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Improving estuarine net flux estimates for dissolved cadmium export at the annual timescale: Application to the Gironde Estuary

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abstract

Dissolved Cd (Cd_D) concentrations along the salinity gradient were measured in surface water of the Gironde Estuary during 15 cruises (2001–2007), covering a wide range of contrasting situations in terms of hydrology, turbidity and season. During all situations dissolved Cd concentrations displayed maximum values in the mid-salinity range, reflecting Cd addition by chloride-induced desorption and complexation. The daily net Cd_D fluxes from the Gironde Estuary to the coastal ocean were estimated using Boyle's method. Extrapolating Cd_D concentrations in the high salinity range to the freshwater end member using a theoretical dilution line produced 15 theoretical Cd concentrations (Cd_D^0), each representative of one distinct situation. The obtained Cd concentrations were relatively similar ($\approx 201 \pm 28 \text{ ng L}^{-1}$) when freshwater discharge Q was $> 500 \text{ m}^3 \text{ s}^{-1}$ ($508 \leq Q \leq 2600 \text{ m}^3 \text{ s}^{-1}$), but were highly variable ($340 \pm 80 \text{ ng L}^{-1}$; $247\text{--}490 \text{ ng L}^{-1}$) for low discharge situations ($169 \leq Q \leq 368 \text{ m}^3 \text{ s}^{-1}$). The respective daily Cd_D net fluxes were $5\text{--}39 \text{ kg day}^{-1}$, mainly depending on freshwater discharge. As this observation invalidates the existing method of estimating annual Cd_D net fluxes, we proposed an empirical model, using representative Cd_D^0 values and daily freshwater discharges for the 2001–2007 period. Subsequent integration produced reliable Cd_D net flux estimates for the Gironde Estuary at the annual timescale that ranged between $3.8\text{--}5.0 \text{ t a}^{-1}$ in 2005 and $6.0\text{--}7.2 \text{ t a}^{-1}$ in 2004, depending on freshwater discharge. Comparing Cd_D net fluxes with the incoming Cd_D fluxes suggested that the annual net Cd_D addition in the Gironde Estuary ranged from 3.5 to 6.7 t a^{-1} , without any clear temporal trend during the past seven years. The annual Cd_D net fluxes did not show a clearly decreasing trend in spite of an overall decrease by a factor ≈ 6 in Cd gross fluxes during the past decade. Furthermore, in six years out of seven (except 2003), the annual Cd_D net fluxes even exceeded river borne total (dissolved + particulate) gross Cd fluxes into the estuary. These observations were attributed to progressive Cd desorption from both suspended particles and bottom sediment during various sedimentation–resuspension cycles induced by tidal currents and/or continuous dredging (navigation channel) and diverse intra-estuarine sources (wet deposition, urban sources, and agriculture). Provided that gross fluxes remain stable over time, dissolved Cd exportation from the Gironde Estuary to the coastal ocean may remain at the present level for the coming decade and the estuarine sedimentary Cd stock is forecast to decrease slowly.

1. Introduction

Hydrological, sedimentological and biogeochemical processes in estuarine systems strongly modify the quantity and the quality of river borne matter transported from the continent to the coastal ocean. Despite the need for reliable annual net flux data, estuarine net flux estimates are only available for the large minority of the global estuaries and cover very few, mostly dissolved, elements.

These estimates are usually valid for one particular day or distinct hydrological situation and, thus, extrapolation from the daily to the annual timescale suffers from many uncertainties due to hydrological variations and complex estuarine processes.

The Gironde Estuary (southwest France) is affected by historic metallic (e.g. Cd, Zn, Hg; Audry et al., 2004a; Schafer et al., 2006) contamination due to former Zn ore treatment in the industrial basin of Decazeville, that was stopped after a major pollution accident in 1986 (Jouanneau et al., 1999). Despite decreasing Cd inputs due to ongoing remediation efforts since the early 1990s, the Lot-Garonne River system still contributes up to 80% to the annual Cd gross fluxes into the Gironde Estuary (Schafer et al., 2002b; Audry et al., 2004a).

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Moreover, important Cd stocks in the Lot River reservoir sediments (≈ 200 t, Lapaquellerie et al., 1995) may be mobilized by flood events and/or dredging activities (e.g. Blanc et al., 1999; Audry et al., 2004b; Coynel et al., 2007). Numerous studies evidenced non-conservative behaviour of Cd along the salinity gradients in estuaries, bays or deltas (Boyle et al., 1982; Elbaz-Poulichet et al., 1987; Windom et al., 1988; Shiller and Boyle, 1991; Boutier et al., 1993; Chiffoleau et al., 1994; Wen et al., 1999; Tang et al., 2002). Typical profiles show a mid-salinity maximum in dissolved Cd (Cd_D) reflecting the balance between chloride-induced desorption/complexation processes and dilution by seawater (e.g. Comans and van Dijk, 1988; Elbaz-Poulichet et al., 1996). Chloride induced-desorption is probably the major phenomenon that govern Cd behaviour in estuaries. However, several studies (e.g. Florence, 1989; Baeyens et al., 1997; Waeles et al., 2005) have shown that organic ligands contribute also to keep Cd in solution in the high-salinity part of the estuaries. Dissolved Cd addition is particularly efficient in macrotidal estuaries due to relatively long particle residence times (Dyer, 1994). Several studies have reported Cd_D addition along the salinity gradient of the Gironde Estuary, suggesting that $\approx 90\%$ of the Cd gross fluxes are dissolved due to estuarine biogeochemical processes, thus increasing Cd availability to aquatic organisms (e.g. Jouanneau et al., 1990; Kraepiel et al., 1997; Boutier et al., 2000; Michel et al., 2000; Audry et al., 2007a,b; Lekhi et al., 2008). Although Cd concentrations in oysters from the Gironde Estuary have decreased from ≈ 100 mg g⁻¹ (dry weight) in the early 1980s to 10–26 mg g⁻¹ d.w. in 2006 (Claisse et al., 1992; RNO, French National Mussel Watch Program, 2006), they still are ≈ 25 times the average level measured along the entire French Atlantic and Mediterranean coasts (Boutier et al., 1989) and clearly exceed the current European consumption safety level for Cd in marine bivalves (5 mg g⁻¹, dry weight; CE No. 466/2001). As a consequence, oyster (bivalve) recovery and production are forbidden in the Gironde Estuary, i.e. this case of compromised/affected ecosystem health goes along with social and economic effects. Furthermore, part of the water and particles exported from the Gironde Estuary to the coastal ocean reach the Marennes-Oléron Bay, i.e. one of Europe's most important oyster production zones, where Cd concentrations in oysters are clearly higher than in other bays along the French coast (Heral et al., 1982; Boutier et al., 2000; IFREMER, 2002). Therefore, precise estimates of annual dissolved Cd export from the Gironde Estuary to the coastal zone are essential to assess Cd sources and inputs into the Marennes-Oléron Bay. Reliable Cd_D net flux estimates need concentration data in the estuarine high-salinity range, where Cd_D values follow a theoretical mixing line. Extrapolation of this dilution line to $S \approx 0$, produces a theoretical freshwater Cd_D concentration (Cd_D^0), that may be multiplied by the daily freshwater discharge (Q) to estimate daily Cd_D net flux (Boyle et al., 1974). Integration of 365 discrete daily Cd_D net fluxes would then provide precise annual flux estimates, but this "ideal" strategy also would be far too expensive and time-consuming. Accordingly, the existing estimates of Cd_D net fluxes in the Gironde Estuary and other estuaries are based on few measurement campaigns, each providing one daily ("snapshot") dataset, i.e. one Cd_D^0 value. Therefore, extrapolation of these daily Cd net flux estimates to the annual timescale may imply important uncertainty. Elbaz-Poulichet et al. (1987) estimated an annual net Cd_D flux of 14 t a⁻¹ (37 kg day⁻¹) based on one measurement campaign in October 1982. Boutier et al. (2000) estimated annual Cd_D net fluxes in 1991 to 11 t a⁻¹ (44 kg day⁻¹), based on one Cd_D^0 value, i.e. one measurement campaign. During five sampling campaigns in the hydrological year 1997–1998, Michel et al. (2000) obtained similar daily Cd_D net fluxes for contrasting seasons/discharges and concluded that daily Cd_D net fluxes (≈ 17 kg day⁻¹) out of the Gironde Estuary would be constant throughout the year. Since the early nineties, Cd gross fluxes into the Gironde estuary have

decreased probably due to ongoing remediation in the watershed and apparently changing hydrological conditions (Schafer et al., 2002a,b; Masson et al., 2006) suggesting profound changes in the Cd budget of the Gironde Estuary. There is, however, at present no estimate of recent annual Cd_D net fluxes in the Gironde Estuary taking into account potential interannual and seasonal variations.

