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Geochemistry, Geophysics, Geosystems

RESEARCH ARTICLE

10.1029/2019GC008668

Q4 Key Points:

- Twenty-three new archaeointensity data were obtained on pottery sherds from Central Mexico in the first millennium BCE
- A new Mexican secular variation curve between 1500 BCE and 200 CE is proposed
- The curve exhibits a fast increase of the intensity (~250 nT/year) during the third century BCE and a large maximum ~65 μT circa 200–100 BCE

Supporting Information:

- Supporting Information S1
- Supporting Information S2

Correspondence to:

G. Hervé,
gwenael.herve@u-bordeaux-mon-taigne.fr

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Secular Variation of the Intensity of the Geomagnetic Field in Mexico During the First Millennium BCE

Gwenaël Hervé¹ , Mireille Perrin¹ , Luis M. Alva-Valdivia², Alejandro Rodriguez-Trejo², Arnaldo Hernandez-Cardona², Mario Córdova Tello³, and Carolina Meza Rodriguez³

¹Aix Marseille University, CNRS, IRD, INRA, Coll France, CEREGE, Aix-en-Provence, France, ²Instituto de Geofísica, Laboratorio de Paleomagnetismo, Universidad Nacional Autónoma de México, Mexico City, Mexico, ³Instituto Nacional de Antropología e Historia, Centro Morelos, Mexico

Abstract In Western Eurasia, the first millennium BCE is characterized by the fastest secular variation of the Earth Magnetic Field observed over the last millennia and by a geomagnetic anomaly centered on the Middle East. On the global scale, the variation of the dipolar field during this period remains poorly constrained because of the lack of data in other geographical areas. Here, we presented 23 new mean archaeointensity data on ceramic sherds dated between 1500 BCE and 200 CE from Chalcatzingo archaeological site in Central Mexico. Archaeointensities were determined using the classical Thellier-Thellier protocol with corrections for TRM anisotropy and cooling rate effects. Our work doubles the number of high-quality archaeointensity data in Mexico during the considered period. Using a Bayesian approach, a new secular variation curve was calculated at Mexico City between 1500 BCE and 200 CE after selection of Mexican archaeointensity data. After a period of oscillations of the intensity between 20 and 40 μT from 1500 to 300 BCE, the curve shows a large maximum ~65 μT in the second century BCE. The corresponding VADM varied between ~4.0 and ~11.0 $\times 10^{22}$ Am², which highlights further that the intensity of the geomagnetic field could vary at regional scale over a larger range as previously thought. However, this amplitude variation may be overestimated, as it does not take into account the fast directional variation observed at this time.

Plain Language Summary Archaeological baked clays are the best material to reconstitute the past secular variation of the Earth magnetic field, because they acquired a generally stable thermoremanent magnetization usually parallel and proportional to the ambient field at the time of their baking. Global modeling of the geomagnetic field requires a spatial and temporal distribution of data as homogeneous as possible, which is still not the case yet. In spite of its rich archaeological heritage, high-quality archaeointensity data are still very few in Mexico. The present study focuses on pottery sherds dated between 1500 BCE and 200 CE from Chalcatzingo archaeological site in Central Mexico. We obtained 21 new mean archaeointensities, which almost doubles the high-quality Mexican dataset for this period. The new secular variation Mexican curve exhibits oscillations of the intensity between 20 and 40 μT from 1500 to 300 BCE, before a large maximum ~65 μT in the second century BCE. This work will help to have a better knowledge of the variation of the dipolar geomagnetic field during the first millennium BCE, when the fastest secular variation over the last millennia was reported in the Middle East (“geomagnetic spikes”).

1. Introduction

Knowing how the geomagnetic field changed through time is crucial to understand the functioning of the geodynamo and the flux dynamic in the Earth's core. Beyond the last centuries, absolute but discrete estimations of the past geomagnetic direction and intensity are recovered from volcanic lava flows and archaeological baked clays. Lacustrine and marine sediments can also be used, but these continuous records provide only relative data that tend to smooth the secular variation of the geomagnetic field. Global models of the field during the Holocene were developed by inversion of absolute and/or relative data using spherical harmonic analysis in space (e.g., Arneitz et al., 2019; Constable et al., 2016; Hellio & Gillet, 2018; Pavón-Carrasco et al., 2014).

In spite of important improvements in the past 10 years, the accuracy of the models is still weakened by the inhomogeneous spatial distribution of worldwide data with around 70% of the data concentrated around

Western Eurasia (Brown et al., 2015). The temporal resolution of the archaeomagnetic records is lower in other regions, such as in Asia (e.g., Cai et al., 2016, 2016), Oceania (e.g., Greve & Turner, 2017), and Northern America (e.g., Hagstrum & Champion, 2002; Lengyel, 2010). One of the main consequences is the poor knowledge of dipole moment variation prior to 1840. Models disagree when to know if the dipole moment decreased regularly since ~750 CE (Poletti et al., 2018) or since ~1700 CE after seven centuries of stability (Hellio & Gillet, 2018). The lack of data also prevents a precise understanding of geomagnetic anomalies at the Earth's surface such as the present South Atlantic anomaly (e.g., Campuzano et al., 2019; Terra-Nova et al., 2017).

Another geomagnetic anomaly centered on the Middle East has recently been proposed at the beginning of the first millennium BCE (Shaar et al., 2016; Shaar et al., 2017). This Levantine Iron Age anomaly (LIAA) was likely at the origin of two short-lived high-intensity events called “geomagnetic spikes” (Ben-Yosef et al., 2009; Shaar et al., 2011). The migration of the LIAA through Europe would be responsible for the large and fast variations of the geomagnetic field observed in this region during the first millennium BCE in both direction (e.g., Hervé et al., 2013; Palencia-Ortas et al., 2017; Shaar et al., 2018) and intensity (e.g., Hervé et al., 2017; Molina-Cardin et al., 2018; Shaar et al., 2016). Besides the role of these nondipolar fields, global reconstructions of data suggest that the influence of the dipolar field should not be underestimated with a maximum of the dipole moment during the first millennium BCE (Usoskin et al., 2016) and a 10–15° dipole tilt during the first half of this millennium (Nilsson et al., 2010). But the knowledge of the respective contributions of different harmonic degrees to the secular variation during the first millennium BCE remains hampered by the lack of data outside Western Eurasia.

