

Lithospheric strength: a potential controlling factor over differential subsidence in the Early Triassic Sonoma Foreland Basin (western USA)?

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Lithospheric strength: a potential controlling factor over differential subsidence in the Early Triassic Sonoma Foreland Basin (western USA)?

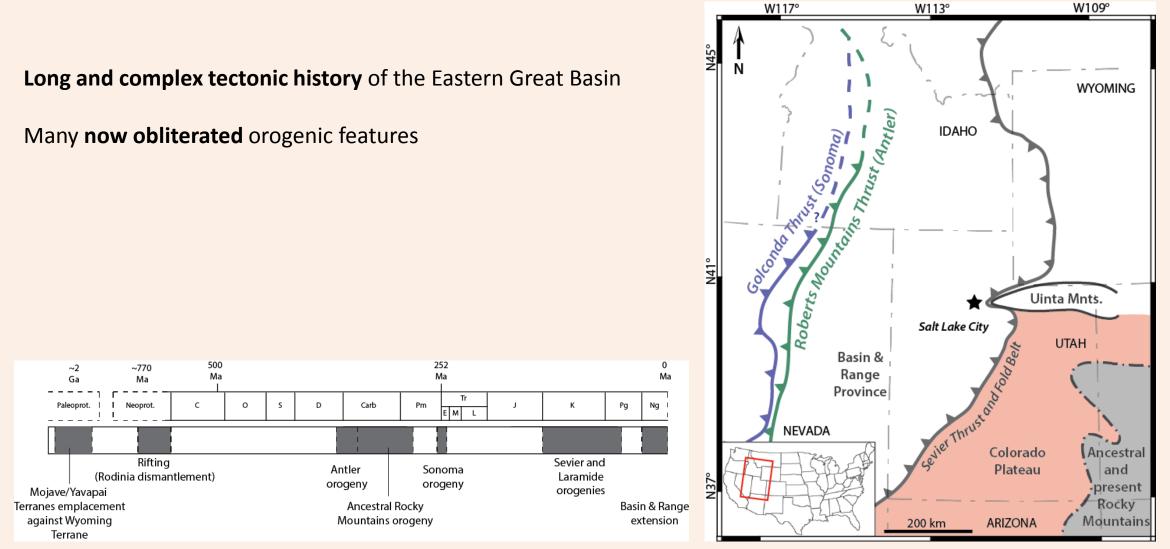
Gwénaël Caravaca*, Arnaud Brayard, Emmanuelle Vennin, Michel Guiraud, Nicolas Olivier, Anne-Sabine Grosjean, Christophe Thomazo, Emmanuel Fara, Gilles Escarguel * <u>gwenael.caravaca@u-bourgogne.fr</u> Université de Bourgogne Franche-Comté, France

CINIS

AGINE NATIONALE DE LA RECHERCHE

BIOGÉOSCIENCES

Introduction and localization



After Caravaca et al., subm.

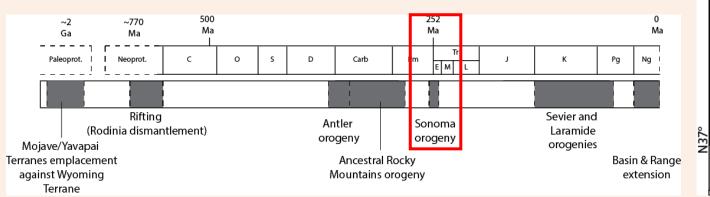
Introduction and localization

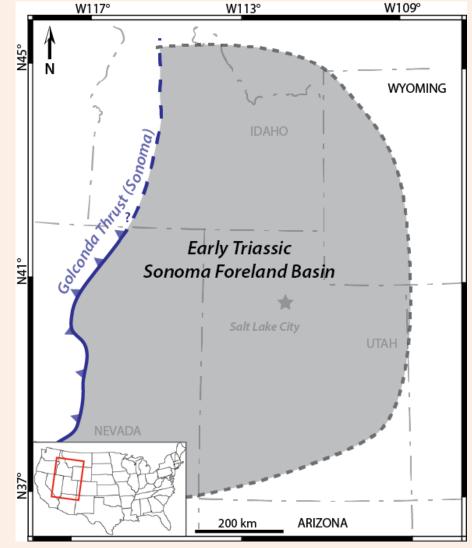
Long and complex tectonic history of the Eastern Great Basin

