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Thomas Gautier, G. Danger, O. Mousis, F. Duvernay, V. Vuitton, et al.. Laboratory experiments to unveil the molecular reactivity occurring during the processing of ices in the protosolar nebula. Earth and Planetary Science Letters, 2020, 531 (1 February), pp.art. 116011. 10.1016/j.epsl.2019.116011. hal-02396048

HAL Id: hal-02396048 https://hal.science/hal-02396048

Submitted on 5 Dec 2019

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1	Laboratory experiments to unveil the molecular reactivity occurring
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16	Running title: Organic Matter Evolution in the protosolar nebula
17	Keywords: Protosolar Nebula, Interstellar Ices, Organic Matter, Laboratory simulations, High
18	Resolution Mass Spectrometry.

19 Abstract

Using laboratory experiments, we investigate the role of photo and thermal degradation in 20 the possible complexification mechanisms of organic matter that may originate from 21 22 interstellar ices prior to, or during the formation of the Solar System. We perform High Resolution Orbitrap Mass Spectrometry on organic residues formed from the photo- and 23 thermochemical alterations of Interstellar Medium (ISM) dirty ice laboratory analogues. We 24 25 probe, at the molecular level, the possible effects within the protosolar nebula on the composition and structure of these organic refractory materials obtained from an initial ice 26 composition representative of astrophysical ices. We show that nitrogen incorporation, by 27 competing with the carbon, has a strong influence on the final composition of the residue. 28 29 NH₃ rich ices lead to a group of unsaturated molecules in the final residue, while H₂O rich ices 30 lead to saturated ones. Finally, we observe and discuss the strong effect of UV irradiation on 31 the decarboxylation on organic matter and discuss potential implications of this result for the protosolar nebula. 32

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40 <u>1.</u> Introduction

The organic content of the solar system small bodies such as comets, asteroids or meteorites 41 found on Earth, has been an ongoing field of study for decades. One of the interest of such 42 study is that the organic matter content of these bodies can be used as a tracer to unveil their 43 history. Such organic matter can be, at least partially, inherited from the organic matter 44 45 formed in the interstellar medium (ISM) prior to the protosolar nebula (PSN) formation. In this 46 regard, understanding the chemistry occurring during the photo and thermal processing of astrophysical ices is a key aspect to apprehend the primitive organic contents of the small 47 48 bodies and the history of the solar system.

Ices are observed in a wide range of objects, from dense molecular clouds (Boogert et al., 49 50 2008) to comets/asteroids or icy satellites of giant planets such as Europa or Enceladus (Clark et al., 2013). The ice organic content can be obtained by their in situ analysis (*e.g.* Goesmann 51 52 et al., 2015), infrared remote sensing (e.g. De Sanctis et al., 2017) or inferred from molecules detected in the gas phase of objects (e.g. Meech and Svoren, 2004). However, such 53 information remains sparse, difficult to obtain and with limited knowledge of the alteration 54 55 processes (aqueous, thermal or by irradiation) undergone by the organic matter analyzed. In contrast, the soluble organic matter (SOM) in meteorites has been extensively studied and 56 displays a wide variety of organic matter, such as hydroxy and amino acids, sugar derivatives 57 58 and hydrocarbons (Burton et al., 2014; Callahan et al., 2014; Cooper et al., 2001; Cronin and 59 Moore, 1971; Martins et al., 2015; Pizzarello and Shock, 2010a).

For a few decades, laboratory simulations have been developed to bridge the gap between the organic matter observed in astrophysical ices and the one observed in meteorites. In this work, we use such an experimental approach to investigate the chemistry of organic

molecules that could be (trans-) formed from interstellar ices in the PSN and later
 incorporated in the meteorites parent bodies.

65 Laboratory ice processing has been shown to form refractory organic residues – also called yellow stuff - if reduced carbon such as methanol or methane is present, in addition to H₂O 66 and NH₃, in the initial ice (Bernstein et al., 1995; Briggs et al., 1992). A wide molecular diversity 67 was inferred in these residues by infrared analyses (Muñoz Caro and Schutte, 2003). Bernstein 68 et al., (1995) showed that infrared bands relative to nitrogen functional groups increase when 69 the quantity of NH₃ increases in the ice, while increasing H₂O in the initial ice results in residues 70 71 presenting similar signatures to the ones of residues formed from NH₃ rich ices (Henderson 72 and Gudipati, 2015; Nuevo et al., 2011). The formation of such organic residue is also of high interest for astrobiology, in particular de Marcellus et al., (2015) and Meinert et al., (2016) 73 74 reported the formation aldehyde and sugars such as ribose and other related sugars in the 75 residue.

A direct snapshot of the residue molecular diversity was obtained by high resolution mass spectrometry (HRMS) (Danger et al., 2013) where thousands of molecules presenting masses up to several thousands of Da were detected. Previous work from our group showed that residues are formed with two distinctive distributions for the organic matter, one enriched in nitrogen, with an average nitrogen to carbon ratio of 1.3, and another one presenting molecules predominantly with a carbon skeleton (Fresneau et al. 2017).

In this contribution, we analyze the molecular content using HRMS and infrared spectroscopy of organic residues formed through the irradiation and warming of several types of H₂O:CH₃OH:NH₃ ices. We mainly focus on the chemistry of ammonia and on the rapid decay

of carboxyls functions under UV irradiation in the protosolar ice, in apparent contradiction
with the large amount of carboxyls found in meteorites.

87 <u>2. Materials and Methods</u>

88 2.1. Sample Synthesis

In this work we investigate the processing of ice layers, analogs to organic rich protosolar 89 grains, into organic residues under photo and thermal alterations. The experimental setup 90 91 used to produce these laboratory analogs has been described in detail in Nuevo et al., 2007. 92 The samples analyzed for the present work were the ones used by Fresneau et al. 2017 to investigate the relationship between cometary materials and interstellar ices at a 93 macromolecular level. The main characteristics of the setup and of the samples are recalled 94 95 here. To produce the analogs, water, ammonia and methanol were mixed in a stainless-steel line before injection into the deposition chamber. The H₂O was purified with a Millipore Direct 96 97 Q5 system. The high purity CH_3OH (99.9%) and NH_3 (99.98%) were purchased from Aldrich and 98 Messer, respectively. To investigate the effect of the original ice composition on the molecular complexification of the organics in the sample, different ratios of water ammonia and 99 100 methanol were used for this study and are given in Table 1. To form the analog samples, ices were deposited under vacuum (10^{-5} Pa) on a MgF₂ substrate cryogenically cooled down to 77K. 101 102 The ices were irradiated with VUV photons (principally Ly- α , at 121.5 nm) simultaneously to its deposition on the substrate using a hydrogen discharge lamp emitting 2 x 10¹⁴ photons.cm⁻ 103 104 2 .s⁻¹. This rate corresponds to an average of 2 Ly α photons interacting with each deposited molecule during the sample production. The photon-molecule interaction occurs 105 106 predominantly in the bulk ice and not before the molecules condensation. The ice deposition 107 and irradiation processes were performed for 48h, after which the deposition and irradiation

108 were stopped and the MgF₂ substrate slowly warmed up to room temperature at 1.7 x 10⁻³ K.s⁻¹ using a thermal regulator (335 Temperature Controller from LakeShore) on the cryogenic 109 head. The sample was monitored during its formation using a FT-IR spectrometer attached to 110 the deposition chamber. The mid-infrared spectra (4000-1000 cm⁻¹, sum of 500 scans and a 111 112 resolution of 2 cm⁻¹) were used to ensure the proper deposition rate and composition of the 113 ice, as well as to give information about the organic residue composition left on the substrate 114 after the ice warm-up (de Marcellus et al., 2017). These residues obtained in the laboratory are considered analogs of the organic residues formed by the processing of icy mantles on 115 116 interstellar grains during the solar system formation. To investigate the effect of UV dose 117 received by the organic residue during its lifetime inside proto-planetary disks, three samples were additionally over-irradiated at room temperature during 48 hours without further 118 119 deposition and after their warm-up to room temperature and 10⁻⁵ Pa, considerably increasing the irradiation dose imparted per bond to around 10keV (de Marcellus et al, 2017). After their 120 synthesis, all samples were retrieved under inert atmosphere (Argon) sample holders to 121 prevent contamination. Given the extremely low amount of material formed during these 122 123 experiments, a quantitative production yield was not determined. However, we can 124 qualitatively assess the production rates relatively to each other by comparing their absorbance spectra, as described in Fresneau et al. 2017. Using this approach, the sample 125 126 with the highest production yield was the H₂O:CH₃OH:NH₃ 3:1:1 sample, followed by the NH₃poor ices 10:1:1 and 3:1:0.2 which have similar residue production yield. However, the NH3-127 rich ice 3:1:5 is by far the one yielding the less organic residue as discussed in Fresneau et al. 128 129 2017.

Table 1: List of samples and their production parameter used for this study. Molecules are
 considered detected in the samples only if they are present in all replicates.

$H_2O:CH_3OH:NH_3$ ratio	Over-Irradiation of the sample	Number of replicates
3:1:1	-	3
3:1:1	48 hrs	3
3:1:5	-	2
3:1:0.2	-	3
10:1:1	-	3

133 2.2. Sample Analysis and data treatment

134 2.2.1. Orbitrap High resolution mass spectrometry

To access their molecular composition, samples were analyzed through high-resolution mass 135 136 spectrometry (HRMS) using an LTQ Orbitrap XL. Residues were dissolved in 250 µL of ultrapure 137 methanol. All residues were totally soluble in methanol except the over-irradiated ones that contained a non-soluble fraction (Fresneau et al. 2017, de Marcellus et al. 2017). Note that 138 only the soluble fraction of these residues in methanol can be analyzed by HRMS with ESI 139 ionization. The exact solubilization protocol can be found in Fresneau et al. 2017. The 140 instrument was equipped with an electrospray ionization source (ESI). Instrument parameters 141 used were: syringe flow rate: 3 µL.min⁻¹; needle voltage: 3.7 kV; capillary temperature: 275°C; 142 143 tube lens voltage: respectively -131 V and +90 V for the negative and positive ionization modes; resolution was set at 100,000 at m/z 400; actual resolution was of 120 000 at m/z 200 144 and reached 220 000 at m/z 400; the mass acquisition range: m/z 200-400. The instrument 145 was externally calibrated using a mix of caffeine, MRFA peptides and Ultramark. 146

147 2.2.2. Data processing

The data were acquired with the XCalibur software and then post-calibrated and processed through Attributor, and in-house software developed at IPAG specifically for the processing and interpretation of high-resolution mass spectrometry data.

