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Compositional variations of chromian spinels from peridotites of the Spontang ophiolite complex, Ladakh, NW Himalayas, India: petrogenetic implications

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ABSTRACT: The Spontang ophiolite complex exposed along the Indus Tsangpo Suture Zone (ITSZ) represents a fragment of oceanic lithosphere emplaced after the closure of the Neo-Tethyan Ocean. The complex lying south of the ITSZ forms the highest tectonic thrust slice above the Mesozoic–Early Tertiary continental margin in the Ladakh-Zaskar Himalaya. The complex consists of a well-preserved ophiolite sequence dominated by peridotites, gabbros and ultramafic cumulates along with highly tectonized sheeted dykes and pillow lavas. The mantle suite of rocks is represented by minor lherzolites, harzburgites and dunites. Chromian spinel is brown to reddish, and its morphology and textural relationship with coexisting silicates varies with strain. Spinel occurs as blebs and vermicular exsolutions within orthopyroxene to spherical inclusions within olivine, characteristic of which is the elongate holly leaf shape. Chrome spinels are characterized by low TiO₂ and high Cr₂O₃ indicative of their depleted nature. Cr# [= atomic ratio Cr/(Cr + Al)] in the studied spinels are characterized by a small decrease in TiO₂ for a larger increase in Cr# consistent with observations for spinels aligned along the Luobusa trend of the Yarlung Zangpo Suture Zone (YZSZ) ophiolites. Variations in Cr-spinel Cr# and Mg# observed in the investigated peridotites may have resulted from a wide range of degrees of mantle melting during their evolution. Mineral and whole-rock chemistry of the Spontang peridotites is characterized by interaction between depleted magma and pre-existing oceanic lithosphere, typical of supra-subduction zone settings. The Spontang peridotites have olivine, clinopyroxene and orthopyroxene compositions similar to those from both abyssal and fore-arc peridotites and display spoon shaped REE profiles characteristic of interaction between LREE-enriched melt, derived from the subducting slab and LREE-depleted mantle residues. Equilibration temperatures calculated for the above rocks indicate that the studied samples represent typical mantle peridotites formed within the spinel stability field.

Key words: Cr-spinel, peridotite, partial melting, ophiolites, Spontang

1. INTRODUCTION

Remnants of a once very extensive ophiolite thrust sheet are observed to be locally preserved along the Indus Tsangpo Suture Zone (ITSZ) (Gansser, 1974; Frank et al., 1977; Srikantia and Razdan, 1980; Honegger et al., 1982; Deitrich et al., 1983; Robertson, 2000; Mahéo et al., 2004; Ahmad et al., 2008). Ophiolites along suture zones in general are observed to occur in two tectonic settings viz. within suture itself (e.g., Xigase ophiolite, south Tibet; Nicolas et al., 1981; Göpel et al., 1984) and those obducted onto a passive continental margin (e.g., Oman ophiolite; Lippard et al., 1986). In the Ladakh-Zaskar region, two groups of ophiolites and related ophiolite mélanges are recognized viz. the North and South Ladakh group of ophiolites (Maheo et al., 2004). The North Ladakh group is represented by the Dras arc, associated with the Ladakh batholith, localized north of the ITSZ. The Spontang, Nidar and Karzog ophiolites constitute the South Ladakh group. Further east, the Indus Suture is offset dextrally

by the Karakorum Fault but correlatives are found in Tibet as part of the Dazhuqu and Zedong terranes (Aitchison et al., 2000). Individually, these include ophiolites recorded at Jungbwa (Miller et al., 2003), Saga and Sangsang (Bédard et al., 2009), Xigaze (Nicolas et al., 1981; Girardeau et al., 1985a; Dubois-Cote et al.,

2005; Dupuis et al., 2005), Dazhuqu (Girardeau et al., 1985b; Xia et al., 2003), Zedong (McDermid et al., 2002; Aitchison et al., 2007) and Luobusa (Zhou et al., 1996). Correlatives to the west in Pakistan include Bela (Sarwar, 1992; Zaigham and Mallick, 2000), Muslim Bagh (Kakar et al., 2013) and Waziristan ophiolites (Jan et al., 1985).

The Spontang ophiolite exposed in the Zaskar mountains of Ladakh is one of the most complete fragments of the Neo Tethys ocean that is preserved along the ITSZ (Fuchs, 1982; Reuber, 1986; Searle, 1986). The Spontang ophiolite comprises the mantle section of lherzolite and harzburgite units along with lower crustal cumulates (Riebel and Reuber, 1982; Reuber, 1986; Reuber et al., 1992). Crustal sections are presented by gabbros, sheeted dykes and pillow lavas (Corfield and Searle, 2000). Detailed field mapping, combined with geochemical analysis carried out by Corfield et al. (2001) has helped define two major units: a full ophiolite sequence (Spontang ophiolite) overlain by an upper unit consisting of 1500-m-thick basalt-andesite volcanic and volcano-sedimentary rocks of island arc affinity (Spong arc). Previous studies carried out on the complex has focused on regional structure and timing of ophiolite obduction (Searle, 1986; Colchen et al., 1987; Corfield and Searle, 2000) along with stratigraphy of underlying passive margin sediments (Fuchrezs, 1982; Gaetani and Garzanti, 1991). Reconnaissance mapping of the high ground of the ophiolite itself was carried out by Reibel and Reuber (1982), followed by more detailed work on the ultramafic ophiolitic rocks (Reuber, 1986) and the ophiolitic mélanges (Reuber et al., 1992; Corfield et al., 1999). U-Pb zircon ages of a diorite yielding 177 ± 1 Ma, and an andesite of 88 ± 5 Ma age have been reported by Pedersen et al. (2001),

where the latter is associated with the Spong arc sequence overlying the ophiolite. Pedersen et al. (2001) interpret the older ages to represent the Neo-Tethyan ocean crust on top of which the Spong arc developed, while the andesitic age of ~88

Ma represents the minimum age of subduction initiation to form the Spong arc.

In this paper, we present for the first-time detailed mineral

chemistry and whole rock analyses of the peridotites of the Spontang ophiolite. In addition to the above, morphological and chemical variations of accessory chromian spinels along with Cr-spinel compositions have been used to provide insights into the genesis and tectonic origin of the studied peridotites.

2. GEOLOGICAL SETTING

2.1. Spontang Massif

The Spontang klippe lying 30 kms to the south of the ITSZ (Fuchs, 1979; Reuber, 1986; Searle, 1986) forms the highest level of the allochthon above the accretionary complex and the marine sediments of the Lamayuru complex. This area is situated between

4000 to 6000 mts elevation near the village Photoksar and is observed to be emplaced above the Zaskar shelf sediments along the passive continental margin of the Indian plate (Corfield et al., 1999; Corfield and Searle, 2000). The complex can be subdivided into two parts viz. a western mantle unit (Reuber et al., 1986a,

1986b; Corfield et al., 2001) and an eastern crustal unit (Corfield et

al., 1999, 2001; Pedersen et al., 2001). The ophiolite sequences tectonically overlie a volcano-sedimentary mélange of late Cretaceous to early Eocene age, which contain ophiolitic blocks, and a serpentinite base (Colchen and Reuber, 1986; Colchen et al., 1987) (Fig. 1). The complete assemblage is observed to be thrust over sedimentary series of the Zaskar unit of lower Eocene age (Colchen and Reuber, 1987; Kelemen et al., 1988). Corfield et al. (2001) were the first to recognize and describe the complete Penrose-style ophiolite suite at Spontang. Although the well-preserved ophiolite suite has been identified, the sequence has been significantly disrupted by faulting. In the present study, field studies have been carried out along 2 major traverses which expose the mantle (the Spong Valley, the Photang Valley) and crustal (the Marling chu and Bumiktse valleys) sections of the ophiolite.

Detailed field studies identify an upper lherzolitic unit containing clinopyroxene and a lower, more depleted harzburgitic unit (Reuber, 1986), well exposed at the river confluence in the Spong Valley near the upper portions of the Sirsir La peak (5550 m) (Fig. 2a). Boulders of peridotite are commonly observed showing presence of chromite veins (Figs. 2b and c). Gabbroic dykes are observed to intrude into the harzburgitic unit consistent with reports by Corfield et al. (2001) (Fig. 2d). Along with several dykes, lenses of dunites are also commonly observed within the peridotite units (Fig. 2e). Structurally, three deformation events as described by Reuber (1986) are observed to affect the peridotites exposed in the Spong valley. The S_1 deformation is seen as a foliation that is ubiquitous in the peridotites, although very weak and often indiscernible in the field. Several sub-vertical NE-SW trending sinistral shear zones define the S_2 deformation. The third deformation event is recorded involving a major sub-horizontal shear zone of 20–70 m width, superposing the upper clinopyroxene rich lherzolitic layer from the lower more depleted harzburgitic unit.

At the base of the massive Photang Kangri peak along the

Fig. 1. (a) Regional geological map of the Himalayas showing location of the ophiolite sequences (after Hébart et al., 2012). The study area has been demarcated by a box. (b) Detailed geological map of the Spontang ophiolite showing the major tectonic units (after Corfield et al., 2001). Stars indicate sample locations of the ultramafic samples.

