66 (11)	1.86**
NOAA-ARL ₁	0.48	-0.33	15	0.35	46	82	200	0.71 (3)	1.78
CMC_{1-2}	0.41	-0.48	19	0.31	205	81	173	0.66 (13)	1.78
IRSN	0.43	-0.13	17	0.40	20	77	165	0.64 (15)	1.76++
CTBTO ₂	0.63	0.54	15	0.68	51	85	403	0.74 (-1)	1.74**
BOKU ₂	0.48	1.29	26	2.65	178	79	1095	0.70 (-2)	1.73
UK-NDC	0.48	-0.31	19	0.25	48	80	198	0.76 (8)	1.69
CMC ₂	0.45	-0.38	15	0.37	199	81	216	0.70 (12)	1.68
BOKU ₃	0.48	1.37	26	2.95	201	78	1230	0.70 (-4)	1.67
NOAA-ARL ₄	0.18	-0.57	16	0.19	34	79	116	0.68 (-6)	1.67
Mean	0.51	0.33	32	0.48	44	80	253	0.72 (6)	2.06+
Median	0.50	0.26	32	0.36	25	80	221	0.72 (7)	2.07 +

Challenge, had to be discarded for the current challenge due to the special nature of the investigated samples. A rigid definition of the KS parameter (i.e., no identical pairs of values, respectively ties, are allowed, only percentiles between 5 and 95% are considered) as in Draxler (2006) does not allow the evaluation of very sparse (i.e., very few samples above MDC) data.

Further, the normalized mean square error NMSE which measures the quadratic difference between paired measured and predicted values in relation to the product of mean measurements and mean predictions as well as the root mean square error, RMSE, were applied. Finally, a further statistical measure, the normalized average absolute deviation, NAAD, was introduced for the purpose of the current challenge, since often the question arises how much simulations deviate on average from the measured samples in terms of %. The score only considers samples were either predictions or measurements are greater than or equal to the MDC and normalizes their absolute average difference with the average measurement values that are greater than or equal to the MDC. A clear disadvantage of this simple-to-understand score, in contrast to FB, is that it favors underestimating model runs over overestimating ones, since the underestimation cannot be bigger than 100%. A more detailed description of the statistical metrics can be found in the Appendix.

One important aspect when calculating the statistics is how to deal with measured samples below the MDC, especially for the current challenge, since samples below MDC are frequently encountered. Here, it was decided to set sample values below the MDC to zero, since the signals at the six selected IMS stations related to the ANSTO emissions are quite distinct in time and it is very likely that the real value, exclusively related to the ANSTO emissions, is closer to zero than to the MDC. Only for samples immediately before or after a measured value above MDC taking the MDC values to calculate the metrics would be probably more appropriate. For the rest of the below MDC samples it is more likely that they reflect unknown minor sources. In fact there exists no optimal homogeneous solution for the treatment of below MDC measurements.

Further, since the 2nd ATM Challenge dealt with long-range transport, the arrival of the plume had to be determined at every station. Otherwise statistical scores would have been artificially impaired by comparing measured samples with simulated ones connected to release times at the ANSTO facility for which no emission values were available (i.e., emissions before May 11th, 2013 00:00 UTC). The time of plume arrival was calculated as the time of the first above-zero median concentration of all involved runs at a given temporal source resolution and at a given station. Of course this procedure does not completely exclude comparing inappropriate measurements with the predicted samples. It has to be assumed that smaller emitters have an additional influence also on the above MDC measurements and that the plume arrival time calculated on the basis of all involved simulations at a station and for a certain time resolution is not perfect. One should mind in addition that the first simulated sample concentrations at plume arrival may lack contributions from stack emissions before the start date of provided releases. However, all major detections at the six IMS stations occur more than two weeks after simulation start.

3. Results and discussion

As can be seen in Fig. 2 (upper panel) within four days after the start of the continuous release from ANSTO using daily emission values two separate branches of the simulated xenon plume reach stations AUX04 and NZX46. Due to prevailing westerlies, eastward dispersion of the plume is much faster than northward dispersion so that stations AUX09 and FRX27 are reached at a similar time. By around May, 25th (Fig. 2, middle panel), the plume hits the remaining four stations, thereby being nearly fully dispersed over the Southern Hemisphere. As can be expected the Inter-Tropical Convergence Zone (ITCZ) acts as natural barrier for the plume and prevents any relevant spread across the equator. Simulated activities in the surface layer (0-100 m) reach maxima of several mBq m⁻³, especially over and around Australia. The lower panel of Fig. 2 demonstrates that around the time of maximum observed and modelled sample concentration the IMS stations (except BRX11) are well immersed in major branches (green and blue colors) of the modelled plume and so that stations are clearly hit.

3.1. Overall statistics

In Table 4, the overall statistics are presented for individual submissions (e.g., BOKU₁), for each organization as a whole (e.g., BOKU₁₋₆) and over all submissions (mean and median in the bottom lines of the table) for the purpose of ranking the submissions similarly to what was done for the 1st ATM Challenge. Scores for submissions are based on up to four runs with different source time resolutions and on up to five different stations and were averaged in order to yield at first a single number per score for every submission and organization. Grouped statistics are presented in the next subsection; detailed statistics per station and emission time resolution are given in the supplementary material. Because of missing results for GBX68 in some of the submissions and because the station is clearly influenced by another major emitter in the period of investigation (see subsection 2.6) the station was also not included in all statistics based on multiple submissions.

The correlation *R*, the fractional bias FB, the fraction within a factor of 5 F5, the root mean square error RMSE, the normalized mean square error NMSE, the accuracy ACC and the normalized average absolute deviation NAAD adopt maximum/minimum values of 0.73, 0.0, 55%, 0.17, 13, 88% and 90%. The maximum time-lagged correlations yield up to 0.80. On average zero shift of measurements against simulations (which means that correlation is equal to the maximum time-lagged correlation) can be found for one submission. Maximum time-lagged correlations adopted for time shifts bigger than one sample period imply that – at least for one involved station – even trying to correct for a reasonable phase mismatch does not improve the forecast. The rank spans from 1.67 in case of the BOKU₃ and the NOAA-ARL₄ submissions up to 2.66 for the ARPANSA submission.

The influence of the four metrics incorporated in the rank on its overall number of 2.06 calculated over all individual organizations is quite diverse. With 69% of its possible range [0,1] (calculated as the difference between the actual maximum and minimum values divided by the difference between the theoretical maximum and minimum values) the absolute number of FB divided by 2 (see equation (1)) exhibits the largest influence, which can be expected since absolute values of measurement and simulation pairs are contrasted to each other. R^2 and F5 with each 50% and 40% are located in the middle field, which is also not unreasonable, because absolute amplitudes play only a subordinate role. With 11% ACC has the weakest influence, because absolute amplitudes and the exact timing of the simulated plumes are only of secondary significance. Apart from the NMSE value statistical scores are quite similar for mean and median calculated over all organizations.

The statistics over all the organizations in terms of mean values can be summarized as follows:

- The correlation of 0.51 is rather moderate.
- According to the fractional bias of 0.33 measured values are on average only slightly overestimated by the predicted ones.
- The fraction within a factor of 5 reflects that on average a moderate fraction of 32% falls within the desired range of simulation versus measurement ratios between 0.2 and 5.0.
- The small root mean square error of 0.48 mBq m⁻³ has to be seen in the light of the larger normalized mean square error of 44.
- The accuracy, which can be somehow considered as most basic metric, yields a rather satisfying value of 80%.
- The normalized average absolute deviation adds up to 253%, which is not uncommon for in-situ model-measurement comparisons.

