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1 **Direct and indirect impact of the bacterial strain *Pseudomonas***
2 ***aeruginosa* on the dissolution of synthetic Fe(III)- and Fe(II)-bearing**
3 **basaltic glasses**
4

5
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Abstract

This study is dedicated to the determination of the respective direct and indirect bacterial contributions during the dissolution process of synthetic basaltic glass. In this regard, three different types of glasses – with or without Fe, in the reduced Fe(II) or oxidized Fe(III) states – were prepared on the basis of a simplified basaltic glass composition. In order to assess the direct/indirect bacterial contributions, the experimental set-up included the use of dialysis bags, which prevent the direct contact between the model siderophore-producing strain *Pseudomonas aeruginosa* and the synthetic glasses. The glasses were immersed at 25°C and pH 6.5 in a growth medium inoculated with the strain and submitted to alteration times from 8 hours to 6 days. Throughout the dissolution experiments, the following parameters were monitored: determination of bacterial growth, quantification of siderophore (*i.e.* pyoverdine) production, microscopic observation of the glass surface and determination of dissolution kinetics. The results were compared to identical experiments performed without dialysis bags and show different behaviors depending on the accessibility of the cells to the glass. For all compositions, isolating the glass from the bacterial suspension triggered the biosynthesis of siderophores. This production appeared to be either continuous in the Fe-free system or stopped as soon as the bacteria entered their stationary phase of growth in the Fe-bearing systems. By contrast, allowing a direct contact between the glass surface and the bacterial suspension triggered a high and continuous production of siderophore in the Fe-free system and only a small production in the presence of the Fe(II)-bearing glass. The dissolution rates were approximately increased by a factor of 3 for Fe-free and Fe(III)-bearing glass and by a factor of 2 for Fe(II)-bearing glass compared to sterile experiments. Pyoverdine concentrations similar to those obtained with dialysis bags were obtained in presence of the Fe(II)-bearing glass and were below detection limit in presence of the Fe(III)-bearing glass. The dissolution rates were correlated to siderophore concentrations as they increased this time with respect to sterile experiments in the order Fe(III)-bearing glass < Fe(II)-bearing glass < Fe-free glass. Finally, in all experimental cases, the bacteria appeared to increase the dissolution kinetics of the three model glasses involved. However, the Fe(III) accessibility to the bacterial cells was shown to be the key parameter regulating the bacterial activity and their efficiency in accelerating the dissolution.

Key-words:

Basaltic glass; *Pseudomonas aeruginosa*; siderophore; bioalteration; dissolution kinetics

66

Introduction

67

68 Silicate glasses are abundant at the Earth's surface and notably near the oceanic ridge axes where
69 large amounts of lava are emitted every year, generating a profusion of both amorphous and
70 crystalline basaltic rocks. The alteration processes of basaltic glasses in permanent contact with
71 seawater are known for partially controlling the composition of this latter but also the composition of
72 the oceanic crust and impacting on the geochemistry of the Earth mantle. The alteration of basaltic
73 glasses could also have a significant impact on the Earth's climate by contributing to the sequestration
74 of atmospheric CO₂ through carbonation processes resulting from the dissolution of the glass and the
75 release of earth alkali elements (Berner et al., 1983). Understanding the alteration mechanisms of the
76 oceanic crust at low temperatures is thus essential to implement the Earth surface mass-balance
77 studies and then improve the geochemical models predicting the evolution of our planet (Hart, 1970;
78 Staudigel and Hart, 1983; Spivack and Staudigel, 1994; Gislason and Oelkers, 2003).

79

80 During the last two decades, the high diversity and the abundance of the microorganisms living at the
81 subsurface of oceanic basalts were evidenced (Thorseth et al., 2001; Santelli et al., 2009; Mason et
82 al., 2009; Orcutt et al., 2011; Henri et al., 2015; Sudek et al., 2017). As basaltic glasses represent one
83 of the most profuse source of Fe and also contain additional elements (Mg, Ca, K...) required for
84 microbial nutrition (Staudigel and Hart, 1983), they are suspected to sustain the development of
85 microbial communities (Sudek et al., 2009). The literature attests to four main potential bacterial sinks
86 for the elements constituting silicate materials submitted to bioalteration: 1/ adsorbed or 2/
87 incorporated inside the bacterial cells, 3/ trapped within the biofilms or 4/ complexed with bacterial
88 exometabolites such as siderophores (e.g. Uroz et al., 2009). However, the role of basaltic glasses as
89 a substantial source of nutrients, the impact of such microbial activity on their dissolution rates and its
90 potential control on the biogeochemical cycling of elements at the rock/water interface is still a matter
91 of debate (Templeton et al., 2005; Cockell et al., 2010; Henri et al., 2016).

92

93 Numerous laboratory studies were dedicated to the understanding of the basaltic glass dissolution
94 processes in a various range of abiotic experimental conditions (e.g. Guy and Schott, 1989; Daux et
95 al., 1997; Techer et al., 2000, 2001; Oelkers, 2001; Oelkers and Gislason, 2001; Stroncik and
96 Schmincke, 2001; Gordon and Brady, 2002; Crovisier et al., 2003; Gislason and Oelkers, 2003; Wolff-
97 Boenisch et al., 2004a; Flaathen et al., 2010; Verney-Carron et al., 2011; Parruzot et al., 2015; Perez
98 et al., 2015). The dissolution of basaltic glass is traditionally defined by two simultaneous mechanisms,
99 whose predominance notably depends on the pH of the aqueous fluid: (1) an ion exchange between
100 network-modifying elements of the glass (Na, K, Ca, Mg) and hydrogenated species from the solution
101 (interdiffusion) and (2) the hydrolysis of the bonds linking oxygen atoms to network-forming elements
102 of the glass (Si, Al, Fe(III)). Dissolution can also be presented as the breaking of all metal-oxygen
103 bonds (from the weakest to the strongest) until the structure is completely destroyed (Oelkers, 2001;
104 Oelkers and Gislason, 2001). As the rate-limiting step of the basaltic glass dissolution process is the
105 detachment of tetrahedral Si near the glass surface, the change of the other network-forming elements

106 (Al, Fe) speciation is assumed to influence the dissolution rate. For these reasons, the impact of the
107 introduction in the liquid medium of biomolecules known for their high affinity with such cations should
108 be investigated.

109
110 Several field studies conducted in deep-sea water and dedicated to the characterization of microbial
111 communities show that if the many cultured organisms are metabolically and functionally diverse, they
112 display at least two of the tested functions: heterotrophy, Fe(II) and Mn(II) oxidation and siderophore
113 production (e.g. Sudek et al., 2009). Siderophores are molecules produced when bacteria are
114 submitted to a nutrient stress, and especially to a high need in Fe. The biosynthesis of siderophores
115 involves several steps catalysed by specific enzymes in different locations from cytoplasm to
116 periplasm, from where they are secreted to the extracellular medium (Schalk and Guillon, 2013).
117 Siderophores have generally a very high affinity for two of the three network-forming elements Al and
118 Fe(III). In two previous studies (Perez et al., 2015, 2016), experiments performed on synthetic Fe(II),
119 Fe(III)-bearing basaltic glasses (MORB2, MORB3) and Fe-free glass (HAPLO) in the presence of
120 siderophores and then in the presence of the model siderophore-producing strain *Pseudomonas* (*P.*)
121 *aeruginosa* exhibited enhanced dissolution rates compared to those performed in sterile controls. The
122 results highlighted the promoting role of organic ligands as complexing agents, and particularly of
123 pyoverdines, which are the siderophores produced by *P. aeruginosa*. These conclusions were in
124 agreement with numerous abiotic studies dealing with dissolution kinetics of minerals/oxides/silicate
125 glasses in the presence of organic molecules (Welch and Ullman, 1992; Watteau and Berthelin, 1994;
126 Stillings et al., 1995; Drever and Stillings, 1997; Kraemer et al., 1999; Liermann et al., 2000; Coccozza
127 et al., 2002; Cheah et al., 2003; Wolff-Boenisch, 2007; Wang et al., 2005; Olsen and Rimstidt, 2008;
128 Martinez-Luevanos et al., 2011; Torres et al., 2014; Akafia et al., 2014; Lazo et al., 2017). This
129 promoting effect is in favour of the occurrence of a bacteria-mediated process, which notably contrasts
130 with several recent studies attesting to a small or negligible impact of the presence of bacteria on
131 dissolution kinetics of minerals (Hopf et al., 2009; Hutchens, 2009; Sverdrup, 2009; Bray et al., 2015)
132 and of basaltic glass (Wolff-Boenisch, 2011; Stockmann et al., 2012). It is therefore important to note
133 that this promoting effect was observed in specific conditions (growth phase and nutrient free systems)
134 in which the acquisition of nutrient was crucial for the bacterial cells. Even though the impact of
135 siderophores on the dissolution kinetics of basaltic model glasses was evidenced, the bacterial-
136 promoted dissolution mechanisms of the glasses remain hard to define as bacteria can also directly
137 interact with glass surfaces or produce other metabolites that could influence the dissolution rates
138 (Ullman et al., 1996; Hutchens, 2009; Shirokova et al., 2012; Ahmed and Holmström, 2015; Cornu et
139 al., 2017; Wang et al., 2018).

