



HAL
open science

Two-Dimensional Spectra of Radar Returns from Sea: 1. Theoretical and Numerical Study of the Group Line

Florestan Platzer, Marc Saillard, Vincent Fabbro

► **To cite this version:**

Florestan Platzer, Marc Saillard, Vincent Fabbro. Two-Dimensional Spectra of Radar Returns from Sea: 1. Theoretical and Numerical Study of the Group Line. *Journal of Geophysical Research. Oceans*, 2019, 10.1029/2019JC015121 . hal-02493149

HAL Id: hal-02493149

<https://hal.science/hal-02493149>

Submitted on 27 Feb 2020

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

RESEARCH ARTICLE

10.1029/2019JC015121

This article is a companion to Platzter et al. (2019), <https://doi.org/10.1029/2019JC015123>.

Key Points:

- Two-dimensional spectra were extracted from radar returns from sea surface
- A pattern is observed at low frequencies in wavenumber-frequency spectra
- The characteristics of this pattern have been explicitly linked to the spectrum of waves

Correspondence to:

F. Platzter, M. Saillard and V. Fabbro, florestan.platzer@onera.fr; marc.saillard@univ-tln.fr; vincent.fabbro@onera.fr

Citation:

Platzter F., Saillard, M., & Fabbro, V. (2019). Two-dimensional spectra of radar returns from sea: 1. Theoretical and numerical study of the group line. *Journal of Geophysical Research: Oceans*, 124, 8767–8776. <https://doi.org/10.1029/2019JC015121>

Received 6 MAR 2019

Accepted 25 AUG 2019

Accepted article online 31 OCT 2019

Published online 11 DEC 2019

Two-Dimensional Spectra of Radar Returns from Sea: 1. Theoretical and Numerical Study of the Group Line

F. Platzter^{1,2} , M. Saillard² , and V. Fabbro¹ 

¹ONERA/DEMR, Université de Toulouse, Toulouse, France, ²Université de Toulon, Aix Marseille Université, CNRS/INSU, IRD, MIO UM 110, Mediterranean Institute of Oceanography, La Garde, France

Abstract The work presented here aims at interpreting the pattern observed at low frequencies in wavenumber-frequency (k, ω) representations (dispersion diagrams) of radar returns from sea surface. This pattern is sometimes referred to as the “group line” in the literature and is supposed to result from the nonlinear behavior of the signal with respect to surface profile. To confirm this assumption, we have calculated the patterns generated by various nonlinear terms of second order in sea surface height. It is shown that the resolution in frequency of most data is not sufficient to allow for an accurate representation of the group line through an FFT. A methodology is proposed to define an average slope of the group line, which can be interpreted in terms of velocity.

1. Introduction

Time series of microwave radar cross section (RCS) from sea surface are almost routinely used to image sea surface and extract informations about surface current or sea state, see for instance the review paper Borge et al. (2013) and references herein. Such radars operate in the microwave range and illuminate the surface under grazing incidence from a vessel or a platform. Data processing is often based on a Fourier transform in both range and time. The resolution in range of the radars are less than a few meters while time series duration is about a few tens of seconds. The resulting 2D (or 3D) spectrum exhibits some characteristic patterns of gravity waves, including a bright line (or surface) in accordance with their dispersion relation, but also some energy along the harmonics and a so-called “group line” at low time and space frequencies. The group line results from the nonlinear dependence of the backscattered field with respect to surface parameters. If coherent radars are used, one can map either the radar cross section or the Doppler velocity, both leading to similar plots Frazier and McIntosh (1996). Interpretation of the group line has been investigated in Frazier and McIntosh (1996), Smith et al. (1996), Plant and Farquharson (2012), with emphasis on the signature of breaking waves or wave groups at the crest. It is also noticed that the group line occurs in slices of the 3D Fourier transform of optical measurements Leckler et al. (2015), but it seems that the authors have not focused their attention on it.

In Frazier and McIntosh (1996), the authors explore many explanations. After suggesting that “there is little doubt that the backscattered power profile has a nonlinear relation to local incidence angle” at grazing incidence, which is the preferred interpretation in the present paper, a detailed discussion about the velocity of non-Bragg contributors that may be linked to breaking waves or scatterers bounded to the crest finally drives the reader to an interpretation in terms of hydrodynamic phenomena. The same year, Smith et al. (1996) also interpret the slope of the group line in terms of “group modulation of wave breaking,” since the associated velocity exceeds what would be expected from the modulation of Bragg wave.

More recently, Plant and Farquharson (2012) have excluded turbulent eddies in the wind, water turbulence advected by currents, second-order hydrodynamic or electromagnetic nonlinearities or shadowing as main contributors, while the latter has been considered as the dominant nonlinear mechanism in Borge et al. (2013). In Plant (2012), Plant establishes a correspondance between the slope of the group line, the velocity derived from Doppler spectra and the peak of Philips' Lambda function, introduced in Phillips (1985). In the wind direction, these quantities are close to each other, suggesting all of them bring the signature of breaking waves. However, numerical simulations of dispersion diagrams based on nonlinear Schrödinger equation do exhibit such features Krogstad and Trulsen (2010), for weakly nonlinear one-dimensional surfaces. This is also the case for the simulations presented here. Therefore, not only the interpretation of the group line

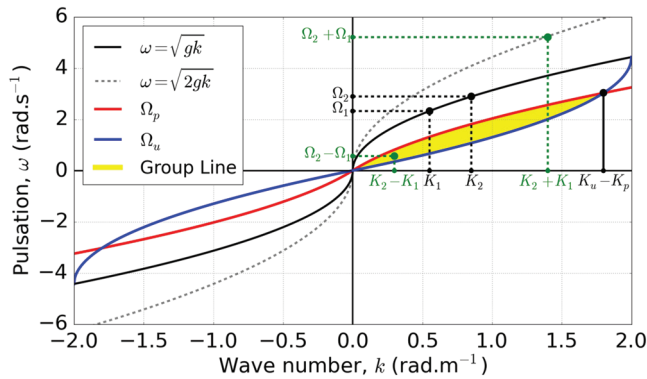


Figure 1. Geometrical description of the group line associated with a band-limited continuous power spectrum.

still deserves to be discussed, but there is also some need for a clear definition of the slope of the group line, since it is associated to the velocity of waves.

