



HAL
open science

Consistently dated Atlantic sediment cores over the last 40 thousand years

Claire Waelbroeck, Bryan C. Lougheed, Natalia Vázquez Riveiros, Lise Missiaen, Joel Pedro, Trond Dokken, Irka Hajdas, Lukas Wacker, Peter Abbott, Jean-Pascal Dumoulin, et al.

► To cite this version:

Claire Waelbroeck, Bryan C. Lougheed, Natalia Vázquez Riveiros, Lise Missiaen, Joel Pedro, et al.. Consistently dated Atlantic sediment cores over the last 40 thousand years. *Scientific Data*, 2019, 6 (1), pp.1-12. 10.1038/s41597-019-0173-8. hal-02292112

HAL Id: hal-02292112

<https://hal.science/hal-02292112>

Submitted on 25 May 2020

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution 4.0 International License

OPEN

DATA DESCRIPTOR

Consistently dated Atlantic sediment cores over the last 40 thousand years

Claire Waelbroeck *et al.*[#]

Rapid changes in ocean circulation and climate have been observed in marine-sediment and ice cores over the last glacial period and deglaciation, highlighting the non-linear character of the climate system and underlining the possibility of rapid climate shifts in response to anthropogenic greenhouse gas forcing. To date, these rapid changes in climate and ocean circulation are still not fully explained. One obstacle hindering progress in our understanding of the interactions between past ocean circulation and climate changes is the difficulty of accurately dating marine cores. Here, we present a set of 92 marine sediment cores from the Atlantic Ocean for which we have established age-depth models that are consistent with the Greenland GICC05 ice core chronology, and computed the associated dating uncertainties, using a new deposition modeling technique. This is the first set of consistently dated marine sediment cores enabling paleoclimate scientists to evaluate leads/lags between circulation and climate changes over vast regions of the Atlantic Ocean. Moreover, this data set is of direct use in paleoclimate modeling studies.

Background & Summary

In order to decipher the mechanisms at play in observed past climate changes, it is necessary to establish a common temporal framework for paleoclimate records from different archives and from different locations. Determining the lead/lag relationships between different climatic and circulation changes can help to identify the underlying causes and foster development of conceptual hypotheses to be tested with climate model simulations. Also, paleoclimate data-model integration studies, such as groundtruthing of transient modeling analyses, timeslice comparisons of proxy data, or data assimilation, necessitate consistent paleoclimate records chronologies in calendar years.

Here we focus on the last 40 ky because it is the time span covered by radiocarbon dating and the sole period for which it is possible to establish calendar age timescales for marine cores with a precision approaching that of ice core or speleothem records.

Radiocarbon dating of marine records is complicated, however, by a difference between the surface water $^{14}\text{C}/^{12}\text{C}$ ratio (expressed as $\Delta^{14}\text{C}$, in ‰) and that of the contemporaneous atmosphere, due to the balance between the input of atmospheric ^{14}C and its removal by radioactive decay in the water column, advection, and mixing with older waters. This difference in $\Delta^{14}\text{C}$ is termed the “reservoir age” of the surface waters. Previous studies have revealed that surface reservoir ages have not remained constant over time at high latitudes of the North Atlantic and Southern Ocean (i.e. poleward of $\sim 38^\circ\text{N}$ and of $\sim 40^\circ\text{S}$) due to changes in the location and vigour of deep-water formation^{1–4}.

In those high-latitude regions, it is thus necessary to use an alternative dating strategy in lieu of ^{14}C dating of marine organisms. Here we adopt a strategy that has been widely applied (e.g. refs^{4–7}) and has been adopted by the INTIMATE (Integration of Ice core, Marine and Terrestrial records of the North Atlantic) group when surface reservoir ages can not be assessed⁸. This strategy consists of synchronizing the sea surface temperature (SST) signal recorded in marine cores with the air temperature signal recorded in polar ice cores. This dating approach is based on the observed thermal equilibrium between the ocean’s surface water and overlying air. Previous studies have demonstrated that changes in air and sea surface temperature were synchronous across the last deglaciation⁹.

[#]A full list of authors and their affiliations appears at the end of the paper. Correspondence and requests for materials should be addressed to C.W. (email: claire.waelbroeck@lsce.ipsl.fr)

and some of the last glacial rapid climate changes¹⁰ over the North Atlantic region. Moreover, modeling studies of the last deglaciation¹¹ or last glacial millennial climate changes^{12,13} show that both increases and decreases in North Atlantic (Southern Ocean) SST and in air temperature above Greenland (Antarctica) are synchronous.

Currently, the Greenland NorthGRIP (NGRIP) ice core can be considered the best-dated continuous continental paleoclimatic archive over the last 50 to 75 ky. The NGRIP Greenland Ice Core Chronology 2005 (GICC05) calendar age scale has been established by annual layer counting with estimated uncertainties of 50 y at 11 calendar ky BP (i.e. calendar ky before 1950, noted ka hereafter), 100 to 450 y for the 11–30 ka interval, and 450 to 800 y for 30–40 ka¹⁴ (y or ky referring to durations and ka to dates). Moreover, a common chronology for Greenland and Antarctica ice cores has been developed based on their records of ¹⁰Be and atmospheric CH₄ concentration^{15,16}. This dating effort yielded the Antarctic AICC2012 age scale for four Antarctic ice cores, which is fully consistent with the GICC05 age scale over the last 60 ky¹⁶. Using the GICC05 and AICC2012 age scales as alignment targets for high latitude SST records of the north and south hemispheres respectively, it is thus possible to directly compare marine records from both hemispheres on a common time frame.

Here, we present the first set of consistently dated Atlantic sediment cores from 92 locations distributed between 68°N and 53°S, and between 400 and 5000 m water depth (Fig. 1, Online-only Table 1, ref. 17), together with consistently derived dating uncertainties. This new data set enables paleoclimate scientists to (i) examine relative phases between Atlantic records (e.g. planktonic and benthic oxygen and carbon isotopes, Pa/Th); and (ii) use the spatial and temporal changes recorded in Atlantic sediments to constrain paleoclimate model simulations.

Methods

We compiled existing paleoceanographic data from Atlantic sediment cores covering part of or the entire 0–40 ka interval, with sedimentation rates of at least 5 cm/ky, for which there exists the following dating means: radiocarbon dates for mid and low latitudes sediment cores, and SST or magnetic records for sediment cores located poleward of ~38°N and ~40°S. New cores were added to fill gaps with respect to the available geographical and water depth coverage, and additional radiocarbon dates were produced to improve the existing age models of some cores (Online-only Table 1).

In mid and low latitudes (i.e. between ~40°S and ~38°N), reservoir ages can be assumed not to have strongly varied in response to ocean circulation changes of the last glacial and deglaciation. The same is true at all latitudes during the Holocene. Thus, in mid and low latitudes, and during the Holocene at higher latitudes, the sediment cores were dated by means of calibrated radiocarbon ages. For this, 1427 published and 104 new radiocarbon dates have been calibrated using the Bayesian calibration program “MatCal”¹⁸, and the IntCal13 and SHCal13 calibration curves^{19,20} for North and South Atlantic cores, respectively.

We accounted for both spatial and temporal variability in ¹⁴C reservoir ages. To estimate spatial variations in reservoir ages we extracted bomb-corrected reservoir ages from the GLObal Ocean Data Analysis Project for Carbon (GLODAP) data set²¹. Prior to extracting these surface reservoir ages, GLODAP data were re-gridded to a 4° × 4° grid, whereby the mean and standard deviation for the GLODAP data points from the upper 250 m for each 4° × 4° grid cell were calculated. The modern surface water reservoir age at a given site is then obtained from the nearest grid node to the core site (Fig. 2). In the case of certain sites that are out of range of the GLODAP grid, such as those in the Gulf of Mexico, we have extrapolated the GLODAP 4° × 4° grid to these areas. This spatially varying component of the reservoir age is subtracted from the laboratory ¹⁴C age before calibration (with error propagation). The error used for this spatial reservoir age component is either the computed GLODAP standard deviation, or 100 ¹⁴C yr, whichever is greater. For pre-Holocene dates, a minimum of 200 ¹⁴C yr is used instead of 100 ¹⁴C yr.

To also consider temporal changes in reservoir age, we further applied a correction to account for the impact of atmospheric CO₂ concentration changes upon surface water ¹⁴C activity. At the Last Glacial Maximum (LGM), the lower atmospheric CO₂ concentration induced an increase in atmospheric Δ¹⁴C of ~30‰ due to the speciation change, everything else being equal²². This ~30‰ increase in atmospheric Δ¹⁴C in turn caused a ~250 y increase in surface water reservoir ages²². To account for this temporal change in surface reservoir age, we linearly scaled a reservoir age correction to atmospheric pCO₂, whereby a correction of 0 ¹⁴C y corresponds to present day pCO₂, and 250 ¹⁴C y to LGM pCO₂. For pCO₂ values, we consulted the composite atmospheric CO₂ record of Antarctic ice cores²³. This age-dependent component of the reservoir age is added to the IntCal13 (or SHCal13) ¹⁴C age record before calibration.

Even in regions where surface reservoir ages can be predicted based on the evolution of atmospheric CO₂, as described above, increased uncertainties in radiocarbon-dated chronologies can still arise from bioturbation biases (e.g. ref.²⁴). Thus, in the best cases, when bioturbation biases and local changes in past surface reservoir ages remain limited, sediment core dating uncertainties mainly arise from the conversion of radiocarbon ages into calendar ages. In these cases, uncertainties are less than 150 y for the time interval 0–11 ka, of about 400 y for the 11–30 ka interval, and of 600 to 1100 y for the 30–40 ka interval¹⁹. In all other cases, dating uncertainties are larger.