140 The present work aims to decipher direct and indirect dissolution mechanisms of basaltic glass by
141 bacteria. For this purpose, bioalteration experiments were performed on HAPLO, MORB2 and MORB3
142 synthetic glasses. In order to assess the direct/indirect bacterial contributions, the experimental set-up
143 included the use of dialysis bags, which prevented the direct contact between the model siderophore-
144 producing strain *P. aeruginosa* and the basaltic glasses tested. Throughout the dissolution
145 experiments the following parameters were monitored: determination of bacterial growth, quantification

146 of siderophore (*i.e.* pyoverdine) production, microscopic observation of the glass surface and
147 determination of dissolution kinetics.

148
149

150 **Material and Methods**

151

152 1. Samples

153

154 The theoretical compositions of the three model glasses used in this study are given in Table 1. The
155 MORB2 and MORB3 glasses were prepared according to a simplified typical *Mid Oceanic Ridge*
156 *Basalt* composition and contain mainly Fe(II) and Fe(III), respectively. The HAPLO glass is an Fe-free
157 equivalent composition in which Fe is mainly substituted by Mg.

158

159 Table 1: Nominal compositions of the synthetic Fe(II)-bearing glass (MORB2), Fe(III)-bearing glass (MORB3) and Fe-free glass
160 (HAPLO) of interest.

	MORB2 & 3	HAPLO
	(wt %)	(wt %)
SiO ₂	48.6	51.6
Al ₂ O ₃	15.7	16.7
FeO	12.5	0.0
CaO	11.1	11.8
MgO	7.7	15.2
Na ₂ O	2.7	2.9
TiO ₂	1.4	1.5
K ₂ O	0.2	0.3

161

162 MORB2, MORB3 and HAPLO glasses were prepared according to the procedure described in the
163 work of [Perez et al. \(2016\)](#). More details concerning the characterization of these glasses (DRX, XAS,
164 BET measurements) are also given in this previous study.

165

166 2. Bioalteration experiments

167

168 Freshly grown cultures of *P. aeruginosa* ATCC® 9027 (Sigma Aldrich) bacteria having identical age
169 and initial nutritional status were used in all bioalteration experiments.

170

171 Two growth media were used in this study. The bacteria were first grown in 2 mL of Lysogeny Broth
172 (LB) medium, a nutritionally rich medium containing: 10 g L⁻¹ tryptone, 5 g L⁻¹ yeast extract, 5 g L⁻¹
173 NaCl.

174

175 The prepared glasses were then immersed in a buffered (pH 6.5) MM9 medium whose composition
176 was adapted from Schwyn and Neilands (1987) and Liermann et al. (2000). Near neutral conditions of
177 pH were chosen to allow for comparison with similar experiments conducted in abiotic conditions
178 (Perez et al., 2015). The medium composition was as follows: 6.06 g L⁻¹ Na₂HPO₄, 0.3 g L⁻¹ KH₂PO₄,
179 0.5 g L⁻¹ NaCl, 1.0 g L⁻¹ NH₄Cl and 50 mM Bis Tris. The pH was adjusted by adding HCl 1 M. The
180 solutions were sterilised and cooled to 25 °C. They were then supplemented with 0.2% (v/v) 1 M
181 MgSO₄ (final concentration 2 mM), 1% (v/v) filter-sterilised (cellulose filter, Ø0.2 µm) 20% (w/v)
182 glucose (final concentration 1 M) and 0.01% (v/v) 1 M CaCl₂ (final concentration 0.1 mM); these
183 solutions were prepared and sterilized separately. No Fe was detected in the prepared growth medium
184 and no precipitate formation was observed. Dissolution experiments were performed in sterile
185 borosilicate erlenmeyers closed with cotton plugs to allow free passage of air over the edge of the
186 vessels and to maintain an aerobic aqueous environment. A volume of 250 mL of MM9 6.5 medium
187 was mixed with 0.625 g of glass powder and inoculated together with 125 µL of the LB culture of *P.*
188 *aeruginosa* after 30 h of incubation at 25 °C and 160 rpm. By contrast with the experiments presented
189 in Perez et al. (2016), the glasses were isolated from the bacterial suspension by sterile dialysis bags
190 (CelluStep T4). Porous dialysis membranes allow small solutes to pass while large species are
191 retained and can thus effectively be used as a separation process based on size rejection. The
192 Molecular Weight Cut Off of the membranes used in our experiments is 12000-14000 Da, which
193 means that: (1) 90% of the species whose molecular weight ranges between 12000 and 14000 Da do
194 not pass the membrane and do not access to the glass surfaces and (2) 100% of the species whose
195 molecular weight exceeds 14000 g/mol do not pass the membrane. In the context of this study, the
196 used membranes allow most exometabolites, in particular siderophores, to interact with the glass
197 surface whereas the bacteria cells are isolated from the solid samples.

198 At each dissolution time, 2.3 mL of the liquid medium were filtered (cellulose acetate filter, Ø0.2 µm)
199 and kept for Inductively Coupled Plasma and Optical Emission Spectroscopy (ICP-OES) analysis
200 whereas 1.7 mL were placed in a spectrophotometer cell for Optical Density (OD) measurements and
201 200 µL were plated right away on LB agar medium for the microbial counting procedure. Sampling
202 times were chosen as a function of the bacteria growth: frequent from the beginning until the end of
203 the exponential phase and then more distant (0.6, 1, 1.6, 2, 3 and 5.7 days). Experiments were
204 extended until 15 days of experiments and 1.7 mL were sampled at days 8, 10, 13 and 15 and
205 introduced in a spectrophotometer cell for additional measurements.

206 For each glass, sterile experiments (without introducing *P. aeruginosa*) were conducted following the
207 same protocol and are mentioned as “abiotic control” in the Results and Discussion section. Moreover,
208 *P. aeruginosa* was cultivated in identical conditions but without adding the glass into the system.
209 These experiments were mentioned as “control” in the following sections.

210 All experiments were performed in duplicates.

211

212 3. Analysis

213 The amounts of dissolved Al, Fe and Si were analyzed using a Perkin Elmer Optima 8300 ICP-OES.
214 Measurements were performed using emission wavelength at 288.16 nm (Si), 308.22 nm (Al),
215 239.56 nm (Fe). Ca, Mg, Na, K were exempted from analysis being initially present in significant
216 amounts in the growth medium.

217 For each element and at each dissolution time, the normalized mass loss NL_i from the glass into the
218 solution was calculated using Eq. (1).

219
$$NL_i = \frac{[i]}{\left(\frac{S}{V}\right) \times x_i} \quad (1)$$

220
221 where $[i]$ is the concentration (mg L^{-1}) of the element i in solution, S the initial surface of the glass
222 powder in contact with the fluid, V the volume of solution and x_i the mass percentage of the element i
223 in the glass.

224 Considering a linear regression, the initial slope of the curve was calculated in order to evaluate the
225 *apparent* initial rate of dissolution r_i as given in Eq. (2):

226
$$r_i = \frac{dNL_i}{dt} \quad (2)$$

227
228 At each sampling time, total biomass growth was estimated by measuring the Optical Density (OD) of
229 the incubated medium at 600 nm and by the Colony Forming Unit (CFU) count procedure. After
230 2.7 days (respectively 2.9 days for MORB3 glass) and 5.7 days, dilutions of the bacterial medium were
231 plated on LB agar and incubated 24 h at 25 °C for the counting step while the sampled solution was
232 also filtered (cellulose acetate filters, $\text{Ø}0.2 \mu\text{m}$) and placed into a spectrophotometric cell. Optical
233 spectra were recorded between 300 and 600 nm. Identical spectrophotometric cells were used for all
234 measurements, with a standard optical path length.

235
236 The samples resulting from the dissolution experiments without dialysis bags (see [Perez et al. \(2016\)](#)
237 for the presentation and discussion of the liquid analysis results of these experiments) were dried
238 following a critical point drying (CPD) protocol in order to avoid the damaging effects of surface tension
239 on bacterial cell and biofilm potentially present at the glass surface. 250 μL of glass in suspension in
240 the bacterial medium were mixed with 4 mL of 0,05 M 4-(2-hydroxyethyl)-1-piperazineethanesulfonic
241 acid (HEPES) solution (pH 8) and then slowly injected in a swinnex® containing a polycarbonate filter
242 ($\text{Ø} 0,22 \mu\text{m}$). The glass grains and bacterial cells were retained on the filter and successively washed
243 with 5 mL of 50%, 70%, 96% and 100% ethanol (vol.). The filter was carefully kept in an absolute
244 ethanol bath then introduced in a CPD chamber (*Leica EM CPD300*) and submitted to 20 cycles (CO_2
245 injection, increase of temperature and pressure and depressurisation) of 4 min each. After being dried,

246 the samples were immediately carbon coated and observed with Scanning Electron Microscopy (*Zeiss*
247 *Ultra55* microscope) with a filament tension of 2 kV.