This work is based on a unique data set of dissolved Cd distributions along the Gironde Estuary salinity gradient, obtained from 15 measurement campaigns over 7 years (2001–2007) covering contrasting hydrological situations (e.g. freshwater discharges, tidal coefficients) and seasons. The aim of this study is to provide recent data on Cd_D net fluxes and to evaluate the variability and long-term evolution of Cd_D export from the Gironde Estuary to the coastal ocean. Based on the relationship between Cd_D distribution along the salinity gradient and freshwater discharge we propose an improved method of extrapolating daily Cd_D net fluxes to the annual timescale. Additionally, we compare recent Cd_D addition and net fluxes to historic data to assess the response of the estuarine Cd budget to generally decreasing gross inputs.

2. Study area

The physical, geochemical and hydrological characteristics of the Gironde Estuary (≈ 170 km length, $\approx 80\,000$ km² watershed surface area; Fig. 1) have been well defined in several studies (Elbaz-Poulichet et al., 1984; Li et al., 1994; Sottolichio and Castaing, 1999; Schafer et al., 2002b). The Gironde Estuary has a mean annual freshwater discharge of ≈ 1000 m³ s⁻¹ with typical water and particle residence times of ≈ 20 –90 days and ≈ 1 –2 years, respectively (Jouanneau and Latouche, 1981). During the tidal cycle, ocean water fluxes at the estuary mouth are 30–40 times higher than fluvial inputs (Allen et al., 1977). Asymmetrical progression of the tidal wave toward the upstream estuary induces a pronounced maximum turbidity zone (MTZ), where concentrations of suspended particulate matter (SPM) exceed 1 g L⁻¹ in surface water and several hundreds of g L⁻¹ in bottom water. This MTZ is typically located in the low salinity region and migrates up and down estuary with seasonal river flow variations (Sottolichio and Castaing, 1999). Sediment resuspension generally occurs during erosion periods at mid-ebb and mid-flood while tidal slacks are sedimentation periods. As a consequence of sedimentation, the navigation channel (≈ 18 km²) of the estuary is continuously dredged by the Bordeaux Autonomous Harbour (P.A.B.). The dredged sediments are either dispersed in the water column on-site or transported downstream within the estuary and deposited in distinct zones, where they are dispersed and/or resuspended by tidal currents (Audry et al., 2007a,b).

3. Material and methods

3.1. Sampling

Samples were collected during 15 cruises within the Gironde Estuary between March 2001 and November 2007 along the salinity gradient onboard the RV "Côtes de la Manche" (INSU). These sampling campaigns cover a wide range of freshwater discharges (Table 1), which were calculated as the sum of the daily flows of the Garonne and Dordogne Rivers (Fig. 2). Sampling of surface and bottom water in the estuarine salinity gradient was performed with acid-cleaned Niskin Bottles, thoroughly rinsed with estuarine water from the site. The samples were immediately filtered through cellulose acetate syringe filters (Sartorius®, 0.2 µm porosity). Aliquots for trace metals measurements were transferred into 125 mL acid-cleaned polypropylene bottles, acidified (pH 1; HNO₃ Baker ultrapure) and stored in the dark at 4 °C.

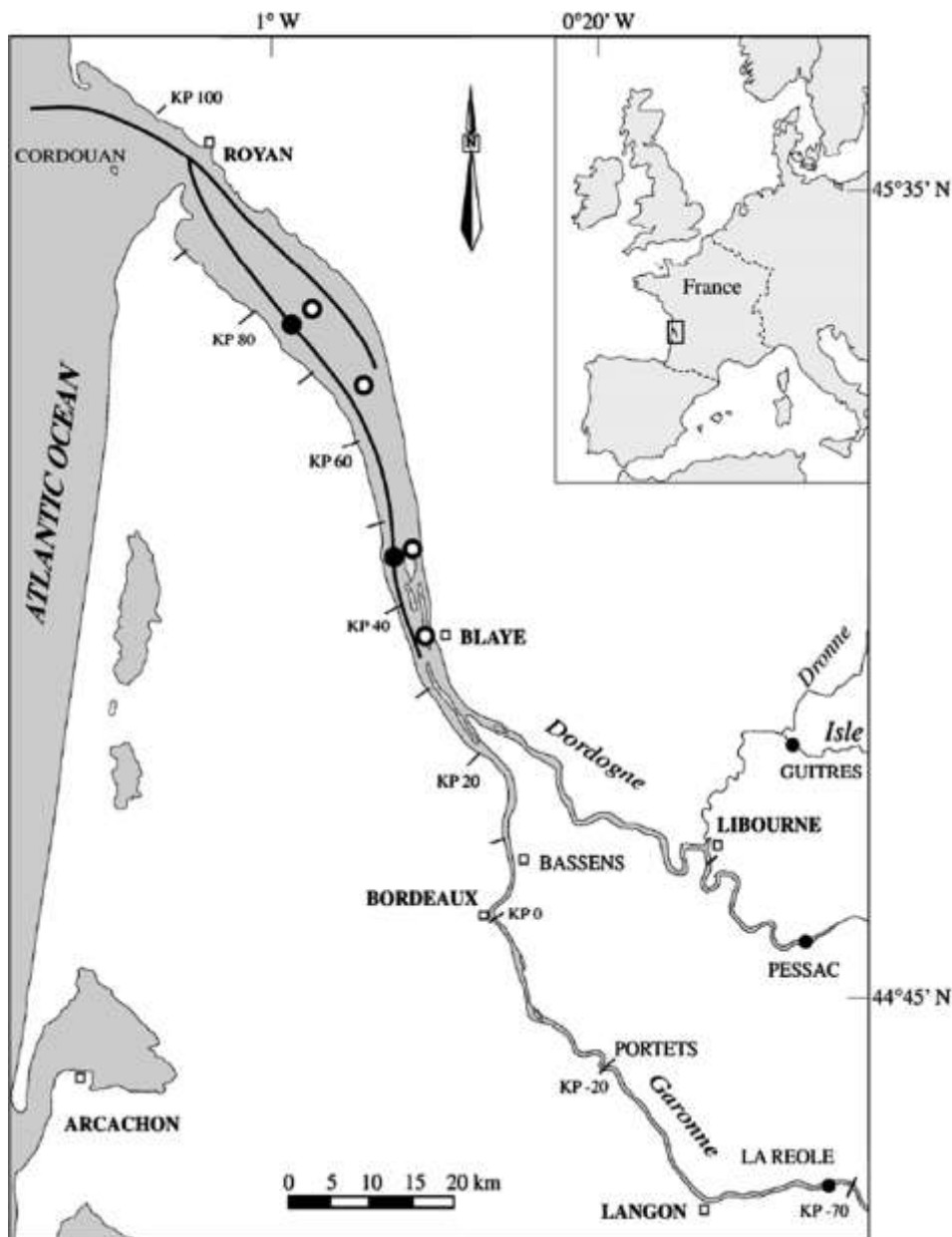


Fig. 1. Map of the Gironde Estuary. Kilometric Points (KP) ¼ distance (km) from the city of Bordeaux.