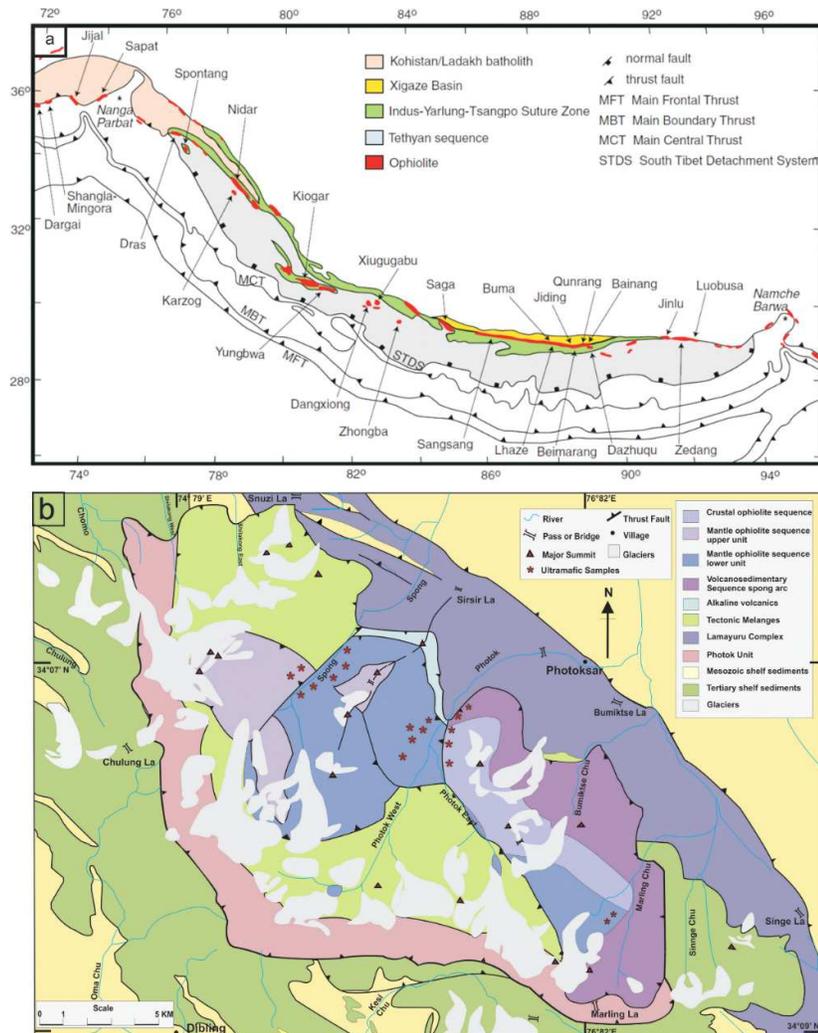
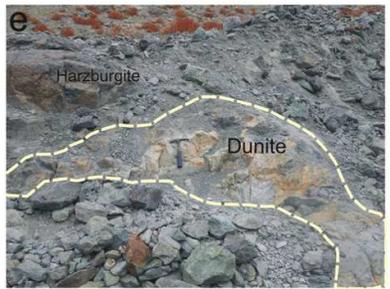


Fig. 2. Field photographs of the Spontang ophiolite: (a) Peridotites exposed along the near the river confluence in the upper Spong Valley. (b) Peridotites occurring as boulders. (c) Note the presence the chromite bands in the peridotites. (d) Gabbroic dykes intruding peridotites exposed in the Spong valley. (e) Dunite bodies are observed intercalated within the harzburgite units. (f) Pillow basalts forming boulder out-crops exposed in the Phu/Photang valley. (g) Tectonized fine grained gabbro outcrops observed in the Bumikste valley. (h) Volcano-clastic chert outcrops exposed in the Bumikste valley.



eastern flank of the Photang valley, lower crustal cumulates

2f). Medium grained basaltic dykes are observed

has reported rare plagiogranites located on the west side of the mid-Photang Valley at a higher elevation, however attempts to locate this body were unsuccessful.

The Marling chu valley and the Bumiktse valley primarily expose lower crustal sequences viz. weathered cumulates of wehrlites and pyroxenites with associated gabbros. The gabbros in the Bumiktse valley exposed below the snout of the glacier are highly tectonized and are thrust directly onto the volcano-sedimentary sequences (Fig. 2g). According to Corfield et al. (2001), a complete sequence through the Moho is preserved in the higher levels between the Bumiktse and Marling valleys forming a distinctive horizon separating the upper crustal cumulate sequence from the lower mantle sequence. Field examination at the above-mentioned location however did not reveal any direct contact between the two units. Exposures in the Marling chu and Bhumiktse valleys revealed only crustal sections continuing across the inaccessible vertical cliffs and glaciers of the Photang Kangri.

2.2. Spong Arc

The volcano-sedimentary sequence of the Spong Island Arc comprising pillow lavas and interbedded volcanoclastic deposits are exposed at the northern edge of the Photang Kangri and in the Bumiktse and Marling chu valleys. No clear defined contact is observed between the arc deposits and the ophiolite sequence however, a transition can be observed based on the onset of volcanoclastic deposits in the field. The sequence comprises several 100 meters thick unit, consisting of angular and poorly sorted clasts, ranging widely in composition with occasional chert is exposed at Bumiktse valley (Fig. 2h). Previous studies have mapped the Spong volcanoclastic rocks as the Dras Volcanics (Srikantia and Razdan, 1981; Fuchs, 1982; Keleman and Sonnenfeld, 1983; Searle et al., 1988) however, detailed studies by Reuber et al. (1992); Corfield et al. (2001), have implied that the Spong volcanic sequence tectonically overlies the Spontang ophiolite and could be related to an immature island arc forming above the oceanic substrate of the Spontang ophiolite.

including wehrlites, pyroxenites, gabbros and pillow lavas are exposed (Fig. exhibiting a gradational contact with the pillow lavas. Corfield et al. (2001)

3. PETROGRAPHY

Detailed petrographic study for the peridotite samples have been carried out and are represented by minor lherzolites, harzburgites and dunites. Textural studies of the studied peridotite rocks were carried out after Mercier and Nicolas (1975) and major textural classification is based on inclusions of accessory spinels in porphyroclasts of major minerals. Modal analysis of the studied peridotites was carried out and is depicted in Figure 3 and provided in Table 1.

Lherzolite samples show textural transition from porphyroclastic

to equigranular mosaic texture with more than 5 vol% clinopyroxene (after Mercier and Nicolas, 1975) and display feeble alteration in the form of presence of serpentine veins and cracks. Olivine occurs as porphyroclasts (0.3 to 1.5 mm) set within groundmass of subhedral to anhedral grains as mosaic crystals with curved boundaries. Both generations show curvilinear, serrated boundaries, recrystallization, and presence of kink bands (Fig. 4a). Clinopyroxene may occur as porphyroblasts or as isolated grains (Fig. 4b) within the olivine matrix with exsolution lamellae of orthopyroxenes.

Fig. 3. Modal classification of ultramafic rocks based on the proportions of olivine (Ol), orthopyroxene (Opx), clinopyroxene (Cpx) (after Streckeisen, 1973).

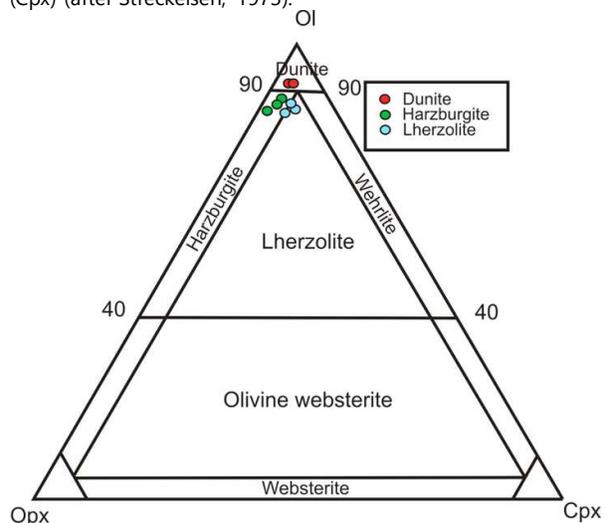


Table 1. Summary of sample details and modal abundances of the peridotites from the Spontang ophiolite complex

Sr.No	Sample Number	Rock Type	Latitude	Longitude	Altitude (m)	Olivine	Orthopyroxene	Clinopyroxene	Spinel	Total
1	B2	Lherzolite	34°06'756"	76°46'512"	4512	76.8	12.9	8.5	1.7	99.9
2	B16	Harzburgite	34°07'412"	76°46'812"	4455	79.7	13.8	3.8	2.5	99.8
3	E1	Harzburgite	34°06'591"	76°46'645"	4526	78.9	13.08	4.26	2.9	99.1
4	E4	Harzburgite	34°06'190"	76°46'044"	4636	80.13	16.41	2.60	0.86	100.0
5	E5	Dunite	34°06'143"	76°45'994"	4640	82.2	14.5	0.75	2.5	99.95
6	E7	Harzburgite	34°06'143"	76°45'994"	4642	83.8	12.3	1.8	2.1	100.0
7	I7	Harzburgite	34°06'210"	34°06'210"	4607	80.2	15.5	3.4	0.8	99.99
8	SP3	Dunite	34°06'410"	76°46'766"	4654	88.2	9.4	1.0	1.3	99.9

Fig. 4. Photomicrographs of mantle peridotites of the Spontang ophiolite: (a) Olivine porphyroblasts showing undulose extinction, curvi-linear and serrated boundaries, recrystallization, and presence of kink bands (BXN). (b) Harzburgite showing porphyroclastic textures. Note the presence of pale green colored clinopyroxene (diopside) exhibiting sutured boundaries (PPL). (c) Orthopyroxene grains showing presence of exsolution lamellae set within an olivine matrix (BXN). (d) Porphyroclasts of orthopyroxene showing exsolution lamellae of spinel (PPL). (e) Dark brown xenomorphic spinel showing characteristic vermicular shape with lobate boundaries in harzburgite (PPL). (f) Holly leaf spinel with opaque boundaries in harzburgite (PPL). (g) Subhedral spinel occurring as inclusions within orthopyroxene porphyroclasts (PPL). (h) euhedral spinel seen in dunite (PPL). (i) X-ray maps for the spinel grains viz. Cr, Al.

