



HAL
open science

Variation of the BrO/SO₂ molar ratio in the plume of Tungurahua volcano between 2007 and 2017 and its relationship to volcanic activity

Simon Warnach, Nicole Bobrowski, Silvana Hidalgo, Santiago Arellano, Holger Sihler, Florian Dinger, Peter Lübcke, Jean Battaglia, Alexander Steele, Bo Galle, et al.

► To cite this version:

Simon Warnach, Nicole Bobrowski, Silvana Hidalgo, Santiago Arellano, Holger Sihler, et al.. Variation of the BrO/SO₂ molar ratio in the plume of Tungurahua volcano between 2007 and 2017 and its relationship to volcanic activity. *Frontiers in Earth Science*, 2019, 7, pp.132. 10.3389/feart.2019.00132 . hal-02376800

HAL Id: hal-02376800

<https://hal.science/hal-02376800>

Submitted on 7 Jan 2021

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License



Variation of the BrO/SO₂ Molar Ratio in the Plume of Tungurahua Volcano Between 2007 and 2017 and Its Relationship to Volcanic Activity

Simon Warnach^{1,2*}, Nicole Bobrowski^{1,2}, Silvana Hidalgo³, Santiago Arellano⁴, Holger Sihler^{1,2}, Florian Dinger^{1,2}, Peter Lübcke², Jean Battaglia⁵, Alexander Steele⁶, Bo Galle⁴, Ulrich Platt^{1,2} and Thomas Wagner¹

¹ Max Planck Institute for Chemistry, Satellite Remote Sensing, Mainz, Germany, ² Institute of Environmental Physics, University of Heidelberg, Heidelberg, Germany, ³ Instituto Geofísico, Escuela Politécnica Nacional, Quito, Ecuador, ⁴ Department of Space, Earth and Environment, Chalmers University of Technology, Gothenburg, Sweden, ⁵ Laboratoire Magmas et Volcans, Observatoire de Physique du Globe de Clermont-Ferrand (OPGC), IRD, CNRS, Université Clermont Auvergne, Aubière, France, ⁶ UCL Hazard Centre, Department of Earth Sciences, University College London, London, United Kingdom

OPEN ACCESS

Edited by:

Alessandro Aiuppa,
University of Palermo, Italy

Reviewed by:

Mattia Pistone,
Université de Lausanne, Switzerland
Stefan Bredemeyer,
Helmholtz Centre Potsdam, Germany

*Correspondence:

Simon Warnach
s.warnach@mpic.de

Specialty section:

This article was submitted to
Volcanology,
a section of the journal
Frontiers in Earth Science

Received: 31 July 2018

Accepted: 14 May 2019

Published: 02 July 2019

Citation:

Warnach S, Bobrowski N, Hidalgo S, Arellano S, Sihler H, Dinger F, Lübcke P, Battaglia J, Steele A, Galle B, Platt U and Wagner T (2019) Variation of the BrO/SO₂ Molar Ratio in the Plume of Tungurahua Volcano Between 2007 and 2017 and Its Relationship to Volcanic Activity. *Front. Earth Sci.* 7:132. doi: 10.3389/feart.2019.00132

Recent long-term observations of the bromine monoxide (BrO) to sulphur dioxide (SO₂) molar ratio in volcanic plumes have suggested a link between changes in the BrO/SO₂ ratio and the volcanic activity. Nevertheless, understanding of the mechanisms determining this link is still limited due to the lack of studies on volcanic bromine release from the melt into the atmosphere. We present the results of 10 years (2007–2017) of observations of the BrO/SO₂ molar ratio in the volcanic plume of Tungurahua volcano, Ecuador. Following the nearly continuous eruptive activity from 1999 to 2008, Tungurahua showed alternating phases of eruptive activity separated by periods of quiescence between late 2008 and March 2016, after which degassing intensity decreased below detection. By comparing the BrO/SO₂ molar ratios collected from 13 eruptive phases to volcanic activity, this study aims to broaden the global observational database investigating their link. For this purpose, we combine three different methods to retrieve the BrO/SO₂ molar ratio to analyse variations over different timescales. We identify a cyclic pattern in BrO/SO₂ molar ratios for 11 of the 13 eruptive phases. The phases are initialised by low BrO/SO₂ molar ratios between 2 and 6 × 10⁻⁵ coinciding with vulcanian-type activity followed by a strong increase to ratios ranging between 4 and 17 × 10⁻⁵ when eruptive dynamism shifts to strombolian. For five phases, we additionally observe a progressive decrease to the initial values of 2 to 5 × 10⁻⁵ toward the end of the phase. This clear pattern indicates a connection between the BrO/SO₂ molar ratio and eruptive dynamics. Based on our new data, we propose a conceptual model of the volcanic processes taking place at Tungurahua during the eruptive phases. Our data furthermore indicate that maximal BrO/SO₂ molar ratios observed during each phase could be related to the input of volatile-rich magma into the active part of the volcanic system of Tungurahua. This study shows that long-term BrO/SO₂ molar ratios can be used as a proxy for the volatile status as well as temporal evolution of the volcanic system.

Keywords: Tungurahua, SO₂, BrO, ratio, degassing, volcano, DOAS, NOVAC

INTRODUCTION

The molar ratios of gas species in volcanic plumes offer crucial information about volcanic processes. Sulphur dioxide (SO₂) emissions have in some cases proven to be a good forecasting tool for volcanic eruptions (e.g. Malinconico, 1979; Olmos et al., 2007; Hidalgo et al., 2018). Moreover, the halogen to sulphur ratios have been used for detecting changes in volcanic activity due to the different melt-fluid partitioning coefficients of these elements (Noguchi and Kamiya, 1963; Carrol and Holloway, 1994).

Changes in the chlorine to sulphur ratio have been suggested to forecast eruptions (Menyailov, 1975) and to distinguish different types of activity (Pennisi and Cloarec, 1998; Allard et al., 2005; Aiuppa, 2009; Spina et al., 2015). Several studies interpreting these measurements advocated preferred release of chlorine and fluorine at low pressures corresponding to shallow depth beneath the volcanic vent (Noguchi and Kamiya, 1963; Villemant and Boudon, 1999; Aiuppa et al., 2002; Burton et al., 2007; Edmonds et al., 2009). For bromine in contrast, the degassing processes are still not fully understood and lack comparison of measurements with model data. Several studies investigated melt-fluid partitioning of bromine with respect to sulphur using artificial melts (Bureau et al., 2000, 2010; Cochain et al., 2015) and concluded that the partition coefficient highly depends on the melt composition. A recent study using natural melts suggested efficient degassing of bromine at shallow depths beneath the volcanic vent similar to chlorine behaviour (Cadoux et al., 2018).

Since the first detection of bromine monoxide (BrO) inside a volcanic plume (Bobrowski et al., 2003), BrO as well as the BrO/SO₂ molar ratio have been retrieved at various volcanoes using ground-based measurements (e.g. Oppenheimer et al., 2006; Bobrowski and Platt, 2007; Kern et al., 2009; Boichu et al., 2011; Kelly et al., 2013) as well as from satellite (Theys et al., 2009; Hörmann et al., 2013). In addition, bromine chemistry in volcanic plumes has also been investigated through plume-chemistry models (e.g. Bobrowski et al., 2007; von Glasow, 2010; Roberts et al., 2014). A review of the current knowledge of bromine chemistry in volcanic plumes as well as a collection of all volcanic bromine measurements since 2003 was reported by Gutmann et al. (2018).

Recent studies of long-term BrO/SO₂ datasets have revealed changes in connection with varying volcanic activity at Mount Etna (Bobrowski and Giuffrida, 2012) and Nevado del Ruiz (Lübcke et al., 2014), or due to external forcing such as Earth tides (Dinger et al., 2018). The advantage of the BrO/SO₂ molar ratio compared to other halogen/sulphur ratios is that both species can be measured simultaneously via automated remote-sensing using scattered sunlight spectra applying the Differential Optical Absorption Spectroscopy (DOAS) technique (Platt and Stutz, 2008). This allows for continuous long-term data retrieval throughout different states of volcanic activity (Galle et al., 2010).

This study focuses on Tungurahua volcano, an andesitic subduction volcano located 120 km south of Quito, Ecuador. The average magma supply rates for this volcano have been estimated at $\approx 1.5 \times 10^6 \text{ m}^3 \text{ yr}^{-1}$ over the last 2300 years

(Hall et al., 1999). Regular eruptions occurred at least once per century (Pennec et al., 2008; Bablon et al., 2018). Hence, Tungurahua is among the most active volcanoes in the northern Andes in recent times (Hall et al., 1999; Wright et al., 2012). The most recent eruptive cycle started in September 1999. From 1999 until late 2008 Tungurahua exhibited long periods of continuous eruptive activity (Arellano et al., 2008; Samaniego et al., 2011; Hidalgo et al., 2015), with typical SO₂ emission rates of the order of 1000 t/d derived from ground-based measurements (Arellano et al., 2008; Galle et al., 2010) and from satellite (Carn et al., 2008; McCormick et al., 2014). From 2008 to 2016, however, the activity has been characterised by alternating eruptive phases and relative quiescence. Eruptive phases have displayed strombolian to vulcanian explosions, which were accompanied by persistent ash and gas emissions, and sporadic pyroclastic flows and lava effusion (Hidalgo et al., 2015).

We investigate the variation of the BrO/SO₂ molar ratio in the plume of Tungurahua from the spectral data recorded by a network of four autonomous scanning-DOAS stations over the period March 2007 to January 2017. To improve readability, we will refer to all ratios by omitting “molar” in the rest of the manuscript, e.g. BrO/SO₂ ratio. The different eruptive phases are decomposed into different stages, in order to identify characteristic patterns in the BrO/SO₂ ratio. We identify a pattern which was repeatedly observed throughout phases over the whole period of investigation. We interpret it with respect to magmatic processes and volcanic activity and derive a conceptual model of the eruptive dynamic representative for each eruptive phase. Lastly, we also discuss the behaviour of the BrO/SO₂ ratios in relation to different types of explosive activity.

MATERIALS AND METHODS

Measurement Setup

At Tungurahua volcano, four stations were run during the study period by the IG-EPN (Instituto Geofísico – Escuela Politécnica Nacional), all of which operated with NOVAC (Network for Observation of Volcanic and Atmospheric Change) version 1 instruments (Galle et al., 2010). The instruments were located 6–11 km from the volcanic vent, with three stations to the west of the crater, in the direction of the prevalent plume direction, and one to the north-east (for a detailed description of the location of the instruments see Hidalgo et al., 2015).

NOVAC currently comprises automatic scanning DOAS instruments at more than 40 volcanoes. The instruments measure volcanic plume constituents by performing a scan across the sky, which is comprised typically of 51 spectra obtained in steps of 3.6°. The scanning plane intersects the volcanic plume under favourable wind conditions. Furthermore, the intersection of simultaneous measurements by different instruments is used to derive plume height and direction, allowing – with knowledge of the wind speed – for the calculation of fluxes (Galle et al., 2010). However, these variables are not needed to extract BrO/SO₂ ratios, which purely depend on the spectral measurements.