In this context, Mexico is a key area to explore the secular variation at low latitudes, thanks to its rich archaeological heritage at this time linked to the powerful Olmec civilization. Mexico is also close to Hall's cave in Texas, where a “geomagnetic spike” seems to exist in two sedimentary sequences at the beginning of the first millennium BCE (Bourne et al., 2016). No absolute data have confirmed this “spike” for the moment.

In Mexico, a large number of archaeomagnetic studies on baked clays or lava flows have been published in the last 10 years (e.g., Fanjat et al., 2013; Mahgoub et al., 2017; Mahgoub et al., 2019; Mahgoub et al., 2019; Pétronille et al., 2012; Rodriguez-Ceja et al., 2009; Rodriguez-Trejo et al., 2019; Terán et al., 2016), but a recent critical analysis of the Mexican dataset (Hervé et al., 2019) showed that around 70% of these data cannot be considered as high quality, according to basic selection criteria. Two local secular variation curves have been recently proposed using only Mexican data (Mahgoub, Juarez-Arriaga, et al., 2019) or also some from the southwestern United States (Goguitchaichvili et al., 2018). The significant differences between the two curves point out that the resolution of the secular variation during the first millennium BCE is still not good. High-quality data are especially lacking between 1500 and 400 BCE and the present study, with 23 new archaeointensities from Chalcatzingo in Central Mexico, doubles the database.

2. Archaeological Context and Sampling

The archaeological site of Chalcatzingo (lat: 18.6766°N, long: 98.7705°W) is located above the fertile Amatzinac valley in the Morelos state at circa 120 km southeast of Mexico City (Figures 1a and 1b). Excavations, carried out since the 30 s, are presently directed by Mario Córdova Tello and Carolina Meza Rodriguez from the *Instituto Nacional de Antropología e Historia* (INAH).

Chalcatzingo was an important regional center in Central Mexico between 1500 and 500 BCE at the Preclassic period. The site was organized in terraces at the base of two volcanic peaks called Cerro Chalcatzingo and Cerro Delgado (Grove, 1987). The remains highlighted an increasing political organization and the development of long-distance trade, especially with the Gulf coast of Mexico that was the heartland of the Olmec civilization at this period. Famous ritual petroglyphs of Chalcatzingo are the most significant witnesses of the Olmec influence in Central Mexico (Cordova Tello & Meza Rodriguez, 2017; Cyphers Guillen, 1982). The stratigraphy divided the occupation in three successive phases, called Amate (~1500 to ~1100 BCE), Barranca (~1100 to ~700 BCE), and Cantera (~700 to ~500 BCE). Each phase was defined by specific pottery types. This relative chronology was fixed to the calendar scale through 57 radiocarbon dates (Cyphers Guillen & Grove, 1987).

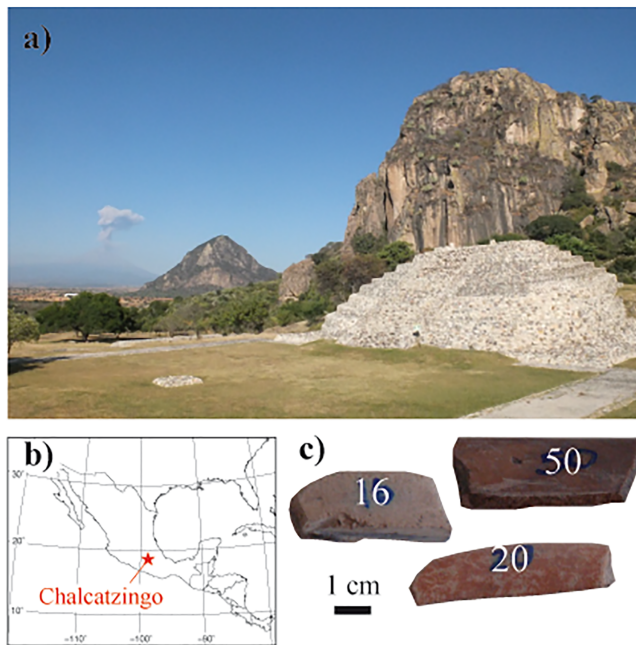


Figure 1. (a) General view and (b) location of Chalcatzingo archaeological site. In (c) are shown representative Preclassic pots of *Blanco Amatzinac* (no. 16), *Anaranjado peralta* (no. 20), and *Cajete concavo abierto* (no. 50) types.

After a peak during the Cantera phase, the site rapidly declined in the Late Preclassic period (400 BCE to 200 CE) but remained inhabited during Classic (200–650 CE), Epiclassic (650–900 CE), and Postclassic (900–1500 CE) periods. Thirteen archaeointensity data from Epiclassic potteries have already been published (Hervé, Chauvin, et al., 2019). Here we focused on Preclassic sherds.

We collected 44 fragments of pots that were discovered in the Terrace 6 “the hunter” or close to the Monument 19 (Table S1 in the supporting information). Pots belong to nine types of pottery: *Tenango café*, *Cuautla rojo*, and *Cuautla café* for the Amate phase; *Blanco Amatzinac*, *Anaranjado peralta*, *Laca*, and *Imitacion Laca* for the Barranca phase; *Atoyac pulido con engobe* for the Cantera phase; and *Cajete concavo abierto* for the Late Preclassic phase. We preferentially sampled these types because they are, at least partly, red-colored (Figure 1c). Archaeointensity experiments are generally more successful on these potteries than on the grey-black one, because they are less sensitive to mineralogical alteration (Osete et al., 2016).

