



Geophysical and geochemical analysis of shallow gas and an associated pockmark field in Bantry Bay, Co. Cork, Ireland

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Sean Jordan, Shane O'Reilly, Daniel Praeg, Dayton Dove, Lorenzo Facchin, et al.. Geophysical and geochemical analysis of shallow gas and an associated pockmark field in Bantry Bay, Co. Cork, Ireland. *Estuarine, Coastal and Shelf Science*, 2019, 225, pp.106232. 10.1016/j.ecss.2019.05.014 . hal-02272568

HAL Id: hal-02272568

<https://hal.science/hal-02272568>

Submitted on 29 Aug 2019

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Manuscript Details

Manuscript number	ECSS_2019_96_R2
Title	Geophysical and geochemical analysis of shallow gas and an associated pockmark field in Bantry Bay, Co. Cork, Ireland.
Article type	Research Paper

Abstract

An integrated geophysical, geological, and geochemical investigation of seabed fluid venting was carried out in upper Bantry Bay, a large marine inlet on the southwest coast of Ireland. The results provide evidence of the seafloor venting of gas rich fluids, resulting in the formation of a pockmark field identified here for the first time. The pockmarks occur in an area where sub-bottom profiles provide evidence of chimney-like features interpreted to record upward gas migration through Quaternary sediments to the seafloor. Three vibrocores up to 6 m long were acquired in water depths of 24-34 m, two from the pockmark field and one from outside. Methane of predominantly biogenic origin was quantified in all three cores by headspace analysis of sediment sub-samples. Well-defined sulfate methane transition zones (SMTZs) were observed in two of the cores, the shallowest (1.25 m) inside the pockmark field and the other (3.75 m) outside. It is likely that an SMTZ occurs at the location of the third core, also within the pockmark field, although beneath the samples obtained during this study. Gas release possibly from a combination of various faulting mechanisms and shallow methanogenesis appears to drive diffuse pore fluid migration across wide areas, while focused flow through the pockmarks may be related to gas originating from the Owenberg River Fault and methanogenesis of pre-glacial lacustrine sediments preserved in a bedrock basin. Analysis of phospholipid fatty acids (PLFAs) and archaeal isoprenoid hydrocarbons was used to investigate the microbial ecology of these sediments. Anaerobic oxidation of methane (AOM) may play a role in controlling release of CH₄ to the water column and atmosphere in this shallow gas setting, potentially mediated by syntrophic sulfate reducing bacteria (SRB) and anaerobic methanotrophic archaea (ANME).

Keywords	Geophysics; biogeochemical processes; pockmarks; fluid migration; anaerobic oxidation of methane (AOM); lipid biomarkers
Taxonomy	Oceanography, Methane
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Suggested reviewers	Alan Judd, Joseph Kelley, Aggeliki Georgiopoulou, Crispin Little

Submission Files Included in this PDF

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Highlights

- Acoustic data provided evidence for widespread fluid migration in a shallow marine bay in Co. Cork, Ireland including shallow gas deposits near the seabed.
- Fluid migration has led to the formation of a previously undescribed pockmark field within the bay.
- Ground-truthing confirmed that the fluid was methane which is likely both thermogenic and biogenic in origin, possibly derived from an underlying fault and methanogenesis of pre-glacial lacustrine sediments.
- Geochemical evidence suggests that microbial anaerobic oxidation of methane (AOM) plays a key role in controlling the release of methane to the atmosphere from the bay.

1 Geophysical and geochemical analysis of shallow gas and an associated
2 pockmark field in Bantry Bay, Co. Cork, Ireland.

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Keywords

Seafloor; pockmarks; biogeochemical processes; fluid migration; anaerobic oxidation of methane (AOM); lipid biomarkers; methane; climate change; geohazards

1. Introduction

Pockmarks are concave depressions within seabed sediments, circular to ellipsoidal in shape, ranging from <1 to 400 m in diameter and up to 20 m deep (Hovland and Judd, 1988;

King and MacLean, 1970), although typically 30 to 40 m wide and 2 to 3 m deep (Acosta et al., 2001). Pockmarks can occur as singular features, in linear patterns known as pockmark trains, or in complex groups known as pockmark fields. The formation and dynamics of these features are still not fully understood, but they are generally considered to be the result of the expulsion of fluids typically including hydrocarbon gases, mainly methane (CH₄), from seafloor sediment (Hovland, 2013; Hovland and Judd, 1988). The emission of fluids containing gas from pockmarks makes them of interest in relation to issues of global carbon cycling and climate change, as well as for seafloor geohazards (Judd and Hovland 2007).

Geologic Emissions of Methane (GEM), which include marine seeps such as pockmarks, have been recognized as a natural source of atmospheric methane second only to wetlands (Etiope et al., 2008). As a greenhouse gas, the warming potential of CH₄ outweighs carbon dioxide (CO₂) by a factor of 25 times per ton, and since pre-industrial times is estimated to have been responsible for approximately 20% of the Earth's warming (Yvon-Durocher et al., 2014). Recent work indicates that contributions from marine sources have been greatly underestimated (Skarke et al., 2014) and there is a need for CH₄ flux revisions in terms of understanding the global carbon cycle (Judd and Hovland, 2009). Seepage sites are globally widespread in shallow water coastal regions and have been suggested to be an important source of CH₄ (Borges et al., 2016; Janssen et al., 2005; Shakhova et al., 2010; Skarke et al., 2014). However, global estimates of the contribution to atmospheric CH₄ concentrations from marine seepage sites are highly uncertain (Römer et al., 2014).

The presence of pockmarks may also be of significance in terms of marine geohazards (Hovland, 1989). Fluid migration through marine sediments, through its influence on pore pressures and sediment strength, is thought to play a key role in slope failure and seabed instability (e.g. Locat and Lee, 2002). Therefore in pockmarked areas the development of offshore infrastructures, such as pipelines, may need to avoid these features (Hovland et al.,

2002). In addition, pockmarks have been suggested as possible indicators of seismic activity (Hovland et al., 2002), based on observations of gas venting from pockmarks before and during earthquakes at sites in California (Field and Jennings, 1987) and Greece (Hasiotis et al., 1996; Soter, 1999). Large-scale multinational monitoring of pockmarks has been advocated (Hovland et al., 2002).

Anaerobic oxidation of methane (AOM) and the microbial consortia involved are important factors in the global methane cycle, and yet they are still poorly understood (Gauthier et al., 2015; Ruff et al., 2016). Although large amounts of CH₄ are transported from deep reservoirs to shallow sediments, it is estimated that <3% reaches the atmosphere due to the AOM performed by microbial communities (Niemann and Elvert, 2008). The predominant mechanism of AOM is thought to be a syntrophic process whereby anaerobic methanotrophic archaea (ANME) and sulfate reducing bacteria (SRB) oxidise CH₄ to CO₂ whilst reducing SO₄²⁻ to H₂S providing energy for both microbial consortia (Boetius et al., 2000; Elvert et al., 2003; Reeburgh, 2007; Valentine and Reeburgh, 2000):



These communities are predominantly found in sediments, however they have also been found in anoxic marine and saline lacustrine water bodies, and in terrestrial mud volcanoes (Alain et al., 2006; Joye et al., 1999; Wakeham et al., 2003). AOM primarily occurs at what is known as the sulfate methane transition zone (SMTZ), where CH₄ diffusion from deeper sediments and SO₄²⁻ penetration from seawater provide optimal conditions for AOM communities (Knittel and Boetius, 2009).

Lipid biomarkers can provide evidence for the role played by archaea and SRB in AOM (Caldwell et al., 2008). Phospholipid fatty acids (PLFAs) are fatty acids chemically cleaved from ester linkage to polar head groups and are a useful tool to provide quantitative measures of viable biomass and microbial community composition (Ringelberg et al., 1997; Zelles,

1997). Phospholipids are rapidly degraded after cell death making them excellent biomarkers for viable microbial cells (Navarrete et al., 2000; White et al., 1997). Certain PLFAs have been used as chemotaxonomic markers for SRB, such as $C_{16:1\omega5c}$ and $cyC_{17:0\omega5,6}$ as indicators of *Desulfosarcina/Desulfococcus* species (Elvert et al., 2003). Archaeal cell membranes are comprised of ether-linked isoprenoid lipids (Schouten et al., 2013). Analysis of these intact lipids or their hydrocarbon skeletons (e.g. phytane, acyclic and cyclic C_{40} isoprenoids) in environmental samples provides a broad measure of archaeal abundance and diversity (e.g. (King et al., 1998). $\delta^{13}C$ values of AOM derived lipids are typically significantly depleted with values $< -50\text{‰}$ (Elvert et al., 2003; Niemann and Elvert, 2008; van Dongen et al., 2007). Isolation of these compounds combined with determination of their $\delta^{13}C$ signatures can help provide an overview of the microbial consortia and their involvement in AOM within cold seep environments (Ge et al., 2015; Pancost et al., 2000).

Pockmark and seepage sites have been reported and investigated at several sites around the coast of Ireland and we are only beginning to understand the dynamics and ubiquity of coastal methane cycling (Crocker et al., 2005, O'Reilly et al., 2014, Szpak et al., 2012, and Szpak et al., 2015). In this paper we present the first description of a pockmark field in the shallow waters (<30 m) of upper Bantry Bay, on the west coast of Ireland. The aim of the study is to characterise CH_4 migration associated with the pockmarks, based on core data acquired during an Irish-led campaign in 2014. The results provide information on the source of the CH_4 and its relation to the microbial ecology of this area, as well as the possible causes of pockmark formation at this site. Our findings contribute to an improved understanding of gas venting features in Irish coastal waters, that may be relevant to environmental planning, economic developments, and global climate change.

