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A vertically-Lagrangian, non-hydrostatic, multilayer model for multiscale free-surface flows

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Abstract

This article presents a semi-discrete, multilayer set of equations describing the three-dimensional motion of an incompressible fluid bounded below by topography and above by a moving free-surface. This system is a consistent discretisation of the incompressible Euler equations, valid without assumptions on the slopes of the interfaces. Expressed as a set of conservation laws for each layer, the formulation has a clear physical interpretation and makes a seamless link between the hydrostatic Saint-Venant equations, dispersive Boussinesq-style models and the incompressible Euler equations. The associated numerical scheme, based on an approximate vertical projection and multigrid-accelerated column relaxations, provides accurate and efficient solutions for all regimes. The same model can thus be applied to study metre-scale waves, even beyond breaking, with results closely matching those obtained using small-scale Euler/Navier–Stokes models, and coastal or global scale dispersive waves, with an accuracy and efficiency comparable to extended Boussinesq wave models. The implementation is adaptive, parallel and open source as part of the Basilisk framework and the documented source codes sufficient to reproduce all results and figures are provided.

Keywords: 3D incompressible Euler; free-surface; approximate projection; adaptive; wave breaking; Boussinesq

1 Introduction

Modelling material motion as the interaction between layers of (possibly different) materials is common in solid mechanics, fluid mechanics and geophysical fluid dynamics. This concept is useful when the system is anisotropic, as layers allow both a simple description of the geometry and motion of interfaces, and simplification of the terms which become negligible due to anisotropy.

A typical example of this approach are the “shallow-water equations” which describe the motion of a single layer bounded vertically by two material interfaces: the bathymetry and free-surface. They were first derived by Saint-Venant using conservation of mass and momentum in vertical slices and the assumption that the vertical pressure profile is essentially hydrostatic. An important limitation of the Saint-Venant equations is that they are strictly consistent only for a vertically-constant velocity profile (a “plug flow”). Lifting this limitation requires the ability to describe the vertical velocity profile, which can be done by considering multiple layers.

Since these additional layers are not generally bounded by material interfaces, their vertical motion can be chosen freely and does not have to follow vertically-Lagrangian dynamics. This vertically-Eulerian point of view is dominant amongst the shallow-water models used to describe oceanic or atmospheric dynamics. It has however the important drawback of introducing vertical advection terms which break the horizontal/vertical decoupling due to anisotropy and thus complicate the equations, their physical interpretation and numerical approximation.

The hydrostatic assumption is another obvious limitation of Saint-Venant systems, which, as a consequence, cannot represent the dispersion of short gravity waves. If this limitation is lifted, multilayer systems could in principle become consistent with layer integration of the (incompressible) Euler equations with a free-surface and bathymetry. The resulting physical, mathematical and numerical model would then be valid without assumptions on the anisotropy of the flow, while preserving the ability to remove terms (and the corresponding numerical code) in anisotropic cases. It would thus provide a seamless link between the (single layer) Saint-Venant equations and a consistent discretisation of the incompressible free-surface Euler equations. This would be of considerable practical interest generally, but in particular for geophysical flows, since the resulting model hierarchy would be applicable and efficient for scales ranging from metres to the global scale.

The goal of this article is to present such a model. We first give its formal description in Section 2 and relate it to existing work in Section 2.1. The associated numerical scheme is described in Section 3, split into hydrostatic and non-hydrostatic parts, with a particular emphasis on accurate dispersion relations (Section 3.5). Section 4 discusses several test cases and applications, for a broad range of spatial scales and with measures of accuracy and performance.

The source code necessary to reproduce all results and figures is freely accessible (GPL license) as part of the Basilisk framework [1].

2 Model equations

We consider n layers of an incompressible (inviscid) fluid bounded vertically by a fixed bottom topography and a free surface (Figure 1). The properties of each layer are functions of time and two-dimensional space only. The resulting model is thus *semi-discrete* i.e. horizontally continuous and two-dimensional, and vertically discrete.

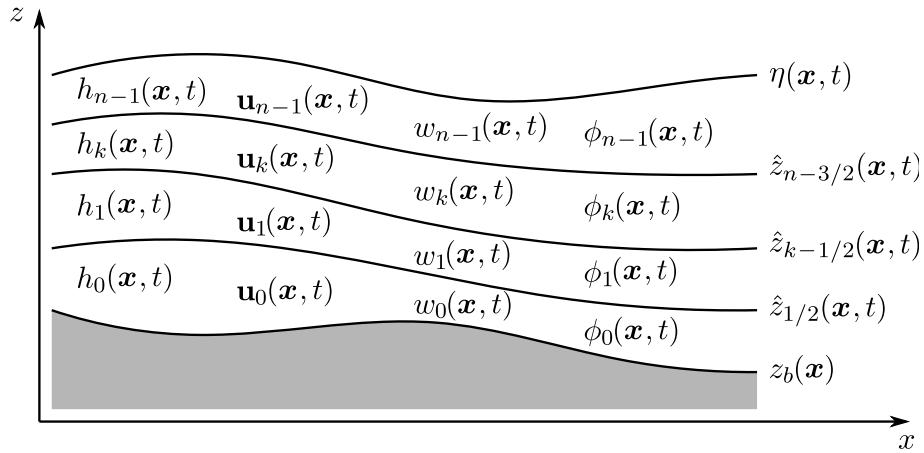


Figure 1. Definition of the n layers. h_k is the layer thickness, $(\mathbf{u}, w)_k$ the velocity vector in the \mathbf{x} - z reference frame, ϕ_k the non-hydrostatic pressure, $\hat{z}_{k-1/2}$ the height of the layer interface and η the free-surface height. Bold letters are used for vector fields.

We propose the following semi-discrete approximation of the incompressible Euler equations with a free-surface and gravity.

$$\partial_t h_k + \nabla \cdot (h \mathbf{u})_k = 0, \quad (1)$$

$$\partial_t (h \mathbf{u})_k + \nabla \cdot (h \mathbf{u} \mathbf{u})_k = -g h_k \nabla \eta - \nabla (h \phi)_k + [\phi \nabla \hat{z}]_k, \quad (2)$$

$$\partial_t (h w)_k + \nabla \cdot (h w \mathbf{u})_k = -[\phi]_k, \quad (3)$$

$$\nabla \cdot (h \mathbf{u})_k + [w - \mathbf{u} \cdot \nabla \hat{z}]_k = 0, \quad (4)$$

with k the index of the layer, h_k its thickness, \mathbf{u}_k, w_k the horizontal and vertical components of the velocity, g the acceleration of gravity, ϕ_k the non-hydrostatic pressure (divided by the fluid density),

$$\eta \equiv z_b + \sum_k h_k, \quad (5)$$

the free-surface height, with z_b the height of the bottom topography,

$$\hat{z}_{k+1/2} \equiv z_b + \sum_{l=0}^k h_l, \quad (6)$$

the height of each layer interface (see Figure 1) and

$$[f]_k \equiv f_{k+1/2} - f_{k-1/2}, \quad (7)$$

the “vertical difference” operator. Note that since all fields are two-dimensional, so is the ∇ operator.

Equation (1) is the layer thickness evolution equation, equations (2) and (3) express conservation of the horizontal and vertical momentum, and (4) is the semi-discrete continuity equation (or semi-discrete incompressibility condition) which expresses volume (and thus mass) conservation.

The system must then be closed by assuming relations between the layer-averaged values ϕ_k, w_k and the layer-interface values $\phi_{k+1/2}, w_{k+1/2}$. We will outline a strategy to optimise this choice in section 3.5.

Note that the formal derivation of this system is not entirely trivial, in particular for the “slope terms” $\nabla \hat{z}$ in (2) and (4). We give further details on their derivation in Appendix A.

Remark 1. One may feel that equations (1) and (4) are redundant (since they appear to both express volume conservation). If this was the case, the system would not be closed since there are four unknowns per layer (h, \mathbf{u}, w and ϕ). Note however that there is no vertical advection term (of the form $\partial_z(h \mathbf{u} w)_k$) in either equations (2) or (3). This is of course because we have assumed that the layer interfaces (as defined by $\hat{z}_{k+1/2}$) are material surfaces which follow the vertical motion of the fluid i.e. the discretisation is Lagrangian vertically. This implies a kinematic condition, relating \hat{z} and w , which must be added to (2) and (3). This condition can be written

$$w = \partial_t \hat{z} + \mathbf{u} \cdot \nabla \hat{z} \quad (8)$$

Using the vertical difference operator, the fact that $h_k = [\hat{z}]_k$ (from the definition (6) of \hat{z}_k) and (4) we get

$$\begin{aligned} [w]_k &= \partial_t [\hat{z}]_k + [\mathbf{u} \cdot \nabla \hat{z}]_k, \\ [w - \mathbf{u} \cdot \nabla \hat{z}]_k &= \partial_t h_k, \\ -\nabla \cdot (h \mathbf{u})_k &= \partial_t h_k, \end{aligned}$$

which is the layer evolution equation (1). Equation (1) is then equivalent to condition (8) and comes from the particular choice of a Lagrangian vertical coordinate system. It does not represent a specific physical property of the system (with the exception of the bottom and top boundary conditions). Equation (4) on the other hand is true irrespective of the choice of vertical coordinate and is necessary and sufficient to enforce volume/mass conservation. This point is further discussed in relation to previous work in Section 2.1.

Remark 2. While the evolution equations for h , \mathbf{u} and w are obvious, the evolution equation for ϕ is not. This is classical for incompressible fluids for which the pressure becomes a gauge field used to project the velocity onto the space of divergence-free solutions [2, 3]. Equation (4) is the equivalent semi-discrete “divergence-free” condition which will be used to build the Poisson-like equation which ϕ must verify.

Remark 3. Setting $\phi = 0$ and dropping the requirement of conservation of the vertical momentum (equation (3)) gives

$$\partial_t h_k + \nabla \cdot (h \mathbf{u})_k = 0, \quad (9)$$

$$\partial_t (h \mathbf{u})_k + \nabla \cdot (h \mathbf{u} \mathbf{u})_k = -g h_k \nabla \eta, \quad (10)$$

$$\nabla \cdot (h \mathbf{u})_k + [w - \mathbf{u} \cdot \nabla \hat{z}]_k = 0, \quad (11)$$

which are the standard hydrostatic, multilayer Saint-Venant (or stacked shallow-water) equations. The vertical velocity w becomes a diagnostic variable which can be obtained from \mathbf{u} and h using equation (11). Note that (8) could be used instead, however the resulting value of w would then not necessarily verify the discrete continuity equation (11).

Remark 4. The system of equations (1), (2), (3) and (4) is consistent even for a single layer for which it can be re-written

$$\partial_t h + \nabla \cdot (h \mathbf{u}) = 0,$$

$$\partial_t (h \mathbf{u}) + \nabla \cdot (h \mathbf{u} \mathbf{u}) = -g h \nabla \eta - \nabla (h \phi) - \phi_{-1/2} \nabla z_b, \quad (12)$$

$$\partial_t (h w) + \nabla \cdot (h w \mathbf{u}) = \phi_{-1/2}, \quad (13)$$

$$\nabla \cdot (h \mathbf{u}) + w_{1/2} - \mathbf{u}_{1/2} \cdot \nabla \eta = 0,$$

where we have used the free-surface boundary condition $\phi_{1/2} = 0$ and bottom boundary impermeability condition $(w - \mathbf{u} \cdot \nabla \hat{z})_{-1/2} = 0$. The system is then closed by assuming functional relationships between the bottom pressure $\phi_{-1/2}$ and free surface velocity $(\mathbf{u}_{1/2}, w_{1/2})$ and the “layer-averaged” pressure ϕ and velocity (\mathbf{u}, w) ; the simplest of which could be $\phi_{-1/2} = \phi$, $(\mathbf{u}_{1/2}, w_{1/2}) = (\mathbf{u}, w)$. Better choices are discussed in section 3.5 in the general multilayer case. For a single layer, they reduce to $\phi_{-1/2} = \frac{8}{3} \phi$, $(\mathbf{u}_{1/2}, w_{1/2}) = (\mathbf{u}, w)$ for the staggered scheme of section 3.6.1, and $\phi_{-1/2} = 2 \phi$, $(\mathbf{u}_{1/2}, w_{1/2}) = (\mathbf{u}, 2w)$ for the Keller box scheme of section 3.6.2.

2.1 Related work

2.1.1 Hydrostatic system

The single layer limit of the hydrostatic system (9), (10) and (11) is the classical Saint-Venant system [4]. Its multilayer generalisation is also well-known and is discussed in many textbooks on geophysical fluid dynamics, in particular in connection with the theoretical study of internal wave dynamics, using the two-layer version [5, 6, 7]. From a modelling perspective, this system is also the basis for isopycnal hydrostatic discretisations of the equations of oceanic motion [8] or isentropic discretisations of the atmospheric equations [9, 10].

Another classical variant of the hydrostatic system (9), (10) and (11) is the corresponding general “quasi Eulerian” vertical discretisation for which the layer positions $\hat{z}_{k+1/2}$ are diagnosed from the other variables rather than computed as part of the solution (using equation (9)). In this case, the vertical advection fluxes must be added to equation (10) and the layer evolution equation (9) is replaced by the evolution equation for the total depth. Classical examples of this type of discretisation are the z - and σ -coordinates models typically used for ocean modelling [11]: for z -coordinates the layer positions are constant i.e. the discretisation is strictly Eulerian vertically, while for σ -coordinates the layer positions are proportional to the total depth and thus vary in both space and time (i.e. quasi-Eulerian).

Interesting overviews of the pros and cons of the purely Lagrangian (i.e. isopycnal) vertical discretisation versus Eulerian or quasi-Eulerian discretisations (z - or σ -coordinates) are given by Griffies et al. (2000) [11], Adcroft & Hallberg (2006) [12] and Chassignet et al. (2011) [13]. Either versions of the multilayer hydrostatic equations have also been studied fairly recently by applied mathematicians in a different context, more closely linked to hydraulic engineering, seeking to clarify the consistency of shock-like solutions (hydraulic jumps) in the presence of viscous dissipation [14, 15, 16].

2.1.2 Non-hydrostatic system

The generalisation of the layered equations to include non-hydrostatic effects is more recent. In ocean modelling this effort has been driven in particular by the increase in computing power which allows to resolve spatial scales for which non-hydrostatic effects are not negligible or even dominant. This “convergence of scales” also means that the approaches developed in scientific communities which were formerly interested in different scales are now also converging.

For example coastal engineers have long been interested in describing oceanic motion (mostly waves) at scales of order of metres. This motivated some of the earliest developments in the non-hydrostatic description of the motion of fluid layers, starting from Boussinesq [17] and much later, Serre (for hydraulic engineering in dams rather than the ocean) [18], Peregrine [19] and Green and Naghdi [20]. Since then these “extended-Boussinesq” wave models have known major developments which are summarised in the interesting review by Brocchini (2013) [21] (see also [22]).

The equations typically used in these models are closely related to the hydrostatic Saint-Venant system, often with a single layer (see Lynett & Liu (2004) [23] or Chazel et al. (2009) [24] for multilayer versions), but without the additional vertical momentum or continuity equations (12) and (13). Rather, the effects of the non-hydrostatic pressure are modelled by including source terms on the right-hand-side of the mass and/or momentum equations (9) and (10). These terms are typically obtained through asymptotic expansion/truncation of the Euler velocity potential. While these equations are thus structurally different from the non-hydrostatic single or multi-layer equations above, a common goal is the accurate description of short waves and in particular their dispersion relation as well as non-linear behaviour.

In parallel with these developments, interest in small-scale hydrodynamics with free surfaces and interfaces also led to extensive work on suitable numerical methods. Key developments in this area are the invention of the staggered (MAC or C-grid) discretisation and associated projection method for incompressible flows. Interestingly these were discovered independently in different communities by Harlow and Welch (1965) [25] (small-scale free-surface incompressible flows) and Arakawa (1966) [26] (large-scale geophysical flows). This gap between communities was only closed relatively recently, for example with the work of Marshall and Adcroft to include non-hydrostatic terms in the MITgcm ocean model, using the MAC projection scheme within a z -coordinate ocean model [27].