3. Archaeomagnetic Experiments

3.1. Rock Magnetism

To investigate the ferromagnetic mineralogy, thermomagnetic curves were measured on powders from nine sherds, one from each pottery type, using the Agico MFK1 apparatus in the CEREGE laboratory. The variation of the susceptibility was measured during heating to 450 or 620°C

and subsequent cooling. Backfield curves of an isothermal remanent magnetization (IRM) from a 1 T magnetic field were acquired on chips from the same sherds with Princeton Micromag Vibrating Sampling Magnetometer at CEREGE.

3.2. Archaeointensity Study

Archaeointensity experiments were performed using the classical Thellier-Thellier method (Thellier & Thellier, 1959) on 135 specimens. Sherds were cut in 0.1 to 4.3 g chips that were embedded in 2 cm cubes of nonmagnetic plaster. Specimens were heated in an ASC TD48-SC furnace with a 30 or 40 μ T laboratory field. Up to 12 temperature steps were performed between 100 and 620°C. Partial thermoremanent magnetization (pTRM) checks every two steps monitored the absence of alteration of the ferromagnetic mineralogy. After each step, remanent magnetizations were measured with a SQUID cryogenic magnetometer (2G Enterprises, model 755R).

For potteries, it is well known that the correction of TRM anisotropy is crucial (Veitch et al., 1984). The most reliable method is the correction by the tensor of TRM anisotropy. The tensor was determined at the specimen level at 530–560°C using six positions (+x, -x, +y, -y, +z, and -z axes) followed by a stability check (Chauvin et al., 2000). The correction was not applied if the difference between the PTRM moments of the first and seventh steps exceeded 10%.

The TRM intensity also depends on the cooling duration, and many archaeointensity studies on baked clays have demonstrated the necessity to correct this cooling rate effect, when the duration of the initial and laboratory cooling is significantly different (e.g., Fox & Aitken, 1980; Hervé et al., 2019). In this study, the cooling of the specimens in the laboratory furnace lasted about 30 min that is much faster than the archaeological cooling. To correct for this effect, the procedure of Gómez-Paccard et al. (2006) was carried out with a slow cooling over 5 hr at the same temperature that the one of the anisotropy correction. This approximation of the archaeological cooling may not be very accurate and was defined mainly for technical reasons. If we cannot totally discard an effect of this possible underestimation of the archaeological cooling on the archaeointensity determinations, it was shown that the TRM intensity increases with the cooling duration following a logarithmic trend (e.g., Genevey et al., 2008; Halgedahl et al., 1980). Hervé, Perrin, et al. (2019) demonstrated that such imprecision on the cooling estimation does not result in a significant

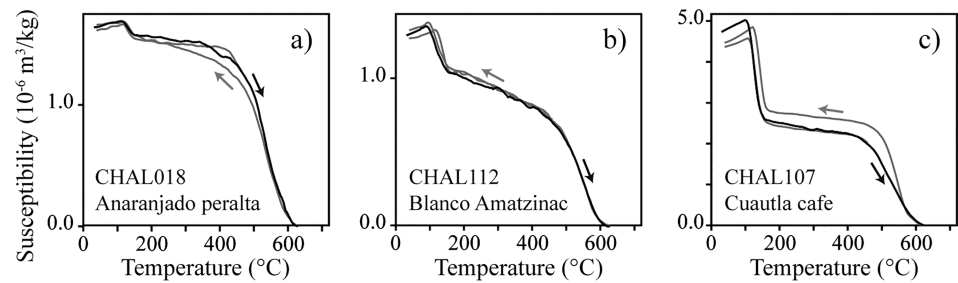


Figure 2. Representative thermomagnetic curves of Preclassic sherds from Chalcatzingo. The black curve is the variation of susceptibility during the heating and the grey curve during the cooling.

inaccuracy of the average archaeointensity. Furthermore, much slower values of cooling rate are not expected at this archaeological period.

4. Results

4.1. Rock Magnetism

Susceptibility curves acquired during successive heating and cooling of the samples (thermomagnetic curves) present a good within 10% reversibility (Figure 2), which highlighted the suitability of Chalcatzingo sherds for archaeointensity experiments. All curves show a ferromagnetic carrier with Curie temperature close to 585°C identified as almost pure magnetite. A second carrier with a Curie temperature around 130–150°C is also seen. This phase represents only a few per cent of the susceptibility in *Anaranjado peralta*, *Atoyac pulido con engobe*, and *Cajete concavo abierto* types (Figure 2a), around 10–20% in *Tenango cafe*, *Cuautla rojo*, and *Blanco Amatzinac* types (Figure 2b) and up to 40–60% in *Cuautla cafe*, *Laca*, and *Imitacion Laca* types (Figure 2c). The low Curie temperature could correspond to a Ti-rich titanomagnetite or to an epsilon iron oxide ($\epsilon\text{-Fe}_2\text{O}_3$) (Lopez-Sanchez et al., 2017). Backfield curves are used to make the difference between the two possibilities. The curves of all specimens are close to saturation at 300 mT magnetization (Figure 3) and do not indicate a significant contribution of high coercivity phases to the remanence. These results exclude the presence of a high-coercivity epsilon iron oxide. The low Curie temperature phase is likely a Ti-rich titanomagnetite.

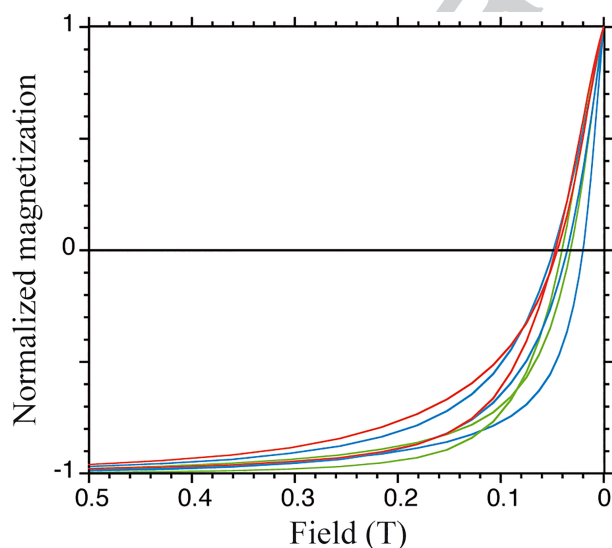


Figure 3. Backfield curves of isothermal remanent magnetization (IRM) normalized to the IRM at 1 T. Samples with red, green and blue curves have thermomagnetic curves showing respectively a low (<5%), medium (10–20%) and high (40–60%) proportion of the low-Curie temperature phase.

