

# In situ X-ray diffraction of silicate liquids and glasses under dynamic and static compression to megabar pressures

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1	In situ X-ray diffraction of silicate liquids and glasses under dynamic and
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# 24 Abstract

25 Properties of liquid silicates under high pressure and high temperature conditions are 26 critical for modeling the dynamics and solidification mechanisms of the magma ocean in 27 the early Earth, as well as for constraining entrainment of melts in the mantle and in the 28 present-day core-mantle boundary. Here, we present *in situ* structural measurements by 29 X-ray diffraction of selected amorphous silicates compressed statically in diamond anvil 30 cells (up to 157 GPa at room temperature) or dynamically by laser-generated shock 31 compression (up to 130 GPa and 6000 K along the MgSiO<sub>3</sub> glass Hugoniot). The X-ray 32 diffraction patterns of silicate glasses and liquids reveal similar characteristics over a 33 wide pressure and temperature range. Beyond the increase in Si coordination observed 34 at 20 GPa, we find no evidence for major structural changes occurring in the silicate melts studied up to pressures and temperatures exceeding Earth's core mantle 35 36 boundary conditions. This result is supported by molecular dynamics calculations. Our 37 findings reinforce the widely-used assumption that the silicate glasses studies are 38 appropriate structural analogs for understanding the atomic arrangement of silicate 39 liquids at these high pressures.

40

## 41 Significance

42 Understanding the structural changes silicate melts undergo over the pressure-temperature 43 range of the Earth's mantle has been a major, longstanding challenge in the geosciences. 44 Experimental studies are extremely difficult, due to required temperatures exceeding 4000 K 45 needed to melt silicates over megabar pressures. To overcome this issue, novel laser-driven 46 shock experiments combined with X-ray Free-Electron Lasers were performed to provide 47 nanosecond resolution on silicate structural transformations. By comparison with statically 48 compressed diamond-anvil cell experiments at ambient temperature, a common high-pressure 49 structural evolution of glasses and liquid silicates was revealed. This supports the concept 50 that silicate glasses of dominant mantle composition are suitable structural analogues for the 51 corresponding liquids at these pressures.

52

# 53 Main text

# 54 Introduction

55 Terrestrial planet interiors have silicate-based mantles and iron-rich cores. It is 56 generally accepted that this structure is inherited from early differentiation stages occurring in 57 concert with radiogenic decay of short-lived nuclides and multiple shock events inducing 58 temperatures high enough to sustain a global magma ocean. Hence, the physical properties of 59 liquid silicates over a wide pressure-temperature range are crucial for understanding early 60 differentiation events that occurred in Earth and other rocky planets (1). Those properties, 61 especially the density, control the extent over which silicate melts can be stably entrained in 62 the mantle (2). In addition, liquid silicates exist in our present-day upper mantle (3–5) and 63 potentially at the core-mantle boundary (CMB) (6, 7). The exact composition and origin of molten or partially-molten domains in the deep mantle remain enigmatic. Direct 64 measurements of molten mantle material properties to relevant pressures and temperatures are 65 required to constrain these geodynamic processes (8, 9) as well as candidate materials 66 67 contributing to mantle heterogeneities (10).

68 MgSiO<sub>3</sub> is the end member composition for the most abundant mineral of the Earth's 69 mantle, i.e. bridgmanite (11). Liquid MgSiO<sub>3</sub> is arguably the closest proxy for liquid silicates 70 in deep Earth. Its simple chemical composition includes two important features of liquid 71 silicates at ambient pressure, with  $SiO_4$  building blocks and octahedral Mg cation sites. In 72 crystalline MgSiO<sub>3</sub>, the pressure-induced transition to bridgmanite above 23 GPa is 73 accompanied by an increase in coordination, from 4- to 6-fold, of Si sites by O atoms. Such 74 dramatic atomic reorganization may also affect the properties of the corresponding liquid (12, 75 13).

76 The experimental study of liquid MgSiO<sub>3</sub> has been a long-term challenge, as MgSiO<sub>3</sub> 77 melts at extremely high temperatures even at moderate pressures (melting temperature 78 exceeds 4000 K at 40 GPa (14–16)). Therefore, MgSiO<sub>3</sub> glass has been extensively studied as 79 a possible structural analog for the liquid state at high pressures (13, 17–19). Vibrational 80 spectroscopy identified structural changes (13, 19–21) which were confirmed by *in situ* X-ray 81 diffraction (XRD) studies (22, 23). This led to early constraints on the density of mantle melts 82 at CMB conditions (18). However, the validity of the liquid/glass structural analogy has been 83 a longstanding question. In order to probe lattice-level structural changes of liquid silicates, 84 XRD from laser-driven shock compression with nanosecond resolution were collected with 85 femtosecond X-ray pulses from a Free Electron Laser. These results were compared to static 86 compression experiments and Molecular Dynamics simulations.

