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## **Variations in benthic foraminiferal assemblages in the Tagus mud belt during the last 5700 years: Implications for Tagus River discharge**

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## 1 **Abstract**

2           We analyzed a 10-m sediment core retrieved at 82 m water depth off the coast of the  
3 Tagus River (Western Iberian Margin, Portugal) to investigate a linkage between variations in  
4 benthic foraminiferal assemblages and Tagus River discharge over the last 5700 years. Benthic  
5 foraminiferal assemblages were studied at high resolution in combination with the stable carbon  
6 and oxygen isotopic composition of fossil shells of *Nonion scaphum*, bulk and molecular  
7 organic matter properties (TOC, TN, C/N ratio,  $\delta^{13}\text{C}_{\text{TOC}}$ ,  $\delta^{15}\text{N}_{\text{bulk}}$ , and BIT index), magnetic  
8 susceptibility, and XRF analyses. Three periods of environmental changes were identified: 1)  
9 high Tagus River discharge in 5750-2200 calendar year before present (cal yr BP), 2) lower  
10 discharge characterized by intense upwelling conditions (2250-1250 cal yr BP), and 3) both  
11 intense upwelling and Tagus River discharge (1250 cal yr BP-present). The data reveal  
12 alternating intense upwelling periods, as shown by the dominance of *Cassidulina carinata*,  
13 *Valvulineria bradyana*, or *Bulimina marginata*, whereas periods of increased river discharge  
14 are indicated by increase of *N. scaphum*, *Ammonia beccarii*, and *Planorbulina mediterraneensis*.  
15 The Tagus River discharge was the strongest during the first period, transporting riverine  
16 material further offshore and preventing the establishment of a mud belt on the mid-shelf  
17 (around 100 m depth). During the second period, a decrease in Tagus River discharge favored  
18 the formation of the Tagus mud belt and strongly influenced the benthic environment by  
19 creating an organic matter stock. During the third period, intense upwelling and increased Tagus  
20 River discharge were recorded by benthic foraminiferal distribution, with an increase of  
21 terrestrial elements present in the mud belt. Furthermore, our results showed that variations in  
22 benthic foraminiferal assemblages corresponded to the well-known climatic periods in the study  
23 area, such as the Roman Period, the Dark Ages, the Medieval Warm Period, and the Little Ice  
24 Age. Our study strongly suggests that benthic foraminiferal assemblages can be used as a bio-  
25 indicator to trace the influence of past river discharge.

26 **Key words:** Holocene, Portuguese Margin, Paleo-reconstruction, North Atlantic Oscillation

27

## 28 **1. Introduction**

29 In the last decades, benthic foraminiferal assemblages have often been used in  
30 paleoceanography to investigate biostratigraphy, paleobathymetry, and abiotic conditions, such  
31 as temperature, salinity, and pH (Jorissen et al., 2007 and references therein). The TROX model  
32 (Jorissen et al., 1995) established that oxygen concentration and organic matter content in the  
33 sediment are the major environmental controls on the distribution of benthic foraminifera in  
34 marine sediments. Subsequent studies emphasized the importance of the quality of the organic  
35 matter for the composition of living benthic foraminifera faunas in various marine environments  
36 (e.g., Goldstein and Corliss, 1994; Suhr et al., 2003; Fontanier et al., 2005, Goineau et al., 2011,  
37 Dessandier et al., 2015). Benthic foraminifera are known to bloom following high marine algal  
38 production (e.g., Kitazato et al., 2000). After a phytoplankton bloom, benthic foraminifera  
39 respond rapidly to the increased influx of fresh organic matter (Fontanier et al., 2003).  
40 Terrestrial input may also impact the organic matter supply and its quality and thereby may  
41 partly control the benthic foraminiferal distribution in continental shelf sediments. This is  
42 particularly true for river-dominated shelves where living benthic foraminifera have been  
43 identified to respond to river inputs (e.g., Mendes et al., 2004; Mojtahid et al., 2009; Goineau  
44 et al., 2011).

45 Dessandier et al. (2016, 2018) have shown through studies on living and dead faunas  
46 that some species could be used as bio-indicators of river discharge and upwelling intensity in  
47 the Portuguese Margin. During the last millennia, these environmental conditions were  
48 influenced by rapid climatic changes (Abrantes et al., 2005, Lebreiro et al., 2006). Within the  
49 Quaternary, the Holocene is characterized by relatively stable climatic conditions with  
50 millennial-scale variability (Bond et al., 1997). The solar variability superimposed on long-term

51 changes in insolation seems to be one of the most likely important forcing mechanisms for the  
52 rapid climate changes over the Holocene (Mayewski et al., 2004). The North Atlantic  
53 Oscillation (NAO) was also described as a major factor for the variability of climate in Europe  
54 (Wanner et al., 2001), which played an important role in the control of the upwelling intensity  
55 and humidity of the Iberian Peninsula (Abrantes et al., 2005; Lebreiro et al., 2006; Fig. 1). The  
56 NAO is an atmospheric process characterized by a seesaw between the Icelandic low and the  
57 Azores high pressures. The positive mode (NAO+) is characterized by well-developed  
58 Icelandic low and Azores high pressures, associated with stronger westerlies over the eastern  
59 North Atlantic and the European continent. The negative mode (NAO-) is characterized by a  
60 rather weak pressure seesaw and reduced westerlies (Wanner et al., 2001). The last millennia  
61 were marked by short-scale climatic changes, mainly controlled by solar activity and NAO-like  
62 processes. Several such periods of short-scale changes were identified, among them the Roman  
63 Period (RP; e.g., Lamb, 1985), the Dark Ages (DA; e.g., Keigwin and Pickart, 1999), the  
64 Medieval Warm Period (MWP), the Little Ice Age (LIA; deMenocal et al., 2000; Trouet et al.,  
65 2009), and a warming during the twentieth century determined by instrumental temperature  
66 measurements of the last two centuries. The major climatic changes, MWP and LIA, have been  
67 described as controlled by NAO+ and NAO-, respectively (Lebreiro et al., 2006; Trouet et al.,  
68 2009). The nature and chronology of these events are still debated (Desprat et al., 2003; Ortega  
69 et al., 2015; Swingedouw et al., 2015), but the influence of these abrupt changes on the rainfall  
70 and upwelling intensity has been observed in the Iberian Margin (Abrantes et al., 2005; Bartels-  
71 Jónsdóttir et al., 2006). During the NAO- mode, the Iberian Margin is marked by increased  
72 humidity, which is responsible for strengthened riverine discharge (Bernárdez et al., 2008). The  
73 NAO+ mode is characterized by an increase in Iberian coastal upwelling (Lebreiro et al., 2006).  
74 Another oscillation, called the Atlantic Multidecadal Oscillation (Kerr, 2000), was also

75 described as a controlling factor of rainfall during the last century that had a greater influence  
76 on the American continent, Sahel, and northern Europe (Knight al., 2006).

77 We tested the applicability of these benthic foraminifera as bio-indicators of past  
78 upwelling intensity and changes in river discharge on a late Holocene sediment sequence from  
79 the Tagus prodelta. We used a multiproxy approach based on organic parameters (total organic  
80 carbon [TOC], total nitrogen [TN], C/N ratio,  $\delta^{15}\text{N}_{\text{bulk}}$ , and  $\delta^{13}\text{C}_{\text{TOC}}$ ), XRF data, benthic  
81 foraminiferal assemblages, and carbon and oxygen stable isotopes of benthic foraminifera to  
82 compare climatic changes of the late Holocene with river regime change. The branched  
83 isoprenoid tetraethers (BIT) index was measured in the core that was previously published in  
84 Warden et al. (2016) and compared with our reconstruction. We applied our reconstruction on  
85 a paleoclimatological record never studied at this water depth, under the influence of the Tagus  
86 River, and provided a new dataset to improve understanding of the effect of NAO-like processes  
87 during the Holocene.

88

## 89 **2. Study area**

90 Our study area was located at the mid-shelf off the Tagus River mouth, on the western  
91 Iberian Margin (Fig. 1) which is a narrow (20–34 km) shelf (Dias et al., 2002) that extends from  
92 37°N to 42°N latitude. The shelf break is located at 140 m depth, on the outer continental shelf  
93 and slope, and three main submarine troughs (the Cascais Canyon, Lisbon Canyon, and Setúbal  
94 Canyon) are related to geological features (Jouanneau et al., 1998). The regional hydrodynamic  
95 regime is driven by the influence of wave action, tidal and rip currents, and storm surges, which  
96 plays a role in determining the particle size distribution on shelf sediments (Oliveira et al.,  
97 2007). However, the Tagus area is protected against swell from the northwest (Jouanneau et al.,  
98 1998). Sandy deposits occur on the inner shelf, where there are high-energy littoral currents.  
99 The dominant regional surface current (the Portugal Coastal Counter Current) flows southward

100 and transports material from the shelf to the deep ocean during winter storms (Vitorino et al.,  
101 2002b). The outer shelf is affected by internal waves, especially during periods when water  
102 masses are stratified, resulting in the deposition of large bodies of sandy and gravelly sands  
103 (Jouanneau et al., 1998). The export of terrestrial sediment and nutrients along the shelf is  
104 predominantly influenced by sediment supplied by three main rivers: the Douro, Tagus, and  
105 Sado (Jouanneau et al., 1998; Dias et al., 2002). At around 100 m water depth, the finest  
106 particles are deposited in accumulation areas, called the mud belt (or mud patch). The mud belt  
107 lies beyond the line where seasonal thermal stratification of surface waters occurs, especially  
108 off the mouths of the Douro and Tagus rivers (Jouanneau et al., 1998; Dias et al., 2002). These  
109 accumulation zones are composed of mixed sources of marine, estuarine, and terrestrial organic  
110 matter (Schmidt et al., 2010). Winter storms can remobilize the sediment and transport it  
111 northward by the action of bottom currents (Dias et al., 2002; Vitorino et al., 2002a), eventually  
112 depositing it on the mid-shelf mud belt between 50 and 130 m water depth (Vitorino et al.,  
113 2002b).