Spectroscopic Retrieval

The amount of SO₂ and BrO in the plume was retrieved via a DOAS analysis comparing the spectra measured in the direction of the volcanic plume with spectra in a direction not intercepting the volcanic plume. To improve the signal-to-noise ratio in this work the co-adding retrieval algorithm was used as described in Lübcke et al. (2014): The algorithm identifies 10 plume spectra (having the highest SO₂ values) and 10 reference spectra (having the lowest SO₂ values) in each scan. The algorithm then co-adds these spectra, respectively from four consecutive scans (i.e. 40 spectra of each type in total). The resulting spectra, which are from now on referred to as “co-added” spectra, are then used as measurement and reference spectrum in the DOAS fitting algorithm, respectively (Platt and Stutz, 2008). This yielded one co-added BrO as well as one SO₂ mean slant column density (SCD) roughly every 30 min.

Fit Parameter

In addition to the co-added plume and reference spectra, the DOAS fitting algorithm also included absorption cross-sections for SO₂ (Vandaele et al., 2009) and O₃ (Burrows et al., 1999) in the SO₂ spectral fit region (314.8–326.8 nm) and SO₂, BrO (Fleischmann et al., 2004), O₄ (Hermans et al., 2003), NO₂ (Vandaele et al., 1998) and CH₂O (Meller and Moortgat, 2000) in the BrO spectral fit region (330.6–352.75 nm) as well as a third-order polynomial to account for broadband absorption structures, Mie, and Rayleigh scattering. Two spectra were included to account for the Ring effect (Grainger and Ring, 1962) – the standard Ring spectrum as well as a wavelength dependent Ring spectrum (derived from the original Ring spectrum by multiplication with λ^4 , where λ represents wavelength; Wagner et al., 2009). The fit settings were chosen as an outcome of the study by Vogel et al. (2013).

The detection limit can be derived as a multiple of the DOAS fit error based on Platt and Stutz (2008). In our study we used three times the fit error – which is usually a quite conservative estimate – yielding here a mean detection limit of 4.9×10^{16} molecules/cm² (SO₂) and 4.9×10^{13} molecules/cm² (BrO). The latter fits very well to the random fluctuations of BrO around zero when SO₂ SCDs are close to zero (i.e. no gas emissions), which similarly span roughly $\pm 5 \times 10^{13}$ molecules/cm². For SO₂ this fluctuation is hardly visible, because the signal is up to two orders of magnitude larger than the detection limit.

Determining the BrO/SO₂ Ratio

The BrO/SO₂ ratio for a series of simultaneous measurements of both species inside a volcanic plume is usually obtained by performing a linear regression for BrO as a function of SO₂ (e.g. Bobrowski et al., 2003, 2007; Bobrowski and Giuffrida, 2012; Hörmann et al., 2013). The linear regression was performed using the error-weighted orthogonal regression algorithm developed by Cantrell (2008).

For this study, it was employed in two ways:

1. Daily regression: This regression takes all co-added BrO and SO₂ data of a particular day to retrieve daily BrO/SO₂

ratios in order to observe short-term changes in the BrO/SO₂ ratio. To ensure the quality of the retrieved BrO/SO₂ ratio, days were considered only if at least one data point exceeded SO₂ SCDs of 5×10^{17} molecules/cm² and the regression fit error was below 1.5×10^{-5} .

2. Period regression: Data acquired over the period of several days/weeks were evaluated using the same methodology as for the daily regression. The beginning and end of these periods were chosen when a change in the daily BrO/SO₂ ratios could be detected. For long periods without significant change, the beginning and end of these periods were additionally chosen when changes in the classification of explosive activity from Hidalgo et al. (2015) occurred. Only those periods where at least five co-added SO₂ SCDs exceeded 5×10^{17} molecules/cm², the maximum SO₂ SCD exceeded 7×10^{17} molecules/cm², and the regression fit error was below 1×10^{-5} were taken into account.

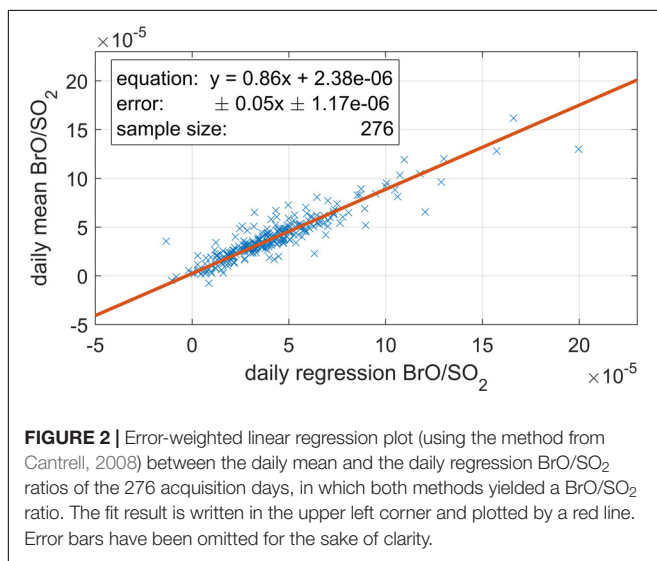
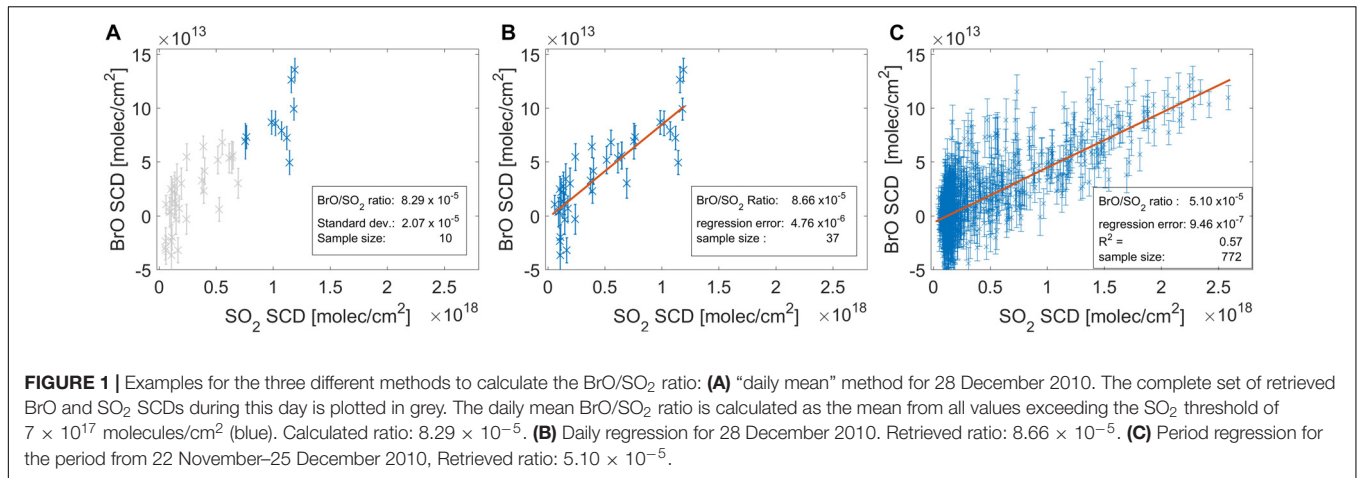
We also calculated the BrO/SO₂ ratio by taking the BrO/SO₂ ratio of every co-added data point with high SO₂ signal (SO₂ SCD > 7×10^{17} molecules/cm²). In a second step, the daily mean BrO/SO₂ ratio was calculated as the error-weighted daily mean over these single BrO/SO₂ ratios. To ensure robust results we calculated the daily mean BrO/SO₂ ratio only if there were at least five data points above the SO₂ threshold. An example of the results obtained by this algorithm for 28 December 2010 is given in **Figure 1A**. This method was used by Lübcke et al. (2014) for NOVAC data and we will refer to it as “daily means” in this study.

In this study, we introduce a scheme including all three methods, the daily means method as well as the two linear regression methods to achieve a sufficient set of BrO/SO₂ ratios. They are presented in **Figure 1**. The daily mean method yielded BrO/SO₂ ratios on 277 days, while the daily regression was successful on 425 days. Of the 276 days, in which both methods yielded a BrO/SO₂ ratio, we observed a good correlation between the two methods. However, the daily mean method consistently yielded slightly lower ratios (we retrieved a slope of 0.86 ± 0.05 ; see **Figure 2**). The results of all daily mean and daily regression BrO/SO₂ ratios as a function of time are plotted in **Figure 3C**.

Volcanic Activity Classification

We used the classification criteria of volcanic activity described by Hidalgo et al. (2015), which includes the eruptive and quiescent phases from 2007 to 2013. We extended this classification in order to cover our study period following the same criteria (i.e. seismo-acoustic records and visual observations).

Three types of activity were distinguished during the study period: (1) Quiescence, (2) Low explosive activity and (3) High explosive activity. Episodes of high explosive activity are those in which discrete vulcanian explosions occurred, with the explosions themselves being classified as vulcanian when their mean acoustic excess pressures were greater than 100 Pa at a distance of 1 km from the vent (Johnson, 2003). Episodes of low explosive activity are those in which no vulcanian explosions were present. These episodes were



instead characterised by Strombolian-type activity, during which the explosions had excess pressures smaller than the 100 Pa threshold (Johnson, 2003). In both cases, explosive activity is commonly accompanied by intense gas and ash emissions. During quiescent episodes, the seismic activity was generally weak and merely a few long period and volcano-tectonic earthquakes were detected.

RESULTS

BrO and SO₂ SCDs From Tungurahua

Over the study period between 2007 and 2017, 18 phases of high SO₂ SCDs (of the order of 10^{18} molecules/cm²) could be identified from the NOVAC data. They were separated by periods of weak SO₂ degassing, where the SO₂ SCDs generally did not exceed 1.6×10^{17} molecules/cm². BrO SCDs exceeded the detection limit of 4.8×10^{13} molecules/cm² only during phases where high SO₂ SCDs were measured. In those phases typical BrO SCDs around 8×10^{13}

molecules/cm² were observed. During the phases of weak SO₂, BrO fluctuated around zero and was below detection. The results of BrO and SO₂ from Tungurahua are displayed in Figures 3A,B.

