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1 **Nitrogen and phosphorus budgets in the Northwestern Mediterranean deep convection**
2 **region**

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15 **Key points up** to three (<140 characters)

16 - Modeling of annual nitrogen and phosphorus cycles in the northwestern Mediterranean
17 deep convection region for the period 2012-2013.

18 - The deep convection area was a sink of inorganic matter and a source of organic matter
19 for the surrounding area over the period 2012-2013.

20 - The N:P ratio in the surface layer is submitted to drastic variations during deep
21 convection and bloom transition periods.

22 **Abstract**

23 The aim of this study is to understand the biogeochemical cycles of the northwestern
24 Mediterranean Sea (NW Med), where a recurrent spring bloom related to dense water formation
25 occurs. We used a coupled physical-biogeochemical model at high resolution to simulate
26 realistic one-year period and analyze the nitrogen (N) and phosphorus (P) cycles. First, the model
27 was evaluated using cruises carried out in winter, spring and summer and a Bio-Argo float
28 deployed in spring. Then, the annual cycle of meteorological and hydrodynamical forcing and
29 nutrients stocks in the upper layer were analyzed. Third, the effect of biogeochemical and
30 physical processes on N and P was quantified. Fourth, we quantified the effects of the physical
31 and biological processes on the seasonal changes of the molar $\text{NO}_3:\text{PO}_4$ ratio, particularly high
32 compared to the global ocean. The deep convection reduced the $\text{NO}_3:\text{PO}_4$ ratio of upper waters,
33 but consumption by phytoplankton increased it. Finally, N and P budgets were estimated. At the
34 annual scale, this area constituted a sink of inorganic and a source of organic N and P for the
35 peripheral area. NO_3 and PO_4 were horizontally advected from the peripheral regions into the
36 intermediate waters (130-800 m) of the deep convection area, while organic matter was exported
37 throughout the whole water column toward the surrounding areas. The annual budget suggests
38 that the NW Med deep convection constitutes a major source of nutrients for the photic zone of
39 the Mediterranean Sea.

40 **Index Terms and Keywords**

41 Mediterranean Sea, physical-biogeochemical modeling, nutrient seasonal dynamics, nitrate and
42 phosphate budget, N:P stoichiometry

43 **1. Introduction**

44 The Mediterranean Sea, often considered as an oligotrophic region [*Antoine et al.*, 1995],
45 exhibits nutrient-depleted surface waters most of the year and low nutrient concentrations in

46 deep waters [*Krom et al., 1991*]. With respect to the global ocean Redfield nitrate (NO_3) to
47 phosphate (PO_4) ratio of 16:1 [*Redfield, 1964*], the Mediterranean Sea is characterized by
48 particularly high values [*Ribera d'Alcalà et al., 2003*]. Moreover, these oligotrophic conditions
49 and high $\text{NO}_3:\text{PO}_4$ (hereafter N:P) ratios show a pronounced east-west gradient [*Pujo-Pay et al.,*
50 2011], with the eastern basin presenting ultra-oligotrophic conditions. The N:P ratios in deep
51 layers were estimated at 20-22 in the western sub-basin [*Béthoux et al. 1998*] and at 27-30 in the
52 eastern sub-basin [*Ribera d'Alcala et al., 2003*] of the Mediterranean.

53 The distribution of nutrients in the semi-enclosed Mediterranean Sea results from internal
54 physical and biogeochemical dynamics, from atmospheric and terrestrial inputs, and from
55 exchanges with the Atlantic Ocean and the Black Sea. Physical dynamics are partly controlled by
56 the formation of intermediate and deep waters as the winter climate of the Mediterranean Sea is
57 characterized by cold, dry local winds blowing from the north [*Hauser et al., 2003*], which
58 induce an increase of the surface layer density by strong evaporation and cooling, and then
59 trigger the deep convection mixing. This process is responsible for the formation of intermediate
60 and deep water masses in the northwestern basin (hereafter NW Med) [*Medoc Group, 1970*], the
61 south Adriatic [*Pollak et al., 1951*], the Aegean [*Nittis et al., 2003*] and the Rhodes Gyre
62 [*Ovchinnikov et al., 1984*]. The high nitrate and phosphate surface concentrations (values close
63 to deep concentrations) observed during convective episodes in these regions [*Yilmaz and*
64 *Tugrul, 1998; Gacic et al., 2002; Santinelli et al., 2012; Severin et al., 2014*] demonstrate that the
65 deep convection process is also responsible for large nutrient enrichments of the upper layers. In
66 addition, a study based on satellite-derived chlorophyll a data [*D'Ortenzio and Ribera d'Alcalà,*
67 2009] has shown that the deep convection regions are characterized by intense to moderate
68 spring blooms, while phytoplankton development in the rest of the open Mediterranean Sea is of
69 low magnitude. This feature of phytoplankton dynamics in convection regions may be explained
70 by the large amounts of nutrient supplied during the intense vertical mixing events [*Lavigne et*
71 *al., 2015*].

72 Among the deep convection regions of the Mediterranean Sea, the NW Med (Fig. 1A) has been
73 identified as the region where the vertical mixing and the associated phytoplankton spring bloom
74 are the most intense and recurrent [Houpert *et al.*, 2015; D'Ortenzio and Ribera d'Alcalà, 2009;
75 Lavigne *et al.*, 2013; Mayot *et al.*, 2016]. The NW Med convection has been shown to largely
76 influence the regional biogeochemical cycles and marine ecosystems, mostly by importing
77 nutrient-enriched deep waters to the surface [Severin *et al.*, 2014; Ulses *et al.*, 2016]. This
78 enrichment changes the biogeochemical characteristics of the surface layers [Durrieu de Madron
79 *et al.*, 2011; Herrmann *et al.*, 2013; Auger *et al.*, 2014], induces a large phytoplankton spring
80 bloom [D'Ortenzio *et al.*, 2009; Lavigne *et al.*, 2013; Estrada *et al.*, 2014], favors high
81 particulate carbon export [Gogou *et al.*, 2014; Ulses *et al.*, 2016] and shifts the zooplankton
82 community toward larger sized organisms [Auger *et al.*, 2014]. Moreover, deep convection in the
83 NW Med has also been shown to promote a change of the surface nutrient ratios [Severin *et al.*,
84 2014; Pasqueron de Fommervault *et al.*, 2015a]. However, the impact of nutrient replenishments
85 on the seasonal cycles of nitrogen and phosphorus has been poorly explored. The few
86 quantifications of the upward flux of nitrate and phosphate that have been performed [Severin *et al.*
87 *et al.*, 2014; Ulses *et al.*, 2016] remain limited in space and/or time. A thorough quantification of
88 the processes linked to the seasonal variations of nitrogen and phosphorus cycles during one
89 complete year has not yet been proposed in this region.

90 The objectives of the DeWEX project (Deep Water formation EXperiment) are (i) to better
91 understand the impact of dense water formation on the marine biogeochemical cycles and (ii) to
92 provide a consistent dataset of hydrological and biogeochemical parameters to improve the
93 numerical modeling of the convection hydrology and coupled biogeochemical processes. To
94 fulfill these objectives, oceanographic research vessels covered the deep convection area during
95 each season from summer 2012 to summer 2013. Once calibrated and validated, numerical
96 modeling is expected to achieve goals at different time scales. In the short term, the model is

97 expected to interpolate between the DeWEX cruises to calculate the nutrient and organic matter
98 budgets over an annual cycle. In the long term, and after complementary validations with multi-
99 year datasets, modeling will be used as an integrative tool to investigate the question of how
100 climate change and anthropogenic activities could impact the cycle of biogenic elements and
101 marine ecosystems.

102 The present study aims to quantify the dynamics of nutrients at the seasonal scale and to estimate
103 an annual budget of nitrogen and phosphorus, in the NW Med dense water formation area. After
104 a brief description of the observation network, the modeling strategy is described (section 2), and
105 then evaluated through comparisons with observations (section 3). The seasonal variabilities of
106 the atmospheric forcing and hydrography are presented (section 4.1.1), together with the annual
107 cycles of nutrient stocks in the upper layer (4.1.2). In section 4.1.3, we discuss the impact of
108 physical and biogeochemical processes on the N:P ratio in the upper layer. Finally, an annual
109 nitrogen and phosphorus budget of the region is proposed (section 4.2).

110 **2. Method**

111 **2.1.Observations**

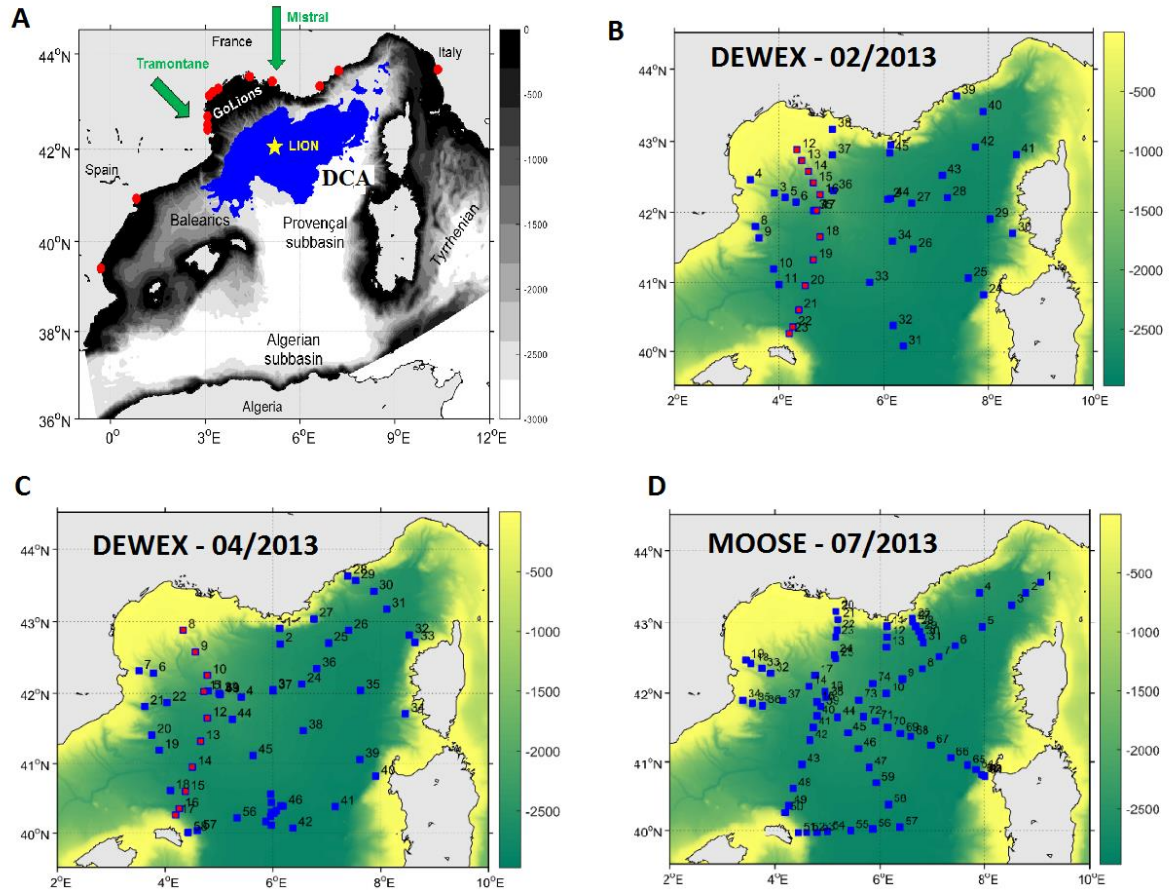
112 During the period from September 2012 to September 2013, a large number of hydrological and
113 biogeochemical observations (>178 profiles) were made in the NW Med in the framework of the
114 MERMEX (Marine Ecosystems Response in the Mediterranean Experiment) and HyMeX
115 (Hydrological cycle in the Mediterranean Experiment) programs. The objective was to better
116 understand the interactions between horizontal and vertical physical processes during a
117 convection event [*Estournel et al., 2016a*] and their impact on nutrient budgets and marine
118 ecosystems.

119 In this study, we used the observations collected during three cruises that covered the NW Med
120 in winter, spring and summer of 2013. During DeWEX Leg 1, conducted from 1 to 21 February

121 2013 on board the R/V Le Suroît, 75 hydrological and 45 biogeochemical profiles were obtained
122 (Fig. 1B) [Testor, 2013; Severin *et al.*, 2017], with the objective of globally mapping the
123 convection area, the distribution of the newly formed deep waters and the distribution of
124 inorganic and organic matter. DeWEX Leg 2 sampled the spring bloom from 5 to 24 April 2013
125 [Conan, 2013; Mayot *et al.*, 2017]. It followed the same sampling network as the winter cruise,
126 with 99 hydrological profiles and 59 biogeochemical profiles (Fig. 1C) [Severin *et al.*, 2017].
127 The third cruise was carried out during the oligotrophic period, between 24 July and 7 August
128 2013, in the framework of the integrated observation network MOOSE (Mediterranean Ocean
129 Observing System for the Environment) [Testor *et al.*, 2013]. 100 biogeochemical profiles were
130 measured (Fig. 1D).

131 Then, data provided by the Bio-Argo float LovBio17b between April and September 2013
132 [Pasqueron de Fommervault *et al.*, 2015b] were used to assess the modeled vertical evolution of
133 the nitrate concentration. LovBio17b provided 73 vertical profiles at a daily frequency during the
134 first 55 days of spring, and then at 5-day frequency for the next summer weeks. Calibration was
135 performed at the deployment using *in situ* observations from 0 to 1000 m depth (about 10 m
136 resolution in the 0-250 m layer, and 30 m below).

137



138

139 **Figure 1.** A/ Bathymetry (m) of the coupled physical-biogeochemical model domain. The red
 140 dots represent the modeled river locations. The blue area DCA (Deep Convection Area)
 141 corresponds to the simulation analysis area. The yellow star corresponds to the LION station
 142 where the heat flux plotted on Fig. 6 was modeled. Maps B, C and D represent the
 143 biogeochemical measurement stations of the DeWEX Leg 1 (February 2013), DeWEX Leg 2
 144 (April 2013) and MOOSE-GE 2013 (July 2013) cruises respectively. Stations in red on Fig. B
 145 and C are located on the transect presented in Fig. 3 and Fig. 4.

146

147 **2.2.Modeling**148 **2.2.1. Hydrodynamics**

149 The SYMPHONIE model used in this study is a 3D primitive equation, with free surface and
150 generalized sigma vertical coordinate, described by *Marsaleix et al.* [2008, 2009, 2011 and
151 2012]. It was previously used for the Mediterranean Sea to simulate convection in the deep sea
152 [*Herrmann et al.*, 2008; *Estournel et al.*, 2016b], coastal dense water formation [*Estournel et al.*,
153 2005; *Ulses et al.*, 2008] and circulation on the continental shelf of the Gulf of Lions [*Petrenko*
154 *et al.*, 2008]. The numerical domain (Fig. 1A) covers most of the western Mediterranean basin,
155 using a curvilinear grid with variable horizontal resolution [*Bentsen et al.*, 1999]. The resolution
156 is 1.4 km to the south and about 0.8 km to the north. The southward decrease of the resolution is
157 intended to cover the W Med basin at a more reasonable cost while considering the northward
158 decrease of the Rossby radius, in particular the need for increased resolution in the winter dense
159 water formation area [*Estournel et al.*, 2016b].

160 Forty vertical levels were used with closer spacing near the surface (15 levels in the first 100
161 meters in the center of the convection zone characterized by depths of ~2500 m). The model was
162 initialized and forced at its lateral boundaries with daily analyses provided by the Mercator-
163 Ocean operational system based on the NEMO ocean model [*Maraldi et al.*, 2013]. The
164 configuration of this model was the PSY2V2R4 prototype based on the NEMO Ocean modeling
165 platform and the SAM data assimilation system [*Lellouche et al.*, 2013] at a resolution of 1/12°
166 over the Atlantic and the Mediterranean from 20° S to 80° N. As in *Estournel et al.* [2016b], the
167 initial field and open boundary conditions were corrected for stratification biases deduced from
168 comparisons between analysis and observations taken during the MOOSE cruise of August 2012.
169 The atmospheric forcing (turbulent fluxes) was calculated using the bulk formulae of *Large and*

170 *Yeager* [2004]. Meteorological parameters, including radiative fluxes, were given by the
171 ECMWF operational forecasts at $1/8^\circ$ horizontal resolution and 3-hour temporal resolution based
172 on daily analysis at 00.00 UTC. Since the underestimation of strong winds is a source of
173 uncertainty in atmospheric forcing [*Herrmann et al.*, 2010], *Estournel et al.* [2016b] performed
174 sensitivity tests to the wind speed. According to those tests, the wind velocity was increased by
175 13% in order to increase the accuracy of the model results in reproducing the convection event
176 [*Estournel et al.*, 2016b] in our hydrodynamic simulation. River runoffs were considered using
177 measured daily values for French rivers (data provided by Banque Hydro,
178 www.hydro.eaufrance.fr) and the Ebro (data provided by SAIH Ebro, www.saihebro.com) and
179 mean annual values for the others.

180 **2.2.2. Biogeochemistry**

181 The Eco3M-S model [*Ulses et al.*, 2016] is a multi-nutrient and multi-plankton functional type
182 model that simulates the dynamics of the biogeochemical decoupled cycles of several biogenic
183 elements (carbon, nitrogen, phosphorus, silicon and oxygen) and of non-Redfieldian plankton
184 groups. The model comprises seven compartments. A first compartment of phytoplankton
185 classified by size is described by the mechanistic formulations of the model Eco3M [*Baklouti et*
186 *al.*, 2006], where pico-phytoplankton [$0.7\text{--}2\ \mu\text{m}$] and nanophytoplankton [$2\text{--}20\ \mu\text{m}$] are
187 composed of dinoflagellates, and micro-phytoplankton [$20\text{--}200\ \mu\text{m}$] is composed of diatoms. A
188 second compartment of zooplankton is composed of nano-zooplankton [$5\text{--}20\ \mu\text{m}$] (small
189 flagellates), micro-zooplankton [$20\text{--}200\ \mu\text{m}$] (ciliates and large flagellates) and meso-
190 zooplankton [$>200\ \mu\text{m}$] (copepods and amphipods). A third compartment, bacteria, is also
191 considered. The behavior of heterotrophic organisms is derived from the model by *Anderson and*
192 *Pondaven* [2003]. The other compartments are dissolved organic matter, particulate organic
193 matter (small and large, differentiated by the settling speed and origin), inorganic nutrients
194 (nitrate, ammonium, phosphate and silicate) and dissolved oxygen. A total of 34 state variables

195 are calculated. The model structure (Fig. 2) used in this study is based on the same pelagic
 196 plankton ecosystem model as the one fully described and used by *Auger et al.* [2011] and *Ulses*
 197 *et al.* [2016]. Here we recall the equations of the rates of change of the nitrate and phosphate,
 198 since we will discuss the time evolution of their stock and of their ratios in section 4.1:

$$199 \quad \frac{\partial[NO_3]}{\partial t} = \text{Nitrification} - \text{PhytoplanktonUptake} \quad (1)$$

$$200 \quad \frac{\partial[PO_4]}{\partial t} = \text{ZooplanktonExcretion} + \text{BacteriaExcretion} - \text{PhytoplanktonUptake} \quad (2)$$

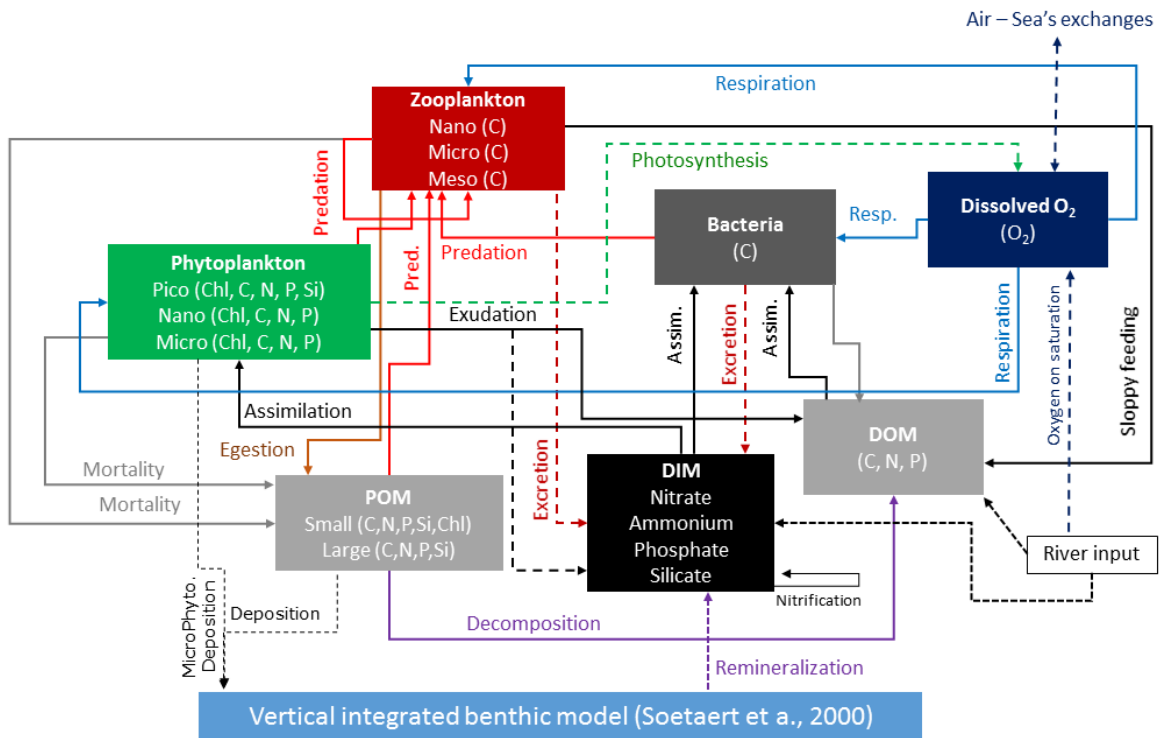
201 where *Nitrification* is the nitrification rate, *PhytoplanktonUptake* is the nutrient uptake by
 202 phytoplankton, and *ZooplanktonExcretion* and *BacteriaExcretion* are the nutrient excretions by
 203 zooplankton and bacteria, respectively. The formulas of these fluxes are given in *Auger et al.*
 204 [2011] (their Table A4). The uptake of nutrients by phytoplankton (*PhytoplanktonUptake*)
 205 depends on the gross primary production rate and on the nutrient concentrations (A19 to A22 in
 206 Table A.4 of *Auger et al.*, 2011). Excretions (*ZooplanktonExcretion* and *BacteriaExcretion*) are
 207 the processes enabling heterotrophs to keep their internal composition constant by releasing the
 208 excess elements in the form of dissolved inorganic matter (A38 to A42 in Table A.4 of *Auger et*
 209 *al.*, 2011). *Nitrification* is the production of nitrate by bacteria, represented here in an implicit
 210 way by a first order function of the ammonium concentration (A55 in Table A.4 of *Auger et al.*,
 211 2011). *Nitrification* is regulated by light intensity and temperature. For the present study of the
 212 NW Med, where winter convection favors strong vertical transport of nutrients in the photic
 213 zone, we neglected the process of nitrogen fixation. This process has been shown to have only a
 214 small effect on stoichiometry in the W Med [*Ribera d'Alcala et al.*, 2003]. Most of the values of
 215 biogeochemical model parameters were based on the previous modeling study by *Ulses et al.*
 216 [2016], but some parameters were re-calibrated using several datasets (MODIS, BOUM,
 217 MOOSE-GE and DeWEX cruises).

218 In this study, we used the “Source Splitting” coupling method [*Butenschön et al.*, 2012]. It consists

219 in an offline forcing of the biogeochemical model by the daily averaged outputs of the physical
220 model. A time step of a few minutes was used for the advection and diffusion of biogeochemical
221 variables, while Eco3M-S computed the biogeochemical fluxes with a time step of about one
222 hour. It was then assumed that biogeochemical properties do not significantly impact the
223 hydrodynamics.

224 The biogeochemical model was downscaled from the Mediterranean basin scale to the regional
225 scale as described hereafter. First, the biogeochemical basin scale model was forced by the daily
226 fields from the NEMO model (PSY2V2R4 analyses), also used for the boundary conditions of
227 the hydrodynamic model as indicated in section 2.2.1. This basin configuration was initialized in
228 June 2010 with climatological nutrient fields from the Medar/MedAtlas database [Manca *et al.*,
229 2004] corresponding to oligotrophic conditions. Daily values of all state variables were extracted
230 from the basin run for the initial and lateral boundary conditions of the regional model. This
231 nesting protocol ensures coherence of the physical and biogeochemical fields at the open
232 boundaries. The regional model was initialized in August 2012.

233 At the mouth of the Rhone river, nitrate, ammonium, phosphate, silicate and dissolved organic
234 carbon concentrations were prescribed using *in situ* daily data [P. Raimbault, personal
235 communication, 2015]. Concentrations of dissolved organic phosphorus and nitrogen and also of
236 particulate organic matter were estimated from these data and the relations deduced from the
237 literature [Moutin *et al.*, 1998; Sempéré *et al.*, 2000] as described in Auger *et al.* [2011]. At the
238 other river mouths, climatological values were prescribed according to Ludwig *et al.* [2010]. The
239 deposition of organic and inorganic matter from the atmosphere was neglected in this study. The
240 benthic fluxes of inorganic nutrients were considered by coupling the pelagic model with a
241 simplified version of the meta-model described by Soetaert *et al.* [2001]. The parameters of this
242 model were set according to the modeling study on the Gulf of Lions shelf performed by Pastor
243 *et al.* [2011].



244

245 **Figure 2.** Eco3m-S biogeochemical model scheme.

246

247 **2.3. Area of study**

248 The analysis of the nitrogen and phosphorus seasonal cycles, and annual budgets, was performed
 249 for the deep convection area of the NW Med, defined here using a mixed layer criterion. This
 250 region (indicated in blue on Fig. 1A), named Deep Convection Area (DCA) hereafter, represents
 251 the area where the mixed layer depth (MLD), defined in section 2.4, exceeded 1000 m for at least
 252 one day during the study period. According to the model, it covered more than 61 700 km² in
 253 2013.

254 For the analysis and budget estimates, we divided the water column into two layers: an upper
255 layer corresponding to the layer where primary production takes place, and a deep layer
256 corresponding to the reservoir of inorganic nutrients. We chose to delimit these two layers with a
257 nutricline criterion. The depth of the nutricline varies over time. As a variable depth would not
258 allow a simple analysis of stock variations and annual budget of processes, we took a constant
259 value corresponding to the maximum depth of the nutricline in the model domain, equal to 130
260 m, as our criterion. As the gradient of nutrient concentration is strong at the nutricline depth, the
261 model estimate of nutrient stock in the upper layer is highly sensitivity to the depth criterion. For
262 instance, an overestimation of the nutricline depth in the model would lead to an underestimation
263 of the modeled nutrient stock. In this study, we focus then on the variation of the nutrient
264 inventory rather than on the value of this inventory. Moreover, we consider the temporal
265 variation of nutricline depth when analyzing that of upper layer nutrient stock.

266 Besides, previous studies [Martin and Pondaven, 2006; Kortzinger et al., 2008] demonstrated
267 that material exported below the productive layer could be resuspended within this zone during
268 mixing periods. Then the maximum yearly mixed layer depth was recommended as a threshold
269 to calculate the export flux or new primary production [Martin and Pondaven, 2006; Palevsky et
270 al., 2016]. The computation of the modeled export of organic and inorganic matter as the sum of
271 sedimentation (for particulate matter) and net flux induced by turbulent mixing and vertical
272 advection in the present study avoids counting the export of organic and inorganic matter that is
273 then reinjected into the upper layer during mixing periods. However, we are aware that if new
274 primary production is considered as the algal growth fueled by nitrate coming in the upper layer
275 for the first time of the studied year, then it could be overestimated due to the recycling of
276 exported organic material in the deep layer whose products would be upwelled back in the upper
277 layer during the mixing period. The estimate of nitrification in the upper layer will be used to
278 estimate this overestimation.