In this paper, after the group line has been defined and exhibited in Section 2, we present in Section 3, for the one-dimensional case, a detailed theoretical analysis of the group line generated by a quadratic function of sea surface height or of its slope. The same calculations can be performed for any quadratic term involving the surface height and its time or spatial derivatives, but only these two have been of interest when compared to the data, in sections 6 and 7 at the end of the paper. This is also consistent with the discussions about the link between sea spikes, fast scatterers, and breaking waves which focus on large slopes and on the crests of high waves. In Section 4, it is shown that the fast Fourier transform (FFT) of experimental time series lacks of resolution in frequency to provide an accurate description of the group line. However, an empirical formula is given in Section 5, which avoids tedious computations for a huge set of frequencies.

2. Group Line

This section focuses on the group line generated by a nonlinear function of a random variable with Gaussian probability density and given power spectrum $S(\mathbf{K})$, typically a linear superposition of gravity waves or the associated distributions of slopes. Since the radar resolution in range is about 1 m, or more, smaller scales are filtered out from S and capillary waves are not considered. On the low frequency side, the spectrum is assumed to vanish very rapidly below the wavenumber K_p associated with the dominant wave. The calculations performed here thus apply to typical spectra of wind waves and do not take into account the superposition of swell.

At large depths and without current, the dispersion relation of gravity waves is $\Omega = \sqrt{gK}$, $K = |\mathbf{K}|$ and Ω denoting the wavenumber and the frequency, respectively.

In the frame of linear approximation, the contribution to surface height of waves longer than the radar resolution in range, characterized by the high-wavenumber cutoff K_u , can be written as

$$\eta_L(\mathbf{x}, t) = \int_{K < K_u} A(\mathbf{K}) \exp(i\mathbf{K} \cdot \mathbf{x} - i\Omega(\mathbf{K})t) d\mathbf{K} + cc, \quad (1)$$

where $|A|$ is the square root of the spectrum of surface gravity waves and $\Omega(\mathbf{K})$ is given by the dispersion relation of gravity waves.

To understand the influence of nonlinearities and the characteristics of the group line, let us first consider the simple case of a quadratic dependence η_L^2 . For simplicity, let us combine two periodic waves propagating in the same direction, $\eta_1(K_1, \Omega_1)$ and $\eta_2(K_2, \Omega_2)$.

The product $\eta_1\eta_2$ generates terms with wavenumbers $K = \pm K_2 \pm K_1$ and associated frequencies $\Omega = \pm \Omega_2 \pm \Omega_1 = \pm \sqrt{gK_2} \pm \sqrt{gK_1}$. If K_1 and K_2 are close to each other, the resulting points in the (k, ω) plane are close either to the second harmonic $\omega = \pm \sqrt{2gk}$ or to the origin. The slope of the straight line joining the origin to the latter is nothing but the group velocity of the wave packet.

Let us now keep $k = K_2 - K_1$ constant and increase K_1 from K_p up to K_u , where K_p is the low frequency cutoff of the spectrum of gravity waves. Since shorter waves than K_u are discarded here, $A(K_1)A^*(K_2)$ vanishes when K_1 lies out of the interval $[K_p, K_u - k]$. $|\Omega_2 - \Omega_1|$ is thus decreasing from $\omega_p = \sqrt{g(K_p + k)} - \sqrt{gK_p}$ down to $\omega_u = \sqrt{gK_u} - \sqrt{g(K_u - k)}$. Consequently, for a wide-band spectrum, the group line fills a part of the (k, ω) plane of which pattern is controlled by the long energetic waves on one side (K_p), the resolution in range on the other side (K_u), as illustrated in Figure 1. Obviously, the same result holds for the square of the slope,

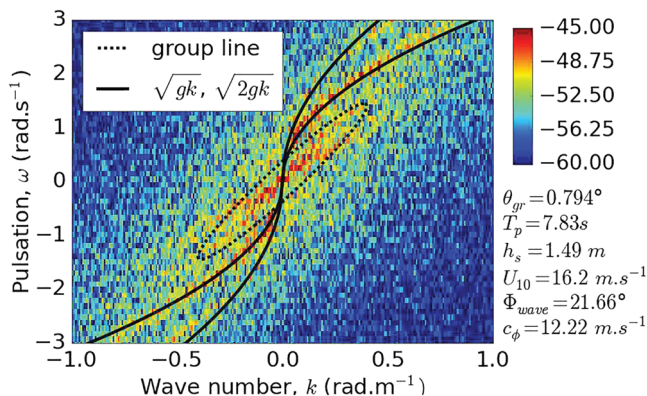


Figure 2. Example of dispersion diagram derived from experimental data Fabbro et al. (2017).

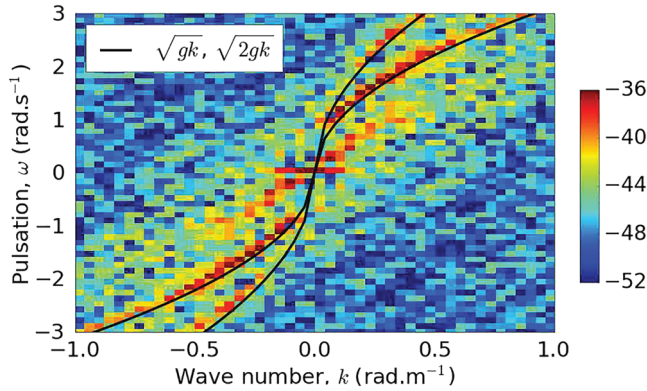


Figure 3. Dispersion diagram of the computed backscattered amplitude for 1D sea surface with wind speed 7 m/s, illuminated under grazing incidence and H polarization at L band. The resolution in range is 1.25 m. Horizontal axis: wavenumber; vertical axis: frequency.

since only $A(\mathbf{K})$ is modified in this case. Diagrams from real data, such as the one in Figure 2 computed through a 2D FFT of range-time series from the database described in Fabbro et al. (2017), do exhibit a cloud of energetic points at low space and time frequencies. The boundaries of the group line do not clearly appear, because of noise, of poor resolution, and of the direction spreading of gravity waves.

