

# **Palaeoecology and palaeoclimatic context of Romanian Carpathian MIS 3 cave bears using stable isotopes (***δ***13C and** *δ***18O)**

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ancient marine and terrestrial communities (e.g. Longinelli, 1984; Clementz, 2012; Matthews

36 et al., 2016; Martin et al., 2017). Carbon and oxygen isotopic abundances (reflected in  $\delta^{13}C$ 37 and  $\delta^{18}$ O values) measured from fossil vertebrate teeth have been used for a long time to decipher the dietary habits and ecology of extinct species (Sullivan and Krueger, 1981; Krueger and Sullivan, 1984; Ambrose and Norr, 1993; Longinelli, 1984; Koch et al., 1989; 40 Reinhard et al., 1996). Carbon isotopic abundances in carbonate hydroxyapatite  $[Ca_{10} (PO<sub>4</sub>,$  $41 \quad \text{CO}_3$ <sub>6</sub> (OH<sub>2</sub>)] reflect the average of the whole diet components (amino acids, carbohydrates and lipids). The oxygen isotopic abundances recorded from the same tissue reflect the 43 variation in the  $\delta^{18}$ O value of body fluids controlled mainly by the isotopic composition of the 44 ingested water. As the ingested water tracks the local precipitation,  $\delta^{18}$ O values obtained from dentine are used to reconstruct palaeoclimatic conditions and sources of water used by animals (Longinelli, 1984; Koch et al., 1989; Reinhard et al., 1996; Dotsika et al., 2011; Bocherens et al., 2011; Clementz, 2012; Krajcarz and Krajcarz, 2014; Krajcarz et al., 2014). Late Pleistocene cave bear fossils are the most abundant remains to be found in the caves of Europe, this extinct species being one of the best studied in Quaternary palaeontology. Nonetheless, there are still many debates within the scientific community regarding the diet, physiology and behavior of these ursids. The permanent canine teeth of adult bears are useful for tracing age-related changes, mostly due to their resilience to weathering. The chronological age of individuals can be accurately determined (Debeljak, 1996) from the annual layers of dentine or cementum, emphasized when odontogenesis is slowed during the annual period of denning (nutritional stress). Although the cementum layers record a longer history of the bear's life, dentine layers are better suited for geochemical analyses as they are thicker and can provide sufficient sampling material. It has been shown that the annual canine dentine layers in adult individuals may vary between five and eight, with the first five layers belonging to the primary dentine, and the others to the secondary dentine. Each dentine layer consists of a thick, light-stained inner zone, deposited during the annual period of activity, and a very thin, dark-stained outer zone related to the period of denning (Rausch, 1961, 1967, 1969; Mundy and Fuller, 1964; Stoneberg and Jonkel, 1966; Sauer, 1966; Sauer et al., 1966; Koch et al., 1989). Knowing that the cyclicity of dentine growth layers in bear canines ends when adulthood is reached (Rausch, 1969), one can trace back both the ecological conditions 65 and the climatic evolution by using a geochemical approach (e.g., using  $\delta^{13}$ C and  $\delta^{18}$ O values) for every identified dentine layer. The carbon isotope analysis of organic tissues of such tooth dentine provides information

reflective of the protein portion (collagen) of an individual's diet (Sullivan and Krueger, 1981;

Krueger and Sullivan, 1984; Ambrose and Norr, 1993; Harrison and Katzenberg, 2003),

