



HAL
open science

A canary in the Southern Ocean

Nicolas Metzl

► **To cite this version:**

Nicolas Metzl. A canary in the Southern Ocean. *Nature Climate Change*, 2019, 9 (9), pp.651-652.
10.1038/s41558-019-0562-1 . hal-02343935

HAL Id: hal-02343935

<https://hal.science/hal-02343935v1>

Submitted on 15 Nov 2019

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

OCEAN CARBON CYCLE

A canary in the Southern Ocean

The Southern Ocean is a major carbon sink, but knowledge of its variability is limited, especially in the coastal Antarctic. Now, results based on 25 years of observations in the West Antarctic Peninsula show that the carbon sink is increasing rapidly, driven by summertime biological production linked to sea ice dynamics.

Nicolas Metz

Two centuries ago, when Captain Nathaniel Brown Palmer and his crew sailed around the Antarctic Peninsula, the concentration of carbon dioxide (CO₂) in the atmosphere was around 280 ppm and the oceans were a net source of CO₂ (ref. ¹). Since then, anthropogenic CO₂ emissions have led to a marked increase of CO₂ in the atmosphere, which reached 410 ppm in 2019. Oceans have taken up a quarter of the anthropogenic CO₂ (ref. ²), slowing the rate of climate change. Without this ocean carbon sink, atmospheric concentrations would now be almost 500 ppm. However, it is uncertain how the marine CO₂ sink will evolve in the future, given the complex coupling between physical, chemical and biological processes not yet well observed or represented in models used to simulate future climate³. To improve the predictions of these models, it is important to obtain long-term, multidisciplinary observations of CO₂ dynamics in the Southern Ocean, where both warming and sea-ice changes alter the ocean stratification and the ecosystems that control the ocean's capacity to absorb carbon. Writing in *Nature Climate Change*, Michael Brown and colleagues⁴ present such observations and find a strong decline in surface water carbon.

The Southern Ocean and coastal Antarctic are particularly sensitive to climate changes that bring about changes in ocean circulation, mixing and phytoplankton productivity. All of these effects lead to variations in surface water CO₂ concentrations and associated air–sea CO₂ exchanges that need to be quantified before we can understand the link between environmental changes and the CO₂ uptake capacity of the ocean. Although regular observations in the Southern Ocean are still sparse, recent studies suggest decadal changes of the Southern Ocean carbon sink at the basin scale⁵. This is, however, not well evaluated in seasonal ice zone and coastal regions⁶, highlighting the need to investigate long-term observations at a regional scale.

Brown et al.⁴ investigate 25 years of data obtained in the West Antarctic Peninsula

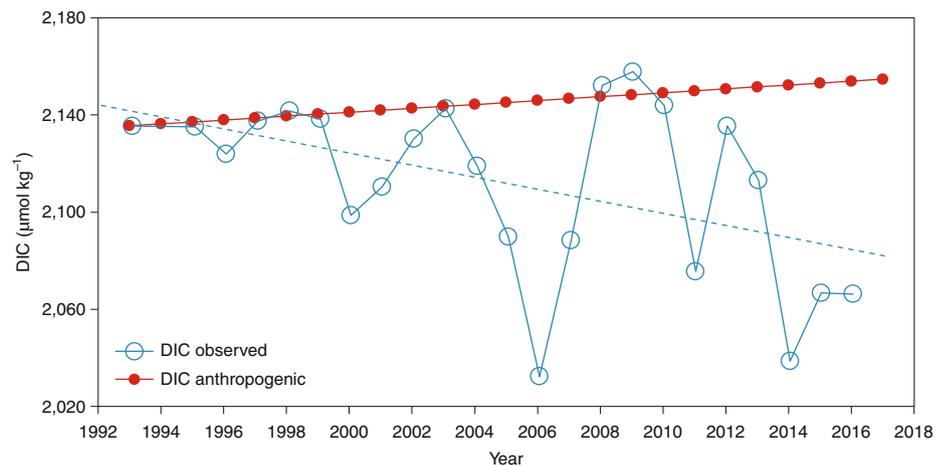


Fig. 1 | Mean dissolved inorganic carbon (DIC) concentration observed in surface waters in the West Antarctic Peninsula. Data for DIC concentrations (blue) are from observations⁴ each summer around 67°–68° S. The observed decrease is in contrast to the increase expected from anthropogenic CO₂ accumulation only (red symbols and line). The strong decrease of DIC is due to phytoplankton productivity leading to an enhanced carbon sink⁴.

in the framework of the Palmer Long-Term Ecological Research (LTER) programme coupled to the Drake Passage Time-series programme. The region has long been recognized as an important CO₂ sink in austral spring–summer, with surface values of CO₂ pressure (*p*CO₂) as low as 100–150 μatm (ref. ⁷). Brown et al. show that this uptake of CO₂ occurs every year from 1993 to 2017 and demonstrate a strong link between this CO₂ sink, ocean stratification and primary productivity. In the Southern Ocean and Antarctic coastal zones, offshore primary production by phytoplankton is not limited by nutrients (except iron). In spring–summer, when the ocean warms and sea ice melts, the ocean stratifies with a shallow mixed layer, so that photosynthesis in this mixed layer is not light-limited. The resulting increase in primary productivity reduces the concentrations of dissolved inorganic carbon (DIC) in surface waters through a process known as the biological carbon pump. This leads to oceanic *p*CO₂ well below atmospheric level, which means

that the ocean can absorb more CO₂. The carbon pump is strengthened when diatom species dominate phytoplankton communities⁴. Although the impact of primary production on the seasonality of the marine CO₂ sink has been recognized in Antarctic coastal zones^{7,8}, little is known about the temporal variability of this process.

