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▶ To cite this version:

Loïc M Martin, Caroline Simonucci, Sétareh Rad, Marc F. Benedetti. Effect of natural organic matter on thallium and silver speciation. Journal of Environmental Sciences, 2020, 93, pp.185-192. 10.1016/j.jes.2020.04.001 . hal-02565435

HAL Id: hal-02565435 https://hal.science/hal-02565435

Submitted on 6 May 2020

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Effect of natural organic matter on thallium and silver speciation

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Donnan membrane detection



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35	Abstract: Natural organic matter (NOM) is known to play an important role in the transport and
36	binding of trace metal elements in aquatic and soil systems. Thallium is a pollutant for which the
37	extent of the role played by NOM is poorly known. Consequently, this study investigates thallium(I)
38	and its complexation to a purified humic substance as proxy for NOM. Experiments were performed
39	with the Donnan Membrane Technique to separate, for the first time, the free Tl ⁺ ion from its
40	complexed form in the bulk solution. Various pH and concentrations were investigated at constant
41	ionic strength and constant NOM proxy concentrations in solution. Experimental results were
42	described with NICA-Donnan model. Thallium complexation was compared to silver complexation
43	using literature data and using the same NICA-Donnan formalism. Parameters for these two cations
44	$(Tl^+ and Ag^+)$ are reported in this article, for the first time. Results display low thallium complexation
45	to the NOM proxy while silver competes with divalent cations for the NOM binding sites. Calculated
46	speciation for dissolved thallium highlights the dominance of free thallium (Tl ⁺) in solution whereas
47	Tl-NOM complexes contribute roughly 15 % to total Tl(I) species in river and lake type waters.
48	Similar results are obtained for soil solutions, Tl-bound to NOM < 30 % of total, from UK soils with
49	different land use and geochemistry.

- 50 Keywords:
- 51 Thallium
- 52 Organic matter
- 53 Speciation
- 54 NICA-Donnan model
- 55 Soil

- 56 Water
- 57
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- 60

61 Introduction

62 Thallium (Tl) is a trace element and a contaminant, mainly byproducts of extractives industries 63 and comes from ferrous and non-ferrous metal and coal mining (Cheam, 2001; Lis et al., 2003; Peter and Viraraghavan, 2005; Casiot, et al., 2011; Campanella et al., 2017). Thallium occurs 64 in two oxidation states, Tl(I) and Tl(III) (Jacobson et al., 2005a; Peter and Viraraghavan, 2005). 65 Tl(I) exhibits either lithophile or chalcophile behavior that is determined by the geological 66 context (Prytulak et al., 2017). In the critical zone, Tl(I) is relatively soluble, mobile and 67 bioavailable (Jacobson et al., 2005a; Coup and Swedlund, 2015; Voegelin et al., 2015). Studies 68 69 demonstrated that Tl(I) is the most dominant and thermodynamically stable form in aquatic 70 systems and soils (Vink, 1993; Xiong, 2009; Casiot, et al., 2011). However, in some environments Tl(III) is also found in variable amounts (Lin and Nriagu, 1999; Peacock and 71 Moon, 2012; Voegelin et al., 2015; Campanella et al., 2017). Tl(III) should be found mostly in 72 highly oxidative environments (Vink, 1993; Watanabe and Takahashi, 2015) but photochemical 73 reactions in surface waters (Karlsson et al., 2006) or microbiological processes (Twining et al., 74 2003) could oxidize Tl(I) into Tl(III) under moderately oxidizing conditions. Stabilization by 75 humic substances was also witnessed at low pH and high redox potential values (i.e. pH < 4 Eh >76 77 0.82 V in Watanabe and Takahashi, (2015)). Recent studies also highlighted the possible role of Fe(III) or As(V) in the formation of Tl(III) compounds in acid mining drainage (Karlsson et 78 79 al., 2006; Campanella et al., 2018).

Processes driving Tl chemical speciation and mobility are well documented but some questions 80 remain open (Twining et al., 2003; Karlsson et al., 2006; Coup and Swedlund, 2015; Voegelin 81 82 et al., 2015; Casiot et al., 2017; Campanella et al., 2017; Campanella et al., 2018). Its interaction 83 with natural organic matter (NOM) is one of them (Jacobson et al., 2005a; Liu et al., 2011). It 84 is known that NOM plays an important part in the fate of metal ions and can control ion mobility, 85 concentration in soils and natural waters (Benedetti et al., 1995; Benedetti et al., 1996a). NOM is a complex mixture of compounds but the most studied and the most reactive fractions are the 86 87 humic substances (HS) consisting of humic (HA) and fulvic (FA) acids (Benedetti et al., 1996a; Koopal et al., 2005). Kaplan and Mattigod (1998) defined two stability constants for Tl(I) 88 89 binding onto FA with $\log K_{TI-Full} = 4.83$ and $\log K_{TI-Ful2} = 3.32$ (no units specified) to predict the

90 impact of NOM on thallium speciation. According to these calculations up to 67% of Tl(I) is bound to organic compounds in bog water despite its very low pH that should not favor Tl(I) 91 binding to NOM (Kaplan and Mattigod, 1998). Thus, NOM could play a non-negligible part in 92 Tl(I) speciation. However, Jacobson et al., (2005a) reported poor Tl(I) complexation within a 93 94 peat soil. Liu et al. (2011) also studied Tl(I) binding using two purified humic acids and their 95 results also suggested weak Tl(I) interactions with these NOM proxies. However, these studies used very high Tl(I) concentrations, which are mostly relevant for highly polluted sites (i.e. 96 97 issued from mining activities (Campanella et al., 2017; Casiot et al., 2017)). Investigations made 98 at lower concentration of Tl(I) will allow a better understanding of NOM role on Tl(I) speciation in the critical zone. It is also assumed that Tl(I) chalcophile behavior accounts for Tl(I) 99 100 complexation to NOM through interactions with the S-rich functional groups of the NOM, such 101 as thiols (Jacobson et al., 2005a, 2005b). However, Tl(I) affinity to sulfur-containing organic 102 compounds, such as cysteine for instance, is still under investigation while no data were found regarding thiol (Garcia Bugarin et al., 1989; Gharib and Shamel, 2009; Rodrigez Mercado et 103 104 al., 2013)

The role of S-rich compounds, such as thiol groups, is often raised when metal complexation 105 106 to humic substances (HS) is studied (Skylberg et al., 2006; Aiken et al., 2011) and there are 107 evidences of S-binding for Hg(II) or As(III) (Haitzer et al., 2002; Skylberg et al., 2006; Aiken 108 et al., 2011; Catrouillet et al., 2015). These divalent or trivalent cations cannot really be 109 compared to Tl(I). Whereas silver, as Ag(I), can (Jacobson et al., 2005a, 2005b), both Ag(I) and Tl(I) are soft, non-essential monovalent cations, (Smith et al., 2002; Pearson, 1968) with 110 both chalcophilic and lithophilic behaviors (Pearson, 1968; Jacobson et al., 2005a, 2005b; 111 112 Aiken et al., 2011). Moreover, Ag(I) has high affinity for thiolate (Smith et al., 2002; Kleja et al., 2016) and Ag-S binding in NOM compounds was already mentioned in a previous study 113 (Kleja et al., 2016). Another remaining question is the abundance of S-rich functional group, 114 especially thiol group, and how to quantify it in NOM (Valravamurthyet al., 1997; Aiken et al., 115 2011; Manceau and Nagy, 2012; Rao et al., 2014). 116