Also in the late 1990s, coastal engineers started adapting these small-scale non-hydrostatic numerical methods to specific aspects of their field. Representative examples of this trend are the work of Casulli [28] and Stelling & Zijlema [29]. Both of these schemes still use a z -coordinate in the vertical so do not differ fundamentally from standard MAC/C-grid incompressible solvers (with a free-surface). Later work by Zijlema & Stelling, however, generalises the numerical scheme to a σ vertical coordinate [30]. The resulting system of equations is thus related to the semi-discrete system (1), (2), (3) and (4), with the fundamental difference of a quasi-Eulerian treatment of the vertical coordinate.

Larger-scale ocean modellers have also been interested in including non-hydrostatic effects, in particular in the σ -coordinate models often used for “regional” ocean studies [31]. These schemes are generally based on the approaches developed by Casulli, Stelling and Zijlema and the resulting set of equations do not differ fundamentally.

Non-hydrostatic, vertically-Lagrangian descriptions of the equations of motion are thus rare, which may be partly due to the confusion between the layer evolution equation and the continuity equation mentioned in Remark 1 above (see in particular [12] and [32] for a clarification). A notable exception is the article by Vitousek & Fringer (2014) [33] which presents a mild slope approximation of the non-hydrostatic equations expressed in isopycnal coordinates. The authors derive an interesting scaling analysis to justify the validity of the mild slope assumption for the oceanic scales they are interested in. We will show later that these terms cannot be neglected for typical coastal-engineering-scale processes, even when the slopes are relatively mild. We will also show how these terms can be retained while keeping the structure and numerical efficiency of the simplified system.

In parallel with these developments, the search for efficient short-wave models applicable to coastal engineering led to the idea that these multilayer non-hydrostatic models, used with as few layers as possible (ideally a single one), could be an efficient alternative to Boussinesq-type models. Zijlema and Stelling showed in particular (somewhat empirically) that a particular version of their numerical scheme had good dispersion properties even for (very) short waves using only two layers [29, 34]. This led Yamazaki et al. (2008) [35] to propose a single-layer model similar to the single-layer model in Remark 4 above, with the notable difference of the absence of the convective term in the vertical momentum equation (12) and neglecting all metric/slope terms (essentially a linearised version). This single-layer model was also re-derived more recently by Bristeau et al. (2015) [36] who pointed out that the inclusion of the convective term is necessary for energy consistency.

2.2 Properties

The semi-discrete system (1), (2), (3) and (4) has a number of favourable features. It only relies on layer-averaging of the incompressible Euler equations, without any other approximation. In particular it is consistent irrespective of the slope of the layers. The vertically-Lagrangian discretisation eliminates any numerical vertical diffusion which is particularly important for long-term integration of ocean models [8]. When using flux-based finite-volume schemes for transport, horizontal momentum as well as tracers are exactly conserved in each layer. Finally, the main advantage of this system is that it provides a clear and natural hierarchy of models, of increasing accuracy (and complexity): single layer Saint-Venant, multilayer (hydrostatic) Saint-Venant, (non-hydrostatic) multilayer free-surface Euler (and Navier–Stokes).

The drawbacks are few. The most obvious one is the addition of the layer thicknesses as prognostic variables. The added complexity and computational cost are low however, since they are passive tracers which will be transported using the same scheme as other tracers.

The multilayer Saint-Venant system is known to be non-hyperbolic [14] with possibly non-unique solutions for the layer thicknesses. This is not surprising and reflects the lack of physical coupling between layers. This is cured in the non-hydrostatic case, due to the coupling between layers induced by the elliptic non-hydrostatic pressure, and/or when using vertical remapping (see section 3.3), due to (numerical) viscous coupling between layers.

3 Numerical scheme

The numerical scheme is designed in a modular fashion, to match the increasing complexity and accuracy of the successive theoretical approximations. The goal is a numerical tool where mathematical terms (and the corresponding code), can be added one-by-one depending on the balance required between accuracy, complexity and numerical efficiency.

3.1 Hydrostatic scheme

We first build a numerical scheme to approximate the multi-layer, hydrostatic system (9), (10) and (11). This system is an hyperbolic system of conservation laws and is often solved based on characteristic (approximate) Riemann solutions. This is important in particular when considering the consistency of shock-like solutions [37, 38]. The resulting numerical schemes are classical but reasonably complex and expensive. They also have the drawback that special measures must be taken to guarantee that the trivial “lake-at-rest” solution is recovered when bathymetry is included [37]. We follow a simpler approach, often taken when building ocean models, which does not explicitly take into account the existence of shock-like solutions.

The time integration of (9), (10) and (11) is split into a first step where all the conserved quantities s_k are advected according to

$$\begin{aligned} \frac{h_{i,k}^{n+1} - h_{i,k}^n}{\Delta t} &= \frac{F_{i-1/2,k}^{n+1/2} - F_{i+1/2,k}^{n+1/2}}{\Delta}, \\ \frac{(h s)_{i,k}^{n+1} - (h s)_{i,k}^n}{\Delta t} &= \frac{(s F)_{i-1/2,k}^{n+1/2} - (s F)_{i+1/2,k}^{n+1/2}}{\Delta}, \end{aligned}$$

where $s_{i,k}$ is either the horizontal velocity, vertical velocity or any additional tracer and $F_{i+1/2,k}$ is the thickness flux $(h u)_{i+1/2,k}$, evaluated using the (Lax-Wendroff-like) Bell–Collela–Glaz (BCG) third-order-upwind advection scheme [39], with a staggered velocity $u_{i+1/2,k}$ to be defined later. Space is discretised in intervals $x_{i+1/2} - x_{i-1/2} = \Delta$. For simplicity, all the conserved variables are collocated and treated as cell-averages over $[x_{i-1/2}, x_{i+1/2}]$. Note that this step is standard and could be performed using a broad range of advection schemes. Desirable properties include: exact conservation (which is guaranteed for flux-based finite-volume schemes as here), monotonicity/positivity which is necessary for consistency in particular when “drying” of the layers occur (i.e. $h_k \rightarrow 0$) (this is guaranteed here by upwinding and min-mod limiting).

In a second step, the pressure gradient term (r.h.s. of (10)) is added to the horizontal momentum, with some care. The free-surface position is first calculated as

$$\eta_i^{n+1} = z_{i,b} + H_i^{n+1} \quad \text{with} \quad H_i^{n+1} = \sum_k h_{i,k}^{n+1}$$

The horizontal staggered acceleration is then obtained as

$$a_{i+1/2,k} = -g \frac{\eta_{i+1,k} - \eta_{i,k}}{\Delta} \quad (14)$$

The staggered velocity/flux field is then computed as

$$(h u)_{i+1/2,k}^{n+1} = \frac{(h u)_{i+1,k}^{n+1} + (h u)_{i,k}^{n+1}}{2} + \Delta t (h a)_{i+1/2,k} \quad (15)$$

where $h_{i+1/2}^{n+1}$ is computed using BCG face reconstruction. Note that the robustness of this step relies on the fact that u and h are collocated and transported using the same advection scheme. Finally the pressure gradient term is added to velocity as

$$u_{i,k}^{n+1} \leftarrow u_{i,k}^{n+1} + \Delta t \frac{(h a)_{i+1/2,k} + (h a)_{i-1/2,k}}{h_{i+1/2,k} + h_{i-1/2,k}} \quad (16)$$

Since time integration is fully explicit, we can expect that the classical stability condition

$$\Delta t < \frac{\Delta}{\sqrt{g H}} \quad (17)$$

applies. This scheme is simple, conservative and robust to wetting and drying. We will also show that it can recover shock-like solutions. It is of course only formally first-order accurate in time and is constrained by the maximum speed of gravity waves. Note however that it could be replaced by more sophisticated schemes, using for example implicit time integration [28] or barotropic/baroclinic mode-splitting [40], without affecting its articulation with the (viscous, non-hydrostatic etc.) extensions presented later.

3.2 Vertical viscosity

In the case of flows with large aspect ratios, diffusion is expected to be controlled essentially by its vertical component which can be added to the momentum equations (2) and (3) as

$$\begin{aligned} \partial_t (h \mathbf{u})_k + \nabla \cdot (h \mathbf{u} \mathbf{u})_k &= -g h_k \nabla \eta - \nabla (h \phi)_k + [\phi \nabla \hat{z}]_k + [\nu \partial_z \mathbf{u}]_k, \\ \partial_t (h w)_k + \nabla \cdot (h w \mathbf{u})_k &= -[\phi]_k + [\nu \partial_z w]_k, \end{aligned}$$

with ν the fluid kinematic viscosity.

Boundary conditions on the top and bottom layers need to be added to close the system for the viscous stresses. We chose to impose a Neumann condition on the free-surface i.e.

$$\partial_z u|_{t=\dot{u}_t}$$

and a Navier slip condition on the bottom i.e.

$$u|_b = u_b + \lambda_b \partial_z u|_b \quad (18)$$

The default boundary conditions are free-slip on the free-surface (i.e. $\dot{u}_t = 0$) and no-slip on the bottom (i.e. $u_b = \lambda_b = 0$).

Note that for practical coastal engineering or ocean modelling applications, vertical diffusion of (ensemble averaged) momentum usually includes subgrid-scale models of turbulent diffusion [41]. The bottom boundary conditions are also chosen to obtain bottom fluxes consistent with turbulence modelling assumptions. The formulation above is general enough to accommodate these turbulence models.

For stability, we discretise the viscous friction term implicitly as

$$\frac{(h u_k)^{n+1} - (h u_k)^*}{\Delta t} = \nu \left(\frac{u_{k+1} - u_k}{h_{k+1/2}} - \frac{u_k - u_{k-1}}{h_{k-1/2}} \right)^{n+1}$$

This can be expressed as the linear system

$$\mathbf{M} \mathbf{u}^{n+1} = \mathbf{rhs},$$

where \mathbf{M} is a tridiagonal matrix. A more detailed description of the general algorithm can be found in [16] and in the documented source code (<http://basilisk.fr/src/layered/diffusion.h>).

3.3 Vertical remapping

Lagrangian methods have the well-known drawback that as times passes discretisation elements become more and more distorted, up to a point where the results loose consistency. A solution is to periodically “remap” the solution onto a “clean” discretisation. The quality of this remapping then largely controls the overall quality of the method. In multiple dimensions this is not a simple operation, in particular when conservation of the remapped quantities is required [42]. In the one-dimensional case we are interested in, this is much simpler and can be done using for example Piecewise Polynomial Reconstruction [43]. In our implementation we have used the quadratic piecewise remapping of the PPR library of Engwirda and Kelley (2016, <https://github.com/dengwirda/PPR>).

The choice of the target remapped discretisation is entirely flexible and could be based on physical considerations (i.e. the wish to resolve certain features: e.g. boundary layers, material interfaces etc.) and/or approximation error analysis [44, 45]. For the moment, we have only applied a simple remapping to “ σ -coordinates” i.e. equal layer thicknesses proportional to the local water depth. This remapping is applied at every timestep just after the advection step of the hydrostatic algorithm.

Note that a subset of the possible remappings can be interpreted as fluxes of the layer-integrated quantities through the layer interfaces (see [45] for an example). The associated “remapping” velocity is the time derivative of the (desired) vertical displacement of the layers. This is in fact a discrete representation of the vertical advection term required when considering the equations in a moving (vertical) reference frame, and is known as an Arbitrary Lagrangian–Eulerian (ALE) (vertical) discretisation. We prefer to avoid this terminology since it is more restrictive than the remapping concept and has historically caused significant confusion (due to an erroneous interpretation of the physical meaning of the remapping velocity: it has none in general). See also [32] for a recent discussion.

As noted above, Lagrangian remapping in arbitrary dimensions is complex and the associated computational cost can be a concern. In the one-dimensional case we are interested in, remapping can be implemented using exactly the same algorithms as for one-dimensional vertical advection [45], and more sophisticated techniques (such as [43]) are also closely related to higher-order Eulerian advection schemes. The computational cost of one-dimensional remapping should thus be comparable to that of one-dimensional Eulerian advection.

3.4 Adaptivity and parallelism

The numerical schemes are implemented within the Basilisk framework [1] which provides transparent quadtree adaptivity and OpenMP/MPI parallelism for Cartesian numerical schemes. A detailed description is beyond the scope of this paper but we refer the interested reader to the web site as well as [46, 47].

3.5 Generic dispersion analysis

The semi-discrete formulation (1), (2), (3) and (4) is well-suited for a generic study of the dispersive properties of various choices of vertical discretisations. Using standard linear perturbation analysis, we look for wave-like solutions of the form

$$\begin{aligned} h_k &= \bar{h}_k + h'_k e^{i(\hat{k}x - \omega t)}, \\ u_k &= u'_k e^{i(\hat{k}x - \omega t)}, \\ w_k &= w'_k e^{i(\hat{k}x - \omega t)}, \\ \phi_k &= \phi'_k e^{i(\hat{k}x - \omega t)}, \end{aligned}$$

with ω the frequency, \hat{k} the wavenumber, \bar{h}_k the unperturbed layer thicknesses and h'_k , u'_k , w'_k and ϕ'_k vanishingly-small perturbations.

The vertical discretisation scheme essentially amounts to the choice of functional relationships between the layer interface values $\phi_{k+1/2}$, $w_{k+1/2}$, ... and the “layer-averaged” values ϕ_k , w_k , ... (as for the single layer model in Remark 4 above).

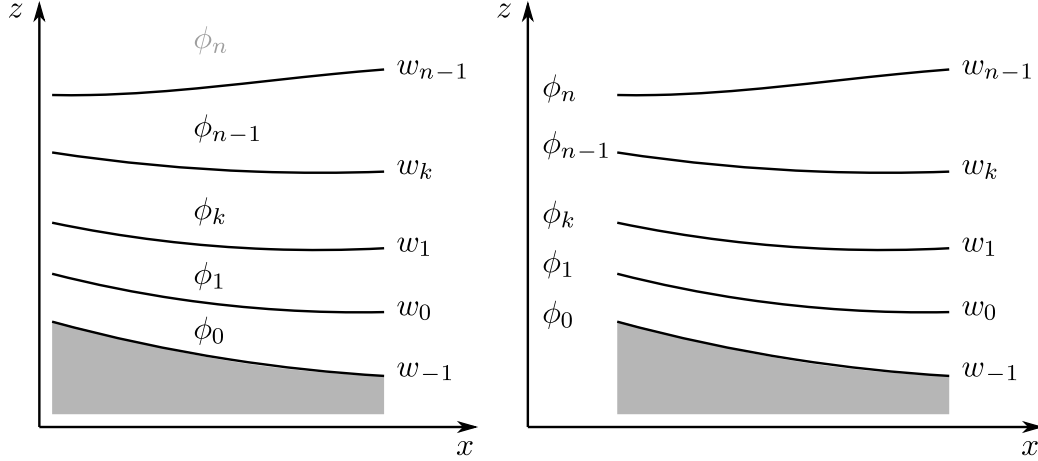


Figure 2. Conventions used for the indices of vertically staggered fields, for the generic dispersion analysis. (left) Staggered vertical discretisation. (right) Keller box scheme.