Many now obliterated orogenic features

- ⇒ Focus on the Sonoma orogeny (~252 Ma at the Permian-Triassic Boundary)
- \Rightarrow Sonoma Foreland Basin

formed by emplacement of Golconda Allochthon



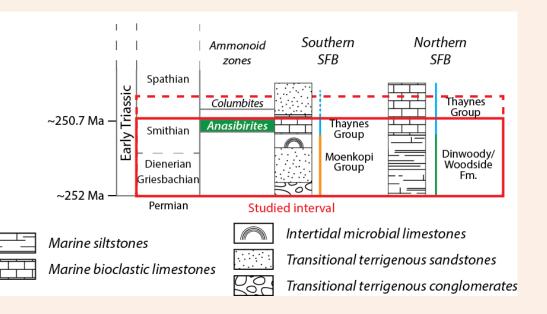


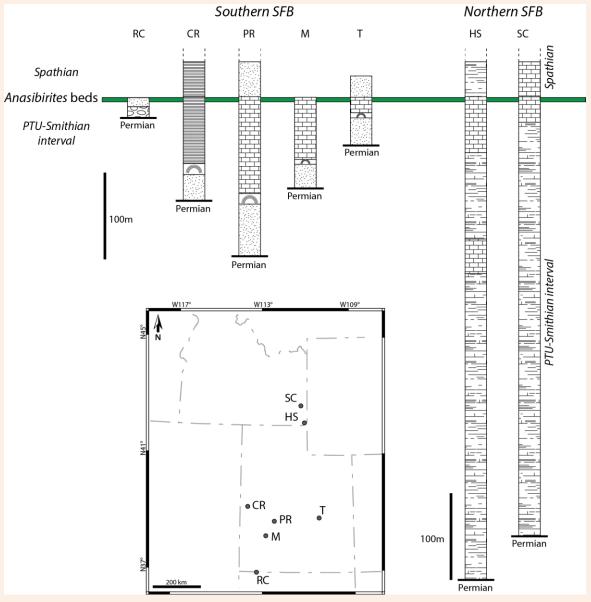
After Caravaca et al., subm.

Introduction and localization

Sonoma Foreland Basin (SFB)

- ⇒ Excellent fossil and sedimentary record of the aftermath of the end-Permian mass-extinction
- ⇒ PT unconformity-Smithian interval (PTU; ~1,3 Myr-long)
- ⇒ Marked discrepancies in term of facies and thicknesses between northern and southern parts





Caravaca et al., subm.

Methods

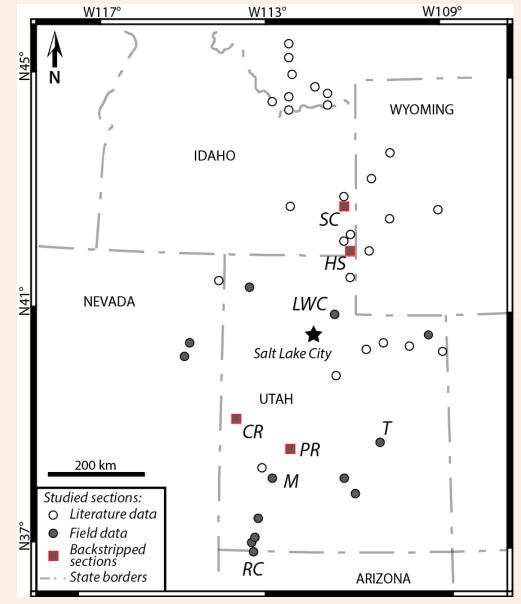
Focus on these record discrepancies and their origin

\Rightarrow Integrated basin-scale study

Using sedimentology, paleontology, cartography, GIS spatialization

Dataset:

- **43 selected sections** after literature and field work
- High time-resolution biostratigraphic framework based on ammonoids biozonation (after Brayard *et al.*, 2013)



After Caravaca et al., subm.