• The overall rank of 2.06 can be considered average, given the fact that this metric can vary between 0 and 4.

At this point it should not be concealed that due to the very sparse nature of the measurement (and also simulation) data shortcomings of the individual statistical metrics, each to a different extent, have to be expected. It is also clear that samples with measurement values set to zero (which are originally values below the MDC) cannot simply be discarded as would be appropriate if lots of data above the MDC would be available.

Nevertheless, trying to compare the overall statistics of the 2nd with those of the 1st ATM Challenge one finds a somewhat reduced correlation (0.57 versus 0.69) a (positive) fractional bias with threefold magnitude (0.82 versus 0.27), a fraction within a factor of 5 which is lower by nearly one half (35% versus 61%) and a normalized mean square error which is bigger by one order of magnitude (31 versus 3.52). Mind that the values for the 2nd ATM challenge were not extracted from Table 4, but - for consistency reasons with the statistics in Eslinger et al. (2016) - calculated for the full ensembles over all submission-IDs and time resolutions per station and finally averaged over the four (AUX04, AUX09, FRX27 and NZX46) involved stations. Unfortunately, for the reasons mentioned, the overall rank is not comparable. In this context it also has to be remembered that the two exercises are completely different in terms of station distances from the source and amplitudes to predict. It is well known from model-measurement inter-comparisons that larger measurements are easier to predict than smaller ones (Arnold et al., 2015). While the maximum measured sample amplitude amounted to around 27 mBq m⁻³ at station Schauinsland (DEX33, Germany) for the previous challenge a maximum value of around 3.6 mBq m^{-3} can be found for Melbourne (AUX04, Australia) with maximum amplitudes at the four remaining stations considered in the statistics being even one order of magnitude lower than that (see Figs. 5, 6, 8 and 9 below).

3.2. Grouped statistics

In the following it was investigated in how far model ranks grouped by common characteristics are comparable to groupings from the 1st ATM-Challenge. In order to end up with balanced values not biased towards certain comprehensive submissions, only the best run per organization and per (set-up) characteristic (e.g., spatial resolution, meteorological input data or IMS station) according to the rank was allowed in the calculation of the average. Since it is evident from Fig. 4 that the model performance using the four different emission time resolutions is not significantly different, all but 7 bars in the plot refer to model runs based on daily emission segments. The IRSN submission (employing the Eulerian model IdX) could naturally not be considered when checking the influence of the number of released particles per hour. As for the overall average model performance given in the bottom lines of Table 4 station GBX68 was neglected especially because of undefined statistical scores or submissions not considering GBX68 at all.

Fig. 4 reveals that some differences in the group averaged ranks can be found, especially according to individual stations (black and grey squares in the right half of the plot). The overall model performance for station NZX46 is reduced by a factor of more than 1/3 compared to that of AUX09. No overall difference is visible for the four different emission segment durations. Even for AUX04, which is by a factor of 4.5 closer to ANSTO than the next nearest station NZX46, only little advantage becomes evident when using hourly resolved emission values (compare the black and grey squares for this station). Besides, AUX04 and NZX46 exhibit quite comparable performance. For AUX09, FRX27, NZX46 and GBX68 with distances to ANSTO greater than or equal to 3,000 km it is comprehensible that the emission segment length has a minor impact. One also has to bear in mind that only two organizations used meteorological input with hourly resolution and that model predictions have to be averaged over at least 12h before being comparable to measured IMS samples.

Whereas for the current challenge a HYSPLIT run using NCEP-GFS input scores best, this was the case for a MLDP run using CMC-GDPS input for the previous challenge. Similarly, HYSPLIT has a slightly better average score contrasted to FLEXPART and the five other models in the 2nd ATM Challenge, whereas HYSPLIT performed worse during the 1st ATM Challenge. NCEP meteorological fields beat ECMWF and other meteorological drivers in the current challenge, but scored lower compared to ECMWF and other meteorological drivers in the previous

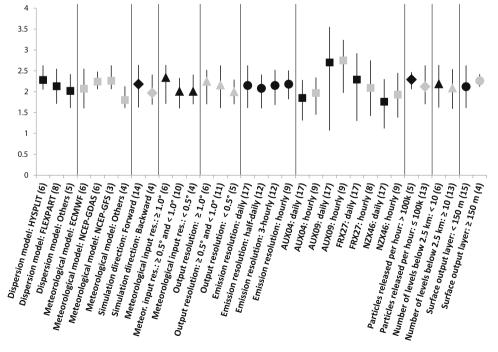
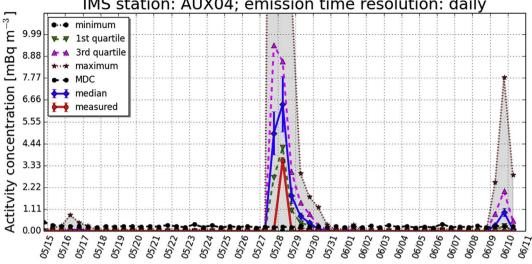


Fig. 4. Average, maximum and minimum ranks for runs grouped by common characteristics. All values pertain to runs based on daily emissions if not indicated otherwise in the bar label. Number in brackets give the number of contributing runs. Boxes denote average values, whiskers minimum and maximum. Vertical lines separate different set-up characteristics.



IMS station: AUX04; emission time resolution: daily

Fig. 5. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (30) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Melbourne (AUX04). First date-time displayed corresponds to the time of the first above-zero median concentration. For better distinguishing the lines in the above and the following legends, the reader is referred to the web version of this article.

challenge. However, the analysis of the influence of the dispersion model and of the meteorology is hampered by the fact that for the current challenge five out of the eight FLEXPART runs evaluated in Fig. 4 have ECMWF fields as input and all HYSPLIT runs evaluated in Fig. 4 have NCEP fields as input. Therefore it is impossible to separate the effect of the model and the meteorological driver. The result for grouping according to models may just reflect those for the meteorological drivers, which may perform differently for different times and regions of the world.

Albeit the 1st and the 2nd ATM Challenges do not allow to deduce generally accepted features (since they are just two case studies), it becomes evident, on the other side, that the two challenges have the following features in common:

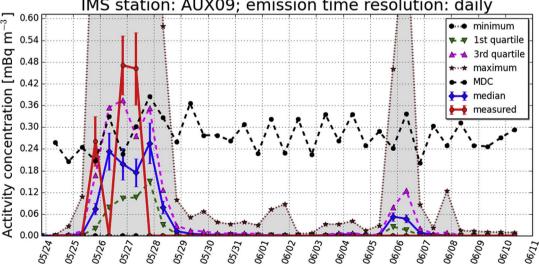
- Daily emissions do not cause a major loss in performance for the purpose of the challenges.
- Using finer (extracted) meteorological field resolutions than 1.0° does not seem to pay off. An equivalent result is found for the current challenge regarding model output grid resolution. However,

model-meteorology combinations used and their specific set-ups differ between the two challenges, which somehow limits comparability.

Due to the fact that the statements listed above hold for both challenges, we have gained at least some evidence that these results are not purely random.

