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## Sidi M'Barek: a representative example of the Moroccan massive sulphide deposits

Abdelhak Outigua<sup>1,2</sup>, Abderrahim Essaifi<sup>2\*</sup>, Michel Corsini<sup>3</sup>, Mohamed Outhounjite<sup>2</sup> & Mohamed Zouhair<sup>2</sup>

<sup>1</sup> *Geology department, Cadi Ayyad University, BP. 2390, Marrakech, Morocco*

<sup>2</sup> *Managem Group, BP. 5199, Casablanca 20100, Morocco*

<sup>3</sup> *Geoazur, CNRS UMR6526, Université Nice Sophia Antipolis, Valbonne, France*

*\* Corresponding author (email: [essaifi@uca.ac.ma](mailto:essaifi@uca.ac.ma))*

abbreviated title: Sidi M'Barek massive sulphide deposit

**Abstract:** In addition to a Zn-rich lens representing the northern prolongation of the Draa Sfar volcanogenic massive sulphide deposit, the Sidi M'Barek deposit includes Cu-rich lenses in which gold is being recovered profitably as a by-product. The Cu-rich mineralization is composed of numerous structurally controlled veins whose morphology suits that of a lode rather than that of a stratiform orebody. Although the Cu-rich and Zn-rich orebodies are dominated by pyrrhotite, pyrrhotite is coarser grained in the Cu-rich lenses and mineralization is polyphase. Pyrrhotite is deformed and recrystallized; it is crosscut and replaced by pyrite infilling veins. Pyrite crystallizes also along the  $S_1$  cleavage and in pressure shadows around bioclasts in the disseminated ore. Chalcopyrite veins postdate both the pyrrhotite and pyrite phases.

The concentration of Co and Au is higher in the Cu-rich lenses relative to the Zn-rich lenses, which have higher concentrations in Pb, Ag and As. Pyrite and pyrrhotite have distinct sulphur and lead isotopic signatures. We propose a polygenetic model in which a pre-existing syngenetic pyrrhotite-rich mineralization is remobilized in shear zones during a major ductile deformation event allowing deposition of pyrite. The shear zones were reopened during a late ductile-brittle deformation event allowing deposition of chalcopyrite.

**Keywords:** Massive sulphide; Pyrrhotite; Draa Sfar; Hercynian; Morocco

Located in the Central Jebilet, the Draa Sfar district contains two deposits: a southern deposit called Tazakourt, and a northern deposit called Sidi M'Barek. These two deposits are separated at the surface by the Tensift River, but there is no structural discontinuity between the orebodies. The Tazakourt orebody consists of one Zn-rich lens that extends N-S for 1.5 km and has been delineated between -50 and -1500 m levels and remains open at depth. It represents the main ore lens at Draa Sfar and has been exploited since the commencement of mining at the Draa Sfar deposit in 2004. At that time the Sidi M'Barek orebody had not yet been delineated, and was thought to be composed of

two massive sulphide lenses, a western or upper Zn-rich lens interpreted to be the continuation of the Tazakourt ore lens, and an eastern or lower Cu-rich lens (Marcoux *et al.* 2008).

The previous studies on the Draa Sfar deposit were concentrated on the Zn-rich lens (Moreno *et al.* 2008; Ben Aissi *et al.* 2005; Hibti 2001; Rziki 2012). This deposit was classified as a VMS deposit (Belkabir *et al.* 2008; Marcoux *et al.* 2008; Moreno *et al.* 2008), but a feeder or stockwork sulphide zone has not been identified. Because the lower or eastern Cu-rich lens contains bismuth selenides, Marcoux *et al.* (2008) suggested that it may represent for the Draa Sfar ore deposit the upper part of a stockwork zone that has been strongly transposed by deformation. Marcoux *et al.* (2008) noticed however that the lead isotopic signature indicate that two types of hydrothermal fluids were involved in the Draa Sfar deposit, a less radiogenic fluid ( $^{206}\text{Pb}/^{204}\text{Pb}\sim 18.30$ ) that formed the Zn-rich lens and a more radiogenic fluid ( $^{206}\text{Pb}/^{204}\text{Pb}\sim 19.01$ ) that formed the Cu-rich lens. They concluded that between the formation of the initial Zn-rich and subsequent more Cu-rich orebodies, hydrothermal circulation may have changed to reflect a different source rock or the fluid may have had a higher temperature increasing its ability to dissolve and carry copper. The differences between the Tazakourt and the Sidi M'Barek deposits are also recorded in the hydrothermal alteration of the host rocks, which is characterized by the dominance of chlorite in the footwall and sericite in the hanging wall of the Tazakourt orebody while the Sidi M'Barek orebody is characterized by a sericite-dominated alteration in both the footwall and the hanging wall (Belkabir *et al.* 2008). These differences led Belkabir *et al.* (2008) to interpret the Tazakourt and Sidi M'Barek orebodies as representing respectively the proximal and the distal volcanic environments of the Draa Sfar VMS deposit.

The commencement of mining at the Sidi M'Barek deposit in 2011 has resulted in increased accessibility to the deposit and its host rocks allowing new data to be obtained from underground and surface mapping and from drill cores. Presently the deposit is composed of a Zn-rich lens and a dozen of Cu-rich lenses, the latter are also rich in cobalt and gold. These new data for the deposit have significant implications on the genesis of the Draa Sfar deposit and other massive sulphide deposits of the Marrakech area. Indeed, these deposits are classified as either Zn-rich deposits (e.g. Tazakourt, Hajjar, Koudiat Aicha) or Cu-rich deposits (e.g. Kettara). There is an ongoing debate on the genesis of these deposits, with a syngenetic hypothesis that classifies the deposits as VMS deposits formed at or below the seafloor in an anoxic basin (Belkabir *et al.* 2008; Marcoux *et al.* 2008; Moreno *et al.* 2008) and an epigenetic hypothesis implying that the deposits formed during regional deformation and metamorphism (Essaifi & Hibti 2008; N'diaye *et al.* 2016). In addition, a multistage evolution of the Moroccan massive sulphide deposits has recently been proposed (Essaifi *et al.* 2019).

In order to better understand the relationship between the Zn-rich and the Cu-Au-rich mineralization, deformation and metamorphism, this paper compares the structures, textures, mineralogy and geochemistry of the Zn-rich and Cu-rich lenses of the Sidi M'Barek orebody, and discusses the temporal relationships between each type of mineralization, alteration, metamorphism and

deformation. The intent is to establish the timing and origin of each mineralization relative to the regional tectono-sedimentary evolution of the Jebilet massif.

### **Geological framework**

The Draa Sfar Cluster is located 15 km to the northwest of Marrakech, along the southern margin of the Jebilet massif, in the southern part of the Moroccan Meseta. The Jebilet massif provides a well-exposed cross-section across the western Meseta (Huvelin 1977; Essaifi *et al.* 2001), which is separated from the eastern Meseta by the middle Atlas (Fig. 1a). The massif is divided in three structural units (Fig. 1b): the Central Jebilet, the Western Jebilet and the Eastern Jebilet.

The Central Jebilet unit comprises sedimentary rocks of Visean age (Huvelin 1977), which have been metamorphosed to low grade and extensively deformed in the Variscan orogeny (Huvelin 1977; Essaifi *et al.* 2001; Delchini *et al.* 2016). The Western Jebilet is separated from the Central Jebilet by a dextral thrust-wrench fault (Mayol & Muller 1985; Le Corre & Bouloton 1987), part of the Western Meseta Shear Zone (Piqué *et al.* 1980). The Western Jebilet unit is part of the Coastal Block, and is composed of Cambro-Ordovician sedimentary rocks that were uplifted in the Late Devonian and only weakly deformed in the Variscan (Huvelin 1977; Piqué & Michard 1989). The Eastern Jebilet is separated from the Central Jebilet by the Marrakech Shear Zone (Lagarde & Choukroune 1982; Essaifi *et al.* 2001). The Eastern Jebilet unit contains a Carboniferous sedimentary sequence, the 'Kharrouba Flysch' (Huvelin 1977) which grades from turbidites to tidally-influenced shallow marine deposits (Graham 1982a, b; Beauchamp 1984; Beauchamp & Izart 1987). The Eastern Jebilet is overlain by tectonically-emplaced nappes of Ordovician to Early Carboniferous sedimentary rocks with a chaotic basal breccia (Graham 1982a, Izart *et al.* 1997). These nappes mark the onset of the closure of the Mesetian basin (Lagarde 1989; Michard *et al.* 2010; Hoepffner *et al.* 2005). Westphalian–Permian continental conglomerates (Huvelin 1977) rest unconformably upon the Hercynian folded sequence in the Eastern and Western Jebilet units.

The Central Jebilet unit, which hosts the Draa Sfar deposit (see Fig. 2), is a thick (>1000 m) largely low-grade metamorphosed Upper Visean metasedimentary sequence called the Sarhleif schists (Huvelin 1977). The Sarhleif schists are dominated by shales with subsidiary limestone and sandstone beds deposited in an anoxic platform, with the interbedded sandstones representing storm deposits (Beauchamp *et al.* 1991; Beauchamp & Izart, 1987). The Sarhleif schists host a bimodal magmatism and have been described as a metamorphosed volcano-sedimentary sequence, with shales interbedded with volcanic and volcanoclastic units (Huvelin 1977; Bordonaro 1983). In fact, the majority of the igneous rocks in the Central Jebilet are intrusive rocks with only minor rhyolitic and rhyodacitic lavas and volcanoclastic units (Aarab & Beauchamp 1987; Essaifi *et al.* 2014). The most significant area of extrusive volcanic rocks forms a rhyolitic dome near the Draa Sfar deposit, on the southern edge of the Central Jebilet exposure (Belkabir *et al.* 2008).

The Variscan magmatism in Central Jebilet can be divided into two groups. The first group is a bimodal association of numerous mafic-ultramafic layered intrusions, sills and dykes with alkaline microgranitic sheet-like intrusions, dykes and domes (Essaifi *et al.* 2014). A microgranitic intrusion (Koudiat Bouzlaf) was dated to  $330.5 \pm 0.68/-0.83$  Ma (U–Pb on zircon; Essaifi *et al.* 2003) while the rhyolitic dome of Draa Sfar gave an age of  $331.7 \pm 7.9$  Ma (Ar–Ar on white mica; Marcoux *et al.* 2008), and more recently a sample from the Koudiat Hamra intrusion has been dated to  $345 \pm 2$  Ma (SHRIMP U/Pb dating; Delchini *et al.* 2018). The second group of intrusions comprises two cordierite-bearing granodioritic plutons, crosscut by leucogranitic sheets and stocks. The granodioritic plutons have been dated to  $327 \pm 4$  Ma and the leucogranites to  $295 \pm 15$  Ma (whole-rock Rb–Sr ages; Mrini *et al.* 1992), although more recent U–Pb SHRIMP dating has suggested somewhat older ages of 358–336 Ma for the granodiorites (Delchini *et al.* 2018). In summary, a sustained and long-lasting magmatic activity occurred in the Jebilet massif between  $358 \pm 7$  Ma and ca. 300 Ma (Mrini *et al.* 1992; Essaifi *et al.* 2003; Delchini *et al.* 2018). The felsic magmatism is characterized by a temporal evolution from magmas of a hybrid origin to magmas of a pure crustal origin reflecting a lithosphere that is heated from below (Essaifi *et al.* 2014).

The NNE-trending dextral thrust-wrench West Moroccan Shear Zone and the NW-trending sinistral Marrakech Shear Zone (Figs 1–2) delimit a triangular block (Central Jebilet) where the Sarhlef schists have undergone low-grade metamorphism and deformation during the Variscan orogeny (Huvelin 1977; Bordonaro 1983; Essaifi *et al.* 2001; Delchini *et al.* 2018). The main regional schistosity ( $S_1$ ) typically trends NNE–SSW and is axial planar to a series of open upright  $F_1$  folds (Bordonaro 1983; Essaifi *et al.* 2001). This schistosity curves into a series of typically N–S trending sinistral shear zones, interconnected by E–NE dextral shear zones (Essaifi *et al.* 2001; N’diaye *et al.* 2016). These minor shear zones are shifted towards the Marrakech Shear Zone, which is the largest shear zone in Central Jebilet. The shear zones are associated with an intensification of foliations, tightening of folds and development of stretching lineations, which can be considered as related to a  $D_2$  phase (Lagarde & Choukroune 1982; Essaifi & Hibti 2008). Brittle shear zones are superimposed on the ductile structures and can be ascribed to a  $D_3$  phase. Overall, these structures reflect coeval transpression, vertical extrusion and south-directed lateral escape that were facilitated by the bounding conjugate WMSZ and MSZ (Essaifi *et al.* 2003). Absolute and relative timing constraints show that transpression was sustained from ca. 330 Ma to 300 Ma (Huon *et al.* 1987; Mrini *et al.* 1992; Essaifi *et al.* 2003; Delchini *et al.* 2018).

An important feature of Central Jebilet is the presence of numerous polymetallic massive sulphide deposits (Fig. 2). These deposits are aligned along three major lineaments trending N–S to NE–SW (Bernard *et al.* 1988; Essaifi & Hibti 2008; Maacha *et al.* 2017). They are typically pyrrhotite-dominated, vertical or steeply dipping, and most of them have been classified as VHMS (Bernard *et al.* 1988; Zouhry 1998; Belkabir *et al.* 2008; Marcoux *et al.* 2008; Moreno *et al.* 2008; Lotfi *et al.*

2008). They are affected by Variscan deformation (Bernard *et al.* 1988; Marcoux *et al.* 2008; N'diaye *et al.* 2016; Admou *et al.* 2018) and show a polyphase mineralization (Lusty *et al.* 2015; Essaifi *et al.*, 2019). According to the geochemical classification of VHMS deposits (Franklin *et al.* 1981; Large 1992), the major massive sulphide deposits of Central Jebilet belong either to the Pb-Zn-Cu type (Tazakourt, Koudiat Aicha) or the Cu-type (Kettara, Sidi M'Barek) (Essaifi *et al.* 2019).