Almost all our age-depth models of low- and mid-latitude cores (51 out of the 92 cores, see Online-only Table 1) are entirely based on calibrated ¹⁴C ages. In three cores (GeoB3910, MD09-3246 and MD09-3256Q), located on the Brazilian margin in a region under the influence of the Intertropical Convergence Zone, it is possible to take advantage of the simultaneous recording of rainfall increases during Greenland stadial periods in the marine cores and in U-Th dated speleothems from the adjacent continent to improve the marine age models. Rainfall increases are recorded both by XRF-Ti/Ca peaks in the marine cores, and by δ¹⁸O decreases in the speleothems²⁵. By aligning the XRF-Ti/Ca in the marine cores to the speleothem δ¹⁸O, it is possible to improve the precision of the marine age models around 40 ka and to extend them beyond the limit of ¹⁴C dating. Importantly, the speleothem record from El Condor cave²⁶, to which we have aligned the three marine cores, has been shown to be in phase, within dating uncertainties, with the NGRIP air temperature record in the GICC05 age scale^{25,27}. Our alignment of GeoB3910, MD09-3246 and MD09-3256Q marine cores to El Condor speleothem is thus consistent with the NGRIP GICC05 age scale.

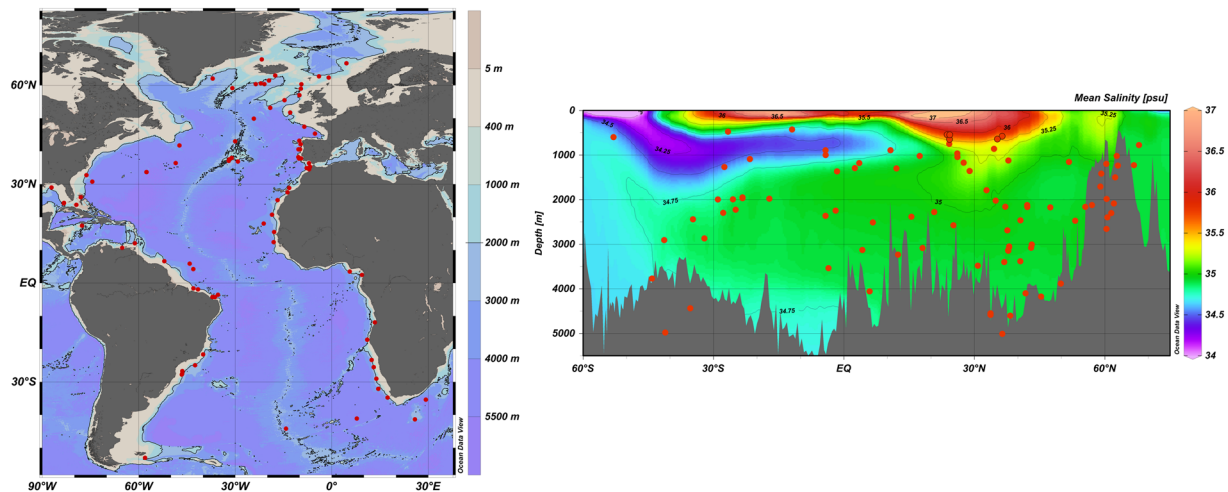


Fig. 1 Location of the 92 dated Atlantic sediment cores (see Online-only Table 1 for precise coordinates and water depths of the cores). The figures were generated using the Ocean Data View software⁵³, the ETOPO bathymetry⁵⁴ (left panel), and the WOA13 mean annual salinity⁵⁵ along a mid-Atlantic north-south section (right panel). The salinity section illustrates the distribution of the cores with respect to the main modern water masses.

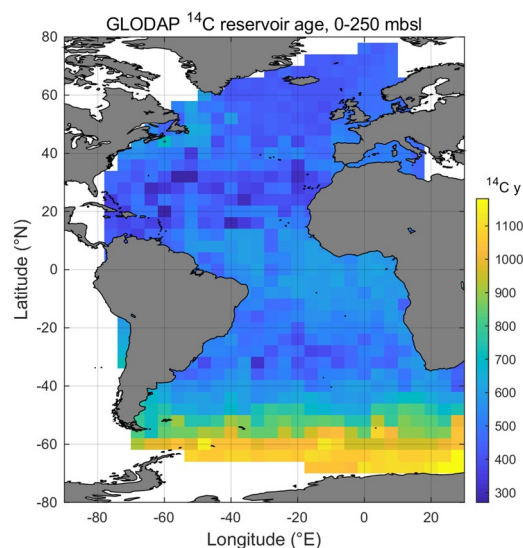


Fig. 2 Average reservoir age extracted from the GLODAP data re-gridded to a $4^\circ \times 4^\circ$ grid and averaged over the upper 250 m of the water column. These values can be downloaded from Figshare⁵⁶.

For cores located north of $\sim 38^\circ\text{N}$ (26 cores) and south of $\sim 40^\circ\text{S}$ (2 cores), and ODP Site 1060 for which there exist no ¹⁴C dates but where planktonic foraminifer census counts exhibit a clear NGRIP signal²⁸, we have used calibrated radiocarbon ages only over the Holocene portion (i.e. after the end of the Younger Dryas, dated at 11.65 ka in the GICC05 age scale²⁹), and aligned their glacial and deglacial portions to NGRIP or EPICA Dronning Maud Land (EDML) air temperature signal. We used different types of chronological markers to derive these 29 age-depth models:

- (1) Tie points defined by aligning high latitude SST records to NGRIP air temperature proxy record on the GICC05 age scale for North Atlantic cores, and to EDML air temperature on the AICC2012 age scale for South Atlantic cores;
- (2) Tie points defined by aligning magnetic properties of northern North Atlantic and Nordic Seas cores to the NGRIP air temperature signal on the GICC05 age scale;
- (3) Dated tephra layers.

The dating procedures (1)–(3) are described in detail below. The alignment procedures (1) and (2) by essence impede the assessment of leads and lags between the aligned records. For instance, leads/lags between SST and polar air temperatures, or among SST records from high latitude marine cores, are by construction not significantly different from zero. In contrast, this dating approach gives access to the relative timing of circulation changes recorded at different water depths in cores located on depth transects.

(1) We aligned SST records to polar ice core air temperature proxy records using the AnalySeries program³⁰. NGRIP alignment targets correspond to the rapid transitions out of and into Greenland stadials, as dated and listed in refs^{29,31} (Online-only Table 2). Tie points were generally defined by aligning rapid warmings recognized in both the ice core and marine core, as recommended in ref.⁸. In rare cases, rapid and well-defined coolings have been aligned. In a few cases, when the SST record resolution was too low or the signal shape ambiguous, maxima or minima have been aligned. Remaining ambiguities in the identification of alignment tie points were solved in most cases by fulfilling the condition that the tie point age is younger or equal to the calibrated ¹⁴C ages obtained by assuming no other change in surface reservoir age than the temporal evolution due to changing atmospheric pCO₂. Not fulfilling this condition would result in negative surface reservoir ages, which is not physically possible (see Supplementary Fig. 1 for an example).

SST alignment to Antarctic temperature variations was made at marked transitions in the temperature record, such as Antarctic Isotopic Maxima³², the onset of the early and late deglacial warming, or the beginning of the Antarctic Cold Reversal.

In addition, we used the following three alignment targets in the North Atlantic:

- (i) A first alignment target is based on the observation that the cooling marking the beginning of Heinrich Stadial 1 in three independently dated North Atlantic cores is synchronous with the sharp increase in dust flux recorded in the Greenland ice cores and dated at $17.48 \text{ ka} \pm 0.21 \text{ ky}$ on the GICC05 age scale³³. This observation is consistent with this cooling being coeval with an increase in dust transport from Asia to Greenland, as observed during other Greenland stadials³⁴.
- (ii) Two other alignment targets correspond to the beginning and the end of the warm event identified in ref.³⁵ within Greenland stadial 3 (GS-3) in several North Atlantic cores between 24 and 25 ka. This warm event within GS-3 is not clearly recorded in Greenland ice ($\delta^{18}\text{O}$) or gas ($\delta^{15}\text{N}$) isotopic records, but corresponds to a marked decrease in dust flux. Here again, we aligned the beginning and end of the warm event to the corresponding changes in the NGRIP dust flux dated on the GICC05 age scale at $25.05 \text{ ka} \pm 0.35 \text{ ky}$ and $24.1 \text{ ka} \pm 0.33 \text{ ky}$, respectively.

For consistency, the alignment tie points in high latitude cores were all defined by the same person. Similarly, one single person defined all the alignment tie points in the three Brazilian cores. Also, the SST records used in the present study are all based on planktonic foraminifer census count data. When SST reconstructions based on full census count data were not available, we used the percentage of the polar species *Neogloboquadrina pachyderma* (left coiling) as a proxy for SST. This approach has been described and validated in a number of studies (e.g. refs^{36–39}). In two North Atlantic cores (ODP Site 1060 and core MD08-3180Q) we used the percentage of warm species instead, because the percentage of *N. pachyderma* was too low. In the particular case of the Iberian margin, it has been shown that *Globigerina bulloides* $\delta^{18}\text{O}$ co-varies with SST^{40,41} and we have used *G. bulloides* $\delta^{18}\text{O}$ as a proxy for SST when no SST estimates were available.

Both age and depth uncertainties are defined for each tie point. The depth uncertainty directly depends on the sampling resolution of the SST curve: it is taken as half of the depth interval corresponding to the rapid warming (or more rarely cooling), or as half of the width of the SST maximum or minimum, when maxima or minima have been aligned. In instances of ambiguities that could not be tested by the constraints provided by ¹⁴C dates, we attributed an uncertainty to the depth of the tie point, large enough to encompass the two events (warmings, or more rarely, coolings or SST maxima or minima) which could both be aligned to the same target. The uncertainty on the tie point ages is the GICC05 dating precision of the transitions between Greenland stadials and interstadials, with one sigma uncertainties defined as half the cumulative ‘maximum counting error’ in the GICC05 age scale^{29,31}. Similarly, the dating uncertainty of the alignment tie points defined with respect to AICC2012 is the dating error given in ref.¹⁶.