151 In order to avoid mis-attribution of peaks and discussing non relevant signal, we chose to be extremely conservative in the attribution of the peaks. We thus keep peaks for which we had 152 153 a completely unambiguous attribution. This lead to discard peaks for which an isobaric 154 ambiguity may have existed and peaks for which the signal-to-noise ratio (i.e. number of ions in the cell) was low enough to affect the mass accuracy of the Orbitrap. This lead us to a sharp 155 decrease between the number of detected peaks and the number of attributed ones. For 156 example, for the 3:1:1 sample, there was ~35000 detected peaks, most of them having a very 157 low signal to noise ratio (S/N<3); amongst the detected peaks, 1168 were at sufficient S/N 158 159 ratio and resolution to limit low ion counts and isobaric effects and were successfully and 160 unambiguously attributed to a molecular formula. We did not detect any significant variation in the ratio of attributed peaks versus detected peaks with the samples studied. All samples 161 162 were treated following exactly the same data treatment to ensure no bias would be introduced. 163

164

2.2.3. Modified van Krevelen Diagrams

van Krevelen diagrams were originally introduced in the 50's for petroleomics and present a scatter plot of the oxygen to carbon ratio versus hydrogen to carbon ratio of the sample elementary composition (van Krevelen, 1950). This visualization is especially adapted to investigate structural changes in complex samples such as petroleum (Kim et al., 2003; Wu et al., 2004) and has been adapted in so-called 'modified van Krevelen' for the study of planetary interest constituted of several thousands of molecules of unknown composition (Danger et

al., 2016; Gautier et al., 2014; Pernot et al., 2010; Ruf et al., 2017; Somogyi et al., 2012). It
should be underlined here that although van Krevelen or modified van Krevelen diagrams
present visible linear alignments of structures, these alignments cannot be directly linked to
the structure of the material but simply reflect the intrinsic order of the C, H, N and O
compositional space (Hertkorn et al., 2008).

176 *2.2.4. Ternary space diagrams*

van Krevelen diagrams allow for the discrimination of samples along two dimensions of the parameter space. This can limit the separation of samples presenting little divergence, in which case a representation using three dimensions could be necessary. A step forward was made using the m/z of the molecules as third dimension in 3D van Krevelen Diagrams (Wu et al., 2004, Pernot et al. 2010, Gautier et al. 2014, Danger et al. 2016) but these 3D diagrams demonstrated to be impractical to use as they require a user-defined orientation, which introduces a bias in the representation.

We introduce here a visualization for complex organic matter using a mix of modified van 184 Krevelen diagrams and ternary space representations. These ternary space representations 185 were introduced in the 18th century for the study of color mixing (Newton, 1704) and are now 186 commonly used in Geology, Chemistry and Material Sciences (Howarth, 1996). This 187 188 visualization can be defined as follows: each of the molecules present in the sample is located in a three-parameter triangular space with non-orthogonal axes. In this work we chose each 189 190 axis to respectively bear information about the oxygen, carbon and nitrogen content, of the molecules. This means that the representation presented here focuses on the large scale 191 192 structure of the macromolecules and does not bear information about the hydrogen content or the double bond equivalent of the molecules. Should such information be needed, a way 193

to mitigate this issue would be to use Hydrogen content as axis instead of carbon, nitrogen oroxygen.

196 The position of a given molecule in the diagram can be defined so that its projection on each 197 axis relatively to an atom α (with α = C, N or O) is be equal to $\frac{\alpha}{C+N+O}$. An example of this 198 representation is given in Figure 1.

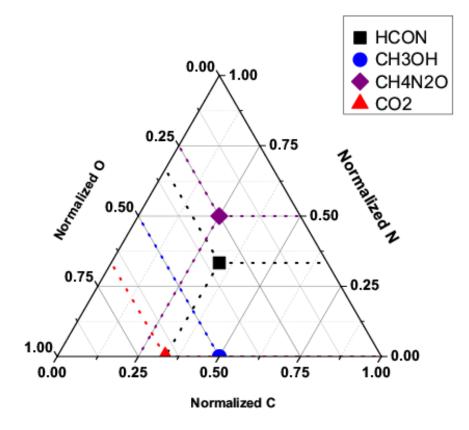


Figure 1 : Example of ternary space molecular diagram and the position of four molecules in this diagram. The black square shows the position of a molecule of composition HCON, the blue circle a molecule of CH_3OH , the red triangle would be CO_2 and the purple diamond a molecule of formula CH_4N_2O . The colored dashed lines are the projection of each point on the three axes to facilitate the interpretation of the diagram.

- 205 The black square shows the position of a molecule of composition HCON which has an equal
- number of carbon, nitrogen and oxygen atoms, each axis thus has a 1/3 weight in the position
- 207 of the molecule. The blue circle shows the position of methanol (CH₃OH, or also CO,
- formaldehyde H₂CO or acetic acid CH₃COOH) where O and C are equals. The red triangle shows

the position of CO_2 and the purple diamond would be the location of a molecule with raw formulae CH_4N_2O . Since this visualization requires the normalization of the molecular ratios to the sum of C+N+O atoms, the position of a point on an axis should not be read as the molecular composition of this molecule.

213 2.2.5. Structural markers

214 The formulas of organic molecules, available thanks to the use of high-resolution mass 215 spectrometry, can be used to infer structural information about the molecules detected in the 216 samples. A structural marker extensively used in the literature is the Double Bond Equivalent (DBE), which gives information regarding the unsaturation (double bonds and cycles) borne 217 by a molecule in its structure. The DBE is calculated as follows: $DBE = \frac{1}{2} (2C - H + N) + 1$ 218 where C, H and N represent respectively the number of carbon, hydrogen and nitrogen atoms 219 within the molecule. The interpretation of DBE in itself is however difficult in the case of large 220 221 molecules and/or molecules with a large number of heteroatoms such as the one present in our samples (Danger et al. 2016). In addition to the DBE we therefore also used the New 222 Aromaticity Equivalent parameter X_c as defined by Yassine et al., 2014 : $X_c = \frac{2C+N-H-2mO}{DBE-mO}$ 223 where *m* is the fraction of oxygen atoms in the conjugated structure of the molecules. Using 224 225 infrared analysis of the residue Danger et al. 2016 estimated that the double bonded O atoms represented approximately 50% of the O atoms in the residue, giving an estimate of m=0.5. 226 227 Should this value of m be slightly higher or lower, this would not affect the classification of the 228 bulk of the molecules but simply shift the X_c distribution toward higher or lower values. The X_c distribution for all samples is given in Figure SI 1. 229

For each molecule detected in the sample we calculated the DBE and X_c . Following Yassine et al. 2014 classification, molecules for which Xc > DBE are by essence aliphatic. If $X_c < 2.5$, the molecules were considered as aliphatic; those with $2.5 \le Xc < 2.7$ were classified as aromatic and those for which X_c is greater than 2.7 as condensed aromatic.

234 3. Results and discussion

235 3.1. Effect of Ammonia and water proportion in ices

236 3.1.1. Influence of ammonia

237 Modified van Krevelen diagram H/C vs. N/C does not give sufficient information on the impact 238 of the initial ammonia ice ratio as they rapidly become congested for such complex mixture of 239 similar properties (Figure SI 8). We therefore used another representation that spreads the sample representation over three dimensions, the ternary space diagrams (Figure 2, 240 respectively with H₂O:CH₃OH:NH₃ ratios of 3:1:5 blue circles and 3:1:0.2 red circles). Each dot 241 242 in this diagram corresponds to the molecules detected in sample either in positive or in negative ionization mode. As inferred from this figure, both samples present a bi-modal 243 244 distribution of their molecules. This bimodality is not related to the positive/negative 245 ionization, but to the presence of a group of nitrogen-rich molecules in both samples. This Nrich fraction of the residue has already been observed by Fresneau et al. 2017 on the average 246 247 composition of the residues. However, using elemental composition only, Fresneau et al. 2017 248 were not able to provide an explanation for this N-rich distribution. We confirm here that this bimodality is clearly detected at the molecular level (bi-modal distributions for both samples 249 250 in Figure 2), and we can use the ternary space diagram to investigate its origin. We can see in 251 Figure 2 that for both residues the N-rich distributions are at higher N-axis and lower C-axis values relatively to the C-rich distribution. This corresponds to molecules with a higher 252

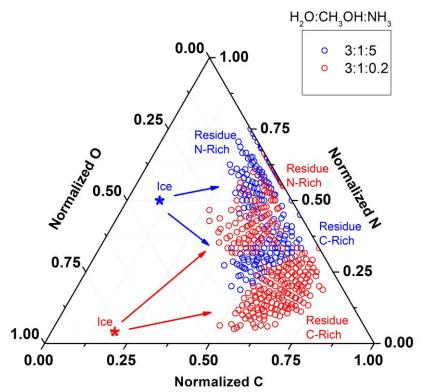


Figure 2 : Ternary space diagrams of the molecules detected in positive and negative ionization mode in the sample produced with $H_2O:CH_3OH:NH_3$ ratios of 3:1:5 (blue circles) and 3:1:0.2 (red circles). The areas at high N ratios correspond to the N-rich distribution identified by Fresneau et al. 2017 in the 3:1:5 and 3:1:0.2 samples respectively. The blue and red stars represent the position in the ternary space of the ice used to form the residue, respectively the $H_2O:CH_3OH:NH_3$ 3:1:5 and 3:1:0.2 ones. Colored arrows show the evolution pathway underwent by the organic matter from the ice to the formation of the residue.