## Methods

This study is based on lithological and structural mapping of the subsurface exploitation levels (underground galleries at -35, -75, -90, -110, -130 m) and logging of drill cores. Samples were collected from both the mine levels and the drill cores and used for petrographical, mineralogical and geochemical studies. Ore mineralogy and textural relationships were studied by transmitted and reflected light microscopy in 70 thin and polished sections, and complemented by Scanning Electron Microscopy (SEM) at Cadi Ayyad University and the Reminex Research Centre (Marrakech).

The chemical compositions of the sulphides were determined at Service CAMPARIS CNRS-UPMC-IPGP Campus Jussieu (Paris, France) using a Cameca SX-100 Electron Probe Microanalyzer (EPMA) equipped with five wavelength-dispersive spectrometers, under a 1µm diameter beam with a voltage of 15 kV, a current of 100 nA, and counting time adjusted to 60 s.

Whole-rock geochemical data were determined at the Reminex Research Centre by ICP-AES method using the same procedures as those described by Oummouch *et al.* (2017). Sulphur isotopes of pyrrhotite, chalcopyrite and pyrite for the Sidi M'Barek samples have been determined by Essaifi *et al.* (2019) while lead isotopes of sulphide samples have been determined by Marcoux *et al.* (2008).

## Geology of the deposit

### *Morphology of the mineralized lenses*

The metasedimentary lithologies in the Draa Sfar district include carbonaceous metapelites to the west and sandy metapelites to the east (Fig. 3). These formations are steeply east-dipping and display a layer-parallel  $S_1$  foliation. In the less deformed metasedimentary rocks, structural relationships indicate that the strata are overturned and younger to the west. The southern part of the district is characterized by the presence, within the sandy metapelites, of a rhyodacitic dome and associated volcanoclastic rocks. These subvolcanic lithologies are folded about a north-trending F1 anticline that places the Draa Sfar district on the west limb of this overturned fold (Belkabir *et al.* 2008; Rziki 2012). A mafic intrusion composed of coarse-grained to porphyritic gabbros occurs in the carbonaceous metapelites in the southern part of the district.

The Draa Sfar Cluster includes a southern deposit called Tazakourt and a northern deposit called Sidi M'Barek, which are separated by the Tensift River. The Tazakourt deposit consists of one Zn-rich lens that has an average thickness of 20 m while the Sidi M'Barek deposit consists of a dozen of small lenses including one Zn-rich lens and several Cu-rich lenses (Figs 3-6). The estimated resources of the

Zn-rich lens are 19 Mt of ore grading 3.8 % Zn, 1.2 % Pb, and 0.35 % Cu; 6 Mt has already been extracted since 2004. The total resources of the Cu-rich lenses of Sidi M'Barek are estimated to 1 Mt of ore grading 0.2 % Zn, 0.25 % Pb and 2% Cu.

The Sidi M'Barek orebody has a well-developed gossan extending to a depth of 20 m. The mineralized lenses reach 400 m in length and 10 m in thickness. These lenses are presently known until a depth of 800 m. The mineralized lenses are hosted by a metasedimentary lithostratigraphic succession of overturned, steeply east-dipping formations composed of carbonaceous metapelites to the west and sandy metapelites containing siltstone and bioclastic beds to the east (Figs 3-6). These sedimentary formations are crosscut by numerous mafic dykes, 0.5 to 3 m thick. The carbonaceous metapelites are characterized by the presence of millimeter thick beds and veins of calcite within a fine-grained schistose assemblage of muscovite, quartz, calcite, chlorite, titanite and leucoxene. The composition of the carbonaceous metapelites is richer in CaO (generally > 1%) than the sandy metapelites (<1%). Most of the Cu-rich lenses are located within the sandy metapelites while the Zn-rich lens is located to the west of the Cu-rich lenses, near the boundary between the sandy metapelites and the carbonaceous metapelites. The carbonaceous metapelites represent a distinct stratigraphic marker at the district of the Draa Sfar. They are located to the top of the Zn-rich lens in both the Koudiat Tazakourt and Sidi M'Barek deposits and thus represent the hanging wall of the Zn-rich massive sulphide lens while the sandy metapelites represent its footwall lithofacies (Belkabir *et al.* 2008; Rziki 2012).

The Zn-rich massive sulphides of Sidi M'Barek form a 400 m-long NNE striking lens that is dismembered by cross-faults into several smaller lenses (Fig. 4). This lens is steeply dipping and has an average thickness of 5 m (Fig. 6). It is composed of massive pyrrhotite containing disseminations of sphalerite, chalcopyrite, galena and arsenopyrite. The Zn-rich lens is located in the northern prolongation of the Tazakourt Zn-rich lens and there is no structural discontinuity between the two lenses (Fig. 3). Like the Tazakourt orebody, the Zn-rich lens is located at the contact between the carbonaceous and the sandstone metapelites and has a sheet-like morphology. Hence it was interpreted to represent the continuation of the Tazakourt orebody (Belkabir *et al.* 2008; Marcoux *et al.* 2008). The Zn-rich lens pinches out 400 m to the north of the Tensift River and exists only as disseminated sulphides within the host schists northwards.

The Cu-rich massive sulphides form several N-S to NNE subvertical small lenses (<150 m long and < 10 m-thick). The number of Cu-rich lenses increases from one principal lens in the southern part to several lenses in the northern part of the deposit (Figs 4-6). The Cu-rich lenses are composed of massive pyrrhotite containing disseminations of chalcopyrite or banded ore showing alternation of chalcopyrite bands with pyrite- and/or pyrrhotite rich-bands. Recent drill cores intercepted Cu-rich lenses to the south of the Tensift River and confirm the structural continuity between the Tazakourt

and the Sidi M'Barek ore deposits. The prolongation of the Cu-rich lenses further north of the explored area is yet to be delineated.

### ***Deformation of the mineralized lenses***

The most prominent structural feature in the host rocks of the Sidi M'Barek deposit is a penetrative N-S trending (N175 to N25°), steeply east-dipping (70-85°) regional S<sub>1</sub> cleavage. Deformation is very heterogeneous and strain intensity variations are recorded along the drill cores where centimeter to meter-scale high strain zones contrast with low strain domains. These strain variations are clearly illustrated in the carbonate beds where highly deformed zones contain echinoderm fossils with boudinage structures and well-developed pressure shadows contrast with low strain zones where fossils are almost spherical without any deformation. The highly deformed zones form anastomosed shear bands that bound lenticular domains of weakly deformed rocks. This structural pattern is similar to that described at different scales in other Moroccan massive sulphide deposits (Admou *et al.* 2018; N'diaye *et al.* 2016) and at the scale of Central Jebilet (Essaifi *et al.* 2001) where anastomosed shear zones bound weakly deformed domains. Locally, the regional S<sub>1</sub> cleavage bears a gently plunging stretching lineation. The Sidi M'Barek deposit is also characterized by the presence of an S<sub>2</sub> crenulation cleavage trending E-W and superimposed on the S<sub>1</sub> penetrative cleavage (Fig. 4).

Numerous brittle-ductile and brittle shear zones are well developed in the Sidi M'Barek deposit. The brittle behavior is indicated by striations along fault planes and by the occurrence of numerous veins. The major faults are N160° to N-S trending dextral strike-slip faults that show striations gently plunging to the north. These faults offset the Zn-rich lens and bound the Cu-rich lenses. The other brittle faults can be grouped into NW dextral faults and NE trending sinistral faults that cut across the mineralized lenses (Fig. 4). These faults are respectively synthetic and antithetic to the major N160° to N-S dextral faults.

The Zn-rich massive sulphides of Sidi M'Barek are highly deformed and are affected by ductile shearing and P<sub>2</sub> folds (Fig. 7b). They include elliptical fragments of the host rocks, which are elongated concurrently to the regional S<sub>1</sub> cleavage (Fig. 7b, g). Centimeter to millimeter scale elliptical nodules of pyrrhotite are common in the host rocks at the boundaries of the zincian lens, especially at its footwall. The thickness of the lens varies along strike and depth from a few centimeters to 10 m. In the zones where the Zn-rich lens pinches out, the sulphides remain as disseminated minerals within the host metapelites. This thickness variation is similar to that described in the Tazakourt orebody, and related to a "swell and pinch geometry" by Belkabir *et al.* (2008) and Marcoux *et al.* (2008).

Ductile deformation in the Zn-rich lens is marked by flattening of the sulphides and associated minerals along a well-defined S<sub>1</sub> cleavage. Asymmetric pressure shadows are developed around sphalerite and arsenopyrite grains (Fig. 7c-f) while chalcopyrite, galena and pyrrhotite are elongated

along the  $S_1$  cleavage (Fig. 7g, h). The schistose fabric is also indicated by the elliptical shape of the non-sulphide minerals and the inclusions of the host rocks (Fig. 7g). The deformation has led to the formation of a mylonitic fabric with development of C and C' shearing planes and  $P_2$  microfolds. A faint but recognizable compositional banding marked by variations in the proportions of sulphide minerals characterizes the Zn-rich lens (Fig. 7a). Although it resembles primary bedding, this mineral banding with development of centimeter scale sphalerite-rich bands alternating with pyrrhotite-rich bands is unlikely to be syngenetic in origin because the layering is parallel to the schistose fabric and primary textures are not observed in the orebodies. Within the pyrrhotite-rich bands, pyrrhotite is fine-grained (mm in size), equigranular and displays polygonal textures with  $120^\circ$  triple junctions that are indicative of recrystallization under low-grade metamorphic conditions (Gilligan & Marshall 1987). These features are similar to those described in the Tazakourt orebody and inferred to be tectonic in origin (Marcoux *et al.* 2008).

The Cu-rich lenses present sigmoidal shapes and sharp contacts with the host rocks. In the extremities of the ore lenses, the sulphides form thin layers that can be followed for more than 100 meters along strike (Fig. 4). Even though these sulphide layers are parallel to the  $S_1$  cleavage, they are not boudinaged. These extremities are marked by a quartz-sulphide association that evolves to disseminated sulphides within the host metapelites (Fig. 8a). They display centimeter-scale lozenge-shaped areas that correspond to right lateral pull-apart (extensional relay) traps where the sulphides are concentrated (Fig. 10c). The superposition of sigmoidal pull-apart formation on earlier shearing deformational event has been clearly observed in these areas. A similar pattern occurs at a larger scale where the central parts of the ore lenses are not affected by ductile deformation and the inclusions of the host rocks within the sulphides are equant and randomly oriented. A subvertical mineral banding marked by alternating bands of chalcopyrite and pyrrhotite is commonly present near the boundaries of the lenses (Fig. 8b). This banding is curvilinear and locally discordant on the  $S_1$  cleavage in the host rocks. As the lenses are bounded by faults, their general sigmoidal shape can be interpreted as a pull-apart trap indicating emplacement between 'en echelon' N-S to NNW dextral faults.

The sulphide textures in the Cu-rich lenses are dominated by open space filling textures. The central parts of the lenses are characterized by massive pyrrhotite with disseminated chalcopyrite, while the outer parts of the lenses present a "mineral banding" with alternation of centimeter-scale chalcopyrite-rich bands with pyrite- and/or pyrrhotite-rich bands. In contrast with the Zn-rich lens, this mineral banding is not related to a tectonic banding because the pyrite and chalcopyrite mineralization lack features of ductile-deformation. Pyrite fills veins that crosscut pyrrhotite while pyrite is in turn crosscut by chalcopyrite (Fig. 8e, f). Chalcopyrite includes arsenopyrite and sphalerite grains and host rock fragments. The host rock inclusions within chalcopyrite are equant and pressure shadows are absent around arsenopyrite or sphalerite grains (Fig. 8g). This indicates that the pyrite- and chalcopyrite-rich bands escaped the ductile deformation recorded in the pyrrhotite-rich bands. Thus

pyrite and chalcopyrite bands represent open-space fillings with crystallization in fluid-filled cavities (Robert & Brown 1986; Touringy *et al.* 1993), most likely by a crack-seal mechanism.

Pyrrhotite in the Cu-rich lenses is typically recrystallized, coarse-grained (1-5 mm long) and displays an elongation most likely representing a tectonic cleavage (Fig. 8c, d). Locally late-stage twins and fractures are also developed in pyrrhotite. The pyrite-bearing veins are almost perpendicular to this tectonic cleavage and they are filled by pyrite along the vein boundaries while carbonates are located in the center of the veins (Fig. 8e). A replacement of pyrrhotite by pyrite occurs along the vein boundaries and progresses along the tectonic cleavage, which indicates that the dilation direction was almost orthogonal to the vein boundaries. The relationship between deformation and pyrite deposition is clearly illustrated at the extremities of the mineralized lenses where the sulphide minerals are disseminated within the host rocks. Here pyrite shows a syn-tectonic growth indicated by stretching along the  $S_1$  cleavage and crystallization in pressure shadows around fossil clasts (Fig. 8h). Chalcopyrite develops as patches and veins invading both pyrite and pyrrhotite, and is crosscut by later veins of carbonates (Fig. 8f). The chalcopyrite veins are mostly parallel to the tectonic cleavage indicated by the elongation of pyrrhotite grains and by the  $S_1$  cleavage in the host rocks. These features indicate coupled dilation and replacement of massive sulphide ores in the Cu-rich lenses.