(2) In high northern latitudes, when SST records are not available, for some cores it is possible to instead use high-frequency variations in magnetic susceptibility (MS) recorded during the last glacial period. The rapid oscillations in magnetic properties in sediment cores on the flow path of North Atlantic Deep Water (NADW) in the Nordic Seas and North Atlantic have indeed been shown to be in phase with the Greenland ice $\delta^{18}\text{O}$ or air temperature signal⁴². Support for this synchronicity comes from tephra and geomagnetic field (Laschamp inclination excursion) marine records. These marine records become aligned with tephra and cosmogenic nuclide Greenland records when the MS tuning to Greenland is applied (e.g. refs^{43–45}).

We dated five cores located north of 62°N by aligning their MS records to the NGRIP ice $\delta^{18}\text{O}$ signal (Online-only Table 1). MS tie points and their associated uncertainties were defined using the same method as described for the alignment of SST signals to ice core records. The MS records of four of these five cores have been previously shown to be in phase with the Greenland air temperature signal⁴². More recently, the identification of tephra layers in core MD99-2284 demonstrated that this core’s MS record is also in phase with the NGRIP $\delta^{18}\text{O}$ record⁴³. This can be explained by the fact that changes in MS arise from changes in the efficiency of the transport of fine grained magnetic particles by deep currents from the source to the site of deposition⁴². The fact that the MS signal is in phase in cores located north and south of the sills separating the Nordic Seas from the North Atlantic basin, suggests that the source of magnetic minerals could be at the sills, with the strength of the overflow from the Nordic Seas directly proportional to the strength of the inflow into the Nordic Seas.

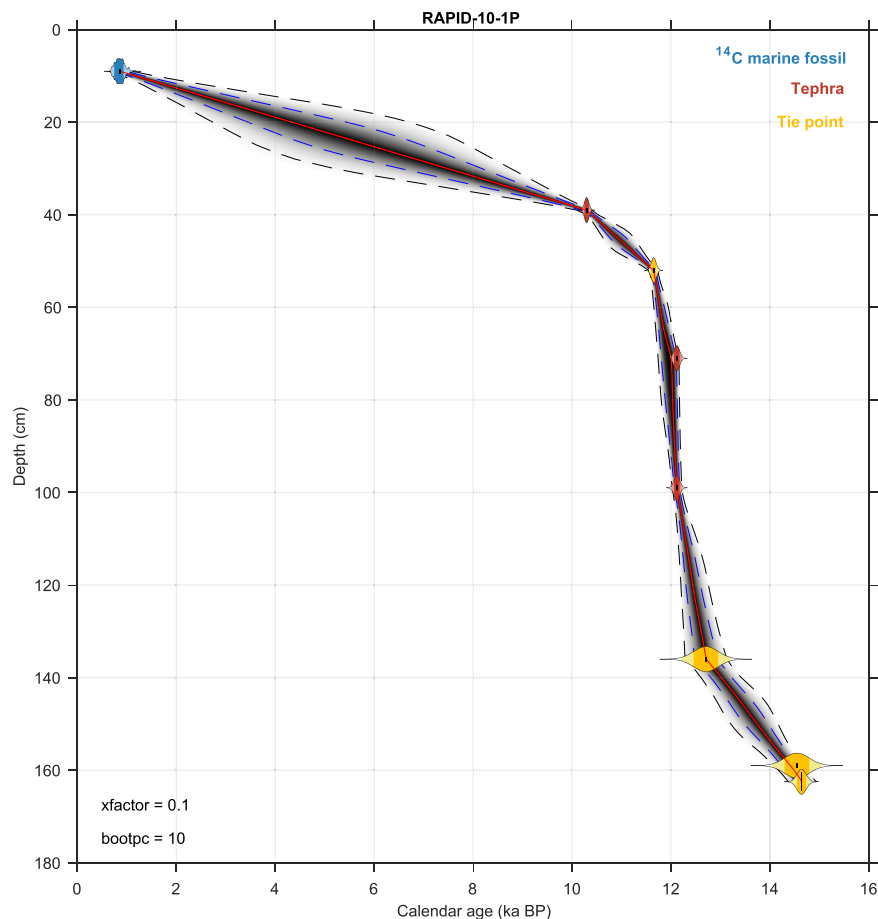


Fig. 3 Example of age-depth plot produced by *Undatable*. Age-depth model produced for North Atlantic core RAPID-10-1P with bootstrapping set to 10% and sedimentation rate uncertainty set to 0.1 (see ref.⁴⁹ for details). Blue, yellow and red probability density functions indicate the radiocarbon and alignment tie points, and tephra age-depth constraints, respectively. The grey cloud indicates the probability density cloud of the age-depth model, whereby darker colors indicate higher age-depth probability. The blue and black broken lines represent 68.27% and 95.45% confidence intervals, respectively. The red line indicates the age-depth model median.

(3) We used dated tephra layers as additional chronological markers over the last 55 ky in 10 of the northernmost cores (Online-only Table 1). The following four tephra layers have been recognized both in Greenland ice cores and in certain North Atlantic and Nordic Seas marine cores: the Saksunarvatn Ash⁴⁶, the Vedde Ash⁴⁶, the Faroe Marine Ash Zone (FMAZ) II^{46,47}, and the widespread rhyolitic component of North Atlantic Ash Zone (NAAZ) II (II-RHY-1)^{46,48} (Online-only Table 2).

Age-depth relationships were built for each core accounting for both the age and depth uncertainties of the ^{14}C dates and chronological markers, using the age-depth modeling routine “*Undatable*”⁴⁹ (Fig. 3). This new rapid age-depth modeling routine was ideal for this project as it allowed us to run and re-run age models for the many sediment cores that we have analyzed. Moreover, this age-depth modeling routine computes a conservative age-depth uncertainty, through the provision of bootstrapping and sediment accumulation rate uncertainty⁴⁹ (Fig. 3). Default values for bootstrapping percentage and sedimentation rate uncertainty were set to 10% and 0.1 respectively. In the presence of age reversals, we progressively increased the bootstrapping percentage in order to make sure that the dating uncertainty computed by *Undatable* was large enough to encompass most calibrated ^{14}C ages, leaving out only outliers beyond 2 sigma dating uncertainty. This way, we take into account increased dating uncertainty associated with the existence of age-depth scatter, which may be related to sedimentation hiatuses, abundance changes, or bioturbation. Also, we considered tephra layers as the most reliable age-depth constraints and, therefore, *a priori* excluded them from the bootstrapping process (e.g. ref.⁵⁰).

In some North Atlantic cores (7 out of 92, cf. Online-only Table 1), we used ^{14}C dates together with SST alignment tie points to NGRIP. These cores are located at the northern edge of the region where surface reservoir ages may be assumed not to have strongly varied in response to ocean circulation changes, and are characterized by large changes in SST which parallel the NGRIP ice $\delta^{18}\text{O}$ signal. In those cores, we used alignment tie points to complement calibrated ^{14}C dates when the latter were too sparse.

Finally, although the focus of this work is the time interval 0–40 ka, we used dating information available beyond 40 ka to ensure the robustness of the computed sedimentation rate and age-depth relationship around 40 ka.

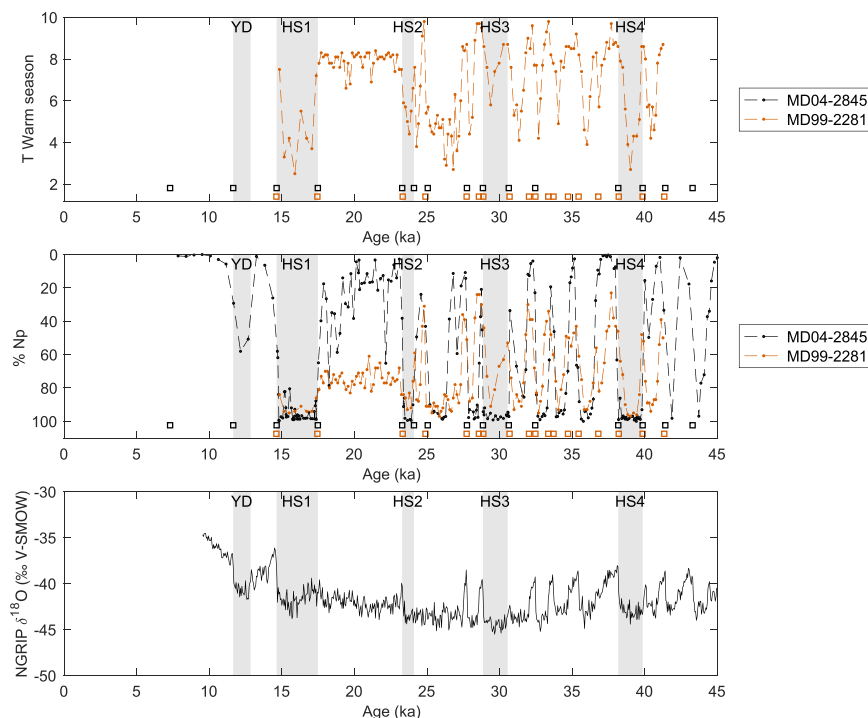


Fig. 4 Example of North Atlantic and Nordic Seas cores dated by alignment of their SST records to the NGRIP ice $\delta^{18}\text{O}$ signal. Top panel: planktic foraminifer-based warm season surface temperature of core MD99-2281^{57,58}; middle panel: % *N. pachyderma* of core MD99-2281 and MD04-2845^{59,60} (both panels: diamonds and squares above the x-axis indicate calibrated ^{14}C ages and alignment tie points, respectively). Bottom panel: NGRIP ice $\delta^{18}\text{O}$ record on the GICC05 age scale⁶¹. Grey bands highlight the Younger Dryas and Heinrich stadials 1–4 chronozones as defined in Online-only Table 2.

Data Records

The present set of age-depth models contains three text files per marine sediment core¹⁷. The first text file (“age depth input”) contains an overview of the ^{14}C ages and other age constraints used in the age-depth model. More specifically, a first section provides all the available ^{14}C raw data, the reservoir age and calibration curve used, as well as the calibrated ages together with the 68.3% highest posterior density interval(s), and specifies which ^{14}C dates have been used to generate the age-depth model. A second section provides the definition of the alignment tie points: the tie points depth and its uncertainty, the tie points age and its uncertainty, the nature of the tie points and the nature of the uncertainty of the tie points age. The second text file (“udinput”) contains the input for the age-depth model in the *Undatable* format. The third text file (“_admodel_ka”) contains the computed age-depth relationship and associated dating uncertainties. In addition to the complete set of data records archived on Seano¹⁷, the 92 “_admodel_ka” text files can be found on Pangaea⁵¹.