Similar plots are also obtained from a numerical solver based on a rigorous boundary integral formalism, devoted to scattering from one-dimensional rough surfaces illuminated under grazing incidence Miret et al. (2014). In Figure 3, sea surface is described by Pierson Moskowitz spectrum and computations are performed for horizontal polarization at L band. Creamer's approach is used to represent hydrodynamic nonlinearities Creamer et al. (1989). The contribution from each range cell is recorded during about 1 min, and a two-dimensional FFT is then performed. It is interesting to notice that the group line is clearly visible, even though hydrodynamic nonlinearities are assumed to be weak and no breaking occurs. This result is consistent with the theoretical predictions of Krogstad and Trulsen (2010).

In the two-dimensional case, the Fourier transform with respect to \mathbf{x} of η_L^2 is nothing but the autoconvolution of $A(\mathbf{K}) \exp(-i\Omega(\mathbf{K})t)$ (defined equation (1)),

$$\begin{aligned} \tilde{\eta}_L^2(\mathbf{k}, t) = & \iint_{K < K_u} A(\mathbf{K})A(\mathbf{k} - \mathbf{K}) \exp(-i(\Omega(\mathbf{K}) + \Omega(\mathbf{k} - \mathbf{K}))t) d\mathbf{K} + cc \\ & + \iint_{K < K_u} A(\mathbf{K})A^*(\mathbf{K} - \mathbf{k}) \exp(-i(\Omega(\mathbf{K}) - \Omega(\mathbf{K} - \mathbf{k}))t) d\mathbf{K} + cc, \end{aligned} \quad (2)$$

where K refers to wavenumber of waves while k refers to the Fourier coordinate. A second Fourier transform with respect to t leads to Dirac distributions $\delta(\omega \pm (\Omega(\mathbf{K}) + \Omega(\mathbf{k} - \mathbf{K})))$ and $\delta(\omega \pm (\Omega(\mathbf{K}) - \Omega(\mathbf{k} - \mathbf{K})))$. Only the second one provides values of ω located below the dispersion relation $\omega = \sqrt{gk}$ and contributes to the group line. If, instead of η_L^2 , a quadratic function involving slopes is considered, the spectrum of gravity waves has to be replaced by the corresponding moment.

In the following, we will focus our attention on the case where the radar beam propagates in the direction of the wind. In these conditions, \mathbf{k} lies along this direction and the quantity $A(\mathbf{K})A^*(\mathbf{K} - \mathbf{k})$ that contributes to the group line takes significant values when \mathbf{K} is also directed in this direction and decays rapidly away from it, as shown in Figure 4 for typical weather conditions encountered during data measurements Fabbro et al. (2017). In addition, since $\omega = \Omega(\mathbf{K}) - \Omega(\mathbf{K} - \mathbf{k})$ is an even function of K_\perp , the component of \mathbf{K} perpendicular to \mathbf{k} , we have $\frac{\partial \omega}{\partial K_\perp}(K_\perp = 0) = 0$. It results that ω is almost unchanged when \mathbf{K} slightly deviates from wind direction. Therefore, it seems reasonable in this case to perform the analytical estimation of the group line in the frame of a 1D problem.

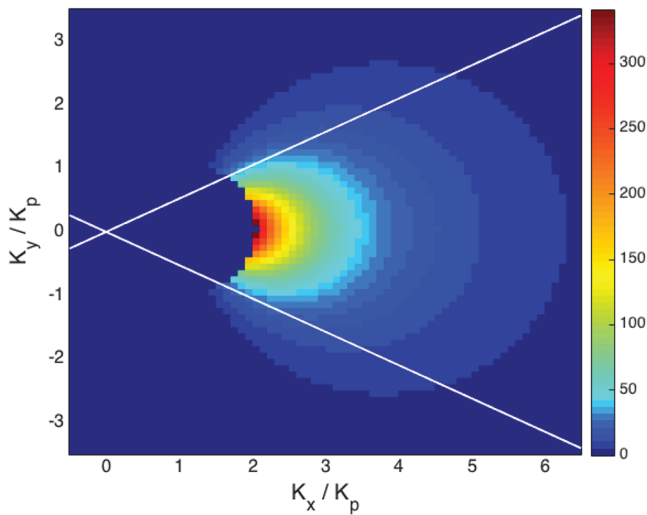


Figure 4. $|A(\mathbf{K})A(\mathbf{K} - \mathbf{k})|$ versus K_x/K_p and K_y/K_p , for $\mathbf{k} = K_p \hat{\mathbf{x}}$ with peak wave $K_p = 0.1$ and phase speed to wind speed ratio $c_p/U = 0.75$. A is assumed to follow the empirical spectrum proposed by Hasselmann et al. (1980); 82% of the integral is concentrated between the white lines with slopes $\pm \tan \pi/6$. The distribution becomes even narrower when k is increased.

3. Analytical Results

Here, it is assumed that all the wavevectors are aligned along the x axis, oriented according to wind direction. Since $\tilde{\eta}_L^2(k, \omega) = \tilde{\eta}_L^{2*}(-k, -\omega)$, we can restrict our investigations to $k \geq 0$. In this case, the 2D FFT $\tilde{\eta}_L^2$ can be calculated exactly. Indeed, only one value of K between $K_p + k$ and K_u satisfies $\omega - \Omega(K) + \Omega(K - k) = 0$, as required by the Dirac distribution.