To conclude three stages of sulphide mineralization can be distinguished in the Sidi M'Barek deposit (Fig. 9): (i) a pre- to syn-tectonic stage during which pyrrhotite precipitated together with small amounts of sphalerite, galena and arsenopyrite. The evidence for pre-tectonic sulphides is indicated by the development of pressure shadows around sphalerite and arsenopyrite grains in the Zn-rich lens and disseminated ore; (ii) a syn-tectonic to late-tectonic pyrite-dominated stage during which pyrite precipitated along the  $S_1$  cleavage plane, in pressure shadows around competent objects and in syn- to late-tectonic veinlets; (iii) a late- to post-tectonic stage during which chalcopyrite precipitated in veins crosscutting both pyrrhotite and pyrite-rich paragenesis.

Pyrrhotite in the Cu-rich lenses is deformed and stretched along the  $S_1$  cleavage. Pyrite fills veins that crosscut pyrrhotite in the massive ore, and crystallizes in pressure shadows and along the  $S_1$  cleavage in the disseminated ore. The pyrite veins are in turn crosscut by chalcopyrite forming veins that are parallel to the  $S_1$  cleavage. Such a pattern indicates that the chalcopyrite parallel veins opened during a subsequent tectonic event with different stress directions as when the pyrite veins formed (Fig. 10a-c). Hence the cleavage serves as a receptacle for chalcopyrite concentration in a late deformation event that causes dilation parallel to a pre-existing cleavage (Chauvet *et al.* 2019). The evidence clearly supports the pyrrhotite-rich mineralization as being overprinted and further remobilized by syn- to late-tectonic pyrite- and chalcopyrite-rich mineralization.

### **Mineralogy of the sulphides**

The massive sulphides are principally composed of pyrrhotite, sphalerite, galena and chalcopyrite, with lesser amounts of magnetite, chlorite, sericite and quartz. At the hand specimen-scale there is a marked variation in texture and mineralogy between the Cu-rich and the Zn-rich ores.

### ***The Zn-rich massive sulphide mineralization***

The massive ore within the Zn-rich lens is composed of pyrrhotite together with lesser amounts of sphalerite, galena, chalcopyrite and arsenopyrite, which form a faint layering marked by alternation of pyrrhotite-rich and sphalerite-rich bands. Pyrrhotite is generally present as grains <1 mm long stretched along the  $S_1$  fabric. Grain boundaries with the other sulphide minerals are often irregular. Recrystallized equant grains of pyrrhotite showing triple junctions are locally present. Sphalerite occurs as subhedral or elliptical grains typically <0.5mm long. Galena is present as anhedral grains roughly 0.2 mm in length showing a spongy texture (i.e. full of micrometer-scale inclusions of other sulphides); the grains are aligned along the  $S_1$  fabric and associated to sphalerite and pyrrhotite. Arsenopyrite occurs as individual euhedral crystals 0.1 to 0.2 mm in diameter. Chalcopyrite is present as elongate anhedral grains < 0.4mm long.

With contents up to 90 vol.%, pyrrhotite is the dominant sulphide in the Zn-rich lens (Fig. 7g, h; Fig. 11a, b). Its zinc content is high (300-6600 ppm) and copper is highly variable (average 700 ppm); the higher values can be due to minute inclusions of chalcopyrite (Table 1). Arsenic and cobalt have not been detected while tin and nickel contents are low (<100 ppm owing to EPMA analyses). Pentlandite has not been observed. Its  $X_{FeS} = 0.92$  and its formula is close to  $Fe_7S_8$ . Three trace elements have been detected in pyrrhotite with highly variable concentrations of cadmium (average 400 ppm), bismuth (average 200 ppm), and antimony (average 100 ppm).

Sphalerite is the second dominant sulphide in the Zn-rich lens where it is concentrated along mm-scale bands (Fig. 11b). Its iron content ranges from 5.77 to 7.09 wt% (=  $X_{FeS}$  10-12 mole%). Copper and cadmium contents are high, 600-800 ppm and 1000-1500 ppm respectively; some very high occurrences of copper values (4500 ppm) could be due to minute inclusions of chalcopyrite. Nickel and cobalt are low (<100 ppm) as well as arsenic (<300 ppm). Bismuth is highly variable (mean 400 ppm); the highest values (up to 1800 ppm) may be due to minute inclusions of bismuth.

Galena is less abundant than sphalerite to which it is generally associated (Fig. 7g; Fig. 11c). It is rich in zinc (1.35-1.45 wt%) and iron (1.26-1.57 wt %) and contains traces of copper (100-700 ppm), arsenic (600-700 ppm) and cadmium (800-1000 ppm).

Less abundant than sphalerite, chalcopyrite occurs as disseminated grains within pyrrhotite and has a homogenous composition. It may be rich in zinc or contain traces of bismuth, tin, cadmium, and arsenic. The high contents of zinc (up to 4200 ppm) probably result from subsurface inclusions of sphalerite.

Pyrite is rare and occurs as isolated grains within a matrix of pyrrhotite, sphalerite and galena. Most commonly pyrite forms the margins of carbonate veins where it replaces pyrrhotite. Porphyroblasts of pyrite, up to 1 cm in diameter, displaying asymmetric pressure shadows were also observed (Fig. 10d). Its zinc and arsenic contents range from 500 to 600 ppm, and from 300 to 800 ppm, respectively.

Arsenopyrite is the most abundant minor sulphide in the Zn-rich lenses. It is enclosed within pyrrhotite and presents fragmented or corroded boundaries indicating a most likely early precipitation. As noted by Marcoux *et al.* (2008), independently from the grain size, arsenopyrite is As-rich (32-34 atom.% As). It contains variable amounts of zinc (mean 2300 ppm), cobalt (mean 300 ppm), nickel (200-3600 ppm) and antimony (mean 1700 ppm), low amounts of cadmium (mean 100 ppm) but bismuth has not been detected.

Stannite was observed in the Zn-rich lens where it occurs as fine clusters within sphalerite (Fig. 11d). Cassiterite was reported from the Zn-rich ore by Barrakad *et al.* (1977).

#### ***The Cu-rich massive sulphide mineralization***

Massive sulphides of the Cu-rich lenses contain essentially pyrrhotite, chalcopyrite, pyrite, sphalerite, arsenopyrite and cobaltite. Native bismuth, Bi-rich sulphides and selenides are present as individual crystals or complex intergrowths within chalcopyrite and pyrrhotite. The associated non-sulphide minerals are chlorite, quartz, muscovite and Fe-rich carbonates. Pyrrhotite occurs as elongate grains, 1-5mm long (Fig. 8c, d), exhibiting straight boundaries between crystals and 120° triple junctions. It includes inclusions of chlorite, sphalerite, cobaltite and arsenopyrite. When present, pyrite occurs within fractures and replaces pyrrhotite. The modal abundance of chalcopyrite reaches 35% contrasting with its low abundance in the Zn-rich lens (<5%). Chalcopyrite occurs as patches invading pyrrhotite or most commonly as regular mm- to cm-scale veins or fracture fillings in pyrrhotite and pyrite. Due to impinging growth of chalcopyrite into pyrrhotite and pyrite, inclusions of pyrrhotite and pyrite are common within and along the boundaries of chalcopyrite. Arsenopyrite is associated with chalcopyrite where it occurs as subhedral to irregular grains < 0.1 mm in diameter (Fig. 8g). Sphalerite occurs as irregular crystals (<1 mm) associated to arsenopyrite clusters or located at the boundaries of chalcopyrite.

Pyrrhotite is the dominant sulphide in the Cu-rich lenses although its proportion is lower and its size is larger than in the Zn-rich lens. Its  $X_{Fe}=0.92$  and its formula is close to  $Fe_7S_8$  (Table 1). It contains Cu (100-500 ppm) and Bi (mean 500 ppm) while its zinc content is low (mean 100 ppm). Arsenic and cobalt are present (mean 100 ppm) together with nickel (mean 300 ppm) and tin (mean 200 ppm). Cadmium and antimony contents are low (<200 ppm).

Sphalerite is a minor phase in the Cu-rich lenses where it is associated with pyrrhotite. Its iron content ranges from 6.37 to 7.29 wt.% (=  $X_{FeS}$  11-13 mole%). Copper and cadmium contents are very high,

ranging between 1800-8100 ppm and 1300-3900 ppm respectively. Bismuth is highly variable (mean 300 ppm) when compared with arsenic (mean 200 ppm) and cobalt (mean 300 ppm) while nickel is almost absent (<100 ppm).

In the Cu-rich lenses the galena is rare, it is fissured and always associated with carbonates.

Chalcopyrite is the most abundant sulphide after pyrrhotite and forms bands that cut across both pyrite and pyrrhotite (Fig. 10a-b; Fig. 11f). Chalcopyrite is the latest precipitated sulphide; its proportion can reach 20% and its composition is homogeneous. It is poor in zinc and nickel (<100 ppm) and contains variable amounts of bismuth (mean 500 ppm), arsenic mean 300 ppm) and tin (mean 200 ppm), and traces of antimony and cadmium.

Pyrite is abundant and forms either veins or fractures crosscutting pyrrhotite or centimeter-scale patches within pyrrhotite (Fig. 8e, f). In both cases, the relationship between the two minerals is clear: pyrite replaces pyrrhotite. The replacement of pyrrhotite by pyrite, involving a decrease in volume, was accompanied by the deposition of other minerals, especially magnetite and siderite. Pyrite contains variable amounts of nickel (100-1500 ppm), copper (mean 300ppm), bismuth, cadmium, and cobalt (mean 200 ppm).

Arsenopyrite in the Cu-lenses is generally associated with chalcopyrite. It is arsenic-rich (32-33 atom. % As) and contains high amounts of copper (250-9700 ppm) and variable amounts of cobalt (mean 200 ppm) and antimony (600-900 ppm). Zinc, nickel and cadmium contents are low (mean 100 ppm) and bismuth has not been detected.

Cobaltite is present in the Cu lenses as 0.025 to 0.25 mm wide crystals within and pyrrhotite (Fig. 11e) and as irregular crystals within chalcopyrite. It contains 2.36-2.48 wt% iron, up to 4% nickel and up to 900 ppm copper. Small inclusions of cassiterite were also observed within pyrrhotite; it is fissural and associated with carbonates.

Native gold, native bismuth, Bi-rich sulphides and selenides are abundant as individual crystals or complex intergrowths within the Cu-rich lenses. Some of the Bi-selenides and sulphides also contain silver and tellurium, and in rare places crystals up to c. 10  $\mu\text{m}$  of native electrum (Fig. 12). Native gold does not exceed 2.5  $\mu\text{m}$  in diameter and occurs as inclusions within pyrrhotite, chalcopyrite and cobaltite. Electrum is rarely intergranular but forms inclusions in all the sulphides of the Cu-rich lenses except pyrite (Fig. 12a-d). A Bi-Se association composed of small (15 to 30  $\mu\text{m}$ ) inclusions of laitakarite,  $\text{Bi}_4(\text{Se,S})_3$  and native bismuth within chalcopyrite was also described in some copper lenses by Marcoux *et al.* (2008) while small (20  $\mu\text{m}$ ) inclusions of paraganajuatite and bismuthinite ( $\text{Bi}_2\text{S}_3$ ), electrum and mackinawite ( $\text{FeS}$ ) were reported by Barrakad *et al.* (1977).

## Geochemistry

Among the drill cores intersecting the Zn-rich and Cu-rich lenses the drill core DF318 is representative of the mineralogy and textures of both ore types and allows comparison of the geochemistry of the Zn- and the Cu-rich lenses (Table 2). The drill core intersects the Zn-rich lens at a depth of 100 m and the Cu-rich lenses at 200 m and 240 m. The plots of Figure 13 show that elements such as lead, silver and arsenic are enriched in the Zn-rich lens (Fig. 13a, b, f), while cobalt and gold show higher concentrations in the Cu-rich lenses (Fig. 13e, g). The elevated cobalt contents in the Cu-rich lenses are attributed to the presence of cobaltite while the high contents of lead and arsenic are due to the presence of galena and arsenopyrite in the Zn-rich lens. The high gold contents in the Cu-rich lenses are mainly attributed to the presence of native gold and electrum while the high silver contents in the Zn-rich lens may be due to high contents of silver in galena as described in the Tazakourt orebody where silver contents in galena reach 0.13 wt% (Ben Aissi 2008). Inclusions of silver or silver-bearing minerals have not been observed within galena in the Draa Sfar district (Barrakad *et al.* 1977; Marcoux *et al.* 2008).

## Discussion

### *Relationships between the Zn-rich and the Cu-rich mineralization*

In VMS deposits, the contents of trace elements in pyrite reflect the composition of the zone in which it crystallizes (Huston *et al.* 1995): it is Cu-rich in the Cu-zone and Zn-rich in the Pb-Zn zone, a fact that has been related to zone-refining processes within the massive sulphides (Elderidge *et al.* 1983). In VMS deposits, cobalt content in pyrite is higher in Cu-rich zones than in Zn-rich zones, a fact that is related to high cobalt concentration in high-temperature hydrothermal fluids (Huston *et al.* 1995 and references therein). Huston *et al.* (1995) noted also that arsenic (and gold) content of pyrite increases from the Cu-rich to the Zn-rich zones of VMS deposits but this pattern disappears in deposits that have undergone metamorphic recrystallization because these elements are expelled from recrystallized pyrite. At Sidi M'Barek, pyrite is rich in copper and cobalt in the Cu-rich lenses, whereas it is rich in zinc and arsenic in the Zn-rich lens. Thus, the copper and cobalt contents of pyrite may be consistent with the interpretation of Marcoux *et al.* (2008) who suggested that the Cu-zone might represent for the Draa Sfar ore deposit the upper part of a stockwork zone that has been strongly transposed by deformation. However the Cu lenses extend to the North of the Zn-rich lens, which is dissimilar to the zoning commonly found in other VMS deposits, with Cu-rich zones towards the central and lower parts of the masses, barren pyrite in the core and a Zn+Pb enrichment towards the upper and lateral zones of the orebodies, (e.g., Large 1977; Lecolle 1977; Sato 1972; Sangster 1972). Moreover, a recent drill hole to the south of the Tensift River intercepted a Cu-rich zone within the Pb-Zn-rich lens.