248

249

250 Results and Discussion

251

252 1. Bacterial growth and siderophore production

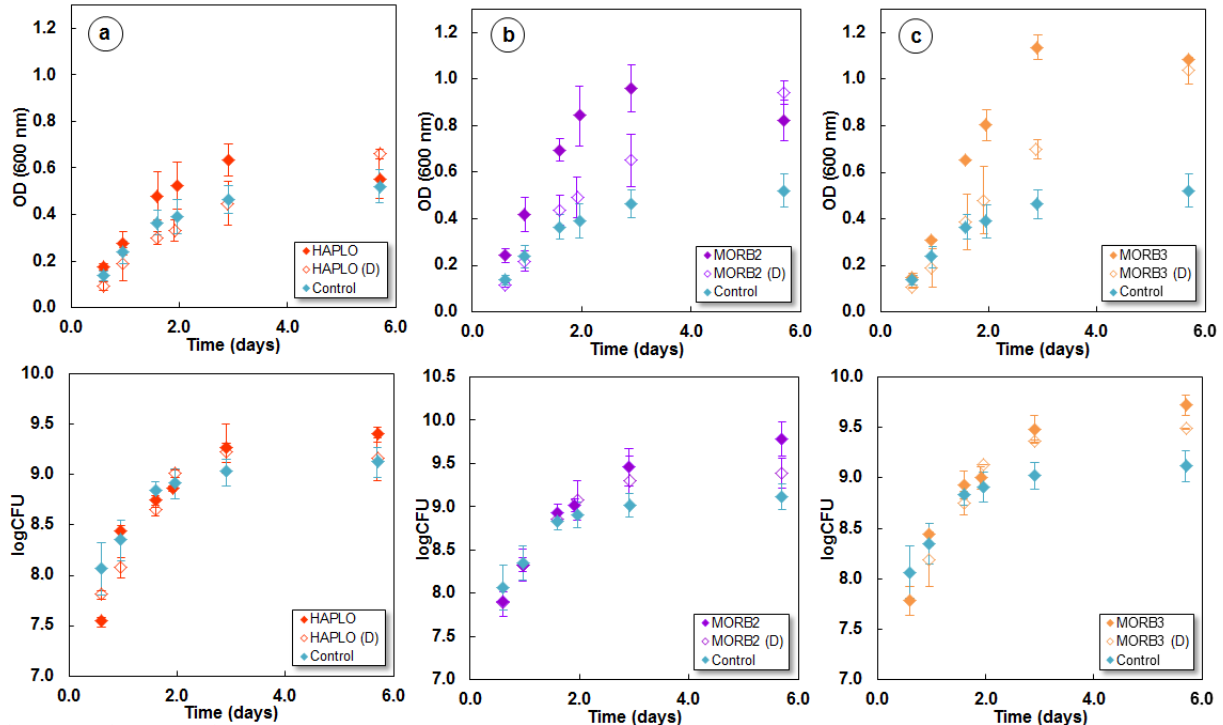
253

254 1.1. Growth

255

256 Optical densities (ODs) at 600 nm of the sampled solutions are plotted versus time in Figure 1 for the
257 three glasses. For each glass composition, the ODs measured in the presence of dialysis bags (D) are
258 compared to the ODs measured in dialysis-free experiments ([Perez et al., 2016](#)). The absorbance of a
259 bacterial medium at 600 nm is correlated to the number of suspended cells within the liquid medium as
260 it characterizes the turbidity of the solution. OD values are in good agreement with the CFU counting.
261 Correlation factors between OD and CFU values have been calculated for each experimental condition
262 and are equal to 0.97 (Control), 0.92 (MORB3), 0.94 (MORB3 D), 0.99 (MORB2), 0.99 (MORB2 D),
263 0.95 (HAPLO), 0.94 (HAPLO D).

264



265

266 Figure 1: (a) Bacterial growth curves (measured optical densities at 600 nm or Colony Forming Units versus time) of *P.*
267 *aeruginosa* culture in the presence of the MORB3 (a), MORB2 (b) and HAPLO (d) glasses, confined (empty symbols) or not (full
268 symbols) in dialysis (D) bags. Blue data points represent the control experiment. Growth experiments in MM9 6.5 were
269 performed in duplicates and error bars correspond to \pm standard deviation (SD) between the 2 measured values.

270

271 The experiments in the presence of the HAPLO glass exhibit growth curves very close to those
 272 obtained in the control experiments (Figure 1a). The bacterial growth is more pronounced when
 273 bacteria are incubated with Fe-bearing glasses, independently of the presence or absence of dialysis
 274 membranes (Figures 1b & 1c). This suggests that Fe-bearing glasses, isolated or not from the
 275 bacterial suspension, might represent substantial nutrient sources for *P. aeruginosa* in exponential
 276 phase of growth. This result is notably in strong agreement with our previous study (Perez et al. 2016)
 277 and with a study by Sudek et al. (2017). In this work, the authors evidenced the elevated growth of
 278 *Pseudomonas stutzeri* VS 10 incubated in the presence of basaltic samples in comparison with basalt-
 279 free experiments.

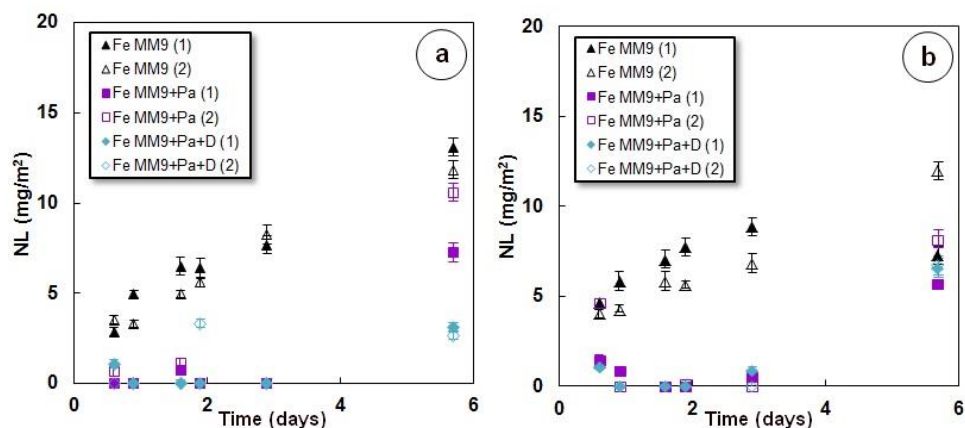
280

281 Moreover, the bacteria concentrations at 5.7 days are similar from one experimental condition (with
 282 dialysis) to another (without dialysis). Bacteria concentrations are nevertheless systematically lower in
 283 dialysis experiments from 0 to 3 days of incubation, suggesting that isolating the bacterial suspension
 284 from the glasses induces a slower cell division process. These results indicate that a direct contact
 285 between basaltic glasses and the bacterial suspension favors the bacterial cell division process,
 286 possibly as a consequence of an improved nutrient supply. On this point too, our results corroborate
 287 the recent findings of Sudek et al. (2017), who observed that a physical contact or at least the
 288 proximity of the cells to basaltic samples was correlated to a more pronounced bacterial growth.

289

290 The NLs of Fe measured for each duplicate in the cultures incubated in the presence of MORB2 and
 291 MORB3 glass (isolated or not from the bacterial suspension by dialysis membranes) are plotted
 292 versus time in Figure 2.

293



294

295 Figure 2: Normalized Loss (NL) profiles of Fe during dissolution experiments of MORB2 (a) and MORB3 (b) glass at 25°C and
 296 in the medium MM9 at pH 6.5 in a bacterial culture of *P. aeruginosa* (Pa) isolated (in blue) or not (in purple) from the glass by
 297 dialysis bags (D). Open symbols relate to duplicate (1) while full symbol relate to duplicate (2). Black data points represent the
 298 NLs measured in sterile controls. Error bars are equal to ± 2SD.

299

300 At the difference with dissolution experiments of those same glasses in UPW (Perez et al., 2015), Fe
 301 concentration levels in solution are most of the time above the ICP-OES detection limit. Basic
 302 speciation calculations (CHESS code) show that, regarding the abundant presence of phosphate salts

303 in the MM9 medium, Fe is mainly present as soluble FeHPO_4^+ (aq), as suggested by Pokrovsky et al.
304 (2009a,b).

305

306 In abiotic conditions (Fe MM9), Fe is continuously released in the medium all along the experiment
307 from both glasses. By contrast, in the presence of the strain (Fe MM9+Pa, Fe MM9+Pa+D), Fe is
308 almost non-detected from 0.6 to 5.7 days independently on the use or not of dialysis membranes. 0.6
309 to 5.7 days also corresponds to a fast cell division step (Figure 1). The few Fe released in solution is
310 probably entirely assimilated by the bacteria and used for the bacterial cell division process. At 5.7
311 days of dissolution experiment, while the bacteria cells are in a stationary phase of growth, Fe is
312 detected in the liquid medium. The presence of Fe at this point shows that Fe is naturally continuously
313 released into the solution, efficiently used by the cells to sustain their growth until they enter a
314 stationary phase and finally remains in solution as it is not needed anymore.