Notably, the fact that the ^{14}C raw data are provided makes the present data set easy to update using a future ^{14}C calibration curve. Also, tie point depths are provided, allowing updates of the age-depth models if higher resolution SST records are produced.

In addition to these three text files, the age-depth model plot produced by the *Undatable* routine (see Fig. 3 for an example) is provided for each core, as well as a plot of the aligned SST or MS record, ice core record, and chosen alignment tie points (see Fig. 4 for an example) for the cores which have been partially or completely dated by alignment to an ice core record¹⁷.

Technical Validation

The information relative to the validation of the age models entirely based on ^{14}C dates can be found in the publications describing the *Undatable* age-depth modeling routine⁴⁹ and the “MatCal” Bayesian calibration program¹⁸. The age-depth model plot provided for each core (e.g. Fig. 3) shows the calibrated ^{14}C dates together with the computed age-depth relationship and dating uncertainty, as well as the bootstrapping percentage and sedimentation rate uncertainty values used in the computation.

Concerning the age models based on the alignment of SST to NGRIP air temperature, a first validation step involved comparing the resulting dated SST signals of different marine cores among themselves and with the NGRIP air temperature signal. An illustration of such a comparison is given in Fig. 4 for North Atlantic core MD04-2845 and Norwegian Sea core MD99-2881. Moreover, available ^{14}C data over glacial and deglacial portions of cores dated by alignment to NGRIP provide a verification of the tie points selection since surface water reservoir ages should not be negative (e.g. Supplementary Fig. 1). Interestingly, around Heinrich stadial 4 (38 to 40 ka), our age-depth models yield ages which are systematically older than the calibrated ages obtained using

IntCal13 and modern surface water reservoir age values, in agreement with the recent findings of ref.⁵² showing that IntCal13 is too young with respect to GICC05 during that time interval.

The age models based on the alignment of MS to NGRIP ice $\delta^{18}\text{O}$ have been validated by comparing the resulting dated MS signals with the NGRIP ice $\delta^{18}\text{O}$ signal (Supplementary Fig. 2). Moreover, these age models have been validated by climate-independent tie points, such as tephra layers in core MD99-2284⁴³, MD95-2010⁴⁴ and ENAM93-21⁴⁵, or changes in the Earth's magnetic field intensity in core MD99-2281⁴⁴.

The age models making use of the alignment of Ti/Ca to speleothem isotopic records have been validated by comparing the radiocarbon-dated upper portion of the cores with the U-Th dated speleothem signal (Supplementary Fig. 3). This validation was the initial step that led to the use of speleothem isotopic records to complement the dating of the three cores from the Brazilian margin since it demonstrates that terrigenous input at these sites is coeval with the precipitation events recorded in the speleothems²⁵.

Code Availability

The *Undatable* software was used to create age-depth models based on the age-depth constraints given in the “udinput” text files for each core. The software and accompanying source code can be downloaded from the Zenodo public archive (<https://doi.org/10.5281/zenodo.2527642>).

References

- Bard, E. *et al.* The North Atlantic atmosphere-sea surface ^{14}C gradient during the Younger Dryas climatic event. *Earth and Planetary Science Letters* **126**, 275–287 (1994).
- Sikes, E. L., Samson, C. R., Guilderson, T. P. & Howard, W. R. Old radiocarbon ages in the southwest Pacific Ocean during the last glacial period and deglaciation. *Nature* **405**, 555–559 (2000).
- Waelbroeck, C. *et al.* The timing of the last deglaciation in North Atlantic climate records. *Nature* **412**, 724–727 (2001).
- Skinner, L. C., Fallon, S., Waelbroeck, C., Michel, E. & Barker, S. Ventilation of the deep Southern Ocean and deglacial CO_2 rise. *Science* **328**, 1147–1151 (2010).
- Govin, A. *et al.* Evidence for northward expansion of Antarctic Bottom water mass in the Southern Ocean during the last glacial inception. *Paleoceanography* **24**, <https://doi.org/10.1029/2008PA001603> (2009).
- Vazquez Riveiros, N. *et al.* Response of South Atlantic deep waters to deglacial warming during Terminations V and I. *Earth and Planetary Science Letters* **298**, 323–333 (2010).
- Waelbroeck, C. *et al.* The timing of deglacial circulation changes in the Atlantic. *Paleoceanography* **26**, <https://doi.org/10.1029/2010PA002007> (2011).
- Austin, W. E. N. & Hibbert, F. D. Tracing time in the ocean: a brief review of chronological constraints (60–8 kyr) on North Atlantic marine event-based stratigraphies. *Quat. Sci. Rev.* **36**, 28–37 (2012).
- Björck, S. *et al.* Synchronized Terrestrial-Atmospheric deglacial records around the North Atlantic. *Science* **274**, 1155–1160 (1996).
- Austin, W. E. N., Wilson, L. J. & Hunt, J. B. The age and chronostratigraphical significance of North Atlantic Ash Zone II. *Journal of Quaternary Science* **19**, 137–146 (2004).
- Liu, Z. *et al.* Transient simulation of last deglaciation with a new mechanism for Bölling-Allerød warming. *Science* **325**, 310–314 (2009).
- Menviel, L., Spence, P. & England, M. Contribution of enhanced Antarctic Bottom Water formation to Antarctic warm events and millennial-scale atmospheric CO_2 increase. *Earth and Planetary Science Letters* **413**, 37–50 (2015).
- Pedro, J. B. *et al.* Beyond the bipolar seesaw: Toward a process understanding of interhemispheric coupling. *Quaternary Science Reviews* **192**, 27–46 (2018).
- Svensson, A. *et al.* A 60 000 year Greenland stratigraphic ice core chronology. *Climate of the Past* **4**, 47–57 (2008).
- Lemieux-Dudon, B. *et al.* Consistent dating for Antarctic and Greenland ice cores. *Quat. Sci. Rev.* **29**, 8–20 (2010).
- Veres, D. *et al.* The Antarctic ice core chronology (AICC2012): an optimized multi-parameter and multi-site dating approach for the last 120 thousand years. *Climate of the Past* **9**, 1733–1748 (2013).
- Waelbroeck, C. *et al.* Consistently dated Atlantic sediment cores over the last 40 thousand years. *SEANO*. <https://doi.org/10.17882/59554> (2019).
- Lougheed, B. & Obrochta, S. MatCal: Open Source Bayesian ^{14}C Age Calibration in Matlab. *Journal of Open Research Software* **4** (2016).
- Reimer, P. *et al.* IntCal13 and Marine13 radiocarbon age calibration curves 0–50,000 years cal BP. *Radiocarbon* **55**, 1869–1887 (2013).
- Hogg, A. G. *et al.* SHCal13 Southern Hemisphere calibration, 0–50,000 years cal BP. *Radiocarbon* **55**, 1889–1903 (2013).
- Key, R. M. *et al.* A global ocean carbon climatology: Results from Global Data Analysis Project (GLODAP). *Global biogeochemical cycles* **18**, 1–23 (2004).
- Galbraith, E. D., Kwon, E. Y., Bianchi, D., Hain, M. P. & Sarmiento, J. L. The impact of atmospheric pCO_2 on carbon isotope ratios of the atmosphere and ocean. *Global Biogeochemical Cycles* **29**, 307–324 (2015).
- Bereiter, B. *et al.* Revision of the EPICA Dome C CO_2 record from 800 to 600 kyr before present. *Geophysical Research Letters* **42**, 542–549 (2015).
- Lougheed, B. C., Metcalfe, B., Ninnemann, U. S. & Wacker, L. Moving beyond the age-depth model paradigm in deep-sea palaeoclimate archives: dual radiocarbon and stable isotope analysis on single foraminifera. *Climate of the Past* **14**, 515–526 (2018).
- Burckel, P. *et al.* Atlantic Ocean circulation changes preceded millennial tropical South America rainfall events during the last glacial. *Geophys. Res. Lett.* **42**, 411–418 (2015).
- Cheng, H. *et al.* Climate change patterns in Amazonia and biodiversity. *Nature communications* **4**, 1411–1417, <https://doi.org/10.1038/ncomms2415> (2013).
- Adolphi, F. *et al.* Connecting the Greenland ice-core and U/Th timescales via cosmogenic radionuclides: testing the synchronicity of Dansgaard-Oeschger events. *Clim. Past* **14**, 1755–1781 (2018).
- Vautravers, M. J., Shackleton, N. J., Lopez-Martinez, C. & Grimalt, J. O. Gulf Stream variability during marine isotope stage 3. *Paleoceanography* **19**, <https://doi.org/10.1029/2003PA000966> (2004).
- Wolff, E. W., Chappellaz, J., Blunier, T., Rasmussen, S. O. & Svensson, A. Millennial-scale variability during the last glacial: The ice core record. *Quaternary Science Reviews* **29**, 2828–2838 (2010).
- Paillard, D., Labeyrie, L. & Yiou, P. Macintosh program performs time-series analysis. *EOS* **77**, 379 (1996).
- Rasmussen, S. O. *et al.* A stratigraphic framework for abrupt climatic changes during the Last Glacial period based on three synchronized Greenland ice-core records: refining and extending the INTIMATE event stratigraphy. *Quaternary Science Reviews* **106**, 14–28 (2014).
- Epica, C. M. One-to-one coupling of glacial climate variability in Greenland and Antarctica. *Nature* **444**, 195–198 (2006).
- Missiaen, L. *et al.* Improving North Atlantic marine core chronologies using ^{230}Th -normalization. *Paleoceanography and Paleoclimatology*, <https://doi.org/10.1029/2018pa003444> (2019).