3.2. Sample analyses

Determination of SPM concentrations was performed by filtration using dry pre-weighed filters (Durieux glass-fibre, 0.7 µm porosity) as described elsewhere (e.g. Audry et al., 2004a; Masson et al., 2006). The filters were dried to constant weight (50 °C) and re-weighed. Analyses of dissolved metals were performed after matrix separation and preconcentration by liquid-solid extraction, to avoid interferences induced by major cations during instrumental analysis. In this work, we used commercially available cartridges (Dig-iSEP Blue®; SCP Science) with amino-diacetate as the functional group that has strong affinities to several divalent transition metal cations (e.g. Cd, Cu, Co, etc.). The samples (50 mL) with pH adjusted to 5.3 ± 0.1 were passed (flowrate 2 mL min⁻¹) through the acid-cleaned and conditioned column (0.1 M ammonium acetate buffer; pH 5.3). Then the cartridges were rinsed with 5 mL of buffer solution and trace metals were eluted with nitric acid (2 M, SCP-Science plasmasurePlus). Dissolved Cd concentrations were measured using

ICP-MS (X7, Thermo) with external calibration. Column blanks were performed by passing 50 mL Milli-Q water under identical conditions, parallel to the samples and certified reference materials. Methodological blanks obtained for Cd were generally lower than 2 ng L⁻¹ and were systematically subtracted from results obtained for the other samples. Detection limit (3.3 times the standard deviation of blank values; $n = 56$) was 6 ng L⁻¹. Reproducibility and accuracy of the method were assessed by analyzing certified reference seawater (CASS-4; Cd: 26 ± 3 ng L⁻¹) and estuarine water (SLEW-3; Cd: 48 ± 4 ng L⁻¹). For both reference materials, reproducibility was generally better than 3% (rsd; $n = 25$) and accuracy was respectively 99% and 98% of the certified values.

3.3. Cd⁰ determination

The commonly applied method for estimating dissolved metal net fluxes (Boyle et al., 1974) is based on the dissolved element distribution in the high salinity zone. According to Boyle et al.

Table 1
Names, dates, area of sampling (KP values) and mean discharges of the 15 cruises. Salinity range used for establishment of the apparent dilution line, number of points used for

each regression, equations of the dilution lines, and the resulting theoretical Cd_D^0 concentrations ($ng\ L^{-1}$; bold numbers) at $S \approx 0$ and daily Cd net flux estimates ($kg\ day^{-1}$). All the relationships are statistically significant at the 95% confidence level.

Cruise		Discharge ($m^3\ s^{-1}$)	PK range	Salinity range	Linear regression	<i>n</i>	<i>R</i> ²	Daily flux ($kg\ day^{-1}$)
<i>Flood events</i>								
Gimet 2	March 2001	2040	0–100	15–29	[Cd] $\frac{1}{4} -6.4$ (T3.4)[S] $\frac{1}{p} 222$ (T79)	8	0.79	39 (T 14)
Gimercad 3	March 2007	2600	0–85	12–26	[Cd] $\frac{1}{4} -3.8$ (T2.9)[S] $\frac{1}{p} 160$ (T56)	4	0.77	36 (T13)
<i>Intermediate freshwater discharge</i>								
Gimet 3	June 2001	546	0–100	17–32	[Cd] $\frac{1}{4} -6.4$ (T2.7)[S] $\frac{1}{p} 230$ (T65)	6	0.92	11 (T3.1)
Girox 1	February 2002	508	0–100	18–32	[Cd] $\frac{1}{4} -4.5$ (T0.7)[S] $\frac{1}{p} 176$ (T19)	9	0.97	8 (T0.8)
Reagir 3	May 2005	687	0–100	15–30	[Cd] $\frac{1}{4} -5.7$ (T0.5)[S] $\frac{1}{p} 219$ (T11)	6	0.99	13 (T0.7)
Gimercad 1	May 2006	705	0–100	19–30	[Cd] $\frac{1}{4} -5.0$ (T1.4)[S] $\frac{1}{p} 202$ (T37)	6	0.96	12 (T2.3)
<i>Low freshwater discharge</i>								
Girox 2	September 2002	368	0–100	23–31	[Cd] $\frac{1}{4} -4.8$ (T1.1)[S] $\frac{1}{p} 247$ (T31)	7	0.96	8 (T1.0)
Girox 4	September 2003	207	0–100	27–33	[Cd] $\frac{1}{4} -13.0$ (T1.2)[S] $\frac{1}{p} 490$ (T38)	6	0.99	9 (T0.7)
Reagir 1	July 2004	310	0–100	20–34	[Cd] $\frac{1}{4} -7.6$ (T1.6)[S] $\frac{1}{p} 296$ (T44)	9	0.96	8 (T1.2)
Reagir 2	October 2004	304	0–100	21–31	[Cd] $\frac{1}{4} -8.5$ (T3.2)[S] $\frac{1}{p} 324$ (T88)	6	0.93	9 (T2.3)
Metogir 1	July 2005	304	0–100	25–34	[Cd] $\frac{1}{4} -10.2$ (T1.5)[S] $\frac{1}{p} 367$ (T46)	7	0.99	10 (T1.2)
Reagir 4	November 2005	319	0–90	17–31	[Cd] $\frac{1}{4} -6.3$ (T5.3)[S] $\frac{1}{p} 264$ (T120)	6	0.93	7 (T3.3)
Gimercad 2	September 2006	169	0–100	23–33	[Cd] $\frac{1}{4} -10.4$ (T1.8)[S] $\frac{1}{p} 399$ (T53)	10	0.95	6 (T0.8)
Gimercad 4	October 2007	305	0–100	24–34	[Cd] $\frac{1}{4} -10.4$ (T2.3)[S] $\frac{1}{p} 400$ (T69)	6	0.98	11 (T1.8)
Gimercad 5	November 2007	209	0–100	23–30	[Cd] $\frac{1}{4} -7.2$ (T3.4)[S] $\frac{1}{p} 273$ (T105)	4	0.98	5 (T1.9)

(1974, 1982), the instant net flux (F_i) of a dissolved element X across an isohaline S at a time t is:

$$F_i \approx Q_i \left[X - X_0 \right]$$

where $[S]$ is salinity, $[X]$ the concentration of the element of interest, and Q_t the freshwater discharge at time t . If the relation between the dissolved element concentration and the salinity is linear (conservative behaviour), $d[X]/d[S]$ is constant and equal to the slope of the segment representing this relation in the concentration-salinity diagram. In this case, for any salinity $[S]$ belonging to the linearity range, $[X]_S - [S] * (d[X]/d[S])$ is constant and equal to the intercept $[X]_0$ of the straight line extrapolated from this segment, and the net instant flux of the element through any isohaline in the linearity range is:

$$F_i \approx Q_i [X]_0$$

As perfect linearity is rarely observed in any salinity range, a linear regression technique is applied to obtain the most probable

intercept, which is then used to estimate the net instant flux of any dissolved element. The salinity range used for this calculation should be representative of the geographic zone selected and should include the isohaline through which the flux estimation is to be made. The salinity range was determined in order to maximize the correlation coefficient. Daily net flux (F_{NET}) can be estimated as:

$$F_{NET} \approx Q_{fw} [X]_0$$

where Q_{fw} is the daily river discharge.

4. Results

4.1. Hydrologic conditions

The fifteen cruises are representative of a 7-years period covering contrasting hydrologic conditions and different seasons (Fig. 2, Table 1). Except for 2003, two campaigns per year were performed. Freshwater discharges during the different sampling cruises ranged from $169\ m^3\ s^{-1}$ (September 2006) to $2600\ m^3\ s^{-1}$ (March 2007).