 whereas stable isotope values from bone, dentine, and enamel apatite carbonate largely reflect the whole diet (*i.e.*, proteins, carbohydrates, and lipids; Ambrose and Norr, 1993; Harrison and Katzenberg, 2003). In general, apatite is more 13C-enriched compared to collagen (by +9,6-13‰; DeNiro and Epstein, 1978a, b; Sullivan and Krueger, 1981; Krueger and Sullivan, 1984; Lee-Thorp and Van der Merwe, 1987; Fahy et al., 2015), their relationship being affected by the trophic level. The utility of bone/tooth apatite for stable carbon isotope analysis in palaeodietary studies stems from the fact that previous results indicate that diagenesis of biogenic carbonates in the mineral over time is unexpectedly limited, and that chemical pretreatment further removes the effects of diagenetic alteration of the biogenic signal (Lee-Thorp et al., 1989). Therefore, palaeodietary inferences could be carried out using 80 this method for dentine apatite for fossil remains up to  $\sim 10^5$  years old (Clementz, 2012). Thus, 81 apatite  $\delta^{13}$ C measurements should not be regarded as an alternative to  $\delta^{13}$ C collagen measurements, but may provide additional useful information about dietary trends. Previous studies have focused on cave bear physiology (e.g. Nelson et al., 1998; Bocherens, 84 2015; Robu et al., 2018) using  $\delta^{13}$ C values from bone collagen and have shown that the isotopic profiles of the four age groups (neonates, juveniles, sub-adults and adults) provide patterns of transfer of stable isotope signatures throughout the immature life of bears. 87 Studying the  $\delta^{13}$ C transfer signature throughout the first five to eight years of life, using  $\delta^{13}$ C data from apatite of each individual, could bring a better understanding of MIS 3 cave bear physiology and a better resolution of this process. Sequential sampling of distinct dentine layers within a single tooth can provide information on dietary and habitat changes over the course of an individual development from juvenile to adult as well as seasonal variations in these ecological parameters later in the animal's life. Previous studies of fossil proboscideans – elephants and their extinct relatives (Koch et al., 1989; Fox and Fisher, 2004; Hope et al., 1999; Rountrey et al., 2007; Clementz, 2012) highlighted how serial sampling of both tooth enamel and dentine can be used to answer different questions about the ecology and life history of these animals. Koch et al. (1989) 97 analyzing the  $\delta^{18}$ O values from dentine layers, have shown that CO<sub>3</sub> fraction of apatite retains a signal in molars (proboscideans and bears), reflecting seasonal changes in composition of drinking water. The Romanian Carpathians host numerous, complex, and rich Late Pleistocene cave bear

sites, suitable for palaeoecological and palaeoclimatological investigations. Three already

famous Late Pleistocene cave bear sites from Romanian Carpathians were chosen for this

103 study – the Ur $\Box$ ilor, Oase and Cioclovina caves – known to the international scientific

- community for their impressive fossil bonebeds, human and faunal remains, and for their
- taphonomic and palaeoichnologic features (e.g. Richards et al., 2008; Robu et al., 2018).
- 106 Analyzing both  $\delta^{13}C$  and  $\delta^{18}O$ , retrieved from dentine apatite serial samples, has not been
- applied systematically to any MIS 3 species as yet. Thanks to their large dimensions,
- analyzing a large set of MIS 3 cave bear canines provides an excellent record of seasonal
- variations in diet and climate during tooth growth and strengthen our understanding of the
- 110 palaeoclimatic and palaeoecological significance of  $\delta^{13}$ C and  $\delta^{18}$ O values retrieved from
- dentine apatite.

#### **MATERIALS AND METHODS**

*Carpathian site-specific samples* 

The Carpathian cave bear specimens derive from excavations in three extensive karst systems:

- Peştera Urşilor de la Chişcău, Peştera Cioclovina Uscată, and Peştera cu Oase (Fig. 1).
- 116 The Pestera Urşilor de la Chişcău (further on: Urşilor;  $46^{\circ} 55 \square 37 \square N$ ,  $22^{\circ} 56 \square 92 \square E$ ) is
- situated in the northwestern part of the Romanian Carpathians (the Apuseni Mountains). The
- analyzed cave bear material was sampled from the palaeontological excavation, situated
- within the Scientific Reserve (lower level of the cave). The analyzed cave bear population
- was AMS 14C dated between *c.*47 and 39 kyr cal BP (Constantin et al., 2014; Fig. 1A).
- 121 The Peștera Cioclovina Uscată (*Dry Cioclovina Cave*; further on: Cioclovina; 45° 34 $\Box$  28 $\Box$
- 122 N,  $23^{\circ}$  08  $\Box$  04  $\Box$  E) (Fig. 1B) is located in the northwestern part of the Southern Carpathians
- (the Şureanu Mountains) (Soficaru et al., 2007; Onac et al., 2011). Results of AMS 14C dating
- indicate that the cave bear population inhabited the cave between *c.*44.2 and 31.4 kyr cal BP
- (Soficaru et al., 2007; Meleg et al., 2018; this study).
- 126 The Pestera cu Oase (*Ponor Plopa Cave*; further on: Oase;  $45^{\circ}$   $05 \square 21 \square N$ ,  $21^{\circ}$   $51 \square 21 \square E$ )
- is situated in the southwestern part of the Southern Carpathians (the Anina Mountains) (Fig.
- 128 1C). The analyzed cave bear samples were excavated from the upper level of the cave (Sala
- Mandibulei and Panta Strămoşilor). The cave bear population of the bone assemblage was
- dated between *c.*50 and 40.4 kyr cal BP (Higham and Wild, 2013; this study).
- 131 <Insert Figure 1>
- *Cave bear samples*
- Teeth of the analyzed cave bears were identified by their morphology (according to Torres *et*
- *al*., 1978; Tsoukala and Grandal-d'Anglande, 2002). DNA analyses have not been completed
- for all studied sites, although some populations were attributed to *U. ingressus* either on
- mtDNA or nuclear DNA basis or dental and metacarpal morphology (Richards et al., 2008;
- Robu, 2016; Meleg et al., 2018).
- 138 A number of 18 fused cave bear lower canines  $(N = 18)$  were analyzed to obtain carbon and
- 139 oxygen isotopic composition from the structural carbonate of their dentine. The canines were
- 140 selected from the analyzed sites ( $N = 9$  females and  $N = 9$  males; Table S1), with six canines
- 141 being sampled from each collection. Hydroxyapatite samples  $(N = 112)$  were drilled from
- 142 series of holes aligned parallel to incremental dentine layers (Fig. 2); the powder from each
- 143 row of holes formed a single sample (*c*. 2 mg). We obtained five to nine non-overlapping
- 144 samples within each inferred year. The samples were analyzed following the method
- 145 described by Koch et al. (1997).
- 146 <Insert Figure 2>
- 147 The isotopic ratios are expressed as  $\delta^{13}$ C and  $\delta^{18}$ O values according to calculation:  $\delta^{A}$ X =
- 148 (R<sub>sample</sub>/R<sub>standard</sub> 1) × 1000, where X is an element (C or O), A is the atomic mass number and
- 149 R is the isotopic ratio  ${}^{13}C/{}^{12}C$  or  ${}^{18}O/{}^{16}O$ . Values are reported in relation to Vienna-PDB
- 150 (VPDB) standard, with  $\delta^{18}$ O values being converted to VSMOW using the formula
- 151  $\delta^{18}O_{VSMOW} = 1.03091 \times \delta^{18}O_{VPPDB} + 30.91$  (Skrzypek et al., 2011). Laboratory standards used
- 152 were CO<sub>2</sub>:He for  $\delta^{13}C$ , and NBS18 and CARRARA 11 for  $\delta^{18}O$ . Both  $\delta^{13}C$  and  $\delta^{18}O$  values
- 153 were measured using a Delta V+ (Thermo-Finnigan, Bremen, Germany) isotope ratio mass
- 154 spectrometer (IRMS), in continuous flow mode (Fry et al., 1992) coupled to an ECS410
- 155 elemental analyzer (Costech Analytical, Valencia, CA, USA) in the Department of Geology,
- 156 University of South Florida.
- 157 *Dating method*
- 158 Ten cave bear samples ( $N = 10$ ) derived from the analyzed canines were AMS <sup>14</sup>C dated at
- 159 Poznan Radiocarbon Laboratory (Poland; Tab. S1). At the Poznan Laboratory, collagen
- 160 extraction was performed using the method described by Pietrowska and Goslar (2002) and
- 161 the extracted collagen was then purified on Elkay 8 mm filters, and ultrafiltered on Vivaspin
- 162 15 MWCO 30 kD filters (Bronk Ramsey et al., 2004). The content of <sup>14</sup>C in carbon samples
- 163 was measured following the protocol described by Goslar et al. (2004).
- 164 *Statistical analyses*
- 165 Statistical analyses were performed using the software PAST (Paleontological Statistics),
- 166 version 3.20 (Hammer *et al*. 2001). Hierarchical clustering was used to explore relationship
- 167 patterns of  $\delta^{18}$ O variability among the studied samples. Euclidean distances were used to
- 168 calculate a similarity matrix based on the average values of  $\delta^{18}$ O for 18 canines from Ur $\Box$ ilor,
- 169 Oase and Cioclovina caves, and Ward's method was selected to define the distance between
- 170 identified classes. To visualize the similarity among studied samples, multidimensional
- 171 scaling and minimum-spanning tree linking samples were computed based on average values
- 172 of  $\delta^{18}O$ ,  $\delta^{13}C$  and median <sup>14</sup>C date for 18 canines from Ur $\Box$ ilor, Oase and Cioclovina caves.