The Cu-Au-Bi association and the high contents of gold in the Cu-lenses at Sidi M'Barek contrast with the Bathrust mining camp, especially the Brunswick No 6 deposit where the gold content is higher in the Zn-Pb zone than in the Cu-zone. In this deposit, the main geochemical and mineralogical

differences observed between the Pb-Zn zone and Cu zone were interpreted to have resulted from primary depositional features and from an almost coeval high-temperature zone refining process (MacLellan *et al.* 2006). Such an interpretation was supported by the  $\delta^{34}\text{S}$  values of the bulk sulphides, which are similar in the Pb-Zn zone and the basal Cu zone. This is not the case at Sidi M'Barek where the sulphur isotopic signature of pyrite ( $-2.5$  to  $-0.8$  ‰) is different from that of pyrrhotite ( $-6.4$  ‰ to  $-7.4$  ‰) and chalcopyrite ( $-6.1$  to  $-5.2$  ‰) indicating that distinct sources of sulphur were involved (Fig. 14a; Essaifi *et al.* 2019). In contrast the  $\delta^{34}\text{S}$  value of pyrrhotite from the Zn-rich lens ( $-7.2$  ‰) is similar to that of pyrrhotite from the Cu-rich lenses indicating that sulphur in pyrrhotite was derived from the same sulphur source, and suggests a local remobilization of pyrrhotite.

The greater proportion of copper, together with gold, and cobalt, and the lower contents of zinc, lead and silver could be consistent with the interpretation that the main control of the formation of the Cu-rich lenses was higher temperatures (Heinrich & Eadington, 1986). Such an interpretation is however not supported by the zonation of the hydrothermal alteration, which is marked by a lateral change from a predominantly chlorite alteration in the footwall of the Tazakourt orebody to sericite alteration in the Sidi M'Barek deposit (Belkabir *et al.* 2008). The differences observed between Cu-rich and Cu-poor zones could be best accounted for by differences in the Se/S ratios of the hydrothermal fluids. Fluids in Zn-rich zones have low Se/S ratios, typical of evolved seawater with minimal magmatic input, while fluids in the Cu-rich zones have high Se/S ratios, which is consistent with a significant magmatic component (Huston *et al.* 1995). The high Se/S ratios indicated by the abundance of the selenides in the Cu-rich lenses and the sulphur isotopic ratios of pyrite, which are close to 0 ‰, are consistent with magmatic derivation, either via exsolution of magmatic fluids or leaching of igneous rocks.

The thin layering developed within the Zn-rich lens by alternation of sphalerite-rich bands and pyrrhotite-rich bands may represent a primary depositional feature, but the absence of preserved primary textures within the orebodies (Moreno *et al.* 2008, Marcoux *et al.* 2008), and the fact that these bands are strictly parallel to the schistose fabric and developed in high strain zones indicate that they may have resulted from local segregation and redistribution of pyrrhotite during prograde metamorphism (cf. McQueen 1987). However, given the low ductility contrast between pyrrhotite, galena, sphalerite and chalcopyrite, the observed zonation and distribution of base and trace metals at Sidi M'Barek cannot be explained by mechanical remobilization only. Elongation of sphalerite, pyrrhotite, chalcopyrite, and galena along the  $S_1$  cleavage in the Zn-rich lens, and the formation of pressure shadows of quartz, muscovite, chalcopyrite, and pyrite around rigid objects in disseminated ore, are strong evidence for pressure solution in the orebodies and enclosing host rocks. Even if fluid-assisted remobilization was active, metamorphic mobilization followed by transport, focusing and precipitation in the emplacement site should be involved because the lead isotope signatures recorded

at Draa Sfar are less radiogenic in the Zn-rich lens than in the Cu-rich lenses and track the existence of lead from two sources (Fig. 14b). A less radiogenic source (with  $^{206}\text{Pb}/^{204}\text{Pb}$  close to 18.28) dominated in the Zn-rich lens, and a more radiogenic source ( $^{206}\text{Pb}/^{204}\text{Pb} \sim 19.01$ ) was predominant in the Cu-rich lenses (Marcoux *et al.* 2008). The most radiogenic lead isotope signatures are recorded in the syn- to late-tectonic pyrite veins crosscutting pyrrhotite at Sidi M'Barek, and mark a contribution of radiogenic lead during the syn-metamorphic mineralization, coeval with low-grade metamorphism (Essaifi *et al.* 2019), while the less radiogenic Pb isotopic values are represented by the Zn-rich ore of Tazakourt, and suggest a major contribution of volcanic rocks as source of metals (Marcoux *et al.* 2008; Bouabdellah *et al.* 2016). The other data points form a mixing trend suggesting that the advent of the Cu-mineralization resulted in a radiogenic Pb component being added to the isotopically less radiogenic Zn-rich mineralization. Hence the Pb isotopic data suggest a significant difference in age between the Zn-Pb mineralization and the introduction of Cu-Au. This is consistent with the superimposed paragenetic successions observed at Sidi M'Barek, which point to a polyphase mineralization rather than a single mineralizing event.

### ***Genetic model***

The difference between the deformational features of the Zn-rich lens and the Cu-rich lenses most likely reflects different times of emplacement. The morphology of the Zn-mineralization, which is composed of one lens located near the boundary between the carbonaceous and the sandy metapelites, suggests an initially stratiform orebody and hence a synsedimentary deposition of the Zn-mineralization (Marcoux *et al.* 2008). Moreover the Zn-rich lens in the Tazakourt deposit is hosted by volcanic and volcanoclastic rocks suggesting an association with a synsedimentary volcanic activity (Marcoux *et al.* 2008; Belkadir *et al.* 2008). However, the morphology of the lens, which is elongated for almost 2 km horizontally and at least 1500 m vertically, shows important variations in thickness suggesting that ductile deformation has accentuated stretching and boudinage of a sheet-like original morphology of the lens.

Well-developed tectono-metamorphic fabrics characterize the Zn-rich lens. It is a highly deformed lens showing mylonitic and recrystallization textures. In this lens most sulphides exhibit recrystallization and remobilization by pressure-solution. Pyrrhotite is stretched and flattened along the cleavage plane or exhibits annealing recrystallization textures. Grain flattening and elongation also characterize galena. Sphalerite is weakly flattened along the cleavage plane and pressure shadows are well developed around the grains. However strain-induced exsolution of chalcopyrite in sphalerite have not been observed. Arsenopyrite is present as euhedral grains showing well developed pressure shadows while chalcopyrite crystallizes in the pressure fringes. These features, which have been interpreted to result from fluid-state remobilization (Gilligan & Marshall 1987; Park 1996; Marshall & Spry 2000), indicate that these sulphide minerals are deformed and most likely their crystallization is pre-tectonic and pre-metamorphic. Hence the Zn-rich lens suffered considerable textural

modification during syn-metamorphic deformation. However, when present in the Zn-rich lens, pyrite is secondary and postdates the pre-tectonic assemblage.

Morphology of the Cu-mineralization, which is composed of numerous vertical veins more or less continuous, suits that of a lode rather than that of a stratiform orebody. The mineralized Cu-lenses show no mesoscopic deformation structures similar to those observed in the Zn-rich lens, which suggest that they have escaped the ductile deformation that affected their host rocks. These lenses are characterized by a sigmoidal-like pull apart shape (Figs 4 & 10c), and by dilation and replacement textures leading to formation of a mineral banding with chalcopyrite-rich bands alternating with pyrite- and/or pyrrhotite-rich bands. A well-defined schistose fabric is lacking in the pyrite- and chalcopyrite-rich bands while in the pyrrhotite-rich bands, pyrrhotite is coarse-grained and elongated suggesting that it underwent recrystallization and ductile deformation. The similarity of the  $\delta^{34}\text{S}$  values of pyrrhotite in the Cu-rich lenses and the Zn-rich lens indicates a common source of sulphur and suggests a remobilization of ductile sulphides at a larger scale (Gilligan & Marshall 1987; Marshall *et al.* 2000).

The pyrrhotite-rich bands are crosscut by pyrite veins, which in turn are crosscut by chalcopyrite. When it occurs as disseminated minerals within carbonaceous beds, pyrite crystallizes in the pressure shadows around echinoderm clasts or as flattened crystals along the cleavage plane (Fig. 8h). As pyrite is normally more competent than the carbonate matrix, the absence of pressure fringes around pyrite indicates that pyrite is not a pre-tectonic mineral. On the contrary, this suggests that pyrite crystallized following the cleavage direction. The presence of pyrite porphyroblasts within deformed pyrrhotite in the Zn-rich lens indicates that ductile deformation of pyrrhotite enhanced porphyroblastic growth of pyrite during greenschist metamorphism (Vokes 2000; Maclellan *et al.* 2006). Because the sulphur isotopic signature of pyrite is different from that of pyrrhotite, the growth of pyrite and the replacement of pyrrhotite by pyrite were assisted by an external source of sulphur. On the contrary, the similarity of the  $\delta^{34}\text{S}$  values of pyrrhotite and chalcopyrite suggests an in-situ remobilization of ductile sulphides. Indeed, chalcopyrite fills extensional relays or pull-aparts related to a late stage of brittle-ductile deformation (Fig. 10c).

The relationships between the main sulphides and ductile deformation allowed us to distinguish an early mineralizing stage that is dominated by pyrrhotite, together with chalcopyrite, sphalerite and galena, accompanied by variable amounts of chlorite, sericite and quartz. A syn-tectonic stage is dominated by pyrite and precipitation of carbonates whereas a late- to post-tectonic stage is dominated by chalcopyrite accompanied by quartz. The presence of asymmetric pressure shadows around arsenopyrite and sphalerite in disseminated ore and the widespread deformation and shearing of the Zn-rich lens indicates that the early mineralizing event is clearly pre- to syntectonic, and supports formation of the Zn-rich precursor mineralization at and below the seafloor in an anoxic basin, during a transtensional deformation event (Belkabir *et al.* 2008; Moreno *et al.* 2008; Marcoux *et*

*al.* 2008). The widespread replacement of pyrrhotite by pyrite and the crystallization of pyrite in the pressure shadows around rigid objects, together with the presence of pyrite porphyroblasts within deformed pyrrhotite, indicate that this mineralizing event is syn- to late-tectonic and contemporaneous with transpression, basin closure and metamorphism. This transpressional event is characterized by formation of N-S sinistral shear zones and associated NE-SW striking  $S_1$  cleavage under the effects of a NW-SE directed shortening (Fig. 15a), which has been previously recognized as the major post-Visean tectonometamorphic event at both the scale of Central Jebilet (Essaifi *et al.* 2001) and the Moroccan Meseta (Lagarde *et al.* 1990). During this event, the pre-existing Zn-rich lens and disseminated sulphides underwent ductile deformation, shearing and remobilization while an external sulphur source allowed deposition of pyrite within structurally controlled veins. A second event of compressional deformation is marked in the Draa Sfar district by an E-W striking  $S_2$  crenulation cleavage (Belkabir *et al.* 2008; Marcoux *et al.* 2008; Rziki 2012). During this N-S shortening event, dextral shearing re-used the previously formed shear zones and allowed sigmoidal extensional relays (pull aparts) to open along the pyrrhotite-rich lenses (Fig. 15b). During this stage chalcopyrite crystallized along parallel veins caused by dilation along the re-opened shear zones. The similarity of the sulphur isotopic signatures of chalcopyrite and pyrrhotite suggests an in situ remobilization of sulphur, while the lead isotopic signature of the Cu-rich lenses, which is significantly more radiogenic than that of the Zn-rich mineralization, points out to the metals being inherited from the hosting black shales. Taken together, these isotopic data would rather indicate interplay between remobilization and mobilization processes during formation of the Cu-rich lenses.

## Conclusion

On the basis of metal contents, the Koudiat Sidi M'Barek deposit is composed of a Zn-rich lens and Cu-rich lenses, and thus constitutes a representative example of the massive sulphide deposits of the Moroccan Meseta, which are composed of either Cu-rich or Zn-rich orebodies. The Zn-rich lens is enriched in silver, lead and arsenic, whereas in the Cu-zone, higher concentrations of cobalt and gold are exhibited. The Zn-rich mineralization forms a stratiform sheet-like lens that was dismembered by superimposed ductile and brittle Variscan deformation. This lens represents the northern prolongation of the Tazakourt deposit, which is associated with felsic volcanic and volcanoclastic rocks, and presents structural and textural features consistent with a VMS type precursor that was further deformed and remobilized by pressure solution during the post-Visean shortening. The Cu-rich mineralization forms several structurally controlled, foliation-parallel veins that display superimposed paragenetic successions and coupled dilation-replacement textures. A coarse grained pyrrhotite is crosscut and replaced by pyrite, which in turn is replaced by chalcopyrite. The Cu-rich mineralized bodies were derived from syn- to post-tectonic mobilization of dispersed metals in the country rocks, probably interacting with metamorphic remobilization of pre-tectonic, massive and disseminated mineralization. The weight of evidence clearly supports Sidi M'Barek as being remobilized,

pre-tectonic Zn-Pb systems that have been partially overprinted and further remobilized by syntectonic Cu-Au systems.

