

# Palaeoecology and palaeoclimatic context of Romanian Carpathian MIS 3 cave bears using stable isotopes ( $\delta$ 13C and $\delta$ 18O)

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1	Palaeoclimatic context and palaeoecology of Romanian Carpathian MIS 3 cave bears
2	using stable isotopes ( $\delta^{13}$ C and $\delta^{18}$ O)
3	
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17	ABSTRACT
18	Millimeter-scale growth rings in canine dentine of MIS 3 cave bears have been interpreted as
19	annual growth hands produced in part by seasonal variation in growth rate. We present new
20	intra-tooth stable carbon ( $\delta^{13}$ C) and oxygen ( $\delta^{18}$ C) isotone profiles in dentine hydroxygenatite
20	of early forming permanent teeth from three famous Late Plaistocene cave bear sites from
21	Demonitor Competitions We accounted \$130 and \$180 afthe CO. for them of denting
22	Romanian Carpathians. We measured $\delta^{13}$ C and $\delta^{13}$ O of the CO <sub>3</sub> fraction of dentine
23	hydroxyapatite from samples covering consecutive inferred years of growth in canines. The
24	$\delta^{13}$ C values retrieved from the dentine layers – from the oldest to the newest – of the same
25	individual, record an increase throughout immature life of bears as has been shown
26	previously, with lower precision, using age categories. For the first time, based on $\delta^{13}C$ data
27	analysis, the weaning process in cave bears was identified. The $\delta^{18}O$ values show substantial
28	variations related, most probably, to seasonal growth of the dentine. Finally, the CO <sub>3</sub> of
29	dentine apatite extracted from cave bear canines proves to be reliable for geochemical
30	analyses, reflecting palaeoclimate conditions, physiology and behavior.
31	Keywords: Geochemistry, dentine, Mammalia, Ursidae, Late Pleistocene
32	INTRODUCTION
33	The discovery of measurable natural variation in the stable isotopic composition of vertebrate
34	fossil remains brought palaeontologists a valuable tool for studying fossil mammals from
35	ancient marine and terrestrial communities (e.g. Longinelli, 1984; Clementz, 2012; Matthews

et al., 2016; Martin et al., 2017). Carbon and oxygen isotopic abundances (reflected in  $\delta^{13}$ C 36 and  $\delta^{18}$ O values) measured from fossil vertebrate teeth have been used for a long time to 37 decipher the dietary habits and ecology of extinct species (Sullivan and Krueger, 1981; 38 Krueger and Sullivan, 1984; Ambrose and Norr, 1993; Longinelli, 1984; Koch et al., 1989; 39 Reinhard et al., 1996). Carbon isotopic abundances in carbonate hydroxyapatite  $[Ca_{10}(PO_4,$ 40  $CO_{3}_{6}$  (OH<sub>2</sub>)] reflect the average of the whole diet components (amino acids, carbohydrates 41 and lipids). The oxygen isotopic abundances recorded from the same tissue reflect the 42 variation in the  $\delta^{18}$ O value of body fluids controlled mainly by the isotopic composition of the 43 ingested water. As the ingested water tracks the local precipitation,  $\delta^{18}$ O values obtained from 44 dentine are used to reconstruct palaeoclimatic conditions and sources of water used by 45 animals (Longinelli, 1984; Koch et al., 1989; Reinhard et al., 1996; Dotsika et al., 2011; 46 Bocherens et al., 2011; Clementz, 2012; Krajcarz and Krajcarz, 2014; Krajcarz et al., 2014). 47 48 Late Pleistocene cave bear fossils are the most abundant remains to be found in the caves of Europe, this extinct species being one of the best studied in Quaternary palaeontology. 49 50 Nonetheless, there are still many debates within the scientific community regarding the diet, physiology and behavior of these ursids. The permanent canine teeth of adult bears are useful 51 52 for tracing age-related changes, mostly due to their resilience to weathering. The chronological age of individuals can be accurately determined (Debeljak, 1996) from the 53 annual layers of dentine or cementum, emphasized when odontogenesis is slowed during the 54 annual period of denning (nutritional stress). Although the cementum layers record a longer 55 history of the bear's life, dentine layers are better suited for geochemical analyses as they are 56 57 thicker and can provide sufficient sampling material. It has been shown that the annual canine dentine layers in adult individuals may vary between five and eight, with the first five layers 58 59 belonging to the primary dentine, and the others to the secondary dentine. Each dentine layer consists of a thick, light-stained inner zone, deposited during the annual period of activity, and 60 a very thin, dark-stained outer zone related to the period of denning (Rausch, 1961, 1967, 61 1969; Mundy and Fuller, 1964; Stoneberg and Jonkel, 1966; Sauer, 1966; Sauer et al., 1966; 62 63 Koch et al., 1989). Knowing that the cyclicity of dentine growth layers in bear canines ends when adulthood is reached (Rausch, 1969), one can trace back both the ecological conditions 64 and the climatic evolution by using a geochemical approach (e.g., using  $\delta^{13}$ C and  $\delta^{18}$ O values) 65 for every identified dentine layer. 66 67 The carbon isotope analysis of organic tissues of such tooth dentine provides information

reflective of the protein portion (collagen) of an individual's diet (Sullivan and Krueger, 1981;