4.2. Archaeointensity Study

Most specimens show a secondary component of magnetization, removed between ~200 and ~500°C, that could have been acquired during the cooking use of the ceramic or after the excavation during the storage close to metallic shelves (Figure 4). These specimens are discarded when the characteristic remanent magnetization (ChRM) carries a fraction of the natural remanent magnetization (NRM) less than 35% (f factor), as in PICRIT-03 (Kissel & Laj, 2004), SELCRIT2 (Biggin et al., 2007), and Thellier Tool TTA (Leonhardt et al., 2004) sets of selection criteria. Specimens are also rejected when they display evidence of mineralogical changes on the temperature interval of the ChRM with negative pTRM checks and/or a concave-up behavior (Figure 4a). Successful specimens, showing linear NRM-TRM diagrams with positive pTRM-checks (Figures 4b and 4c), have a quality factor (q) between 5 and 58 (Table S1). Around 80% of the specimens have a maximum angular deviation (MAD) lower than 5°, a deviation angle (DANG) lower than 5°, and a ratio of the standard error of the slope to the absolute value of the slope (β) lower than 0.05. The acceptance rate reaches 58% with 78 accepted specimens from 23 pots. Except for two pots (CHAL019 and CHAL132), at least three specimens are accepted per pot. No pots from *Cuautla cafe*, *Imitacion Laca* types, and only one from *Laca* type provided a good

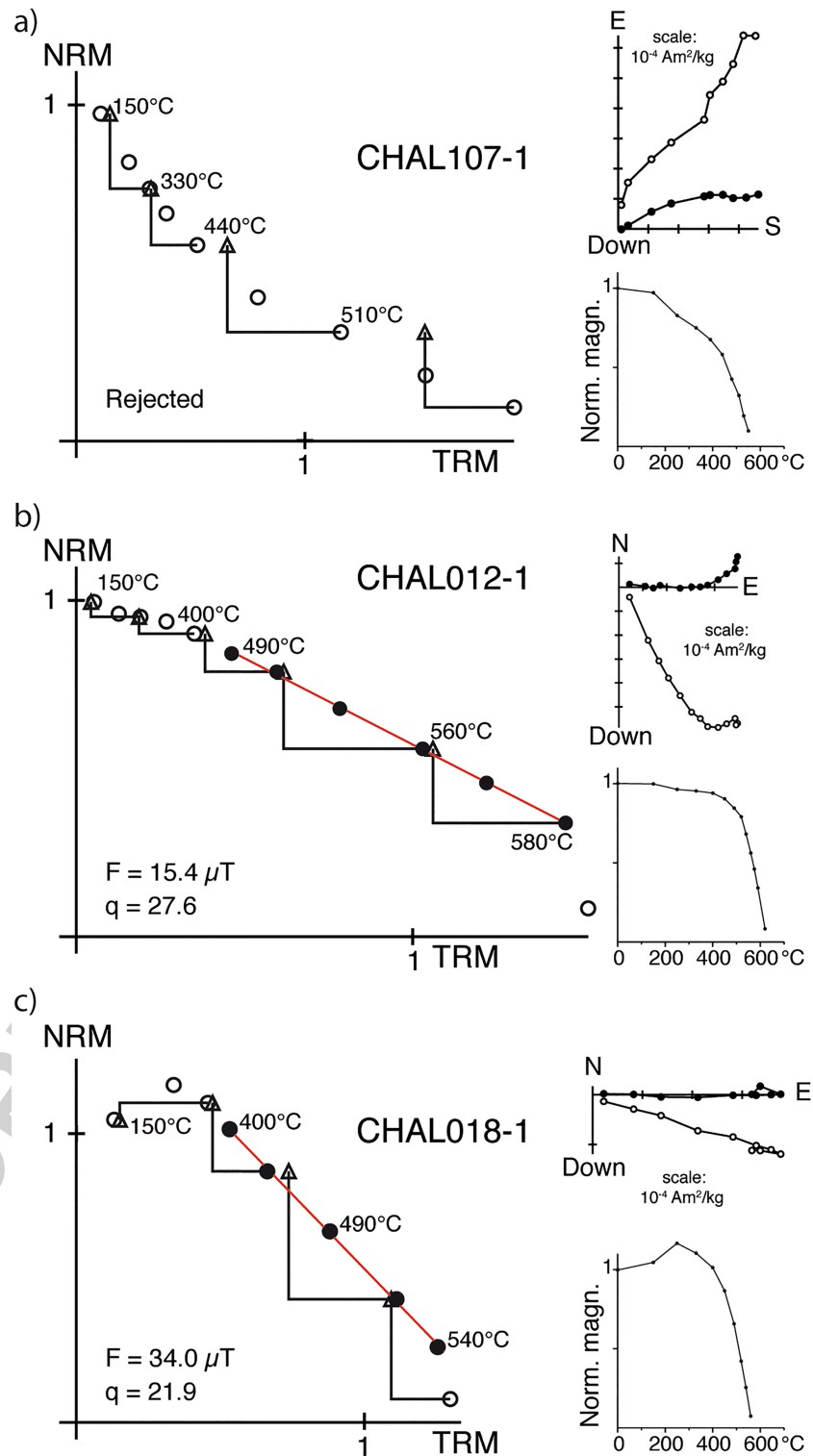


Figure 4. (a) Rejected and (b and c) accepted archaeointensity results with NRM-TRM diagrams on the left and orthogonal plots and demagnetization curves on the right. The solid circles on NRM-TRM diagrams indicate the temperature steps used in the intensity determination. On orthogonal plots, the open (solid) circles denote the projection on the vertical (horizontal) plane.