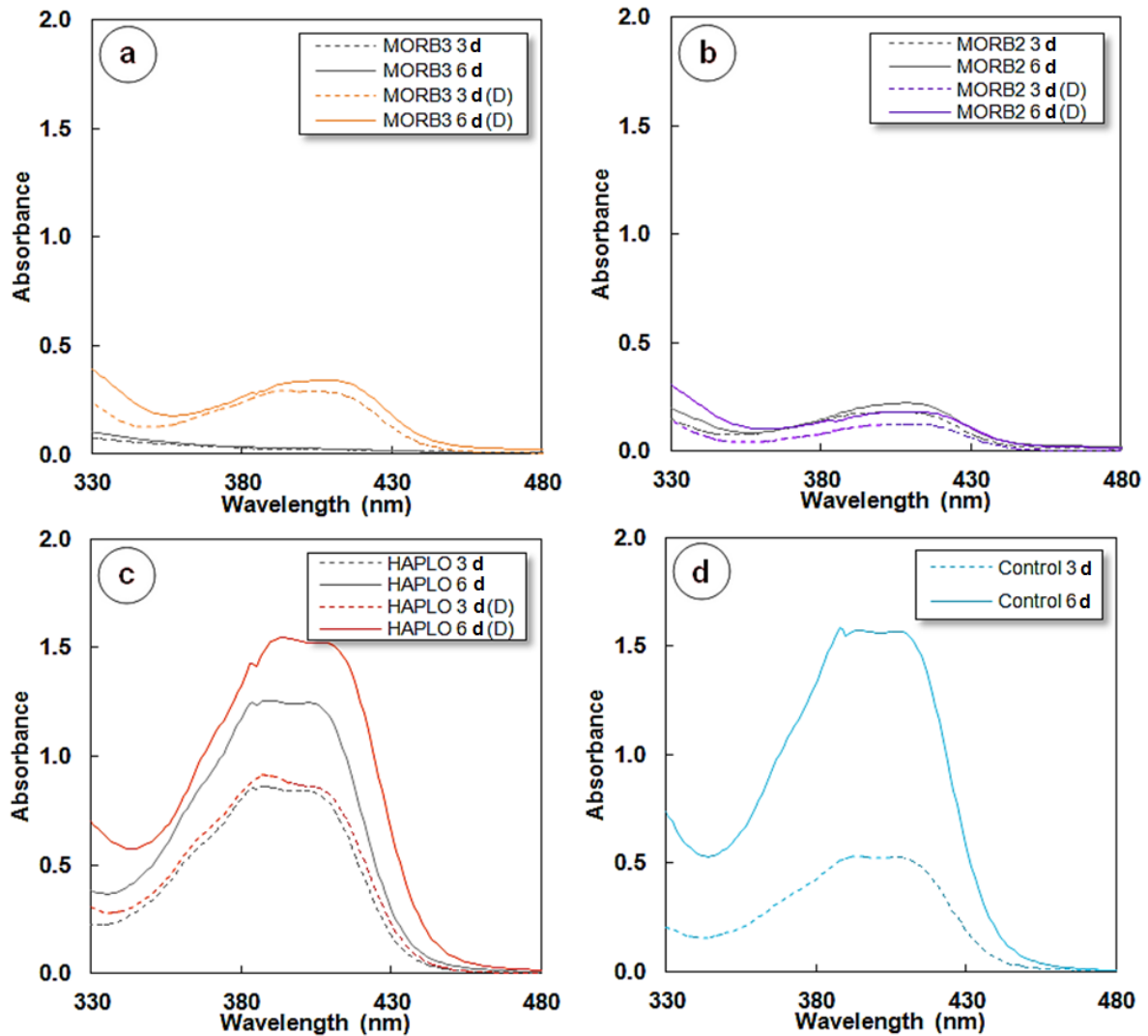
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316 Coupling bacterial growth and Fe concentrations finally highlights the high-affinity transport
317 mechanisms that *P. aeruginosa* presents in order to acquire Fe. The following section will be
318 dedicated to the identification of the strategies used by the strain to acquire this nutrient.

319 1.2. Siderophore production

320

321 Pyoverdines can be detected between 380 and 420 nm with a UV-Vis spectrophotometer through the
322 observation of a two-shoulder characteristic absorption peak (Meyer and Abdallah, 1978). The
323 absorbance spectra of the bacterial cultures, recorded in the UV absorption region characteristic of
324 pyoverdine, are plotted in Figure 3 and compared with the spectra obtained in no dialysis-using
325 experiments.



326
 327 Figure 3: UV Absorption spectra of the MM9 6.5 bacterial medium in the absence (Control) (d) or the presence of the MORB3
 328 (a), MORB2 (b), and HAPLO (c) glasses contained (colored lines) or not (greys lines) in dialysis bags, recorded at 3 days (3d)
 329 and 6 days (6d) of the experiments.
 330

331 According to the shape of the recorded spectra, pyoverdine molecules are detected in all experiments
 332 except when bacteria are incubated in contact with the MORB3 glass (no membranes) (Figure 3a).
 333

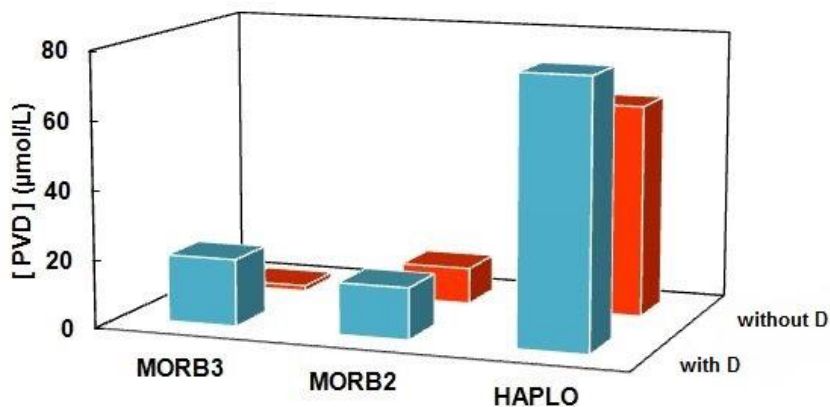
334 On the basis of the measured absorbance value at 405 nm (A_{405}), an estimated free pyoverdine
 335 concentration can be calculated according to the Beer Lambert law:
 336

337
$$[PVD] = \frac{A_{405}}{19000} \times 10^6 \quad (3)$$

 338

339 The pyoverdine is characterized by different molar extinction coefficients depending on its form in
 340 solution. The coefficient of free pyoverdine (19000 L/mol/cm) is lower than that of Fe^{3+} -pyoverdine
 341 complexes (26000 L/mol/cm) and greater than that of Al^{3+} -pyoverdine complexes (18000 L/mol/cm).
 342 Considering that in our systems, it was not possible to experimentally distinguish between the free
 343 pyoverdine, Fe^{3+} -pyoverdine and Al^{3+} -pyoverdine fractions, approximate values of total pyoverdine

344 concentrations in solution were calculated by using the molar extinction coefficient of free pyoverdine
345 given in Equation (3). These concentrations are plotted in function of the experimental conditions on
346 Figure 4.
347



348
349 Figure 4: Apparent concentrations in pyoverdine in HAPLO, MORB2 and MORB3 systems after 6 days of experiments
350 with/without dialysis (D) membranes.
351

352 In the case of the HAPLO glass (Figure 3c, Figure 4), high amounts of pyoverdine are detected in the
353 incubated medium, independently of the presence or absence of dialysis bags, with pyoverdine
354 concentrations after 6 days greater than or equal to 80 µM and 70 µM, respectively. These results are
355 very close to those obtained in the control experiments (Figure 3d), as both control and HAPLO
356 systems are Fe-free systems.

357
358 In the case of the MORB2 glass, the absorbance spectra presented in Figure 3b show similar
359 intensities (corresponding to pyoverdine concentration of ~10 µM, Figure 4) either at 3 and 6 days
360 suggesting that isolating this glass from the bacterial suspension does induce any significant changes
361 regarding the siderophore production. The spectra recorded at days 3 and 6 mainly differ by a change
362 in the shape of the absorption peak of pyoverdine species (Figure 3b). In order to highlight this
363 evolution, more absorbance spectra of the cultures recorded until 15 days of incubation for MORB2
364 experiments are represented in Figure 5.

365

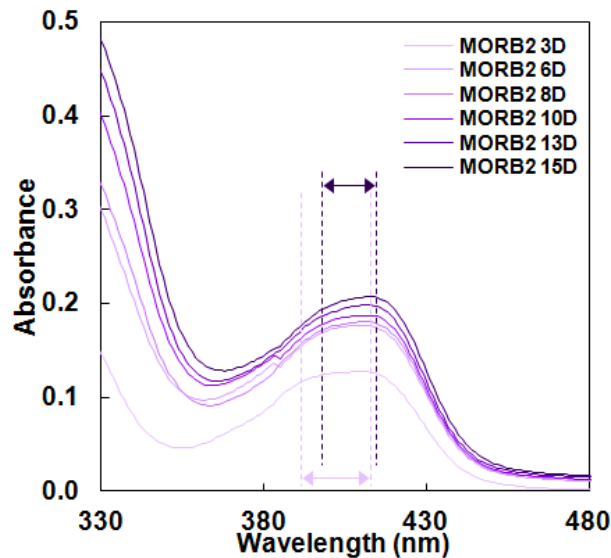


Figure 5: Absorption spectra of the MM9 6.5 bacterial medium in the presence of the MORB2 glass contained in dialysis bags, recorded at 3, 6, 8, 10, 13 and 15 days (D) of the experiments. The arrows materialize the shift of the pyoverdine absorption region toward higher wavelengths with time.