69 Krueger and Sullivan, 1984; Ambrose and Norr, 1993; Harrison and Katzenberg, 2003),

whereas stable isotope values from bone, dentine, and enamel apatite carbonate largely reflect 70 71 the whole diet (*i.e.*, proteins, carbohydrates, and lipids; Ambrose and Norr, 1993; Harrison and Katzenberg, 2003). In general, apatite is more <sup>13</sup>C-enriched compared to collagen (by 72 73 +9,6-13‰; DeNiro and Epstein, 1978a, b; Sullivan and Krueger, 1981; Krueger and Sullivan, 1984; Lee-Thorp and Van der Merwe, 1987; Fahy et al., 2015), their relationship being 74 affected by the trophic level. The utility of bone/tooth apatite for stable carbon isotope 75 analysis in palaeodietary studies stems from the fact that previous results indicate that 76 diagenesis of biogenic carbonates in the mineral over time is unexpectedly limited, and that 77 78 chemical pretreatment further removes the effects of diagenetic alteration of the biogenic 79 signal (Lee-Thorp et al., 1989). Therefore, palaeodietary inferences could be carried out using this method for dentine apatite for fossil remains up to  $\sim 10^5$  years old (Clementz, 2012). Thus, 80 apatite  $\delta^{13}$ C measurements should not be regarded as an alternative to  $\delta^{13}$ C collagen 81 82 measurements, but may provide additional useful information about dietary trends. 83 Previous studies have focused on cave bear physiology (e.g. Nelson et al., 1998; Bocherens, 84 2015; Robu et al., 2018) using  $\delta^{13}$ C values from bone collagen and have shown that the isotopic profiles of the four age groups (neonates, juveniles, sub-adults and adults) provide 85 patterns of transfer of stable isotope signatures throughout the immature life of bears. 86 Studying the  $\delta^{13}$ C transfer signature throughout the first five to eight years of life, using  $\delta^{13}$ C 87 data from apatite of each individual, could bring a better understanding of MIS 3 cave bear 88 physiology and a better resolution of this process. 89 Sequential sampling of distinct dentine layers within a single tooth can provide information 90 on dietary and habitat changes over the course of an individual development from juvenile to 91 92 adult as well as seasonal variations in these ecological parameters later in the animal's life. Previous studies of fossil proboscideans – elephants and their extinct relatives (Koch et al., 93 1989; Fox and Fisher, 2004; Hope et al., 1999; Rountrey et al., 2007; Clementz, 2012) 94 95 highlighted how serial sampling of both tooth enamel and dentine can be used to answer different questions about the ecology and life history of these animals. Koch et al. (1989) 96 analyzing the  $\delta^{18}$ O values from dentine layers, have shown that CO<sub>3</sub> fraction of apatite retains 97 a signal in molars (proboscideans and bears), reflecting seasonal changes in composition of 98 drinking water. 99 The Romanian Carpathians host numerous, complex, and rich Late Pleistocene cave bear 100

sites, suitable for palaeoecological and palaeoclimatological investigations. Three already

102 famous Late Pleistocene cave bear sites from Romanian Carpathians were chosen for this

103 study – the Ur lilor, Oase and Cioclovina caves – known to the international scientific

- 104 community for their impressive fossil bonebeds, human and faunal remains, and for their
- taphonomic and palaeoichnologic features (e.g. Richards et al., 2008; Robu et al., 2018).
- 106 Analyzing both  $\delta^{13}$ C and  $\delta^{18}$ O, retrieved from dentine apatite serial samples, has not been
- applied systematically to any MIS 3 species as yet. Thanks to their large dimensions,
- analyzing a large set of MIS 3 cave bear canines provides an excellent record of seasonal
- variations in diet and climate during tooth growth and strengthen our understanding of the
- palaeoclimatic and palaeoecological significance of  $\delta^{13}$ C and  $\delta^{18}$ O values retrieved from
- 111 dentine apatite.

#### 112 MATERIALS AND METHODS

- 113 *Carpathian site-specific samples*
- 114 The Carpathian cave bear specimens derive from excavations in three extensive karst systems:
- 115 Peștera Urșilor de la Chișcău, Peștera Cioclovina Uscată, and Peștera cu Oase (Fig. 1).
- 116 The Peştera Urşilor de la Chişcău (further on: Urşilor;  $46^{\circ} 55 \square 37 \square N$ ,  $22^{\circ} 56 \square 92 \square E$ ) is
- situated in the northwestern part of the Romanian Carpathians (the Apuseni Mountains). The
- analyzed cave bear material was sampled from the palaeontological excavation, situated
- 119 within the Scientific Reserve (lower level of the cave). The analyzed cave bear population
- was AMS  $^{14}$ C dated between *c*.47 and 39 kyr cal BP (Constantin et al., 2014; Fig. 1A).
- 121 The Peştera Cioclovina Uscată (*Dry Cioclovina Cave;* further on: Cioclovina; 45° 34 🗆 28 🗆
- 122 N, 23° 08 $\square$  04 $\square$  E) (Fig. 1B) is located in the northwestern part of the Southern Carpathians
- 123 (the Şureanu Mountains) (Soficaru et al., 2007; Onac et al., 2011). Results of AMS <sup>14</sup>C dating
- indicate that the cave bear population inhabited the cave between c.44.2 and 31.4 kyr cal BP
- 125 (Soficaru et al., 2007; Meleg et al., 2018; this study).
- 126 The Peştera cu Oase (*Ponor Plopa Cave*; further on: Oase;  $45^{\circ} 05 \square 21 \square N$ ,  $21^{\circ} 51 \square 21 \square E$ )
- 127 is situated in the southwestern part of the Southern Carpathians (the Anina Mountains) (Fig.
- 128 1C). The analyzed cave bear samples were excavated from the upper level of the cave (Sala
- 129 Mandibulei and Panta Strămoșilor). The cave bear population of the bone assemblage was
- dated between *c*.50 and 40.4 kyr cal BP (Higham and Wild, 2013; this study).
- 131 <Insert Figure 1>
- 132 *Cave bear samples*
- 133 Teeth of the analyzed cave bears were identified by their morphology (according to Torres et
- al., 1978; Tsoukala and Grandal-d'Anglande, 2002). DNA analyses have not been completed
- for all studied sites, although some populations were attributed to *U. ingressus* either on
- 136 mtDNA or nuclear DNA basis or dental and metacarpal morphology (Richards et al., 2008;
- 137 Robu, 2016; Meleg et al., 2018).