Table 1
Average Archaeointensities

Date (years CE)	Sherd	Type of pottery	N_{acc}/N_{meas}	$F_{ATRM+CR} \pm SD$ (μT)	$F_{Mexico\ City}$ (μT)	VADM (10^{22} A.m ²)
[-1500; -1100] (Amate phase)	CHAL102	Tenango cafe	3/5	28.7 ± 1.7	29.0	6.5 ± 0.4
	CHAL103	Tenango cafe	3/4	37.7 ± 1.8	38.0	8.5 ± 0.4
	CHAL104	Cuautla rojo	3/5	25.9 ± 0.7	26.1	5.9 ± 0.2
	CHAL105	Cuautla rojo	3/3	26.0 ± 1.2	26.2	5.9 ± 0.3
	CHAL106	Cuautla rojo	4/4	30.6 ± 2.6	30.9	6.9 ± 0.6
[-1100; -700] (Barranca phase)	CHAL012	Blanco Amatzinac	3/5	15.1 ± 0.5	15.2	3.4 ± 0.1
	CHAL013	Blanco Amatzinac	4/4	29.5 ± 2.1	29.8	6.7 ± 0.5
	CHAL015	Blanco Amatzinac	3/5	27.7 ± 1.6	27.9	6.3 ± 0.4
	CHAL016	Blanco Amatzinac	4/6	26.8 ± 1.6	27.0	6.1 ± 0.3
	CHAL112	Blanco Amatzinac	3/4	25.3 ± 1.1	25.5	5.7 ± 0.2
	CHAL113	Blanco Amatzinac	4/4	22.5 ± 2.8	22.7	5.1 ± 0.6
	CHAL018	Anaranjado peralta	4/4	29.9 ± 2.6	30.2	6.8 ± 0.6
	<i>CHAL019</i>	<i>Anaranjado peralta</i>	2/5	31.3 ± 2.5	31.6	7.1 ± 0.6
	CHAL020	Anaranjado peralta	3/5	30.8 ± 2.5	31.1	7.0 ± 0.6
	CHAL116	Anaranjado peralta	5/5	37.1 ± 3.0	37.4	8.4 ± 0.7
[-1100; -600] (Barranca and Cantara temprana phases)	CHAL117	Laca	4/4	26.9 ± 1.3	27.1	6.1 ± 0.3
[-700; -500] (Cantara phase)	CHAL126	Atoyac pulido	3/3	42.2 ± 1.2	42.6	9.5 ± 0.3
	CHAL127	Atoyac pulido	3/4	38.0 ± 2.2	38.3	8.6 ± 0.5
	CHAL128	Atoyac pulido	3/4	24.8 ± 2.0	25.0	5.6 ± 0.5
	CHAL130	Atoyac pulido	4/5	32.3 ± 2.3	32.6	7.4 ± 0.5
	CHAL131	Atoyac pulido	4/4	27.6 ± 1.3	27.8	6.3 ± 0.3
	<i>CHAL132</i>	<i>Atoyac pulido</i>	2/4	31.1 ± 0.5	31.5	7.1 ± 0.1
[-400; +200] (Late Preclassic)	CHAL050	Cajete concavo abierto	4/4	42.9 ± 2.1	43.3	9.7 ± 0.5

Note. Columns from left to right: date of the pottery type, sherd number, pottery type, number of accepted over measured specimens, average archaeointensity corrected for the effects of TRM anisotropy and cooling rate with its standard deviation, and average value relocated to Mexico City, Virtual Axial Dipole Moment. The two samples in italic are not taken into account in the calculation of the secular variation curve because the number of accepted specimens is lower than 3.

result. These three types have the largest Ti-rich titanomagnetite content, suggesting that this magnetic carrier was likely the main responsible of the thermal instability.

The 23 new average archaeointensities have a good internal consistency with experimental uncertainty from 0.5 to 3.0 μT (Table 1). We do not observe a significant difference between samples studied with 30 or 40 μT . The uncertainty decreases after anisotropy and cooling rate corrections (see Table S1). The average values range between 15 and 43 μT . No clear difference is observed between pots from the Terrace 6 and from the Monument 19. In the Barranca phase between 1100 and 700 BCE, pots from *Blanco Amatzinac* type provide lower archaeointensities than those of *Anaranjado peralta* type, suggesting a noncontemporaneity of these two types (Figure 5a).

5. Discussion

5.1. Comparison With Other Mexican Data

A recent critical analysis of the Mexican intensity dataset showed that the data quality is unequal (Hervé, Chauvin, et al., 2019). In this study, the 194 available Mexican data were classified according to their cooling unit consistency and the quality of their palaeointensity protocol: (i) a number of specimens higher than three, (ii) a standard deviation lower than 15%, (iii) a palaeointensity acquired by Thellier-Thellier or microwave protocols and calculated over a well-defined characteristic remanent magnetization, (iv) the presence of the cooling rate correction for archaeological baked clays, and (v) the presence of an anisotropy correction by the TRM tensor for generally anisotropic material such as potteries.

Over the 194 available in Mexico, only 14 data, eight from Duran et al. (2010), three from Rodriguez-Ceja et al. (2009), and three from Rodriguez-Ceja et al. (2012), fulfilled the quality criteria for the period between 2000 BCE and 200 CE. Intensities of Rodriguez-Ceja et al. (2009, 2012) calculated on a secondary or two