Methods

Focus on these record discrepancies and their origin

\Rightarrow Integrated basin-scale study

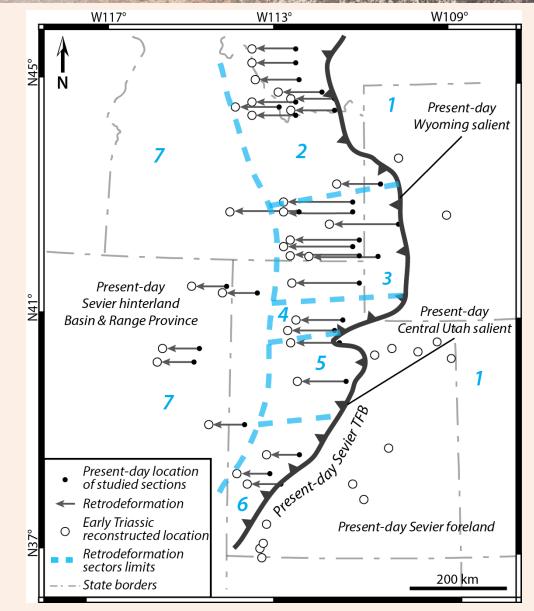
Using sedimentology, paleontology, cartography, GIS spatialization

Dataset:

- 43 selected sections after literature and field work
- High time-resolution biostratigraphic framework (after Brayard *et al.*, 2013)

Palinspastic reconstructions:

- Necessary to obtain an unbiased view of the area during the Early Triassic
- **Original location** of the sections within the basin



After Caravaca et al., subm.

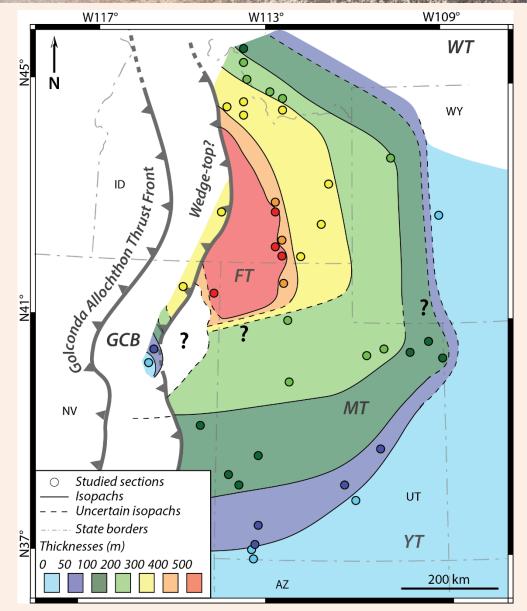
Results: thickness variations in the basin

Isopach map of the PTU-Smithian interval

 \Rightarrow Spatial distribution of sedimentary thickness

2 different areas in the basin:

- Northern high thickness zone (>300 up to ~550m thick)
- **Southern low thickness** zone (~10 up to 250m thick)



Caravaca et al., subm.

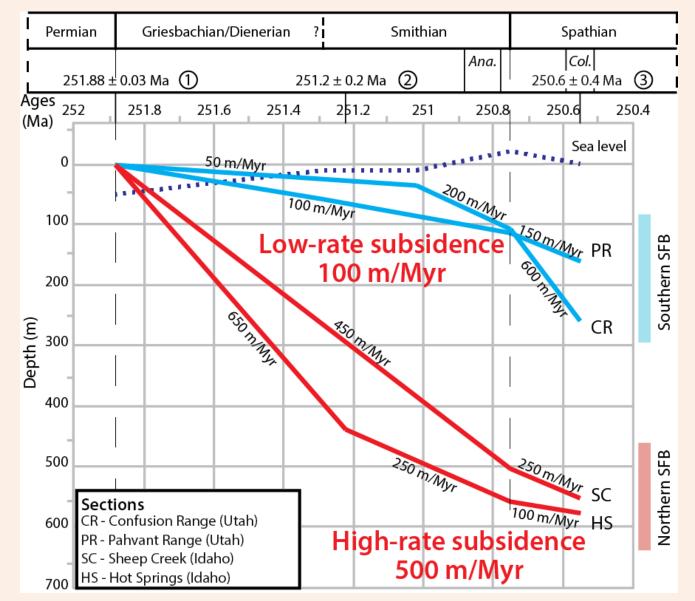
Results: subsidence variations in the basin

Backstripping analysis of the PTU-Smithian interval

 $\Rightarrow \textbf{Tectonic subsidence is preponderant} \\ \Rightarrow \textbf{Shape: characteristic of a foreland basin}$

Differential subsidence in the basin:

- Northern high-rate subsidence
- Southern low-rate subsidence

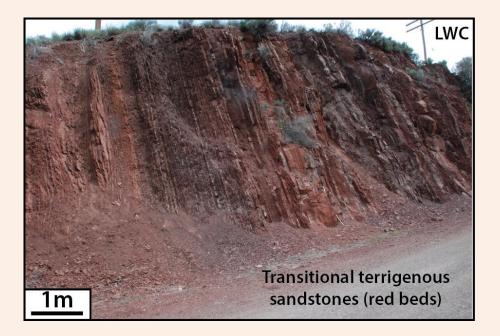


After Caravaca et al., subm.

