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# 1 Assessment of a high-level spent nuclear fuel disposal model

2  
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11  
12 **Abstract:** This paper assesses a model for the full-scale engineered barrier experiment  
13 (Febex) in situ test, simulating the disposal of heat-emitting, high-level spent nuclear fuel.  
14 The model has been developed using Code\_Bright and is based on previous modelling  
15 efforts. The model focuses on the thermo-hydro-mechanical behaviour of a bentonite  
16 barrier surrounding a heater that simulates the heat production of radioactive spent fuel.  
17 The new model incorporates the new dismantling operations and has been compared with  
18 new in situ measurements. Once calibrated, the long-term response of the model has been  
19 analysed. Furthermore, the implementation of additional processes and their impact on  
20 hydration and stress development have been assessed. In this assessment, it has been  
21 found that the bentonite barrier does not fully saturate near the canister in any case, which  
22 is an important goal for the isolation of the spent nuclear fuel. The intrinsic permeability  
23 has been found to give results closer to measurements when its value is doubled. The  
24 double-structured model has given interesting results, but the authors found that it does  
25 not have a great influence in the general behaviour of the model.

26 **Keywords:** FEBEX, CODE\_BRIGHT, numerical modeling, nuclear waste disposal.

## 27 28 1. Introduction

29 Deep geological storage or disposal remains the preferred option for waste management  
30 of heat-emitting, high-level radioactive nuclear waste (HLW) in several countries.

31 Excavation of deep underground repositories is limited to rock units that are reasonably  
32 stable and without major groundwater flow, and to depths of a few hundred metres  
33 below the surface.

34 The required degree of waste isolation needed for HLW is provided by a combination of  
35 engineered and natural barriers placed between the potentially harmful radionuclides  
36 and the biosphere. The natural barrier is fundamentally the host rock, and the artificial  
37 barriers are the solid matrix of the waste itself, the metallic canister enclosing the waste,  
38 and the backfill –the sealing material placed around the canisters. This backfill, also  
39 called an ‘engineered barrier’, is often constructed using compacted expansive clay.  
40 Bentonite has generally been chosen because of its high swelling capacity, low  
41 permeability and favorable retardation properties (Gens et al., 2009).

42 The bentonite barrier fulfills several important functions. In the first instance, a very  
43 low hydraulic conductivity restricts water penetration and retards significantly solute  
44 transport due to its low diffusion coefficient and to additional sorption effects. It should  
45 also provide a favorable chemical environment and be able to self-heal if subjected to  
46 physical perturbation such as cracking and fissuring events (Gens, 2003). The  
47 engineered clay barrier and adjacent host rock (usually called the ‘near field’) will be  
48 subjected to the heating effect of the nuclear waste, and also to various associated  
49 hydraulic and mechanical phenomena that interact in a complex way.

50 In addition, compacted bentonite is initially unsaturated, and will therefore be subjected  
51 to hydration from the surrounding rock, triggering further coupled thermo-hydro-  
52 mechanical (THM) phenomena. In order to achieve a safe and robust repository design,  
53 it is necessary to have a good understanding of the processes that occur in the near field,  
54 and of their evolution over time (Gens et al., 2009).

55 Given the importance of the problem, it has proved useful to perform large-scale and  
56 medium-scale heating tests simulating repository conditions in underground laboratories  
57 around the world (Dixon et al., 2002; Pusch et al., 1985; Selvadurai, 1997; Volckaert et  
58 al., 1996). Because of the low permeability of the materials involved, such experiments  
59 usually require long testing times, measured in years, in order to obtain meaningful  
60 results. The complexity of the phenomena involved, and of their interactions, implies  
61 that process understanding and interpretation of results are limited unless supported by  
62 suitable numerical models able to reproduce the main features of the test. Owing to the  
63 coupled nature of the THM interactions, coupled THM formulations are inevitably  
64 required.

65 The present paper contains the description of an updated model for the large scale in  
66 situ heating test FEBEX. In this test, after five years of heating, one of the heaters was  
67 switched off and the experiment was partially dismantled, allowing the final state of the  
68 barrier to be observed directly. In this way, very valuable information on the state of the  
69 bentonite at the end of the test was obtained. The test has received attention during the  
70 initial (Gens et al., 1998) and intermediate stages (Alonso & Alcoverro, 2005). Finally,  
71 Gens et al. (2009) discussed the thermal, hydraulic and mechanical observations in the  
72 bentonite barrier and in the host rock, paying special attention to the progress of  
73 hydration in the barrier, the effects of heating and vapor transport, and the development  
74 of swelling pressures in the barrier.

75 In the FEBEX experiment, heaters are emplaced in the axis of a tunnel excavated in  
76 granite to simulate the heat production of radioactive waste. The test is fully  
77 instrumented, and attention is focused on the thermo-hydro-mechanical (THM)  
78 behavior of the near-field region constituted by the compacted bentonite barrier  
79 surrounding the heater and the immediately adjacent rock. Interpretation of the test is  
80 assisted by the performance of a coupled numerical analysis based on a formulation that  
81 incorporates the relevant THM phenomena. Further description of the FEBEX in situ  
82 test can be found in Gens et al. (2009).

83 The model of this test has been developed using CODE\_BRIGHT (Olivella et al.,  
84 1996), a Finite Element Method program that enables coupled thermo-hydro-  
85 mechanical analysis in geological media. This is a code version that incorporates  
86 customized pre- and post- process interfaces so that simulation models that use  
87 CODE\_BRIGHT –available from: [https://deca.upc.edu/en/projects/code\\_bright](https://deca.upc.edu/en/projects/code_bright)– can be  
88 developed more easily. It allows easy modification of parameters, boundary conditions,  
89 excavation protocols, meshing and organization of calculation intervals. The reason is

90 that it has been developed using the GiD interface (<http://www.gidhome.com>) that  
91 permits pre- and post- process of data in a user-friendly way.

92 The updated model for FEBEX is based on the geometry and material properties of the  
93 previous modeling efforts (Gens et al., 2009). The main objective of this work is to  
94 develop a new model of the well-known FEBEX experiment that can be further  
95 exploited in terms of two main goals: (1) comparison with new measurements and (2)  
96 implementation of additional processes such as coupled flows and double structure  
97 models to investigate their impact of hydration and stress development.

98 This updated model is intended to be complementary with the work done by Sanchez et  
99 al. (2012) who also made long term predictions about the FEBEX model but didn't  
100 contemplate the introduction and assessment of the additional processes that are  
101 included in this article.

102

## 103 **2. The FEBEX updated model**

104 Figure 1 shows the model geometry and materials considered in this model. The host  
105 rock, the bentonite buffer, and the concrete plug are the geomaterials considered. In  
106 addition, a steel liner confines the bentonite and plug. The model assumes axisymmetry  
107 along the tunnel axis. The geometry and material distribution are prepared to simulate  
108 the partial dismantling of the test which was carried out during the performance of the  
109 test in such a way that one of the heaters has continued operating. Boundary conditions  
110 are modified accordingly for the excavation of the plug and part of the bentonite is  
111 removed from the model thus simulating excavation.

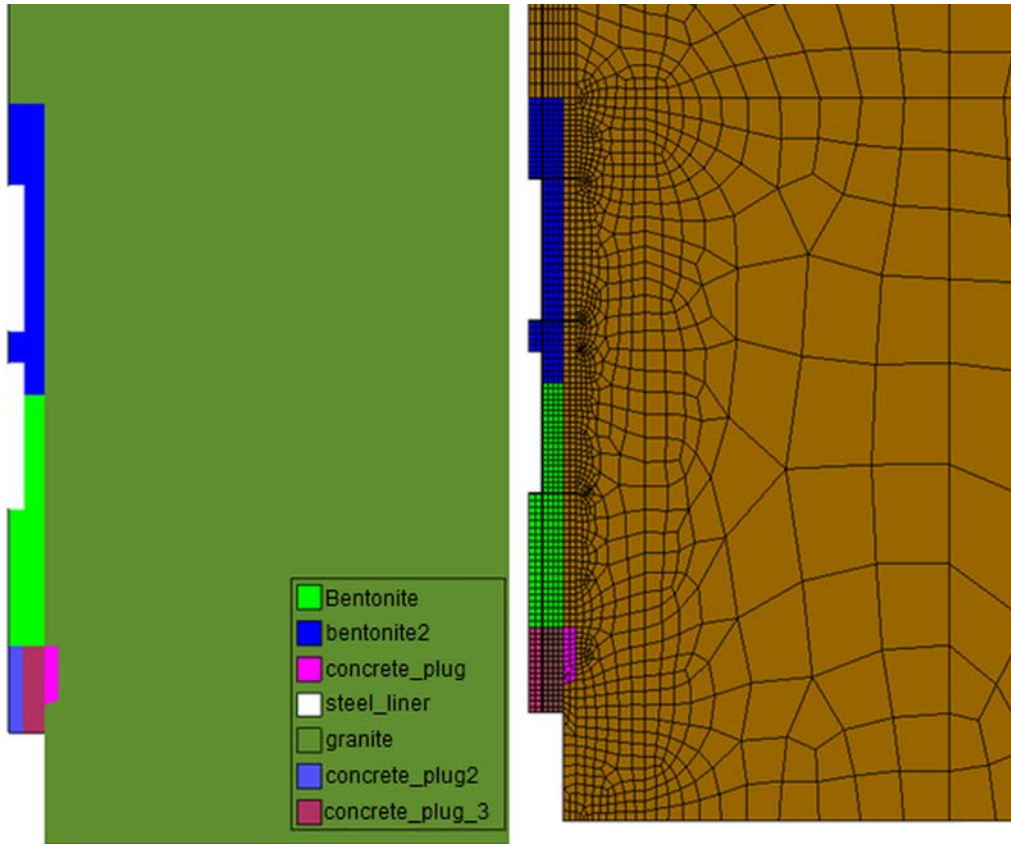
112 This model takes advantage of the unstructured mesh development for the discretization  
113 of the host rock. However, the buffer is discretized with structured mesh as this is more  
114 adequate in the zone of interest to reduce numerical errors and to facilitate the  
115 monitoring of calculated variables that should be compared with measurements.