253	nitrogen content and lower carbon content (hence their names). But we can also see that
254	these distributions have a similar position on the O-axis, implying that the competition for
255	inclusion in the residue does not happen between N and O but between N and C. This indicates
256	that the difference between the carbon rich and the nitrogen rich group would be due to the
257	existence of nitrogen bearing functional groups with no oxygen or fewer carbon atoms, such
258	as cyclic or aromatic structures including nitrogen. See part 1.3 for further discussion on the
259	ice reactivity

For both initial ice mixtures, it appears that the residue formed contained far less oxygen relatively to carbon and nitrogen than the original ice. This could be due to a large use of

carbon and nitrogen for the growth of the macromolecular structure, while oxygen would tend
to inhibit polymeric or cyclic molecules formation. An unsurprisingly low reactivity of H₂O in
the residue formation could also explained these observations, implying that most of the O
atoms in the residue comes from the methanol.

Interestingly, while there is almost no species in common in the distribution of the molecules detected in the residue formed from $H_2O:CH_3OH:NH_3$ ratios of 3:1:5 and 3:1:0.2, they both overlap with the distribution of molecules from the sample produced with a $H_2O:CH_3OH:NH_3$ ratio of 3:1:1 as shown in Figure 3.

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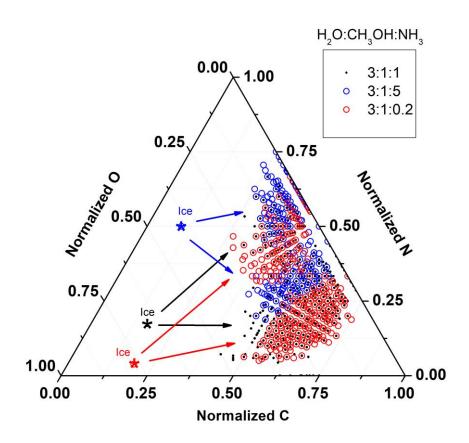


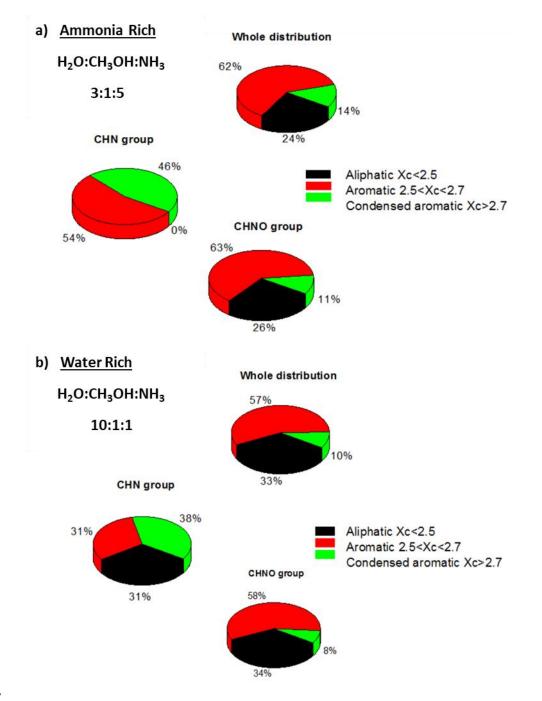
Figure 3: Superposition of the ternary space diagrams of the molecules detected in the sample produced with H₂O:CH₃OH:NH₃ ratios of 3:1:1 (black dots) compared to the one detected in the 3:1:5 (blue circles) and 3:1:0.2 (red circles) samples. Black, blue and red stars indicate the position of the initial ices on the ternary diagrams, with a ratio H₂O:CH₃OH:NH₃ of 3:1:1, the 3:1:5 and the 3:1:0.2 respectively.

277 As visible in this figure, a major fraction of the molecules detected in the 3:1:1 sample are 278 indeed present either in the 3:1:5 or in the 3:1:0.2 sample. This implies that although the 279 chemistry leading to the formation of the organic residue can be strongly driven in distinguishable pathways depending on the ammonia content, the 3:1:1 sample represents an 280 281 intermediate case with both pathways starting to emerge (however, both pathways 282 incomplete in the 3:1:1 sample as the most extreme points of the distribution - high nitrogen 283 and/or low carbon and/or high oxygen contents - are only populated by the 3:1:0.2 and 3:1:5 284 samples). This can be confirmed also in the mid-infrared signature of the residue (see spectra in Fresneau et al. 2017), where the absorption bands in the 1000-1700 cm⁻¹ wavenumber 285 range $(5.9 - 10 \,\mu\text{m})$ of the 3:1:1 residue correspond to the bands detected either in the 3:1:0.2 286 or the 3:1:5 samples. 287

288 3.1.2. Ammonia and water competition

289 Fresneau et al. 2017 reported that the average composition of samples produced with 290 H₂O:CH₃OH:NH₃ ratios of 3:1:5 and 10:1:1 were surprisingly similar and confirmed the results of Henderson and Gudipati 2015 that "an increase of water or ammonia in the initial ice leads 291 292 to the same effect: a higher incorporation of nitrogen in the residue". This first order similarity 293 between the residues formed from water and ammonia rich mixtures can be confirmed by their highly similar infrared signatures (Fresneau et al. 2017) and strong correlations in 294 295 modified van Krevelen and ternary space diagrams as visible in Figure SI 2 and Figure SI 3. In 296 this work we investigate this phenomenon further by comparing, at a molecular level, the two samples respectively enriched in water and ammonia using the aromaticity criteria 297 298 classification (Yassine et al. 2014, Danger et al. 2016).

Figure 4 a) and b) present the division of different classes of molecules respectively in the 299 ammonia rich (3:1:5) and water rich (10:1:1) samples. Each molecule detected in these 300 samples was classified amongst three categories depending on its aromaticity equivalent. 301 Molecules for which $0 < X_c < 2.5$ or a double bond equivalent (DBE) <4 are classified as having 302 303 an aliphatic structure with a few unsaturation carried by heteroatoms, molecules for which $2.5 < X_c < 2.7$ are classified as aromatic based structures and finally molecules with a X_c higher 304 than 2.7 are considered to be condensed aromatics. For more information, the histogram of 305 306 the complete distribution for all samples is given in Figure SI 1.



308 Figure 4 : a) Repartition of the molecules detected in the Ammonia rich H₂O:CH₃OH:NH₃ 3:1:5 sample 309 depending on their level of unsaturation based on the evolution of the aromaticity equivalent (Xc) for the totality of the molecules detected in the sample (whole distribution, 711 molecules, top 310 311 chart), the molecules containing C,H,N and O atoms (CHNO group, 652 molecules, bottom chart) and 312 the molecules solely containing C, H and N atoms (CHN group, 59 molecules, left chart). No molecules solely comprised of only C, H and O were identified. b) : Repartition of the molecules detected in the 313 314 water rich H₂O:CH₃OH:NH₃ 10:1:1 sample depending on their level of unsaturation based on the 315 evolution of the aromaticity equivalent (Xc) for the totality of the molecules detected in the sample 316 (whole distribution, 831 molecules, top chart), the molecules containing C,H,N and O atoms (CHNO 317 group, 775 molecules, bottom chart) and the molecules solely containing C, H and N atoms (CHN 318 group, 56 molecules, left chart). No molecules solely comprised of only C, H and O were identified. 319 The fraction of molecules classified as aliphatic is shown in black, red represent the fraction classified as aromatic and green the fraction considered to be condensed aromatic molecules. The percentage corresponds to the one of the molecules classified according to a given aromaticity relatively to the entire distribution. Given the number of classified molecules and their variability in the replicates, we can make a conservative estimate of $\pm 10\%$ as the 2σ uncertainties associated with this classification for the CHN groups. Uncertainties on the CHNO and whole distributions percentages are much lower.

When looking at the whole distribution (top charts in Figure 4a and b), corresponding to the 326 327 molecules identified in the ammonia rich and water rich samples respectively, the major fraction of the molecules detected in these samples fits into the aromatic category (roughly 328 60%). The second most predominant class corresponds to the aliphatic molecules (24% and 329 330 33% respectively for the 3:1:5 and 10:1:1: sample) and the least present group is the one comprised of condensed aromatics (14% and 10% respectively). Given the associated error 331 bars of these distributions, we can conclude that, taken as a whole, both samples have a 332 practically identical molecular content as previously observed. This is actually driven by the 333 similarity of the CHNO groups (bottom charts) of 3:1:5 and 10:1:1 that comprise ~90% of the 334 335 molecules detected in these samples.

336 Differences between the ammonia and the water driven residues become noticeable when it 337 comes to the molecules containing solely carbon, hydrogen and nitrogen atoms, the CHN group (left charts in Figure 4a and 4b). However, these effects are visible on a minor fraction 338 of the total attributed peaks (59 and 56 molecules in the CHN group for 3:1:5 and 10:1:1 339 samples respectively), potentially rendering the statistics poorly representative of the samples 340 341 and the following paragraph discussion presents only hypotheses that would require further investigated in the future: for the sample produced from an ammonia rich ice, the CHN group 342 343 does not contain any aliphatic molecules. This indicates that the ammonia could tend to drive the molecular growth leading to the formation of the CHN fraction toward aromatization, 344 probably due to the likelihood of nitrogen atom inclusion in aromatic structures. On the other 345 346 hand, the CHN fraction of the sample produced from water rich ice has the highest aliphatic

content of all the samples investigated in this work (see Figure SI 4 and Figure SI 5 for the other samples distributions). This effect does not appear to be correlated with the length of the chains as both distributions have a similar average m/z (266.04 for the ammonia rich sample and 267.02 for the water rich sample). For further discussion on the aromaticity of the residues, readers can direct themselves to Fresneau et al. 2017 and its Figure 5 presenting the DBE vs N/C ratio in both positive and negative mode of the 3:1:1 sample.