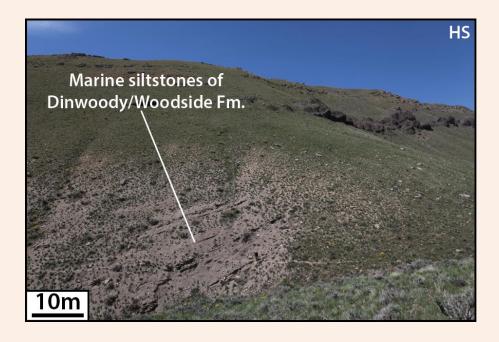
What are the mechanisms controlling this differential subsidence observed in the Early Triassic Sonoma Foreland Basin?

- Sedimentary overload?

d? Sedimentary loading should be in southern SFB ⇒ at variance with our results Northern SFB



Coarse terrigenous of the Moenkopi Group (conglomerates and sandstones) (d~2.5 to 2.8 kg/cm³) Topped by metric beds of microbial limestones (d~2.6 to 2.8 kg/cm³)



Dominated by **fine marine siltstones** of the Woodside and Dinwoody Fm. (d~2.3 to 2.7 kg/cm³)

- Differential topographic loading?
- Golconda Allochthon remains in central Nevada:
- \Rightarrow Presence of coarse conglomerates in West-central SFB
- \Rightarrow Presence of Koipato volcanics in Southwest SFB

Rhyolitic volcanism Syn-volcanic normal-faulting Sealed by Middle Triassic series

 \Rightarrow Characteristic of late-orogenic volcanism Partial melting caused by asthenosphere shallowing during crustal thinning

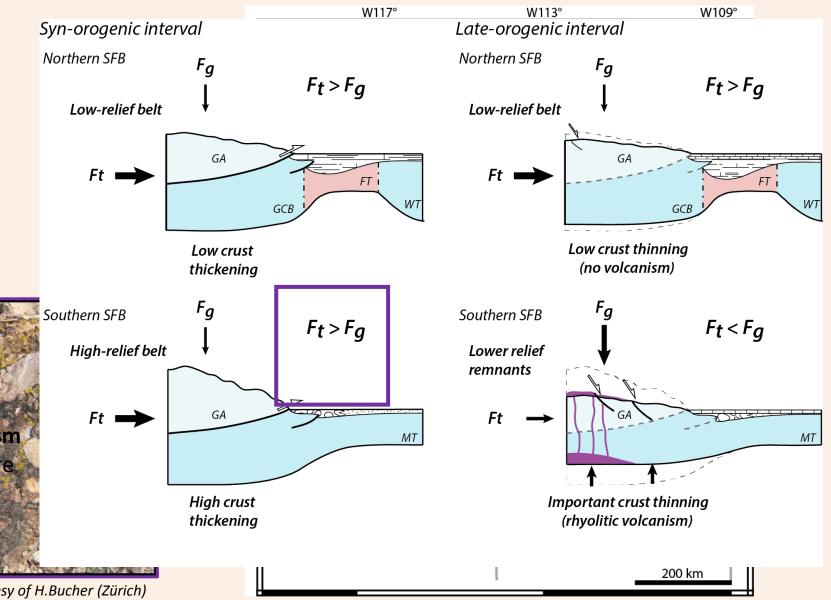


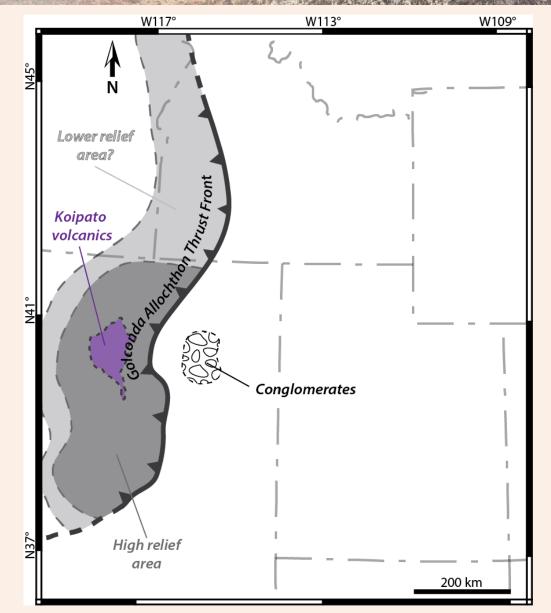
Photo courtesy of H.Bucher (Zürich)

Caravaca et al., subm.

