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The Reaction of Peroxy Radicals with OH Radicals

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Abstract

The reaction of peroxy radicals, RO₂, with OH radicals has long been ignored to play any role in atmospheric chemistry. Recent experimental and modeling studies show however that these reactions are fast and can play a role in remote atmospheres with low NO concentrations when the major fate of peroxy radicals, its reaction with NO, slows down. The present article summarizes recent work on this class of reactions and its implications in different environments.

Introduction

Peroxy radicals, RO_2 , are key species in atmospheric and low temperature combustion chemistry. They are reactive intermediates formed during the oxidation of most hydrocarbons:

$$\mathbf{RH} + \mathbf{OH} \rightarrow \mathbf{R} + \mathbf{H}_2\mathbf{O} \tag{R1}$$

$$\mathbf{R} + \mathbf{O}_2 \to \mathbf{R}\mathbf{O}_2 \tag{R2}$$

In the atmosphere, their subsequent reaction depends on the environment. A detailed review on their chemistry has been given by G. Tyndall and collegues [1,[2]. Briefly, in polluted environments, *i.e.* in the presence of high NO_x concentrations, peroxy radicals will rapidly react with NO to form NO_2 and alkoxy radicals, RO.

$$RO_2 + NO \rightarrow RO + NO_2$$
 (R3)

The NO₂ will be photolysed and leads to an increase in O₃ concentration.

$$NO_2 + h\nu \rightarrow NO + O(^3p) \xrightarrow{O_2} O_3$$
 (R4)

The alkoxy radicals react further with O_2 , leading to a carbonyl species and HO_2 , which oxidizes another NO to NO_2 and finally leads to formation of a second O_3 molecule:

$$RO + O_2 \rightarrow R_{-H}O + HO_2 \tag{R5}$$

$$HO_2 + NO \rightarrow OH + NO_2$$
 (R6)

With decreasing NO concentration, other reaction pathways become competitive for peroxy radicals. Until recently, only the reactions with HO₂ or the self- and cross reactions with other peroxy radicals, leading mostly to stable products, were considered as fate for RO₂ radicals under low NO conditions in atmospheric chemistry models such as MCM [3,[4,[5]. For the simplest RO₂, CH_3O_2 , this leads to:

$$CH_3O_2 + HO_2 \rightarrow CH_3OOH + O_2 \tag{R7}$$

$$CH_3O_2 + CH_3O_2 \rightarrow 2 CH_3O + O_2 \xrightarrow{O_2} CH_2O + HO_2$$
 (R8a)

$$CH_3O_2 + CH_3O_2 \rightarrow CH_3OH + CH_2O + O_2$$
(R8b)

In 2009, Archibald *et al.* [6] proposed for the first time an alternative reaction pathway for RO_2 in low NO environments: the reaction of RO_2 radicals with OH radicals. In a pure modeling study they investigated the possible impact of this class of reaction on the composition of the atmosphere. They simulated the impact of such a reaction in the Marine Boundary Layer (MBL) by adding to their mechanism the reaction of OH with peroxy radicals up to C_4 and compare the concentrations with the

base case, *i.e.* without $RO_2 + OH$. Three different possible reaction paths were simulated, leading for the example of the most simple peroxy radical, CH_3O_2 , to the following products:

$OH + CH_3O_2 \rightarrow CH_2O_2 + H_2O$	(R9a)
$OH + CH_3O_2 \rightarrow CH_3O + HO_2$	(R9b)
$OH + CH_3O_2 \rightarrow CH_3OH + O_2$	(R9c)

They run the three different product scenarios by varying the rate constants for all RO_2 + OH reactions between 0.5, 1.0 and 1.5×10^{-10} cm³s⁻¹. For all scenarios they found only a small, negligible effect on the mixing ratios of O_3 , NO_x , OH, and other trace gas species in the marine boundary layer. However, a substantial increase in the mixing ratios of HCOOH was observed for all simulated rate constant scenarios, if the reaction pathway would be formation of the Criegee radical (R9a). A strong increase in the mixing ratio of CH₃OH was observed if the major pathway would be (R9c). The impact on the RO₂ and HO₂ radical budget was very minor for all scenarios.

Even though this study pointed towards a possible, non-negligible impact of this class of reaction on the composition of remote atmospheres, no further attention was paid to this reaction in the following years and it was only in 2014 that Bossolasco *et al.* [7] measured for the first time the rate constant of a reaction between a peroxy radical and an OH radical. The rate constant turned out to be fast enough to make this reaction non-negligible under remote environments and since then, several aspects of this class of reaction have been studied experimentally and theoretically by different research groups.

In the following article, a résumé of the findings on the reaction of RO_2 radicals with OH radicals and the impact of this class of reactions on the composition of the atmosphere in different environments will be summarized. The article is separated in two parts: the reaction of the most simple peroxy radicals, CH_3O_2 , with OH radicals will be discussed separately in the first part, and the reaction of larger RO_2 radicals with OH will be discussed in the second part. Each section will first discuss the current knowledge on the rate constants and in a second subsection summarizes what is known about the product yields and what impact this reaction might have on the composition of the troposphere.

The reaction of $CH_3O_2 + OH$

Rate constant

The first measurement of a rate constant for the reaction of a peroxy radical with OH radicals was reported in 2014 by Bossolasco *et al.* [7]. Using laser photolysis coupled to cw-cavity ring down spectroscopy for a quantification of CH_3O_2 [8] and high repetition rate LIF for a time resolved detection of OH radicals [9] they investigated the reaction of the simplest peroxy radical

$$CH_3O_2 + OH \rightarrow \text{ products}$$
 (R9)

The schematic view of the experimental set-up used for most of the studies so fare published on RO_2 + OH reactions is shown Figure 1.



Figure 1: Schematic view of experimental set-up: excimer laser generates simultaneously OH and RO_2 radicals from appropriate precursors, 2 symmetric cw-CRDS paths for detection of absolute, time resolved traces of HO_2 , CH_3O_2 or OH cross the photolysis pulse in a small angle, leading to an overlap of 37 cm between photolysis and detection beam, high repetition rate laser induced fluorescence for detection of relative OH concentration time profiles is installed perpendicular to the photolysis beam. APD: Avalanche Photo Diode, AOM: Acousto-Optic Modulator, M: Mirror, L: Lens, PM: photomultiplier. Separate, identical trigger circuits and data acquisition systems are used for both cw-CRDS systems, but are not shown in the figure for clarity.