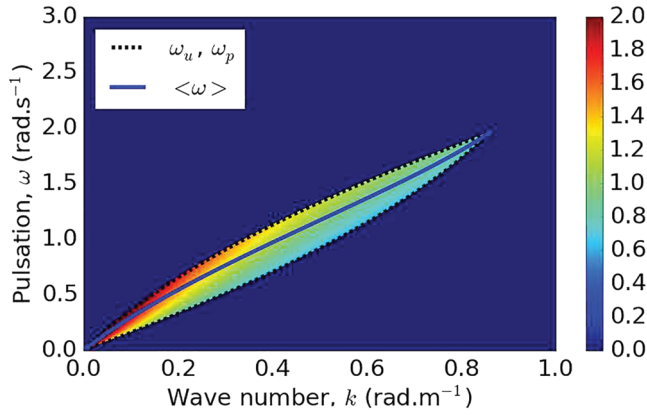


Figure 5. Group line associated with η_L^2 (log scale) for a band-limited K^{-3} wave spectrum. The low-frequency cutoff is $K_p = 0.14 \text{ rad/m}$ and the radar range resolution, acting as high-frequency cutoff, is $K_u = 1 \text{ rad/m}$. The blue line represents the weighted average value, $\langle \omega \rangle$ (k).

From, $\omega = \sqrt{gK} - \sqrt{g(K-k)}$, it comes once squared,

$$\sqrt{g(K-k)} = \frac{\sqrt{gk}}{2} \left(\frac{\sqrt{gk}}{\omega} - \frac{\omega}{\sqrt{gk}} \right) \quad (3)$$

$$\sqrt{gK} = \frac{\sqrt{gk}}{2} \left(\frac{\sqrt{gk}}{\omega} + \frac{\omega}{\sqrt{gk}} \right). \quad (4)$$

Therefore, if $S(K)$ is a power-law spectrum which behaves as $K^{-\alpha}$ and with $A(K) \propto \sqrt{K^{-\alpha}}$,

$$|A(K)A^*(K-k)| \propto k^{-\alpha} \left(\frac{gk}{\omega^2} - \frac{\omega^2}{gk} \right)^{-\alpha}. \quad (5)$$

In addition, since $\omega(K)$ is a single valued invertible function, the contribution of $\tilde{\eta}_L^2(k, \omega)$ to the group line is

$$\left| \frac{dK}{d\omega} \right| A(K)A^*(k-K) \quad (6)$$

with $|A(K)A^*(k-K)|$ given by equation (5). The derivative of equation (4) with respect to ω gives

$$\left| \frac{dK}{d\omega} \right| = \frac{k}{2\omega} \left(\frac{gk}{\omega^2} - \frac{\omega^2}{gk} \right) \quad (7)$$

resulting in

$$\left| A(K)A^*(K-k) \frac{dK}{d\omega} \right| \propto \frac{k^{1-\alpha}}{\omega} \left(\frac{gk}{\omega^2} - \frac{\omega^2}{gk} \right)^{1-\alpha}. \quad (8)$$

An example is given in Figure 5, which exhibits the group line derived from equation (6), for a band-limited K^{-3} spectrum. The plot has been restricted to the upper right quadrant, since the spectrum is centro-symmetric and all the waves go in the same direction here. To translate this 2D map in terms of “velocity,” the weighted average $\langle \omega \rangle$ is calculated for each k ,

$$\langle \omega \rangle = \frac{\int_{\omega_u}^{\omega_p} \omega \left| \tilde{\eta}_L^2(k, \omega) \right| d\omega}{\int_{\omega_u}^{\omega_p} \left| \tilde{\eta}_L^2(k, \omega) \right| d\omega}, \quad (9)$$

and the mean slope $\langle \omega \rangle / k$ between 0 and k is considered. In Figure 5, $\langle \omega \rangle$ has been plotted, according to the following formula, derived after tedious calculations

$$\langle \omega \rangle_{\eta_L^2} = \sqrt{gk} \frac{\left[\frac{\frac{\omega}{\sqrt{gk}}}{1 - \left(\frac{\omega^2}{gk}\right)^2} - \frac{1}{4} \ln \frac{1 + \frac{\omega}{\sqrt{gk}}}{1 - \frac{\omega}{\sqrt{gk}}} - \frac{1}{2} \arctan \frac{\omega}{\sqrt{gk}} \right]_{\omega_p}^{\omega_u}}{\left[\frac{1}{1 - \left(\frac{\omega^2}{gk}\right)^2} \right]_{\omega_p}^{\omega_u}}, \quad (10)$$

where $\omega_p = \sqrt{g(K_p + k)} - \sqrt{gK_p}$ and $\omega_u = \sqrt{gK_u} - \sqrt{g(K_u - k)}$ are the upper and lower boundaries of the domain covered by the group line, respectively.

The same kind of calculations can be performed for the square of the slope, η_{Lx}^2 ($\eta_x = \frac{\partial \eta}{\partial x}$). In this case, since the spectrum has a slower decay, as K^{-1} , equation (8) takes a very simple form, namely, $1/\omega$. As a result the energy density of the group line spreads out in a more homogeneous way, leading to smaller values of $\langle \omega \rangle$ than for η_L^2 .

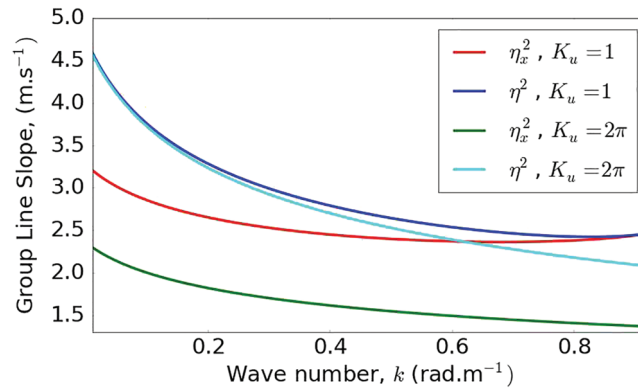


Figure 6. Average slope of the group line $\langle \omega \rangle / k$ versus k for a band-limited K^{-3} spectrum with $K_p = 0.14 \text{ rad/m}$. Blue line: slope of η_L^2 for $K_u = 1$; cyan line: the same for $K_u = 2\pi$; red line: slope of η_{Lx}^2 for $K_u = 1$; green line: the same for $K_u = 2\pi$.

$$\langle \omega \rangle_{\eta_{Lx}^2} = \frac{[\omega]_{\omega_p}^{\omega_u}}{[\ln \omega]_{\omega_p}^{\omega_u}}. \quad (11)$$

The difference between the group lines associated with η_L^2 and η_{Lx}^2 is illustrated in Figure 6, which compares both average slopes $\langle \omega \rangle / k$. The high sensitivity of the group line of η_{Lx}^2 to the high-frequency cutoff is also exhibited in Figure 6, where both cases $K_u = 1$ and $K_u = 2\pi$ have been considered. The higher K_u , the smaller slope. This results from the slow decay of the spectrum of η_{Lx} , of which rms depends on K_u .