114         The Eastern North Atlantic Central Water (ENACW) is a slope current characterized by  
115 a decrease of salinity under the surface currents, with a minimum of 35.6 at 450–500 m water  
116 depth (Oliveira et al., 2007). The ENACW is upwelled during summer. Between May and  
117 September, the Azores high-pressure system is driven closer to the coast. Together with the  
118 associated northerly winds, this atmospheric system makes the colder, less salty, and nutrient-  
119 enriched subsurface water (60–120 m water depth) rise to the surface along the Iberian Margin  
120 (Fiúza, 1983). This upwelling leads to an increased productivity in the summer along a 50-km-  
121 wide zone. The main upwelling front is oriented along the bathymetry off the 100 m isobath in  
122 the northern part of the Iberian Peninsula; it then turns slightly offshore and southward (Peliz  
123 et al., 2002). Most of the primary producers, especially cyanobacteria and diatoms, that are  
124 dominant in the Iberian Margin increase during upwelling events (Tilstone et al., 2003). Active

125 upwelling periods have a large impact on marine trophic conditions in this area (Fiúza, 1983)  
126 and are characterized by maximal organic carbon exports to the seafloor (Jouanneau et al.,  
127 1998). Periods of strong Iberian river discharge that occur during phases when upwelling is  
128 reduced are characterized by a substantially increased export of continental nutrients, which  
129 triggers phytoplankton production (Prieto et al., 2009; Rodrigues et al., 2009). To a lesser  
130 extent, phytoplanktonic blooms occur in November and between April and May (Ferreira and  
131 Duarte, 1994). In winter, the Azores high-pressure system moves south, which results in  
132 southerly winds and downwelling conditions that lead to the deposition of sediments on the  
133 shelf (Frouin et al., 1990). Upwelling activity and fluvial discharge are consequently the two  
134 major parameters controlling the marine biology on the Iberian Shelf (Lebreiro et al., 2006) due  
135 to their impact on the deposition of organic matter, which is important for benthic organisms.

136         The Tagus is the longest (1,008 km) Iberian river in the central part of the Peninsula and  
137 has a large mesotidal estuary with an area of 340 km<sup>2</sup> (Vale and Sundby, 1987). The Tagus  
138 watershed is about 80,600 km<sup>2</sup> and has an annual mean water discharge of 360 m<sup>3</sup> s<sup>-1</sup> (Jouanneau  
139 et al., 1998), with strong seasonal changes in discharge from 1 to 2,200 m<sup>3</sup> s<sup>-1</sup> (Atlas Nacional  
140 de España), controlled by maximal rainfall in winter (Aguar and Ferreira, 2005; Azevedo et  
141 al., 2008). The Tagus River flows were largely modulated by the NAO during the last century  
142 (Trigo et al., 2004) and the last millennia (e.g., Abrantes et al., 2005; Bartels-Jónsdóttir et al.,  
143 2006; Lebreiro et al., 2006). The river discharge also controls the input of chlorophyll  
144 associated with phytodetritic material in the marine environment and the water column  
145 stratification (Relvas et al., 2007).

146

### 147 **3. Material and methods**

148         Core 64PE332-30-2 was obtained in March 2011 during the Pacemaker 64PE332 cruise  
149 on board the R/V *Pelagia* (38°39'04''N, 9°28'13''W). This 978-cm Kullenberg piston core



150 was retrieved from the Tagus mud belt at 82 m water depth (Fig. 1). Sediment slices 1 or 2 cm  
151 thick were sampled every 10 cm, dried, weighted, and washed through 63 and 150  $\mu\text{m}$  sieves.  
152 For this study, 101 samples of  $>150 \mu\text{m}$  benthic foraminifera were handpicked and placed in  
153 Chapman cells before taxonomic identification under a stereomicroscope. After splitting using  
154 an Otto microsplitter, a minimum of 250 specimens were counted. Diversity indices (Shannon  
155 [S] and Evenness indices) were calculated using the PAleontological STatistics (PAST)  
156 software (Version 2.14; Hammer et al., 2001). The benthic foraminiferal number (BFN), which  
157 represents the number of individuals per analyzed dry sediment mass, was calculated for all  
158 samples.

159 Isotopic analyses were performed on monospecific samples of the benthic foraminiferal  
160 species *Nonion scaphum*, which is present all along the core and is typical of the Iberian mud  
161 belt (Dessandier et al., 2015; 2016), at the EPOC laboratory, University of Bordeaux. This  
162 species was observed alive below the oxygen penetration, between 1 and 2 cm depth in the mud  
163 belt area (Dessandier et al., 2016), suggesting that it may reflect subsurface sediment pore water  
164 rather than bottom water conditions. However, we compared only the data measured on the  
165 same species in every sample, avoiding any bias from the early diagenesis effect for  
166 environmental reconstruction. For each sample, three or four specimens were handpicked and  
167 dissolved in acid via the Micromass Multiprep autosampler system. The resulting carbon  
168 dioxide gas was analyzed against the international reference standard NBS 19 ( $\delta^{13}\text{C} = + 1.96$   
169  $\text{‰}$  / PDB and  $\delta^{18}\text{O} = - 2.20 \text{‰}$  / PDB) using an Optima Micromass mass spectrometer.  
170 Measurements were taken for each depth horizon (1 cm) in triplicate to reduce uncertainties.  
171 The analytical precision was better than 0.05  $\text{‰}$  for  $\delta^{18}\text{O}$  and 0.03  $\text{‰}$  for  $\delta^{13}\text{C}$ .

172 The age model for the sediment core, modified according to Warden et al. (2016), was  
173 based on the magnetic susceptibility (MS) and seven accelerated mass spectrometry (AMS)  $^{14}\text{C}$   
174 radiocarbon dates (Table 1). The MS record of Core 64PE332-30-2 was compared with that of

175 Core GeoB 8903 (Abrantes et al., 2008), which was also retrieved in the Tagus prodelta at ~100  
176 m water depth (Fig. 1). Table 1 summarizes the  $^{14}\text{C}$  AMS dating points of the two cores  
177 considered in this study. Because of the low number of dating points in the upper part of Core  
178 64PE332-30-2, we compared our record of MS and age model with Core GeoB 8903, which  
179 had a substantial number of  $^{14}\text{C}$  AMS dating points in the upper section (Figs. 2 and 3). The  
180 MS was measured on board at 5-cm intervals using a Bartington MS meter with a 12-cm  
181 diameter loop. The final age model was achieved by a linear interpolation between each AMS  
182  $^{14}\text{C}$  age. The  $^{14}\text{C}$  data calibration was made via the program CALIB V0.6 with the Marine13  
183 calibration curve (Stuiver and Reimer, 1993), using the common reservoir age of 400 years  
184 because no regional effect on reservoir age is known in our sampling area (Abrantes et al.,  
185 2005). AMS  $^{14}\text{C}$  ages and the dating points of GeoB8903 were converted to cal yr BP.

186 Core 64PE332-30-2 was scanned with an Avaatech XRF core scanner at NIOZ at 1-cm  
187 resolution. Detailed bulk-chemical composition records acquired by XRF core scanning allow  
188 accurate determination of stratigraphical changes and assessment of the contribution of the  
189 various components in lithogenic and marine sediments (Stuut et al., 2014). The XRF core  
190 scanner uses energy dispersive fluorescence radiation to measure the chemical composition of  
191 the sediment as element intensities in total counts or counts per second (Tjallingii et al., 2007).  
192 After cleaning and preparation of the archive-half core surface and covering with SPEX Certi  
193 Ultralene<sup>®</sup> foil, the core was measured at both 10 kV and 30 kV. Element intensities are  
194 presented as log ratios that are normally distributed and linearly related to log ratios of element  
195 concentration (Weltje and Tjallingii, 2008). Terrestrial exports, such as metals or contaminants,  
196 are indicated mainly by Fe/Ti and Pb/Ti ratios. The Zr/Rb ratio serves as a grain size indicator  
197 (Taylor, 1965), and the Br/Cl ratio indicates organic sediment (Ziegler et al., 2008).

198 Sediments were freeze-dried and ground before the geochemical analyses. TN and  $\delta^{15}\text{N}$   
199 were measured with a Thermo-Scientific Flash 2000 Elemental Analyzer interfaced at NIOZ.

200 The analyses were determined at least in duplicate and the analytical error was, on average,  
201 smaller than 0.1 wt. % for the TN content. The TOC content, the stable carbon isotopic  
202 composition of TOC ( $\delta^{13}\text{C}_{\text{TOC}}$ ), and BIT data were previously published by Warden et al.  
203 (2016). The C/N ratio was calculated as the division of TOC/TN.

204 A principal component analysis (PCA) was performed on 22 samples of normalized  
205 environmental and faunal data, using PRIMER version 6.0 software (Clarke and Warwick,  
206 1994) to compare the response of the faunal and environmental parameters. The two major PCA  
207 scores were plotted to define the different phases of the reconstruction.

208

## 209 **4. Results**

### 210 4.1. Sedimentological features and age model

211 The final age model of Core 64PE332-30-2 (Fig. 3) revealed that the sedimentation rates  
212 were increasing over time, with a first phase of  $\sim 0.06 \text{ cm/yr}^{-1}$  from 5750 to 2200 cal yr BP. In  
213 a second phase, 2200-present, the sedimentation rate was  $\sim 0.52 \text{ cm yr}^{-1}$ . The MS record and  
214 grain size distribution of the Core GeoB 8903 were plotted as a function of age, and the MS  
215 and Ca/Ti record of Core 64PE332-30-2 were plotted as a function of core depth (Fig. 2).  
216 Similar trend signals in the MS records were identified in the two cores, and the Ca/Ti record  
217 measured by XRF showed a similar trend to the grain size record of core GeoB 8903. The MS  
218 record showed three phases, with the first characterized by an increase of MS at 770 cm  
219 sediment core depth, the second a larger increase at around 400 cm, and the third a more stable  
220 trend until the end. This trend was opposite that of the Ca/Ti content, which revealed two  
221 successive decreases of values at the same depths.

222

### 223 4.2. Environmental change phases

224 XRF, organic matter, and benthic foraminiferal isotopes data are plotted in Fig. 4. Three  
225 phases appeared following the environmental changes, mainly determined by the XRF ratios.  
226 Phase I (575-2250 cal yr BP) was characterized by high Zr/Rb and Ca/Ti counts, while the  
227 Br/Cl and Fe/Ti counts were low. The BIT index was low and stable (below  $\sim 0.04$ ), the TOC  
228 content was also stable at around 0.9 wt. %,  $\delta^{13}\text{C}_{\text{TOC}}$  was about  $-24.3$  ‰, and  $\delta^{15}\text{N}_{\text{bulk}}$  was  
229 about  $+3.9$  ‰ (Warden et al., 2016). The TN was the only organic parameter that slightly  
230 increased during this phase, from 0.04 to 0.08 wt. % (Warden et al., 2016). The C/N ratio  
231 showed an opposite trend, with a clear decrease from 18 to 11. Both the carbon and oxygen  
232 stable isotope composition of *N. scaphum* slightly decreased during this interval,  $\delta^{18}\text{O}_{\text{bf}}$  from  $+$   
233  $1.7$  to  $+1.4$  ‰ and  $\delta^{13}\text{C}_{\text{bf}}$  from  $-1.1$  to  $-1.2$  ‰.