The 248nm photolysis of CH₃I in the presence of O₂ was used as a precursor for CH₃O₂ radicals, while OH radicals were simultaneously generated by photolysis of H_2O_2 [10]. Using this method, they found a very fast rate constants of $k_9 = (2.8 \pm 1.4) \times 10^{-10} \text{ cm}^3 \text{s}^{-1}$. With such a fast rate constant, this reaction could indeed become competitive to the other reaction paths. CH_3O_2 is an important radical, particularly over the tropical oceans where, in absence of other hydrocarbons emitted by flora or human activity, OH radicals process a large fraction of the global CH₄ emission. Neglecting an important reaction path for this radical might therefore lead to a bias in HO_x concentration, CH_4 lifetime or O₃ formation and destruction rate. Fittschen et al. [11] have integrated (R9) into the detailed MCM model [4] and have determined the importance of (R9) as a sink of CH₃O₂ radicals in remote marine environments. Conditions such as found during a field campaign at Cape Verde Island [12] in 2007, representative for the tropical remote ocean, were simulated using the rate constant from Bossolasco *et al.* [7] and it has been shown that up to 30% of all CH_3O_2 radicals would decay through (R9). Following the work of Bossolasco et al., Yan et al. [13] have re-measured the rate constant of the reaction of OH radicals with CH_3O_2 using laser photolysis coupled to UV-absorption spectroscopy, radicals were generated by simultaneous 193nm photolysis of CH₃COCH₃ as precursor for CH₃O₂ radicals and N₂O/H₂O as precursor for OH. The absorption-time profiles were fitted to a complex mechanism containing 46 reactions, with many of them being radical-radical reactions. A rate constant of $k_9 = (8.7 \pm 1.7) \times 10^{-11} \text{ cm}^3 \text{s}^{-1}$ at 298K was obtained, more than a factor of 3 slower than Bossolasco *et* al. In order to elucidate such a large difference, Assaf et al. [14] have measured again the rate constant of (R9) using the same set-up as Bossolasco *et al.* [7]. In this work, the photolysis of XeF_2 in the presence of CH₄ and H₂O was used to simultaneously generate CH₃O₂ and OH radicals instead of the co-photolysis of CH₃I/H₂O₂. A rate constant of $k_9 = (1.6\pm0.4)\times10^{-10}$ cm³s⁻¹ was obtained, nearly a factor of 2 slower than the earlier determination by the same group [7], but still nearly a factor of 2 higher compared to the result of Yan et al. [13]. The high rate constant of Bossolasco et al. is now thought of having been induced by the presence of I-atoms (or accumulated I_2 molecules) when using CH_3I as precursor for CH_3O_2 radicals: a follow-up study of the reaction of larger peroxy radicals with OH radicals [15] showed the same behavior, i.e. a much faster rate constants when using the corresponding alkyl-iodides as precursor for the peroxy radicals (for more details see below). However, the rather large difference between the results of Assaf et al. [14] and Yan et al. [13] could not be explained. The measurements of Assaf et al. [14] are in principle more direct, *i.e.* less secondary reactions take place in their system, compared to the measurements of Yan et al. The rate constants obtained in both studies rely directly on the absorption cross sections of CH₃O₂ radicals used to transform the absorption time profiles into absolute CH₃O₂ concentration and any error in the absorption cross section would directly translate into a proportional error in the rate constant. Yan et al. used UV-absorption spectroscopy, a wavelength region where the absorption cross sections of peroxy radicals are well studied and believed to be well-known, while Assaf et al. used cw-CRDS in the near IR-region, a wavelength range much less explored. Indeed, the absorption cross section of CH₃O₂ in the near IR region is still discussed controversially: the value such as determined by Farago *et al.* [8] $(\sigma_{7489.16 \text{ cm}^{-1}} = (3.4 \pm 0.4) \times 10^{-20} \text{ cm}^2)$ and used by Assaf *et al.* is 2 to 3 times higher than the only two other determinations of this absorption cross sections that were published at that time [16,[17]. A too large absorption cross section would return a too small CH_3O_2 concentration and thus a too high rate constant. Even though Farago et al. produced possible reasons for the differences in their absorption cross sections compared to literature values (disregard of diffusion by Atkinson an Spillman [17] and disregard of secondary reactions by Pushkarsky et al. [16]), the doubt persisted. In 2016, Assaf *et al.* [14] re-determined the CH_3O_2 absorption cross section using not by using the kinetic method, i.e. measuring the CH_3O_2 decay and taking advantage of the known rate constant to retrieve the initial CH₃O₂ concentration, but relative to the cross section of HO₂ radicals, which is believed to be well known [18,[19,[20,[21]: they photolysed XeF₂ in the presence of either CH₃OH/O₂ or CH₄/O₂, and generated this way either HO₂ or CH₃O₂. Assuming on the basis of the known chemistry the nearcomplete conversion of F-atoms into either HO₂ radicals or CH₃O₂ radicals, the absorption cross section is then measured relative to the well-know HO₂ value, and the initially obtained value by Farago et al. [8] was confirmed. However, in 2018 Assaf et al. have carried out a detailed investigation of the reaction system of $F + CH_3OH$ [22] and they were able to determine experimentally for the first time the rate constant of the reaction

$$CH_3O + HO_2 \rightarrow products$$
 (R10)

The rate constant was found to be very fast ($k_{10} = 1.1 \times 10^{-10} \text{ cm}^3 \text{s}^{-1}$), in contrary to the only available estimations [23,[24] that predicted rate constants orders of magnitude slower ($10^{-13} \text{ cm}^3 \text{s}^{-1}$). With such a fast rate constant, the reaction of CH₃O+HO₂ is not anymore negligible under the condition of the redetermination of the absorption cross section of CH₃O₂ by Assaf *et al.* in 2016. The same is true for the CH₃O self-reaction:

$$CH_3O + CH_3O \rightarrow products$$
 (R11)