In this study, the main objective is to evaluate the importance and strength of Tl complexation by a NOM proxy. In order to do that, complexation of Tl(I) with a purified humic acid (HA) is investigated at multiple pH and concentrations. Experimentations are conducted with Donnan Membrane Technique (Temminghoff et al., 2000) to determine the amount of ion bound to the NOM proxy. These observations are used to calibrate the NICADonnan model to predict Tl(I) behavior during complexation. In addition, Tl(I) results are compared to Ag(I) literature data (Sikora and Stevenson, 1988) and the potential role of S-rich functional groups is discussed. Ag(I) specific parameters for NICA-Donnan model are not yet available, thus data of Ag(I) sorption to various HA is described with NICA-Donnan model for the first time in this study (Sikora and Stevenson, 1988; Chen et al., 2012; Kleja et al. 2016). Finally, the role of NOM Tl(I) cycle is also investigated by revisiting existing data from natural waters (Kaplan and Mattigod, 1998) and simulating Tl(I) speciation in soil solutions under different geochemical contexts.

130

131 **1 Materials and methods**

132 **1.1 Experimental approach**

Donnan Membrane Technique (DMT) was first described by Temminghoff et al. (2000). 133 Additional developments were performed to fully constrain the kinetics aspect of the technique 134 (Weng et al., 2005) and to adapt the geometry of the cells (Kalis et al., 2006; Kalis et al., 2007). 135 136 In this study, the DMT device, used for the experiments, is similar to the one used in Kalis et al. (2006, 2007) and its design shown in Appendix A (Fig. S1). Prior to experiments with humic 137 138 acids, the time to reach Donnan equilibrium for Tl(I) was evaluated. Experiments with only free Tl(I) ions in a 50 mmol/L KNO₃ background electrolyte were performed and showed that 139 140 equilibrium was reached after 24 hr (Appendix A, Fig. S3).

141

142 **1.2 Reagents and chemicals**

In order to have Tl(I) at the beginning of all experiments, the source of thallium is a Tl(I)NO₃ 143 144 salt (99.9%, Sigma Aldrich). Tl(I) should be the only stable thallium species under our experimental conditions (Vink, 1993; Xiong, 2009; Campanella et al., 2017; Casiot et al., 2017). 145 Tl(I) oxidation to Tl(III) happens in highly acidic conditions or on mineral surfaces which 146 stabilize Tl(III), either on vacant sites or as oxides like Tl₂O₃ (Watanabe and Takahashi, 2015; 147 Campanella et al., 2017; Prytulack et al., 2017). Consequently, free Tl(III) should not be present 148 149 in our solutions because of the chosen conditions i.e. pH > 4 and moderate Eh. Background electrolyte solutions were prepared with KNO3 (VWR Chemical) and Ca(NO3)2 (VWR 150 Chemical) dissolved in ultrapure water (milli Q, Millipore). pH was adjusted with distilled 0.1 151 152 mol/L HNO₃ and 0.5 mol/L NaOH (VWR Chemical). A peat purified humic acid from the Mogi river region of Ribeirão Preto, São Paulo State, Brazil was used in all experiments. Detailed of 153 purification procedures and chemical compositions are described in Botero et al. (2017). 154 Thalium calibration standards were diluted from a thallium ICP standard solution (TraceCERT[®], 155 Sigma Aldrich). Chemical blanks were analyzed with HR-ICP-MS Element 2 (Thermo 156 157 Scientific) and no thallium was detected.

159 **1.3 Tl isotherms.**

160 Experiments were performed at room temperature $(23 \pm 1^{\circ}C)$ open to atmosphere with a unique concentration of soluble HA. Previous studies reported low Tl(I) complexation to humic acid 161 162 (Jacobson et al., 2005; Liu et al., 2011). Then, 150 mg/L was used during experiments (pH and concentration isotherms) to maximize Tl complexation to humic acid and to derive NICA-163 Donnan parameters for Tl⁺. The effects of Tl(I) concentration and pH were tested. For sorption 164 edge experiments pH ranged from 3.98 to 7.89 ± 0.05 and the total Tl(I) initial concentration 165 166 was 200 pmol/L. In the case of concentration isotherms, the pH was fixed at 7.35 ± 0.05 and initial concentrations ranged from 20 pmol/L to 2.0 µmol/L. The use of this range of 167 168 concentrations was reported for Tl in the literature for freshwater (pristine and polluted). 20 169 pmol/L correspond to Tl concentration in most of the world big rivers (Kaplan and Mattigod, 170 1998; Nielsen et al., 2005) and 2.0 µmol/L is close to the highest Tl values measured in polluted rivers in mining areas (Casiot et al., 2011; Campanella et al., 2017). The background electrolyte 171 172 in the donor and acceptor sides was KNO₃ 50 mmol/L. DMT were run in duplicates for each experimental data point. 173

174 Prior to isotherm experiments, all cells, bottles and beakers used for the donor solution were 175 washed with Suprapur HNO₃ (Merck) 2% and ultrapure water (milliQ). The cation exchange 176 membrane (551652U, VWR Chemicals) was prepared by successive washing as in previous studies (Temminghoff et al., 2000; Weng et al., 2005; Kalis et al., 2006; Chen et al., 2012). At 177 each pH or concentration steps, two new DMT cells were added in the donor solution whereas 178 179 the two at equilibrium with the solution conditions were removed from the later and both 180 acceptor and donor solutions were sampled. The donor solution corresponded to the sample with Tl(I) ions, HA and the background electrolyte while the acceptor solution was the one 181 182 inside the DMT device and contained the electrolyte and also the free Tl(I) ions after the Donnan 183 equilibrium was reached in 24 hr (Appendix A, Fig. S3). Preliminary modeling suggested that only free Tl(I) ions were stable in the acceptor solution (Appendix A, Table S1). Every 184 185 dissolved species in solution were sum up in Appendix A (Table S2). For more details on the experimental design readers are invited to refer to supporting information (Appendix A, Fig. S1 186 and S2). Finally, collected samples were acidified to pH~1 with distilled HNO₃ 2%, diluted 187 when necessary and stored for further analysis. 188

189

190 **1.4 Measurements of cations in solution.**

Thallium was measured with an HR-ICP-MS Element 2 (Thermo scientific, USA) installed in 191 a clean room with a solution of ¹¹⁵In (5 ppb) as internal standard. For each sample, 9 192 measurements were done (3 x 3 runs, Appendix A, Table S3). Then, the average was used as 193 194 concentration value and the standard deviation was used as measurement error. Concentrations 195 were corrected from signal drift with the internal standard. Certified materials TM23.4 and TM24.4 (Environment Canada) validated Tl(I) analysis (Appendix A, Table S4). Limits of 196 quantification (LOQ) ranged from 2.5 pmol/L to 7.5 pmol/L for Tl(I). Major cations (K⁺ and 197 Na⁺) were analyzed with an ICP-AES (ICAP 6200, Thermo scientific, USA). More details of 198 199 HR-ICP-MS measurement for thallium can be found in Appendix A (Table S3).