234 Phase II (2250-1250 cal yr BP) was characterized by a rapid decrease in Ca/Ti and Zr/Rb  
235 counts, a strong increase in Br/Cl counts, and a slight increase in Fe/Ti. The BIT index increased  
236 slightly, to 0.06. The TOC,  $\delta^{13}\text{C}_{\text{TOC}}$ , and C/N ratio reached high values at the end of the period  
237 (around 1.5 wt. %,  $-23.0$  ‰, and 16, respectively) after a large increase. TN and  $\delta^{15}\text{N}_{\text{bulk}}$   
238 showed a smaller increase; TN moved from 0.9 to 0.13 wt. % and  $\delta^{15}\text{N}_{\text{bulk}}$  from  $+3.6$  to  $+3.9$   
239 ‰. Isotopes measured on benthic foraminifera became more variable at the beginning of the  
240 period, and this trend continued until the end of the record. The decrease in  $\delta^{13}\text{C}_{\text{bf}}$  observed in  
241 phase I was stronger, and  $\delta^{18}\text{O}_{\text{bf}}$  remained fairly constant.

242 Phase III (1250 cal yr BP-present) was characterized by a continued decrease of Ca/Ti  
243 counts and a large decrease of Zr/Rb counts. Br/Cl counts were also unstable and lacking any  
244 clear trend. The Fe/Ti and Pb/Ti counts from Phase III showed an increase until the end of the  
245 record, especially during the last 500 years for Pb/Ti, after a stable trend during the two previous  
246 phases. They both peaked at 250 cal yr BP, when Br/Cl decreased. The TOC content,  $\delta^{13}\text{C}_{\text{TOC}}$ ,  
247 and C/N ratio decreased, to 1.2 wt. %,  $-24$  ‰, and 8, respectively. The  $\delta^{13}\text{C}_{\text{TOC}}$  and TOC  
248 increased during the last 200 years, reaching  $-28.8$  ‰ and 1.2 wt. %, respectively. TN and

249  $\delta^{15}\text{N}_{\text{bulk}}$  were roughly constant before a large increase at the end of the period, reaching 0.17 ‰  
250 and + 4.7 ‰, respectively. The  $\delta^{18}\text{O}_{\text{bf}}$  increased at the beginning from + 1.3 to + 1.8 ‰ and  
251 decreased until + 1.5 ‰ at the end;  $\delta^{13}\text{C}_{\text{bf}}$  showed the opposite trend, with a large decrease from  
252 – 1.5 to – 3.5 ‰. The C/N ratio was slightly decreasing from the start of this period until the  
253 present. Conversely, the BIT index was progressively increasing, reaching 0.12, but dropped in  
254 the most recent sediment horizon analyzed.

255

#### 256 4.3. Benthic foraminiferal distribution over the last 5750 years

257 Fig. 5 shows percentages of the major species (> 5 %). Based on the distribution of these  
258 dominant taxa, three different phases could be identified. The first phase does not map to the  
259 environmental phases and ends at 2500 cal yr BP. During this phase (5750-2500 cal yr BP), *N.*  
260 *scaphum* and *Ammonia beccarii* dominated, making up 30 and 18 % of the total species,  
261 respectively. *Planorbulina mediterraneensis* and *Bolivina spathulata* were relatively abundant  
262 as well, with each ~10 % of the population. *A. beccarii* was particularly dominant (10-20 %)  
263 between 5750 and 4750 cal yr BP and then quickly decreased to ~ 5 %. The epibenthic species  
264 *Cibicides lobatulus* was only > 5 %, while *Hyalinea balthica* and *Uvigerina bifurcata* increased  
265 until the end of the period. This first period was marked by a progressive increase of specific  
266 richness; the S index increased from 28 to 36, and H' from 2.3 to 2.8. The foraminiferal density  
267 (BFN) was relatively low during this period.

268 Between 2500 and 1250 cal yr BP, *Cassidulina carinata* increased sharply, while *N.*  
269 *scaphum* clearly decreased in relative abundance. *Valvulineria bradyana* was nearly absent  
270 during the first phase but became abundant from 2250 cal yr BP and reached 10 % of the fossil  
271 assemblage during the latter period of this second phase. Smaller variation in the relative  
272 abundances of the other dominant species, such as *P. mediterranea*, *A. beccarii*, *H. balthica*,  
273 and *U. bifurcata*, was observed without a clear trend. *V. bradyana* and *Bulimina marginata*

274 increased, and *C. lobatulus* and *B. spathulata* diminished. The highest percentages of *C.*  
275 *carinata* (40 %) were observed between 1750 and 1500 BP, whereas *Bulimina marginata*, *P.*  
276 *mediterraneensis*, and *B. spathulata* decreased. The BFN increased as both specific richness and  
277 the S index began an initial decline; specific richness was decreasing until 30 taxa, and H' until  
278 2.5.

279 The last phase (1250 cal yr BP-present, Phase III) was characterized by higher relative  
280 abundance of *V. bradyana*, dominance of *C. carinata*, and a progressive decrease of *P.*  
281 *mediterraneensis*, *A. beccarii*, and *B. spathulata*. The relative abundance of *N. scaphum* slightly  
282 increased from 1250 cal yr BP, compared to 2600-1250 cal yr BP, but decreased again from  
283 500 cal yr BP. After 1000 cal yr BP, a strong increase was recorded for *C. carinata*, to more  
284 than 40 % at 500 cal yr BP and in the modern period. Deep infaunal species (i.e., *Chilostomella*  
285 *oolina* and *Globobulimina affinis*), *B. marginata*, and *Eggerelloides scaber* also increased  
286 during this phase. The last 200 years showed strong abundances of *C. carinata* (44 %) and a  
287 large loss of both *V. bradyana* and *N. scaphum*, which decreased to ~ 10 %. *Bolivina spathulata*  
288 and *H. balthica* totally disappeared, while *U. bifurcata*, *A. beccarii*, and *B. marginata* declined.  
289 With *C. carinata*, only *B. aculeata* and *E. scaber* are increasing during this modern period. The  
290 last 750 years showed a large decline of BFN, specific richness (~ 20), and S index (~ 2).

291

#### 292 4.4. Multiproxy approach

293 A PCA was performed on the major environmental parameters (TOC, TN,  $\delta^{13}\text{C}_{\text{TOC}}$ ,  
294  $\delta^{15}\text{N}_{\text{bulk}}$ , BIT index, Ca/Ti, Fe/Ti, Br/Cl, Zr/Rb, and benthic foraminiferal stable isotopes) and  
295 on the relative abundances of the major benthic foraminiferal species (*C. carinata*, *N. scaphum*,  
296 *V. bradyana*, *A. beccarii*, *P. mediterraneensis*, *B. spathulata*, *E. scaber*, *H. balthica*, *U.*  
297 *bifurcata*, *B. marginata*, and deep infaunas) (Fig. 6A). PC1 and PC2 explained 64 % (52 and  
298 12 %, respectively) of the total variance observed in the dataset. The relative abundance of *N.*

299 *scaphum*, *A. beccarii*, and *P. mediterranensis* loaded positively on PC1 and negatively on PC2,  
300 together with  $\delta^{13}\text{C}_{\text{bf}}$ , Ca/Ti, and Zr/Rb. Most organic compounds, such as TOC, TN,  $\delta^{13}\text{C}_{\text{TOC}}$ ,  
301 and  $\delta^{15}\text{N}_{\text{bulk}}$ , loaded negatively on PC1 and PC2 with *C. carinata*, Br/Cl, and Pb/Ti. The deep  
302 infaunas and *B. marginata*, *V. bradyana*, and *E. scaber* loaded negatively on PC1 and positively  
303 on PC2, together with Fe/Ti and the BIT index. *Uvigerina bifurcata*, *H. balthica*, and *B.*  
304 *spatulata* loaded positively on PC1 and PC2 with  $\delta^{18}\text{O}_{\text{bf}}$ .

305 The scores of the different samples on PC1 and PC2 were plotted as a function of age  
306 in Fig. 6B. The score of PC1 was stable during the first phase and then showed a slight increase  
307 from 2500 to 2250 cal yr BP. At the start of Phase II (2250 cal yr BP), the score of PC1  
308 decreased until 1500 cal yr BP, then increased again. At the start of Phase III (1250 cal yr BP),  
309 a second progressive decrease started and characterized this phase until the present. The plot of  
310 PC2 showed an increase toward positive loading until the end of phase I. The second phase  
311 (2250-1250 cal yr BP) was characterized by a sharp decrease, before an increase at the  
312 beginning of Phase III that continued until 750 cal yr BP. PC2 decreased, reaching 0 at 500 cal  
313 yr BP, showed a slight increase between 500 and 250 cal yr BP, and decreased again during the  
314 last period of Phase III.

315

## 316 **5. Discussion**

### 317 5.1. Benthic foraminiferal response to environmental and climatic changes

318 Climatic changes during the late Holocene have been well studied in the Portuguese  
319 Margin, especially for the last 3000 years (e.g., Desprat et al., 2003; Abrantes et al., 2005;  
320 Bartels-Jónsdóttir et al., 2006; Alt-Epping et al., 2009). This rendered the Holocene a suitable  
321 period to test benthic foraminifera as bio-indicators for past Tagus River discharge and  
322 environmental changes based on the living foraminiferal calibration developed by Dessandier  
323 et al. (2016, 2018) in the same location. Previous sediment cores in the area extended back to

324 2000 (Abrantes et al., 2005; Bartels-Jónsdóttir et al., 2006) or 3000 cal yr BP (Alt-Epping et  
325 al., 2009). The 10-m core retrieved in our study provided a record that dated to 5750 cal yr BP.  
326 Next, we present the three phases defined by environmental changes related to sediment  
327 supplies, as shown by the XRF ratios and visible in the scores of the PCA (Fig. 6B).