3.5.1 Staggered vertical discretisation

We first consider the simple choice

$$\begin{aligned} \phi_{k-1/2} &= \phi_k, \\ w_{k+1/2} &= w_k, \end{aligned}$$

which can be interpreted as a vertically-staggered discretisation of the velocity (i.e. pressure at the centre of the layer and vertical velocity on the layer interface, see Figure 2, left). The perturbations are then solutions of the linearised version of (1), (2), (3) and (4) (for a flat bathymetry)

$$-\omega h'_k + \bar{h}_k u'_k \hat{k} = 0, \quad (19)$$

$$-\bar{h}_k u'_k \omega = -g \hat{k} \bar{h}_k \sum h'_k - \bar{h}_k \hat{k} \phi'_k, \quad (20)$$

$$-\bar{h}_k w'_k \omega i = \phi'_k - \phi'_{k+1}, \quad (21)$$

$$\bar{h}_k u'_k \hat{k} i + w'_k - w'_{k-1} = 0 \quad (22)$$

The system must be closed with the (linearised) bottom impermeability condition

$$w'_{-1} = 0,$$

and a condition for the top pressure ϕ_n . With this choice of variable staggering, the top pressure is above the free surface and we must use an approximation of the free-surface boundary condition $\phi=0$, such as the third-order accurate interpolation (valid for identical layer depths)

$$\phi_n = \frac{\phi_{n-2} - 6\phi_{n-1}}{3} \quad (23)$$

The dispersion relation can then be obtained by considering non-trivial solutions of (19), (20), (21) and (22) i.e. solutions for which the determinant of the system vanishes. For a single layer, this gives

$$\omega_1^2(\hat{k}) = 8g \frac{h_0 \hat{k}^2}{3h_0^2 \hat{k}^2 + 8} \quad (24)$$

(where we have dropped the $-$ for convenience) and for two layers one finds

$$\omega_2^2(\hat{k}) = 8g \frac{h_0 h_1^2 \hat{k}^4 + (h_1 + h_0) \hat{k}^2}{3h_0^2 h_1^2 \hat{k}^4 + 3(3h_0^2 + h_0 h_1 + h_1^2) \hat{k}^2 + 8}$$

The calculation for an increasing number of layers becomes rapidly cumbersome but can be carried out using a symbolic algebra package. We give explicit formulae for a larger number of layers in Appendix C. The general form is the rational function

$$\omega_n^2(\hat{k}) = g \frac{\sum_{i=1}^n a_{2i}(h_k) \hat{k}^{2i}}{\sum_{i=0}^n b_{2i}(h_k) \hat{k}^{2i}},$$

with a_{2i} and b_{2i} polynomial functions of h_k .

The discrete dispersion relations above are an approximation of the exact linear dispersion relation for waves in arbitrary water depth

$$\omega^2(\hat{k}) = g \hat{k} \tanh(\hat{k} h)$$

It is clear that a larger number of layers should improve the approximation of the hyperbolic function since the degree of the rational function increases with the number of layers. The details of this convergence will matter in practice however, and will depend on the details of the discretisation. The analysis of this convergence is thus a good way of choosing between different potential schemes.

For example, the single-layer system approximates the hyperbolic function as (according to (24))

$$\tanh(x) \approx \frac{x}{\frac{3}{8}x^2 + 1}, \quad (25)$$

which can be compared with the Padé [1/2] approximant

$$\tanh(x) \approx \frac{x}{\frac{3}{9}x^2 + 1}$$

The choice of the second-order top pressure interpolation

$$\phi_n = -\phi_{n-1}, \quad (26)$$

instead of (23) gives a different discrete dispersion relation and the approximation

$$\tanh(x) \approx \frac{x}{\frac{1}{2}x^2 + 1},$$

which is obviously less accurate than (25).

3.5.2 Keller box scheme

Stelling and Zijlema (2003) [29] noted that their z -coordinate implementation of the “Keller box scheme” [48] had much better dispersion characteristics than another implementation (similar to the one above but with a different treatment of boundary conditions). Zhu et al. (2014) [34] more recently derived an analytical dispersion relation for this scheme and proposed optimised layer thicknesses. We repeat briefly this analysis here, but using the simpler framework presented above.

The box scheme relies on a collocated discretisation of the vertical velocity w and non-hydrostatic pressure ϕ , both defined on the layer interfaces (see Figure 2, right). The layer-averaged horizontal pressure gradient required in the horizontal momentum equation (2) is then obtained by averaging the corresponding layer-interface gradients. In a similar manner, the evolution equation (3) for the vertical momentum in each layer is written as the average of the evolution equations for the layer-interface momenta (see Section 3.6.2 for a detailed description). Using the notations above, the corresponding linearised perturbation system can be written

$$\begin{aligned} -\omega h'_k + \bar{h}_k u'_k \hat{k} &= 0, \\ -\bar{h}_k u'_k \omega &= -g \hat{k} \bar{h}_k \sum h'_k - \bar{h}_k \hat{k} \frac{\phi'_{k+1} + \phi'_k}{2}, \\ -\bar{h}_k \frac{w'_k + w'_{k-1}}{2} \omega i &= \phi'_k - \phi'_{k+1}, \\ \bar{h}_k u'_k \hat{k} i + w'_k - w'_{k-1} &= 0, \end{aligned}$$

with the (linearised) bottom impermeability condition

$$w'_{-1} = 0,$$

and the (exact) top pressure boundary condition

$$\phi'_n = 0$$

Computing the determinant then gives the dispersion relations (see Appendix C). For a single layer, the corresponding approximation is

$$\tanh(x) \approx \frac{x}{\frac{1}{4}x^2 + 1} \tag{27}$$

which seems worse than (25) (using the Padé approximant as reference).

To better appreciate the difference between schemes, we display in Figure 3 the ratio of the discrete phase velocity $c = \omega / \hat{k}$ to the exact Stokes phase velocity

$$c_e = \sqrt{\frac{g}{\hat{k}} \tanh(\hat{k} h)}$$

as functions of $\hat{k} h$. We have also included the dispersion relations of the extended Boussinesq two-layer model of Chazel et al. (2009) [24] and of the optimised one-parameter Serre–Green–Naghdi model [49].

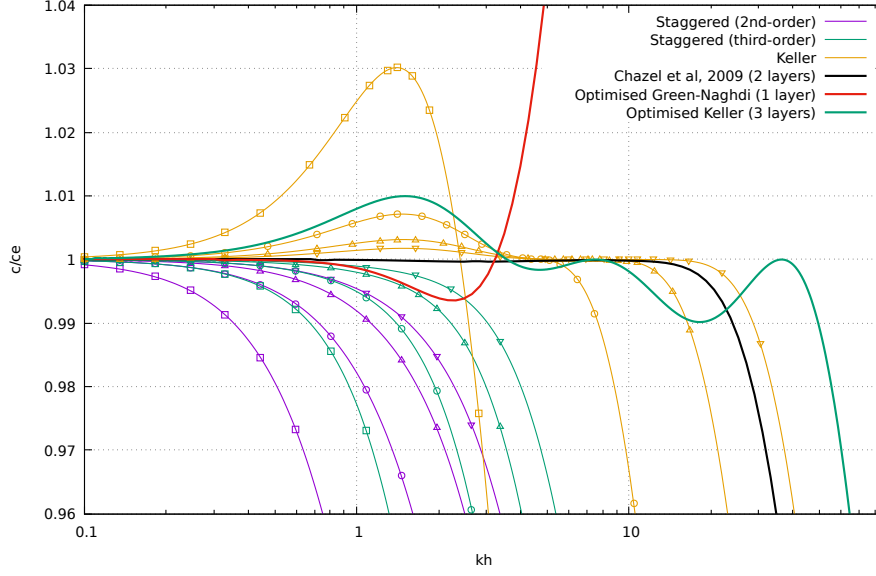


Figure 3. Ratio of the discrete phase velocity to the exact phase velocity for a range of numerical schemes and number of layers. The symbols indicate the number of layers: \square , one layer; \circ , two layers; \triangle , three layers; ∇ , four layers.

The two-layer model of Chazel et al. (2009) has excellent dispersion properties, which is expected since Boussinesq models are designed based on their dispersion characteristics. The range of validity of the staggered schemes of section 3.5.1 and of the Keller box scheme both increase with an increasing number of layers, as expected. Using the third-order free-surface pressure condition (23) instead of the second-order condition (26) also significantly improves the range of validity of the staggered scheme. The range of validity of the Keller box scheme is however much broader than that of the staggered schemes. One can note in particular that the two-layers Keller scheme is superior to the single-parameter optimised Green-Naghdi model and that the three- and four-layers (uniform thicknesses) Keller schemes are competitive with the two-layer Boussinesq model of Chazel et al. (2009). Using an “optimal” layer thickness distribution of 68%–26.5%–5.5%, as in [34], gives a three-layer Keller model with less than 1% error for a very broad range of $\hat{k}h$.

It is also interesting to note that the one-layer Keller box scheme, with a dispersion relying on approximation (27), has a much broader range of validity than the one-layer, third-order, staggered scheme, even though (25) is closer to the Padé approximant. Of course, this increase in the range of validity comes at the expense of a lesser accuracy for small values of $\hat{k}h$.

3.6 Non-hydrostatic scheme

The non-hydrostatic extension is derived from the classical “projection method” [2, 3]. Note that when using Eulerian discretisations and in particular z -coordinates, the natural projection scheme is a straightforward extension of the staggered MAC or C-grid scheme, applied to the vertical coordinate [27] (also known as a “Lorenz grid” in atmospheric models [50]). This corresponds to the linearised scheme studied in section 3.5.1 and Figure 2 (left). For general coordinates, including the Lagrangian formulation we use here, the slope of the layers must be taken into account.

3.6.1 Staggered vertical discretisation

To illustrate the principle of the method, we start with the simpler staggered vertical discretisation of section 3.5.1. Using the semi-discrete notation introduced above, the non-hydrostatic scheme can be derived as the following extension of the hydrostatic scheme:

1. Compute h_k^{n+1} , $(h u)_k^*$, $(h w)_k^*$ using h_k^n , u_k^n , w_k^n and the hydrostatic scheme described in the previous section. Note that the vertical momentum $(h w)$ has been added to the list of advected tracers. The \star denotes intermediate values which do not include the hydrostatic and non-hydrostatic pressure terms.
2. Project the $(h u_k^*, h w_k^*)$ vector using ϕ_k so that the incompressibility condition (4) is verified. Following (2) and (3), the momentum at $n+1$ is obtained as

$$(h u)_k^{n+1} = (h u)_k^* - \Delta t (\partial_x(h \phi) - [\phi \partial_x \hat{z}])_k, \quad (28)$$

$$(h w)_k^{n+1} = (h w)_k^* - \Delta t [\phi]_k. \quad (29)$$

Replacing the corresponding expressions in (4) then gives the Poisson-like equation

$$\left[\frac{[\phi]}{h} - \left(\frac{\partial_x(h \phi) - [\phi \partial_x \hat{z}]}{h} \right) \partial_x \hat{z} \right]_k = \frac{\partial_x(h u)_k + [w^* - \textcolor{red}{u}^* \partial_x \hat{z}]_k}{\Delta t}, \quad (30)$$

with u_k^* computed using (15), $w_k^* = (h w)_k^* / h_k^{n+1}$ and where the terms in colour depend on the slope of the layers. If these terms are neglected, the standard structure of a discrete Poisson problem is recovered (the first term in the l.h.s. is the horizontal second derivative, the second term is the vertical second derivative and the r.h.s. is the divergence of the intermediate velocity field).

3. Use (28) and (29) to compute the final values of $(h u)_k^{n+1}$ and $(h w)_k^{n+1}$ (see also Appendix B).

As for any incompressible flow solver, the overall efficiency of the scheme depends on the linear solver used to invert the Poisson-like system (30). We use the generic geometric multigrid solver built into Basilisk, which also works in parallel and on quadtree-adaptive grids. A more detailed description of this solver is given in [51] where it is used to invert the pseudo-elliptic Green–Naghdi operator. To improve robustness and convergence rate, column relaxation is used in the vertical direction: in the discretised version of the l.h.s. of (30), the off-diagonal values of ϕ in the horizontal directions, $\phi_{i+1,k}$ and $\phi_{i-1,k}$ are supposed to be known, which then gives a linear, tri-diagonal system of equations for $\phi_{i,k}$ which can be inverted efficiently using Thomas algorithm. This vertical column relaxation is combined with multigrid acceleration to obtain the solution to within a specified tolerance. The value of this tolerance is also important for efficiency. A good balance between accuracy and efficiency is obtained using the following criterion

$$\frac{|h_k (\nabla \cdot (h \mathbf{u})_k + [w - \mathbf{u} \cdot \nabla \hat{z}]_k)|_\infty}{|h_k|_1^2} \Delta t < \epsilon, \quad (31)$$

where the max norm $||_\infty$ and average $||_1$ are computed over all layers and all elements. The tolerance ϵ can be interpreted as the maximum relative change of volume allowed for any discretisation element during one timestep. Unless otherwise indicated it will be set to 10^{-3} in the numerical applications below.

This scheme works well in practice but is reasonably complicated to implement, in particular due to the slope terms in (30). Its extension to the more complicated Keller box scheme is thus not straightforward. Thankfully all the slope terms in (30) are not equally important. The terms in blue and green in the l.h.s. are in particular $O(\Delta t)$ less important than the term in red in the r.h.s. Indeed neglecting them would only contribute to increasing the existing $O(\Delta t)$ time-splitting error of the projection scheme. Furthermore the blue term depends on the vertical difference of the layer slopes (i.e. a second-derivative), while the green terms depend on even higher-order derivatives.

To test these assumptions, we compared the results obtained using the full scheme and a simplified scheme neglecting the slope terms (green and blue) in (30), for all the test cases presented in section 4 and found negligible differences. As expected, ignoring the blue term has the largest (but still negligible) impact. Note that all other slope terms must be retained, in particular in the r.h.s. of (30) and in (28).

Remark 5. It is important to note that this approximation is possible only because of the choice of a Lagrangian vertical discretisation. In the case of an Eulerian or quasi-Eulerian discretisation the vertical velocity is used to transport tracers in the vertical direction. Volume, mass and tracer conservation at the discrete level thus rely on verifying *exactly* the discrete non-divergence condition. The corresponding discrete Laplacian operator, used in the l.h.s. of the Poisson equation, must then be the exact composition of the discrete divergence and discrete gradient operators and not just an approximation.

For the Lagrangian vertical discretisation on the other hand, volume, mass and tracer conservation do not depend on the vertical velocity field (which is not used for transport) but is guaranteed by the consistent advection of the layer thicknesses and of the tracers. Using an “approximate vertical projection” may thus affect the accuracy of the non-hydrostatic terms but not the conservation properties.