200

201 **1.5 Modeling.**

This experimental approach was combined with NICA-Donnan modeling within ECOSAT 202 203 software (Keiser and Van Riemsdijk, 1994). Here, metal ions binding to humic substances (HS, mix of humic and fulvic acids) was assumed to occur through specific interactions between 204 205 cations and reactive organic moieties of the NOM and by nonspecific Coulombic binding to any residual negative charge (Benedetti et al., 1995; Benedetti et al., 1996a) The model merges 206 207 the Non-ideal Competitive Adsorption (NICA) isotherm model with the Donnan model. The first (NICA isotherm) describes competitive binding of protons and cations to NOM including 208 209 binding site heterogeneity and the second (Donnan model) details the electrostatic interactions within the structure of humic substances considered as a water-permeable gel (Benedetti et al., 210 1996a; Benedetti et al., 1996b) The model accounts for binding site heterogeneity by including 211 two different groups of binding sites each with a continuous affinity distribution: type 1 and 2, 212 corresponding to low affinity and high affinity sites, respectively (Benedetti et al., 1995). The 213 model equation is (Eq. 1): 214

215
$$Q_{i} = \frac{n_{i,1}}{n_{H,1}} Q_{\max 1,H} \frac{\left(\tilde{\kappa}_{i,1}c_{i}\right)^{n_{i,1}}}{\sum_{i}\left(\tilde{\kappa}_{i,1}c_{i}\right)^{n_{i,1}}} \frac{\left[\sum_{i}\left(\tilde{\kappa}_{i,1}c_{i}\right)^{n_{i,1}}\right]^{p_{1}}}{1 + \left[\sum_{i}\left(\tilde{\kappa}_{i,1}c_{i}\right)^{n_{i,1}}\right]^{p_{1}}} + \frac{n_{i,2}}{n_{H,2}} Q_{\max 2,H} \frac{\left(\tilde{\kappa}_{i,2}c_{i}\right)^{n_{i,2}}}{\sum_{i}\left(\tilde{\kappa}_{i,2}c_{i}\right)^{n_{i,2}}} \frac{\left[\sum_{i}\left(\tilde{\kappa}_{i,2}c_{i}\right)^{n_{i,2}}\right]^{p_{2}}}{1 + \left[\sum_{i}\left(\tilde{\kappa}_{i,2}c_{i}\right)^{n_{i,1}}\right]^{p_{2}}}$$
(1)

where the amount of bound ion i, Q_i (mol/kg) at solution concentration c_i (mol/L) is given by the sum of two identical binding expressions, the first one for low affinity type of site (subscript 1) and the second one for high affinity type of site (subscript 2) distributions. Four parameters describe intrinsic heterogeneity of HS ($Q_{max1,H}$, $Q_{max2,H}$, p_1 and p_2) and four ion specific parameters ($\tilde{K}_{i,1}$, $\tilde{K}_{i,2}$, $n_{i,1}$ and $n_{i,2}$) portray ion binding. p_1 and p_2 define the width of the affinity distributions and condense the intrinsic heterogeneity of HS. $Q_{max1,H}$ and $Q_{max2,H}$ (mol/kg) are

- the maximum proton binding capacity of HS. $\tilde{K}_{i,1}$ and $\tilde{K}_{i,2}$, are the median values of the affinity distributions for ion *i*, and $n_{i,1}$ and $n_{i,2}$ express the non-idealities of the ion-binding for each distribution. For further NICA-Donnan theory, the reader has to refer to the following studies (Benedetti et al., 1995; Benedetti et al., 1996a and b; Kinniburgh et al., 1996; Milne et al., 2001; Koopal et al., 2005; Milne et al., 2003).
- Input data were total concentrations of K and Tl, or Ag in solution, pH and ionic strength. 228 Tl and K ($\tilde{K}_{i,1}, \tilde{K}_{i,2}, n_{i,1}$ and $n_{i,2}$) were adjust together by fitting Tl experimental data from 229 concentration isotherms only. Parameters specific to the used purified humic acid and proton 230 binding were from Botero et al. (2017) and summarized in Appendix A (Table S5). 231 Experimental errors were calculated according to the propagation of uncertainties theory 232 233 (details in Appendix A, Eq. 1) while Root Mean Square Error (*RMSE*, Appendix A, S-Eq. 2) was used to estimate the gap between experimental and modelled data. In multiple studies 234 assumptions were made of similar behaviour between K and Tl(I) (Jacobson et al., 2005; Peter 235 and Viraraghavann 2005; Voegelin et al., 2015). While no specific parameters existed for K and 236 no isotherm for K were made, two modelling hypothesis were tested: (i) equal ion specific 237 parameters for both Tl(I) and K (i.e. $\tilde{K}_{i,1}, \tilde{K}_{i,2}, n_{i,1}$ and $n_{i,2}$) are chosen; (ii) two different sets of 238 parameters specific to each element are considered. K parameters will be discussed in paragraph 239 240 2.1. and values for K parameters for which the fit was the closest to the experimental data were kept to model Ag and Tl final data (reader has to refer to Appendix A, Fig. S4 for more details). 241 242

243 **1.6 Ag NICA-Donnan parameter derivation.**

- 244 No experimental data were newly acquired for Ag in this study, previously published data from Sikora and Stevenson (1988) were used and only Ag(I) NICA-Donnan parameters were 245 optimized. Sikora and Stevenson (1988) used three purified humic acids in their experiments, 246 247 HA1 from a A horizon taken from sandy silt loam, HA3 from Rifle peat, Ohio and HA8 from a A horizon sample from the Hapster silt loam (Sikora and Stevenson, 1988). Binding isotherms 248 constructed at pH 6.5 with 0.1 mol/L KNO3 and 0.5 g/L of humic acid (Sikora and Stevenson, 249 1988). Ag specific parameters were adjust to fit experimental data. Q_{max1, H} for HA1, HA3 and 250 HA8 were directly taken from Sikora and Stevenson (1988) while the remaining specific 251 parameters for protons ($Q_{max2,H}$, p_1 , p_2 , $\log \tilde{K}_{H,1}$, $\log \tilde{K}_{H,2}$) were optimized during modeling and 252 kept within the range of parameters defined for humic acids in Milne et al., (2001). Details are 253 254 summarized in Appendix A (Table S6).
- 255

256 2 Results and discussion

257 2.1 Tl(I) and Ag(I) complexation with humic acids.