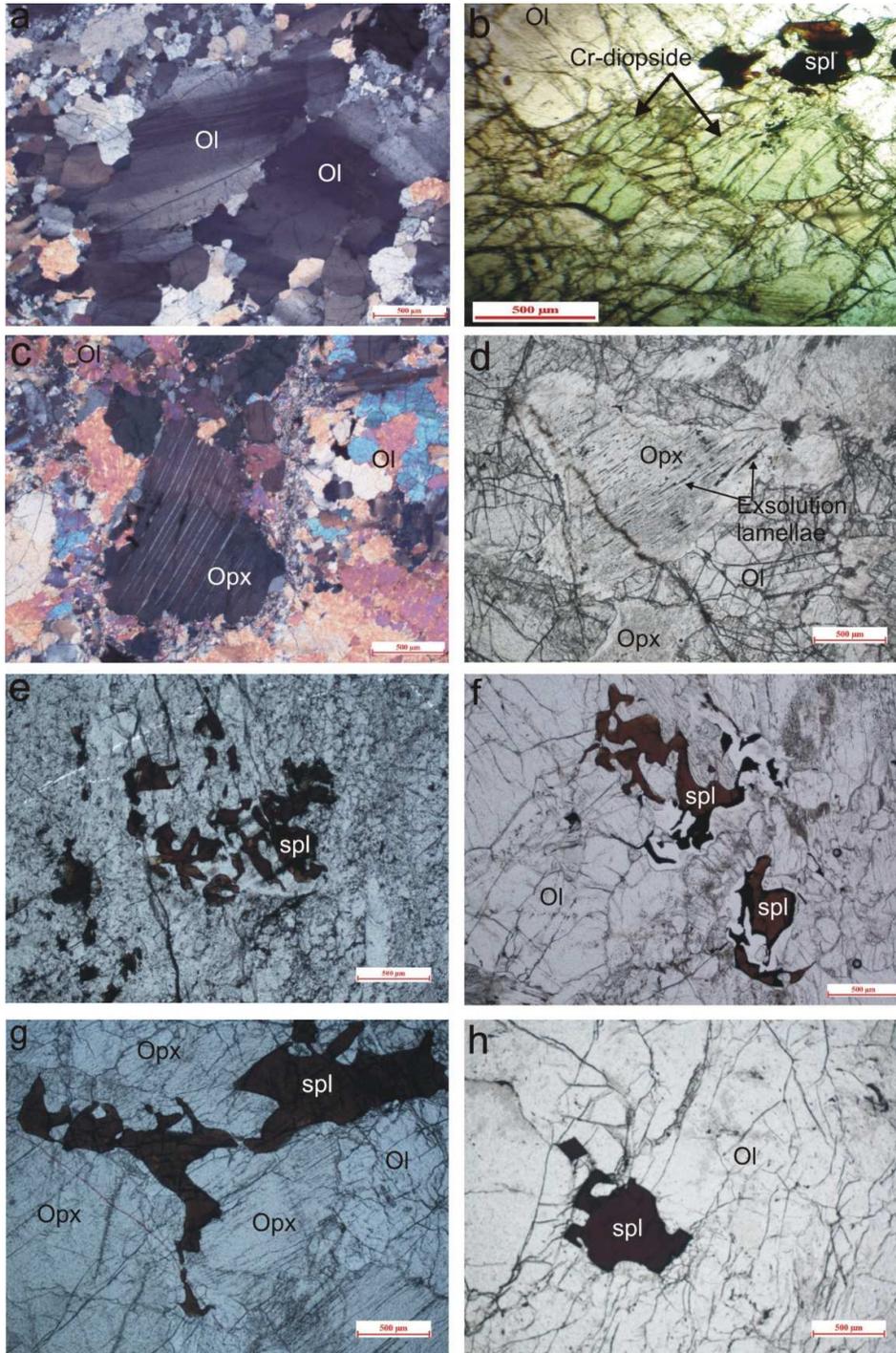
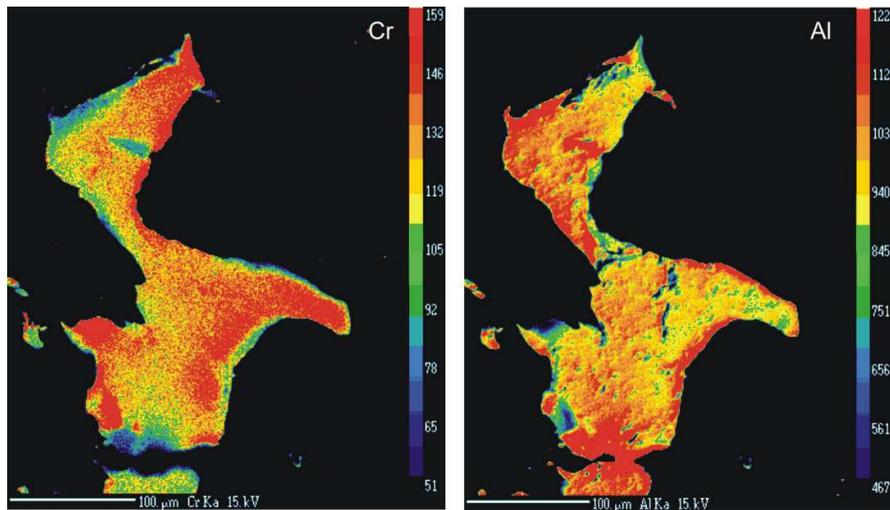


Fig. 4. (continued).



Exsolution lamellae of clinopyroxene within the orthopyroxenes are observed at places (Fig. 4c).

Harzburgite samples show protogranular to porphyroclastic textures. Their primary modal mineralogy includes olivine, orthopyroxenes, minor relicts of clinopyroxene and accessory spinel. Olivine grains are coarse, exhibit undulose extinction surrounded by small grained olivine crystals and display curvilinear and serrated boundaries meeting at triple junctions at 120° . Zoning, kink bands and recrystallization are commonly observed. Orthopyroxenes are colorless, usually occur as large prismatic crystals, 0.2 to 3.5 mm in size with curvilinear boundaries and often display diopside exsolution lamellae, kink bands and gliding (Fig. 4d). Minor clinopyroxenes are small, 0.1 to 0.3 mm in size and display pale green color in plane polarized light exhibiting sutured boundaries.

Dunites are fresh unaltered with granular texture. Olivine grains are large, porphyroclastic, elongated and display undulose extinction and intense zoning. They range in size from 0.9 to 4.00 mm and are surrounded by smaller subhedral grains. Matrix of olivines are subhedral to anhedral and 0.5 to 2 mm in size, display polygonization and serrated boundaries. Euhedral spinels are commonly observed along olivine borders and as inclusions within olivines.

Spinels are light brown picotites to dark brown colored chromium

rich and constitute up to 0.5 vol% to 3 vol% of the studied rocks. Grain sizes vary within samples, but most are between 0.2 mm and 1 mm. Morphologically spinel is observed to exhibit various shapes in relationship with coexisting silicates and these shapes

vary with the levels of strain. Spinel in the harzburgite samples occur as xenomorphic, vermicular intergrowths with lobate boundaries (Fig. 4e). The characteristic holly leaf shape is also observed in these samples (Fig. 4f). Intergrowths of orthopyroxene and spinel can be sporadically seen in these samples. At other places, subhedral spinel occurs as inclusions within orthopyroxene porphyroclasts (Fig. 4g). Spinels in dunites are mainly anhedral and occur at olivine boundaries and are rarely seen as inclusions within olivines. Euhedral idiomorphic spinels are occasionally seen to occur as inclusions within olivine grains (Fig. 4h). Under the microscope, few grains show irregular zoning due to replacement by an opaque phase along grain boundaries (Fig. 4f). This zoning is apparent in back-scattered electron images (BSE), in which the edges of the zoned Cr-spinel grains appear to be bright and separated from the inner part of the grains by sharp contacts. Mapping of a single Cr-spinel grain shows Al-Cr zoning with distinct Al rich-Cr poor and Al poor-Cr rich regions (Fig. 4i). The observed zoning in the spinel grain is interpreted as a possible result of deformation explained by stress-directed lattice diffusion of Al and Cr (Ozawa, 1989).

4. ANALYTICAL METHODS

Representative mineral analyzes of spinels, pyroxenes and olivines of the investigated peridotites (harzburgites, lherzolites and dunites) of the Spontang ophiolite are given in Table 2. A Cameca SX-50 microprobe at Department of Geology, Institute of Science, Banaras Hindu University was employed to obtain

Table 2. Representative electron microanalysis (wt%) and atomic proportions of Cr-spinels, olivines and pyroxenes from peridotites of the Spontang ophiolite complex