2. Regional setting

Bantry Bay is the largest marine inlet in the southwest of Ireland, spanning an offshore area of 300 km² (Fig. 1A). It is approximately 40 km long, narrowing in width from 10 km at its mouth, where water depths are up to 60 m, to 5 km at its head. The bay contains two large islands; Bere Island in the outer bay and Whiddy Island in the inner bay. The Melagh, Owvane, Coomhola, Glengarriff, and Adrigole rivers all drain into Bantry Bay. Geologically, the bay lies within the South Munster Basin, comprising Devonian strata dominated by the Old Red Sandstone beneath uppermost Devonian and Carboniferous marine sandstones and mudstones (Plets et al., 2015; Vermeulen et al., 2000). Several fault lines are inferred to run through the bay offsetting the Old Red Sandstone (Fig. 1B): the Bantry Fault runs from the southeast of Whiddy Island, continuing along the centre of the bay; the Owenberg River Fault lies north east of Whiddy Island before meeting the Bantry Fault; while northeast of Whiddy Island are the Glengarriff Harbour and Coolieragh Faults (Szpak et al., 2015).

The sedimentary infill of the Bay was described by (Plets et al. 2015), based on sub-bottom profiles tied to shallow sediment cores, who recognized bedrock to be overlain by up to six units, interpreted to record deposition prior to and since the last glacial maximum (LGM). The oldest unit corresponds to stratified sediments infilling bedrock depressions, correlated to pre-LGM lacustrine sediments reported in the upper Bay by Stillman (1968). This is overlain by glacial sediments, truncated by tidal to estuarine units recording the inundation of the Bay and capped by a seafloor unit recording the establishment of fully marine conditions after 11 ka BP. In the inner Bay, the upper stratified marine unit is underlain by a unit of strong discontinuous reflections described as ‘turbid’, that cores show to correspond to estuarine deposits, laminated sands and muds containing organic matter, suggested on the basis of its acoustic character to also contain pockets of gas (Plets et al. 2015). In addition, in the upper Bay above at least 65 m water depth, the sediment column is crossed by vertical, pillar-like

acoustic turbidity zones (ATZs) that rise to within several metres of seafloor; although not interpreted by Plets et al. (2015), these appear typical of gas chimneys (Dondurur et al. 2011). Seabed classification maps based on backscatter and particle size analysis (PSA) show that the sediment type is predominantly mud to fine sand with increasing medium to coarse sand towards the mouth of the bay. There are areas of medium to coarse sand, coarse sand to gravel, and rock throughout the bay primarily along the perimeter (INFOMAR, 2011).

3. Materials and methods

The data used in this study were acquired during campaigns undertaken as part of the INFOMAR (Integrated Mapping for the Sustainable Development of Ireland's Marine Resources) programme. Acoustic datasets including multibeam bathymetric and backscatter coverage of all of Bantry Bay were obtained during INFOMAR campaigns from 2004-2007 (see Plets et al. 2015), while the sediment cores and sub-bottom profiles used in this study were acquired as part of the GATEWAYS campaign of the Celtic Explorer in February 2014 (CE14003).

3.1. Acoustic data

Seafloor bathymetric and backscatter data were collected using two Kongsberg Simrad multibeam systems, an EM1002 (95 kHz) and an EM3002D (200 kHz). The multibeam data were processed using QTC Multiview software to generate bathymetric terrain models of 2 x 2 m grid size. No multibeam water column data were available for this study.

Sub-bottom profiles were acquired in 2014 using a heave-corrected SES Probe 5000 pinger with a 4x4 transducer array (hull-mounted) and a CODA DA2000 acquisition system.

Frequency content of 2.5 kHz corresponds to decametric vertical resolution. Acquisition parameters, data logging, and interpretation were performed using the CODA Geokit suite.

3.2. Sediment cores

Three sediment cores were obtained in 2014 using a 6 m pneumatic vibrocorer, deployed in water depths of 24-34 m. Recorded positions are those of the ship, which may differ from the corer by up to 30 m. Two cores were obtained from within the pockmark field and one core was taken from outside the field. Once on deck, cores were cut into 1 m sections and capped. Core sections were split and the archive halves were photographed and logged. Sediment porewaters were sampled downcore using Rhizons (Rhizosphere Research Products) for analysis of SO_4^{2-} distribution. These were attached to 10 mL plastic syringes to create vacuum pressure. The sampled porewater was placed in a plastic vial and preserved with 10 μL CHCl_3 for sulfate analysis. All porewater samples were refrigerated at 4°C onboard for the duration of the cruise and back in the laboratory prior to analysis.

Gas samples were immediately taken from the vibrocore sections to determine gas composition and distribution. Two 10 cm^3 sediment plugs were sampled using plastic syringes with tips removed and transferred to 50 mL glass headspace vials containing 20 mL 2 M sodium hydroxide. Vials were sealed, homogenised, and stored upside-down in the dark at 4 °C for the duration of the cruise.

Sediment sub-samples were taken immediately after porewater and gas sub-samples. Particle size analysis (PSA) samples were placed in ziplock bags and stored at room temperature. Samples for lipid biomarker analysis were wrapped in fired Al foil, placed in ziplock bags, and stored at -20 °C.

3.3. Porewater and gas analysis

SO₄²⁻ concentration in porewater was determined by the turbidimetric method. 10 mL of sample was stirred constantly and 2-3 drops of glycerol were added. Crushed BaCl₂, approximately 50 mg, was added to the mixture and stirring was continued for 1 minute after which an aliquot was taken and the absorbance measured at 420 nm on a Shimadzu UV Mini 1240. Further aliquots were taken after 2, 2.5, and 3 minutes and an average reading was calculated and used to determine concentration by extrapolation from a calibration curve. The calibration curve was prepared with Na₂SO₄ standards in a range of 10 to 100 ppm.

CH₄ analysis was performed on an Agilent 7820A GC-FID with a 30 m HP-PLOTQ column (Agilent, Santa Clara, USA). Column conditions were isothermal (50 °C). CH₄ was quantified using calibration standards prepared from a 99.995% CH₄ standard (Sigma Aldrich, Dorset, UK).

3.4. Bulk physical and chemical analysis

PSA and total organic carbon (TOC) data were obtained from sub-samples taken surrounding the SMTZ locations which were determined by CH₄ and SO₄²⁻ analyses. PSA was determined by laser granulometry using a Mastersizer 2000 particle size analyser (Malvern, Worcestershire, UK). Organic carbon (OC) was removed using 30% hydrogen peroxide (H₂O₂) prior to analysis. Elemental analysis was performed in triplicate using a Fisons NCS 1500 NA elemental analyser. Samples were treated with 1 N HCl in Ag capsules following the procedure of Verardo et al. (1990) to remove carbonate. After drying overnight, the capsules were wrapped in Sn boats and combusted in the presence of O₂. The CO₂ evolved was measured and the TOC content (%) calculated by comparison with the certified reference standard acetanilide.

3.5. Lipid biomarker analysis

Sediment samples were selected from sub-samples associated with the SMTZs. These were freeze-dried and homogenized and lipid compounds were extracted from 30 g of powdered sediment using a modified Bligh-Dyer extraction (White et al., 1997). Total lipid extracts (TLEs) were concentrated and elemental S was removed by reaction with activated Cu. TLEs were fractionated into neutral, glyco-, and polar lipids using Bond-Elut SPE columns packed with an aminopropylsilica solid phase (5mm diameter, PE, 500mg Ultra-Clean NH₂, Agilent Technologies) as outlined by Pinkart et al. (1998). A portion of each polar lipid fraction was subjected to acid methanolysis (0.5 M sodium methoxide, 50 °C, 30 min) to transmethylate ester-linked fatty acids. Double-bond positions of monounsaturated PLFAs were determined by the formation of dimethyl disulfide (DMDS) adducts as described by Nichols et al., (1986). Archaeal isoprenoid lipids were separated from polar head groups by cleavage of their ether linkages following the method of Trent et al. (2003). 100 ppm 5 α cholestane was added to all derivatised fractions as an internal standard prior to analysis.

Aliquots (1 μ l) of samples were injected in triplicate onto an Agilent model 7890N gas chromatograph coupled to an Agilent 5973N mass selective detector operating in electron impact mode at 70 eV. The column was a 30 m HP-5MS column (0.25 mm i.d., 1 μ m film thickness). Each sample (1 μ l) was injected with a 2:1 split ratio. The GC inlet temperature was 280 °C and the oven programme was 65 °C (held 2 min) to 300 °C (held 20 min) at 6 °C/min. Individual compounds were assigned from comparison with mass spectral library databases (NIST and Wiley) and comparison of MS patterns with published spectra and authentic standards. Analytes were quantified from total ion peak area using multiple-point calibration curves of representative standards (methyl tetradecanoate and squalane). Percentage recovery was measured using an internal standard added prior to extraction and was found to be > 95%. Procedural blanks were run to monitor background interferences.

1 µl aliquots of samples were injected in triplicate onto an Agilent model 7890N gas chromatograph coupled to an IsoPrime 100 isotope ratio mass spectrometer. The $\delta^{13}\text{C}$ values were measured against a CO_2 reference gas of known $\delta^{13}\text{C}$ value and are reported vs. a stable isotope reference standard (*n*-alkanes mixture B2, Indiana University, USA). Reproducibility was better than $\pm 0.5\%$ and only well resolved major analytes are reported here.

Lipid nomenclature is according to xCy ω z, where x refers to the number of carbon atoms present, y refers to the number of double bonds on the carbon chain and z refers to the position of the first double bond from the methyl end. Iso and anteiso branching is denoted by 'i' and 'ai' respectively whilst the presence of the cyclopropane ring in a compound is denoted by 'cy'.

4. Results

4.1. Geophysical analyses

Multibeam morpho-bathymetric data provide evidence of an elongate pockmark field north of Whiddy Island (Fig. 2). This is a narrow (max width ca. 275 m) pockmark field of approximately 2.4 km in length, covering an area of ca. 0.5 km². Interestingly, this field coincides with part of the Owenberg River Fault (Fig. 1B). The data show that the pockmarks average ca. 10 m in diameter and are of low relief, with some features near the core locations as shallow as ca. 0.3 m in depth (Fig. 2B). Recorded GPS position onboard the vessel may differ from the actual sample location by up to 20–30 m. Therefore, although both VC24 and VC25 were taken within the pockmark field, it is not possible to be sure whether either core penetrated directly into a single pockmark feature.