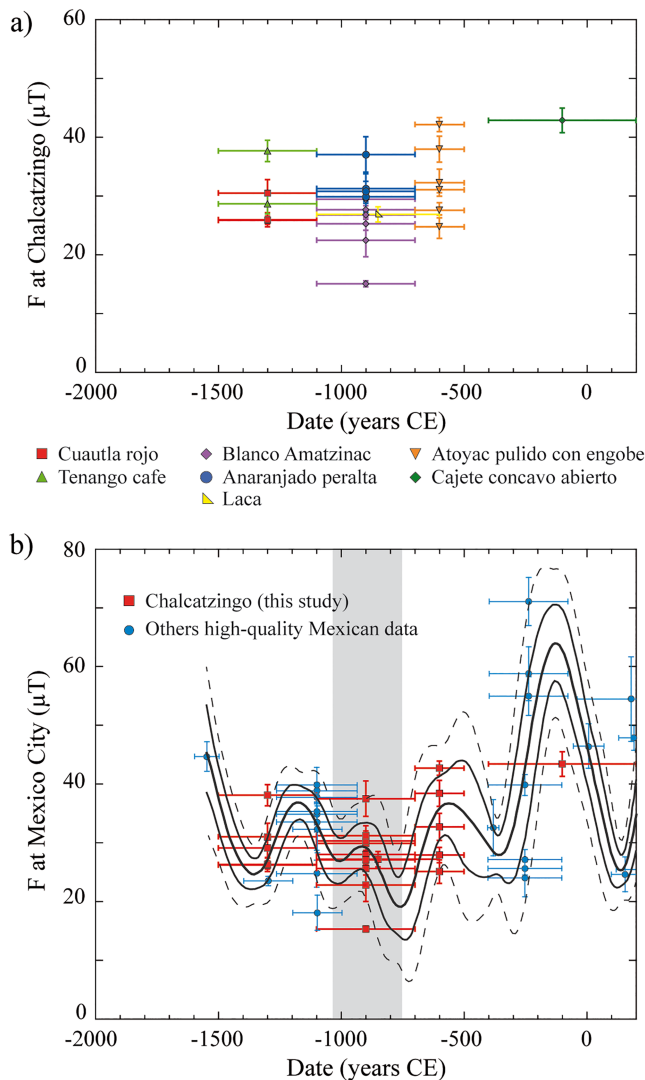


Figure 5. Secular variation of the intensity in Mexico between 2000 BCE and 200 CE. (a) All new Chalcatzingo data classified per pottery type. (b) Comparison of Chalcatzingo data with high-quality data from Mexico and new secular variation curve plotted with its 68% and 95% confidence envelope. All data are relocated to Mexico City. The grey shaded area highlights the period of the geomagnetic “spike” found in sediments of Hall’s cave (Bourne et al., 2016).

in which the apparent dispersion of archaeointensity data, as between 1500 and 500 BCE, likely reflects the secular variation within this period. It makes more difficult the reconstitution of the secular variation with a high resolution. Another outstanding example is Cuanalan site dated between ~400 and ~100 BCE, for which the seven intensity values from Rodriguez-Ceja et al. (2011) and Mahgoub, Juárez-Arriaga, et al. (2019) range from ~25 to ~70 μT , highlighting a fast increase of the geomagnetic field strength during the lifetime of the site.

5.2. Secular Variation Curve of the Intensity in Mexico Between 2000 BCE to 200 CE

From the selected dataset, we calculated a new secular variation curve with the same Bayesian framework (Lanos, 2004) as the one used in recent studies for Austria (Schnepp et al., 2015), Bulgaria (Kovacheva et al., 2014), France (Hervé & Lanos, 2018), and Hawaii (Tema et al., 2017). The inversion process takes into account experimental and age uncertainties on data and minimized the misfit of each data with the curve by exploring the multi-dimensional space of probability densities using Monte Carlo Markov chains (MCMC). The secular variation is given as a smooth continuous curve obtained by averaging cubic splines. The degree

overlapping components of magnetization were not considered. Since the publication of Hervé, Chauvin, et al. (2019), the short high-quality dataset was updated with four new data from Mahgoub, Juárez-Arriaga, et al. (2019) and seven from Mahgoub, Juárez-Arriaga, et al. (2019). Concerning our new Chalcatzingo data, 21/23 respect these criteria; the two sherds (CHAL019 and CHAL132) are discarded because only two successful archaeointensities were obtained by sherd

Most other Mexican data were corrected for anisotropy effects with the Mean (XYZ) method, averaging the archaeointensities of six specimens from the same sample, for which the laboratory field during the Thellier experiments was applied along $+x$, $-x$, $+y$, $-y$, $+z$, or $-z$ axis, respectively. As demonstrated by the experimental tests of Hervé, Chauvin, et al. (2019) and Poletti et al. (2016), this method leads to systematic larger imprecisions and to possible inaccuracies of more than 10 μT . In Hervé, Chauvin, et al. (2019), we decided to retain such archaeointensities if they were calculated with six specimens, because the balanced participation of the three x , y , and z axes reduces the risk of inaccuracy. It is worth pointing out that this was a temporary solution, while waiting for publication of more archaeointensity data corrected with the TRM tensor. Now that new data from Chalcatzingo and Mahgoub et al. (2019a, b) are available between 2000 BCE and 200 CE, we discard all data corrected with the Mean (XYZ) method.

On the Figure 5b, the 25 selected high-quality Mexican data (listed in Table S2) are compared to the 21 new high-quality Chalcatzingo data. All are relocated to Mexico City (19.4°N; 260.9°E) using the Virtual Axial Dipole Moment (VADM) correction. The new Chalcatzingo data almost doubled the number of high-quality data between ~2000 BCE and ~200 CE and especially filled up the temporal gap between 1000 and 400 BCE. They are consistent with the previously published ones.

Another drawback of the Mexican dataset is the imprecision of the archaeological chronology during the Preclassic period (from ~2000 BCE to ~200 CE) with phases (e.g., Amata, Barranca, and Cantara at Chalcatzingo) often covering several centuries. If the age of the boundaries between phases is relatively well controlled, thanks to radiocarbon dates, each phase is rarely subdivided on most sites as in Chalcatzingo. Therefore, there is a succession of wide time slices,

of the smoothing/fitting of the cubic spline function to the data is controlled by a “shrinkage” prior probability (Congdon, 2010). Here, the average curve with the 68% and 95% confidence envelopes is calculated after 50,000 MCMC iterations (Figure 5b).