- 138 A number of 18 fused cave bear lower canines (N = 18) were analyzed to obtain carbon and
- 139 oxygen isotopic composition from the structural carbonate of their dentine. The canines were
- selected from the analyzed sites (N = 9 females and N = 9 males; Table S1), with six canines
- being sampled from each collection. Hydroxyapatite samples (N = 112) were drilled from
- series of holes aligned parallel to incremental dentine layers (Fig. 2); the powder from each
- 143 row of holes formed a single sample (c. 2 mg). We obtained five to nine non-overlapping
- samples within each inferred year. The samples were analyzed following the method
- 145 described by Koch et al. (1997).
- 146 <Insert Figure 2>
- 147 The isotopic ratios are expressed as  $\delta^{13}$ C and  $\delta^{18}$ O values according to calculation:  $\delta^{A}X =$
- 148  $(R_{sample}/R_{standard} 1) \times 1000$ , where X is an element (C or O), A is the atomic mass number and
- 149 R is the isotopic ratio  ${}^{13}C/{}^{12}C$  or  ${}^{18}O/{}^{16}O$ . Values are reported in relation to Vienna-PDB
- 150 (VPDB) standard, with  $\delta^{18}$ O values being converted to VSMOW using the formula
- 151  $\delta^{18}O_{VSMOW} = 1.03091 \times \delta^{18}O_{VPDB} + 30.91$  (Skrzypek et al., 2011). Laboratory standards used
- 152 were CO<sub>2</sub>:He for  $\delta^{13}$ C, and NBS18 and CARRARA 11 for  $\delta^{18}$ O. Both  $\delta^{13}$ C and  $\delta^{18}$ O values
- 153 were measured using a Delta V+ (Thermo-Finnigan, Bremen, Germany) isotope ratio mass
- spectrometer (IRMS), in continuous flow mode (Fry et al., 1992) coupled to an ECS410
- elemental analyzer (Costech Analytical, Valencia, CA, USA) in the Department of Geology,
- 156 University of South Florida.
- 157 *Dating method*
- 158 Ten cave bear samples (N = 10) derived from the analyzed canines were AMS <sup>14</sup>C dated at
- 159 Poznan Radiocarbon Laboratory (Poland; Tab. S1). At the Poznan Laboratory, collagen
- 160 extraction was performed using the method described by Pietrowska and Goslar (2002) and
- the extracted collagen was then purified on Elkay 8 mm filters, and ultrafiltered on Vivaspin
- 162 15 MWCO 30 kD filters (Bronk Ramsey et al., 2004). The content of <sup>14</sup>C in carbon samples
- 163 was measured following the protocol described by Goslar et al. (2004).
- 164 *Statistical analyses*
- 165 Statistical analyses were performed using the software PAST (Paleontological Statistics),
- version 3.20 (Hammer *et al.* 2001). Hierarchical clustering was used to explore relationship
- 167 patterns of  $\delta^{18}$ O variability among the studied samples. Euclidean distances were used to
- 168 calculate a similarity matrix based on the average values of  $\delta^{18}$ O for 18 canines from Ur $\Box$ ilor,
- 169 Oase and Cioclovina caves, and Ward's method was selected to define the distance between
- identified classes. To visualize the similarity among studied samples, multidimensional

- scaling and minimum-spanning tree linking samples were computed based on average values
- 172 of  $\delta^{18}$ O,  $\delta^{13}$ C and median <sup>14</sup>C date for 18 canines from Ur $\Box$ ilor, Oase and Cioclovina caves.

#### 173 **RESULTS**

- 174 The results of radiocarbon dating are given in Table S1 (N = 10 samples). All ages were
- 175 calibrated using the OxCal 4.3.2 program (Bronk Ramsey, 2017) and the IntCal13 calibration
- 176 curve (Reimer et al., 2013). Radiocarbon ages (Fig. 3) of the 10 cave bear bone samples
- 177 collected from the analyzed sites range between c. 45 and 31 kyr cal BP (median values) for
- both surface and excavated finds. The majority of the dated samples of bears belong to 39-45
- 179 kyr cal BP interval known for its faunal richness across caves from both the Carpathians and
- the rest of Eurasia and to a last pulsation of this species, around 31 kyr cal BP (N = 2; from
- 181 Cioclovina Cave) at a time believed to be close to its extinction (Pacher and Stuart, 2009).
- 182 <Insert Figure 3>
- 183  $\delta^{13}C$  values
- 184 The obtained  $\delta^{13}$ C values of all 112 dentine apatite samples range from -17.5 to -9‰ (average
- of all samples -12.8%; Fig. 4A-C). All  $\delta^{13}$ C profiles are recording an increment of carbon
- values throughout the deposition of cave bear dentine until adulthood. Ur  $\Box$  ilor samples (N =
- 187 36 dentine layers) have extreme values between -16.7 and -9.2‰, Oase samples (N = 38
- dentine layers) spread between -14.9 and 9,0% and, Cioclovina samples (N = 38 dentine
- layers) spread between -17.5 and -9.9‰.
- 190 <Insert Figure 4>
- 191 Comparing the  $\delta^{13}$ C values of the oldest dentine layers of the analyzed canines (precluding
- adulthood in bears), with the average of the  $\delta^{13}C$  collagen values obtained for the same
- 193 populations (e.g. Robu et al., 2013, 2018), we found that the lowest  $\Delta \delta^{13}C_{collagen-apatite}$
- 194 difference ( $\delta^{13}C_{collagen} \delta^{13}C_{apatite}$ ) was recorded for Cioclovina, +7,9‰ ( $\delta^{13}C_{collagen} = -21.1$
- and  $\delta^{13}C_{apatite} = -13.2\%$ ), while the Ur  $\Box$  ilor population had a  $\Delta\delta^{13}C_{collagen-apatite}$  value of +9.6‰
- 196  $(\delta^{13}C_{\text{collagen}} = -21.6 \text{ and } \delta^{13}C_{\text{apatite}} = -12.0\%)$ . The highest average difference recorded in the
- 197 Oase population between the two parameters  $\Delta \delta^{13}C_{collagen-apatite}$  was +11,3‰ ( $\delta^{13}C_{collagen} = -$
- 198 21,4 and  $\delta^{13}C_{apatite} = -10.1\%$ , respectively).
- 199 Figure 5 shows the variation of  $\delta^{13}$ C values between sexes among the studied Carpathian
- 200 populations: for females, it appears that the highest average is recorded for the Oase and
- 201 Urşilor females (-11.9‰, and -12‰, respectively), and the lowest for Cioclovina, -14.6‰
- 202 (Fig. 5A). For males, the highest average value was measured in the Oase population: -
- 11.9%, and lowest values in Ur $\Box$ ilor (-13.9‰) and Cioclovina (-13.7‰) (Fig. 5B). Both
- graphs suggest differences among sexes within all three cave bear populations: for Ur□ilor,