279

280 **2.4. Derived variables**

281 Variables used for the analysis of the annual cycles and budgets were derived from the coupled
 282 model. The modeled MLD is defined as the depth where the potential density exceeds its value at
 283 10 m below the surface by 0.01 kg m^{-3} . We calculated the depths of the 1 mmolN m^{-3} and 0.05
 284 mmolP m^{-3} iso-concentrations of nitrate and phosphate, as a proxy of the depths of the nitracline
 285 and phosphacline, respectively, according to *Lavigne et al.* [2013] and *Lazzari et al.* [2012]
 286 studies. Both depths are significantly correlated in time in this region ($R=0.84$, $p<0.01$) and the
 287 maximum difference between the depths of the nutriclines is 10 m. For this reason, the nutricline
 288 refers to both the nitracline and the phosphacline hereafter. The stock of nutrients in the surface
 289 layer [0-130 m] was calculated in the simulation. Moreover, we computed a second estimate of
 290 the nutrient stock in the Deep Convection Area using the observations from the three cruises
 291 described in section 2.1. The nutrient concentrations measured during the cruises were
 292 interpolated over the model grid and then integrated in the 0-130 m layer as in the model. The
 293 interpolation method considers the notion of distance in space but also the deviation from
 294 selected physical fields (temperature, salinity). In practice, the nitrate and phosphate fields are
 295 interpolated from the cruise data (74 vertical nutrient profiles for MOOSE, 104 nutrient profiles
 296 for DEWEX cruises). It is therefore possible to estimate the deviation of temperature and salinity
 297 between all points of the numerical grid and the points corresponding to the position of the
 298 nutrient profiles to be interpolated. Since the bathymetry of a region such as the Mediterranean
 299 has a great influence on the organization of the circulation, the bathymetry deviation was also
 300 considered in a similar way. Finally, the concentration of the nutrients C at grid point (x, y, z)
 301 was obtained in the following way:

$$302 \quad C(x, y, z) = \frac{\sum_q W_q C(q, z)}{\sum_q W_q} \quad (3)$$

303 where q is the nutrient profile number, $C(q, z)$ the concentration of nutrients at the point

304 corresponding to the position of the nutrient profile q and depth z , and W_q the weight depending
 305 on the tracers and bathymetry deviations according to:

$$306 \quad w_q = e^{-\left(\frac{S_{x,y,z}-S_{q,z}}{\delta S_0}\right)^2} \times e^{-\left(\frac{T_{x,y,z}-T_{q,z}}{\delta T_0}\right)^2} \times e^{-\left(\frac{H_{x,y}-H_q}{\delta H_0}\right)^2} \times e^{-\left(\frac{Dq}{D_0}\right)^2} \quad (4)$$

307 where S is the salinity, and δS_0 is an empirically determined scale of salinity variation (the terms
 308 related to temperature (T) and bathymetry (H) are obtained in a similar manner), Dq is the
 309 horizontal distance between the considered grid point and the nutrient profile q , D_0 is an
 310 empirically determined scale of distance.

311 Finally, we deduced a value of the N:P ratio from the ratio of nitrate to phosphate stocks in the
 312 upper layer, INO_3 and IPO_4 (in mmol m^{-2}), respectively. The contribution of biogeochemical or
 313 physical processes to the change of N:P ratio, was computed with Eq. 5:

$$\frac{\partial(N:P)}{\partial t} = \frac{IPO_4 \frac{\partial INO_3}{\partial t} - INO_3 \frac{\partial IPO_4}{\partial t}}{(IPO_4)^2}$$

314 where $\frac{\partial INO_3}{\partial t}$ and $\frac{\partial IPO_4}{\partial t}$ (in $\text{mmol m}^{-2} \text{d}^{-1}$) are the rates of change of upper layer nitrate and
 315 phosphate stocks, respectively:

$$316 \quad \frac{\partial INO_3}{\partial t} = VertNO_3 + HorAdvNO_3 + I\left(\frac{\partial [NO_3]}{\partial t}\right)_{bio} \quad (6)$$

$$317 \quad \frac{\partial IPO_4}{\partial t} = VertPO_4 + HorAdvPO_4 + I\left(\frac{\partial [PO_4]}{\partial t}\right)_{bio} \quad (7)$$

318 where $VertNO_3$ and $VertPO_4$ are the net upward fluxes of nitrate and phosphate, respectively, at
 319 130 m depth, due to vertical advection and turbulent mixing, $HorAdvNO_3$ and $HorAdvPO_4$ are
 320 the net horizontally advected inputs of nitrate and phosphate, respectively, in the upper layer
 321 from surrounding regions, $I\left(\frac{\partial [NO_3]}{\partial t}\right)_{bio}$ and $I\left(\frac{\partial [PO_4]}{\partial t}\right)_{bio}$ are the upper layer integrated

322 biogeochemical processes (see Eq. 1 and Eq. 2).

323 **2.5 Statistical analysis of the model results**

324 A point-to-point approach was used to quantify the performance of the model in its ability to
 325 represent the dynamics of inorganic nutrients and chlorophyll over the study period. The model
 326 results were compared with the observations at the same dates and positions. Following the
 327 recommendations of *Allen et al.* [2007] we calculated four metrics: the standard deviation ratio
 328 ($r_\sigma = \frac{\sigma_m}{\sigma_o}$ where σ_m and σ_o are the standard deviation of model outputs and observations,
 329 respectively); the Pearson correlation coefficient ($R = \frac{\frac{1}{k} \sum_{n=1}^N (m_n - \bar{m}) - (o_n - \bar{o})}{\sigma_m \times \sigma_o}$), where K is the
 330 number of observations, m_n is the model output that corresponds to the observation n , o_n ,
 331 \bar{m} and \bar{o} are the mean of model outputs and observations respectively); the Nash-Sutcliffe
 332 efficiency ($NS = 1 - \frac{\frac{1}{N} \sum_{n=1}^N (o_n - m_n)^2}{\sum_{n=1}^N (o_n - \bar{o})^2}$) and the percentage bias ($PB = \frac{\bar{m} - \bar{o}}{\bar{o}}$). The calculation of
 333 these four-complementary metrics enabled different aspects of the model results to be assessed.
 334 The standard deviation ratio assessed the variability in the model results compared to that in
 335 observations, a value of $r_\sigma > 1$ indicating that the variability was stronger in the model outputs
 336 than in observations. The correlation coefficient assessed the similarity between model and
 337 observations and significant correlations ($p < 0.01$) were obtained for all R calculations
 338 presented in section 3. The Nash-Sutcliffe efficiency assessed the difference between model and
 339 observations compared to the variability in the observations. According to *Allen et al.* [2007] a
 340 value > 0.65 indicates excellent performance, $[0.65, 0.5]$ is very good, $[0.5, 0.2]$ is good, and < 0.2
 341 indicates poor performance. Finally, the percentage bias indicated a general overestimation in the
 342 model results when it was positive and conversely. According to *Allen et al.* [2007] the
 343 performance of the model can be considered as excellent if the percentage bias is $< 10\%$.

344

345 **3. Comparison of model results with observations**

346 An evaluation of the hydrodynamic simulation is described in *Estournel et al.* [2016b] where
347 modeled temperature and salinity were compared with different available sets of observations..
348 The present study focuses on evaluating the model's ability to reproduce the spatial and temporal
349 variability of inorganic nutrients.

350 **3.1 Comparison with winter, spring and summer cruise observations**

351 To evaluate the model's performance to reproduce the horizontal and vertical
352 biogeochemical patterns in relation with the hydrology, the simulation was compared to the
353 observations from cruises that took place during three key periods for the ecosystem (winter,
354 spring and summer). For winter and spring periods, first, we compared the surface concentrations
355 of nutrients. Then the model/observations comparison was made over the water column along a
356 transect across the deep convection area, from the Balearic Islands to the Gulf of Lions shelf
357 (measurement locations are indicated by the red dots on Fig. 1B for winter and 1C for spring).
358 For both comparisons, the modeled concentrations of nutrients were averaged over the period of
359 sampling: 21 and 20 days for the surface observations, and 2 and 3 days for the transect
360 observations, for the winter and spring cruises, respectively.

361 **3.1.1 Winter**

362 In winter, during the winter cruise, modeled and observed surface concentrations of nitrate and
363 phosphate (Fig. 3A and 3B) show some heterogeneity. The model indicates that the patch of high
364 surface nutrient concentrations ($> 7.5 \text{ mmol m}^{-3}$ for nitrate and $> 0.3 \text{ mmol m}^{-3}$ for phosphate)
365 was located between 40.5°N and 43°N and between 3°E and 8°E . These characteristics and
366 values match the observations, which show surface concentrations $>7.64 \text{ mmol m}^{-3}$ for nitrate
367 and $>0.35 \text{ mmol m}^{-3}$ for phosphate [*Severin et al.*, 2017]. The highest values corresponded to the
368 center of the Deep Convection region where intense convection brought nutrients to the surface

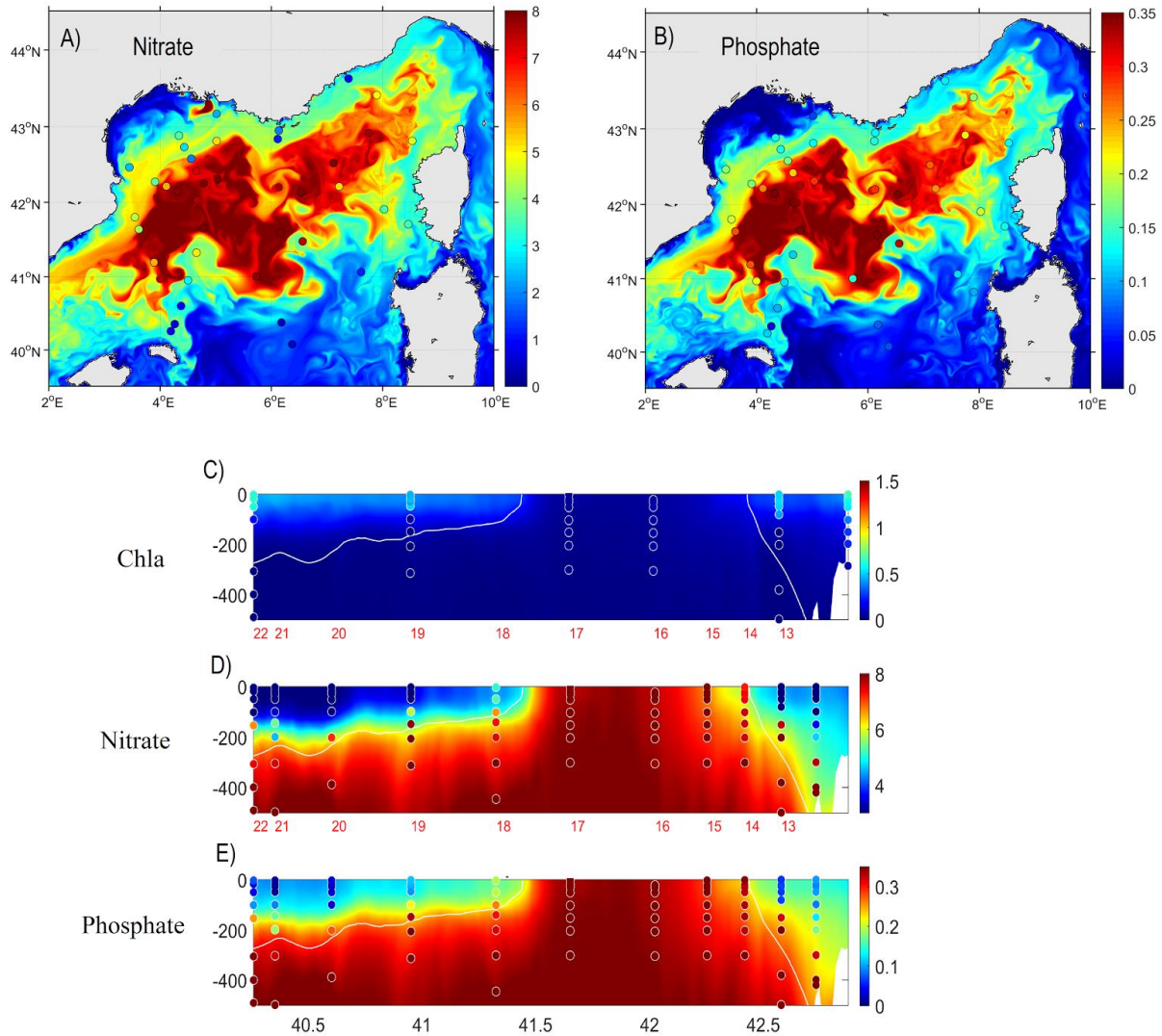
369 layer [Severin *et al.*, 2017]. Outside the convective mixed patch, the surface nutrient
370 concentrations were very low ($<1 \text{ mmol m}^{-3}$ for nitrate and $<0.05 \text{ mmol m}^{-3}$ for phosphate) as
371 detected in the observations south of latitude 41°N . Between these two regions, a strong gradient
372 was observed and modeled on a narrow band. Around the convective patch, in the Balearic Sea
373 and in the Gulf of Lions, the observations show values close to 2.95 mmol m^{-3} for nitrate and
374 0.11 mmol m^{-3} for phosphate, while the model overestimates these values by 1 mmol m^{-3} for
375 nitrate and by 0.4 mmol m^{-3} for phosphate. The strong gradient between the deep convection
376 patch and the surrounding areas was a location of complex frontal dynamics. The greatest biases
377 could originate from vertical motions at very fine scale on these fronts. Nevertheless, it is
378 interesting to note the fine scale structures in the model, which could be explained by the
379 enrichment of deep nutrients induced by the small-scale dynamics.

380 Figures 3C-E give complementary information on the vertical distribution of chlorophyll and
381 nutrient concentrations in the model compared with the observations. Between latitudes 41.5°N
382 and 42.5°N , the modeled and observed water column was homogeneous in biogeochemical
383 variables. Both model and observations show that the surface chlorophyll a concentration was
384 minimum in the deep convection area (lower than 0.10 mg m^{-3} in the model, lower than 0.16 mg
385 m^{-3} in the observations). This feature was previously highlighted using satellite data and 3D
386 coupled modeling [Auger *et al.*, 2014; Houpert *et al.*, 2016; Herrmann *et al.* 2017]) and was
387 explained by a low phytoplankton growth and/or the vertical dilution of phytoplankton cells
388 throughout the whole mixed layer that can reach the ocean bottom ($\sim 2500 \text{ m}$). The maximum
389 vertical gradient of nutrients corresponded to the 29.05 kg m^{-3} isopycnal (superimposed on
390 vertical sections; Fig. 3C-E) at the northern and southern limits of the Deep Convection Area.
391 The isopycnal depth varies between 200 m to the south and 400 m to the north. Beyond this
392 threshold, modeled nitrate concentration increased from 6 to 8 mmol m^{-3} and phosphate
393 concentration from 0.2 to 0.35 mmol m^{-3} .

394 Vertical homogeneity, at least over several hundred meters in the center of the Deep Convection
395 Area, in temperature, salinity [*Estournel et al.*, 2016b] and nutrient concentrations, was also
396 observed during winter 2011 as reported by *Severin et al.* [2014] and *D'Ortenzio et al.* [2014].
397 Other authors [*Yilmaz and Tugrul*, 1998; *Gacic et al.*, 2002; *Dilek et al.*, 2005; *Santinelli et al.*,
398 2012] have reported some homogeneous biogeochemical profiles and high surface inorganic
399 matter enrichment by convective mixing in the eastern Mediterranean basin as mentioned before.

400 When comparing modeled nitrate and phosphate concentrations (mean values and standard
401 deviations: 6.35 ± 1.92 , 0.26 ± 0.09 mmol m⁻³), respectively with the whole winter cruise
402 observations (mean values and standard deviations: 6.45 ± 2.66 , 0.28 ± 0.13 mmol m⁻³), we find
403 a Pearson correlation coefficient of 0.87 and 0.89 (p-value<0.01), respectively, and a Nash-
404 Sutcliffe efficiency of 0.73 and 0.75, respectively, both indicating excellent performance of the
405 coupled simulation. The standard deviation ratio of 0.72 and 0.73, for nitrate and phosphate
406 concentrations respectively, indicates a lower variability in model results than in observations,
407 and the percentage biases of -1.5% and -5.5%, respectively, an underestimation of nutrient
408 concentrations, visible around the convective region on Figure 3. Regarding the modeled
409 chlorophyll concentrations (mean value and standard deviation: 0.12 ± 0.16 mg m⁻³) compared to
410 that observed on the cruise period (mean value and standard deviation: 0.18 ± 0.19 mg m⁻³), the
411 Pearson coefficient correlation of 0.81 and the Nash-Sutcliffe efficiency of 0.56 show that the
412 model reproduces very well the spatial phytoplankton distribution. However, the percentage bias
413 shows that the model underestimates the observations by 32% in average.

414



415

416 **Figure 3.** Comparison of modeled and observed surface (A) nitrate and (B) phosphate
 417 concentrations (mmol m^{-3}) during the winter cruise. Vertical section of (C) chlorophyll (mg m^{-3}),
 418 (D) nitrate (mmol m^{-3}) and (E) phosphate (mmol m^{-3}) concentrations across the Deep Convection
 419 Area, indicated by red dots on Fig. 1B. The model is represented by background colors and
 420 observations are indicated in colored circles (DeWEX Leg1, February 2013). White lines in the

421 vertical sections (C to E) represent the 29.05 kg m^{-3} isopycnal potential density anomaly. The
422 maps (A and B) correspond to an average modeled concentration over the period of the cruise (1
423 to 21 February) while the vertical cross section represents a model average over the transect
424 observation period (12 to 13 February). The numbers of the measurement stations along the
425 vertical section are indicated on the x-axis of panels. C and D. Latitudes are indicated on the x-
426 axis of panel. E.

427 **3.1.2 Spring**

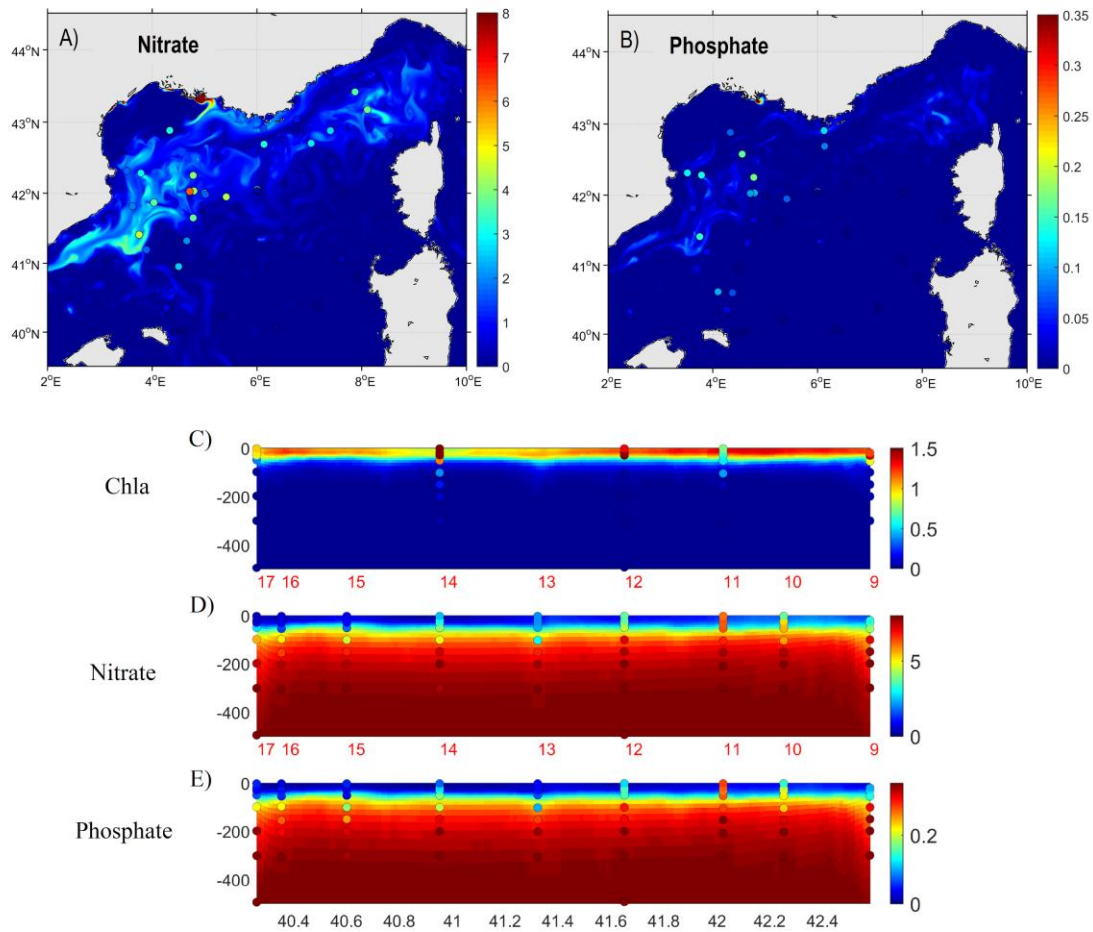
428 In spring, nutrient and chlorophyll patterns (Fig. 4) changed drastically compared to February
429 2013. *Severin al.* [2017] attributed these changes to a strong phytoplankton development, that
430 was considered as a bloom by *Mayot et al.* [2017]. In this study, the onset of the bloom was
431 considered as the day when the net accumulation rate of the carbon biomass integrated through
432 the water column [*Behrenfeld, 2010*] exceeded 50% of its maximum yearly value. Surface
433 nutrient concentrations were low (Fig. 4A and 4B) but remained higher in the Deep Convection
434 Area than outside. In the observations, as in the model, surface values of $2 \text{ to } 3 \text{ mmol m}^{-3}$ of
435 nitrate and $0.08 \text{ to } 0.12 \text{ mmol m}^{-3}$ of phosphate were observed in the center of the basin, while
436 oligotrophic conditions were becoming established outside the Deep Convection Area, with
437 surface nutrient concentrations down to 0 mmol m^{-3} in some areas away from the convection
438 zone. In general, the model appears to underestimate the surface nutrient concentrations.

439 Figures 4D and 4E give an estimation of the nutricline depth, located at 60 m in average
440 (standard deviation: 20 m) in the model, and following the density gradient. A surface
441 chlorophyll development was observed and modeled (Figure 4C) in conjunction with the nutrient
442 depletion (Fig. 4D and 4E), with a concentration exceeding 1.5 mg m^{-3} in the first 30 m, then 0.4
443 $\pm 0.2 \text{ mg.m}^{-3}$ from 30 to 60 m and zero below 60 m. A homogeneous physical, biogeochemical
444 profile was observed at the station $42^\circ\text{N } 5^\circ\text{E}$ in April 8. Although the model reproduces this

445 homogeneous feature, the modeled nutrient concentrations are underestimated and the
446 chlorophyll concentration is overestimated in the surface layer at this station (not shown on the
447 three-days average in Fig. 4C, 4D and 4E). The model simulates during April, several short wind
448 gust events that interrupted the bloom development and mixed the upper layer. The strong gusts
449 of wind associated with negative heat flux that occurred on April 6 and 19 are examples, visible
450 on Figure 7. These events resulted in a temporary decrease of the chlorophyll surface
451 concentrations, and an increase of nutrients in mesoscale structures. Restratification followed the
452 wind events, and reconstructed the bloom conditions. The comparisons of model results with
453 observations suggest that the location of these mesoscale structures is not correctly reproduced in
454 the model.

455 The calculation of metrics indicates that the model reproduces very well the spatial distribution
456 of the whole nitrate and phosphate concentrations observed during the spring cruise (mean
457 values and standard deviations in the model: 5.07 ± 2.92 , 0.21 ± 0.13 mmol m⁻³, respectively,
458 mean values and standard deviations in observations: 5.44 ± 3.09 , 0.23 ± 0.14 mmol m⁻³,
459 respectively): the correlation coefficients are 0.91 and 0.90 (p-value<0.01), respectively, the
460 Nash-Sutcliffe efficiencies are 0.80 and 0.77, respectively. The model presents lower variability
461 by 5 and 6% and underestimates by 6 and 8 % the nitrate and phosphate concentrations,
462 respectively. Regarding the chlorophyll concentrations (mean value and standard deviation in the
463 model: 0.38 ± 0.45 mg m⁻³, mean value and standard deviation in observations: 0.56 ± 0.79 mg
464 m⁻³), a good correlation (0.77) and a good Nash-Sutcliffe efficiency (0.50) between model and
465 observations are found. The percentage bias indicates that the model underestimates observed
466 chlorophyll concentration by 32%.

467



468
 469 **Figure 4.** Same as Figure 3 but for spring cruise (DeWEX Leg 2) in April 2013. Model outputs
 470 are averaged over the period 5-24 April for panel. A and B, and over the period 8-10 April for
 471 panels C, D, and E. The numbers of the measurement stations along the vertical section are
 472 indicated on the x-axis of panels C and D. Latitudes are indicated on the x-axis of panel E.

473 **3.1.3 Summer**

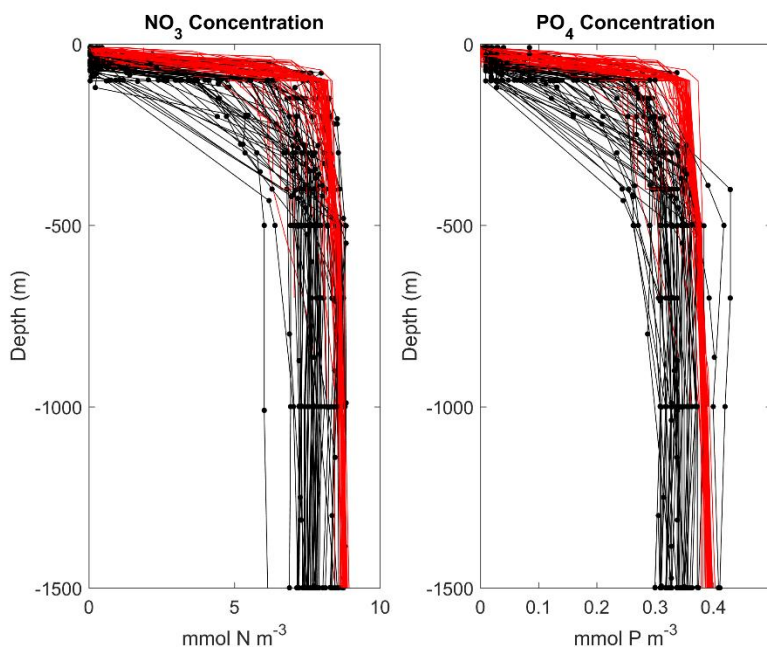
474 Comparisons of modeled nitrate and phosphate concentrations in summer conditions (mean

475 values and standard deviations: 6.24 ± 3.10 , 0.15 ± 0.25 mmol m⁻³, respectively) with the whole
476 summer cruise observations (mean values and standard deviations: 5.02 ± 3.28 , 0.15 ± 0.19
477 mmol m⁻³, respectively), at the same dates and locations, are shown in Fig. 5. From the bottom to
478 300 m depth, the modeled nutrient concentrations fall in the observed range. However, one can
479 notice that most of the modeled nutrient concentrations in deep layers are in the upper range of
480 the observed profiles. The nutrient gradient at 100 m depth is well reproduced by the model.
481 However, higher values of modeled phosphate are visible at 50, 150 and 200 m depth. The model
482 reproduces then correctly the depletion of nutrients in the 0-50 m surface layer.