### 173 **RESULTS**

- 174 The results of radiocarbon dating are given in Table S1 (*N* = 10 samples). All ages were
- 175 calibrated using the OxCal 4.3.2 program (Bronk Ramsey, 2017) and the IntCal13 calibration
- 176 curve (Reimer et al., 2013). Radiocarbon ages (Fig. 3) of the 10 cave bear bone samples
- 177 collected from the analyzed sites range between *c*. 45 and 31 kyr cal BP (median values) for
- 178 both surface and excavated finds. The majority of the dated samples of bears belong to 39-45
- 179 kyr cal BP interval known for its faunal richness across caves from both the Carpathians and
- 180 the rest of Eurasia and to a last pulsation of this species, around 31 kyr cal BP (*N* = 2; from
- 181 Cioclovina Cave) at a time believed to be close to its extinction (Pacher and Stuart, 2009).
- 182 <Insert Figure 3>
- 183 *δ13C values*
- 184 The obtained  $\delta^{13}$ C values of all 112 dentine apatite samples range from -17.5 to -9‰ (average
- 185 of all samples -12.8‰; Fig. 4A-C). All  $\delta^{13}$ C profiles are recording an increment of carbon
- 186 values throughout the deposition of cave bear dentine until adulthood. Ur**cultum** samples ( $N =$
- 187 36 dentine layers) have extreme values between -16.7 and -9.2‰, Oase samples (*N* = 38
- 188 dentine layers) spread between -14.9 and 9,0‰ and, Cioclovina samples (*N* = 38 dentine
- 189 layers) spread between -17.5 and -9.9‰.
- 190 <Insert Figure 4>
- 191 Comparing the  $\delta^{13}$ C values of the oldest dentine layers of the analyzed canines (precluding
- 192 adulthood in bears), with the average of the  $\delta^{13}$ C collagen values obtained for the same
- 193 populations (e.g. Robu et al., 2013, 2018), we found that the lowest  $\Delta \delta^{13}C_{\text{collagen-anatite}}$
- 194 difference  $(\delta^{13}C_{\text{collagen}} \delta^{13}C_{\text{anaitte}})$  was recorded for Cioclovina, +7,9‰  $(\delta^{13}C_{\text{collagen}} = -21.1)$
- 195 and  $\delta^{13}C_{\text{anative}} = -13.2\%$ , while the Ur $\Box$ ilor population had a  $\Delta \delta^{13}C_{\text{collagen-anative}}$  value of +9.6‰
- 196  $(\delta^{13}C_{\text{collagen}} = -21.6$  and  $\delta^{13}C_{\text{anaitte}} = -12.0\%$ . The highest average difference recorded in the
- 197 Oase population between the two parameters  $\Delta \delta^{13}C_{\text{collagen-apative}}$  was  $+11,3\%$  ( $\delta^{13}C_{\text{collagen}}$  = -

198 21,4 and  $\delta^{13}C_{\text{anaitie}} = -10.1\%$ , respectively).