116 Table 1 shows the intervals considered in the analysis up to 2500 days, i.e. before  
117 dismantling operations. In order to match the measured evolution of temperature during  
118 the experiment (see more details about this in Gens et al., 2009), the heat power was  
119 progressively modified (see table 1) until the 100°C target temperature was reached.  
120 From that point on, the temperature was prescribed at the value of 100°C in both  
121 heaters. Later, one of the heaters was removed but the other continued heating.

122 A mechanical boundary condition restraining the normal displacements has been  
123 applied all along the external boundary. Also, a constant pressure of 0.9 MPa is imposed  
124 on the external boundary. As indicated above, during intervals 3 to 5 the power of the  
125 heaters is prescribed and therefore the temperature increases. Once the temperature  
126 reaches the target value on the surface canister (100°C) the boundary condition on the  
127 heater is changed from constant power to constant temperature (more information about  
128 boundary condition implementation in the Code\_Bright User's Guide, downloadable  
129 from the Code\_Bright web page – [https://deca.upc.edu/en/projects/code\\_bright](https://deca.upc.edu/en/projects/code_bright)).

130 Tables 2 and 3 show the constitutive laws used in the model and the properties  
131 corresponding to the FEBEX bentonite as considered with single porosity and  
132 considering that it can be represented with the Barcelona Basic Model or BBM (Alonso  
133 et al., 1990), van Genuchten model, Darcy's law, power relative permeability and  
134 geometric mean for thermal conductivity. The parameters are divided in hydraulic and  
135 mechanical parameters and correspond to the Code\_Bright materials input window

136 (more information about this parameters in Code\_Bright User's Guide). More precise  
 137 information may be obtained from the model, which is available by request to the  
 138 authors.



139  
 140 Figure 1. Model geometry, materials and mesh considered for FEBEX updated  
 141 modeling (some materials are differentiated to simulate excavation only).

142  
 143 Table 1. Time intervals considered for the simulation up to 2500 days

| Time interval          | Initial Time (days) | Final Time (days) | Prescribed heat flow (W/heater) | Prescribed Temperature (°C) |
|------------------------|---------------------|-------------------|---------------------------------|-----------------------------|
| 1 (stress equilibrium) | 0                   | 20                | -                               | 12                          |
| 2 (construction)       | 20                  | 155               | -                               | 12                          |
| 3                      | 155                 | 176               | 1200                            | -                           |
| 4                      | 176                 | 197               | 2000                            | -                           |
| 5                      | 197                 | 207               | 2400                            | -                           |
| 6                      | 207                 | 2500              | -                               | 100                         |

144  
 145 Figure 2a shows the distribution of temperature around the 2 heaters at 21 days, 1.5  
 146 years and 5.5 years, i.e. during normal operation of the test and thus before dismantling.  
 147 It can be observed that the temperature in the vicinity of the heaters rises till it reaches a  
 148 value close to 100°C and the temperature of the surrounding host rock continues  
 149 increasing after that.

150 Figure 2b shows the distribution of degree of saturation. In general, the rock remains  
 151 saturated except at early times. As the bentonite is installed in the tunnel with a high  
 152 initial suction, some desaturation of the rock takes place. This zone is re-saturated very

153 soon as the water from the host rock is mobilized. At 5.5 years, the bentonite buffer is  
 154 not fully saturated, but the drying near the canister is less intense than at 1.5 years,  
 155 indicating that the saturation process continues.

156 Figure 2c shows the distribution of porosity in the bentonite. Clearly, swelling of the  
 157 bentonite takes place near the hydration boundary (i.e. the contact with the rock) while  
 158 contraction takes place in the zone of drying (i.e. the contact with the heaters). As it can  
 159 be seen in the figure, the range of porosity variation after 5.5 years is 0.378 to 0.413,  
 160 with an average very close to the initial value of 0.4.

161

162 Table 2. Hydraulic parameters for bentonite. See Code\_Bright User's Guide for further  
 163 details.

| Retention curve   |         |   |
|---|---------|---|
| P1: $P_o$ (MPa)   | 20      | Van Genuchten model:<br>$S_e = \frac{S_l - S_{rl}}{S_{ls} - S_{rl}} = \left[ 1 + \left( \frac{P_g - P_l}{P} \right)^{\frac{1}{1-\lambda}} \right]^{-\lambda} \quad P = P_o \frac{\sigma}{\sigma_o}$ |
| P2: $\sigma_o$ (Nm <sup>-1</sup> )                          | 0.072   |   |
| P3: $\lambda$   | 0.18    |   |
| P4: $S_{rl}$  | 0.01    |   |
| Intrinsic permeability                                      |         |   |
| P1: $(k_{11})_o$ (m <sup>2</sup> )                          | 1.9E-21 | Darcy's law: $\mathbf{q}_l = -\frac{\mathbf{k}k_{rl}}{\mu_l} (\nabla P_l - \rho_l \mathbf{g})$  |
| P2: $(k_{22})_o$ (m <sup>2</sup> )                          | 1.9E-21 |   |
| P3: $(k_{33})_o$ (m <sup>2</sup> )                          | 1.9E-21 | Kozeny's model: $\mathbf{k} = \mathbf{k}_o \frac{\phi^3}{(1-\phi)^2} \frac{(1-\phi_o)^2}{\phi_o^3}$   |
| P4: $\phi_o$  | 0.4     |   |
| P5: $\phi_{min}$  | 0.001   |   |
| Liquid phase relative permeability                          |         |   |
| ITYCL   | 6       | $k_{rl} = AS_e^\lambda$   |
| P2: A   | 1       |   |
| P3: $\lambda$   | 3       |   |
| Diffusive flux of vapor                                     |         |   |
| ITYCL   | 1       | Fick's law for molecular diffusion:<br>$\mathbf{i}_g^w = -(\tau \phi \rho_g S_g D_m^w \mathbf{I}) \nabla \omega_g^w D_m^{vapor} = D \left[ \frac{(273.15+T)^n}{P_g} \right]$                        |
| P1: $D$ (m <sup>2</sup> s <sup>-1</sup> K <sup>-n</sup> Pa) | 5.9E-6  |   |
| P2: $n$   | 2.3     |   |
| P3: $\tau_o$  | 0.8     |   |
| $\tau = \text{constant} = \tau_o$                           |         |   |
| Conductive flux of heat 1                                   |         |   |
| ITYCL   | 1       | Fourier's law: $\mathbf{i}_c = -\lambda \nabla T$<br>$\lambda = \lambda_{sat}^{S_l} \lambda_{dry}^{(1-S_l)}$  |
| P1: $\lambda_{dry}$ (WmK <sup>-1</sup> )                    | 0.47    |   |
| P2: $\lambda_{sat}$ (WmK <sup>-1</sup> )                    | 1.15    |   |

164

165 Figure 3 shows the variation of temperature at some points within the bentonite and the  
 166 comparison between the Operational Base Case (OBC) model calculations –  
 167 corresponding to the analysis reported in Gens et al. (2009)– and the current model. It  
 168 can be seen that the temperature is reproduced in the same way by the model as it was  
 169 obtained with the OBC model, so this new model is equivalent to that one. It is  
 170 important to recall that one of the objectives of this contribution was to prepare a new  
 171 geometry and mesh (GiD environment) to be able to consider additional cases.

172 A similar response is obtained in terms of relative humidity (Figure 4) and the  
 173 correspondence of the present model with the OBC model shows that no significant  
 174 variation has occurred. Actually, small variations on the parameters or shapes of the  
 175 constitutive equations do not produce a significant variation of relative humidity. This

176 shows that it is possible that some additional mechanisms of water flow (in addition to  
 177 Darcy and Fick equations) may play a role and may help to explain the different  
 178 hydration rates.

179 The evolution of stresses in the bentonite is represented in Figure 5. The response shows  
 180 the same limitations as observed in the OBC especially concerning the overestimation  
 181 of stress development at early stages of hydration. The low measured stresses can be  
 182 interpreted qualitatively in different ways, one of them being the possible hydration of  
 183 the bentonite aggregates controlled by double structure effects.