Specific parameters fitted for K, Ag and Tl are given in Table 1. The hypothesis where K(I)
parameters are different from Tl(I) parameters gave the best description of the complexation of
Tl(I) by the HA.

Thallium complexation to HA increases with increasing concentrations (Fig. 1a). However, the 261 amount of complexed Tl(I) is low and always below 50% of the total thallium in solution. Ion 262 specific NICA-Donnan parameters were adjusted to fit Tl(I) concentration isotherm values (Fig. 263 264 1a). Then, the robustness of the model parameters was tested by comparing measured free Tl(I) to predicted free Tl(I) obtained from sorption edge and fixed pH isotherm experiments (Fig. 1b) 265 266 not used during the calibration procedure. Fig. 1b also shows that the predictions were satisfactory for the pH sorption edge data (i.e. within a 5% confidence area) notwithstanding 267 268 the parameter's optimization made with data corresponding to one pH value (7.35 \pm 0.05). Similar conclusions are reached when comparing measured bound-Tl(I) to simulated bound-269 270 Tl(I) (Fig. 1c). The slope of the concentration isotherm (Fig. 1a), close to one (i.e. 0.86), suggests that Langmuirian behaviour dominates the process and that monodentate binding 271 272 controls Tl(I) sorption onto HA (Koopal et al., 1994). This is confirmed by model parameters 273 since non-ideality parameters partly related to the type of binding (mono vs polydentate) $(n_{i,1})$ 274 and n_{i,2}, Benedetti et al., 1995; Benedetti et al., 1996a) optimized for Tl(I) are close to 1 (Table 1). Additional data for Tl(I) are presented in Appendix A (Table S7). 275

276 Fig. 2 displays Ag(I) modelling results whereas optimized Ag(I) specific parameters are given in **Table 1**. For Ag(I), non-ideality parameters were substantially different from 1, suggesting 277 278 more chemical ion specific binding heterogeneity as defined in NICA-Donnan papers (see above for references) during the complexation of Ag(I) when compared to Tl(I). Indeed at pH 279 280 below 5, the small amount of bound-Tl(I) is controlled by low affinity type of sites. At pH > 5281 and independently of thallium concentrations, high affinity sites dominate the binding, accounting for more than 65 % of total bound-Tl(I). These results derive from the high values 282 283 found for the median affinity constant associated to the high affinity sites (**Table 1**). At low pH, 284 15% of total bound-Tl(I) is located in the Donnan gel via Coulombic interactions. In Sikora's study (Sikora and Stevenson, 1988) with pH = 6.5, Ag(I) is almost exclusively bound to high 285 affinity sites and only a small amount is bound to low affinity sites or are in the Donnan gel (i.e. 286 less than 5% of total bound silver). Ag(I) binding to humic acid is strongly controlled by high 287 288 affinity sites conversely to Tl(I) and the amount of Ag(I) bound is also much higher (Fig. 3).

Ag has a higher affinity for organic matter than Tl(I) and its mobility is strongly influenced by this type of ligand (Sikora and Stevenson, 1988; Jacobson et al., 2005a). **Fig. 3** also suggests that Ag(I) could compete with divalent cations for binding to NOM reactive sites. The pHdependency of Ag complexation is smaller than the one of divalent ions. At neutral pH, Ag may compete with cobalt for instance (**Fig. 3**) and at acidic pH, Ag(I) has more affinity for NOM proxies than Co^{2+} , Mn²⁺ and Cd²⁺ (**Fig. 3**).

295

296 2.2 The S-rich functional groups.

297 Two recent studies evaluated the thiol group concentration in natural organic matter (NOM; Rao et al., 2014) calculated a concentration of 0.7 µmol/g for Suwanee River NOM and 3.6 298 299 µmol/g for a soil HA. Joe-Wong et al., (2012) estimated thiol group concentration at 535 µmol/g of dissolved organic matter (DOM). Regardless of the huge discrepancy between these three 300 301 numbers, in our experiments, the amount of bound Tl(I) at pH 7.35 varied from 68.4 pmol/g of HA to 2.57 µmol/g of HA for the lowest and highest total Tl concentration, respectively. In 302 Sikora's data, (Sikora and Stevenson, 1988) bound silver ranges between 87 µmol/g and 920 303 304 µmol/g of HA. The amounts of bound Tl(I) and Ag(I) are lower than the range of thiol group 305 concentration. Ag(I) has a high affinity for thiolate (Smith et al., 2002; Kleja et al. 2016) and since Ag-S binding was evidenced (Kleja et al., 2016) thiol groups could definitely play a role 306 307 in its speciation but other uncharacterized S-rich functional groups cannot be excluded. A similar assumption could be relevant for Tl(I) despite limited access to data investigating the 308 309 role of thiol groups for Tl(I) biogeochemistry (Garcia Bugarin et al., 1989; Jacobson et al., 310 2005b; Rodrigez-Mercado et al., 2013).

311 However, due to the low abundance of thiol groups and their expected pK_a value around 8

- 312 (Smith et al., 2002; Rao et al., 2014), they are probably accounted non-specifically within the
- high affinity sites continuous distribution of the NICA-Donnan model (Benedetti et al., 1995;
- Benedetti et al., 1996a, 1996b). The $\log \tilde{K}_{i,2}$ values 4.88 and 2.40, obtained for Ag(I) and Tl(I),
- respectively (Table 1), are higher than the values reported in Milne et al.'s study (2003) and
- 316 should be used now to simulate Ag or Tl interaction with NOM regardless of the assumptions
- 317 on the role of S-rich functional groups.
- 318

319 **2.3.** Consequences for Tl(I) geochemical cycle in waters and soils.

320 In this study, Tl(I) aqueous speciation (dissolved phase), in various natural water systems from

321 Kaplan and Mattigod (1998) is revisited. Data on groundwater, river water, eutrophic lake water,

322 bog water and seawater are gathered together with constants used in Tl(I) speciation modeling

