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► **To cite this version:**

R. Pellen, D. Aslanian, Marina Rabineau, J.P. Suc, C. Gorini, et al.. The Messinian Ebro River incision. *Global and Planetary Change*, 2019, 181, pp.102988. 10.1016/j.gloplacha.2019.102988 . hal-02185361v1

HAL Id: hal-02185361

<https://hal.science/hal-02185361v1>

Submitted on 5 Jan 2021 (v1), last revised 11 Sep 2020 (v2)

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The Messinian Ebro River incision

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Keywords: NW Mediterranean Sea; Segmentation; Knickpoint; Ebro fluvial system; Messinian Salinity Crisis; Incised-valley system.

Abstract

Morphological sills condition sedimentary, water and fauna exchanges between different domains. In particular, sills are crucial factors to consider during the Messinian Salinity Crisis (MSC) palaeogeographic evolution (5.97-5.33 Ma) of the NW Mediterranean area. Here we focus on the Ebro River and its up to now unexplained short Messinian onshore length (~ 100 km) compared to that of the Messinian Rhone River (~ 480 km) despite similar present-day drainage basins. Thanks to an extensive seismic and borehole dataset, we present a new interpretation of a complete 270 km long Messinian Ebro incised-valley system course underneath the present-day continental margin and bathyal basin and its related distal detrital deposits. These results favour a syn-MSC or pre-MSC opening of the (endorheic) Ebro Basin to the Mediterranean. We propose a mechanism of retrogressive erosional process, localized at structural knickpoints that shift seaward through time. This mechanism resulted in the development of the complete incised-valley system and the falling stage system tract (FSST) during the MSC sea level fall with negligible or even null retrogressive inland erosion beyond the Catalan Coastal Range. The shifting of erosion-deposition is controlled by the pre-Messinian stepwise morphology and segmentation in the Valencia, Menorca and Liguro-Provence Basins. By comparison, the Rhone system is simpler, characterized by the key role of a single knickpoint (at the shelf-break) and a steeper continental slope. Both cases highlight the relationship between kinematic, segmentation and their relative morphologies, base-level fall and erosional/depositional response particularly well expressed during the outstanding MSC associated with a huge relative sea-level drop that we measured down to -1100m below present day sea-level.

37 **Abbreviations****General**

ESP Expended Spread Profile
Km Messinian Knickpoint
Ka Present-day knickpoint
MSC Messinian Salinity Crisis
Twt Two-Way Travel time

Geography

CCR Catalan Coastal Range
CFZ Central Fracture Zone

EB Ebro Basin
GOL Golf of Lion
LPB Liguro-Provence Basin
MB Menorca Basin
NBFZ North Balearic Fracture Zone
VB Valencia Basin

FAD First Appearance Datum
FSST Falling Stage System Tract
HST Highstand System Tract
LAD Last Appearance Datum
LST Lowstand System Tract
LU Lower Unit
MES Messinian Erosional Surface
MRS Maximum Regression Surface
MLM Messinian Lower Megasequence
MUM Messinian Upper Megasequence
MU Mobile Unit
RSME Regressive Surface of Marine Erosion
TST Transgressive System Tract
UU Upper Unit

Stratigraphy

1

2 **1. Introduction**

3 Morphological sills, whether related to sedimentary, inherited structural reliefs or originating from the
4 kinematic history, play a crucial role in the sedimentary, water and fauna exchanges between basins through
5 time. These sills often fit areas of different crustal nature and different subsidence histories, as observed in the
6 Mediterranean-Paratethys area (e.g. [Leever et al., 2010](#); [Leroux et al., 2015a](#); [Palcu et al., 2017](#); [Pellen et al.,](#)
7 [2017](#)). In the case of huge relative sea-level variations, as during the MSC (5.97-5.33 Ma) ([Manzi et al., 2013](#)),
8 consideration of these barriers is of primary importance to understand the morphological and sedimentary
9 evolution of the different basins (e.g. [Leever et al., 2010](#); [Flecker et al., 2015](#); [Palcu et al., 2017](#); [Suc et al., 2015](#);
10 [Balázs et al., 2017](#)).

11 **1.1. The Ebro Paradox**

12 The Rhone and Ebro fluvial systems are excellent examples of major rivers in the Mediterranean Sea (the 2nd
13 and 3rd respectively after the Nile) that experienced the same crisis in response to the isolation of the
14 Mediterranean Sea. Present-day drainage areas are of similar size (nearly 99,000 km² and 86,000 km²,
15 respectively) ([Babault et al., 2006](#)) ([Fig. 1](#)). However direct comparisons between the Messinian palaeo-Ebro
16 and Messinian palaeo-Rhone systems point out the so-called “Ebro paradox” ([Babault et al., 2006](#)): despite
17 similar present-day drainage areas, the length of the Messinian subaerial incision of the Ebro river seems much

18 shorter than that of the Rhone river. In the case of the Messinian palaeo-Rhone, a total 480 km long fluvial
19 incision network was described along the Alps (Clauzon, 1982, 1999) to the pre-Messinian Platform edge
20 (Bache, 2008) (Fig. 1). In the case of the Messinian palaeo-Ebro erosion system, the incision is limited to about
21 100 km from the Catalan Coastal Range (CCR) edges to the former pre-Messinian platform edge (Fig. 1)
22 (Urgeles *et al.*, 2011) with no major incision described in the endorheic Ebro Basin. Why such a difference?
23 Babault *et al.* (2006) proposed that the Ebro River did not exist before the Messinian Salinity Crisis and that
24 the connection between the endorheic Ebro Basin to the Mediterranean Sea occurred after the MSC (or
25 potentially during this phase).