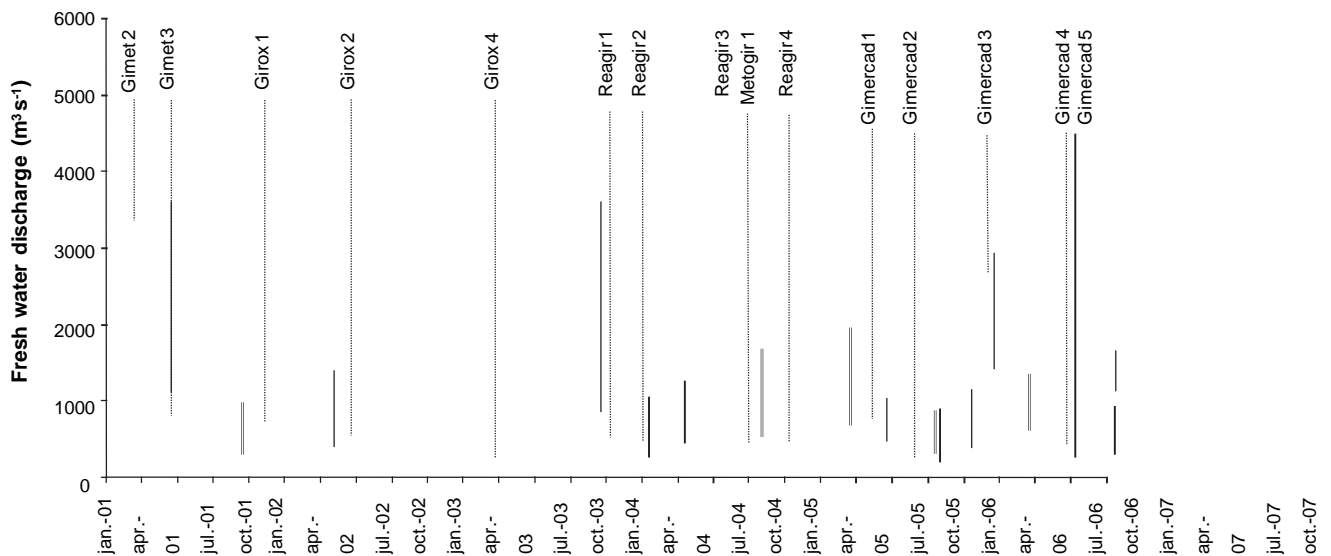


Fig. 2. Freshwater discharges of the Gironde Estuary during the 2001–2007 period.

Based on freshwater discharge data, the studied hydrological situations may be classified as follows: (a) low freshwater discharge during summer and autumn ($<500 \text{ m}^3 \text{ s}^{-1}$; September 2002, September 2003, July 2004, October 2004, July 2005, November 2005, October 2007 and November 2007), (b) intermediate discharge during the winter/spring period ($500\text{--}700 \text{ m}^3 \text{ s}^{-1}$; June 2001, February 2002, May 2005 and June 2006) and (c) the flood events ($>2000 \text{ m}^3 \text{ s}^{-1}$; March 2001, March 2007). Although the sampled periods cover a very wide range of discharge situations in the Gironde Estuary, the discharge range from 750 to $2000 \text{ m}^3 \text{ s}^{-1}$ has not been sampled. During the 2001–2007 period, hydrological situations with freshwater discharges comprised between 750 and $2000 \text{ m}^3 \text{ s}^{-1}$ only occurred during $\approx 25\%$ of the time. Accordingly, the sampled periods represent hydrological situations covering $\approx 75\%$ of the time during 2001–2007.

4.2. Dissolved Cd distribution along the salinity gradient

The Cd_D concentrations of the first 2 cruises (Gimet 2, Gimet 3) have already been presented in an earlier study with a different focus (Audry et al., 2007a,b) and are presented here together with original data obtained during 13 additional cruises to increase the representativity of the present work. Dissolved Cd (Cd_D) concentrations in surface waters along the salinity gradient displayed a similar distribution during all 15 cruises (Fig. 3). In the freshwater reaches, Cd_D concentrations were generally close to 20 ng L^{-1} (Fig. 3; Audry et al., 2007a,b). In the low-salinity range ($S < 15$), dissolved Cd concentrations increased with salinity reaching maximum values of 98 ng L^{-1} (November 2007) to 166 ng L^{-1} (September 2003) in the mid-salinity range ($S \approx 15\text{--}20$). In contrast, in the high-salinity range ($S > 20$), Cd_D behaved conservatively, i.e. showed a linear decrease with salinity (Fig. 3). The dissolved Cd concentrations obtained for bottom water profiles fit with the concentrations in surface waters, except during the flood event (March 2007) when bottom water concentrations were higher and more variable.

5. Discussion

5.1. Cadmium chlorocomplexation and variations of the maximum Cd_D values

The Cd_D concentrations observed for the 15 profiles ($8\text{--}166 \text{ ng L}^{-1}$; Fig. 3) were in the same range as those measured during the previous decade (Kraepiel et al., 1997; Boutier et al., 2000; Michel et al., 2000). The maximum values were higher than those reported for heavily industrialised/urbanised estuaries, such as the Seine (Chiffolleau et al., 1999), the Loire (Waeles et al., 2004) and the Scheldt Estuaries (Zwolsman et al., 1997), but they were clearly lower than those of in the Rio Tinto, which is severely impacted by acid-mine drainage (Braungardt et al., 2003) or in the Mersey Estuary, which drains a heavily populated and highly industrialised area (Martino et al., 2002).

The observed dissolved Cd distributions of all the campaigns showed typical non-conservative behaviour of Cd as observed in other macrotidal estuaries, e.g. the Tay (Owens and Balls, 1997), the Charente (Boutier et al., 2000), the Loire (Boutier et al., 1993; Waeles et al., 2004), the Scheldt (Zwolsman et al., 1997) and the Seine Estuaries (Chiffolleau et al., 1994, 1999). The commonly observed dissolved Cd addition in the low salinity range ($S < 15$) has been attributed to chloride-induced desorption of particle-bound Cd and formation of stable dissolved chlorocomplexes (e.g. Elbaz-Poulitchet et al., 1987; Comans and van Dijk, 1988; Turner et al., 1993; Turner, 1996). The apparently conservative behaviour in the mid- to high-salinity range (\approx linear decrease in Cd_D

concentrations with increasing salinity) may be attributed to dilution of mid-salinity high Cd_D waters by oceanic, low Cd_D water. This, in turn, suggests that the suspended particles in the high salinity range are depleted in easily desorbable Cd.

Like the low- to intermediate discharge situations, both flood situations (March 2001 and March 2007) showed a clear increase in the low-salinity range, but the mid-salinity maximum was much less defined, i.e. the data were more scattered than during low discharge situations (Fig. 3). This may suggest that floods modify the typical Cd_D distribution along the salinity gradient, eventually due to (a) local resuspension of bottom sediment (Sottolichio and Castaing, 1999) enriched in dissolved and particulate Cd (Audry et al., 2006) and/or (b) very short water residence times (≈ 20 days, Jouanneau and Latouche, 1981) with incomplete mixing of “fresh” Cd-rich and “old” Cd-depleted particles. In fact, fluvial inputs of dissolved and particulate Cd by the Garonne, Dordogne and Isle Rivers are at maximum during floods (Schafer et al., 2002b; Masson et al., 2006). The particle fluxes from the watershed to the estuary during major flood events (e.g. $\approx 0.8\text{--}1 \text{ Mt}$ in 6 days in February 2003; Coynel et al., 2007) may be equivalent to $20\text{--}25\%$ of the particle mass in the MTZ ($4\text{--}5 \text{ Mt}$; Jouanneau et al., 1990). This may modify the composition of the estuarine particle pool, especially for reactive carrier phases and associated elements that undergo profound changes in the estuarine geochemical gradients, inducing local heterogeneities in particle composition and reactivity. This hypothesis may be supported by the similarity between Cd_D concentrations in surface and bottom water profiles during low- and intermediate discharge (Fig. 3), suggesting that the water column was well-mixed under these hydrologic conditions. In contrast, during the floods, clear Cd_D concentration differences between surface and bottom water were observed and attributed to (a) the impact of sediment remobilisation including porewater release into the water column and/or (b) incomplete mixing and variable residence times of different water masses and particles.