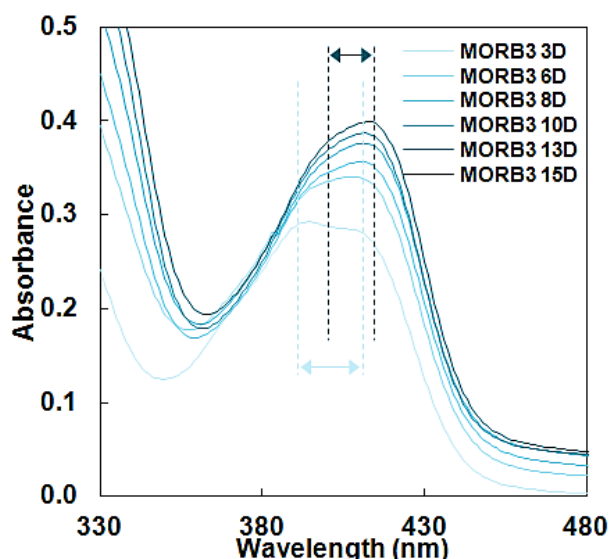
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The spectra show a constantly-evolving signal with time (shifted toward higher wavelengths and changing into a one-shoulder peak). At near neutral pH condition, this particular shape characterizes the presence of Fe^{3+} -pyoverdine complexes (Albrecht-Gary et al., 1994). Such an evolution could attest the continuous formation of Fe^{3+} -pyoverdine complexes in solution. The increase in Fe concentration after 5.7 days of MORB2 experiment (Figure 2) was suggesting that no more Fe was incorporated inside the cells once Fe homeostasis conditions were reached. This result shows now that the Fe fraction continuously released from the glass over the course of the experiments (from day 6 to 15) must be, once oxidized into Fe^{3+} in contact with the oxygenated medium, progressively chelated by the free-pyoverdine molecules remaining in solution.

380

Finally, Figures 3a and 4 show that the presence of dialysis membranes in the MORB3 experiments impacts on siderophore production. If pyoverdine is not detected in the no-dialysis systems, isolating the glass from the bacterial suspension results in a significant production of siderophore in the medium over the course of the experiments. Pyoverdine concentration after 6 days of experiments in the incubated medium is estimated to be lower than or equal to $20 \mu\text{M}$. Figure 6 illustrates the change in the shape and the position of the absorption peak of pyoverdine thorough 15 days of experiments.

387



388 Figure 6: Absorption spectra of the MM9 6.5 bacterial medium in the presence of the MORB2 (a) and MORB3 (b) glasses
 389 contained in dialysis bags, recorded at 3, 6, 8, 10, 13 and 15 days (D) of the experiments. The arrows materialize the shift of the
 390 pyoverdine absorption region toward higher wavelengths with time.
 391

392
 393 As observed for the MORB2 glass, the two-shoulder peak characteristic of free-pyoverdine
 394 continuously evolve into the single absorption peak of Fe³⁺-pyoverdine complexes, leading to the
 395 same conclusion than those made for the reduced glass.

396
 397 A previous study performed in identical experimental conditions showed that without any separation
 398 between the glass and the bacterial suspension, the Fe fraction abiotically released in solution was
 399 lower than the Fe fraction measured within the bacterial cells (Perez et al., 2016). This result suggests
 400 that, in the presence of the strain, the Fe fraction naturally released from the glass into the liquid
 401 medium is not sufficient to sustain bacterial needs and that, in consequence, more Fe is mobilized
 402 from the glass through a non-identified biological process. In the same previous study, the absence of
 403 siderophore in the system indicated that pyoverdine molecules were not responsible for enhancing the
 404 Fe release. In this work, the results highlight that separating the glass from the bacterial suspension
 405 induces a high pyoverdine production by the cells. This production of pyoverdine in presence of
 406 dialysis membranes is a good indicator of a high bacterial need in Fe, not observed when the bacteria
 407 can access to the glass. This demonstrates that the biological process that enhance the Fe release
 408 consists in direct glass surface/bacteria interactions.

409 1.3. Evidence of direct bacteria/glass interactions

410
 411
 412 The SEM images of the surface of a MORB3 glass sample incubated in the presence of *P. aeruginosa*
 413 without dialysis bags (direct contact) are presented in Figure 7.

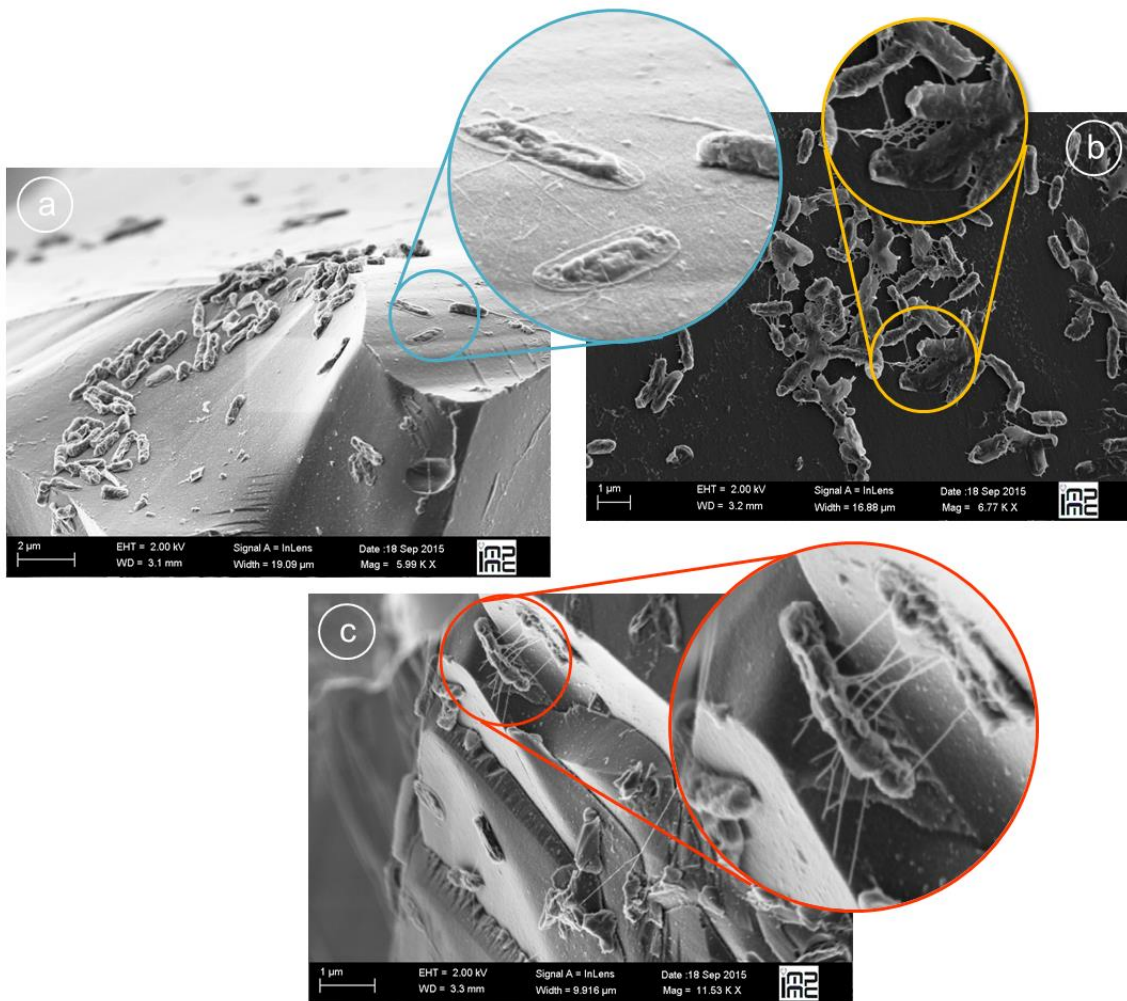


Figure 7: SEM images of the MORB3 glass surface after 15 days of experiments (without dialysis bags).

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If bacteria do not seem to massively colonize the surface, they are however homogeneously present on the glass grains. The images show that the cells are strongly attached to the glass, as dead cell residues remain at the surface after the CPD protocol (Figure 7a). They also show that biofilm communities start to develop: bacteria produce filamentous structures that allow them to progressively connect to each other (Figure 5b) and to the surface of the glass (Figure 7c). Very similar biofilm patterns were presented in the work on *Pseudomonas stutzeri* of Sudek et al. (2017) between 4 and 7 days of incubation with basaltic glass samples.