BrO/SO₂ Ratios

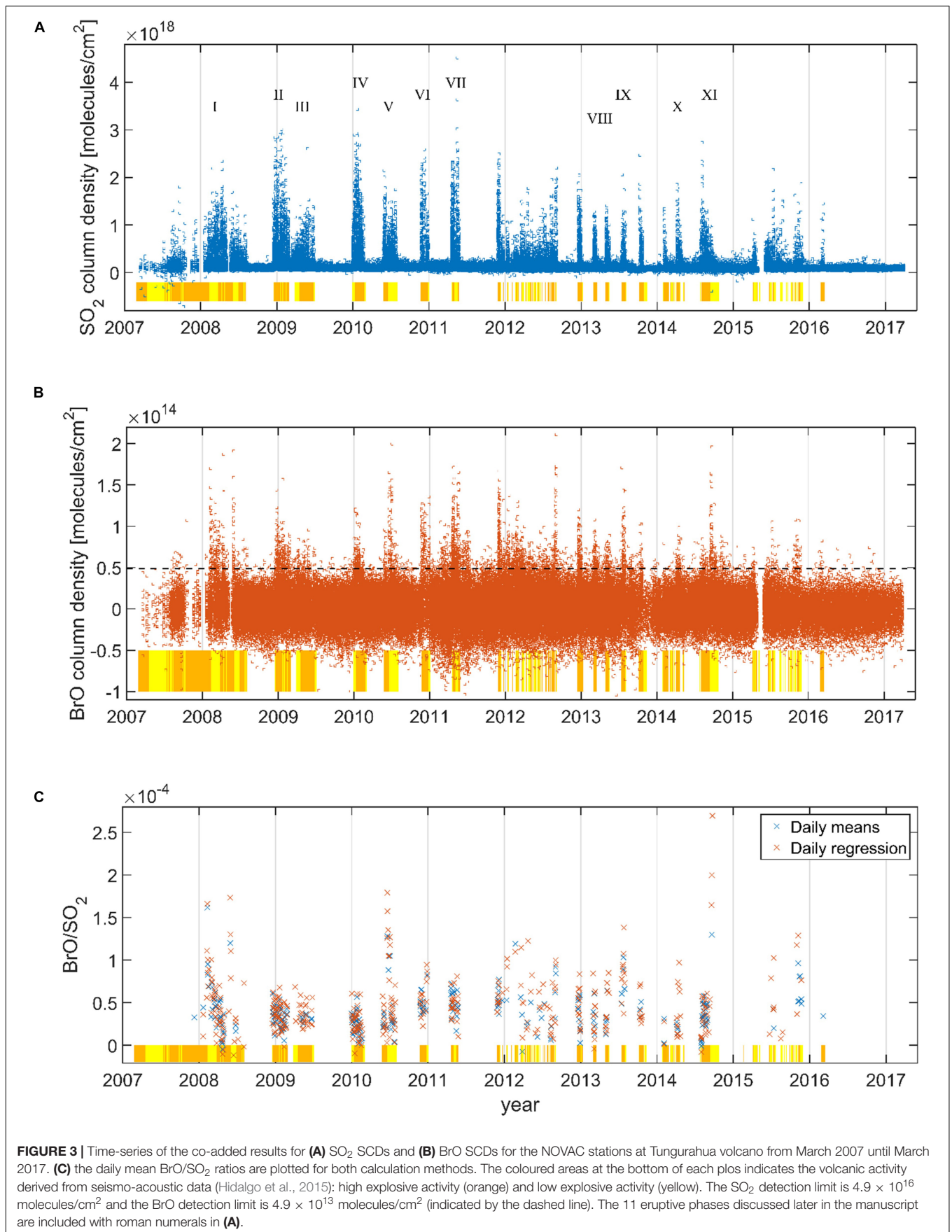
BrO/SO₂ ratios typically ranged between 2 and 8×10^{-5} , which is similar to ratios observed at other arc volcanoes in the Northern Volcanic Zone of the Andes, such as Nevado del Ruiz (Lübcke et al., 2014) and Cotopaxi (Dinger et al., 2018). In 2009 and the first half of 2010, relatively constant BrO/SO₂ ratios of 2 to 6×10^{-5} were observed. This contrasts with periods in 2008 and after 2010, where a larger range of ratios between 2 and 20×10^{-5} were detected. During the intermitting periods of weak SO₂ degassing, we could only infer an upper limit for the BrO/SO₂ ratio, calculated as the ratio between the BrO detection limit and the maximal SO₂ SCD during these periods:

$$\left(\frac{\text{BrO}}{\text{SO}_2}\right)_{\text{upper limit}} = \frac{\text{BrO SCD}_{\text{detection limit}}}{\text{SO}_2 \text{ SCD}_{\text{max, during low SO}_2 \text{ periods}}} \approx \frac{4.8 \times 10^{13}}{1.6 \times 10^{17}} = 3 \times 10^{-4}$$

Since this value far exceeds even the largest BrO/SO₂ ratios observed during phases of high SO₂ SCDs, we are not able to draw any conclusion from BrO/SO₂ ratios during these phases of low SO₂ at Tungurahua.

Dynamic of the BrO/SO₂ Ratio at Tungurahua

We identified changing BrO/SO₂ ratios during 11 phases of high SO₂ SCDs, which we number with roman numerals in chronological order (see Figure 3A). We identified a characteristic pattern in the BrO/SO₂ ratio over the course of each phase. For two phases (I and VII) the pattern is observed more than once and we separate this by adding the suffix a and b to the phase numbering, thus adding up to 13 phases, where we observe the BrO/SO₂ pattern. This pattern can be divided into



four stages, consisting of two initial stages (0 and 1), followed by an intermediate stage 2 and an ending stage 3. The separation of the initial stages into stage 0 and 1 could only be derived for four phases.

Stage 0 exclusively occurred in phases II, III, IV, and VII and typically lasted for periods of time between 2 and 21 days. We observed BrO and SO₂ degassing and BrO/SO₂ ratios between 3 and 6×10^{-5} , with a mean ratio of 4.31×10^{-5} .

Stage 1 was present in every phase. We observed a constant baseline BrO/SO₂ ratio of typically less than 5×10^{-5} (only for phase IX was it significantly higher at 6.20×10^{-5}). Mean ratio: 3.58×10^{-5} .

Stage 2 generally started with a strong increase in bromine emissions, which resulted in a BrO/SO₂ ratio of up to 17×10^{-5} , often twice as high as the ratio during stage 1. If stage 2 continued for more than a week, the ratio progressively decreased until it reached the former baseline level from stage 1. We will denote the latter part of this stage as stage 2b and the former as stage 2a. Stage 2 was observed in 11 phases (see **Table 1**). Mean ratio: 9.22×10^{-5} .

Stage 3 started when the BrO/SO₂ ratio decreased to low BrO/SO₂ ratios, similar to the ones observed in stage 1 ($2\text{--}5 \times 10^{-5}$). It could be seen in 5 phases (Ia, Ib, II, V, and X). Mean ratio: 3.44×10^{-5} .

The most significant feature is the strong increase in the BrO/SO₂ ratios from stage 1 to stage 2. This increase was lowest for phases II and VII and highest for phases Ia, Ib, V, and XI. The evolving pattern of BrO/SO₂ ratios through the four stages can be clearly recognised for every phase (see **Figure 4** and **Table 1**).

Evolution of the BrO/SO₂ Ratio Over the Course of Each Phase

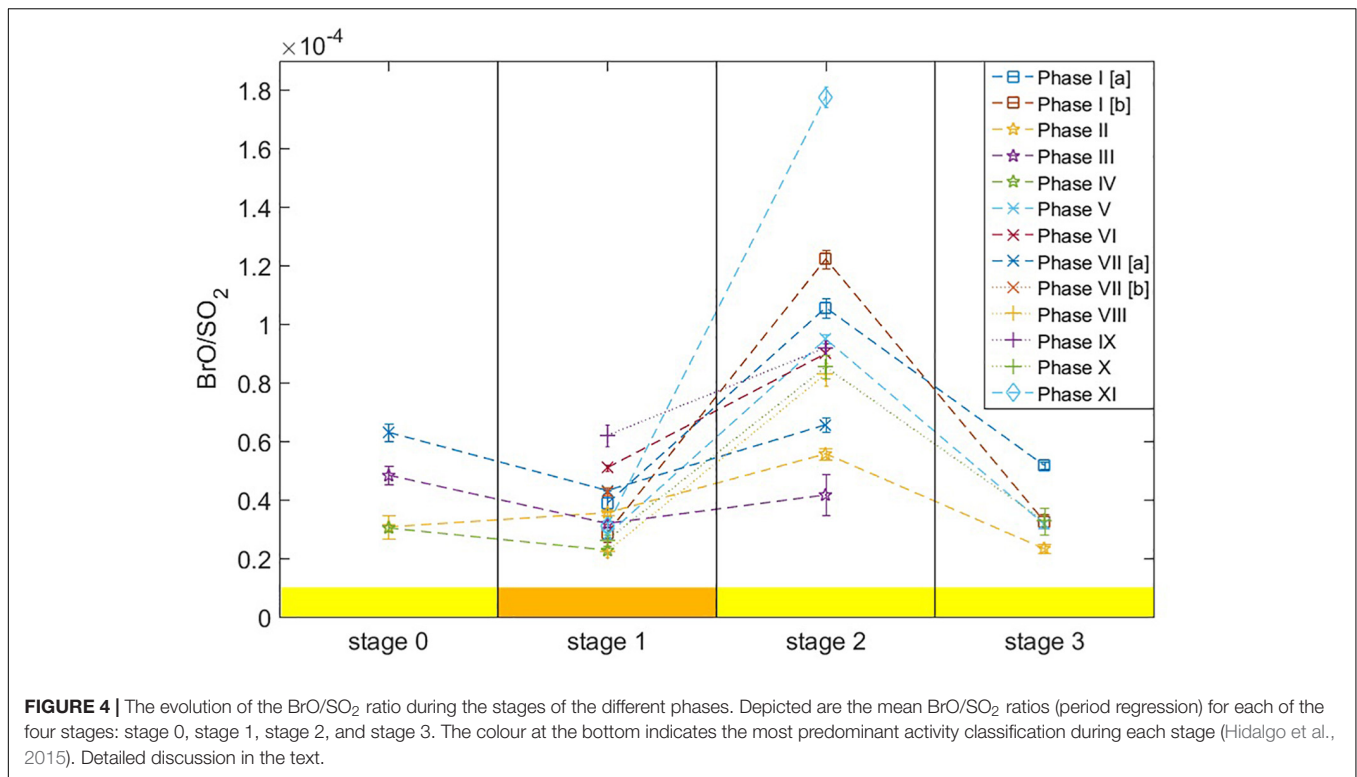
Here, we discuss in detail the evolution of the BrO/SO₂ ratio through every phase in chronological order. Phase V – serving as an example – will be discussed in more detail.

Phase I: During phase I (from 19 January to 22 July 2008), we observed the pattern twice. We will refer to them as phase Ia and phase Ib. Phase Ia started with a BrO/SO₂ ratio of 4.54×10^{-5} during the initialisation from 19 January 2008 (stage 1), lasting until an increase to 1.7×10^{-4} was observed on 8 February (stage 2a, lasting for 5 days). The BrO/SO₂ ratio decreased over the following weeks (stage 2b) until baseline levels were reached in March/April 2008 (stage 3, see **Figure 5**). Phase Ib started on 28 April 2008 with a ratio of 2.86×10^{-5} , reaching a short intermediate stage 2 on 28 May 2008, lasting only 6 days. It then decreased to 3.26×10^{-5} until 27 July 2008.