Although freshwater discharges during the different cruises varied by a factor 16, maximum Cd_D concentrations in the mid-salinity range of each profile showed only little variability, ranging from 98 ng L^{-1} to 166 ng L^{-1} (Fig. 3). Interestingly, there was no systematic difference in the maximum Cd_D for the different types of hydrological situations. This observation is not what would have been expected, presuming dilution of dissolved Cd in the estuary by freshwater inputs during floods. In fact, one would have expected maximum Cd_D mid-salinity levels during low water discharge, when long residence times of water (up to 3 months), and particles (>2 years) should account for most efficient Cd desorption (Jouanneau and Latouche, 1981). Furthermore, neither the position nor the SPM concentration level in the MTZ seem to control maximum Cd_D concentrations in the Gironde Estuary (Fig. 3).

5.2. Temporal variations in the theoretic Cd_D concentrations at zero salinity

The theoretic Cd_D concentrations at zero salinity (Cd_D^0) obtained from applying Boyle's method to the different longitudinal profiles are reported in Table 1. During the flood events, Cd_D^0 concentrations were 222 ng L^{-1} in March 2001 ($2040 \text{ m}^3 \text{ s}^{-1}$) and 160 ng L^{-1} in March 2007 ($2600 \text{ m}^3 \text{ s}^{-1}$). This is in good agreement with Cd_D^0 (206 ng L^{-1}) obtained from previously reported data on dissolved Cd concentrations in the Gironde Estuary during a comparable hydrologic situation in 1994 ($Q \approx 2100 \text{ m}^3 \text{ s}^{-1}$; Kraepiel et al., 1997). For intermediate freshwater discharge conditions (2001–2007 period), Cd_D^0 concentrations ranged from 176 ng L^{-1} (February 2002; $Q \approx 508 \text{ m}^3 \text{ s}^{-1}$) to 230 ng L^{-1} (June 2001; $Q \approx 546 \text{ m}^3 \text{ s}^{-1}$), i.e. were rather similar to those observed during floods.

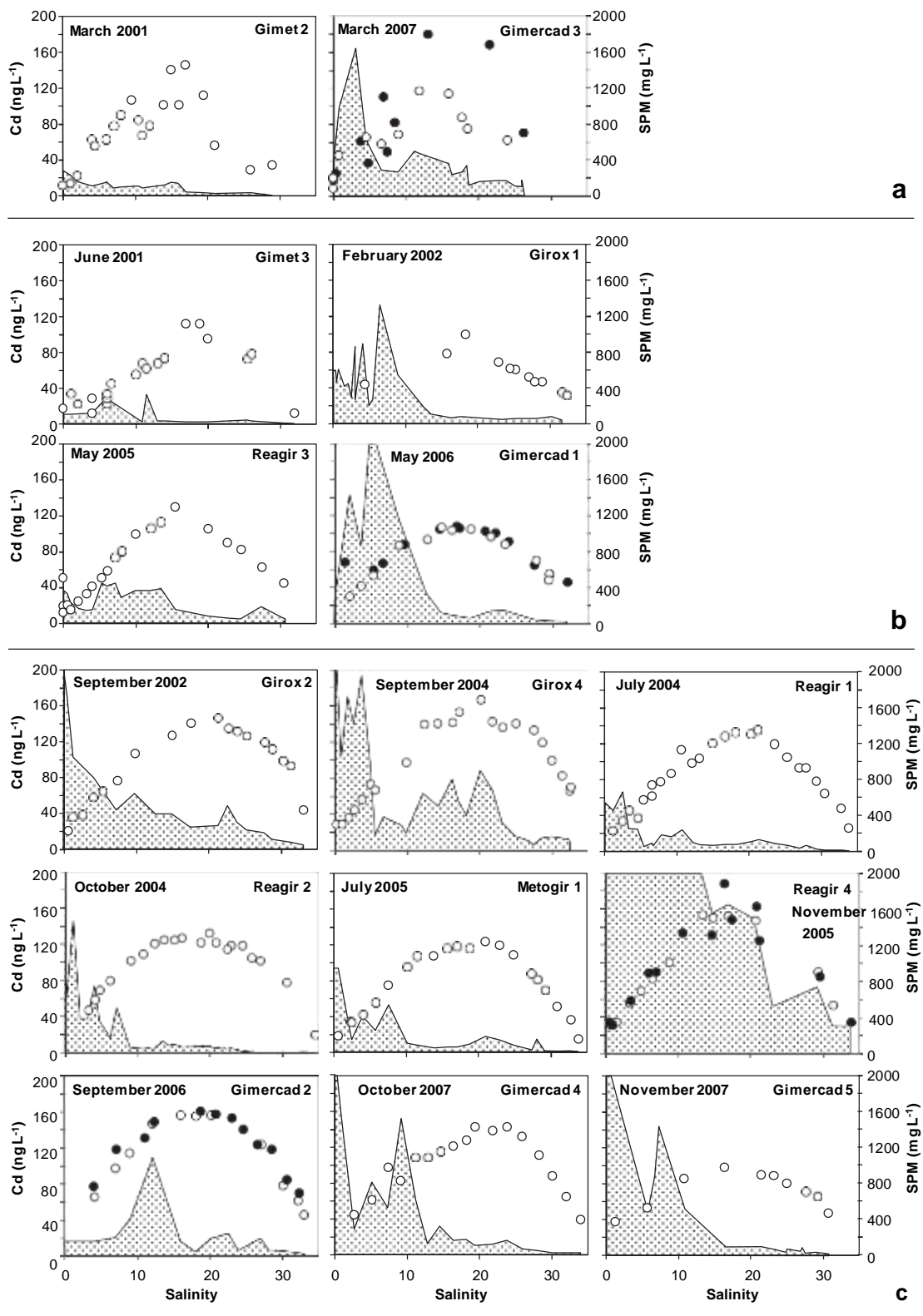


Fig. 3. Distribution of dissolved Cd and suspended particulate matter along the salinity gradient for (a) flood events, (b) intermediate discharge and (c) low discharge. Dotted areas represent surface water SPM concentrations, open circles represent Cd_D concentrations in surface water and full circles represent Cd_D in bottom water. Note that the Cd_D values for the Gimet 2 and 3 cruises (March and June 2001) are from Audry et al. (2007a).

In contrast, clearly higher Cd_D values may be deduced from older data measured during intermediate freshwater discharge in 1984 ($Cd_D \approx 1000 \text{ ng L}^{-1}$, $Q \approx 600 \text{ m}^3 \text{ s}^{-1}$; Jouanneau et al., 1990) and in 1991 ($Cd_D \approx 504 \text{ ng L}^{-1}$, $Q \approx 1010 \text{ m}^3 \text{ s}^{-1}$; Boutier et al., 2000). These

results are consistent with strongly reduced emissions after the stop of the ore treatment activity (1987) and gradually decreasing particulate Cd concentrations of SPM in the fluvial-estuarine system, which is further supported by sedimentary records in Lot River

reservoir lakes (e.g. Audry et al., 2004b) and in the flood-tidal dock at Bordeaux (Grousset et al., 1999).