26 Conversely, several authors consider that seismic data provide evidence of well-developed sedimentation and
27 deltaic progradation during the Tortonian and the development of a main incised-valley system during the
28 MSC, implying a Tortonian connection between the Ebro Basin (EB) and the Valencia Basin (VB) (e.g.
29 Comeselle *et al.*, 2014). Recent modeling and field data on the foreland EB further argued that the transition
30 from endorheic to exorheic of Ebro is pre-Messinian, although the precise timing remains to be established (12
31 to 7.5 Ma) (Fillon *et al.*, 2013; Garcia-Castellanos and Larrasoana, 2015). However, the absence of major fluvial
32 incision toward the onshore EB, the short incision of ~ 100 km and the absence of well-described detrital syn-
33 MSC deposits in the VB remain unanswered. Indeed, Messinian clastic deposits were only partially described
34 and mainly on the (present-day) outer shelf. These deposits were described either on top (Escutia and
35 Maldonado, 1992; Maillard *et al.*, 2006a,b, or below the so-called Upper Unit (UU, Garcia *et al.*, 2011;
36 Comeselle *et al.*, 2014, 2015) of the well-known Messinian trilogy (LU-MU-UU) synthesised by Lofi *et al.* (2011)
37 (Fig. 1). This peculiar morphology strongly contrasts with that of the Western, Gulf of Lion, and Eastern, Nile
38 region, Mediterranean that both show long fluvial system incisions and large detrital deposits during the
39 Messinian sea-level fall (Gorini *et al.*, 2015) and remain to be explained.

41 1.2. Geodynamic and stratigraphic framework

42
43 The Neogene palaeoenvironmental framework of the NW Mediterranean region is largely condition by
44 geodynamic history which controls the vertical motion and sedimentation in each sub-basin (Leroux *et al.*,
45 2015a, 2015b; Pellen *et al.*, 2016). The geometry and physical link between the evaporitic and detritic
46 successions, and major tectonic hinge lines are well known in the Liguro-Provence Basin (LPB, Bache *et al.*,
47 2015; Leroux *et al.*, 2015a) (Fig. 1). However, a complete regional stratigraphic and morphological seismic study
48 linking the CCR edges to the deep LPB is still lacking.

49 A NW-SE segmentation distinguishes the MB from the Mesozoic VB (Roca, 1992; Etheve *et al.*, 2016, 2018)
50 and the deep LPB (Fig 1): the Central and North Balearic Fracture Zones (CFZ and NBFZ) represent two major
51 sills that created a stepwise accommodation between the three domains (Pellen *et al.*, 2016). The geodynamic
52 process leading to the formation of these sills is rooted in the counter-clockwise rotation of Corsica-Sardinia
53 block during the early Miocene which drove the movement of Menorca Island, and non-kinematic motion of
54 Ibiza-Majorca islands (Pellen *et al.*, 2016). These motions distinguish early Miocene LPB and MB basins from
55 the Mesozoic VB basin (e.g. Olivet, 1996; Gueguen *et al.*, 1998; Bache *et al.*, 2010; Pellen *et al.*, 2016, Etheve
56 *et al.*, 2016). One of the main consequences is a different subsidence history for each basin, with purely vertical
57 Neogene “sag” subsidence for the Mesozoic VB (Pellen *et al.*, 2016).

58 A NW-SE segmentation shapes the GOL-LPB with five differentiated crustal domains on the French side
59 (Moulin *et al.*, 2015; Afilhado *et al.*, 2015; Jolivet *et al.*, 2015): unthinned continental crust, thinned
60 continental crust, highly thinned continental crust, exhumed lower continental crust and proto-oceanic crust.
61 Three domains of subsidence were found defined by three hinge lines (Fig.1, 2). On the platform and slope,
62 the subsidence takes the form of a seaward tilt with different amplitudes, whereas the deep basin (with exhumed
63 lower continental crust and proto-oceanic crust) subsides purely vertically (Leroux *et al.*, 2015).

64 The shelves of the MB and the GOL (domain 2a) (Fig. 1) first record a late Oligocene-Aquitania marine influx
65 (Roca and Guimera, 1992; Roca *et al.*, 2001; Bache *et al.*, 2010), followed by Burdigalian-Tortonian clay-marl
66 then sand deposits associated with a prograding trend. The outermost position of this Tortonian shelf is shown
67 on Figure 1 (from Bache, 2008). The deep part of the LPB and MB are usually associated with condensed
68 Burdigalian-Tortonian deposits (Roca *et al.*, 2001). The Mesozoic VB (Roca, 1992; Etheve *et al.*, 2016, 2018)
69 experienced the first marine ingression during the late Burdigalian-Langhian followed by an enhanced
70 Tortonian prograding trend in the Tortosa area (Clavell and Berastegui, 1992; Cameselle *et al.*, 2014; Pellen *et al.*,
71 *et al.*, 2016).

72

73 1.3. Ebro versus Rhone inheritances

74 The initiation of the endorheic sedimentation in the EB has been estimated at ~35 Ma (Costa *et al.*, 2010),
75 the final preserved lacustrine fill -stage is proposed to be at least 13.5 and 12 Ma (in Urgeles *et al.*, 2011) and
76 up to Tortonian (e.g. Vasquez-Urbez *et al.*, 2013). At the time of the basin’s opening towards the Mediterranean
77 Sea, an elevation of 535-750 m above sea level has been estimated (Garcia-Castellanos and Larrasoña, 2015)
78 which places the Catalan Coastal Range (CCR) as an important morphological barrier between the VB and
79 EB. At the dawn of the MSC, a stepped morphology is observed for the VB-MB where offshore knickpoints

are located at the edge of the Miocene platform or at major geodynamic boundaries (Fig. 2). On the contrary, the GOL is characterized by a single continental slope (Guennoc *et al.*, 2000). Furthermore, no major geodynamical sill is observed along the onshore Rhone drainage system (Séranne, 1999; Guennoc *et al.*, 2000; Sissingh, 2001).