A sub-bottom profile for VC27 was not prepared as the data was obstructed by sideswipe from a rocky outcrop. Sub-bottom profiles across the sites of VC24 and VC25

provide acoustic evidence of gas migration through the sediment column (Fig. 3). The sedimentary succession is crossed by columnar or conical zones of blanking (AB on Fig. 3), most of which underlie strong reflector segments that lie at varying depths of ca. 4-10 metres below sea floor (mbsf) (Fig. 3). Similar ‘pillar-like’ acoustic zones were previously described on sub-bottom profiles across upper Bantry Bay by Plets et al. (2015). On high frequency seismic data, such effects may arise due to overlap with the resonance frequencies of gas bubble populations, resulting in energy loss by attenuation (reverberation and scattering) as well as changes in P-wave velocity (Mathys et al. 2005). Gas concentrations as low as 0.5% may result in a range of possible amplitude and coherence effects described as acoustic turbidity (Abegg and Anderson 1997; Fleischer et al. 2001; Judd and Hovland 2007). We interpret the vertical acoustic zones observed in Bantry Bay to be typical chimney structures, recording the upward migration of gas-rich fluids through the sediment column (e.g. Dondurur et al., 2011).

On Fig. 3, the tops of the chimneys are seen to lie at varying stratigraphic levels, the shallowest within an interval of strong discontinuous reflections of varying thickness. This interval corresponds to unit III of Plets et al. (2015), which their cores showed to comprise organic-rich laminated sands/muds of estuarine origin, hypothesised to contain gas pockets due to their acoustically ‘turbid’ character. This unit was also penetrated by our cores, which provide no evidence that its acoustic character can be correlated to higher gas content. We suggest instead that the reflective character is likely to reflect the unit’s distinctive lithology, comprising sand and mud laminae capable of generating strong impedance contrasts (SI Fig. S1).

4.2. Gas and porewater geochemistry

All measured CH₄ values are provided in the supporting information (SI Table S1). The highest concentrations of CH₄ were observed in VC24, taken from the pockmark field (Fig. 4).

Values fluctuated between 2.62 and 3.57 mM rising through the core before steadily decreasing from 3.28 mbsf (3.68 mM) to the surface sample at 0.01 mbsf (0.002 mM), the minimum overall value for VC24. SO_4^{2-} concentrations for VC24 ranged from 7.0 to 26.8 mM displaying an overall decreasing trend from the surface, opposite to that of CH_4 (Fig. 4). A minimum value was observed at 2.12 mbsf from which concentrations remain relatively constant through to the bottom of the core.

Overall CH_4 concentrations detected within VC25 were the lowest of the three vibrocores analysed with a maximum observed at 5.23 mbsf (0.018 mM) and a minimum observed at 1 mbsf (Fig. 4). Concentrations decrease gradually from the base of the core to the sediment surface from 0.016 mM to 0.003 mM. Concentrations of porewater SO_4^{2-} were relatively high throughout VC25 compared to VC24 and VC27 (Fig. 4). Values were gradually depleted from the seafloor (0.17 mbsf) with a concentration of 22.1 mM to the deepest sample from the core (5.66 mbsf) with a concentration of 12.0 mM.

In VC27, outside the pockmark field, CH_4 concentrations decreased from 3.66 mM at the base of the core (4.96 mbsf) to 0.97 mM at 4.08 mbsf before falling sharply to 0.07 mM at 3.6 mbsf (Fig. 4). Depletion gradually continued from this depth to 0.001 mM at the surface of the core (0.02 mbsf). SO_4^{2-} concentrations followed an opposing trend with a maximum of 23.9 mM at 0.02 mbsf decreasing to a minimum of 7.1 mM at 4.08 mbsf and remained at similar concentration to the base (4.96 mbsf) (Fig. 4).

4.3. PSA and elemental analysis

The overall sediment type for the three cores taken from Bantry Bay was poorly to very poorly sorted sandy mud. All values for mean particle size, percentage clay, silt, sand, and gravel are provided in table 1. Mud percentages (clay and silt) ranged from 69.4 to 92.3% in VC24, from 42.2 to 81.2% in VC25, and from 29.3 to 84.0% in VC27. The 42.2% value from

VC25 was obtained at 4.99 mbsf, a sample comprised of poorly sorted muddy sand due to its high sand content (57.8%). The 29.3% value in VC27 was obtained at 1.93 mbsf where sediment type can be described as very poorly sorted, slightly gravelly, muddy sand due to its gravel (4.9%) and sand (65.8%) content. This gravel-containing layer had the largest mean particle size of 0.8 phi whereas the lowest value of 5.3 phi was observed in VC24 at 0.77 mbsf, the layer with the highest overall mud content (92.3%). The mean particle size for the remaining samples ranged between 4.5 and 3.3 phi.

Total organic carbon (TOC) content was low throughout all cores with an average overall value of 0.6% (Table 2). The highest observed values were 2 and 1.2% for VC24 0.025 and 0.27 mbsf respectively. No other sample had a value greater than 0.7%. In VC24, TOC decreased from 0.025 to 1.93 mbsf (2 to 0.3%) before increasing slightly to 0.5% at 2.92 mbsf and decreasing again to 0.3% at 3.9 mbsf. VC25 values were relatively constant. The TOC content of VC27 at 1.93 and 2.96 mbsf was 0.5%. This decreased to 0.4% at 3.98 mbsf and 0.3% at 4.97 mbsf.

Table 1. PSA results for all vibrocores.

Core	Depth (mbsf)	Mean (phi)	Clay (%)	Silt (%)	Sand (%)	Gravel (%)
VC24	0.08	4.5	10.6	69.4	20	0
	0.33	3.3	10.1	62.8	27.1	0
	0.72	5.3	15.3	77	7.7	0
	1.88	4.0	9.7	68.9	21.4	0
	2.97	3.7	6.2	77.2	16.6	0
	3.96	3.8	5.5	63.9	30.6	0
VC25	0.81	3.7	2.3	63.5	34.2	0
	2.93	3.4	6.1	62.8	31.1	0
	3.93	4.3	8.4	72.7	18.8	0
	4.99	3.6	1.8	40.4	57.8	0
VC27	0.93	0.8	2.7	26.6	65.8	4.9
	1.96	4.1	11.9	72.1	16	0
	2.98	4.1	10.9	69.1	20	0

4.4.4. Lipid biomarkers

A summary of key lipid biomarker concentrations is provided in table 2. The highest overall concentrations of PLFAs in all three vibrocores were observed in VC24. 310.1 and 235.2 $\mu\text{g gOC}^{-1}$ were detected at 0.03 and 0.27 mbsf respectively, the largest quantities of PLFAs in all analysed samples. The remaining depths of VC24 contained between 31.1 (0.77 mbsf) and 90.1 $\mu\text{g gOC}^{-1}$ (1.93 mbsf). Saturated fatty acids (SATFAs) and monounsaturated fatty acids (MUFAs) were the dominant PLFAs at 0.03 and 0.27 mbsf whilst SATFAs and branched fatty acids (brFAs) were dominant from 0.77 to 3.9 mbsf. Polyunsaturated fatty acids (PUFAs) were not found at 1.93 or 2.92 mbsf and were the smallest class of PLFAs at all other depths. Total PLFA concentrations ranged from 49.7 to 73.9 $\mu\text{g gOC}^{-1}$ (5.96 and 2.93 mbsf respectively) in VC25. SATFAs were the dominant compounds throughout the core with concentrations approximately 10 times greater than MUFAs and brFAs. There were no PUFAs observed in any VC25 samples. The highest concentration of PLFAs in VC27 was 77.4 $\mu\text{g gOC}^{-1}$ observed at 3.98 mbsf. The lowest concentration was 51.2 $\mu\text{g gOC}^{-1}$ which was observed at 1.93 mbsf. Similar to VC25, total SATFA concentrations were significantly greater than other PLFA classes. There was little variation in total concentrations of other PLFA classes throughout the core.

Five archaeal ether (AE) lipids were isolated from each sample taken from VC24, VC25, and VC27. These were; phytane, acyclic biphytane ($\text{cyC}_{40:0}$), and three cyclic biphytanes ($\text{cyC}_{40:1}$, $\text{cyC}_{40:2}$, and $\text{cyC}_{40:3}$). The $\text{cyC}_{40:0}$ was the major isoprenoid in all samples whilst the $\text{cy}_{40:1}$ was the minor isoprenoid.

4.5. Carbon isotope values of individual PLFAs

364 $\delta^{13}\text{C}$ values could not be obtained for most lipid compounds identified in the three
365 Bantry Bay vibrocores. This was due to a combination of low abundance in polar lipid extracts
366 and low sensitivity of the GC-IRMS instrument. $\delta^{13}\text{C}$ values for three PLFAs were measured
367 in the VC24 0.27 sample which contribute to the study of these sediments. The MUFAs $\text{C}_{16:1\omega7}$
368 and $\text{C}_{16:1\omega5}$ provided $\delta^{13}\text{C}$ values of -31.0‰ and -46.1‰ respectively, and the SATFA $\text{C}_{16:0}$
369 $\delta^{13}\text{C}$ value was -27.7‰.

Table 2. TOC values (%) and Biomarker concentrations ($\mu\text{g gOC}^{-1}$) for each vibrocore taken in Bantry Bay. Total overall concentrations and total concentrations for certain groups of biomarkers are included (indicated by the prefix Σ). Individual concentrations of selected biomarker compounds from within each group are also shown. Depth in mbsf.