87

#### 88 **Results**

We performed in situ XRD measurements on three silicate glasses: MgSiO<sub>3</sub>, 89 90 Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub>, and SiO<sub>2</sub>. XRD patterns were collected up to 160 GPa at room temperature under 91 static compression using diamond-anvil cells (DACs) and up to 130 GPa and ~6000 K under 92 dynamic compression. Dynamic compression experiments reached the liquid phase along the 93 principal Hugoniot. Integrated XRD patterns for MgSiO<sub>3</sub> glasses measured under dynamic 94 and static compression, as well as synthetic spectra obtained from molecular dynamics (MD) 95 simulations, are shown in Figure 1. Similar features were observed for these different 96 approaches as well as for previous results from large volume experiments (17) (Figure 1). Amorphous MgSiO<sub>3</sub> displays a first large diffuse feature at 20 nm<sup>-1</sup> (Q1) below a few GPa; 97 98 this is clearly visible in the static compression data at 2.9 GPa, while in the dynamic 99 experiment reference spectrum Q1 is mixed with the signal from the unshocked polystyrene 100 ablator at lower Q-values (Figure 1). The signal from the ablator disappears upon shock 101 compression and the recorded patterns capture contributions from both the shocked and

102 unshocked parts due to experimental geometry. Careful background subtraction allows the 103 separation of these two contributions (SI Appendix, Figure S2). A second diffuse diffraction 104 peak (Q2) arises under compression in both static and dynamic data around Q-values of 30-35 105  $nm^{-1}$  (Figure 1) (17, 24) and marks the transition towards a dense amorphous structure. This spectral evolution is related to the transition from 4- to 6-fold oxygen coordination around the 106 107 Si atoms (17, 24). As the pressure is further increased, both diffuse scattering peaks shift 108 toward higher Q. The relative intensity of the second (Q2) compared to the first peak (Q1) 109 increases, as qualitatively illustrated by intensity plots (Figure 1) and quantitatively confirmed 110 by the structure factors S(Q) (SI Appendix, Figure S3).

111 Shock-compressed MgSiO<sub>3</sub> glass does not show evidence for recrystallization at the 112 nanosecond timescale up to 129( $\pm$ 35) GPa and 6000 K, well above the melting along the 113 Hugoniot (SI Appendix, Figure S5). This observation highlights a fundamental difference 114 with fused silica SiO<sub>2</sub> which is found to crystallize into stishovite in a few- to 100s of 115 nanoseconds (25, 26) under shock compression. The absence of recrystallization of MgSiO<sub>3</sub> 116 glass along the Hugoniot may be related to the presence of Mg affecting the degree of 117 polymerization, thus leading to an increased stability of the glass structure at nanosecond 118 timescales. The dense disordered structures obtained in both static and dynamic compressions 119 can therefore be compared along the entire pressure range reached in the shock experiment, 120 which includes equilibrium liquid conditions. Indeed, according to previous works, melting of 121 MgSiO<sub>3</sub> glass samples along the principal Hugoniot is expected to occur between 90 and 110 122 GPa (15, 16) (SI Appendix, Figure S5). In our dynamic experiments, we used velocimetry 123 records via a velocity interferometer system for any reflector (VISAR) diagnostic to 124 determine the peak pressure in our shots. Four out of the six datasets presented in Figure 1 can be assigned to pressures below 90 GPa, considering the existing measurements of the 125 126 Hugoniot of liquid MgSiO<sub>3</sub> from shocked MgSiO<sub>3</sub> glass (15, 16, 27–29) and the Hugoniot of 127 MgSiO<sub>3</sub> glass as calculated from density functional theory (DFT) MD simulations (30), and 128 are therefore in a glassy solid state. The two data points at 124 and 129 GPa are above the 129 melting curve and are representative of a high-pressure liquid (SI Appendix, Figure S5). To a 130 first approximation, the data show no significant difference between the diffraction of the 131 glass acquired along the Hugoniot and the corresponding liquid at similar pressure conditions.