On the contrary, the slope of the group line of η_L^2 is hardly changed, because most of the energy is concentrated along the line $\omega = \omega_p(k)$. It is also worth to notice that the limit when k tends to 0 of $\langle \omega \rangle$ in equation (10) is $0.4\sqrt{g/K_p}k$, leading to slope at origin $0.4c_\phi(K_p)$ for η_L^2 in Figure 6, where $c_\phi(K_p) = \sqrt{g/K_p}$ denotes the phase velocity of the most energetic wave. This is a little less than its group velocity $c_g = c_\phi/2$, because of the weight of slower waves of sea spectrum. The 0.4 coefficient is thus the signature of the K^{-3} decay of the spectrum. Let us also point out that the slope decreases with k .

Similar analytical calculations can be done for other types of second-order terms, such as η_{Lt}^2 or $\eta_L\eta_{Lx}$, but such calculations become too complicated for higher order nonlinearities. However, it has been observed that the dispersion diagrams derived from experimental data present very small amount of energy in the vicinity of higher harmonics (Figure 2), suggesting that second-order terms are the leading nonlinearities.

4. Synthetic Data

Range-time series of $\eta_L(x, t)$ have also been generated numerically through the superposition of time-harmonic waves with random phases. It appears that for usual durations of recording, typically a minute, the dispersion plots derived from 2D FFT significantly differs from the previous analytical results. The discrepancy is reduced when very long time series are considered.

To understand the origin of this behavior, let us come back to Figure 1 in Section 2, where we have pointed out that, for given k , $\omega = \sqrt{g(k+K)} - \sqrt{gK}$ is slowly decreasing toward zero when K increases. More precisely, since the derivative of equation 4 gives

$$\frac{dK}{d\omega} = -\frac{k}{2\omega} \left(\frac{gk}{\omega^2} - \frac{\omega^2}{gk} \right), \quad (12)$$

the interval dK which feeds the range of frequencies $[\omega - d\omega, \omega]$ behaves as ω^{-3} for small values of ω . Therefore, when performing a time FFT, with resolution $\Delta\omega = 2\pi/T$ where T denotes the duration of the time series, the contributions from large wavenumbers, for instance $K_1 = K_u$ and $K_2 = K_u - k$, can be separated if $T > \frac{8\pi}{\sqrt{gk}} \left(\frac{K_u}{k} \right)^{3/2}$. With typical values such as $K_u = 2\pi$ and $k = 0.1$, we get $T > 12600$ s. Numerical

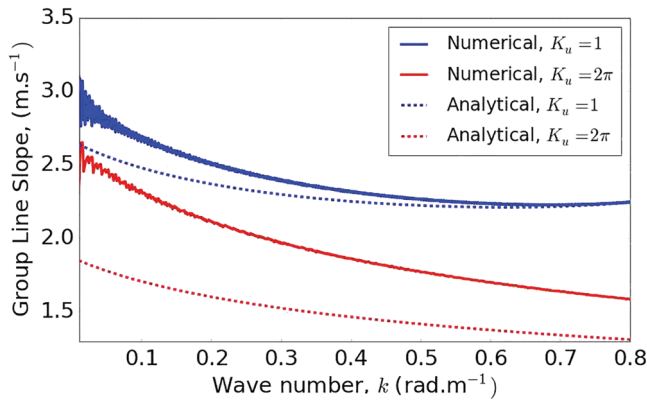


Figure 7. Comparison of theoretical and computed average slopes of the group line $\langle \omega \rangle / k$ versus k , for η_{Lx}^2 (blue: $K_u = 1$; red: $K_u = 2\pi$).

with analytical formula. For instance, when considering the group line associated with η_{Lx}^2 , the average frequency $\langle \omega \rangle (k)$ follows with very good accuracy

$$\langle \omega \rangle = \frac{3 \left[\omega^{5/2} \right]_{\omega_u}^{\omega_p}}{5 \left[\omega^{3/2} \right]_{\omega_u}^{\omega_p}} \quad (13)$$

instead of equation (11). Though we have no mathematical proof, everything occurs as if $dk/d\omega$ behaves as $\omega^{-3/2}$ instead of ω^{-3} (equation 12). This rule also works for η_L^2 , leading to replace equation (10) by the formula given in Appendix A

As an example, Figure 8 exhibits the agreement between the empirical formula and computations of the 2D FFT of short synthetic range-time series. Displacements of 500 m long surface profiles have been recorded during 400 s. Averaging is performed over 200 of such surfaces. To sum up the results presented in the last two sections, Figure 9 presents how the predictions of the 2D FFT evolves from the empirical formula at low recording durations down to the theoretical result of Fourier analysis at very long ones. Accuracy at very low values of k is hard to achieve, since a small fluctuation of $\langle \omega \rangle$ generates a large deviation of $\langle \omega \rangle / k$.

Hence, we have now at hand a way to transform the average value of ω derived from FFT of range-time series with poor resolution in frequency into the one that should result from Fourier transform of an infinitely long time series, through multiplication by the ratio equation (11) /equation (13). This step is essential if one aims at giving physical interpretation of the data.