184

185 Table 3. Mechanical parameters for bentonite. Thermo-elasto-plastic (TEP) model. See  
 186 Code\_Bright User's Guide for further details.

| Elastic parameters               |        |  |
|----------------------------------|--------|--|
| ITYCL                            | 1      |  |
| P1: $k_{io}$                     | -0.05  |  |
| P2: $k_{so}$                     | 0.25   | $d\varepsilon_v^e = \frac{k_i(s)}{1+e} \frac{dp'}{p'} + \frac{k_s(p',s)}{1+e} \frac{ds}{s+0.1} + (\alpha_o)dT$ |
| P3: $K_{min}$ (MPa)              | 0.1    |  |
| P5: $\nu$                        | 0.4    |  |
| P8: $\alpha_i$                   | -0.003 |  |
| P9: $\alpha_{sp}$                | -0.161 |  |
| P10: $p_{ref}$ (MPa)             | 0.01   |  |
| Thermal parameters               |        |  |
| ITYCL                            | 1      | $k_i(s) = k_{io}(1 + \alpha_i s)$  |
| P1: $\alpha_o$                   | 1.5E-4 | $k_s(p',s) = k_{so}(1 + \alpha_{sp} \ln p'/p_{ref})$   |
| P5: $T_{ref}$ (°C)               | 20     |  |
| Plastic parameters 1             |        |  |
| ITYCL                            | 1      |  |
| P1: $\lambda(0)$                 | 0.15   | $p_o = p^c \left( \frac{p_o^*(T)}{p^c} \right)^{\frac{\lambda(o)-k_{io}}{\lambda(s)-k_{io}}}$                  |
| P2: $r$                          | 0.925  |  |
| P3: $\beta$ (MPa <sup>-1</sup> ) | 0.05   | $p_o^*(T) = p_o^*$   |
| P4: $\rho$ (°C <sup>-1</sup> )   | 0.2    |  |
| P5: $k$                          | 0.1    |  |
| P6: $p_{so}$ (MPa)               | 0.1    | $\lambda(s) = \lambda(o)[(1-r)\exp(-\beta s) + r]$   |
| Plastic parameters 2             |        |  |
| ITYCL                            | 1      |  |
| P1: $p^c$ (MPa)                  | 0.5    | $p_s = p_{so} + k s \exp(-\rho \Delta T)$  |
| P2: $M$                          | 1      |  |
| P3: $\alpha$                     | 0.53   | $\Delta T = T - T_{ref}$   |
| P4: $e_o$                        | 0.6    |  |
| P5: $p_o^*$ (MPa)                | 12     |  |

187

188 Finally, the evolution of water pressure in the rock is represented in Figure 6. This  
 189 model uses 0.9 MPa as a boundary condition in the outer boundary. From the  
 190 observation of the results (overestimated pressure in the rock roughly by 0.05 MPa), it  
 191 is probable that 0.85 MPa would be more appropriate. This will be modified in future  
 192 models but the effect on the other variables is expected to be very small taking into  
 193 account the large suction gradients in the bentonite.

194

195 **3. Long term results**

196 Once the presented model was calibrated with real measurements, a long term  
197 prediction has been performed. This section contains some results corresponding to the  
198 model considering a maximum calculation time of 20 years and the incorporation of the  
199 operations of shut down and dismantling of one canister, concrete abutment and  
200 bentonite. The distribution of temperatures is shown in Figure 7a. It can be observed  
201 that the operations of dismantling affect the canister close to the access drift together  
202 with part of the barrier, while the second canister heating continues. Figure 7b shows  
203 the distribution of degree of saturation and Figure 7c the distribution of mean stresses.  
204 Degree of saturation indicates that at 20 years, full saturation has not been achieved yet.

205 The evolution of temperatures is presented in Figure 8. Temperatures near the  
206 dismantled heater show that the cooling takes place quite fast and that it has only a  
207 small influence on section D1. In section F2 and G, the effect is very important and the  
208 model captures well the rapid fall of the temperature rapid down to the measured lower  
209 values due to the new configuration of the test after partial dismantling.

210 Pore pressures are represented in Figure 9. The negative pore pressures correspond  
211 actually to suction. The strong evaporation induced by heating near the canister  
212 produces a large increase of suction. The maintained heating in section F2 delays the  
213 pore pressure increase near the canister. This is because the evaporation is strong and  
214 vapor migration compensates the water flow from the rock. In section G, which  
215 undergoes cooling caused by the canister removal, the water pressure recovers more  
216 rapidly at the same time that cooling takes place. The rate of water pressure recovery  
217 increases rapidly.

218 Relative humidity (HR) depends on pressure and temperature and can be calculated  
219 from these variables at any point. This is necessary in order to compare the calculated  
220 variables with the measurements of relative humidity at the sensors. The comparison is  
221 shown in Figure 10. Note that there is not real data for  $r = 0.52$  m at section F2. While  
222 in section F2, near the heater, relative humidity does not reach 100%, which would  
223 correspond to full saturation, the 100% relative humidity is reached after 20 years at  
224 section H, indicating that when the heating stops, full saturation is reached more  
225 rapidly.

226 Figure 10 also shows the evolution of stresses at sections F2 and G. Note that sections H  
227 and G can be considered equivalent sections, since they are very close to each other (see  
228 Figure 3) and since neither there is data for stress at section H nor there is data for  
229 relative humidity at section G. The stress development shows also the effect of the  
230 dismantling but it is less pronounced in the central zone of the prevailing heater.  
231 Stresses reach a value in the range of 4 to 5 MPa according to the model calculations.  
232 The correspondence with the measurements is quite good at F2. In contrast, at section G  
233 in which the measurements show a delay in the response, the sensor seems to have been  
234 reactivated when dismantling took place. This model will be considered the base case  
235 for the additional processes assessment that will be described next.

236

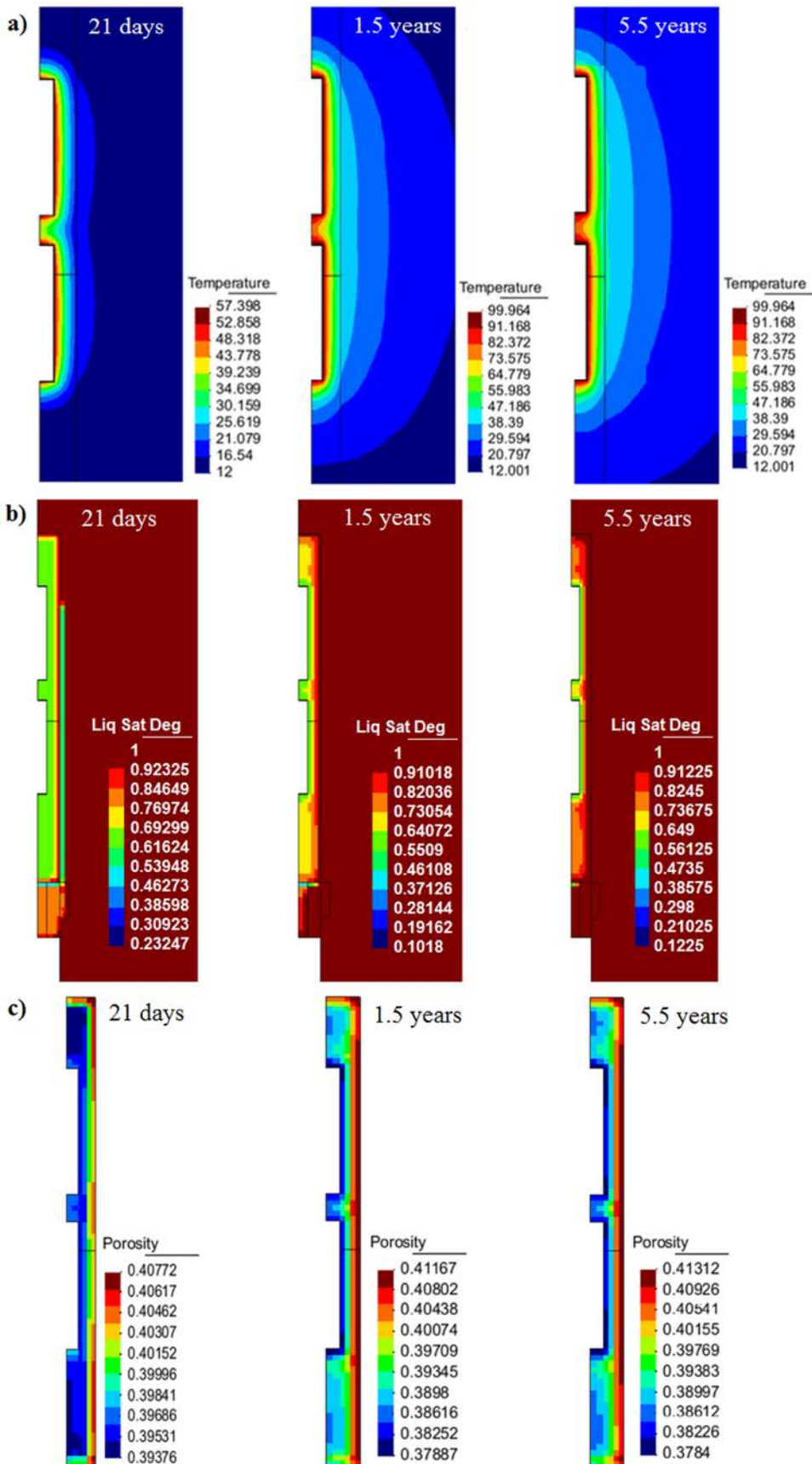
237 **4. Additional processes assessment**

238 The model presented above will be considered from now on as a base case for the  
239 comparison with the results from alternative models including additional processes such  
240 as the incorporation of a coupled flow induced by a thermal gradient, the variation of



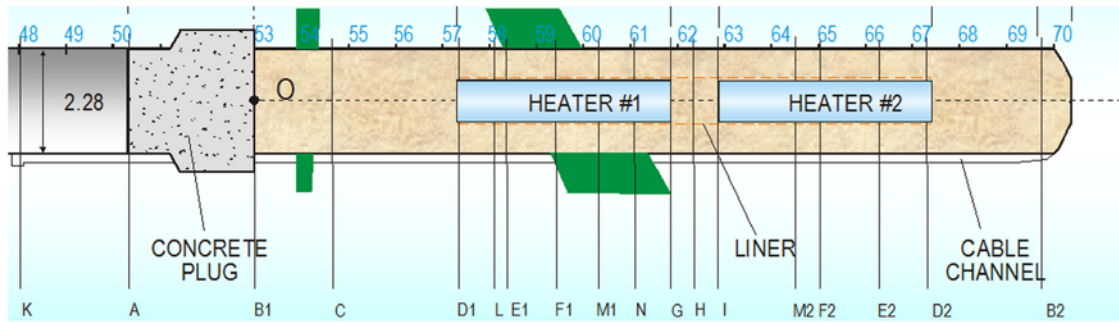
241 the intrinsic permeability or the introduction of a double porosity model, for which the  
 242 impact on hydration and stress development has been assessed.