3.6.2 Keller box scheme

We have seen in section 3.5 that the staggered vertical scheme is not competitive with extended Boussinesq models, even when a (prohibitively) large number of layers is used. This motivates the generalisation of the non-hydrostatic algorithm to the Keller scheme. The horizontal momentum and vertical velocity are now given by

$$\begin{aligned} (h u)_k^{n+1} &= (h u)_k^* - \Delta t (\partial_x(h \phi) - [\phi \partial_x \hat{z}])_k, \\ \left(\frac{w_{k-1/2} + w_{k+1/2}}{2} \right)^{n+1} &= w_k^* - \Delta t \frac{[\phi]_k}{h_k^{n+1}}, \end{aligned}$$

and the pressure gradient in the middle of the layer is approximated as

$$\partial_x(h \phi)_k = \frac{\partial_x(h_k \phi_{k-1/2}) + \partial_x(h_k \phi_{k+1/2})}{2} \quad (32)$$

Using the incompressibility condition (4), the pressures $\phi_{k+1/2}$ and velocities $w_{k+1/2}$ are then solutions of the system

$$\begin{aligned} \frac{w_{k-1/2} + w_{k+1/2}}{2} &= w_k^* - \Delta t \frac{\phi_{k+1/2} - \phi_{k-1/2}}{h_k}, \\ \partial_x(h u)_k^* - [u^* \partial_x \hat{z}]_k - \Delta t \frac{\partial_x \partial_x(h_k \phi_{k-1/2}) + \partial_x \partial_x(h_k \phi_{k+1/2})}{2} + [w]_k &= 0, \end{aligned} \quad (33)$$

where we have neglected the terms equivalent to the **green** and **blue** terms in (30), and dropped the $n + 1$ indices for convenience. To obtain a system of equations for ϕ , we need to express $[w]_k$ as a function of ϕ . This is readily done using the recursion relation (derived from (33))

$$w_{k+1/2} = 2 \left(w_k^* - \Delta t \frac{\phi_{k+1/2} - \phi_{k-1/2}}{h_k} \right) - w_{k-1/2},$$

which gives

$$w_{k+1/2} = 2 \sum_{l=0}^k (-1)^{k+l} \left(w_l^* - \Delta t \frac{\phi_{l+1/2} - \phi_{l-1/2}}{h_l} \right),$$

and

$$[w]_k = 2 \left(w_k^* - \Delta t \frac{\phi_{k+1/2} - \phi_{k-1/2}}{h_k} \right) + 4 \sum_{l=0}^{k-1} (-1)^{k+l} \left(w_l^* - \Delta t \frac{\phi_{l+1/2} - \phi_{l-1/2}}{h_l} \right)$$

The pressure equation is then

$$\begin{aligned} & \partial_x \partial_x (h_k \phi_{k-1/2}) + \partial_x \partial_x (h_k \phi_{k+1/2}) + \\ & 4 \frac{\phi_{k+1/2} - \phi_{k-1/2}}{h_k} + 8 \sum_{l=0}^{k-1} (-1)^{k+l} \frac{\phi_{l+1/2} - \phi_{l-1/2}}{h_l} = \\ & \frac{2}{\Delta t} \left(\partial_x (h u)_k^* - [u^* \partial_x \hat{z}]_k + 2 w_k^* + 4 \sum_{l=0}^{k-1} (-1)^{k+l} w_l^* \right) \end{aligned} \quad (34)$$

We see that the value of the pressure in a given layer is now dependent on the values in all the layers below it. Compared to the staggered scheme derived previously (equation (30)), the structure of the linear problem has thus changed from a sparse, multi-diagonal system to a dense, block-Hessenberg system. This has important implications for the expected performance of the linear solver and lead Stelling & Zijlema [30] to propose an approximation of their quasi-Eulerian box scheme which neglects this vertical coupling. We tried their approach without success. In particular, all our implementations of their scheme were unstable for a number of layers larger than five. Similar unstable results were obtained when using their open-source implementation SWASH [52].

It turns out that system (34) can be solved efficiently using a multigrid-accelerated column relaxation technique similar to that used to solve (30). Assuming known values of $\phi_{i+1,k+1/2}$ and $\phi_{i-1,k+1/2}$ leads to an Hessenberg system for $\phi_{i,k+1/2}$ in each column i . An Hessenberg matrix is the sum of a triangular and a tri-diagonal matrix. The resulting system can be solved efficiently using the algorithm of Henry (1994) [53]. Note however that the cost of solution for each column relaxation will scale like $O(n^2)$ rather than $O(n)$ for a tridiagonal system. We will demonstrate later that this does not preclude efficient solutions.

The resulting algorithm can be summarised as follows:

1. Compute h_k^{n+1} , $(h u)_k^*$, $(h w)_k^*$ using h_k^n , u_k^n , w_k^n and the hydrostatic scheme (without pressure gradients).
2. Compute $\phi_{k+1/2}$ using multigrid-accelerated Hessenberg column-relaxations to solve (34).

3. Update the staggered accelerations and velocities

$$\begin{aligned}\alpha_{i+1/2,k} &\leftarrow \partial_x(h\phi)_{i+1/2,k} - [\phi \partial_x \hat{z}]_{i+1/2,k} \\ (ha)_{i+1/2,k} &\leftarrow (ha)_{i+1/2,k} - \alpha_{i+1/2,k} \\ (hu)_{i+1/2,k} &\leftarrow (hu)_{i+1/2,k} - \alpha_{i+1/2,k} \Delta t\end{aligned}$$

where $\partial_x(h\phi)_{i+1/2,k}$ is approximated using (32).

4. Compute $(hu)_{i,k}^{n+1}$ and $(hw)_{i,k}^{n+1}$ using (16) and (29).

The corresponding documented code is available on the Basilisk web site (<http://basilisk.fr/src/layered/nh.h>).

3.6.3 Stability of the non-hydrostatic scheme

Since we use explicit timestepping, we can expect the timestep to be limited according to

$$\Delta t < \frac{\Delta}{|\mathbf{u}| + c},$$

with c the phase speed of gravity waves. Note that $|\mathbf{u}|$ does not include the vertical velocity component since vertical advection is Lagrangian and thus not subject to a CFL constraint. The phase speed should take into account dispersion and thus be $\sqrt{g/\hat{k} \tanh(\hat{k}H)}$ rather than just \sqrt{gH} (as for the hydrostatic case). The wavenumber \hat{k} can be taken as the inverse of the smallest possible wavelength i.e. $1/\Delta$. Re-arranging the terms then gives the condition

$$\Delta t < \frac{\Delta}{|\mathbf{u}| + \sqrt{gH} \sqrt{(\Delta/H) \tanh(H/\Delta)}} \quad (35)$$

The difference with the hydrostatic case can be better understood using the approximation

$$\sqrt{(\Delta/H) \tanh(H/\Delta)} \approx \begin{cases} 1 & \text{if } \Delta \geq H \\ \sqrt{\Delta/H} & \text{if } \Delta < H \end{cases}$$

which shows that the increase in the timestep is proportional to $\sqrt{H/\Delta}$ whenever the mesh size is smaller than the water depth. We will see later that this relaxed stability condition can lead to large gains for practical cases.

Note that this condition is based on the exact dispersion relation, not its numerical approximation (as displayed on Figure 3). We can expect the schemes for which the exact dispersion relation is an (approximate) upper bound to respect this condition: that is all schemes on Figure 3 with the exception of the optimised Green–Naghdi model.

3.6.4 Wave breaking

Since $\eta(\mathbf{x}, t)$ describes the free surface as the graph of a (uni-valued) function, it cannot represent the overturning motion characteristic of breaking waves. This is a classical issue for Boussinesq models of wave motion [54]. In the case of the Saint-Venant equations, the corresponding limiting case is the formation of an hydraulic jump (i.e. a shock-like solution) which is known to be a remarkably good model of breaking (given its simplicity) [55]. The classical approach in Boussinesq models is thus to turn off dispersive terms when breaking is detected [51]. This approach does not seem to generalise easily to the present model. We found instead that the somewhat *ad hoc* following recipe works well.

Breaking is characterised both by steep free-surface slopes and large vertical velocities. Dissipation, linked to the unresolved turbulent motions occurring during breaking, can be introduced by limiting the maximum vertical velocity. This is done after application of the vertical non-hydrostatic pressure gradient (equation (29)) using

$$w_k^{n+1} \leftarrow \text{sign}(w_k^{n+1}) \min(|w_k^{n+1}|, b \sqrt{g |H|_\infty}), \quad (36)$$

where $\sqrt{g |H|_\infty}$ is assumed to be the characteristic horizontal velocity scale of the breaking wave and the *breaking parameter* b is a number smaller than one. In the case of strong breaking (see the example in section 4.6) this is not sufficient to guarantee the stability of the non-hydrostatic solver and the following slope-limiting is added to $\partial_x \hat{z}$ in equations (28) and (34)

$$\partial_x \hat{z} \leftarrow \begin{cases} \partial_x \hat{z} & \text{if } |\partial_x \hat{z}| < \tan(30^\circ) \\ \text{sign}(\partial_x \hat{z}) & \text{otherwise} \end{cases} \quad (37)$$

3.6.5 Wetting and drying

Although wetting and drying, and the associated “coastline representation”, has often been considered a difficult issue in ocean models, it is now well-known that the Saint-Venant equations are physically and mathematically well-posed even in the limit of vanishing water depth [37]. The idea of a “coastline” separating dry from wet (or “topography” from “bathymetry”) thus appears to be a misleading concept, as demonstrated in particular by tsunami/flood models based on the Saint-Venant equations [46, 56]. The same approach is taken here. The only adaptation specific to the wet/dry transition is for the solution of the Poisson equation (30) or (34) which is first multiplied by h_k to avoid any division by zero.

4 Numerical results

We present a range of numerical results which illustrate the generality and efficiency of the model. When not otherwise indicated, results have been obtained using the non-hydrostatic Keller box scheme and a timestep given by the stability condition (35). The layer-slope limiting (37) is always turned on but the vertical-velocity limiting (36) is turned off by default. The performances indicated were obtained on an 8 cores Intel® Core™ i7-8550U CPU @ 1.80GHz with 8 GB of memory. The links to the source codes sufficient to reproduce all the results and figures are given in the captions. Note also that other tests and examples are available at <http://basilisk.fr/src/layered/hydro.h#usage>

4.1 Dispersion relation

This first case verifies that the numerical dispersion relation follows the linear analysis of section 3.5 and Figure 3. A standing wave is initialised in a periodic domain and its frequency is measured using the zero-crossings of the wave amplitude over ten periods of oscillation. To ensure spatially-converged results, the wavelength is resolved using 128 grid points. The total water depth H is varied to obtain results for a broad range of $\hat{k} H$. The convergence criterion (31) is decreased to $\epsilon = 10^{-6}$. This is necessary to obtain accurate results at large $\hat{k} H$. Figure 4 illustrates the results obtained for a range of numerical schemes and a varying number of layers. The Green–Naghdi results were obtained using the numerical scheme described in [51]. Close agreement is obtained with the theoretical dispersion relations.

All results were obtained using a timestep given by the stability condition (35), with the exception of the Green–Naghdi model which proved unstable, as expected, and had to use the more restrictive hydrostatic condition (17). Note that for the maximum value of $\hat{k}H \approx 60$ displayed on the figure and for 128 grid points per wavelength, the ratio of the non-hydrostatic timestep to hydrostatic timestep is approximately $\sqrt{H/\Delta} = \sqrt{128 \hat{k}H/2\pi} \approx 35$.

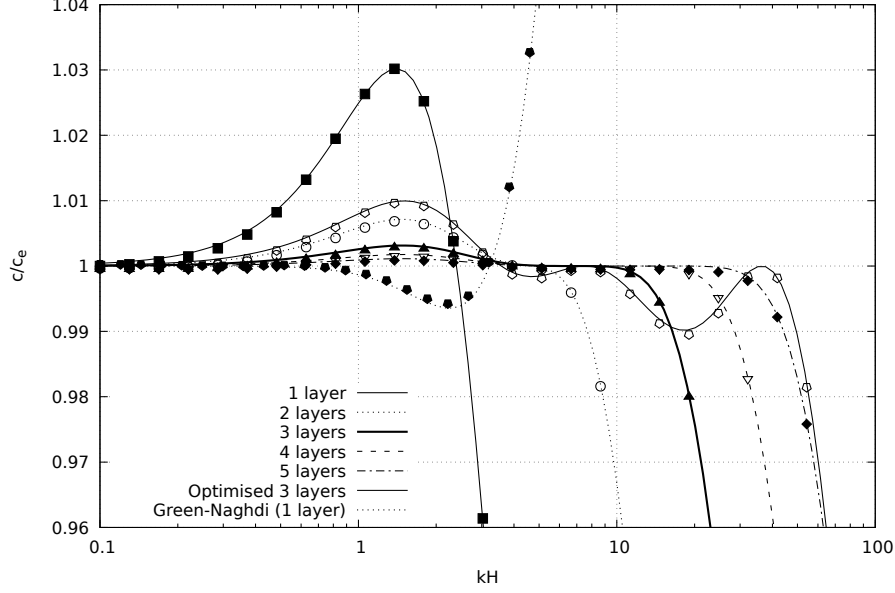


Figure 4. Dispersion relation. The lines are the analytical solutions and the symbols the numerical results (<http://basilisk.fr/src/test/dispersion.c>).

The same test case is used to estimate the numerical total energy variation per period, computed as the sum of the discrete potential and kinetic energies. The results are displayed in Figure 5. All schemes appear to conserve energy very well.

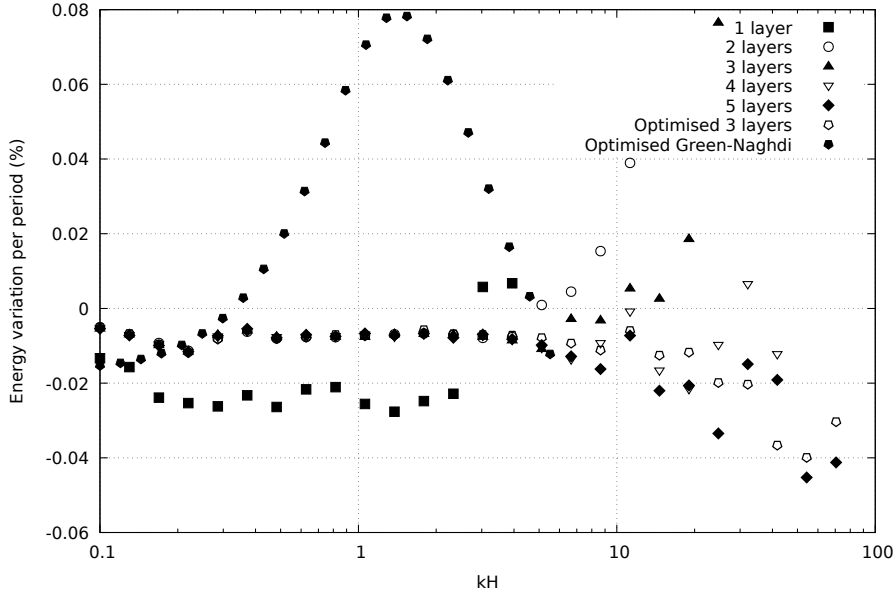


Figure 5. Relative variation in total energy per oscillation period (<http://basilisk.fr/src/test/dispersion.c>).

4.1.1 Profile of vertical velocity

The theoretical vertical profile of the vertical velocity component under a linear wave is given by Stokes' theory as [24]

$$w_s(z) = \frac{A g k}{\omega} \frac{\sinh(k z)}{\cosh(k H)} \sin(\omega t - k x),$$

with A the amplitude of free-surface displacement. It reaches its maximum for multiples of $T/4$ with T the Stokes wave period. Figure 6 displays the normalised profiles of vertical velocity obtained numerically at $t = 37 T/4$, for different values of $k H$. As the wave shortens, its velocity field becomes shallower and thus harder to resolve with a fixed number of equidistant layers.

It is interesting to note that the behaviour of the Keller box scheme used here is quite different from that of the staggered scheme (not shown). As illustrated in Figure 3, the staggered scheme would have much larger dispersion errors. This would lead to large phase errors at $t = 37 T/4$ and it would not be possible to compare the velocity profiles with linear theory, unless a very large number of layers was used.

Remarkably, with the Keller scheme the phase is accurate, even though the number of layers and their distribution are insufficient to estimate an accurate vertical velocity profile, as illustrated for $k H = 24.7065$ in Figure 6 (triangular symbols). This profile also suffers from a clear grid-scale oscillation which is probably due to the vertical collocation of pressure and vertical velocity. This could be a concern but has not caused problems in any of the cases we present. Furthermore, the amplitude of these spurious oscillations vanishes with increasing vertical resolution.

As expected, these grid-scale oscillations are absent when using the vertically-staggered scheme (not shown), however the velocity profile extends much deeper than predicted by theory, which also explains the corresponding dispersion/phase error.