328         The first phase was characterized by positive values of PC1 and a continuous increase  
329 in PC2, starting from negative values. The Zr/Rb ratio indicates a bigger grain size, which had  
330 a strong positive loading on PC1 together with Ca/Ti. Ca was previously linked to grain size in  
331 the Portuguese Margin and thought to be associated with coarse reworked shells of  
332 macrobenthic organisms (Martins et al., 2007; Abrantes et al., 2008; Alt-Epping et al., 2009).  
333 The two different increases of Zr/Rb and Ca/Ti were two successive phases in grain size  
334 decrease from Phase I to Phase III (Fig. 4). During Phase I, the TOC content of the sediment  
335 and the Br/Cl were very low; both indicated low amounts of organic compounds, likely because  
336 of this coarse grain size. This phase was characterized by river-influenced species, such as *N.*  
337 *scaphum*, *A. beccarii*, and *P. mediterraneis*. This first group of species consequently  
338 represents the river discharge bio-indicator in this study. Among them, *N. scaphum* was  
339 interpreted as an indicator of active upwelling conditions in previous studies of the Portuguese  
340 Margin (Bartels-Jónsdóttir et al., 2006). However, the results from late winter showed that this  
341 species was clearly dominant in the living community of the Portuguese inner shelf during  
342 winter. By contrast, during active upwelling context, it was present only in the dead community,  
343 reflecting lower relative abundances (Dessandier et al., 2016; 2018). The large occurrences of  
344 *A. beccarii* and *P. mediterraneis* during this first phase indicated strong bottom currents and  
345 coarse sediments (e.g., Murray, 2006; Schönfeld, 2002), whereas *C. lobatulus* was previously  
346 believed not to be endemic to this study area (Dessandier et al., 2018). The presence of the latter  
347 was probably an indicator of transport from the estuary to the shelf, as this species is typically  
348 found in Portuguese estuaries (Martins et al., 2015). The dominance of *N. scaphum* and *A.*



349 *beccarii* further suggested inputs from the estuary of phytoplankton or nutrients that may boost  
350 local marine productivity, such as coccolithophores, as has been previously observed where the  
351 Douro River flows into the ocean, where these species dominated the living fauna in late winter  
352 (Dessandier et al., 2015). This may indicate an influx of relatively labile organic matter,  
353 consistent with the progressive increase of faunal diversity during this period.

354 The second phase began with the Tagus mud belt establishment at around 2250 cal yr  
355 BP, which was represented by the plot of PC1, when the Ca/Ti and Zr/Rb counts clearly  
356 decreased and terrestrial elements (Fe/Ti, Pb/Ti, and BIT index, Fig. 4) increased. This suggests  
357 a deposition of finer sediments of terrestrial origin in the Tagus prodelta during this period. This  
358 important change in sedimentary conditions toward muddier and organic-rich sediment  
359 (Martins et al., 2006) was observed in the Portuguese Margin earlier, at around 2000 cal yr BP  
360 (Alt-Epping et al., 2009; Martins et al., 2007), and indicated the onset of the Tagus mud belt,  
361 which was mainly composed of sediment exported from the Tagus estuary (Jouanneau et al.,  
362 1998). This increase of organic compounds was also highlighted by the large increase in the  
363 Br/Cl ratio during this phase. The reduction of bottom water currents driven by decreasing wind  
364 and more humid conditions have been postulated as the physical processes that triggered a  
365 strong Tagus River export of sediment and led to the formation of the mud belt around this time  
366 (Alt-Epping et al., 2009). The mud belt formation was responsible for the increased  
367 accumulation of organic matter, especially during Phase II, corresponding to the increased  
368 levels of Br/Cl, TOC, and TN content (Fig. 4). The strong negative loading of both  $\delta^{13}\text{C}_{\text{TOC}}$  and  
369 TOC on PC2 might suggest that the major supply of organic matter was driven by marine  
370 production, particularly during the strong upwelling conditions that characterized Phase II. The  
371  $\delta^{18}\text{O}_{\text{bf}}$  in this area was primarily controlled by salinity, with a reduced temperature effect,  
372 showing an increase in salinity during upwelling events (Lebreiro et al., 2006). This phase was  
373 characterized by a large increase of *C. carinata*, *V. bradyana*, and *B. marginata* abundances,

374 indicating that higher trophic levels occurred in the sediments, caused by high organic matter  
375 content in the mud belt. In this area, the active upwelling period corresponds to the most  
376 eutrophic conditions, mainly highlighted by the abundance of *C. carinata*, associated with  
377 marine organic matter (Br/Cl and  $\delta^{13}\text{C}_{\text{TOC}}$ ; Fig. 6). *Cassidulina carinata* was already interpreted  
378 as highly dominant in an active upwelling context and adapted to cold, nutrient-rich waters  
379 (Bartels-Jónsdóttir et al., 2006; Martins et al., 2006). Together with this species, *V. bradyana*,  
380 *B. marginata*, *H. balthica*, and *U. bifurcate* have essentially been found only in the dead  
381 community and with almost no occurrence in the late winter (Dessandier et al., 2016). The  
382 increase in *V. bradyana* abundance correlated with the onset of Phase II and the mud belt. This  
383 result corroborated this species' need for rich trophic conditions and suggested that the second  
384 phase stabilized with the onset of strong upwelling conditions. *Nonion scaphum* and *V.*  
385 *bradyana* were highly dominant during this interval, although *N. scaphum* was less abundant  
386 than in Phase I. These two species live under organic-rich conditions and can tolerate anoxic  
387 sediments (Fontanier et al., 2002; Barras et al., 2014), which may suggest that large terrestrial  
388 inputs in this area led to periodic anoxia. The upwelling events create ideal environmental  
389 conditions for diatom blooms that resulted in the presence of other species during summer  
390 periods (Dessandier et al., 2016). This seasonal production indicated a preference of these  
391 opportunistic species for summer periods, when the influence of upwelling is at its maximum.  
392 During upwelling periods, diatoms are the major phytoplankton group responding to cold and  
393 nutrient-rich ENACW. However, the preservation of this group as fossils is dependent on high  
394 fluxes of individuals to the seabed (Abrantes et al., 1988), which constrains the use of this proxy  
395 for environmental changes. Conversely, winter periods characterized by maximum continental  
396 runoff had evidence of coccolithophore blooms, which responded to stratified waters (Abrantes  
397 and Moita, 1999). Since phytoplanktonic groups favor certain environmental conditions, these  
398 phytodetrital sources for benthic foraminifera may partially explain the seasonal variation in

399 species composition. This result may be the consequence of preferences for the specific  
400 component of seasonally deposited phytodetritus, as has been observed in the Antarctic (Suhr  
401 et al., 2003).

402 During Phases II and III,  $\delta^{13}\text{C}_{\text{bf}}$ , representing the exported phytodetritus to the seafloor  
403 (Curry et al., 1988), indicated higher primary production, whereas the increase of  $\delta^{15}\text{N}_{\text{bulk}}$  may  
404 illustrate nutrient degradation or a stronger influence of estuarine sources of organic matter  
405 (Owens, 1985). Phase III was characterized by a change in organic matter sources from  
406 predominantly marine origin (negative loads on PC2) to predominantly terrestrial origin  
407 (positive loads on PC2). This period showed a drop in foraminiferal diversity corresponding to  
408 the dominance of *C. carinata* (up to 43 %). The increase in abundance of the deep infaunal  
409 species at around 1000 cal yr BP suggested this is when the enrichment in organic matter of the  
410 prodelta occurred and might indicate episodic periods of dysoxia or even anoxia related to  
411 potential eutrophication, such as the occurrence of *G. affinis* and *C. oolina*, well known to live  
412 in highly eutrophic conditions, often below the oxic sediments (Jorissen et al., 1998; Mojtahid  
413 et al., 2010a). The faunal evolution, in terms of assemblages and diversity within Phases II and  
414 III, revealed other controlling factors independent of the presence of the Tagus mud belt,  
415 represented by the PC2 (Fig. 6B). An increase in pollutants during this period was thought to  
416 occur concomitantly with the disappearance of *H. balthica* and appearance of *E. scaber*, as has  
417 been observed off Iberian rivers (Diz et al., 2002; Bartels-Jónsdóttir et al., 2006). The increase  
418 of pollutants indicated by the appearance of *E. scaber* also fits the results of Alve (1995), who  
419 described *E. scaber* as a pollution-tolerant species under aquaculture influence.

420

## 421 5.2. Reconstruction of environmental evolution in the Tagus prodelta during the last 5750 years

### 422 5.2.1. First phase (5750-2250 cal yr BP): High Tagus River discharge

423           The modern sediment cover in the Tagus prodelta is mainly supplied by terrestrial silts  
424 and clays exported by Tagus River discharge (Jouanneau et al., 1998). The influence of sea  
425 level changes on the Iberian Margin is no longer significant on the sedimentation after ~ 7000  
426 yr BP (Vis et al., 2008). The sedimentary evolution of the Tagus prodelta was mainly controlled  
427 by Tagus River discharge and impacted by climatic changes after this period. This area was  
428 particularly characterized by very humid conditions between 6500 and 5500 yr BP,  
429 corresponding to the African Humid Period (deMenocal et al., 2000; Renssen et al., 2006; Vis  
430 et al., 2010). The coarse sediments observed during Phase I, highlighted by both environmental  
431 and faunal evidence, are in good agreement with conditions described by Rodrigues et al.  
432 (2009), who interpreted intense deforestation and soil destabilization as factors responsible for  
433 the increased current velocity of the Tagus River. These coarse sediments with low organic  
434 content and of marine origin were similar to sediments of modern inner shelf conditions  
435 observed down to 50 m water depth in the Portuguese Margin (Schmidt et al., 2010). These  
436 conditions are synthesized in Fig. 7A and demonstrate that when the Tagus River discharge was  
437 large, it prevented any deposition of fine material on the mud belt, resulting in low  
438 sedimentation rates. The fine sediment deposits of terrestrial origin were probably transported  
439 further offshore during this phase. The position of Core 64PE332-30-2 might have been a zone  
440 of fine sediment bypass similar to what is currently observed on the inner shelf. This explains  
441 the low organic content, the low BIT index, and the  $\delta^{13}\text{C}_{\text{TOC}}$  that signaled the marine origin of  
442 the sediments. Similar conditions were observed for this period in the Galicia mud deposit  
443 (Martins et al., 2007; Bernárdez et al., 2008) and off the Guadiana River (Mendes et al., 2010)  
444 and were believed to indicate a strong hydrodynamic regime. Tagus River floods have been  
445 reconstructed in the lower Tagus River valley between 4900 and 3500 yr BP and were  
446 associated with strong deforestation during this period (Vis et al., 2010).