483 We find a Pearson correlation coefficient of 0.82 and 0.82 (p-value<0.01), and a Nash-Sutcliffe
484 efficiency of 0.52 and 0.46, for nitrate and phosphate concentrations respectively, both metrics
485 indicating excellent performance of the coupled simulation. The standard deviation ratio of 0.94
486 and 1.04, for nitrate and phosphate concentrations respectively, indicates small differences in the
487 variability in model results compared to observations, and the percentage bias of 24% and 32%,
488 respectively, an overestimation of nutrient concentrations in the model.

489

490



491
 492 **Figure 5:** Comparison of modeled (red) and observed (black) nitrate and phosphate
 493 concentration (mmol m^{-3}) profiles during MOOSE cruise in July 2013.

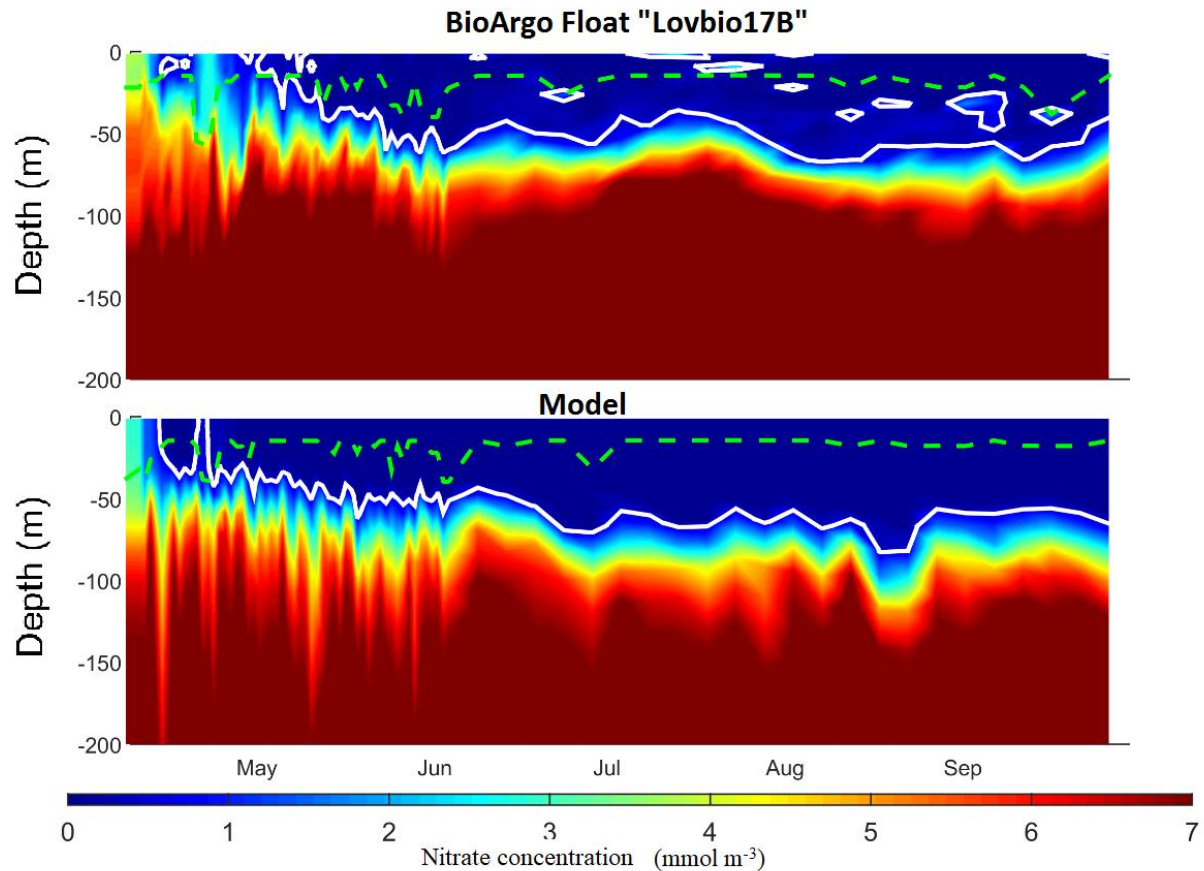
494 3.2 Variability between April and September 2013

495 The objective of this section is to evaluate the ability of the model to reproduce the variability of
 496 the biogeochemistry during a transitional phase. The Bio-Argo float Lovbio17B deployed in the
 497 NW Med in April 2013 during the bloom presents a time-series of the nitrate concentration that
 498 describes the transition in the surface layer from the bloom to the oligotrophic phase.

499 Figure 6 shows an evolution of nitrate concentration measured by the float and simulated by the
 500 model at the same points. The deepening of the nitracline occurs earlier in the model than in the
 501 observations. The nitrate concentration measured by the float in April is 1 mmol m^{-3} higher than
 502 in the model but measured and modeled concentrations are identical in May, with a nutricline

503 located at the same depth (~ 50 m). Variations of nitrate concentrations from 0 to 7 mmol m⁻³ in
504 less than 30 m were observed in the surface layer and were correlated to high frequency motions
505 distinctly detected in both observations and model in April and May. These motions seem to
506 supply nutrients into the mixed layer (the depth of which is indicated by the green dashed line on
507 Fig. 6) more efficiently in the observations. This could be explained by an underestimation of the
508 vertical gradient at 100 m in the model, as shown also in Fig. 4D and 4E. However, in the model
509 results and observations, after each pulse of nutrient injection, the nitrate depletion in the surface
510 layer was fast and the nutricline deepened with time from 50 m at the beginning of June to 70 m
511 during the whole summer. At the same time, a shallow mixed layer was observed around 10-15
512 m.

513 The calculated metrics show that the model can correctly reproduce the transitional phase
514 observed by the float. The Pearson correlation coefficient (R=0.94) and the Nash-Sutcliffe
515 efficiency (NS=0.82) indicate excellent performance of the model. However, slightly higher
516 variability in the observation is indicated by the standard deviation ratio of 0.95. The float
517 measured concentrations higher than 1 mmol m⁻³ in the upper layer from May to October, when
518 the model's concentrations were very low above the nitracline. A general underestimation by the
519 model is indicated by the percentage bias of -16%.



520

521 **Figure 6.** Hovmöller diagram of the nitrate concentration (mmol m^{-3}) along the float track from
522 April to September 2013. Data are extracted from the float Lovbio17B (upper panel) and from
523 the model (lower panel). White line represents the depth of the nitracline (position of the 1 mmol
524 m^{-3} concentration) and the green dashed line represents the MLD.

525

526 **4. Results and Discussion**

527 The comparison exercise reveals that the coupled model is able to reproduce the general seasonal

528 dynamics of inorganic nutrients and phytoplankton growth deduced from the cruise observations.
529 The spatial extension of the convection area, where deep inorganic nutrients were upwelled and
530 phytoplankton were diluted over the whole water column, appears to be in good agreement with
531 cruise observations. The model reproduces the strong inorganic nutrient decrease between winter
532 and spring cruises. Finally, the modeled depletion of nutrients of the upper layer in summer
533 agrees with MOOSE and Lovbio17B observations. However, the model appears to generally
534 underestimate the observed surface nutrient concentrations in spring and to anticipate the
535 deepening of the nutricline observed by the Bio-Argo float.

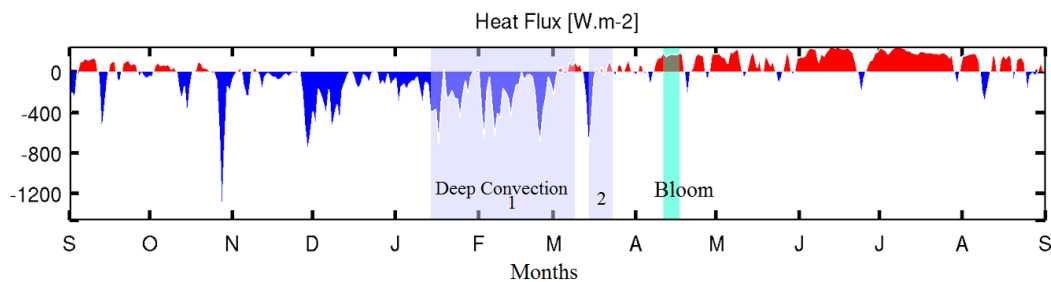
536 Aware of the weaknesses and the strengths of the present coupled simulation, in the following
537 sections, we use the model outputs to disentangle the impact of the physical and biogeochemical
538 processes on the seasonal variations of the nutrient stock and stoichiometry in the upper layer of
539 the Deep Convection Area, from September 2012 to September 2013. Finally, the model is also
540 used to estimate a budget of nitrogen and phosphorus in the upper and deeper layers over the
541 one-year study period.

542 4.1. Annual cycle of nutrients

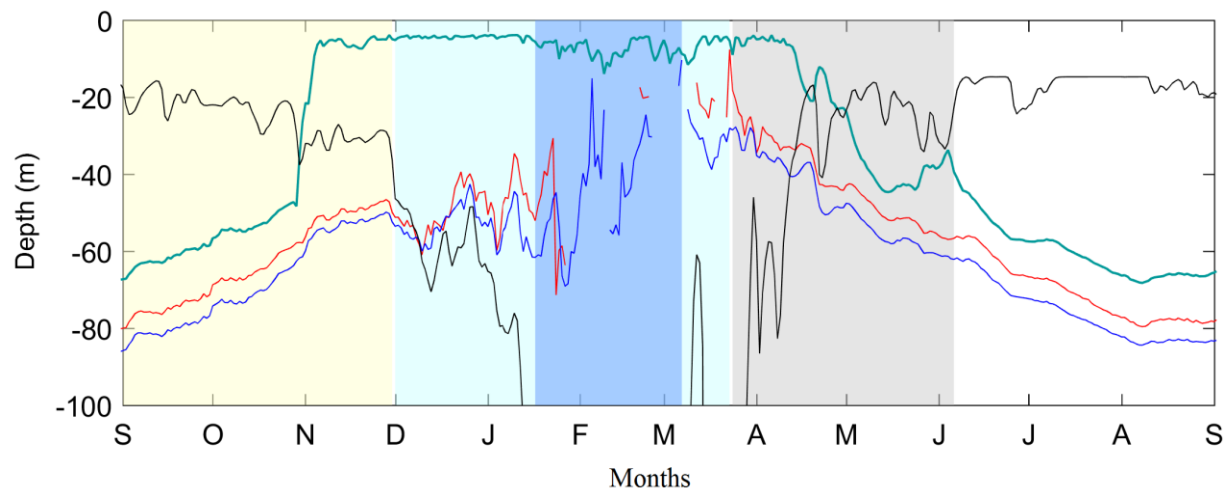
543 **4.1.1 Atmospheric forcing and ocean vertical mixing**

544 Figure 7 presents the time series of the modeled heat flux at the LION station (42°N, 5°E),
545 located in the center of the Deep Convection Area, and Fig. 8 presents the time series of the
546 modeled MLD averaged over the Deep Convection Area, from September 2012 to September
547 2013. Seasonal variability of the atmospheric forcing is clearly visible. In September and
548 October 2012, positive heat fluxes were regularly interrupted by short, strong heat loss events
549 corresponding to wind gusts from the north. In late October, an extreme heat loss event (-1300 W
550 m⁻²) severely weakened the stratification of the upper layer and the spatially-averaged MLD
551 deepened by 20 m (Fig.8). In late November/early December, a strong wind episode lasted three

552 weeks. Then a succession of heat loss events followed from mid-January to the end of February.
 553 From December, each event produced a pronounced increase of the MLD. At the beginning of
 554 February, the ML reached the seafloor in the deep convection center (not shown), when the
 555 spatially averaged MLD reached 1500 m (maximum deepening not shown in Fig. 8). A final
 556 short, strong wind episode occurred between the second and the third weeks of March. This
 557 event caused a breakdown of the weak restratification established during the calm second week
 558 of March. Winter 2013 was thus characterized by a first and main convection event, from mid-
 559 January to early March, and a secondary convection event occurring during the third week of
 560 March. During the last week of March, the heat flux became durably positive in general. The
 561 frequency of the short, intense wind events decreased from three per month in April to one per
 562 month in summer. After March, the heat gain was a component of the progressive surface layer
 563 restratification, the development of instabilities around the mixed patch being the other
 564 component [Estournel *et al*, 2016b]. In spring, large variations (40 m) of the MLD (Fig. 8) can
 565 be noted because of the wind effect on a weak stratified surface layer, as during the third week of
 566 April.



567
 568 **Figure 7.** Time series of modeled total heat flux (W m^{-2}) at LION (indicated on Fig. 1A) from
 569 September 2012 to September 2013. Blue bars represent the period when the MLD was higher
 570 than 1000 m (convective mixing), while the green bar represents the restratification period
 571 coinciding with the phytoplankton development maximum.



572
 573 **Figure 8.** Time series of the depth (m) of the modeled mixed layer (in black), nitracline (in red),
 574 phosphocline (in blue) and deep chlorophyll maximum (in green) averaged over the deep
 575 convection region from September 2012 to September 2013. The background color stands for the
 576 season: yellow for autumn, blue for winter, with the dark blue representing the period when the
 577 MLD was higher than 1000 m (convective mixing), gray for spring and white for summer.

578 4.1.2 Stock of nutrients in the upper layer

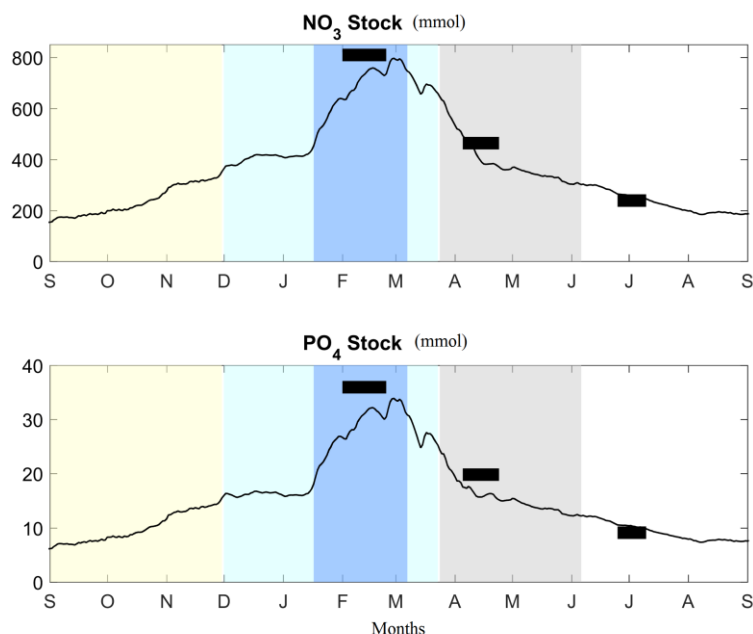
579 Figure 9 presents the annual cycle of the simulated nitrate and phosphate stocks in the upper 130
 580 m of the Deep Convection Area. Stocks deduced from the cruise observations (see section 2.4 for
 581 details) are shown as horizontal bars overlaying the model output. The model reproduces the
 582 trend and orders of magnitude of the nutrient stock as estimated from the observations made
 583 during the three cruises, in winter (800 mmolN m^{-2} and 35 mmolP m^{-2}), spring ($<450 \text{ mmolN m}^{-2}$
 584 and 20 mmolP m^{-2}) and summer ($<200 \text{ mmolN m}^{-2}$ and 10 mmolP m^{-2}). An underestimation of
 585 10 mmolN m^{-2} (7% of the stock) for the stock of nitrate and 4 mmolP m^{-2} for the stock of
 586 phosphate (16% of the stock) is, however, visible in the model estimates for February 2013 when
 587 they are compared with the observations. In this section, we use the model outputs to analyze the

588 impact of the physical and biogeochemical processes on the seasonal variations of the nutrient
589 stock. The biogeochemical interfaces (nitracline, phosphacline and deep chlorophyll maximum –
590 DCM) are plotted in Fig. 8. The year is divided into four periods corresponding to the seasons
591 analyzed individually (they are indicated by colored vertical bars in Figs. 8, 9, 10 and 11). The
592 first season is autumn, from 1 September to 1 December 2012. Winter was chosen to start when
593 both strong negative heat fluxes (from 1 December 2012 to 23 March 2013) (Fig. 7) and
594 relatively deep mixed layer (> 50 m, Fig. 8) conditions co-occurred. Spring, which begins on 24
595 March here, corresponds to the period where the mixed layer became shallower than the
596 epipelagic layer, which was at 120 m in average (standard deviation: 30 m). It ends on 6 June
597 when the mixed layer stabilized at very shallow depth (< 20 m) and the DCM became established
598 below 40 m. Summer is the last stage, which starts on 6 June and lasts until 1 September 2013,
599 the end of the time series.

600

601

602



603

604 **Figure 9.** Time evolution of modeled nitrate (upper panel) and phosphate (lower panel) stocks
 605 (in mmol m^{-2}) in the upper layer [0-130 m] of the Deep Convection Area from September 2012
 606 to September 2013. Stock estimates derived from observations are shown as horizontal bars
 607 overlaying the model outputs. The four seasons identified in section 4.1.2 are indicated with
 608 background colors: yellow represents autumn, blue is for winter - with the dark blue representing
 609 the period when the MLD was higher than 1000 m (deep convective mixing), gray is for spring
 610 and white is for summer.

611 Figure 10 presents the annual cycle of the biogeochemical sink and source terms of nitrate and
 612 phosphate as well as the trends induced by vertical hydrodynamics, for the period 2012-2013.
 613 From early September to the end of October 2012, a shallowing of the nutricline from 80 to 60 m
 614 was modeled (standard deviation of 20 m). This upward displacement of the nutricline (Fig. 8)
 615 could be induced by the isopycnal doming during the preconditioning phase of the deep
 616 convection process in the NW Med [Marshall and Schott, 1999] and/or by the variation of light

617 intensity [Mignot *et al.*, 2014]. This led to a low but continuous increase of the nutrient stock in
618 the upper layer (Fig. 9). At the same time, the uptake of nutrients decreased continuously (Fig.
619 10B). From October to November 2012, uptake of nitrate (respectively phosphate) became lower
620 than nitrification (respectively excretion of phosphate). Therefore, the net effects of both
621 physical and biogeochemical processes led to an increase of the nutrient stock in the upper layer
622 (Fig. 9).

623 The sudden northern wind event occurring at the end of November - beginning of December
624 (Fig. 7), thickening the ML to 50-60 m depth on average, close to the depth of the nutricline (Fig.
625 8), led to an abrupt import of nutrients into the surface layer (Fig. 10A), which were rapidly
626 assimilated by phytoplankton (Fig. 10B). This was followed by a progressive increase of
627 phosphate excretion (Fig. 10C), reflecting the triggering of the development of heterotrophs.
628 Regarding the surface nutrients stock, after the wind gust of early December, the balance of
629 source (vertical import and nitrification/phosphate excretion) and sink (mainly uptake) terms
630 explains the nearly constant stocks (Fig. 9).

631 The deep convection mixing, beginning in mid-January when the MLD exceeded 100 m (Fig. 8),
632 led to a large upwelling of inorganic nutrients to the surface layer (Fig. 10A). The deep
633 convection event also caused a sudden reduction of nitrate and phosphate uptake which,
634 nevertheless, remained non-negligible (Fig 10B). The phytoplankton biomass was vertically
635 mixed, and preys and predators were then decoupled as shown by *Auger et al.* [2014]. This
636 resulted in an increase of the zooplankton mortality as shown by the intense drop of excretion
637 rates during this period (Fig 10C). The decrease of ammonium and phosphate excretion rates was
638 higher than the decrease of nitrate and phosphate uptake rates. It is noteworthy that the uptake of
639 nitrate (mean value: $6.4 \text{ mmolN m}^{-2} \text{ d}^{-1}$) was higher than the uptake of ammonium (mean value:
640 $1.9 \text{ mmolN m}^{-2} \text{ d}^{-1}$) during deep convection (Fig. 10B and 10D). Also, nitrification rate in the
641 upper layer (mean value: $1.2 \text{ mmolN m}^{-2} \text{ d}^{-1}$) was low during this intense mixing period (Fig.

642 10C). Thus, during deep convection, nitrate and phosphate stock in the surface layer increased in
643 the model (Fig. 9) because of the nutrient convective supply and the reduced nutrient uptake.
644 They reached their maximum in early March (Fig. 9), when the main deep convection mixing
645 stopped (Fig. 8).

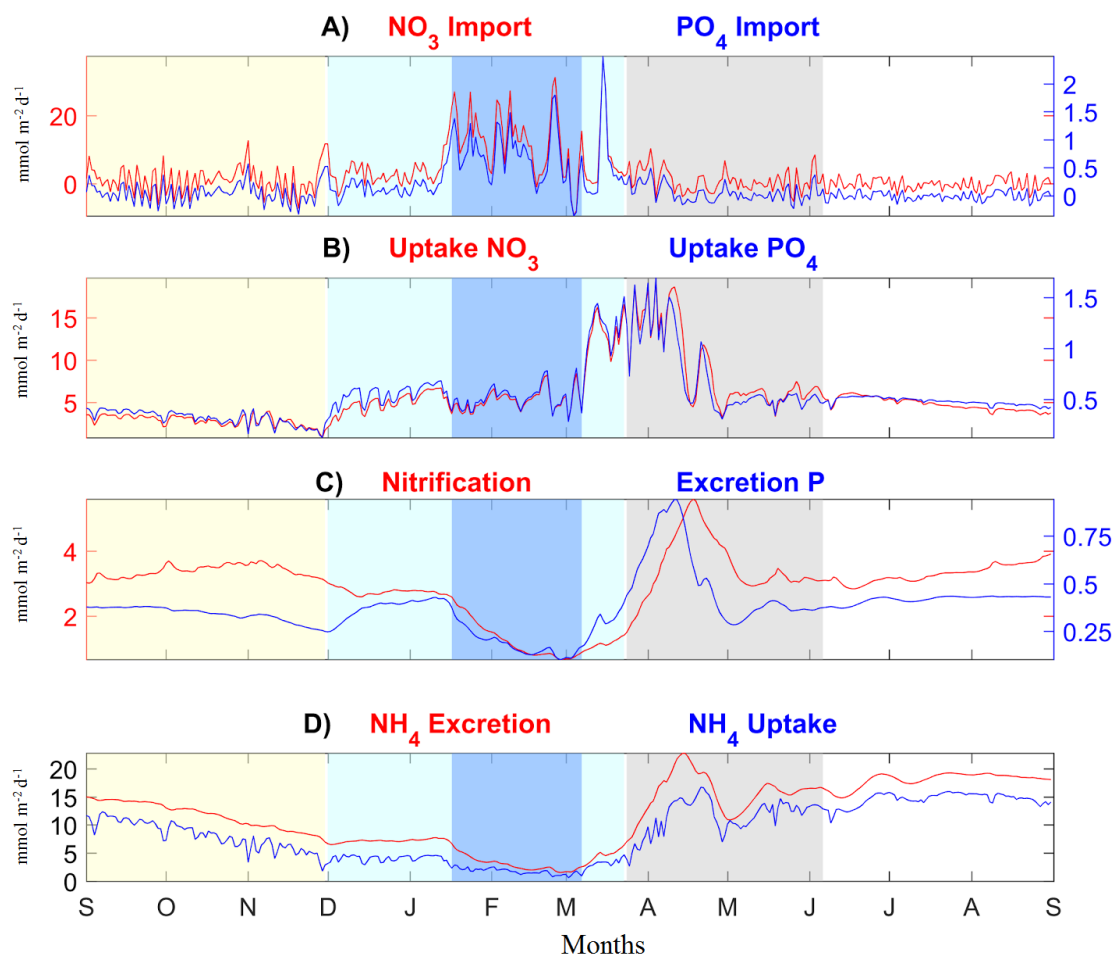
646 An intense phytoplankton development began during the second week of March, favored by the
647 weakening of vertical mixing (not shown). This bloom was characterized by a sudden, strong
648 consumption of nutrients (Fig. 10B) that provoked a first sharp decrease of their surface stocks
649 (Fig. 9).

650 From the end of March to mid-April, nitrate and phosphate uptake remained strong (Fig. 10B)
651 (phytoplankton biomass reached its maximum in mid-April, not shown). In parallel, excretion
652 increased rapidly (Fig. 10C and 10D). In mid-April, the reduction of vertical mixing (Fig. 10A)
653 led to a clear decrease of the vertical enrichment of the upper layer in nutrients. At this period, a
654 strong decrease of nutrient uptake occurred. At the end of April, a new wind gust induced a brief
655 deepening of the mixed layer (Fig. 8) that enriched the surface layer in nutrients. A last strong,
656 short pulse of nutrient uptake (Fig. 10B) was then observed. Afterwards, the nitrate and
657 phosphate stock continuously decreased (Fig. 9), and the nutricline and DCM depth increased
658 (Fig. 8). It is important to note that, after this last strong wind event, ammonium uptake
659 dominated the total nitrogen uptake (Fig. 10B and 10D). Regarding the nutrient stock, it can be
660 concluded that, during the spring period, the strong rate of uptake (higher than the excretion and
661 nitrification rate) explains the strong decreasing trend (Fig. 9).

662 From June to August, the surface layer was depleted in nutrients (Fig. 8 and 9). The model
663 indicates an increase of the phosphate and ammonium excretion as well as of the nitrification
664 rate (Fig. 10C and 10D). The vertical physical flux was very weak (Fig. 10A) and the net
665 biogeochemical flux resulted in a low, continuous decrease of the nitrate and phosphate surface
666 layer stock until the beginning of August, when it reached the value of November 2012 (Fig. 8).

667 In August, as the DCM and nutricline depths, the nutrient stock stabilized.

668 The trend of the contribution of nitrate and ammonium uptake in the total nitrogen uptake, shown
669 in late spring, remained throughout the summer as the trophic regime was dominated by
670 regenerated production, ammonium being continuously produced (Fig. 10D), while nitrate was
671 weakly imported from intermediate layers (Fig. 10A). During summer, nitrification increased
672 slightly but continuously (from 3 to 4 mmol m⁻² d⁻¹, Fig. 10C). It constituted from ~60 to 100%
673 of nitrate uptake. The simulated nitrification rate presents the same order of magnitude as the
674 measurements of *Bianchi et al.* [1999] who reported a rate of 1.4 mmol m⁻² d⁻¹ within the DCM
675 in the NW deep sea. In the model results, we found a value of 1.8 mmol m⁻² d⁻¹ (25% higher than
676 in the observations of *Bianchi et al.* [1999]) by integrating nitrification only within the DCM in
677 the Deep Convection Area.



678

679 **Figure 10.** Time series of physical and biogeochemical fluxes that impact the stock of the
 680 inorganic nitrogen and phosphorus in the surface layer (0-130 m) from September 2012 to
 681 September 2013. These fluxes are inferred from the model and averaged over the open Deep
 682 Convection Area. (A) Net import due to vertical advection and turbulent mixing of nitrate (red)
 683 and of phosphate (blue) into the surface layer, (B) uptake of nitrate (red) and phosphate (blue),
 684 (C) nitrification (red) and inorganic phosphorus excretion rates (blue), (D) ammonia excretion
 685 (red) and uptake (blue). Units: mmol m⁻² d⁻¹.

686 **4.1.3 N:P ratio in the upper layer**

687 In marine ecology, the nitrate to phosphate ratio was determined to be universally 16:1 because
688 of the phytoplankton requirements which control the marine biogeochemistry [Redfield *et al.*,
689 1963]. Even if this averaged ratio has been reappraised several times [Banse, 1994; Geider and
690 La Roche, 2002], its dynamics can still be used to understand the nitrogen and the phosphorus
691 reactivity [Arrigo, 2005]. Indeed, the N:P stoichiometry reflects relations between several
692 processes: nutrients uptake by phytoplankton and small heterotrophs, rate of mineralization of
693 organic matter, inorganic matter excretion by heterotrophs and finally the effects of physical
694 processes (mixing and advection).