34. Ruth, U. *et al.* Ice core evidence for a very tight link between North Atlantic and east Asian glacial climate. *Geophysical Research Letters* **34** (2007).
35. Austin, W. *et al.* The synchronization of palaeoclimatic events in the North Atlantic region during Greenland Stadial 3 (ca 27.5 to 23.3 kyr b2k). *Quaternary Science Reviews* **36**, 154–163 (2012).
36. Bé, A. W. H. & Tolderlund, D. S. In *Micropaleontology of the oceans*, 105–149 (Cambridge University Press, 1971).
37. Knutz, P. C. *et al.* Multidecadal ocean variability and NW European ice sheet surges during the last deglaciation. *Geochemistry, Geophysics, Geosystems* **3**, 1–9 (2002).
38. Govin, A. *et al.* Persistent influence of ice sheet melting on high northern latitude climate during the early Last Interglacial. *Climate of the Past* **8**, 483–507 (2012).
39. Peck, V. L. *et al.* High resolution evidence for linkages between NW European ice sheet instability and Atlantic Meridional Overturning Circulation. *Earth and Planetary Science Letters* **243**, 476–488 (2006).
40. Voelker, A. H. & de Abreu, L. A review of abrupt climate change events in the Northeastern Atlantic Ocean (Iberian Margin): Latitudinal, longitudinal, and vertical gradients. *Geophysical Monograph Series* **193**, 15–37 (2011).
41. Landais, A., Waelbroeck, C. & Masson-Delmotte, V. On the limits of Antarctic and marine climate records synchronization: lag estimates during marine isotopic stages 5d and 5c. *Paleoceanography* **21**, <https://doi.org/10.1029/2005PA001171> (2006).
42. Kissel, C. *et al.* Rapid climatic variations during marine isotopic stage 3: magnetic analysis of sediments from Nordic Seas and North Atlantic. *Earth and Planetary Science Letters* **171**, 489–502 (1999).
43. Dokken, T. M., Nisancioglu, K. H., Li, C., Battisti, D. S. & Kissel, C. Dansgaard-Oeschger cycles: Interactions between ocean and sea ice intrinsic to the Nordic seas. *Paleoceanography* **28**, 491–502 (2013).
44. Wary, M. *et al.* Regional seesaw between the North Atlantic and Nordic Seas during the last glacial abrupt climate events. *Climate of the Past* **13**, 729–739 (2017).
45. Hoff, U., Rasmussen, T. L., Stein, R., Ezat, M. M. & Fahl, K. Sea ice and millennial-scale climate variability in the Nordic seas 90 kyr ago to present. *Nature communications* **7**, 12247 (2016).
46. Blockley, S. P. *et al.* Tephrochronology and the extended intimate (integration of ice-core, marine and terrestrial records) event stratigraphy 8–128 ka b2k. *Quaternary Science Reviews* **106**, 88–100 (2014).
47. Abbott, P. M. & Davies, S. M. Volcanism and the Greenland ice-cores: the tephra record. *Earth-Science Reviews* **115**, 173–191 (2012).
48. Austin, W. E. & Abbott, P. M. Comment: Were last glacial climate events simultaneous between Greenland and France? A quantitative comparison using non-tuned chronologies. M. Blaauw, B. Wohlfarth, J. A. Christen, L. Ampel, D. Veres, K. Hughen, F. Preusser and A. Svensson (2009). *Journal of Quaternary Science* **25**, 1045–1046 (2010).
49. Loughheed, B. C. & Obrochta, S. A Rapid, Deterministic Age-Depth Modeling Routine for Geological Sequences With Inherent Depth Uncertainty. *Paleoceanography and Paleoclimatology* **34**, 122–133 (2019).
50. Obrochta, S. *et al.* Mt. Fuji Holocene eruption history reconstructed from proximal lake sediments and high-density radiocarbon dating. *Quaternary Science Reviews* **200**, 395–405 (2018).
51. Waelbroeck, C. *et al.* Consistently dated Atlantic sediment cores. PANGAEA. <https://doi.org/10.1594/PANGAEA.900073> (2019).
52. Cheng, H. *et al.* Atmospheric 14C/12C changes during the last glacial period from Hulu Cave. *Science* **362**, 1293–1297 (2018).
53. Schlitzer, R. Ocean Data View, <http://odv.awi.de> (2007).
54. Amante, C. & Eakins, B. ETOPO 1 Arc-Minute Global Relief Model: Procedure. *Data Source and Analysis, National Oceanic and Atmospheric Administration NOAA* (2009).
55. Locarnini, R. A. *et al.* World ocean atlas 2013. Volume 1, Temperature (2013).
56. Waelbroeck, C. *et al.* Gridded GLODAP surface mean reservoir age. *figshare*. <https://doi.org/10.6084/m9.figshare.c.4563890> (2019).
57. Zumaque, J. *et al.* An ocean–ice coupled response during the last glacial: a view from a marine isotopic stage 3 record south of the Faeroe Shetland Gateway. *Climate of the Past* **8**, 1997–2017 (2012).
58. Wary, M. *et al.* Stratification of surface waters during the last glacial millennial climatic events: a key factor in subsurface and deep-water mass dynamics. *Climate of the Past* **11**, 1507–1525 (2015).
59. Sanchez Goni, M. F. *et al.* Contrasting impacts of Dansgaard–Oeschger events over a western European latitudinal transect modulated by orbital parameters. *Quaternary Science Reviews* **27**, 1136–1151 (2008).
60. Daniiau, A.-L., Goñi, M. F. S. & Duprat, J. Last glacial fire regime variability in western France inferred from microcharcoal preserved in core MD04-2845, Bay of Biscay. *Quaternary Research* **71**, 385–396 (2009).
61. Seierstad, I. K. *et al.* Consistently dated records from the Greenland GRIP, GISP2 and NGRIP ice cores for the past 104 ka reveal regional millennial-scale $\delta^{18}\text{O}$ gradients with possible Heinrich event imprint. *Quaternary Science Reviews* **106**, 29–46 (2014).
62. Voelker, A. H. L. Dansgaard-Oeschger events in ultra-high resolution sediment records from the Nordic Seas, PhD thesis, Kiel University, Germany (1999).
63. Voelker, A. H. & Hafliðason, H. Refining the Icelandic tephrochronology of the last glacial period—the deep-sea core PS2644 record from the southern Greenland Sea. *Global and Planetary Change* **131**, 35–62 (2015).
64. Dokken, T. M. & Jansen, E. Rapid changes in the mechanism of ocean convection during the last glacial period. *Nature* **401**, 458–461 (1999).
65. Abbott, P. M., Griggs, A. J., Bourne, A. J., Chapman, M. R. & Davies, S. M. Tracing marine cryptotephra in the North Atlantic during the last glacial period: Improving the North Atlantic marine tephrostratigraphic framework. *Quaternary Science Reviews* **189**, 169–186 (2018).
66. Thornalley, D. J. R., Elderfield, H. & McCave, I. N. Intermediate and deep water paleoceanography of the northern North Atlantic over the past 21,000 years. *Paleoceanography* **25**, <https://doi.org/10.1029/2009PA001833> (2010).
67. Rasmussen, T. L., van Weering, T. C. E. & Labeyrie, L. High resolution stratigraphy of the Faeroe-Shetland channel and its relation to North Atlantic paleoceanography: the last 87 kyr. *Marine Geology* **131**, 75–88 (1996).
68. Rasmussen, T. L., Thomsen, E. & van Weering, T. C. E. In *Geological processes on continental margins: Sedimentation, mass-wasting and stability* Vol. 129 (eds Stoker, M. S., Evans, D. & Cramp, A.) 255–267 (Geological Society, Special Publications, 1998).
69. Sadatzki, H. *et al.* Sea ice variability in the southern Norwegian Sea during glacial Dansgaard-Oeschger climate cycles. *Science advances* **5**, eaau6174 (2019).
70. Elliot, M., Labeyrie, L. & Duplessy, J. C. Changes in North-Atlantic deep-water formation associated with the Dansgaard-Oeschger temperature oscillations (60–10 ka). *Quaternary Science Reviews* **21**, 1153–1165 (2002).
71. Manthé, S. *Variabilité de la circulation thermohaline glaciaire et interglaciaire en Atlantique Nord tracée par les foraminifères planctoniques et la microfaune benthique*, PhD thesis, University of Bordeaux I, France (1998).
72. Raymo, M. E. *et al.* Stability of North Atlantic water masses in face of pronounced climate variability during the Pleistocene. *Paleoceanography* **19**, <https://doi.org/10.1029/2003PA000921> (2004).
73. Barker, S. *et al.* Icebergs not the trigger for North Atlantic cold events. *Nature* **520**, 333 (2015).
74. Oppo, D. W. & Lehman, S. J. Suborbital timescale variability of North Atlantic deep water during the past 200,000 years. *Paleoceanography* **10**, 901–910 (1995).
75. van Kreveld, S. *et al.* Potential links between surging ice sheets, circulation changes and the Dansgaard-Oeschger cycles in the Irminger Sea, 60–18 kyr. *Paleoceanography* **15**, 425–442 (2000).