Phases II and III: The pattern is less clear for these phases, where only a small enhancement in the BrO/SO₂ ratio could

TABLE 1 | Overview over all 13 phases: Name, period, stages observed and their respective duration (in days), and the BrO/SO₂ ratio for each of the stages (if present). The evolution of the BrO/SO₂ ratio during the phases is also depicted in **Figure 4**.

Phase	Period (dd.mm.yy)	Stages observed (duration in d)	BrO/SO ₂ stage 0	BrO/SO ₂ stage 1	BrO/SO ₂ stage 2	BrO/SO ₂ stage 3
Ia	19.01.08–25.03.08	1- 2- 3 (20- 18- 29)	Below detection	3.88×10^{-5}	1.06×10^{-4}	5.18×10^{-5}
Ib	28.04.08–04.08.08	1- 2- 3 (30- 6- 26)	Below detection	2.86×10^{-5}	1.22×10^{-4}	3.26×10^{-5}
II	16.12.08–01.03.09	0- 1- 2- 3 (2- 32- 8- 5)	3.07×10^{-5}	3.56×10^{-5}	5.57×10^{-5}	2.12×10^{-5}
III	28.03.09–03.07.09	0- 1- 2 (21- 69- 8)	4.84×10^{-5}	3.19×10^{-5}	4.17×10^{-5}	Below detection
IV	30.12.09–04.03.10	0- 1 (12- 45)	3.04×10^{-5}	2.28×10^{-5}	Below detection	Below detection
V	26.05.10–04.08.10	1- 2- 3 (24- 22- 25)	Below detection	2.80×10^{-5}	9.49×10^{-5}	3.16×10^{-5}
VI	22.11.10–25.12.10	1- 2 (34- 9)	Below detection	5.10×10^{-5}	9.00×10^{-5}	Below detection
VIIa	20.04.11–08.05.11	0- 1- 2 (2- 12- 5)	6.31×10^{-5}	4.32×10^{-5}	6.57×10^{-5}	below detection
VIIb	16.05.11–26.05.11	1 (11)	Below detection	4.36×10^{-5}	Below detection	Below detection
VIII	28.04.13–16.05.13	1- 2 (8- 11)	Below detection	2.24×10^{-5}	8.29×10^{-5}	Below detection
IX	14.07.13–05.08.13	1- 2 (6- 17)	Below detection	6.20×10^{-5}	9.19×10^{-5}	Below detection
X	02.04.14–24.04.14	1- 2- 3 (13- 6- 4)	Below detection	2.63×10^{-5}	8.54×10^{-5}	3.27×10^{-5}
XI	02.08.14–25.10.14	1- 2 (41- 44)	Below detection	3.11×10^{-5}	1.78×10^{-4}	Below detection



be observed on 19 January 2009 (phase II) and on 26 June 2009 (phase III).

Phase IV: This phase exclusively consisted of the initial stages 0 and 1 with low BrO/SO₂ ratio ($2\text{--}3 \times 10^{-5}$). After stage 1, the gas signal decreased below the detection limit and no BrO/SO₂ ratio could be retrieved for stage 2.

Phase V: This phase took place from 29 May until 4 August 2010 and comprised stages 1 to 3. Stage 1, spanning from 29 May until 18 June, was characterised by high SO₂ emissions with SO₂ SCDs of up to 2×10^{18} molecules/cm² while BrO SCDs did not exceed 1×10^{14} molecules/cm² (see **Figure 6**, bottom left). During this stage, a mean BrO/SO₂ ratio of 2.80×10^{-5} was retrieved. Stage 2 started on 19 June with a strong increase in the BrO/SO₂ ratio by more than 1×10^{-4} to 1.7×10^{-4} as a result of a sharp rise in BrO SCDs up to 2×10^{14} molecules/cm² (see **Figure 6**, bottom middle). As stage 2 progressed, the BrO/SO₂ ratio slowly decreased (stage 2b), reaching previous baseline levels of roughly 3×10^{-5} by 11 July (after 22 days), where it remained throughout stage 3 until the gas emission dropped below the detection limit on 4 August. Both the BrO/SO₂ ratio as well as the linear regression results for the three stages of phase V are plotted in **Figure 6**. The period regression of each of these three stages reveals that the gas composition was similar during stages 1 and 3. While the SO₂ SCDs were similar for all three stages, BrO SCDs were higher during stage 2, if compared to stages 1 and 3 (**Figure 6**, bottom).

Phases VI to IX and XI exhibited only the initial stages and the intermediate stage 2, with a varying increase in strength from initialisation to stage 2 of 2.3×10^{-5} (during phase VII) to 1.4×10^{-4} (during phase XI).

Phase X: This phase behaved similar to phases VI–IX, but also had a short stage 3.

The detailed depiction of the co-added SO₂ and BrO SCDs as well as the BrO/SO₂ ratios for every phase can be found in the **Supplementary Materials**.

In 2012, in the second half of 2015, and for six phases between 2013 and 2016, only short periods with few BrO SCDs above the detection limit could be observed and longer period of BrO/SO₂ ratios could not be identified. For the six phases between 2013 and 2016 constant BrO/SO₂ ratio were calculated, with values ranging from 1.8 to 5.2×10^{-5} , similar to the range during stages 0 and 1.

Relation of the BrO/SO₂ Ratio With Explosive Activity

The observed changes in the BrO/SO₂ ratio between the different stages in the cycle observed during each phase at Tungurahua can be linked to changes in the explosive activity classification. The classification is closely related to the pattern of the BrO/SO₂ ratio in the following way: Stage 0 was typically related to low explosive activity. During Stage 1 – characterised by low BrO/SO₂ ratios – explosive activity was always defined as high. The subsequent increase in BrO/SO₂ ratios between stage 1 and 2 coincides with a change from high to low explosive activity. Only phases VIII, IX, and X (all after 2012), which are the shortest phases, did not exhibit this change in the classification. For these phases the activity classification remained high throughout stage 2. Stage 3 was generally governed by low explosive activity.

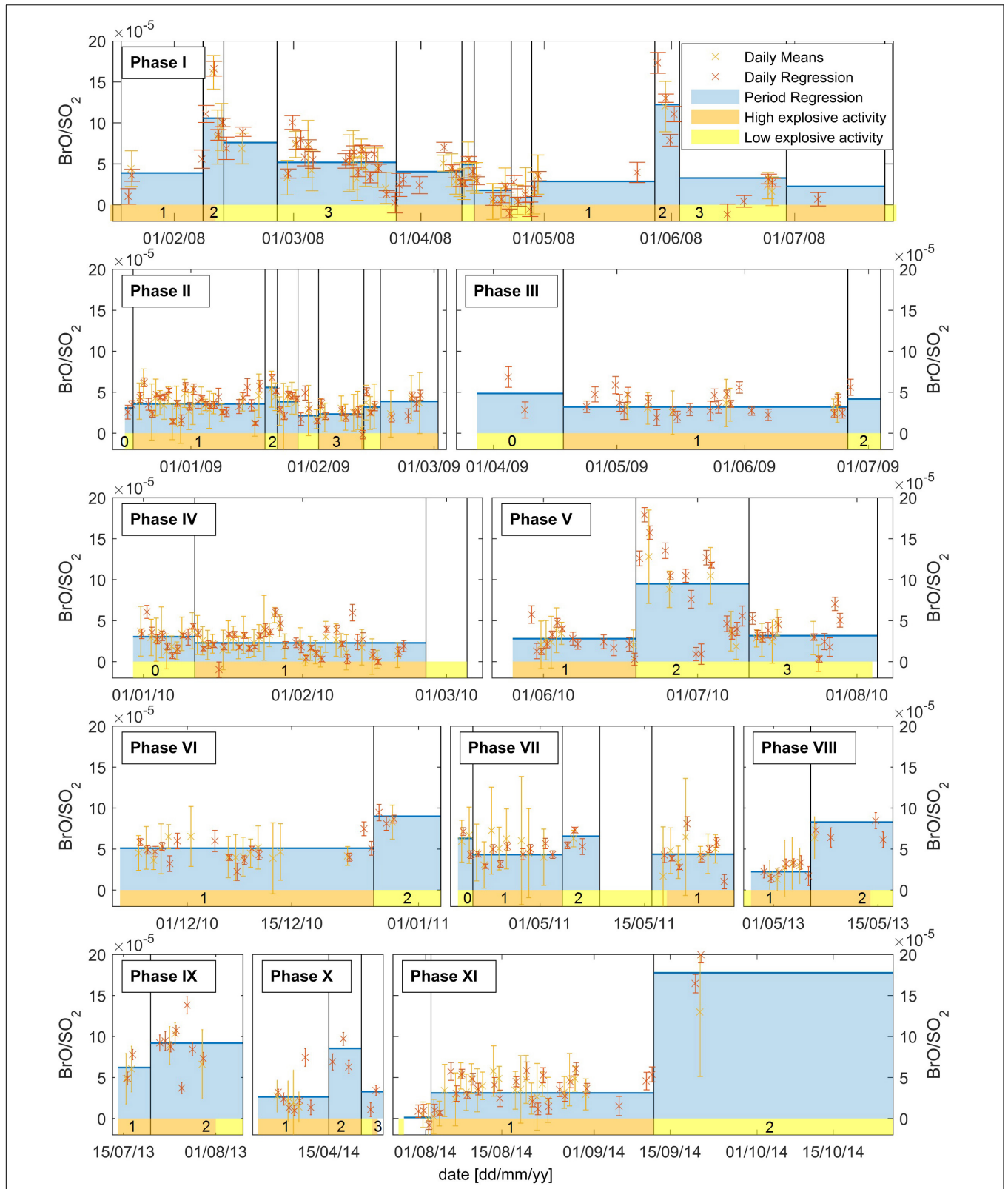
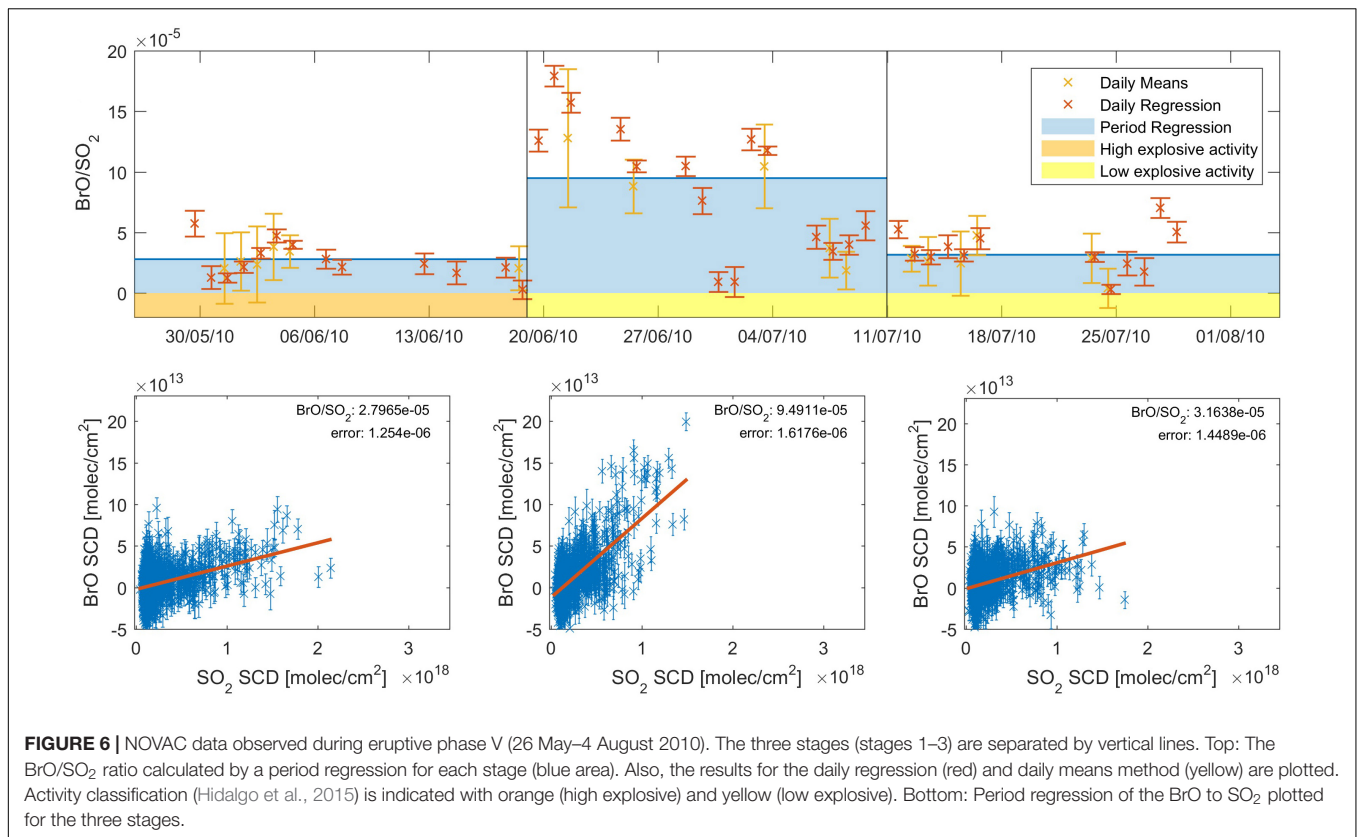