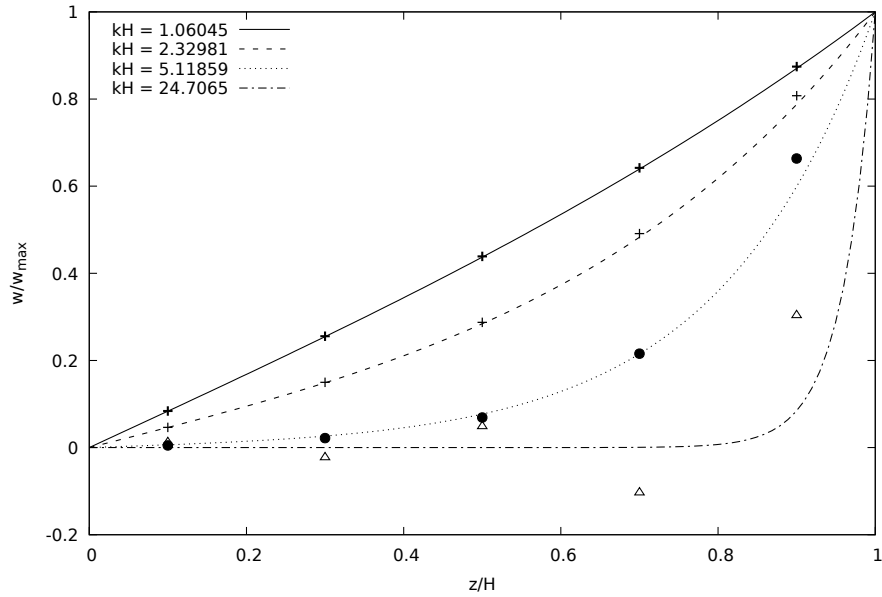


Figure 6. Normalised profiles of the vertical component of velocity under a linear wave, for different values of $k H$. The numerical results (symbols) are obtained using the Keller box scheme with five layers (<http://basilisk.fr/src/test/dispersion.c>).

4.2 Sinusoidal wave propagation over a bar

We next consider a classical test case for the Boussinesq models used in coastal engineering. Beji and Battjes [57] and Luth et al. [58] studied experimentally the transformation of sinusoidal waves propagating over a submerged bar (or reef). This is a good test case for dispersive models as higher harmonics are non-linearly generated and released with phase shifts corresponding to the dispersion relation. Note that wave breaking also occurs experimentally on top of the submerged bar. See section 4.2 of [35] for a detailed description of the setup and results using a one-layer model.

The results displayed on Figure 7 were obtained using the two-layer Keller scheme with a breaking parameter $b = 0.1$ (see (36)) and 2048 grid points (for a domain length of 50 metres). The agreement with measurements is excellent, in particular for gauge eleven which is the farthest from the reef and thus most affected by dispersion errors. Using only a single layer (not shown), the results are comparable up to gauge nine but dispersion errors clearly affect the results for gauges ten and eleven. The results of Figure 7 can be compared to Figure 4 of [35] and are of comparable quality up to gauge eight, but much better for the last three gauges.

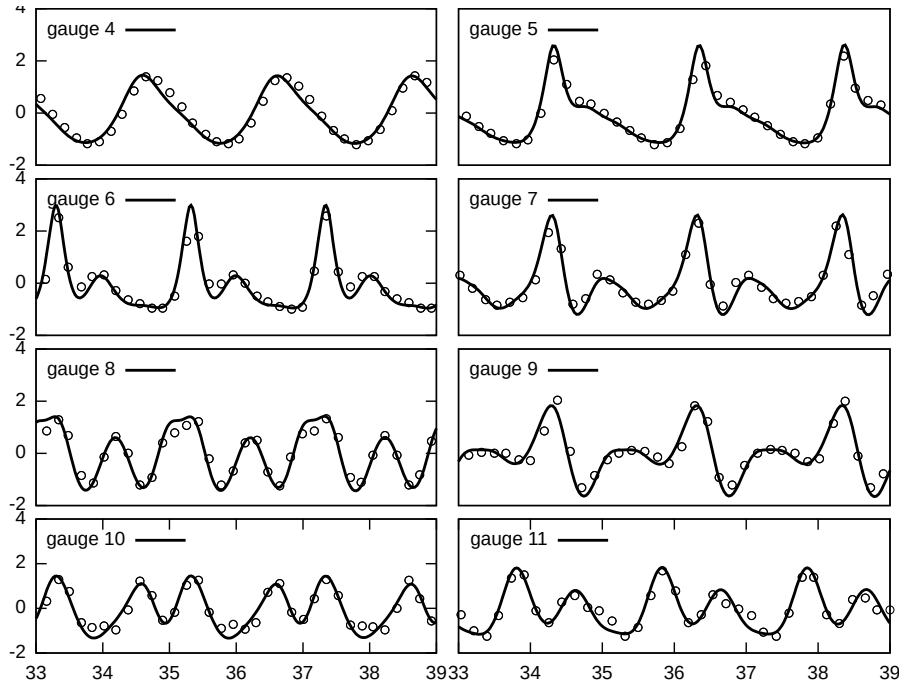


Figure 7. Computed (solid line) and measured (symbols) free-surface elevations for the wave propagation over a bar obtained experimentally by Beji and Battjes [57]. The numerical results are obtained with two layers (<http://basilisk.fr/src/test/bar-ml.c>).

4.3 Dispersive tsunami

While tsunamis are often modelled using the non-dispersive Saint-Venant equations, it is now well-known that dispersive effects are often not negligible [59, 60, 51]. To illustrate the applicability and robustness of the layered model to such a case, we consider the Japanese

Tohoku tsunami of 11th March 2011, which we already studied using both the Green–Naghdi [51] and Saint-Venant [56] models. The setup is identical to these previous papers where a more detailed description and discussion can be found. The results obtained using the one-layer Keller scheme are displayed on Figure 8. This simulation combines wetting and drying, dispersion, quadtree-adaptivity and complex bathymetry (reconstructed dynamically using the algorithm described in the Appendix of [46]). The results of Figure 8 can be compared with Figure 12 of Popinet (2015) [51].

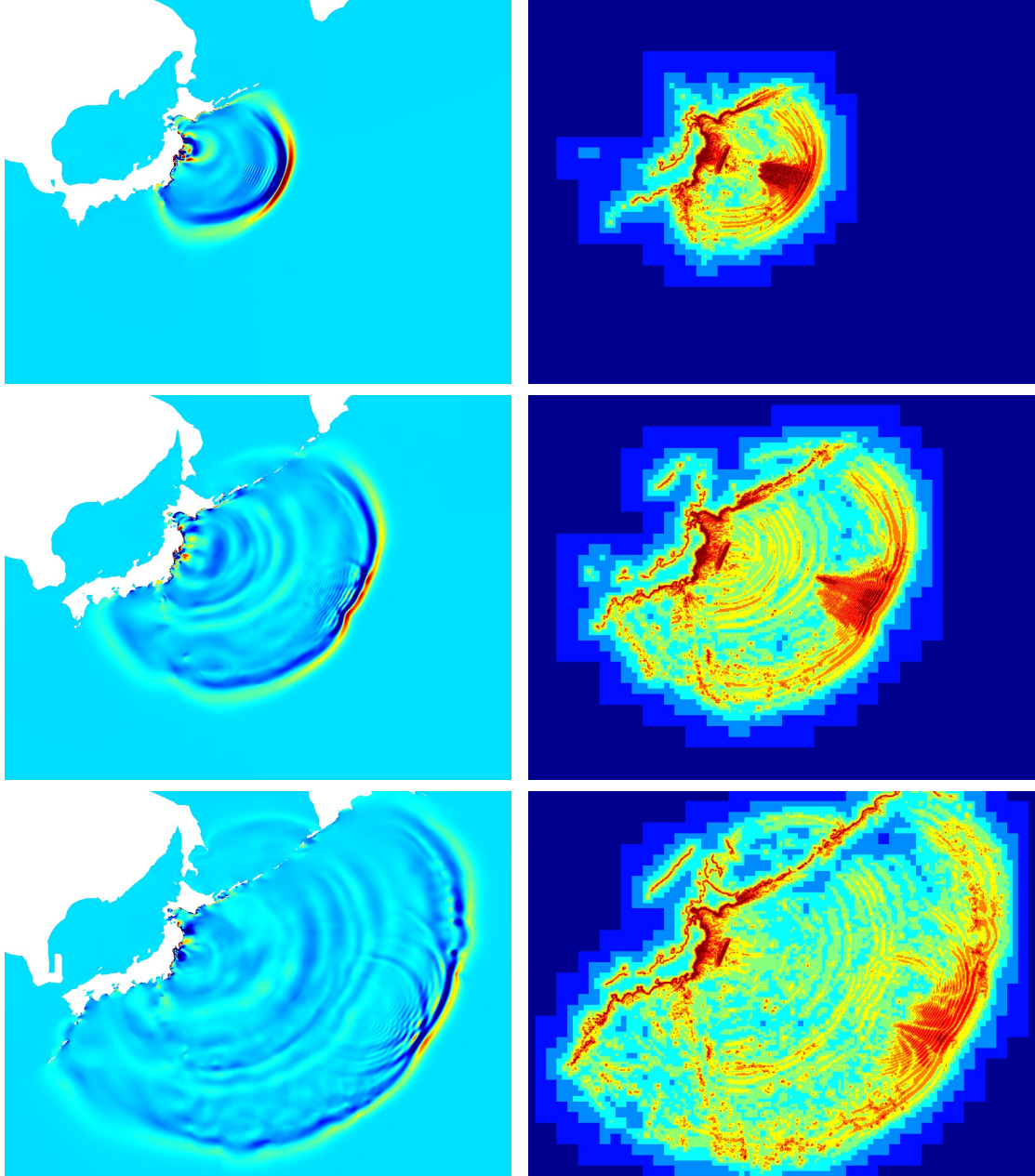


Figure 8. Tohoku tsunami. Evolution of the free-surface elevation (left column) and adaptive grid (right column). The elevation scale ranges from -1 metre (dark blue) to $+2$ metres (dark red). The level of refinement ranges from 5 (≈ 250 km, dark blue) to 13 (≈ 1 km, dark red). The times are from top to bottom: 1, 2 and 3 hours after the fault rupture (<http://basilisk.fr/src/examples/tohoku.c>).

The importance of dispersion on long-distance wave propagation is highlighted on Figure 9 where three different models are compared. The Serre–Green–Naghdi results (obtained using the solver described in [51]) are clearly close to those obtained using the one-layer non-hydrostatic model, the most significant difference being the longer “tail” of dispersive waves obtained using the one-layer model. This reflects the lower numerical dissipation of this scheme, compared to that of the second-order MUSCL/Riemann scheme used in [51]. Besides the absence of a dispersive tail, the hydrostatic scheme also produces a much steeper leading wave: a result consistent with previous studies and the expectation that the lack of dispersive terms leads to over-steepening of wave fronts (and eventually the formation of hydraulic jumps).

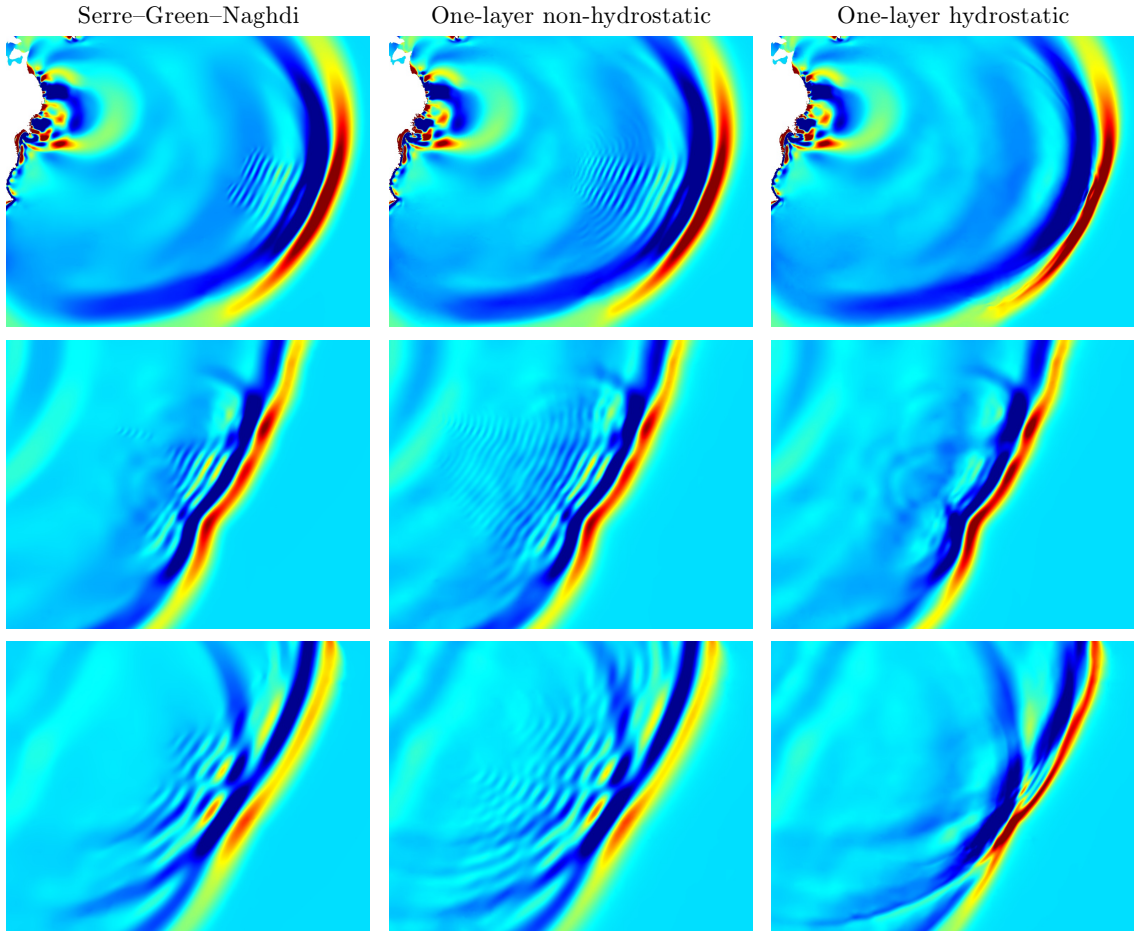


Figure 9. Close-up views of the leading wave front for the times of Figure 8, obtained using different models. Left-column: Serre–Green–Naghdi; Central-column: non-hydrostatic one-layer model; Right-column: hydrostatic one-layer model (<http://basilisk.fr/src/examples/tohoku.c>).

The ability of the model to deal robustly with wetting and drying is illustrated on Figure 10 which shows a detail of the maximum flood level. The area displayed is approximately 140×200 km and includes the extensively-flooded Sendai plain as well as extremely high runoff on the Oshika peninsula. These results are consistent with both satellite data and previous simulations (right panel), although the resolution here is significantly lower (a

maximum of 1 km compared to 250 metres)

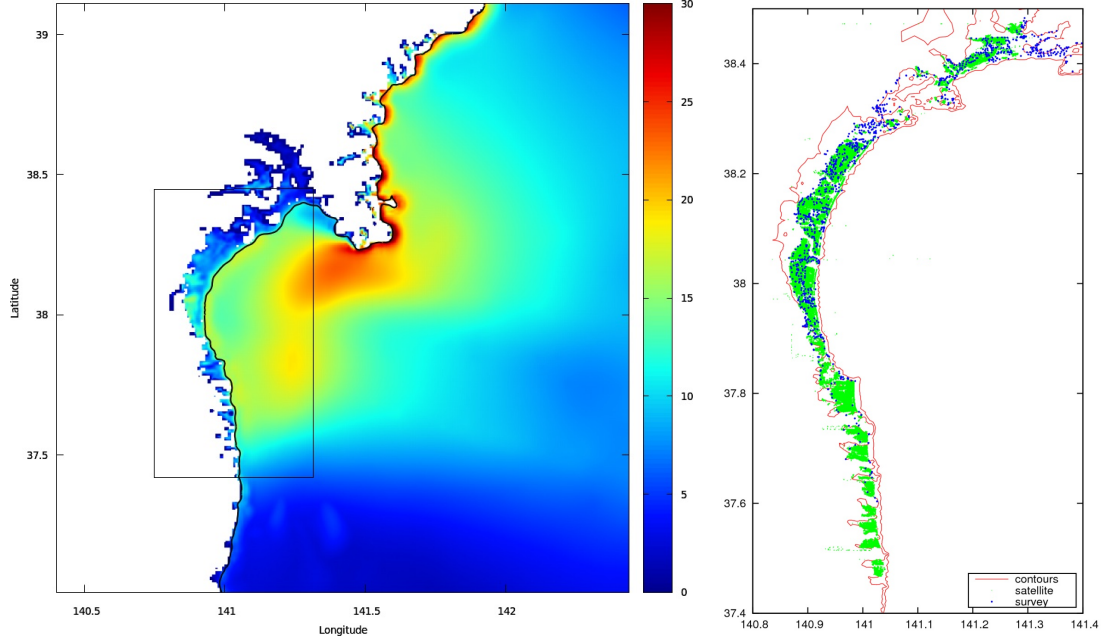


Figure 10. (left) Detail of flooding in the Ofunato, Miyagi and Fukushima districts. The black line indicates the normal coastline. The colorscale is the maximum flood level (maximum water elevation above normal) in meters (<http://basilisk.fr/src/examples/tohoku.c>). (right) Maps of inundation extent for the Miyagi prefecture area: zero and ten meters water level contours (red), satellite estimate of inundation extent on 12 March 2011 (green), field survey data points (blue). The corresponding area is shown by the black box on the left panel. See also [56].