During low freshwater discharge ($Q < 500 \text{ m}^3 \text{ s}^{-1}$), Cd^0 concentrations varied by a factor of 2 and ranged from 247 ng L^{-1} (September 2002; $Q \approx 368 \text{ m}^3 \text{ s}^{-1}$) to 490 ng L^{-1} (September 2003; $Q \approx 207 \text{ m}^3 \text{ s}^{-1}$; Table 1; Fig. 4a). To our knowledge, this is the most complete (9 campaigns) data set on Cd^0 values covering most low freshwater discharge situations in the same estuary during consecutive years (2002–2007). This dataset is the first to show the great variability of the Cd^0 concentration during low freshwater discharge ($Q < 500 \text{ m}^3 \text{ s}^{-1}$; Fig. 4a). The observed variations probably reflect the balance of desorption efficiency, residence time of water and particles in the estuary and mixing of fresh- and seawater. Such non-steady state conditions may eventually result from several factors. First, variations in the input of freshwater particles (rich in desorbable Cd) or urban particles from Bordeaux may enhance addition, especially during low discharge, when particle residence time in the estuary is long (Sottolichio and Castaing, 1999). Second, variable Cd_D release may be due to sediment and/or fluid mud resuspension by tidal currents and/or dredging (Robert et al., 2004; Audry et al., 2007a,b). Third, O_2 -depletion in the bottom of the water column during low freshwater discharge (Abril et al., 1999) could eventually induce removal processes, e.g. by sulphide precipitation, as reported for the Scheldt Estuary (Zwolsman et al., 1997) temporarily and locally limiting Cd addition. However, the well-mixed Gironde Estuary generally displays relatively high oxygen concentrations in most of the water column

(too high for sulphide precipitation), despite seasonal hypoxic conditions in the fluid mud layer during pronounced low discharge conditions, (IFREMER, 1994; Abril et al., 1999). This is consistent with sulphide oxidation in the bottom of the water column induced by resuspension of anoxic sediments (Audry et al., 2007b).

Plotting Cd^0 concentrations over daily water discharge separates the recent (2001–2007) data into two major domains (Fig. 4a). The first relates to the low discharge period ($Q < 500 \text{ m}^3 \text{ s}^{-1}$) for which Cd^0 values were very variable ($247\text{--}490 \text{ ng L}^{-1}$). The second domain ($Q > 500 \text{ m}^3 \text{ s}^{-1}$) regroups the flood events and intermediate freshwater discharge periods for which Cd^0 values were less variable ($201\text{--}28 \text{ ng L}^{-1}$).

Plotting daily Cd_D net fluxes over daily water discharge shows how much Cd_D net fluxes are controlled by discharge (Fig. 4b). These results are clearly different from those of Michel et al. (2000) reporting rather constant daily Cd_D net fluxes in the Gironde Estuary whatever the freshwater discharge. This fundamental difference may eventually be attributed to different observation periods. The results of Michel et al. (2000) are based on one relatively dry hydrological year (1997/1998) following a succession of very wet years with numerous intense floods (1992, 1994, 1996; Schafer et al., 2002b). Assuming that (a) a major part of the MTZ had been expelled out of the estuary in 1996 and the preceding wet years and (b) the MTZ may not have been well installed/renewed in 1997/1998 due to low discharge, one cannot exclude that at that time the desorbable Cd stock in estuarine suspended particles was too low to buffer discharge-induced variations, e.g. due to dilution. Accordingly, this observation period may be considered as probably not representative of present day (2001–2007) conditions that correspond to a series of dry to average years with few isolated intense floods (Masson et al., 2007), which is not favourable to particle expulsion out of the estuary. Although the existing dataset does not allow a definitive conclusion on this point, it is evident that the hypothesis of constant daily Cd_D net fluxes cannot be applied to obtain reliable flux estimates at the annual timescale.

5.3. Daily Cd_D net fluxes

During 2001–2007, daily Cd_D net fluxes, classically estimated from the product of Cd^0 and the respective daily freshwater discharge, ranged from 5 kg to 39 kg (Fig. 4b). These values are of the same order as net Cd_D net fluxes reported for the Seine (11 kg day^{-1} ; Chiffolleau et al., 1999), one of the most metal contaminated estuaries in the world (Meybeck et al., 2004) and the Loire Estuaries (14 kg day^{-1} ; Waeles et al., 2004). Comparing the recent daily Cd_D net flux estimates for the Gironde Estuary to those obtained in 1982 (37 kg day^{-1} for $400 \text{ m}^3 \text{ s}^{-1}$; Elbaz-Poulichet et al., 1987), 1984 (54 kg day^{-1} for $Q \approx 600 \text{ m}^3 \text{ s}^{-1}$; Jouanneau et al., 1990) and 1992 (44 kg day^{-1} for $Q \approx 1000 \text{ m}^3 \text{ s}^{-1}$; Boutier et al., 2000) shows that daily Cd_D net fluxes have decreased by a factor 3–5 during the past two decades. This probably reflects stepwise reduction of emissions in the Decazeville watershed and is in good agreement with progressive decontamination of the system discussed above. Although during the nineties alternation of dry and wet hydrological years induced some variability of particulate Cd inputs into the Gironde Estuary ($2.5\text{--}23.0 \text{ t a}^{-1}$ with a mean of 12.4 t a^{-1} ; 1991–1999 period; Schafer et al., 2002b), frequency of these variations was similar to the particle residence time in the estuary (1–2 years; Castaing and Jouanneau, 1979), which may have “smoothed” Cd_D addition resulting in more or less constant Cd_D net fluxes. In contrast, since 2000 dissolved Cd fluxes from the source to the estuary have decreased due to remediation (Audry et al., 2004b) and particulate Cd transport in the fluvial-estuarine system clearly changed due to increasingly frequent and long low discharge periods. For example, succession of 2–3 dry years

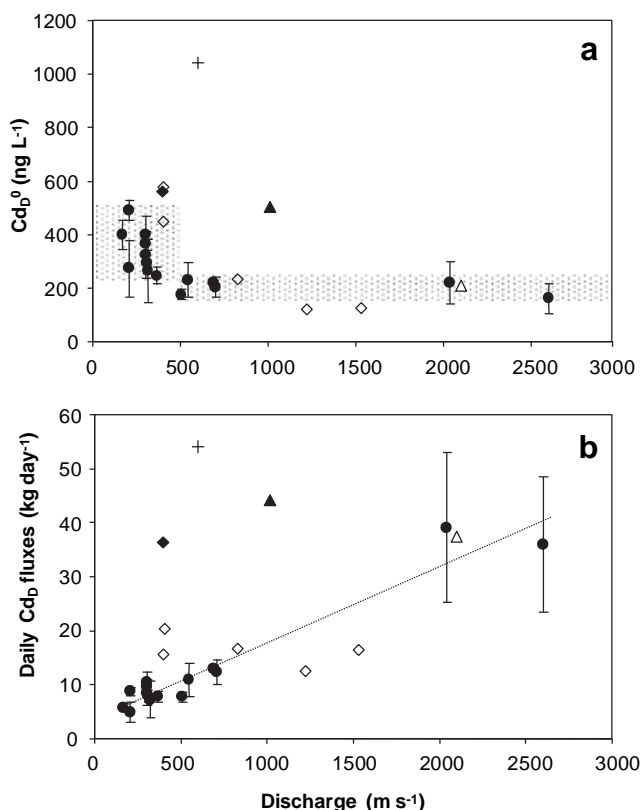


Fig. 4. Theoretic Cd^0 concentrations (a) and daily Cd_D flux (b) over the whole range of discharges observed. Full circles correspond to the data of this study (2001–2007), full diamond corresponds to 1982 data (Elbaz-Poulichet et al., 1987), the cross corresponds to 1984 data (Jouanneau et al., 1990), full triangle corresponds to 1991 data (Boutier et al., 2000), open triangle corresponds to 1994 data (Kraepiel et al., 1997), and open diamonds correspond to 1997–1998 data (Michel et al., 2000). The linear correlation refers only to the data of this study (2001–2007).

followed by extreme flood events resulted in erosion and destabilisation of ancient Lot River sediment, enhanced by flood management (e.g. dam opening; Coynel et al., 2007). These processes contributed to temporarily very high Cd transport into the estuary followed by relatively long periods with very low Cd_p fluxes (Schafer et al., 2002b; Coynel et al., 2007; Masson et al., 2007). Extremely long low discharge periods without renewal/expulsion of MTZ particles (e.g. July 2004–November 2006; Fig. 2) one might expect the estuarine SPM in the MTZ to become depleted in desorbable Cd, resulting in decreasing addition and temporarily reduced net fluxes. Replenishment and/or exchange of estuarine SPM during the following flood would then restore the estuarine stock in desorbable Cd. Accordingly, the point of time of a sampling campaign with respect to depletion/replenishment periods may eventually explain the observed variability of Cd_p values during the studied low discharge situations.