447           The transition between Phases I and II was marked by a shift between the faunal and  
448 environmental signals, between 2600 and 2250 cal yr BP (Fig. 5). The species that first reacted  
449 to this, *C. carinata*, started to increase in abundance, resulting in a decrease in the relative  
450 abundances of the other species, including the taxa indicators of river discharge. This reflected  
451 the very opportunistic nature of *C. carinata* and provided a more accurate signal of the  
452 environmental change than the geochemical parameters, which merely recorded the mud belt  
453 conditions. A progressive decrease of fluvial influence was visible following the strong decline  
454 of *A. beccarii* and *N. scaphum*, which has also been observed during the last 3000 years within  
455 the Ría de Vigo (Diz et al., 2002). This may be the transition between colder and wetter  
456 Subboreal conditions and warmer, dryer Sub-Atlantic conditions at 3000 yr BP (Alt-Epping et  
457 al., 2009; Bernárdez et al., 2008). The Sub-Atlantic period was associated with the reduced  
458 influence of winds on the Iberian Margin (Martins et al., 2007), a reduction in hydrodynamic  
459 marine currents, and the collection of fine sediments in the Galicia mud deposit. These  
460 conditions were responsible for the lateral movement of the mud deposit on the shelf, with the  
461 construction of the mud belt at the beginning of Phase II (2250 cal yr BP). The accumulation  
462 of muddy sediments on the shelf off the Tagus River at ~ 2000 years BP was synchronous with  
463 the establishment of the mud belt off the Douro River (Drago et al., 1998). This coordination  
464 suggests a response to regional rather than local change in the climate, responsible for a dryer  
465 period and an increase of clay accumulation closer to the Tagus River mouth, due to the  
466 decrease in Tagus discharge.

467

#### 468 5.2.2. Second phase (2250-1250 cal yr BP): High upwelling intensity

469           After the onset of the Tagus mud belt at ~ 2250 cal yr BP, two alternative regimes  
470 prevailed, one characterized by intense upwelling and the other by strong river discharge, as  
471 summarized in Fig. 7B. During this phase, TN,  $\delta^{13}\text{C}_{\text{bf}}$ , and  $\delta^{15}\text{N}_{\text{bulk}}$  showed an increase in

472 productivity, probably due to a change in Tagus River flux conditions. However, this was not  
473 recorded by all environmental parameters. The faunal results could indicate that the organic  
474 matter supply during Phase II was related to upwelling events more than Tagus River inputs.  
475 The decrease in the C/N ratio was not in agreement with  $\delta^{13}\text{C}_{\text{TOC}}$  data, making it difficult to  
476 evaluate the organic matter source. The  $\delta^{13}\text{C}_{\text{TOC}}$  and the C/N ratio are often used to determine  
477 the sources of organic matter (e.g., Hedges and Parker, 1976; Peters et al., 1978; Alt-Epping et  
478 al., 2007). However, the C/N ratio is affected by the preferential remineralization of nitrogen in  
479 marine sediments or nitrogen sorption onto clay minerals (Schubert and Calvert, 2001), and the  
480  $\delta^{13}\text{C}_{\text{TOC}}$  signal could be from a mixture of C3 and C4 plants, mimicking the isotopic signal of  
481 marine algae (e.g., Goñi et al., 1998). Both of these indicators in the paleorecords might be  
482 affected by organic matter degradation. Additionally,  $\delta^{15}\text{N}_{\text{bulk}}$  could be affected in the study  
483 area by a higher influence of agriculture and pollution (Alt-Epping et al., 2009). Nevertheless,  
484  $\delta^{13}\text{C}_{\text{TOC}}$  and the C/N ratio were useful to discriminate the marine and terrestrial sources of  
485 organic matter in another study on the Portuguese Margin (Schmidt et al., 2010). In the present  
486 study, only the C/N ratio seemed to be clearly affected by the early diagenesis. These organic  
487 parameters were essential to compare the environmental signal with benthic foraminifera, and  
488 their disagreement confirms the importance of using a multiproxy approach that included bio-  
489 indicators that are less affected by organic matter degradation.

490 Phase II began at 2250 cal yr BP with a low amount of terrestrial input (indicated by the  
491 negative loading of PC2; Fig. 6) and lasted until 1800 cal yr BP, which corresponded to the RP  
492 (Lamb, 1985; Bernárdez et al., 2008). Despite the increase in eutrophy-tolerant species, the  
493 faunal diversity and BFN were low at the beginning of the RP. The deposition of contaminants  
494 and increased sediment accumulation rate may have limited faunal production during this  
495 period. Lebreiro et al. (2006) also observed a large export of terrestrial particles into the Tagus  
496 prodelta during the RP. This export was interpreted as a consequence of anthropogenic

497 activities, such as Roman gold mining, along the Tagus River. This export of terrestrial material  
498 may have been enhanced by the NAO- phase that occurred during that period and led to  
499 intensified rainfall in the Portuguese Margin (Abrantes et al., 2005). Periods of NAO- have  
500 strongly influenced Iberian river discharge, especially on the Tagus River (Trigo et al., 2004).  
501 The northern part of the Iberian Margin is marked by varying rainfall responses. This zone,  
502 which is very close to the limit of the NAO influence, is known to alternate between positive  
503 and negative correlation with the NAO and humid conditions (Alvarez and Gomez-Gesteira,  
504 2006). Desprat et al. (2003) also identified warm and relatively humid conditions during the RP  
505 in the Ría de Vigo. A strong river regime such as this one was not observed off the Capbreton  
506 in the bay of Biscay (Mojtahid et al., 2009), suggesting an anti-correlation of the NAO phases  
507 between the two study areas.

508         The end of Phase II showed a decrease in the Tagus River influence that corresponded  
509 with an increase in positive loading of PC2, characterized by high marine organic matter content  
510 likely brought by active upwelling conditions. The increased abundance of *C. carinata*  
511 corroborated this context in the middle of Phase II, corresponding to a period of intense  
512 upwelling activity, which has also been described in the same area in a study using benthic  
513 foraminifera (Bartels-Jónsdóttir et al., 2006). The dominance of the opportunistic species *C.*  
514 *carinata* constrained faunal diversity, as observed in other studies (Fontanier et al., 2003;  
515 Dessandier et al., 2018). However, *C. carinata* did not increase until the end of Phase II,  
516 suggesting that the intensity of upwelling slowed down at the end of this phase. By contrast,  
517 the Tagus mud belt built up, allowing an increased deposition of marine organic matter until  
518 the end of Phase II. This difference could be the result of biased organic parameters (e.g., caused  
519 by early diagenesis). In addition, this study area was characterized by organic matter from  
520 different sources, which were recorded as a mixture of marine and terrestrial compounds by the  
521 organic parameters. What was revealed by the organic parameters as a record of upwelling

522 activity could instead result from high productivity throughout the entire year. Therefore,  
523 benthic foraminiferal species recorded the organic matter sources more accurately.

524

525 5.2.3. Third phase (1250 cal yr BP-present): Alternating upwelling and Tagus discharge  
526 influence

527 The beginning of Phase III (Fig. 7C), which corresponds to the DA, was followed by  
528 the MWP, which occurred between 1100 and 600 yr BP, as reported in several studies  
529 performed on the Portuguese Margin (e.g., Desprat et al., 2003; Rosa et al., 2007). The DA was  
530 described as a period characterized by strong upwelling activity off the Douro River (Rosa et  
531 al., 2007). The MWP is also well known in the Portuguese Margin as a period characterized by  
532 active upwelling conditions triggered by NAO+ conditions (Abrantes et al., 2005; Bartels-  
533 Jónsdóttir et al., 2006; Rosa et al., 2007; Rodrigues et al., 2009). The high primary productivity  
534 was linked to upwelling events and created eutrophic conditions in sediments that allowed for  
535 less competition among *B. spathulata*, *P. mediterraneensis*, *Criboelphidium gerthi*, and *H.*  
536 *balthica*. This period, characterized by a warm and dry climate, was influenced by  
537 anthropogenic activities and soil erosion (Rodrigues et al., 2009) that caused Fe and Pb to be  
538 massively transported via Tagus River runoff to the prodelta. Dry soils restrained infiltration  
539 and possibly triggered large flooding of the Tagus River (Benito et al., 2003). The increased  
540 BIT index (Warden et al., 2016) and Fe/Ti and Pb/Ti ratios during the beginning of the MWP  
541 suggested increased terrestrial input, corresponding with the increased loading of PC2 in the  
542 first part of Phase III and the accumulation of finer sediments.

543 The LIA, between 600 and 100 yr BP, was characterized by the increase in terrestrial  
544 material, as shown by the abrupt increase in the BIT index and Fe/Ti and Pb/Ti ratios at 250 cal  
545 yr BP. This increase was also highlighted by the sharp decrease in PC2 (Fig. 6B), which could  
546 be linked to the Lisbon earthquake in 1755 AD (200 cal yr BP; Abrantes et al., 2008) or to



547 massive exports of fine sediments from the Tagus River (low Zr/Rb). This period has been  
548 described as characterized by abrupt cooling and wet conditions (Bradley, 2000). The LIA was  
549 affected by high-frequency episodic Tagus River paleo-floods (Benito et al., 2003) and has been  
550 associated with the transport of fine sediments from the continent via discharge from the Tagus  
551 River under NAO- conditions (Abrantes et al., 2005; Bartels-Jónsdóttir et al., 2006). During the  
552 LIA, an abrupt decrease in the abundance of *C. carinata* may indicate that a reduction in  
553 upwelling conditions occurred. *V. bradyana*, *E. scaber*, *A. beccarii*, and deep infaunas  
554 increased, possibly as a consequence of episodic anoxia and the presence of contaminants. The  
555 presence of the deep infaunas that can tolerate refractory organic matter (Murray, 2006)  
556 suggests a decrease in organic matter quality, as was also observed in the Galicia mud deposit  
557 during this period (Martins et al., 2006). This decrease in quality may be the major cause of the  
558 decreased infaunal diversity that occurred between 1000 cal yr BP and the present.  
559 Nevertheless, the faunal results suggested that intense upwelling periods also occurred during  
560 the LIA, as demonstrated by the increased abundance of *C. carinata*. This highlighted that this  
561 period was unstable and characterized by several environmental changes. The results from this  
562 study were in good agreement with previous studies in the same area for the last 2000 years  
563 (Abrantes et al., 2005; Bartels-Jónsdóttir et al., 2006; Alt-Epping et al., 2009), except for small  
564 time shifts in the upwelling versus Tagus River discharge periods corresponding to the  
565 transition between Phases I and II. These time shifts could be the consequence of different age  
566 models. Bartels-Jónsdóttir et al. (2006) demonstrated that an “intense upwelling period”  
567 occurred, followed by a “very intense upwelling period” during the MWP, and finally a large  
568 Tagus River discharge during the LIA, all determined through the analysis of benthic  
569 foraminiferal distribution.

570           There were some differences during the MWP and the LIA between the reconstructions  
571 from this study and the results from Bartels-Jónsdóttir et al. (2006). These variations could be

572 due to a different interpretation of certain species, specifically *N. scaphum*, which was  
573 determined to be controlled by upwelling in previous studies; by contrast, it was used as a river  
574 discharge proxy in this study, based on the ecological results from the study area (Dessandier  
575 et al., 2016). The use of environmental data combined with the composition of the major species  
576 in this study may allow a better interpretation of the mud belt onset than merely using benthic  
577 foraminiferal data. Our results were also in good agreement with the reconstruction of the Tagus  
578 River discharge and the upwelling strength reported in Abrantes et al. (2009).