Mineral Sample	Spinel							
	B2	B16	E1	E4	E5	E7	I7	SP3
Rock Type	Lherzolite n = 7	Harzburgite n = 5	Harzburgite n = 15	Harzburgite n = 16	Dunite n = 3	Harzburgite n = 11	Harzburgite n = 3	Dunite n = 10
SiO ₂	2.47	0.01	-	-	-	0.03	0.12	0.17
TiO ₂	0.09	0.07	0.07	0.11	0.05	0.02	0.13	0.11
Al ₂ O ₃	27.12	39.62	42.34	40.40	22.63	45.76	29.90	29.76
Cr ₂ O ₃	40.07	27.68	23.92	26.20	42.76	21.90	38.82	34.78
FeO	15.98	15.77	14.63	13.78	19.56	13.80	15.58	18.49
MgO	11.96	15.19	16.30	16.56	11.95	16.92	14.56	13.35
MnO	0.32	0.14	0.27	0.26	0.42	0.12	0.23	0.24
NiO	0.14	-	0.16	0.15	0.29	-	-	0.34
CaO	0.16	0.03	-	-	0.06	0.03	0.01	0.13
SrO	1.63	-	0.89	1.04	1.72	-	-	1.44
Total	99.88	98.69	99.49	99.52	99.44	98.72	99.49	99.01
Ti	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	1.01	1.34	1.42	1.36	0.84	1.49	1.05	1.03
Cr	1.00	0.63	0.54	0.59	1.06	0.49	0.91	0.87
Fe ³⁺	0.07	0.03	0.05	0.05	0.09	0.03	0.04	0.09
Fe ²⁺	0.43	0.35	0.30	0.29	0.42	0.31	0.35	0.40
Mn	0.01	0.00	0.01	0.01	0.01	0.69	0.01	0.01
Mg	0.56	0.65	0.69	0.71	0.56	0.00	0.64	0.58
Total	3.00	3.00	3.00	3.00	3.00	3.00	3.00	3.00
Cr#	0.50	0.62	0.27	0.30	0.53	0.24	0.46	0.44
Mg#	0.57	0.34	0.69	0.71	0.57	0.69	0.65	0.59
Mineral Sample	Olivine							
	B2	B16	E1	E4	E7	I7	SP3	
Rock Type	Lherzolite n = 7	Harzburgite n = 3	Harzburgite n = 11	Harzburgite n = 11	Harzburgite n = 11	Harzburgite n = 3	Dunite n = 21	
SiO ₂	42.61	41.45	45.74	41.41	47.00	43.35	44.11	
FeO	8.07	8.78	7.94	9.06	8.03	8.21	8.88	
MgO	47.61	48.96	44.05	48.66	42.84	47.28	46.17	
MnO	0.19	0.09	0.17	0.11	0.10	0.10	0.10	
NiO	0.39	-	0.21	0.26	-	-	0.31	
Cr ₂ O ₃	0.19	0.15	0.24	0.14	0.34	0.14	0.17	
CaO	-	0.03	-	-	0.21	0.10	-	
Total	99.64	99.53	99.47	100.12	99.45	99.50	100.02	
Si	1.06	1.02	1.15	1.02	1.19	1.07	1.10	
Fe ³⁺	0.02	-	0.02	0.04	-	-	0.02	
Fe ²⁺	0.15	0.18	0.15	0.15	0.17	0.17	0.16	
Mn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
Mg	1.75	1.79	1.64	1.78	1.60	1.74	1.70	
Ca	0.01	0.00	-	-	0.01	0.00	0.00	
%Fo	90.61	90.73	90.62	90.40	89.96	90.82	90.09	
%Fa	8.65	9.13	9.18	9.49	9.52	8.89	9.73	
Ca-Ol	0.53	0.04	-	-	0.40	0.17	0.07	

mineral chemistry data of the studied rocks. The accelerating voltage and beam current were 15 kV and 10 nA respectively. Polished thin sections were coated with 20 nm thin layer of carbon for electron probe micro analyses using LEICA-EM ACE200 instrument. Natural (fluorite, albite, halite, periclase, peridotite, corundum, wollastonite, apatite pyrite, orthoclase, rutile, chromite,

Table 2. (continued)

Mineral Sample	Clinopyroxene					Orthopyroxene			
	B2	B16	E7	I7	SP3	B16	E7	I7	SP3
Rock Type	Lherzolite n = 14	Harzburgite n = 5	Harzburgite n = 9	Harzburgite n = 10	Dunite n = 10	Harzburgite n = 1	Harzburgite n = 2	Harzburgite n = 11	Dunite n = 7
SiO ₂	53.14	52.81	52.88	53.04	52.74	56.48	56.21	55.90	56.12
TiO ₂	0.11	0.11	0.12	0.13	0.01	0.04	0.03	0.03	1.24
Al ₂ O ₃	1.52	3.03	2.90	3.12	5.27	2.35	2.67	3.11	0.56
Cr ₂ O ₃	0.65	1.03	0.78	1.27	0.96	0.56	0.71	0.91	0.46
FeO	1.67	2.25	2.04	2.03	3.68	6.07	6.22	6.09	6.32
MnO	0.05	0.05	0.06	0.08	0.06	0.12	0.15	0.15	2.22
MgO	17.44	16.62	16.55	16.40	23.90	33.14	32.92	32.61	19.86
CaO	24.09	23.09	23.91	23.25	12.37	0.57	0.68	0.83	13.12
K ₂ O	0.04	0.00	0.00	0.01	0.00	–	0.01	0.01	0.01
Na ₂ O	0.18	0.27	0.13	0.42	0.33	0.01	0.01	0.01	0.07
NiO	0.07	–	–	–	0.06	–	–	–	0.12
Total	99.38	99.33	99.44	99.81	99.47	99.34	99.64	99.68	100.24
Si	1.95	1.93	1.93	1.93	1.87	1.97	1.95	1.94	1.94
Ti	–	–	–	–	–	–	–	–	– Al
Cr	0.07	0.13	0.13	0.13	0.24	0.10	0.11	0.13	0.07
Cr	0.02	0.03	0.02	0.04	0.03	0.02	0.02	0.02	0.01
Fe ³⁺	0.03	0.00	–	–	0.09	–	–	–	0.03
Fe ²⁺	0.03	0.07	0.06	0.06	0.12	0.18	0.18	0.18	0.16
Mn	–	0.00	–	–	–	–	–	–	–
Mg	0.95	0.91	0.90	0.89	1.26	1.72	1.71	1.69	1.70
Ca	0.95	0.91	0.94	0.91	0.47	0.02	0.03	0.03	0.11
Na	0.01	0.02	0.01	0.03	0.03	–	–	–	0.01
Total	4.00	4.00	4.00	4.00	4.00	4.00	4.00	4.00	4.00
Wo	0.49	0.48	0.49	0.48	0.70	0.01	0.02	0.02	0.02
En	0.50	0.48	0.47	0.47	0.07	0.90	0.89	0.89	0.90
Fs	0.01	0.04	0.03	0.33	0.26	0.09	0.09	0.09	0.08
Mg#	0.98	0.92	0.94	0.94	0.94	0.91	0.91	0.91	0.92

rhodonite, hematite, celestine, zircon and barite) and synthetic standards (Ni & Nb and synthetic glass standard YAG supplied by CAMECA-AMETEK) were used for calibration of the machine. Data processing was carried out using SxSAB version 6.1 and SX-Results softwares of CAMECA. Major oxides were determined by X-ray fluorescence spectroscopy (XRF) on pressed pellets at the Department of Geology, Savitribai Phule Pune University. The analytical uncertainty is estimated to be $\pm 1\%$ for all major oxides. Trace and Rare Earth Elements (REE's) were analyzed following detailed sample preparation and analytical procedures described by Barrat et al. (1996, 2012) and Rospabé et al. (2017) at GET-OMP, Toulouse, France. About 0.200 g of sample powder was accurately weighed and transferred to 22 ml Savillex PFA beakers. An amount (between 1.5–15 ng) of Tm (in solution) was added to each sample, except to the blank and to the reference material BHVO-2. Sample digestion procedure was adopted following method C of Yokoyama et al. (1999). After total dissolution,

an aliquot of the final solution was taken and diluted ($F = 10000$) for ICP-MS measurements on the ELEMENT XR, Thermo-Scientific and proceed following the direct measurement protocol from Rospabé et al. (2017). International geostandard UBN and blanks were regularly run during the set (see Table 3). The blanks were subtracted to the sample and standard signals before processing the data according to Rospabé et al. (2017).

5. RESULTS

5.1. Whole Rock Data

Whole-rock major elements were carried out on 6 peridotite samples and are given in Table 3. Loss on ignition (LOI) values are $< 1\%$ indicating the freshness of the samples. The Mg# varies from 92 to 93 similar to that of modern oceanic depleted and residual peridotites (Bonatti and Michal, 1989; Bodinier and

Table 3. Representative whole-rock, trace and REE compositions of ultramafic rocks of the Spontang ophiolite complex

Samples	B2	B16	E1	E4	E5	E7	L13	L14
Rock Type	Lherzolite	Harzburgite	Harzburgite	Harzburgite	Dunite	Harzburgite	UBN	UBN
SiO ₂	45.06	46.04	44.78	45.09	44.4	44.57		
TiO ₂	nd	0.02	nd	0.02	nd	nd		
Al ₂ O ₃	1.26	1.28	1.41	1.49	0.66	1.18		
MnO	0.14	0.14	0.14	0.14	0.14	0.15		
Fe ₂ O ₃	9.9	9.48	9.88	9.81	9.76	10.08		
CaO	1.69	1.19	1.2	2.08	0.77	1.15		
MgO	40.8	40.88	41.51	39.66	43.21	41.52		
Cr ₂ O ₃	0.47	0.43	0.38	0.46	0.44	0.33		
NiO	0.34	0.32	0.34	0.33	0.36	0.34		
LOI	0.35	0.23	0.35	0.93	0.26	0.67		
TOTAL	100.1	100	99.99	100.01	99.99	99.98		
Rb	0.020	0.024	0.021	0.011	0.062	0.017	3.34	3.48
Sr	0.122	0.173	0.203	0.361	2.541	0.215	4.0	4.1
Y	0.180	0.368	0.272	0.696	0.184	0.395	2.65	2.76
Zr	0.023	0.037	0.040	0.049	0.494	0.068	3.3	3.5
Nb	0.008	0.013	0.013	0.023	0.125	0.019	0.04	0.04
Mo	0.025	0.022	0.023	0.012	0.307	0.022	0.33	0.33
Ba	0.178	0.282	0.257	0.217	0.788	0.203	27.2	28.4
La	0.003	0.015	0.0048	0.006	0.363	0.007	0.33	0.35
Ce	0.008	0.006	0.0116	0.014	0.692	0.016	0.85	0.89
Pr	0.001	0.001	0.0011	0.002	0.065	0.002	0.12	0.13
Nd	0.003	0.004	0.0054	0.010	0.207	0.009	0.60	0.64
Sm	0.002	0.005	0.0035	0.009	0.030	0.005	0.22	0.23
Eu	0.001	0.002	0.0014	0.005	0.006	0.003	0.088	0.092
Gd	0.006	0.017	0.0109	0.036	0.030	0.017	0.280	0.298
Tb	0.002	0.004	0.0034	0.009	0.004	0.006	0.060	0.063
Dy	0.019	0.040	0.0308	0.093	0.028	0.050	0.443	0.465
Ho	0.006	0.014	0.0094	0.025	0.007	0.014	0.098	0.105
Er	0.024	0.052	0.0415	0.089	0.024	0.055	0.297	0.312
Yb	0.039	0.074	0.0548	0.112	0.034	0.074	0.305	0.324
Lu	0.007	0.013	0.0100	0.020	0.006	0.012	0.046	0.050
Hf	0.002	0.004	0.0023	0.008	0.018	0.006	0.156	0.171
Ta	nd	0.000	nd	0.008	0.033	0.001	0.015	0.016
Pb	0.026	0.010	0.016	0.012	0.286	0.017	16.6	16.9
Th	0.001	0.001	0.0010	0.001	0.125	0.002	0.070	0.072
U	0.003	0.002	0.0040	0.003	0.011	0.003	0.056	0.061
Sc	9.61	12.03	11.98	3.87	7.41	16.66	12.5	13.0
Ti	41.59	128.02	59.89	128.92	56.52	88.82	582	603
V	33.98	49.17	44.21	62.07	23.19	38.25	62	65
Cr	1,750.69	2,184.52	2,228.96	2,292.79	1,529.09	1,226.96	2516	2733
Mn	851.75	1,004.95	968.75	1,007.97	931.31	891.33	936	983
Co	103.44	125.64	121.31	128.45	125.87	109.40	98	102
Ni	2,076.25	2,352.54	2,254.92	2,304.09	2,470.31	2,121.24	2026	2109
Cu	9.26	8.98	12.16	22.48	4.96	8.06	23.7	23.7
Zn	36.34	51.40	44.54	49.10	40.77	40.91	82	86
Ga	0.59	1.04	0.88	1.36	0.46	0.92	2.32	2.38