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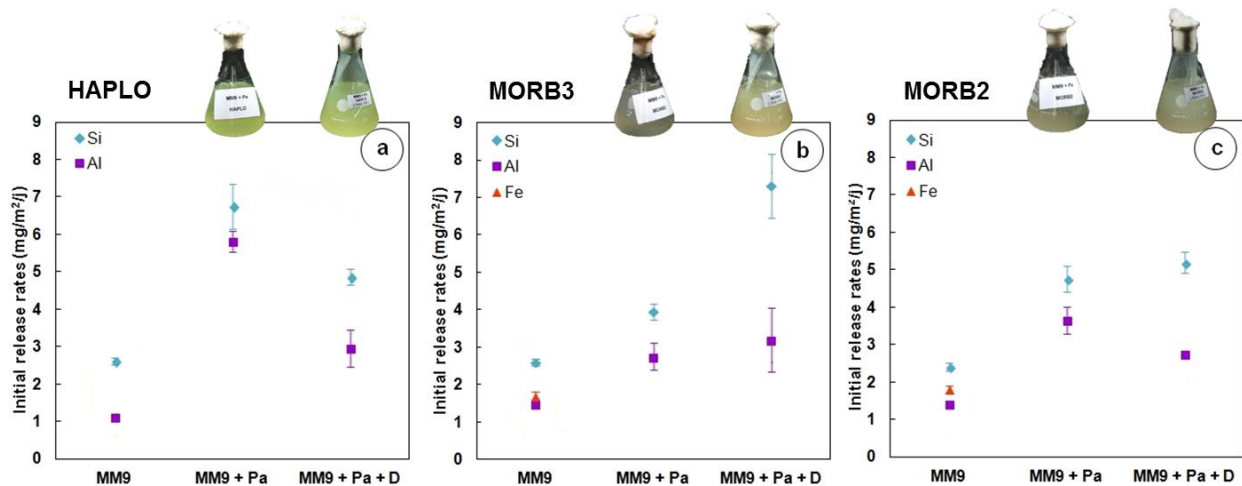
Impact on dissolution kinetics

426

Apparent dissolution rates calculated in abiotic control experiments and in *P. aeruginosa* (with glass confinement or not) based on Si, Al and Fe concentrations measured in the liquid medium are plotted in Figure 8. As free pyoverdine molecules are fluorescent and the formation of complexes between pyoverdine and Fe^{3+} cations is known to induce fluorescence quenching (Albrecht-Gary et al., 1994), the color of the medium after 6 days of experiments is visually represented on the graphics in order to

431

432 help the reader to remember the pyoverdine production in each biotic system and its possible
433 interactions with Fe.
434



435
436 Figure 8: Initial release rates of the HAPLO (a), MORB3 (b) and MORB2 (c) glass confined in dialysis bags and immersed in *P.*
437 *aeruginosa* culture (MM9 + Pa + D) in comparison with initial rates determined in sterile MM9 and in *P. aeruginosa* cultures
438 (MM9 + Pa) (Perez et al., 2016). The rates were calculated using the average of NL values for both sets of duplicates
439 experiments. Error bars are equal to $\pm 2SD$.
440

441 Considering that a fraction of dissolved Al or Fe from the glasses is likely to be trapped within the
442 bacterial cells or adsorbed at their surface (and in consequence not taken in account by ICP-OES
443 measurements), the dissolution rates of the three glasses will be mainly discussed on the basis on Si
444 data.

445 First of all, the presence of bacterial life in the aqueous medium (MM9) has a positive impact on the
446 global dissolution of the three glasses: the Si release rates are enhanced by a factor of 1.5 to 3
447 depending on the experimental conditions, whether the glasses are contained in dialysis bag or not.

448 The dissolution rates calculated from Si concentration of the three glasses are enhanced by
449 comparison with abiotic systems in the order: MORB3 < MORB2 \approx MORB2-dialysis < HAPLO-dialysis
450 < HAPLO < MORB3-dialysis.

451 All these results show that:

452 1/ the presence of pyoverdine molecules has the most significant catalyzing effect on the dissolution.
453 This effect is notably evidenced by the comparison between dialysis and no dialysis-using MORB3
454 experiments. Without dialysis membranes, the lowest bacterial impact on the dissolution kinetics
455 among all the tested glasses ($4.0 \pm 0.2 \text{ mg.m}^{-2}.\text{d}^{-1}$ for Si, Figure 8b) is correlated with the non-detection
456 of pyoverdine in the system (Figure 4). These experiments show that the moderate enhancement
457 relative to the experiment without bacteria is caused by a direct bacteria/glass interaction. By contrast,
458 the dialysis-membrane-using experiments, in which low but significant amounts of pyoverdine (around
459 $20 \mu\text{M}$) are detected in solution, are characterized by the highest Si release rates ($7.3 \pm 0.8 \text{ mg.m}^{-2}.\text{d}^{-1}$
460 for Si, Figure 8b).

461 2/ the relative effect of pyoverdine correlates well with the complex stability constants of the
462 siderophore with the cations in presence ($K^{\text{Fe}^{2+}} < K^{\text{Al}^{3+}} < K^{\text{Fe}^{3+}}$) (Albrecht-Garyet al., 1994; Chen et al.,
463 1994; Hernlem et al., 1996; Szabo et Karkas, 2011). In particular, the siderophore has the most
464 significant effect on the dissolution rates when the cations that are the most likely to form complexes
465 with it are involved as network-forming elements in the glass structure (Al^{3+} , Fe^{3+}). This confirms that
466 complexation reactions between pyoverdine and structural Al and/or Fe promote the dissolution either
467 by a surface-controlled mechanism (Cocozza et al., 2003) or by increasing the solubility of these
468 metallic elements in solution (Oelkers and Schott, 1998). Such siderophore-promoted dissolution has
469 already been well investigated in studies dedicated to the interactions between desferrioxamine-B and
470 Fe(III)-bearing minerals/glass (Reichard et al., 2005; Wolff-Boenisch and Traina, 2007; Cervini-Silva,
471 2008; Haack et al., 2008; Dehner et al., 2010; Perez et al., 2015). However, there is only a limited
472 number of existing studies exploring the promotion of mineral/glass dissolution by other siderophores
473 and notably pyoverdine (Ferret et al., 2014; Parello et al., 2016; Perez et al., 2016). The results of this
474 work contribute to show the efficiency of pyoverdine in promoting the dissolution of an Fe(III) or Al(III)-
475 bearing glass. For the MORB2 glass, the very similar results in terms of siderophore production
476 (Figure 4) and dissolution rates (Figure 8) from dialysis-membrane to no dialysis-membrane-using
477 experiments confirm that the bacteria-promoting process mainly takes place in solution. This
478 corroborates what was assessed in a previous study by Perez et al. (2016): during the MORB2 glass
479 dissolution, the release of Fe first as Fe^{2+} cations in solution induces the production of pyoverdine
480 molecules. The Fe^{2+} cations being quickly oxidized into Fe^{3+} cations in contact with the oxygenated
481 liquid medium are preferentially chelated by the pyoverdine molecules. Although the change of Fe
482 speciation in solution favors the release of Fe^{2+} , its structural role (network-modifier cation) can explain
483 the lower impact of this preferential release on the global dissolution kinetics of the glass.

484 3/ the correlation between the effect of pyoverdine on the dissolution kinetics and the complex stability
485 constants does not seem to depend on the concentrations of pyoverdine in the systems. As an
486 example, pyoverdine has the highest impact on dissolution rates of the MORB3 glass with a maximum
487 concentration in the medium equal to 20 μM (Figure 4). By contrast, it has a lower impact on the
488 dissolution kinetics of the HAPLO glass although it is continuously produced by the cells and its
489 concentration approximates 70-80 μM after 6 days of dissolution (Figure 4). Moreover, isolating the
490 HAPLO glass from the bacterial suspension induces a more intense production of pyoverdine (80 μM
491 versus 70 μM) but these experiments are characterized by lower global dissolution rates of the glass
492 ($4.9 \pm 0.2 \text{ mg.m}^{-2}.\text{d}^{-1}$ versus $6.8 \pm 0.6 \text{ mg.m}^{-2}.\text{d}^{-1}$ for Si, Figure 8a). These results echo to several
493 studies evidencing a potential “threshold effect” on dissolution rates concerning the abundance of
494 chelating molecules in solution (Cheah et al., 2003; Wolff-Boenisch and Traina, 2006). A slight
495 decrease in the dissolution rate of basalts was even observed by Neaman et al. (2005) by increasing
496 the concentration of organic ligands in solution.

497

498

499

500 **Conclusion**

501 This work attests of the impact of bacterial life on the dissolution kinetics of three synthetic basaltic
502 model glasses. Isolating the samples from the bacterial suspension was essential to target more
503 precisely the bacterial impact on the dissolution mechanisms/kinetics of basaltic glass and quantifying
504 the different types of bacterial contributions. A direct access to the surface of an Fe(III)-bearing glass
505 is shown to inhibit the production of siderophore by the bacteria and consequently to minimize the
506 bacteria-promoted effect, as the dissolution rates are increased by 1.5 times compared to sterile
507 experiments. By contrast, isolating the glass from the bacterial suspension with dialysis bags triggers
508 the biosynthesis of siderophore, stopped as soon as the bacteria reach a stationary phase of growth.
509 This limited quantity of pyoverdine molecules in solution (around 20 μM) is sufficient to significantly
510 improve the dissolution rate of the MORB3 glass, which is increased by 3 times compared to sterile
511 experiments. By contrast, the dissolution kinetics of an Fe(II)-bearing glass incubated with *P.*
512 *aeruginosa* are not impacted by the limited access of bacteria to the glass in dialysis experiments:
513 small amounts of siderophore are produced in both cases and the dissolution rate is increased by 2
514 times compared to sterile experiments. Finally, the Fe-starved bacteria incubated in the presence of
515 the HAPLO glass are shown to continuously produce pyoverdine. As most siderophores, pyoverdine
516 has also a high affinity with Al^{3+} and the dissolution seems to be promoted by the preferential release
517 of Al from the glass induced by the formation of Al^{3+} -pyoverdine complexes.