In this study, based on a large set of seismic and borehole data, we made a detailed synthesis of the offshore physiography during Messinian, by correlating the VB, MB, and LPB with seismic stratigraphic and geomorphological methods tied to existing borehole data with revisited detailed biostratigraphic analysis to propose a new understanding of Ebro and Rhone systems evolution during the MSC.

2. Data and Method

2.1. Dataset

Our study is based on a large set of seismic and borehole data collected during the extensive French academia-industry programs (Actions-Marges, GRI Tethys Sud), in close collaboration with the Total and Schlumberger groups as well as data available from the Spanish SIGEOF database. This compilation provided for the correlation of seismic markers from the VB towards the LPB (as shown already in Pellen *et al.*, 2016). Twenty-four petroleum wells drilled between 1969 and 2004 along the NE Iberian margin have been used to calibrate seismic interpretations, although the stratigraphic information was often sparse or vague. We used the detailed stratigraphic, sedimentological and micropaleontological information available from three wells located on the present-day Ebro platform to constrain ages and palaeoenvironments of identified surfaces and units: Fornax 1 (Bailey *et al.*, 2008; Urgeles *et al.*, 2011; Cameselle *et al.*, 2014; Mauffrey *et al.*, 2017), Benicarlo C1 (this study) and Tarragona E2 (Evans *et al.*, 1978 ; Bessais & Cravatte, 1988).

In this paper the detailed lithological, environmental and biostratigraphic information from the Benicarlo C1 are provided for Miocene-Messinian-Pliocene series (based on Evans *et al.*, 1978) that we have revisited in the light of recent biostratigraphic charts, a modern geological time scale (Gradstein *et al.*, 2012) and correlation with more regional information (see chapters 3.2 to 4.2).

2.2. Sequence stratigraphy

Conventional seismic stratigraphic method was used (Mitchum & Vail, 1977; Vail *et al.*, 1977) to identify seismic discontinuities and stratigraphic surfaces and units based on reflection configurations and facies characteristics (See the Table in Supplementary Material 3 for details of surfaces and units). However, identified surfaces and units were interpreted using principles of sequence stratigraphy in their revised most recent

112 version i.e. including interpretation of Falling Stage System Tract (FSST) (e.g. [Hunt and Tucker, 1992](#);
113 [Catuneanu et al., 2011](#)), forced regression concept (e.g. [Plint et al., 1988](#)) with Maximum Regression surface
114 (MRS) and Regressive Surface of Marine erosion (RSME) (e.g. [Helland-Hansen and Hampton, 2009](#)). Ages of
115 these sequences were provided by well data. Our interpretations were then integrated into the 'Kingdom Suite'
116 software to insure coherency at all crossing lines. For the Messinian units, we used the now well established so-
117 called « Messinian-trilogy » Lower Unit (LU)/Mobile Unit (MU)/ Upper Unit (UU) as synthesized by [Lofi et al.](#)
118 ([2011](#)). Nevertheless, we provide special focus on units and sub-units related to clastic units (see results below).

121 2.3. Time to depth conversion and backstripping

122 Our seismic interpretations were converted from time to depth using average sediments velocities (layer cake
123 model), to obtain the geometries in depths (meters) of surfaces and sedimentary units ([Fig. 3, Transect01](#)).
124 Seismic velocities strongly vary in space within the same stratigraphic interval, with significant variations
125 between the shelf, the base of slope, and the basin. Compiled borehole velocities on the VB platform and on
126 the foot slope, as well as ESP velocities located in the deep LPB ([Leroux et al., 2016](#)) were used to estimate an
127 average velocity for each layer (see velocity model in [suppl. mat. 1](#)). The depth converted profile allowed us to
128 measure the difference in altitudes between key points along the profiles, and sedimentary unit thicknesses.
129 A simple 2-D backstripping was performed to reconstruct geometries through time (in particular at early
130 Zanclean, at Messinian low stand incised-system and related deposits (see details in [suppl. mat. 2](#)). Input
131 parameters for the VB domain include paleobathymetry, age, and porosity of sedimentary packages based on
132 [Urgeles et al. \(2011\)](#). An estimated total subsidence value up to 960 m/Ma for the deep LPB domain was also
133 applied to the 2-D backstripping for Pliocene-Quaternary times as quantified by [Rabineau et al., \(2014\)](#).

135 2.4. Knickpoint definition and chosen approach

136 Our study used the identification of knickpoints, highlighted either by morphological change or new
137 sedimentary sequences development along the seismic profile. Different processes are associated with onshore
138 and offshore development of such morphological features. On land, knickpoints are interpreted as resulting
139 from renewed erosion of a rejuvenated river that propagates upstream and influences the development of
140 landscapes and source-to-sink systems ([Grimaud et al., 2016](#)). They are usually triggered by a relative fall in river
141 base-level (either through eustatic sea-level drop, tectonic faulting or uplift) and can also highlight the onshore-
142 offshore position ([Goswani et al., 2016](#)). Submarine knickpoints are often located at the shelf-break creating

143 an offlap-break either (i) at the transition between the outer shelf and the continental slope (e.g. Rabineau *et al.*, 2014), or (ii) at the boundary of domains where tectonic motion displaces the seafloor (Amblas *et al.*, 2011)
144 or (iii) at the site where channel levees are breached. Knickpoints are key elements to take into account in
145 when considering segmented domains, we therefore paid a particular attention to their identification.
146

147 Below, we present our results from the onshore domain towards the deeper offshore domains.
148

149 3. Results

150 3.1. Morphology from seismic interpretation

151 **Figure 3** presents our stratigraphic interpretation across the VB-MB along six seismic profiles perpendicular to
152 the main erosional incision. Transect 01 has been converted to depth. Eight major seismic surfaces are
153 illustrated, S0 (Tertiary basement), MES, S11, S20, S21, S22, S24, and S30 which respectively delimit four
154 seismic units: SU12a-c, LU, MU and UU and three major knickpoints (Km1-3). Three secondary surfaces
155 (S26a, b, c) are also identified in the UU. Relationships, age and interpretations of surfaces and units are
156 presented in the table in [suppl. mat. 3](#).