243

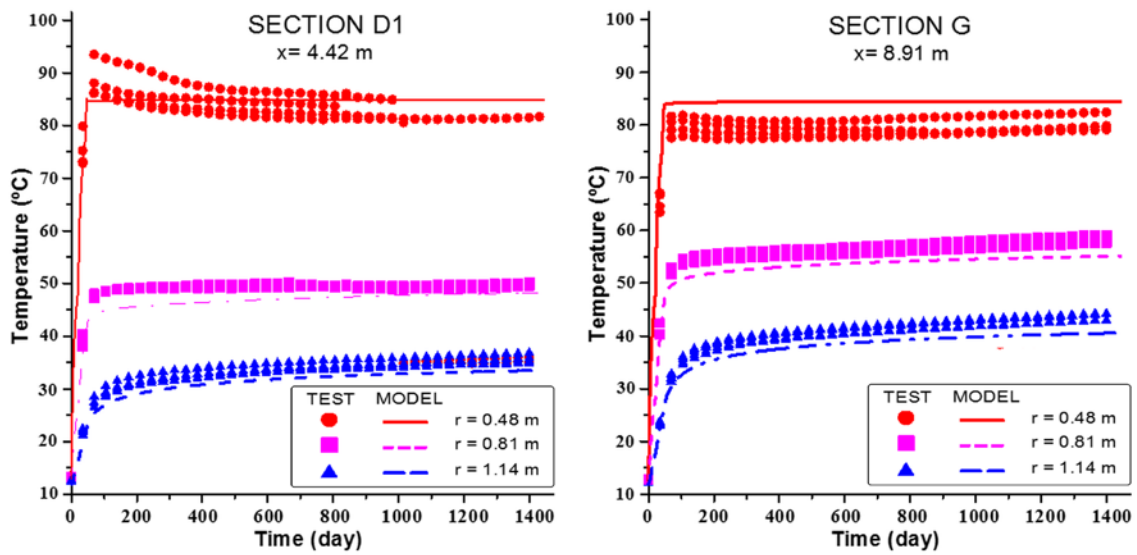


244

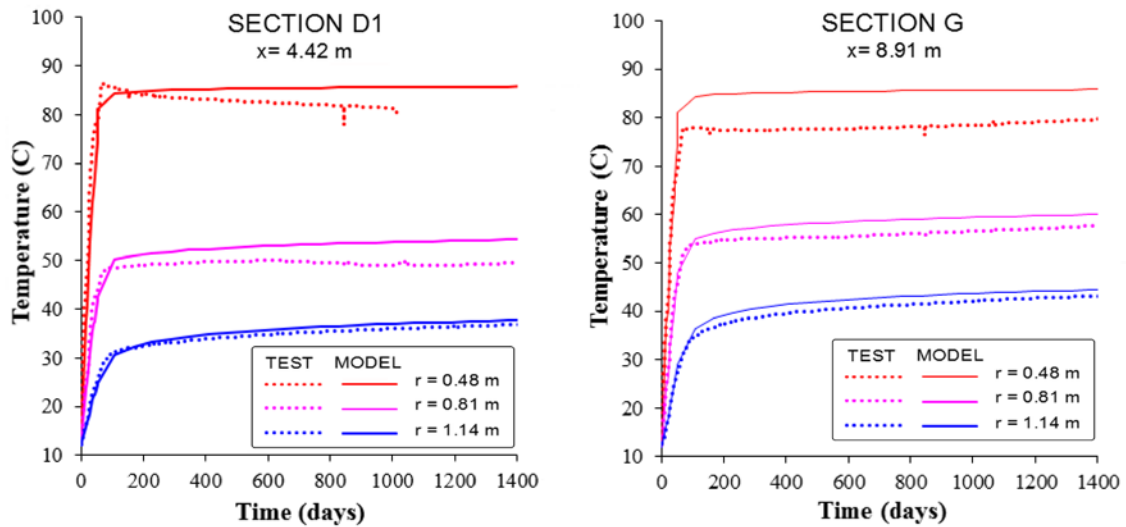
245 Figure 2. Distributions at 21 days, 1.5 years and 5.5 years of: a) temperature; b) degree  
 246 of saturation; c) porosity.  
 247



a) Results from OBC case as compared with measurements



b) Results from the current new model



248  
 249 Figure 3. Temperature evolution at different points.  
 250

251 **4.1. Thermal gradient induced water flow**

252 The water flow induced by a thermal gradient can be expressed as an additional flux to  
 253 the usual Darcy's flux. The generalized advective flux including thermal osmosis is  
 254 expressed as in Eq. (1):

$$255 \quad \mathbf{q}_l = -\frac{\mathbf{k}k_{rl}}{\mu_l}(\nabla p + \rho_l g \nabla z) - k_{rl}k_T \nabla T \quad (1)$$

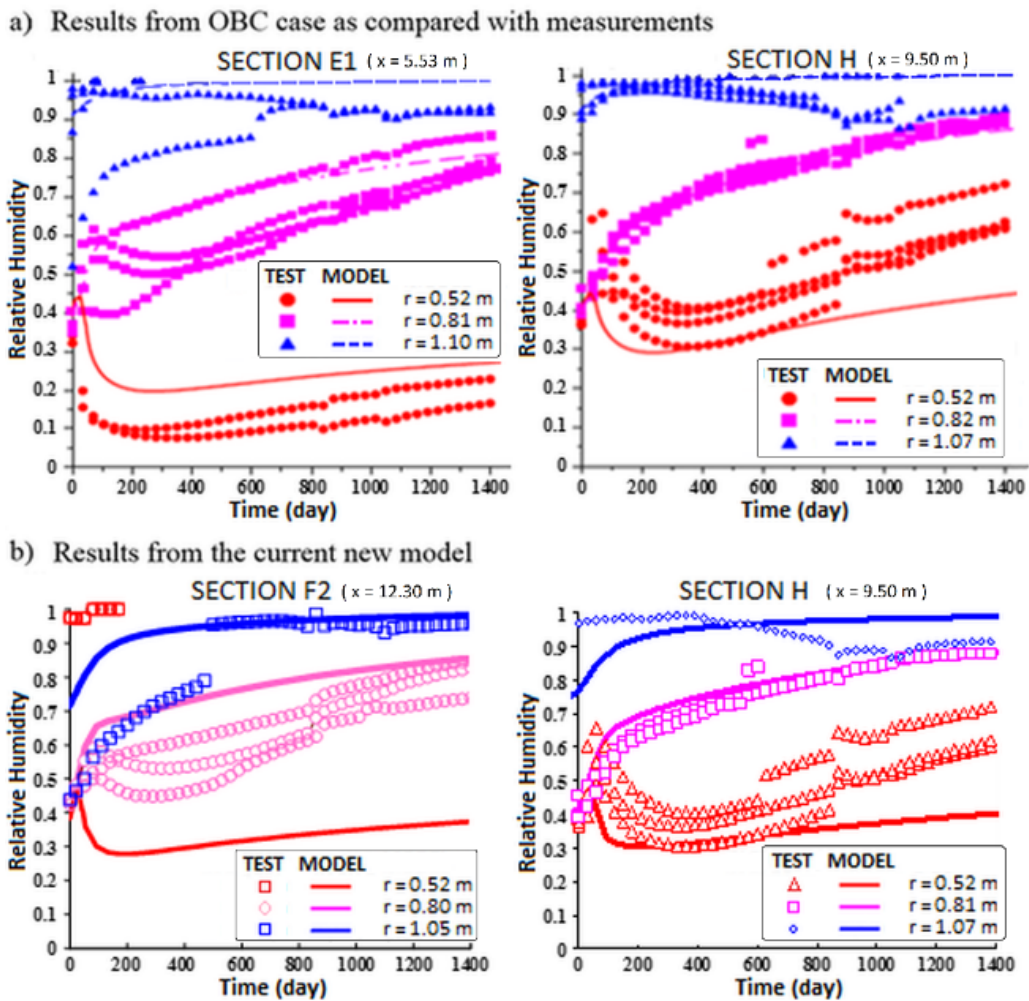
256 Where  $\mathbf{k}$  [ $\text{m}^2$ ] denotes the intrinsic permeability tensor,  $k_{rl}$  [-] denotes the relative  
 257 permeability of the liquid,  $g$  is the gravity vector,  $\mu_l$  is the dynamic viscosity,  $p$  [MPa]  
 258 is the hydraulic pressure,  $T$  [K] is temperature,  $\rho_l$  is liquid density, and  $k_T$  [ $\text{m}^2/\text{K}/\text{s}$ ] is  
 259 the thermo-osmotic permeability.

260 This is the flux of liquid phase induced by both pressure and temperature gradients. It is  
 261 an advective flux in the sense that it drags the water. The total flux of water is then  
 262 written as in Eq. (2):

$$263 \quad \mathbf{j}_l^w = -D\rho_l \nabla w_l^w + \rho_l \mathbf{q}_l \quad (2)$$

264 Where  $D$  [ $\text{m}^2/\text{s}$ ] is the diffusion coefficient, which actually corresponds to the dissolved  
 265 species, in this case the air, which is dissolved in the water.