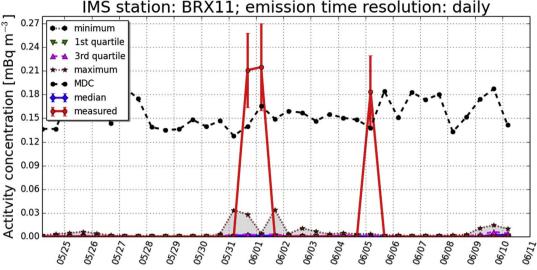
3.3. Individual station series

Ensemble plots based on daily emissions are discussed in the following paragraphs. Plots for individual runs can be found in the supplementary material. For displaying the error bars for the median of all the individual simulations a more conservative estimate of 20% in the daily emission strengths was adopted for simplicity. Only in case error bars of simulated and measured samples overlap the dispersion model and/or the underlying meteorology is likely not the cause of discrepancy. However, this is rarely the case. An alternative, probably more sophisticated way - at least in theory - would be starting with the



IMS station: AUX09; emission time resolution: daily

Fig. 6. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (30) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Darwin (AUX09). First date-time displayed corresponds to the time of the first above-zero median concentration



IMS station: BRX11; emission time resolution: daily

Fig. 7. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (30) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Rio de Janeiro (BRX11). First date-time displayed corresponds to the time of the first above-zero median concentration (ARPANSA submitted a second result for station BRX11 (ARPANSA₂), which is only considered here. CTBTO₂ run not available for this station).

errors of the provided, more highly resolved stack measurements and calculating the error for the coarser emission resolutions on the basis of Gaussian error propagation, i.e., computing the total error as square root of the sum of squared individual errors. Gaussian error propagation, however, assumes that errors are independent from each other; an assumption which may not be fulfilled. Likewise, when summing the SRS values scaled with the (daily) releases with 20% uncertainty to yield concentrations for every collection period, errors are simply added giving the maximum possible error related to the emissions for every sample.

Fig. 5 shows that for the station AUX04, closest to the source, the timing of the only one measured sample above MDC is roughly captured with by far more than 75% of simulations being advanced, some others delayed by one sampling period (i.e., 12 h). The median deviates by around 100% from the above MDC sample whereas the 1st quartile is in agreement with the observation. At the beginning, but even more at the end of the time period of investigation, we can see further peaks in a lot of the simulations not reflected in the measurements.

For the ZAMG-FLEXPART run as an example, given a daily source resolution, the most relevant release for the simulated maximum of 5.7 mBq m^{-3} resulting in a contribution of 3.4 mBq m^{-3} for the collection period starting on May, 27th, 18:53 UTC (with no above MDC measurement), occurs on May, 26th. For the contribution of the next release on May, 27th, to the next collection period starting on May, 28th, 6:53 UTC (with an actual above MDC measurement of 3.6 mBq m^{-3}), a value similar in magnitude of 2.2 mBq m⁻³, is obtained. The maximum contribution from hourly releases of 1.9 mBq m^{-3} originates on May, 27th, between 2:00 and 3:00 UTC, and adds up together with the other hourly release contributions for that day to 3.2 mBq m^{-3} , to be contrasted to the overall simulated value of 3.6 mBq m^{-3} , for the collection period starting on May, 28th, 6:53 UTC. For the contribution of the hourly releases from May, 26th, to the previous collection period starting on May, 27th, 18:53 UTC, one obtains a total value of 3.4 mBq m^{-3} . One notices that for the collection period starting on the evening of May, 27th, the daily as well as the hourly emission segments result in exactly the same improper contribution of 3.4 mBq m⁻³. No

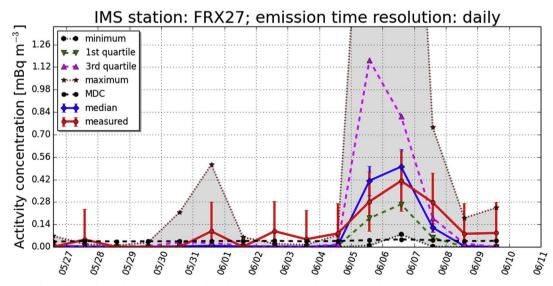


Fig. 8. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (29) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Papeete/Tahiti (FRX27). First date-time displayed corresponds to the time of the first above-zero median concentration (CTBTO₂ run not available for this station).

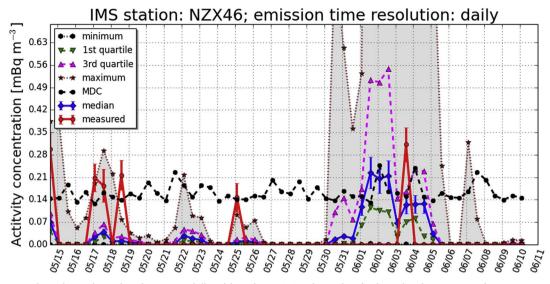


Fig. 9. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (30) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Chatham Island (NZX46). First date-time displayed corresponds to the time of the first above-zero median concentration.

concentration above the MDC of around 0.20 mBq m⁻³ was observed for that period. So there is no benefit from a high temporal resolution for the ZAMG-run, a feature which matches the overall picture.

For station AUX09 (Fig. 6) the starting time of the plume passage is captured by more than 75% of the runs, but often not the double-peak structure of the measured signal. For some runs (see also supplementary material) a double-peak structure can be found, but not in phase with the measurements. The median underestimates the measured peak amplitudes by around 50%. Here it is the 3rd quartile which comes closest to the peak measurements. As for AUX04 the simulated signal is too broad in time. A second peak in the simulations is visible also for this station, however, hardly exceeding the MDCs (3rd quartile well below the MDC).

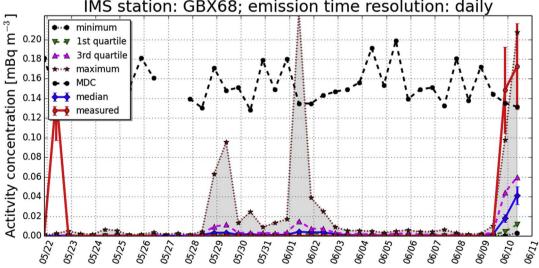
For station BRX11 (Fig. 7) all simulated values stay far below the MDC and nothing relevant is predicted by all the different model-meteorology combinations, despite a relatively low MDC of only around 0.15 mBg m^{-3} . According to sensitivity studies (see subsection 2.6) the two samples above MDC could be related to the CNEA facility in Ezeiza.

Station FRX27 (Fig. 8) is an example of what ATM can achieve even

more than 6,000 km away from the source. The plume timing of the main measured samples around June, 6th, is perfectly reproduced by all runs. The median deviates at maximum by around 50% from the measured samples.

Station NZX46 (Fig. 9) makes visible considerable deficiencies in the simulations. Whereas there are indications that some measurements just above MDC are reflected by some model runs (although with values mainly below MDC) the main peak at the beginning of the month of June 2013 is not at all correctly depicted. We can see a broad spread of the simulations around the actual event, which hints at an extended plume passage in upper layers together with inappropriate downward mixing (see subsection 3.4 and Fig. 14). It is evident from the error bars related to measurement and stack emission detection uncertainties that the meteorology and/or model errors are the main drivers of the mismatch between measured and predicted samples.

Finally, for station GBX68 (Fig. 10), we find another result in favor of the capability of ATM. Even around 17,000 km away from the source the measured signal is correctly depicted with regard to time by all the model runs. However, for all but the NOAA-ARL1 run (see brown line



IMS station: GBX68; emission time resolution: daily

Fig. 10. Minimum, 1st quartile, median, 3rd quartile and maximum of all model simulations (28) with error bars for the median due to errors in the measurement system at the stack. Measurements with error bars due to errors in the measurement system at the station and MDCs for IMS station Tristan da Cunha (GBX68). First date-time displayed corresponds to the time of the first above-zero median concentration (CTBTO₁₋₂ runs not available for this station).