Godard, 2007). All peridotites are strongly depleted in major elements with SiO₂ < 45 wt%, TiO₂ < 0.02 wt%, CaO < 2.08 wt%

and enriched in transition elements viz. Ni in the range of 2076 to 2470 ppm and Cr in the range of 1126 to 2292 ppm. MgO

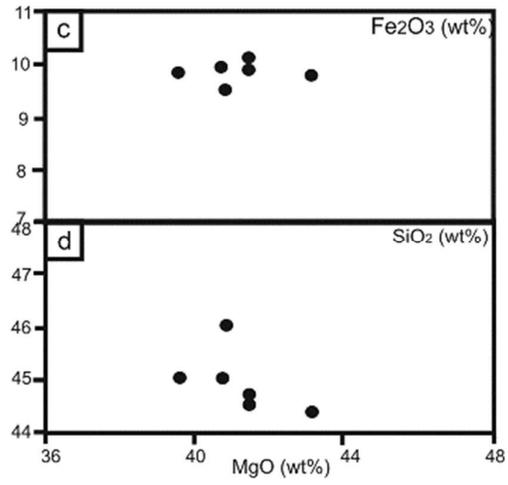
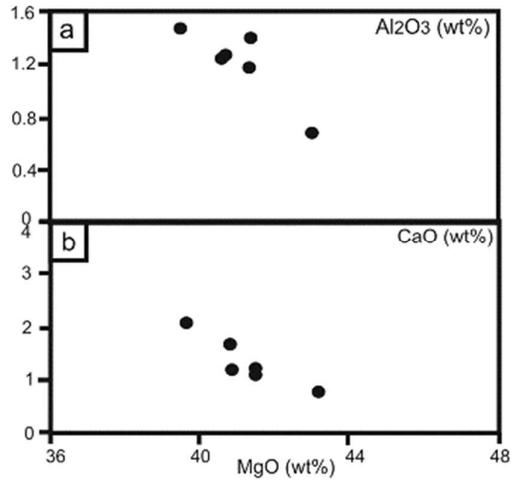


Fig. 5. Primary bulk composition of peridotites of the Spontang ophiolite complex as a function of MgO (a) Al_2O_3 , (b) CaO, (c) Fe_2O_3 , and (d) SiO_2 .

contents increase systematically from lherzolite to harzburgite and dunite considered as an index of melt depletion (Parkinson and Pearce, 1998). Al_2O_3 and CaO contents range from 0.66 wt% to 1.49 wt% and 0.77 wt% to 2.08 wt% respectively and show an inverse correlation when plotted against MgO (Fig. 5). In addition to this, bulk SiO_2 shows a negative correlation against MgO and is consistent with trends displayed by abyssal peridotites (Baker and Beckett, 1999). FeO-MgO trends are essentially horizontal and are probably attributed to abyssal peridotites having experienced extensive olivine addition by upwelling mantle melts (Kelemen et al., 1997; Niu et al., 1997; Baker and Beckett, 1999). Chondrite normalized rare earth element patterns are illustrated in Figure 6 and are characterized by extremely low REE concentrations

ranging from 0.01 to 0.12 times the chondritic value. The studied samples show smooth LREE depleted patterns to positive MREE to HREE slopes indicating high degree of partial melting. The Spontang peridotites have spoon-shaped REE profiles; characteristic of interaction between tholeiitic or calc-alkaline melts and REE- depleted mantle residues. It is important to note that the dunite displays the most LREE enriched profile, that is characteristic of reactive melting (Dick, 1977a; Quick, 1981; Edwards, 1990; Kelemen, 1990; Kelemen et al., 1992; Zhou et al., 2005) and typically found in dunites sitting at mantle transition zone levels (Godard et al., 2000; Koga et al., 2001; Rospabé et al., 2018). Lherzolite sample shows the most depleted MREE and HREE patterns as compared to the harzburgite samples. The absence of Eu anomaly indicates

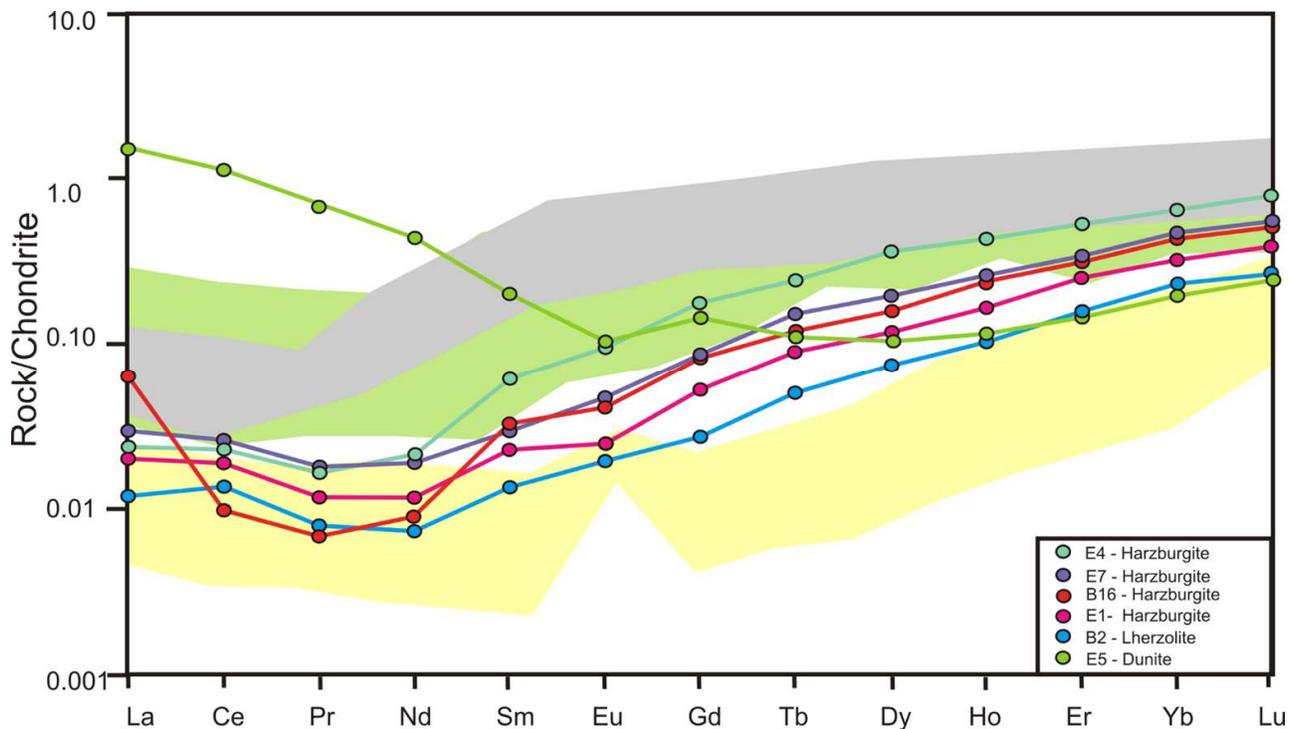


Fig. 6. Chondrite normalized REE patterns for the peridotites of the Spontang ophiolite complex. Normalizing values are after Sun and McDonough, (1989). Grey band represents composition of YSZS peridotites and are taken after Dupuis et al. (2005). Yellow and green bands represent compositions of fore-arc (after Parkinson and Pearce, 1998) and abyssal peridotites, respectively (after Singh, 2013).

absence of plagioclase accumulation in the source.

5.2. Spinel

Analyzed spinels in the studied peridotites are fresh, unaltered with low SiO₂ and TiO₂ contents and high Cr₂O₃. Dunite samples contain Cr-spinels with high Cr# values ranging between 0.31 to as high as 0.73. Cr₂O₃ and Al₂O₃ values vary between 28 wt% to 51 wt% and 9 wt% to 39 wt%. FeO contents vary between 18 wt% to 20 wt% with low TiO₂ contents (0.05 wt% to 0.11 wt%). The harzburgites host Cr-spinels with lower Cr# (0.21 to 0.54) and Mg# (0.42 to 0.70) relative to dunites and are classified as spinels with one sample falling on the spinel-hercynite solid solution boundary (Fig. 7a). Cr₂O₃ ranges between 19 wt% and 43 wt%, Al₂O₃ between 24 wt% and 48 wt%. FeO contents are lower than dunites and range between 14 wt% and 16 wt%. TiO₂ ranges between 0.02 wt% and 0.13 wt%. Lherzolites display Cr-spinels

with the highest Mg# (0.70 to 0.79) and lowest Cr# (0.10 to 0.34), indicating a moderately fertile character (Fig. 7a). Overall it can be said that the Cr# values of Cr-spinel increase and Mg# values decrease simultaneously from lherzolite to harzburgite and dunite and plot along a trend, defined by a large increase in Cr# for a relatively small decrease in Mg#. All chromian spinels have Cr# and Mg# similar to those of mantle peridotites and plot in the mantle array field on the Cr₂O₃ vs. Al₂O₃ diagram (Fig. 7b). Samples when plotted in the Cr-Al-Fe³⁺ triangular diagram (Fig. 7c) of Jan and Windley (1990) fall in the residual peridotite or ophiolite fields consistent with the low TiO₂ contents (0.1 wt% to 0.3 wt%) of the analyzed Cr-spinels.