76. Brendryen, J., Hafliðason, H. & Sejrup, H. P. Non-synchronous deposition of North Atlantic Ash Zone II in Greenland ice cores, and North Atlantic and Norwegian Sea sediments: an example of complex glacial-stage tephra transport. *Journal of Quaternary Science* **26**, 739–745 (2011).
77. Knutz, P. C., Zahn, R. & Hall, I. R. Centennial-scale variability of the British Ice Sheet: Implications for climate forcing and Atlantic meridional overturning circulation during the last deglaciation. *Paleoceanography* **22**, <https://doi.org/10.1029/2006PA001298> (2007).
78. Vidal, L. *et al.* Evidence for changes in the North Atlantic Deep Water linked to meltwater surges during the Heinrich events. *Earth and Planetary Science Letters* **146**, 13–26 (1997).
79. Duplessy, J. C. *et al.* Changes in surface salinity of the North Atlantic Ocean during the last deglaciation. *Nature* **358**, 485–487 (1992).
80. Vogelsang, E., Sarnthein, M. & Pflaumann, U. $\delta^{18}\text{O}$ Stratigraphy, chronology, and sea surface temperatures of Atlantic sediment records (GLAMAP-2000 Kiel). Report 13 (Universität Kiel, Kiel, 2001).
81. Weinelt, M. *et al.* Variability of North Atlantic heat transfer during MIS 2. *Paleoceanography* **18**, <https://doi.org/10.1029/2002PA000772> (2003).
82. Peck, V. L., Hall, I. R., Zahn, R. & Scofield, J. D. Progressive reduction in NE Atlantic intermediate water ventilation prior to Heinrich events: Response to NW European ice sheet instabilities? *Geochemistry Geophysics Geosystems* **8**, <https://doi.org/10.1029/2006GC001321> (2007).
83. Peck, V., Hall, I. R., Zahn, R. & Elderfield, H. Millennial-scale surface and subsurface paleothermometry from the northeast Atlantic, 55ka BP. *Paleoceanography* **23**, <https://doi.org/10.1029/2008PA001631> (2008).
84. Broecker, W. S. Salinity history of the northern Atlantic during the last deglaciation. *Paleoceanography* **5**, 459–467 (1990).
85. Bond, G. *et al.* Correlations between climate records from North Atlantic sediments and Greenland ice. *Nature* **365**, 143–147 (1993).
86. Eynaud, F. *et al.* Deglacial laminated facies on the NW European continental margin: The hydrographic significance of British-Irish Ice Sheet deglaciation and Fleuve Manche paleoriver discharges. *Geochemistry, Geophysics, Geosystems* **8**, <https://doi.org/10.1029/2006GC001496> (2007).
87. Toucanne, S. *et al.* Millennial-scale fluctuations of the European Ice Sheet at the end of the last glacial, and their potential impact on global climate. *Quaternary Science Reviews* **123**, 113–133 (2015).
88. Salgueiro, E. *et al.* Temperature and productivity changes off the western Iberian margin during the last 150 ky. *Quaternary Science Reviews* **29**, 680–695 (2010).
89. Naughton, F. *et al.* Present-day and past (last 25000 years) marine pollen signal off western Iberia. *Marine Micropaleontology* **62**, 91–114 (2007).
90. Naughton, F. *et al.* Climate variability across the last deglaciation in NW Iberia and its margin. *Quaternary International* **414**, 9–22 (2016).
91. Salgueiro, E. *et al.* Past circulation along the western Iberian margin: a time slice vision from the Last Glacial to the Holocene. *Quaternary Science Reviews* **106**, 316–329 (2014).
92. Naughton, F. *et al.* Wet to dry climatic trend in north-western Iberia within Heinrich events. *Earth and Planetary Science Letters* **284**, 329–342 (2009).
93. Labeyrie, L. *et al.* In *Geophysical Monograph Series 112, Mechanisms of global climate change at millennial time scales* (eds Clark, P., Webb, R. S. & Keigwin, L. D.) 77–98 (AGU, 1999).
94. Schönfeld, J., Zahn, R. & de Abreu, L. Surface and deep water response to rapid climate changes at the Western Iberian Margin. *Global and Planetary Change* **36**, 237–264 (2003).
95. de Abreu, L., Shackleton, N., Schönfeld, J., Hall, M. & Chapman, M. Millennial-scale oceanic climate variability of the Western Iberian margin during the last two glacial periods. *Marine Geology* **196**, 1–20 (2003).
96. Lebreiro, S. *et al.* Sediment instability on the Portuguese continental margin under abrupt glacial climate changes (last 60 kyr). *Quaternary Science Reviews* **28**, 3211–3223 (2009).
97. Repschläger, J., Garbe-Schönberg, D., Weinelt, M. & Schneider, R. Holocene evolution of the North Atlantic subsurface transport. *Climate of the Past* **13**, 333–344 (2017).
98. Schwab, C., Kinkel, H., Weinelt, M. & Repschläger, J. Coccolithophore paleoproductivity and ecology response to deglacial and Holocene changes in the Azores Current System. *Paleoceanography* **27**, <https://doi.org/10.1029/2012PA002281> (2012).
99. Sarnthein, M., Balmer, S., Grootes, P. M. & Mudelsee, M. Planktic and benthic 14 C reservoir ages for three ocean basins, calibrated by a suite of 14 C plateaus in the glacial-to-deglacial Suigetsu atmospheric 14 C record. *Radiocarbon* **57**, 129–151 (2015).
100. Voelker, A., De Abreu, L., Schönfeld, J., Erlenkeuser, H. & Abrantes, F. Hydrographic conditions along the western Iberian margin during marine isotope stage 2. *Geochemistry, Geophysics, Geosystems* **10**, <https://doi.org/10.1029/2009GC002605> (2009).
101. Shackleton, N. J., Hall, M. A. & Vinent, E. Phase relationships between millennial-scale events 64 000–24 000 years ago. *Paleoceanography* **15**, 565–569 (2000).
102. Bard, E. *et al.* Radiocarbon calibration/comparison records based on marine sediments from the Pakistan and Iberian margins. *Radiocarbon* **55**, 1999–2019 (2013).
103. Chabaud, L., Sánchez Goñi, M. F., Desprat, S. & Rossignol, L. Land–sea climatic variability in the eastern North Atlantic subtropical region over the last 14,200 years: atmospheric and oceanic processes at different timescales. *The Holocene* **24**, 787–797 (2014).
104. Skinner, L., Shackleton, N. & Elderfield, H. Millennial-scale variability of deep-water temperature and $\delta^{18}\text{O}_{\text{dw}}$ indicating deep-water source variations in the Northeast Atlantic, 0–34 cal. ka BP. *Geochemistry Geophysics Geosystems* **4**, <https://doi.org/10.1029/2003GC000585> (2003).
105. Skinner, L. & Shackleton, N. Rapid transient changes in northeast Atlantic deep water ventilation age across Termination I. *Paleoceanography* **19**, PA2005, <https://doi.org/10.1029/2003PA000983> (2004).
106. Bard, E. *et al.* Retreat velocity of the North Atlantic polar front during the last deglaciation determined by 14C accelerator mass spectrometry. *Nature* **328**, 791–794 (1987).
107. Richter, T. *Sedimentary fluxes at the mid-atlantic ridge - sediment sources, accumulation rates, and geochemical characterisation*. GEOMAR Report 73 (Christian Albrechts University in Kiel, Kiel, 1998).
108. Labeyrie, L., Waelbroeck, C., Cortijo, E., Michel, E. & Duplessy, J.-C. Changes in deep water hydrology during the Last Deglaciation. *C. R. Geoscience* **337**, 919–927 (2005).
109. Gherardi, J.-M. *et al.* Glacial-interglacial circulation changes inferred from 231Pa/230Th sedimentary record in the North Atlantic region. *Paleoceanography* **24**, <https://doi.org/10.1029/2008PA001696> (2009).
110. Keigwin, L. D. & Swift, S. A. Carbon isotope evidence for a northern source of deep water in the glacial western North Atlantic. *Proceedings of the National Academy of Sciences* **114**, 2831–2835 (2017).
111. Cacho, I. *et al.* Variability of the western Mediterranean Sea surface temperature during the last 25,000 years and its connection with the Northern Hemisphere climatic changes. *Paleoceanography* **16**, 40–52 (2001).
112. Eynaud, F. *et al.* Position of the Polar Front along the western Iberian margin during key cold episodes of the last 45 ka. *Geochemistry Geophysics Geosystems* **10**, Q07U05, <https://doi.org/10.1029/2009GC002398> (2009).
113. Delivet, S. *Sedimentary expression of internal waves on Quaternary contouritic processes along the Irish and Moroccan Atlantic margins*. Ph.D. thesis, Ghent University, Belgium (2016).
114. Sarnthein, M. *et al.* Changes in east Atlantic deepwater circulation over the last 30 000 years: eight time slice reconstructions. *Paleoceanography* **9**, 209–267 (1994).