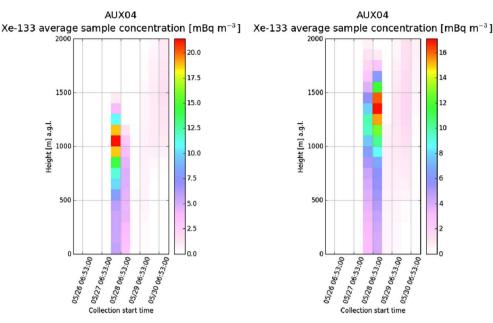


Fig. 11. Time-height cross sections of average concentrations for sample collection times for station Melbourne (AUX04) of the ZAMG-FLEXPART (left) and NOAA-ARL-HYSPLIT (right) run. Please mind the difference in the scale depicted.

for the ensemble maximum) predicted samples stay well below the MDC, with the median underestimating the true maximum by around 75%.

The last three stations cited (FRX27, NZX46 and GBX68) all owe their distinct maximum peaks to the exceptional release on May, 29th, between 9:00 and 10:00 UTC, which demonstrates the need to work with real emission data rather than disaggregated averages.

3.4. Mismatch between the measured and predicted samples

An attempt is made in the following to explain at least some incorrectly predicted samples by investigating the vertical distribution of Xe-133 concentrations for selected model outputs.

Figs. 11–15 show the time-height cross sections (on the basis of 20 model levels) of average Xe-133 concentrations over the individual station collection times for the ZAMG-FLEXPART and the NOAA-ARL₂

HYSPLIT-NCEP run (see Table 2) using daily resolved source estimates. However, in order to optimize comparability, the HYSPLIT output grid was adapted to those of the FLEXPART run, i.e., to 0.5° and 100 m vertical intervals. Thus, the values for the lowest layer are only exactly equal to those depicted in supplementary material section 1.17 for ZAMG. Whereas the sample on May, 28th, with collection start at 6:53 UTC is correctly reproduced for Melbourne (AUX04) by both runs, the modelled value of around 6 mBq m^{-3} , 4 mBq m^{-3} respectively, just one sample before (May, 27th, collection start at 18:53 UTC) is not present at all in the measurements. Looking at the time-height cross sections in Fig. 11 one finds a modelled average concentration of over 20 mBq m⁻³ at an altitude of around 1,000 m a.g.l. for the FLEXP-ART-run; thus having more than a triplication of concentrations from the ground up to 1,000 m a.g.l. For HYSPLIT amplitudes for the first of the two discussed samples at around 1,000 m a.g.l. are about half as big as for the FLEXPART run. Likely as a consequence, the wrongly

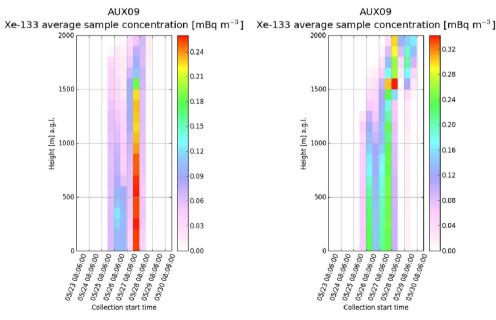


Fig. 12. Same as Fig. 11, but for station Darwin (AUX09).

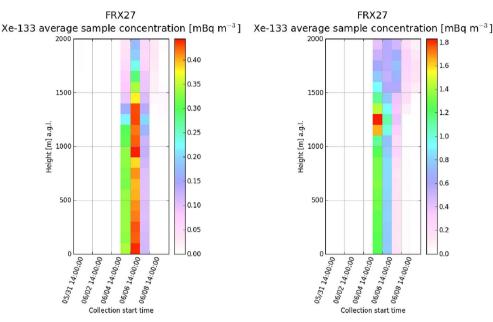


Fig. 13. Same as Fig. 11, but for station Papeete/Tahiti (FRX27).

predicted Xe concentration at the surface is smaller than for FLEXPART and is nearly the same as for the second sample. It can be assumed that in this specific case more vertical mixing occurs in the models than is the case in reality and that the FLEXPART run is more affected than the HYSPLIT run.

This hypothesis gets confirmed when checking the vertical profiles for May, 28th, 00 UTC, as measured by a radiosonde and supplied by the University of Wyoming (http://weather.uwyo.edu/upperair/ sounding.html) and the average vertical profiles of the four grid points in the ECMWF (0.5°) and NCEP (1.0°) input data nearest to Melbourne for virtual potential temperature and both horizontal wind components (Fig. 16). All three quantities are crucial parameters for determining the height and the characteristics of the planetary boundary layer (PBL). In FLEXPART the PBL height is parameterized via the Richardson number (Vogelezang and Holtslag, 1996), whereas for the HYSPLIT run shown it comes as direct (parametrized) output from the numerical weather prediction model, but could also be computed from the profile of potential temperature. Although a double inversion structure is observed between 950 and 850 hPa in Fig. 16 for the virtual potential temperature, there is no sign of such a structure (or even one inversion) in the analyses. On the other side, for May, 28th, 12 UTC, when the average sample concentration was properly reproduced by the models, the profiles of virtual potential temperature match quite good. For the third investigated date, June, 10th, 00 UTC, which co-incides with considerable above MDC model predictions in contrast to a below MDC measurement, one can detect again a stable layer in the sounding close to the surface not reproduced by the ECMWF and NCEP analyses. Looking at the wind components also quite some deviations in the v-component between analyses and soundings become apparent for May, 28th, 00 UTC, and June, 10th, 00 UTC.

A similar, but even more striking situation is found for station Chatham Island (NZX46, Fig. 14). On June, 1st, and collection start 14:15 UTC a modelled average value in the surface layer of around 0.3 mBq m⁻³ occurs for the FLEXPART run, which increases tenfold up to

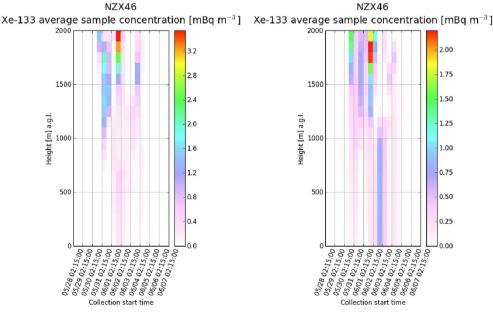


Fig. 14. Same as Fig. 11, but for station Chatham Island (NZX46).

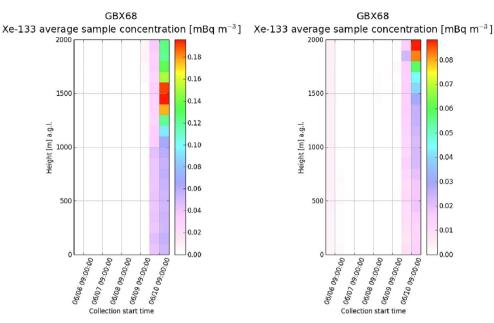


Fig. 15. Same as Fig. 11, but for station Tristan da Cunha (GBX68).

over 3 mBq m⁻³ at an altitude of around 2,000 m a.g.l. The situation is very similar for the HYSPLIT run. Four collection periods later, when an observed sample of about 0.3 mBq m⁻³ is encountered, FLEXPART model values of around 1.5 mBq m⁻³ are present at around 1,500 m a.g.l. not reaching the ground to the necessary extent. Although HYSPLIT better reproduces the one and only observed sample, for both models the concentrations are peaking some collection periods before the actual observation. In general a broad plume spanning several collection periods is visible for both simulations with modelled concentrations often adding up to more than 0.5–1 mBq m⁻³ at altitudes between 1,000 and 2,000 m a.g.l. Such a situation must be considered prone to yield wrong modelled surface concentrations. A problem

which clearly does affect also other participants' contributions.