The rate constant had been estimated by Tsang and Hampson [24] to lead to the formation of CH₂O and CH₃OH with a very fast rate constant of k_{II} =1×10⁻¹⁰ cm³s⁻¹ while the few experimental determinations [25,[26,[27,[28,[29,[30] obtained lower values between (1.1–3.8)×10⁻¹¹ cm³s⁻¹, mostly from fitting reaction systems to complex mechanisms. Assaf *et al.* [22] determined k_{II} =7×10⁻¹¹ cm³s⁻¹, making this reaction also significant under their conditions. **Figure 2** shows a re-evaluation of the experimental data from Figure 3 in Assaf *et al.* [14] that had been used to re-determine the absorption cross section of CH₃O₂: the magenta line shows the model of the HO₂ profile without (R10) and (R11) and leads to an initial F-atom concentration of 4.4×10^{12} cm⁻³, converting the initial CH₃O₂ absorption coefficient (red symbols) into σ = 3.4×10^{-20} cm². The black line show the model including (R10) (F-atom concentration converted to products of this reaction as dashed green line in **Figure 2**) and (R11) (F-atom concentration converted to products of this reaction as dashed blue line in **Figure 2**), now yielding an initial F-atom concentration of 6.8×10^{12} cm⁻³. This higher initial F-atom concentration results in a proportionally lower absorption cross section for CH₃O₂: $\sigma = 2.2 \times 10^{-20}$ cm².



Figure 2: Data from Figure 3 in Assaf *et al.* 2016 [14]. Concentration time profile of HO₂ (black symbols, left y-axis) and absorption coefficient of CH_3O_2 (red symbols, right y-axis) as a function of time. 248nm photolysis of XeF₂ was the radical source, and experiments have been carried out back-to-back by using either 2.8×10^{15} cm⁻³ CH₃OH / 4.7×10^{17} cm⁻³ O₂ to generate HO₂ or 8.5×10^{15} cm⁻³ CH₄ / 4.7×10^{17} cm⁻³ O₂ to generate CH₃O₂. H₂O concentration was below 3×10^{14} cm⁻³, OH was below

detection limit. Full black line corresponds to a model taken from Assaf *et al.* 2018 [22], dashed magenta line presents the model neglecting (R10) and (R11).

This value is in good agreement with a very recent determination by Onel et al. [31], who finds 1.49×10^{-20} cm² at 7487.98 cm⁻¹, translating to 1.9×10^{-20} cm² at the wavelength used by Assaf *et al.* (7489.16 cm⁻¹). Assaf *et al.* had also applied the kinetic method to their experiments using XeF₂ as precursor (black symbols in **Figure 3**). Even though the initial radical concentrations using XeF_2 were much lower than for the experiments Farago *et al.* using CH_3I (red symbols in **Figure 3**), and with this the uncertainty of the extrapolation of the linear regression to I = 0 (black line in Figure 3, giving $m_{I=0}$ = $(2.9\pm0.4)\times10^7$ cm s⁻¹) was higher (*i.e.* the loss of CH₃O₂ due to diffusion was much more important than for Farago *et al.* : red line in Figure 3 with $m_{I=0} = (2.8\pm0.2)\times10^7$ cm s⁻¹), this method also confirmed the absorption cross section in agreement with the value of Farago et al.. The blue line in Figure 3 represents a regression with the intercept forced to a value $(4.36 \times 10^7 \text{ cm s}^{-1})$, dotted line in **Figure 3**) corresponding to $\sigma = 2.2 \times 10^{-20} \text{ cm}^2$. It can be seen that all decays from Assaf *et al.* are above the intercept corresponding to $\sigma = 2.2 \times 10^{-20} \text{ cm}^2$, *i.e.* in agreement with a decay through CH₃O₂ selfreaction plus an increasing loss through diffusion with decreasing concentration, while this is not true for Farago *et al*: all decays are *slower* than would be expected for CH_3O_2 decay only, and they become even slower with increasing concentration. This disagreement with the work of Farago et al. is currently not understood, maybe I-atom chemistry somehow perturb the CH₃O₂ decay [32,[33].



Figure 3: Data from Farago *et al* [8] (red symbols) and from Assaf *et al.* [14] (black symbols). Black line corresponds to a linear regression of the values of Assaf *et al.*, blue line is a regression forced to an intercept corresponding to $\sigma = 2.2 \times 10^{-20} \text{ cm}^2$.

Because the rate constant of (R9) such as determined by Assaf *et al.* (and Bossolasco *et al.*) directly depends on the absorption cross section of CH_3O_2 , this now lower value entails a proportionally lower value of the rate constant: 1.0×10^{-10} cm³s⁻¹. The new value is in good agreement with the value of Yan *et al.* [13] and it seems that near IR absorption cross section of CH_3O_2 and rate constant for (R9) are

settling towards a final value. This long story confirms the rule that rate constants decrease with time.... The results of the three experimental studies as well as the re-evaluated rate constant from Assaf *et al.* [14] using the re-evaluated absorption cross section for CH_3O_2 are summarized in **Table 1**.

Method	Rate constant / cm ³ s ⁻¹	Remark	Reference	
248 nm Laser photolysis of CH ₃ I		Rate constant relies on near IR	Bossolasco	
/ H_2O_2 , detection of CH_3O_2 by	$(2.8\pm1.4) imes 10^{-10}$	absorption cross section of CH ₃ O ₂ ,	et al. 2014	
cw-CRDS, OH by LIF		rate constant probably biased by I/I_2	[7]	
193 nm Laser photolysis of		Profiles fitted to complex mechanism,	Yan <i>et al</i> .	
CH ₃ COCH ₃ / N ₂ O, detection of	$(8.4\pm1.7) imes 10^{-11}$	high initial radical concentrations up	2016	
CH ₃ O ₂ and OH by UV absorption		to $10^{14} \mathrm{cm}^{-3}$	[13]	
		Rate constant relies on near IR	Assaf <i>et al</i> .	
248 nm Laser photolysis of XeF ₂	$(1.6\pm0.4) \times 10^{-10}$	absorption cross section of CH ₃ O ₂	2016	
/ CH ₄ / H ₂ O, detection of CH ₃ O ₂		$(3.4 \times 10^{-20} \text{ cm}^{-2})$	[14]	
by cw-CRDS, OH by LIF	$(1.0\pm0.4) \times 10^{-10}$	Rate constant using absorption cross	This work	
		section re-evaluated $(2.2 \times 10^{-20} \text{ cm}^{-2})$		

Table 1: Summary of rate constant measurements for the reaction of CH₃O₂ + OH radicals