The limited Q range available for dynamic compression XRD patterns precludes a more detailed analysis, specifically preventing a real space analysis of the diffraction patterns. This limitation was intrinsic to the experimental setup available (see Methods). The partial structure factors for MgSiO<sub>3</sub> glass highlight the fact that this portion of the total structure

136 factor is mainly related to contributions from cation-anion (Mg-O and Si-O) bonds (22, 31). A 137 further consequence of the limited Q-range intrinsic to the dynamic experiments is the 138 inability to accurately track structural differences between statically and dynamically 139 compressed samples at the atomic level. However, differences in atomic arrangements should 140 be reflected in a change in the structure factor, with the appearance of new diffuse peaks (e.g. 141 between 2.5 and 20 GPa in Figure 1) or by a difference in the O1/O2 ratio. In this regard, we 142 find that the XRD signal obtained from shock compression of MgSiO<sub>3</sub> glasses follows the 143 same Q1/Q2 trend as the one from DAC-compressed glasses (Figure 2). Moreover, another 144 silicate glass with an Mg-enriched composition (Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub>) also follows a similar Q1/Q2 145 trend (Figure 2, SI Appendix, Figure S7). Our experiment also provides evidence that liquid 146  $SiO_2$  has a similar structure because the Q1/Q2 ratio lies on the trend found for statically 147 compressed SiO<sub>2</sub> glass (Figure 2). The difference in absolute position of the two features of 148 amorphous MgSiO3 and SiO2 samples can be attributed to the presence of Mg. Based on this 149 data, it can be stated that silicate glasses at high pressures and room temperature have 150 remarkably similar structures to their corresponding liquids at high temperatures. In addition, 151 no modification of the amorphous structure with increasing SiO2 content (from 50% 152  $(MgSiO_3)$  to 60%  $(Mg_3Si_2O_7)$ ) was observed, which agrees with reported results of glasses at 153 ambient pressure (32).

154 The pressure dependence of the first sharp diffraction peak (FSDP) Q1 is often used as 155 a proxy for pressure induced changes in coordination number (33-36). For example, 156 correlation between pressure increase in Q1 and evolution of oxygen-packing fraction is a 157 common feature for a large variety of glasses (33). Figure 3 illustrates the Q1 dependence on 158 pressure for our data, without evidence for sharp discontinuities in both the liquids and 159 glasses, indicating monotonic structural changes (24) with increasing pressure up to the liquid 160 state. In particular, the MD simulations (see Methods for details) reproduce the correlation 161 between Q1 and Q2 values found in our experiment, for both intensity distribution (Figure 1) 162 and quantitative Q1/Q2 ratios (Figure 2). The simulations show a correlation between Q1 and 163 Q2 values very similar to the experimental one, and the coordination number of O around Si 164 smoothly increases from 6 to 7 (between 6.1 and 6.7; SI Appendix, Table S8).

165 It is challenging to discriminate between temperature effects and structural 166 rearrangements, when comparing the pressure dependent peak positions of Q1 and Q2 under 167 static compression at room temperature and dynamic compression along the Hugoniot. 168 Therefore, peak positions are also plotted as a function of density (Figure 3 and 4). For static 169 compression, density values were extracted from recent experimental studies of amorphous

170 SiO2 and MgSiO3 (18, 37). For dynamic compression, they were deduced from VISAR 171 measurements and a known equation of state (see Methods). The error propagation involved 172 in this analysis results in large uncertainties (up to 30%) for the density of the shocked state. 173 The plots clearly show no significant difference in peak position between statically 174 compressed glass and high temperature disordered silicates with similar densities, suggesting 175 that no structural changes occur over the entire examined density range. This analysis can be 176 extended to the second diffraction peak Q2, which appears above 10 GPa (Figure 4). 177 Theoretical calculations (31), indicate that Q2 is related to the evolution of cation-anion bonds 178 and the appearance of 5- and 6-fold coordinated Si-O and Mg-O polyhedra. The pressure 179 dependence of the Q2 position is similar for dynamic and static compression. Moreover, the 180 SiO<sub>2</sub> and MgSiO<sub>3</sub> samples show comparable Q2 features as a function of pressure, suggesting 181 similar packing mechanisms.

182

# 183 Conclusions

184 Dynamic compression experiments allowed us to probe the structure of liquid silicates 185 with compositions relevant to the Earth's lower mantle (over 130 GPa and 6000 K). Although 186 the limited Q range does not allow for a full structural description, no evidence for major 187 structural changes between liquid and cold compressed silicate glass was found. In particular, 188 the ratio of Q1/Q2 measured for cold compressed MgSiO<sub>3</sub> glass at 90 GPa and room 189 temperature is very similar to that of liquid MgSiO<sub>3</sub> along the Hugoniot at 130 GPa and 6000 190 K. Similar conclusions can be drown for liquid and statically compressed SiO<sub>2</sub>. Therefore, 191 XRD measurements of cold compressed silicate glasses can be used to infer structural 192 properties of liquid silicates at Earth's interior conditions. Our experiments further 193 demonstrate the possibility of high-quality XRD diffraction measurements of liquid silicates 194 to previously inaccessible pressure-temperature regimes. In addition, signatures of chemical 195 reactions between the shocked liquids and the target chamber are minimized in the in situ 196 XRD spectra due to the nanosecond timescale of the shock compression experiment (or the 197 femtosecond probing timescale of the XFEL) (38).