5.3. Olivine and Pyroxenes

Analyzed olivines are fairly homogenous in composition and do not show any distinct zoning from core to rim. No major

Fig. 7. (a) Cr spinel compositions from the Spontang ophiolite complex, in terms of Cr# [Cr/(Cr + Al)] versus Mg# [Mg/(Mg + Fe²⁺)] overlain on the classification of the composition of Cr-spinel and ferrian chromite in terms of Cr# versus Mg# diagram. Data for spinel in modern abyssal peridotites is from Dick and Bullen (1984) and Juteau et al. (1990). Field for spinel in boninites is taken from Dick and Bullen (1984). Data for spinel in fore-arc peridotites are from Ishii et al. (1992) and Ohara and Ishii (1998). Cr-spinel composition is also contoured at a nominal temperature of 1200 °C for olivine compositions from Fo₉₀ to Fo₉₆ (quantitatively computed by Dick and Bullen, 1984). Arrows with ticks represent the percentage of melting of the host peridotite after Hirose and Kawamoto, (1995). Black, pink and orange samples are from the Ker-guelen harzburgites (after Gregoire et al., 1997), South Sandwich peridotites (after Pearce et al., 2000) and Mariana Trough peridotites (after Ohara et al., 2002) respectively. (b) Cr₂O₃ versus Al₂O₃ in the spinels to determine the spinel origin. All samples plot in the mantle array field. The fields are from Conrad and Kay (1984), Haggerty (1989) and Kepezhinskas et al. (1995). (c) Cr, Fe³⁺ and Al relations in spinels of the Spontang peridotites. The superimposed fields are taken from Jan and Windley (1990). All analysed samples plot in the residual peridotites or ophiolites field in this diagram. (d) Compositional variations of orthopyroxene and clinopyroxene projected in the Di-En-Hd-Fs quadrilateral of the Spontang peridotites (after Morimoto et al., 1988).

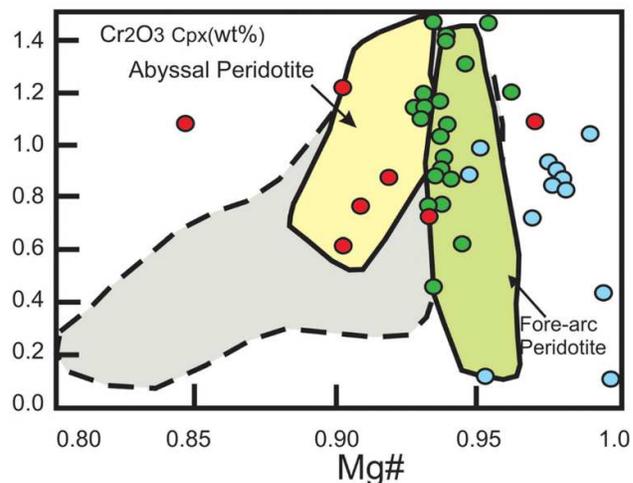
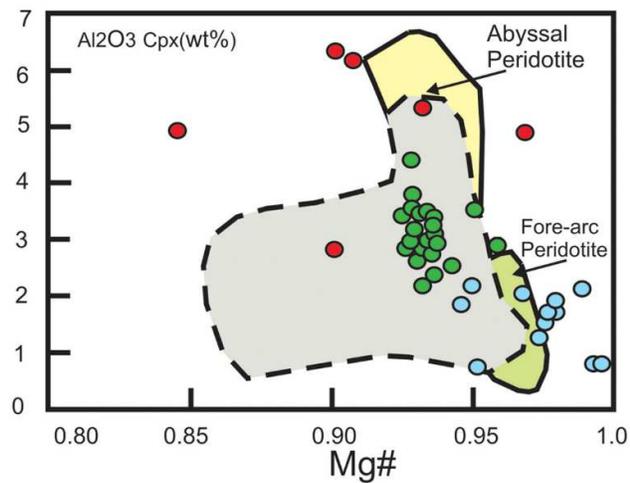
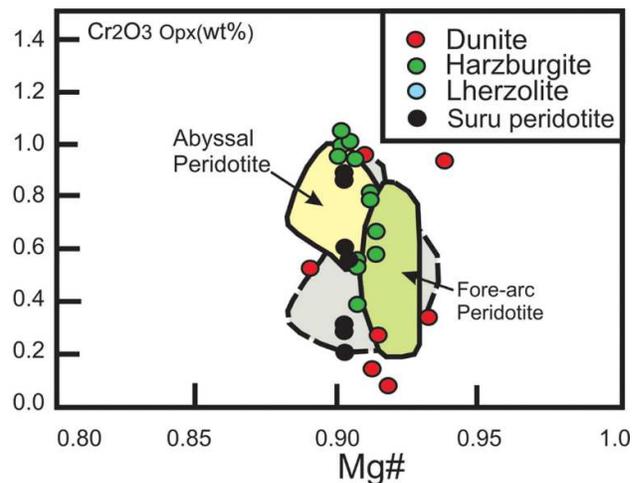
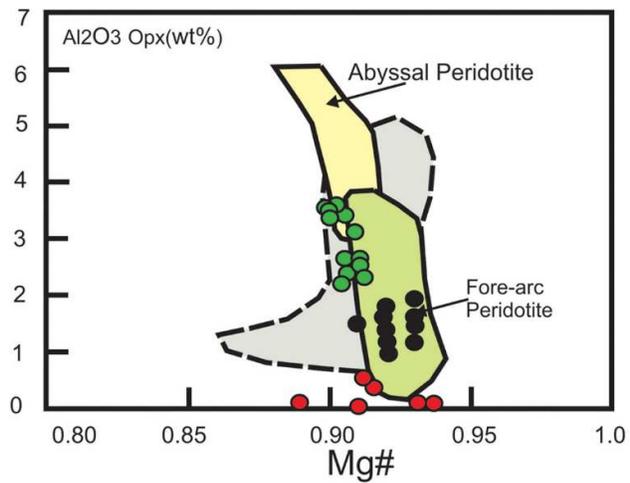
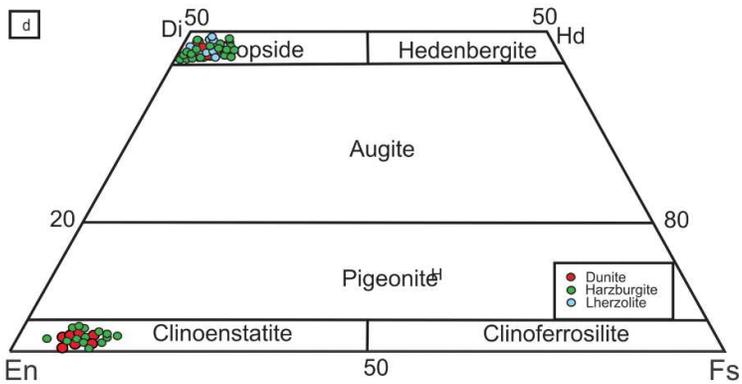
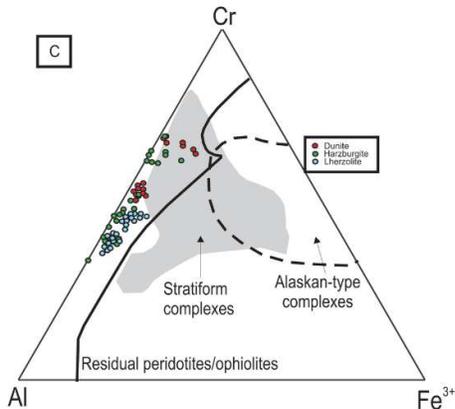
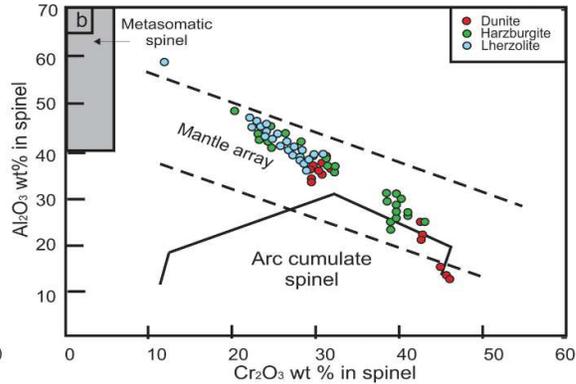
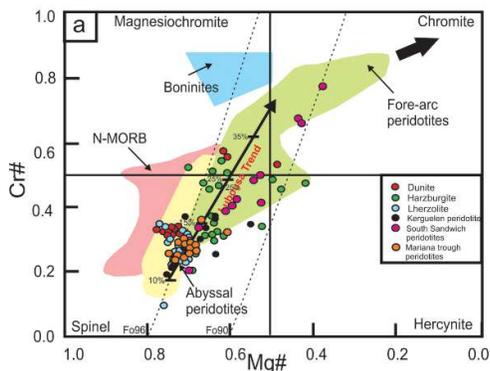


Fig. 8. Compositional variations of pyroxenes from peridotites of the Spontang ophiolite complex. Cr_2O_3 and Al_2O_3 vs. Mg#. Fields outline for clinopyroxene and orthopyroxene compositions in abyssal peridotites (Johnson et al., 1990) and forearc peridotites (Ishii et al., 1992). Dashed field outlines compositions in peridotites from the YZSZ ophiolites (Hébert et al., 2003; Dubois-Côté, 2004). Black dots are Suru peridotites after Bhat et al. (2018).

variation in the composition is observed among the studied peridotites. Forsterite content (Mg#) in the analyzed crystals has a restricted compositional range between 87% and 92% in the lherzolite and harzburgite samples and between 89% and 91% in the dunite which is close to that of olivine in abyssal peridotites (average Fo content: 91%; Dick and Bullen, 1984). NiO contents range between 0.21 wt% and 0.39 wt%.