323 in Appendix A (Table S8). The following assumptions are made to simulate the Tl(I) speciation in the above natural water systems. For groundwater, river water, eutrophic lake and sweater it 324 is assumed that DOM was mainly FA and that it is HA for the bog water. 80% in mass of the 325 DOM was assumed to be reactive according to Tipping (2002). In order to get HA or FA 326 327 concentrations in natural waters, DOM was calculated from DOC values considering a ratio DOM to DOC of 2:1 (by mass of C) (Ren et al., 2015). Constants for other complexes than Tl(I) 328 aqueous species are summarized in Appendix A (Table S9). The parameter values for FA were 329 estimated by making the following hypothesis. The heterogeneity parameters $n_{i,1}$ and $n_{i,2}$ in 330 NICA-Donnan are very similar for HA and FA (Milne et al., 2003). The values obtained here 331 for a HA can thus be used for FA in combination with the other generic binding parameters. 332 Conversely, in Milne et al., (2003) empirical relationship linking $n_{i,1}$, $\tilde{K}_{i,1}$ and K_{OH} for FA and 333 HA are similar. Therefore, combining those two observations, we assumed that the derived Tl(I) 334 HA parameters can be used in combination with the generic FA parameters from Milne et al.'s 335 336 study (2003) to simulate the Tl(I) interaction to the low affinity site of the DOM. A similar reasoning is made for the high affinity sites to derive $\tilde{K}_{i,2}$. The generic parameters used to 337 describe the FA and HA interaction with Tl(I) are given in Appendix A (Table S10). With the 338 339 following assumptions, we are now able to simulate the different aquatic ecosystems and the results are given in Appendix A (Table S11) and presented in Fig. 4. This figure shows that 340 341 only two dissolved aqueous species of Tl(I) would exist in natural waters, Tl(I) represents 342 between 76 and 98% of total dissolved Tl while Tl bound to NOM analogues represents the remaining thallium, except for the seawater system where TlCl and TlCl2 account for 46% of 343 the total thallium. However, unlike Kaplan and Mattigod's results (1998) it seems that there is 344 not a major contribution of Tl-NOM species in bog water (i.e. TlHA is < 10%). This is the result 345 of the competition between protons and Tl(I) to the same type of sites. Under the low pH 346 347 conditions of the peat bog the protons compete very efficiently and prevent Tl(I) binding to the bog NOM. In this study, Tl(I)-NOM complex is however slightly more important for all other 348 conditions if compared to Kaplan and Mattigod's results (1998) (Table S11). Discrepancies 349 with Kaplan and Mattigod's study (1998) are probably due to the different type of modelling 350 351 performed in both studies as well as the available binding data. Indeed, NICA-Donnan or 352 WHAM 7 models were proven to give much more robust prediction of metal ion speciation even outside of the range of metal ion concentrations used for their calibration (Benedetti et al., 353 354 1996b; Tipping, 2002). These findings support that our modelling predictions could be more 355 robust than previous ones as far as concentration is concerned. In addition, the range of concentration used in this study, to derive parameters, is always difficult to simulate for models (Benedetti et al., 1996b; Tipping, 2002). Most models always performed well at high concentration but their extrapolations to much lower concentration ranges are always much more problematic (Benedetti et al., 1996b; Tipping, 2002). As expected, no Tl(I)-NOM complexes are found in seawater and half of thallium is free Tl(I) ions and half is chloride complex species (TlCl and TlCl₂, **Table 1**).

Thallium speciation in soil solution was modelled. Compositions of soil solutions are taken 362 from Kinniburgh et al., (1983). Various concentrations relevant to pristine or polluted systems 363 364 are taken into consideration. Many organic molecules are regrouped as NOM in soil solutions. According to Ren et al. (2015), approximately 29% are humic substances, 66% are hydrophilic 365 366 acids and 5% are non-reactive organic components. During modelling, hydrophilic acids are 367 equally divided in acetic acid, citric acid and oxalic acid, three low molecular weight organic 368 acids (LMWOAs). Humic substances are modelled as FA since Ren et al., (2015) and van Zomeren and Comans, (2007) revealed that they are FA and not HA as previously imagined. 369 370 Details of used binding parameters, compositions and results are presented in Appendix A (Tables S10 and S12). Results are gathered in Fig. 5. Thallium does not strongly complex with 371 372 such LMWOA as shown in Vanek et al., (2012). Therefore, like for natural water systems, only two species control Tl(I) speciation: Tl(I) ions and Tl-FA. At low Tl(I) concentrations (≤ 1 373 374 nmol. L^{-1}), up to 35% of the thallium is bound to NOM (Fig. 5). Conversely, at higher Tl(I) concentrations ($\geq 1 \text{ nmol}.L^{-1}$), less than 10% of Tl(I) is bound to NOM (Fig. 5). 375

376 Wick et al., (2018) gathered data on Tl binding to a number of phases from clay minerals, to manganese oxides. When compared to these results, our data suggest that similar amounts of 377 378 Tl(I) are complexed by humic substance as ferrihydrite, illite and smectite (Wick et al., 2018; Martin et al., 2018). In addition, our results are close to the ones from Liu et al., (2011). Despite 379 380 similar affinity for organic matter and illitic clays, observations in soils showed a limited role 381 of organic matter on thallium uptake (Vanek et al., 2009; Vanek et al., 2013; Voegelin et al., 2015). Thus, thallium seems to be highly mobile and labile when it is observed in organic rich 382 383 soil horizons (Vanek et al., 2009; Vanek et al., 2013) in agreement with results showed in Fig. 384 5. Furthermore, natural organic matter could also lead to release thallium in soil by enhancing silicate alteration/weathering as suggested in some studies (Vanek et al., 2009; Ren et al., 2015). 385 Tl(I) in critical zone conditions will behave more like a lithophile element close to alkali metals 386 (K⁺, Rb⁺ and Cs⁺). This conclusion is supported by observations made that illitic clay systems 387 control the mobility of Tl(I) like for K⁺ and Cs⁺ (Jacobson et al., 2005; Voegelin et al., 2015; 388 Wick et al., 2018; Martin et al., 2018). These results also point out that in rivers, Tl(I) is 389

predominantly a highly bioavailable free ion (Tl⁺). Indeed, various studies highlighted that Tl(I) 390 poorly reacts with particulate phases and remains mainly in the dissolved fraction as a free ion 391 (Fig. 4) (Law and Turner, 2011; Boning et al., 2017; Casiot et al., 2017). However, several 392 questions regarding thallium remain open, especially the role of Tl(III) and its 393 394 abundance/stability in the environment (Lin and Nriagu, 1999; Twining et al., 2003; Karlsson 395 et al., 2006; Watanabe and Takahashi, 2015). Only one paper reported the binding of Tl(III) to HA (Watanabe and Takahashi, 2015), the Tl(III) species would be stabilized by HA but only at 396 397 low pH and high redox potential values. The reactivity of Tl(III) with NOM needs more studies 398 to fully picture the Tl environmental biogeochemistry. The combination of HPLC-ICP-MS 399 system designed based on previous studies (Campanella et al., 2017; Casiot et al., 2017) and 400 the presently used DMT approach would greatly help to make progress. Finally, in order to 401 complete the database, Tl(I/III) interactions with other analogues of NOM should be studied.

402

403 **3 Conclusion**

404 Thallium complexation was compared to silver complexation using literature data and using the 405 same NICA-Donnan formalism. We report here, for the first time, parameters for these two 406 cations. Results display low thallium complexation to the NOM proxy while silver competes with divalent cations for the NOM binding sites. Calculated speciation for dissolved thallium 407 408 highlights the dominance of free thallium (Tl⁺) in solution whereas Tl-NOM complexes 409 contribute roughly 15% to total Tl(I) species in river and lake type waters. Similar results are 410 obtained for soil solutions, Tl-bound to NOM < 30% of total, from UK soils with different land 411 use and geochemistry. However, several questions regarding thallium remain open, especially 412 the role of Tl(III) and its abundance/stability in the environment.