695 In the Mediterranean Sea, the ratio N:P is higher than the canonical value of 16:1. It was
696 estimated on average at 24 [Pujo-Pay *et al.*, 2011] with a pronounced horizontal gradient from
697 west to east. To the west, the ratio is 22 in the Levantine Intermediate Water and the Western
698 Mediterranean Deep Water while in the Levantine Basin, the ratio reaches $27 \pm 3:1$ below 200 m
699 [Ribera d'Alcalà *et al.*, 2003]. The N:P ratio can be much higher above the thermocline
700 [Raimbault and Coste, 1990] because of the preference of phytoplankton for phosphate
701 compared to nitrate. These ratios confirm that the Mediterranean waters are limited by
702 phosphate, especially in the eastern basin.

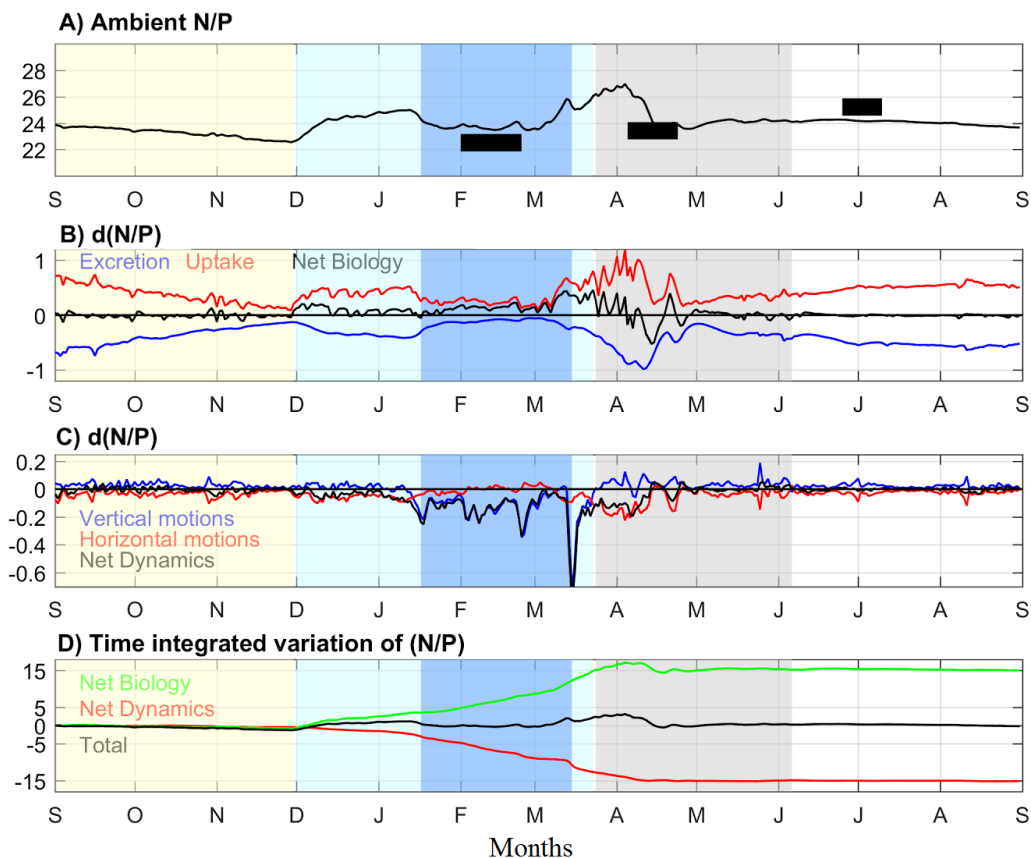
703 In this section, we use our simulation to characterize the time evolution of the stoichiometry in
704 the upper layer of the Deep Convection Area, playing a key role in the water mass
705 biogeochemical characteristics in the Mediterranean Sea, and to quantify the impact of the
706 physical and biogeochemical processes responsible for their seasonal variations, over the period
707 September 2012 to September 2013. The N:P ratio generally presents a negative gradient in the
708 water column as observed in previous observational studies [Pujo-Pay *et al.*, 2011; Huertas *et*
709 *al.*, 2012; Severin *et al.*, 2014; Pasqueron de Fommervault *et al.*, 2015a]. The deep value
710 (averaged between 130 m and the seafloor in the Deep Convection Area, not shown) is 22.4 on

711 average (standard deviation: 0.5) in the model results (it is 21.7 on average, standard deviation:
712 0.6 according to the interpolation of Dewex cruise observations). The maximum N:P vertical
713 gradient is located below the nutricline (not shown). The N:P ratio presents higher variability in
714 the surface layer than in deep layers [*Pasqueron de Fommervault et al.*, 2015].

715 Figure 11A shows the time series of the spatially averaged N:P ratio in the upper layer of the
716 Deep Convection Area. N:P ratio deduced from the three cruise observations are also indicated
717 by horizontal bars. Since the underestimation of phosphate stock is more pronounced than that of
718 nitrate stock in the model during the winter cruise (Figure 9), the model overestimates the N:P
719 ratio during this period. Besides, the overestimation of phosphate stock during MOOSE cruise
720 period lead to an underestimation of N:P ratio during the summer period.

721 A first general remark about the effect of physical and biogeochemical processes on N:P is that
722 uptake increases seawater N:P ratio while release of nitrate and phosphate decreases N:P (Fig.
723 11B). The N:P ratio of uptake is lower than the ratio in seawater, which presents an excess of
724 nitrate compared to phosphate. The effect of uptake is then to increase the water N:P ratio. For
725 the same reason, the release of nutrients by the ecosystem produces a decrease of water N:P. A
726 second remark is that convective mixing events characterized by depths > 200 m produce a
727 decrease of surface N:P ratio because of the general negative gradient of N:P.

728



729

730 **Figure 11.** (A) Time series of the modeled N:P ratio (molN molP^{-1}) averaged in the upper layer
 731 [0-130 m] of the Deep Convection Area from September 2012 to September 2013. Estimates
 732 from observations are shown as horizontal bars overlaying the model output. (B)
 733 Biogeochemical and (C) physical processes ($\text{molN molP}^{-1} \text{d}^{-1}$) influencing the N:P dynamics and
 734 (D) their net time integrated resultant.

735

736 At the beginning of September 2012, the spatial mean N:P ratio was 24.1 (standard deviation:
737 0.5) according to the model (Fig. 11A). From September to December 2012, the influence of
738 each physical process on the N:P ratio was lower than the effect of the biogeochemical processes
739 (Fig. 11C and 11D). Yet, both physical and biogeochemical processes had a negative impact on
740 N:P (-0.7 for the net effect of the biogeochemical processes and -0.5 for the net effect of physical
741 processes), explaining its slow and regular decrease.

742 From December 2012 to mid-January 2013, when nitrate and phosphate uptake rates were higher
743 than the nitrification and phosphate excretion, respectively, the biogeochemical processes tended
744 to strongly increase the N:P ratio (Fig. 11B and 11D). In parallel, the horizontal dynamics tended
745 to slightly decrease the N:P ratio. The resulting evolution was a rapid increase of the N:P ratio
746 from 22.8 to 24.2 in the first 15 days of the autumnal bloom (1 December to 15 December), and
747 a second slow increase from 24.2 to 25 during the following month (15 December to 15
748 January).

749 During the main convective period (16 January to 7 March 2013), the contributions of individual
750 biogeochemical processes to N:P variations were low (Fig 11B). Besides, the strong deepening
751 of the mixed layer in mid-January induced a rapid decrease of surface N:P, due to vertical
752 imports of lower N:P deep waters (Fig. 11C). During the rest of the convective period, the N:P
753 ratio reached a relatively stationary phase: the positive trend of the nutrient uptake that
754 dominated the biology was balanced by the negative trend of the vertical mixing. During this
755 period, the model N:P ratio was around 23, which is close to the mean value of 22.5 (standard
756 deviation: 0.6) deduced from the observations of the February cruise. This difference between
757 the two ratios could be explained by an underestimation of the N:P ratio of the deep waters at the
758 model initialization.

759 Then, the sudden and strong consumption of nutrients at the onset of the winter/spring bloom
760 induced a pronounced increase of the N:P ratio to 26 (Fig. 11A and 11B). This increase was

761 interrupted during the secondary convective mixing of the third week of March, inducing a small
762 decrease of the N:P ratio of 0.3 (Fig. 11C).

763 The period with the highest N:P ratio corresponds to the transition between winter and spring
764 (late March - early April) when the phytoplankton growth and nutrient uptake were strong (Fig.
765 10B and 11A). In early April, phosphate excretion increased proportionally more rapidly than the
766 production of nitrate by nitrification (Fig. 10C). This induced a decrease of the N:P ratio (Fig.
767 11B and 11D), intensified by the effect of vertical physical processes.

768 From mid-April, the contribution of physical processes to the variations of N:P was very low,
769 N:P being mostly affected by the balance between excretion and uptake processes. The time lag
770 between excretion of phosphate, nitrification and uptakes of nitrate and phosphate, as well as the
771 difference in the rate of these processes, produced a biogeochemical contribution that was
772 alternately positive and negative, with a decreasing trend until the end of spring.

773 From early June to early August, the effect of vertical motions balanced the effect of horizontal
774 motions, and, likewise, the effect of uptake processes balanced the effect of production
775 processes. Overall, the variations of N:P were negligible. For the period of the MOOSE summer
776 cruise, the modeled N:P (24.2) is close to the ratio deduced from observations (25.1).

777 Our results shows that the N:P ratio was submitted to strong changes in the upper layer at the
778 beginning of phytoplankton blooms. Those rapid modeled increases of N:P ratio, reflecting a
779 phosphorus limitation for phytoplankton growth, are in agreement with the increase measured by
780 Séverin et al. [2014] in March 2011 after restratification. This is also coherent with the results of
781 *Lazzari et al.* [2016], based on 3D modelling on the whole Mediterranean sea, indicating a
782 limitation of primary production by phosphate in our study area. Two other rapid changes
783 (decrease) were also modeled during the studied period. The first modeled decrease of N:P ratio
784 in the upper layer occurring at the beginning of the deep convection event is coherent with the
785 negative vertical gradient observed in the study region. The second modeled decrease occurred

786 when uptake of nitrate by phytoplankton started to decrease after the peak of the spring bloom.
787 This should be addressed and confirmed with further observations.

788

789 **4.2 Annual budgets of nitrogen and phosphorus**

790 The annual nitrogen and phosphorus budgets of the deep convection area were calculated with
791 the model in the upper layer (0-130 m) and in the deep layer (130 m - seafloor) (Fig. 12). The
792 total stocks of nitrogen and phosphorus in this area exhibit a variation of -0.37% for nitrogen and
793 -1.37% for phosphorus between September 2012 and September 2013.

794 The annual net vertical import of nitrate and phosphate is estimated at $1029 \text{ mmolN m}^{-2} \text{ y}^{-1}$ and
795 $59 \text{ mmolP m}^{-2} \text{ y}^{-1}$ respectively. The nutrient vertical supply occurring over the convection period
796 (17 January - 17 March) corresponds to 68% of the annual vertical supply. According to the
797 comparison exercise (section 3), it may have been underestimated. The simulated vertical
798 nutrients import is of the same order of magnitude as previous estimates for the NW Med [*Pujo-*
799 *Pay and Conan, 2003; Severin et al., 2014*] (Table 1). It is also of the same order of magnitude
800 as the estimation of $1400 \text{ mmolN m}^{-2} \text{ y}^{-1}$ in the convective Atlantic subpolar region [*Williams et*
801 *al., 2000*]. The annual lateral exports of nitrate and phosphate from the surface layer of the
802 convection zone, to the surrounding areas, are low and are estimated at $5 \text{ mmolN m}^{-2} \text{ y}^{-1}$ and 2
803 $\text{ mmolP m}^{-2} \text{ y}^{-1}$ respectively.

804 Regarding the biogeochemical processes budget in the surface layer, the consumption of nitrate
805 and ammonium makes up 39% and 61%, respectively, of the total nitrogen uptake. The annual
806 production of nitrate by nitrification is 103% of the annual nitrate supply from deep waters and
807 50% of the nitrate uptake. Therefore, the new production as defined by *Dugdale and Goering*
808 [1967] (primary production fueled by non-regenerated nutrients) could be overestimated by
809 100% if nitrification rates are taken into account in the calculation. Our results are coherent with

810 those of *Raimbault et al.* [1998], who showed a 20% to 100% overestimation of new primary
811 production in the Pacific Ocean when nitrification was not taken into account in the calculation.
812 They agree also with those of *Dore and Karl* [1996], who established that 47 to 142% of the
813 nitrate assimilated by primary producers was produced by nitrification at the ALOHA
814 oligotrophic station, and those of *Martin and Pondaven* [2006] who found that uptake of
815 regenerated nitrate through nitrification constituted 22 to 57% of the total nitrate uptake. Our
816 results are also close to the results of *Gentilhomme* [1992] who found that nitrification
817 represented 23% of total nitrogen uptake (18% in this study) in the Algerian basin. From a
818 quantitative point of view, using the Redfield ratio to facilitate comparison with previous studies,
819 we have obtained a new production ranging between 77 and 162 gC m⁻² y⁻¹, depending on
820 whether nitrification is considered in the calculation. As mentioned in section 2.3 the choice of a
821 constant depth corresponding to the base of the productive layer could lead to an overestimation
822 of the new production. Indeed, as demonstrated by previous studies [*Martin and Pondaven*,
823 2006; *Kortzinger et al.*, 2008; *Palevsky et al.*, 2016], nitrate originating from the recycling of
824 organic matter exported downward during deep mixing periods could have been reinjected in the
825 upper layer. We estimate nitrification in the deep layer during deep convection period at 158
826 mmol m⁻² (28% of the annual deep nitrification). If the whole nitrate produced by nitrification in
827 the deep layer during convection period is reinjected in the upper layer and used by
828 phytoplankton for its growth, then our estimate of new production is overestimated by 13 gC m⁻²
829 y⁻¹ (using a Redfield ratio). The lower values of our estimate of new production (65-77 gC m⁻² y⁻¹)
830 are in the range of a previous estimate for the offshore NW Med (Table 1), a region that
831 exhibits the highest values in the open Mediterranean Sea [*Béthoux*, 1989].

832 The net amount of organic matter produced in the surface layer is mostly exported towards the
833 deep layer. The amount of produced organic matter advected horizontally toward the adjacent
834 surface regions represents a small percentage of the net amount produced: 11% of the total
835 organic nitrogen and 9% of the total organic phosphorus. The vertical export to the deep layer of

836 dissolved and particulate organic nitrogen and phosphorus is estimated at $724 \text{ mmolN m}^{-2} \text{ y}^{-1}$ and
837 $52 \text{ mmolP m}^{-2} \text{ y}^{-1}$ respectively. These estimates are in the same order of magnitude as the
838 estimates in *Pujo-Pay and Conan* [2003] (Table 1). During the convection period (17 January to
839 17 March), the export represents 55% of the annual amount of organic nitrogen and 58% of the
840 annual amount of organic phosphorus. Two thirds of the total exported organic matter is
841 remineralized by heterotrophs and one third is advected to the surrounding deep regions. With
842 respect to the deep-water biogeochemical nitrogen cycle, the ammonium produced by
843 remineralization is mostly transformed into nitrate by nitrification. The difference in the
844 nitrification:uptake of ammonium ratio (<1 in the upper layer and >1 in the deep layer) in upper
845 and deep layer is explained by both higher ammonium uptake rates in the upper layer and lower
846 nitrification rates due to an inhibition effect by light in the upper layer. An annual import of
847 nitrate ($260 \text{ mmolN m}^{-2} \text{ y}^{-1}$) and phosphate ($6 \text{ mmolP m}^{-2} \text{ y}^{-1}$) is obtained by horizontal advection
848 in the deep layer. It is worth noting that this net horizontal supply results from two strong
849 opposite fluxes: an import of $936 \text{ mmolN m}^{-2} \text{ y}^{-1}$ for nitrate and of $39 \text{ mmolP m}^{-2} \text{ y}^{-1}$ for
850 phosphate in the intermediate waters (between 130 m and 800 m) and an export of 676 mmolN
851 $\text{m}^{-2} \text{ y}^{-1}$ for nitrate and of $33 \text{ mmolP m}^{-2} \text{ y}^{-1}$ for phosphate in the deep waters (below 800 m).

852 The model results suggest that, at the annual scale, the Deep Convection Area appears as an
853 autotrophic system (uptake rates dominate excretion rates) where deep inorganic nutrients
854 injected in the surface layer are almost entirely consumed by phytoplankton. Our estimates for
855 the period 2012-2013 suggest that, at the scale of the Mediterranean Sea, the total vertical supply
856 by physical processes in the NW deep convection region could constitute a major source of
857 nutrients in the surface layer compared to the terrestrial and atmospheric inputs [*Ribera d'Alcala*
858 *et al.*, 2003; *Krom et al.*, 2004] (Table 2).

859 Atmospheric deposition and nitrogen fixation were not taken into account in the simulation.
860 Results of previous studies suggested that they would have little influence on the annual N and P

861 budget in the NW Med deep convection region, where there is a large vertical supply of
 862 nutrients. Aphotic nitrogen fixation was shown to contribute 0.5% of the total nitrogen inputs in
 863 the whole Mediterranean basin [*Benavides et al.*, 2016]. Regarding the atmospheric deposition,
 864 *Lazzari et al.* [2012], using coupled modeling, showed that atmospheric and terrestrial inputs led
 865 to a 5% increase in the net primary production. The experiments performed by *Ridame and*
 866 *Guieu* [2002] showed that, in the western Mediterranean, the inputs of Saharan dissolved
 867 inorganic phosphorus had a negligible impact on the new production integrated over the
 868 productive layer, at the scale of the oligotrophic period, when the contribution of atmospheric
 869 deposition to the total nutrient input to the euphotic layer is largest.

870 The budget of the present study was obtained using a given coupled model and for a one-year
 871 period. Multi-year simulations would allow the variability of these terms in response to the
 872 interannual variability of atmospheric and ocean conditions to be evaluated. Moreover,
 873 uncertainties on the various terms of this budget could be estimated from additional experiments
 874 based on other coupled models.

875

Process	Unit	Reference	Period	Values	
Vertical import of inorganic nutrients	$\text{mmol m}^{-2} \text{y}^{-1}$			NO3	P04
		<i>Pujo-Pay and Conan</i> [2003]	1993	433-650	-

		<i>Severin et al.</i> [2014]	Feb./March 2011 convection events	760/800	36/3 6
		This study	2013 deep convection event	674	40
		This study	Sept. 2012-Sept. 2013	1029	59
New production	mgC m ⁻² y ⁻¹				
		<i>Severin et al.</i> [2014]	Feb./March 2011	46/63	
		This study	Sept. 2012-Sept. 2013	65-77	
Export of organic matter	mmol m ⁻² y ⁻¹				
		<i>Pujo-Pay and Conan</i> [2003]	1993	274-814	
		This study	Sept. 2012-Sept. 2013	724	

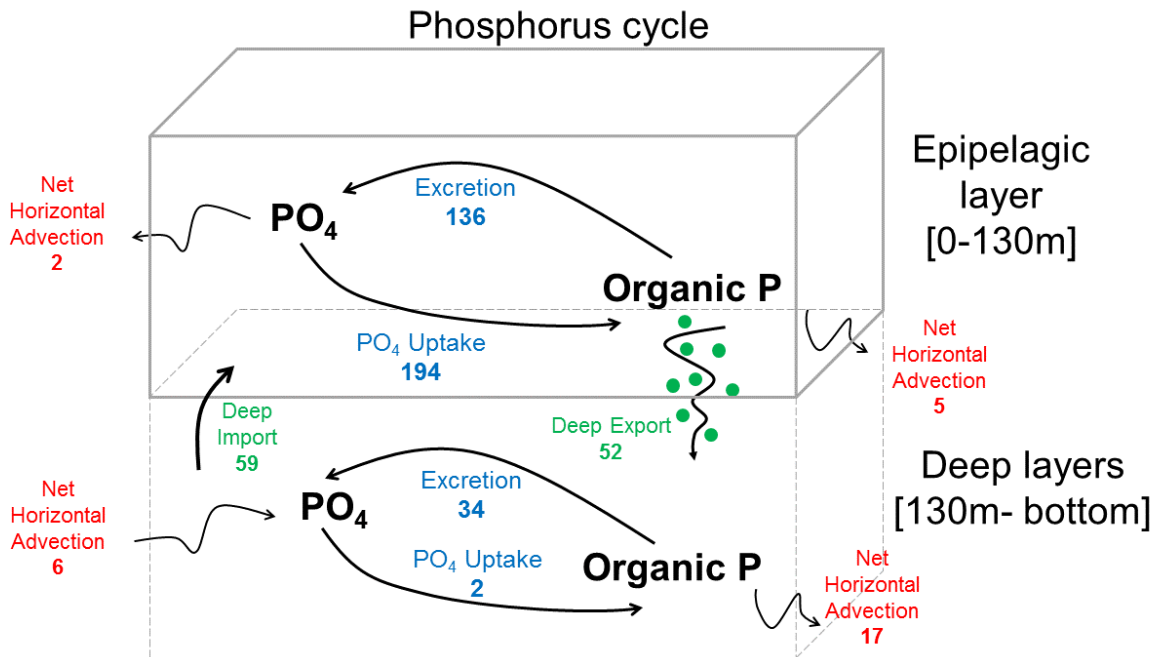
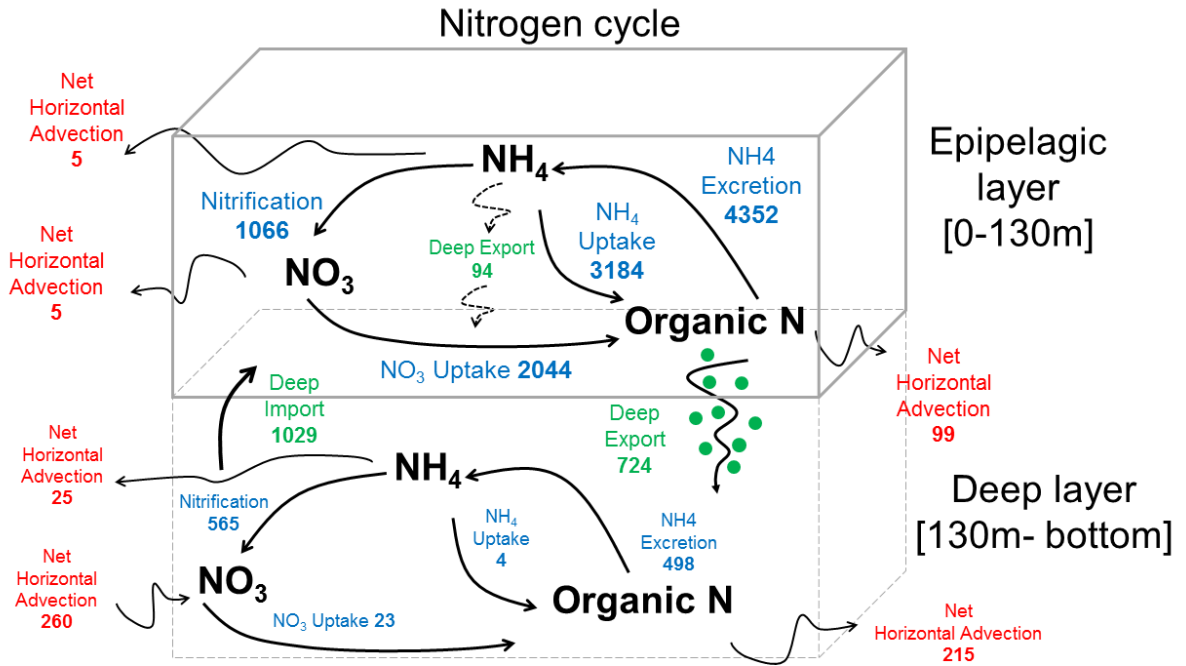
876 **Table 1.** Compilation of nutrient flux estimates in the northwestern Mediterranean deep sea.

877

Sources in the surface layer	N	P	Reference
Atmospheric inputs in the western basin	10,042- 72,825	357-697	Ribera d'Alcalà et al. [2003]
Terrestrial inputs in the western basin	7,840-18,020	168- 1,258	Ribera d'Alcalà et al. [2003]
Atmospheric inputs in the eastern basin	6,064-73,621	379-957	Ribera d'Alcalà et al. [2003]
	111,000	950	Krom et al. [2004]
Terrestrial inputs in the eastern basin	11,592- 77,519	248- 1,541	Ribera d'Alcalà et al. [2003]
	63,000	2,400	Krom et al. [2004]
Vertical physical inputs in the NW Med deep convection area	63,474	3,640	This study

878 **Table 2.** Estimates of the net inputs of nitrogen and phosphorus in the surface layer in the
879 Mediterranean Sea (in 10^6 mol y^{-1}).

880



882 **Figure 12.** Modeled annual fluxes of nitrogen (top panel) and phosphorus (bottom panel) in the
883 deep convection region (in $\text{mmol m}^{-2} \text{y}^{-1}$) over the period from September 2012 to September
884 2013.

885

886 **5. Conclusion and perspectives**

887 A coupled physical-biogeochemical model was compared to a high-resolution dataset of three
888 seasons of the same year (2013) and Bio-Argo float data, which document a strong convective
889 winter, a phytoplankton bloom and an oligotrophic period in the NW Med. Despite some limits
890 highlighted through the comparisons with observations, the model was shown to be able to
891 reproduce the general observed seasonal dynamics of the physical and biogeochemical events
892 correctly.

893 Our study shows that injections of nutrients during the wind intensification period triggered the
894 autumn bloom. Then, convection in winter upwelled large amounts of nutrients into the upper
895 layer. Decoupling of phytoplankton and zooplankton due to dilution as suggested by the collapse
896 of heterotrophs when the mixed layer deepened strongly (mid-January), reduced the grazing
897 pressure. When the conditions for phytoplankton development were all present (reduction of
898 vertical mixing, low grazing pressure), a bloom was triggered with a massive consumption of
899 nutrients during more than one month, resulting in a depletion of nutrients at the surface by the
900 end of April. Nutrient consumption continued to deplete nutrients at increasing depths, inducing
901 a deepening of the nutricline and DCM. This finally led to the summer oligotrophy of the water
902 column.

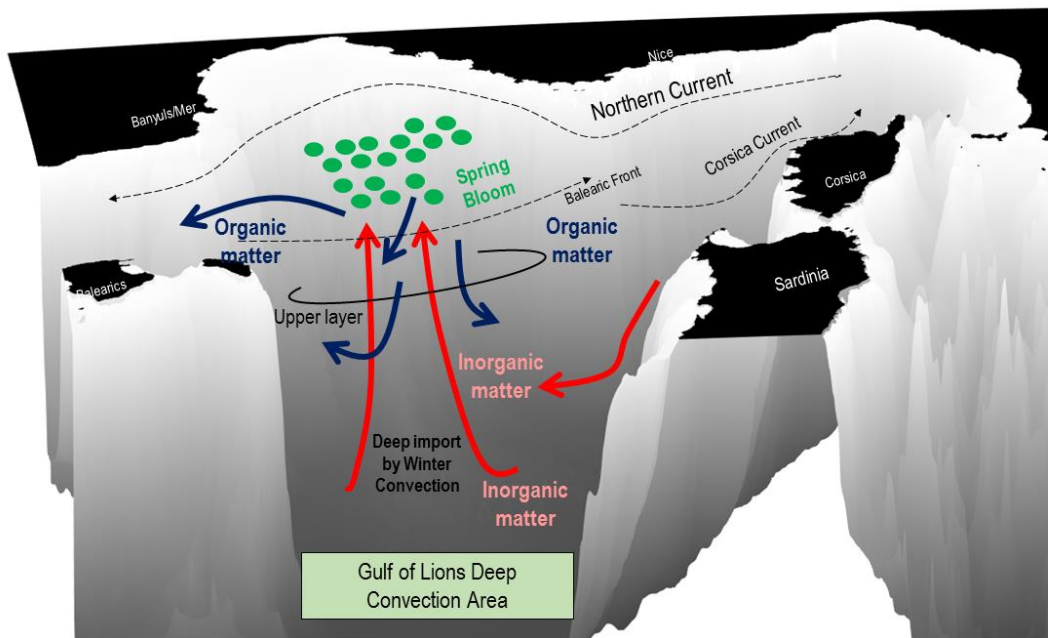
903 The N:P ratio changed mainly through physical or biogeochemical processes during transition
904 periods. Firstly, during the intense convection mixing, the large enrichment of surface waters in
905 deep nutrients suddenly reduced the surface N:P ratio. Secondly, during the beginning of the

906 bloom, the predominance of high nutrient uptake over low excretion induced a strong increase of
907 the N:P. Thirdly, during the peak of the bloom when the nutrient stock decreased strongly,
908 uptake decreased, and excretion was maximum. Nitrification created a decoupling in the
909 recycling of nutrients, decreasing the ratio N:P. At the annual scale, the surface N:P ratio did not
910 appear to be impacted by the intense vertical mixing and spring bloom induced by the deep
911 convection.