FIGURE 5 | BrO/SO₂ ratio over the course of each phase. Mean BrO/SO₂ ratio of the different periods (blue area, separated by black lines), the corresponding stage number is marked below the blue area. Furthermore, daily BrO/SO₂ from daily regression (red crosses) as well as from daily mean method (yellow crosses) are included. Explosivity classification is marked by yellow (low explosive) and orange area (high explosive).



The fact that for most of the phases the change in activity from high to low explosive activity was linked to an increase in the BrO/SO₂ ratio suggests that this might be inherent to the volcanic system of Tungurahua during the study period. In order to compare the explosive activity classification with the BrO/SO₂ ratio throughout the whole dataset, we performed period regressions for every period of high and low explosive activity. The results are shown in **Figure 7**. While the mean of all BrO/SO₂ ratio during high explosive activity (3.87×10^{-5}) is slightly lower than the mean of all BrO/SO₂ ratios during low explosive activity (5.19×10^{-5}), there is large variance caused by other parameters and thus this difference seems to be not statistically significant.

DISCUSSION

In **Figure 4**, a recurring temporal pattern in the eruptive phases of Tungurahua can be observed. We can summarise our observations as follows:

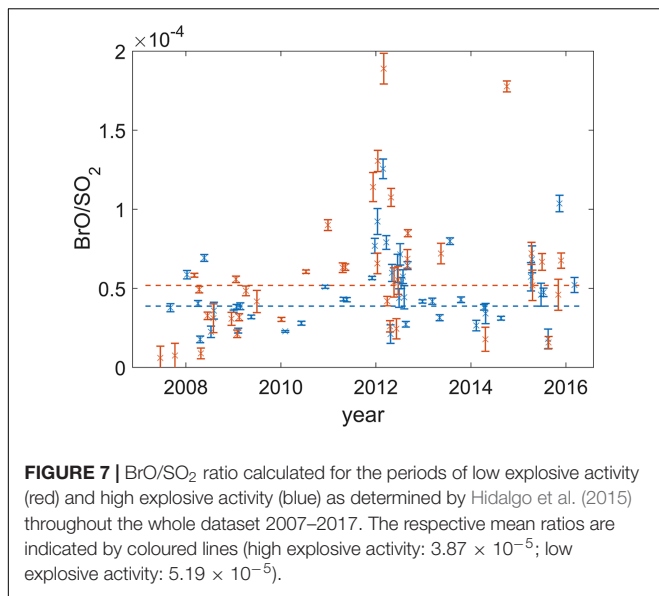
1. Initiation of degassing with low BrO/SO₂ ratios is observed during the initial stages. These are associated with high SO₂ emissions and strong seismicity, i.e. vulcanian explosions, occurring at the beginning of the eruptive phases.
2. After the initial stages, increased BrO values and nearly unchanged SO₂ emissions are observed in the plume, leading to an increased BrO/SO₂ ratio (observed for 11

of 13 eruptive phases). This typically coincided with the termination of high explosive activity.

3. For five phases, BrO/SO₂ decreased afterward and remained low until SO₂ SCDs decreased below detection. For the other eight phases, degassing had decreased already to a point that SO₂ and BrO were below the limit to calculate meaningful BrO/SO₂ ratios.

Changes of the BrO/SO₂ ratio in the plume of a single volcano have previously been linked to changes in volcanic processes (Bobrowski and Giuffrida, 2012; Lübcke et al., 2014; Dinger et al., 2018), when observing 3 to 20 min old volcanic plumes (e.g. Platt and Bobrowski, 2015; Gutmann et al., 2018). This was assumed to be caused by different melt-fluid partitioning of bromine and sulphur in magmatic melts. Moreover, differences in the BrO/SO₂ ratio among different volcanoes have been attributed to differences in magmatic composition (Platt and Bobrowski, 2015).

The melt-fluid partitioning of halogens (especially of chlorine and fluorine) with respect to sulphur has been the subject of multiple studies (e.g. Spilliaert et al., 2006; Burton et al., 2007). They found lower melt-fluid partitioning of chlorine compared to sulphur and thus concluded that (1) sulphur is usually released from magma earlier than chlorine (Carroll and Holloway, 1994; Spilliaert et al., 2006; Aiuppa, 2009; Webster et al., 2009) and (2) that the Cl/S ratio decreases with increasing degassing source depth (e.g. Burton et al., 2007). Measurements of chlorine and sulphur suggest that chlorine is more significantly released only



at shallow depth (typically less than 2 km; Allard et al., 2005; Spina et al., 2015).

However, the melt-fluid partitioning of bromine with respect to sulphur is still not very well constrained. Laboratory studies using natural silicate glasses suggested lower melt-fluid partitioning coefficients for bromine compared to sulphur also (Cadoux et al., 2018). They found melt-fluid partitioning coefficients between 3.8 and 20.8, which are similar to estimates derived from melt inclusions (0.3 to 21; Kutterolf et al., 2015) and from experimental data (2.18 to 9.2 and 17.5; Bureau et al., 2000, 2010).

This is similar to chlorine, where melt-fluid partitioning coefficients were found between 0.3 and 50 (Alletti et al., 2009; Webster et al., 2009; Zajacz et al., 2012). Thus, Cadoux et al. (2018) argue that bromine should behave similarly to chlorine and should be released predominantly during later degassing stages in comparison to sulphur. They assumed sulphur melt-fluid partitioning to be ≈ 15 –20 times larger compared to bromine, based on sulphur melt-fluid partitioning ranging between 3 and 236 (Beermann et al., 2015), and a factor of ≈ 9 between the chlorine and sulphur partitioning coefficients (Aiuppa, 2009).

Following these studies, and assuming no change in magmatic composition, higher BrO/SO₂ ratios should be associated with degassing from magma at shallow depths and lower BrO/SO₂ ratios released from magma degassing at deeper levels in the volcanic system.

Conceptual Model of the Evolution of BrO/SO₂ During Tungurahua's Eruptive Phases

Based on the BrO/SO₂ patterns observed in our study and on the recent studies of Cadoux et al. (2018), we propose a conceptual model that aims to explain the eruptive phases at Tungurahua.

Increasing SO₂ and BrO SCDs in combination with low BrO/SO₂ ratios are observed at the beginning of the explosive phases (stage 0 and 1). In agreement with increased SO₂ fluxes calculated by Hidalgo et al. (2015), we therefore interpret the initialisation of the phase as being driven by the intrusion of gas-rich magma into the upper part of the volcanic system with degassing signature dominated by deep degassing sources.

The increase in the BrO/SO₂ ratios on the transition to stage 2 indicates that bromine can partition from the melt more efficiently. SO₂ SCDs are similar during stages 1 and 2 (also observed in the SO₂ fluxes Hidalgo et al., 2015), implying that the increase in the BrO/SO₂ ratio was unlikely to be caused by depletion of SO₂ in the melt, but rather caused by a change of the degassing source depth. We thus propose that the magma was degassing at shallower depths during stage 2 (compared to stage 1), promoting bromine partitioning from the melt.