413

414 Acknowledgments

The authors thank IRSN, CEA, BRGM and IPGP for funding this research. Parts of this work
were also supported by IPGP multidisciplinary program PARI and by Region Ile de France
SESAME Grant no. 12015908.

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419 Appendix A. Supplementary dara

420 Supplementary data associated with this article can be found in the online version at xxxxxx.

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- 422 **References**

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Table 1 Ion specific parameters (ISP) for the NICA-Donnan model selected after adjustment with experimental data. *RMSE* were calculated only with data used for ISP adjustments. For Tl(I) it means only data from the concentration isotherm. Subscript 1 and 2 correspond respectively to low affinity and high affinity sites of the humic substances as defined in NICA-Donnan model theory (Benedetti et al., 1995; Benedetti et al., 1996a)

Element	$\log \widetilde{K}_{i,1}$	n _{<i>i</i>,1}	$\log \widetilde{K}_{i,2}$	n _{<i>i</i>,2}	RMSE for free	RMSE for
					ion	bound ion
Tl(I)	0.20	0.88	2.40	0.84	0.04 (n = 9)	0.16 (n = 9)
Ag(I)	-0.10	0.59	4.88	0.71	-	0.02 (n = 62)
K(I)	0.2	0.88	1.14	0.84	-	-

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Fig. 1 (a) Evolution of bound thallium as function of free thallium. pH for measured data was 7.35 \pm 0.05 with [HA] = 150 mg/L. For free TI: *RMSE* = 0.04 (n = 9) and for bound Tl(I): *RMSE* = 0.16 (n = 9). (b) Calculated free thallium after the optimization of NICA-Donnan parameters with the data set given in (a) vs. measured free Tl(I). (c) Measured vs. calculated for complexed/bound thallium. For b and c, red lines are regression lines whereas dashed red lines correspond to the confidence area (5%) of the regressions. The dashed black lines are one to one lines. Error bars are $\pm 2\sigma$.

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Fig. 2 Bound silver versus free silver for three humic acids (HA1, 3 and 8). Experimental points
are from Sikora and Stevenson (1988). Background electrolyte was KNO₃ 0.1 mol/L, pH was
6.5 and [HA] = 500 mg/L. Dashed black lines corresponding to NICA Donnan simulation
obtained after adjustment of NICA-Donnan model parameters values



640

Fig. 3 Log [metal bound] calculated with NICA-Donnan generic parameters for humic acids from Milne et al., (2003) for metal ions represented by black dots. Tl and Ag parameters come from the present study. The total metal concentration is equal to 10 nmol/L of a given metal and a background concentration of 1 mmol/L of Ca^{2+} is used for each calculations.





647 **Fig. 4** Modeled Tl speciation in various natural water systems from Kaplan and Mattigod (1998)





Fig. 5 Modeled Tl speciation in three soil solutions at relevant environmental concentrations.

- 651 More details are summarized in Appendix A (Tables S10 to S12 from Kinniburgh and Miles,
- 652 1983)[.]

Appendix A. Supplemenary data
Effect of natural organic matter on thallium and silver speciation
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1 Experimental setup with Field-Donnan Membrane Technique (Kalis et al. 2007)



Fig. S1 Field-DMT from Kalis et al. 2007 (a) Front view; (b) top view; (c) DMT cell.



- **Fig. S2** Experimental set-up used for Tl(I) experiments. *Tl solution'' = Donor solution is a
- 24 mixture that contains $Tl(I)^+$, K^+ , NO_3^- and humic acid.
- 25



26 2 Time to equilibrium for Tl⁺ (Donnan Membrane Technique)

Fig. S3 Time to reach equilibrium for TI^+ at $pH = 4.0 \pm 0.1$, between the DMT compartments (i.e. acceptor vs. donor sides) in absence of humic acid. Error bars are $\pm 2\sigma$ of the experimental data. Time to reach equilibrium is **24 hrs** for TI^+ Dashed lines correspond to the modeled variation of cation concentrations as function of time during experiments. Model equation is given in Temminghoff et al., 2000.

33

34 **3** Composition and calculated speciation of experimental solutions

Table S1 Calculated Tl speciation in the acceptor side of the DMT. Modeled with CHESS 2.0.

36 Only species corresponding to more than 1% of total concentration are presented here.

37 Saturations indices indicate that under-saturation towards Tl bearing minerals (TlOH, TlNO₃,)

is expected during the experiments. Additional constants came from Casiot et al. (2011).

	Total				Abunda	nce in %	6 of tota	1	
Element	concentration	Species	pH =	pH =	pH =	pH =	pH =	pH =	pH =
	(mol/L)		3	4	5	6	7	8	9
Tl	2. 10 ⁻¹⁰	Tl^+	100	100	100	100	100	100	99.99
	2. 10 ⁻⁶	$T1^+$	100	100	100	100	100	100	99.99

39

40 **Table S2** Compositions of donor and acceptor solutions at t=0. TINO₃ is not present because it

41 only exists as a solid specie.

	Tl isotherms
Donor	Tl ⁺ , K ⁺ , NO ₃ ⁻ , Humic Acids, Na ⁺ (after basic
	pH)
Acceptor	K^+, NO_3^-

42

43 4 ²⁰⁵Tl measurements with HR-ICP-MS Element 2 (Thermo Scientific)

44 **Table S3** Operating conditions and data acquisition parameters for HR-ICP-MS Element 2 used

45 for ²⁰⁵Tl measurements. Thallium was measured only in low-resolution mode.

<u>HR-ICP-MS</u>	1

	Rf power (W)	1350
¹)	Sample uptake rate (mL.min ⁻	0.2
	Argon flow rates (L.min ⁻¹)	
	Cool	16
	Auxiliary	1.0
	Sample	Between 0.9 and 1.0 (depends of sensibility)
	Torch	Quartz torch with a separate quartz injector tube
	Nebulizer	PFA ST microflow nebulizer

Spray chamber	Quartz Cyclo spray chamber
Sampler cone	Nickel
Skimmer cone	Nickel
Data acquisition	Low Resolution
No. of passes	3
Mass window (%)	20
Search window (%)	0
Integration window (%)	20
Samples per peak	25
Sample time (s)	1.0 (²⁰⁵ Tl)

47 **Table S4** Measured Tl values for TM23.4 and TM24.4 certified materials (Environment

48 Canada).

	ТМ	23.4	TM	[24.4
	Expected	Average	Expected	Average
	Expected	Measured	Expected	Measured
[Tl] µg/L	3.97	4.08 (n=6)	4.20	4.27 (n = 4)
Error or SD	0.33	0.10	0.34	0.06
Difference to expected value (%)	-	2.67	-	1.55

⁴⁹

50 **5 NICA-Donnan model parameters**

- 51 Table S5 Purified humic acid specific parameters used during NICA-Donnan modeling are
- 52 from Botero et al., 2017.