157 The first knickpoint Km₁ is located at the transition between the former foreland EB (~ 535 m above sea level)
158 and the VB at the edges of the CCR (**Figs. 2 and 3**). Below this knickpoint, the Marginal Erosional Surface
159 (MES, Lofi *et al.*, 2011) shows a highly rugged and dissected surface with frequent incisions organized in a
160 dendritic pattern (also interpreted on 3D seismic data, e.g. Urgeles *et al.*, 2011, and Comeselle *et al.*, 2014)
161 (**Figs. 2 and 3, section a**). Four palaeo-drainage systems are observed (**Fig. 2**) that present numerous incised
162 fluvial tributaries with rough badland-like morphology of the MES (Urgeles *et al.*, 2011) above the isocontour
163 of 1.6 sec TWT (~ 2.5 km depth) below present-day sea level. It is worth noting that the MES truncates the
164 former Tortonian prograding clinoforms which form the Miocene platform in the Tortosa-Ebro domain (**Figs.**
165 **3 and 4**). This is also confirmed by the Tarragona E2 well which sampled Tortonian deposits just below the
166 MES (Bessais & Cravatte, 1988). A wide dendritic palaeo-drainage system is mapped for the Tortosa-Ebro
167 domain (**Fig. 2**) which is the main and deepest NW-SE river channel interpreted as the Messinian Ebro fluvial
168 system (Urgeles *et al.*, 2011) (**Fig. 2, 3 section a & b**). A ~ 61 km long main channel incision can be measured
169 from the pre-MSC coast to the Tortosa-Ebro lower slope.

170 The erosive surface S11 extends seaward of the MES at the position of the Km₂ (**Fig. 3, Transect01 - km 130;**
171 **Fig. 4**) at the Miocene platform edge, delimiting the base of SU12a showing downlapping prograding
172 clinoforms geometries. Only the bottomset part of the clinoforms has been preserved on both sides of the main
173 Messinian river channel. The basal erosive surface S11 also marks a change in the progradational trend along

174 the Miocene platform edge (Fig.4, section h) of the Tarragona area: A major change in direction of sediment
175 input from the morphological domain of Barcelona to the Tortosa-Ebro domain is highlighted by the basal
176 erosive surface S11.

177 S11 shows a generally smoother morphology towards the VB foot-slope domain. In this area the MES subtracts
178 into two erosive surface: the basal erosive surface S20 which truncate the seismic unit SU12a while the top
179 erosive surface S30 bound the UU seismic unit (Fig. 3, sections c, d). Along the former Miocene outer shelf,
180 the MES-S20 surface shows a smoother erosive surface below 2.5 km depth (Fig. 4, sections g, h) with the
181 disappearance of badland morphology except in the axis of the main channel incision (Fig. 3, section b; Fig.
182 4). The extension of the S20 surface across the VB foot-slope domain highlights the lateral continuation of an
183 incised channel along the VB domain. A slight NE-ward inclination is observed along the incised-valley profile
184 with a mean present-day depth position ranging from 2.8 km below sea level at the Tortosa domain foot-slope
185 to 3.2 km depth at its outlet (Fig. 3, sections c-f). The morphology shows an evolution from a ~6 km wide with
186 an interfluvial/thalweg height of 370-400 m at the toe of the proto-Ebro slope offshore the Tortosa-Tarragona
187 domain (Fig. 3, sections a-c) to a ~16 km wide, 100-130 m deep incision at the transition towards the MB (Fig.
188 3, sections d, e). Bedrock morphology partly deflects the path of the incised-valley and slightly influences its
189 sinuosity (Fig.2; Fig. 3, section c).

190 On both sides of this SW-NE valley, which totals a length of ~134 km in the deep VB domain, the surface is
191 relatively conformable compared to the underlying deposits. The main incised-valley is joined by two smaller
192 tributaries from the SW part of the VB and already mapped by Cameselle *et al.* (2015).

193 At the transition between the VB and MB (CFZ area in Figs. 2, 3 and 5, section i) (Maillard and Mauffret,
194 1999; Pellen *et al.*, 2016), S20 shows another knickpoint Km₃ (~3.2 km depth, at Transect01, km 270 in Fig.
195 3). It marks: 1) the deepening of the basal surface S11 and its overlying unit SU12c – as the lateral correlation
196 of SU12a; 2) the last S20 toplap features; 3) the formation of two new surfaces, S21 – as the lateral correlation
197 of S20 – and S22 at the top and 4) the thickening of the UU.

198 The S12 and S21 surfaces delimitates the development of SU12c with the NE-ward preserved prograding
199 clinofolds (Fig. 5, section j). S21 at the top of SU12c also shows some widespread erosion and becomes
200 smoother and conformable towards the MB-LPB transition (Fig. 3, section f). The conformable and non-incised
201 S21 development also highlights the formation of the overlying LU. Both observations attest the termination
202 of the trans-VB incised-system thus marking its outlet at the VB-MB transition.

203 The continuity of the erosional system from the VB platform to the Km₃ knickpoint therefore points to a 270
204 km long system that we identify as the trans-VB Messinian valley system (Figs. 2 and 3, Transect 01 and sections
205 b-d). The MB-LPB transition marked by the North Balearic Fracture Zone (NBFZ) corresponds to the full

206 development of the Messinian trilogy units (LU-MU-UU) as identified in the deep LPB domain (Figs. 3 and
207 9).