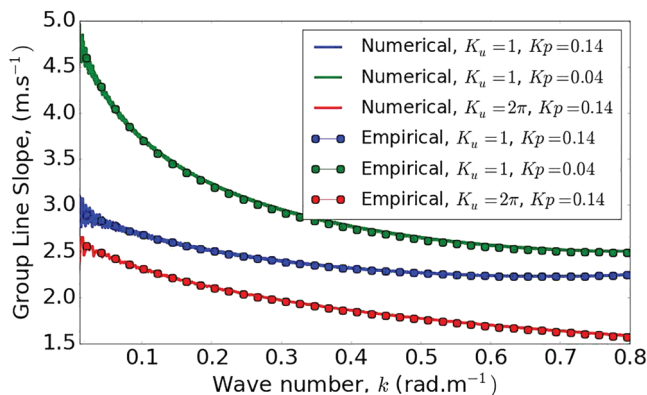


Figure 8. Comparison of empirical formula and computed average slopes of the group line $\langle \omega \rangle / k$ versus k , for η_{Lx}^2 . The solid lines represent the computed average slopes while the circle markers represent the empirical formulas.

experiments have shown that an accurate computation of the integral is even more demanding. So a few minutes long time series do not allow an accurate representation of the group line generated by high-resolution radars. Obviously, the 2D FFT underestimates the contribution from low frequencies. This explains why the numerical estimation of $\langle \omega \rangle$ is greater than the exact value and why the discrepancy is greater for $K_u = 2\pi$ than for $K_u = 1$ in Figure 7, where 400 s long time series have been considered. In the same way, the error is much lower when considering η_L^2 than η_{Lx}^2 , since the weight of low frequencies is smaller in the former case.

Hence, we can conclude that the 2D FFT is not suited to study the group line generated from real data. As a consequence, quantitative estimation of physical parameters such as velocity of waves or breakers from such a representation is not relevant.

5. Empirical Formula to Correct 2D FFT Output

For such “short” time series, numerical experiments have shown that the group line derived from the 2D FFT presents a behavior that can be fitted

6. Comparison with Data from Numerical Solution of the Scattering Problem

To check the validity of the previous results, we have tested them against the group line derived from numerical solution of the scattering problem, as already exhibited in Figure 3. The computation provides time series of the backscattered amplitude from each radar cell. A dispersion diagram is obtained through a 2D FFT and the average value $\langle \omega \rangle (k)$ is computed from equation (9), where the square height has been replaced by the radar cross section. In Figure 10, the group line associated with the backscattered field from 100 radar cells, each 1.25 m in length, recorded during 1 min, is presented. The radar frequency is 1.2 GHz and its polarization is horizontal. In Miret et al. (2014) Tatarskii and Charnotskii (1998), it has been demonstrated that the backscattered amplitude from a rough

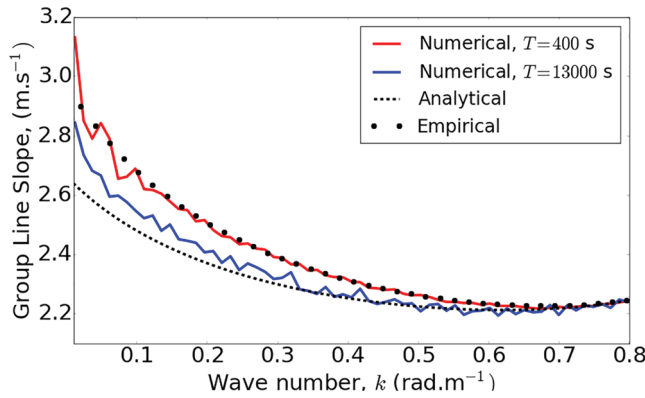


Figure 9. Comparison of theoretical, empirical, and computed average slopes of the group line $\langle \omega \rangle / k$ versus k , for η_{Lx}^2 ($K_u = 1$). (blue: $T_{int} = 13,000$ s; red: $T_{int} = 400$ s).

surface separating two homogeneous media behaves as the square of the vertical component of the incident wavevector when the grazing incidence angle tends toward zero. In the present model, the hydrodynamic nonlinearities are assumed to be weak and the surface profile remains smooth. The assumption of a boundary separating two homogeneous media is thus fulfilled. If one extends the limiting behavior of the scattering amplitude for each radar cell of sea surface, even if it is much smaller than the wavelength of long gravity waves, it comes that the fluctuations of the scattering amplitude with range follow that of the square of the vertical component of the local incidence wavevector. Keeping in mind that the slope of gravity waves remain low, the local incidence angle remains low as well. Therefore, denoting by θ_0 the grazing incident angle, the backscattered amplitude from a radar cell is proportional to the square of the local grazing angle

$$s^0(x) \propto (\sin \theta_0 + \eta_{Lx}(x))^2 \quad (14)$$

of which first nonlinear term with respect to sea surface profile is η_{Lx}^2 .

Therefore, since the time series is short, the average frequency derived from the scattering computations (blue solid line) has to be compared with the one predicted by the empirical formula established for the square of surface slope η_{Lx}^2 (equation (13)). A very good agreement is observed.

7. Comparison of the Slope of the Group Line with the Velocity of Breakers

Coherent radars permit one to get also range-time maps of the velocity, defined as the first moment of the Doppler spectrum. Their frequency-wavenumber spectra have been studied in Plant (2012), where it is noticed that the slope of the group line compares with the most probable velocity of breaking gravity waves c_0 . In Irisov and Plant (2016), the authors have fitted the set of c_0 data with a second-order polynomial in $c_\phi(K_p)$, the phase velocity of the dominant wave,

$$c_0 = 0.22 + 0.39c_\phi - 0.008c_\phi^2. \quad (15)$$

It is interesting to notice that the first order coefficient is very close to the slope at $k = 0$ of $\langle \omega \rangle$ in equation (10), which is equal to 0.4. This suggests to compare c_0 with the slope associated with η_{Lx}^2 . Since the second-order derivative of $\langle \omega \rangle$ at $k = 0$ is negative, we have found a small value of k which allows us to fit the polynomial, as shown in Figure 11. Although choosing $k = 0.11$ is not based on physical considerations, noticing that the group line in Figure 2 or in figure 1 of Plant (2012) concentrates energy between $k = -0.2$ and $k = 0.2$, it makes sense to estimate the slope in this range. But the most important here is that we get the same behavior and the same order of magnitude over a very wide range of values of $c_\phi(K_p)$. Therefore, if the radar echo is dominated by the contribution of breakers and if c_0 is a good approximation of their velocity, it may be conjectured that the group