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## Figure Captions

**Fig. 1.** The geological map of the Jebilet massif showing the location of Draa Sfar, Kettara and Koudiat Aicha massive sulphide deposits (modified from Huvelin, 1977). Inset shows the Jebilet massif in the frame of the Variscan outcrops of Morocco (in brown). EM: Eastern Meseta, WM: Western Meseta.

**Fig. 2.** Geological setting of the Central Jebilet Unit showing the schistose pelite of Sarhlef series, the outcrops of the magmatic rocks and the general distribution of the massive sulphide ore deposits (modified after Huvelin, 1977). The white color represents undifferentiated post-Paleozoic sedimentary units. Location of figure 3 is indicated.

**Fig. 3.** Geological map of the Draa Sfar mine area showing the surface projection of Koudiat (Kt.) Tazakourt and Koudiat Sidi M'Barek massive sulphide ore bodies, the host rocks and the selected drill holes illustrated in figures 5, 6 and 13. Location of figure 4 is indicated.

**Fig. 4.** Structural map of the Kouidiat Sidi M'barek (underground mapping of level -75m supplemented by interpretation of drill cores). **(a)** stereonet showing equal area, lower hemisphere projection of the  $S_1$  cleavage. **(b)** orientation of the fault planes measured in the mine level -75m.

**Fig. 5.** East-west borehole sections at Kouidiat Sidi M'Barek, drawn from detailed logging of the DS170, DS173, DS175 and DS223 drill cores shown in figure 3.

**Fig. 6.** East-west vertical cross sections across the zinc- and copper rich-lenses.

**Fig. 7.** The main features of the zinc-rich lens illustrated by drill core photographs and photomicrographs in transmitted and reflected light. **(a)** Banded ore with alternation of pyrrhotite-rich and sphalerite-rich bands. **(b)** Massive pyrrhotite with folded inclusions of metapelites (HR). **(c-d)** Disseminated Zn-rich ore showing pyrrhotite (Po) elongated along the  $S_1$  cleavage while pressure shadows are developed around arsenopyrite (Apy); the main non-metallic minerals are chlorite (Chl), sericite (Ser) and quartz (Qtz); **(c)** reflected light, **(d)** transmitted light. **(e-f)** Pressure shadows developed around weakly flattened sphalerite (Sp); **(e)** reflected light, **(f)** transmitted light. **(g-h)** Zn-rich massive ore with the main sulphides elongated along the  $S_1$  cleavage (Ga: Galena, Ccp: chalcopyrite); reflected light, PPL **(g)** and XPL **(h)**.

**Fig. 8.** The main features of the copper-rich lenses illustrated by underground photographs and photomicrographs. **(a)** Photograph in the underground gallery showing unconformity between the  $S_1$  cleavage in the metapelites (HR) and the extremity of a copper-rich lens composed of quartz (Qtz) and disseminated pyrrhotite (Po) and chalcopyrite. **(b)** Sample showing alternation of pyrrhotite-rich and chalcopyrite-rich bands. **(c-d)** Coarse-grained massive pyrrhotite showing triple junctions. Reflected light, XPL. **(e)** Pyrite (Py) and carbonate (Cb) infilling veins crossing massive pyrrhotite. **(f)** Chalcopyrite (Ccp) crosscutting pyrrhotite, pyrite and magnetite (Mag), and crosscut by a late carbonate vein. **(g)** Dissemination of arsenopyrite (Apy) and sphalerite (Sp), together with inclusions of pyrrhotite (Po) and metapelites (HR) in the chalcopyrite-rich band. **(h)** Syntectonic pyrite crystallizing along the  $S_1$  cleavage marked by flattening of chlorite, sericite and quartz in a sample from the metapelites. Reflected light, PPL **(e-h)**.

**Fig. 9.** Paragenetic successions of the mineralization at Kouidiat Sidi M'barek.

**Fig. 10.** Photograph **(a)** and sketch **(b)** of the massive ore in the Cu-rich lenses showing the superimposition of pyrrhotite (Po), pyrite (Py) and chalcopyrite (Ccp). **(c)** Extensional relay filled by chalcopyrite and quartz. **(d)** Pyrite porphyroblast within pyrrhotite.

**Fig. 11.** Representative plane-polarised photomicrographs of the Zn-rich and the Cu-rich lenses. **(a)** Massive pyrrhotite with chalcopyrite and sphalerite in the Zn-rich lens. **(b)** Banding in the Zn-rich lens marked by pyrrhotite-rich and sphalerite-rich bands parallel to the  $S_1$  tectonic cleavage. **(c)** Pyrrhotite (Po), sphalerite (Sp), galena (Gn), Chalcopyrite (Ccp) and inclusions of metapeltes (HR) flattened along the  $S_1$  cleavage in the Zn-rich lens. **(d)** Inclusions of stannite within sphalerite in the Zn-rich lens. **(e)** Cobaltite within pyrrhotite in a Cu-rich lens. **(f)** Chalcopyrite replacing pyrrhotite in a Cu-rich lens.

**Fig. 12.** Scanning Electron Microscope photographs showing native gold, electrum and bismuth inclusions within sulphides in the Cu-rich lenses. Native gold within pyrrhotite **(a-b)**, cobaltite **(c)**, and chalcopyrite **(d)**. Electrum within chalcopyrite **(e)**, pyrrhotite **(f)**, sphalerite **(g)**, arsenopyrite **(h)**, cobaltite **(i)**, and between pyrrhotite and cobaltite **(j)**. po: pyrrhotite, Ccp: chalcopyrite, Apy: arsenopyrite.

**Fig. 13.** Zn, Pb, Cu, Ag, Co, As and Fe element distribution in the Zn-rich lens and the copper-rich lens in the DF 318 drill core, showing high contents of Zn, Pb, Ag and As in the Zn-rich lens and low contents in the Cu-rich lenses **(a, b, d, f)**. Cu, Co and Au contents are high in the Cu-rich and low in the Zn-rich lens **(c, e, g)**.

**Figure 14.** Lead and sulphur isotopic compositions of the Draa Sfar ores (data from Marcoux et al. 2008; Essaifi *et al.* 2019). The lead isotopic compositions of the IPB VMS deposits (Blue field; Marcoux 1998) are also shown for comparison. S&K: average growth curve of Stacey and Kramers (1975), C&R: upper crust growth curve of Cumming and Richards (1975). Evolution curves for Upper Crust and Orogene (Zartman & Doe 1981) are shown for reference.

**Fig. 15.** Two stage tectonic model for the formation of the Cu-rich lenses of Sidi M'Barek.

(a) Stage 1, NW-SE shortening generating the  $S_1$  cleavage, remobilization of pyrrhotite-rich mineralization and deposition of pyrite within syntectonic veins associated to N-S left lateral shear zones.

(b) Stage 2, Reopening of the shear zones by right lateral shearing during a subsequent N-S shortening generating the  $S_2$  crenulation cleavage and deposition of chalcopyrite.

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**Table 1.** Electron microprobe analysis of sulphide minerals of Koudiat Sidi M'Barek

**Pyrrhotite**

	Zn-rich lens DF311-122	Zn-rich lens DF311-122	Zn-rich lens DF311-122	Zn-rich lens DF276-67	Zn-rich lens DF276-67	Cu-rich lens DS161-90.5	Cu-rich lens DS161-90.5	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py
S	39.84	39.34	39.34	39.58	39.71	39.18	39.43	39.13	39.59	39.65	39.55
Fe	58.27	57.30	57.75	59.51	59.55	58.70	58.60	58.45	58.02	58.08	57.86
As	-	-	-	-	-	0.02	-	0.02	-	-	0.03
Co	-	-	-	-	-	0.03	0.03	-	-	-	0.01
Cu	-	0.16	0.20	0.01	-	0.01	0.01	0.05	0.03	0.01	0.02
Zn	0.44	0.06	0.66	0.05	0.03	0.01	-	0.01	-	0.02	-
Bi	-	0.06	0.02	0.04	-	-	-	0.13	-	-	0.15
Ni	-	0.01	-	0.01	-	0.04	0.05	-	0.01	0.03	0.03
Cd	-	0.07	0.03	-	0.08	-	-	-	-	0.02	-
Sn	0.01	0.01	-	0.01	-	-	-	0.04	0.02	0.02	0.02
Sb	-	0.01	0.05	-	-	-	0.02	0.01	-	-	0.02
Total wt.%	98.56	97.02	98.05	99.21	99.36	97.99	98.13	97.84	97.66	97.82	97.69
S cat	54.20	54.35	53.93	53.64	53.71	53.72	53.92	53.79	54.29	54.29	54.30
Fe cat	45.50	45.44	45.45	46.30	46.24	46.20	46.00	46.12	45.68	45.66	45.60
XFeS	0.91	0.91	0.91	0.93	0.92	0.92	0.92	0.92	0.91	0.91	0.91

**Sphalerite**

	Zn-rich lens DF311-122	Zn-rich lens DF311-122	Zn-rich lens DF311-122	Zn-rich lens DF311-122	Zn-rich lens DF276-67	Zn-rich lens DF276-67	Zn-rich lens DF276-67	Cu-rich lens DS161-90.5	Cu-rich lens DS161-90.5	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py	Cu-rich lens L2-90-Cpy-Py
S	33.1	33.1	33.0	33.4	32.8	33.0	32.3	33.2	33.3	33.3	32.7	33.1	32.9
Fe	6.65	6.70	6.95	6.71	6.68	5.77	7.09	6.37	6.49	6.66	7.29	7.07	7.00
As	0.03	0.01	-	0.01	-	-	0.03	0.09	-	-	-	0.04	-
Co	-	-	-	0.01	-	-	0.01	0.04	0.05	0.02	-	0.01	0.03
Cu	0.06	0.07	0.08	0.06	0.07	0.07	0.45	0.56	0.66	0.25	0.81	0.18	0.30
Zn	59.5	59.6	59.7	59.8	60.6	61.0	59.2	59.2	59.7	60.1	58.5	59.6	59.3
Cd	0.10	0.10	0.13	0.13	0.12	0.15	0.11	0.39	0.29	0.13	0.15	0.22	0.15
Ni	0.01	-	-	-	-	-	0.01	-	0.01	-	-	-	-
Bi	-	0.18	0.08	-	-	-	-	-	-	-	0.12	0.04	0.02
Total wt.%	99.5	99.8	100.	100.	100.	100.	99.2	99.8	100.	100.	99.5	100.	99.7
mole %	8	8	03	25	34	03	6	9	53	51	8	37	8
FeS	11.5	11.5	12.0	11.5	11.5		12.3	11.0	11.1	11.4	12.6	12.1	12.1
	3	9	2	4	3	9.97	6	2	6	5	6	8	3

### Chalcopyrite

	Zn-rich lens	Cu-rich lens											
	DF311-122	DF311-122	DF311-122	DF311-122	DF276-67	DS161-90.5	DS161-90.5	DS170-E7	DS170-E7	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py
S	35.0	34.7	34.7	34.9	35.2	34.4	34.6	34.8	34.7	34.7	34.7	34.9	34.8
	1	5	0	0	5	3	3	7	7	4	4	1	8
Fe	28.8	29.1	28.9	28.8	29.5	29.1	29.6	29.0	29.5	29.4	29.3	29.4	29.4
	1	7	8	3	3	3	7	9	8	1	1	1	3
As	0.04	0.01	0.04	-	-	0.04	-	0.04	0.03	0.06	0.07	-	0.02
Co	-	-	-	-	-	-	-	-	-	-	-	-	-
Cu	34.6	34.8	34.7	34.8	34.9	34.0	34.9	34.9	34.9	34.8	34.9	34.9	35.0
	6	3	8	6	5	9	4	8	5	5	1	2	2
Zn	-	0.05	0.42	0.06	0.01	-	-	-	-	-	-	0.01	-
Bi	-	-	-	-	-	0.03	0.10	0.02	-	0.11	-	0.01	0.11
Cd	-	0.04	0.02	-	0.03	-	-	0.01	0.03	-	0.02	-	-
Sn	-	0.02	-	0.04	0.02	0.01	0.07	0.02	-	0.03	-	-	-
Total wt.%	98.5	98.8	98.9	98.6	99.8	97.7	99.4	99.0	99.3	99.1	99.0	99.2	99.4
	1	7	5	9	0	2	2	1	5	9	5	6	5

### Pyrite

	Zn-rich lens	Cu-rich lens								
	DF276-67	DS170-54.2	DS170-54.2	DS170-54.2	DS170-E7	DS170-E7	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py
S	54.55	52.27	51.92	52.05	53.52	52.80	51.93	51.33	52.19	52.63
Fe	46.18	45.67	45.12	45.13	45.73	46.09	45.89	45.76	45.72	44.89
As	0.03	-	-	0.03	0.02	-	-	-	-	-
Co	-	0.03	0.03	0.07	-	-	-	-	0.01	0.02
Cu	0.01	0.02	-	-	0.05	-	0.01	0.01	0.06	0.11
Zn	0.06	0.03	-	0.01	-	-	-	-	-	0.01
Bi	-	0.04	0.03	-	-	0.07	-	-	-	0.08
Ni	-	0.12	0.11	0.15	0.02	0.03	0.02	0.04	0.01	0.04
Cd	-	-	-	0.06	0.03	0.03	-	-	0.03	-
Sn	-	-	0.01	0.02	-	-	0.02	0.02	0.01	0.03
total	100.83	98.18	97.22	97.52	99.36	99.02	97.87	97.16	98.04	97.80