After a fast decrease from ~ 45 to ~ 25 μT circa 1500 BCE, the average curve exhibits a succession of minima and maxima between 20 and 40 μT until circa 300 BCE. Then the geomagnetic field strength increases rapidly during the third century BCE and reaches a high maximum close to 65 μT in 150–100 BCE. The following intensity decrease during the first centuries BCE and CE seems also rapid. The 150–100 BCE maximum is constrained by three high-quality data on potteries from Mahgoub, Juárez-Arriaga, et al. (2019). The effect of a nonlinear TRM acquisition behavior (Selkin et al., 2007) was not tested in Mahgoub, Juárez-Arriaga, et al. (2019) but is likely not significant, because the laboratory field was close to the archaeointensity estimates.

The oscillatory behavior between 1500 and 300 BCE is the solution found by the modeling to fit the different intensity values observed between data dated in the same archaeological phase. Even though these intensity differences clearly point out that the intensity of geomagnetic field was not stable between 1500 and 300 BCE, one cannot exclude that some current oscillations, as maybe the 800 BCE minimum related to only CHAL012 Chalcatzingo data point, are a statistical artifact due to the specific distribution of ages in the dataset. New archaeomagnetic data with intermediate ages are required to confirm the limits of these small minima and maxima between 1500 and 300 BCE.

The present curve indicates a maximal secular variation rate of ~ 250 nT/year during the third century BCE. This rate is five times faster than the average variation rate of 55 nT/year since 1900 CE at Mexico City (Thébault et al., 2015) but is similar to the fastest rate observed in Europe during the last 3,500 years (e.g., Genevey et al., 2016; Hervé et al., 2017). Our curve does not show any evidence for the “geomagnetic spike” seen in 893 ± 135 BCE in two sedimentary sequences of Hall’s cave in Texas, only 1,200 km away from Mexico City (Bourne et al., 2016) (Figure 5b). Unlike palaeointensities higher than 100 μT , two data (CHAL102, this study; EB-229, Mahgoub, Juárez-Arriaga, et al., 2019) possibly point out an intensity low around 15–20 μT at the same period. These low values will have to be confirmed but already shows that the intensity of the geomagnetic field varied at the regional scale over a larger range than previously thought. Therefore, the hypothesis of Bourne et al. (2016) should be taken with caution, as the relative palaeointensity records from sediments are less reliable than the absolute estimations from archaeological baked clays.

5.3. Comparison With Mexican Master Curves and Global Geomagnetic Models

Two other secular variation curves of the intensity in Mexico have recently been published by Goguitchaichvili et al. (2018) and Mahgoub, Juárez-Arriaga, et al. (2019), both calculated using a bootstrap approach (Thébault & Gallet, 2010) after different data selection (Figure 6a). Mahgoub, Juárez-Arriaga, et al. (2019) curve differs by the inversion process based on the full geomagnetic vector taking into account directional and intensity data at the same time. Whereas Mahgoub, Juárez-Arriaga, et al. (2019) curve is close to our, although smoother, with an intensity maximum at the second century BCE, the curve of Goguitchaichvili et al. (2018) shows a minimum at this period. This inconsistency is likely related to the differences in the selection of the input dataset. Goguitchaichvili et al. (2018) study used a less up-to-date dataset with less stringent selection criteria, where data corrected with the Mean (XYZ) anisotropy method were also taken into account. In Mahgoub, Juárez-Arriaga, et al. (2019) study, selection criteria are very similar to ours and all our selected data, except the new ones from Chalcatzingo, are included in their curve. The higher temporal resolution of our curve between 1500 and 500 BCE illustrates the impact of the Chalcatzingo data.

Contrary to our Bayesian curve, the predictions at Mexico City of the most recent geomagnetic global models, BIGMUDI4k.1 (Arneitz et al., 2019), CALS10k.2 (Constable et al., 2016), COV-ARCH, and COV-LAKE (Hellio & Gillet, 2018), strongly smooth the secular variation (Figure 6a). Our Mexican curve is almost systematically below the models between 1500 and 300 BCE. Models are also unable to reproduce the high-intensity maximum of the second century BCE. It is well known that these models, calculated using spherical harmonic analysis, are limited by the uneven spatial and temporal distribution of the archaeomagnetic database (e.g., Brown et al., 2015). In Mexico, the unequal data quality also plays a major role (Hervé,