912 At the annual scale and for the studied year, the dense water formation area was a sink of nitrate
913 and phosphate in intermediate layers, and a source of organic nitrogen and phosphorus
914 throughout the water column, for the surrounding regions (Fig. 13). The upward nutrient fluxes
915 into the surficial NW Med deep convection area is estimated to be in the same order of
916 magnitude as the atmospheric and terrestrial inputs of each of the western and eastern sub-basins
917 of the Mediterranean Sea.

918 Further multi-year simulations from the coupled model will be used to better understand and
919 quantify the impact of atmospheric and ocean interannual variability of the pelagic ecosystem
920 and associated carbon, nitrogen and phosphorus cycles in this region. Previous studies suggested
921 that deep convection could be strongly weakened [Somot et al., 2006; Adloff et al., 2015] or
922 strengthened [Adloff et al., 2015] by the end of XXI century. We will then use the coupled
923 model to explore and quantify how climate change and anthropogenic forcing may alter the NW
924 Med deep convection biogeochemical cycles, in particular the nutrient vertical supply in the
925 euphotic layer and the new production, in the near future.

926



927
 928 **Figure 13.** Conceptual model of the main processes driving the organic and inorganic matter
 929 dynamics in the NW Mediterranean Sea over the period from September 2012 to September
 930 2013.

931
 932 **Additional Supporting Information**

933

Symbol	Description	Unit	Value		
			Phy ₁	Phy ₂	Phy ₃

T_{Phy}^{REF}	Reference temperature	$^{\circ}C$	14	14	14
τ_{mort,Phy_i}	Natural mortality rate	d^{-1}	0.09	0.03	0.05
Φ_{max,Phy_i}	Maximum quantum yield	$mmolC J^{-1}$	$2.6 \cdot 10^{-4}$	$2.87 \cdot 10^{-4}$	$3.85 \cdot 10^{-4}$
			Zoo ₁	Zoo ₂	Zoo ₃
T_{Zoo}^{REF}	Reference temperature	$^{\circ}C$	18	18	18
τ_{mort,Zoo_1}	Natural mortality rate	d^{-1}	0.01	$0.008 d^{-1}$	-
	Non-living matter remineralization rate				
$\tau_{rem,CDet}$	Detritus remineralisation rate, C	d^{-1}	0.17		
$\tau_{rem,NDet}$	Detritus remineralisation rate, N	d^{-1}	0.2		
$\tau_{rem,PDet}$	Detritus remineralisation rate, P	d^{-1}	0.18		
$\tau_{rem,ChlDet}$	Detritus remineralisation rate, Si	d^{-1}	0.4		
$\tau_{rem,SiDet}$	Detritus remineralisation rate, Chl	d^{-1}	0.02		

935 Table S1. List of recalibrated parameters of the biogeochemical model compared to *Ulses et al.*
936 [2016]

937

938

939

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944 <http://dx.doi.org/10.17600/13450110>), particularly DeWEX and MOOSE projects project of the
945 MISTRALS international program. Numerical simulations were performed using the computing
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948 d'Avenir », ANR J11R107-F).

949 **Figures and tables captions**

950 **Figure 1.** A/ Bathymetry (m) of the coupled physical-biogeochemical model domain. The red
951 dots represent the modeled river locations. The blue area DCA (Deep Convection Area)
952 corresponds to the simulation analysis area. The yellow star corresponds to the LION station
953 where the heat flux plotted on Fig. 6 was modeled. Maps B, C and D represent the
954 biogeochemical measurement stations of the DeWEX Leg 1 (February 2013), DeWEX Leg 2
955 (April 2013) and MOOSE-GE 2013 (July 2013) cruises respectively. Stations in red on Fig. B
956 and C are located on the transect presented in Fig. 3 and Fig. 4

957 **Figure 2.** Eco3m-S biogeochemical model scheme.

958 **Figure 3.** Comparison of modeled and observed surface (A) nitrate and (B) phosphate
959 concentrations (mmol m^{-3}) during the winter cruise. Vertical section of (C) chlorophyll (mg m^{-3}),
960 (D) nitrate (mmol m^{-3}) and (E) phosphate (mmol m^{-3}) concentrations across the Deep Convection
961 Area, indicated by red dots on Fig. 1B. The model is represented by background colors and
962 observations are indicated in colored circles (DeWEX Leg1, February 2013). White lines in the
963 vertical sections (C to E) represent the 29.05 kg m^{-3} isopycnal potential density anomaly. The
964 maps (A and B) correspond to an average modeled concentration over the period of the cruise (1
965 to 21 February) while the vertical cross section represents a model average over the transect
966 observation period (12 to 13 February). The numbers of the measurement stations along the
967 vertical section are indicated on the x-axis of panels. C and D. Latitudes are indicated on the x-
968 axis of panel. E

969 **Figure 4.** Same as Figure 3 but for spring cruise (DeWEX Leg 2) in April 2013. Model outputs
970 are averaged over the period 5-24 April for panel. A and B, and over the period 8-10 April for
971 panels C, D, and E. The numbers of the measurement stations along the vertical section are
972 indicated on the x-axis of panels C and D. Latitudes are indicated on the x-axis of panel E.

973 **Figure 5:** Comparison of modeled (red) and observed (black) nitrate and phosphate
974 concentration (mmol m^{-3}) profiles during MOOSE cruise in July 2013.

975 **Figure 6.** Hovmöller diagram of the nitrate concentration (mmol m^{-3}) along the float track from
976 April to September 2013. Data are extracted from the float Lovbio17B (upper panel) and from
977 the model (lower panel). White line represents the depth of the nitracline (position of the 1 mmol
978 m^{-3} concentration) and the green dashed line represents the MLD.

979 **Figure 7.** Time series of modeled total heat flux (W m^{-2}) at LION (indicated on Fig. 1A) from
980 September 2012 to September 2013. Blue bars represent the period when the MLD was higher
981 than 1000 m (convective mixing), while the green bar represents the restratification period
982 coinciding with the phytoplankton development maximum.

983 **Figure 8.** Time series of the depth (m) of the modeled mixed layer (in black), nitracline (in red),
984 phosphacline (in blue) and deep chlorophyll maximum (in green) averaged over the deep
985 convection region from September 2012 to September 2013. The background color stands for the
986 season: yellow for autumn, blue for winter, with the dark blue representing the period when the
987 MLD was higher than 1000 m (convective mixing), gray for spring and white for summer.

988 **Figure 9.** Time evolution of modeled nitrate (upper panel) and phosphate (lower panel) stocks
989 (in mmol m^{-2}) in the upper layer [0-130 m] of the Deep Convection Area from September 2012
990 to September 2013. Stock estimates derived from observations are shown as horizontal bars
991 overlaying the model outputs. The four seasons identified in section 4.1.2 are indicated with
992 background colors: yellow represents autumn, blue is for winter - with the dark blue representing
993 the period when the MLD was higher than 1000 m (deep convective mixing), gray is for spring
994 and white is for summer.

995 **Figure 10.** Time series of physical and biogeochemical fluxes that impact the stock of the
996 inorganic nitrogen and phosphorus in the surface layer (0-130 m) from September 2012 to
997 September 2013. These fluxes are inferred from the model and averaged over the open Deep
998 Convection Area. (A) Net import due to vertical advection and turbulent mixing of nitrate (red)
999 and of phosphate (blue) into the surface layer, (B) uptake of nitrate (red) and phosphate (blue),
1000 (C) nitrification (red) and inorganic phosphorus excretion rates (blue), (D) ammonia excretion
1001 (red) and uptake (blue). Units: $\text{mmol m}^{-2} \text{d}^{-1}$.

1002 **Figure 11.** (A) Time series of the modeled N:P ratio (molN molP^{-1}) averaged in the upper layer
1003 [0-130 m] of the Deep Convection Area from September 2012 to September 2013. Estimates
1004 from observations are shown as horizontal bars overlaying the model output. (B)
1005 Biogeochemical and (C) physical processes ($\text{molN molP}^{-1} \text{d}^{-1}$) influencing the N:P dynamics and
1006 (D) their net time integrated resultant.

1007 **Figure 12.** Modeled annual fluxes of nitrogen (top panel) and phosphorus (bottom panel) in the
1008 deep convection region (in $\text{mmol m}^{-2} \text{y}^{-1}$) over the period from September 2012 to September
1009 2013.

1010 **Figure 13.** Conceptual model of the main processes driving the organic and inorganic matter
1011 dynamics in the NW Mediterranean Sea over the period from September 2012 to September
1012 2013.

1013 **Table 1:** Compilation of nutrient flux estimates in the northwestern Mediterranean open sea.

1014 **Table 2:** Estimates of the net inputs of nitrogen and phosphorus in the surface layer in the
1015 Mediterranean Sea (in 10^6 mol y^{-1}).

1016 **Table S1.** List of recalibrated parameters of the biogeochemical model compared to Ulses et al.
1017 [2016]

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1019 **References**

- 1020 - Adloff, F., Somot, S., Sevault, F., Jordà, G., et al. (2015) Mediterranean Sea Response to
1021 Climate Change in an Ensemble of Twenty First Century Scenarios. *Climate Dynamics*,
1022 45,2775-2802. <http://dx.doi.org/10.1007/s00382-015-2507-3>
- 1023 - Allen, J. I., Somerfield, P., and Gilbert, F. (2007), Quantifying uncertainty in high-resolution
1024 coupled hydrodynamic-ecosystem models, *J. Marine Syst.*, 64, 3–14,
1025 doi:10.1016/j.jmarsys.2006.02.010, 2007.
- 1026 - Anderson, T. R., and P. Pondaven (2003), Non-redfield carbon and nitrogen cycling in the
1027 Sargasso Sea: Pelagic imbalances and export flux, *Deep Sea Res., Part I*, 50, 573–591.
- 1028 - Antoine, D., A. Morel, H.R. Gordon, V.F. Banzon, and R.H. Evans (2005), Bridging ocean
1029 color observations of the 1980s and 2000s in search of long-term trends, *J. Geophys. Res.*,

- 1030 *110*, C06009, doi:10.1029/2004JC002620.
- 1031 - Arrigo K.R. (2005), Review Marine microorganisms and global nutrient cycles. *Nature* **437**,
1032 349-355 (15 September 2005) | doi:10.1038/nature04159.
- 1033 - Auger, P., C. Ulses, C. Estournel, L. Stemmann, S. Somot, and F. Diaz (2014), Interannual
1034 control of plankton communities by deep winter mixing and prey/predator interactions in the
1035 NW Mediterranean: Results from a 30-year 3D modeling study, *Prog. Oceanogr.*, 124, 12–
1036 27, doi:10.1016/j.pocean.2014.04.004.
- 1037 - Auger, P., F. Diaz, C. Ulses, C. Estournel, J. Neveux, F. Joux, M. Pujo-Pay, and J.-J. Naudin,
1038 (2011), Functioning of the planktonic ecosystem on the Gulf of Lions shelf (NW
1039 Mediterranean) during spring and its impact on the carbon deposition: A field data and 3-D
1040 modelling combined approach, *Biogeosciences*, 8(11), 3231–3261, doi:10.5194/bg-8-3231-
1041 2011.
- 1042 - Baklouti, M., V. Faure, L. Pawlowski, and A. Sciandra (2006), Investigation and sensitivity
1043 analysis of a mechanistic phytoplankton model implemented in a new modular numerical
1044 tool (Eco3M) dedicated to biogeochemical modelling, *Prog. Oceanogr.*, 71(1), 34–58,
1045 doi:10.1016/j.pocean.2006.05.003.
- 1046 - Banse, K., and D. C. English (1994), Seasonality of coastal zone color scanner phytoplankton
1047 pigment in the offshore oceans, *J. Geophys. Res.*, 99(C4), 7323–7345,
1048 doi:10.1029/93JC02155.
- 1049 - Behrenfeld, M.J. (2010), Abandoning Sverdrup’s critical depth hypothesis on phytoplankton
1050 blooms. *Ecology*, 91: 977–989.
- 1051 - Benavides, M., et al. (2016), Basin-wide N₂ fixation in the deep waters of the Mediterranean

- 1052 Sea, *Global Biogeochem. Cycles*, 30, 952–961, doi:10.1002/ 2015GB005326.
- 1053 - Bentsen, M., G. Evensen, H. Drange, and A. D. Jenkins (1999), Coordinate Transformation
1054 on a Sphere Using Conformal Mapping. *Monthly Weather Review*, 2733-2740
- 1055 - Béthoux, J. P., P. Morin, C. Chaumery, O. Connan, B. Gentili, D. Ruiz-Pino (1998).
1056 Nutrients in the Mediterranean Sea, mass balance and statistical analysis of concentrations
1057 with respect to environmental change. *Marine Chemistry* 63 1998. 155–169
- 1058 - Béthoux, J. P., (1989). Oxygen consumption, new production, vertical advection and
1059 environmental evolution in the Mediterranean Sea. *Deep Sea Research*, 36, 5, 769-781.
1060 doi:10.1016/0198-0149(89)90150-7
- 1061 - Bianchi, M, C.Fosset, and P. Conan (1999), Nitrification rates in the NW Mediterranean Sea,
1062 *Aquatic Microbial Ecology*, 17, 267-278.
- 1063 - Butenschön, M., M. Zavatarelli, and M. Vichi (2012), Sensitivity of a marine coupled
1064 physical biogeochemical model to time resolution, integration scheme and time splitting
1065 method, *Ocean Modell.*, 52–53, 36–53, doi:10.1016/j.ocemod.2012.04.008.
- 1066 - Conan, P. (2013), DEWEX-MERMEX 2013 LEG2 cruise, RV Le Suroît, edited by SISMER,
1067 doi:10.17600/13020030.
- 1068 - D'Ortenzio, F. et al. (2014), Observing mixed layer depth, nitrate and chlorophyll
1069 concentrations in the northwestern Mediterranean: A combined satellite and NO₃ profiling
1070 floats experiment, *Geophys. Res. Lett.*, 41, 6443–6451, doi:10.1002/2014GL061020.
- 1071 - D'Ortenzio, F., and M. Ribera d'Alcala (2009), On the trophic regimes of the Mediterranean
1072 Sea: a satellite analysis. *Biogeosciences* 6, 139–148.
- 1073 - Dilek E., A. Yilmaz and S. Tugrul (2005), Vertical profiles of particulate organic matter and

- 1074 its relationship with chlorophyll-a in the upper layer of the NE Mediterranean Sea. *J. Mar.*
1075 *Syst.* 55 (2005) 311– 326.
- 1076 - Dore J. E., and D.M. Karl (1996), Nitrification in the euphotic zone as a source for nitrite,
1077 nitrate, and nitrous oxide at Station ALOHA. *Limnol Oceanogr*, 41,1619-1628
- 1078 - Dugdale R.C., and J.J. Goering J J (1967), Uptake of new and regenerated forms of nitrogen
1079 in primary productivity. *Limnol. Oceanogr*, 12, 196-206.
- 1080 - Durrieu de Madron, X. and Mermex Group (2011), Marine ecosystems' responses to
1081 climatic and anthropogenic forcings in the Mediterranean. *Prog. Oceanogr.*, 91(2), 97–166.
1082 doi:10.1016/j.pocean.2011.02.003.
- 1083 - Estournel, C., P. Testor, I. Taupier-Letage, M.N. Bouin, L. Coppola, P. Durand, P. Conan,
1084 A. Bosse, P.E. Brilouet, L. Beguery, S. Belamari, K. Béranger, J. Beuvier, D. Bourras, G.
1085 Canut, A. Doerenbecher, X. Durrieu de Madron, F. D'Ortenzio, P. Drobinski, V. Ducrocq, N.
1086 Fourrié, H. Giordani, L. Houpert, L. Labatut, C. Lebeaupin Brossier, M. Nuret, L. Prieur, O.
1087 Roussot, L. Seyfried, and S. Somot (2016a), HyMeX-SOP2, the field campaign dedicated to
1088 dense water formation in the north-western Mediterranean. *Oceanography*, 29,4.
- 1089 - Estournel, C., P. Testor , P. Damien , F. D'Ortenzio, P. Marsaleix, P. Conan, F. Kessouri, X.
1090 Durrieu de Madron, L. Coppola, J-M. Lellouche, S. Belamari, L. Mortier, C. Ulses, M-N.
1091 Bouin and L. Prieur (2016b), High resolution modeling of dense water formation in the
1092 north-western Mediterranean during winter 2012-2013: Processes and budget. *J. Geophys.*
1093 *Res. Oceans* DOI: 10.1002/2016JC011935.
- 1094 - Estournel, C., V. Zervakis, P. Marsaleix, A. Papadopoulos, F. Auclair, L. Perivoliotis, and E.
1095 Tragou (2005), Dense water formation and cascading in the Gulf of Thermaikos (North

- 1096 Aegean) from observations and modelling, *Cont. Shelf Res.*, 25, 2366-2386
1097 doi:10.1016/j.csr.2005.08.014.
- 1098 - Estrada, M., M. Latasa, M. Emelianov, A. Gutiérrez-Rodríguez, B. Fernández-Castro, J.
1099 Isern-Fontanet, B. Mouriño-Carballido, J. Salat, M. Vidal (2014), Seasonal and mesoscale
1100 variability of primary production in the deep winter-mixing region of the NW Mediterranean.
1101 *Deep Sea Res. Part I Oceanogr. Res. Pap.*, 94, 45–61.
1102 doi:http://dx.doi.org/10.1016/j.dsr.2014.08.003
- 1103 - Gačić, M., Civitarese, G., Miserocchi, S., Cardin, V., Crise, A., Mauri, E. (2002), The open-
1104 ocean convection in the Southern Adriatic: a controlling mechanism of the spring
1105 phytoplankton bloom. *Cont. Shelf Res.*, 22, 1897–1908. doi:10.1016/S0278-4343(02)00050-
1106 X
- 1107 - Geider R. and La Roche J. (2002), Redfield revisited: variability of C:N:P in marine
1108 microalgae and its biochemical basis, *European Journal of Phycology*, 37:1, 1-17, DOI:
1109 10.1017/S0967026201003456
- 1110 - Gentilhomme, V. (1992), Quantification des flux d'absorption et de régénération de l'azote
1111 minéral (nitrate, nitrite, ammonium) et organique (urée) dans la couche euphotique des
1112 océans oligotrophes. Thèse de l'Université Aix-Marseille
- 1113 - Gogou, A., A. Sanchez-Vidal, X. Durrieu de Madron, S. Stavrakakis, A. M. Calafat, M.
1114 Stabholz, S. Psarra, M. Canals, S. Heussner, I. Stavrakaki, and E. Papathanassiou (2014),
1115 Reprint of: Carbon flux to the deep in three open sites of the Southern European Seas (SES),
1116 *J. Mar. Syst.*, 135, 170–179, doi:10.1016/j.jmarsys.2014.04.012.
- 1117 - Hauser, D., H. Branger, S. Bouffies-Cloch e, S. Despiau, W. Drennan, H. Dupuis, P. Durand,

- 1118 X. Durrieu de Madron, C. Estournel, L. Eymard, C. Flamant, H. Graber, C. Guérin, K.
1119 Kahma, G. Lachaud, J.-M. Lefèvre, J. Pelon, H. Pettersson, B. Piguet, P. Queffeuilou, D.
1120 Tailliez, J. Tournadre, and A. Weill (2003), The FETCH experiment: an overview, *Journal of*
1121 *Geophysical Res.*, 108, doi:10.1029/2001JC001202
- 1122 - Herrmann, M., F. Diaz, C. Estournel, P. Marsaleix, and C. Ulses (2013), Impact of
1123 atmospheric and oceanic interannual variability on the Northwestern Mediterranean Sea
1124 pelagic planktonic ecosystem and associated carbon cycle. *J. Geophys. Res. Oceans*, 118,
1125 5792–5813.
- 1126 - Herrmann, M., S. Somot, F. Sevault, C. Estournel, and M. Déqué (2008), Modeling the deep
1127 convection in the Northwestern Mediterranean sea using an eddy-permitting and an eddy-
1128 resolving model: Case study of winter 1986–87, *J. Geophys. Res.*, 113, C04011,
1129 doi:10.1029/2006JC003991.
- 1130 - Houpert L., P. Testor, X. Durrieu de Madron, S. Somot, F. D’Ortenzio, C. Estournel, H.
1131 Lavigne (2015), Seasonal cycle of the mixed layer, the seasonal thermocline and the upper-
1132 ocean heat storage rate in the Mediterranean Sea derived from observations. *Progress in*
1133 *Oceanography* 132 (2015) 333–352.
- 1134 - Huertas, E., A. F. Ríos, J. García-Lafuente, G. Navarro, A. Makaoui, A. Sánchez-Román, S.
1135 Rodriguez-Galvez, A. Orbi, J. Ruíz, and F. F. Pérez (2012), Atlantic forcing of the
1136 Mediterranean oligotrophy. *Global Biogeochem. Cycles*, Vol. 26, GB2022, doi
1137 :10.1029/2011GB004167.
- 1138 - Krom M. D., Herut B., Mantoura R. F. C., (2004), Nutrient budget for the Eastern
1139 Mediterranean: Implications for phosphorus limitation, *Limnology and Oceanography*, 5,

- 1140 doi: 10.4319/lo.2004.49.5.1582.
- 1141 - Körtzinger, A., U. Send, D. W. R. Wallace, J. Kartensen, and M.
1142 DeGrandpre (2008), Seasonal cycle of O₂ and pCO₂ in the central Labrador Sea:
1143 Atmospheric, biological, and physical implications, *Global Biogeochem. Cycles*, 22,
1144 GB1014, doi:[10.1029/2007GB003029](https://doi.org/10.1029/2007GB003029).
- 1145 - Krom, M. D., Kress, N., Brenner, S., Gordon, L.I., (1991), Phosphorus limitation of primary
1146 productivity in the eastern Mediterranean Sea. *Limnol. Oceanogr.* 36, 424–432.
- 1147 - Large, W. G. and S. G. Yeager (2004), Diurnal to decadal global forcing for ocean and sea-
1148 ice models: the data sets and flux climatologies. NCAR Technical Note NCAR/TN-
1149 460+STR, DOI: 10.5065/D6KK98Q6.
- 1150 - Lavigne, H., D'Ortenzio, F., M. Ribera D'Alcala, H. Claustre, R. Sauzede and M. Gacic
1151 (2015), On the vertical distribution of the chlorophyll-a 1 concentration in the Mediterranean
1152 Sea: a basin scale and seasonal approach. *Biogeoscience*, 12, 5021–5039, doi:[10.5194/bg-12-](https://doi.org/10.5194/bg-12-5021-2015)
1153 [5021-2015](https://doi.org/10.5194/bg-12-5021-2015)
- 1154 - Lavigne, H. (2013), Impact of mixed layer depth seasonal variations on the phytoplankton
1155 phenology in the Mediterranean Sea, PhD thesis, University Pierre et Marie Curie, Paris,
1156 France.
- 1157 - Lazzari P., C. Solidoro, S. Salon, G. Bolzon. (2016),
1158 Spatial variability of phosphate and nitrate in the Mediterranean Sea: A modeling approach.
1159 *Deep Sea Research Part I: Oceanographic Research Papers*. Volume 108, February 2016,
1160 Pages 39-52. <https://doi.org/10.1016/j.dsr.2015.12.006>
- 1161 - Lellouche, J.-M., O. Le Galloudec, M. Drévillon, C. Régnier, E. Greiner, G. Garric, N. Ferry,

- 1162 C. Desportes, C.E. Testut, C. Bricaud, R. Bourdallé-Badie, B. Tranchant, M. Benkiran, Y.
1163 Drillet, A. Daudin, and C. De Nicola (2013), Evaluation of global monitoring and forecasting
1164 systems at Mercator Océan, *Ocean Sci.*, 9, 57-81, doi:10.5194/os-9-57-2013.
- 1165 - Ludwig, W., A. F. Bouwman, E. Dumont, and F. Lespinas (2010), Water and nutrient fluxes
1166 from major Mediterranean and Black Sea rivers: Past and future trends and their implications
1167 for the basin-scale budgets, *Global Biogeochem. Cycles*, 24, GB0A13,
1168 doi:10.1029/2009GB003594.
- 1169 - Manca, M., M. Burca, A. Giorgetti, C. Coatanoan, M.J. Garcia, and A. Iona (2004), Physical
1170 and biochemical averaged vertical profiles in the Mediterranean regions: an important tool to
1171 trace the climatology of water masses and to validate incoming data from operational
1172 oceanography, *J. Mar. Syst.*, 48, 1–4, 83-116.
- 1173 - Maraldi C., J. Chanut, B. Levier, N. Ayoub, P. De Mey, G. Reffray, F. Lyard, S. Cailleau,
1174 M. Drevillon, E.A. Fanjul, M.G. Sotillo, P. Marsaleix, and the Mercator Research and
1175 Development Team (2013), NEMO on the shelf: assessment of the Iberia–Biscay–Ireland
1176 configuration. *Ocean Science*, 9, 745–771. <http://dx.doi.org/10.5194/os-9-745-2013>
- 1177 - Marshall, J., and F. Schott (1999), Open-ocean convection: Observations, theory, and
1178 models, *Rev. Geophys.*, 37(1), 1–64, doi:10.1029/98RG02739.
- 1179 - Marsaleix, P., F. Auclair, J.W. Floor, M.J. Herrmann, C. Estournel, I. Pairaud, and C. Ulses
1180 (2008), Energy conservation issues in sigma-coordinate free-surface ocean models. *Ocean*
1181 *Modell.* 20, 61-89. <http://dx.doi.org/10.1016/j.ocemod.2007.07.005>
- 1182 - Marsaleix, P., C. Ulses, I. Pairaud, M.J. Herrmann, J.W. Floor J. W., C. Estournel, and F.
1183 Auclair (2009), Open boundary conditions for internal gravity wave modelling using

- 1184 polarization relations. *Ocean Modell.*, 29, 27-42.
1185 <http://dx.doi.org/10.1016/j.ocemod.2009.02.010>
- 1186 - Marsaleix, P., F. Auclair, C. Estournel, C. Nguyen, and C. Ulses (2011), An accurate
1187 implementation of the compressibility terms in the equation of state in a low order pressure
1188 gradient scheme for sigma coordinate ocean models. *Ocean Modell.*, 40, 1-13
1189 <http://dx.doi.org/10.1016/j.ocemod.2011.07.004>
- 1190 - Marsaleix, P., F. Auclair, T. Duhaut, C. Estournel, C. Nguyen, and C. Ulses (2012),
1191 Alternatives to the Robert-Asselin filter. *Ocean Modell.*, 41, 53-66
- 1192 - Martin, A.P.; Pondaven, P. (2006), New primary production and nitrification in the western
1193 subtropical North Atlantic: a modeling study. *Global Biogeochemical Cycles*, 20 (4).
1194 [GB4014.10.1029/2005GB002608](https://doi.org/10.1029/2005GB002608)
- 1195 - Marty, J-C., J. Chiavérini, M-D.Pizay, and B. Avril (2002), Seasonal and interannual
1196 dynamics of nutrients and phytoplankton pigments in the western Mediterranean Sea at the
1197 DYFAMED time-series station (1991–1999). *Deep-Sea Res. II*, 49, 1965–1985.
- 1198 - Mayot, N., D'Ortenzio, F., Ribera d'Alcalà, M., Lavigne, H., and H. Claustre (2016), The
1199 Mediterranean trophic regimes from ocean color satellites: a reappraisal. *Biogeoscience*, 13,
1200 1901–1917, [doi:10.5194/bg-13-1901-2016](https://doi.org/10.5194/bg-13-1901-2016)
- 1201 - Mayot, N., D'Ortenzio, F., Taillandier, V., Prieur, L., Pasqueron de Fommervault, O.,
1202 Claustre, H., Bosse, A., Testor, P and Conan, P. (2017) Physical and biogeochemical controls
1203 of the phytoplankton blooms in North Western Mediterranean Sea: a multiplatform approach
1204 over a complete annual cycle (2012-2013 DEWEX experiment), *J. Geophys. Res.*
- 1205 - MEDOC Group (1970), Observation of formation of deep water in the Mediterranean Sea,

- 1206 Nature, 227(1037), 1040.
- 1207 - Mignot, A., H. Claustre, J. Uitz, A. Poteau, F. D'Ortenzio, and X. Xing (2014),
1208 Understanding the seasonal dynamics of phytoplankton biomass and the deep chlorophyll
1209 maximum in oligotrophic environments: A Bio-Argo float investigation, *Global*
1210 *Biogeochem. Cycles*, 28, doi:10.1002/2013GB004781.
- 1211 - Moutin, T., P. Raimbault, H.L. Golterman, and B. Coste (1998), The input of nutrients by the
1212 Rhône river into the Mediterranean Sea: recent observations and comparison with earlier
1213 data, *Hydrobiologia*, 373: 237. doi:10.1023/A:1017020818701
- 1214 - Nittis, K., A. Lascaratos, A. Theocharis (2003), Dense water formation in the Aegean Sea:
1215 Numerical simulations during the Eastern Mediterranean Transient. *Journal of Geophysical*
1216 *Research*. Vol.. 108, No. c9, 8120, doi:10.1029/2002JC001352, 2003
- 1217 - Ovchinnikov, I. M. (1984). The formation of Intermediate Water in the Mediterranean.
1218 *Oceanology* 24: 168–173.
- 1219 - Pasqueron de Fommervault, O., C. Migon, F. D'Ortenzio, M. Ribera d'Alcalà and L. Coppola
1220 (2015a) , Temporal variability of nutrient concentrations in the northwestern Mediterranean
1221 sea (DYFAMED time-series station), *Deep Sea Res. I*, 100, 1-12.
- 1222 - Pasqueron de Fommervault, O., F. D'Ortenzio, A. Mangin, R. Serra, C. Migon, H. Claustre,
1223 H. Lavigne, M. Ribera d'Alcalà, L.Prieur, V. Taillandier, C. Schmechtig, A. Poteau, E.
1224 Leymarie, A. Dufour, F. Besson, G. Obolensky (2015b), Seasonal variability of nutrient
1225 concentrations in the Mediterranean Sea: Contribution of Bio-Argo floats. *J. Geophys. Res.*
1226 *Oceans*, 120, 8528–8550, doi:10.1002/2015JC011103.
- 1227 - Pastor, L., C. Cathalot, B. Deflandre, E. Viollier, K. Soetaert , F. J. R. Meysman, C. Ulises, E.