	VC24						VC25				VC27			
Depth	0.025	0.270	0.768	1.933	2.920	3.895	2.930	3.925	4.980	5.960	1.930	2.960	3.980	4.970
TOC	2.0	1.2	0.6	0.3	0.5	0.3	0.3	0.3	0.3	0.3	0.5	0.5	0.4	0.3
Σ PLFA	310.14	235.24	31.07	90.12	51.21	44.02	73.90	57.15	56.90	49.68	51.24	68.89	77.36	63.01
Σ SATFA	133.28	118.72	20.43	64.94	33.06	30.99	52.11	41.72	42.73	43.24	36.15	42.50	50.35	46.84
Σ MUFA	91.12	51.97	2.61	2.25	3.40	1.94	5.18	3.86	1.35	8.54	4.22	5.89	5.99	5.78
Σ PUFA	12.92	17.73	0.00	0.00	0.00	1.16	0.00	0.00	0.00	0.00	1.15	1.66	1.73	0.00
Σ brFA	64.58	32.62	4.31	4.44	12.15	7.62	7.19	5.09	5.17	4.38	4.91	4.90	4.50	5.25
Σ AE	4.67	5.40	7.39	13.37	19.30	11.75	11.35	4.25	9.18	6.93	11.20	10.28	9.37	5.98
SATFA														
12:0	0.78	0.65	0.10	-	-	-	0.25	-	0.19	-	0.16	0.14	0.18	-
13:0	0.35	0.27	0.05	-	-	-	0.07	-	0.08	-	0.11	0.13	0.12	0.11
14:0	7.34	6.85	1.36	0.98	1.43	0.94	2.00	1.06	1.60	1.12	1.81	1.82	1.58	1.78
15:0	2.69	2.27	0.63	0.54	0.77	0.54	0.82	0.60	0.73	0.61	0.91	0.83	0.63	0.71
16:0	28.08	20.90	6.38	5.77	8.21	5.73	8.65	7.04	6.14	7.29	8.32	6.46	8.03	8.82
17:0	2.52	1.16	0.33	0.47	1.58	0.40	0.53	0.52	0.46	0.53	0.52	0.62	0.57	0.65
18:0	10.40	2.49	1.75	2.55	3.68	2.30	3.32	2.89	3.53	3.09	2.69	3.48	3.97	4.33
19:0	2.70	5.54	0.28	0.50	0.56	0.93	0.71	0.90	0.51		0.60	0.62	0.73	0.76
MUFA														
16:1 ω ?	16.26	4.27	0.32	0.39	0.48	0.41	0.35	0.37	0.40	0.31	0.36	0.44	0.47	0.54

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16:1ω7	2.67	0.88	-	-	-	-	-	-	-	-	-	-	-	-	-
16:1ω5	5.72	2.40	0.19	-	-	-	-	-	-	-	-	0.21	0.20		
18:1ω9	12.09	0.83	0.48	1.28	0.73	1.00	0.92	0.99	0.95	1.25		0.91	0.98	0.67	0.78
18:1ω7	24.80	9.96	0.40	0.59	0.41	-	-	-	-	-		0.45		0.84	0.86
19:1ω?	10.61	3.05	1.22	4.46	1.78	1.09	3.91	2.50	2.90	-		2.29	4.27	4.01	3.60
PUFA															
18:2ω?	-	-	-	-	-	-	-	-	-	-		0.44	0.56	0.67	-
20:4ω6	3.77	3.05	-	-	-	-	-	-	-	-		-	-	-	-
20:5ω3	3.71	5.54	-	-	-	-	-	-	-	-		-	-	-	-
brFA															
i13:0	0.43	0.24	0.06	-	-	-	0.09	-	0.06	-		0.10	0.10	0.07	-
ai13:0	0.58	0.40	0.05	-	-	-	0.13	-	0.08	-		0.16	0.15	0.10	-
i15:0	7.68	3.37	0.47	0.61	0.77	0.75	0.79	0.55	0.53	0.45		0.63	0.66	0.55	0.60
ai15:0	14.39	7.55	0.90	1.17	0.54	0.54	1.69	1.12	1.24	1.17		1.56	1.41	1.14	1.21
3Me15:0	-	-	-	-	-	-	0.25	-	-	-		0.29	0.52	0.35	0.36
i16:0	3.18	2.40	-	-	-	0.53	0.70	0.78	-	-		0.56	0.67	0.51	0.47
i17:0	2.13	1.36	0.18	0.43	0.27	0.38	0.29	0.32	0.39	0.85		0.30	0.29	0.27	0.26
ai17:0	1.72	2.46	0.28	0.56	0.34	0.40	0.50	0.39	0.47	0.53		0.47	0.50	0.41	0.40
cyFA															
cy17:0	-	4.36	-	-	-	-	-	-	-	-		-	-	-	-
AE															
Phytane	0.76	0.64	0.72	1.43	1.63	0.86	1.24	0.47	0.85	0.81		0.94	1.51	1.30	0.98
C40:0	2.28	2.72	3.47	5.80	8.25	4.75	5.04	1.79	4.24	3.19		5.62	4.40	4.26	2.95
C40:1	0.36	0.51	0.44	1.67	1.74	1.44	0.92	0.45	0.71	0.51		0.81	0.76	0.75	0.31

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C40:2	0.70	0.81	1.45	2.31	3.69	2.50	2.20	0.88	1.74	1.23	2.03	1.86	1.41	0.88
C40:3	0.57	0.72	1.31	2.16	3.99	2.21	1.94	0.67	1.64	1.19	1.79	1.75	1.65	0.87

5. Discussion

Methane is widespread within upper Bantry Bay, as shown here through both acoustic evidence and millimolar concentrations of CH₄ in core samples. Migration of gas-rich fluids towards the seafloor is interpreted to have led to the formation of pockmarks, which we describe here for the first time. Detailed geochemical analysis of porewater samples coupled to results from the gas analysis of sediment plugs, depict strong SMTZs occurring in two of three sediment core locations. Results from the third core, VC25, suggest that a similar SMTZ likely occurs below the maximum penetration depth of the vibrocorer. Lipid biomarker analysis provides evidence of the presence of active communities of both SRB and archaea within these sediment cores. These archaea are potentially anaerobic methanotrophs (ANME) which are likely involved in AOM, contributing to the prevention of regular methane seepage above the seafloor as evidenced by the distinct SMTZs.

Sub-bottom profiles provide evidence of vertical gas migration through the sediments of upper Bantry Bay, although no gas signals were observed within the water column, geochemical data provide evidence of low concentrations of gas just beneath the seafloor. In the area of the pockmark field where our sediment cores were obtained, we observed vertical zones of acoustic blanking (AB) beneath strong reflectors at varying depths below the seafloor, which we interpret as typical gas chimneys (Fig. 3). The observation of chimney-like features as well as of blanking below enhanced reflectors suggest upward fluid migration is predominant at this location (Szpak et al., 2012). Similar acoustic chimneys were observed to rise to within a few metres of the seafloor across upper Bantry Bay above water depths of at least 65 m by Plets et al. (2015, their Fig. 10a), which we interpret to indicate the upward migration of gas from depth over wide areas beneath the pockmark field. However,

our results do not support the suggestion of Plets et al. (2015) that the presence of gas may account for the reflective character of their unit III, penetrated by our cores at ca. 2-6 mbsf (Fig. 3), which we instead suggest is due to its laminated lithological character (SI Fig. S1).

All gas headspace samples yielded undetectable amounts of C₂-C₄ hydrocarbons, indicating a biogenic source, rather than thermogenic source for gas in Bantry Bay (Faber and Stahl, 1984; Floodgate and Judd, 1992). The likely origins of this biogenic gas are microbial decomposition of buried organic matter and methanogenesis (Antler et al., 2014; Froelich et al., 1979). River run-off likely delivers a significant amount of OM to the bay. However, the majority of this terrestrially derived OM is likely consumed in the surface sediments as seen from the TOC results obtained from VC24. As such, this OM is probably not a large contributor to CH₄ generation within Bantry Bay. Previous work in Bantry Bay encountered black lacustrine sediments at ca. 57 m water depth (ca. 25 mbsf) which were dated to 13-14 ka cal. BP within a borehole off Whiddy Island (Stillman, 1968). Plets et al. (2015) found that these deposits occurred within Unit 2 of their assigned seismo-stratigraphy profile. They suggested that the material was likely older than the value provided by Stillman (1968) as they were situated below acoustic Unit 4 which was described as a possible glacial till, thereby placing Unit 2 in the position of a pre-Last Glacial Maximum (LGM) deposit. The LGM is defined as 26.5-19 ka BP (Clark et al., 2009). These sediments likely undergo enhanced anaerobic decomposition and methanogenic activity due to their high organic content which makes them favourable candidates for the source of the gas observed in this area, however this awaits further investigation.

The pockmark field north of Whiddy Island is comprised of very shallow depressions of ca. 0.3 m depth. Due to the substantial gas activity observed in these

sediments it is likely that biogenic CH₄ resulting from the decomposition of organic material, possibly from ancient lacustrine deposits buried deep beneath the seafloor, was the primary cause of pockmark formation. Although no active seepage from pockmarks to the water column was observed in this study it is still possible that some of them are actively venting. Wheeler (2002) determined that significant currents regularly resuspend the seabed near Whiddy island. If the pockmarks were inactive, this could suggest that they have been filled in by fresh sediment. However, recent work suggests that inactive pockmarks can in fact be kept open by ocean currents (Hammer et al., 2009; Pau et al., 2014). Many studies have proposed that accumulation of large volumes of gas below the seafloor followed by periodic large expulsions is the predominant cause of pockmark formations (Cole et al., 2000; Dondurur et al., 2011; Gay et al., 2007; Hovland et al., 2002; Hovland and Judd, 1988). As such, these events likely reform the present features in Bantry Bay and potentially form new features as well. Further bathymetric analysis of this site is required to determine the precise layout of this field and the exact number of pockmarks within it, as well as regular monitoring of this area to determine the level of gas seepage activity which may represent a potential hazard to any planned economic activity in the bay.

A pockmark field in a similar setting has been described in Dunmanus Bay, south of Bantry Bay (Szpak et al., 2015). These authors showed that the pockmarks were associated with CH₄ emissions and argued that the source of the gas was an underlying Dunmanus Fault, via a venting mechanism involving seal failure-renewal cycles. The CH₄ from Dunmanus contained only trace levels of C₂-C₄ hydrocarbons and it was suggested that methanogenesis also contributed to the gas in this location. The Bantry Bay pockmark field overlies the Owenberg River Fault, which runs along the northwest of Whiddy Island. It is possible that venting in Bantry Bay is controlled

449 by a similar bedrock faulting mechanism as proposed in Dunmanus. Both bays lie in
450 the South Munster Basin and are similar in their geology (Vermeulen et al., 2000).
451 However, a higher contribution of low molecular weight hydrocarbons would be
452 expected if this gas was predominantly thermogenic. Therefore, it is most likely that
453 the gas observed at both of these sites is a combination of thermogenic gas release
454 from underlying faults and biogenic gas produced by methanogenic communities
455 feeding on deeply buried organic material.