208 Using the backstripped section of Transect01 at 5.3 Ma (i.e. at the end of the MSC), we measure the general
209 slope of the MES-S20-S21 from 90 to 430 km as equal to 0.55 % (i.e 0.32 °). However, individual segments in
210 between knickpoints show a slope even smaller than 0.4 % (0.23°) and 0.45% (0.26°), between 90-280 km and
211 310-430 km respectively. Conversely, surface S11 at the base of SU12c shows a slope of 1.58% (0.9 °), between
212 310 and 430 km.

213 214 215 **3.2. Borehole data**

216 **3.2.1. Biostratigraphy**

217 The Benicarlo C1 well (Fig. 6) provided planktonic foraminifers recorded from cuttings (Evans *et al.*, 1978).
218 These data have been re-visited and evaluated in light of the most recent biostratigraphic zonation in the
219 Mediterranean (Iaccarino *et al.*, 2007) and the modern geological time scale (Gradstein *et al.*, 2012).

220 The top of Cretaceous limestones (Upper-Middle Cenomanian) is well identified at 4330 m depth. With
221 respect to several Mediterranean long records (Suc *et al.*, 1992; Lirer and Iaccarino, 2011), the base of
222 Pleistocene deposits (i.e. the base of the Gelasian, at 2.6 Ma) can be assessed between 2330 m depth (uppermost
223 occurrence of *Sphaeroidinellopsis subdehiscens*, usually estimated at 3.2 Ma) and 2060 m (lowermost occurrence
224 of *Globorotalia inflata*, dated at 2.09 Ma in the Mediterranean Sea; Iaccarino *et al.*, 2007). Consequently, our
225 biostratigraphic re-evaluation focuses on the depth interval 4300–2300 m, which encompasses the Middle-Late
226 Miocene and whole Pliocene (Fig. 6). Examining the vertical distribution of the species with a biostratigraphic
227 significance along the studied interval, we paid particular attention to discontinuous occurrences, some
228 punctual records may result from cavings or reworkings.

229 *Orbulina universa* is present from 4240 m depth that suggests an age younger than 14.10 Ma (First Appearance
230 Datum, FAD, of the species in the Mediterranean; Iaccarino *et al.*, 2007). *Neogloboquadrina acostaensis* has been
231 continuously recorded from 4240 to 4180 m, then intermittently at 3830 m, and continuously from 3710 m
232 to 2040 m. The FAD of *N. acostaensis* in the Mediterranean Sea is dated at 11.90 Ma. Two interpretations can
233 be discussed: (1) to place the FAD of *N. acostaensis* at 4240 m or (2) to place it at 3710 m. Hypothesis 1 implies
234 very thick deposits to be ascribed to Tortonian while hypothesis 2 suggests that deposits from the base of the
235 studied interval partly belong to the Serravallian and maybe to the Langhian. Caliper log (Fig. 6) suggests the
236 presence of numerous caving effect between 3600 m and 4240 m which would affect the record of *N. acostaensis*
237 or *Orbulina universa* below 3710 m. As well, the biostratigraphic interpretation of the nearby Tarragona E2 well

(Cravatte, 1980; Bessais, 1984; Bessais & Cravatte, 1988) would be more consistent with the second hypothesis. Accordingly, sediments between 4240 and 3710 m depth could belong to a time-interval running from biozone MMi5 to biozone MMi10 (Fig. 6).

Above, the continuous occurrence of *Neogloboquadrina humerosa* is recorded from 3230 m. The FAD of *N. humerosa* is placed within biozone MMi12 (Iaccarino *et al.*, 2007), in agreement with the identification of this biozone in the Benicarlo C1 well (Fig. 6). A similar comment can be expressed for *Globigerinoides extremus*, which has been continuously recorded in the Benicarlo C1 well between 3220 m and 2900 m, that is also conceivable with the FAD of the species at the base of biozone MMi12 (Iaccarino *et al.*, 2007).

The occurrence of *Globorotalia conomiozea* from 3110 to 2640 m characterizes the MMi13 biozone ending at 5.97 Ma (Iaccarino *et al.*, 2007). Then, *Globorotalia margaritae* and *Spheroidinellopsis paenedehiscens* are recorded together from 2535 m, indicating the biozones MPI1-MPI2 up to the first record of *Globorotalia puncticalata* at 2511 m indicating the base of biozone MPI3 (Iaccarino *et al.*, 2007). Finally, the uppermost record of *Globorotalia margaritae* at 2370 m is used to delimit the top of biozone MPI3, dated at 3.80 Ma. These data are transferred to Fig. 6 where the ages of the biozone boundaries are also indicated according to Iaccarino *et al.* (2007). It is thus obvious that an important portion of the deposits drilled at Benicarlo C1, more precisely those between 3110 m and 2511 m depth, belongs to the pre-evaporitic phase of the Messinian.

3.2.2. Lithologies and Environments

The very thick sequence of volcanic tuffs and lavas between 4315 m - 3730 m were deposited in a fairly shallow marine environment during late Serravallian-Early Tortonian time as shown by the thin interbeds of soft limestones containing rare *Amphistegina* sp. and planktonic foraminifera (Evans *et al.*, 1978).

The sequence between 3695-3016 m, of Tortonian age is of open marine, outer sublittoral to bathyal nature. The onset of colder water conditions over this interval is shown by the presence of sinistral coiled specimens of *Globorotalia acostaensis* (since around 3400m) before the warmer Messinian period (Suc *et al.*, 1995; Evans *et al.*, 1978).

The interval 2810-2525 m represents a shallowing upward phase, with the probable formation of underwater supra-littoral conditions at the top. The reduction in planktonic foraminifera and the upward coarsening of the sediments suggests a general shallowing of the basin at this time.