Product yields and atmospheric implications

Following the first measurement of the rate constant k_9 by Bossolasco *et al.* [7], Fittschen *et al.* [11], came to the conclusion, that this reaction is a non-negligible sink for CH₃O₂ radicals over tropical oceans with up to 30% of CH₃O₂ reacting through this pathway. Even though the rate constant of (R1) has been revised since then and it can be estimated that (R1) will be proportionally less important as sink for CH₃O₂ (i.e. around 10%), it becomes important to determine the reaction products in order to appreciate the impact of this reaction on the composition of the atmosphere. Müller *et al.* [34] have investigated this question by means of detailed theoretical calculations as well as by integrating this reaction with different product branching scenarios into a global model. By high level *ab-initio* calculations they characterized the triplet and singlet potential energy surfaces (PES) and deduced that the triplet entrances are negligible: the dominant initial product is the singlet trioxide intermediate, CH₃OOOH

$$OH + CH_3O_2 \rightarrow CH_3OOOH$$
(R9d)

which can either be stabilized or rapidly converted to a pre-product complex (CH₃O...HOO). The coupling of the triplet and singlet surfaces via intersystem crossing can affect the product distribution and makes a prediction of product distributions through RRKM calculations less reliable. The singlet state undergoes rapid bond scission and yields the bimolecular products CH₃O + HO₂ (R9b) with a small contribution to the channel CH₃OH + O₂(¹ Δ) (R9c), while the triplet surface mostly yields CH₃O + HO₂ (R9b) with small contribution to CH₃OH + O₂ (${}^{3}\Sigma_{g}$ ⁻). Formation of the Criegee intermediate through pathway (R9a) has been excluded, which had also been found experimentally by Yan *et al.*

[13] and later by Assaf *et al.* [35]. Müller *et al.* [34] proposed that this reaction might help resolving a strong missing source of methanol over tropical oceans [36]. From running different product branching scenarios in their global model, they deduced that, when using the rate constant of Bossolasco *et al.*, a yield of up to 30% CH₃OH would be needed as product of (R9) to bring into agreement measurements and models, while their RRKM calculations predicted only a yield of 7% for CH₃OH. Besides, it is now clear that the rate constant k_9 from Bossolasco *et al.* is too high by a factor of 3, which leads to an even higher yield that would be necessary for bringing into agreement model and measurements. Ferracci *et al.* [37] have run a global model to investigate the impact of (R9) on the composition of the atmosphere, now using the revised rate constant from Assaf *et al.* [14]. They found that an unreasonably high yield of $\phi_{CH_3OH} = 0.8$ would be needed to reconcile model and measurements.

A first experimental indication on possible product yields of (R9) was given by Yan et al. [13] who deduced from modeling their UV absorption-time profiles to a complex mechanism an upper limit of 5% for the formation of the Criegee intermediate (R9a). This upper limit was later confirmed by Assaf et al. [35] using broadband cavity enhanced absorption spectroscopy [38]. Assaf et al. [35] could by simultaneous measurement of absolute HO2 [18]and OH [39] concentration-time profiles determine the HO₂ yield (R9b) at a total pressure of 50 Torr Helium to be $\phi_{HO_2} = 0.8\pm0.2$. This yield was slightly revised in a later work [40] to $\phi_{HO_2} = 0.9 \pm 0.1$ because it was found in the meantime, that (R10) and (R11) are fast and played some role in the reaction system of Assaf et al. [35]. This high HO₂-yield left little room for a large CH₃OH yield, in agreement with the predictions of Müller *et al.* [34]. A very recent work by Caravan et al. [41] was then intended to a direct determination of the CH₃OH yield. Three different experimental set-ups where used: a low pressure (30 Torr) or a high pressure (740 Torr) laser photolysis reactor, both coupled to time resolved VUV photoionization mass spectrometry [42] as well as an atmospheric simulation chamber coupled to a detection of reaction products by PTR-ToFMS. From simulation of the laser photolysis experiments, a yield of $\phi_{CH_2OH} = 9\pm5$ % and 6 ± 2 % was obtained from experiments in the low and high pressure reactor, respectively. Experiments in the high pressure reactor also found evidence for a non-negligible yield of the stabilized trioxide, CH₃OOOH (R9d), no signal however could be detected when using the low pressure reactor. This is in good agreement with the calculations of Müller et al. [34], who predicted a yield for the stabilization of $\phi_{CH_2OOOH} = 7$ % at 760 Torr and $\phi_{CH_2OOOH} = 0.02$ % at 30 Torr. The experiments in the atmospheric simulation chamber resulted in a markedly higher yield of $\phi_{CH_2OH} = 17\pm3$ %, not in agreement with the laser photolysis experiment, even considering the reciprocal error bars, while no CH₃OOOH signal could be detected in the simulation chamber / PTR ToFMS experiments, even though carried out at 760 Torr. It was therefore speculated that the trioxide, CH₃OOOH, is detected by the PTR-ToFMS at the mass of CH₃OH, thus increasing the apparent CH₃OH yield. It could not be decided if the

CH₃OOOH decomposes into CH₅O⁺ upon protonisation or if CH₃OOOH is already decomposing heterogeneously into CH₃OH, either on the Teflon walls of the simulation chamber or on the walls of the sampling tubes. This hypothesis brings into excellent agreement theoretical calculations, Teflon chamber and laser photolysis experiments, as the CH₃OH yield determined in the Teflon chamber corresponds to the predicted sum of CH₃OH and CH₃OOOH. This now experimentally confirmed low CH₃OH yield results in an even decreased concentration of CH₃OH due to the decreased concentration of CH₃O₂ radicals (from its reaction with OH), whose self-reaction is the major source of CH₃OH in the remote atmosphere.