456 Both VC24 and VC27 yielded clear SMTZs, where CH₄ diffusing upwards
457 from depth first encounters SO₄²⁻ diffusing downwards from the ocean, which reflect
458 the depth of maximal anaerobic oxidation (Antler et al., 2014; Lin et al., 2016;
459 Valentine, 2002). The decreasing trend of SO₄²⁻ in VC25 suggests complete depletion
460 coinciding with a SMTZ at ca. 10 mbsf. The sub-bottom profile suggests that there is
461 no significant gas penetration into this core whereas there is gas penetration observed
462 in the core location of VC24. This is consistent with the significantly lower CH₄
463 concentrations within the VC25 samples. Analysis of sediments from the deeper
464 SMTZ in VC25 would likely yield similar CH₄ concentrations to that of VC24 and
465 VC27. Thus the three cores are indicative of variable rates of upward penetration of
466 gas-rich fluids towards the seafloor.

467 These SMTZs suggest that microbial communities are consuming CH₄ rising
468 from depth as well as SO₄²⁻ diffusing downward from the seafloor above. This
469 signature represents the metabolic pathways of microorganisms involved in the AOM,
470 namely ANME and SRB. At present, it appears that the activity of these microbial
471 communities aids in preventing the release of CH₄ to the water column and potentially
472 the atmosphere on a regular basis, reducing the potential impact of this powerful
473 greenhouse gas on global climate. However, as previously mentioned, the pockmark

474 features are indicative of possible recurring episodic expulsions of gas from these
475 sediments and as such the overall CH₄ flux from this site is poorly constrained. This
476 is a scenario which is observed in shallow marine seepage environments around the
477 world. It is important for these unique environments to be monitored so that their
478 potential contribution to climate change can be better understood.

479 PLFA biomarker results provide further evidence of this ongoing microbial
480 activity. High levels of MUFAs and low levels of PUFAs are an indication of the
481 dominant contribution of bacterial communities to sediment biomass (Rajendran et al.,
482 1995, 1992; Taylor and Parkes, 1983; Volkman et al., 1980). Bacteria appear to
483 dominate the microbial ecology in all three vibrocores in this study. Abundances of
484 PUFAs are increased in the surface sediments of VC24, however MUFA abundances
485 are still higher. Interestingly, at 0.8 and 3.9 mbsf in VC24, contributions of MUFAs
486 and PUFAs are similar although MUFAs remain dominant. Comparison of MUFAs
487 (<C₁₉) with total brFAs provides an insight to the aerobic/anaerobic conditions in the
488 sediment. Values less than 1 indicate an anaerobic environment whereas values greater
489 than 1 are representative of aerobic conditions (Rajendran et al., 1992). Only the
490 shallower sediments of VC24 (1.12 and 1.08 for 0.03 and 0.27 mbsf respectively) are
491 classified as aerobic using this approach, therefore the overall conditions observed
492 here are anaerobic.

493 Mid-chain brFAs in marine sediments are often produced by SRB and are used
494 as chemotaxonomic markers for these microorganisms (Dowling et al., 1986; Li et al.,
495 2007). *iC*_{15:0}, *aiC*_{15:0}, *iC*_{16:0}, *iC*_{17:0}, and *aiC*_{17:0} are all reported biomarkers for the
496 *Desulfovibrio* species of SRB (Dowling et al., 1986; Findlay et al., 1990; Li et al.,
497 2007; Rajendran et al., 1995; Taylor and Parkes, 1983). These compounds were
498 present throughout all three vibrocores taken in Bantry Bay, suggesting a significant

499 contribution of SRB to the microbial ecosystem here. SRB tend to display a higher
 500 ratio of $iC_{15:0}$ to $aiC_{15:0}$ in their PLFA profiles (Dowling et al., 1986). Applying this
 501 ratio to the Bantry Bay sediments resulted in substantially higher values observed in
 502 VC24 than VC25 or VC27, both of which had similar values. This suggests that SRB
 503 are substantially higher contributors to the microbial community of VC24. Elvert et al
 504 determined that $C_{16:1\omega5}$, $C_{17:1\omega6}$, and $cyC_{17:0\omega5,6}$ were specific membrane fatty acids for
 505 SRB of the *Desulfosarcinia/Desulfococcus* group which were involved in AOM
 506 (Elvert et al., 2003). $C_{16:1\omega5}$ was identified in four samples from the Bantry Bay
 507 vibrocores; VC24 0.03 mbsf, VC24 0.27 mbsf, VC27 1.93 mbsf, and VC27 2.96 mbsf.
 508 Isotope ratio analysis provided a depleted $\delta^{13}C$ value of -46.1‰ for this compound at
 509 0.27 mbsf in VC24, compared with those obtained from more ubiquitous bacterial
 510 PLFAs; $16:1\omega7$ and $16:0$ at -31.0‰ and -27.7‰ respectively. While lipid abundances
 511 are low, these depleted values still indicate the possible incorporation of CH_4 derived
 512 C into the membranes of these SRB providing evidence of their involvement in AOM
 513 at this site.

514 Archaeal biomarkers are present at all depths in all three vibrocores taken in
 515 Bantry Bay. Archaea involved in AOM typically belong to three major anaerobic
 516 methanotroph (ANME) consortia (Caldwell et al., 2008). It is generally assumed that
 517 ANMEs oxidize and assimilate CH_4 , following which CH_4 -derived C is consumed by
 518 the SRB as CO_2 or a partially oxidized intermediate completing the syntrophic reaction
 519 (Alperin and Hoehler, 2009). It has been suggested that C_{20} isoprenoids derived from
 520 archaeal ether lipids may be specific biomarkers for ANME-2 archaea whilst C_{40}
 521 isoprenoids may be specific for the ANME-1 type (Blumenberg et al., 2004; Brocks
 522 and Pearson, 2005). The presence of these compounds within the gas-rich sediments
 523 in Bantry Bay further suggests the involvement of AOM mediated by ANME and SRB

524 in limiting gas release at this site. More detailed biogeochemical analysis of these
525 sediments could shed more light on the composition of this particular microbial
526 community with isotopic analysis determining their contribution to AOM. Due to the
527 importance of understanding the microbial community structure at shallow gas
528 seepage sites like Bantry Bay, a more detailed phylogenetic study of this site is
529 recommended.

530

531 **6. Conclusions**

532

533 The upward migration of gas-rich fluids through the sediment column appears
534 to be widespread in upper Bantry Bay, as inferred from chimney-like acoustic zones
535 on sub-bottom profiles and confirmed by shallow SMTZs within sediment cores.
536 Shallow SMTZs are observed both within and outwith a newly identified pockmark
537 field, suggesting that diffuse pore fluid upwelling over wide areas is only locally
538 accompanied by focused flow within conduits. Methanogenesis is taking place within
539 organic-rich Quaternary sediments deposited across the upper Bay prior to and since
540 the last deglaciation. The presence of pockmarks off Whiddy Island may be explained
541 by enhanced gas flux from the underlying Owenberg River Fault and methanogenesis
542 of organic-rich lacustrine sediments pre-dating the LGM that are preserved in bedrock
543 basins.

544 Fluid flow affects not only the physical nature of the sea-floor in the bay but
545 also the microbial ecosystem. The gas is CH₄ with a predominantly biogenic signature.
546 As CH₄ flows upwards from its origin it provides a substrate for certain
547 microorganisms to thrive in the shallower sediments above. Archaea, possibly
548 ANMEs, are present in these shallower sediments as are SRB. The CH₄ is steadily

depleted before it reaches the seafloor and SO_4^{2-} concentrations also become depleted in the opposite direction providing a well-defined SMTZ. This is likely due to AOM carried out by these two groups of microorganisms in a syntrophic relationship, however further work is needed to confirm this pathway. This study suggests that AOM in Bantry Bay is important in limiting CH_4 emissions from the seafloor preventing the potential climatic implications of a release of this powerful greenhouse gas to the atmosphere. Similar conditions have been observed in a pockmark field in Dunmanus Bay, to the east of Bantry Bay (Szpak et al. (2015)) and on the Malin Shelf off the north coast of Ireland (Szpak et al. (2012)). This indicates that marine CH_4 production may be common around the island of Ireland.

Global estimates of the contribution of CH_4 from marine seepage sites are highly uncertain (Römer et al., 2014). Release of CH_4 to the atmosphere has been observed in Arctic regions, areas particularly vulnerable to climate change, and these releases have been attributed to rising temperatures (Shakhova et al., 2010; Westbrook et al., 2009). As CH_4 is a potent greenhouse gas, these releases serve only to increase rates of global climate change. AOM and the microbial consortia involved are important factors in the global methane cycle (Gauthier et al., 2015). For these reasons further study of these sites and their microbial ecology should be prioritised.

Acknowledgements

The authors would like to thank the INFOMAR program, joint programme between the Geological Survey Ireland and the Marine Institute funded by DCCAE. We also thank the captain and crew of the R.V. *Celtic Explorer* during the GATEWAYS II campaign (CE14003), which was funded by a ship-time award to

Stephen McCarron under the Sea Change strategy with the support of the Marine Research Sub-programme of the Irish National Development Plan 2007-2013. We would like to thank the Irish Research Council for supporting the work of those researchers based at Dublin City University. The participation of OGS in the GATEWAYS II campaign was supported by Italian PNRA project IPY GLAMAR (grant number 2009/ A2.15), and Daniel Praeg also acknowledges funding from the European Union's Horizon 2020 research and innovation program under the Marie Skłodowska-Curie grant agreement No 656821 (project SEAGAS, 2016-2020). Bulk physical analyses of core samples was carried out at the Coastal and Marine Research Centre (CMRC) at University College Cork (UCC). We also thank Dr. Bart van Dongen and the anonymous reviewers whose input significantly improved this manuscript.