The absence of major structural change under pressure up to  $\sim$ 130 GPa in liquid MgSiO<sub>3</sub> is expected to be important for modeling partial melts at depth in the Earth. Similar mechanisms are likely to occur in rocky exoplanets with Earth-like compositions (39). Finally, our results indicate that combined static and dynamic datasets could be used to infer silicate liquid properties. Future dynamic compression experiments combined with XFEL sources are complementary tools suitable for probing liquid silicates under extremeconditions.

- 205
- 206 Methods

# 207 Target preparation and characterization

208 Targets for shock compression consisted of a 50 µm-thick polystyrene ablator glued to 209 a silicate sample (between 35 to 90 µm thickness) that was polished on both sides, and coated 210 with 300 nm Ti on the drive laser side prior to gluing (SI Appendix, Figure S9). Glass starting 211 materials with MgSiO<sub>3</sub> and Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub> compositions were synthetized by laser levitation in an 212 inert gas flux of argon. Fused silica starting material was manufactured by Asphera company. 213 The thickness of each sample as well as its planarity were individually measured using a 2D 214 interferometer. This characterization has been used for a relevant estimate of the mean shock 215 velocity and hence of the induced pressure.

216

# 217 Static compression

218 MgSiO<sub>3</sub> glass was synthetized by laser levitation of MgSiO<sub>3</sub> powder in an inert gas 219 flow of Ar. Flakes of the glass were placed in a pre-indented and drilled Re gasket. For the 220 lower pressure (LP) run, a gas driven membrane Le Toullec type DAC was used to compress 221 the sample up to 100 GPa. A small piece of KCl was placed on the side of the sample in order 222 to calculate pressure from the equation of state (EoS) (40). The experimental setup available 223 on the High Pressure beamline ID27 (ESRF, Grenoble, France) (41) included KB mirrors to focus the X-ray beam to  $3x3 \mu m^2$ , a MAR CCD detector to collect the diffraction pattern, and 224 225 a multichannel collimator slit system to reduce the inelastic X-ray contribution from the 226 diamond anvils (42, 43). High-quality data were obtained with this set-up, enabling the 227 extraction of a high Q range structure factor (SI Appendix, Figure S3).

228 The background contribution from the diamond anvils was first measured by placing 229 an empty DAC in the X-ray beam. Then the background was subtracted from the sample 230 diffraction patterns. The diffraction was normalized to obtain the structure factor S(Q)231 following the previously described method (44, 45). For the high pressure (HP) run up to 160 232 GPa, a different experimental setup was used (SI Appendix, Figure S4). The background from 233 the diamonds was collected on the empty cell and removed from the raw signal. The Re-234 gasket EoS was used to determine the pressure, which is well correlated with the EoS of KCl 235 used in the LP run (46).

236

#### 237 **Dynamic compression**

High pressure and temperature conditions in the MgSiO<sub>3</sub> glass samples were obtained via laser-driven shock compression using the two frequency-doubled Nd-glass laser beams available at the Matter in Extreme Conditions (MEC) end-station of the Linac Coherent Light Source (LCLS) at SLAC National Acceleratory Laboratory. Various thermodynamic conditions could be generated by choosing the laser energy for a 10ns square pulse on a 250um focal spot, smoothed using a continuous phase plate.

244 In situ XRD measurements were performed using quasi-monochromatic (dE/E =0.2–0.5%), 9.0 keV X-ray pulses of 60 fs duration with an average of ~  $10^{12}$  photons per 245 246 pulse. The LCLS X-ray free electron laser spot size was set to 50 µm diameter, much smaller 247 than the flat compressed region so as to avoid detecting pressure and temperature gradients. 248 The measurement of uniform thermodynamic conditions during shock propagation were 249 ensured by optimizing the target design and probe timing. Example patterns collected on the 250 CSPADs of the typical diffuse scattering of amorphous silicates under high pressures and 251 temperatures are shown in SI Appendix, Figure S8. The accessible energy and position of the 252 CSPAD were limited at the MEC end-station at the time of the experiment.

253 The time-resolved velocity of the reflecting interfaces in the sample were measured by 254 two Doppler interferometers (VISARs) operating at 532 nm whose VPF were 4.658 km/s and 255 1.814 km/s. For most of the targets, time-resolved particle velocity inside the sample could be 256 measured from the VISAR fringe shift if the sample remained transparent after shock loading. 257 The mean shock velocity inside the sample was instead measured from the shock arrival and 258 exit times recorded on the VISAR images. For some of the targets, the time-resolved material 259 velocity inside the rear-side lithium fluoride window was also measured to investigate the 260 shock stationarity.