- Differential topographic loading?

Highest relief in South-central Golconda front

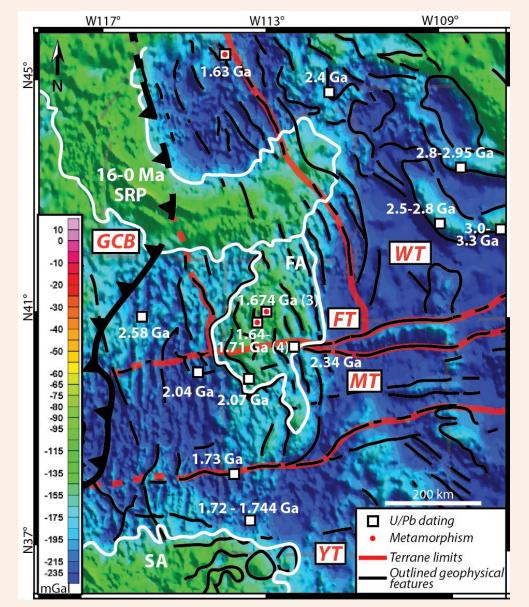
- \Rightarrow Cannot rule out impact of the topographic load
- \Rightarrow Not the main controlling factor



Caravaca et al., subm.

- Differential rheological resistance?
- \Rightarrow New local terrane maps from literature and geophysical data

- Five terranes identified, with different ages:
- Archean Wyoming Terrane (WT)
- Archean Grouse Creek Block (GCB)
- Paleoproterozoic Mojave Terrane (MT)
- Mesoproterozoic Yavapai Terrane (YT)
- Mesoproterozoic Farmington Terrane (FT)



Caravaca et al., subm.

- Differential rheological resistance?

⇒Strong lithospheres

Oldest (>1.7 Ga), coldest and more rigid, resistant to flexuration

- WT, GCB, MT

\Rightarrow Weak lithosphere

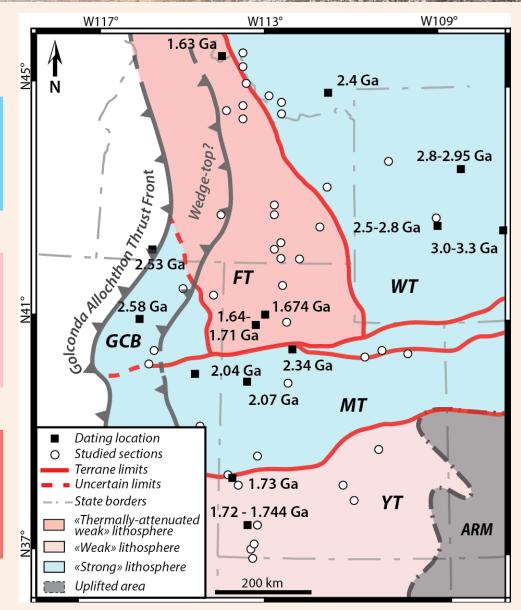
Juvenile (<1.6 Ga), warmer and **less rigid**, more prone to deformation and flexuration

- YT

⇒Thermally-attenuated weak lithosphere

Underwent at least one **intense thermal metamorphism event** (~1,6 Ga)

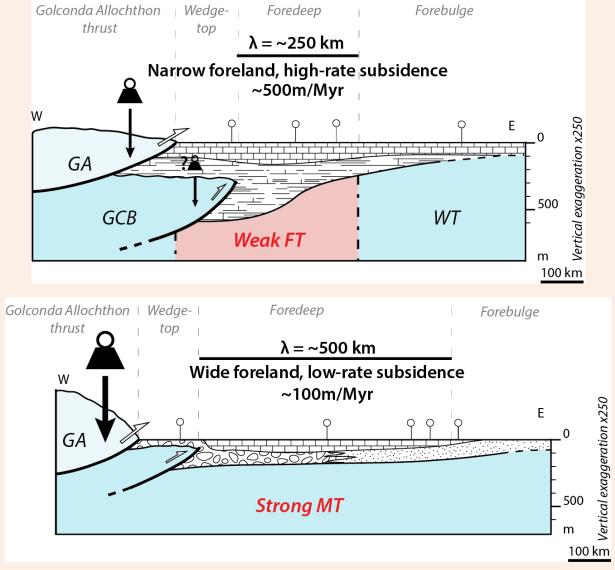
- FT (Mobile belt)



Caravaca et al., subm.