- 199 Figure 5 shows the variation of  $\delta^{13}$ C values between sexes among the studied Carpathian
- 200 populations: for females, it appears that the highest average is recorded for the Oase and
- 201 Urşilor females (-11.9‰, and -12‰, respectively), and the lowest for Cioclovina, -14.6‰
- 202 (Fig. 5A). For males, the highest average value was measured in the Oase population: -
- 203 11.9‰, and lowest values in Ur $\Box$ ilor (-13.9‰) and Cioclovina (-13.7‰) (Fig. 5B). Both
- 204 graphs suggest differences among sexes within all three cave bear populations: for  $Ur \Box$ ilor,
- 205 females show  $\delta^{13}$ C values enriched by +1.9‰ when compared with males, while for Oase and
- 206 Cioclovina, the situation is opposite, males have enriched  $\delta^{13}$ C values by +0.5 and +0.9‰,
- 207 respectively. The differences in  $\delta^{13}$ C values among sexes could be due to either random and
- 208 size sampling (only three males and females analyzed/site), to different intra-population
- 209 dietary habits, or a combination of both.
- 210 <Insert Figure 5>
- 211 *δ18O values*
- 212 The obtained  $\delta^{18}$ O values of the 112 dentine apatite samples range from 16.4 to 29.5‰
- 213 (average of all samples 23.1‰; Fig. 4D-F) and are in agreement with data obtained elsewhere
- 214 for MIS 3 cave bears (e.g. Krajkarz et al., 2014). All  $\delta^{18}O$  profiles show high variability
- 215 throughout the deposition of cave bear dentine until adulthood. Ur**Cilor** samples ( $N = 36$ )
- 216 layers) have extreme values ranging between 19.3 and 29.5‰ (average 23.4‰), Oase (*N* = 38
- 217 layers), range between 16.4 and 25.8‰ (average 23.1‰) and Cioclovina (*N* = 38 layers),
- 218 range between 17.9 and 26.6‰ (average 22.7‰).
- 219 The intra-tooth  $\delta^{18}$ O variability of the analyzed bears ranges from 1 to 8.2‰. Among the three
- 220 populations, Oase has the highest intra-tooth variability for all  $\delta^{18}$ O profiles, while Cioclovina
- 221 has the lowest. Nonetheless, all sites have bears with both types of  $\delta^{18}$ O profiles, suggesting
- 222 intra-populational variability.
- 223 <Insert Figure 6>

#### 224 **DISCUSSION**

- 225 *Physiology and behavior using δ13C data*
- 226 The large range of  $\delta^{13}$ C values measured for the MIS 3 Carpathian cave bears span all dietary
- 227 range, from herbivores to carnivores, taking into account the values for dentine apatite
- 228 published so far (e.g. Bocherens et al., 1994). While the few data obtained from dentine
- 229 apatite for Pleistocene bears (adults) were recorded around -16 to -14‰ (Bocherens et al.,
- 230 1991), our mean values of the  $\delta^{13}$ C show a large inter-population variation (-16 to -9,0‰, for
- 231 the last layer of the analyzed specimens), with values typical for both carnivores and
- 232 herbivores (Fig. 7). The large variation in  $\delta^{13}$ C values of the Carpathian cave bears may
- 233 indicate ecological differences, linked to different dietary and/or habitat preferences as
- 234 previously shown by Robu et al. (2013, 2018).
- 235 <Insert Figure 7>
- 236 The  $\delta^{13}$ C records of studied cave bears show a similar pattern among them, in which  $\delta^{13}$ C
- 237 values are increasing from juveniles to adults, in accordance to previous research (e.g. Nelson
- 238 et al., 1998; Robu et al., 2018), indicating a change of diet across time (Fig. 4A-C).