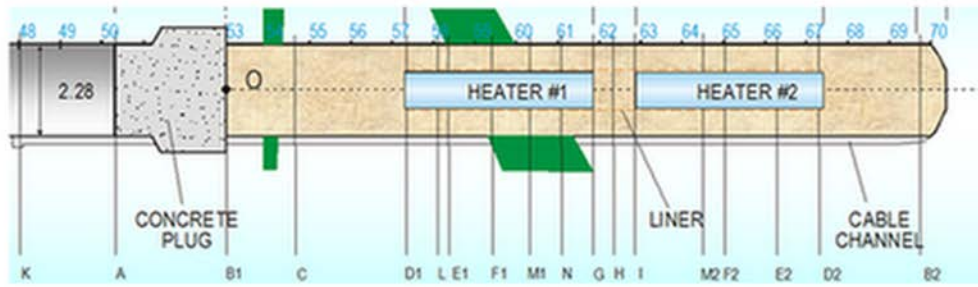
266



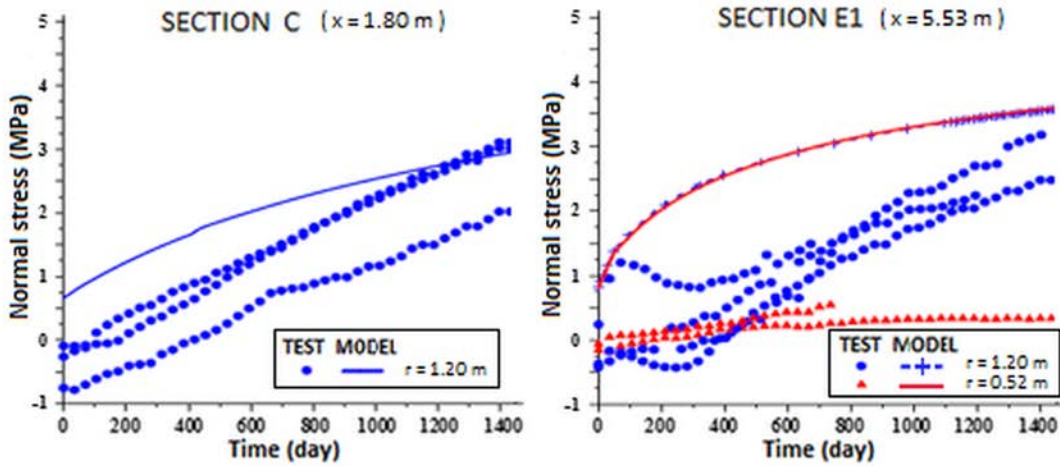
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268 Figure 4. Relative humidity evolution.

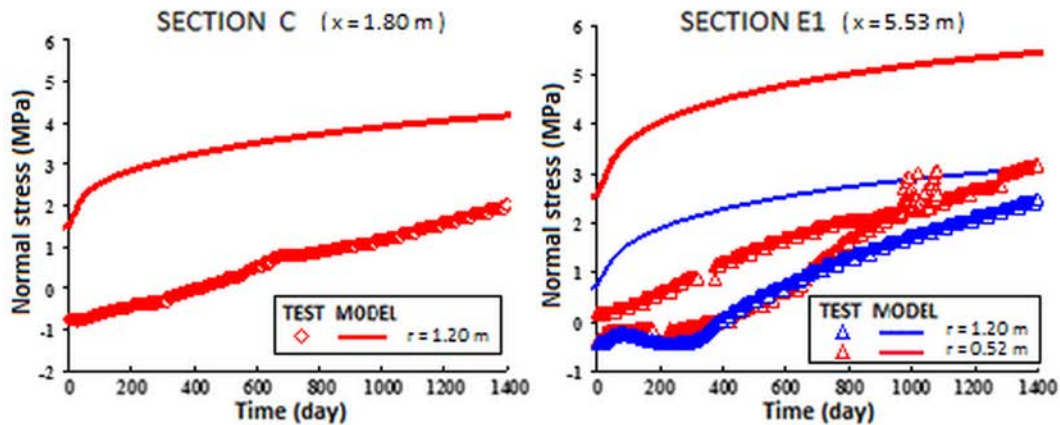
269



a) Results from OBC case as compared with measurements



b) Results from the current new model

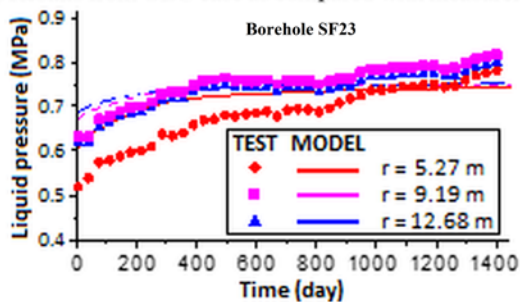


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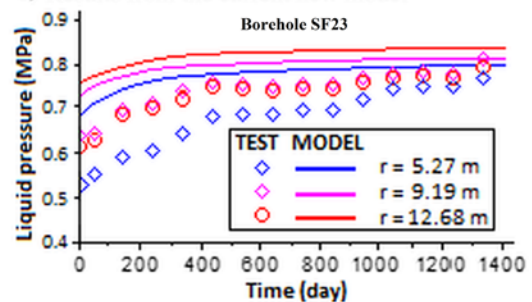
271 Figure 5. Evolution of stresses

272

a) Results from OBC case as compared with measurements



b) Results from the current new model

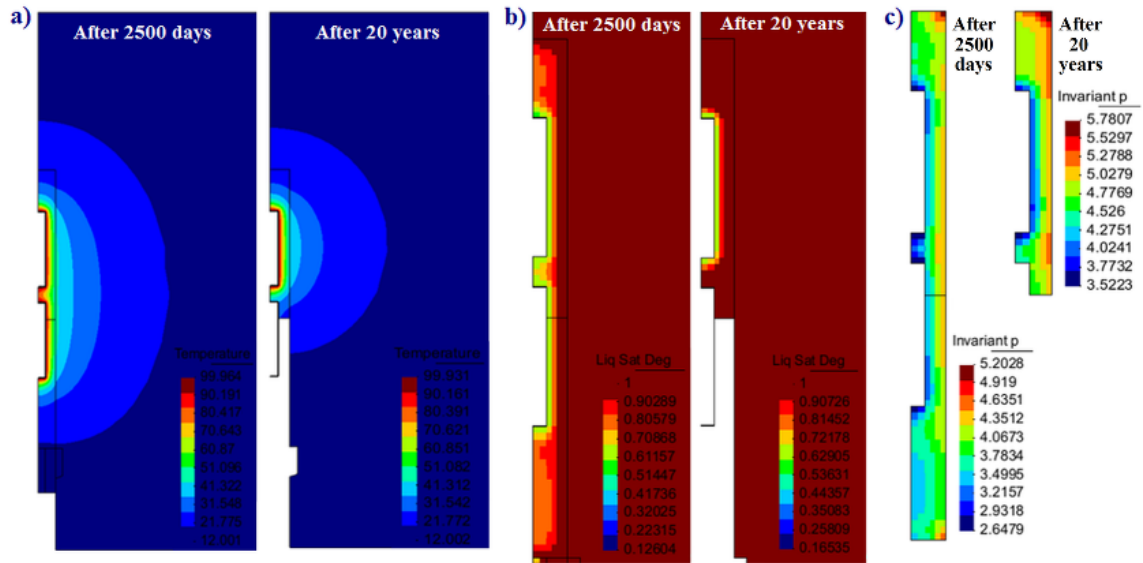


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274 Figure 6. Evolution of pressure in the rock.