- 1228 Metzger, and C. Rabouille (2011), Modeling biogeochemical processes in sediments from the
1229 Rhone River prodelta area (NW Mediterranean Sea), *Biogeosciences*, 8, 1351-1366
- 1230 - Petrenko, A., C. Dufau, and C. Estournel (2008), Barotropic eastward currents in the western
1231 Gulf of Lion, north-western Mediterranean Sea, during stratified conditions, *J. Mar. Syst.*, 74,
1232 406-428 doi:10.1016/j.jmarsys.2008.03.004
- 1233 - Pollak, M. (1951), The sources of deep water of the Eastern Mediterranean Sea. *J. Mar. Res.*
1234 10: 128–152
- 1235 - Pujo-Pay, M., P. Conan, L. Oriol, V. Cornet-Barthaux, C. Falco, J.-F. Ghiglione, C. Goyet,
1236 T. Moutin, and L. Prieur (2011), Integrated survey of elemental stoichiometry (C, N, P) from
1237 the western to eastern Mediterranean Sea. *Biogeosciences*, 8(4), 883–899. doi:10.5194/bg-8-
1238 883-2011.
- 1239 - Pujo-Pay, M. and Conan, P. (2003), Seasonal variability and export of dissolved organic
1240 nitrogen in the northwestern Mediterranean Sea. *Journal of Geophysical Research* 108: doi:
1241 10.1029/2000JC000368. issn: 0148-0227.
- 1242 - Raimbault, P., Coste, B., (1990), Very high values of nitrate:phosphate ratio (>30) in the
1243 subsurface layers of the western Mediterranean Sea. *Rapp. Comm. Int. Mer. Mediterr.* 32 (1),
1244 C-18.
- 1245 - Raimbault, P., G. Slawyk, B. Boudjellal, C. Coatanoan, P. Conan, B. Coste, N. Garcia, T.
1246 Moutin, M. Pujo-Pay (1998), Carbon and nitrogen uptake and export in the equatorial Pacific
1247 at 150°W: evidence of an efficient regenerated production cycle. *J Geophys Res* 104:3341-
1248 3356.
- 1249 - Redfield, A. C., B. H. Ketchum and F. A. Richards (1963), The influence of organisms on

- 1250 the composition of sea-water. Contribution No 1113, Woods Hole Oceanographic Institution
1251 and Contribution No. 238, Department of Oceanography of the University of Washington.
- 1252 - Ribera d'Alcalà, M., G. Civitarese, F. Conversano, and R. Lavezza (2003), Nutrient ratios
1253 and fluxes hint at overlooked processes in the Mediterranean Sea, *J. Geophys. Res.*, 108,
1254 8106, doi:10.1029/2002JC001650, C9.
- 1255 - Ridame, C., and Guieu, C. (2002). Saharan input of phosphate to the oligotrophic water of
1256 the open western Mediterranean Sea. *Limnol. Oceanogr.* 47, 856–869. doi:
1257 10.4319/lo.2002.47.3.0856
- 1258 - Santinelli, C., V. Ibello, R. Lavezza, G. Civitarese, and A. Seritti (2012), New insights into
1259 C, N and P stoichiometry in the Mediterranean Sea: The Adriatic Sea case. *Cont. Shelf Res.*,
1260 44 (2012) 83–93
- 1261 - Sempéré, R., B. Charrière, F. Van Wambeke, and G. Cauwet (2000), Carbon Inputs of the
1262 Rhône River to the Mediterranean Sea: Biogeochemical Implications, *Global Biogeochem.*
1263 *Cycles*, 14, 2, 669–681.
- 1264 - Severin, T., et al. (2017), Open-ocean convection process: A driver of the winter nutrient
1265 supply and the spring phytoplankton distribution in the Northwestern Mediterranean Sea,
1266 *J. Geophys. Res. Oceans*, 122, doi:10.1002/2016JC012664.
- 1267 - Severin, T., P. Conan, X. Durrieu de Madron, L. Houpert, M.J. Oliver, L. Oriola, J. Caparros,
1268 J.F. Ghiglione, and M. Pujó-Pay (2014), Impact of open-ocean convection on nutrients,
1269 phytoplankton biomass and activity, *Deep Sea Res. I*, doi: 10.1016/j.dsr.2014.07.015
- 1270 - Soetaert, K., P.M.J.
- 1271 - Herman, J.J. Middelburg, C. Heip, C.L. Smith, P. Tett, and K. Wild-Allen (2001), Numerical

- 1272 modelling of the shelf break ecosystem: reproducing benthic and pelagic measurements.
1273 *Deep-Sea Res. II*, 48, 3141 – 3177.
- 1274 - Somot, S., F. Sevault, and M. Déqué, 2006: Transient climate change scenario simulation of
1275 the Mediterranean Sea for the 21st century using a high-resolution ocean circulation model.
1276 *Climate Dyn.*, 27, 851–879
- 1277 - Testor, P. (2013), DEWEX-MERMEX 2013 LEG1 cruise, RV Le Suroît, edited by SISMER,
1278 doi:10.17600/13020010.
- 1279 - Testor, P., L. Coppola, and L. Mortier (2013), MOOSE-GE 2013 cruise, RV Téthys II, edited
1280 by SISMER, doi:10.17600/13450110.
- 1281 - Ulses, C., C. Estournel, P. Puig , X. Durrieu de Madron, and P. Marsaleix (2008), Dense
1282 shelf water cascading in the northwestern Mediterranean during the cold winter 2005:
1283 Quantification of the export through the Gulf of Lion and the Catalan margin. *Geophys. Res.*
1284 *Lett.* 35.
- 1285 - Ulses, C., P.-A. Auger, K. Soetaert, P. Marsaleix, F. Diaz, L. Coppola, M. Herrmann, F.
1286 Kessouri, and C. Estournel (2016), Budget of organic carbon in the North-Western
1287 Mediterranean Open Sea over the period 2004-2008 using 3D coupled physical
1288 biogeochemical modeling, *J. Geophys. Res.* , doi:10.1002/2016JC011818.
- 1289 - Williams, R. G., and A. J. McLaren (2000), Estimating the convective supply of nitrate and
1290 implied variability in export production over the North Atlantic, *Global Biogeochemical*
1291 *Cycles*. Vol. 14. , NO 4, Pages 1299-1313.
- 1292 - Yilmaz, A. and S. Tugrul (1998), The effect of cold- and warm-core eddies on the
1293 distribution and stoichiometry of dissolved nutrients in the northeastern Mediterranean. *J.*

1294 Mar. Syst., 16, 3-4, 253–268.

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Figure 1.

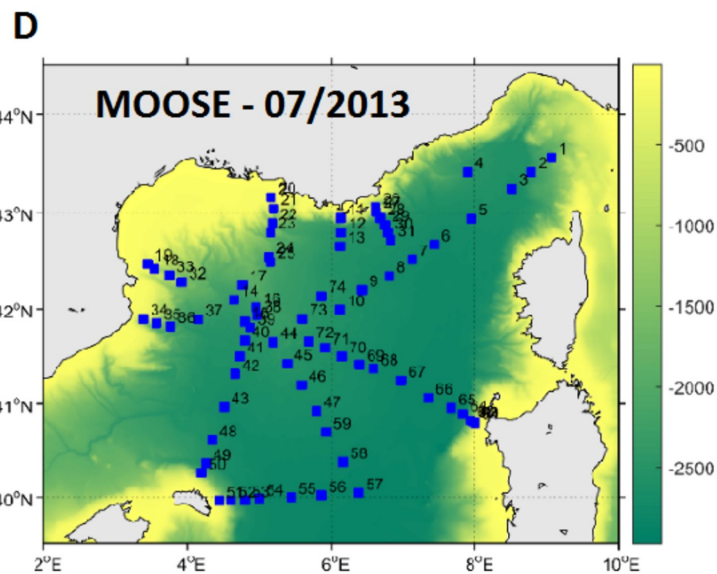
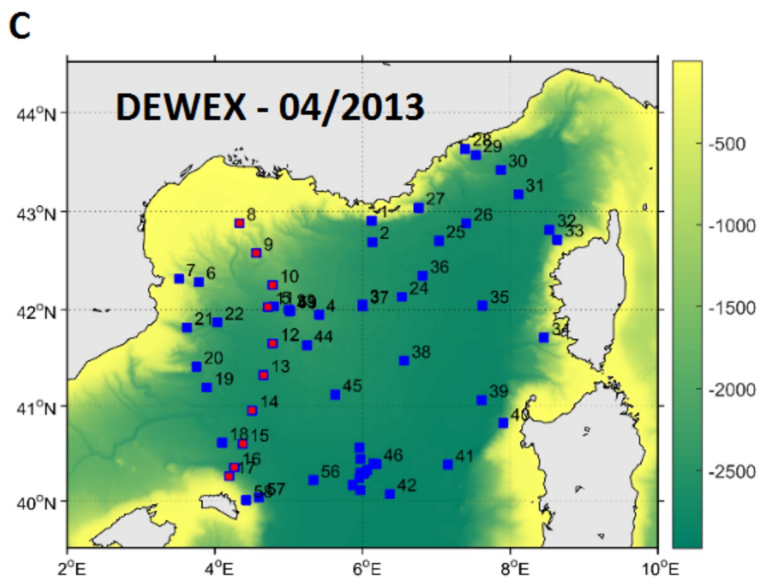
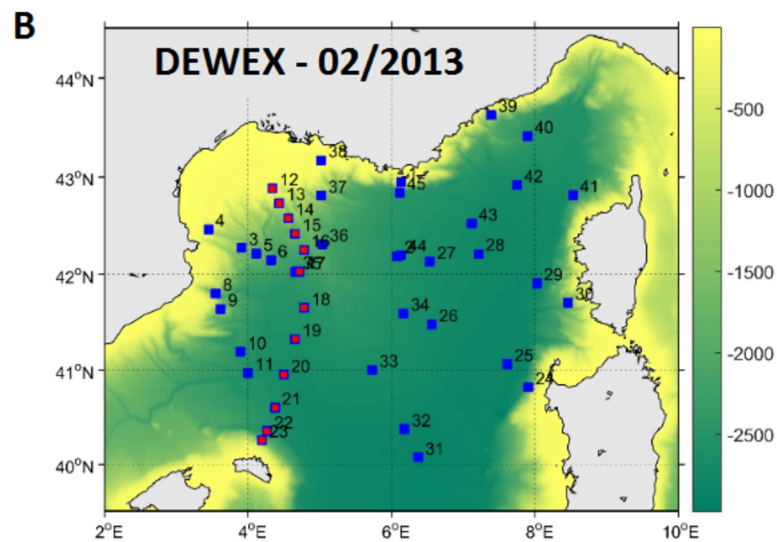
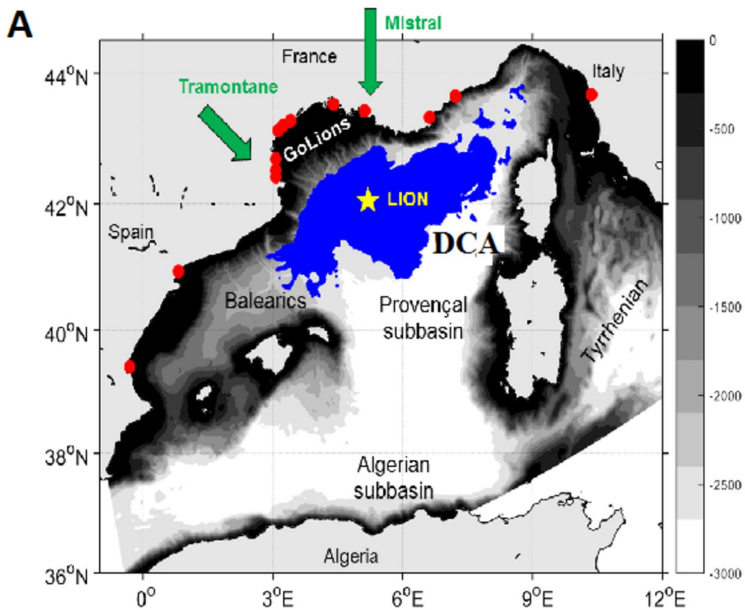


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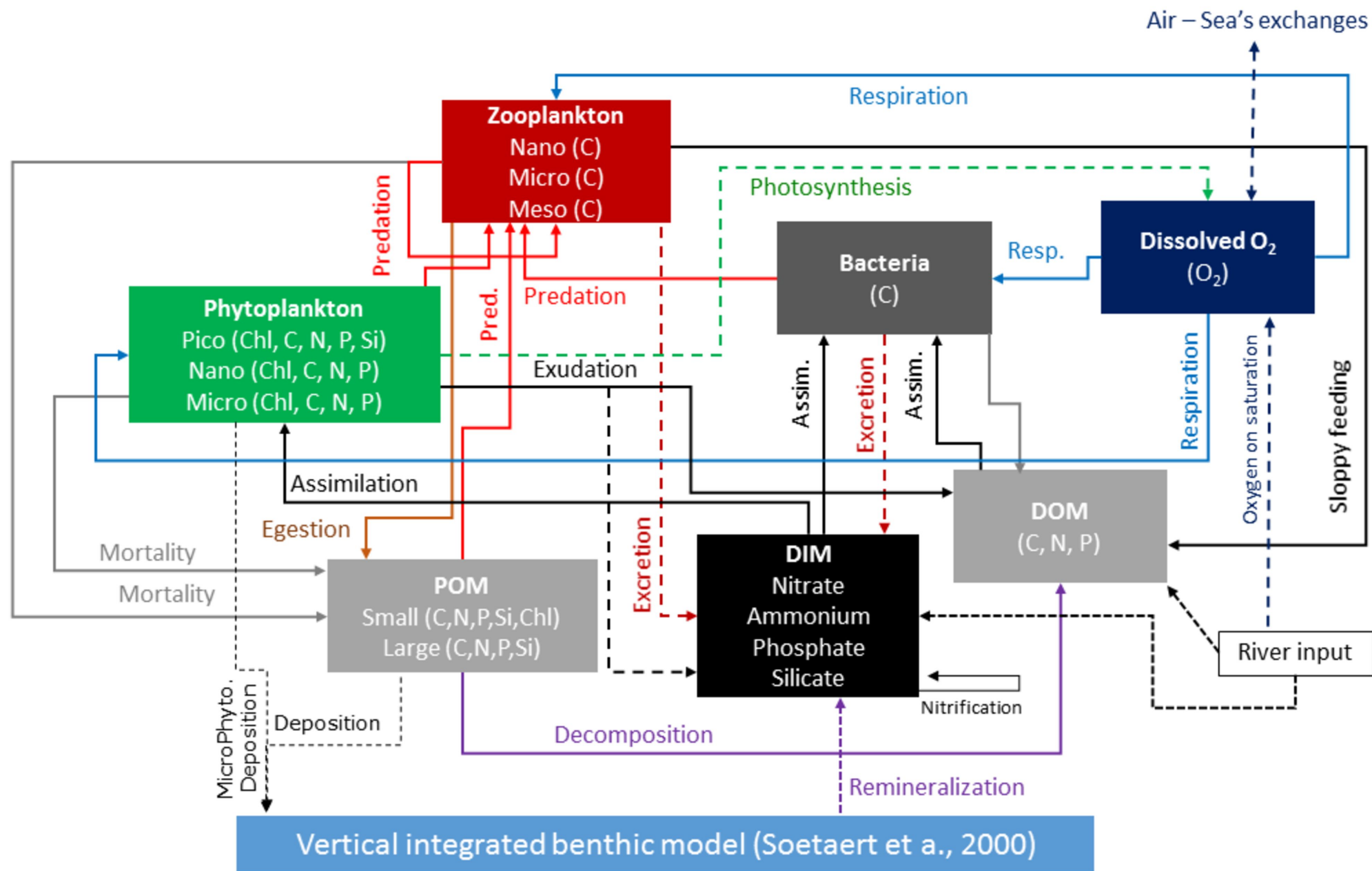


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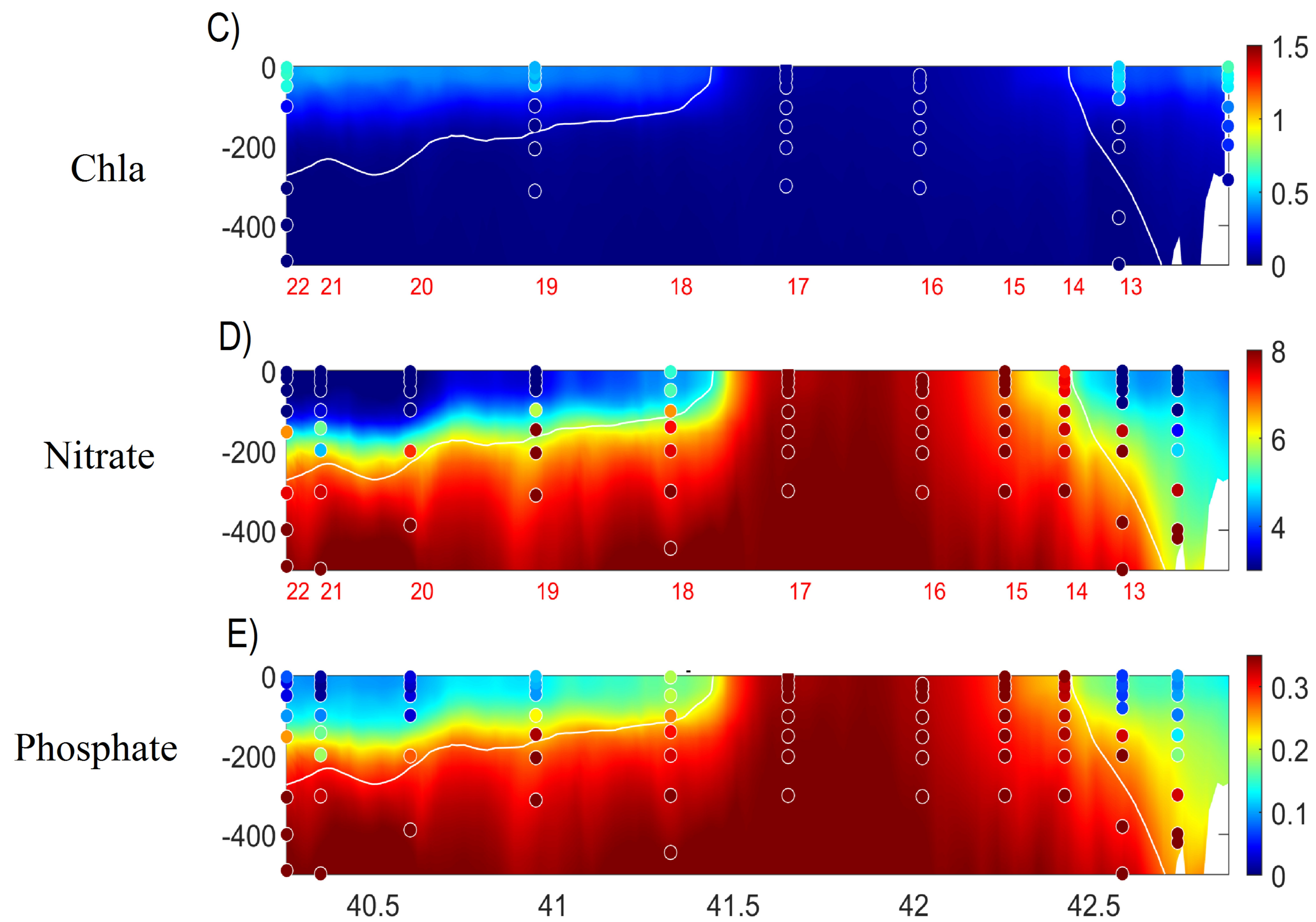
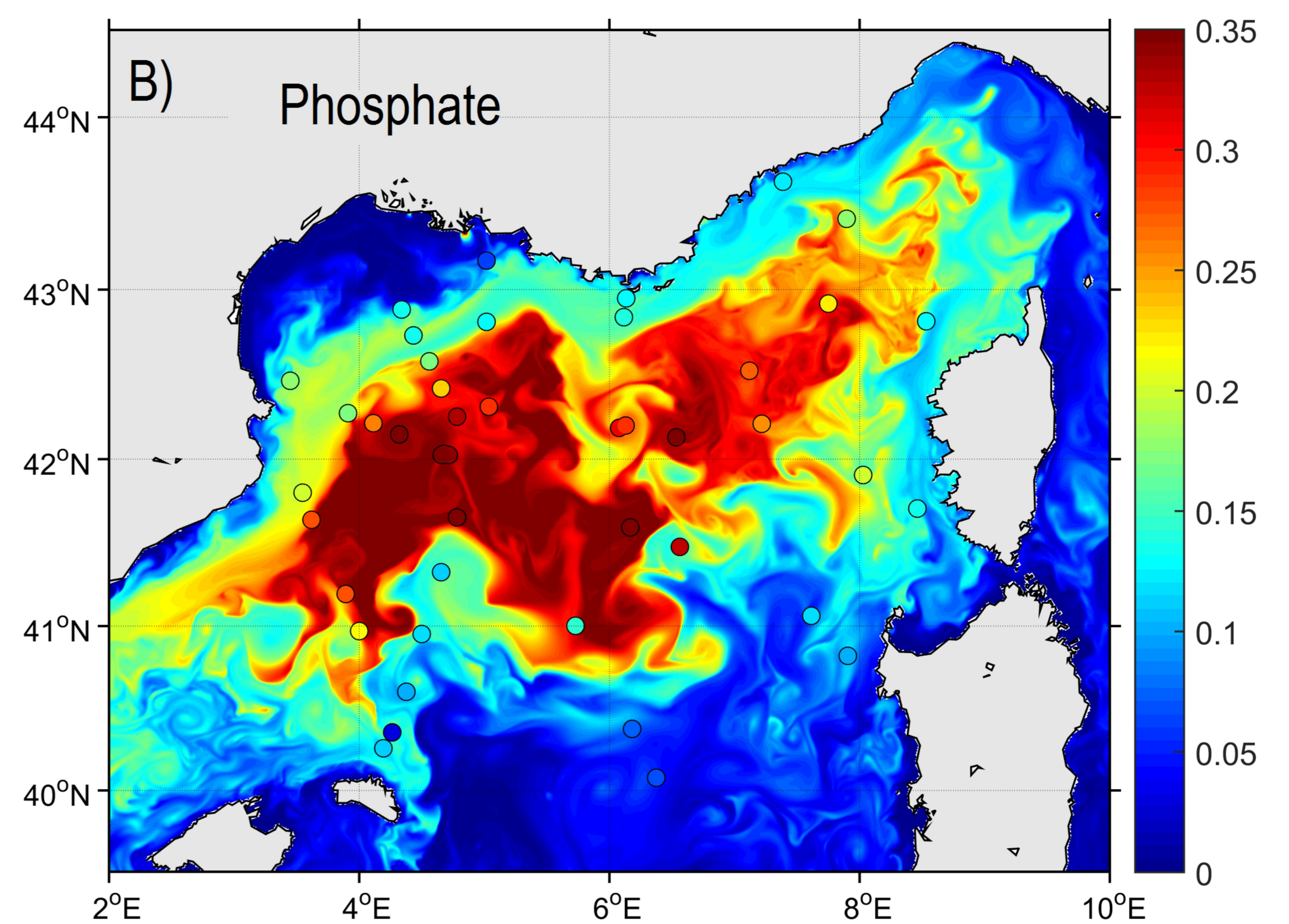
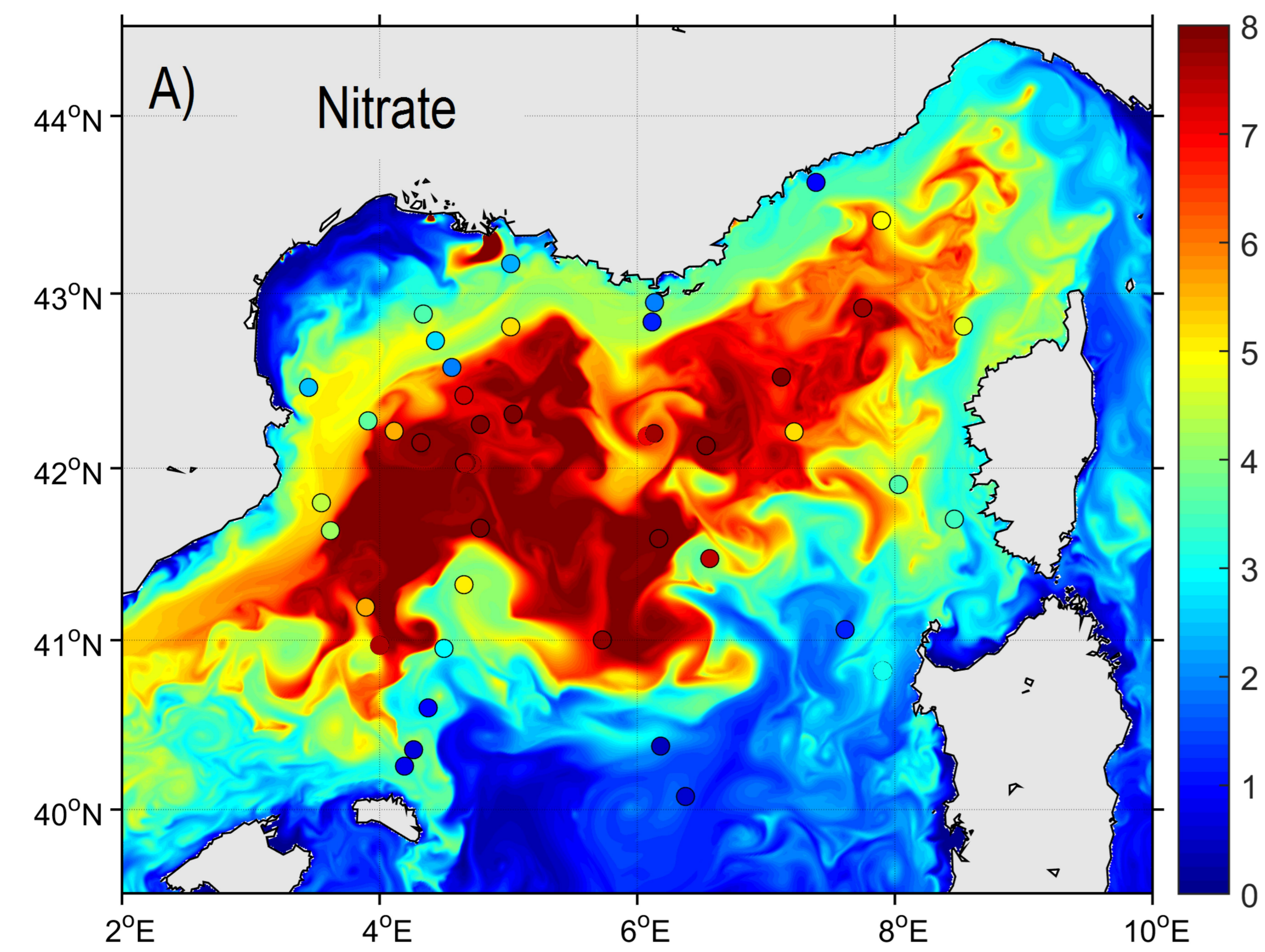