The interval between 2525 - 2500 m can be separated in two parts: the topmost part of this interval between approximately 2510 - 2500 m as well as interval 2525 - 2524 m contain thin beds of white algal limestone interbedded with white sandstone and yellow to ochre marl/claystone. The lower limestone contains anhydrite.

270 Additionally, the occurrence of a single cyprid ostracod, algae and bryozoan in the 2515 and 2525-2520 m
271 ditch cuttings sample is taken to indicate both a stratigraphic and environmental break at this level and possibly
272 reflects the establishment of brakisch - non-marine conditions (Evans *et al.*, 1978).

273 The interval 2500 - 2370 m is characterized by soft, light grey calcareous shales and marls/claystone with minor
274 sand stringers throughout. This interval is assigned to the Ebro Group, Early Pliocene, deposited in an open
275 marine environment.

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278 4. Discussion

279 4.1. Age assignment and sequence stratigraphy interpretation

280 Combining the results from seismic interpretation in correlation with biostratigraphy from Benicarlo C1 well,
281 we consider that the Messinian is represented mostly by marine sediments from 3016 to 2525 m depths, i.e.
282 491 m in thickness with a shallowing upward trend (Fig. 6). The proximal to distal correlation of seismic units
283 SU12a to SU12c (both topped by the MES-S20) also shows a downward-shift of uppermost surface (as seen on
284 depth-converted seismic transect shown in Fig. 3 and in backstripped section (Supp. Material 2).

285 The MES(-S20) is depicted as the maximum regressive surface (MRS in the sense of Catuneanu *et al.*, 2011)
286 and is located at 2525 m in the Benicarlo C1 well. This MRS (MES) would be dated at 5.6 Ma according to
287 Clauzon *et al.* (1996, 2015) and CIESM (2008). The underlying Messinian sediments (SU12a and SU12c)
288 therefore correspond to the remaining sediments of the Falling Stage System Tract (FSST). According to our
289 seismic interpretation, the erosive surface at the base (S11) may correspond to the regressive surface of marine
290 erosion during the falling stage of the MSC crisis, also described in Quaternary sequences related to 100 000
291 years cycles (Rabineau *et al.*, 2005). However, some early pre-MSC Messinian sediments seem to be preserved
292 as biostratigraphic data suggest with the presence of *Globorotalia Cibaoensis* around 3000 m but considering the
293 relative Messinian sea level fall velocity and amplitude, earlier pre-MSC deposits are very fortunate to be
294 reworked and preserved in MSC related deposits.

295 This interpretation of the FSST for SU12a is supported by the fact that Benicalo C1- shows sediments of a
296 younger age than those found in both Fornax-1 and Tarragona E2 wells (Evans *et al.*, 1978 ; Bessais & Cravatte,
297 1988; Urgeles *et al.*, 2011) as well as a progressive downward shift of MES-S20 surface. Our seismic
298 interpretation and correlation suggest that SU12c is also part of this FSST and in fact represent the end of it
299 with the most distal regressive deposits. Development of SU12c unit is followed by the initiation of the LU as

lowstand system tract (in the sense of [Boyd et al., 2006](#); [Gorini et al., 2015](#)) (Fig. 8, stage 3). This FSST and LST would correspond to the MLM (Messinian Lower Megasequence) of [Gorini et al. \(2015\)](#).

Unfortunately, no well exists to sample this unit SU12c, so we have no direct related biostratigraphic information, which could confirm or invalidate this hypothesis.

Above the MES at Benicarlo C1, the occurrence of a single cyprid ostracod, algae and brozoan in the 2515 m and 2525-2520 m cuttings suggest that this thin unit going from probable emersion to shallow water environments records the end of the MSC. Following [CIESM \(2008\)](#) or [Bache et al., \(2012\)](#) scenarios, this unit could be interpreted as corresponding to those deposited during the early marine reflooding of the Mediterranean Basin (Transgressive System Tract) that preceded the base of the Zanclean, and estimated between 5.46 - 5.33 Ma. This TST is overlain by a clear open-marine high stand System Tract (HST) with warm conditions in early Zanclean Pliocene. (5.33-3.85 Ma) ([Suc et al., 1995](#)).

4.2. Subaerial versus submarine erosion and position of the Messinian shoreline

Erosional surfaces may develop in subaerial or submarine environments. However, process and formation of erosional surfaces differ in the air and under water, due to the varying physics of the different fluids. In subaerial environments topography is highly variable. Landscapes are carved by rivers and strongly dependent on precipitation that will erode and transport sediment as a function of gravity and slopes as the primary controls. In submarine environments, widespread erosion also occurs: both in proximal environments (at the shoreline and on the shelf) with a strong impact of waves, tides and currents but also in deep environments such as submarine canyons, slopes and deep-sea areas with processes such as turbidity currents, landslides, debris flows (e.g. [Amblas et al., 2006](#)) and contour currents (e.g. [Rebesco et al., 2014](#)). In shallow environments (between 0 to 100 - 150m) a very efficient erosion process is related to wave action. The main result of this wave action is the creation of large smooth morphologies and wave-cut terraces that are recognized worldwide throughout the geological time scale (e.g. [Cattaneo and Steel, 2003](#); [Catuneanu 2006](#); [Bache et al., 2014](#)).

In our case study, S11 marks the first MSC regressive surface of marine erosion (RSME) prior to the MES development (see also above) and highlights Km₂ at an onshore-offshore position ([Goswami et al., 2017](#)). The overlying U12a clinofolds (VB area) correspond to the Messinian forced-regressive prisms (FSST), with erosion and redeposition of sediments in the forced-regressive prisms (Fig. 4, section g; Fig. 6) and a MRS at the top. The subaerial development of the Trans-VB incised-valley system is attested across the Tortosa-Ebro domain with rugged badland-type morphology, terraces, fluvial meanders, oxidation surfaces (Fornax-1) as fully explained in ([Urgeles et al., 2011](#); [Cameselle et al., 2014](#)).