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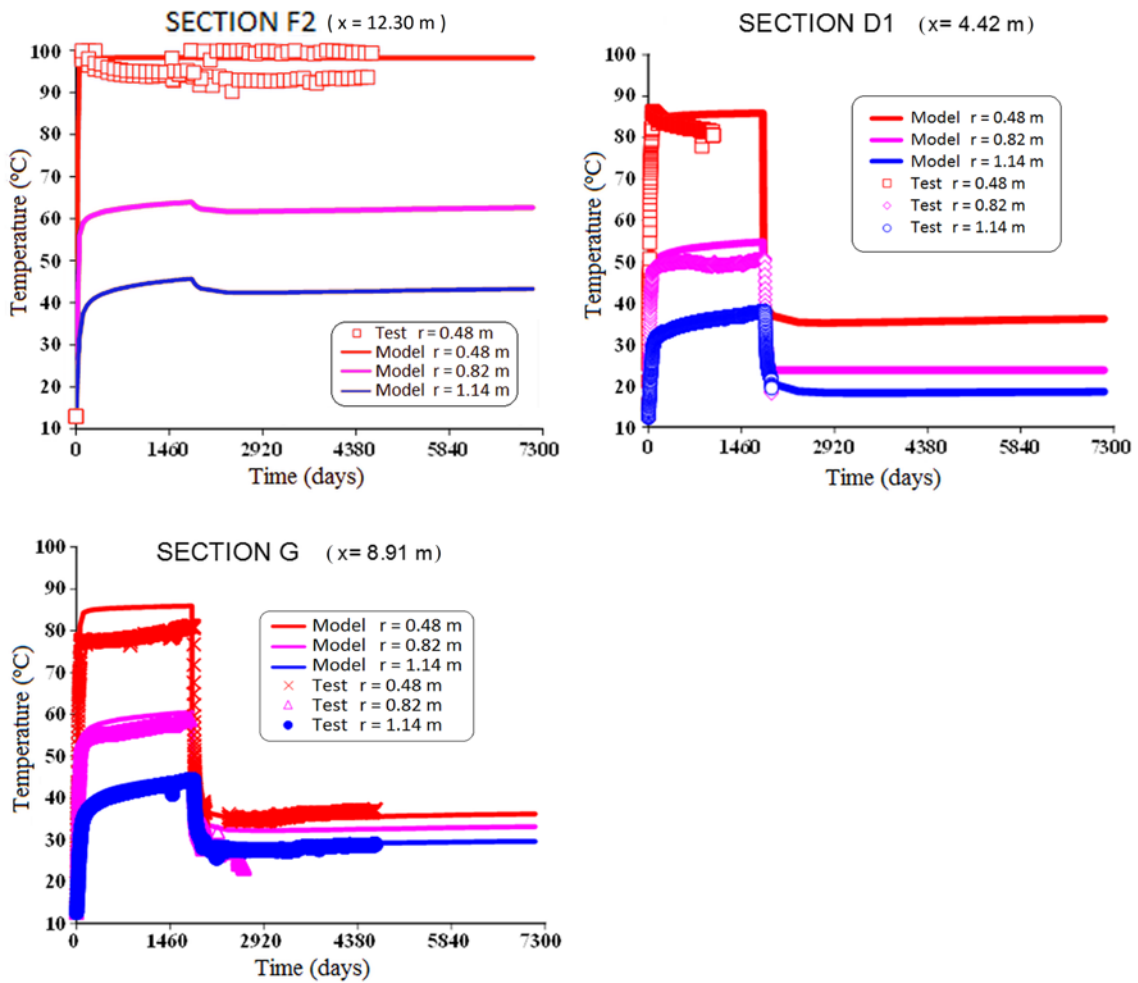




276

277 Figure 7. Distributions for 2500 days and for 20 years of: a) temperature; b) distribution  
 278 of degree of saturation; c) distribution of mean stress.

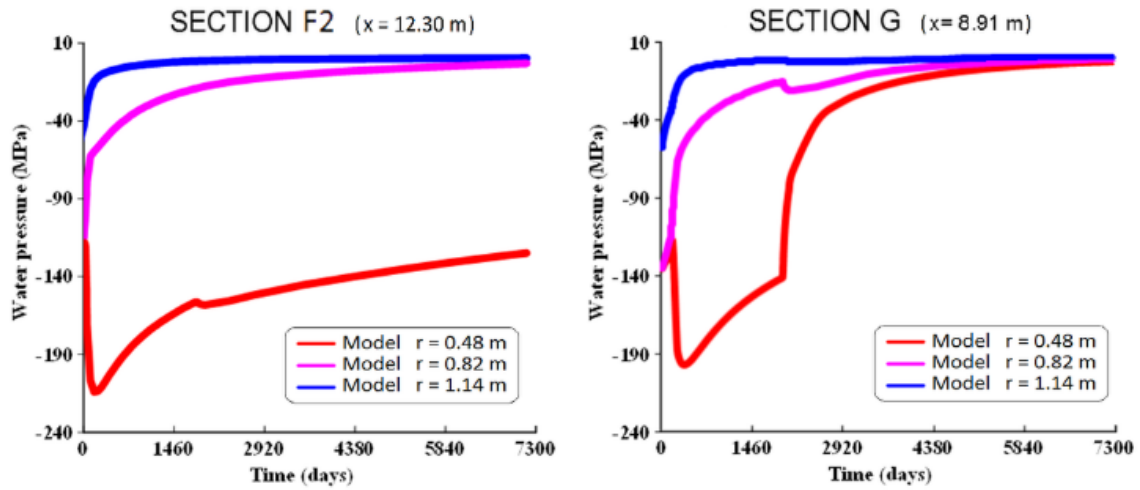
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281 Figure 8. Evolution of temperature at sections D1, F2 and G.

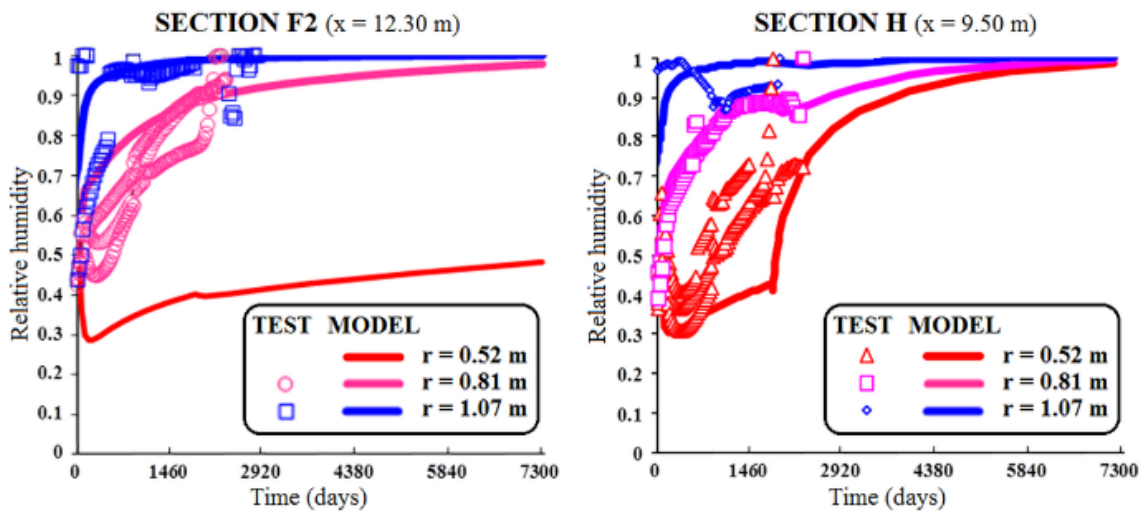
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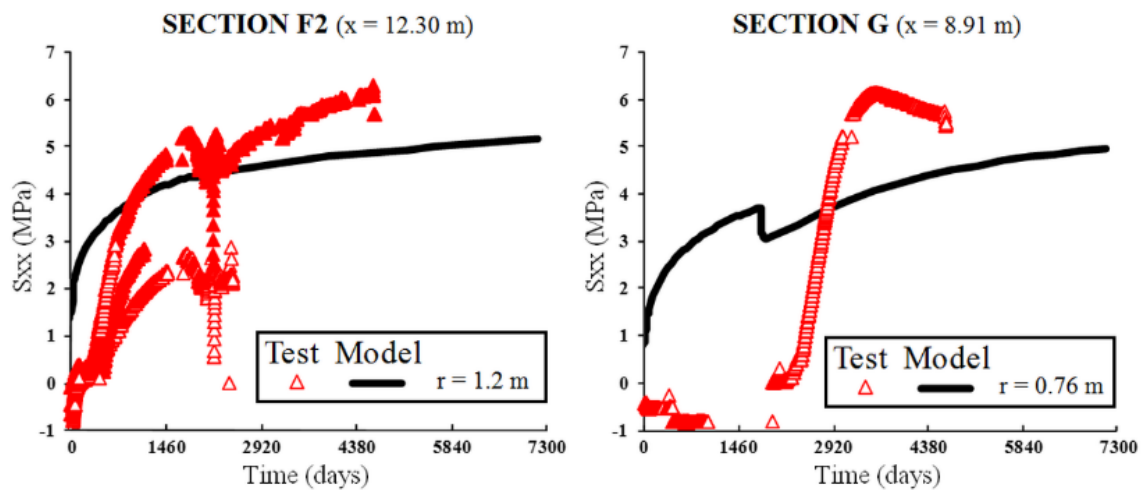
283

284 Figure 9. Evolution of pressure at sections F2 and G.

285



286



287

288 Figure 10. Evolution of relative humidity (up) and stress (down) at different sections for  
 289 the base case.

290

291 Soler (2001) describes the variations of thermo-osmotic permeability in the following  
 292 way: "Reported values of  $k$  for Na-saturated kaolinite and Na bentonite (Dirksen, 1969),

293 at an average temperature of 25 °C and various temperature gradients and porosities,  
 294 range between  $10^{-14}$  and  $3 \cdot 10^{-13}$  m<sup>2</sup>/K/s. In another study [11] thermal osmosis across a  
 295 kaolinite membrane was investigated. An estimate of  $k$  from the data of this study  
 296 (Srivastava, 1975), corresponding to conditions similar to the ones giving the high  $k$   
 297 values (Dirksen, 1969), yields a value of  $2.6 \cdot 10^{-10}$  m<sup>2</sup>/K/s, which is three orders of  
 298 magnitude greater. As a first approximation, the values of  $k$  mentioned above can be  
 299 used to define a range of values  $k_T = 10^{-14} - 10^{-10}$  m<sup>2</sup>/K/s”.

300 In order to compare this value with the value corresponding to the pressure gradient (i.e.  
 301 the dominant effect), it is necessary to calculate the ratio between intrinsic permeability  
 302 and viscosity. For the intrinsic permeability value of the Febex bentonite a value of  
 303  $1.9 \cdot 10^{-21}$  m<sup>2</sup> which divided by the viscosity of water 0.001 Pa·s, leads to:  $1.9 \cdot 10^{-12}$   
 304 m<sup>2</sup>/MPa/s. These two values are comparable.