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3.1.3. Effects on the ice reactivity

354 The chemistry driving the residue formation is largely dependent on radical chemistry at low temperature (77K) (Danger et al., 2013). During the UV irradiation, the initial molecules in the 355 ices are photo-activated leading to the formation of radicals. In case of H₂O, CH₃OH and NH₃ 356 357 ices, H[•] and OH[•] are formed from H₂O, H[•] and NH₂[•] for NH₃ and various radicals from CH₃OH such as CH₃O[•], HCO[•], CH₂OH[•], CH₃[•] and OH[•]. Therefore, the radical recombination leading to 358 359 the first molecules at low temperature is highly impacted by the initial ice composition. 360 Furthermore, this reactivity probably occurs concomitantly to the evolution of morphological and physical properties of organic residues, as observed by Piani et al. (2017). In the case of 361 362 ices rich in H₂O, water photochemistry increases the amount of radical hydrogen (H[•]), which rapidly migrates through the ice at 77K, potentially increasing a rapid recombination with 363 other radicals. This recombination would be enhanced depending on the temperature, since 364 365 Tachibana et al., (2017) observed a liquid-like behavior at 50-150 K in such ice. This 366 observation would imply a higher radical diffusion in UV-irradiated ice leading to an effective formation of organic molecules. This would imply a fast-paced saturation of the alkyl chains 367 368 during the molecular growth of the residue. Furthermore, H₂O can act as a diluting agent spatially separating the different radicals at low temperature and limiting their direct 369

370 recombination for those that do not migrate at this temperature. Therefore, a higher content
371 of H₂O in the initial ice could limit the incorporation of nitrogen on the molecular skeleton,
372 which is not the case when NH₃ is dominant in the initial ice. This effect is actually visible in
373 the higher N/C ratio detected in the ammonia rich sample (1.8 on average for molecules of
374 the CHN group) versus the one of the water rich sample (0.8 on average for the CHN group).
375 For the ammonia rich mixture, nitrogen radicals quickly saturate carbon-bearing radicals,
376 implying a fast incorporation of nitrogen atoms into the molecular skeletons.

377 These results clearly demonstrate that nitrogen plays an important role in the chemistry of 378 the residue formation. Next to carbon, it drives the chemistry and directly impacts the global 379 molecular skeleton of the residues formed. At high amounts of NH₃ in the initial ice, nitrogen competes with carbon, implying the formation of molecules-bearing skeletons mixing nitrogen 380 381 and carbon. This leads to the formation of higher unsaturated structures such as nitrogen 382 heterocycles. Interestingly, as observed for residues produced from a H₂O:CH₃OH:NH₃ 3:1:1, 383 with the same amount of C vs. N in the initial ice, a well-balanced chemistry is driven by 384 nitrogen and carbon leading to a molecular diversity fading of the dichotomy observed between 3:1:5 and 3:1:0.2 ices. Nitrogen is thus not only a heteroatom allowing the formation 385 of specific chemical functions such as amines, nitriles or amides. Nitrogen incorporation also 386 387 probably enhances the formation of polar structure in the residue, reacting easily with radicals 388 when the ice is warmed. Nitrogen is thus an important reactant impacting the chemistry driving the residues formation, at the difference of oxygen as discussed in the next section. It 389 is also interesting to note that without NH₃ in the mixture almost no residue is obtained. 390

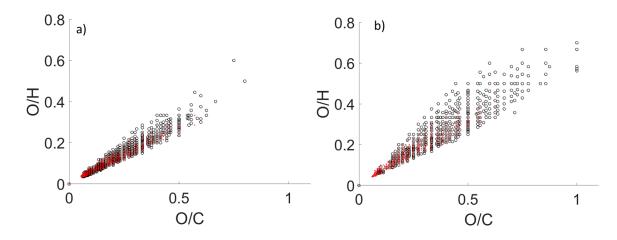
Furthermore, the observed reactivity in the residue formation is directly correlated to molecules detected in the gas phase while the ice is warmed up to 300 K. Abou Mrad et al., 2017 have shown that CHO molecule abundances in the gas phase are highly influenced by

the presence of ammonia in the initial ice. Indeed, the presence of ammonia reduces the amount of aldehyde observed in the gas phase by recombination of the aldehyde precursor radicals with radicals produced by NH₃ photodestruction. The same behavior between radicals is observed for ethers. Therefore, at least in the laboratory, a direct relation can be drawn between the residue composition and the molecular abundances in the gas phase. Such investigation is beyond the scope of this manuscript, but the monitoring of the gas phase during the residue irradiation could help identifying processes at act during this phase.

401 3.2. Residue irradiation

402 3.2.1. Evidence of molecular decarboxylation

To visualize the effect of sample over-irradiation at the molecular level we used a modified van Krevelen representation as presented in Figures 5, SI 1 and SI 2. Black circles represent the molecules detected in the sample before over-irradiation in their O/C vs O/H (Figure 5), H/C vs O/H (Figure SI 6) and H/C vs O/C (Figure SI 7) parameter spaces. Green dots represent the molecules detected in the sample after over-irradiation.



408

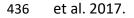
Figure 5: O/C vs O/H modified van Krevelen diagram of the molecules detected in the sample in
 positive (a) and negative (b) ionization mode before (black circles) and after (green dots) the sample
 irradiation. See Figures SI 6 and SI 7 for the O/H vs. H/C and H/C vs. O/C representation.

412 From these visualizations, one can see that a substantial fraction of the molecules initially present in the soluble fraction of the 3:1:1 sample are not detectable anymore after over-413 irradiation of the sample. These molecules degraded during irradiation do have the following 414 compositional parameters: (i) an oxygen to hydrogen ratio greater than 0.3 which is 415 significantly higher than the O/H ratio of most of the molecules detected in the over-irradiated 416 417 sample; (ii) an oxygen to carbon ratio greater than 0.4; and (iii) a hydrogen to carbon ratio 418 lower than 1.7. We can reasonably exclude this being solely due to a detection effect (such as 419 an intensity threshold) since (1) this effect is visible in all replicates of the experiments and (2) these molecules are significantly located in the same area of the parameter space and not 420 421 randomly dispersed in Figures 5, SI 1 and SI 2.

422 The loss of molecules upon over-irradiation also corresponds to molecules detected in the 423 negative ionization mode. This implies that chemical functions bearing acidic hydrogen such 424 as carboxylic acids are strongly affected (Holčapek et al., 2007; Kruve et al., 2014). 425 Concomitantly to this loss in the over-irradiated samples, new molecules are observed, mainly 426 in the negative ionization mode, with lower O/H and O/C ratios (see Figure SI 7). The residue 427 over-irradiation seems to induce a decarboxylation process, clearly visible in the negative mode, which presents a disappearance of molecules at high O/C and O/H in total concordance 428 429 to a loss of COOR (R: H or alkyl) groups, and in counterpart, the appearance of new molecules at low O/C and O/H. However, the whole mechanism is not entirely observable by HRMS 430 analyses, since a part of the molecules modified by the over-irradiation phase is actually 431 insoluble (de Marcellus et al., 2017) and thus not taken into account in this type of analysis. 432

This decarboxylation was thus confirmed by the infrared absorption spectra of the residue
before and after irradiation. These spectra are presented in Figures 6 and 7 together with their

435 Gaussian deconvolution. Bands attributions and positions are the ones described in Fresneau



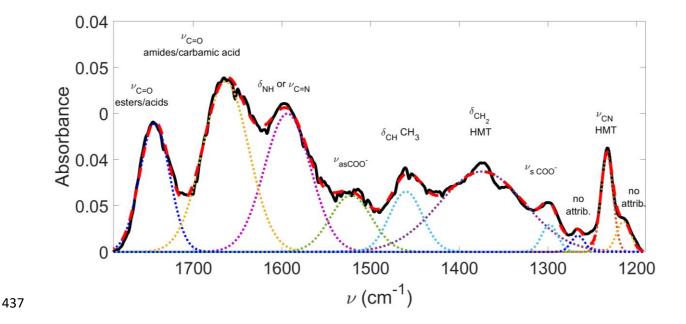


Figure 6: Gaussian deconvolution after baseline subtraction of the 1200-1850 cm⁻¹ region of the IR absorption spectrum of the residue produced from 3:1:1 ices before its irradiation. The measured spectrum is in full black, the deconvoluted Gaussians in dotted color lines and the dashed red line corresponds to the reconstructed fit after deconvolution. For the Gaussian fit, band positions were taken from Fresneau et al. 2017 and their intensities and HWHM were fitted iteratively using Fityk software.

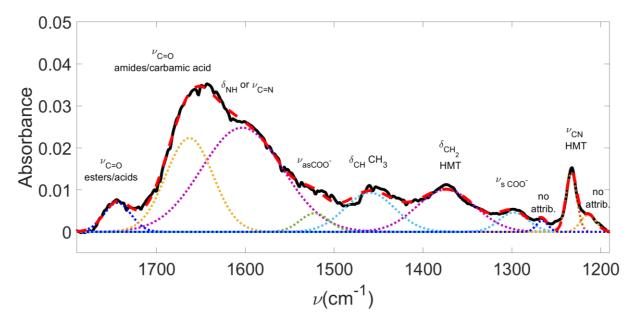


Figure 7: Gaussian deconvolution after baseline subtraction of the 1200-1850 cm⁻¹ region of the IR absorption spectrum of the residue produced from 3:1:1 ices after its irradiation. The measured spectrum is in full black, the deconvoluted Gaussians in dotted color lines and the dashed red line

448 corresponds to the reconstructed fit after deconvolution. For the Gaussian fit, band positions were 449 taken from Fresneau et al. 2017 and their intensities and HWHM were fitted iteratively using Fityk 450 software.