115. Penaud, A. *et al.* Contrasting paleoceanographic conditions off Morocco during Heinrich events (1 and 2) and the Last Glacial Maximum. *Quaternary Science Reviews* **29**, 1923–1939 (2010).
116. McManus, J. F., Francois, R., Gherardi, J.-M., Keigwin, L. D. & Brown-Leger, S. Collapse and rapid resumption of Atlantic meridional circulation linked to deglacial climate changes. *Nature* **428**, 834–837 (2004).
117. Carlson, A. E. *et al.* Subtropical Atlantic salinity variability and Atlantic meridional circulation during the last deglaciation. *Geology* **36**, 991–994 (2008).
118. Keigwin, L. D. & Jones, G. A. Western North Atlantic evidence for millennial-scale changes in ocean circulation and climate. *Journal of Geophysical Research* **99**, 12397–12410 (1994).
119. Keigwin, L. D., Jones, G. A. & Lehman, S. J. Deglacial meltwater discharge, North Atlantic deep circulation and abrupt climate change. *Journal of Geophysical Research* **96**, 16811–16826 (1991).
120. Keigwin, L. D. Radiocarbon and stable isotope constraints on Last Glacial Maximum and Younger Dryas ventilation in the western North Atlantic. *Paleoceanography* **19**, <https://doi.org/10.1029/2004PA001029> (2004).
121. Rasmussen, T. & Thomsen, E. Changes in planktic foraminiferal faunas, temperature and salinity in the Gulf Stream during the last 30,000 years: influence of meltwater via the Mississippi River. *Quaternary Science Reviews* **33**, 42–54 (2012).
122. Hoogakker, B., McCave, I. & Vautravers, M. Antarctic link to deep flow speed variation during Marine Isotope Stage 3 in the western North Atlantic. *Earth and Planetary Science Letters* **257**, 463–473 (2007).
123. Ziegler, M., Nürnberg, D., Karas, C., Tiedemann, R. & Lourens, L. J. Persistent summer expansion of the Atlantic Warm Pool during glacial abrupt cold events. *Nature Geoscience* **1**, 601 (2008).
124. Nürnberg, D., Ziegler, M., Karas, C., Tiedemann, R. & Schmidt, M. W. Interacting Loop Current variability and Mississippi River discharge over the past 400 kyr. *Earth and Planetary Science Letters* **272**, 278–289 (2008).
125. Henderiks, J. *et al.* Glacial–interglacial variability of particle accumulation in the Canary Basin: a time-slice approach. *Deep Sea Research Part II: Topical Studies in Oceanography* **49**, 3675–3705 (2002).
126. Plewa, K., Meggers, H. & Kasten, S. Barium in sediments off northwest Africa: A tracer for paleoproductivity or meltwater events? *Paleoceanography* **21**, <https://doi.org/10.1029/2005PA001136> (2006).
127. Kuhlmann, H., Freudenthal, T., Helmke, P. & Meggers, H. Reconstruction of paleoceanography off NW Africa during the last 40,000 years: influence of local and regional factors on sediment accumulation. *Marine Geology* **207**, 209–224 (2004).
128. Came, R. E., Oppo, D. W. & McManus, J. F. Amplitude and timing of temperature and salinity variability in the subpolar North Atlantic over the past 10 ky. *Geology* **35**, 315–318 (2007).
129. Marchitto, T. M. Jr., Curry, W. B. & Oppo, D. W. Millennial-scale changes in North Atlantic circulation since the last glaciation. *Nature* **393**, 557 (1998).
130. Curry, W. B., Marchitto, T. M., McManus, J. F., Oppo, D. W. & Laarkamp, K. L. In *Geophysical Monograph Series 112, Mechanisms of Global Climate Change at Millennial Time Scales* (eds Clark, P., Webb, R. S. & Keigwin, L. D.) 59–76 (AGU, 1999).
131. Came, R. E., Oppo, D. W. & Curry, W. B. Atlantic Ocean circulation during the Younger Dryas: Insights from a new Cd/Ca record from the western subtropical South Atlantic. *Paleoceanography* **18**, 1086, [10.1029/2003PA000888](https://doi.org/10.1029/2003PA000888) (2003).
132. Zahn, R., Winn, K. & Sarinthein, M. Benthic foraminiferal $\delta^{13}C$ and accumulation rates of organic carbon: *Uvigerina peregrina* group and *Cibicides wuellerstorfi*. *Paleoceanography* **1**, 27–42 (1986).
133. Lynch-Stieglitz, J., Schmidt, M. W. & Curry, W. B. Evidence from the Florida Straits for Younger Dryas ocean circulation changes. *Paleoceanography* **26**, <https://doi.org/10.1029/2010PA002032> (2011).
134. Lynch-Stieglitz, J. *et al.* Muted change in Atlantic overturning circulation over some glacial-aged Heinrich events. *Nature Geoscience* **7**, 144 (2014).
135. Valley, S., Lynch-Stieglitz, J. & Marchitto, T. M. Timing of Deglacial AMOC Variability From a High-Resolution Seawater Cadmium Reconstruction. *Paleoceanography* **32**, 1195–1203 (2017).
136. Came, R. E., Oppo, D. W., Curry, W. B. & Lynch-Stieglitz, J. Deglacial variability in the surface return flow of the Atlantic meridional overturning circulation. *Paleoceanography* **23**, <https://doi.org/10.1029/2007PA001450> (2008).
137. Collins, J. *et al.* Interhemispheric symmetry of the tropical African rainbelt over the past 23,000 years. *Nature Geoscience* **4**, 42–45 (2011).
138. Papenfuss, T. Glazial-Interglaziale Variation von Planktonproduktivität und Nährstoffgehalten im tropisch-subtropischen Ostatlantik im Abbild der Barium-Gehalte. Report 5 (Institut für Geowissenschaften, Universität Kiel, Kiel, Germany, 1999).
139. Jullien, E. *et al.* Low-latitude “dusty events” vs. high-latitude “icy Heinrich events”. *Quaternary Research* **68**, 379–386 (2007).
140. Skonieczny, C. *et al.* Monsoon-driven Saharan dust variability over the past 240,000 years. *Science advances* **5**, eaav1887 (2019).
141. Mulitza, S. *et al.* Sahel megadroughts triggered by glacial slowdowns of Atlantic meridional overturning. *Paleoceanography* **23**, <https://doi.org/10.1029/2008PA001637> (2008).
142. Zarriess, M. & Mackensen, A. The tropical rainbelt and productivity changes off northwest Africa: A 31,000-year high-resolution record. *Marine Micropaleontology* **76**, 76–91 (2010).
143. Vink, A. *et al.* Shifts in the position of the North Equatorial Current and rapid productivity changes in the western Tropical Atlantic during the last glacial. *Paleoceanography* **16**, 479–490 (2001).
144. Hughen, K., Southon, J., Lehman, S., Bertrand, C. & Turnbull, J. Marine-derived ^{14}C calibration and activity record for the past 50,000 years updated from the Cariaco Basin. *Quaternary Science Reviews* **25**, 3216–3227 (2006).
145. Voigt, I. *et al.* Variability in mid-depth ventilation of the western Atlantic Ocean during the last deglaciation. *Paleoceanography* **32**, 948–965 (2017).
146. Zhang, Y. *et al.* Origin of increased terrigenous supply to the NE South American continental margin during Heinrich Stadial 1 and the Younger Dryas. *Earth and Planetary Science Letters* **432**, 493–500 (2015).
147. Rühlemann, C. *et al.* Late Quaternary productivity changes in the western equatorial Atlantic: Evidence from ^{230}Th -normalized carbonate and organic carbon accumulation rates. *Marine Geology* **135**, 127–152 (1996).
148. Vidal, L. *et al.* Link between the North and South Atlantic during the Heinrich events of the last glacial period. *Climate Dynamics* **15**, 909–919 (1999).
149. Pastouret, L., Chamley, H., Delibrias, G., Duplessy, J.-C. & Thiede, J. Late Quaternary climatic changes in western Tropical Africa deduced from deep-sea sedimentation of the Niger delta. *Oceanologica Acta* **1**, 217–232 (1978).
150. Weldeab, S., Lea, D. W., Schneider, R. R. & Andersen, N. 155,000 years of West African monsoon and ocean thermal evolution. *Science* **316**, 1303–1307 (2007).
151. Mulitza, S. *et al.* Synchronous and proportional deglacial changes in Atlantic Meridional Overturning and northeast Brazilian precipitation. *Paleoceanography* **32**, 622–633 (2017).
152. Burckel, P. *et al.* Changes in the geometry and strength of the Atlantic meridional overturning circulation during the last glacial (20–50 ka). *Climate of the Past* **12**, 2061 (2016).
153. Freeman, E. *et al.* An Atlantic–Pacific ventilation seesaw across the last deglaciation. *Earth and Planetary Science Letters* **424**, 237–244 (2015).
154. Arz, H. W., Gerhardt, S., Pätzold, J. & Röhl, U. Millennial-scale changes of surface- and deep-water flow in the western tropical Atlantic linked to Northern Hemisphere high-latitude climate during the Holocene. *Geology* **29**, 239–242 (2001).
155. Jaeschke, A., Rühlemann, C., Arz, H., Heil, G. & Lohmann, G. Coupling of millennial-scale changes in sea surface temperature and precipitation off northeastern Brazil with high-latitude climate shifts during the last glacial period. *Paleoceanography* **22**, <https://doi.org/10.1029/2006PA001391> (2007).

156. Kim, J. H., Schneider, R. R., Mulitza, S. & Müller, P. J. Reconstruction of SE trade-wind intensity based on sea-surface temperature gradients in the Southeast Atlantic over the last 25 kyr. *Geophysical Research Letters* **30**, <https://doi.org/10.1029/2003GL017557> (2003).
157. Kim, J.-H. *et al.* Pronounced subsurface cooling of North Atlantic waters off Northwest Africa during Dansgaard–Oeschger interstadials. *Earth and Planetary Science Letters* **339**, 95–102 (2012).
158. Behling, H., Arz, H. W., Pätzold, J. & Wefer, G. Late Quaternary vegetational and climate dynamics in southeastern Brazil, inferences from marine cores GeoB 3229-2 and GeoB 3202-1. *Palaeogeography, Palaeoclimatology, Palaeoecology* **179**, 227–243 (2002).
159. Balmer, S., Sarnthein, M., Mudelsee, M. & Grootes, P. M. Refined modeling and 14C plateau tuning reveal consistent patterns of glacial and deglacial 14C reservoir ages of surface waters in low-latitude Atlantic. *Paleoceanography* **31**, 1030–1040 (2016).
160. Collins, J., Schefuss, E., Govin, A., Mulitza, S. & Tiedemann, R. Insolation and glacial–interglacial control on southwestern African hydroclimate over the past 140 000 years. *Earth and Planetary Science Letters* **398**, 1–10 (2014).
161. Santos, T. P. *et al.* Prolonged warming of the Brazil Current precedes deglaciations. *Earth and Planetary Science Letters* **463**, 1–12 (2017).
162. Farmer, E. C., Demenocal, P. B. & Marchitto, T. M. Holocene and deglacial ocean temperature variability in the Benguela upwelling region: Implications for low-latitude atmospheric circulation. *Paleoceanography* **20**, <https://doi.org/10.1029/2004PA001049> (2005).
163. Portillo-Ramos, R. *et al.* Methane release from the southern Brazilian margin during the last glacial. *Scientific reports* **8**, 5948 (2018).
164. Hoffman, J. & Lund, D. Refining the stable isotope budget for Antarctic Bottom Water: New foraminiferal data from the abyssal southwest Atlantic. *Paleoceanography* **27**, <https://doi.org/10.1029/2011PA002216> (2012).
165. Dickson, A. J. *et al.* Oceanic forcing of the Marine Isotope Stage 11 interglacial. *Nature Geoscience* **2**, 428–433 (2009).
166. Martínez-Méndez, G. *et al.* Contrasting multiproxy reconstructions of surface ocean hydrography in the Agulhas Corridor and implications for the Agulhas Leakage during the last 345,000 years. *Paleoceanography* **25**, <https://doi.org/10.1029/2009PA001879> (2010).
167. Barker, S. & Diz, P. Timing of the descent into the last Ice Age determined by the bipolar seesaw. *Paleoceanography* **29**, 489–507 (2014).
168. Ziegler, M., Diz, P., Hall, I. R. & Zahn, R. Millennial-scale changes in atmospheric CO₂ levels linked to the Southern Ocean carbon isotope gradient and dust flux. *Nature Geoscience* **6**, 457–461 (2013).
169. Gottschalk, J. *et al.* Abrupt changes in the southern extent of North Atlantic Deep Water during Dansgaard–Oeschger events. *Nature Geoscience* **8**, 950–954 (2015).
170. Roberts, J. *et al.* Evolution of South Atlantic density and chemical stratification across the last deglaciation. *Proceedings of the National Academy of Sciences* **113**, 514–519 (2016).
171. Kindler, P. *et al.* Temperature reconstruction from 10 to 120 kyr b2k from the NGRIP ice core. *Climate of the Past* **10**, 1–16 (2014).