Analyzed pyroxenes in the investigated harzburgites consist of enstatite orthopyroxene ($\text{En}_{89-90}\text{Wo}_{1-2}\text{Fs}_{8-9}$). Clinopyroxene in the lherzolites and harzburgites is diopsidic with compositions $\text{En}_{50}\text{Wo}_{49}\text{Fs}_1$ and $\text{En}_{47}\text{Wo}_{48}\text{Fs}_9$ respectively (Fig. 7d) and do not exhibit any systematic zoning and are chemically homogenous. Clinopyroxenes from lherzolites show high Mg# averaging 0.978. Harzburgite clinopyroxene shows Mg# ranging between 0.926 and 0.959 averaging 0.935 when compared with orthopyroxene (Av: 0.905; 0.899–0.912). Cr# values in enstatite are lower (0.100–0.180) as compared with diopside (0.119–0.245). Dunite samples contain orthopyroxene and clinopyroxene with the following compositions $\text{En}_8\text{Wo}_{88}\text{Fs}_6$ and $\text{En}_7\text{Wo}_{70}\text{Fs}_{26}$ respectively. Clinopyroxene in the dunites (Av: 0.942; 0.847–1.028) shows higher Mg# as compared with orthopyroxene (Av: 0.9015; 0.890–0.937) and could

be a result of reactive melt percolation (Abily and Ceuleneer,

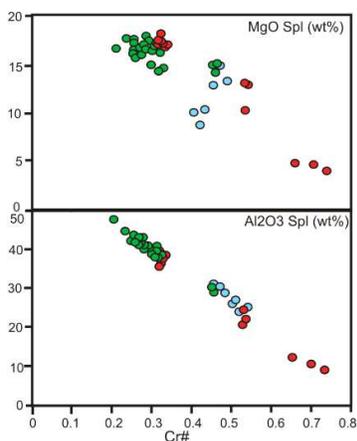
2013; Rospabe et al., 2017). Overall clinopyroxenes have Al_2O_3 contents that range from 1.52 wt% to 5 wt%, Cr_2O_3 in the range of 0.65 wt% to 1.27 wt%. Orthopyroxenes have Al_2O_3 and Cr_2O_3 contents varying between 0.56 wt% to 3 wt% and 0.46 wt% to 0.56 wt% respectively. Orthopyroxene and clinopyroxene compositions for the harzburgites when plotted in the Mg# vs. Al_2O_3 and Cr_2O_3 diagrams (Fig. 8), are observed to fall intermediate to the fields of abyssal and fore-arc peridotites. Dunite samples however, do not show any consistent trends. Lherzolite samples on the contrary show few samples falling in the fore-arc region. Chemical compositions of pyroxenes from the Spontang peridotites are compared with peridotites along the Indus Suture Zone viz. Suru peridotites (orthopyroxenes; Bhat et al., 2018) and the peridotites along the Yarlung Zangpo Suture Zone (Hébert et al., 2003; Dubois-Côté, 2004).

5.4. Geothermobarometry

Equilibration temperatures for the studied peridotite samples have been calculated using compositions from co-existing mantle phases and detailed calculations are given in Supplementary Tables S1 and S2. Care was taken to analyse orthopyroxene porphyroclasts free of clinopyroxene exsolutions lamellae. Using the partitioning of Mg and Fe^{2+} between olivine and spinel of Fabries (1979) the calculated temperatures range from 646 °C to 795 °C. Application of the two-pyroxene geothermometer of Brey and Kohler (1990), yielded temperatures varying between 718 °C and 896 °C, assuming a pressure of 1.5 GPa typical of upper mantle peridotites within the spinel stability field (Gasparik, 1987). The low temperature values yielded by olivine-spinel geothermometer may have resulted due to recrystallisation of different rates of diffusion between the mineral pairs used in the calculation.

6. DISCUSSION

Spinel habits and their compositions can be used to understand the genesis of ultramafic rocks. Spinel's originating from the deeper upper mantle has been attributed to multiple origins ranging from products of partial melting to metasomatic processes and silicate decomposition (Haggerty, 1991). Ophiolitic spinels are widely considered to be derived from the upper mantle or as a product of crystal fractionation in cumulate sequences in the lithosphere (Paktunc, 1990). The spinels from the Spontang ophiolite along the Indus suture zone are compositionally consistent with the exception of some minor increase and decrease of certain elements. Morphological changes are observed in the spinel grains from symplectitic intergrowths in the harzburgites to more idiomorphic grains in the dunites. The Cr# in the spinels from the harzburgites and dunites are observed to show an inverse relationship with Al and Mg, i.e., with increasing Cr content the spinel gets depleted in Al and Mg, whereas Cr#



versus Mg in lherzolites show a positive correlation (Fig. 9) indicating that the lherzolites represent the fertile mantle. Spinel's evolved from an Al rich phase to Cr rich-Al poor phase which are expected as a result of partial melting (Karmalkar et al., 1997) coinciding with the above-mentioned morphological changes. Spinel-orthopyroxene intergrowth observed in the harzburgites could be result of garnet

+ olivine reaction indicating garnet to spinel peridotite transition during partial melting (Karmalkar et al., 1995, 1997).

6.1. Interpretation of Cr-spinel Data

Compositions of accessory Cr-spinel in peridotites is regarded as a useful tool for revealing melting processes in mantle (e.g., Okamura et al., 2006; Uysal et al., 2007). Cr-spinel is known to be extremely sensitive to bulk composition mineralogy and petrogenesis of the host rock (Irvine, 1965, 1967; Evans and Frost, 1975; Dick, 1977b; Fisk and Bence, 1980). Cr-spinel compositions help determine the degree of partial melting of the mantle source and/or composition of produced mafic melt (e.g., Dick and Bullen, 1984; Arai, 1992; Zhou et al., 1998; Proenza et al., 1999; Barnes and Roeder, 2001; Hellebrand et al., 2001; Kamenetsky et al., 2001; Zhou et al., 2005; González-Jiménez et al., 2011) as well as processes involved in the evolution of upper mantle rocks i.e., mantle metasomatism (e.g., Kubo, 2002; Arif and Jan,

2006). It is known that Al and Cr# of spinel is sensitive to mantle melting processes and that Al systematically decreases whereas Cr# increases with the degree of peridotite depletion (e.g., Arai, 1994; Zhou et al., 2005; Zaccarini et al., 2011; Uysal et al., 2012). In the studied samples, the Cr# of chromian spinel is observed to be lower in spinel from lherzolitic and harzburgitic rocks compared with dunites. In additions to this, TiO₂ and MnO contents are low typical of unaltered Cr-spinel in ultramafic rocks (e.g., Barnes, 2000; Singh and Singh, 2013).

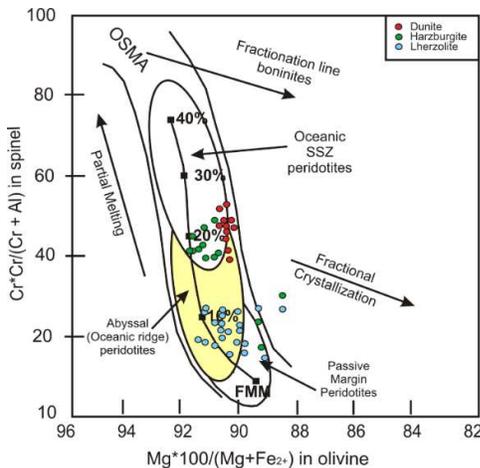
Differences among the studied peridotites can be observed in the plot of Cr# in spinel versus Fo content in olivine (Fig. 10). The spinel Cr# is observed to increase as the Fo content in olivine increases such that the spinel-olivine pairs from the studied peridotites lie within the olivine-spinel mantle array (Arai, 1994). This correlation between the spinel and olivine compositions confirms the mantle residue origin of the investigated samples. In the Cr# vs. Mg# plot, Cr-spinels from the investigated peridotites fall mainly along a trend defined by a large increase in Cr# for a small decrease in Mg# (Hirose and Kawamoto, 1995) and follow the Luobusa trend defined by spinel of Yarlung Zangpo

chemical variations observed in spinel caused by different degrees of apparent partial melting on the host peridotites. Spinel aligned along the Luobusa trend have a low TiO_2 content (< 0.14 wt%) and are characterized by a small decrease in TiO_2 for a larger increase in Cr# and are consistent with observations made in the studied samples (Fig. 7a). Cr-spinel hosted in the lherzolites are represented by higher Mg# and lower Cr# suggesting up to 10–12% partial melting (Fig. 10). On the other hand, harzburgites contain Cr-spinel with lower Mg# and higher Cr# up to 18–22% partial melting, whereas the dunites with highest Cr# are observed to form a separate cluster indicating higher (25–28%) degrees of partial melting consistent with those provided by Hirose and Kawamoto (1995) (Fig. 7a). Based on the above, the studied peridotites from the Spontang ophiolite are inferred to be produced by variable degrees of mantle melting and that the dunites represent a mantle residue resulting from higher degree of partial melting.