451

As visible in these figures, the C=O stretch in esters or acids (blue Gaussian curves) is strongly 452 453 depleted by the irradiation of the sample compared to C=O stretch in amides (orange Gaussian curves). As a counterpart, the NH bendings or C=N bands appear stronger in the over-454 irradiated sample. 455

456 Together with the HRMS analyses we can thus conclude that the residue over-irradiation induces a strong decarboxylation process. This result would imply a fast decay of carboxylic 457 functions in the protosolar organic matter analog upon UV irradiation. Regarding residue 458 irradiation, in addition to the chemical changes we investigate in our work, it should also be 459 460 noted that Piani et al. (2017) have also reported significant evolution of morphological and physical properties of organic residues formed from the evolution of H₂O-CH₃OH-NH₃ ices due 461 462 to UV irradiation.

463

3.2.2. Possible implications of this decarboxylation process on the -COOR molecules in meteorites 464

Fresneau et al. 2017 have defined an α parameter as $\alpha = \frac{A_{1745 cm^{-1}}}{A_{1665 cm^{-1}}}$ where $A_{1745 cm^{-1}}$ is the 465 area of the $v_{C=0}$ band of ester or acids and $A_{1665 cm^{-1}}$ is the area of the $v_{C=0}$ band of amide or 466 467 carbamic acid. Using the area of the deconvoluted Gaussian presented in Figure 6 and

Figure 7, we can retrieve the value of α to be 0.54 in the residue before irradiation, while it 468 plummets to 0.09 after the residue irradiation. Given the 48 hours of the over-irradiation of 469 our sample and that our lamp has a fluence of 2 x 10^{14} photons.cm⁻².s⁻¹, the total dose 470

deposited in our sample is of 3 .5 x 10¹⁹ photons.cm⁻². The experimental conditions used here 471 472 can be compared with the model developed by Ciesla and Sandford, 2012, which describes the exposure of ice grains to the UV flux during their dynamical evolution within the protosolar 473 nebula (PSN). In their nominal model, these authors find that a typical dose for particles with 474 1 μ m radii is about 5 x 10¹² photons after 1 Myrs of radial and vertical transport within the 475 disk, and assuming a nominal interstellar UV flux of 10⁸ photons cm⁻².s⁻¹ (Habing, 1968). This 476 would give a dose for such particles of 1.6 x 10²⁰ photons cm⁻² in 1 Myrs. In comparison, the 477 total dose from our experiment of 3.5 x 10¹⁹ photons.cm⁻² would represent ~220 kyrs of the 478 evolution of Ciesla and Sandford (2012)'s PSN model for such particles. 479

480 Our results show that these 220 kyrs would be enough to induce a dramatic decrease in the -COOR or -COOH functional group of the organic matter in the grains. If organic rich particles 481 482 indeed spent ~1 Myr in the PSN, they would be completely depleted in such functional groups. 483 This conclusion is not in agreement with measurement performed on chondritic meteorites, where carboxylic acids are a major component of the soluble organic matter. For example, in 484 485 the Murchison meteorite (CM2), the total abundance of carboxylic acids has been estimated to be around 330 µg g⁻¹, roughly 40% of the identified organic content and more than any 486 487 other organic groups besides the macromolecular structure (Lawless and Yuen, 1979; Pizzarello and Shock, 2010b; Sephton, 2002; Yuen et al., 1984). 488

The reader should keep in mind that our results are based on only one kinetic irradiation point (48h in the laboratory – 220 kyrs in the PSN) and that further experiments should be dedicated to the study of the kinetics of carboxylic functions degradation. However, our results suggest a fast decay of the carboxylic acids and carboxylic esters functional groups upon UV

493 irradiation. This fast decay could be in contradiction with their high abundances in CM. This
494 may have at least three different explanations, and possibly a combination of them;

495 1- The organic matter could be shielded in the protosolar nebula either by silicates or self-496 shielded. This phenomenon may be strong enough to, at least partially, shield the carboxylic 497 function until their inclusion in CM parent bodies (Greenberg, 1993; Herbst, 2001). However 498 self-shielding by organic matter in only 1 μm size grains as proposed by Ciesla and Stanford 499 2012 seems unlikely.

500 2- The timescale on which the organic residue would be irradiated may be much shorter than the irradiation time of ~220 kyrs investigated in this paper, due for example to shorter lifetime 501 502 of the PSN or more effective convection in the disk. The advantage of irradiation timescales 503 much shorter than 1 Myrs is that they are consistent with dynamical scenarios predicting the 504 growth of grains and their decoupling from gas within fifty thousands of years of the disk 505 evolution (Weidenschilling and Cuzzi, 1993). It also should be noted that a similar dose could 506 be attained in much shorter irradiation timescales in the PSN if one assumes an enhanced interstellar UV flux. This interstellar flux can be higher by a factor of 10³ compared to the 507 508 nominal one, depending on the possible presence of nearby stars at that time as supported 509 by the presence of certain radio elements in meteorites (Adams et al., 2006; Gounelle and Meynet, 2012). 510

511 3- The carboxylic functions detected in CM meteorites may just simply not be inherited from 512 the organic matter formed in the ISM and altered solely by photoprocesses within the edge of 513 the protosolar nebula as described in Ciesla and Sandford (2012). Instead such material may 514 have been, at least partially, reprocessed during later stages of the solar system formation. A 515 potential pathway for the reprocessing of carboxylic acid could be via the hydrolysis of nitriles

516 functions by aqueous alteration in meteorites parent body (Cronin and Chang, 1993). In regard 517 to these hypotheses it has to be noted that CM meteorites aliphatic hydrocarbons have a 518 consistently lighter δ^{13} C compared to carboxylic groups, pointing either toward an isotopically selective process in the transformation of these materials from ISM to CM parent bodies, or 519 toward the fact that the carboxylic and aliphatic groups of the SOM are derived from 520 521 isotopically-distinct reservoirs (Aponte et al., 2019; Sephton, 2002; Yuen et al., 1984). However, studies performed on the Insoluble Organic Matter of CM meteorites have shown 522 523 that the carboxylic group only represents a minor fraction of this material and is narrowly affected by aqueous alteration (Cody and Alexander, 2005; Orthous-Daunay et al., 2013). 524 Aponte et al., (2019) also extensively discussed the carbon isotopic fractionation of carboxylic 525 526 acid in chondrites. They found decreasing concentrations of monocarboxylic acid with 527 increasing processing (aqueous and thermal) and that the least oxidized CR chondrites contain higher concentration of carboxylic acids that CM, CO, CV and CK chondrites suggesting that if 528 formed in parent bodies "The accretion of aliphatic precursors of monocarboxylic acids was 529 not homogeneous throughout different parent bodies, or that the abundance of 530 531 monocarboxylic acids was mostly affected by parent body processes". Again we emphasize on 532 the need for further experimental data on the kinetics of carboxylic function degradation in the PSN until the accretion in parent bodies. 533

534 <u>4.</u> Conclusions

535 We used laboratory experiments to obtain deeper information on the reactivity occurring 536 during the processing of icy grains. Ice analogs are submitted to processes endured by ices 537 when they evolve within a proto-planetary disk, meaning in our experimental conditions UV 538 irradiation at Lyman α and heating. Organic residues obtained in vacuum at 300 K are then

539 analyzed by very high-resolution mass spectrometry (HRMS) using a LTQ-XL-orbitrap. HRMS 540 analyses provide invaluable information on the molecular content of each residue. Data obtained allow comparing the residue evolution for residues formed from different ice 541 compositions (variability of H₂O and NH₃ in the initial ice). We particularly show that nitrogen 542 543 atoms take an important role in the reactivity by competing with the carbon, the resulting 544 composition of residues depending on the initial amount of ammonia in the ice. Higher relative 545 amounts of ammonia in the ices for a same ratio H₂O/CH₃OH leads to residue enrichment in 546 nitrogen and to the formation of higher unsaturated structures (nitrogen in cycles). It was previously shown with FT-IR analyses and elemental analyses by HRMS that a higher amount 547 of water in the initial ice also leads to an enrichment of residues in nitrogen. We confirm these 548 549 results and by going deeper in the molecular content of the residues, we emphasize that this 550 similarity is only the tip of the iceberg. In fact, similarities between residues coming from nitrogen or water rich ices are due to the prevalence of CHNO molecules forming residues that 551 552 are indeed similar. However, our present investigation shows that a difference exists in molecules bearing only CHN atoms. In ammonia rich ices, this group is only unsaturated while 553 554 in water rich ones, it is mainly saturated. Here again, the reactivity of ammonia and water 555 explains these results, since ammonia drives the residue formation toward aromatization of CHN molecules while water rich ice tends to form aliphatic rich residues. Finally, we 556 557 demonstrate that residue over-irradiation at 300 K leads to a quick decarboxylation of some 558 of the molecules. This result could have an important impact in our understanding of 559 protosolar nebula timescales and of the molecular processes leading to the organic matter 560 detected in meteorites.

561

562 Acknowledgments

The authors wish to thank P.A. Girardi for proof-reading the language of this manuscript. T.G. thanks CNES for a postdoctoral grant during part of this work. G.D. thanks the ANR support from the grant RAHIIA_SSOM (ANR-16-CE29-0015), the CNES from its exobiology program, as well as the CNRS from the PCMI and PNP programs. L.L.S.H. thanks the French national space agency (CNES) for its constant support in the MICMOC experiment at the origin of these residues.