### Arsenopyrite

	Zn-rich lens	Cu-rich lens	Cu-rich lens	Cu-rich lens	Cu-rich lens					
	DF311-122	DF311-122	DF311-122	DF311-122	DF276-67	DF276-67	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py	L2-90-Cpy-Py
S	19.98	20.02	19.92	19.68	20.04	20.14	19.58	19.24	19.50	19.85
Fe	33.27	33.37	33.47	33.68	34.31	34.22	34.28	34.19	33.74	33.71
As	44.45	44.25	44.45	44.58	44.27	43.81	44.71	43.80	45.01	44.36

Co	0.02	0.03	0.05	0.06	-	-	-	-	-	0.07
Cu	-	-	0.01	0.04	-	-	0.30	0.25	0.97	0.80
Zn	0.10	-	0.02	-	0.56	0.69	0.01	-	0.01	0.02
Ni	0.21	0.36	0.04	0.18	0.02	0.03	-	-	0.02	0.01
Cd	0.01	0.02	-	0.01	-	0.01	-	-	-	0.05
Sb	0.04	-	0.02	0.14	0.60	0.24	0.07	0.07	0.06	0.09
total	98.08	98.04	97.97	98.37	99.82	99.13	98.94	97.55	99.32	98.97
As cat	32.63	32.46	32.67	32.73	32.04	31.80	32.67	32.44	32.84	32.34

### **Cobaltite**

	Cu-rich lens DS161- 90.5	Cu-rich lens DS161- 90.5	Cu-rich lens DS161- 90.5
S	19.47	20.14	19.24
Fe	2.48	2.42	2.36
As	44.08	44.10	44.64
Co	33.23	29.83	32.87
Cu	-	0.09	0.02
Zn	-	-	-
Bi	-	-	-
Ni	-	4.09	0.50
Total	99.27	100.67	99.62
As cat	32.61	32.05	33.02

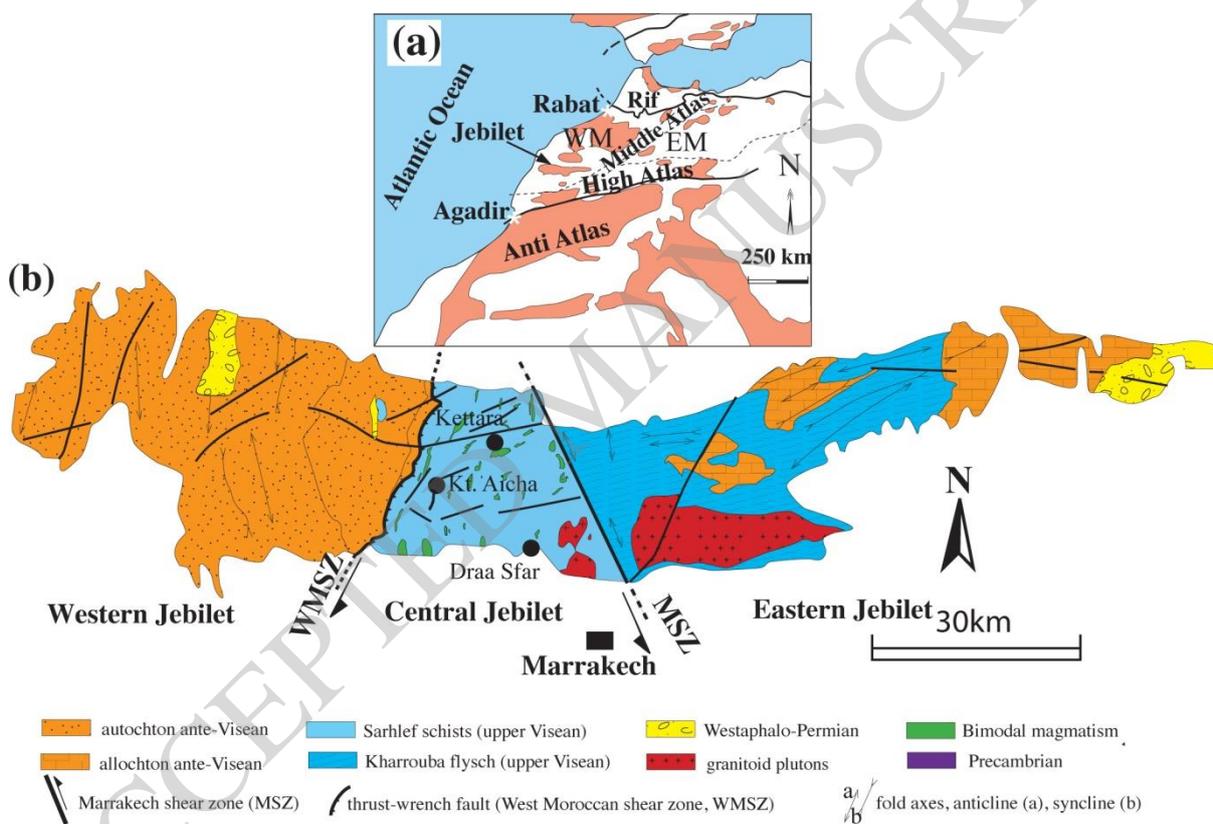
**Table 2.** Atomic absorption analysis of drill core DF318

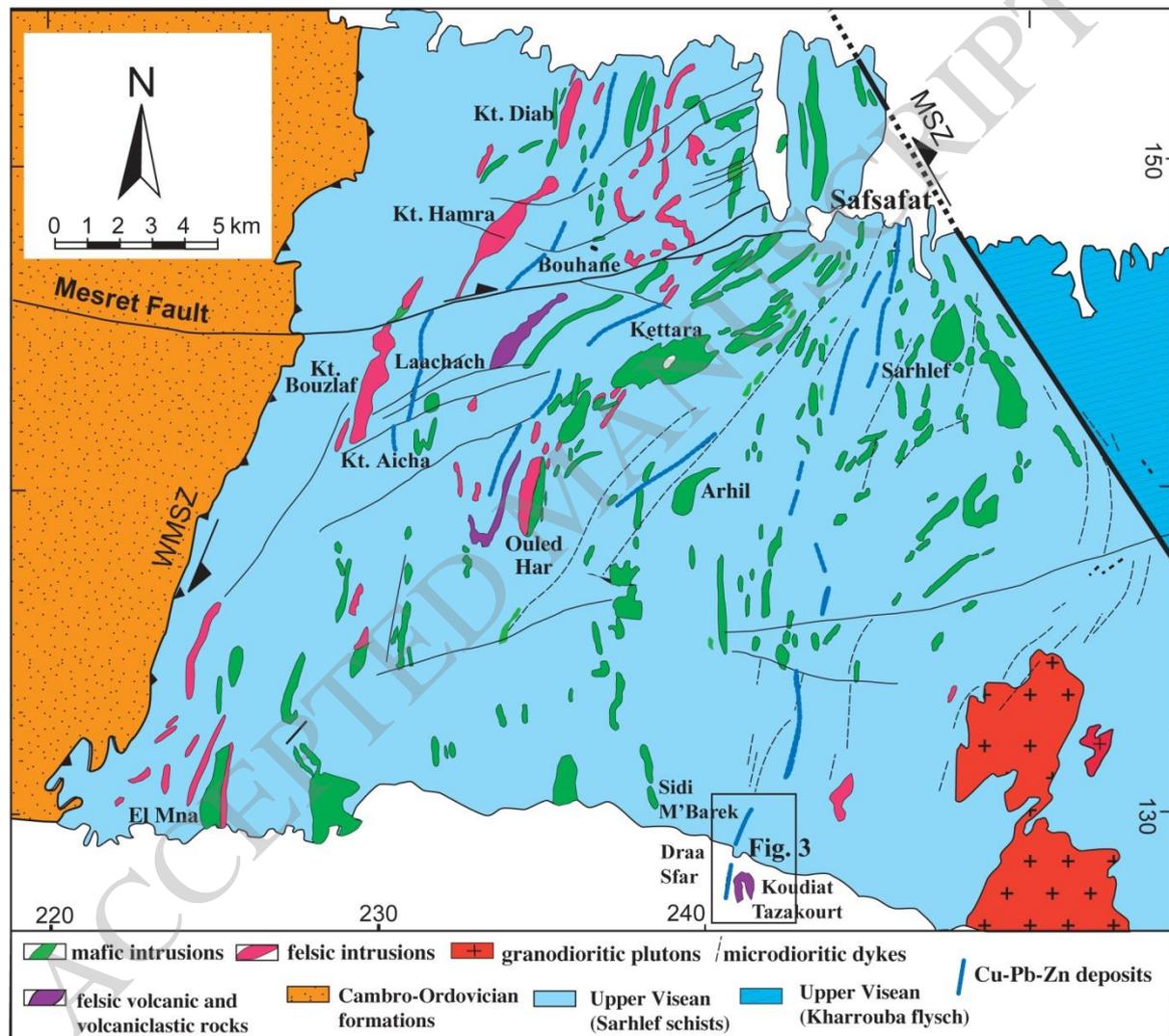
Depth (m)	Zn wt. %	Pb wt. %	Cu wt. %	Fe wt. %	Ag ppm	Co ppm	As %	Au ppm
87.7	0.03	0.074	0.009	na*	208	22	0.025	<dl
88.4	10.26	3.34	0.22	na	234	31	1.05	<dl
88.9	0.03	0.02	0.094	na	78	11	0.014	<dl
89.8	0.39	0.27	0.66	na	16	21	0.04	<dl
90.3	12.65	4.99	0.19	na	40	13	0.92	0.08
91.3	11.05	1.74	0.29	na	32	28	0.87	<dl
92.3	10.38	0.64	0.29	na	32	30	0.04	<dl
92.65	9.18	0.52	0.29	na	28	<dl	0.09	<dl
93	12.1	3.3	0.37	na	35	26	0.96	<dl
94	11.27	2.27	0.38	na	35	33	1.08	<dl
95	12.3	4.45	0.34	na	35	21	1.4	<dl
96	11.28	5.55	0.33	na	42	17	0.75	<dl
97	7.44	5.66	0.28	na	45	24	1.09	<dl
97.4	0.89	0.68	0.075	na	17	26	0.059	<dl
98	1.44	0.48	0.037	na	16	18	0.013	<dl
99	1.94	0.58	0.041	8.92	6	22	0.024	<dl
100	2.24	0.48	0.041	9.51	8	16	0.019	0.09
101	1.88	0.29	0.041	8.66	8	24	0.055	<dl
102	1.52	0.16	0.073	7.64	7	42	0.421	<dl
102.5	12.26	0.77	0.27	40.39	20	41	0.91	0.06
103.5	12.43	2.96	0.32	39.82	25	45	0.8	<dl
104.5	12.04	7.66	0.3	38.77	28	57	1.03	<dl
105.7	1.07	0.18	0.024	7.8	8	114	0.004	<dl
106.7	0.85	0.066	0.044	8.16	12	74	0.007	<dl
107.95	10.98	3.19	0.11	40.12	24	63	0.68	<dl
108.3	0.29	0.014	0.062	7.84	8	55	<dl	0.07
109.3	0.98	0.008	0.019	8.08	7	64	0.004	<dl
110	13.72	0.52	0.17	31.51	13	143	0.24	<dl
110.43	0.41	0.012	0.22	10.11	7	71	0.01	<dl
110.85	10.09	0.012	0.62	22.12	14	137	0.14	0.09
111	1.12	0.005	0.06	7.7	10	85	<dl	0.08
181	0.017	0.004	0.009	11.15	8	368	0.062	<dl
182	0.031	0.011	0.51	43.46	5	599	0.04	0.19
182.3	0.022	0.011	0.73	58.39	6	601	0.04	0.31
183.5	<dl4	0.019	0.98	58.22	8	605	0.04	0.29
184.5	0.19	0.045	0.67	58.41	8	611	0.04	0.31
185.5	0.12	<dl4	0.84	58.2	7	614	0.05	0.59
186.53	0.04	0.022	0.69	59.23	5	702	0.04	0.93
187.5	0.017	0.01	0.83	57.62	4	670	0.04	0.56
188.5	0.012	0.008	0.66	58.67	3	662	0.04	0.36
189.5	0.016	0.01	1.13	58.74	5	729	0.05	0.57