261 The shock pressure was deduced by inserting the measured shock and particle 262 velocities into the Rankine-Hugoniot equations, in case they were recorded. For the shots 263 where only the shock velocity could be measured, the known principal Hugoniot curves 264 (fused silica (SESAME 7386), MgSiO<sub>3</sub> glass (27, 29)) were used to estimate the pressure. An 265 independent pressure estimate could be obtained by applying the impedance matching 266 technique between the sample and the lithium fluoride window. SESAME 7270 was used 267 extract the lithium fluoride Hugoniot curve. The uncertainties on pressure estimates were 268 determined via a Monte-Carlo routine which repeated the analysis 5000 times. Each analysis 269 was performed extracting the inputs from a Gaussian distribution whose mean were the

measured value and the standard deviation were the estimated uncertainty. The standard
deviation of the output distribution has been taken as the uncertainty in pressure. An extensive
discussion on the VISAR analysis is presented in the Supplementary information.

273

### 274 Molecular dynamics simulations

275 MD simulations were performed using the LAMMPS code (47). Pair potentials 276 constructed in Oganov et al. 2000 (48) were used to model the interactions between Mg, Si 277 and O atoms. The simulations contained a total of 112000 atoms and were performed in the 278 isobaric-isothermal (NPT) ensemble with a time step of 1 fs. The first melt structure was 279 created by heating an orthoenstatite structure at ambient pressure up to 4000 K, and 280 equilibrating the resulting melt during 100 ps. Then, we compressed and heated the system to 281 the target pressure-temperature conditions in 100 ps (SI Appendix, Table S6). The cells were 282 then equilibrated during an additional 100 ps at the considered high pressure-temperature 283 conditions. XRD intensities were then calculated using the USER-DIFFRACTION package 284 implemented in LAMMPS (49).

285

#### 286 Data availability

The different data presented in this manuscript are available in the SupplementaryInformation.

289

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304			
305	References		
306	1.	Elkins-Tanton LT (2012) Magma Oceans in the Inner Solar System. Annu Rev Earth	
307		<i>Planet Sci</i> 40(1):113–139.	
308	2.	Labrosse S, Hernlund JW, Coltice N (2007) A crystallizing dense magma ocean at the	
309		base of the Earth's mantle. Nature 450(7171):866-869.	
310	3.	Sanloup C (2016) Density of magmas at depth. Chem Geol 429:51-59.	
311	4.	Schmandt B, Jacobsen SD, Becker TW, Liu Z, Dueker KG (2014) Dehydration melting	
312		at the top of the lower mantle. Science (80-) 344(6189):1265–1268.	
313	5.	Tauzin B, Debayle E, Wittlinger G (2010) Seismic evidence for a global low-velocity	
314		layer within the Earth's upper mantle. Nat Geosci 3(10):718-721.	
315	6.	Lay T, Williams Q, Garnero EJ (1998) The core-mantle boundary layer and deep Earth	
316		dynamics. Nature 392(2):461-468.	
317	7.	Fiquet G, et al. (2010) Melting of peridotite to 140 gigapascals. Science	
318		329(5998):1516–1518.	
319	8.	Wen L, Helmberger D V (1998) Ultra-Low Velocity Zones Near the Core-Mantle	
320		Boundary from Broadband PKP Precursors. Science 279(March):1701–1704.	
321	9.	Garnero EJ, McNamara AK, Shim S-H (2016) Continent-sized anomalous zones with	
322		low seismic velocity at the base of Earth's mantle. Nat Geosci 9(June):481.	
323	10.	Trønnes RG, et al. (2019) Tectonophysics Core formation, mantle differentiation and	
324		core-mantle interaction within Earth and the terrestrial planets. Tectonophysics	
325		(October 2018):1–34.	
326	11.	Tschauner O, Beckett JR, Prescher C, Prakapenka VB, Rossman GR (2014) Discovery	
327		of bridgmanite, the most abundant mineral in Earth, in a shocked meteorite. Science	
328		346(6213):1100–1103.	
329	12.	Stixrude L, Karki B (2005) Structure and Freezing of MgSiO 3 Liquid in Earth 's	
330		Lower Mantle. Science 185(October):297–300.	
331	13.	Murakami M, Bass JD (2011) Evidence of denser MgSiO3 glass above 133 gigapascal	
332		(GPa) and implications for remnants of ultradense silicate melt from a deep magma	
333		ocean. Proc Natl Acad Sci 108(42):17286–17289.	
334	14.	Shen G, Lazor P (1995) Measurement of melting temperatures of some minerals under	
335		lower mantle pressures. J Geophys Res 100(B9):17699–17713.	
336	15.	Akins JA, Luo SN, Asimow PD, Ahrens TJ (2004) Shock-induced melting of MgSiO3	
337		perovskite and implications for melts in Earth's lowermost mantle. Geophys Res Lett	

338 31(14):2–5.