6.2. Petrogenesis and Tectonic Setting

Traditionally, ophiolites were believed to form along mid ocean ridge spreading centres however, recent studies of modern oceanic basins indicate that subduction is an important process involved in the formation of these ophiolites more commonly known as supra-subduction (SSZ) ophiolites. Globally, most well-preserved ophiolites appear to show mantle compositions of both MOR and SSZ (Pearce et al., 1984; Stern, 2004). Mantle

Fig. 10. Cr# in spinel versus Fo content of olivine in peridotites of Spontang ophiolite. Fields for spinels occurring in abyssal (and ocean ridge), oceanic SSZ and passive margin peridotites are after Dick and Bullen (1984) and Pearce et al. (2000). OSMa means Olivine Spinel Mantle Array and fractionation line of boninites are after Arai (1994a, 1994b).



peridotites are known to contain both SSZ and abyssal peridotites in the fore-arc regions (Parkinson and Pearce, 1998; Pearce et al., 2000). Composition of Cr-spinel in addition to the elemental and modal composition of the mineral constituents of mantle peridotites is considered to be a good indicator of the tectono- magmatic history of the host rock. (e.g., Dick and Bullen, 1984; Arai, 1992; Zhou et al., 2005; Arai et al., 2011; Ahmed et al., 2012; Uysal et al., 2012), and can be used to determine the degree of partial melting (Dick and Bullen, 1984; Arai, 1994; Zhou et al., 1996; Hellebrand et al., 2002; Aswad et al., 2011 and references therein). SSZ peridotites are characterized by spinels with high Cr#s ranging between 38 to 80 (e.g., Dick and Bullen, 1984; Juteau et al., 1990; Ishii et al., 1992; Ohara and Ishii, 1998) and pyroxenes with extremely low Al, Ti and Na indicating significantly higher degrees of partial melting ($> 15\%$) as compared to abyssal peridotites with low Cr# in spinel = 38 to 58 (Arai, 1994; Gaetani and Grove, 1998; Choi et al., 2008).

Petrography and geochemistry of the studied samples provide clear evidence of the Spontang peridotites having a multi-process history. The Spontang peridotites are characterized by increase in Fo and NiO contents of olivine (Fig. 11), Mg# and Cr_2O_3 content of pyroxenes (Fig. 8), and Cr# of spinel with decreasing Al_2O_3 (Fig. 9) and TiO_2 (Fig. 12) content in spinel and bulk rock as melting progresses. Similar petrological and geochemical observations are reported in upper mantle peridotites of the YZSZ (Dupuis et al., 2005 and references therein) along with peridotites from other Himalayan ophiolites (Fig. 8; Bhat et al., 2018); mantle xenoliths of the Kerguelen archipelago (Grégoire et al., 1997), and in peridotites of the South Sandwich arc-basin (Pearce et al., 2000) and Mariana Trough (Ohara et al., 2002) (Fig. 7a). The spinels compositions from studied harzburgites fall in the fore-arc region with lherzolites and dunites falling in the abyssal peridotite fields (Fig. 7a). The Spontang peridotites have olivine

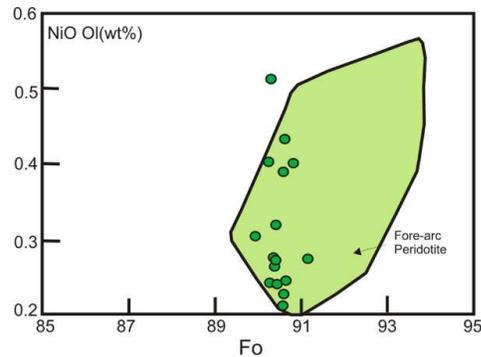


Fig. 11. Variation of NiO versus Fo in olivine from peridotites of the Spontang ophiolite complex. Field outline is olivine compositions in fore-arc peridotites (after Ishii et al., 1992).

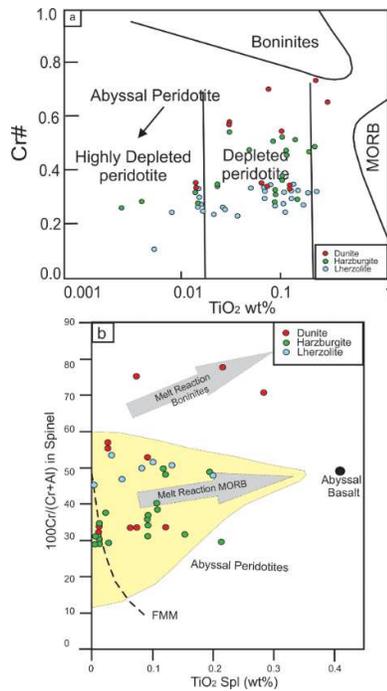


Fig. 12. (a) TiO_2 vs. Cr# diagram in spinels of Spontang ophiolite. Fields are after Dick and Bullen (1984); Jan and Windley (1990); Arai (1992). (b) Dotted line field of abyssal peridotite spinels (after Dick and Bullen, 1984); Dotted arrow shows effect of MORB melt on refractory abyssal peridotite spinels. Grey arrow shows effect of boninite melt reaction on refractory supra-subduction zone peridotite spinels. FMM = Fertile MORB Mantle.

clinopyroxene and orthopyroxene compositions similar to those from both abyssal and fore-arc peridotites (Fig. 6). The Cr-spinels when plotted in the TiO_2 vs. Cr# diagram (Fig. 12a) fall within the depleted peridotite field to highly depleted mantle fields typical of SSZ settings (Ohara et al., 2002) away from the MORB field. The Al_2O_3 vs. TiO_2 diagram (Fig. 13) the samples plot between the MOR and SSZ peridotites. Overall, the mineral chemistry of the Spontang peridotites are characterized by interaction between depleted magma and ultra-depleted melts of pre-existing oceanic lithosphere, typical of supra-subduction zone settings (Fig. 12b; Choi et al., 2008).

Whole rock REE signatures also suggest formation of the Spontang peridotites in a supra-subduction zone setting (Fig. 7). In comparison with the Spontang peridotites, Figure 7 shows peridotites found along the Indus-Yarlung Suture

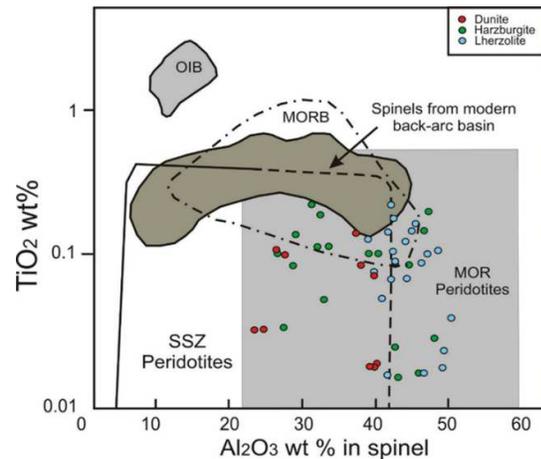


Fig. 13. TiO_2 vs. Al_2O_3 of spinels from peridotites of the Spontang ophiolite complex. Fields are after Kamenetsky et al. (2001)

zone (Fig. 7; Dupuis et al., 2005 and references therein) and fore-arc peridotites from the Leg 125 peridotites from Western Pacific region (after Parkinson and Pearce, 1998) along with abyssal peridotites from the Manipur ophiolitic complex (after Singh, 2013). The studied samples have depleted LREE profiles similar to fore-arc peridotites in comparison to LREE's from abyssal settings that exhibit relatively flat to upward-inflected patterns (Fig. 6). LREE enrichment in sample E-5 could be attributed to interaction of melt or aqueous fluids (Rospabé et al., 2017, 2018). Overall MREE-HREE profiles display positive slopes comparable with abyssal peridotites. Overall the studied samples display spoon-shaped REE profiles characteristic of interaction between LREE-enriched melt, derived from the subducting slab and LREE-depleted mantle residues (Edwards, 1990; Kelemen et al., 1992; Zhou et al., 2005). Overall, based on the mineral and whole rock geochemistry, the Spontang ophiolite complex shows evidence of a mid-ocean ridge history prior to a supra-subduction history.

7. CONCLUSIONS

1) The ultramafic bodies of the Spontang ophiolite complex are presented by dunites, harzburgites and minor lherzolites displaying porphyroclastic to granular textures. Morphologically spinels in the studied samples occur as xenomorphic, vermicular intergrowths displaying characteristic holly leaf shapes. Euhedral spinel inclusions within olivine at places are also observed. Morphological changes in the spinel grains range from symplectitic intergrowths in the harzburgites to more idiomorphic grains in the dunites. Spinel's

evolved from an Al rich phase to Cr rich-Al poor phase which are expected as a result of partial melting coinciding with the above-mentioned morphological changes.

2) The Spontang peridotites are characterized by increase in Fo and NiO contents of olivine, Mg# and Cr₂O₃ content of pyroxenes, and Cr# of spinel with decreasing Al₂O₃ and TiO₂ content in spinel and bulk rock as melting progresses. Cr# values of Cr-spinel increase and Mg# values decrease from lherzolite to harzburgite and dunite. Studied spinels are characterized by a small decrease in TiO₂ for a larger increase in Cr# consistent with observations for spinels aligned along the Luobusa trend of the YZSZ ophiolites.

3) Mineral and whole-rock chemistry of the Spontang peridotites is characterized by interaction between magma and pre-existing oceanic lithosphere, typical of supra-subduction zone settings. The Spontang peridotites have olivine, clinopyroxene and orthopyroxene compositions similar to those from both abyssal and fore-arc peridotites with most spinels from harzburgites falling in the fore-arc region with lherzolites and dunites falling in the abyssal peridotite fields. Overall the studied samples display spoon shaped REE profiles characteristic of interaction between LREE-enriched melt, derived from the subducting slab and LREE-depleted mantle residues.

4) Equilibration temperatures for the studied peridotites calculated using the two-pyroxene geothermometer and the partitioning of Mg and Fe²⁺ between olivine and spinel yield temperatures varying between 707 °C and 883 °C and 646 °C to 795 °C respectively typical of mantle peridotites within spinel stability field.

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