Finally, Table 1 gives the runtimes of the different models, corresponding to six and a half hours (390 minutes) of physical time. Since the mesh is adaptive, the number of grid points indicated is the average over the entire simulation time. Eight cores were used with OpenMP parallelism. While the highest computation speed is obtained with the hydrostatic solver, the shortest computation time is given by the non-hydrostatic solver. This reflects the significant gain in number of timesteps due to the relaxed stability condition (35). The Serre–Green–Naghdi solver [51] is substantially slower, even though it uses less grid points (due to increased numerical dissipation of small waves).

Model	# Timesteps	# Points	Runtime	Speed
Hydrostatic one-layer	17 532	500 666	318	460 000
Non-hydrostatic one-layer	7945	503 230	258	258 000
Serre–Green–Naghdi	17 168	330 161	448	211 000

Table 1. Runtimes (minutes) for the dispersive tsunami case. The speeds are given in points×timesteps/sec. The number of grid points is a time average.

4.4 Large-amplitude standing wave

After this validation of the numerical scheme in the context of dispersive waves, resolved

with only a few layers, we now move to applications illustrating the consistency of the model for the incompressible Euler/Navier-Stokes equations with a free-surface. A simple initial condition is given by a large amplitude sinusoidal perturbation of the free-surface

$$\eta(x, 0) = 0.5 + 0.07 \cos(2 \pi x),$$

and zero initial velocity, in a one meter wide domain with free-slip side walls. The acceleration of gravity is set to 9.81 m/s^2 . The horizontal dimension is discretised with 512 grid points and 20 layers are used in the vertical. The evolution of the free surface is illustrated in Figure 11. Strong non-linearity and vertical accelerations give rise to a vertical “jet” of liquid which then falls back due to gravity. The evolution with time of the maximum vertical velocity and free-surface slope is illustrated in Figure 12. The vertical acceleration is of the order of the acceleration of gravity, i.e. the flow is strongly non-hydrostatic, and the maximum slope is of order one. Clearly in this case the metric/slope terms cannot be neglected, which is confirmed when comparing the results obtained when neglecting the slope terms $\partial_x \hat{z}$ in expressions (28) and (34) (dotted line on Figure 11).

Since analytical solutions are not available for this non-linear problem, a calculation is also done using a completely different solver: the incompressible, two-phase, Volume-Of-Fluid, Euler/Navier-Stokes solver described in [61]. This scheme, also implemented in Basilisk, uses a fixed regular Cartesian grid and a VOF description of the interface separating the two phases. The equations of motion are solved both in the dense phase below the interface and the light phase above the interface. The density ratio is set to $1/1000$ which is small enough to ensure that the motion of the light phase does not affect the results. The 1×1 metre domain is resolved using 256×256 grid points. The corresponding interface evolution is displayed in Figure 11 (dashed lines). The agreement with the layered model is clearly excellent at all times.

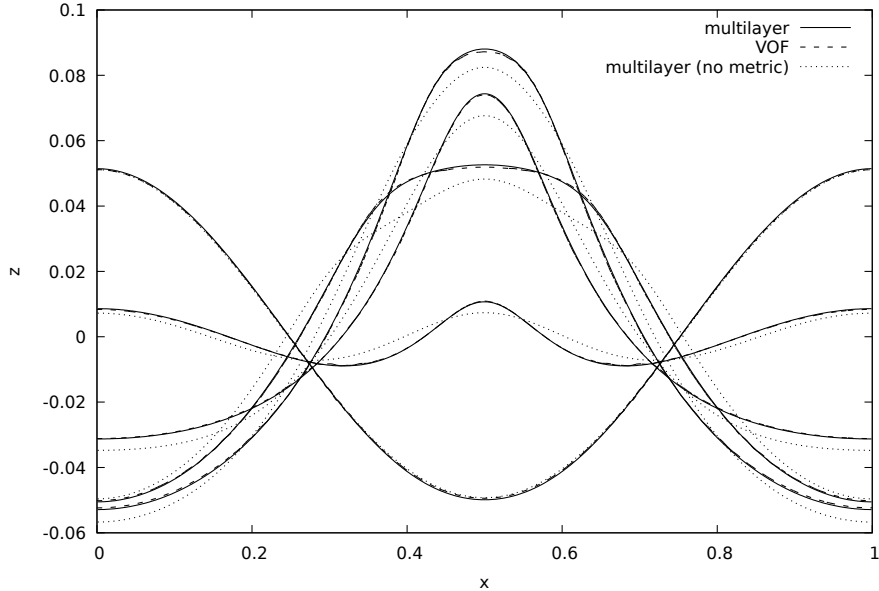


Figure 11. Free-surface evolution. The free-surface is displayed at times $t = 0.1$ sec to 0.5 sec in 0.1 sec intervals (<http://basilisk.fr/src/test/large.c>).

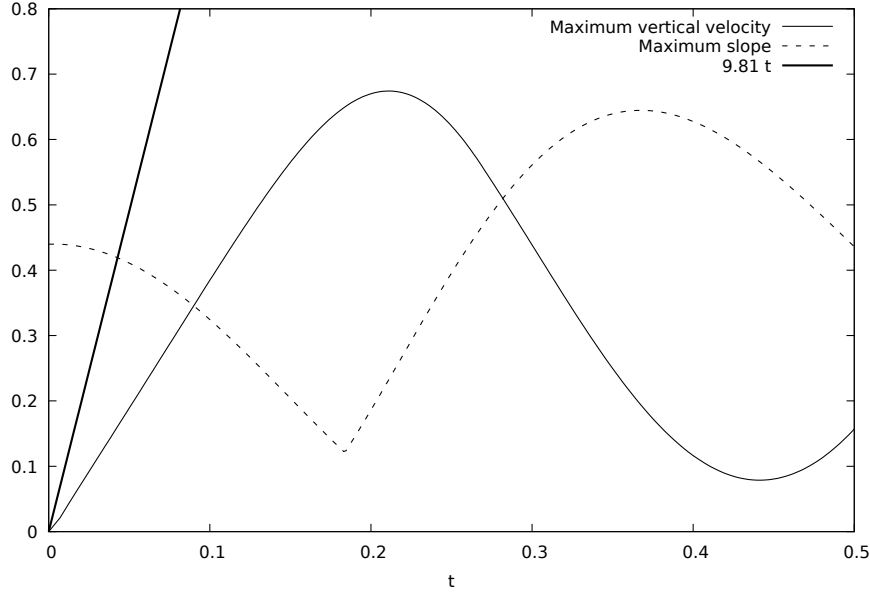


Figure 12. Timeseries of maximum vertical velocity and maximum slope (<http://basilisk.fr/src/test/large.c>).

4.5 Viscous hydraulic jump

Shocks are a well-known characteristic of the hyperbolic nature of the Saint-Venant equations. It is also well-known that they can be a useful theoretical model of a wide range of real features observed in geophysical flows and/or hydraulic engineering e.g. wave breaking [55], bores and hydraulic jumps [16]. The inclusion of viscous and non-hydrostatic effects into this theoretical model is difficult however, and numerical models (combined with experiments) are a useful way to improve our understanding of these complex processes.

4.5.1 Hydrostatic

We first consider the viscous, hydrostatic transcritical flow over a bump, driven by a parabolic velocity profile at inflow and an outflow at a given depth. The inflow rate and water depth at the outflow are chosen so that the Froude number U/\sqrt{gH} is smaller than one at inflow, larger than one over the bump and smaller than one at the outflow (i.e. the flow is transcritical). The domain is 30 metres long and roughly one metre deep, so not unlike a typical channel flow one would encounter in hydraulic engineering. The bump geometry is given by the Gaussian function

$$z_b(x) = 0.4 e^{-\frac{(x-10)^2}{5}}$$

The outflow depth is set to 0.6 metres and the inflow profile is given by

$$u(0, z) = \frac{3}{2} \frac{Q}{H(0)} \left[1 - \left(\frac{z}{H(0)} - 1 \right)^2 \right],$$

with the flow rate Q set to one m^2/s . The (vertical) kinematic viscosity is set to $10^{-2} \text{ m}^2/\text{s}$ and the acceleration of gravity is $g = 9.81 \text{ m/s}^2$. Note that viscosity drastically changes the expected solution, since it controls the position of the (stationary) hydraulic jump: in the absence of viscosity the hydraulic jump is simply advected out of the domain.

The horizontal is discretised with 512 grid points and the vertical with 20 layers. We consider two different numerical models: the present multilayer model run in hydrostatic mode with vertical viscosity and remapping and the shock-capturing, Riemann, quasi-

Eulerian multilayer model described in [16], itself inspired from [15].

The simulation is started from rest with a constant water level of 0.6 metres. The corresponding evolution of the free surface is illustrated on Figure 13. As fluid flows in from the left and hits the fixed-depth right boundary, an hydraulic jump forms just downstream of the bump and eventually reaches a stationary position (after about 60 seconds). The agreement between the two models is clearly very good, both for the time evolution and the stationary solution, with a slightly sharper shock for the Riemann-based scheme.

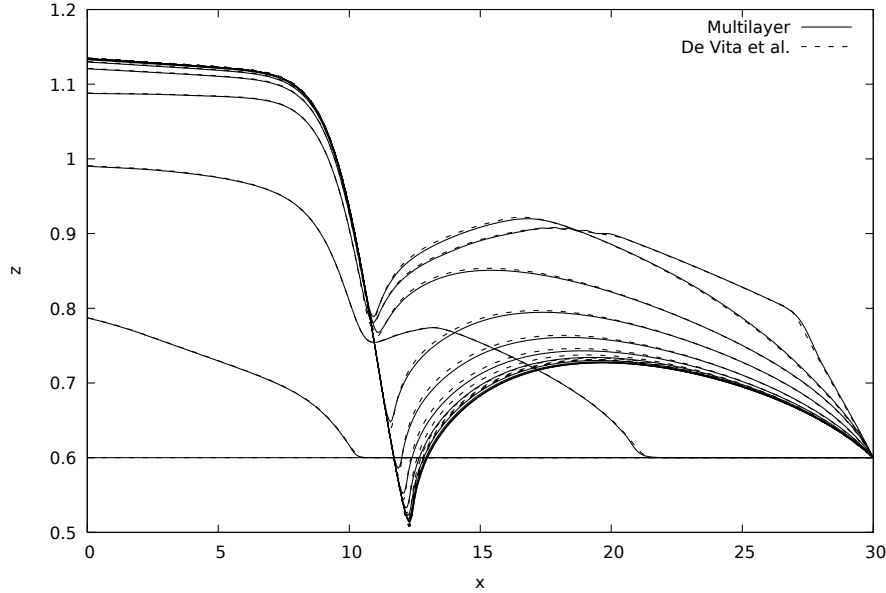


Figure 13. Evolution of the free-surface obtained with the two models indicated in the legend. The time interval between profiles is 5 seconds (<http://basilisk.fr/src/test/gaussian.c>).

This agreement is confirmed by the comparison of the stationary horizontal velocity fields displayed on Figure 14. The white isoline indicates the presence of a (barely resolved) weak recirculation zone behind the bump.

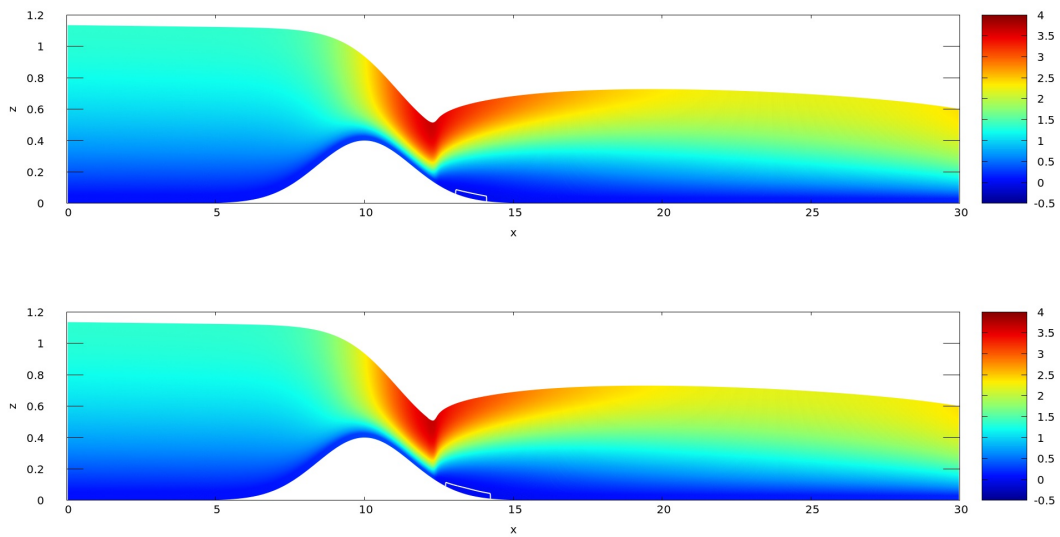


Figure 14. Horizontal velocity. (top) Hydrostatic multilayer model. (bottom) De Vita et al. The white isoline indicates $u = 0$ (<http://basilisk.fr/src/test/gaussian.c>).

4.5.2 Non-hydrostatic

We repeat the same simulation using the non-hydrostatic multilayer solver. This leads to a very different stationary solution, as illustrated in Figures 15, 16 and 17. Vertical inertia is not negligible and leads to the formation of a dispersive wave train rather than a shock-like solution. This is a much more realistic solution than the hydrostatic version and recalls the waves first studied by Favre [62] in the context of dam overflows. The non-hydrostatic pressure field displayed in Figure 18 further confirms that non-hydrostatic effects are important (i.e. the non-hydrostatic pressure divided by the density is of order one, so is comparable to the hydrostatic pressure).