239 For all studied cave bears, a negative shift of  $\delta^{13}$ C profiles was recorded during the second 240 (mainly at  $Ur\Box$ ilor) or the third (mainly at Oase and Cioclovina caves) year of life. 241 Apparently, the "second year drop" profile is present as well at Oase and Cioclovina caves, 242 while the "third year drop" profile was recorded at  $Ur$  lilor as well, to a lesser extent. Although it seems that these records are site-characteristic, in fact, it is a matter of random 244 sampling and it is very likely for the  $\delta^{13}$ C negative shifts to be related to the weaning completion in cave bears, a critical period for the future adults, reflected in low lipids and 246 energy input in their diet. It was shown in some studies that the positive trend in  $\delta^{13}C$  associated with weaning could be the result of decreasing proportion of lipids in the diet brought about by the transition from high to low dependence on milk (e.g. Hobson and Sease, 249 1998). Thus, when fat represents an important component in the animal diet, carbonate  $\delta^{13}C$  should have more negative values than when fat is less important (Rountrey et al., 2007). In modern brown bear, *Ursus arctos*, it was documented that the weaning period takes place during the first or second year of life, while the separation of cubs from their mothers occurs at 1–3 (occasionally 4) years of age (McLellan, 1994; Dahle and Swenson, 2003), a fact 254 supported as well by our results. The same  $\delta^{13}$ C variation related to the weaning process have been reported for human canines retrieved from archaeological context, by Sandberg et al. (2014).

### *Paleoclimatic context using δ18O data*

259 The similarity of  $\delta^{18}$ O patterns between our fossil data and modern bears published by Koch et al. (1989), in both amplitude and relation to growth layers, indicate that the isotopic signal recorded in our fossil samples has no diagenetic influence. The lower values corroborated with the dark bands, could be, as in the case of modern black bears (Stoneberg and Jonkel, 1966) the isotopic expression of the late fall/winter denning (Fig. 4D-F).

264 The  $\delta^{18}$ O values of dentine retrieved from the structural carbonate for the MIS 3 Carpathian cave bears are in agreement with the results published elsewhere on dental enamel apatite 266 (e.g. Bocherens et al., 1994; Krajkarz et al., 2014). The amplitude of isotopic variation of  $\delta^{18}O$  values (max. 8.2‰) is the same as the one observed for the Younger Dryas North American 268 mammoths and bears (Koch et al., 1989). Large  $\delta^{18}$ O intra-population variations have been recorded as well for modern mammal population, ranging from 5.2‰ for wild horses, to 8.5‰ for yaks and 10.8‰ for horses from the Tibetan Plateau (Wang et al., 2008). On the 271 other hand, no sex-related differences in  $\delta^{18}$ O values among the analyzed cave bear

populations was recorded, with high and low variation patterns being obtained for both

females and males.

 With some exceptions it appears that bears, during the first two dentine growth layers (first 275 two years of life) have similar  $\delta^{18}$ O values suggesting a low mobility whereas starting with the third growing layers, the values may show significant shifts, sometimes up to 6‰. This 277 pattern may indicate a growing mobility and could be related with the  $\delta^{13}C$  data that show a significant change in bears physiology between the second and the third year of life. Modern brown bears, the closest relative to the MIS 3 cave bears, have a home-range size according to recent studies (Krofel et al., 2010, 2012) between 300 and 1600 km, with an average of 600- 700 km. It is therefore plausible that during the feeding and mating period, bears could drink 282 water with different  $\delta^{18}O$  signature than usual, a notion reflected in the  $\delta^{18}O$  isotopic record. Thus, the high variations could be a reflection of a broader territorial range. On the other 284 hand, the large intra-species  $\delta^{18}$ O variations seen in cave bears could be related as well to the fact that a large part of their body water was obtained from plant sources (*e.g*., fruits) in the context of the unstable climate of MIS 3. Both males and females show high variability of  $287 \delta^{18}$ O intra-tooth values. However, it must be kept in mind as well that during the period 288 between 30 and 50 kyr BP many short-term climatic changes have occurred and the  $\delta^{18}O$ signal could be biased (Shackleton and Hall, 2001; Van Meerbeeck et al., 2009; Bocherens et

al., 2011; Bocherens, 2015).