Experiment and theory corroborate the formation of the trioxide, CH₃OOOH, from the reaction between CH₃O₂ and OH radicals. The atmospheric fate of this new species is therefore of interest. Anglada and Solé [43] have recently employed high level theoretical methods to investigate the oxidation of CH₃OOOH by OH radicals. They found that the reaction can proceed by abstraction of either the terminal hydrogen atom of OOOH group leading to formation of CH₃O, O₂ and H₂O, or of one of the hydrogen atoms of the CH₃ group producing H₂CO, HO₂ and H₂O. They found that the abstraction of the terminal hydrogen atom of the OOH group is with $k = 4.7 \times 10^{-11}$ cm³s⁻¹ about 22 times faster than the abstraction of a hydrogen atom of the CH₃ group ($k = 2.1 \times 10^{-12}$ cm³s⁻¹). The lifetime of CH₃OOOH in the troposphere has been predicted to be about 3.9 hours at 275 K and decreasing to 0.2 hours at 310 K. Such a lifetime could be sufficient to allow accumulation of sizeable concentrations of CH₃OOOH under very low NO concentrations. **Figure 4** summarizes the current knowledge on the branching of reaction products for the reaction of CH₃O₂ + OH.





The reaction of larger $RO_2 + OH$

Rate constants

While the reaction of CH_3O_2 radicals with OH radicals plays an important role over remote tropical oceans, the reactions of larger peroxy radicals with OH could play a role in remote biogenic environments. Therefore, the rate constant of OH radicals with the next larger peroxy radical has been measured by Farago *et al.* [44]:

$OH + C_2H_5O_2 \rightarrow products$ (R12)

In that work, OH radicals have been generated by photolysis of H_2O_2 , i.e. the same way as in Bossolosco et al. [7], while the peroxy radicals were generated from the reaction of Cl-atoms (generated by simultaneous 248 nm photolysis of $(COCl)_2$) with C_2H_6 . This method was not applicable for the generation of CH₃O₂: the reaction of CH₄ with Cl-atoms is only 16 times faster than the reaction of OH radicals with CH_4 , while this ratio is around 250 in the case of C_2H_6 . Therefore, the CH₄ concentrations necessary for rapidly converting Cl-atoms into CH₃ radicals would at the same time result in a fast reaction of OH radicals with CH₄ (and thus only a small fraction of OH radicals would react with CH_3O_2). Another difference between the two studies of CH_3O_2 radicals [7,[14] and the $C_2H_5O_2$ study was the quantification method of the absolute peroxy radical concentrations: in the case of CH₃O₂, the radicals were directly quantified by cw-CRDS absorption spectroscopy, and thus the rate constant directly relies on the absorption cross section of CH_3O_2 used to transform the absorption-time profiles into concentration-time profiles (see first part of this paper). In the case of $C_2H_5O_2$, this method was not applicable, because the absorption cross section of $C_2H_5O_2$ in the near infrared region is not well known. Therefore, the initial Cl-atom concentration was first quantified by transforming them into HO₂ radicals through their fast reaction with excess CH₃OH, which can reliably be quantified [18,[19]. Thereafter, CH₃OH was switched for C₂H₆, and it was supposed that the resulting peroxy radical concentration was equal to the HO₂ concentration. Figure 5 shows typical experimental results for the reaction of C₃H₇O₂ with OH radicals.



Figure 5: Left graph: HO₂ concentration time profiles used to quantify the initial Cl-atom concentration through exptrapolation of the decay to t=0: $[CH_3OH] = 2 \times 10^{14} \text{ cm}^{-3}$, $[O_2] = 2.7 \times 10^{17} \text{ cm}^{-3}$, $[(COCl)_2]$ increased from $(0.4 - 3.5) \times 10^{14} \text{ cm}^{-3}$, from bottom to top. Right graph: OH decays for the reaction of $C_3H_7O_2$ with OH, *i.e.* same conditions as left graph, but replacing CH₃OH by $[C_3H_8] = 3.1 \times 10^{14} \text{ cm}^{-3}$. Figure adapted from Figure 1 in [15].

This way, a rate constant of $k_{12} = (1.2\pm0.3)\times10^{-10} \text{ cm}^3\text{s}^{-1}$ was obtained, slightly faster than the rate constants of (R9). In a follow-up work, Assaf *et al.* [15] have also determined the rate constants for the two next-larger peroxy radicals, C₃H₇O₂ and C₄H₉O₂, with OH. Both radicals have been generated the same way as described above for C₂H₅O₂: the reaction of Cl-atoms with the corresponding hydrocarbon, with the Cl-atoms being quantified in back-to-back experiments with CH₃OH. Rate constants of $(1.4\pm0.3)\times10^{-10} \text{ cm}^3\text{s}^{-1}$ and $(1.5\pm0.3)\times10^{-10} \text{ cm}^3\text{s}^{-1}$ have been determined for propyl- and butylperoxy, respectively. All results are summarized in **Table 2**. Some test experiments were carried out using the photolysis of the corresponding alkyl iodides as precursor for the larger peroxy radicals, and again rate constants in the range of $3\times10^{-10} \text{ cm}^3\text{s}^{-1}$ have been found, supporting the hypothesis that the presence of I-atoms or accumulated I₂ molecules systematically biased the experiments of Bossolasco *et al.* [7].

Product yields and atmospheric implications

The reaction of RO₂ radicals with OH radicals might become competitive at decreasing NO concentrations. In the case of CH₃O₂, this will be the case in remote marine boundary layers, while the reaction of peroxy radicals generated from larger hydrocarbons might become competitive in remote biogenic environments. Therefore, it is interesting to determine the product yields for the reaction of larger peroxy radicals. A first theoretical investigation of the reaction of C₂H₅O₂ with OH radicals was carried out by Liu *et al.* [45], who determined the PES at the CCSD(T)/6-311++G(2df,2p)//B3LYP/6-311+G(d,p) level of theory. The results showed that, same as for CH₃O₂, the C₂H₅O₂ + OH reaction proceeds on both the singlet and the triplet PES. On the singlet PES, the favorable pathway is the addition of OH radical to the terminal oxygen atom of the C₂H₅O₂ radical, leading to the formation of trioxide C₂H₅O₃H through a barrierless process. The trioxide then directly decomposes to the products C₂H₅O and HO₂ radicals. On the triplet PES, the predominant pathways are the abstraction of a hydrogen atom forming biradical products (CH₃CHO₂) and (CH₂CH₂O₂) with barriers of 12.02 and 19.19 kJ mol⁻¹, respectively. The properties of the trioxide C₂H₅O₃H were investigated and the results indicate that the trioxide can exist stably in the atmosphere owing to a significantly large and negative enthalpy of formation (-118.44 kJ/mol).