518 This work is a new contribution regarding the high affinity of microorganisms for basaltic glasses as an
519 Fe-source. In this case Fe is the most likely used by the ferri-reductive model strain *P. aeruginosa* in
520 its Fe^{3+} form as an electron sink. Its quick uptake by *P. aeruginosa* is shown to significantly carry
521 benefit to the growth of the bacteria. Moreover, this study brings new evidences of the positive impact
522 on the dissolution kinetics of these glasses, of the strategies developed by the bacteria to acquire Fe,
523 depending on the form under which it is primarily available to them: 1/ as structural Fe(II) or Fe(III)
524 within the glass network, directly accessible (or not) to them depending on the presence of dialysis-
525 membranes or 2/ in solution first as Fe^{2+} ions or forming soluble complexes with phosphates initially
526 present within the liquid medium. For each Fe source, one bacterial response was identified and each
527 response was shown to impact differently on the dissolution kinetics. Among these bacterial
528 responses, siderophore production was shown to particularly regulate the dissolution rates. In a
529 basaltic environment where Fe solubility is low and where siderophore-producing strains are
530 commonly detected, the biosynthesis of these complexing agents operates as a mechanism
531 developed to ensure nutritional Fe-availability.

532

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535

538 Akafia M.M., Harrington J.M., Bargar J.R. and Duckworth O.W. (2014) Metal oxyhydroxide dissolution
539 as promoted by structurally diverse siderophores and oxalate. *Geochim. Cosmochim. Acta* **141**, 258-
540 269.

541 Ahmed E. and Holmström S. J. M. (2015) Microbe-mineral interactions: The impact of surface
542 attachment on mineral weathering and element selectivity by microorganisms. *Chem. Geol.* **403**, 13-
543 23.

544 Albrecht-Gary A., Blanc S., Rachel N., Ocaktan A., Mohamed A. and Abdallah A. (1994) Bacterial iron
545 transport: coordination properties of pyoverdine PaA, a peptidic siderophore of *Pseudomonas*
546 *aeruginosa*. *Inorg. Chem.* **33**, 6391-6402.

547 Berner R.A., Lasaga A.C., Garrels R.M. (1983) The carbonate-silicate geochemical cycle and its effect
548 on atmospheric carbon dioxide over the past 100 million years. *Am. J. Sci.* **283**, 641-683.

549 Bray A.W., Oelkers E.H., Bonneville S., Wolff-Boenisch D., Potts N.J., Fones G. and Benning L.G.
550 (2015) The effect of pH, grain size, and organic ligands on biotite weathering rates. *Geochim.*
551 *Cosmochim. Acta* **164**, 127-145.

552 Callow B., Falcon-Suarez I., Ahmed S. and Matter J. (2018) Assessing the carbon sequestration
553 potential of basalt using X-ray micro-CT and rock mechanics. *International Journal of Greenhouse Gas*
554 *Control* **70**, 146-156.

555 Cheah S., Kraemer S., Cervini-Sila J. and Sposito G. (2003) Steady-state dissolution kinetics of
556 goethite in the presence of desferrioxamine-B and oxalate ligands: implications for the microbial
557 acquisition of iron. *Chem. Geol.* **198**, 63-75.

558 Cockell C. S. et al. (2010) Microbial endolithic colonization and the geochemical environment in young
559 seafloor basalts. *Chem Geol* **279**, 17–30.

560 Cocozza C., Tsao C.C.G., Cheah S.F., Kraemer S.M., Raymond K.N., Miano T.M. and Sposito G.
561 (2002) Temperature dependence of goethite dissolution promoted by trihydroxamate siderophores.
562 *Geochim. Cosmochim. Acta* **66**, 431-438.

563 Cornu J.Y., Huguenot D., Jézéquel K., Lollier M. and Lebeau T. (2017) Bioremediation of copper-
564 contaminated soils by bacteria. *World Journal of Microbiology and Biotechnology* **33**.

565 Crovisier J., Advocat T. and Dussossoy J.L. (2003) Nature and role of natural alteration gels formed
566 on the surface of ancient volcanic glasses, natural analogs of waste containment glasses. *Journal of*
567 *Nuclear Materials* **321**, 91-109.

568 Daux V., Guy C., Advocat T., Crovisier J. L. and Stille P. (1997) Kinetic aspects of basaltic glass
569 dissolution at 90°C: role of aqueous silicon and aluminium. *Chem. Geol.* **142**, 109-26.

570 Dessert C., Dupré B., Gaillardet J., François L.M. and Allègre C.J. (2003) Basalt weathering laws and
571 the impact of basalt weathering on the global carbon cycle. *Chem. Geol.* **202**, 257-273.

572 Drever J. and Stillings L. (1997) The role of organic acids in mineral weathering. *Colloids and Surfaces*
573 **120**, 167-181.

574 Ferret C. (2014) Siderophore-promoted dissolution of smectite by fluorescent *Pseudomonas*.
575 *Environmental Microbiology Reports* **6**, 459-647.

- 576 Flaathen T., Gislason S. and Oelkers E. (2010) The effect of aqueous sulphate on basaltic glass
577 dissolution rate. *Chem. Geol.* **8**, 345-354.
- 578 Gislason S. R. and Oelkers E. H. (2003) Mechanism, rates and consequences of basaltic glass
579 dissolution: II. An experimental study of the dissolution rates of basaltic glass as a function of pH and
580 temperature. *Geochim. Cosmochim. Acta* **67**, 3817-3832.
- 581 Gordon S.J. and Brady P.V. (2002) In situ determination of long-term basaltic glass dissolution in the
582 unsaturated zone. *Chem. Geol.* **190**, 113-122.
- 583 Guy C. and Schott J. (1989) Multisite surface reaction versus transport control during the
584 hydrolysis of a complex oxide. *Chem. Geol.* **78**, 181-204.
- 585
586 Hart R. (1970) Chemical exchange between sea water and deep ocean basalts. *Earth Planet Sci Lett*
587 **9**, 269-279.
- 588
589 Henri P., Rommevaux-Jestin C., Lesongeur F. and Ménez B. (2016) Structural iron (II) of basaltic
590 glass as an energy source for Zetaproteobacteria in an Abyssal Plain Environment, off the Mid Atlantic
591 Ridge. *Frontiers in Microbiology* **6**, 1518.
- 592 Hopf J., Langenhorst F., Pollok K., Merten D. and Kothe E. (2009) Influence of microorganisms on
593 biotite dissolution: an experimental approach. *Chemie der Erde – Geochemistry* **69**, 45-56.
- 594 Hutchens E. (2009) Microbial selectivity on mineral surfaces: possible implications for weathering
595 processes. *Fungal Biology Reviews* **23**, 115-121.
- 596 Kraemer S., Cheah S., Zapf R., Xu J., Raymond K. and Sposito G. (1999). Effect of hydroxamate
597 siderophores on Pb(II) adsorption and Fe release from goethite. *Geochim. Cosmochim. Acta* **63**, 3003-
598 3008.
- 599
600 Lazo D.E., Dyer L.G. and Alorro R.D. (2017) Silicate, phosphate and carbonate mineral dissolution
601 behavior in the presence of organic acids: a review. *Mineral Engineering* **100**, 115-123.
- 602 Liermann L. J., Kalinowski B. E., Brantley S. L. and Ferry J. G. (2000) Role of bacterial siderophores in
603 dissolution of hornblende. *Geochim. Cosmochim. Acta* **64**, 587-602.
- 604 Martinez-Luevanos A., Rodriguez-Delgado M., Uribe-Salas A., Carillo-Pedroza F. and Osuna-Alarcon
605 J. (2011) Leaching kinetics of iron from low grade kaolin by oxalic acid solutions. *Applied Clay Science*
606 **51**, 473-477.
- 607 Mason O., Di Meo-Savoie C., van Nostrand J., Zhou J., Fisk M. and Giovannoni J. (2009) Prokaryotic
608 diversity, distribution and insights into their role in biogeochemical cycling in marine basalts, *ISME J* **2**,
609 231-242.
- 610 Meyer J. M. and Abdallah M. A. (1978) The fluorescent pigment of *Pseudomonas fluorescens*:
611 biosynthesis, purification and physicochemical properties. *J. Gen. Microbiol.* **107**, 319-328.
- 612 Morin G.P., Vigier N., Verney-Carron A. (2015) Enhanced dissolution of basaltic glass in brackish
613 waters: Impact on biochemical cycles. *Earth and Planetary Science Letters* **417**, 1-8.
- 614 Oelkers E. (2001) General kinetic description of multioxide silicate mineral and glass dissolution.
615 *Geochim. Cosmochim. Acta* **65**, 3703-3719.
- 616 Oelkers E. H. and Gislason S. R. (2001) The mechanism, rates and consequences of basaltic glass
617 dissolution: I. An experimental study of the dissolution rates of basaltic glass as a function of aqueous
618 Al, Si and oxalic acid concentration at 25°C and pH 3 and 11. *Geochim. Cosmochim. Acta* **65**, 3671-
619 3681.