291 Bocherens et al. (2011) suggested that warmer time periods led to more positive  $\delta^{18}$ O values

292 retrieved from cave bear tissues, due to temperature effects on  $\delta^{18}$ O values of meteoric water.

 Therefore, values >22‰ would correspond to warmer periods, values <21.3‰ correspond to a colder period, and those between 21.3 - 22‰ indicate intermediate climatic periods. Figure

295 8 shows how the average  $\delta^{18}$ O values of Romanian bears cluster, indicating two main sets of

values: first one, with highest values, situated between 25.4 and 25.5‰ (two individuals from

297 Urilor and Cioclovina, a male and a female, respectively;  $N = 2$ ), and the second one,

situated between 21.3 and 24‰ (*N* = 16). The second set of data has two subsets, one that

299 ranges between 22.6 and 24‰  $(N = 12)$ , while the second one is situated between 21.3 and

22.1‰ (*N* = 4). According to Bocherens et al. (2011) criteria, the vast majority of the studied

bears lived during a warmer climate (*N* = 14 bears; 22.6 - 25.5‰), while only four of them

during a colder stage (21.3 - 22.1‰; *N* = 4).

<Insert Figure 8>

304 Corroborating the  $\delta^{13}$ C and  $\delta^{18}$ O values of the studied bears (Fig. 9A), it appears that there is a

high variability and affinity among the three sites. Nonetheless, the core of the data shows a

- 306 slightly stronger similarity between Ur $\Box$ ilor and Oase. When adding an additional parameter
- 307 to this analysis, *i.e.*, the absolute age of bears (Fig. 9B), two "different" cave bear populations
- emerge: (i) one dated at around 45-39 kyr cal BP, with numerous individuals from all
- 309 analyzed sites, and with similar high  $\delta^{13}$ C and  $\delta^{18}$ O values among the bears and (ii) a second
- one, derived only from Cioclovina (two individuals), dated at 34.5-31.5 kyr cal BP.
- Interestingly, within the first population, two individuals, a male (PO/42) and a female
- 312 (PU/072), from Oase and Urilor, respectively, show similar data in terms of age and  $\delta^{13}C$
- 313 and  $\delta^{18}$ O variation, suggesting a similar environment and physiology.
- 314 <Insert Figure 9>

#### **CONCLUSIONS**

- The CO<sub>3</sub> component of dentine apatite retrieved from cave bear canines proves to be a reliable
- proxy for reconstructing paleoclimatic conditions, physiology and behavior.
- 318 Based on  $\delta^{13}$ C data profiles, our study shows that the weaning process in cave bears can be
- identified and attributed to the first and second year of life, while the mother-offspring
- separation takes place during the second and the third year. Serial dentine samples display
- 321 large intra-tooth  $\delta^{18}$ O variation of up to 8,2 ‰, reflecting seasonal variations in the  $\delta^{18}$ O of
- meteoric water and a variable mobility of bears, regardless of sex. This study shows that bear
- ontogeny when analyzing the dentine layers encompasses the neonate period and
- 324 adulthood, thus providing more data of  $\delta^{13}$ C and  $\delta^{18}$ O for a single individual and
- 325 palaeoecological insights than before. Corroborating the  $\delta^{13}$ C and  $\delta^{18}$ O values with
- radiocarbon data may lead to correlations with past climatic and habitat changes during the
- lifetime of the cave bears.