The only experimental investigation of the product yields for the reaction of larger peroxy radicals with OH radicals was carried out by Assaf *et al.* [40]. Using the same technique as in the case of CH₃O₂ [35], *i.e.* the simultaneous quantitative detection of OH and HO₂ radicals by cw-CRDS, coupled to a simultaneous generation of OH and RO₂ radicals by 248nm photolysis of XeF₂ in presence of H₂O and different hydrocarbons, they determined the HO₂ yield for C₁ – C₄ peroxy radicals. It was found that at 50 Torr total pressure the HO₂ yield strongly decreased with increasing

size of the alkyl moiety of the peroxy radical: $\phi_{CH_3O_2} = (0.90\pm0.1)$, $\phi_{C_2H_5O_2} = (0.75\pm0.15)$, $\phi_{C_3H_7O_2} = (0.41\pm0.08)$, $\phi_{C_4H_9O_2} = (0.15\pm0.03)$. An accompanying theoretical study optimized the geometries and rovibrational characteristics of the critical points on the PES at the M06-2X/aug-*cc*pVTZ level of theory and further refined the relative energies at the CCSD(T)/aug-*cc*-pVTZ level of theory. RRKM master equation calculations could explain the clear decrease in HO₂ yield with increasing size of the alkyl moiety by an increased stabilization of the corresponding trioxide, ROOOH. Extrapolation of the experimental results to atmospheric conditions showed that the stabilized adduct, ROOOH, would be the nearly exclusive product of the reaction between OH radicals and peroxy radicals containing more than 3 C-atoms such as peroxy radicals obtained from the oxidation of biogenic hydrocarbons.

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Reaction	Precursor	$k / \text{cm}^3 \text{s}^{-1}$	HO_2 yield / %	Reference
$CH_3O_2 + OH$	XeF_2+CH_4/H_2O	$(8.4\pm1.7)\times10^{-11}$ $(1.0\pm0.4)\times10^{-10}$	90 ±10	Yan <i>et al.</i> [13], Assaf <i>et al.</i> [14,[35], this work
$C_2H_5O_2 + OH$	$XeF_2+C_2H_6/H_2O$	$(1.2\pm0.3)\times10^{-10}$ $(1.3\pm0.3)\times10^{-10}$	75±15	Farago <i>et al</i> . [44] Assaf <i>et al</i> . [15,[40]
$C_3H_7O_2 + OH$	$\begin{array}{c} XeF_2+n\text{-}C_3H_8/\\ H_2O \end{array}$	(1.4±0.3)×10 ⁻¹⁰	41±8	Assaf et al. [15,[40]
$C_4H_9O_2 + OH$	$XeF_2 + n-C_4H_{10} / H_2O$	(1.5±0.3)×10 ⁻¹⁰	15±3	Assaf <i>et al</i> . [15,[40]

Table 2: Summary of rate constant and HO_2 yield measurements for the reaction of different RO_2 + OH radicals

It can thus be supposed, that ROOOH is formed in remote biogenic environments and non-negligible concentrations can possibly accumulate, depending on the lifetime of this class of species. In a recent work, Fittschen et al. [46] suggested that this class of species might be responsible for an unidentified interference in OH quantification in remote biogenic environments. Indeed, several field campaigns in such environments [47,[48,[49], all carried out by laser induced fluorescence after gas expansion (FAGE), have measured OH concentrations up to a factor of 10 higher than predicted by models. To explain such high OH concentrations, the oxidation mechanisms of biogenic VOCs, and especially of isoprene, have since been investigated in detail by theory e.g. [50,[51] and experiment e.g. [52,[53] in order to unravel unknown mechanism of OH-recycling in absence of NO and has been included in models since then [54]. And even though new OH recycling chemistry has been discovered, the impact on the steady-state OH concentration is not high enough to explain the large disagreement between measurements and models. W. Brune and coworkers [55] suspected interferences in FAGE measurements to be at least partially responsible for the high OH concentrations measured under remote biogenic conditions. They upgraded their FAGE instrument to unravel such interferences by mixing high concentrations of C_3H_6 or C_3F_6 into the airflow just before intake into the FAGE instrument. This way, all ambient OH would react away, and any remaining signal would be identified as interference. This way, they observed strong interferences in OH measurements with their FAGE instrument [55,[56], with the intensity increasing with decreasing NO concentration. Other groups also found interferences using the same technique of pre-injection [57,[58,[59]. The decomposition of Criegee intermediates within the FAGE cell has been suspected to be at the origin of this interference [57], however other groups could not confirm this hypothesis [59]. Currently there is no consensus on the origin of this interference. Fittschen *et al.* [46] have recently shown that ROOOH, generated from the reaction of butylperoxy or isoprenehydroxyperoxy radicals with OH radicals, leads in their FAGE instrument to OH interference. However, the mechanism of this OH generation is currently not clear and might be due to heterogeneous decomposition of ROOOH within their FAGE cell. As all FAGE instruments have different designs, other FAGE instruments might behave differently and more work is needed to investigate if this class of species could explain the high OH concentrations observed in earlier field campaigns in remote forested areas.

In the recent attempt of designing the next generation of explicit mechanism, based on a combination of the Generator for Explicit Chemistry and Kinetics of Organics in the Atmosphere, GECKO-A [60], and the Master Chemical Mechanism, MCM [4], the reactions of RO₂ radicals with OH radicals are taken into account [61], and the next few years will probably bring more knowledge about this class of reactions through well-designed experiments.

Conclusion

The reaction of peroxy radicals with OH radicals as a possible reaction pathway for peroxy radicals in low NO environments has been ignored in atmospheric chemistry models until recently. Its possible role in clean atmospheres was evoked for the first time in 2009 by Archibald et al. [6], but it was only since Bossolasco *et al.* [7] measured for the first time in 2014 a rate constant for the simplest reaction $CH_3O_2 + OH$, that common interest in this class of reaction grew. The current manuscript summarizes the findings that have been achieved since then: rate constants have been measured for peroxy radicals up to C4, and some product yields for the simplest RO₂ radicals have also been measured. The role of the reaction of $CH_3O_2 + OH$ over tropical oceans has already been included in models, the major product of this reaction is $CH_3O + HO_2$, and it is clear now that this reaction is not a source for atmospheric methanol. The possible role of the reactions of larger peroxy radicals with OH in remote biogenic environments however has not yet been investigated. The major product of this reaction is the trioxide, ROOOH, a new class of species that can accumulate to sizeable concentrations in remote environments. Their fate will probably attract some interest in the coming years.