A	β	Q _{max,1} (mol/kg)	$\log \widetilde{K}_{H,1}$	$\mathbf{n}_{1,\mathrm{H}}$	Q _{max,2} (mol/kg)	$\log \widetilde{K}_{H,2}$	n _{2,H}	p_1	<i>p</i> ₂
0.60	-0.27	3.18	3.65	0.84	3.02	8.07	0.65	0.79	0.44

⁵³

6 Estimation of experimental errors and *RMSE* **calculations**

55 Uncertainties reported in this section are calculated from the propagation of known 56 measurements or lab errors following the same approach as Reinoso-Maset et al (2014). These 57 did not take into account the humic acid variability. Experimental errors were hence calculated for each experimental point following the propagation errors theory. The variance of a G function of different x_i variable can be calculated from the variances of the variables x_i using the following expression:

$$\sigma_G^2 = \sum_{i=1}^n \left(\frac{\partial G}{\partial x_i}\right)^2 \sigma_i^2 + 2\sum_i \sum_j \frac{\partial G}{\partial x_i} \frac{\partial G}{\partial x_j} \sigma_{ij}$$
(S Eq. 1)

62 where $(\delta G/\delta x_i)$ is the partial derivative of G with respect to x_i , σ^2 is the variance of xi and 63 σ_{ij} the covariance of the x_i and x_j variables. If these variables are independents, the covariance 64 term is then equal to zero.

65 Root Mean Square Error (*RMSE*, S Eq. 2) was calculated to assess the gap between 66 experimental and modelled values for each series.

67
$$RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^{n} \left(\log[Me^{z_{+}}]_{i,measured} - \log[Me^{z_{+}}]_{i,modelled} \right)^{2}}$$
(S Eq. 2)

68

69 **7 Experimental data and additional modeled data**



70

Fig. S4 Free Tl concentration over total Tl concentration function of pH. When Tl = K is mentioned (green dashed line), it means that ion-specific parameters, $\tilde{K}_{i,1}$, $\tilde{K}_{i,2}$, $n_{i,1}$ and $n_{i,2}$ for both Tl and K were the same during modeling. On the opposite, when Tl \neq K is mentioned (red dashed line), those parameters were different. Experimental error bars are $\pm 2\sigma$.

75

76 8 Silver complexation with humic acids

77 **Table S6** Parameters used during modeling with ECOSAT software for experimental data of

silver complexation to three humic acids from Sikora et al. (1988). In all simulations, pH=6.5,

	HA1	HA3	HA8
В	0.49	0.49	0.49
$Q_{max1,\mathrm{H}}$	4.43 (from Sikora <i>et al.</i> , 1988)	3.95 (from Sikora <i>et al.</i> , 1988)	3.95
p_1	0.62	0.62	0.62
$\log \widetilde{K}_{H,1}$	3.00	3.00	3.00
$\mathbf{n}_{\mathrm{H,1}}$	0.80	0.80	0.80
Qmax2, H	3.30	3.30	3.30
p ₂	0.41	0.41	0.41
$\log \widetilde{K}_{H,2}$	7.00	7.60	7.25
n _{H,2}	0.63	0.63	0.63
$\log \widetilde{K}_{K,1}$	0.20	0.20	0.20
n _{K,1}	0.88	0.88	0.88
$\log \widetilde{K}_{K,2}$	1.14	1.14	1.14
n _{K,2}	0.84	0.84	0.84
$\log \widetilde{K}_{Ag,1}$	-0.10	-0.10	-0.10
n _{Ag,1}	0.59	0.59	0.59
$\log \widetilde{K}_{Ag,2}$	4.88	4.88	4.88
nAg,2	0.71	0.71	0.71
m_1	0.50	0.50	0.50
m ₂	0.26	0.26	0.26

79	background electrolyte was KNO ₃ 0.1 mol/L and HA concentration was 0.5 g/L. Values in bold
80	and italic are those that were adjusted. $m_i = n_i p_i$ (Benedetti et al., 1995).

Table S7 Experimental data acquired for Tl(I) and used for calculations. n.s. not significant

[T1]donor, eq	[T1]acceptor,	[T1] _{free,don, eq}	[T1] _{complexed}	log[T1] _{free}	log[Tl] _{complexed}	T (°C)	ъU
(mol/L)	_{eq} (mol/L)	(mol/L)	(mol/kg)	(mol/L)	(mol/kg)	Tavg (C)	рп
2.24E-11	1.67E-11	1.210E-11	6.84E-08	-10.92	-7.17	24.4	7.39
2.24E-11	1.68E-11	1.16E-11	7.17E-08	-10.94	-7.14	24.4	7.39
2.11E-10	3.08E-10	1.27E-10	5.59E-07	-9.90	-6.25	23.5	7.40
1.73E-09	2.59E-09	1.47E-09	1.74E-06	-8.83	-5.76	24.2	7.35
1.73E-09	2.57E-09	1.46E-09	1.81E-06	-8.84	-5.74	24.2	7.35

2.03E-07	2.78E-07	1.72E-07	2.05E-04	-6.76	-3.69	23.1	7.31
2.03E-07	2.76E-07	1.67E-07	2.39E-04	-6.78	-3.62	23.1	7.31
1.82E-06	2.79E-06	1.60E-06	1.43E-03	-5.79	-2.84	23.4	7.31
1.82E-06	2.62E-06	1.43E-06	2.57E-03	-5.84	-2.59	23.4	7.31
1.71E-10	1.81E-10	1.73E-10	n.s.	-9.76	n.s.	20.9	3.98
1.71E-10	1.84E-10	1.75E-10	n.s.	-9.76	n.s.	20.9	3.98
1.57E-10	1.57E-10	1.69E-10	n.s.	-9.77	n.s.	21.6	4.87
1.57E-10	1.55E-10	1.65E-10	n.s.	-9.78	n.s.	21.6	4.87
1.45E-10	1.39E-10	1.43E-10	1.88E-08	-9.85	-7.73	22.3	5.78
1.48E-10	1.28E-10	1.33E-10	9.49E-08	-9.88	-7.02	20.9	6.77
1.48E-10	1.29E-10	1.34E-10	8.93E-08	-9.87	-7.05	20.9	6.77
1.74E-10	1.22E-10	1.38E-10	2.38E-07	-9.86	-6.62	22.5	7.75
1.34E-10	1.13E-10	1.04E-10	1.96E-07	-9.98	-6.71	22.0	7.89
3.17E-11	2.89E-11	2.84E-11	2.16E-08	-10.55	-7.67	22.3	7.5
3.17E-11	2.75E-11	2.78E-11	2.60E-08	-10.56	-7.59	22.3	7.5

84 9 Kaplan and Mattigod's original data

Table S8 Concentrations reported by Kaplan and Mattigod (1998) for Tl(I) speciation calculation in the dissolved phase in selected natural water systems. Authors reported the following references for water composition: Groundwater and lake water (Stumm and Morgan 1981); river water (Hem 1985); bog water (Thurman 1985); seawater (Nordstrom *et al.*, 1979);

e	,				
	Groundwater	River water	Eutrophic Lake Water	Bog water	Seawater
[Tl(I)] _{tot} (mol/L)	3.55.10-8	9.79.10-11	9.79.10-11	9.79.10-11	6.36 .10-11
[Ca] (mg/L)	59	15	40	0.2	422
[Mg] (mg/L)	26	4.1	10	0.19	1,322
[Na] (mg/L)	22	6.3	9	1.5	11,020
[K] (mg/L)	4	2.3	2	0.31	408
[CO ₃] (mg/L)	266	57	122	0.06	145
[SO ₄] (mg/L)	108	11	24	0.53	2,775
[Cl] (mg/L)	11	7.8	9	0.99	19,805
[F] (mg/L)	0.1	1	1	0.1	1.4

89 organic carbon (Thurman 1985).