332 Moreover, as shown is the results section: the general slope of backstripped MES-S20-S21 erosional
333 surface is only 0.55% (0.32°) or even less in segments between knickpoints (0.23° to 0.26 °). These slope
334 gradients are much smaller than any present-day well-defined submarine canyon. In the GOL, the shelf-break
335 occurs when slope gradients reach 1%. Gradients < 1% delimit the continental shelf, whereas gradients > 1%
336 correspond to the upper slope. The values obtained for the Messinian MES-S20-S21 are typical of low shelf
337 gradients and not slope gradients where submarine canyons occur, they are generally between 2 and 5 ° in
338 canyon head areas in the GOL (Baztan *et al.*, 2005). Another example is provided by present-day submarine
339 canyons in the VB that all show slopes around 10° in upper slope submarine canyons to 2-4° at the base of
340 slope canyons (Amblas *et al.*, 2006). Those values are more than one order of magnitude higher than what we
341 measured here. MES-S20 general shape cannot be related to any submarine canyon mechanism.
342 If MES-S20 might have been also initiated as submarine one, this composite erosive surface was then re-worked
343 as a subaerial surface of erosion related to the Messinian lowermost sea level (MRS) with a coastline that
344 reached at least km₃ (270 km; Fig. 3 Transect 01). Its outlet is associated with the development of SU12c after
345 the Km₃ knickpoint and the start of S11c erosional surface at its base. The limit of prograding SU12c unit
346 would thus correspond to the most distal detritic products related to the relative MSC sea level fall stage.

347
348 The Messinian shoreline (Fig. 3 Transect01; Fig. 5, section I and Supp. Mat 2) existed at least up to km 270.
349 In any event, the minimum lowest MSC sea level (at km₃) was around -1100 m below present day sea-level
350 according to the backstripped section, SU12c corresponds to the final development of the detritic product
351 related to the MSC sea-level fall. The values we give here for sea-level reconstruction are probably maximum
352 values as backstripping was conducted up to 5.3 Ma after the UU deposition (and not 5.6 Ma at the maximum
353 sea level drawdown).

354 355 **4.3. A model for the Messinian Ebro incised-valley formation: Segmentation, knickpoint role, and** 356 **palaeogeographic implications.**

357
358 The model in Fig. 7 attempts to explain the Ebro paradox and the following scenario, which highlights the
359 multiple versus single knickpoint model implication in the fluvial dynamic during relative sea-level change.
360 From this simplified model, identification of a 270 km long Messinian Ebro incised-valley system and its
361 infilling by detrital FSST deposits allows us to propose an explanation clarifying the relationship between
362 drainage areas and the Messinian fluvial incision.

363 The existence of a pre-Messinian physiography and segmentation in the Valencia Basin (Pellen *et al.*, 2016)
364 plays a crucial role in its subsidence history, with an induced pre-Messinian step-by-step seafloor deepening
365 between the sagged VB, the MB and the deep LPB (Figs. 7 & 8a, stage 1).

366 Prior to the MSC sea level fall, the gradual restriction of the marine connections (CIESM, 2008; Roveri *et al.*,
367 2014) would imply relative minor sea level variation. As a fact a first retrogressive erosion Km_1 is located at the
368 CCR area, following the principle of the river equilibrium profile (Miall, 2013). The result is the initiation of
369 S11 basal erosive surface at the platform edge (Km_2) (Fig. 8a, Stage 1) and the detrital sediment deposition
370 (SU12a) offshore fluvial systems.

371 The progressive limitation of exchanges followed by the abrupt break of water exchanges through the Rifian
372 strait (Krijgsman *et al.*, 1999) at 5.60 Ma resulted in a rapid sea-level fall within the Mediterranean Basin. The
373 sea level quickly reached Km_2 (Fig. 8a, Stage 2a) implying a shift to the outermost limit of the Miocene platform
374 of the main retrogressive erosion. The eroded products were gradually deposited along the former Miocene
375 slope area, forming the SU12a unit (as drilled in Benicarlo well and showing ages of at least 6.14 to 5.72 Ma).
376 With the sea-level fall continuing, up to 5.6 - 5.45 Ma depending on the chronostratigraphy of the discussed
377 scenarios (i.e. CIESM, 2008; Bache *et al.*, 2012, 2015; Gorini *et al.*, 2015), the main regressive erosion shifted
378 again and the upper part of S12a was partially reworked and redeposited in the offshore domain as in FSST or
379 Forced regression models (Gorini *et al.*, 2015). Little detrital SU12a formation is observed at the VB base of
380 slope because, between stage 2a and 2b (Fig. 8a), the accommodation space in the Valencia foot slope area was
381 too small to accumulate thick series. This implies intense sedimentary transfer from the subaerial eroded
382 platform domain directly to the deep marine Menorca and Provence domains. When the sea level reached
383 Km_3 , the fluvial base-level and main retrogressive erosion migrated to the Valencia-Menorca transition with
384 clear erosional truncations (Fig. 8a, stage 2b) and the detrital sedimentation was transferred to the Menorca
385 and Provence deep domain, as shown by the development of the SU12c prograding clinoforms. At the lowest
386 sea-level, our interpretations locate the shoreline at Km_3 , observed on the CFZ axis (Fig. 3).