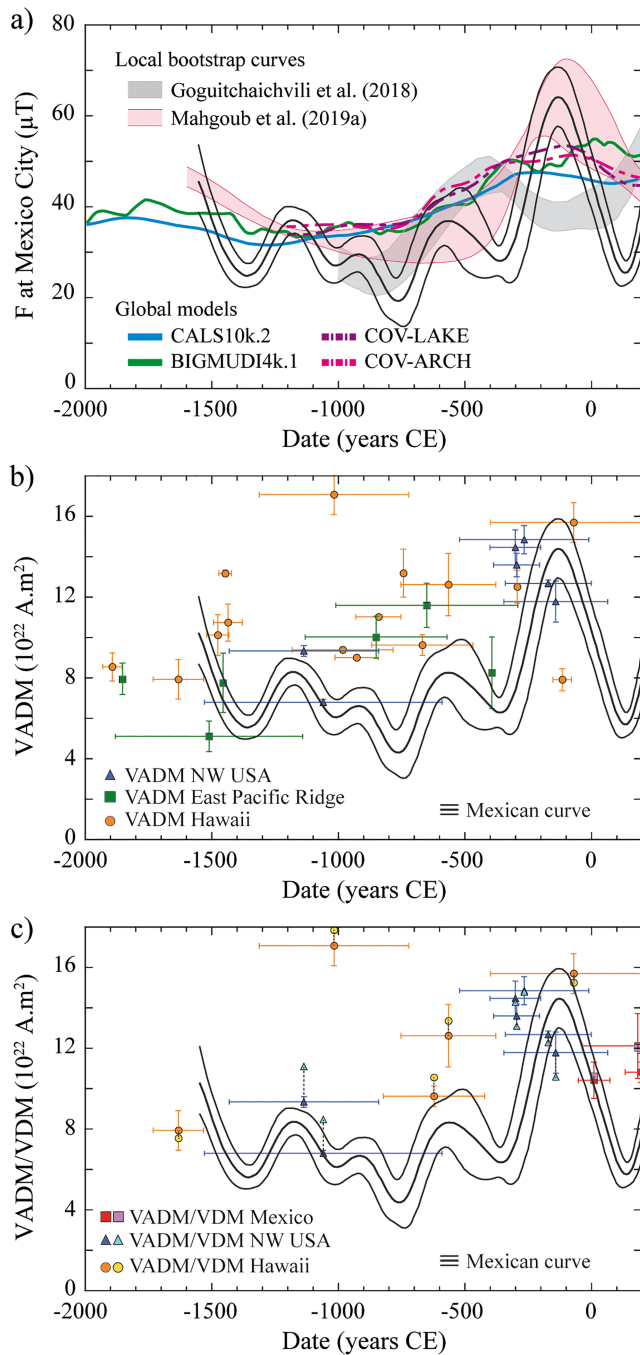


Figure 6. (a) Comparison of the new secular variation curve with the previously published local bootstrap curves for Mexico (Goguitchaichvili et al., 2018; Mahgoub, Juarez-Arriaga, et al., 2019) and the predictions of global models BIGMUDI4k.1 (Arneitz et al., 2019), CALS10k.2 (Constable et al., 2016), COV-ARCH, and COV-LAKE (Hellio & Gillet, 2018). (b) Comparison of the curve of the Virtual Axial Dipole Moment (VADM) in Mexico with data from north-western USA (Champion, 1980), the East Pacific Ridge (Gee et al., 2000; Pick & Tauxe, 1993) and Hawaii (Cromwell et al., 2018; Mankinen & Champion, 1993; Pressling et al., 2006; Pressling et al., 2007; Tema et al., 2017; Valet et al., 1998). (c) Comparison of the VADMs with the Virtual Dipole Moments (VDMs) of full vector data from Mexico, Hawaii and north-western USA.

Chauvin, et al., 2019). The models were calculated using all Mexican data present in GEOMAGIA50.v3 database whatever their quality (Brown et al., 2015). Yet most low-quality data, which were discarded in our local curve, were not adequately corrected for TRM anisotropy and cooling rate effects, which explains the smoothing of the variation by the models and the tendency to overestimate the intensity between 1500 and 300 BCE (Hervé, Chauvin, et al., 2019).

5.4. Comparison With Other Regional Records

In Figure 6b, the secular variation curve of the Mexican VADMs, derived from our intensity curve, was compared to data from surrounding areas selected with the same selection criteria as the Mexican data. There are presently no high-quality archaeointensities on archaeological baked clays in areas surrounding Mexico (southern United States, Central America, and northern South America). The closest are volcanic lava flows from north-western USA around 3000 km away from Mexico City (Champion, 1980) (Figure 6b). The East Pacific ridge data on submarine basaltic glasses (Gee et al., 2000; Pick & Tauxe, 1993), located in the same longitudinal band at ~4,000–5,000 km off the coast of Peru, are also considered. Finally, we compare with the Hawaiian dataset ~6,000 km away at 55–60° of longitude (Cromwell et al., 2018; Tema et al., 2017). All selected data from the three regions are listed in Table S3.

While the VADMs from NW United States agree with the Mexican curve, the Hawaiian and East Pacific Ridge datasets show consistently higher values between 1000 and 300 BCE. This trend may indicate a nondipolar field feature close to Mexico during the first millennium BCE. Even though this hypothesis has clearly to be confirmed by new high-quality data from the Americas, it is interesting to note that CALS10k.2 model predicts a reverse flux patch of the radial component of the geomagnetic field at the core-mantle boundary under Central America around 1000 BCE (Davies & Constable, 2017).

This discussion has also to take into account the fast directional variation observed in these regions during the first millennium BCE (e.g., Hagstrum & Champion, 2002; Tema et al., 2017). Actually, the Virtual Geomagnetic Poles derived from NW U.S. palaeodirections on lava flows were located up to ~70°N from ~60°E to ~0°E between 700 and 400 BCE. As the seven intensities that we selected from NW USA are also full vector, VDMs can be calculated in addition to the VADMs. VDMs are up to $\sim 1.2 \times 10^{22}$ A.m² lower than the VADMs during the second half of the first millennium BCE and $\sim 1.7 \times 10^{22}$ A.m² higher between 1500 and 500 BCE (Figure 6c). The same trend is observed in the few full vector Hawaiian data. These differences between the VDMs and the VADMs minimize the amplitude of the secular variation of the VDM during the first millennium BCE in North America. The same observation is likely also valid for Mexico, where there are only three younger high-quality full vector data from Mahgoub, Juarez-Arriaga, et al. (2019) (Figure 6c). This clearly underlines the necessity of full vector determination for a good resolution of the secular variation of the Earth Magnetic field.

6. Conclusion

The study of potteries from the archaeological site Chalcatzingo provides 23 new archaeointensity data for Mexico between 1500 BCE and 200 CE. Archaeointensities were determined using the classical Thellier-Thellier protocol with corrections for TRM anisotropy and cooling rate effects. With at least three specimens per pot, 21/23 data can be considered as high-quality, which almost doubles the number of archaeointensity data in Mexico during the considered period.

Using a Bayesian framework, a new secular variation curve was calculated at Mexico City between 1500 BCE and 200 CE after selection of Mexican archaeointensity data. After a period of oscillations of the intensity between 20 and 40 μT from 1500 to 300 BCE, the curve shows a large maximum $\sim 65 \mu\text{T}$ in the second century BCE. The curve does not exhibit the “geomagnetic spike” seen in Texas (only $\sim 1,200$ km away from Mexico City) on relative sedimentary data at the beginning of the first millennium BCE. The corresponding VADM varied between ~ 4.0 and $\sim 14.0 \times 10^{22}$ A.m², but this variation amplitude would likely be lower if the directional variation is considered. This highlights the critical need of full vector data in Mexico.

Our new data fill a geographic gap in the global archaeomagnetic database between 1000 and 400 BCE, important to better reconstitute the variation of the dipolar field during this period in which “geomagnetic spikes” were reported in the Middle East. In this manner, the Levantine Iron Age anomaly can be better put in its general context.

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