[NO ₃] (mg/L)	39	1	3.4	1	0.3
[PO ₄] (mg/L)	0.1	0.0767	0.6	0.0767	0.0614
[H4SiO4] (mg/L)	48	20.8	2	20.8	4.4
Org. Carbon (mg/L)	0.7	5.0	10.0	30.0	0.5
pH	7.14	8.01	7.70	3.60	8.22

10 Constants and associated references used in this paper

Table S9 Table of constant used in Tl(I) speciation modeling with ECOSAT software.

Reaction	log <i>K</i> (298.15K)	Reference	
$Tl^+ + OH^- \leftrightarrow TlOH$	0.63	This study	
$Tl^+ + SO_4^{2-} \leftrightarrow TlSO_4^{-}$	1.37	a	
$Tl^+ + Cl^- \leftrightarrow TlCl(aq)$	0.51	a	
$Tl^+ + 2Cl^- \leftrightarrow TlCl_2^-$	0.28	a	
$Tl^+ + CO_3^{2-} \leftrightarrow TlCO_3^{-}$	2.16	a	
$Tl^+ + 2CO_3^{2-} \leftrightarrow Tl(CO_3)_2^{3-}$	0.11	a	
$Tl^+ + HCO_3^- \leftrightarrow TlHCO_3$	0.90	a	
$Tl^+ + HPO_4^2 \leftrightarrow TlHPO_4^-$	1.27	b	
$Tl^+ + PO_4^{3-} \leftrightarrow TlPO_4^{2-}$	3.54	c	
$Tl^+ + F^- \leftrightarrow TlF$	0.10	c	
$Tl^+ + Cit^{3-} \leftrightarrow TlCit^{2-}$	2.00	c	
$Tl^+ + Ox^{2-} \leftrightarrow TlOx^-$	1.39	b	
$Tl^+ + Ac^- \leftrightarrow TlAc^0$	-0.11	c	
$Tl^+ + 2SO_4^{2-} + 3Fe^{3+} + 6OH$	2.26	b	
$\leftrightarrow TlFe_3(SO_4)_2(OH)_6(s)$	3.30	, ,	
$Tl^+ + 2SO_4^{2-} + 3Al^{3+} + 24H^+$	7 46	с	
$\leftrightarrow TlAl(SO_4)_2 \cdot 12H_2(s)$	/.40	-	
$Tl^+ + OH^- \leftrightarrow TlOH(s)$	1.08	с	
$Tl^+ + Cl^- \leftrightarrow TlCl(s)$	-3.65	b	
$2Tl^+ + CO_3^{2-} \leftrightarrow Tl_2CO_3(s)$	3.84	d	
$2Tl^{+} + SO_{4}^{2\text{-}} \leftrightarrow Tl_2SO_4(s)$	3.79	a	
$Tl^+ + NO_3^- \leftrightarrow TINO_3(s)$	1.06	b	

93 * Calculated from Gibbs free energy of formation

94 a: Lin and Nriagu, (1999); b: Vink, (1983); c: Xiong, (2009) and d: Casiot et al., (2011)

Table S10 Table of generic NICA-Donnan constants used in Tl speciation modeling with
 ECOSAT software for the description of NOM in various aquatic systems or soil solutions.

	Q _{max,1} (mol/kg)	$\log \widetilde{K}_{H,1}$	n _{1,H}	Q _{max,2} (mol/kg)	$\log \widetilde{K}_{H,2}$	n _{2,H}	p_1	<i>p</i> ₂
FA	5.88	2.34	0.88	1.86	8.6	0.84	0.59	0.70
HA	3.15	2.93	0.81	3.15	8.0	0.68	0.62	0.41

98 Parameters for other elements are taken form Milne et al. (2003)

99

100 Table S11 Calculated distribution of Tl(I) species for the dissolved phase in selected natural

101 water systems according to Kaplan and Mattigod (1998). Species not reported here are102 negligible.

	Groundwater	River water	Eutrophic Lake Water	Bog water	Seawater
Tl^+	90.4	82.7	76.8	32.4	51.9
TlHCO ₃ °	4.4	1.2	2.0	-	0.5
TlCO ₃ -	-	-	-	-	0.1
T1SO4 ⁻	3.6	0.4	0.8	-	11.2
TlCl°	0.1	0.1	0.1	-	30.7
TlCl ₂ -	-	-	-	-	5.4
Tl-Fulvate°	1.4	15.6	20.3	67.6	0.2

103

104 **11 Modeling with soil solutions**

Table S12: Calculated dissolved Tl(I) speciation in soil solution with ECOSAT. Compositions 105 of solutions and soils characteristics are in Kinniburgh et al. (1983). DOM (Dissolved organic 106 matter) is composed of 29.3% of humic substances, 66.2% of hydrophilic acids and 4.5% is 107 considered non-reactive according to average from Ren et al. (2015) and Groenenberg et al. 108 109 (2010). Humic substances are simplified as only fulvic acids and hydrophilic acids are simplified (one third for each) as combination of citric acid (Cit), oxalic acid (Ox) and acetic 110 111 acid (Ac). Tl minerals taken in consideration are lanmuchangite, dorallcharite, TlCl(s), TINO₃(s) and Tl₂SO₄(s). No precipitation occurred. Only species with above 1 % are reported. 112 Tl-Cit²⁻, Tl-Ox⁻, Tl-Ac, TlCl(aq), TlOH and TlSO₄⁻ are negligible. Constants (minerals and 113 dissolved complexes) used in calculation are summarized in Table S9. 114

G . :1			[Tl]total	Tl-FA	Tl+ (% Tl
5011	рн	HA (mg/L)	(nmol/L)	(% Tl total)	total)
Icknield	7.8	27.0	1.00.10-2	34.7	65.0
			0.16	25.5	74.2
			1.26	19.9	79.8
			10.00	15.3	84.4
			158.49	10.6	89.0
			1258.90	8.0	91.6
<u>Denchworth</u>	7.0	32.8	1.00.10-2	16.3	83.0
			0.16	11.6	87.7
			1.26	8.9	90.4
			10.00	6.8	92.4
			158.49	4.8	94.5
			1258.90	3.7	95.6
Southampton	4.3	50.4	1.00.10-2	4.8	94.8
			0.16	3.5	96.0
			1.26	2.8	96.7
			10.00	2.3	97.2
			158.49	1.8	97.8
			1258.90	1.5	98.1

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