In contrast to the two stations just mentioned, IMS stations Darwin (AUX09, Fig. 12), Papeete (FRX27, Fig. 13) and Tristan da Cunha (GBX68, Fig. 15) do not display sharp concentrations gradients within a few hundreds of meters of altitude. Apparently for two of them, FRX27 and GBX68, the FLEXPART and the HYSPLIT run feature quite good plume timing results (bearing also in mind the considerable distance between ANSTO and GBX68). Amplitudes, however, are better captured by the FLEXPART run. A proper timing can once again also be found for other submissions. For both stations time-height cross-sections (Figs. 13 and 15) exhibit well mixed modelled average concentrations up to above 1,000 m a.g.l.

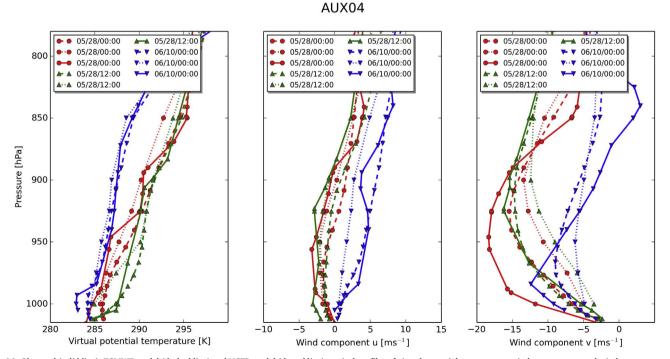


Fig. 16. Observed (solid line), ECMWF model (dashed line) and NCEP model (dotted line) vertical profiles of virtual potential temperature, wind component u and wind component v for the three sample periods with the biggest ZAMG-FLEXPART and NOAA-ARL₂-HYSPLIT model concentrations in the surface layer (0–100 m a.g.l.) at station Melbourne (AUX04). Circles: May, 28th, 00 UTC; triangles up: May, 28th, 12 UTC; triangles down: June, 10, 00 UTC.

3.5. Ensembles

Having at hand a multitude of different simulations an ensemble approach was analyzed as for the previous ATM Challenge. However, for the current analysis the rank, as defined in eq. (1), instead of the mean square error as in Eslinger et al. (2016) was employed. This was motivated by the intention to maximize overall model performance rather than performance related to a specific statistical metric. The procedure was such that for every possible ensemble size, the rank based on ensemble mean concentrations for all possible run combinations (having all again daily resolved emissions as basis) was determined and finally the combination with the maximum rank got selected. The ensemble size varies between one (which means only one model is involved) and, depending on the station, up to 24. The latter number follows from discarding four BOKU runs and two NOAA-ARL runs. For these two submissions only the best two runs per station could be considered due to computational constraints. Nevertheless, up to $2^{24} - 1 = 16,777,215$ combinations had to be evaluated.

Fig. 17 shows as an example the result together with the NAAD corresponding to the optimized rank values for station NZX46. It is a striking feature that the rank performance decreases substantially by around 1.0 if all runs are averaged. Performance decreases for all four considered stations for large ensemble sizes, at least to a certain degree. This is a well-known feature of ensemble modelling (Solazzo and Galmarini, 2014, 2015; Kioutsioukis and Galmarini, 2014). Further, for all stations the ensemble modelling exhibits the desired feature, i.e. an optimum rank at an ensemble size greater than one. However, the desired feature is not very pronounced (increase in rank for NZX46 from 2.30 for the SCKCENRMI₁ run only up to 2.63). Finally, it has to be pointed out that the optimum ensemble size of 2 for the previous challenge falls within the range of 2–6 members found here.

The NAAD for all considered stations shows an optimum which roughly coincides with the rank optimum, although the NAAD was not directly involved in the optimization process. For the two stations with the substantial decrease in rank (AUX04 and NZX46) the NAAD exhibits a steep increase, especially for an ensemble size greater than 20. Nevertheless, the NAAD values of around 450 and 225% are still far below the biggest NAAD values of individual runs based on daily emissions, namely 3879% and 1051%.

4. Conclusions

After a first ATM Challenge in 2015, a second, more comprehensive and technically demanding challenge was conducted within the CTBT context in 2016. One aim of this exercise was again to ascertain the level of agreement one can achieve between real IMS measurements and those simulated using only stack release data of Xe-133 and ATM. Another aim consisted in gaining further evidence of an optimal parameter setting (like temporal resolution of emissions or spatial resolution of meteorological input data) for predicting industry related radioxenon samples at IMS stations.

Several statistic metrics were calculated, including a rank measure, for four out of the six stations. These stations were found to be very likely influenced at least only by one main emitter, i.e. ANSTO, in a prestudy. The following conclusions based on the results of the 2nd ATM Challenge itself and on comparisons to the conclusions of the 1st ATM Challenge can be drawn:

- The performance of individual submissions at individual stations is quite diverse. There exists no single model-meteorology combination which performs best for all stations.
- Using coarser (e.g., daily) resolved emission values does not seem to result in an overall disadvantage compared to highly (e.g., hourly) resolved ones. Although this is an outcome for a special scenario, it was already stated in 2015's Challenge. For atmospheric transport and dispersion problems on a scale of 1,000 km or more this feature

got confirmed just recently by De Meutter et al. (2018).

- Likewise, using meteorological and output field (the latter was only investigated for the current challenge) resolution smaller than 1.0° does not result in better overall predictions. However, it has to be remembered that even when using the coarse resolution, smoothed input data, benefits from the underlying high-resolution NWP model are reaped.
- ATM performance varies quite some from station to station. The station statistics do not depend on the distance between the source and the individual stations. Remote stations can have better statistics than close ones (e.g., Papeete/Tahiti versus Melbourne).
- Checking the influence of meteorological input and dispersion models yields a reversed picture compared to the previous exercise. This behavior clearly demonstrates that no specific model or meteorological driver should be preferred, especially if the effect of meteorology and dispersion model cannot be disentangled.
- Assuming a more conservative uncertainty of around 20% in the daily stack emissions does not account for most of the observed deficiencies in the predictions. Therefore there is currently no need to costly reduce the uncertainty of stack emission measurements. For some samples there is a clear indication that an inappropriate depiction of boundary layer processes causes wrong surface concentrations.
- An ensemble of model runs gives a crude impression about the uncertainty of the model predictions and using the ensemble mean of all involved models is beneficial if one has no idea about the single model performances a priori. No or only little benefit unfolds from combining two or some selected models compared to just considering the best model-meteorology combination. This is likely because the multi-model ensembles biased towards certain model-meteorology combinations (i.e., FLEXPART-ECMWF and HYSPLIT-NCEP) do not fully sample the uncertainties. Further, any practical use is limited by the fact that one would need to know a robust optimum ensemble configuration a priori. But training a robust optimum ensemble configuration which outperforms the best single run clearly requires more comprehensive emission (and related above MDC measurement) data, which, however, may become accessible in the near future.