579         Finally, during the last century, this area was contaminated by anthropogenic pollution,  
580 including numerous trace metals that were measured in excess in the Tagus estuary, such as  
581 AS, Pb, Zn, Cu, and Cd (Caçador et al., 1996; Jouanneau et al., 1998). The catchment area was  
582 also influenced by anthropogenic contaminants, such as domestic sewage and industrial wastes  
583 (e.g., petrochemistry, fertilizers, smelters; Carvalho, 1997). Our analyses were not performed  
584 to identify the sources of anthropogenic pollution; however, the XRF signal shows a clear  
585 increase of several contaminants (such as Pb) in the most recent interval. We do not have  
586 enough faunal resolution and environmental data to investigate the anthropogenic influence on  
587 benthic ecosystems, but we assume that the faunal assemblages respond to this anthropogenic  
588 activity, as shown in other environments (e.g., Alve, 1995). The faunal distribution shows  
589 increased abundances of deep infaunal species and of *E. scaber* but the disappearance of other  
590 species, such as *H. balthica*, as has already been observed as a consequence of anthropogenic  
591 pollution in the Tagus prodelta (Bartels-Jónsdóttir et al., 2006). The last century was also  
592 characterized by the construction of dams, which have likely influenced the sequestration of  
593 organic matter in the estuary and changed the influence of the Tagus River discharge on the  
594 shelf through increased input of finer material (Jouanneau et al., 1998) and increased correlation  
595 with NAO phases (Trigo et al., 2004), all of which could have significant environmental  
596 consequences that will be crucial to understand in the future.

597

## 598 **6. Conclusions**

599         The results of this study demonstrate the validity of using benthic foraminifera as bio-  
600 indicators of past river discharge and upwelling intensity in the Tagus prodelta during the last  
601 5750 years. Major environmental changes, linked with late Holocene climatic variations, are  
602 summarized by key benthic foraminiferal taxa data in Fig. 8. This study showed an additional  
603 phase, between 5750 and 2250 cal yr BP, that has not yet been investigated in the Portuguese  
604 shelf. This phase was characterized by coarser sediment cover and dominated mostly by *N.*  
605 *scaphum* and *A. beccarii*, suggesting a very dynamic Tagus River system that facilitated the  
606 dispersion of coarse terrigenous particles onto the prodelta.

607         The progressive decrease of the Tagus River flow, resulting in the establishment of the  
608 modern Tagus mud belt, was the major process explaining the environmental changes that  
609 occurred before Phase II. The presence of a transition period (2500-2250 cal yr BP) at the end  
610 of Phase I was only supported by the composition of the benthic foraminifera; geochemical data  
611 did not reveal any environmental changes. This might be due to the extremely poor preservation  
612 of organic matter in coarse sediments (in which benthic foraminifera are fairly well preserved).

613         Phase II (2250-1250 cal yr BP) was characterized by fine continental deposits,  
614 responsible for a better food stock for benthic organisms throughout the record. This phase was  
615 marked by the increase of *C. carinata*, *V. bradyana*, and *B. marginata*, opportunistic species  
616 that mark the increase in upwelling intensity in this area, revealing the strongest upwelling  
617 activity during the DA and the MWP. The high organic matter stocks in the Tagus prodelta  
618 during this phase created more refractory organic matter, responsible for a faunal diversity  
619 decrease from 2250 cal yr BP to present.

620         Phase III (1250 cal yr BP-present) showed the disappearance of *B. spathulata*, *P.*  
621 *mediterraneensis*, and *H. balthica* and the decrease of *A. beccarii* and *N. scaphum*, which may

622 be a consequence of a decrease in the organic matter quality and/or the occurrence of  
623 anthropogenic pollution, as indicated by benthic foraminiferal assemblages and the XRF data.  
624 Conversely, *E. scaber* increased, probably because of its pollution tolerance, and *C. carinata*  
625 remained dominant, likely due to the strong upwelling conditions. Benthic foraminifera  
626 responded accurately to record environmental and climate changes in the North Atlantic  
627 continental shelf and therefore could be used as a bio-indicator of environmental changes, such  
628 as changes in upwelling activity and river discharge, that were directly linked with the NAO.

629

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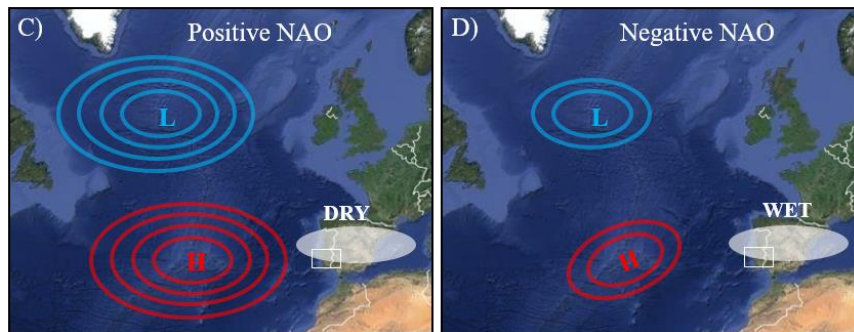
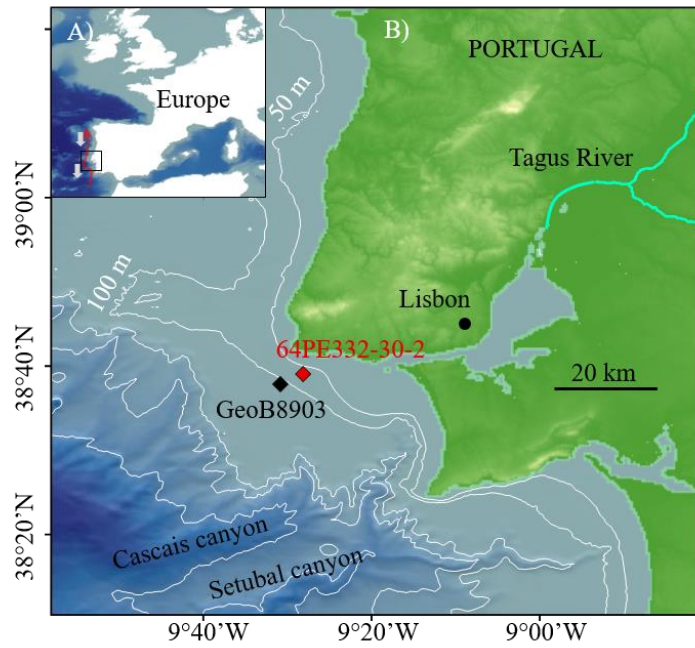
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902

903 **Figures**

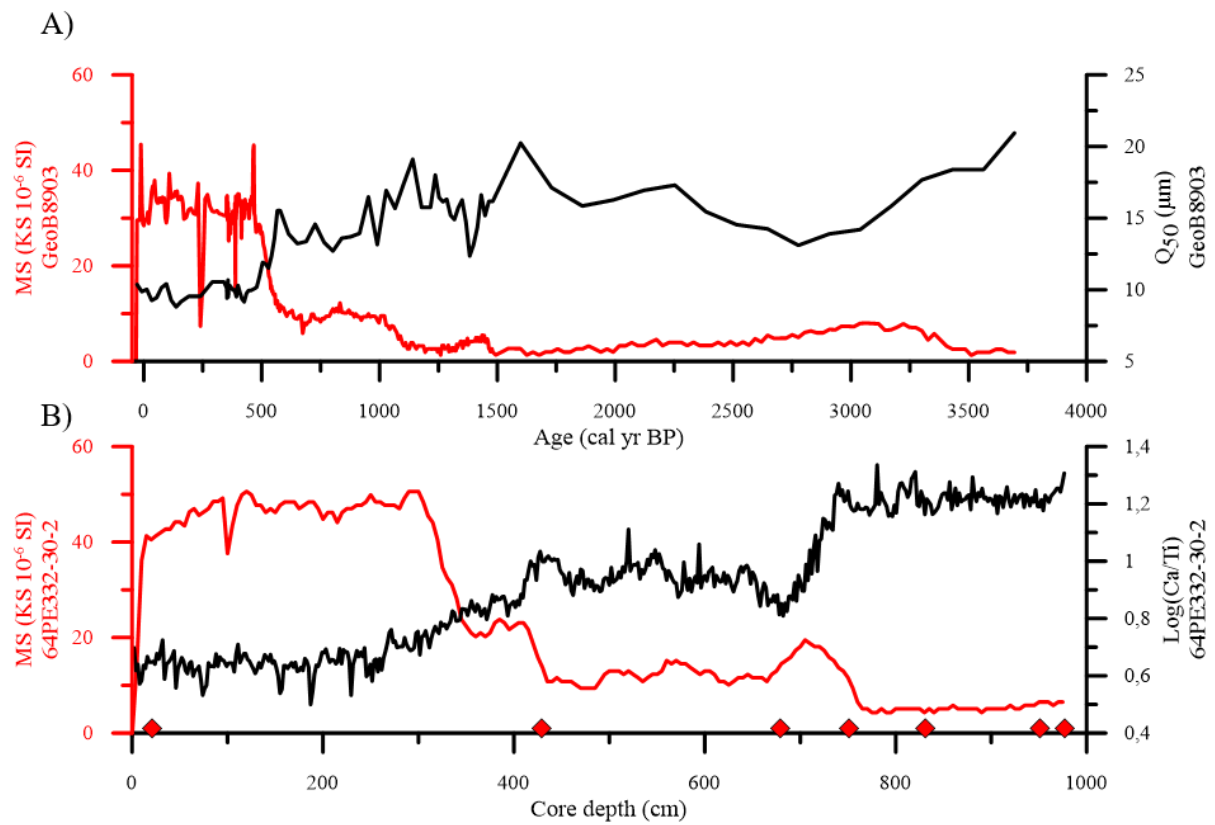


904

905 Fig. 1. A) General location map of the study area, with a schematic representation of the  
 906 Portugal Coastal Counter Current (red arrow) and the ENACW (grey arrows). B) Detailed map  
 907 of the core locations considered in this study. Red diamond: piston cores 64PE332-30-2, this  
 908 study; black diamond: GeoB8903 (Abrantes et al., 2008). C) Positive and D) Negative phase of  
 909 NAO with L = Iceland low-pressure system and H = Azores high-pressure system.