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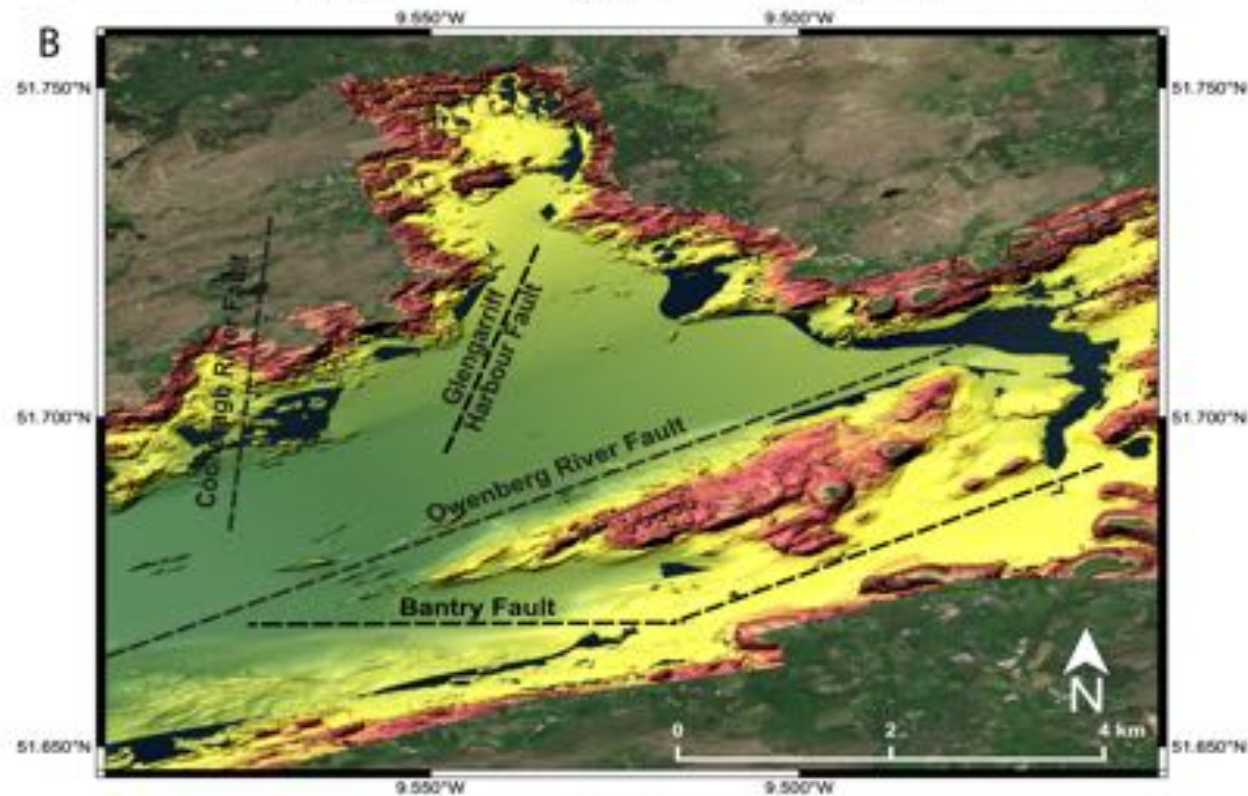
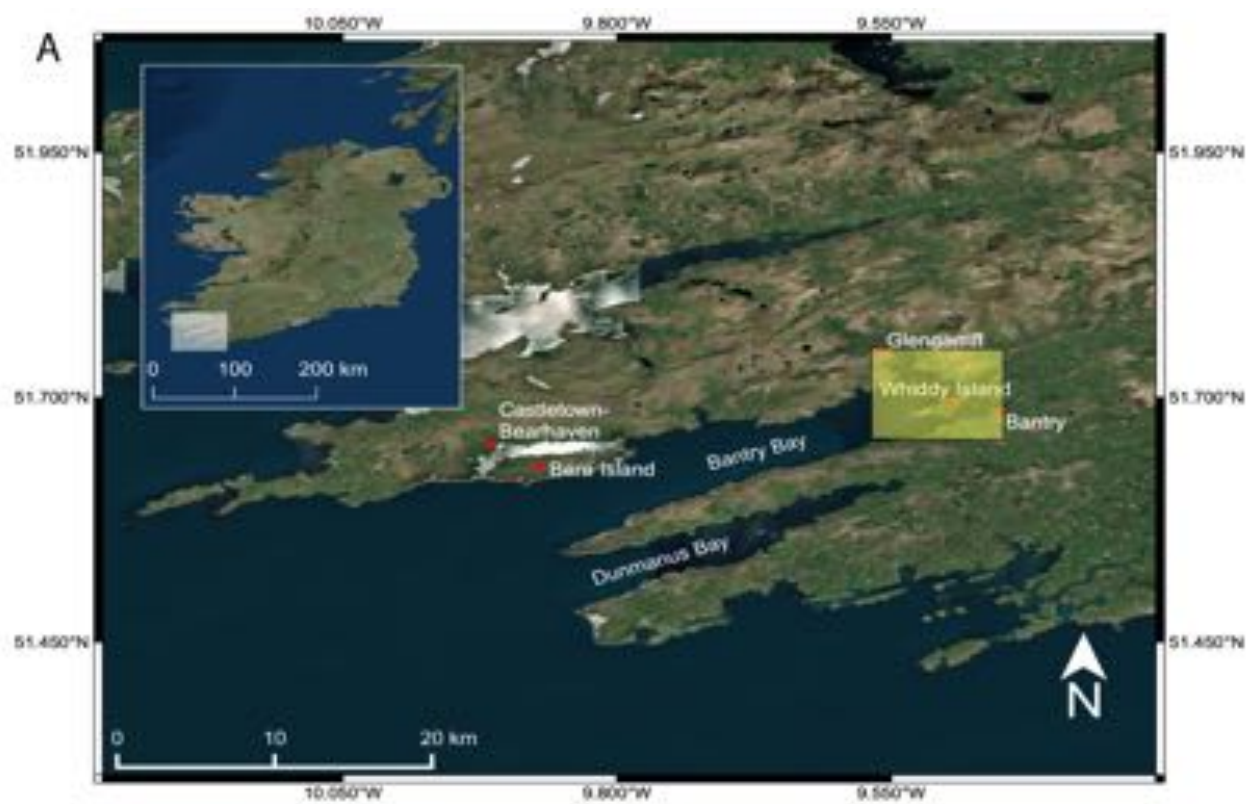
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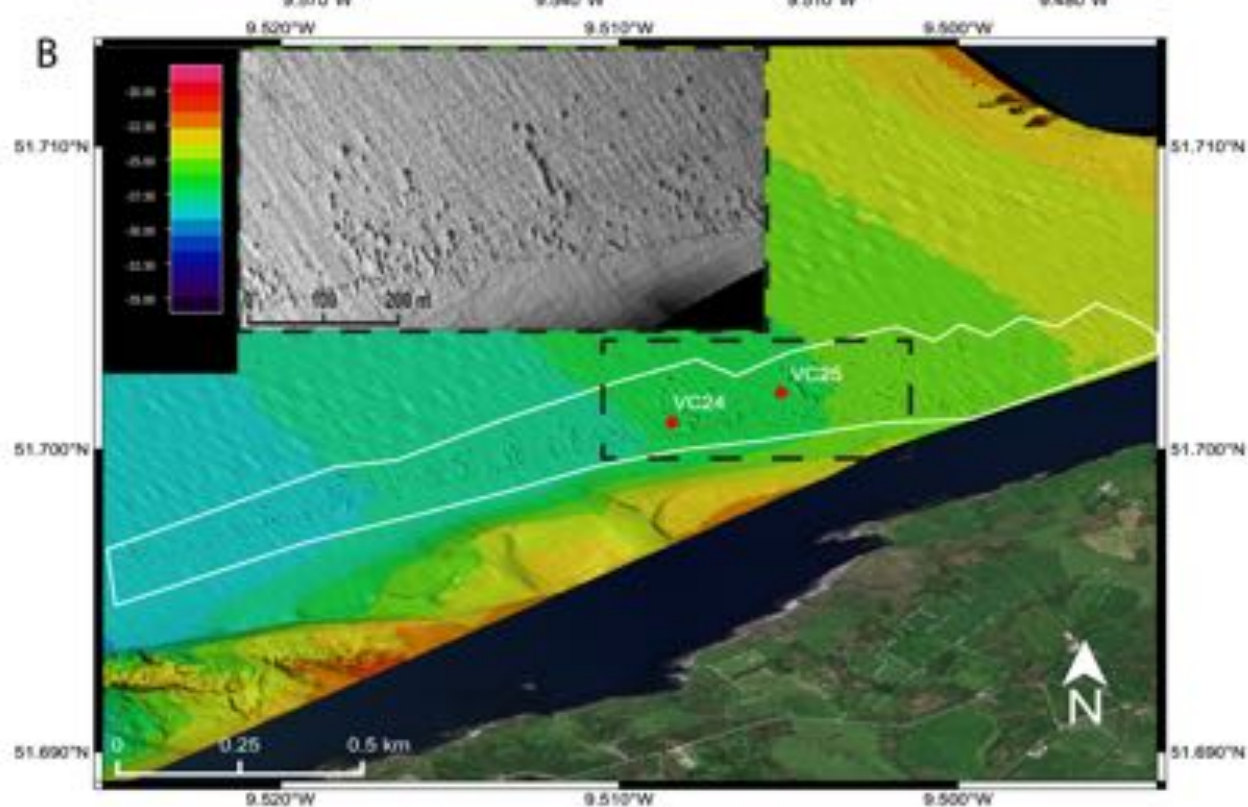
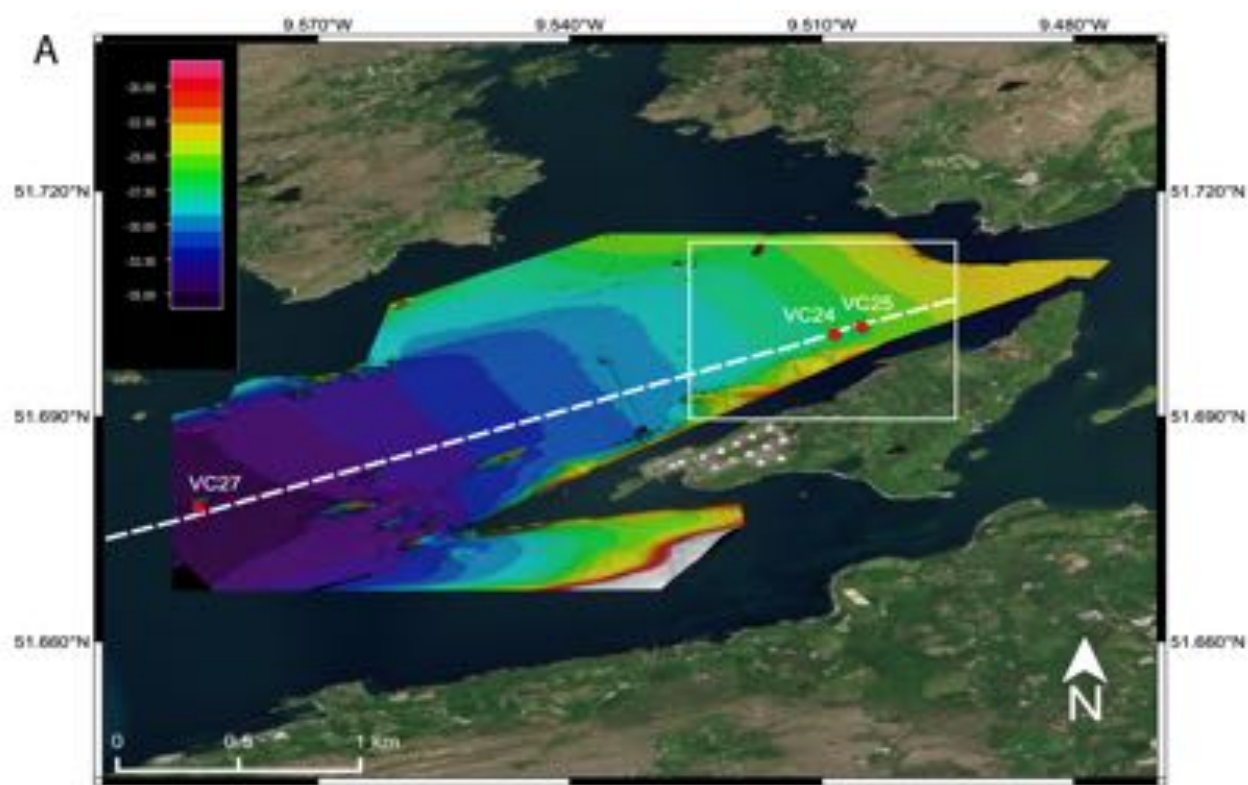
Figure 1. (A) Map of Bantry Bay and surrounding area, location of Bantry Bay within Ireland (inset). (B) Bathymetric map of inner Bantry Bay showing locations of underlying faults.