- females show  $\delta^{13}$ C values enriched by +1.9‰ when compared with males, while for Oase and
- 206 Cioclovina, the situation is opposite, males have enriched  $\delta^{13}$ C values by +0.5 and +0.9‰,
- 207 respectively. The differences in  $\delta^{13}$ C values among sexes could be due to either random and
- size sampling (only three males and females analyzed/site), to different intra-population
- 209 dietary habits, or a combination of both.
- 210 <Insert Figure 5>
- 211  $\delta^{18}O$  values
- The obtained  $\delta^{18}$ O values of the 112 dentine apatite samples range from 16.4 to 29.5‰
- 213 (average of all samples 23.1%; Fig. 4D-F) and are in agreement with data obtained elsewhere
- for MIS 3 cave bears (e.g. Krajkarz et al., 2014). All  $\delta^{18}$ O profiles show high variability
- throughout the deposition of cave bear dentine until adulthood. Ur  $\Box$  ilor samples (N = 36
- layers) have extreme values ranging between 19.3 and 29.5‰ (average 23.4‰), Oase (N = 38
- layers), range between 16.4 and 25.8‰ (average 23.1‰) and Cioclovina (N = 38 layers),
- 218 range between 17.9 and 26.6‰ (average 22.7‰).
- 219 The intra-tooth  $\delta^{18}$ O variability of the analyzed bears ranges from 1 to 8.2‰. Among the three
- populations, Oase has the highest intra-tooth variability for all  $\delta^{18}$ O profiles, while Cioclovina
- has the lowest. Nonetheless, all sites have bears with both types of  $\delta^{18}$ O profiles, suggesting
- 222 intra-populational variability.
- 223 <Insert Figure 6>

#### 224 DISCUSSION

- 225 Physiology and behavior using  $\delta^{13}C$  data
- 226 The large range of  $\delta^{13}$ C values measured for the MIS 3 Carpathian cave bears span all dietary
- range, from herbivores to carnivores, taking into account the values for dentine apatite
- published so far (e.g. Bocherens et al., 1994). While the few data obtained from dentine
- apatite for Pleistocene bears (adults) were recorded around -16 to -14‰ (Bocherens et al.,
- 1991), our mean values of the  $\delta^{13}$ C show a large inter-population variation (-16 to -9,0‰, for
- the last layer of the analyzed specimens), with values typical for both carnivores and
- herbivores (Fig. 7). The large variation in  $\delta^{13}$ C values of the Carpathian cave bears may
- 233 indicate ecological differences, linked to different dietary and/or habitat preferences as
- previously shown by Robu et al. (2013, 2018).
- 235 <Insert Figure 7>
- 236 The  $\delta^{13}$ C records of studied cave bears show a similar pattern among them, in which  $\delta^{13}$ C
- values are increasing from juveniles to adults, in accordance to previous research (e.g. Nelson
- et al., 1998; Robu et al., 2018), indicating a change of diet across time (Fig. 4A-C).

For all studied cave bears, a negative shift of  $\delta^{13}$ C profiles was recorded during the second 239 (mainly at Ur lior) or the third (mainly at Oase and Cioclovina caves) year of life. 240 Apparently, the "second year drop" profile is present as well at Oase and Cioclovina caves, 241 while the , third year drop" profile was recorded at Ur lilor as well, to a lesser extent. 242 Although it seems that these records are site-characteristic, in fact, it is a matter of random 243 sampling and it is very likely for the  $\delta^{13}$ C negative shifts to be related to the weaning 244 completion in cave bears, a critical period for the future adults, reflected in low lipids and 245 energy input in their diet. It was shown in some studies that the positive trend in  $\delta^{13}C$ 246 associated with weaning could be the result of decreasing proportion of lipids in the diet 247 brought about by the transition from high to low dependence on milk (e.g. Hobson and Sease, 248 249 1998). Thus, when fat represents an important component in the animal diet, carbonate  $\delta^{13}C$ 250 should have more negative values than when fat is less important (Rountrey et al., 2007). In 251 modern brown bear, Ursus arctos, it was documented that the weaning period takes place during the first or second year of life, while the separation of cubs from their mothers occurs 252 253 at 1-3 (occasionally 4) years of age (McLellan, 1994; Dahle and Swenson, 2003), a fact supported as well by our results. The same  $\delta^{13}$ C variation related to the weaning process have 254 255 been reported for human canines retrieved from archaeological context, by Sandberg et al. (2014). 256

257

#### 258 Paleoclimatic context using $\delta^{18}O$ data

The similarity of  $\delta^{18}$ O patterns between our fossil data and modern bears published by Koch et al. (1989), in both amplitude and relation to growth layers, indicate that the isotopic signal recorded in our fossil samples has no diagenetic influence. The lower values corroborated with the dark bands, could be, as in the case of modern black bears (Stoneberg and Jonkel, 1966) the isotopic expression of the late fall/winter denning (Fig. 4D-F).