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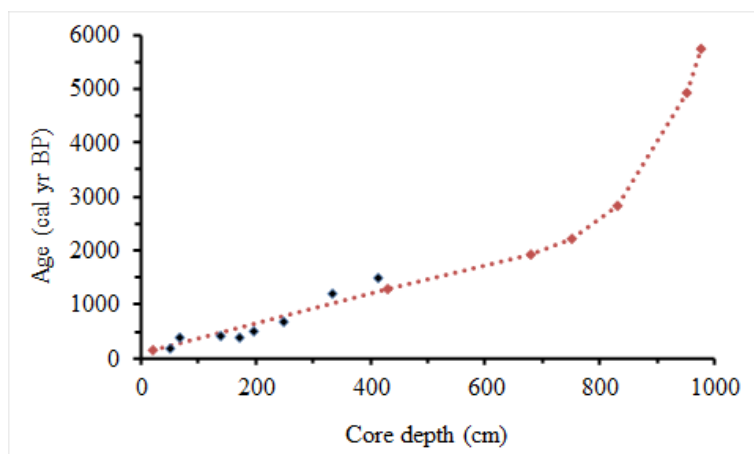


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913 Fig. 2. Comparison of Ca/Ti and magnetic susceptibility (MS) of Core 64PE332-30-2 with grain  
 914 size (Q<sub>50</sub>) and MS from Core GeoB 8903 (Abrantes et al., 2008; Alt-Epping et al., 2009). Red  
 915 diamonds represent <sup>14</sup>C data points.

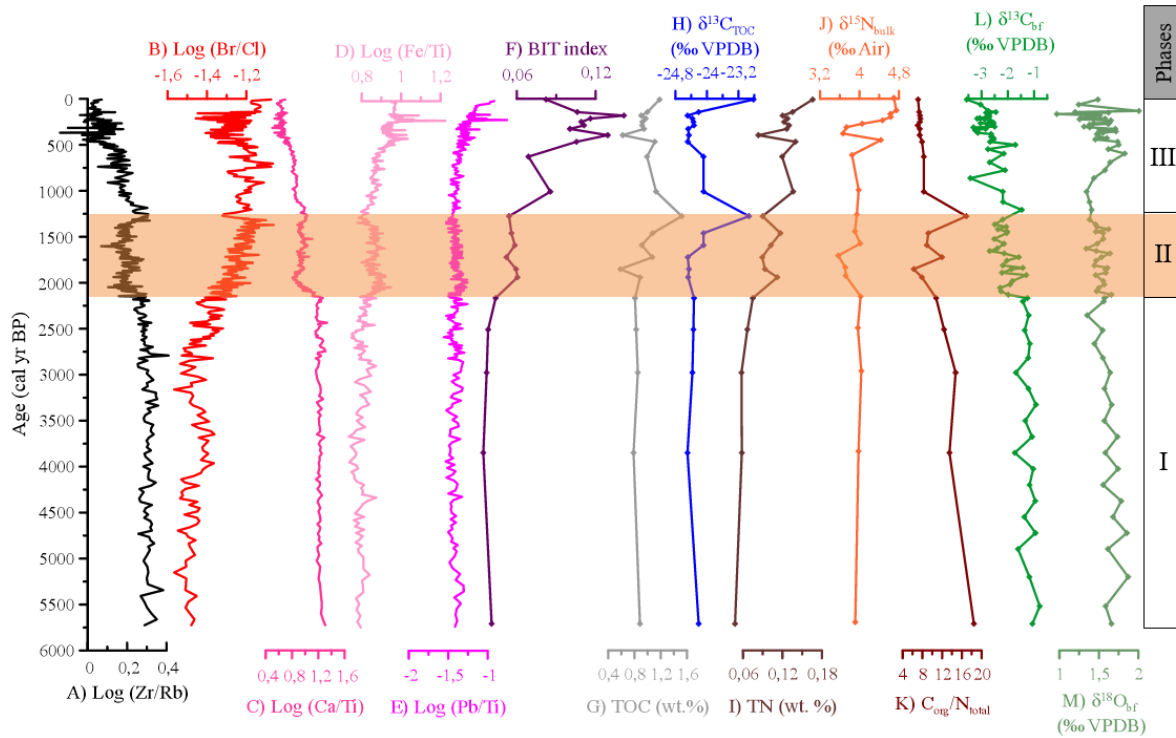
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918 Fig. 3. Age model of Core 64PE332-30-2 based on AMS <sup>14</sup>C data (Table 1). Blue diamonds  
 919 represent the <sup>14</sup>C data points of GeoB 8903.

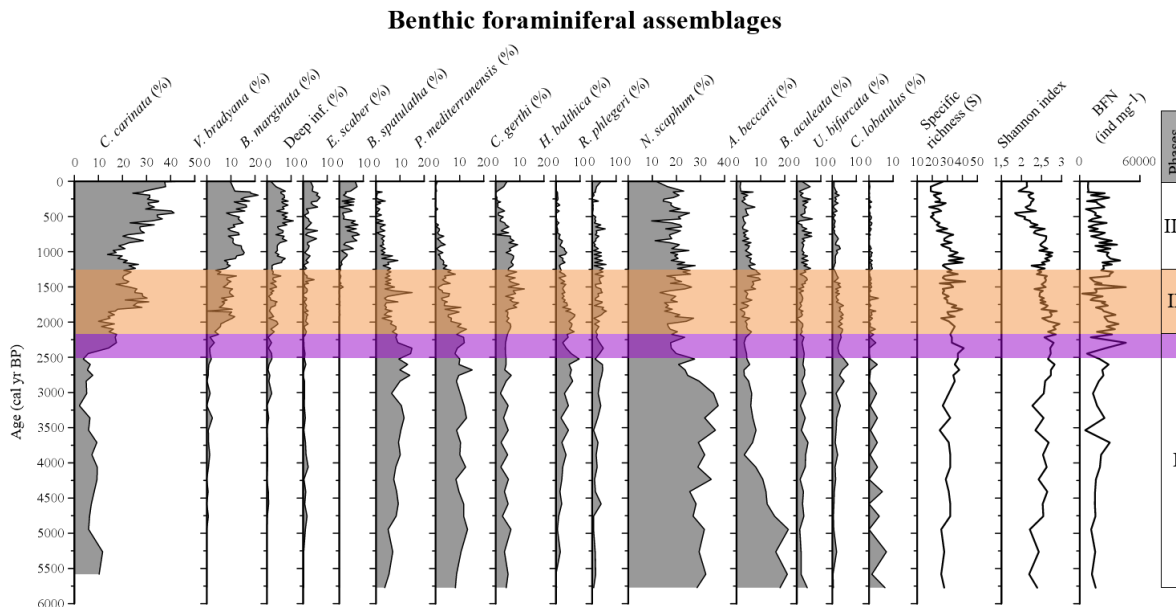
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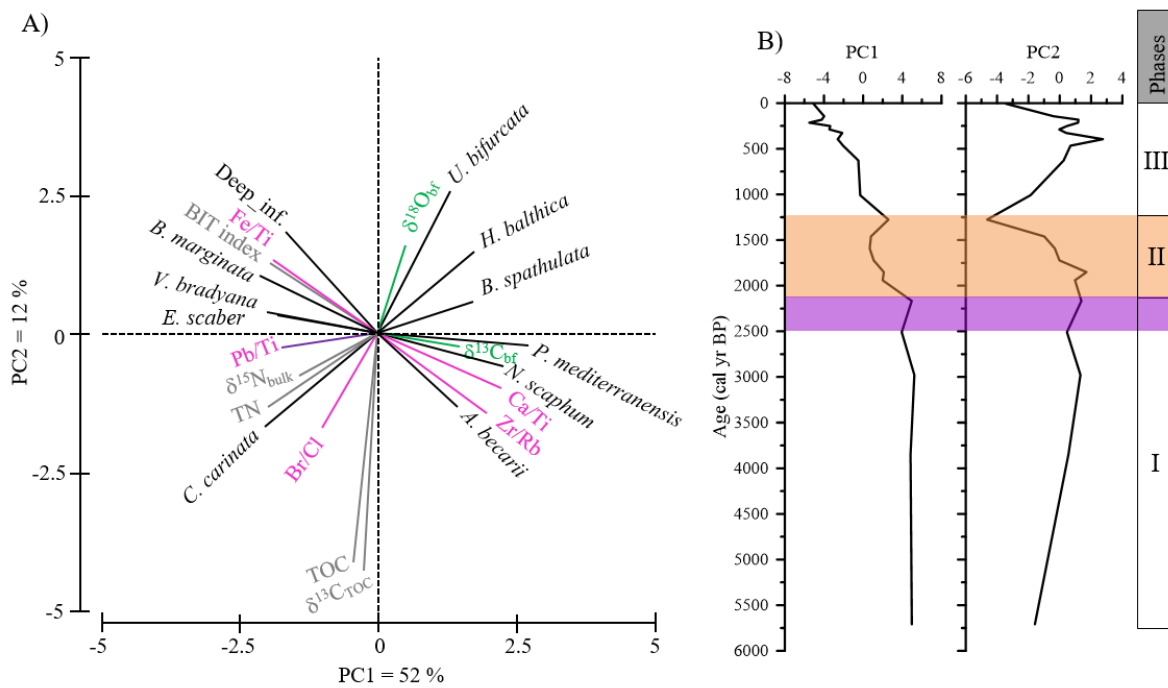
922 Fig. 4. Distribution of environmental parameters for the last 5750 years. XRF data: A) Zr/Rb,  
 923 B) Br/Cl, C) Ca/Ti, D) Fe/Ti and E) Pb/Ti. Organic measures: F) TN, G) TOC, H) C/N ratio, I)  
 924  $\delta^{15}\text{N}_{\text{bulk}}$ , J)  $\delta^{13}\text{C}_{\text{TOC}}$ , and K) BIT index. Carbonate isotopes L)  $\delta^{13}\text{C}_{\text{bf}}$  and M)  $\delta^{18}\text{O}_{\text{bf}}$  analyzed  
 925 for the species *N. scaphum*. The orange square indicates the limits of the three phases.