The 1st and the 2nd ATM Challenge are the first two in a row of exercises that will likely continue in the coming years. Intensive discussion is taking place regarding optimal settings. Ideally one would like to have a scenario with multiple IMS stations hit regularly by known emitters over an extended period in order to end up with significant statistics and thus to demonstrate the practical benefit of analyzing radioxenon background levels at IMS stations via ATM. The current challenge suffered from selecting IMS stations with sparse observed time series, which was the consequence of having only at hand emission data from one main emitter and trying to avoid measured

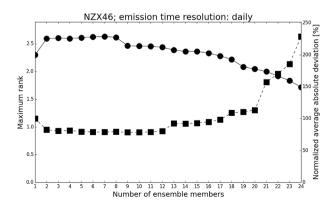


Fig. 17. Maximum rank (circles) and corresponding normalized average absolute deviation (squares) as function of ensemble size for NZX46.

Journal of Environmental Radioactivity 192 (2018) 667-686

samples related to other main emitters for which no emission data is available. The situation is clearly going to improve for upcoming challenges. The ultimate goal of all these exercises is to provide an ensemble analysis of radioxenon background levels at IMS stations frequently hit by industrial emissions. Analysts at the NDCs should in the end be able to immediately sort out suspicious events which can be clearly traced back to radiopharmaceutical isotope production facilities and/or nuclear power plants.

Different model parameters (like the meteorological driver or the meteorological model and output grid resolution) should be explored in a more coherent manner in following exercises. This means that, e.g., all participants should be encouraged to test a defined output grid resolution. Prescribing emission segments was a first step to overcome the risk of lacking comparability. As was revealed by one participant using high resolution ECMWF input data extracted at the native 0.125° resolution for the scenario of the last challenge in a thorough analysis, high resolution meteorological input can be disadvantageous compared to coarser, down-sampled input data in certain situations and may thus distort the picture when it comes to evaluating different dispersion models. Specifically, it was found that realistic topographical conditions in combination with suspected sub-grid scale processes not present in the NWP model may prevent the tracer reaching the 1,190 m high station Schauinsland east of the Rhine valley. Using the same ECMWF input data at 1.0° resolution with smoothed topography could thus yield results matching the observed samples better for a certain

event specific episode. Apart from enhancing comparability for modelinter-comparison enhancing diversity by pre-scribing parameters (e.g., different parameterizations or different meteorological drivers) in the context of ensemble modelling is another aim for a next challenge.

Disclaimer

The views expressed in this study are those of the authors and not necessarily represent the views of their affiliated institutions.

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Appendix A. Supplementary data

Supplementary data related to this article can be found at http://dx.doi.org/10.1016/j.jenvrad.2018.01.030.

Appendix

Given N predictions P_j and measurements M_i at times t_j and t_i with mean values \overline{P} and \overline{M} as well as minimum detectable concentrations MDC_j and MDC_i the statistical scores in subsection 2.8 are formally defined as:

$$FB = 2\frac{(\overline{P} - \overline{M})}{(\overline{P} + \overline{M})}$$
(A.1)

$$R = \frac{\sum (M_i - \overline{M})(P_i - \overline{P})}{\sqrt{(M_i - \overline{M})^2(P_i - \overline{P})^2}}$$
(A.2)

$$TLR(\Delta t_{ij}) = \frac{\sum (M_i - \overline{M})(P_j - \overline{P})}{\sqrt{\sum (M_i - \overline{M})^2(P_j - \overline{P})^2}}$$
(A.3)

with Δt_{ij} being the time lag between the P_j and M_i . If $\Delta t_{ij} = 0$: TLR = R. If $\Delta t_{ij} < or > 0$: Simulations are advanced/delayed with regard to measurements by amount $\hat{t} = \Delta t_{ij}$.

$$RMSE = \sqrt{\frac{1}{N} \sum (M_i - P_i)^2}$$
(A.4)

$$NMSE = \frac{RMSE^2}{\overline{MP}}$$
(A.5)

$$NAAD = \frac{\frac{1}{N}\sum |P_i|_{P_i \ge MDC_i \lor M_i \ge MDC_i} - M_i|_{M_i \ge MDC_i \lor P_i \ge MDC_i}|}{\frac{1}{N}\sum M_i|_{M_i \ge MDC_i \lor P_i \ge MDC_i}} 100$$
(A.6)

Let T denote the number of fractions satisfying:

$$0.2 \le \frac{P_i}{M_{i|M_i > 0.0}} \le 5.0 \tag{A.7}$$

and Q the number of pairs with $P_{i|P_i \ge MDC_i}$ and/or $M_{i|M_i \ge MDC_i}$ then F5 is defined as:

$$F5 = \frac{T}{Q}100$$
(A.8)

Given the number of correctly forecasted above MDC values *A* (true positives) and below MDC values *B* (true negatives) as well as the number of not correctly forecasted above MDC values (false positives) *C* and below MDC values *D* (false negatives), the Accuracy is given as:

$$ACC = \frac{A+B}{A+B+C+D}100$$

- Achim, P., Generoso, S., Morin, M., Gross, P., Le Petit, G., Moulin, C., 2016.
- Characterization of Xe-133 global atmospheric background: Implications for the International Monitoring System of the Comprehensive Nuclear-Test-Ban Treaty. J. Geophys. Res. Atmos. 121, 4951–4966.
- Arnold, D., Maurer, C., Wotawa, G., Draxler, R., Saito, K., Seibert, P., 2015. Influence of the meteorological input on the atmospheric transport modelling with FLEXPART of radionuclides from the Fukushima Daiichi nuclear accident. J. Environ. Radioact. 139, 212–225.
- Auer, M., Kumberg, T., Sartorius, H., Wernsperger, B., Schlosser, C., 2010. Ten years of development of equipment for measurement of atmospheric radioactive xenon for the verification of the CTBT. Pure Appl. Geophys. 167, 471–486.
- Becker, A., Wotawa, G., De Geer, L.E., Seibert, P., Draxler, R.R., Sloan, C., D'Amours, R., Hort, M., Glaab, H., Heinrich, P., Grillon, Y., Shershakov, V., Katayama, K., Zhang, Y., Stewart, P., Hirtl, M., Jean, M., Chen, P., 2007. Global backtracking of anthropogenic radionuclides by means of a receptor oriented ensemble dispersion modelling system in support of Nuclear-Test-Ban Treaty verification. Atmos. Environ. 41, 4520–4534.
- Bowyer, T.W., Kephart, R., Eslinger, P.W., Friese, J.I., Miley, H.S., Saey, P.R.J., 2013. Maximum reasonable radioxenon releases from medical isotope production facilities and their effect on monitoring nuclear explosions. J. Environ. Radioact. 115, 192–200.
- Bowyer, T.W., Schlosser, C., Abel, K.H., Auer, M., Hayes, J.C., Heimbigner, T.R., McIntyre, J.I., Panisko, M.E., Reeder, P.L., Satorius, H., Schulze, J., Weiss, W., 2002. Detection and analysis of xenon isotopes for the comprehensive nuclear-test-ban treaty international monitoring system. J. Environ. Radioact. 59, 139–151.
- Buehner, M., McTaggart-Cowan, R., Beaulne, A., Charette, C., Garand, L., Heilliette, S., Lapalme, E., Laroche, S., Macpherson, S.R., Morneau, J., Zadra, A., 2015. Implementation of Deterministic Weather Forecasting Systems Based on Ensemble-Variational Data Assimilation at Environment Canada. Part I: the global system. Mon. Weather Rev. 143, 2532–2559.
- Buehner, M., Morneau, J., Charette, C., 2013. Four-dimensional ensemble-variational data assimilation for global deterministic weather prediction. Nonlinear Process Geophys. 20, 669–682.
- Chang, J.C., Hanna, S.R., 2004. Air quality model performance evaluation. Meteorol. Atmos. Phys. 87, 167–196.
- Charron, M., Polavarapu, S., Buehner, M., Vaillancourt, P.A., Charette, C., Roch, M., Morneau, J., Garand, L., Aparicio, J.M., MacPherson, S., Pellerin, S., St-James, J., Heilliette, S., 2012. The Stratospheric Extension of the Canadian Global Deterministic Medium-Range Weather Forecasting System and Its Impact on Tropospheric Forecasts. Mon. Weather Rev. 140, 1924–1944.
- CTBT, 1996. Text of the Comprehensive Nuclear-test-ban Treaty. Online. Accessed 24th January 2017.
- CTBTO, 2016. WEB-GRAPE 1.8.2. Technical Report. CTBTO Preparatory Commission, International Data Center (IDC).
- CTBTO, 2017. Verification Regime. Online. Accessed 25th April 2017.
- D'Amours, R., Malo, A., Flesch, T., Wilson, J., Gauthier, J.P., Servranckx, R., 2015. The Canadian Meteorological Centre's atmospheric transport and dispersion modelling suite. Atmos.-Ocean 53, 176–199.
- D'Amours, R., Malo, A., Servranckx, R., Bensimon, D., Trudel, S., Gauthier-Bilodeau, J.P., 2010. Application of the atmospheric Lagrangian particle dispersion model MLDP0 to the 2008 eruptions of Okmok and Kasatochi volcanoes. J. Geophys. Res. Atmos. 115, 1–11.
- Davies, T., Cullen, M.J.P., Malcolm, A.J., Mawson, M.H., Staniforth, A., White, A.A., Wood, N., 2005. A new dynamical core for the Met Offices global and regional modelling of the atmosphere. Q. J. R. Meteorol. Soc. 131, 1759–1782.
- De Meutter, P., Camps, J., Delcloo, A., Deconninck, B., Termonia, P., 2016. On the capability to model the background and its uncertainty of CTBT-relevant radioxenon isotopes in Europe by using ensemble dispersion modeling. J. Environ. Radioact. 164, 280–290.
- De Meutter, P., Camps, J., Delcloo, A., Deconninck, B., Termonia, P., 2018. Time resolution requirements for civilian radioxenon emission data for the CTBT verification regime. J. Environ. Radioact. 182, 117–127.
- Déqué, M., Dreveton, C., Braun, A., Cariolle, D., 1994. The ARPEGE/IFS atmosphere model: a contribution to the French community climate modelling. Clim. Dynam. 10, 249–266.
- Déqué, M., Piedelievre, J.P., 1995. High resolution climate simulation over Europe. Clim. Dynam. 11, 321–339.
- Done, J., Davis, C.A., Weisman, M., 2004. The next generation of NWP: explicit forecasts of convection using the Weather Research and Forecasting (WRF) model. Atmos. Sci. Lett. 5, 110–117.
- Draxler, R.R., 2006. The use of global and mesoscale meteorological model data to predict the transport and dispersion of tracer plumes over Washington, D.C. Weather Forecast. 21, 383–394.
- Dubasov, Y.V., Popov, Y.S., Prelovskii, V.V., Donets, A.Y., Kazarinov, N.M., Mishurinskii, V.V., Popov, V.Y., Rykov, Y.M., Skirda, N.V., 2005. The ARIKS – 01 automatic facility for measuring concentrations of radioactive xenon isotopes in the atmosphere. Instrum. Exp. Tech. 48, 373–379.
- Ermak, D.L., Nasstrom, J.S., 2000. A Lagrangian stochastic diffusion method for

(9)

inhomogeneous turbulence. Atmos. Environ. 34, 1059-1068.

- Eslinger, P.W., Bowyer, T.W., Achim, P., Chai, T., Deconninck, B., Freeman, K., Generoso, S., Hayes, P., Heidmann, V., Hoffman, I., Kijima, Y., Krysta, M., Malo, A., Maurer, C., Ngan, F., Robins, P., Ross, J.O., Saunier, O., Schlosser, C., Schoeppner, M., Schrom, B.T., Seibert, P., Stein, A.F., Ungar, K., Yi, J., 2016. International challenge to predict the impact of radioxenon releases from medical isotope production on a comprehensive nuclear test ban treaty sampling station. J. Environ. Radioact. 157, 41–51.
- Eslinger, P.W., Friese, J.I., Lowrey, J.D., McIntyre, J.I., Miley, H.S., Schrom, B.T., 2014. Estimates of radioxenon released from Southern Hemisphere medical isotope production facilities using measured air concentrations and atmospheric transport modeling. J. Environ. Radioact. 135, 94–99.
- Ferber, G.J., Heffter, J.L., Draxler, R.R., Lagomarsino, R.J., Thomas, F.L., Deitz, R.N., Benkovitz, C.M., 1986. Cross-appalachian Tracer Experiment (CAPTEX-83) Final Report. Report NOAATech. Memo. ERL ARL-142. NOAA/Air Resources Laboratory Online, Accessed date: 12 June 2017.
- Fontaine, J.P., Pointurier, F., Blanchard, X., Taffary, T., 2004. Atmospheric xenon radioactive isotope monitoring. J. Environ. Radioact. 72, 129–135.
- Gudiksen, P.H., Ferber, G.J., Fowler, M.M., Eberhard, W.L., Fosberg, M.A., Knuth, W.R., 1984. Field studies of transport and dispersion of atmospheric tracers in nocturnal drainage flows. Atmos. Environ. 18, 713–731.
- Gueibe, C., Kalinowski, M.B., Baré, J., Gheddou, A., Krysta, M., Kusmierczyk-Michulec, J., 2017. Setting the baseline for estimated background observations at IMS systems of four radioxenon isotopes in 2014. J. Environ. Radioact. 178–179, 297–314.
- Hoffman, I., Ungar, K., Bean, M., Yi, J., Servranckx, R., Zaganescu, C., Ek, N., Blanchard, X., Le Petit, G., Brachet, G., Achim, P., Taffary, T., 2009. Changes in Radioxenon Observations in Canada and Europe during Medical Isotope Production Facility Shut Down in 2008. J. Radioanal. Nucl. Chem. 282, 767–772.
- Hoffmann, E.L., Loosz, T., Mokhber-Shahin, L., 2001. Environmental and Effluent Monitioring at ANSTO Sites, 2000. Technical Report. Online. , Accessed date: 22 January 2018.
- Hoffmann, E., 2017. Personal Communication Regarding the Uncertainty of Stack Measurements.
- ISO, 2010. ISO 2889:2010 Sampling airborne radioactive materials from the stacks and ducts of nuclear facilities. Technical Report. Online. , Accessed date: 12 June 2017. Jones, A., Thomson, D., Hort, M., Devenish, B., 2007. The U.K. Met Office's Next-
- Generation Atmospheric Dispersion Model. NAME III. Springer, pp. 580–589.
- Kalinowski, M.B., Axelsson, A., Bean, M., Blanchard, X., Bowyer, T.W., Brachet, G., Hebel, S., McIntyre, J.I., Peters, J., Pistner, C., Raith, M., Ringbom, A., Saey, P.R.J., Schlosser, C., Stocki, T.J., Taffary, T., Ungar, R.K., 2010. Discrimination of Nuclear Explosions against Civilian Sources Based on Atmospheric Xenon Isotopic Activity Ratios. Pure Appl. Geophys. 167, 517–539.
- Kalinowski, M.B., Becker, A., Saey, P.R.J., Tuma, M.P., Wotawa, G., 2008. The complexity of CTBT verification. Taking noble gas monitoring as an example. J. Complex 14, 89–99.
- Kioutsioukis, I., Galmarini, S., 2014. De praeceptis ferendis: good practices in multimodel ensembles. Atmos. Chem. Phys. 14, 11791–11815.
- Kuśmierczyk-Michulec, J., Krysta, M., Kalinowski, M., Hoffmann, E., Baré, J., 2017. Longrange transport of Xe-133 emissions under convective and non-convective conditions. J. Environ. Radioact. 175–176, 135–148.
- Larson, D.J., Nasstrom, J.S., 2002. Shared- and distributed-memory parallelization of a Lagrangian atmospheric dispersion model. Atmos. Environ. 36, 1559–1564.
- Michalakes, J., Chen, S., Dudhia, J., Hart, L., Klemp, J., Middlecoff, J., Skamarock, W., 2001. Development of a next generation regional Weather Research and Forecast Model. World Scientific Publishing, Singapore, pp. 269–276.
- NCEP, 2003. Environmental Modeling Center: The GFS Atmospheric Model. NOAA/ NCEP, Environmental Modeling Center Office Note 442. Technical Report. Online. Accessed date: 12 June 2017.
- Peykov, P., Cameron, R., 2014. Medical Isotope Supply in the Future: Production Capacity and Demand Forecast for the 99Mo/99mTc Market, 2015-2020. Report NEA/SEN/ HLGMR (2014) 2. Online. Organisation for Economic Co-Operation and Development, Nuclear Energy Agency, Accessed date: 12 June 2017.
- Ringbom, A., Axelsson, A., Aldener, M., Auer, M., Bowyer, T.W., Fritioff, T., Hoffman, I., Khrustalev, K., Nikkinen, M., Popov, V., Popov, Y., Ungar, K., Wotawa, G., 2014. Radioxenon detections in the CTBT international monitoring system likely related to the announced nuclear test in North Korea on February 12, 2013. J. Environ. Radioact. 128, 47–63.
- Ringbom, A., Larson, T., Axelsson, A., Elmgren, K., Johansson, C., 2003. SAUNA a system for automatic sampling, processing, and analysis of radioactive xenon. Nucl. Instrum. Meth. A 508, 542–553.
- Saey, P.R., 2009. The influence of radiopharmaceutical isotope production on the global radioxenon background. J. Environ. Radioact. 100, 396–406.
- Saey, P.R., Schlosser, C., Achim, P., Auer, M., Axelsson, A., Becker, A., Blanchard, X., Brachet, G., Cella, L., De Geer, L.E., Kalinowski, M., Le Petit, G., Peterson, J., Popov, V., Popov, Y., Ringbom, A., Sartorius, H., Taffary, T., Zaehringer, M., 2010. Environmental Radioxenon Levels in Europe: a Comprehensive Overview. Pure Appl. Geophys. 167, 499–515.
- Saha, S., Moorthi, S., Wu, X., Wang, J., Nadiga, S., Tripp, P., Behringer, D., Hou, Y., Chuang, H., Iredell, M., Ek, M., Meng, J., Yang, R., Mendez, M.P., van den Dool, H., Zhang, Q., Wang, W., Chen, M., Becker, E., NCEP Climate Forecast System Version 2 (CFSv2) 6-hourly Products, 2011. Research Data Archive at the National Center for