- 339 16. Mosenfelder JL, Asimow PD, Frost DJ, Rubie DC, Ahrens TJ (2009) The
- 340 MgSiO3system at high pressure: Thermodynamic properties of perovskite,
- 341 postperovskite, and melt from global inversion of shock and static compression data. J

342 *Geophys Res Solid Earth* 114(1):1–16.

- Kono Y, Shibazaki Y, Kenney-Benson C, Wang Y, Shen G (2018) Pressure-induced
  structural change in MgSiO 3 glass at pressures near the Earth's core–mantle boundary. *Proc Natl Acad Sci*:201716748.
- 346 18. Petitgirard S, et al. (2015) Fate of MgSiO <sub>3</sub> melts at core–mantle boundary conditions.
  347 *Proc Natl Acad Sci* 112(46):14186–14190.
- 348 19. Sanchez-Valle C, Bass JD (2010) Elasticity and pressure-induced structural changes in
  349 vitreous MgSiO3-enstatite to lower mantle pressures. *Earth Planet Sci Lett* 295(3–
  350 4):523–530.
- 351 20. Shim S, Catalli K (2009) Compositional dependence of structural transition pressures
  352 in amorphous phases with mantle-related compositions. *Earth Planet Sci Lett* 283(1–
  353 4):174–180.
- Kubicki JD, Hemley RJ, Hofmeister AM (1992) Raman and infrared study of pressureinduced structural changes in MgSiO3 CaMgSi2O6, and CaSiO3 glasses. *Am Mineral*77:258–269.
- Wilding MC, Benmore CJ, Tangeman JA, Sampath S (2004) Coordination changes in
   magnesium silicate glasses. *Europhys Lett* 67(2):212–218.
- 23. Cormier L, Cuello GJ (2013) Structural investigation of glasses along the MgSiO3CaSiO3join: Diffraction studies. *Geochim Cosmochim Acta* 122(December):498–510.
- 361 24. Prescher C, et al. (2017) Beyond sixfold coordinated Si in SiO <sub>2</sub> glass at ultrahigh
  362 pressures. *Proc Natl Acad Sci*:201708882.
- 363 25. Gleason AE, et al. (2017) Time-resolved diffraction of shock-released SiO2and
  364 diaplectic glass formation. *Nat Commun* 8(1):8–13.
- 365 26. Gleason AE, et al. (2015) Ultrafast visualization of crystallization and grain growth in
  366 shock-compressed SiO2. *Nat Commun* 6:1–6.
- 27. Luo SN, Akins JA, Ahrens TJ, Asimow PD (2004) Shock-compressed MgSiO3 glass,
  enstatite, olivine, and quartz: Optical emission, temperatures, and melting. *J Geophys Res Solid Earth* 109(5):1–14.
- Spaulding DK, et al. (2012) Evidence for a phase transition in silicate melt at extreme
  pressure and temperature conditions. *Phys Rev Lett* 108(6):1–4.

- 372 29. Bolis RM, et al. (2016) Decaying shock studies of phase transitions in MgO-SiO2
  373 systems: Implications for the super-Earths' interiors. *Geophys Res Lett* 43(18):9475–
  374 9483.
- 375 30. de Koker N, Stixrude L (2009) Self-consistent thermodynamic description of silicate
  376 liquids, with application to shock melting of MgO periclase and MgSiO3 perovskite.
  377 *Geophys J Int* 178(1):162–179.
- 378 31. Ghosh DB, Karki BB, Stixrude L (2014) First-principles molecular dynamics
  379 simulations of MgSiO 3 glass : Structure , density , and elasticity at high pressure. *Am*380 *Mineral* 99:1304–1314.
- 381 32. Wilding MC, Benmore CJ, Tangeman JA, Sampath S (2004) Evidence of different
  382 structures in magnesium silicate liquids: Coordination changes in forsterite- to
  383 enstatite-composition glasses. *Chem Geol* 213(1–3):281–291.
- 384 33. Zeidler A, Salmon PS (2016) Pressure-driven transformation of the ordering in
  amorphous network-forming materials. *Phys Rev B* 93(21):1–5.
- 386 34. Meade C, Hemley RJ, Mao HK (1992) High-Pressure X-Ray Diffraction of Si02 Glass.
  387 *Phys Rev Lett* 69(9):1387–1391.
- 388 35. Benmore CJ, et al. (2011) High pressure x-ray diffraction measurements on Mg2SiO 4
  389 glass. *J Non Cryst Solids* 357(14):2632–2636.
- 36. Sato T, Funamori N (2008) Sixfold-Coordinated Amorphous Polymorph of SiO2 under
  High Pressure. *Phys Rev Lett* 101(25):255502.
- 392 37. Petitgirard S, et al. (2017) SiO2 Glass Density to Lower-Mantle Pressures. *Phys Rev*393 *Lett* 119(November):215701.
- 394 38. Morard G, et al. (2018) Solving controversies on the iron phase diagram under high
  395 pressure. *Geophys Res Lett* 45:1–9.
- 396 39. Stixrude L (2014) Melting in super-earths. *Philos Trans R Soc A*
- 397 372(March):20130076.
- 398 40. Dewaele A, et al. (2012) High-pressure-high-temperature equation of state of KCl and
  399 KBr. *Phys Rev B Condens Matter Mater Phys* 85(21):1–7.
- 400 41. Mezouar M, et al. (2005) Development of a new state-of-the-art beamline optimized
- 401 for monochromatic single-crystal and powder X-ray diffraction under extreme
- 402 conditions at the ESRF. *J Synchrotron Radiat* 12(5):659–664.
- 403 42. Morard G, et al. (2011) High efficiency multichannel collimator for structural studies
- 404 of liquids and low-Z materials at high pressures and temperatures. *Rev Sci Instrum*405 82(2):2–7.