Proposed model



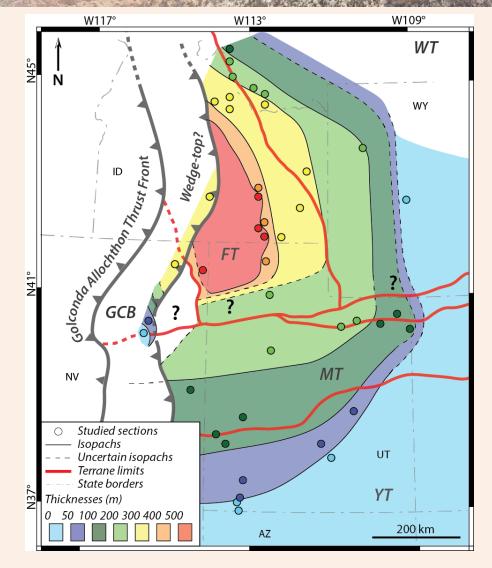


⇒Lowest thicknesses and accommodation rates in southern and eastern SFB above Strong terranes

After Caravaca et al., subm.

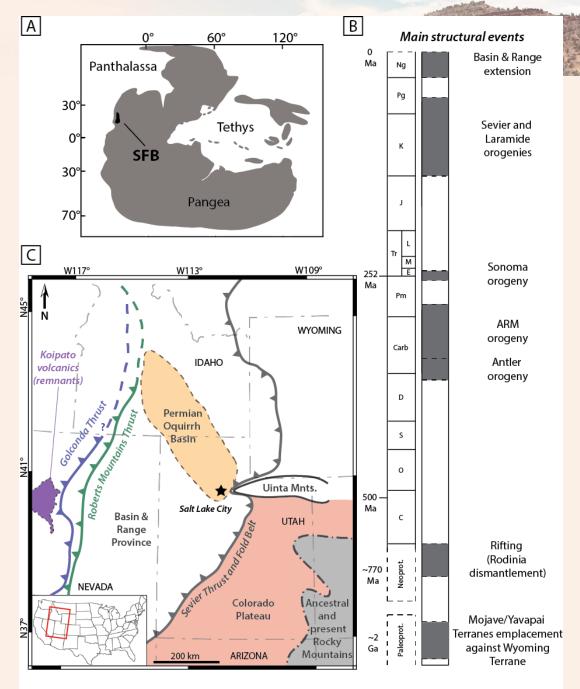
Conclusion

- Early Triassic Sonoma Foreland Basin: two parts with differential tectonic subsidence
- Controlling mechanisms?
 - Negligible sedimentary overloading
 - Weak impact of the topographic loading
 - Rheological behaviour of the basement is likely the main controlling factor
- Lithospheric strength: a major controlling factor on flexuration, and consequently on the differential tectonic subsidence in the basin