The  $\delta^{18}$ O values of dentine retrieved from the structural carbonate for the MIS 3 Carpathian 264 cave bears are in agreement with the results published elsewhere on dental enamel apatite 265 (e.g. Bocherens et al., 1994; Krajkarz et al., 2014). The amplitude of isotopic variation of  $\delta^{18}$ O 266 values (max. 8.2‰) is the same as the one observed for the Younger Dryas North American 267 mammoths and bears (Koch et al., 1989). Large  $\delta^{18}$ O intra-population variations have been 268 recorded as well for modern mammal population, ranging from 5.2% for wild horses, to 269 8.5% for yaks and 10.8% for horses from the Tibetan Plateau (Wang et al., 2008). On the 270 other hand, no sex-related differences in  $\delta^{18}$ O values among the analyzed cave bear 271

272 populations was recorded, with high and low variation patterns being obtained for both

females and males.

With some exceptions it appears that bears, during the first two dentine growth layers (first 274 two years of life) have similar  $\delta^{18}$ O values suggesting a low mobility whereas starting with 275 the third growing layers, the values may show significant shifts, sometimes up to 6%. This 276 pattern may indicate a growing mobility and could be related with the  $\delta^{13}C$  data that show a 277 significant change in bears physiology between the second and the third year of life. Modern 278 brown bears, the closest relative to the MIS 3 cave bears, have a home-range size according to 279 recent studies (Krofel et al., 2010, 2012) between 300 and 1600 km, with an average of 600-280 281 700 km. It is therefore plausible that during the feeding and mating period, bears could drink 282 water with different  $\delta^{18}$ O signature than usual, a notion reflected in the  $\delta^{18}$ O isotopic record. 283 Thus, the high variations could be a reflection of a broader territorial range. On the other hand, the large intra-species  $\delta^{18}$ O variations seen in cave bears could be related as well to the 284 fact that a large part of their body water was obtained from plant sources (e.g., fruits) in the 285 286 context of the unstable climate of MIS 3. Both males and females show high variability of  $\delta^{18}$ O intra-tooth values. However, it must be kept in mind as well that during the period 287 288 between 30 and 50 kyr BP many short-term climatic changes have occurred and the  $\delta^{18}$ O signal could be biased (Shackleton and Hall, 2001; Van Meerbeeck et al., 2009; Bocherens et 289

al., 2011; Bocherens, 2015).

Bocherens et al. (2011) suggested that warmer time periods led to more positive  $\delta^{18}$ O values

retrieved from cave bear tissues, due to temperature effects on  $\delta^{18}$ O values of meteoric water.

Therefore, values >22% would correspond to warmer periods, values <21.3% correspond to a colder period, and those between 21.3 - 22‰ indicate intermediate climatic periods. Figure

295 8 shows how the average  $\delta^{18}$ O values of Romanian bears cluster, indicating two main sets of

values: first one, with highest values, situated between 25.4 and 25.5% (two individuals from

297 Ur $\square$ ilor and Cioclovina, a male and a female, respectively; N = 2), and the second one,

situated between 21.3 and 24‰ (N = 16). The second set of data has two subsets, one that

ranges between 22.6 and 24‰ (N = 12), while the second one is situated between 21.3 and

22.1% (N = 4). According to Bocherens et al. (2011) criteria, the vast majority of the studied

bears lived during a warmer climate (N = 14 bears; 22.6 - 25.5‰), while only four of them

302 during a colder stage (21.3 - 22.1%; N = 4).

303 <Insert Figure 8>

304 Corroborating the  $\delta^{13}$ C and  $\delta^{18}$ O values of the studied bears (Fig. 9A), it appears that there is a

high variability and affinity among the three sites. Nonetheless, the core of the data shows a

- 306 slightly stronger similarity between Ur□ilor and Oase. When adding an additional parameter
- 307 to this analysis, *i.e.*, the absolute age of bears (Fig. 9B), two "different" cave bear populations
- 308 emerge: (i) one dated at around 45-39 kyr cal BP, with numerous individuals from all
- analyzed sites, and with similar high  $\delta^{13}$ C and  $\delta^{18}$ O values among the bears and (ii) a second
- one, derived only from Cioclovina (two individuals), dated at 34.5-31.5 kyr cal BP.
- Interestingly, within the first population, two individuals, a male (PO/42) and a female
- 312 (PU/072), from Oase and Ur $\Box$ ilor, respectively, show similar data in terms of age and  $\delta^{13}C$
- and  $\delta^{18}$ O variation, suggesting a similar environment and physiology.
- 314 <Insert Figure 9>

## 315 CONCLUSIONS

- 316 The CO<sub>3</sub> component of dentine apatite retrieved from cave bear canines proves to be a reliable
- 317 proxy for reconstructing paleoclimatic conditions, physiology and behavior.
- Based on  $\delta^{13}$ C data profiles, our study shows that the weaning process in cave bears can be
- 319 identified and attributed to the first and second year of life, while the mother-offspring
- 320 separation takes place during the second and the third year. Serial dentine samples display
- large intra-tooth  $\delta^{18}$ O variation of up to 8,2 ‰, reflecting seasonal variations in the  $\delta^{18}$ O of
- 322 meteoric water and a variable mobility of bears, regardless of sex. This study shows that bear
- 323 ontogeny when analyzing the dentine layers encompasses the neonate period and
- adulthood, thus providing more data of  $\delta^{13}C$  and  $\delta^{18}O$  for a single individual and
- palaeoecological insights than before. Corroborating the  $\delta^{13}$ C and  $\delta^{18}$ O values with
- radiocarbon data may lead to correlations with past climatic and habitat changes during the
- 327 lifetime of the cave bears.