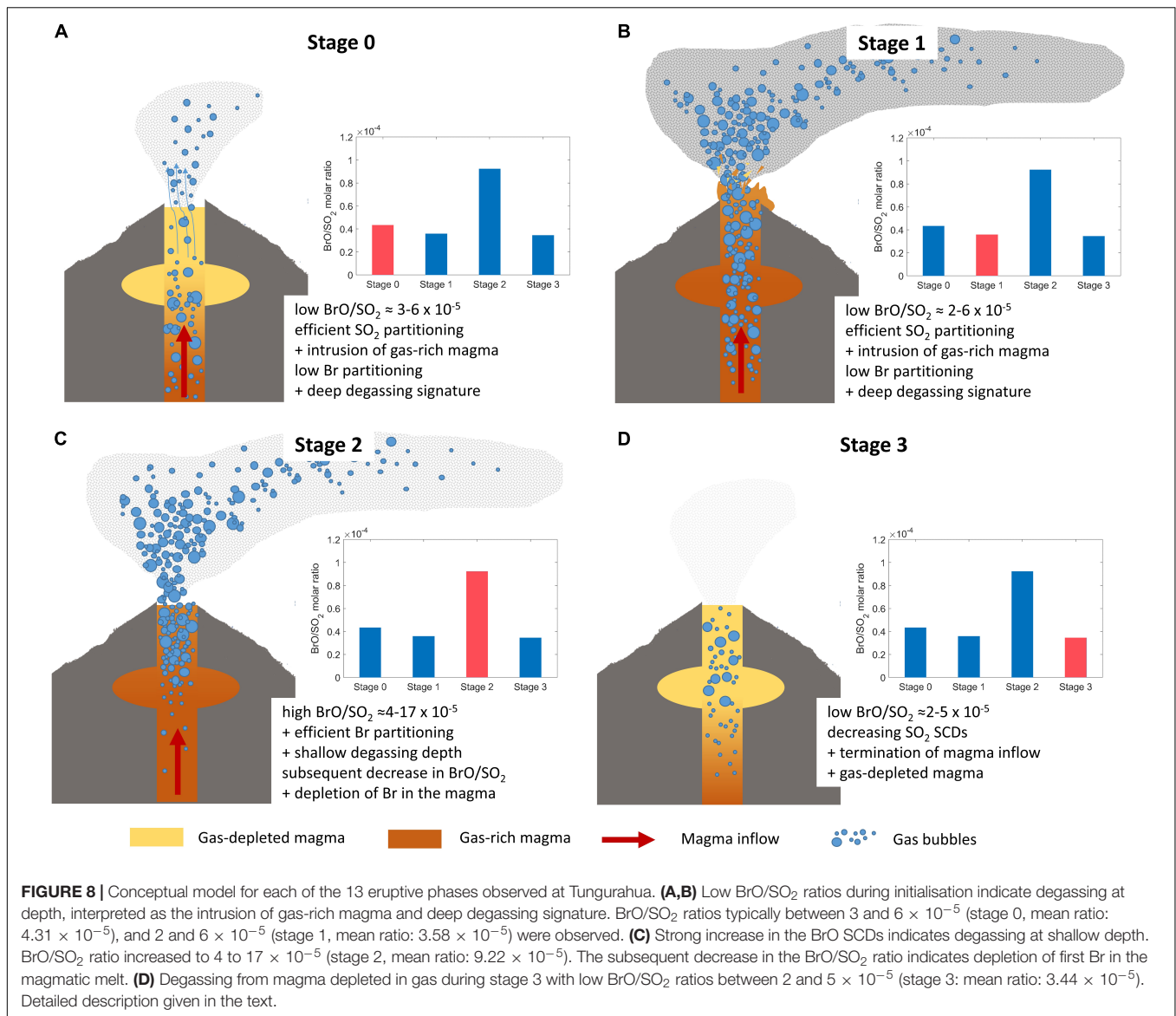
We further propose that the efficient degassing over the course of stage 2 resulted in a faster Br depletion due to its lower abundance in the melt. This led to the decrease in the BrO/SO₂ ratios toward the end of stage 2. Together with the corresponding decrease in the SO₂ SCDs this implies the interruption of magma intrusion. Stage 3 was thus governed by the degassing from gas-depleted magma until the gas emissions decreased below detection and the phase ends. The proposed conceptual model for the four stages is illustrated in **Figures 8A–D**, respectively.

Based on our observations and data from previous studies (Cadoux et al., 2018), we therefore suggest that each eruptive phase is driven by an intrusion of gas-rich magma, which subsequently degasses under different conditions. However, the amount of magma driving the phase might vary between the different eruptive phases.

Comparison of the Conceptual Model to Petrological Studies

Petrological studies showed that major eruptive phases at Tungurahua, like August 2006 (VEI 3) and May 2010 (phase V in this study), have been triggered by the input of new, hot, volatile-rich magma into a long-lived magmatic reservoir (Samaniego et al., 2011; Myers et al., 2014). Both studies propose an andesitic magma reservoir located at 8–10 km depth, which is coupled to a deeper mafic reservoir, probably located at depth of 15–16 km (Andújar et al., 2017). Myers et al. (2014) suggest that phase V was driven by the ascent of volatile-rich magma from the deeper basaltic andesite magma reservoir resulting in mixing with the upper andesitic, degassed, reservoir. This mechanism can account for the major eruptive phases documented at Tungurahua, and is in agreement with our conceptual model, which proposes the intrusion of volatile-rich magma (at least sulphur- and bromine-rich) as the main triggering process.

The lowest BrO/SO₂ ratios are observed for phases III (June 2009) and IV (February 2010), where the BrO/SO₂ ratios varied only between 2 and 4×10^{-5} . Thus, the high ratio recorded for phase V (18×10^{-5} at the beginning of stage 2) could indicate a comparatively high input of mafic magma leading toward this eruption (as proposed by Myers et al., 2014) and reciprocally the low BrO/SO₂ ratios during phase III and IV indicate low



or no input of volatile-rich mafic magma. Therefore, we can potentially link the BrO/SO₂ ratios observed during efficient degassing (stage 2) of the eruptive phases to the respective volatile status of the degassing magma.

Furthermore, it is important to note that during phase V, the energy of the vulcanian explosions was extremely high (Hidalgo et al., 2015) and larger than for the other eruptive phases between 2007 and 2016.

Comparison to Other Studies With Respect to Volcanic Activity

There are other studies similarly discussing long-term trends of BrO/SO₂ ratios with respect to changes in volcanic activity. Similar to our study, Bobrowski and Giuffrida (2012) found lower BrO/SO₂ ratios during periods of higher volcanic activity at Mount Etna during an observation period of more than

3 years including two long lasting periods of effusive eruptive activity and several more violent events (e.g. lava fountaining). Lübcke et al. (2014) obtained a continuous 3-year time-series of BrO/SO₂ ratios at the andesitic arc volcano Nevado del Ruiz. They observed low BrO/SO₂ ratios of $\approx 2 \times 10^{-5}$ during a period of high seismic activity and an increase of the BrO/SO₂ ratios to $\approx 5 \times 10^{-5}$ when the seismicity later decreased. This is similar to the behaviour which we observed during the eruptive phases of Tungurahua between the high explosive stage 1 and the low explosive stage 2.

In contrast to the laboratory measurements from Cadoux et al. (2018), Bobrowski and Giuffrida (2012) assumed an earlier bromine melt-fluid partitioning in comparison to sulphur as an interpretation of the spectral measurements, i.e. higher BrO/SO₂ ratios during the intrusion of gas-rich magma into the system. A similar hypothesis was proposed by Lübcke (2014), who associated the low BrO/SO₂ ratios to shallow degassing during

a period of increased seismicity. The increase in BrO/SO₂ ratios during the activity of Cotopaxi in 2015 (Dinger et al., 2018) were, like our study, similarly interpreted in terms of efficient degassing of rising magma in the conduit following the opening of the volcanic system by four phreato-magmatic explosions (Hidalgo et al., 2018).

In contrast to our measurements, the data of Bobrowski and Giuffrida (2012) and Lübcke et al. (2014) additionally enabled to determine BrO/SO₂ ratios during non-eruptive periods and to include those in their interpretation, which could be a reason for the divergence in the interpretation to our study. In order to resolve this divergence in the interpretation of the BrO/SO₂ ratio with respect to magmatic processes and partitioning of gas constituents from the melt, further field studies as well as laboratory studies should be undertaken. For such field studies, we therefore recommend even more comprehensive data collection, including seismic tomography to better determine source regions from a geophysical standpoint.

Multiple studies have investigated the behaviour of other halogens with respect to volcanic activity, predominantly through Cl/S ratio or HCl/SO₂ ratios. Measurements at Mount Etna revealed a decrease in the Cl/S ratio when strombolian activity followed a period of violent lava fountaining, and even lower Cl/S during a phase of passive degassing (Aiuppa et al., 2004). Furthermore, Spina et al. (2015) suggested a correlation with eruption intensity derived from FTIR measurements of HCl/SO₂ ratios at Mount Etna. This contrasts to what we found for the BrO/SO₂ ratio, which could potentially be an indication that bromine and chlorine do not behave similarly as has been assumed here and in several previous studies (Seo and Zajacz, 2016; Cadoux et al., 2018). However, other measurements at Mount Etna did display an increase in the HCl/SO₂ ratio, when explosive activity subsided (Allard et al., 2005).

To our knowledge there is only one study with simultaneous measurements of BrO/SO₂ and Cl/S (Bobrowski et al., 2017). In this work the authors find a positive correlation of both ratios during a one-week period at Nyiragongo volcano in June 2011. Both ratios change concurrently with a significant decrease (several tens of metre) in the height of the lava lake. Another study by Aiuppa et al. (2005), investigating Br/S and Cl/S ratios during the transition from quiescence to eruption at Mount Etna starting in October 2004, does not demonstrate a clear change of either ratio, both staying within the variance of the measurements. Further studies will be essential to confirm or disprove the assumptions made regarding the similarities of fluid-melt partitioning of chlorine and bromine.

CONCLUSION

We successfully retrieved BrO/SO₂ ratios for 13 eruptive phases at Tungurahua volcano between 2007 and 2017. The data show typical values of BrO/SO₂ ratios between 2 and 8×10^{-5} , similar to other andesitic arc volcanoes.

For 11 of the 13 eruptive phases, we identified an evolutionary pattern in the BrO/SO₂ ratio, which we interpret as characteristic of the volcano's eruptive phases. We propose

a conceptual eruptive model for Tungurahua linking this pattern in the BrO/SO₂ ratio to the eruptive dynamics of the volcanic system. The appearance of low BrO/SO₂ ratios in the plume at the beginning of the eruptive phases is directly linked to a volatile-rich magma intrusion, which acts as a trigger for the eruptive activity and is initially characterised by high explosivity. A following increase in molar ratios is associated with a shift to shallower depth at which efficient bromine degassing occurs, and is typically associated with a sustained period of low explosive activity. The subsequent depletion in volatiles in the degassing magma body finally leads to a decrease in BrO/SO₂ ratios altogether, and an end to the eruptive phase. The evolutionary pattern of the BrO/SO₂ ratios is thus linked to the explosive activity of the volcano, with low BrO/SO₂ ratios associated to high explosive activity later evolving to high ratios during a period of low explosive activity.

High BrO/SO₂ ratios observed after a reported high input of volatile-rich magma into the upper volcanic system could infer a relationship between bromine degassing to the amount of volatiles in the active system. The BrO/SO₂ can thus potentially be used as a proxy for the degree of depletion of volatiles. This study therefore demonstrates the increasing potential of long-term measurements of the BrO/SO₂ molar ratio to reveal important information about status as well as temporal evolution of the volcanic system.

AUTHOR CONTRIBUTIONS

NB, SW, PL, and UP contributed conception and design of the study. SW performed the data evaluation. NB, HS, and SW performed the statistical analysis. SW wrote the draft of the manuscript. NB, SH, FD, SA, UP, TW, and SW contributed to the interpretation of the data. All authors contributed to manuscript revision, read and approved the submitted version.

FUNDING

The publication costs of this study were kindly provided by the Max Planck Institute for Chemistry (MPIC).