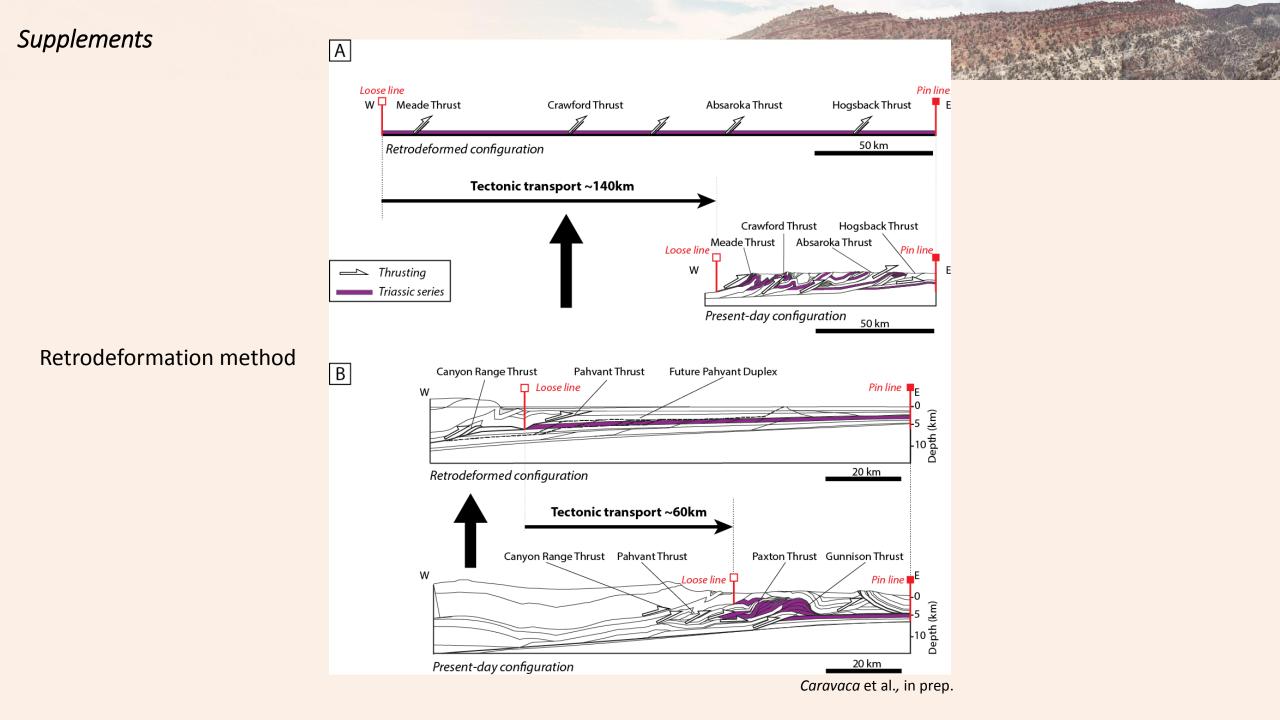


Thank you for paying attention!

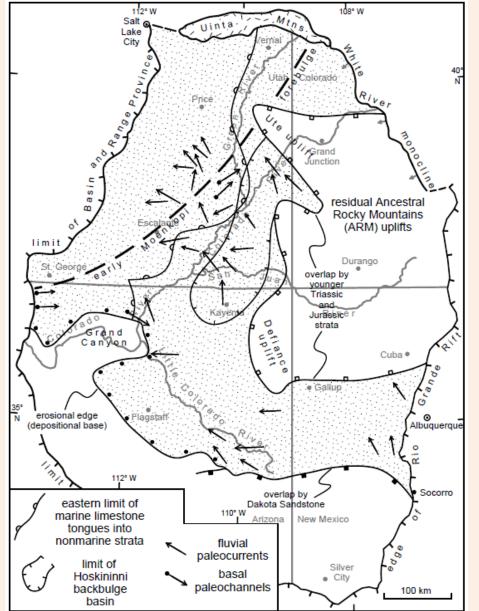




Caravaca_et_al_figure_1



Perspectives: depositional settings



W.R. Dickinson, Pers. comm.

Oldest lithospheres are:

- Thicker and colder than juvenile counterparts (due to R* elements depletion
- More buoyant (d°<d° juvenile lithosphere)
- \Rightarrow More rigid and resistant to flexuration

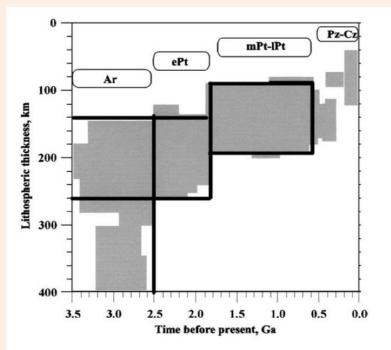


Fig. 2. Lithospheric thermal thickness versus geologic age of the continental lithosphere. The Archean lithosphere has bimodal thickness distribution centered at \sim 350 and \sim 220 km. Gray area shows the lithospheric thickness estimates derived from thermal data (Artemieva and Mooney, 2001). Key: Ar—Archean; ePt, mPt, lPt—early, middle and late Proterozoic, respectively; Pz—Paleozoic; Cz—Cenozoic.