Following recent work investigating long-term changes in the oceanic carbonate system in this region⁹, Brown et al. shed new light on the questions of how and why the ocean CO₂ sink changes over time. In contrast to previous work, they separate the northern and southern regions, which are influenced by different forcing, and find significant differences in the regional responses — linked to the southward extent of climate forcing along the West Antarctic Peninsula. In the northern region, there was a small decrease in DIC (about $-0.16 \mu\text{mol kg}^{-1} \text{yr}^{-1}$), similar to previous studies⁹. However, a much greater decrease in DIC — up to $-2 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ — was

identified in the southern region (Fig. 1), highlighting the contrasting response in the two subregions. This implies a strengthening of the carbon sink over more than two decades in the southern region during the austral summer.

Interestingly, the decrease in DIC is in contrast to the increase expected from anthropogenic CO₂ accumulation (+0.8 μmol kg⁻¹ yr⁻¹ in surface waters at these latitudes (Fig. 1)). It is worth recalling that DIC concentrations in surface waters range between 1,900 and 2,200 μmol kg⁻¹, so detecting a trend of around 1–2 μmol kg⁻¹ yr⁻¹ needs very accurate measurements. The data presented⁴ were qualified at international standards^{10,11} and leave no doubt that the summer ocean carbon sink has increased rapidly since 1993 in the investigated region, driven by an increase in phytoplankton production, particularly from diatoms, related to ocean stratification.

Surface pCO₂ generally increases at a rate of 1–2 μatm yr⁻¹ in most oceanic regions¹² and the observed long-term pCO₂ decrease of –2.8 μatm yr⁻¹ evaluated by Brown et al.

is the most rapid negative trend recorded in the ocean. It will be important to follow this signal in the next decade to confirm the reaction of the carbon sink in Antarctic coastal regions to future declines in sea ice and shifts in phytoplankton communities.

This new analysis⁴ is an important step forward in our understanding of the complex interplay of physical, biological and anthropogenic processes that govern ocean carbon uptake in a region particularly sensitive to changes in local or remote forcing. Such a long-term study also presents a challenge for sensitivity testing and validation of models that attempt to reproduce ocean biogeochemical cycles and predict future climate changes.

The data from the long-term observational project Palmer-LTER, including the presence of certain phytoplankton species, could now serve to detect the impact of climate change in the Southern Ocean, like the canaries used in the coal mines of previous centuries. It is crucial to maintain such time-series and promote observations in other Antarctic sectors to compare responses in different regions

where ice-melt or warming present different changes. Ship-based data should now be associated with new data from autonomous floats¹³ to better document the variations of the ocean carbon sink in all seasons. □

Nicolas Metzler

LOCEAN-IPSL (CNRS/SU/IRD/MNHN), Sorbonne Université, Paris, France.

e-mail: nicolas.metzler@locean-ipsl.upmc.fr

Published online: 26 August 2019

<https://doi.org/10.1038/s41558-019-0562-1>

References

1. Brovkin, V. et al. *Biogeosciences* **16**, 2543–2555 (2019).
2. Le Quéré, C. et al. *Earth Syst. Sci. Data* **10**, 2141–2194 (2018).
3. Pilcher, D. J. et al. *J. Geophys. Res. Oceans* **120**, 4625–4637 (2015).
4. Brown, M. S. et al. *Nat. Clim. Change* <https://doi.org/10.1038/s41558-019-0552-3> (2019).
5. Landschützer, P. et al. *Science* **349**, 1221 (2015).
6. Laruelle, G. G. et al. *Nat. Commun.* **9**, 454 (2018).
7. Karl, D. M. et al. *Deep Sea Res. A* **38**, 1097–1126 (1991).
8. Legge, O. J. et al. *Deep Sea Res. II* **139**, 167–180 (2017).
9. Hauri, C. et al. *Biogeosciences* **12**, 6761–6779 (2015).
10. Dickson, A. G., Sabine, C. L. & Christian, J. R. *Guide to Best Practices for Ocean CO₂ Measurements* (PICES Special Publication 3, 2007).
11. Bakker, D. C. E. et al. *Earth Syst. Sci. Data* **8**, 383–413 (2016).
12. Takahashi, T. et al. *Deep-Sea Res II* **56**, 554–577 (2009).
13. Gray, A. et al. *Geophys. Res. Lett.* **45**, 9049–9057 (2018).