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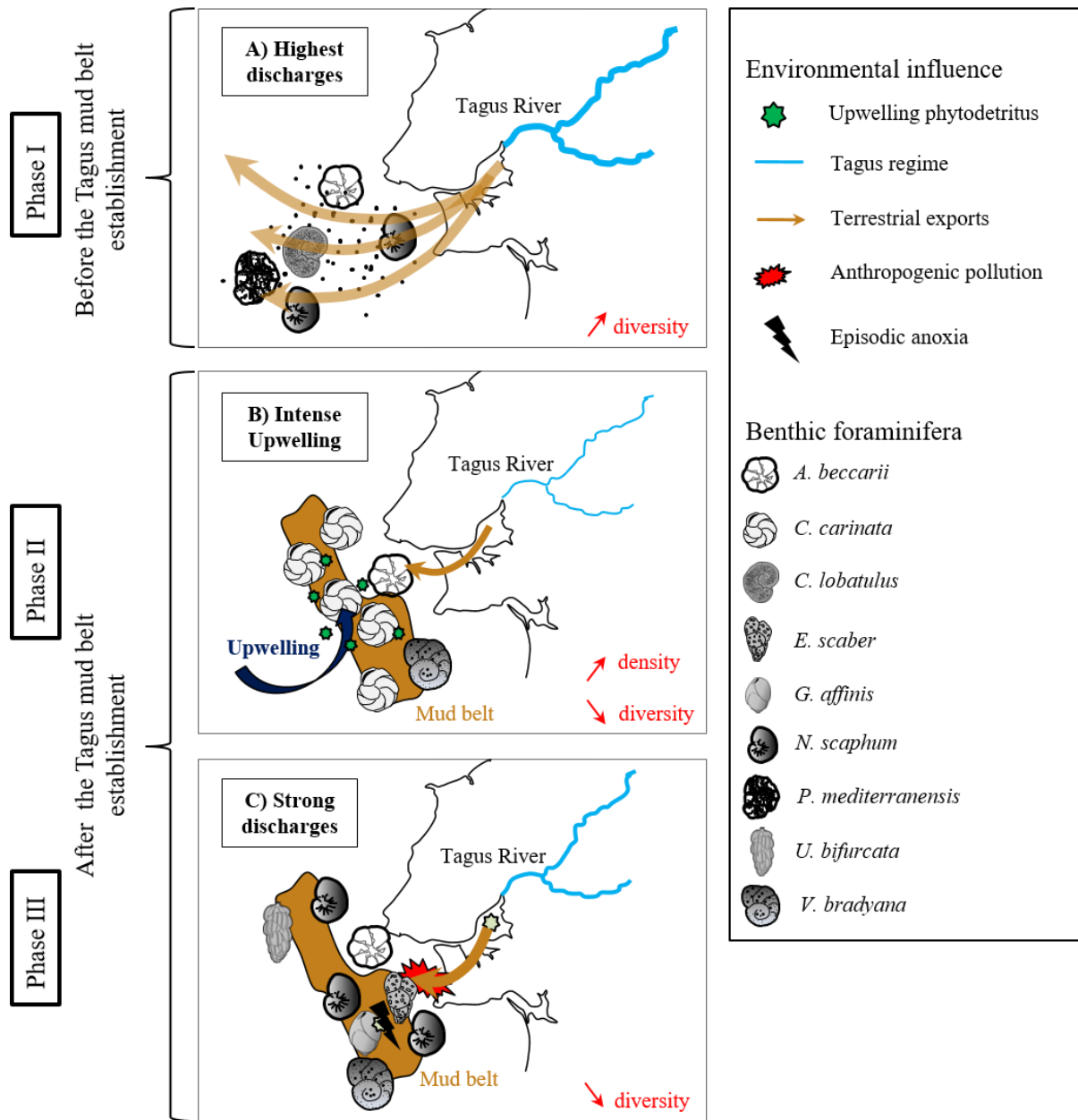
927

928 Fig. 5. Distribution of major species of benthic foraminifera (>5%) and specific richness from  
 929 the 101 samples of Core 64PE332-30-2 for the last 5750 years. Deep inf. = *G. affinis* + *C.*  
 930 *oolina*. The orange square indicates the limits of the three phases, and the purple square follows  
 931 the main change of the faunal distribution.



933

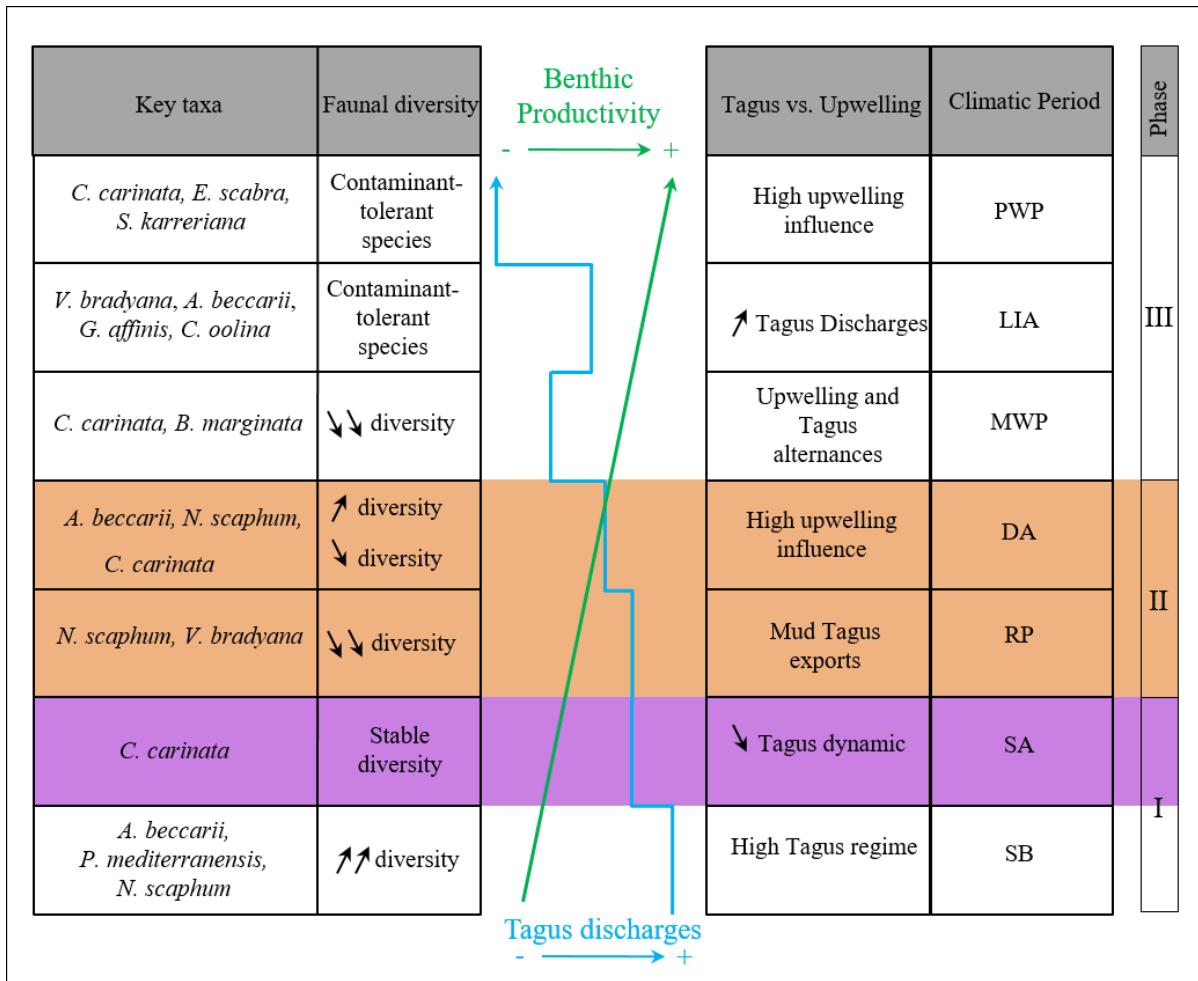
934 Fig. 6.A) PCA based on the major species percentages (in black), XRF data (in pink), organic  
 935 compounds (in grey), and benthic foraminiferal isotopes (in green). B) Score profiles of PC1  
 936 and PC2 as a function of time. The orange square indicates the limits of the three phases, and  
 937 the purple square follows the main change of the faunal distribution.



938

939 Fig. 7. Schematic representation of the three modes associated with the Tagus mud belt  
 940 establishment, Tagus River regime, and upwelling intensity. A) First phase, characterized by  
 941 very high Tagus River discharge. B) Second phase: mud belt establishment and very intense  
 942 upwelling activity. C) Alternating periods of strong upwelling activity and high Tagus River  
 943 discharge.

944



945

946 Fig. 8. Summary of the paleoenvironmental evolution of the Tagus River during the last 5750  
 947 years. RP = Roman Period, DA = Dark Ages, MWP = Medieval Warm Period, LIA = Little Ice  
 948 Age, PWP = Present Warm Period, SA = Sub-Atlantic Period, SB = Subboreal Period. The  
 949 orange square indicates the limits of the three phases, and the purple square follows the main  
 950 change of the faunal distribution.

951

952

953 Table 1. <sup>14</sup>C AMS dates of cores GeoB 8903 and 64PE332-30-2.

Sediment core	Lab no.	Core depth interval [cm]	Mean depth in core [cm]	Uncorrected AMS <sup>14</sup> C ages [cal yr BP]	Analytical error ( $\pm 1\sigma$ ) [yrs]	Ages ( $\Delta R = 0$ yr)( $\pm 2\sigma$ ) [cal yr BP]	Ages [cal yr BP]	Analyzed material	Reference
GeoB 8903	KIA30888	52-53	51	210	35	138-223	182	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	KIA30890	65-70	69	335	55	300-501	394	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	139-141	140	360	25	349-456	425	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	171-173	172	285	30	314-408	381	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	198-199	198	360	45	423-498	487	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	248-249	248	735	30	657-726	679	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	333-334	333	1260	35	1121-1282	1210	Foraminifera	Alt-Epping et al. (2009)
GeoB 8903	-	413-414	413	1600	40	1390-1567	1478	Foraminifera	Alt-Epping et al. (2009)
64PE332-30-2	BETA 348791	20-22	21	500	30	41 - 235	138	Gastropod	Warden et al. (2016)
64PE332-30-2	BETA 348792	428-430	429	1730	30	1219 - 1350	1284.5	Foraminifera	Warden et al. (2016)
64PE332-30-2	BETA 348793	678-680	679	2320	30	1848 - 2033	1940.5	Gastropod	Warden et al. (2016)
64PE332-30-2	VERA-51394	750-752	751	2530	70	2122-2293	2207.5	Gastropod	This study
64PE332-30-2	VERA-51395	830-832	831	3060	70	2749-2902	2825.5	Gastropod	This study
64PE332-30-2	VERA-51396	950-952	951	4690	70	4831-5007	4919	Gastropod	This study
64PE332-30-2	BETA 317911	976-978	977	5370	30	5644 - 5850	5747	Shell fragments	Warden et al. (2016)

954 Note that <sup>14</sup>C data from GeoB 8903 were from Alt-Epping et al. (2009) and reconverted using the CALIB V0.6 with the Marine13 calibration  
 955 curve (Stuiver and Reimer, 1993).