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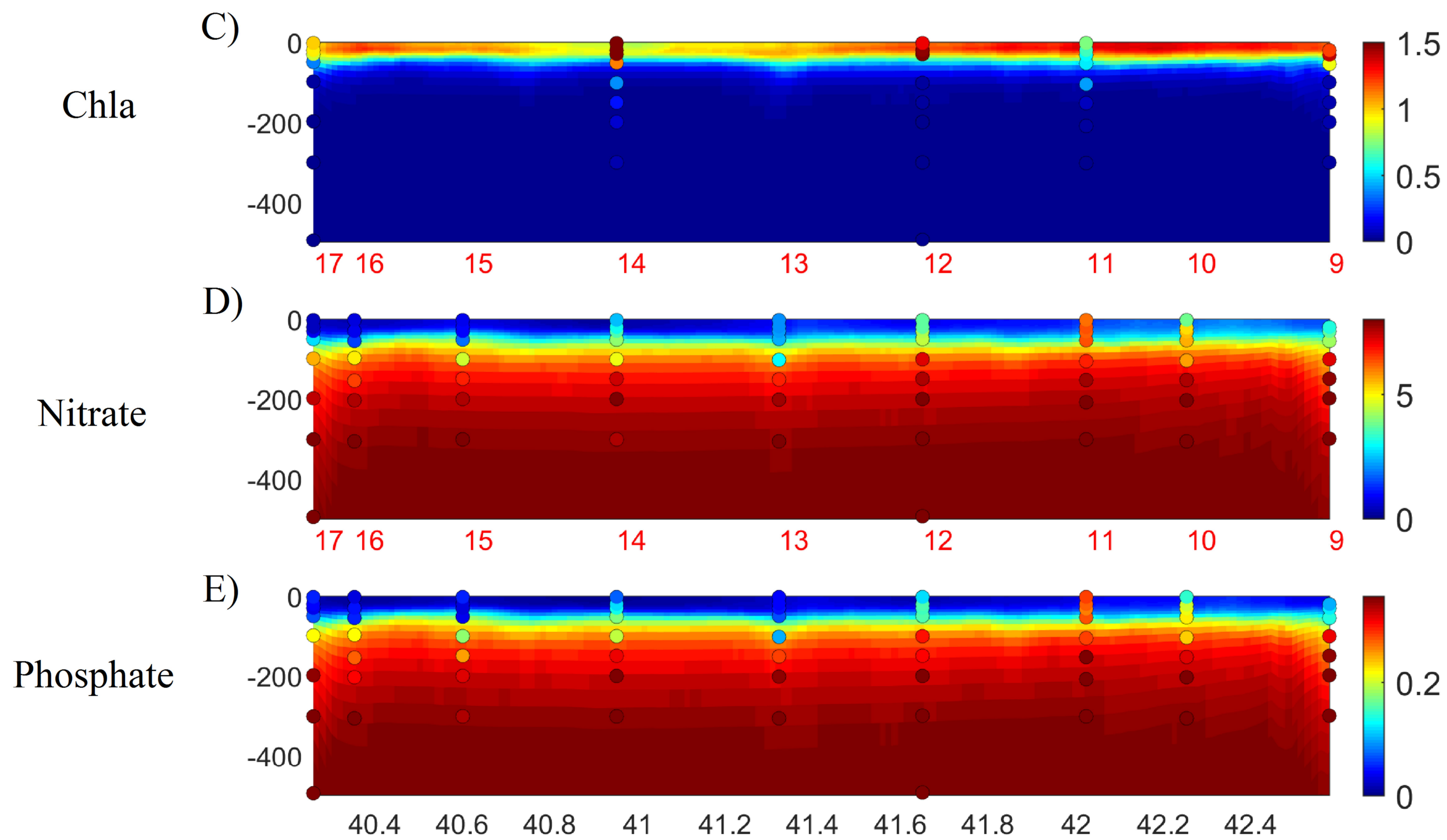
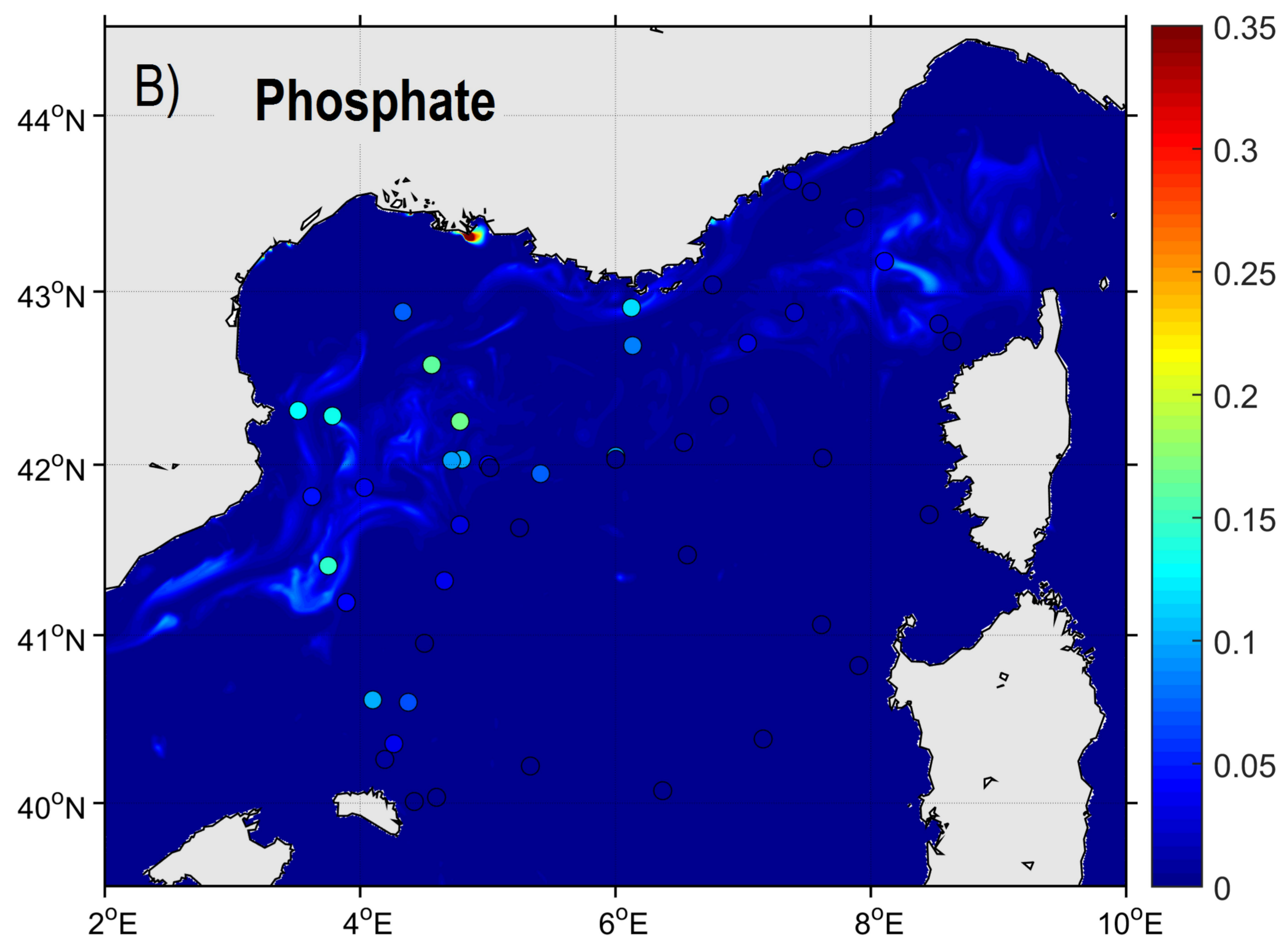
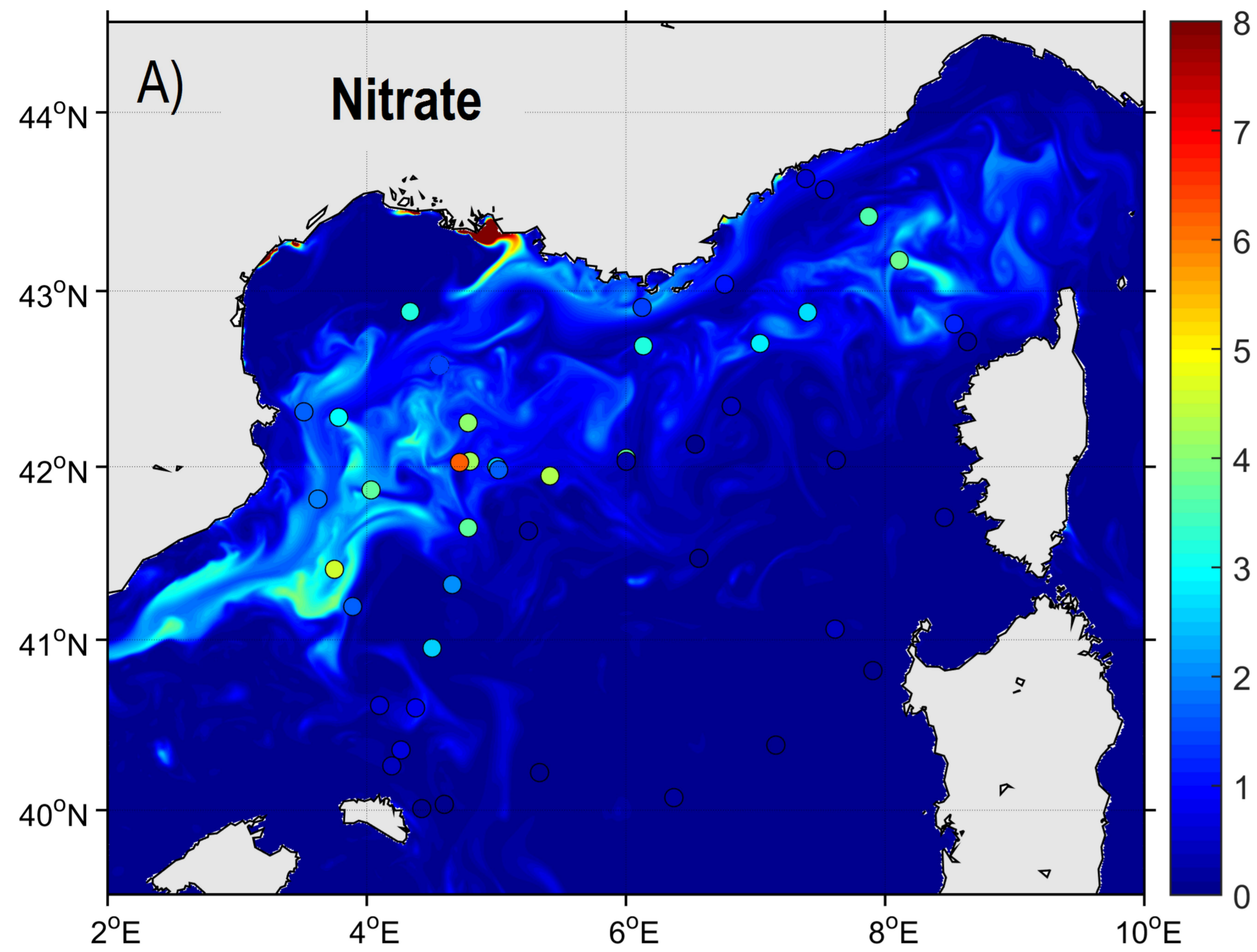


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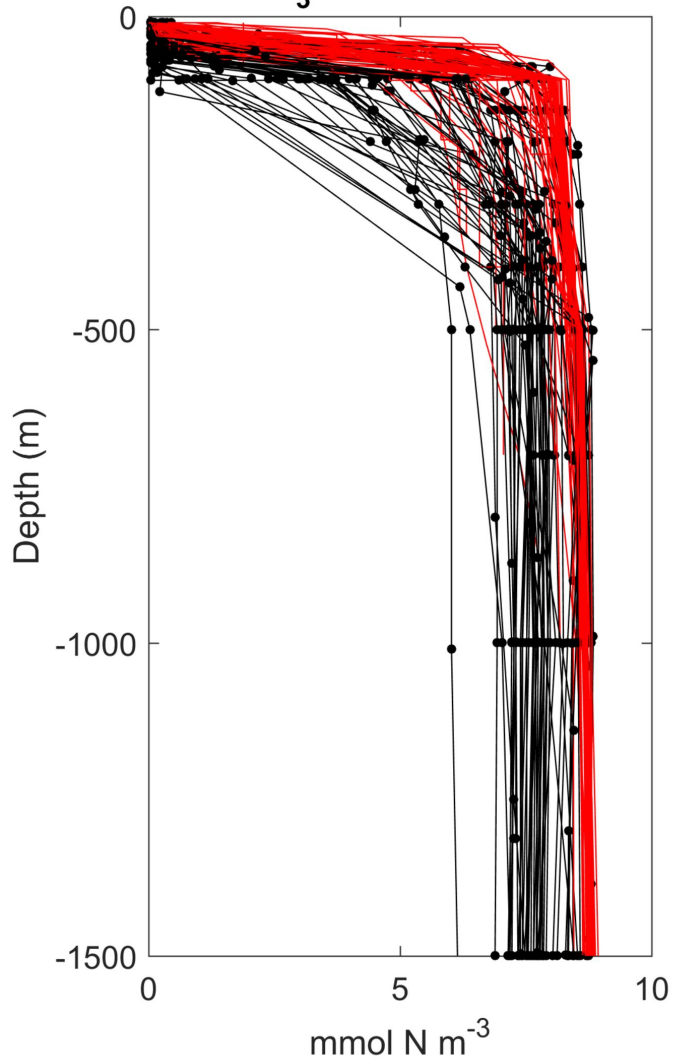
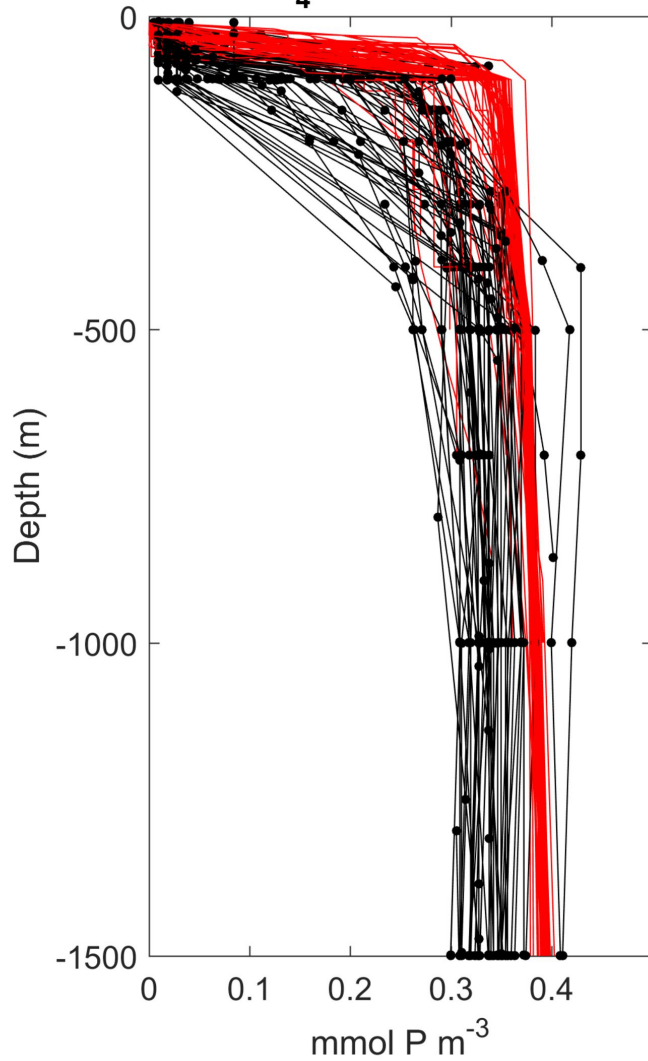
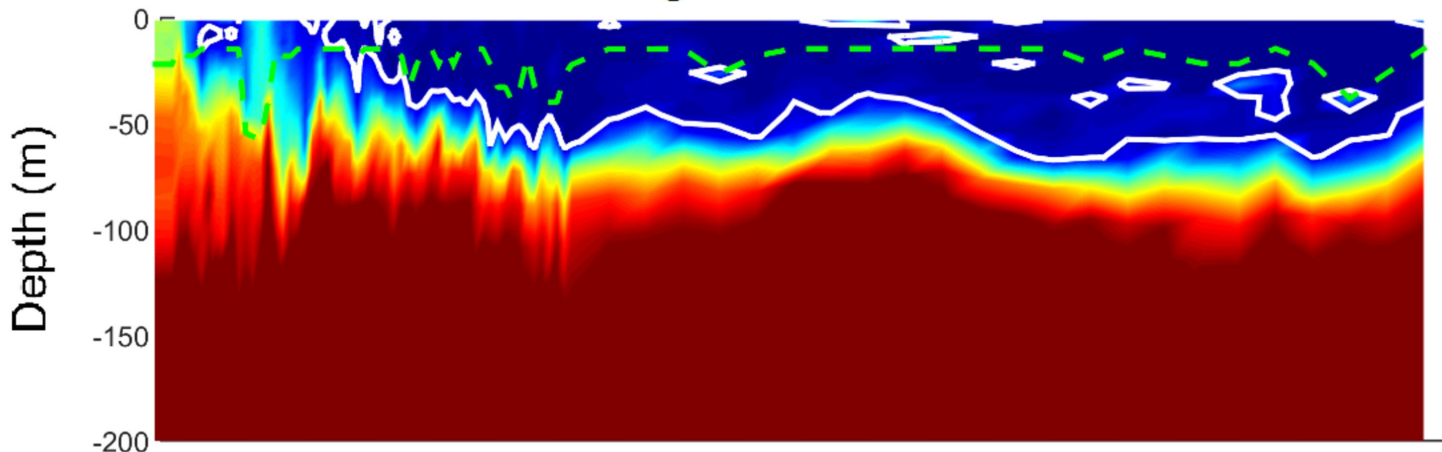
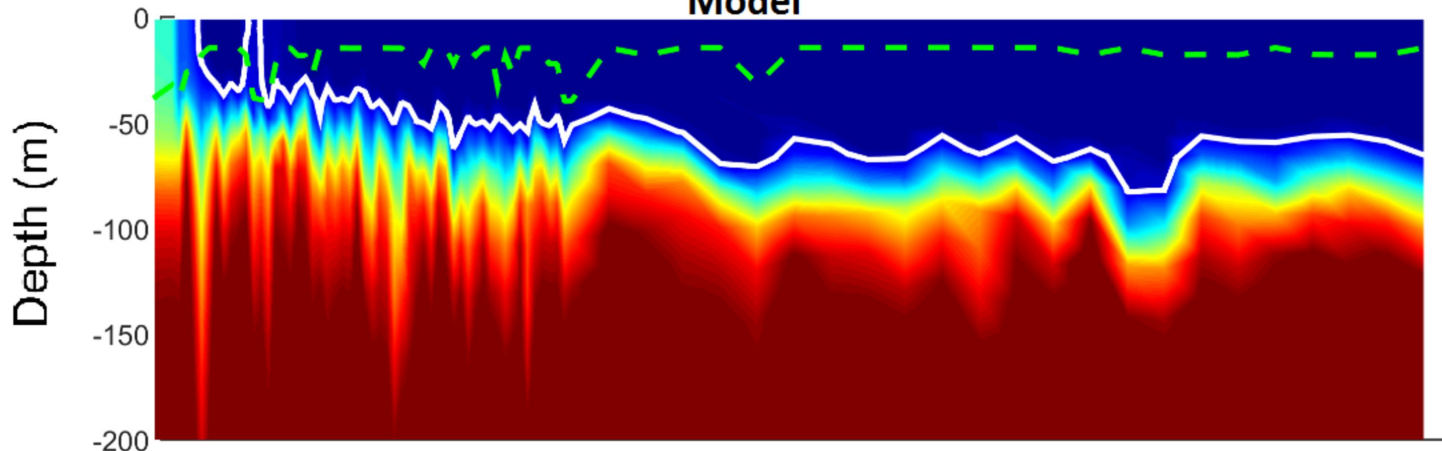
NO₃ Concentration**PO₄ Concentration**

Figure 6.

BioArgo Float "Lovbio17B"



Model



May

Jun

Jul

Aug

Sep

0

1

2

3

4

5

6

7

Nitrate concentration (mmol m^{-3})

Figure 7.

Heat Flux [W.m-2]

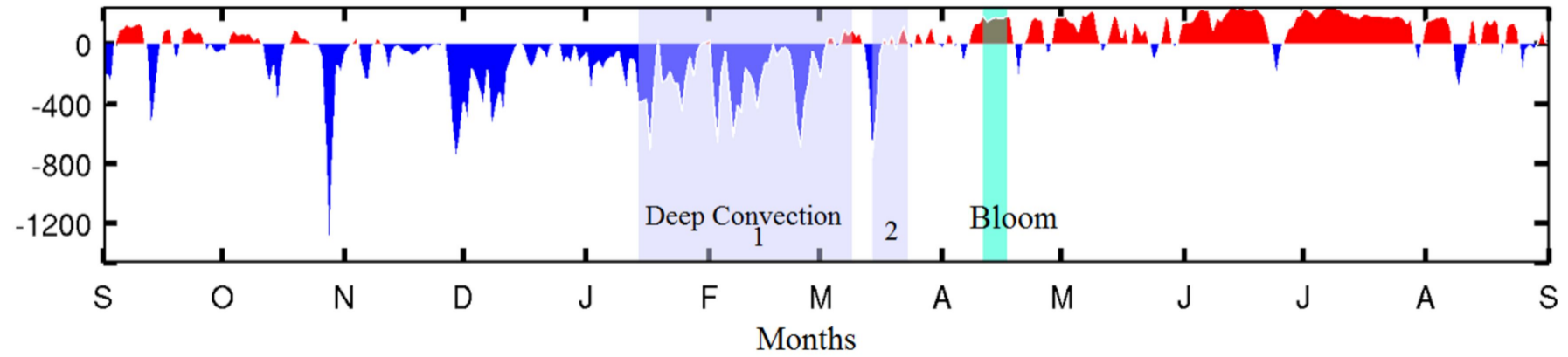


Figure 8.

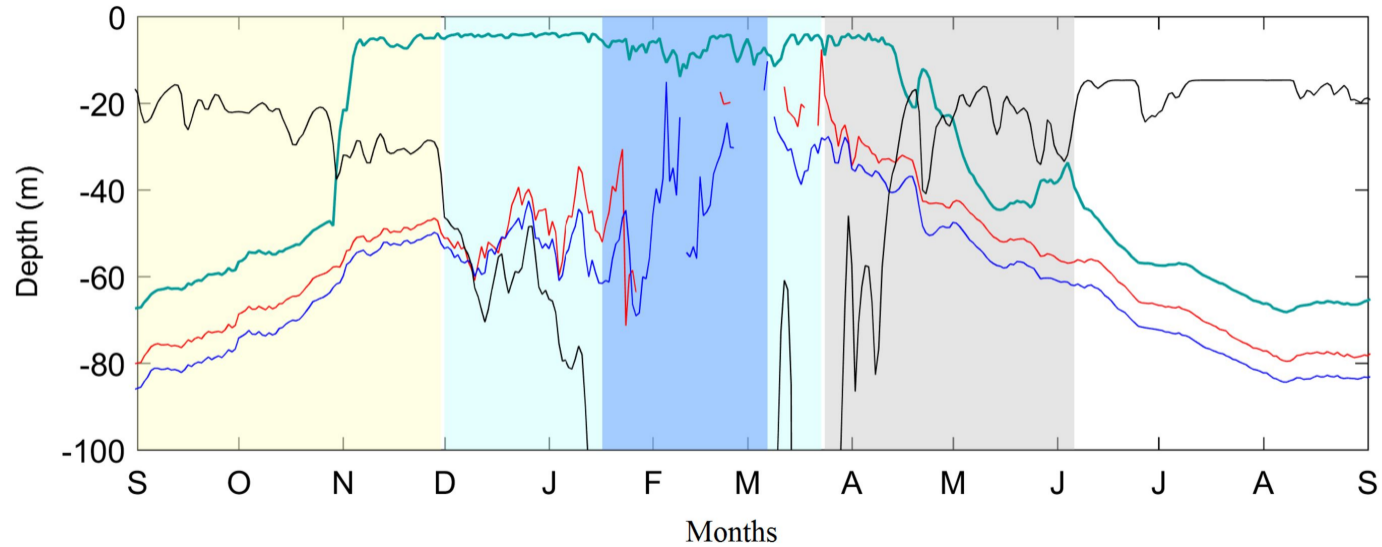
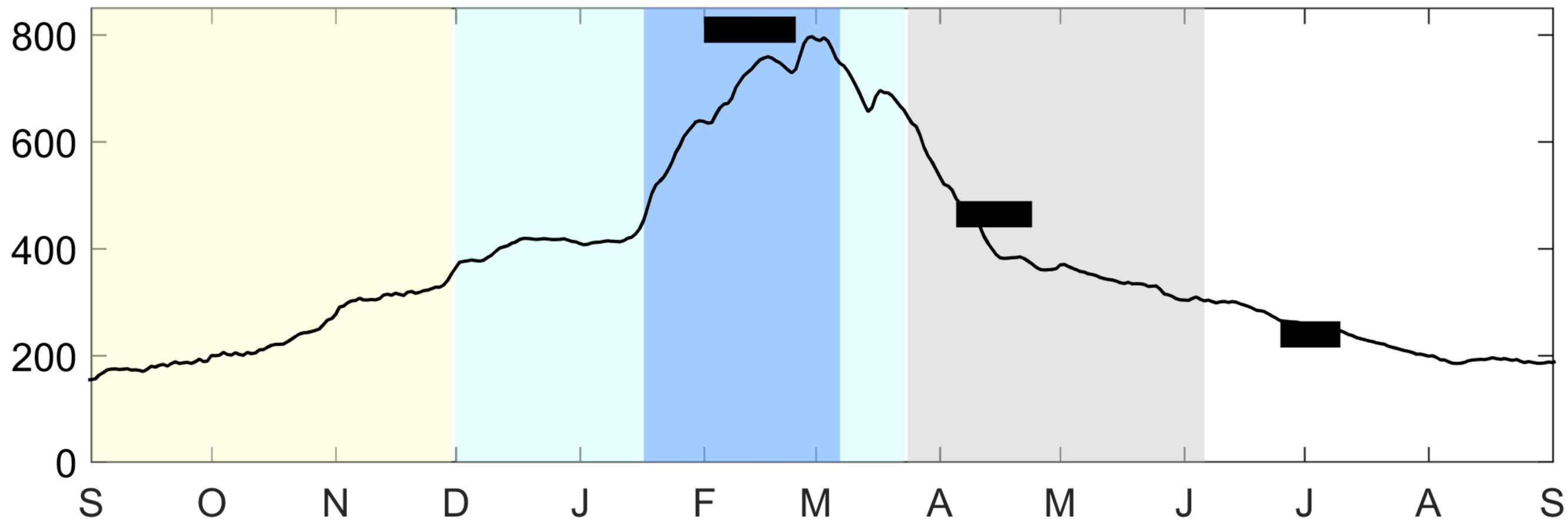


Figure 9.