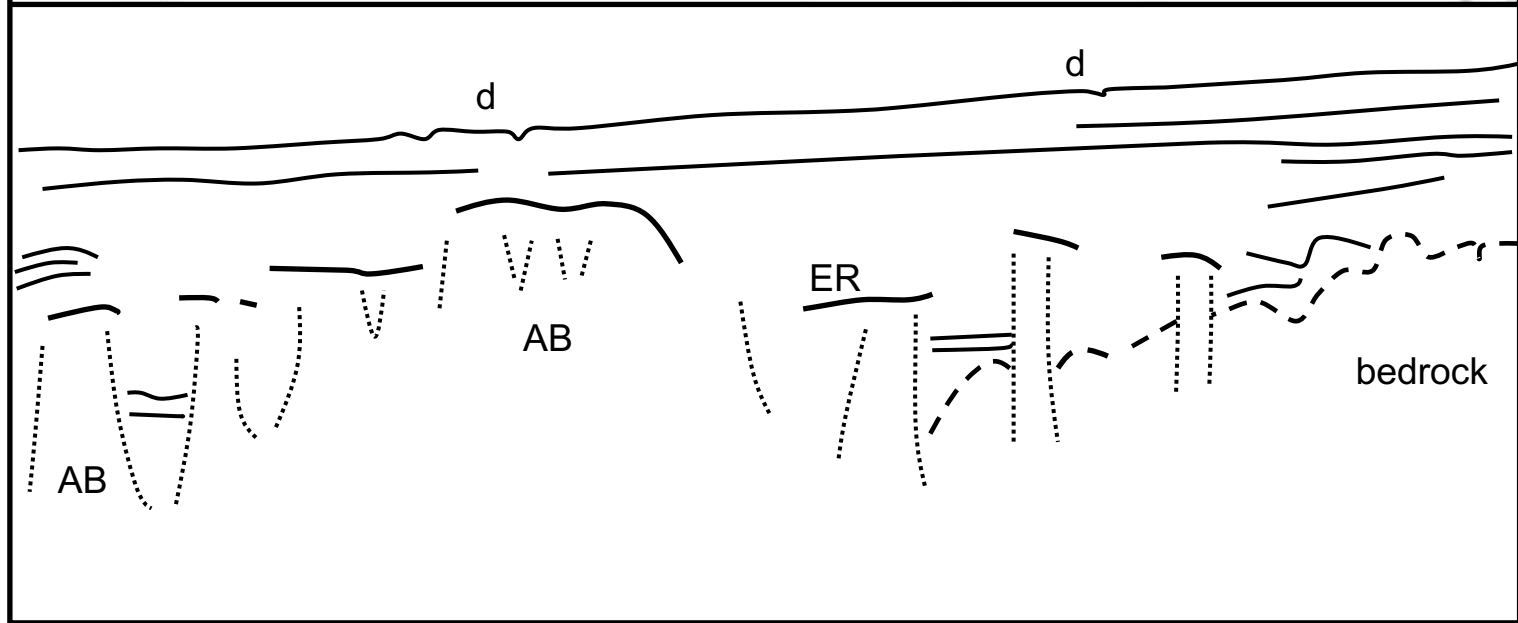
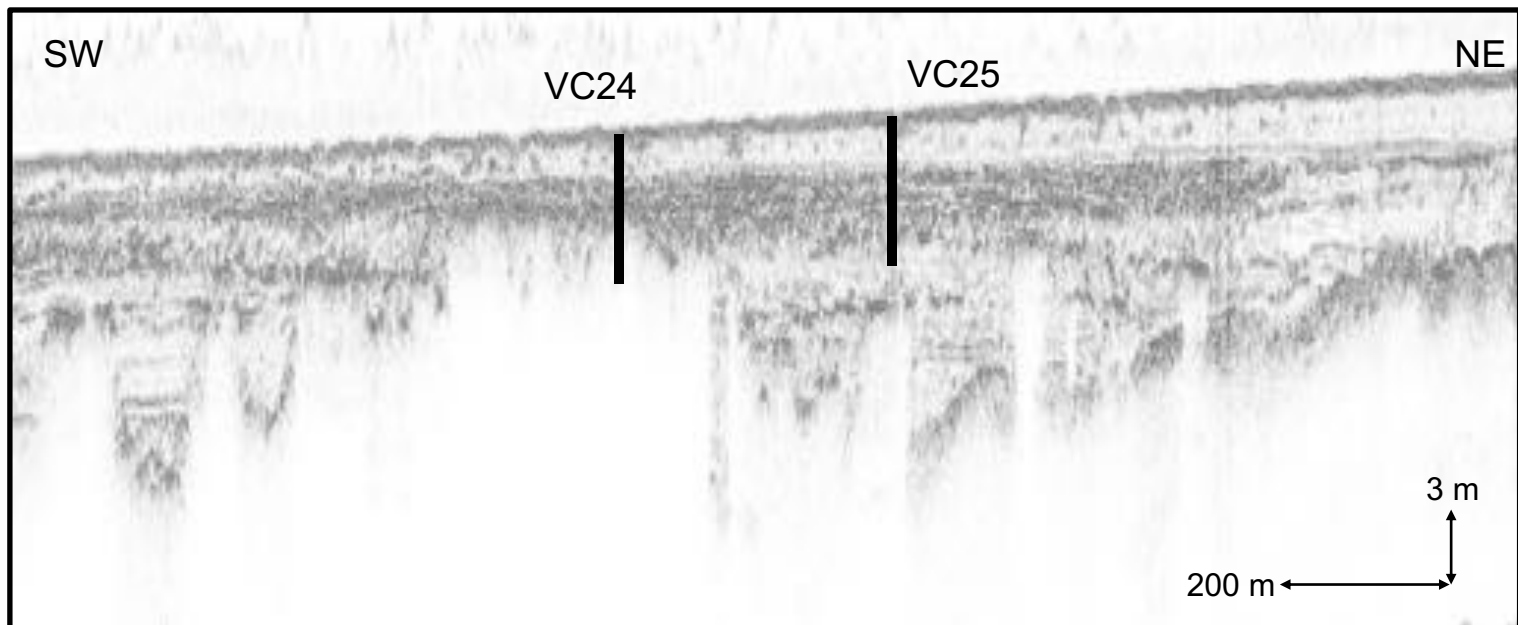
Figure 2. (A) Bathymetric map of inner Bantry Bay, vibrocore locations marked with red dots. The pockmark field north of Whiddy Island is located within the white box. (B) Close up of pockmark field from (A) with the entire field highlighted by a white outline and vibrocore locations marked with red dots. A close up of the section of the pockmark field within the black dashed rectangle is also depicted (inset).