ACKNOWLEDGMENTS

We thank the reviewers for their valuable time, input, and critical remarks. We thank the reviewer, MP, for the persistence in criticism which helped to make the discussion more concise and coherent. We would especially like to thank the reviewer, SB, for the detailed remarks and suggestions, which improved the manuscript significantly both with respect to overall message and conciseness of the interpretation. Both reviewers helped to improve the manuscript significantly. We thank the editor, AA, for his input and managing during the publishing process. We kindly thank the staff of the Instituto Geofísico – Escuela Politécnica Nacional (IG-EPN) – especially the Observatory at

Tungurahua – for maintaining the NOVAC instruments and for providing the spectral data. We thank the European Commission for funding of the NOVAC project under Framework 6 Research Program and DFG for financial support under project DFG PL 193/14-1.

REFERENCES

- Aiuppa, A. (2009). Degassing of halogens from basaltic volcanism: insights from volcanic gas observations. *Chem. Geol.* 263, 99–109. doi: 10.1016/j.chemgeo.2008.08.022
- Aiuppa, A., Federico, C., Antonio, P., Giovannella, P., and Mariano, V. (2002). S, Cl and F degassing as an indicator of volcanic dynamics: the 2001 eruption of Mount Etna. *Geophys. Res. Lett.* 29, 54–1–54-4. doi: 10.1029/2002GL015032
- Aiuppa, A., Federico, C., Franco, A., Giudice, G., Gurrieri, S., Inguaggiato, S., et al. (2005). Emission of bromine and iodine from Mount Etna volcano. *Geochem. Geophys. Geosyst.* 6:Q08008. doi: 10.1029/2005GC000965
- Aiuppa, A., Federico, C., Giudice, G., Gurrieri, S., Paonita, A., and Valenza, M. (2004). Plume chemistry provides insights into mechanisms of sulfur and halogen degassing in basaltic volcanoes. *Earth Planet. Sci. Lett.* 222, 469–483. doi: 10.1016/j.epsl.2004.03.020
- Allard, P., Burton, M., and Muré, F. (2005). Spectroscopic evidence for a lava fountain driven by previously accumulated magmatic gas. *Nature* 433, 407–410. doi: 10.1038/nature03246
- Alletti, M., Baker, D., Scaillet, B., Aiuppa, A., Moretti, R., and Ottolini, L. (2009). Chlorine partitioning between a basaltic melt and H₂O–CO₂ fluids at Mount Etna. *Chem. Geol.* 263, 37–50. doi: 10.1016/j.chemgeo.2009.04.003
- Andújar, J., Martel, C., Pichavant, M., Samaniego, P., Scaillet, B., and Molina, I. (2017). Structure of the plumbing system at Tungurahua volcano, Ecuador: insights from phase equilibrium experiments on July–August 2006 eruption products. *J. Petrol.* 58, 1249–1278. doi: 10.1093/ptrology/egx054
- Arellano, S., Hall, M., Samaniego, P., Pennec, J.-L. L., Ruiz, A., Molina, I., et al. (2008). Degassing patterns of Tungurahua volcano (Ecuador) during the 1999–2006 eruptive period, inferred from remote spectroscopic measurements of SO₂ emissions. *J. Volcanol. Geotherm. Res.* 176, 151–162. doi: 10.1016/j.jvolgeores.2008.07.007
- Bablon, M., Quidelleur, X., Samaniego, P., Le Pennec, J.-L., Lahitte, P., Liorzou, C., et al. (2018). Eruptive chronology of Tungurahua volcano (Ecuador) revisited based on new K–Ar ages and geomorphological reconstructions. *J. Volcanol. Geotherm. Res.* 357, 378–398. doi: 10.1016/j.jvolgeores.2018.05.007
- Beermann, O., Botcharnikov, R., and Nowak, M. (2015). Partitioning of sulfur and chlorine between aqueous fluid and basaltic melt at 1050°C, 100 and 200 MPa. *Chem. Geol.* 418, 132–157. doi: 10.1016/j.chemgeo.2015.08.008
- Bobrowski, N., and Giuffrida, G. (2012). Bromine monoxide / sulphur dioxide ratios in relation to volcanological observations at Mt. Etna 2006 - 2009. *Solid Earth* 3, 433–445. doi: 10.5194/se-3-433-2012
- Bobrowski, N., Giuffrida, G., Yalire, M., Lübcke, P., Arellano, S., Balagizi, C., et al. (2017). Multicomponent gas emission measurements of the active lava lake of Nyiragongo, DR Congo. *J. Afr. Earth Sci.* 134, 856–865. doi: 10.1016/j.jafrearsci.2016.07.010
- Bobrowski, N., Hönniger, G., Galle, B., and Platt, U. (2003). Detection of bromine monoxide in a volcanic plume. *Nature* 423, 273–276. doi: 10.1038/nature0162
- Bobrowski, N., and Platt, U. (2007). SO₂/BrO ratios studied in five volcanic plumes. *J. Volcanol. Geotherm. Res.* 166, 147–160. doi: 10.1016/j.jvolgeores.2007.07.003
- Bobrowski, N., von Glasow, R., Aiuppa, A., Inguaggiato, S., Louban, I., Ibrahim, O. W., et al. (2007). Reactive halogen chemistry in volcanic plumes. *J. Geophys. Res.* 112:D06311. doi: 10.1029/2006JD007206
- Boichu, M., Oppenheimer, C., Roberts, T. J., Tsanev, V., and Kyle, P. R. (2011). On bromine, nitrogen oxides and ozone depletion in the tropospheric plume of Erebus volcano (Antarctica). *Atmos. Environ.* 45, 3856–3866. doi: 10.1016/j.atmosenv.2011.03.027
- Bureau, H., Foy, E., Raepsaet, C., Somogyi, A., Munsch, P., Simon, G., et al. (2010). Bromine cycle in subduction zones through in situ Br monitoring in diamond anvil cells. *Geochim. Cosmochim. Acta* 74, 3839–3850. doi: 10.1016/j.gca.2010.04.001
- Bureau, H., Keppler, H., and Métrich, N. (2000). Volcanic degassing of bromine and iodine: experimental fluid/melt partitioning data and applications to stratospheric chemistry. *Earth Planet. Sci. Lett.* 183, 51–60. doi: 10.1016/S0012-821X(00)00258-2
- Burrows, J., Richter, A., Dehn, A., Deters, B., Himmelman, S., Voigt, S., et al. (1999). Atmospheric remote-sensing reference data from GOME-2. Temperature-dependent absorption cross sections of O₃ in the 231–794 nm range. *J. Quant. Spectrosc. Radiat. Transf.* 61, 509–517. doi: 10.1016/S0022-4073(98)00037-5
- Burton, M., Allard, P., Muré, F., and La Spina, A. (2007). Magmatic gas composition reveals the source depth of slug-driven strombolian explosive activity. *Science* 317, 227–230. doi: 10.1126/science.1141900
- Cadoux, A., Iacono-Marziano, G., Scaillet, B., Aiuppa, A., Mather, T. A., Pyle, D. M., et al. (2018). The role of melt composition on aqueous fluid vs. silicate melt partitioning of bromine in magmas. *Earth Planet. Sci. Lett.* 498, 450–463. doi: 10.1016/j.epsl.2018.06.038
- Cantrell, C. A. (2008). Technical note: review of methods for linear least-squares fitting of data and application to atmospheric chemistry problems. *Atmos. Chem. Phys.* 8, 5477–5487. doi: 10.5194/acp-8-5477-2008
- Carn, S., Krueger, A., Arellano, S., Krotkov, N., and Yang, K. (2008). Daily monitoring of Ecuadorian volcanic degassing from space. *J. Volcanol. Geotherm. Res.* 176, 141–150. doi: 10.1016/j.jvolgeores.2008.01.029
- Carroll, M. R., and Holloway, J. R. (1994). *Volatiles in Magmas*, Vol. 30. Washington, D.C: Mineralogical Society of America.
- Cochain, B., Sanloup, C., de Grouchy, C., Crépeisson, C., Bureau, H., Leroy, C., et al. (2015). Bromine speciation in hydrous silicate melts at high pressure. *Chem. Geol.* 404, 18–26. doi: 10.1016/j.chemgeo.2015.03.015
- Dinger, F., Bobrowski, N., Warnach, S., Bredemeyer, S., Hidalgo, S., Arellano, S., et al. (2018). Periodicity in the BrO/SO₂ molar ratios in the volcanic gas plume of Cotopaxi and its correlation with the Earth tides during the eruption in 2015. *Solid Earth* 9, 247–266. doi: 10.5194/se-9-247-2018
- Edmonds, M., Gerlach, T. M., and Herd, R. A. (2009). Halogen degassing during ascent and eruption of water-poor basaltic magma. *Chem. Geol.* 263, 122–130. doi: 10.1016/j.chemgeo.2008.09.022
- Fleischmann, O. C., Hartmann, M., Burrows, J. P., and Orphal, J. (2004). New ultraviolet absorption cross-sections of BrO at atmospheric temperatures measured by time-windowing Fourier transform spectroscopy. *J. Photochem. Photobiol. A Chem.* 168, 117–132. doi: 10.1016/j.jphotochem.2004.03.026
- Galle, B., Johansson, M., Rivera, C., Zhang, Y., Kihlman, M., Kern, C., et al. (2010). Network for observation of volcanic and atmospheric change (NOVAC) - a global network for volcanic gas monitoring: network layout and instrument description. *J. Geophys. Res.* 115:D05304. doi: 10.1029/2009JD011823
- Grainger, J. F., and Ring, J. (1962). Anomalous Fraunhofer line profiles. *Nature* 193, 762–762. doi: 10.1038/193762a0
- Gutmann, A., Bobrowski, N., Roberts, T. J., Rüdiger, J., and Hoffmann, T. (2018). Advances in bromine speciation in volcanic plumes. *Front. Earth Sci.* 6:213. doi: 10.3389/feart.2018.00213
- Hall, M. L., Robin, C., Beate, B., Mothes, P., and Monzier, M. (1999). Tungurahua volcano, Ecuador: structure, eruptive history and hazards. *J. Volcanol. Geotherm. Res.* 91, 1–21. doi: 10.1016/S0377-0273(99)00047-5
- Hermans, C., Vandaele, A., Fally, S., Carleer, M., Colin, R., Coquart, B., et al. (2003). “Absorption cross-section of the collision-induced bands of oxygen from the UV to the NIR,” in *Weakly Interacting Molecular Pairs: Unconventional Absorbers of Radiation in the Atmosphere*, Vol. 27, eds C. Camy-Peyret and A. Viggas (Netherlands: Springer), 193–202. doi: 10.1007/978-94-010-0025-3_16
- Hidalgo, S., Battaglia, J., Arellano, S., Sierra, D., Bernard, B., Parra, R., et al. (2018). Evolution of the 2015 Cotopaxi eruption revealed by combined geochemical & seismic observations. *Geochem. Geophys. Geosyst.* 19, 2087–2108. doi: 10.1029/2018GC007514

SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2019.00132/full#supplementary-material>

- Hidalgo, S., Battaglia, J., Arellano, S., Steele, A., Bernard, B., Bourquin, J., et al. (2015). SO₂ degassing at Tungurahua volcano (Ecuador) between 2007 and 2013: transition from continuous to episodic activity. *J. Volcanol. Geotherm. Res.* 298, 1–14. doi: 10.1016/j.jvolgeores.2015.03.022
- Hörmann, C., Sihler, H., Bobrowski, N., Beirle, S., Penning de Vries, M., Platt, U., et al. (2013). Systematic investigation of bromine monoxide in volcanic plumes from space by using the GOME-2 instrument. *Atmos. Chem. Phys.* 13, 4749–4781. doi: 10.5194/acp-13-4749-2013
- Johnson, J. (2003). Generation and propagation of infrasonic airwaves from volcanic explosions. *J. Volcanol. Geotherm. Res.* 121, 1–14. doi: 10.1016/S0377-0273(02)00408-0
- Kelly, P. J., Kern, C., Roberts, T. J., Lopez, T., Werner, C., and Aiuppa, A. (2013). Rapid chemical evolution of tropospheric volcanic emissions from Redoubt volcano, Alaska, based on observations of ozone and halogen-containing gases. *J. Volcanol. Geotherm. Res.* 259, 317–333. doi: 10.1016/j.jvolgeores.2012.04.023
- Kern, C., Sihler, H., Vogel, L., Rivera, C., Herrera, M., and Platt, U. (2009). Halogen oxide measurements at Masaya volcano, Nicaragua using active long path differential optical absorption spectroscopy. *Bull. Volcanol.* 71, 659–670. doi: 10.1007/s00445-008-0252-8
- Kutterolf, S., Hansteen, T. H., Freundt, A., Wehrmann, H., Appel, K., Krüger, K., et al. (2015). Bromine and chlorine emissions from Plinian eruptions along the Central American Volcanic Arc: from source to atmosphere. *Earth Planet. Sci. Lett.* 429, 234–246. doi: 10.1016/j.epsl.2015.07.064
- Lübcke, P. (2014). *Optical Remote Sensing Measurements of Bromine and Sulphur Emissions: Investigating their Potential as Trace of Volcanic Activity*. Ph.D. thesis. Heidelberg: Universität Heidelberg.
- Lübcke, P., Bobrowski, N., Arellano, S., Galle, B., Garzón, G., Vogel, L., et al. (2014). BrO/SO₂ molar ratios from scanning DOAS measurements in the NOVAC network. *Solid Earth* 5, 409–424. doi: 10.5194/se-5-409-2014
- Malinconico, L. L. (1979). Fluctuations in SO₂ emission during recent eruptions of Etna. *Nature* 278, 43–45. doi: 10.1038/278043a0
- McCormick, B., Herzog, M., Yang, J., Edmonds, M., Mather, T. A., Carn, S. A., et al. (2014). A comparison of satellite- and ground-based measurements of SO₂ emissions from Tungurahua volcano, Ecuador. *J. Geophys. Res. Atmos.* 119, 4264–4285. doi: 10.1002/2013JD019771
- Meller, R., and Moortgat, G. K. (2000). Temperature dependence of the absorption cross sections of formaldehyde between 223 and 323K in the wavelength range 225 - 375 nm. *J. Geophys. Res. Atmos.* 105, 7089–7101. doi: 10.1029/1999JD901074
- Menyailov, I. A. (1975). Prediction of eruptions using changes in composition of volcanic gases. *Bull. Volcanol.* 39, 112–125. doi: 10.1007/BF02596951
- Myers, M. L., Geist, D. J., Rowe, M. C., Harpp, K. S., Wallace, P. J., and Dufek, J. (2014). Replenishment of volatile-rich mafic magma into a degassed chamber drives mixing and eruption of Tungurahua volcano. *Bull. Volcanol.* 76, 872–888. doi: 10.1007/s00445-014-0872-0
- Noguchi, K., and Kamiya, H. (1963). Prediction of volcanic eruption by measuring the chemical composition and amounts of gases. *Bull. Volcanol.* 26, 367–378. doi: 10.1007/BF02597298
- Olmos, R., Barrancos, J., Rivera, C., Barahona, F., López, D. L., Henriquez, B., et al. (2007). Anomalous emissions of SO₂ during the recent eruption of Santa Ana volcano, El Salvador, Central America. *Pure Appl. Geophys.* 164, 2489–2506. doi: 10.1007/s00024-007-0276-6
- Oppenheimer, C., Tsanev, V. I., Braban, C. F., Cox, R. A., Adams, J. W., Aiuppa, A., et al. (2006). BrO formation in volcanic plumes. *Geochim. Cosmochim. Acta* 70, 2935–2941. doi: 10.1016/j.gca.2006.04.001
- Pennec, J.-L. L., Jaya, D., Samaniego, P., Ramón, P., Yáñez, S. M., Egred, J., et al. (2008). The AD 1300–1700 eruptive periods at Tungurahua volcano, Ecuador, revealed by historical narratives, stratigraphy and radiocarbon dating. *J. Volcanol. Geotherm. Res.* 176, 70–81. doi: 10.1016/j.jvolgeores.2008.05.019
- Pennisi, M., and Cloarec, M.-F. L. (1998). Variations of Cl, F, and S in Mount Etna's plume, Italy, between 1992 and 1995. *J. Geophys. Res.* 103, 5061–5066. doi: 10.1029/97JB03011
- Platt, U., and Bobrowski, N. (2015). “Quantification of volcanic reactive halogen emissions,” in *Volcanism and Global Environmental Change*, eds A. Schmidt, K. Fristad, and L. Elkins-Tanton (Cambridge: Cambridge University Press), 115–132. doi: 10.1017/cbo9781107415683.011
- Platt, U., and Stutz, J. (2008). *Differential Optical Absorption Spectroscopy: Physics of Earth and Space Environments*. Berlin: Springer. doi: 10.1007/978-3-540-75776-4
- Roberts, T. J., Martin, R. S., and Jourdain, L. (2014). Reactive bromine chemistry in Mount Etna's volcanic plume: the influence of total Br, high-temperature processing, aerosol loading and plume-air mixing. *Atmos. Chem. Phys.* 14, 11201–11219. doi: 10.5194/acp-14-11201-2014
- Samaniego, P., Pennec, J.-L. L., Robin, C., and Hidalgo, S. (2011). Petrological analysis of the pre-eruptive magmatic process prior to the 2006 explosive eruptions at Tungurahua volcano (Ecuador). *J. Volcanol. Geotherm. Res.* 199, 69–84. doi: 10.1016/j.jvolgeores.2010.10.010
- Seo, J. H., and Zajacz, Z. (2016). Fractionation of Cl/Br during fluid phase separation in magmatic-hydrothermal fluids. *Geochim. Cosmochim. Acta* 183, 125–137. doi: 10.1016/j.gca.2016.04.009
- Spilliaert, N., Métrich, N., and Allard, P. (2006). S-Cl-F degassing pattern of water-rich alkali basalt: modelling and relationship with eruption styles on Mount Etna volcano. *Earth Planet. Sci. Lett.* 248, 772–786. doi: 10.1016/j.epsl.2006.06.031
- Spina, A. L., Burton, M., Allard, P., Alparone, S., and Muré, F. (2015). Open-path FTIR spectroscopy of magma degassing processes during eight lava fountains on Mount Etna. *Earth Planet. Sci. Lett.* 413, 123–134. doi: 10.1016/j.epsl.2014.12.038
- Theys, N., Roozendaal, M. V., Dils, B., Hendrick, F., Hao, N., and Mazière, M. D. (2009). First satellite detection of volcanic bromine monoxide emission after the Kasatochi eruption. *Geophys. Res. Lett.* 36:L03809.
- Vandaele, A., Hermans, C., and Fally, S. (2009). Fourier transform measurements of SO₂ absorption cross sections: II: temperature dependence in the 29 000 - 44 000 cm⁻¹ (227 - 345 nm) region. *J. Quant. Spectrosc. Radiat. Transf.* 110, 2115–2126. doi: 10.1016/j.jqsrt.2009.05.006
- Vandaele, A., Hermans, C., Simon, P., Carleer, M., Colin, R., Fally, S., et al. (1998). Measurements of the NO₂ absorption cross-section from 42 000 cm⁻¹ to 10 000 cm⁻¹ (238-1000 nm) at 220K and 294 K. *J. Quant. Spectrosc. Radiat. Transf.* 59, 171–184. doi: 10.1016/S0022-4073(97)00168-4
- Villemant, B., and Boudon, G. (1999). H₂O and halogen (F, Cl, Br) behaviour during shallow magma degassing processes. *Earth Planet. Sci. Lett.* 168, 271–286. doi: 10.1016/S0012-821X(99)00058-8
- Vogel, L., Sihler, H., Lampel, J., Wagner, T., and Platt, U. (2013). Retrieval interval mapping: a tool to visualize the impact of the spectral retrieval range on differential optical absorption spectroscopy evaluations. *Atmos. Meas. Tech.* 6, 275–299. doi: 10.5194/amt-6-275-2013
- von Glasow, R. (2010). Atmospheric chemistry in volcanic plumes. *Proc. Natl. Acad. Sci. U.S.A.* 107, 6594–6599. doi: 10.1073/pnas.0913164107
- Wagner, T., Beirle, S., and Deutschmann, T. (2009). Three-dimensional simulation of the Ring effect in observations of scattered sun light using Monte Carlo radiative transfer models. *Atmos. Meas. Tech.* 2, 113–124. doi: 10.5194/amt-2-113-2009
- Webster, J., Sintoni, M., and Vivo, B. D. (2009). The partitioning behavior of Cl, S, and H₂O in aqueous vapor-saline-liquid saturated phonolitic and trachytic melts at 200 MPa. *Chem. Geol.* 263, 19–36. doi: 10.1016/j.chemgeo.2008.10.017
- Wright, H. M., Cashman, K. V., Mothes, P. A., Hall, M. L., Ruiz, A. G., and Le Pennec, J. (2012). Estimating rates of decompression from textures of erupted ash particles produced by 1999–2006 eruptions of Tungurahua volcano, Ecuador. *Geology* 40, 619–622. doi: 10.1130/G32948.1
- Zajacz, Z., Candela, P. A., Piccoli, P. M., and Sanchez-Valle, C. (2012). The partitioning of sulfur and chlorine between andesite melts and magmatic volatiles and the exchange coefficients of major cations. *Geochim. Cosmochim. Acta* 89, 81–101. doi: 10.1016/j.gca.2012.04.039

Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

The handling Editor is currently editing a Research Topic with one of the authors NB, and confirms the absence of any other collaboration.

Copyright © 2019 Warnach, Bobrowski, Hidalgo, Arellano, Sihler, Dinger, Lübcke, Battaglia, Steele, Galle, Platt and Wagner. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner(s) are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.