## 328 ACKNOWLEDGEMENTS

- 329 Thirty years ago, Paul L. Koch, Daniel C. Fisher and David Dettman published in *Geology*
- 330 (1989) a study where they highlighted, for the first time, how dentine serial sampling of bear's
- canine can reveal paleoecological and paleoenvironmental insights from the past. Their study
- 332 was an inspiration to us and therefore, this paper is dedicated to them. The authors are
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### 543 Figure captions

Figure 1. The studied Romanian MIS 3 cave bears sites: A – Ur□ilor Cave, B – Cioclovina
Cave, C – Oase Cave (cave maps redrawn after Vălena□, 1979; Onac et al., 2011; Constantin
et al., 2013).

Figure 2. Sampling methodology within a cross-section of a cave bear canine. Dentine layers 1 $\rightarrow$ 6 layers were sampled for stable isotope analysis; the 1<sup>st</sup> layer is the earliest to form while the 6<sup>st</sup> is the last to form.

550 Figure 3. Calibrated <sup>14</sup>C AMS ages of the cave bear samples (N = 10).

551 Figure 4.  $\delta^{13}C$  and  $\delta^{18}O$  data obtained from the dentine layers of MIS 3 cave bears from

analyzed sites (six canines analyzed/each site). A – C:  $\delta^{13}$ C results from Ur $\Box$ ilor (N = 36

layers), Oase (N = 38) and Cioclovina caves (N = 38), respectively; D – E:  $\delta^{18}$ O results from

- 554 Ur  $\square$  ilor (N = 36), Oase (N = 38) and Cioclovina caves (N = 38), respectively. X axis 1  $\rightarrow$  8:
- 555 dentine layers.
- 556 Figure 5: Boxplots of  $\delta^{13}$ C data from Ur $\Box$ ilor, Oase and Cioclovina caves. A: females (N = 9557 canines, 52 layers); B: males (N = 9 canines, 60 layers).
- 558 Figure 6: Boxplots of  $\delta^{18}$ O data from Ur $\Box$ ilor, Oase and Cioclovina caves. A: females (N = 9559 canines, 52 layers); B: males (N = 9 canines, 60 layers).

560 Figure 7: Jitter plot of  $\delta^{13}$ C values from dentine apatite from various MIS 3 mammals across

561 Europe. Cave bears: CBUR – Ur $\Box$ ilor (N = 6); CBOA – Oase (N = 6); CBCI – Cioclovina (N = 6); CBAL – Aldéne (N = 5); CBMI – Mialet (N = 4). Herbivores: CAMI – Mialet (N = 3);

563 CEES – Escale (N = 2) (data for Aldéne. Mialet and Escale from Bocherens et al., 1994).

Figure 8: Hierarchical clustering of average values of  $\delta^{18}$ O of all dentine layers of each analyzed canine (N = 18) data from Ur $\Box$ ilor, Oase and Cioclovina caves. Red lines indicate warm periods, blue lines indicate colder periods.

- 567 Figure 9: Multidimensional scaling and minimum spanning tree in a two-dimensional visual
- representation constructed for the MIS cave bears from the analyzed sites based on:  $A CS^{12}C$  and  $CS^{12}C$  and  $CS^{12$
- average values of  $\delta^{13}$ C,  $\delta^{18}$ O data for each canine (N = 18) (first two components explain 95% of the variance; stress = 0,002); B –  $\delta^{13}$ C,  $\delta^{18}$ O (average value for each canine) and <sup>14</sup>C AMS
- 571 data (N = 10) (first two components 88% of the variance; stress = 0,05). Red dots Oase,
- 572 yellow dots Cioclovina, blue dots Ur $\Box$ ilor.





