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References

- [1] J. J. Orlando, G. S. Tyndall, Chem. Soc. Rev. 41 (2012) 6294-6317.
- G. S. Tyndall, R. A. Cox, C. Granier, R. Lesclaux, G. K. Moortgat, M. J. Pilling, A. R.
 Ravishankara, T. J. Wallington, J. Geophys. Res. 106 (2001) 12157-12182.
- [3] S. M. Saunders, M. E. Jenkin, R. G. Derwent, M. J. Pilling, Atmos. Environ. 31 (1997) 1249-1249.
- [4] S. M. Saunders, M. E. Jenkin, R. G. Derwent, M. J. Pilling, Atmos. Chem. Phys. 3 (2003) 161-180.
- [5] C. Bloss, V. Wagner, A. Bonzanini, M. E. Jenkin, K. Wirtz, M. Martin-Reviejo, M. J. Pilling, Atmos. Chem. Phys. 5 (2005) 623-639.
- [6] A. T. Archibald, A. S. Petit, C. J. Percival, J. N. Harvey, D. E. Shallcross, Atmos. Sci. Lett. 10 (2009) 102-108.
- [7] A. Bossolasco, E. P. Faragó, C. Schoemaecker, C. Fittschen, Chem. Phys. Lett. 593 (2014) 7-13.
- [8] E. P. Faragó, B. Viskolcz, C. Schoemaecker, C. Fittschen, J. Phys. Chem. A 117 (2013) 12802-12811.
- [9] A. Parker, C. Jain, C. Schoemaecker, C. Fittschen, React. Kinet. Catal. Lett. 96 (2009) 291-297.
- [10] J. Thiebaud, A. Aluculesei, C. Fittschen, J. Chem. Phys. 126 (2007) 186101.
- [11] C. Fittschen, L. K. Whalley, D. E. Heard, Environ. Sci. Technol. 118 (2014) 7700–7701.
- [12] L. K. Whalley, K. L. Furneaux, A. Goddard, J. D. Lee, A. Mahajan, H. Oetjen, K. A. Read, N. Kaaden, L. J. Carpenter, A. C. Lewis, J. M. C. Plane, E. S. Saltzman, A. Wiedensohler, D. E. Heard, Atmos. Chem. Phys. 10 (2010) 1555-1576.
- [13] C. Yan, S. Kocevska, L. N. Krasnoperov, J. Phys. Chem. A 120 (2016) 6111–6121.
- [14] E. Assaf, B. Song, A. Tomas, C. Schoemaecker, C. Fittschen, J. Phys. Chem. A 120 (2016) 8923-8932.
- [15] E. Assaf, S. Tanaka, Y. Kajii, C. Schoemaecker, C. Fittschen, Chem. Phys. Lett. 684 (2017) 245-249.
- [16] M. B. Pushkarsky, S. J. Zalyubovsky, T. A. Miller, J. Chem. Phys. 112. (2000) 10695-10698.
- [17] D. B. Atkinson, J. L. Spillman, J. Phys. Chem. A 106 (2002) 8891-8902.
- [18] J. Thiebaud, S. Crunaire, C. Fittschen, J. Phys. Chem. A 111 (2007) 6959-6966.
- [19] Y. Tang, G. S. Tyndall, J. J. Orlando, J. Phys. Chem. A 114 (2010) 369-378.
- [20] L. Onel, A. Brennan, M. Gianella, G. Ronnie, A. Lawry Aguila, G. Hancock, L. Whalley, P. W. Seakins, G. A. D. Ritchie, D. E. Heard, Atmos. Meas. Tech. Discuss. 10 (2017) 4877-4894.
- [21] E. Assaf, L. Liu, C. Schoemaecker, C. Fittschen, Journal of Quantitative Spectroscopy & Radiative Transfer 211 (2018) 107-114.
- [22] E. Assaf, C. Schoemaecker, L. Vereecken, C. Fittschen, Phys. Chem. Chem. Phys. 20 (2018) 8707.