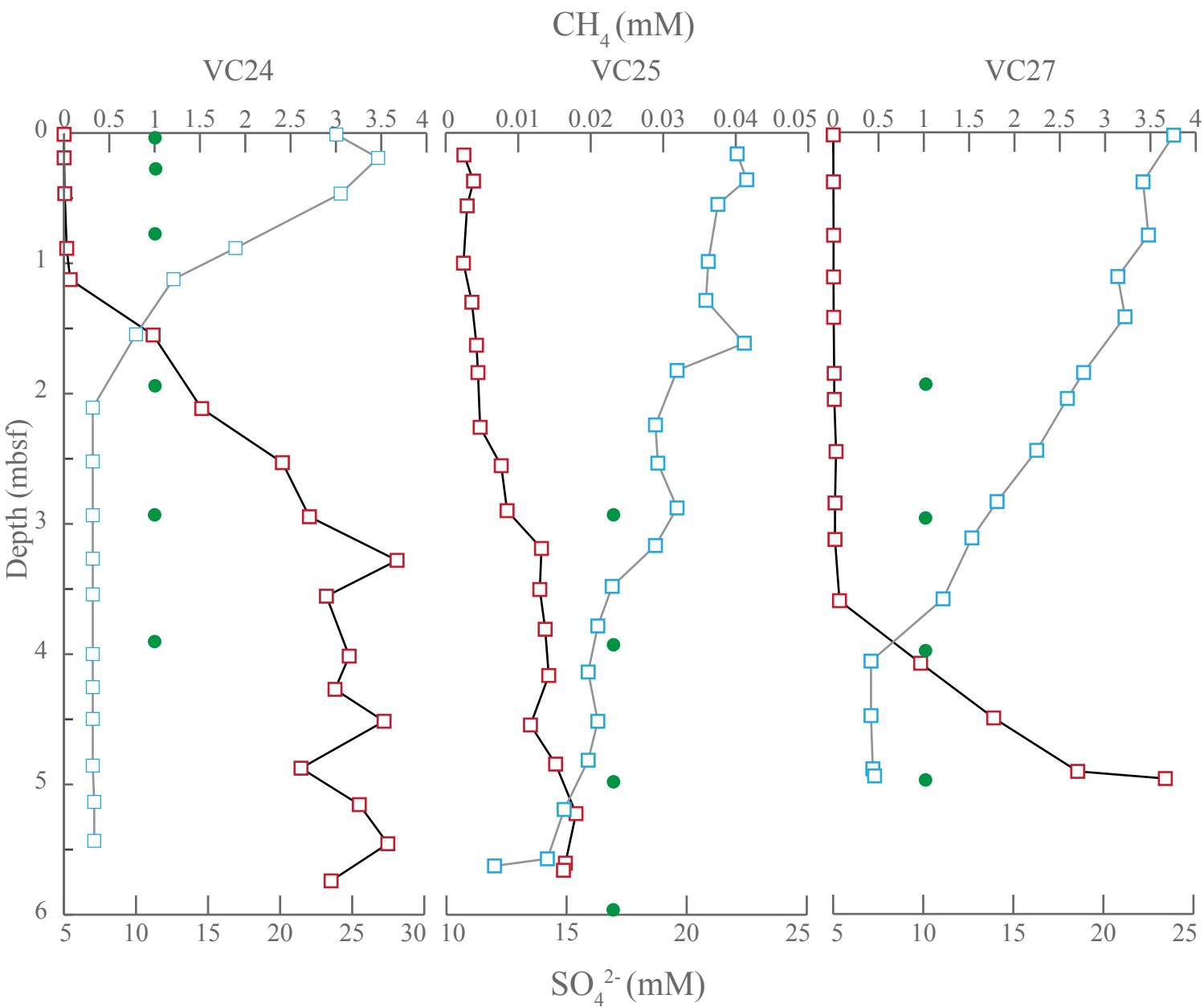
Figure 3. Sub-bottom profiles taken at the site of VC24 and VC25 vibrocores showing sampling locations (black line), enhanced reflectors (ER), and acoustic blanking (AB).

Figure 4. CH₄ (mM) and SO₄²⁻ (mM) profiles for each core. Green dots represent sub-sampling locations for lipid biomarker analysis.









- CH₄
- SO₄²⁻
- Biomarker Sample

Supporting Information

Geophysical and geochemical analysis of shallow gas and an associated pockmark field in Bantry Bay, Co. Cork, Ireland.

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Results

Table S1. CH₄ (μM) and SO₄²⁻ (mM) data from geochemical analysis of vibrocores.

VC24			VC25			VC27		
Depth (mbsf)	CH ₄ (μM)	SO ₄ ²⁻ (mM)	Depth (mbsf)	CH ₄ (μM)	SO ₄ ²⁻ (mM)	Depth (mbsf)	CH ₄ (μM)	SO ₄ ²⁻ (mM)
0.01	1.5	23.9	0.17	2.5	22.1	0.02	0.9	23.9
0.19	1.8	26.8	0.37	3.8	22.5	0.38	2.6	22.2
0.47	8.6	24.2	0.56	3.0	21.3	0.79	4.5	22.5
0.89	30.5	16.9	1.00	2.5	20.9	1.11	3.8	20.8
1.13	68.9	12.6	1.30	3.6	20.8	1.42	5.5	21.2
1.55	983.5	10.0	1.63	4.2	22.4	1.85	11.2	18.9
2.12	1519.6	7.0	1.84	4.4	19.6	2.05	12.4	18.0
2.53	2409.5	7.0	2.26	4.8	18.7	2.45	31.9	16.3
2.95	2707.7	7.0	2.56	7.7	18.8	2.85	20.3	14.1
3.28	3674.9	7.0	2.90	8.5	19.6	3.13	19.9	12.7
3.56	2895.5	7.0	3.19	13.2	18.7	3.60	69.6	11.1
4.02	3146.6	7.0	3.51	13.0	16.9	4.08	964.9	7.1
4.27	2989.9	7.0	3.81	13.7	16.3	4.50	1770.4	7.1
4.52	3531.9	7.0	4.17	14.2	15.9	4.91	2697.4	7.2
4.88	2615.2	7.0	4.55	11.7	16.3	4.96	3664.2	7.3
5.16	3258.2	7.1	4.85	15.2	15.9			
5.46	3573.0	7.1	5.23	18.0	14.9			
5.74	2946.9		5.61	16.5	14.2			
			5.66	16.2	12.0			

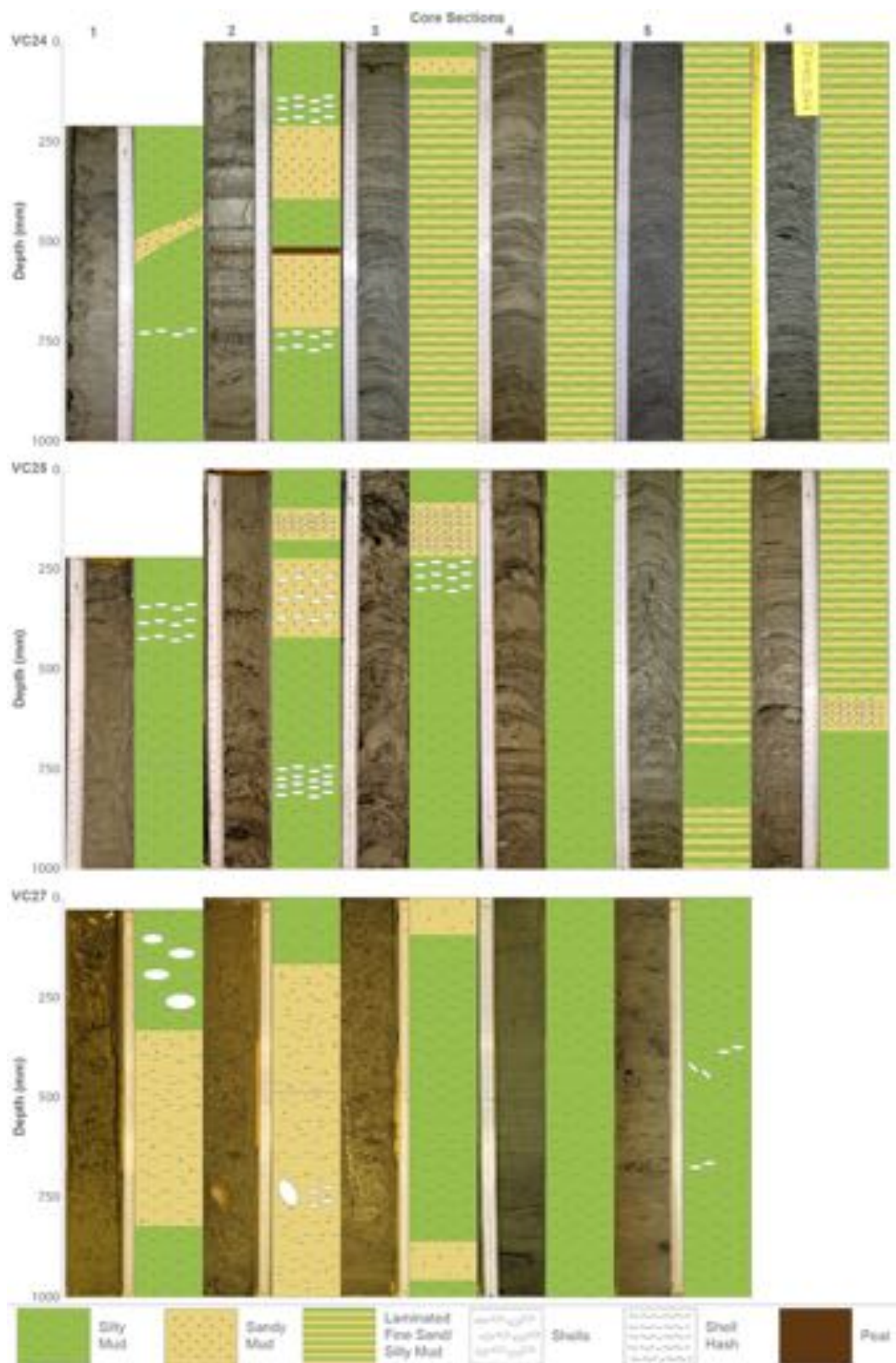


Fig. S1. Photographs of vibrocore sections taken onboard the *RV Celtic Explorer* during research cruise CE14003. Graphical depictions of sediment type from core logs are displayed alongside relevant sections.