305 In order to discriminate the relative dominance of fluxes in this case, it is also necessary  
 306 to introduce the gradients of pressure and temperature. While the temperature gradient  
 307 can be estimated easily and does not change very much with time (although temperature  
 308 and water content change), the pressure gradient varies more, in space and time.  
 309 Therefore, depending on the zone in the bentonite and the time, the advective flux may  
 310 be influenced by thermal osmosis and this should be investigated.

311 In order to implement in an easy way the thermal gradient term, Equation (1) can be  
 312 written as:

$$313 \quad \mathbf{q}_l = -\frac{k k_{rl}}{\mu_l} (\nabla p + \alpha \nabla T + \rho_l g \nabla z) \quad (3)$$

314 In FEBEX, the pressure and the temperature gradients are opposite in sign. Actually the  
 315 thermo-osmotic flux delays hydration and therefore it may explain some delayed  
 316 effects. The parameter  $\alpha$  in this equation is the ratio between the thermo osmotic  
 317 conductivity and the hydraulic conductivity following Equation (3). As we have seen  
 318 that the order of magnitude can be similar, it follows that  $\alpha=1$  gives a condition in  
 319 which the thermal osmosis induced flux is comparable to the hydraulic flux. This  
 320 situation occurs because the bentonite has very low hydraulic conductivity and therefore  
 321 it is possible to find a value of the conductivity for thermal osmosis which leads to  $\alpha=1$   
 322 (i.e.  $1.9 \cdot 10^{-12}$  m<sup>2</sup>/K/s, the value calculated above but in the units of  $k_T$ ).

323 The evolution of relative humidity in the case in which the thermal gradient induced  
 324 flux is incorporated is shown in Figure 11 (upper left) and Figure 12 (upper left), for  
 325 section F2. The water flow induced by the temperature gradient delays hydration.  
 326 Actually, for the values of the parameters considered, it manages to compensate the  
 327 water inflow up to a point that for steady state temperature conditions, the relative  
 328 humidity stabilizes and does not tend to 100%. This means that the bentonite barrier is  
 329 not fully saturated and hence the isolation would be achieved.

330 The evolution of the stress development is also shown in Figure 11 (lower left) and  
 331 Figure 12 (lower left). Since the hydration reaches a point that does not progress, and  
 332 relative humidity stabilizes, the stress also stabilizes below the swelling pressure.

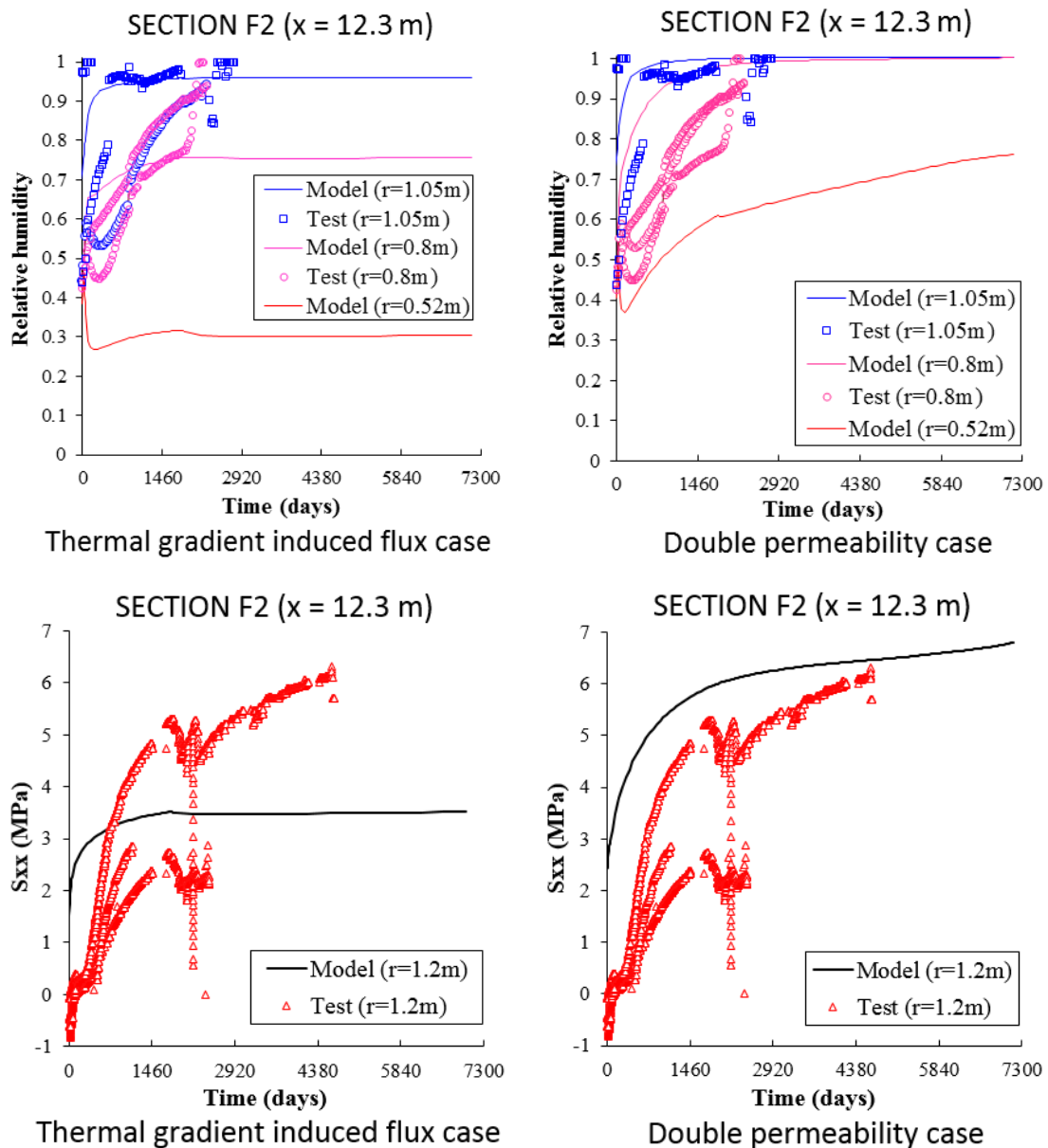
333 Compared to the base case (see Figure 10), we can observe that the results of the base  
 334 case are closer to the real data than those obtained considering this thermal gradient  
 335 induced flux. In fact, considering this induced flux, the relative humidity seems to stay  
 336 constant with time, at least at points not far from the heater, which is not the result that  
 337 the real measurements are given.

338

339 **4.2. Sensitivity to permeability**

340 An analysis about the sensitivity of the model to permeability changes has also been  
 341 performed, since permeability is the variable subjected to bigger variations due to  
 342 changes in density, structure and construction of the barrier, measurement difficulties,  
 343 etc. For the base case presented in this paper, permeability of the bentonite has been  
 344 increased by a factor of 2, which is a reasonable variation. This accelerates somewhat  
 345 the hydration as the permeability of the buffer controls the hydration rate in FEBEX.  
 346 Moreover, as liquid pressure increases faster due to permeability increase, the stress  
 347 developed also increases faster leading to values which are still comparable to the  
 348 measurements (Figures 11 and 12).