Artemieva & Mooney, 2001

Carpathians, differential subsidence after rheology

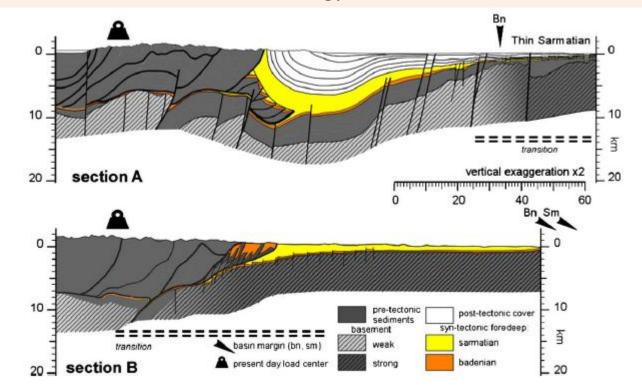


Fig. 11. (a) Badenian (16.4–12.5 Ma) and (b) Sarmatian (12.5–10.5 Ma) isopach maps of the Romanian East Carpathian Foreland-Focşani Depression. (c) cross sections. Section A (after Matenco et al., submitted for publication) crosses the Focsani Depression. Note the large thickness of post orogenic sediments, beyond the scope of this paper. Section B (after section A11 from Stefanescu et al.; Matenco and Bertotti, 2000).

Magallanes basin, differential subsidence after rheology (attenuated lithosphere)

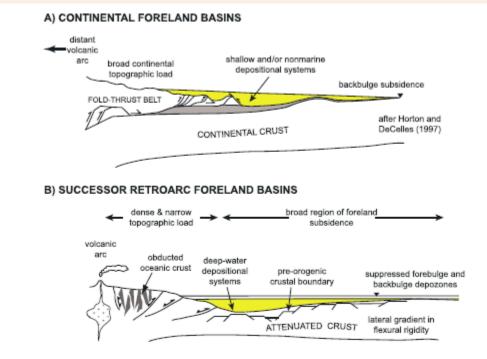


Figure 12. Conceptual cross sections depicting the major structural and stratigraphic differences between foreland basin systems that develop atop (a) continental and cratonal lithosphere (modified from *Horton and DeCelles* [1997]) and (b) attenuated rift and back-arc basin lithosphere in successor retroarc foreland basins (*this study*; with aspects from *Romans et al.* [2011]). In particular, topographic loading from a mafic thrust belt, a lateral gradient in flexural rigidity, and sediment loading of the foredeep promote a broad and deep region of subsidence.

Fosdick et al., 2014

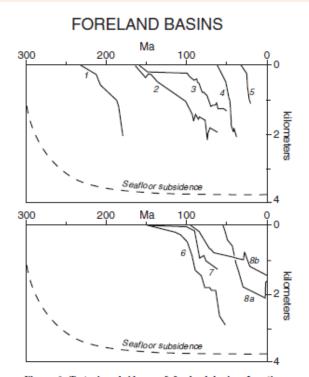
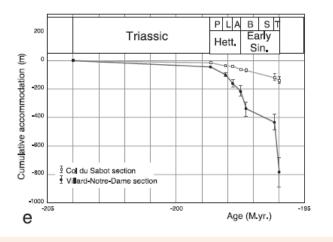
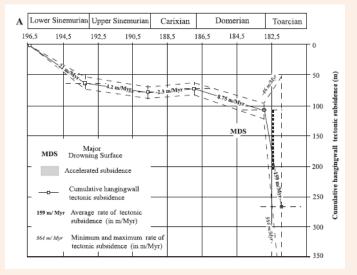


Figure 6. Tectonic subsidence of foreland basins. Locations shown in Figure 1. Thermal decay curve (dashed) for subsidence of cooling seafloor (Stein and Stein, 1992), minus 1500 m, is shown for comparison. 1—Eastern Avalonia, Anglo-Brabant fold belts (van Grootel et al., 1997); 2—Southern Alberta Basin (Gillespie and Heller, 1995); 3—San Rafael Swell, Utah (Heller et al., 1986); 4—Pyrenean foreland basin, Gombrèn (Vergés et al., 1998); 5—Swiss Molasse basin (Burkhard and Sommaruga, 1998) modified from total subsidence using water:sediment density contrast); 6—Hoback Basin, Wyoming (Cross, 1986); 7—Green River Basin, Wyoming (Cross, 1986; Heller et al., 1986); 8—Magallanes Basin (Biddle et al., 1986).



Chevalier et al, 2003



Perspectives: depositional settings

Depositional settings distributions between the *Owenites* and *Anasibirites* ammonoids biozones (middle Smithian)

Striking **differences between southern and northern parts** of the basin:

- ⇒ Differences between paleontological and sedimentological record
- ⇒ Paleogeography to be reconstructed using this integrative set of data
- \Rightarrow Impact of lithospheric control

Depositional settings:

- 📃 Deep subtidal
- Shallow subtidal
- Subtidal to intertidal (microbial carb.)
- Transitional/coastal (red beds)