569 References

- 570 Abou Mrad, N., Duvernay, F., Isnard, R., Chiavassa, T., Danger, G., 2017. The Gaseous Phase
- as a Probe of the Astrophysical Solid Phase Chemistry. Astrophys. J. 846, 124.
- 572 https://doi.org/10.3847/1538-4357/aa7cf0
- Adams, F.C., Proszkow, E.M., Fatuzzo, M., Myers, P.C., 2006. Early Evolution of Stellar Groups
- and Clusters: Environmental Effects on Forming Planetary Systems. Astrophys. J. 641,
- 575 504–525. https://doi.org/10.1086/500393
- 576 Aponte, J.C., Woodward, H.K., Abreu, N.M., Elsila, J.E., Dworkin, J.P., 2019. Molecular
- 577 distribution, ¹³ C-isotope, and enantiomeric compositions of carbonaceous chondrite
- 578 monocarboxylic acids. Meteorit. Planet. Sci. 54, 415–430.
- 579 https://doi.org/10.1111/maps.13216
- 580 Bernstein, D.I., Sandford, S.A., Allamandola, L.J., Chang, S., Bernstein, M., Scharberg, M.A.,
- 581 1995. Organic compounds produced by photolysis of realistic interstellar and cometary
- ice analogs containing methanol. Astrophys. J. 454, 327–344.
- 583 Bernstein, M.P., Sandford, S.A., Allamandola, L.J., Chang, S., Scharberg, M.A., 1995. Organic

584	Compounds Produced by Photolysis of Realistic Interstellar and Cometary Ice Analogs
585	Containing Methanol. Astrophys. J. 454, 327. https://doi.org/10.1086/176485
586	Boogert, A.C.A., Pontoppidan, K.M., Knez, C., Lahuis, F., Kessler-Silacci, J., van Dishoeck, E.F.,
587	Blake, G.A., Augereau, JC., Bisschop, S.E., Bottinelli, S., Brooke, T.Y., Brown, J., Crapsi,
588	A., Evans II, N.J., Fraser, H.J., Geers, V., Huard, T.L., Jørgensen, J.K., Öberg, K.I., Allen,
589	L.E., Harvey, P.M., Koerner, D.W., Mundy, L.G., Padgett, D.L., Sargent, A.I., Stapelfeldt,
590	K.R., 2008. The c2d Spitzer Spectroscopic Survey of Ices around Low-Mass Young Stellar
591	Objects. I. H $_2$ O and the 5–8 μm Bands1,2. Astrophys. J. 678, 985–1004.
592	https://doi.org/10.1086/533425
593	Briggs, R., Ertem, G., Ferris, J.P., Greenberg, J.M., McCain, P.J., Mendoza-Gomez, C.X.,
594	Schutte, W., 1992. Comet Halley as an aggregate of interstellar dust and further
595	evidence for the photochemical formation of organics in the interstellar medium. Orig.
596	Life Evol. Biosph. 22, 287–307. https://doi.org/10.1007/BF01810858
597	Burton, A.S., Glavin, D.P., Elsila, J.E., Dworkin, J.P., Jenniskens, P., Yin, QZ., 2014. The amino
598	acid composition of the Sutter's Mill CM2 carbonaceous chondrite. Meteorit. Planet.
599	Sci. 49, 2074–2086. https://doi.org/10.1111/maps.12281
600	Callahan, M.P., Martin, M.G., Burton, A.S., Glavin, D.P., Dworkin, J.P., 2014. Amino acid
601	analysis in micrograms of meteorite sample by nanoliquid chromatography-high-
602	resolution mass spectrometry. J. Chromatogr. A 1332, 30–34.
603	https://doi.org/10.1016/j.chroma.2014.01.032
604	Ciesla, F.J., Sandford, S.A., 2012. Organic Synthesis via Irradiation and Warming of Ice Grains
605	in the Solar Nebula. Science (80). 336.

606	Clark, R.N., Carlson, R., Grundy, W., Noll, K., 2013. Observed Ices in the Solar System.
607	Springer, New York, NY, pp. 3–46. https://doi.org/10.1007/978-1-4614-3076-6_1
608	Cody, G.D., Alexander, C.M.O. 'D., 2005. NMR studies of chemical structural variation of
609	insoluble organic matter from different carbonaceous chondrite groups. Geochim.
610	Cosmochim. Acta 69, 1085–1097. https://doi.org/10.1016/J.GCA.2004.08.031
611	Cooper, G., Kimmich, N., Belisle, W., Sarinana, J., Brabham, K., Garrel, L., 2001. Carbonaceous
612	meteorites as a source of sugar-related organic compounds for the early Earth. Nature
613	414, 879–883. https://doi.org/10.1038/414879a
614	Cronin, J.R., Chang, S., 1993. Organic Matter in Meteorites: Molecular and Isotopic Analyses
615	of the Murchison Meteorite, in: The Chemistry of Life's Origins. Springer Netherlands,
616	Dordrecht, pp. 209–258. https://doi.org/10.1007/978-94-011-1936-8_9
617	Cronin, J.R., Moore, C.B., 1971. Amino Acid Analyses of the Murchison, Murray, and Allende
618	Carbonaceous Chondrites. Science (80). 172, 1327–1329.
619	https://doi.org/10.1126/science.172.3990.1327
620	Danger, G., Fresneau, A., Abou Mrad, N., de Marcellus, P., Orthous-Daunay, FR., Duvernay,
621	F., Vuitton, V., Le Sergeant d'Hendecourt, L., Thissen, R., Chiavassa, T., 2016. Insight into
622	the molecular composition of laboratory organic residues produced from
623	interstellar/pre-cometary ice analogues using very high resolution mass spectrometry.
624	Geochim. Cosmochim. Acta 189, 184–196. https://doi.org/10.1016/j.gca.2016.06.014
625	Danger, G., Orthous-Daunay, FR., de Marcellus, P., Modica, P., Vuitton, V., Duvernay, F.,
626	Flandinet, L., Le Sergeant d'Hendecourt, L., Thissen, R., Chiavassa, T., 2013.
627	Characterization of laboratory analogs of interstellar/cometary organic residues using

- 628 very high resolution mass spectrometry. Geochim. Cosmochim. Acta 118, 184–201.
- 629 https://doi.org/10.1016/j.gca.2013.05.015
- de Marcellus, P., Fresneau, A., Brunetto, R., Danger, G., Duvernay, F., Meinert, C.,
- 631 Meierhenrich, U.J., Borondics, F., Chiavassa, T., d'Hendecourt, L.L.S., 2017. Photo and
- 632 thermochemical evolution of astrophysical ice analogues as a source for soluble and
- 633 insoluble organic materials in Solar system minor bodies. Mon. Not. R. Astron. Soc.
- 634 https://doi.org/10.1093/mnras/stw2292
- de Marcellus, P., Meinert, C., Myrgorodska, I., Nahon, L., Buhse, T., d'Hendecourt, L.L.S.,
- 636 Meierhenrich, U.J., 2015. Aldehydes and sugars from evolved precometary ice analogs:
- 637 Importance of ices in astrochemical and prebiotic evolution. Proc. Natl. Acad. Sci.
- 638 https://doi.org/10.1073/pnas.1418602112
- 639 De Sanctis, M.C., Ammannito, E., McSween, H.Y., Raponi, A., Marchi, S., Capaccioni, F.,
- 640 Capria, M.T., Carrozzo, F.G., Ciarniello, M., Fonte, S., Formisano, M., Frigeri, A.,
- Giardino, M., Longobardo, A., Magni, G., McFadden, L.A., Palomba, E., Pieters, C.M.,
- Tosi, F., Zambon, F., Raymond, C.A., Russell, C.T., 2017. Localized aliphatic organic
- 643 material on the surface of Ceres. Science 355, 719–722.
- 644 https://doi.org/10.1126/science.aaj2305
- Gautier, T., Carrasco, N., Schmitz-Afonso, I., Touboul, D., Szopa, C., Buch, A., Pernot, P., 2014.
- 646 Nitrogen incorporation in titan's tholins inferred by high resolution orbitrap mass
- 647 spectrometry and gas chromatography-mass spectrometry. Earth Planet. Sci. Lett. 404,
- 648 33–42. https://doi.org/10.1016/j.epsl.2014.07.011
- 649 Goesmann, F., Rosenbauer, H., Bredehoeft, J.H., Cabane, M., Ehrenfreund, P., Gautier, T.,
- 650 Giri, C., Krueger, H., Le Roy, L., MacDermott, A.J., McKenna-Lawlor, S., Meierhenrich,

651	U.J., Munoz Caro, G.M., Raulin, F., Roll, R., Steele, A., Steininger, H., Sternberg, R.,
652	Szopa, C., Thiemann, W., Ulamec, S., 2015. Organic compounds on comet
653	67P/Churyumov-Gerasimenko revealed by COSAC mass spectrometry. Science (80).
654	https://doi.org/10.1126/science.aab0689
655	Gounelle, M., Meynet, G., 2012. Solar system genealogy revealed by extinct short-lived
656	radionuclides in meteorites. https://doi.org/10.1051/0004-6361/201219031
657	Greenberg, J.M., 1993. Physical and Chemical Composition of Comets — From Interstellar
658	Space to the Earth, in: The Chemistry of Life's Origins. Springer Netherlands, Dordrecht,
659	pp. 195–207. https://doi.org/10.1007/978-94-011-1936-8_8
660	Habing, H.J., 1968. The interstellar radiation density between 912 A and 2400 A. Bull. Astron.
661	Institutes Netherlands 19, 421.
662	Henderson, B.L., Gudipati, M.S., 2015. DIRECT DETECTION OF COMPLEX ORGANIC PRODUCTS
663	IN ULTRAVIOLET (Ly α) AND ELECTRON-IRRADIATED ASTROPHYSICAL AND COMETARY
664	ICE ANALOGS USING TWO-STEP LASER ABLATION AND IONIZATION MASS
665	SPECTROMETRY. Astrophys. J. 800, 66. https://doi.org/10.1088/0004-637X/800/1/66
666	Herbst, E., 2001. The chemistry of interstellar space. Chem. Soc. Rev. 30, 168–176.
667	https://doi.org/10.1039/a909040a
668	Hertkorn, N., Frommberger, M., Witt, M., Koch, B.P., Schmitt-Kopplin, P., Perdue, E.M., 2008.