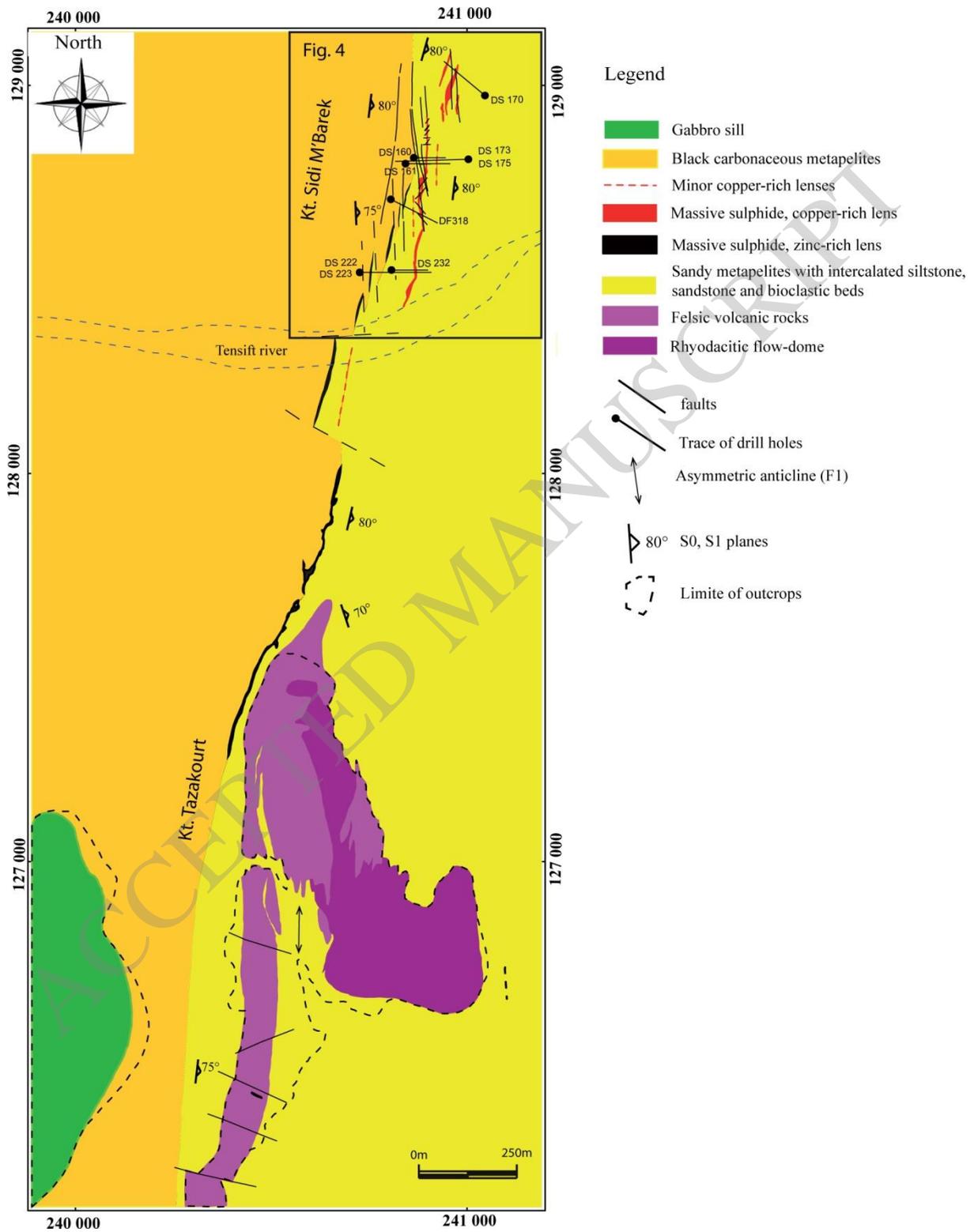
190.5	<dl	0.018	2.45	54.74	4	906	0.07	0.48
191.5	0.048	0.022	1.76	55.38	5	1200	0.1	0.32
192.52	0.08	0.037	1	53.74	4	804	0.05	0.33
193.5	0.088	0.043	0.97	48.48	6	1300	0.11	0.19
194.5	1.62	0.55	1.33	44.45	4	687	0.14	<dl
194.8	0.024	0.004	0.004	16.58	3	<dl	0.012	0.25
231	0.027	0.011	0.21	20.34	5	73	0.01	<dl
231.9	0.008	0.011	0.24	44.01	8	616	0.02	<dl
233.4	0.014	0.01	0.082	25.35	7	211	<dl	<dl
234.1	0.012	0.014	0.7	53.16	9	967	0.05	0.44
234.7	0.018	0.025	0.32	25.35	9	200	<dl	0.2
236.4	0.009	0.005	0.019	11.47	5	39	<dl	<dl
237.3	0.06	0.041	0.51	47.76	13	751	0.02	<dl
238.5	0.006	0.009	0.024	1<dl	6	41	0.015	0.08

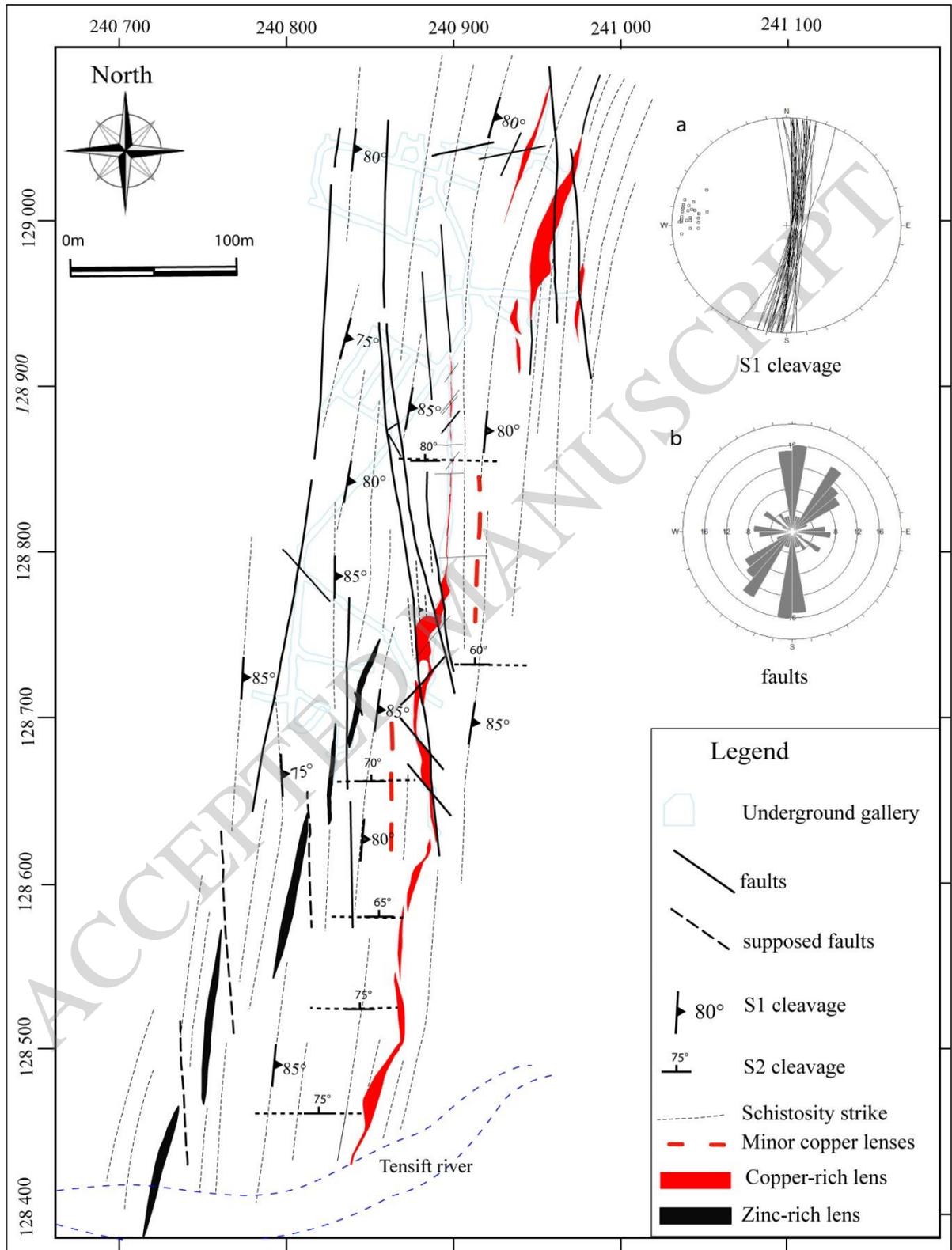
\* not analysed

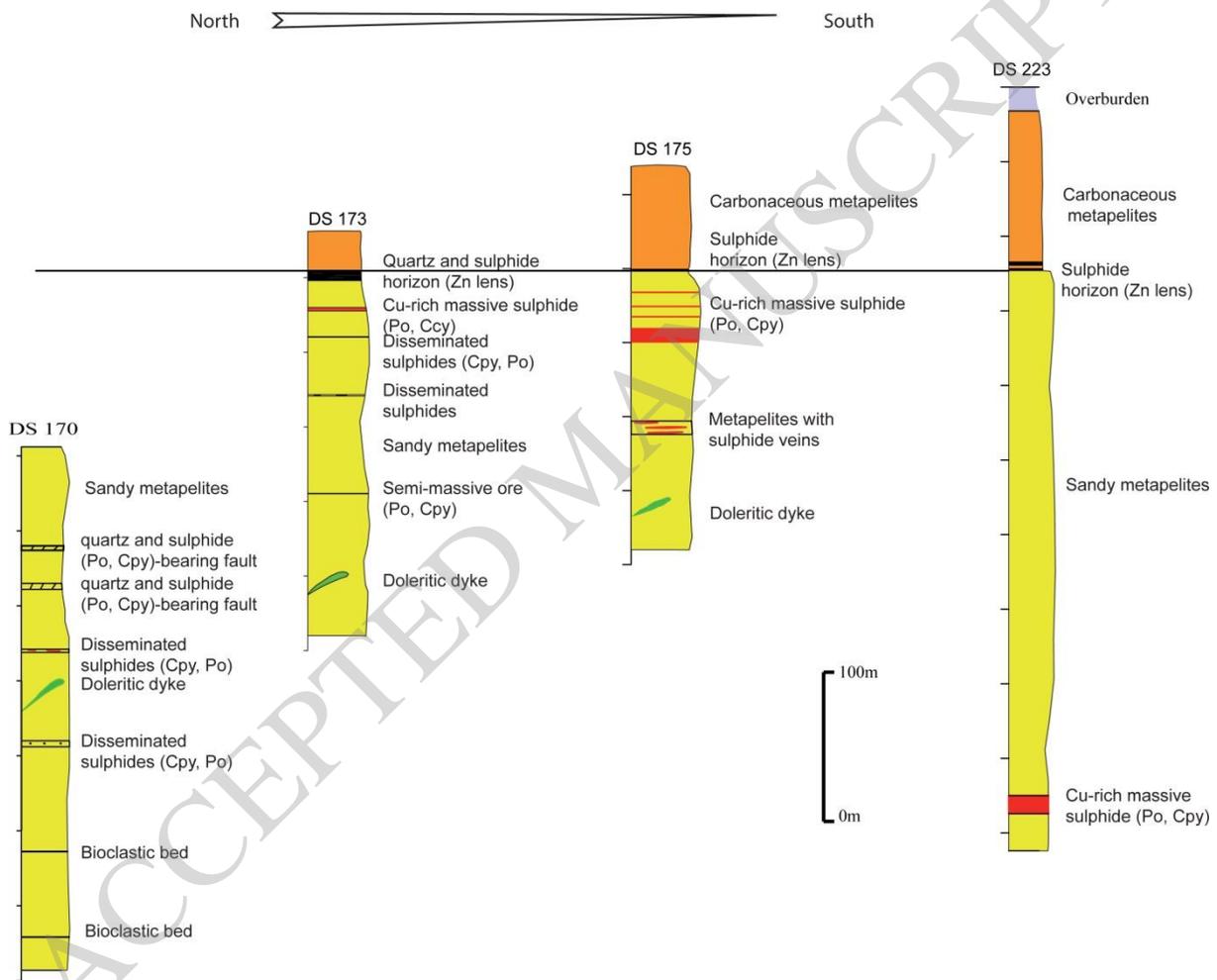
Detection limits (dl) 0.001% for Zn, Pb, Cu and Fe, 0.003% for As, 1 ppm for Ag, 10 ppm for Co and 0.05 ppm for Au

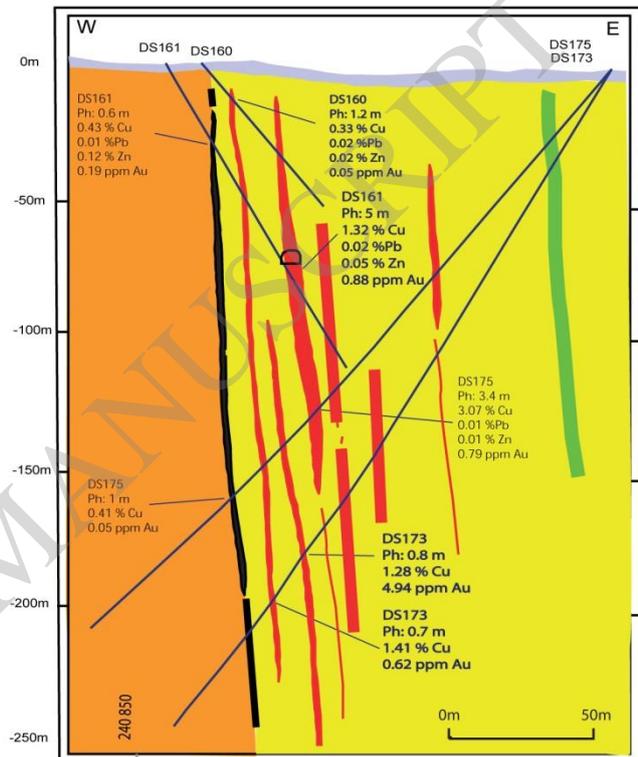
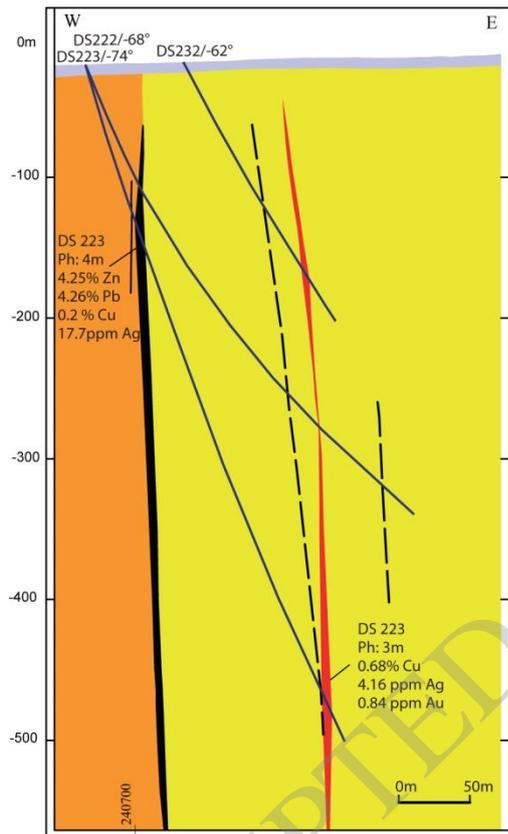