5.4. Annual Cd_D net flux estimation

The major part (up to 95%, unpublished data) of particulate Cd (Cd_p) exposed to the estuarine salinity gradient is transferred to the dissolved phase before/while exported to the coastal ocean (Kraepiel et al., 1997). In fact, typical annual average Cd_p concentrations (weighted by SPM fluxes) at the entry of the Gironde Estuary during the 1990–2005 period were 4–4.5 mg kg⁻¹ (after Schafer et al., 2002b; Masson et al., 2006), whereas typical particulate Cd concentration in MTZ particles are 0.4–0.5 mg kg⁻¹ (Kraepiel et al., 1997; Audry et al., 2007a). Assuming that Cd_D may easily enter the aquatic food chain and potentially impact the marine resources (e.g. oysters; Lekhi et al., 2008), it is important to quantify Cd_D net fluxes. Ideally, annual net flux estimates should result from integration of daily net fluxes, but in reality only few daily flux estimates exist even for the most intensively observed estuaries. Thus, daily net flux estimates have to be extrapolated and the choice of the appropriate extrapolation method will determine the quality of annual net flux estimates. Some pioneer studies on net Cd_D fluxes in the Gironde Estuary used one Cd_p value multiplied by the mean annual freshwater discharge (Elbaz-Poulichet et al., 1982) or by the annual freshwater discharge of the respective year (e.g. Boutier et al., 2000). This in turn requires low seasonal variability of Cd_p compared to the variability of the river water discharge. However, both, the observation that Cd_p values varied with freshwater discharge during 1997/1998 (Michel et al., 2000) and the present dataset show that random Cd_p values may be too variable to produce reliable Cd_D net flux estimates at the annual scale. In fact, multiplying the three different Cd_p values reported by Michel et al. (2000) in 1997 by annual discharge would produce Cd_D net flux estimates of 2.5–12.2 t a⁻¹, depending on the Cd_p value used. Such high uncertainty seems to invalid the method proposed by Boutier et al. (2000). This and the high variability of Cd_p values obtained from our 2001–2007 dataset supports the idea that reliable annual Cd_D net flux estimates cannot be obtained from random single Cd_p values especially, when they were obtained during low

discharge (Fig. 4a). However, the dataset also suggests that for freshwater discharge >500 m³ s⁻¹ in the Gironde Estuary, Cd_D values were relatively stable (i.e. 201 T 28 ng L⁻¹) and did not seem to change much during the past decade (Fig. 4a). Extrapolation of Cd_D net fluxes to the annual scale needs taking into account these discharge-dependent features. Accordingly, we suggest using different Cd_p for different discharge situations. For discharge >500 m³ s⁻¹, Cd_p value of 201 T 28 ng L⁻¹ should provide satisfactory daily Cd_D net flux estimates. For discharge <500 m³ s⁻¹, a range of Cd_D net fluxes could be estimated, by using the lowest and the highest Cd_p value observed 201 T 31 ng L⁻¹ and 201 T 490 T 38 ng L⁻¹. This would produce an estimated range of the annual Cd_D net fluxes as follows:

$$\begin{aligned} \text{Min } F_{\text{NET}} \text{ Cd}_D &= \sum_{Q_i > 500} \text{Cd}_{D, Q_i > 500}^0 \times Q_i \\ &+ \sum_{Q_i < 500} \text{Cd}_{D, Q_i < 500}^{\text{Min}} \times Q_i \\ \text{Max } F_{\text{NET}} \text{ Cd}_D &= \sum_{Q_i > 500} \text{Cd}_{D, Q_i > 500}^0 \times Q_i \\ &+ \sum_{Q_i < 500} \text{Cd}_{D, Q_i < 500}^{\text{Max}} \times Q_i \end{aligned}$$

The annual Cd_D net flux estimates for the 2001–2007 period obtained by this method suggest that annual Cd_D net flux ranged from 3.8–5.0 t a⁻¹ in 2005 to 6.0–7.2 t a⁻¹ in 2004, depending on interannual variations in freshwater inputs (Table 2; Fig. 5). The differences between the minimum and maximum annual Cd_D net flux estimates ranged through 0.8–1.3 t a⁻¹, with uncertainties mainly related to the duration of low freshwater discharge situations.

5.5. Annual net Cd addition

The annual Cd_D fluxes at the upstream limit of the salinity gradient (F_I : incoming fluxes; Audry et al., 2007a,b) were estimated as follows:

$$F_I \text{ Cd}_{S/0} = \sum Q_i \text{ Cd}_{S/0}$$

where Q_i represents the daily freshwater discharges and $\text{Cd}_{S/0}$ represents the average Cd_D concentration measured at $S/0$ for the different campaigns ($\text{Cd}_{S/0} = 17 \text{ T } 5 \text{ ng L}^{-1}$). The annual net Cd_D addition was estimated from the difference between F_{NET} and F_I for each year (Table 2). The resulting annual Cd_D addition in the salinity gradient of the Gironde Estuary ranged through 3.5–5.7 t a⁻¹ using $\text{Min } F_{\text{NET}}$ and through 4.7–6.7 t a⁻¹ using $\text{Max } F_{\text{NET}}$. This suggests that Cd_D addition provides Cd amounts 11–16-fold higher than the incoming flux (F_I), depending on freshwater discharge. Furthermore, Cd_D addition did not seem to show a clearly decreasing trend during the observation period, despite of the generally decreasing contamination of the Lot-Garonne-Gironde fluvial-estuarine system (Fig. 5).

Dredging activities (Audry et al., 2007a,b), ongoing reload of the MTZ by Cd-rich particles from sedimentary stocks in the watershed

Table 2
Annual Cd_D net flux, net addition and annual freshwater inputs in the Gironde Estuary during the 2001–2007 period. Annual Cd_D dissolved net fluxes are statistically significant at the 95% confidence level.

	2001	2002	2003	2004	2005	2006	2007
$\text{Min } F_{\text{NET}} \text{ Cd}_D \text{ (t a}^{-1}\text{)}$	6.0 (T0.8)	4.4 (T0.6)	5.0 (T0.7)	6.0 (T0.8)	3.8 (T0.5)	4.3 (T0.6)	4.4 (T0.6)
$\text{Max } F_{\text{NET}} \text{ Cd}_D \text{ (t a}^{-1}\text{)}$	7.0 (T0.8)	5.7 (T0.6)	5.8 (T0.7)	7.2 (T0.8)	5.0 (T0.5)	5.4 (T0.6)	5.6 (T0.6)
Min Net addition	5.7	4.0	4.6	5.5	3.5	3.9	4.0
Max Net addition	6.6	5.3	5.4	6.7	4.7	5.0	5.2
Annual freshwater inputs (km ³)	28.8	20.8	24.3	28.0	17.7	20.6	21.0