- 406 43. Weck G, et al. (2013) Use of a multichannel collimator for structural investigation of
  407 low-Z dense liquids in a diamond anvil cell: Validation on fluid H2 up to 5 GPa. *Rev*408 *Sci Instrum* 84(6).
- 409 44. Eggert J, Weck G, Loubeyre P, Mezouar M (2002) Quantitative structure factor and
  410 density measurements of high-pressure fluids in diamond anvil cells by x-ray
- 411 diffraction: Argon and water. *Phys Rev B* 65(17):1–12.
- 412 45. Morard G, et al. (2014) Density measurements and structural properties of liquid and
  413 amorphous metals under high pressure. *High Press Res* 34(1).
- 414 46. Anzellini S, Dewaele A, Occelli F, Loubeyre P, Mezouar M (2014) Equation of state of
  415 rhenium and application for ultra high pressure calibration. *J Appl Phys* 115(4).
- 416 47. Plimpton S (1995) Fast parallel algorithms for short range molecular dynamics. J
  417 *Comput Phys* 117:1–19.
- 418 48. Oganov AR, Brodholt JP, Price GD (2000) Comparative study of quasiharmonic lattice
  419 dynamics , molecular dynamics and Debye model applied to MgSiO3 perovskite. *Phys*420 *Earth Planet Int* 122:277–288.
- 421 49. Coleman SP, Spearot DE, Capolungo L (2013) Virtual diffraction analysis of Ni [ 0 1 0
  422 ] symmetric tilt. *Model Simul Mater Sci Eng*:055020.
- 423 50. Militzer B (2013) Ab initio investigation of a possible liquid-liquid phase transition in
  424 MgSiO3 at megabar pressures. *High Energy Density Phys* 9(1):152–157.
- 425 51. Fratanduono DE, et al. (2018) Thermodynamic properties of MgSiO3 at super-Earth
  426 mantle conditions. *Phys Rev B* 97(21):214105.
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428

- 429 Figure Captions
- 430

431 Figure 1: Diffraction patterns of compressed amorphous MgSiO<sub>3</sub> obtained under static 432 and dynamic compression, and from molecular dynamics simulations. Top left: XRD 433 patterns for MgSiO<sub>3</sub> glass collected during static compression in a DAC. Full S(Q), up to 80 434 nm<sup>-1</sup>, are presented in SI Appendix, Figure S3. Diffraction patterns for MgSiO<sub>3</sub> up to 160 435 GPa are presented in SI Appendix, Figure S4. Top right: XRD patterns collected for MgSiO<sub>3</sub> 436 amorphous targets compressed along the Hugoniot during dynamic compression experiments. 437 The reported pressures, hence the liquid or glassy solid state of the sample, are determined from VISAR measurements (see Supplementary Materials). Diffraction peaks around 24 and 438 439 28 nm<sup>-1</sup> come from the shocked Ti layer located between the ablator and the sample (see SI 440 Appendix, Figure S1). The signal from the cold part of the sample has been removed 441 following the procedure presented in Figure S2. There is a clear evolution toward higher Q-442 values for the two main sample contributions (Q1 and Q2). Bottom left: XRD patterns 443 calculated from MD simulations. Details for the calculations are given in the Methods section. 444 Different pressure-temperature conditions are presented in SI Appendix, Figure S5, 445 corresponding to two Hugoniots representing extreme cases (50, 51). Bottom right: Comparison of different XRD patterns at similar density, 4.5 g.cm<sup>-3</sup>. The XRD patterns 446 447 collected in a previous large volume experiment have been recalculated from the dataset 448 presented (17). The dashed vertical lines serve as guides for the eyes.