- 669 Natural Organic Matter and the Event Horizon of Mass Spectrometry. Anal. Chem. 80,
- 670 8908–8919. https://doi.org/10.1021/ac800464g
- Holčapek, M., Volná, K., Vaněrková, D., 2007. Effects of functional groups on the
- 672 fragmentation of dyes in electrospray and atmospheric pressure chemical ionization

673 mass spectra. Dye. Pigment. 75, 156–165.

- 674 https://doi.org/10.1016/J.DYEPIG.2006.05.040
- Howarth, R.J., 1996. Sources for a history of the ternary diagram. Br. J. Hist. Sci. 29, 337.
 https://doi.org/10.1017/S000708740003449X
- 677 Kim, S., Kramer, R.W., Hatcher, P.G., 2003. Graphical Method for Analysis of Ultrahigh-
- 678 Resolution Broadband Mass Spectra of Natural Organic Matter, the Van Krevelen
- 679 Diagram. Anal. Chem. 75, 5336–5344. https://doi.org/10.1021/ac034415p
- 680 Kruve, A., Kaupmees, K., Liigand, J., Leito, I., 2014. Negative Electrospray Ionization via
- 681 Deprotonation: Predicting the Ionization Efficiency. Anal. Chem. 86, 4822–4830.
- 682 https://doi.org/10.1021/ac404066v
- Lawless, J.G., Yuen, G.U., 1979. Quantification of monocarboxylic acids in the Murchison
- carbonaceous meteorite. Nature 282, 396–398. https://doi.org/10.1038/282396a0
- Martins, Z., Modica, P., Zanda, B., d'Hendecourt, L.L.S., 2015. The amino acid and
- 686 hydrocarbon contents of the Paris meteorite: Insights into the most primitive CM
- 687 chondrite. Meteorit. Planet. Sci. 50, 926–943. https://doi.org/10.1111/maps.12442
- Meech, K.J., Svoren, J., 2004. Using cometary activity to trace the physical and chemical
 evolution of cometary nuclei. Comets II 317.
- 690 Meinert, C., Myrgorodska, I., de Marcellus, P., Buhse, T., Nahon, L., Hoffmann, S. V,
- d'Hendecourt, L.L.S., Meierhenrich, U.J., 2016. Ribose and related sugars from
- 692 ultraviolet irradiation of interstellar ice analogs. Science 352, 208–12.
- 693 https://doi.org/10.1126/science.aad8137
- 694 Muñoz Caro, G.M., Schutte, W.A., 2003. UV-photoprocessing of interstellar ice analogs: New

infrared spectroscopic results. Astron. Astrophys. https://doi.org/10.1051/0004-

696 6361:20031408

- Newton, I., 1704. Opticks: Or, a treatise of the reflexions, refractions, inflexions and colours
 of light. R. Br. Soc. https://doi.org/10.1007/s13398-014-0173-7.2
- 699 Nuevo, M., Meierhenrich, U.J., d'Hendecourt, L., Muñoz Caro, G.M., Dartois, E., Deboffle, D.,
- 700 Thiemann, W.H.-P., Bredehöft, J.-H., Nahon, L., 2007. Enantiomeric separation of
- 701 complex organic molecules produced from irradiation of interstellar/circumstellar ice
- analogs. Adv. Sp. Res. 39, 400–404. https://doi.org/10.1016/j.asr.2005.05.011
- Nuevo, M., Milam, S.N., Sandford, S.A., De Gregorio, B.T., Cody, G.D., Kilcoyne, A.L.D., 2011.
- 704 XANES analysis of organic residues produced from the UV irradiation of astrophysical

705 ice analogs. Adv. Sp. Res. 48, 1126–1135. https://doi.org/10.1016/j.asr.2011.05.020

- 706 Orthous-Daunay, F.-R., Quirico, E., Beck, P., Brissaud, O., Dartois, E., Pino, T., Schmitt, B.,
- 707 2013. Mid-infrared study of the molecular structure variability of insoluble organic
- matter from primitive chondrites. Icarus 223, 534–543.
- 709 https://doi.org/10.1016/J.ICARUS.2013.01.003
- 710 Pernot, P., Carrasco, N., Thissen, R., Schmitz-Afonso, I., 2010. Tholinomics—Chemical

711 Analysis of Nitrogen-Rich Polymers. Anal. Chem. 82, 1371–1380.

- 712 https://doi.org/10.1021/ac902458q
- Piani, L., Tachibana, S., Hama, T., Tanaka, H., Endo, Y., Sugawara, I., Dessimoulie, L., Kimura,
- 714 Y., Miyake, A., Matsuno, J., Tsuchiyama, A., Fujita, K., Nakatsubo, S., Fukushi, H., Mori,
- 715 S., Chigai, T., Yurimoto, H., Kouchi, A., 2017. Evolution of Morphological and Physical
- 716 Properties of Laboratory Interstellar Organic Residues with Ultraviolet Irradiation.

- 717 Astrophys. J. 837, 35. https://doi.org/10.3847/1538-4357/aa5ca6
- 718 Pizzarello, S., Shock, E., 2010a. The organic composition of carbonaceous meteorites: the
- evolutionary story ahead of biochemistry. Cold Spring Harb. Perspect. Biol. 2, a002105.
- 720 https://doi.org/10.1101/cshperspect.a002105
- 721 Pizzarello, S., Shock, E., 2010b. The organic composition of carbonaceous meteorites: the
- evolutionary story ahead of biochemistry. Cold Spring Harb. Perspect. Biol.
- 723 https://doi.org/10.1101/cshperspect.a002105
- Ruf, A., Kanawati, B., Hertkorn, N., Yin, Q.-Z., Moritz, F., Harir, M., Lucio, M., Michalke, B.,
- 725 Wimpenny, J., Shilobreeva, S., Bronsky, B., Saraykin, V., Gabelica, Z., Gougeon, R.D.,
- 726 Quirico, E., Ralew, S., Jakubowski, T., Haack, H., Gonsior, M., Jenniskens, P., Hinman,
- 727 N.W., Schmitt-Kopplin, P., 2017. Previously unknown class of metalorganic compounds
- revealed in meteorites. Proc. Natl. Acad. Sci. U. S. A. 114, 2819–2824.
- 729 https://doi.org/10.1073/pnas.1616019114
- 730 Sephton, M.A., 2002. Organic compounds in carbonaceous meteorites. Nat. Prod. Rep. 19,
- 731 292–311. https://doi.org/10.1039/b103775g
- 732 Somogyi, Á., Smith, M.A., Vuitton, V., Thissen, R., Komáromi, I., 2012. Chemical ionization in
- the atmosphere? A model study on negatively charged "exotic" ions generated from
- 734 Titan's tholins by ultrahigh resolution MS and MS/MS. Int. J. Mass Spectrom. 316–318,
- 735 157–163. https://doi.org/10.1016/J.IJMS.2012.02.026
- 736 Tachibana, S., Kouchi, A., Hama, T., Oba, Y., Piani, L., Sugawara, I., Endo, Y., Hidaka, H.,
- 737 Kimura, Y., Murata, K., Yurimoto, H., Watanabe, N., 2017. Liquid-like behavior of UV-
- 738 irradiated interstellar ice analog at low temperatures. Sci. Adv. 3, eaao2538.

https://doi.org/10.1126/sciadv.aao2538

740	van Krevelen, D.W., 1950. Graphical-statistical method for the study of structure and
741	reaction processes of coal. Fuel. https://doi.org/10.1520/D0850-11.1
742	Weidenschilling, S.J., Cuzzi, J.N., 1993. Formation of Planetesimals in the Solar Nebula.
743	Protostars Planets III 1031.
744	Wu, Z., Rodgers, R.P., Marshall, A.G., 2004. Two- and Three-Dimensional van Krevelen
745	Diagrams: A Graphical Analysis Complementary to the Kendrick Mass Plot for Sorting
746	Elemental Compositions of Complex Organic Mixtures Based on Ultrahigh-Resolution
747	Broadband Fourier Transform Ion Cyclotron Resonance Mass Measurements. Anal.
748	Chem. 76, 2511–2516. https://doi.org/10.1021/ac0355449
749	Yassine, M.M., Harir, M., Dabek-Zlotorzynska, E., Schmitt-Kopplin, P., 2014. Structural
750	characterization of organic aerosol using Fourier transform ion cyclotron resonance
751	mass spectrometry: Aromaticity equivalent approach. Rapid Commun. Mass Spectrom.
752	28, 2445–2454. https://doi.org/10.1002/rcm.7038
753	Yuen, G., Blair, N., Des Marais, D.J., Chang, S., 1984. Carbon isotope composition of low
754	molecular weight hydrocarbons and monocarboxylic acids from Murchison meteorite.
755	Nature 307, 252–254. https://doi.org/10.1038/307252a0
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757	Supplementary information

758 Attributor Data Processing

759 The first processing step is to remove the instrument noise by deleting the less intense peaks 760 until reaching a convex intensity distribution. Radio interference and low level noises are also reduced by isolating structures of interest for the sample in a mass defect versus exact mass 761 762 (MDvEM) space representation as depicted in Danger et al. 2013. This representation is also 763 used to detect and eliminate the ringing effects, visible as vertical lines in the MDvEM. Once 764 this is done, the software locates the centroids of each Gaussian peak in the spectra and 765 reduces each peak to its centroid, greatly reducing the size of the data. The same process is 766 applied to the blank spectra (*i.e.* same experimental protocol with no sample in the solvent). 767 Peaks found to be present in the blank spectra are then removed from the sample mass 768 spectra.