Atmospheric Research. Computational and Information Systems Laboratory, Dataset. Technical Report. Online , Accessed date: 26 July 2017.

Schoeppner, M., 2017. Performance Assessment of the CTBTO Noble Gas Network to Detect Nuclear Explosions. Pure Appl. Geophys. 174, 2161–2171.

- Schoeppner, M., Plastino, W., Hermanspahn, N., Hoffmann, E., Kalinowski, M., Orr, B., Tinker, R., 2013. Atmospheric transport modelling of time resolved 133Xe emissions from the isotope production facility ANSTO. Australia. J. Environ. Radioact. 126, 1–7.
- Simmons, A.J., Burridge, D.M., Jarraud, M., Girard, C., Wergen, W., 1989. The ECMWF medium-range prediction models development of the numerical formulations and the impact of increased resolution. Meteorol. Atmos. Phys. 40, 28–60.
- Skamarock, W., Klemp, J., Dudhia, J., Gill, D.O., Barker, D., Duda, M.G., Huang, X., Wang, W., 2008. A Description of the Advanced Research WRF Version 3. Report NCAR/TN-475 + STR. National Center for Atmospheric Research.
- Solazzo, E., Galmarini, S., 2014. A science-based use of ensembles of opportunities for assessment and scenario study: a re-analysis of HTAP-1 ensemble. Atmos. Chem. Phys. Discuss. 14, 30523–30546.
- Solazzo, E., Galmarini, S., 2015. The Fukushima Cs-137 deposition case study: properties of the multi-model ensemble. J. Environ. Radioact. 139, 226–233.
- Stein, A.F., Draxler, R.R., Rolph, G.D., Stunder, B.J.B., Cohen, M.D., Ngan, F., 2015. NOAA's HYSPLIT atmospheric transport and dispersion modeling system. Bull. Am. Meteorol. Soc. 96, 2059–2077.
- Stohl, A., Forster, C., Frank, A., Seibert, P., Wotawa, G., 2005. Technical note: the Lagrangian particle dispersion model FLEXPART version 6.2. Atmos. Chem. Phys. 5, 2461–2474.
- Stohl, A., Hittenberger, M., Wotawa, G., 1998. Validation of the Lagrangian particle

dispersion model FLEXPART against large-scale tracer experiment data. Atmos. Environ. 32, 4245–4264.

- Swets, J.A., 1988. Measuring the Accuracy of Diagnostic Systems. Science 240, 1285–1293.
- Tinker, R., Orr, B., Grzechnik, M., Hoffmann, E., Saey, P., Solomon, S., 2010. Evaluation of radioxenon releases in Australia using atmospheric dispersion modelling tools. J. Environ. Radioact. 101, 353–361.
- Tombette, M., Quentric, E., Quèlo, D., Benoit, J.P., Mathieu, A., Korsakissok, I., Didier, D., 2014. C3X: a software platform for assessing the consequences of an accidental release of radioactivity into the atmosphere. In: Poster Presented at Fourth European IRPA Congress, June 2014, Geneva, pp. 23–27.
- Vogelezang, D.H.P., Holtslag, A.A.M., 1996. Evaluation and model impacts of alternative boundary-layer height formulations. Bound. Layer Meteor 81, 245–269.
- Wotawa, G., Becker, A., Kalinowski, M., Saey, P., Tuma, M., Zaehringer, M., 2010. Computation and analysis of the global distribution of the radioxenon isotope 133Xe based on emissions from nuclear power plants and radioisotope production facilities and its relevance for the verification of the Nuclear-Test-Ban Treaty. Pure Appl. Geophys. 167, 541–557.
- Wotawa, G., De Geer, L.E., Denier, P., Kalinowski, M., Toivonen, H., D'Amours, R., Desiato, F., Issartel, J.P., Langer, M., Seibert, P., Frank, A., Sloani, C., Yamazawa, H., 2003. Atmospheric transport modelling in support of CTBT verification: Overview and basic concepts. Atmos. Environ. 37, 2529–2537.
- Zaehringer, M., Becker, A., Nikkinen, M., Saey, P., Wotawa, G., 2009. CTBT radioxenon monitoring for verification: today's challenges. J. Radioanal. Nucl. Chem. 282, 737–742.