#### **ACKNOWLEDGEMENTS**

- Thirty years ago, Paul L. Koch, Daniel C. Fisher and David Dettman published in *Geology*
- (1989) a study where they highlighted, for the first time, how dentine serial sampling of bear's
- canine can reveal paleoecological and paleoenvironmental insights from the past. Their study
- was an inspiration to us and therefore, this paper is dedicated to them. The authors are
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## 543 **Figure captions**

544 Figure 1. The studied Romanian MIS 3 cave bears sites:  $A - Ur$  lior Cave,  $B - Cioclovina$ 545 Cave, C – Oase Cave (cave maps redrawn after Vălena $\Box$ , 1979; Onac et al., 2011; Constantin 546 et al., 2013).

547 Figure 2. Sampling methodology within a cross-section of a cave bear canine. Dentine layers 548 1 $\rightarrow$  1  $\rightarrow$  6 layers were sampled for stable isotope analysis; the 1<sup>st</sup> layer is the earliest to form while

549 the  $6^{st}$  is the last to form.

550 Figure 3. Calibrated <sup>14</sup>C AMS ages of the cave bear samples  $(N = 10)$ .

551 Figure 4.  $\delta^{13}$ C and  $\delta^{18}$ O data obtained from the dentine layers of MIS 3 cave bears from

552 analyzed sites (six canines analyzed/each site).  $A - C$ :  $\delta^{13}C$  results from Ur $\Box$ ilor ( $N = 36$ )

553 layers), Oase ( $N = 38$ ) and Cioclovina caves ( $N = 38$ ), respectively; D – E:  $\delta^{18}$ O results from

- 554 Urilor ( $N = 36$ ), Oase ( $N = 38$ ) and Cioclovina caves ( $N = 38$ ), respectively. X axis  $1 \rightarrow 8$ :
- 555 dentine layers.
- 556 Figure 5: Boxplots of  $\delta^{13}C$  data from Urilor, Oase and Cioclovina caves. A: females ( $N = 9$ ) 557 canines, 52 layers); B: males (*N* = 9 canines, 60 layers).
- 558 Figure 6: Boxplots of  $\delta^{18}$ O data from Ur**uilor**, Oase and Cioclovina caves. A: females ( $N = 9$ ) 559 canines, 52 layers); B: males (*N* = 9 canines, 60 layers).

560 Figure 7: Jitter plot of  $\delta^{13}$ C values from dentine apatite from various MIS 3 mammals across

561 Europe. Cave bears: CBUR – Ur $\Box$ ilor ( $N = 6$ ); CBOA – Oase ( $N = 6$ ); CBCI – Cioclovina ( $N$ 562 = 6); CBAL – Aldéne (*N* = 5); CBMI – Mialet (*N* = 4). Herbivores: CAMI – Mialet (*N* = 3);

563 CEES – Escale (*N* = 2) (data for Aldéne. Mialet and Escale from Bocherens et al., 1994).

564 Figure 8: Hierarchical clustering of average values of  $\delta^{18}$ O of all dentine layers of each 565 analyzed canine  $(N = 18)$  data from Ur**lilor**, Oase and Cioclovina caves. Red lines indicate

- 566 warm periods, blue lines indicate colder periods.
- 567 Figure 9: Multidimensional scaling and minimum spanning tree in a two-dimensional visual
- 568 representation constructed for the MIS cave bears from the analyzed sites based on:  $A -$
- 569 average values of δ<sup>13</sup>C, δ<sup>18</sup>O data for each canine ( $N = 18$ ) (first two components explain 95%
- 570 of the variance; stress = 0,002);  $B \delta^{13}C$ ,  $\delta^{18}O$  (average value for each canine) and <sup>14</sup>C AMS
- 571 data ( $N = 10$ ) (first two components 88% of the variance; stress = 0,05). Red dots Oase,
- 572 yellow dots Cioclovina, blue dots Ur $\Box$ ilor.





















Fig. 4



PC/057







Distance













**Table S1**. Romanian/Carpathian MIS 3 cave bear carbon (δ13C) and oxygen (δ18O) stable isotope samples and results.