Acknowledgements

The research leading to these results has received funding from the European Research Council under the European Union's Seventh Framework Program (FP7/2007–2013 Grant agreement n° 339108). New ¹⁴C dates for cores EW9209-1JPC and V29-202 were funded by NSF OCE grants to DWO, FN, ES and AV acknowledge FCT funding support through project UID/Multi/04326/2019. We thank T. Garlan and P. Guyomard for having given us access to cores from the Service Hydrographique et Océanographique de la Marine. We acknowledge N. Smialkowski for help with formatting the data into text files, and L. Maclair, L. Leroy and G. Isguder for the picking of numerous foraminifer samples for radiocarbon dating. We are grateful to S. Obrochta, E. Cortijo, E. Michel, F. Bassinot, J.C. Duplessy, and L. Labeyrie for advice and fruitful discussions. This paper is LSCE contribution 6572.

Author contributions

C.W. initiated and assembled the data set with the help of N.V.R. and W.F.; C.W., B.C.L. and N.V.R. defined the approach for the various methods used; B.C.L. processed GLODAP ¹⁴C reservoir ages; C.W., N.V.R. and L.M. defined the alignment targets; C.W. defined the alignment tie points for high latitude cores; J.B.P. assisted with alignment tie points for Nordic Seas cores; N.V.R. defined the alignment tie points for the Brazilian cores and checked consistency for all South Atlantic cores; each author specifically checked the dating information for the core(s) on which he/she has led previous studies. B.C.L. ran the *Undatable* age-depth models; C.W. drafted the first version of the manuscript and all authors contributed to drafting and editing.

Additional Information

Supplementary information is available for this paper at <https://doi.org/10.1038/s41597-019-0173-8>.

Competing Interests: The authors declare no competing interests.

Publisher's note: Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.



Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons license, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons license and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this license, visit <http://creativecommons.org/licenses/by/4.0/>.

The Creative Commons Public Domain Dedication waiver <http://creativecommons.org/publicdomain/zero/1.0/> applies to the metadata files associated with this article.

© The Author(s) 2019

Claire Waelbroeck¹, Bryan C. Loughheed¹, Natalia Vazquez Riveiros^{1,2}, Lise Missiaen¹, Joel Pedro³, Trond Dokken³, Irka Hajdas⁴, Lukas Wacker⁴, Peter Abbott^{5,6}, Jean-Pascal Dumoulin^{1,7}, François Thil¹, Frédérique Eynaud⁸, Linda Rossignol⁸, Wiem Fersi¹, Ana Luiza Albuquerque⁹, Helge Arz¹⁰, William E. N. Austin¹¹, Rosemarie Came¹², Anders E. Carlson¹³, James A. Collins¹⁴, Bernard Dennielou², Stéphanie Desprat^{8,15}, Alex Dickson¹⁶, Mary Elliot¹⁷, Christa Farmer¹⁸, Jacques Giraudeau⁸, Julia Gottschalk¹⁹, Jorijntje Henderiks²⁰, Konrad Hughen²¹, Simon Jung²², Paul Knutz²³, Susana Lebreiro²⁴, David C. Lund²⁵, Jean Lynch-Stieglitz²⁶, Bruno Malaizé⁸, Thomas Marchitto²⁷, Gema Martínez-Méndez²⁸, Gesine Mollenhauer²⁸, Filipa Naughton^{29,30}, Silvia Nave³¹, Dirk Nürnberg³², Delia Oppo²¹, Victoria Peck³³, Frank J. C. Peeters³⁴, Aurélie Penaud³⁵, Rodrigo da Costa Portilho-Ramos³⁶, Janne Repschläger³⁷, Jenny Roberts³⁸, Carsten Rühlemann³⁹, Emilia Salgueiro^{29,30}, Maria Fernanda Sanchez Goni^{8,15}, Joachim Schönfeld³², Paolo Scussolini³⁴, Luke C. Skinner⁴⁰, Charlotte Skonieczny⁴¹, David Thornalley⁴², Samuel Toucanne², David Van Rooij⁴³, Laurence Vidal⁴⁴, Antje H. L. Voelker^{29,30}, Mélanie Wary⁴⁵, Syee Weldeab⁴⁶ & Martin Ziegler⁴⁷

¹LSCE/IPSL, Laboratoire CNRS-CEA-UVSQ, 91191, Orme des Merisiers, France. ²Ifremer, Unité de Geosciences Marines, 29280, Plouzané, France. ³Uni Research, Nygårdsgaten 112, 5008, Bergen, Norway. ⁴Laboratory of Ion Beam Physics, ETH Zürich, 8093, Zürich, Switzerland. ⁵School of Earth and Ocean Sciences, Cardiff University, CF10 3AT, Cardiff, UK. ⁶Institute of Geological Sciences and Oeschger Center for Climate Change Research, University of Bern, 3012, Bern, Switzerland. ⁷LMC14, Université Paris-Saclay, F-91191, Gif-sur-Yvette, France. ⁸EPOC, Université Bordeaux, Allée Geoffroy St Hilaire, 33615, Pessac, France. ⁹LOOP, Universidade Federal Fluminense, Niterói, RJ, Brazil. ¹⁰Leibniz-Institute for Baltic Sea Research Warnemünde, Seestrasse 15, 18119, Rostock, Germany. ¹¹University of St Andrews, St Andrews, Scotland, KY16 9AL, UK. ¹²University of New Hampshire, 56 College Road, Durham, NH, 03824, USA. ¹³Oregon State University, Corvallis, OR, 97331, USA. ¹⁴GeoForschungsZentrum, 14473, Potsdam, Germany. ¹⁵Ecole Pratique des Hautes Etudes (EPHE, PSL), 4-14 rue Ferrus, 75014, Paris, France. ¹⁶Royal Holloway University of London, Egham, Surrey, TW20 0EX, UK. ¹⁷LPG-Nantes, Université de Nantes, 44300, Nantes, France. ¹⁸Hofstra University, Hempstead, NY, 11549-1140, USA. ¹⁹Lamont-Doherty Earth Observatory, Columbia University, 61 Route 9W - PO Box 1000, Palisades, NY, 10964-1000, USA. ²⁰Uppsala University, Geocentrum, Villavägen 16, SE-752 36, Uppsala, Sweden. ²¹Woods Hole Oceanographic Institution, 266 Woods Hole Rd., Woods Hole, MA, 02543-1050, USA. ²²University of Edinburgh, School of Geosciences, Edinburgh, EH9 3FE, UK. ²³Geological Survey of Denmark and Greenland, Øster Voldgade 10, 1350, Copenhagen, Denmark. ²⁴IGME - Instituto Geológico y Minero de España, Calle Ríos Rosas, 23, 28003, Madrid, Spain. ²⁵University of Connecticut, 1080 Shennecossett Road, Groton, CT, 06340, USA. ²⁶Georgia Institute of Technology, 311 Ferst Drive, Atlanta, GA, 30332-0340, USA. ²⁷INSTAAR, University of Colorado, Boulder, CO, 80303, USA. ²⁸AWI, Am Alten Hafen 26, 27568, Bremerhaven, Germany. ²⁹IPMA-DivGM, Avenida Doutor Alfredo Magalhães Ramalho, 6, 1495-165, Alges, Portugal. ³⁰CCMAR, Universidade do Algarve, Campus de Gambelas, 8005-139, Faro, Portugal. ³¹LNEG, Bairro do Zambujal, 2610-999, Amadora, Portugal. ³²GEOMAR, Wischhofstrasse 1-3, 24148, Kiel, Germany. ³³UK British Antarctic Survey, Madingley Road, Cambridge, CB3 0ET, UK. ³⁴Vrije Universiteit Amsterdam, De Boelelaan 1087, 1081 HV, Amsterdam, Netherlands. ³⁵Université de Bretagne Occidentale, Technopôle Brest-Iroise, 29280, Plouzané, France. ³⁶MARUM, University of Bremen, D-28359, Bremen, Germany. ³⁷Max Planck Institute, Hahn- Meitner- Weg 1, 55128, Mainz, Germany. ³⁸Thermo Fisher Scientific, Hanna-Kunath Straße 11, Bremen, 28199, Germany. ³⁹Geozentrum Hannover, Stilleweg 2, 30655, Hannover, Germany. ⁴⁰University of Cambridge, Godwin Laboratory for Palaeoclimate Research, Cambridge, CB2 3EQ, UK. ⁴¹GEOPS, Université de Paris Sud, 91405, Orsay, France. ⁴²University College London, Gower Street, London, WC1E 6BT, UK. ⁴³Ghent University, Krijgslaan 281, 9000, Gent, Belgium. ⁴⁴Aix-Marseille Université, CNRS, IRD, INRA, Coll France, CEREGE, Europole de l'Arbois, 13545, Aix-en-Provence, France. ⁴⁵ICTA, Universitat Autònoma de Barcelona, 08193, Bellaterra, Barcelona, Spain. ⁴⁶University of California Santa Barbara, Santa Barbara, 1006 Webb Hall, CA, 93106-9630, USA. ⁴⁷University of Utrecht, Princetonlaan 8a, 3584 CB, Utrecht, Netherlands.