449

450 Figure 2: Correlation between First (Q1) and Second (Q2) Diffraction peaks for 451 MgSiO<sub>3</sub>, Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub>, and SiO<sub>2</sub> glasses under static and dynamic compression (SI 452 Appendix, Tables S1-S6). Round symbols represent data acquired during static compression 453 using DACs (MgSiO<sub>3</sub> glass: with multichannel collimator (MCC) experimental set-up (43)454 (black circles) and without (red circles); SiO<sub>2</sub> glass: data from literature (24, 34)). Square 455 symbols represent diffraction experiments recorded during dynamic compression (starting 456 with SiO<sub>2</sub> glass, MgSiO<sub>3</sub> glass, Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub> glass, and MgSiO<sub>3</sub> enstatite crystals). In addition, 457 results from our MD simulations of MgSiO<sub>3</sub> glass are shown as filled blue diamonds.

458

Figure 3: Position of the first sharp diffraction peak (FSDP) Q1 for MgSiO<sub>3</sub> and SiO<sub>2</sub>
compressed amorphous samples as a function of density (SI Appendix, Tables S1-S6).
Density of cold compressed glasses is calculated following recent experimental measurements
(18, 37). Density of shocked samples is calculated following VISAR measurements (see

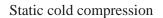
- Methods). Round, square and diamond symbols are similar to Figure 2. Inset: peak position as
  a function of pressure. Pressures for shock compressed samples are calculated from VISAR
  measurements (see Methods).
- 466

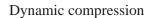
# 467 Figure 4: Position of the second diffraction peak (Q2) for MgSiO<sub>3</sub> and SiO<sub>2</sub> compressed

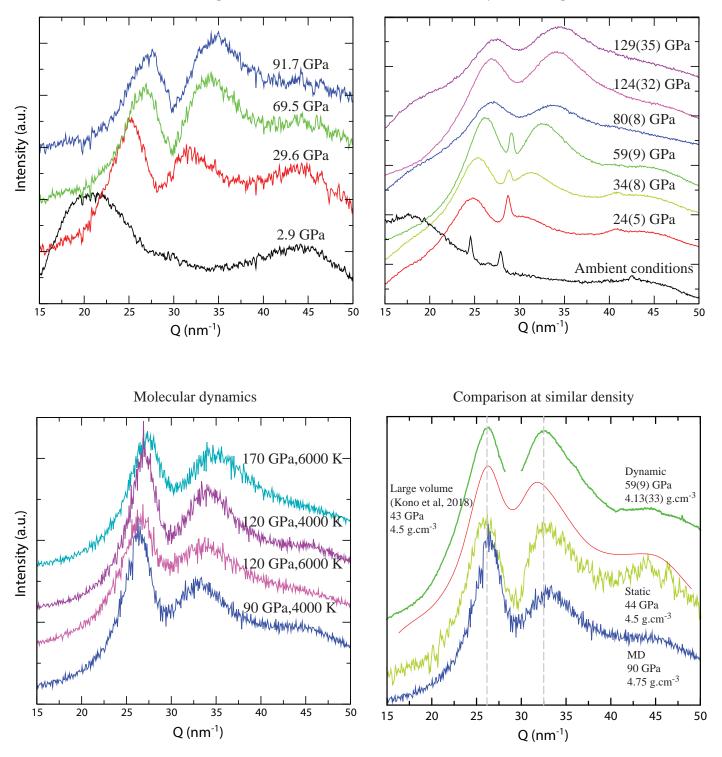
468 **amorphous samples as a function of density (SI Appendix, Tables S1-S6).** This second 469 peak appears in the diffraction pattern above 10 GPa. Symbols used in this figure are similar

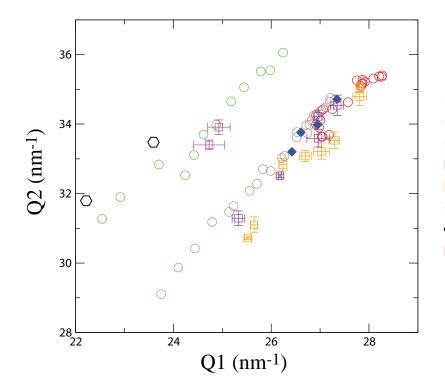
470 to those in Figure 3. Inset: peak position as a function of pressure, following the method used

- 471 in Figure 3.
- 472









- MgSiO<sub>3</sub> glass, molecular dynamics
- O MgSiO<sub>3</sub> glass, static DAC (Run HP)
- MgSiO<sub>3</sub> glass, static DAC (Run LP)
- □ MgSiO<sub>3</sub> glass, dynamic compression
- Mg<sub>3</sub>Si<sub>2</sub>O<sub>7</sub> glass, dynamic compression
- $\bigcirc$  SiO<sub>2</sub> glass, static compression (Prescher et al, 2017)
- $\bigcirc$  SiO<sub>2</sub> glass, static compression (Meade et al, 1992)
- $\Box$  SiO<sub>2</sub> glass, dynamic compression