- [23] S. H. Mousavipour, Z. Homayoon, J. Phys. Chem. A 115 (2011) 3291-3300.
- [24] W. Tsang, R. F. Hampson, J. Phys. Chem. Ref. Data 15 (1986) 1087.
- [25] E. Hassinen, J. Koskikallio, Acta Chemica Scandinavica Series a-Physical and Inorganic Chemistry 33 (1979) 625-630.
- [26] J. Weaver, R. Shortridge, J. Meagher, J. Heicklen, Journal of Photochemistry 4 (1975) 109-120.
- [27] M. J. Yee Quee, J. C. J. Thynne, Berichte der Bunsengesellschaft für physikalische Chemie 72 (1968) 211-217.
- [28] M. J. Y. Quee, J. C. J. Thynne, Transactions of the Faraday Society 62 (1966) 3154-3161.
- [29] P. Biggs, C. E. Canosa-Mas, J.-M. Fracheboud, D. E. Shallcross, R. P. Wayne, J. Chem. Soc., Faraday Trans. 93 (1997) 2481-2486.
- [30] U. Meier, H. H. Grotheer, G. Riekert, T. Just, Berichte der Bunsengesellschaft für physikalische Chemie 89 (1985) 325-327.
- [31] L. Onel, A. Brennan, M. Gianella, J. Hooper, N. Ng, G. Hancock, L. Whalley, P. W. Seakins, G. A. D. Ritchie, D. E. Heard, Atmos. Meas. Tech. to be submitted (2019).
- [32] T. J. Dillon, M. E. Tucceri, J. N. Crowley, Chemphyschem 11 (2010) 4011-4018.
- [33] T. J. Dillon, M. E. Tucceri, J. N. Crowley, Phys. Chem. Chem. Phys. 8 (2006) 5185-5198.
- [34] J.-F. Müller, Z. Liu, V. S. Nguyen, T. Stavrakou, J. N. Harvey, J. Peeters, Nat. Commun. 7 (2016) 13213.
- [35] E. Assaf, L. Sheps, L. Whalley, D. Heard, A. Tomas, C. Schoemaecker, C. Fittschen, Environ. Sci. Technol. 51 (2017) 2170-2177.
- [36] D. J. Jacob, B. D. Field, Q. Li, D. R. Blake, J. de Gouw, C. Warneke, A. Hansel, A. Wisthaler, H.
 B. Singh, A. Guenther, J. Geophys. Res.-Atmos. 110 (2005) D08303.
- [37] V. Ferracci, I. Heimann, N. L. Abraham, J. A. Pyle, A. T. Archibald, Atmos. Chem. Phys. 18 (2018) 7109-7129.
- [38] L. Sheps, J. Phys. Chem. Lett. 4 (2013) 4201-4205.
- [39] E. Assaf, C. Fittschen, J. Phys. Chem. A 120 (2016) 7051-7059.
- [40] E. Assaf, C. Schoemaecker, L. Vereecken, C. Fittschen, Int. J. Chem. Kinet. 50 (2018) 670–680.
- [41] R. L. Caravan, M. A. H. Khan, J. Zádor, L. Sheps, I. O. Antonov, B. Rotavera, K. Ramasesha, K. Au, M.-W. Chen, D. Rösch, D. L. Osborn, C. Fittschen, C. Schoemaecker, M. Duncianu, A. Grira, S. Dusanter, A. Tomas, C. J. Percival, D. E. Shallcross, C. A. Taatjes, Nat. Commun. 9 (2018) 4343.
- [42] D. L. Osborn, P. Zou, H. Johnsen, C. C. Hayden, C. A. Taatjes, V. D. Knyazev, S. W. North, D. S. Peterka, M. Ahmed, S. R. Leone, Rev. Sci. Instrum. 79 (2008) 104103.
- [43] J. M. Anglada, A. Solé, Phys. Chem. Chem. Phys. 20 (2018) 27406-27417.
- [44] E. P. Faragó, C. Schoemaecker, B. Viskolcz, C. Fittschen, Chem. Phys. Lett. 619 (2015) 196-200.
- [45] Y. Liu, L. Chen, D. Chen, W. Wang, F. Liu, W. Wang, Chemical Research in Chinese Universities 33 (2017) 623-630.
- [46] C. Fittschen, M. Al Ajami, S. Batut, V. Ferracci, S. Archer-Nicholls, A. T. Archibald, C. Schoemaecker, Atmos. Chem. Phys. 19 (2019) 349-362.
- [47] A. Hofzumahaus, F. Rohrer, K. Lu, B. Bohn, T. Brauers, C. C. Chang, H. Fuchs, F. Holland, K. Kita, Y. Kondo, X. Li, S. Lou, M. Shao, L. Zeng, A. Wahner, Y. Zhang, Science 324 (2009) 1702-1704.
- [48] L. K. Whalley, P. M. Edwards, K. L. Furneaux, A. Goddard, T. Ingham, M. J. Evans, D. Stone, J. R. Hopkins, C. E. Jones, A. Karunaharan, J. D. Lee, A. C. Lewis, P. S. Monks, S. J. Moller, D. E. Heard, Atmos. Chem. Phys. 11 (2011) 7223-7233.
- [49] J. Lelieveld, T. M. Butler, J. N. Crowley, T. J. Dillon, H. Fischer, L. Ganzeveld, H. Harder, M. G. Lawrence, M. Martinez, D. Taraborrelli, J. Williams, Nature 452 (2008) 737-740.
- [50] J. Peeters, T. L. Nguyen, L. Vereecken, Phys. Chem. Chem. Phys. 11 (2009) 5935-5939.
- [51] J. Peeters, T. L. Nguyen, J. Phys. Chem. A (2012).

- [52] P. O. Wennberg, K. H. Bates, J. D. Crounse, L. G. Dodson, R. C. McVay, L. A. Mertens, T. B. Nguyen, E. Praske, R. H. Schwantes, M. D. Smarte, J. M. S. Clair, A. P. Teng, X. Zhang, J. H. Seinfeld, Chem. Rev. 118 (2018) 3337-3390.
- [53] H. Fuchs, A. Hofzumahaus, F. Rohrer, B. Bohn, T. Brauers, H. P. Dorn, R. Haseler, F. Holland,
 M. Kaminski, X. Li, K. Lu, S. Nehr, R. Tillmann, R. Wegener, A. Wahner, Nature Geosci advance online publication (2013).
- [54] M. E. Jenkin, J. C. Young, A. R. Rickard, Atmos. Chem. Phys. 15 (2015) 11433-11459.
- J. Mao, X. Ren, W. H. Brune, D. M. Van Duin, R. C. Cohen, J.-H. Park, A. H. Goldstein, F. Paulot, M. R. Beaver, J. D. Crounse, P. O. Wennberg, J. P. DiGangi, S. B. Henry, F. N. Keutsch, C. Park, G. W. Schade, G. M. Wolfe, J. A. Thornton, Atmos. Chem. Phys. - ACP 12 (2012) 8009-8020.
- P. A. Feiner, W. H. Brune, D. O. Miller, L. Zhang, R. C. Cohen, P. S. Romer, A. H. Goldstein, F. N. Keutsch, K. M. Skog, P. O. Wennberg, T. B. Nguyen, A. P. Teng, J. DeGouw, A. Koss, R. J. Wild, S. S. Brown, A. Guenther, E. Edgerton, K. Baumann, J. L. Fry, Journal of the Atmospheric Sciences 73 (2016) 4699-4710.
- [57] A. Novelli, K. Hens, C. Tatum Ernest, M. Martinez, A. C. Nölscher, V. Sinha, P. Paasonen, T.
 Petäjä, M. Sipilä, T. Elste, C. Plass-Dülmer, G. J. Phillips, D. Kubistin, J. Williams, L. Vereecken,
 J. Lelieveld, H. Harder, Atmos. Chem. Phys. 17 (2017) 7807-7826.
- [58] A. Novelli, K. Hens, C. T. Ernest, D. Kubistin, E. Regelin, T. Elste, C. Plass-Dulmer, M. Martinez, J. Lelieveld, H. Harder, Atmospheric Measurement Techniques 7 (2014) 3413-3430.
- [59] P. Rickly, P. S. Stevens, Atmospheric Measurement Techniques 11 (2018) 1-16.
- [60] B. Aumont, S. Szopa, S. Madronich, Atmos. Chem. Phys. 5 (2005) 2497-2517.
- [61] M. E. Jenkin, R. Valorso, B. Aumont, A. R. Rickard, Atmos. Chem. Phys. Discuss. 2019 (2019) 1-46.