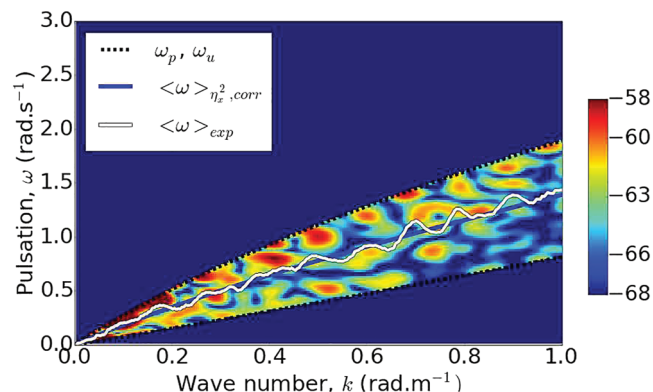


Figure 10. Focus on the group line computed from the solution of the scattering problem. White solid line: average frequency from scattering data. Blue solid line: average frequency for η_{Lx}^2 , according to equation (13).

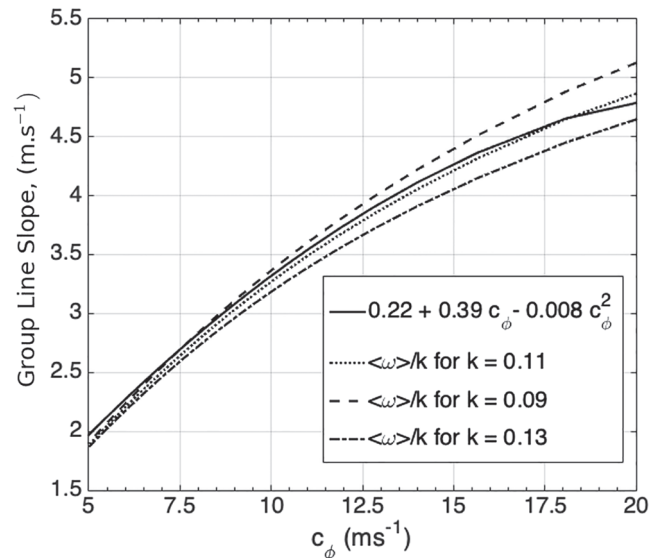


Figure 11. Slope of the group line $\langle \omega \rangle / k$ (equation (10)) versus phase velocity of the dominant wave $c_\phi(K_p)$ (red), for k around 0.11, versus the empirical formula proposed in Irisov and Plant (2016) for the most probable velocity of breakers (solid line).

line of the dispersion diagram of the average Doppler velocity should behave as the one derived from η_L^2 .

8. Conclusion

The characteristics of the group line have been explicitly linked to the spectrum of waves, at least for simple nonlinear functions of η . Since the group line is not a single line but an area of the (k, ω) plane, a systematic procedure has been proposed to give a mathematical sense to the “slope of the group line,” to allow its interpretation in terms of velocity. Applying numerical 2D fast Fourier transform to data with coarse frequency grid leads to overestimate the slope and to misinterpret the associated velocity. A correction must be applied to make the numerical result consistent with that derived from a Fourier transform.

Since the amplitude of the backscattered electromagnetic field is a nonlinear function of surface profile, the 2D dispersion diagram derived from the Fourier transform in range and time always exhibit a group line, whatever the sea surface state. In the particular case of low grazing incidence, when the asymptotic regime predicted in Tatarskii and Charnotskii (1998) is reached, numerical simulations provide consistent results, the slope of the group line fitting the one generated by the square of the slope of long gravity waves η_{Lx}^2 .

Finally, it has been noticed that the behavior of the slope of the group line associated with η_L^2 , with respect to phase velocity of the dominant wave, fits well the most probable velocity of breakers, as estimated in Irisov and Plant (2016). Since the computations performed here are restricted to weak hydrodynamic nonlinearities, we have not been able to exhibit group lines with such high slopes. This will be investigated in a companion paper focusing on in situ data analysis Platzer, Fabbro, & Saillard, 2019. (submitted).

Appendix A

Equations (9, 11, 13) require analytical calculation of integrals of the type

$$I(\beta, n) = \int \omega^\beta \left(\frac{gk}{\omega^2} - \frac{\omega^2}{gk} \right)^{-n} d\omega, \quad (\text{A1})$$

with $\beta = -1, 0, \frac{1}{2}, \frac{3}{2}$ and $n = 0$ or 2 .

Below is given the average value of ω for the group line associated with η_L^2 when a 2D FFT is used to compute the dispersion diagram.

$$\langle \omega \rangle = N/D \quad (\text{A2})$$

with $N = I(\frac{3}{2}, 2)$ and $D = I(\frac{1}{2}, 2)$.

$$\begin{aligned}
 N = \sqrt{gk} & \left[8 \frac{\left(\frac{\omega}{\sqrt{gk}}\right)^{5/2}}{1 - \left(\frac{\omega}{\sqrt{gk}}\right)^4} + 5 \ln \frac{1 - \frac{\omega}{\sqrt{gk}}}{1 + \frac{\omega}{\sqrt{gk}}} \right]_{\omega_u}^{\omega_p} \\
 & + \sqrt{gk} \left[\frac{10}{\sqrt{2}} \arctan \left(\frac{\sqrt{\frac{2\omega}{\sqrt{gk}}}}{1 - \frac{\omega}{\sqrt{gk}}} \right) - 10 \arctan \sqrt{\frac{\omega}{\sqrt{gk}}} \right]_{\omega_u}^{\omega_p} \\
 & + \sqrt{gk} \left[\frac{5}{\sqrt{2}} \ln \frac{1 + \sqrt{2\frac{\omega}{\sqrt{gk}} + \frac{\omega}{\sqrt{gk}}}}{1 - \sqrt{2\frac{\omega}{\sqrt{gk}} + \frac{\omega}{\sqrt{gk}}}} \right]_{\omega_u}^{\omega_p}
 \end{aligned} \tag{A3}$$