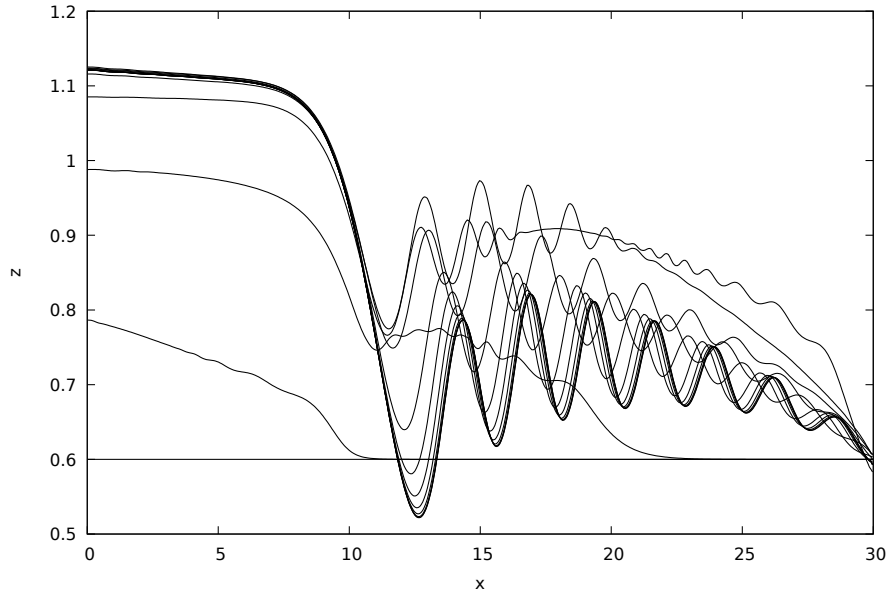


Figure 15. Evolution of the free-surface. The time interval between profiles is 5 seconds (<http://basilisk.fr/src/test/gaussian.c>).

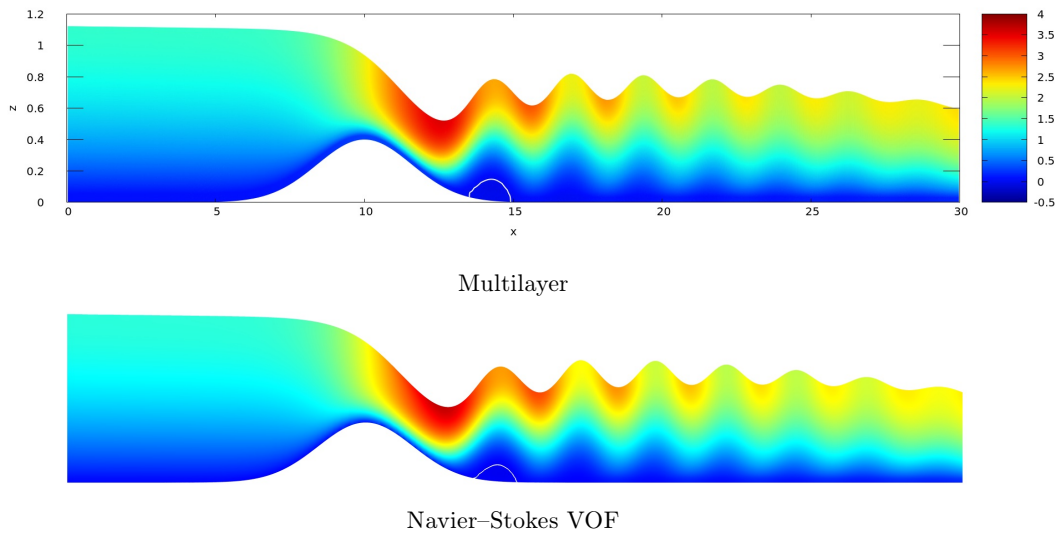


Figure 16. Horizontal velocity field computed with the multilayer solver (<http://basilisk.fr/src/test/gaussian.c>) and the VOF Navier-Stokes solver (<http://basilisk.fr/src/examples/gaussian-ns.c>). The white isoline indicates $u = 0$.

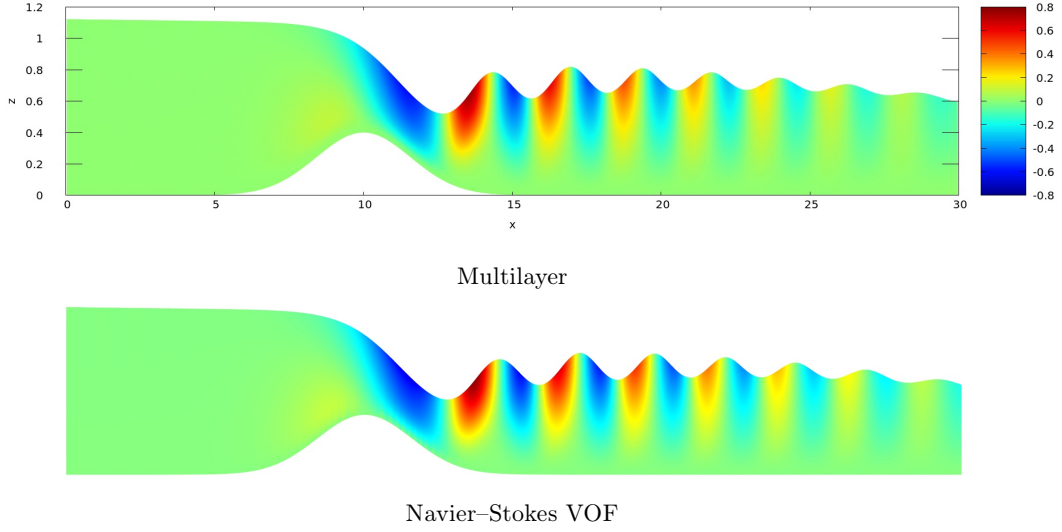


Figure 17. Vertical velocity field computed with the multilayer solver and the VOF Navier–Stokes solver.

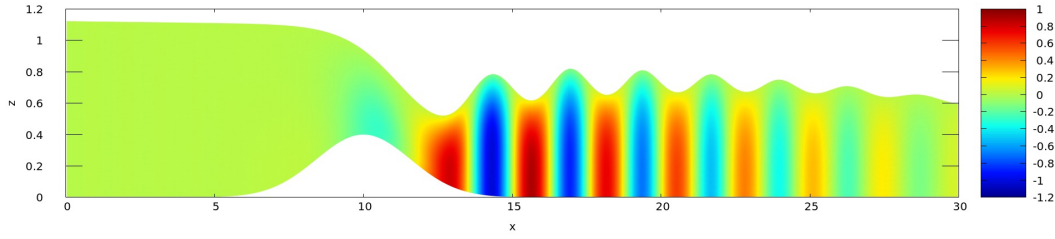


Figure 18. Non-hydrostatic pressure field (divided by the fluid density) computed with the multilayer solver (<http://basilisk.fr/src/test/gaussian.c>).

In Figures 16 and 17, the horizontal and vertical components of the velocity are compared with solutions obtained using the Navier–Stokes/VOF solver introduced in section 4.4. The solution is significantly more difficult to compute in the present case, in particular due to the large horizontal to vertical aspect ratio. Converged results required an equivalent resolution of 1024×50 grid points for the Navier–Stokes solver. The agreement between the two solutions is clearly qualitatively good, including for the weak recirculation zone behind the bump, which is now properly resolved. Figure 19 gives a more quantitative comparison of the final free-surface profiles. The importance of including the metric/slope terms is clear, even for this configuration with a relatively mild maximum slope of 0.25. The source of the remaining discrepancies between the Navier–Stokes and multilayer solutions is unclear. An obvious candidate could be the lack of horizontal viscosity in the multilayer model, but its addition does not seem to significantly change the results, which is not so surprising given the $1/30$ aspect ratio of the flow. An explanation could be some remaining numerical errors in the Navier–Stokes VOF solution.

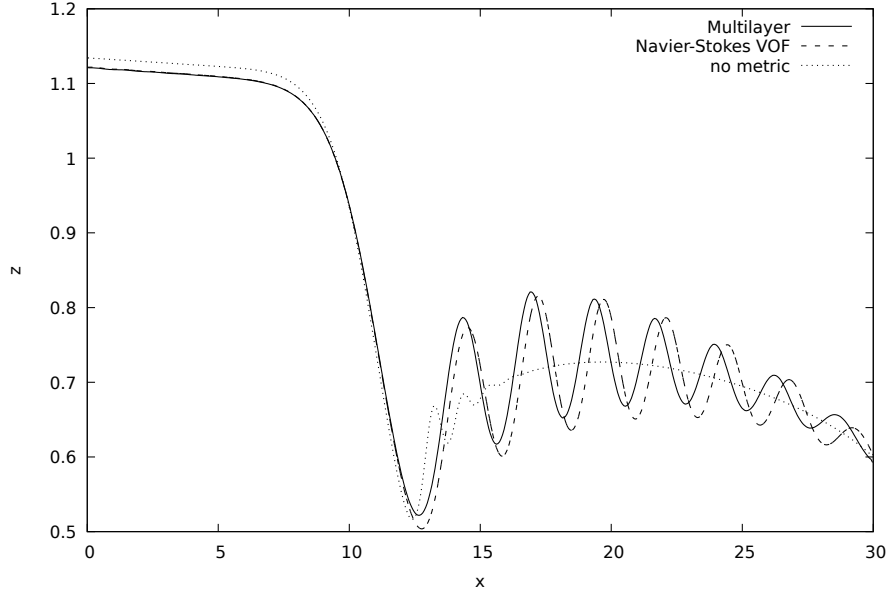


Figure 19. Final free-surface profiles (<http://basilisk.fr/src/test/gaussian.c>).

Finally Table 2 gives a summary of the performances of all solvers for this case. The gain in number of timesteps due to the relaxed stability condition (35) is significant again, but not sufficient to offset the overhead due to the resolution of the Poisson equation (when comparing the runtimes of the hydrostatic and non-hydrostatic multilayer solvers). The hydrostatic multilayer solver is significantly slower than the quasi-Eulerian, Riemann solver of [16]. This is due to the added cost of vertical remapping and reflects the poor performance of the PPR library we are using. The gain in runtime between the multilayer and Navier–Stokes/VOF solutions is remarkable, even if the comparison is somewhat unfair since the VOF solver cannot at the moment use anisotropic meshes which would allow to optimise the discretisation.

Model	# Timesteps	Runtime	Speed	Remap
Hydrostatic (De Vita et al.)	13078	28	4 782 811	n.a.
Hydrostatic (multilayer)	12945	41	3 233 092	47.6%
Non-hydrostatic (multilayer)	8448	147	588 486	10.4%
Non-hydrostatic (NS/VOF)	14754	4448	186 000	n.a.

Table 2. Runtimes (seconds) for the viscous hydraulic jump case. The speeds are given in points×timesteps×layers/sec. NS/VOF run on eight cores, others run on a single core. The “Remap” column gives the proportion of the total runtime taken by vertical remapping.

4.6 Breaking Stokes wave

Since the study and prediction of (very) short waves is one of the primary motivation for the development of dispersive wave models, we now choose an example which illustrates that the multilayer model can handle robustly and accurately the detailed modelling of very steep water waves up to, and even beyond their threshold of stability. In previous work we have studied in detail the evolution of unstable Stokes waves, including breaking and bubble cloud formation [63, 64]. We reuse this setup here, with an initial free-surface and velocity field given by a third-order Stokes wave solution [65]. Since we do not include surface tension, there are only two parameters: the Reynolds number $\text{Re} = \sqrt{g\lambda^3}/\nu$ and

the initial steepness $a\hat{k}$, with $\lambda = 2\pi/\hat{k}$ the wavelength and a the wave amplitude. The water depth is set to half the wavelength and is assumed to be deep enough to have a negligible influence on the results. The solution is obtained using both the two-phase Navier–Stokes/VOF solver (with density and viscosity ratios set to air/water) and the free-surface, non-hydrostatic multilayer solver.

The Reynolds number is set to $\text{Re} = 40000$ and the steepness to $a\hat{k} = 0.35$, which corresponds to a spilling/overturning configuration [63]. The multilayer solver uses 256 grid points in the horizontal and 60 vertical layers. The Navier–Stokes/VOF solver uses an adaptive mesh with a maximum equivalent resolution of 512^2 . Figure 20 illustrates the free-surface evolution for both solvers. The time interval between profiles is $T/4$ with $T = \hat{k}/\sqrt{g\hat{k}}$ the Stokes wave period. The agreement between the two solutions is clearly excellent right up to the point of breaking at $t \approx 5T/4$. The multilayer code cannot be expected to match the VOF solution after breaking, since the interface becomes multivalued, however it is still able to robustly compute a solution which is in remarkable qualitative agreement with the VOF solution. The runtimes and speeds given in Table 3 are also favourable.

Note that this test case only touches upon the potential of the model for more sophisticated parameterisations of breaking than the simple approach presented in section 3.6.4. This should only be seen as a starting point, which already gives results similar to those of Boussinesq-style models for a range of classical test cases (section 4.2 and [66, 67]). The fact that the model can produce very accurate results (when compared to DNS) right up to the point of breaking opens up perspectives for parameterisations of breaking which go beyond what would be doable with Boussinesq-style models.

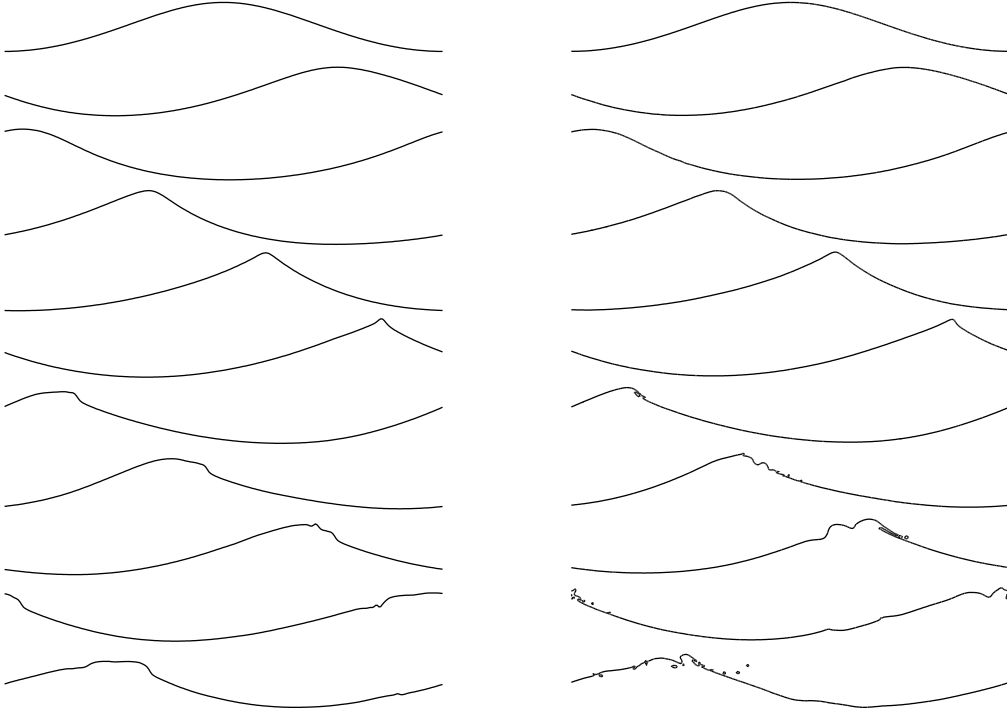


Figure 20. Free-surface evolution. Time increases from top to bottom in intervals of $T/4$ with T the Stokes wave period. Left column: layered model. Right column: Navier–Stokes/VOF (<http://basilisk.fr/src/test/stokes.c>).

Model	# Timesteps	Runtime	Speed	Remap
Multilayer	1489	69 332 400	6.4%	
Navier-Stokes/VOF	3079	215 97 200	n.a.	

Table 3. Runtimes (seconds) for the breaking Stokes wave case. The speeds are given in points \times layers \times timesteps/sec/core. Run on a single core. The “Remap” column gives the proportion of the total runtime taken by vertical remapping.

5 Conclusion and perspectives

The semi-discrete system (1), (2), (3) and (4) is a consistent discretisation of the incompressible Euler equations with topography and a free-surface, valid without assumptions on the slopes of the interfaces. Expressed as a set of conservation laws for each layer, the formulation has a clear physical interpretation and makes a seamless link between the hydrostatic Saint-Venant equations, dispersive Boussinesq-style models and the incompressible Euler equations.