769 The data is then post-calibrated. In Attributor, this is done by looking at the distribution of the 770 mass differences between every single peak of the spectra. Typically, for organic samples, one 771 expects to find pattern (ϵ) differences at Δm 2, 12 or 14, corresponding to a difference in H₂, 772 C or CH_2 in the composition of the molecules detected. For a given pattern ε and in an ideal 773 case the mass difference between molecules X and X+ ε should not depend of the mass of X. 774 In reality, when looking at the spectra, we can see multiple pairs of points for which the mass 775 differences are close to ε but differ with the masses of X, meaning a bias on the linearity of 776 the measure. To correct this bias, Attributor will check for two correction points corresponding 777 to trusted molecules. These molecules correspond to 2 peaks chosen by attributor for which a molecular formula can be unambiguously attributed. Typically, these correspond to intense 778 peaks for which 1/ the isotopic patterns with ¹³C are clearly identified in the mass spectra and 779 780 2/ have univocal possible attribution (the second closest possible attribution would have an 781 error in ppm far superior to the closest possible attribution). Attributor then give a list of possible attributions for these two trusted molecules that can be checked by the operator 782

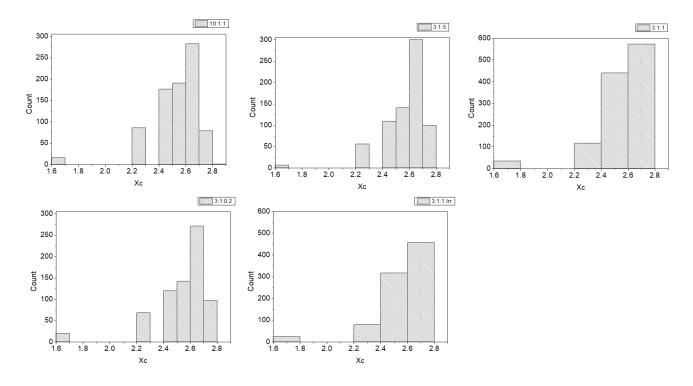
before proceeding to the data recalibration. Once these two trusted molecules identified, a recalibration coefficient is applied to all the peaks in the mass spectra. If needed, this operation can be repeated until all attributions in the mass spectra are within a 2ppm error window.

Once the data calibrated, the software attributes if possible each peak to a molecular formula. 787 788 The matrix of list of atoms, their isotopes, and number boundaries to be considered for the 789 final attribution are chosen by the operator; for the present work, given our sample composition and the mass range investigated we used only the major isotopes (¹H, ¹²C, ¹⁴N, 790 791 ¹⁶O) and arbitrarily limited the number of C atom to be between 1 and 30, the number of H 792 atoms between 3 and 90, the number of N atoms between 0 and 15, and the number of O 793 atoms between 0 and 15. Once the constraints given to the software, it computes for each 794 peak m/z a list of combinations that span at least m/z-1 to m/z+1 of selected blocks so that 795 the nitrogen rule is verified and keep attributions. Attribution also takes into account the 796 charge of the ions and mass of the hydrogen added or loss during the ESI ionization. The 797 matches within a 5 ppm windows are kept for further investigation. Once the attribution done, the software checks for the presence of isotopic pattern. For this work we considered every 798 799 isotopic peaks expected above a hundredth of the intensity of the base peak and excludes all 800 possible attribution that do not pass this criterion. The remaining best match (i.e. with the 801 nominal mass the closest of the measured mass) is kept and considered attributed if the error between nominal and measured mass is under 5 ppm (typical maximal differences in our 802 measurements were ±2 ppm). For each peak whose attribution has been completed, the 803 804 software outputs the number of C, H, N and O atoms and the error in ppm between the 805 nominal mass of the attributed formula and the measured one. If calibration has been successful, the error in ppm between attribution and nominal masses should be centered 806

around 0 and not depend of the mass of the molecules. The program thus checks for a tendency in the error, and if a bias is observed, a new calibration-attribution cycle is started.

This step allowed to attribute molecular formulae to compound detected in the mass spectra up to m/z 400. A molecule was considered to be present in the sample only if it was detected in all replicates of the experiment. The list of molecules detected in the sample and used to produce figures presented in this paper includes both the molecules detected in positive and negative ionization modes.

814 Supplementary Figures:



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816 Figure SI 1 : Histogram of Xc for all the molecules detected in the samples.

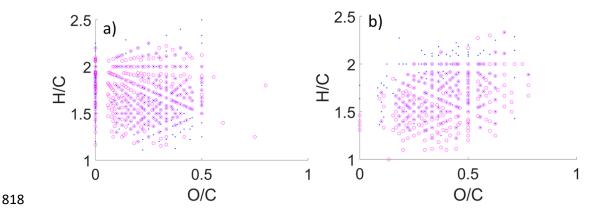
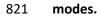


Figure SI 2 :H/C vs O/C van Krevelen plot of the residues formed with 10:1:1 (pink circles) and 3:1:5 (blue dots) H₂O:CH₃OH:NH₃ ices for molecules detected in positive (a) and negative (b) ionization



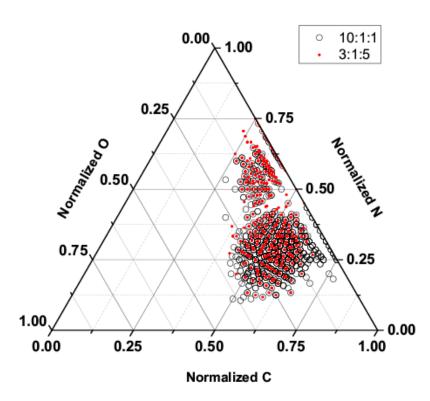
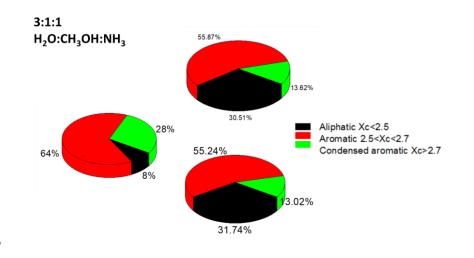
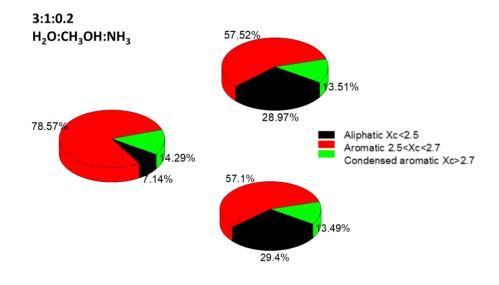


Figure SI 3: Ternary plot of the residues formed with 10:1:1 (black circles) and 3:1:5 (red dots) H₂O:CH₃OH:NH₃ ices

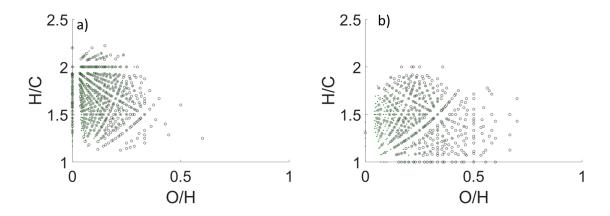


826 Figure SI 4 : Repartition of the molecules detected in the 3:1:1 sample depending on their level of 827 unsaturation based on the evolution of the aromaticity equivalent (Xc) for the totality of the 828 molecules detected in the sample (whole distribution, top chart), the molecules containing C,H,N 829 and O atoms (CHNO group, bottom chart) and the molecules solely containing C, H and N atoms 830 (CHN group, left chart). The fraction of molecules classified as aliphatic is shown in black, red 831 represent the fraction classified as aromatic and green the fraction considered to be condensed 832 aromatic molecules. The percentage corresponds to the percentage of the molecules classified 833 according to a given aromaticity relatively to the entire distribution. Given the number of classified 834 molecules, we can make a conservative estimate of $\pm 5\%$ as the 2σ uncertainties associated with this classification for the whole distribution. 835



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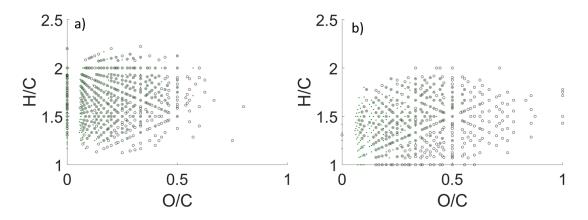
837 Figure SI 5 Repartition of the molecules detected in the 3:1:0.2 sample depending on their level of 838 unsaturation based on the evolution of the aromaticity equivalent (Xc) for the totality of the 839 molecules detected in the sample (whole distribution, top chart), the molecules containing C,H,N 840 and O atoms (CHNO group, bottom chart) and the molecules solely containing C, H and N atoms 841 (CHN group, left chart). The fraction of molecules classified as aliphatic is shown in black, red 842 represent the fraction classified as aromatic and green the fraction considered to be condensed 843 aromatic molecules. The percentage correspond to the percentage of the molecules classified 844 according to a given aromaticity relatively to the entire distribution. Given the number of classified 845 molecules, we can make a conservative estimate of $\pm 5\%$ as the 2σ uncertainties associated with this 846 classification for the whole distribution



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Figure SI 6: O/H vs H/C modified van Krevelen diagram of the molecules detected in the 3:1:1
 sample in positive (a) and negative (b) ionization mode before (black circles) and after (green points) the sample irradiation.





- 853 Figure SI 7: H/C vs O/C van Krevelen diagram of the molecules detected in the 3:1:1 sample in
- positive (a) and negative (b) ionization mode before (black circles) and after (green dots) the
- 855 sample irradiation.

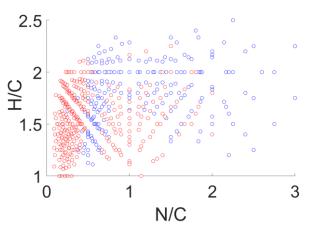


Figure SI 8 : H/C vs N/C modified van Krevelen diagram of the molecules detected in the 3:1:5 (blue circles) and 3:1:0.2 (red circles) samples.