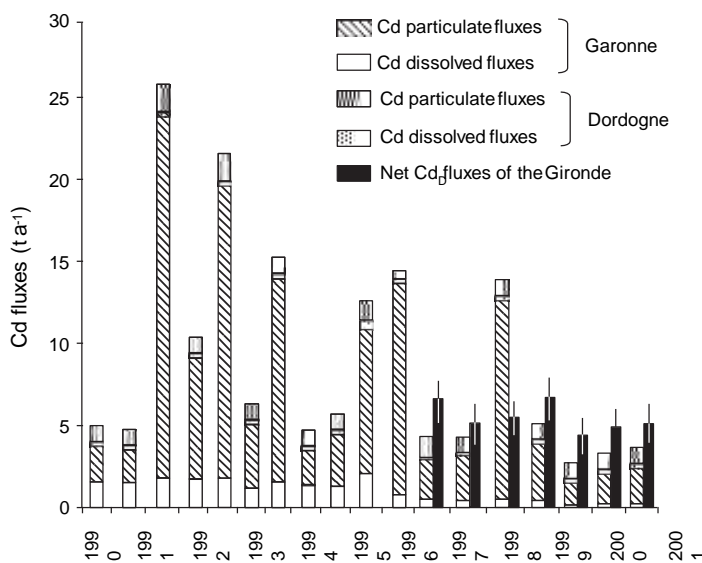


Fig. 5. Dissolved and particulate annual Cd fluxes in the Garonne and Dordogne River for the 1990–2007 period (1990–1999: Schafer et al., 2002a,b; 2000–2002: Masson et al., 2006; 2003–2007: unpublished data) and dissolved annual Cd_D net fluxes in the Gironde Estuary for the 2001–2007 period. The errors bars represent the minimum and maximum flux estimates taking into account the 95% confidence intervals in each of these fluxes.

during flood events (Coynel et al., 2007) and/or progressive release of the estuarine Cd stock due to tidal resuspension of the bottom sediment (Robert et al., 2004) could eventually explain this relatively constant dissolved Cd net addition in the Gironde Estuary.

5.6. Comparison between the annual Cd_D net fluxes and the gross total Cd inputs

Estuarine mass balances need reliable estimates of gross fluxes, i.e. fluxes derived from the upstream watershed and net fluxes, i.e. fluxes leaving the estuary towards the coastal ocean. The existing observation methods allow precise and reliable estimates of gross fluxes, which are at the origin of net fluxes, but it is widely accepted that estuarine processes widely decouple gross and net fluxes (Zwolsman et al., 1997; Monbet, 2006). Accordingly, comparison of gross and net fluxes reasonably integrates larger (e.g. annual) timescales.

The annual freshwater fluxes entering the Gironde Estuary are derived from the sum of dissolved and particulate fluxes in the major tributaries, i.e. the Garonne, the Dordogne and the Isle Rivers. However, the Garonne River accounts for the majority of the highly variable and discharge-dependent SPM and Cd fluxes entering the Gironde Estuary, whereas the Dordogne River typically contributes $1.2 \text{ T } 0.3 \text{ t a}^{-1}$ (Schafer et al., 2002b; Fig. 5). For the 2001–2007 period, the total annual Cd gross fluvial inputs were clearly lower and less variable than during 1990–2000 ($2.8\text{--}5.1 \text{ t a}^{-1}$; Fig. 5) except for 2003 (13.9 t a^{-1}), when a major flood remobilised important stocks of Cd-polluted sediment in the Lot River contributing $\approx 90\%$ to the high annual gross fluxes into the estuary (Coynel et al., 2007). The annual gross Cd_D fluxes also tended to decrease during the observation period (2001–2007: from 0.5 to 0.2 t a^{-1}). In contrast, annual Cd_D net fluxes did not show any decreasing trend and were higher than annual gross total Cd fluxes in six years out of seven (Fig. 5), with a difference of $\approx 1.4 \text{ T } 0.4 \text{ t a}^{-1}$. This difference has become visible only recently, due to the decreasing overall Cd budget of the Gironde Estuary and may reflect different sources of increasing relative importance. Atmospheric Cd wet deposition ($\approx 7 \text{ ng cm}^{-2} \text{ a}^{-1}$; Maneux et al., 1999) would contribute $\approx 0.05 \text{ t a}^{-1}$

to the estuary surface and $\approx 0.2 \text{ t a}^{-1}$ to the adjacent watersheds. Diffuse inputs (e.g. by agriculture), urban wastewater, runoff and aerosols would also be expected to represent non-negligible Cd sources to the estuary.

The average particulate Cd concentration ($0.57 \text{ T } 0.18 \text{ mg kg}^{-1}$, $n = 91$; unpublished data) in the muddy sediments (silt and clay $>80\%$) of the Gironde Estuary covering $\approx 400 \text{ km}^2$ (Kapsimalis et al., 2004), tend to be higher than that of SPM in the water column of the downstream estuary ($0.44 \text{ T } 0.04 \text{ mg kg}^{-1}$, $n = 546$, unpublished data; Kraepiel et al., 1997). Assuming that tidal currents, dredging, etc. may eventually erode the uppermost 10 cm of these sediments and using the average concentration values in sediment and SPM, the expected Cd desorption in the salinity gradient would suggest a potential release of $\approx 2.6 \text{ t}$ of Cd_D. However, erosion of deeper sediment layers with higher Cd concentrations (Robert et al., 2004) cannot be excluded, which would imply a higher sedimentary Cd stock potentially recycled. Without fundamental changes in estuarine functioning, recycling of the sedimentary Cd stock in the Gironde Estuary and diverse Cd inputs from diffuse sources may be sufficient to maintain the Cd budget of the Gironde Estuary at the present level for another decade. Assuming that the present annual Cd budget of the Gironde Estuary is close to steady state, one may reasonably expect that (a) the present empirical model for Cd_D net flux estimates at the annual timescale will still be valid during this period and (b) Cd_D net fluxes to the coastal ocean are forecast to decrease slowly, even when the ongoing inputs of Cd from the watershed continue to decrease.

6. Conclusion

The present work provides a very complete and recent dataset on dissolved Cd concentrations in the salinity gradient of the Gironde Estuary covering seven years and contrasting hydrological situations. Although Cd_D concentrations along the salinity gradient followed the classical pattern, the data support substantial improvement in quantitative understanding of Cd behaviour in this system. Maximum Cd_D concentrations in the mid-salinity range were independent from discharge and average SPM concentration. For a given cruise and salinity, Cd_D concentrations in surface bottom water were similar, except during floods, suggesting limited mixing and heterogeneous water and particle residence times. Comparing recent and historic data clearly indicated a decrease by a factor 3–5 in daily Cd_D net fluxes during the past two decades, mainly reflecting emission control in the Decazeville watershed. The present data revealed that during 2001–2007 daily Cd_D net fluxes depended strongly on discharge, which is the contrary of the hypothesis (Cd_D net fluxes independent from discharge; Michel et al., 2000) established for the previous decade. Therefore, the commonly applied method of estimating annual Cd_D net fluxes for the Gironde Estuary is not valid anymore for the present conditions. Taking into account discharge-dependent daily Cd_D net fluxes, we proposed a new empirical model for Cd_D net flux estimates in the Gironde Estuary at the annual timescale.

Comparing annual gross Cd fluxes to annual Cd_D net fluxes obtained from the model revealed that neither annual Cd_D addition nor Cd_D net fluxes showed a clear trend during 2001–2007 despite of decreasing gross fluxes. This and the fact that Cd_D net fluxes systematically exceeded total gross fluxes imply inputs by various diffuse sources. These inputs were negligible in former Cd budgets, but have become significant due to decreasing fluxes from the major source in the Decazeville basin. These sources and Cd release from the sedimentary stock may keep the Cd budget of the Gironde Estuary at the present level for at least another decade, suggesting slow decontamination of the estuary, i.e. decreasing sedimentary Cd stocks. If Cd_D fluxes to the coastal ocean are forecast to decrease

slowly, further work is needed to (a) quantify inputs from intra-estuarine sources and (b) optimize management and control of total Cd inputs to accelerate the clean-up of this system and minimize future impact on the coastal zone.

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