The associated numerical scheme follows this consistent hierarchy and provides accurate and efficient solutions for all regimes. For surface gravity waves the same model can then be used to study metre-scale waves, even beyond breaking, with results closely matching those obtained using small-scale Euler/Navier–Stokes models, and coastal or global scale dispersive waves, with an accuracy and efficiency comparable to extended Boussinesq wave models.

This efficiency is obtained through the introduction of an “approximate vertical projection” which does not compromise conservation properties due to the decoupling provided by the vertically-Lagrangian discretisation. Accurate discrete dispersion relations are obtained thanks to the multigrid-accelerated, Hessenberg column-relaxation solution of a Keller box scheme vertical projection.

The method allows simple and robust wetting and drying on complex topography as demonstrated by adaptive solutions of the dispersive propagation and inundation during the Japanese Tohoku tsunami. Remarkably, for this type of application the introduction of a relaxed stability condition, based on the dispersive properties of non-hydrostatic waves, leads to shorter runtimes for the dispersive model than for the hydrostatic version.

This solid basis opens a number of perspectives. From a computational point of view, the performance of the vertical remapping should be improved. Since vertical remapping is comparable to one-dimensional advection (see Section 3.3), there is no reason for its computational cost to be three times as large, as is the case in our implementation using the PPR library. The model is already very efficient but this optimisation should further increase the overall performance.

The flexibility of vertical remapping should also be explored with the goal of developing “vertically adaptive” methods able to automatically track material interfaces, steep gradients and boundary layers [44].

The ability of the model to provide accurate solutions right up to the point of breaking allows a much deeper exploration of wave breaking parameterisations than what is done in Section 3.6.4, or than what is possible with Boussinesq-style models. A first step could be detailed comparisons with available experimental and DNS data [68, 64] of the profiles of velocity and associated energy dissipation before, during and after breaking.

Finally, as suggested in the introduction, an obvious extension is the inclusion of variable material properties, starting with density. We envisage to take into account both large density differences and Boussinesq-style small density variations. This would allow to efficiently model both multi-material flows, such as wind-driven waves or submarine

avalanches and the more classical temperature/salinity stratified flows typical for oceanic motion. This will be the topic of a follow-up publication.

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Appendix A Derivation of slope terms

The set of equations (1), (2), (3) and (4) can be derived using different methods: for example one could start from the incompressible Euler equations, written in Cartesian coordinates and use the general vertical coordinate transformation proposed by Kasahara (1974) [9], followed by formal vertical integration over the layer thicknesses (see also [12]).

Another approach is to apply conservation principles directly to thin vertical slices for each of the layers, a technique close to that used by Saint-Venant to derive the single-layer system. We will assume here that the hydrostatic equations (9) and (10) are given (see e.g. [7]) and will first deal with the non-hydrostatic pressure terms in the horizontal momentum equation (2).

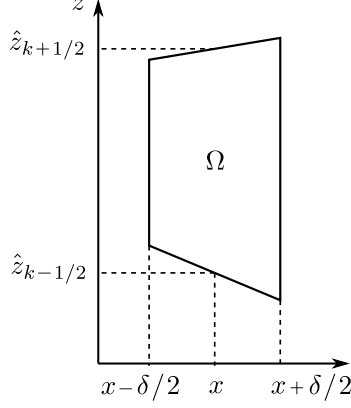


Figure 21. Thin vertical slice through a layer k .

If we consider a thin vertical slice Ω of width δ through a layer bounded vertically by $[\hat{z}_{k-1/2}, \hat{z}_{k+1/2}]$ (Figure 21), the variation in the total horizontal momentum of this slice due to pressure can be written

$$\partial_t \int_{\Omega} \rho u d\Omega = - \int_{\Omega} \partial_x p d\Omega = - \int_{\partial\Omega} p dz,$$

where $\partial\Omega$ denotes the boundary of Ω . This can be further developed as

$$\begin{aligned} \int_{\partial\Omega} p dz &= p_k^+ (\hat{z}_{k+1/2} - \hat{z}_{k-1/2})^+ - p_k^- (\hat{z}_{k+1/2} - \hat{z}_{k-1/2})^- \\ &\quad - p_{k+1/2} (\hat{z}^+ - \hat{z}^-)_{k+1/2} + p_{k-1/2} (\hat{z}^+ - \hat{z}^-)_{k-1/2} \\ &= p_k^+ h_k^+ - p_k^- h_k^- - [p (\hat{z}^+ - \hat{z}^-)]_k, \end{aligned}$$

where we have used the definitions of the layer thicknesses $h_k = \hat{z}_{k+1/2} - \hat{z}_{k-1/2}$ and of the vertical difference operator. The $+$ and $-$ superscripts denote average values on the $x \pm \delta/2$ vertical boundaries. For a vanishing δ , we have $f^+ - f^- = \delta \partial_x f$, where f is a differentiable function. This readily gives

$$\int_{\partial\Omega} p dz = \delta \partial_x (p h)_k - \delta [p \partial_x \hat{z}]_k$$

Since the integrated horizontal momentum can be written

$$\int_{\Omega} \rho u d\Omega = \delta \rho h_k u_k,$$

with u_k the layer-averaged horizontal velocity, we get

$$\begin{aligned} \partial_t \int_{\Omega} \rho u d\Omega &= - \int_{\partial\Omega} p dz, \\ \partial_t (\rho h_k u_k) &= - \partial_x (p h)_k + [p \partial_x \hat{z}]_k, \end{aligned}$$

which, for ρ constant and $\phi = p/\rho$, corresponds to the non-hydrostatic pressure terms in equation (2). Note that the corresponding derivation for the vertical momentum equation (3) is trivial since the $x \pm \delta/2$ boundaries are strictly vertical.

We next consider the continuity equation (4). Conservation of volume can be written

$$\begin{aligned}
0 &= \int_{\Omega} \nabla \cdot \mathbf{u} \, d\Omega = \int_{\partial\Omega} \mathbf{u} \cdot \mathbf{n} \, d\partial\Omega \\
&= u_k^+ (\hat{z}_{k+1/2} - \hat{z}_{k-1/2})^+ - u_k^- (\hat{z}_{k+1/2} - \hat{z}_{k-1/2})^- \\
&\quad - u_{k+1/2} \delta \partial_x \hat{z}_{k+1/2} + u_{k-1/2} \delta \partial_x \hat{z}_{k-1/2} \\
&\quad + w_{k+1/2} \delta - w_{k-1/2} \delta \\
&= \delta \partial_x (u h)_k - \delta [u \partial_x \hat{z}]_k + \delta [w]_k \\
0 &= \partial_x (h u)_k + [w - u \partial_x \hat{z}]_k,
\end{aligned}$$

which is the one-dimensional version of (4).

Appendix B Colocated horizontal pressure gradient

While the pressure gradient term in (29) is clearly defined (due to the vertically staggered discretisation), the discretisation of the horizontal pressure gradient in (28) must be described. The classical projection scheme uses the MAC or C-grid staggered velocity/pressure arrangement to avoid velocity/pressure decoupling. On the other hand, as discussed previously, the natural discretisation of the (hyperbolic) hydrostatic system of conservation laws relies on a colocated discretisation of the velocity components (i.e. Arakawa A- or B-grid). We solve this problem using colocated pressure and horizontal velocity/momentum components and an “approximate projection” method [69], which matches the computation of the hydrostatic pressure gradient in (14) and (16). The non-hydrostatic pressure gradient is first added to the staggered acceleration as

$$(h a)_{i+1/2,k} \leftarrow (h a)_{i+1/2,k} - \frac{h_{i+1,k}^{n+1} \phi_{i+1,k} - h_{i,k}^{n+1} \phi_{i,k}}{\Delta} + [\phi \partial_x \hat{z}]_{i,k},$$

the updated acceleration is then used to compute $(h u)_{i+1/2,k}^{n+1}$ and $(h u)_{i,k}^{n+1}$ according to (15) and (16).

Appendix C Discrete dispersion relations

The discrete dispersion relation is given by

$$\omega_n^2(\hat{k}) = g \frac{\sum_{i=1}^n a_{2i} \hat{k}^{2i}}{\sum_{i=0}^n b_{2i} \hat{k}^{2i}}$$

The following coefficients can be computed using the Maxima [70] script available at <http://basilisk.fr/src/test/dispersion.mac>

C.1 Keller box scheme

One layer:

$$\begin{aligned}
a_2 &= 4 h_0 \\
b_0 &= 4 \\
b_2 &= h_0^2
\end{aligned}$$

Two layers:

$$\begin{aligned}
 a_2 &= 16 (h_1 + h_0) \\
 a_4 &= 4 h_1 (h_0^2 + h_0 h_1) \\
 b_0 &= 16 \\
 b_2 &= 4 (h_1^2 + 4 h_0 h_1 + h_0^2) \\
 b_4 &= h_0^2 h_1^2
 \end{aligned}$$

Three layers:

$$\begin{aligned}
 a_2 &= 64 (h_2 + h_1 + h_0) \\
 a_4 &= (16 h_1 + 16 h_0) h_2^2 + (16 h_1^2 + 64 h_0 h_1 + 16 h_0^2) h_2 + 16 h_0 h_1^2 + 16 h_0^2 h_1 \\
 a_6 &= (4 h_0 h_1^2 + 4 h_0^2 h_1) h_2^2 + 4 h_0^2 h_1^2 h_2 \\
 b_0 &= 64 \\
 b_2 &= 16 h_2^2 + (64 h_1 + 64 h_0) h_2 + 16 h_1^2 + 64 h_0 h_1 + 16 h_0^2 \\
 b_4 &= (4 h_1^2 + 16 h_0 h_1 + 4 h_0^2) h_2^2 + (16 h_0 h_1^2 + 16 h_0^2 h_1) h_2 + 4 h_0^2 h_1^2 \\
 b_6 &= h_0^2 h_1^2 h_2^2
 \end{aligned}$$

Four layers:

$$\begin{aligned}
 a_2 &= 256 (h_3 + h_2 + h_1 + h_0) \\
 a_4 &= (64 h_2 + 64 h_1 + 64 h_0) h_3^2 + (64 h_2^2 + (256 h_1 + 256 h_0) h_2 + 64 h_1^2 + 256 h_0 h_1 + 64 h_0^2) h_3 + \\
 &\quad (64 h_1 + 64 h_0) h_2^2 + (64 h_1^2 + 256 h_0 h_1 + 64 h_0^2) h_2 + 64 h_0 h_1^2 + 64 h_0^2 h_1 \\
 a_6 &= ((16 h_1 + 16 h_0) h_2^2 + (16 h_1^2 + 64 h_0 h_1 + 16 h_0^2) h_2 + 16 h_0 h_1^2 + 16 h_0^2 h_1) h_3^2 + ((16 h_1^2 + \\
 &\quad 64 h_0 h_1 + 16 h_0^2) h_2^2 + (64 h_0 h_1^2 + 64 h_0^2 h_1) h_2 + 16 h_0^2 h_1^2) h_3 + (16 h_0 h_1^2 + 16 h_0^2 h_1) h_2^2 + \\
 &\quad 16 h_0^2 h_1^2 h_2 \\
 a_8 &= ((4 h_0 h_1^2 + 4 h_0^2 h_1) h_2^2 + 4 h_0^2 h_1^2 h_2) h_3^2 + 4 h_0^2 h_1^2 h_2^2 h_3 \\
 b_0 &= 256 \\
 b_2 &= 64 h_3^2 + (256 h_2 + 256 h_1 + 256 h_0) h_3 + 64 h_2^2 + (256 h_1 + 256 h_0) h_2 + 64 h_1^2 + 256 h_0 h_1 + \\
 &\quad 64 h_0^2 \\
 b_4 &= (16 h_2^2 + (64 h_1 + 64 h_0) h_2 + 16 h_1^2 + 64 h_0 h_1 + 16 h_0^2) h_3^2 + ((64 h_1 + 64 h_0) h_2^2 + (64 h_1^2 + \\
 &\quad 256 h_0 h_1 + 64 h_0^2) h_2 + 64 h_0 h_1^2 + 64 h_0^2 h_1) h_3 + (16 h_1^2 + 64 h_0 h_1 + 16 h_0^2) h_2^2 + (64 h_0 h_1^2 + \\
 &\quad 64 h_0^2 h_1) h_2 + 16 h_0^2 h_1^2 \\
 b_6 &= ((4 h_1^2 + 16 h_0 h_1 + 4 h_0^2) h_2^2 + (16 h_0 h_1^2 + 16 h_0^2 h_1) h_2 + 4 h_0^2 h_1^2) h_3^2 + ((16 h_0 h_1^2 + \\
 &\quad 16 h_0^2 h_1) h_2^2 + 16 h_0^2 h_1^2 h_2) h_3 + 4 h_0^2 h_1^2 h_2^2 \\
 b_8 &= h_0^2 h_1^2 h_2^2 h_3^2
 \end{aligned}$$

C.2 Staggered scheme (third-order)

One layer:

$$\begin{aligned}
 a_2 &= 8 h_0 \\
 b_0 &= 8 \\
 b_2 &= 3 h_0^2
 \end{aligned}$$

Two layers:

$$\begin{aligned}
 a_2 &= 8(h_1 + h_0) \\
 a_4 &= 8h_0^2 h_1 \\
 b_0 &= 8 \\
 b_2 &= 3(h_1^2 + h_0 h_1 + 3h_0^2) \\
 b_4 &= 3h_0^2 h_1^2
 \end{aligned}$$

Three layers:

$$\begin{aligned}
 a_2 &= 8(h_2 + h_1 + h_0) \\
 a_4 &= 8((h_1^2 + h_0 h_1 + h_0^2) h_2 + h_0^2 h_1) \\
 a_6 &= 8h_0^2 h_1^2 h_2 \\
 b_0 &= 8 \\
 b_2 &= 3h_2^2 + (3h_1 + 3h_0) h_2 + 9h_1^2 + 9h_0 h_1 + 8h_0^2 \\
 b_4 &= (3h_1^2 + 3h_0 h_1 + 3h_0^2) h_2^2 + 3h_0^2 h_1 h_2 + 9h_0^2 h_1^2 \\
 b_6 &= 3h_0^2 h_1^2 h_2^2
 \end{aligned}$$

Four layers:

$$\begin{aligned}
 a_2 &= 8(h_3 + h_2 + h_1 + h_0) \\
 a_4 &= 8((h_2^2 + (h_1 + h_0) h_2 + h_1^2 + h_0 h_1 + h_0^2) h_3 + (h_1^2 + h_0 h_1 + h_0^2) h_2 + h_0^2 h_1) \\
 a_6 &= 8(((h_1^2 + h_0 h_1 + h_0^2) h_2^2 + h_0^2 h_1 h_2 + h_0^2 h_1^2) h_3 + h_0^2 h_1^2 h_2) \\
 a_8 &= 8h_0^2 h_1^2 h_2^2 h_3 \\
 b_0 &= 8 \\
 b_2 &= 3h_3^2 + (3h_2 + 3h_1 + 3h_0) h_3 + 9h_2^2 + (9h_1 + 9h_0) h_2 + 8h_1^2 + 8h_0 h_1 + 8h_0^2 \\
 b_4 &= (3h_2^2 + (3h_1 + 3h_0) h_2 + 3h_1^2 + 3h_0 h_1 + 3h_0^2) h_3^2 + ((3h_1^2 + 3h_0 h_1 + 3h_0^2) h_2 + 3h_0^2 h_1) h_3 + (9h_1^2 + 9h_0 h_1 + 9h_0^2) h_2^2 + 9h_0^2 h_1 h_2 + 8h_0^2 h_1^2 \\
 b_6 &= ((3h_1^2 + 3h_0 h_1 + 3h_0^2) h_2^2 + 3h_0^2 h_1 h_2 + 3h_0^2 h_1^2) h_3^2 + 3h_0^2 h_1^2 h_2 h_3 + 9h_0^2 h_1^2 h_2^2 \\
 b_8 &= 3h_0^2 h_1^2 h_2^2 h_3^2
 \end{aligned}$$