NO₃ Stock (mmol)



PO₄ Stock (mmol)

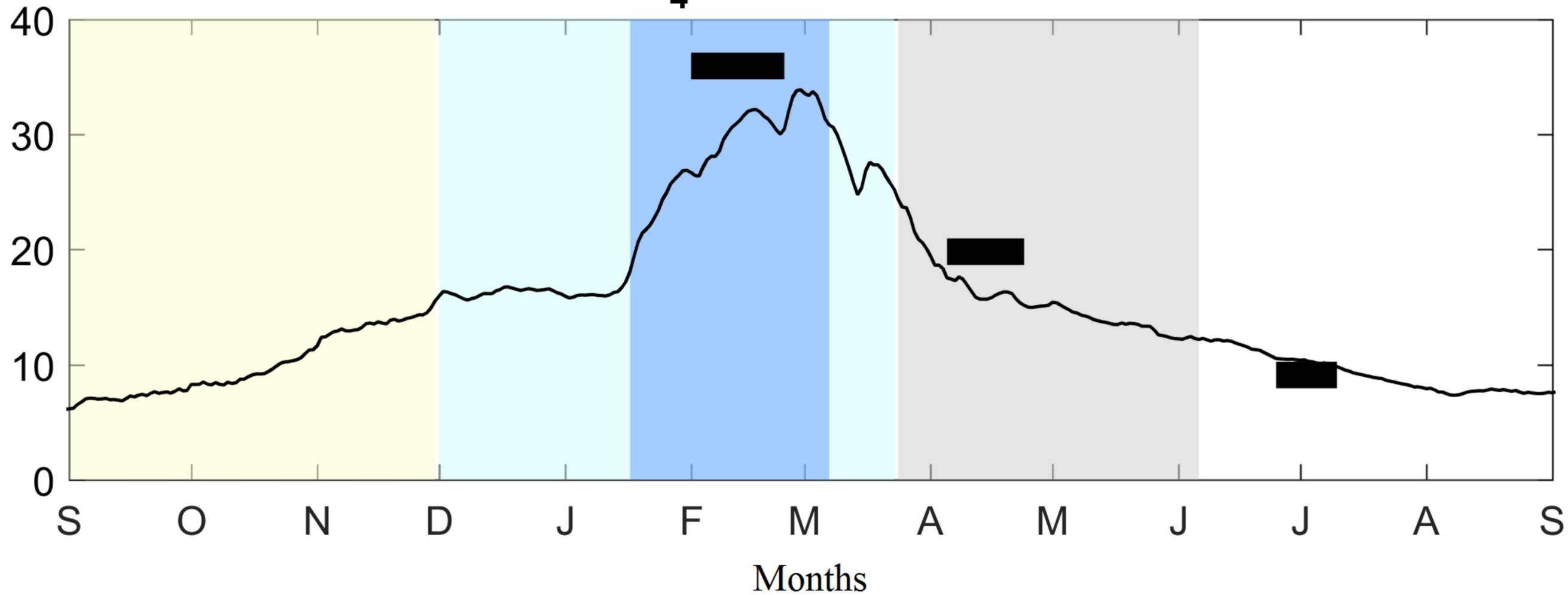


Figure 10.

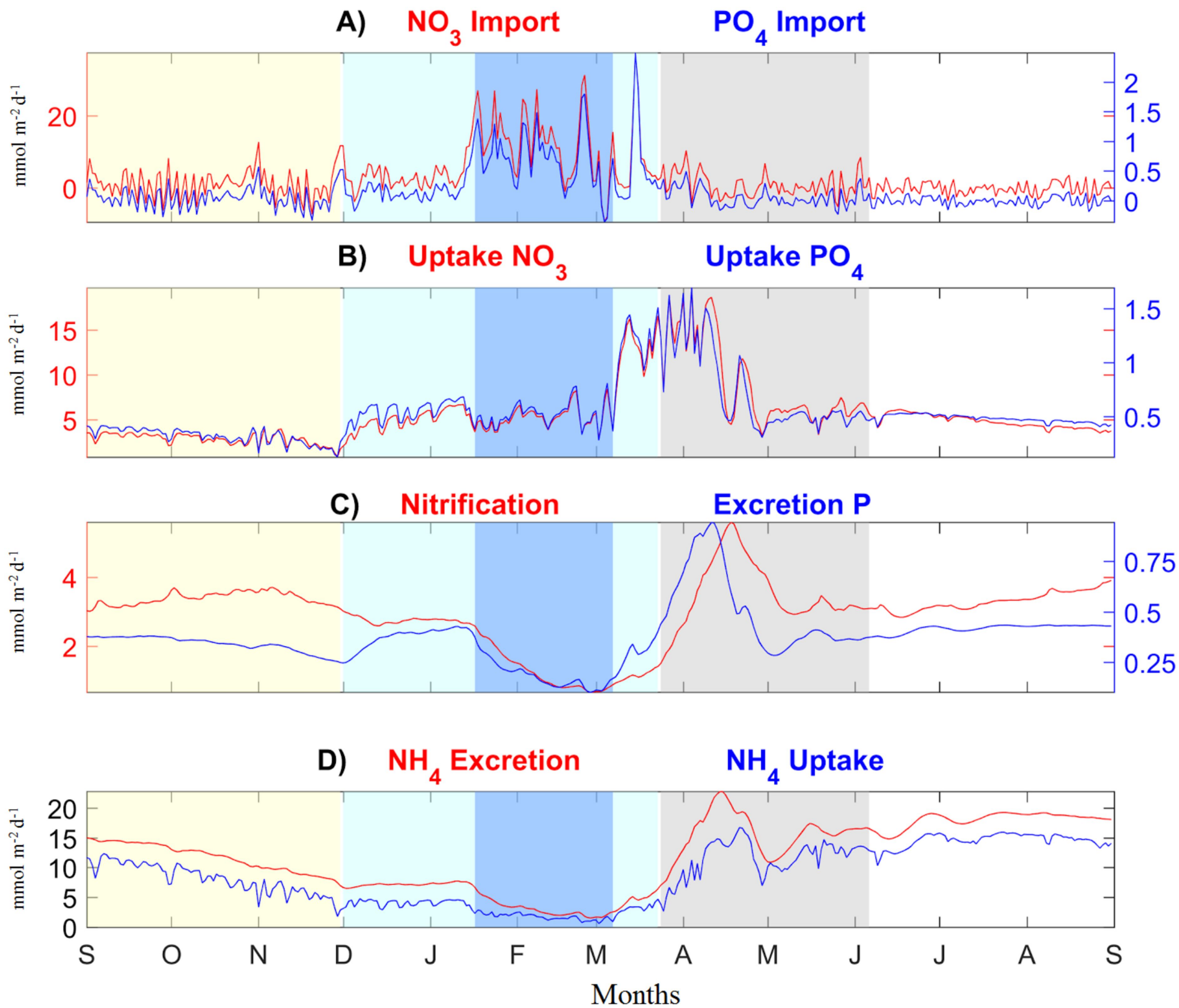
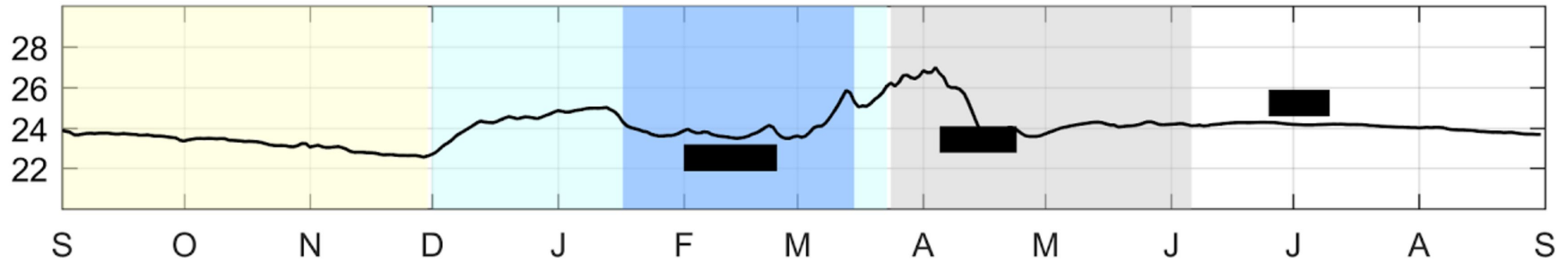
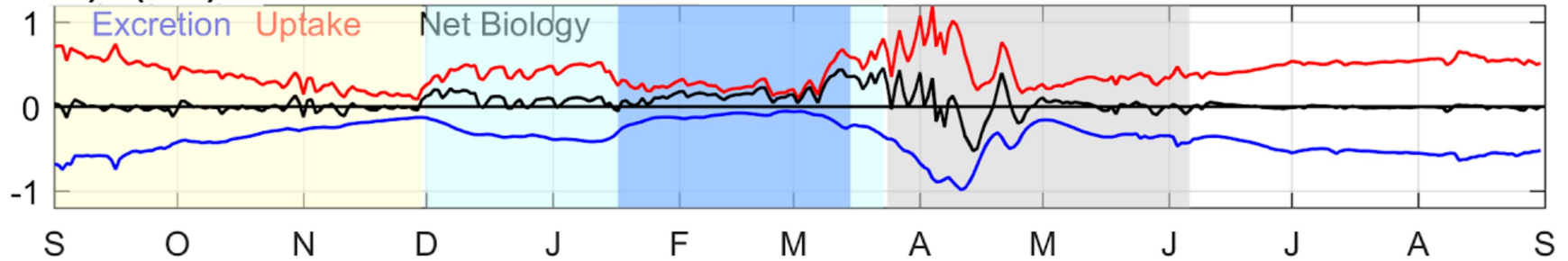


Figure 11.

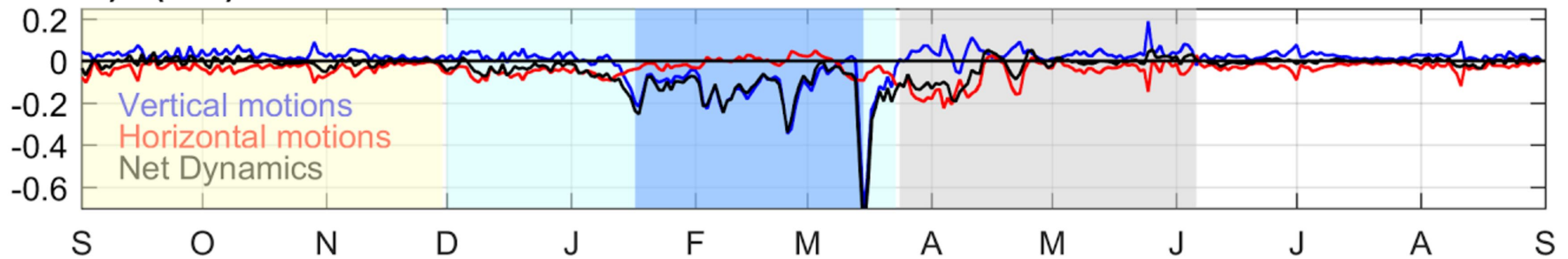
A) Ambient N/P



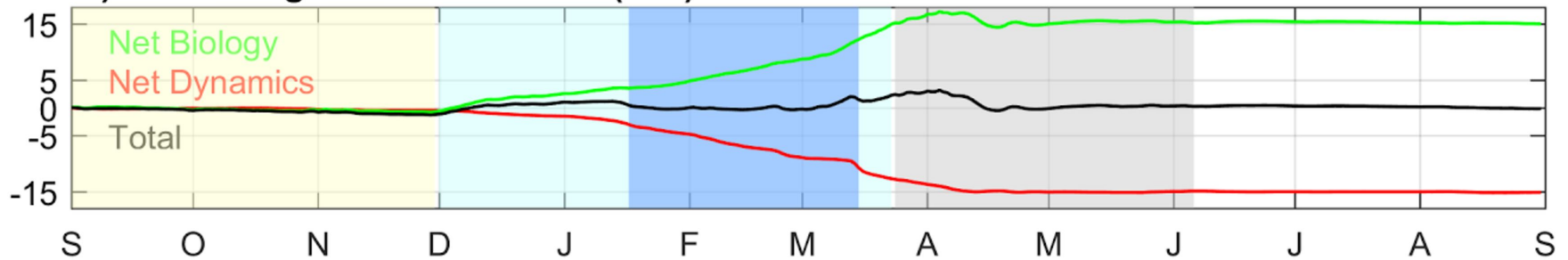
B) $d(N/P)$



C) $d(N/P)$



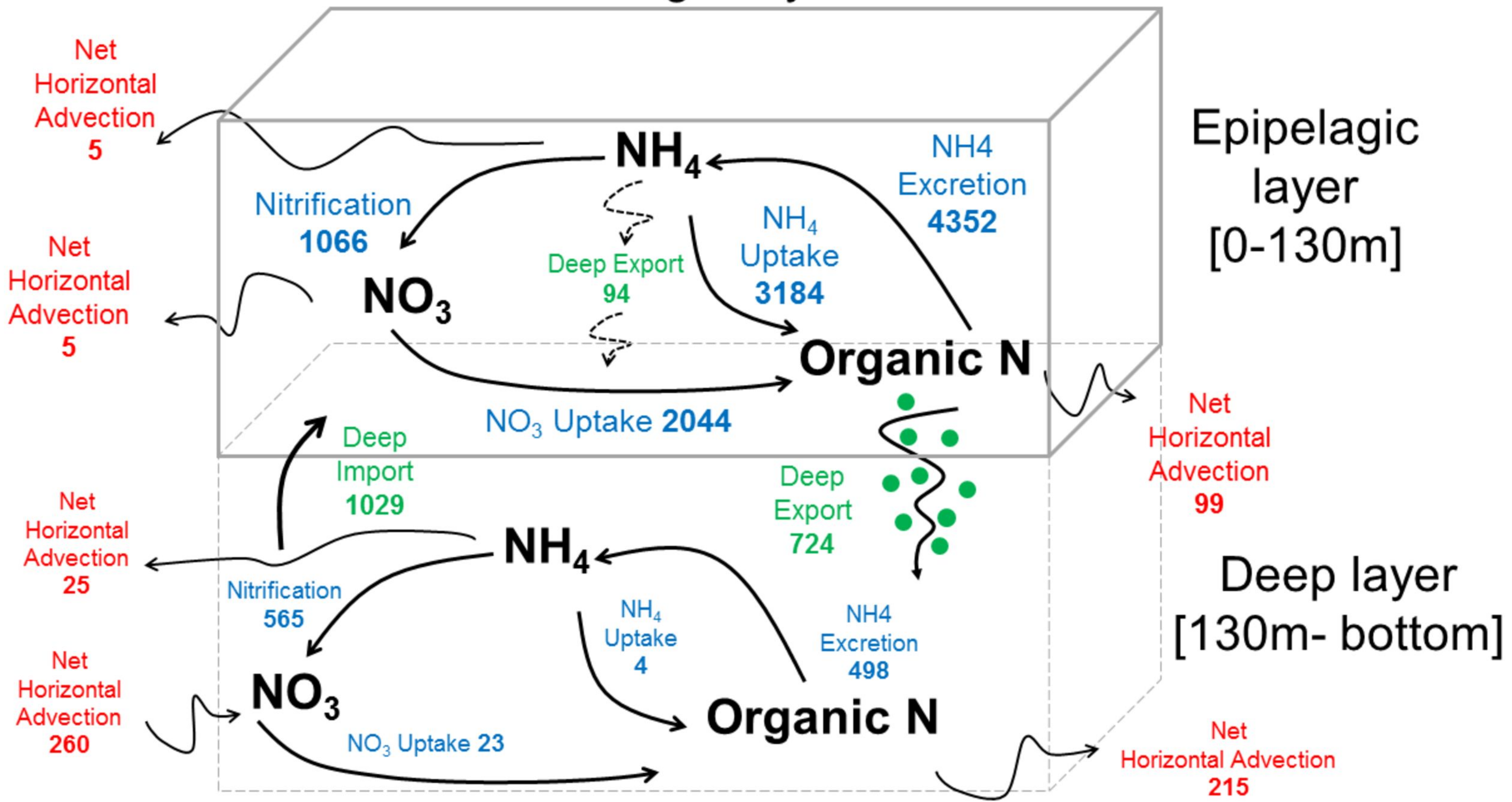
D) Time integrated variation of (N/P)



Months

Figure 12.

Nitrogen cycle



Phosphorus cycle

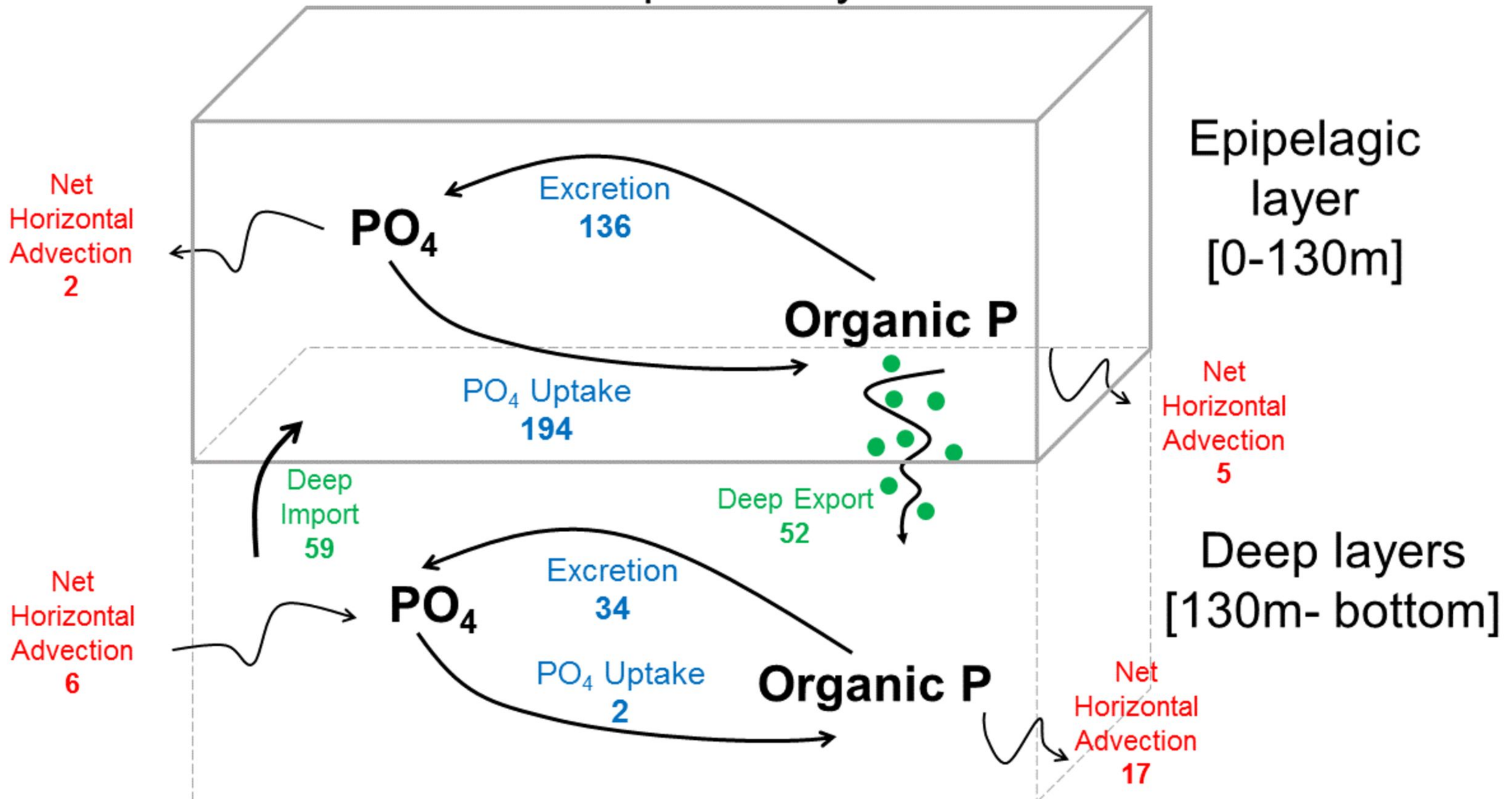
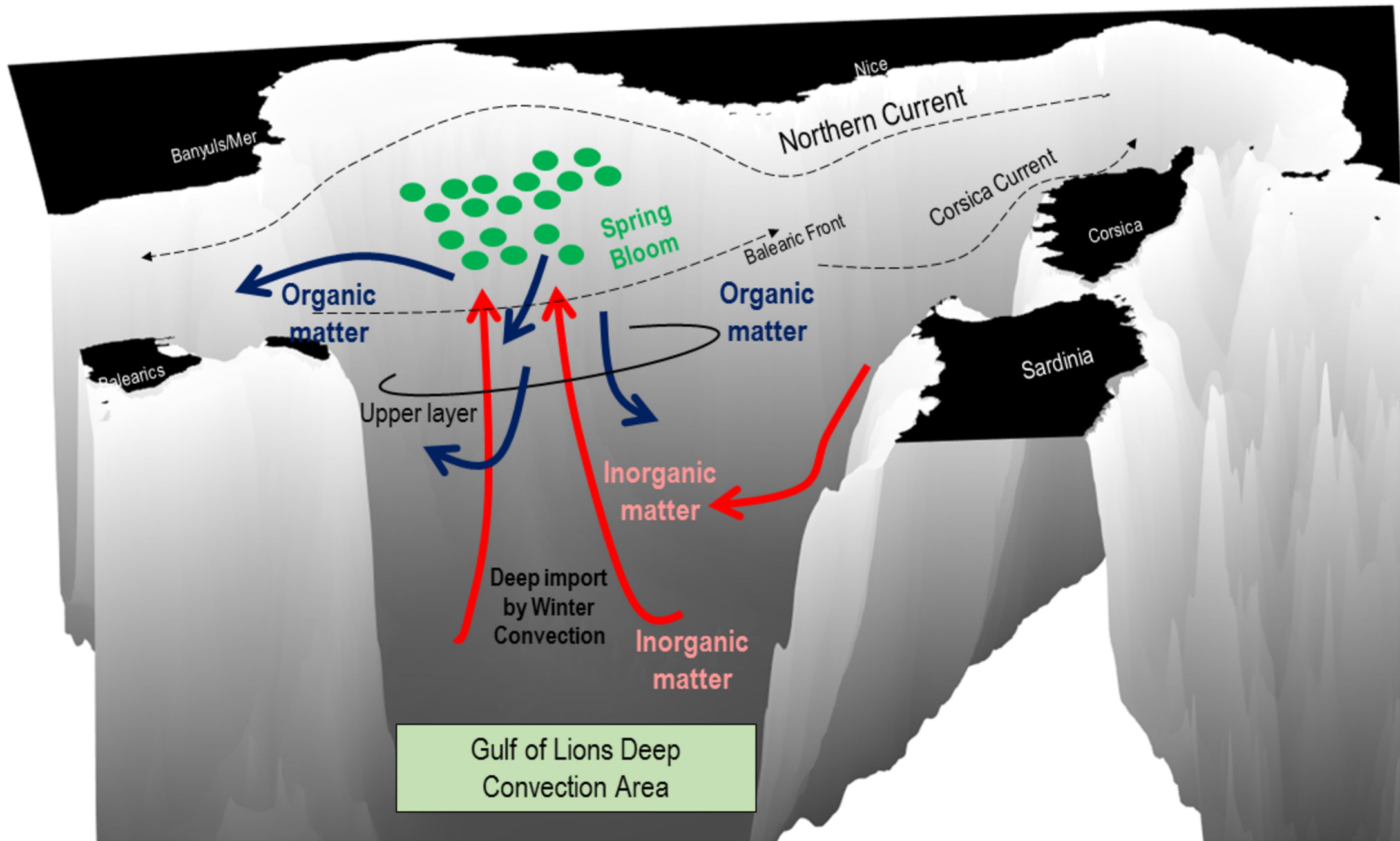


Figure 13.



Process	Unit	Reference	Period
Vertical import of inorganic nutrients	mmol m ² y ⁻¹		
		<i>Pujo-Pay and Conan</i> [2003]	1993
		<i>Severin et al.</i> [2014]	Feb./March 2011 convection events
		This study	2013 deep convection event
		This study	Sept. 2012-Sept. 2013
New production	mgC m ² y ⁻¹		
		<i>Severin et al.</i> [2014]	Feb./March 2011
		This study	Sept. 2012-Sept. 2013
Export of organic matter	mmol m ² y ⁻¹		
		<i>Pujo-Pay and Conan</i> [2003]	1993
		This study	Sept. 2012-Sept. 2013

values	
NO3	P04
433-650	-
760/800	36/36
674	40
1029	59
46/63	
	77
274-814	
	724

Sources in the surface layer	N	P
Atmospheric inputs in the western basin	10,042-72,825	357-697
Terrestrial inputs in the western basin	7,840-18,020	168-1,258
	6,064-73,621	379-957
Atmospheric inputs in the eastern basin	111,000	950
	11,592-77,519	248-1,541
Terrestrial inputs in the eastern basin	63,000	2,400
Vertical physical inputs in the NW Med deep convection area	63,474	3,640

Reference
Ribera d'Alcalà et al. [2003]
Ribera d'Alcalà et al. [2003]
Ribera d'Alcalà et al. [2003]
Krom et al. [2004]
Ribera d'Alcalà et al. [2003]
Krom et al. [2004]
This study