Fig. 4



PC/057







Distance





					δ <sup>13</sup> C	δ <sup>18</sup> Ο		δ <sup>18</sup> O	Radiocarbon age
Sample	Site	Species	Sex	Age	(‰V-PDB)	(‰V-PDB)	%CO3	(‰SMOW)	(calibrated; 95.4% probability)
PU/068 1	Ur□ilor	U. spelaeus	Male	Adult	-12,4	-7,1	6,0	23,5	N/A
PU/068 2	Ur□ilor	U. spelaeus	_	_	-13,8	-6,4	3,5	24,3	N/A
PU/068 3	Ur□ilor	U. spelaeus	_	_	-13,2	-7,5	4,5	23,1	N/A
PU/068 4	Ur□ilor	U. spelaeus	_	_	-13,3	-6,8	5,9	23,8	N/A
PU/068 5	Ur□ilor	U. spelaeus	_	_	-11,7	-9,7	5,3	20,8	N/A
PU/068 6	Ur□ilor	U. spelaeus	_	_	-12,5	-7,6	8,2	23,0	N/A
PU/068 7	Ur□ilor	U. spelaeus	_	_	-12,3	-8,1	3,2	22,5	N/A
PU/069 1	Ur□ilor	U. spelaeus	Male	Adult	-13,7	-6,4	5,0	24,3	N/A
PU/069 2	Ur□ilor	U. spelaeus	_	_	-14,2	-6,5	5,2	24,2	N/A
PU/069 3	Ur□ilor	U. spelaeus	_	_	-13,9	-2,1	3,1	28,7	N/A
PU/069 4	Ur□ilor	U. spelaeus	_	_	-14,4	-7,3	5,0	23,4	N/A
PU/069 5	Ur□ilor	U. spelaeus	_	_	-12,5	-6,8	5,9	23,8	N/A
PU/069 6	Ur□ilor	U. spelaeus	_	_	-13,6	-1,4	4,4	29,5	N/A
PU/069 7	Ur□ilor	U. spelaeus	_	_	-12,7	-6,5	4,2	24,1	N/A
PU/070 1	Ur□ilor	U. spelaeus	Male	Adult	-15,9	-5,4	N/A	25,3	N/A
PU/070 2	Ur□ilor	U. spelaeus	_	_	-15,6	-5,6	N/A	25,1	N/A
PU/070 3	Ur□ilor	U. spelaeus	_	_	-16,1	-6,8	N/A	23,8	N/A
PU/070 4	Ur□ilor	U. spelaeus	_	_	-16,7	-8,2	N/A	22,4	N/A
PU/070 5	Ur□ilor	U. spelaeus	_	_	-16,3	-11,2	N/A	19,3	N/A
PU/070 6	Ur□ilor	U. spelaeus	_	_	-14,1	-5,9	N/A	24,7	N/A
PU/070 7	Ur□ilor	U. spelaeus	_	_	-13,8	-6,4	N/A	24,3	N/A
PU/071 1	Ur□ilor	U. spelaeus	Female	Adult	-14,7	-7,6	5,1	23,0	42,707 - 37,913 cal yr BP
PU/071 2	Ur□ilor	U. spelaeus	_	_	-15,8	-7,6	2,9	23,0	N/A
PU/071 3	Ur□ilor	U. spelaeus	_	_	-14,5	-8,8	4,4	21,8	N/A
PU/071 4	Ur□ilor	U. spelaeus	_	_	-13,3	-6,9	4,1	23,7	N/A
PU/071 5	Ur□ilor	U. spelaeus	_	_	-12,6	-9,1	3,6	21,4	N/A
PU/072 1	Ur□ilor	U. spelaeus	Female	Adult	-12,2	-7,3	6,6	23,4	42,144 - 38,286 cal yr BP
PU/072 2	Ur□ilor	U. spelaeus	_	_	-13,6	-7,6	7,6	23,1	_
PU/072 3	Ur□ilor	U. spelaeus	_	_	-12,5	-8,0	6,9	22,6	_
PU/072 4	Ur□ilor	U. spelaeus	_	_	-12,2	-7,3	5,6	23,3	_
PU/072 5	Ur□ilor	U. spelaeus	_	_	-11,6	-6,0	5,2	24,7	_
PU/073 1	Ur□ilor	U. spelaeus	Female	Adult	-9,5	-8,5	7,1	22,1	41,490 - 37,276 cal yr BP
PU/073 2	Ur□ilor	U. spelaeus	_	_	-9,8	-10,9	6,0	19,6	_

PU/073 3	Ur□ilor	U. spelaeus	_	_	-9,9	-11,0	5,7	19,5	_
PU/073 4	Ur□ilor	U. spelaeus	_	_	-9,2	-7,4	6,2	23,3	_
PU/073 5	Ur□ilor	U. spelaeus	_	_	-9,2	-7,8	5,0	22,8	_
PO/038 1	Oase	U. spelaeus	Female	Adult	-11,0	-8,0	4,8	22,6	N/A
PO/038 2	Oase	U. spelaeus	_	_	-10,6	-7,1	3,9	23,6	N/A
PO/038 3	Oase	U. spelaeus	_	_	-11,4	-7,9	4,9	22,7	N/A
PO/038 4	Oase	U. spelaeus	_	_	-11,0	-7,3	4,0	23,3	N/A
PO/038 5	Oase	U. spelaeus	_	_	-9,6	-7,2	3,9	23,4	N/A
PO/039 1	Oase	U. spelaeus	Female	Adult	-14,9	-6,4	2,2	24,3	N/A
PO/039 2	Oase	U. spelaeus	_	_	-13,9	-6,6	3,9	24,0	N/A
PO/039 3	Oase	U. spelaeus	_	_	-12,5	-6,6	2,5	24,1	N/A
PO/039 4	Oase	U. spelaeus	_	_	-11,8	-5,9	2,5	24,8	N/A
PO/039 5	Oase	U. spelaeus	_	_	-12,1	-7,3	2,6	23,4	N/A
PO/039 6	Oase	U. spelaeus	_	_	-11,6	-11,7	3,4	18,8	N/A
PO/039 7	Oase	U. spelaeus	_	_	-9,3	-7,3	1,9	23,3	N/A
PO/040 3	Oase	U. spelaeus	Female	Adult	-13,4	-5,14	N/A	25,6	45,911 - 40,842 cal yr BP
PO/040 4	Oase	U. spelaeus	_	_	-13,7	-11,6	N/A	18,9	_
PO/040 5	Oase	U. spelaeus	_	_	-13,7	-5,5	N/A	25,2	_
PO/040 6	Oase	U. spelaeus	_	_	-11,8	-4,9	N/A	25,8	_
PO/040 7	Oase	U. spelaeus	_	_	-11,2	-6,3	N/A	24,4	_
PO/041 1	Oase	U. spelaeus	Male	Adult	-13,1	-7,8	6,3	22,8	N/A
PO/041 2	Oase	U. spelaeus	_	_	-14,3	-8,9	6,2	21,7	N/A
PO/041 3	Oase	U. spelaeus	_	_	-12,7	-6,6	6,1	24,1	N/A
PO/041 4	Oase	U. spelaeus	_	_	-13,1	-7,0	6,3	23,7	N/A
PO/041 5	Oase	U. spelaeus	_	_	-13,3	-9,0	4,7	21,6	N/A
PO/041 6	Oase	U. spelaeus	_	_	-11,2	-7,9	5,5	22,7	N/A
PO/041 7	Oase	U. spelaeus	_	_	-10,1	-6,4	5,5	24,3	N/A
PO/042 1	Oase	U. spelaeus	Male	Adult	-12,9	-8,0	6,4	22,7	42,230 - 38,560 cal yr BP
PO/042 2	Oase	U. spelaeus	_	_	-11,4	-6,7	5,4	24,0	_
PO/042 3	Oase	U. spelaeus	_	_	-12,0	-7,5	5,1	23,2	_
PO/042 4	Oase	U. spelaeus	_	_	-12,0	-9,3	5,7	21,3	_
PO/042 5	Oase	U. spelaeus	_	_	-11,4	-8,9	4,5	21,7	_
PO/042 6	Oase	U. spelaeus	_	_	-10,6	-8,7	4,9	21,9	_
PO/042 7	Oase	U. spelaeus		_	-11,2	-7,3	5,2	23,3	_
PO/043 1	Oase	U. spelaeus	Male	Adult	-10,8	-6,2	N/A	24,4	49,347 - 42403 cal yr BP