349

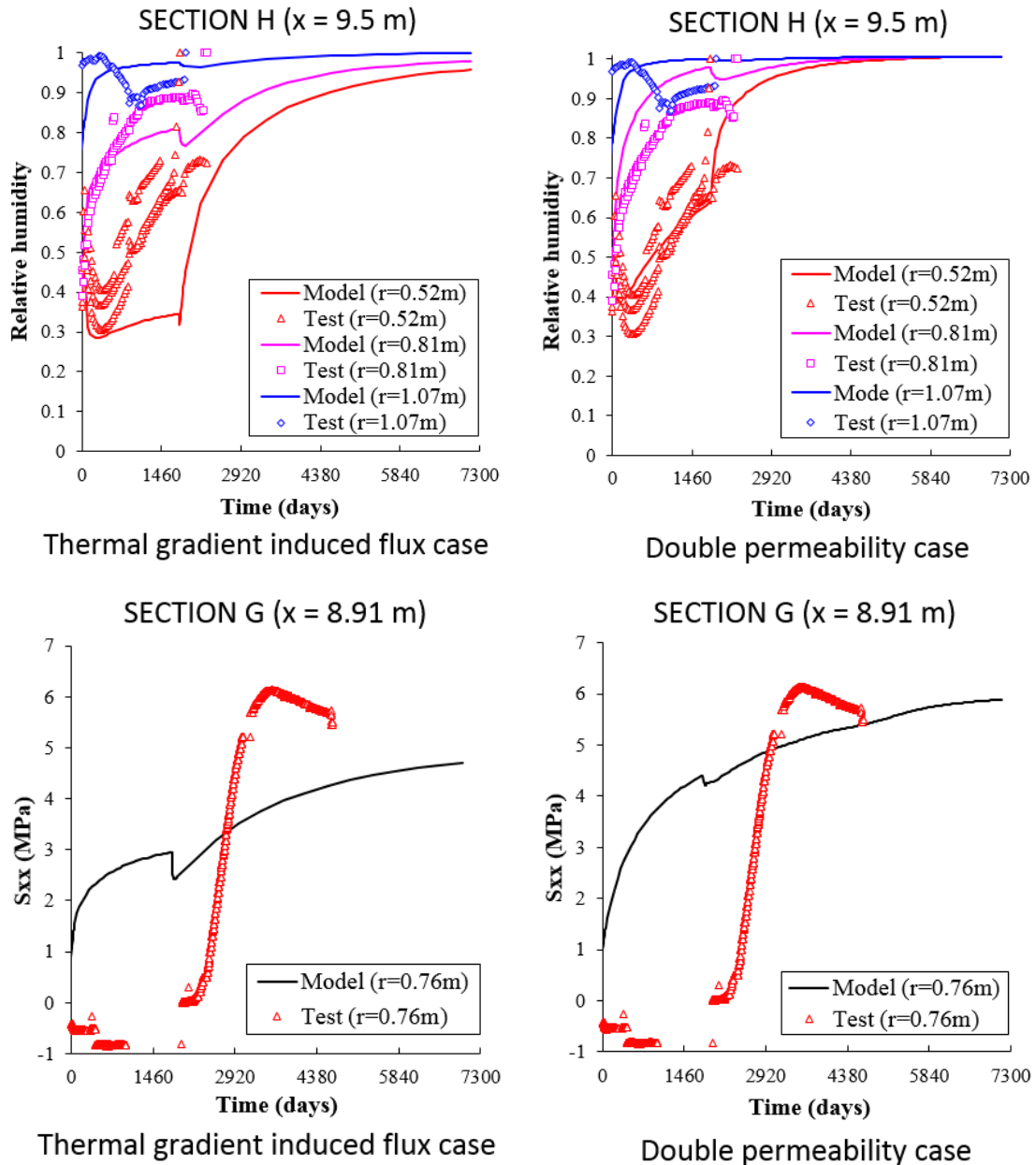


350 Figure 11. Evolution of relative humidity (up) and stress (down) at section F2 for two  
 351 additional cases: thermal gradient induced flux (left) and double permeability (right).

352



353 Actually, the stress development in the case of double permeability is quite consistent  
 354 with measurements, giving closer values to the real data than the base case. It can also  
 355 be observed that even doubling the permeability, full saturation of the bentonite barrier  
 356 is not achieved, even after 20 years.  
 357



358 Figure 12. Evolution of relative humidity (up) and stress (down) at section H for relative  
 359 humidity and G for stresses for two additional cases: thermal gradient induced flux (left)  
 360 and double permeability (right).

361

### 362 4.3. Double porosity model

363 The effect on the hydration and swelling including a double structure model is considered  
 364 in this section. The double porosity model is an extension of the BBM approach presented  
 365 above. The double structure effects are considered in the following way: two structural

366 levels are considered with different suctions. The two suctions are related by an  
 367 interaction term.

368 Equation 4 shows the simplified mass balance equations incorporating the microstructure.  
 369 It is assumed that porosity can be decomposed into two adding contributions.

$$\frac{\partial}{\partial t} \left( \phi_M \rho_{IM}^w S_{IM} + \phi_M \rho_{gM}^w S_{gM} \right) + \nabla \cdot \left( \mathbf{j}_{IM}^w + \mathbf{j}_{gM}^w \right) = -\Gamma^w \quad (M) \quad (4)$$

$$\frac{\partial}{\partial t} \left( \phi_m \rho_{lm}^w S_{lm} \right) = \Gamma^w \quad (m)$$

370 Where an interchange term has been introduced. This term is calculated as a function of  
 371 the suctions in the micro pores and the macro pores (Eq. 5).

$$\Gamma^w = \alpha_{micro-macro} (s_m - s_M) \quad (5)$$

372 Where  $s_m$  is the suction micro and  $s_M$  corresponds to the suction macro. The macro and  
 373 micro porosities are defined following Eq. 6.

$$\phi = \phi_M + (1 - \phi_M) \phi_m = \phi_M + \phi_m \quad (6)$$

374 Where the porosity of the microstructure may be defined in two different ways, i.e. with  
 375 respect to the volume of the microstructure or with respect to the total volume. The  
 376 backfill material in the FEBEX is composed by blocks. The macro porosity is assumed  
 377 to be 0.40 while the micro porosity is set to 0.15 in these model calculations.

378 The combination of the double structure and the BBM model consists –in the model used  
 379 here– in using the suction in the microstructure for the mechanical calculations. For  
 380 instance, for the swelling terms, the volumetric elastic deformation is calculated as in Eq.  
 381 7.

$$d\varepsilon_v = \frac{\kappa_p}{1+e} \frac{dp'}{p'} + \frac{\kappa_s}{1+e} \frac{ds_m}{(s_m + p_{atm})} \quad (7)$$

382 In addition, the micropores contribute to water storage and a specific retention curve is  
 383 considered, which essentially imposes that the micropores remain saturated (Eq. 8).

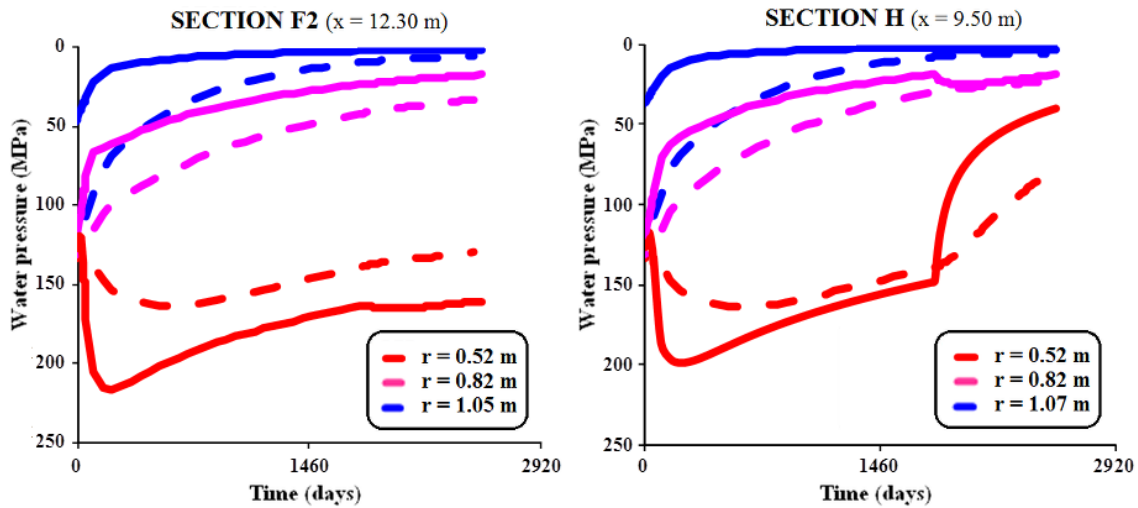
$$S_{lm} = \exp \left( -\frac{(s_m)^{\beta_{micro}}}{P_{o\ micro}} \right) \quad (8)$$

384 Where  $\alpha_{micro-macro} = 5 \times 10^{-12} \text{ kg/MPa/m}^3/\text{s}$ ,  $\beta_{mic} = 1.150$  and  $P_{o\ micro} = 700 \text{ MPa}$ .

385 The results of suction development are shown in figure 13. The micro suction shows a  
 386 smoothed response when sharp variations of macro suction take place. This is an expected  
 387 response as the water flow from macro to micro and vice-versa is delayed by the  
 388 interaction parameter  $\alpha$ .

389 Moreover, the results of the stress development when the double porosity is considered  
 390 are quite similar to the results obtained in the base case, and thus are not shown in this  
 391 paper for the sake of brevity. The swelling stress in the FEBEX experiment is an  
 392 integrated response of expansion of the backfill in a heterogeneous way as the process of  
 393 hydration is modified by the heating. The suction in the micro and macro pores is not  
 394 equal but it may happen that the averaged values are not so different. Actually, this is  
 395 controlled by the capacity to hydrate and this, in turn, is controlled by the permeability of  
 396 the backfill.

397



398

399 Figure 13. Evolution of macro and micro suction. Dashed line is considered for micro  
 400 suction while the continuous line corresponds to the macro suction.

401

## 402 5. Conclusions

403 This paper presents a model that has been prepared based on previous modeling efforts  
 404 performed for FEBEX. The model produces similar results as the previous models  
 405 taking into account that some slight modifications have been introduced on the  
 406 constitutive equations (for instance, the retention curve). Once calibrated, the long term  
 407 response of the model predictions up to 20 years has also been performed, taking in  
 408 consideration the operations of shut down and dismantling of the first canister, concrete  
 409 abutment and bentonite. The evolution of temperature, degree of saturation, stress,  
 410 water pressure, relative humidity have been analyzed.

411 The model has been used to analyze the effect of various processes, such as the  
 412 incorporation of thermal osmosis, the variation of intrinsic permeability and the  
 413 introduction of a double porosity model. The incorporation of thermal-osmosis leads to  
 414 the possibility of developing a steady state situation in which the water flow induced by  
 415 the thermal gradient is compensated by the water flow induced by the hydraulic  
 416 gradient. This implies that the clay will not reach a full saturation even for long term  
 417 conditions.

418 The vapor flux, although it can also be related to a thermal gradient is no able to reach  
 419 this condition as it vanishes when the gas degree of saturation reduces. As the material  
 420 progressively saturates, the vapor flux reduces and finally is zero at full saturation. In  
 421 contrast, the liquid water flow induced by a thermal gradient increases with water  
 422 saturation and reaches its maximum for a given thermal gradient, at full saturation.

423 The intrinsic permeability has been shown to be a parameter that permits a significant  
424 improvement simply by varying it by a factor of two. In fact, the results obtain for the  
425 double permeability case are closer to real data than the base case, meanwhile the  
426 thermal gradient induced flux case gives worse results than the base case, if we compare  
427 them with the real data.

428

429 Finally, the double structure model has given different results for micro and macro  
430 suction but we found that does not have great influence in the general behavior of the  
431 model, as the stress development obtained is quite similar to the base case.

432

433

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