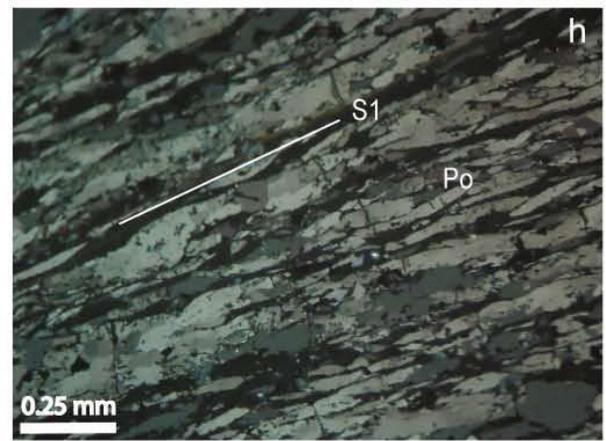
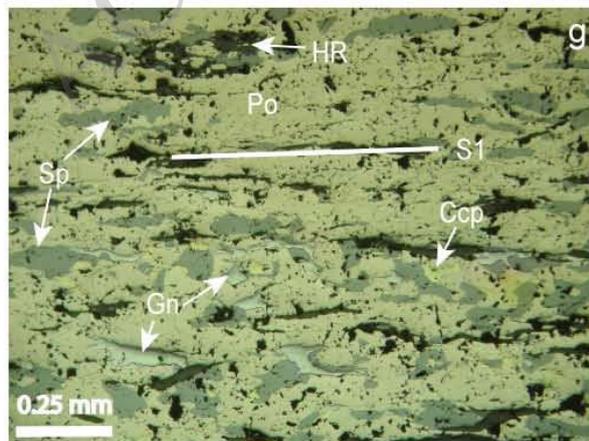
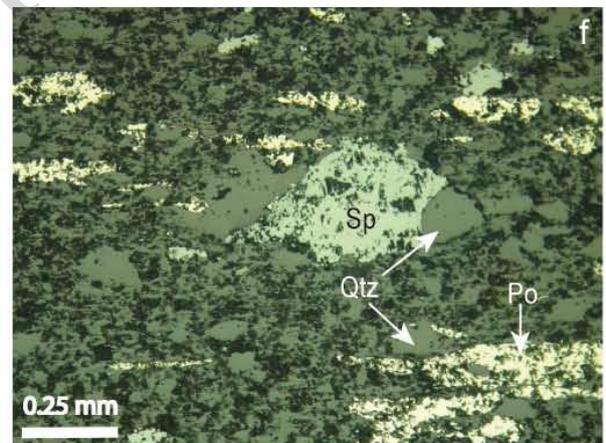
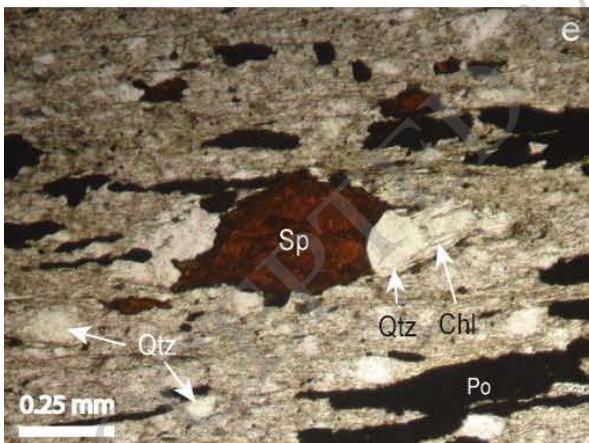
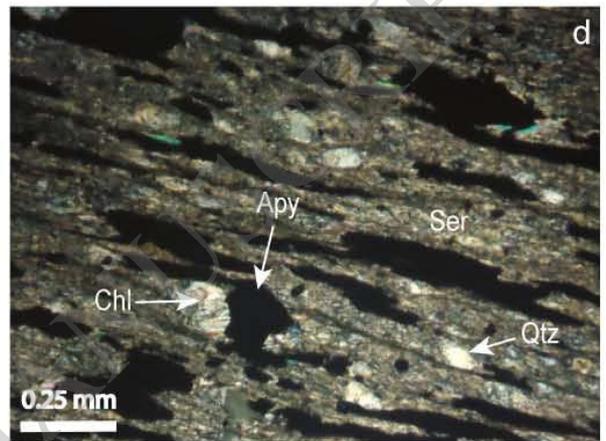
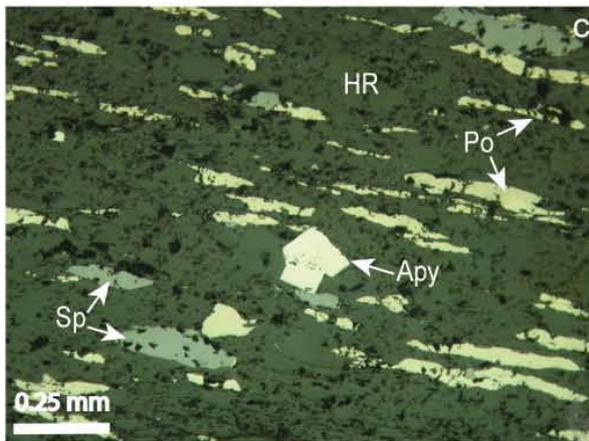
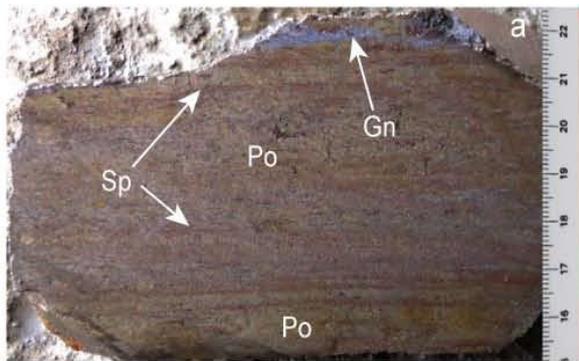


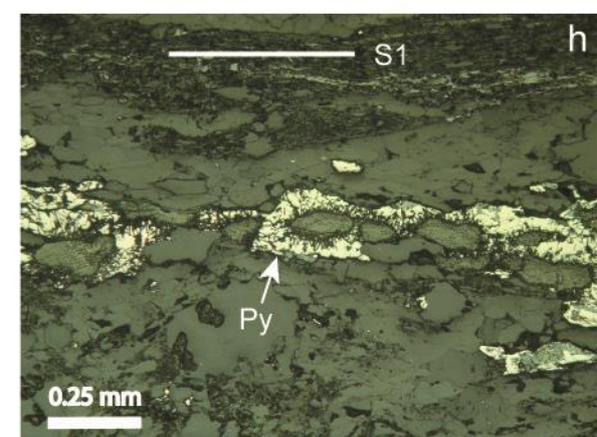
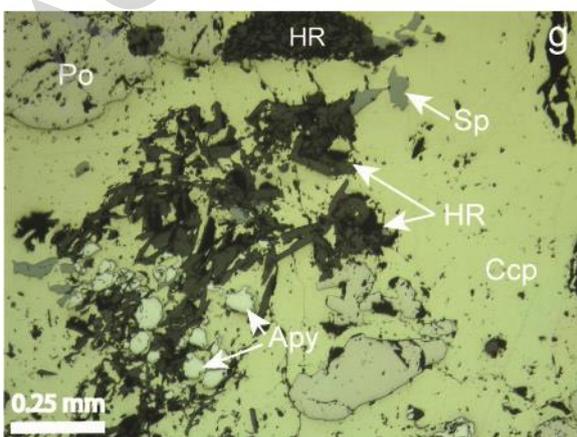
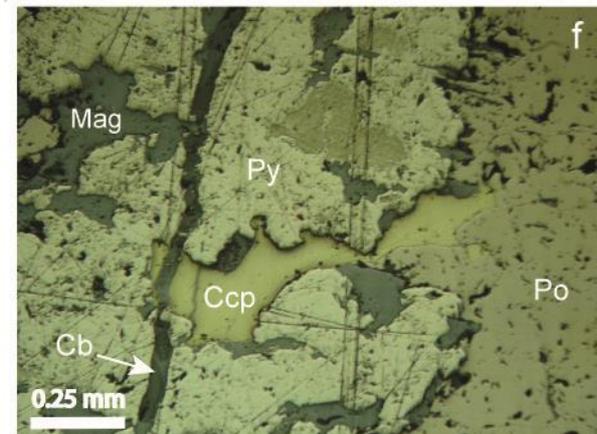
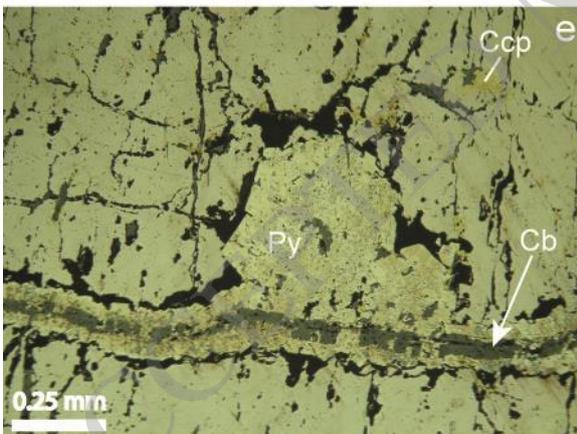
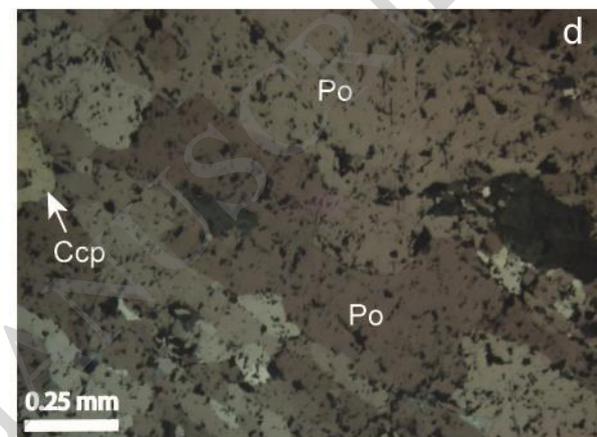
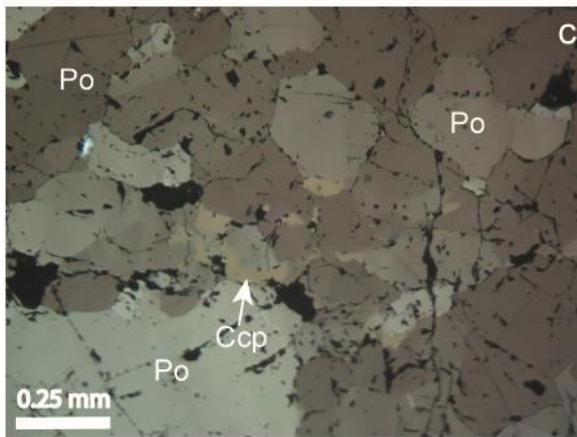
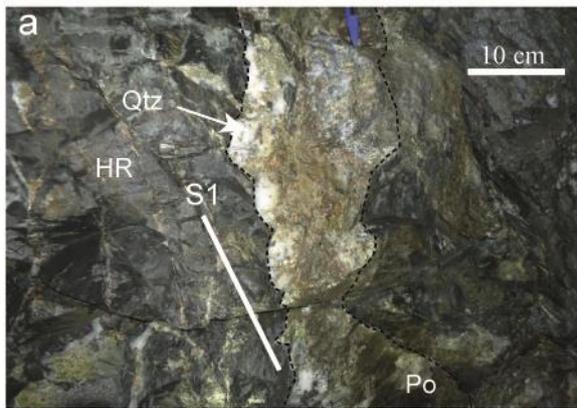






- Overburden
- Doleritic dyke
- Black carbonaceous metapelites
- Massive sulphide, copper-rich lens
- Massive sulphide, zinc-rich lens
- Sandy metapelites with intercalated siltstone, sandstone and bioclastic beds
- Fault
- Trace of drill holes
- Underground galleries





	Pre- to syn-tectonic	Syn- to late-tectonic	late- to post-tectonic
• Pyrrhotite		---	—
• Chalcopyrite		-----	
• Magnetite	----		
• Sphalerite		--	—
• Arsenopyrite			
• Cobaltite			-----
• Galena			-----
• Bismuth	--		-----
• Gold			-----
• Pyrite		 ?	
• Marcassite			
• Hematite		-----	
• Quartz			
• Chlorite			
• Sericite			
• Talc			
• Carbonates			-----
			
	<5 vol %	>5 and < 10 vol %	>20 vol %