$$\begin{aligned}
 D = & \left[8 \frac{\left(\frac{\omega}{\sqrt{gk}}\right)^{3/2}}{1 - \left(\frac{\omega}{\sqrt{gk}}\right)^4} + 3 \ln \frac{1 - \frac{\omega}{\sqrt{gk}}}{1 + \frac{\omega}{\sqrt{gk}}} \right]_{\omega_u}^{\omega_p} \\
 & + \left[6 \arctan \sqrt{\frac{\omega}{\sqrt{gk}}} - \frac{6}{\sqrt{2}} \arctan \left(\frac{\sqrt{\frac{2\omega}{\sqrt{gk}}}}{1 - \frac{\omega}{\sqrt{gk}}} \right) \right]_{\omega_u}^{\omega_p} \\
 & + \left[\frac{3}{\sqrt{2}} \ln \frac{1 + \sqrt{2\frac{\omega}{\sqrt{gk}} + \frac{\omega}{\sqrt{gk}}}}{1 - \sqrt{2\frac{\omega}{\sqrt{gk}} + \frac{\omega}{\sqrt{gk}}}} \right]_{\omega_u}^{\omega_p}
 \end{aligned} \tag{A4}$$

with $\omega_p = \sqrt{g(K_p + k)} - \sqrt{gK_p}$ and $\omega_u = \sqrt{gK_u} - \sqrt{g(K_u - k)}$.

Acknowledgments

The authors acknowledge the Délégation Générale de l'Armement (DGA) for financial support and Gabriel Soriano, at Institut Fresnel, for valuable discussions and for providing the computer code devoted to solving the scattering problem. The data used in the paper are not available publicly. These are the property of Fraunhofer FHR, WTD71, and ONERA. Please contact the authors at vincent.fabbro@onera.fr for more information.

References

- Borge, J. N., Reichert, K., & Hessner, K. (2013). Detection of spatio-temporal wave grouping properties by using temporal sequences of x-band radar images of the sea surface. *Ocean Modelling*, *61*, 21–37. <https://doi.org/10.1016/j.ocemod.2012.10.004>
- Creamer, D. B., Henyey, F., Schult, R., & Wright, J. (1989). Improved linear representation of ocean surface waves. *Journal of Fluid Mechanics*, *205*, 135–161. <https://doi.org/10.1017/S0022112089001977>
- Fabbro, V., Biegel, G., Förster, J., Poisson, J. B., Danklmayer, A., Böhrer, C., et al. (2017). Measurements of sea clutter at low grazing angle in mediterranean coastal environment. *IEEE Transactions on Geoscience and Remote Sensing*, *55*(11), 6379–6389. <https://doi.org/10.1109/TGRS.2017.2727057>
- Frazier, S. J., & McIntosh, R. E. (1996). Observed wavenumber-frequency properties of microwave backscatter from the ocean surface at near-grazing angles. *Journal of Geophysical Research*, *101*(C8), 18,391–18,407. <https://doi.org/10.1029/96JC01685>
- Hasselmann, D. E., Dunckel, M., & Ewing, J. A. (1980). Directional wave spectra observed during jonswap 1973. *Journal of Physical Oceanography*, *10*(8), 1264–1280. [https://doi.org/10.1175/1520-0485\(1980\)010<1264:DWSODJ>2.0.CO;2](https://doi.org/10.1175/1520-0485(1980)010<1264:DWSODJ>2.0.CO;2)
- Irisov, V., & Plant, W. (2016). Phillips' lambda function: Data summary and physical model. *Geophysical Research Letters*, *43*, 2053–2058. <https://doi.org/10.1002/2015GL067352>
- Krogstad, H. E., & Trulsen, K. (2010). Interpretations and observations of ocean wave spectra. *Ocean Dynamics*, *60*(4), 973–991. <https://doi.org/10.1007/s10236-010-0293-3>
- Leckler, F., Arduin, F., Peureux, C., Benetazzo, A., Bergamasco, F., & Dulov, V. (2015). Analysis and interpretation of frequency-wavenumber spectra of young wind waves. *Journal of Physical Oceanography*, *45*, 2484–2496. <https://doi.org/10.1175/JPO-D-14-0237.1>
- Miret, D., Soriano, G., & Saillard, M. (2014). Rigorous simulations of microwave scattering from finite conductivity two-dimensional sea surfaces at low grazing angles. *IEEE Transactions on Geoscience and Remote Sensing*, *52*(6), 3150–3158. <https://doi.org/10.1109/TGRS.2013.2271384>
- Phillips, O. M. (1985). Spectral and statistical properties of the equilibrium range in wind-generated gravity waves. *Journal of Fluid Mechanics*, *156*, 505–531. <https://doi.org/10.1017/S0022112085002221>
- Plant, W. J. (2012). Whitecaps in deep water. *Geophysical Research Letters*, *39*, L16601. <https://doi.org/10.1029/2012GL052732>
- Plant, W. J., & Farquharson, G. (2012). Origins of features in wave number-frequency spectra of space-time images of the ocean. *Journal of Geophysical Research*, *117*(C6), C06015. <https://doi.org/10.1029/2012JC007986>

- Platzer, F., Fabbro, V., & Saillard, M. (2019). Two-dimensional spectra of radar returns from sea: Analysis of the group line from experimental data. *Journal of Geophysical Research: Oceans*, *124*. <https://doi.org/10.1029/2019JC015123>
- Smith, M., Poulter, E., & McGregor, J. (1996). Doppler radar measurements of wave groups and breaking waves. *Journal of Geophysical Research, Oceans*, *101*(C6), 14,269–14,282. <https://doi.org/10.1029/96JC00766>
- Tatarskii, V. I., & Charnotskii, M. I. (1998). On the universal behavior of scattering from a rough surface for small grazing angles. *IEEE Transactions on Antennas and Propagation*, *46*(1), 67–72. <https://doi.org/10.1109/8.655452>