620 Oelkers E. and Schott, J. (1998) Does organic acid adsorption affect alkali feldspar dissolution rates?
621 *Chem. geol.* **151**, 235-245.
622

623 Olsen A. and Rimstidt D. (2008) Oxalate-promoted forsterite dissolution at low pH. *Geochim.*
624 *Cosmochim. Acta* **72**, 1758-1766.
625

626 Orcutt B., Sylvan J., Knab N. and Edwards K. (2011) Microbial ecology of the dark ocean above, at
627 and below the seafloor. *Microbiology and Molecular Biology Review* **75**, 361-422.
628

629 Parrello D., Zegeye A., Mustin C. and Billard P. (2016) Siderophore-mediated iron dissolution from
630 nontronites is controlled by mineral crystallochemistry. *Front. Microbiol.* **7**, Article 423.
631

632 Parruzot B., Jollivet P., Rébiscoul D. and Gin S. (2015) Long-term alteration of basaltic glass :
633 Mechanisms and rates. *Geochim. Cosmochim. Acta* **154**, 28-48.
634 Techer I., Advocat T., Lancelot J. and Liotard J. (2000) Basaltic glass : Alteration mechanisms and
635 analogy with nuclear waste glasses. *J. Nucl. Mater.* **282**, 40-46.

636 Perez A., Rossano S., Trcera N., Verney-Carron A., Huguenot D., van Hullebusch E. D., Catillon G.,
637 Razafitianamaharavo A. and Guyot F., (2015) Impact of iron chelators on short-term dissolution of
638 basaltic glass. *Geochim. Cosmochim. Acta* **162**, 83-98.

639 Perez A., Rossano S., Trcera N., Huguenot D., Fourdrin C., Verney-Carron A., van Hullebusch E. and
640 Guyot F. (2016) Bioalteration of synthetic Fe(III)-, Fe(II)-bearing basaltic glasses and Fe-free glass in
641 the presence of the heterotrophic bacteria strain *Pseudomonas aeruginosa*: impact of siderophores.
642 *Geochimica et Cosmochimica Acta* **188**, 147-162.
643

644 Santelli C., Edgcomb V., Bach W. and Edwards K. (2009) The diversity and abundance of bacteria
645 inhabiting seafloor lavas positively correlates with rock alteration. *Environmental Microbiology* **11**, 86-
646 98.
647

648 Schalk IJ and Guillon L (2013) Fate of ferrisiderophores after import across bacterial outer
649 membranes: different iron release strategies are observed in the cytoplasm or periplasm depending on
650 the siderophore pathways. *Amino Acids* **44**, 1267-1277.

651 Shirokova L. S., Bénézech P., Pokrovsky O. S., Gerard E., Menez B. and Alfredsson H. (2012) Effect
652 of the heterotrophic bacterium *Pseudomonas reactans* on olivine dissolution kinetics and implications
653 for CO₂ storage in basalts. *Geochim. Cosmochim. Acta* **80**, 30–35.

654 Schwyn B. and Neilands J. B. (1987) Universal Chemical-Assay for the Detection and Determination
655 of Siderophores. *Anal. Biochem.* **160** 47–56.

656 Spivack A. J. and Staudigel H (1994) Low-temperature alteration of the upper oceanic crust and the
657 alkalinity budget of seawater. *Chem Geol* **115**, 239–247.

658 Staudigel H., Hart S.R. (1983) Alteration of basaltic glass: mechanisms and significance for the
659 oceanic crust seawater budget. *Geochim. Cosmochim. Acta* **47**, 337–350.

660 Stillings L. L., Drever J.I., Brantley S. L and Sun Y. (1995) Rates of feldspars dissolution at pH 3-7 with
661 0-8 mM oxalic acid. *Chem. geol.* **132**, 79-89.

662 Stockmann G. J., Shirokova L. S., Pokrovsky O. S., Bénézech P., Bovet N., Gislason S. R. and
663 Oelkers E. H. (2012) Does the presence of heterotrophic bacterium *Pseudomonas reactans* affect
664 basaltic glass dissolution rates? *Chem. geol.* **296-297**, 1-18.
665

666 Stroncik N. and Schmincke H. (2001) Evolution of palagonite : Crystallization, chemical changes, and
667 element budget. *Geochem. Geophys. Geosy.* **2**, 365-370.
668

669 Sudek L.A., Templeton A.S., Tebo B.M. and Staudigel H. (2009). Microbial ecology of Fe(hydroxide)
670 mats and basaltic rock from Vailulu'u Seamount, American Samoa. *Geomicrobiol. J.* **26**, 581-596.
671

672 Sudek L.A., Wanger G., Templeton A.S., Staudigel H. and Tebo B.M. (2017) Submarine Basaltic glass
673 colonization by the heterotrophic Fe(II)-oxidizing and siderophore-producing deep-sea bacterium
674 *Pseudomonas stutzeri* VS-10: The potential role of basalt in enhancing growth. *Front. Microbiol.*

675 Sverdrup H. (2009) Chemical weathering of soil minerals and the role of biological processes. *Fungal*
676 *Biology Reviews* **23**, 94-100.

677 Techer I., Lancelot J., Clauer N., Liotard J. and Advocat T. (2001b) Alteration of a basaltic glass in an
678 argillaceous medium : The Salagou dike of the Lodève Permian Basin (France). Analogy with an
679 underground nuclear waste repository. *Geochim. Cosmochim. Acta* **65**,1071-1086.
680

681 Templeton A. S., Staudigel H. and Tebo B. M. (2005) Diverse Mn(II)-Oxidizing Bacteria Isolated from
682 Submarine Basalts at Loihi Seamount. *Geomicrobiol* **22**, 127–139.

683 Thorseth I., Torsvik T., Torsvik V., Daae F. and Pedersen R. (2001) Diversity of life in ocean floor
684 basalt. *Earth and Planetary Science Letters* **194**, 31-37.
685

686 Torres M.A., West A.J. and Nealson K. (2014) Microbial acceleration of olivine dissolution via
687 siderophore production. *Procedia Earth and Planetary Science* **10**, 118-122.
688

689 Ullman W. J., Kirchman D. L., Welch S. A. and Vandevivere P. (1996) Laboratory evidence for
690 microbially mediated silicate mineral dissolution in nature. *Chem. Geol.* **132**, 11-17.

691 Verney-Carron A., Viguiet N. and Millot R. (2011) Experimental determination of the role of diffusion
692 on Li isotope fractionation during basaltic glass weathering. *Geochim. Cosmochim. Acta* **75**, 3452-
693 3468.

694 Wang Q., Gao S., Ma X., Mao X., He L. and Sheng X. (2018) Distinct mineral weathering effectiveness
695 and metabolic activity between mineral-weathering bacteria *Burkholderia metallica* F22 and
696 *Burkholderia phytofirmans* G34. *Chemical Geology*, in press.

697 Wang X., Li Q., Hu H., Zhang T. and Zhou Y. (2005) Dissolution of kaolinite induced by citric, oxalic
698 and malic acids. *Journal of Colloids and Interface Sciences* **290**, 481-488.

699 Watteau F. and Berthelin J. (1994) Microbial dissolution of iron and aluminium from soil minerals:
700 Efficiency and specificity of hydroxamate siderophores compared to aliphatic acids. *Soil Biology* **30**, 1-
701 9.

702 Welch S. and Ullman W. (1992) The effect of organic acids on plagioclase dissolution rates and
703 stoichiometry. *Geochim. Cosmochim. Acta* **57**, 2725-2736.

704

705 Wolff-Boenisch D., Gislason S. R., Oelkers E. H., and Putnis C. V. (2004a) The dissolution rates of
706 natural glasses as a function of their composition at pH 4 and 10.6, and temperatures from 25 to
707 74°C. *Geochim. Cosmochim. Acta* **68**(23), 4843-4858.

708 Wolff-Boenisch D., Gislason S.R. and Oelkers E.H. (2004b) The effect of fluoride on the dissolution
709 rates of natural glasses at pH 4 and 25°C. *Geochim. Cosmochim. Acta* **68**, 4571-4582.

710 Wolff-Boenisch D. and Traina S.J. (2007) The effect of desferrioxamine B, enterobactin, oxalic acid,
711 and Na-alginate on the dissolution of uranyl-treated goethite at pH 6 and 25°C. *Chem. Geol.* **243**, 357-
712 368.

713 Wolff-Boenisch D., Wenau S., Gislason S. R. and Oelkers E. H. (2011) Dissolution of basalts and
714 peridotite in seawater, in the presence of ligands, and CO₂: implications for mineral sequestration of
715 carbon dioxide. *Geochim. Cosmochim. Acta* **75**, 5510-5525.

716