PO/043 2	Oase	U. spelaeus	_	_	-10,2	-6	N/A	24,6	_
PO/043 3	Oase	U. spelaeus	_	_	-10,9	-6,3	N/A	24,4	_
PO/043 4	Oase	U. spelaeus	_	_	-10,2	-7,8	N/A	22,8	_
PO/043 5	Oase	U. spelaeus	_	_	-9,8	-8	N/A	22,6	_
PO/043 6	Oase	U. spelaeus	_	_	-9,2	-14,1	5,2	16,4	_
PO/043 7	Oase	U. spelaeus	_	_	-9	-6,3	N/A	24,4	_
PC/056 1	Cioclovina	U. spelaeus	Male	Adult	-11	-6,8	N/A	23,8	34,166 - 32,924 cal yr BP
PC/056 6	Cioclovina	U. spelaeus	_	_	-9,9	-6,7	N/A	23,9	_
PC/057 1	Cioclovina	U. spelaeus	Male	Adult	-17,5	-8,7	3,0	21,9	40,888 - 38,215 cal yr BP
PC/057 3	Cioclovina	U. spelaeus	_	_	-16,7	-7,5	2,4	23,1	_
PC/057 4	Cioclovina	U. spelaeus	_	_	-15,8	-8,0	2,5	22,6	_
PC/057 5	Cioclovina	U. spelaeus	_	_	-16,1	-8,0	2,7	22,7	_
PC/057 6	Cioclovina	U. spelaeus	_	_	-16,1	-9,2	3,4	21,4	_
PC/057-7	Cioclovina	U. spelaeus	_	_	-15,8	-8,5	3,9	22,1	_
PC/057 8	Cioclovina	U. spelaeus	_	_	-15,1	-7,8	4,0	22,8	_
PC/057 9	Cioclovina	U. spelaeus	_	_	-14,6	-10,2	1,6	20,3	_
PC/058 1	Cioclovina	U. spelaeus	Male	Adult	-15,1	-7,4	2,5	23,2	N/A
PC/058 2	Cioclovina	U. spelaeus	_	_	-13,4	-8,1	3,7	22,5	N/A
PC/058 3	Cioclovina	U. spelaeus	_	_	-16,0	-8,6	3,3	22,0	N/A
PC/058 4	Cioclovina	U. spelaeus	_	_	-15,1	-8,5	3,7	22,1	N/A
PC/058 5	Cioclovina	U. spelaeus	_	_	-14,2	-7,5	2,4	23,1	N/A
PC/058 6	Cioclovina	U. spelaeus	_	_	-14,4	-7,2	2,5	23,4	N/A
PC/058 7	Cioclovina	U. spelaeus	_	_	-13,7	-8,7	3,1	21,8	N/A
PC/058 8	Cioclovina	U. spelaeus	-	-	-14	-6,6	N/A	24,0	N/A
PC/059 1	Cioclovina	U. spelaeus	Female	Adult	-16,2	-7,4	4,6	23,3	43,046 - 40,446 yr cal BP
PC/059 2	Cioclovina	U. spelaeus	_	_	-14,5	-7,5	N/A	23,2	-
PC/059 3	Cioclovina	U. spelaeus	_	_	-15,5	-11,2	4,4	19,3	_
PC/059 4	Cioclovina	U. spelaeus	_	_	-15,6	-8,2	4,1	22,4	-
PC/059 5	Cioclovina	U. spelaeus	_	_	-13,9	-8,5	4,6	22,1	-
PC/059 6	Cioclovina	U. spelaeus	_	_	-13,9	-10,5	1,9	20,0	-
PC/059 7	Cioclovina	U. spelaeus	_	_	-13,6	-11,5	2,3	19,0	-
PC/060 2	Cioclovina	U. spelaeus	Female	Adult	-15,9	-7,9	3,2	22,7	31,860 - 30,974 cal yr BP
PC/060 3	Cioclovina	U. spelaeus	_	_	-15,5	-8,1	2,0	22,5	-
PC/060 4	Cioclovina	U. spelaeus	_	_	-15,3	-7,8	3,6	22,8	-
PC/060 5	Cioclovina	U. spelaeus	_	_	-15,1	-8,2	3,2	22,4	_

PC/060 6	Cioclovina	U. spelaeus	_	_	-14,2	-12,6	2,2	17,9	_
PC/060 7	Cioclovina	U. spelaeus	_	_	-14,6	-9,7	1,9	20,8	_
PC/061 1	Cioclovina	U. spelaeus	Female	Adult	-14,9	-4,4	_	26,3	N/A
PC/061 2	Cioclovina	U. spelaeus	_	_	-15,1	-5,4	_	25,3	N/A
PC/061 3	Cioclovina	U. spelaeus	_	_	-14,5	-4,1	_	26,6	N/A
PC/061 4	Cioclovina	U. spelaeus	_	_	-13,9	-5,7	_	25,0	N/A
PC/061 5	Cioclovina	U. spelaeus	_	_	-13,2	-5	_	25,7	N/A
PC/061 6	Cioclovina	U. spelaeus	_	_	-13,2	-5,7	_	25,0	N/A
PC/061 7	Cioclovina	U. spelaeus	_	-	-12,2	-5,8	_	24,9	N/A

**Table S1**. Romanian/Carpathian MIS 3 cave bear carbon ( $\delta^{13}$ C) and oxygen ( $\delta^{18}$ O) stable isotope samples and results.