387 This evolution of downward shift erosion and the clear shift of the depocentres towards the deep basin
388 is typical of a falling stage system tract that occurs on rather flat areas separated by steps which involve a
389 proximal regressive erosion together with a distal retrogressive erosion at the step (knickpoint). As the sea-level
390 fall continues, we observe a quick shift of this double erosion from Km_1 to Km_3 . With this mechanism, the
391 river equilibrium profile has to be well developed at Km_3 so as to observe major retrogressive erosional process
392 at Km_2 area. Depending of the inertia of the retrogressive erosion dynamics at each knickpoint, there is no
393 increased fluvial erosion nor propagation towards the onshore Ebro areas as in the case of the Rhône (see

394 below). The existence of the CCR may also further prevent the propagation of the fluvial erosion towards the
395 inner parts of the EB.

396 In the Rhone system, earlier work showed the existence of three hinge lines in between different
397 structural domains related to the opening of the LPB. The model in Fig. 8b (stage 1-4) illustrates the
398 development of Messinian Pyrenean-Languedocian and Rhone fluvial systems during the MSC sea-level drop,
399 based on the work of Bache *et al.* (2009, 2010, 2012, 2015) and Gorini *et al.* (2015). Domain 2b corresponds
400 to the main geodynamic sill which records the gradual transition between marine and aerial domain during
401 the MSC sea-level fall (5.6-5.55 Ma) i.e. Gorini *et al.*, 2015) and slow sea-level rise (formation of wave
402 ravinement erosion surface between 5.55-5.46 Ma if we follow Bache *et al.*, 2012 scenario). Only minor onshore
403 knickpoints are mapped along the Rhone fluvial profile (Clauzon, 1982), and the main knickpoint is therefore
404 located at the hinge line 2 (the shelf-break) which also corresponds to the first development of forced falling
405 stage system tract of the MSC (Leroux *et al.*, 2015a, b). The model highlights the initiation of retrogressive
406 erosion at the second geodynamic knickpoint (the shelf break), which quickly progresses landward from the
407 former Miocene shelf-edge to the edge of the Central Massif (Pyrenean-Languedocian rivers) or along the
408 Rhone-Bresse valley (Fig. 8b, stage 1-3) up to Lyon city.

409 Domain 2a (GOL) can be compared to the VB basin area/VB-MB transition, which records the lowest
410 MSC base-level and early relative sea-level rise. Around the Km3 area, a minimum of four successive marine
411 incisions are observed in the UU deposits (Fig. 3, section e), implying (1) a composite refilling for the trans-VB
412 valley system and (2) preservation of the latter in the sense of Boyd *et al.*, (2006). If the UU are mainly developed
413 along the main channel (~0.3 twt thick) few deposits are observed on both sides of the channel, involving a
414 partial/complete reshaping of S20 erosional surface. This observation may be related to other morphological
415 features, such as the general extent of the smooth surface along the Catalan and domain 2b of the LPB (Bache
416 *et al.*, 2009, 2012; Garcia *et al.*, 2011; this study) (Fig. 6). This tends to favour a similar genesis during a relative
417 slow sea-level rise (Fig. 8, stage 4), the rough surface being eroded by wave action (Cattaneo and Steel, 2003;
418 Bache *et al.*, 2012, 2015) and only preserved upstream by the later rapid sea-level rise closing the Messinian
419 event.

420 We thus propose that domain 2b and VB basin area/VB-MB transition are equivalent morphological domains
421 which both record the lowest MSC sea-level. Significant lateral spacing between each morphological knickpoint
422 has provided better stratigraphic and morphological observation for the VB. As the VB record the whole
423 Messinian palaeo-Ebro incised valley, the Ebro paradox disappears if the length of Messinian rivers are
424 measured from MSC paleo-shorelines at the CFZ area, and not from pre-Messinian or present-day shorelines
425 (Fig. 7).

426 At the MSC paroxysm, the trans-VB incised valley system watershed would include the Valencia Basin and its
427 borders (55 000 km²; including the Betic-influenced southwest area, its mountainous borders (southern part
428 of the Catalan Range) and at least the eastern part of the Ebro Basin (? < Ebro watershed < 85 000 km²).
429 However, our result cannot confirm or infirm a pre-Messinian proto-Ebro hypothesis. Stratigraphic and
430 thermochronological interpretations path suggest a pre-Messinian opening of the Ebro Basin (Fillon *et al.*,
431 2012; Cameselle *et al.*, 2013) with a single proto-Ebro river or multiple mountainous rivers cutting across the
432 CCR. These models need to be confirmed by an assessment of offshore sedimentary volumes during Miocene,
433 Messinian and Pliocene-Quaternary times that would lead to a better estimate of the paleo-drainage area
434 evolution before/after the opening of the EB.

435 The NBFZ sill is overloaded by igneous pulses starting during the LU deposition and during Pliocene-
436 Quaternary time (Fig. 9; this was also observed onshore by Marti *et al.*, 1992; Araña *et al.*, 1983). Recent studies
437 have linked this Messinian magmatic pulse to the Messinian desiccation (Sternai *et al.*, 2017) or to a more
438 global kinematic re-organization (Leroux *et al.*, 2018). Whatever the lithospheric processes implied, the Ebro
439 incised-fluvial system constrained by the CCR and by the CFZ confirms the strong influence of VB-MB-LPB
440 margin structural segmentation on the development of the Messinian Ebro fluvial system, and more generally
441 on the whole Neogene sedimentological evolution.

443 5. Conclusion

444 Extending the Cameselle *et al.* (2014) results, we have identified and characterized along the Valencia, Menorca
445 and towards Liguro-Provence basins a 270 km long Messinian fluvial valley-incision - from the Messinian
446 shoreline to the CCR - as well as the spatial distribution of the related distal detrital deposits. The Ebro
447 Messinian erosional system now displays the same order of incised river length as the Rhone, resolving the
448 paradoxal relationship between the present-day fluvial drainage area and the Messinian fluvial system of
449 Babault *et al.*, (2006). Following our interpretation, the Ebro Basin and Mediterranean Sea were connected at
450 least at the MSC sea-level drop. Only detailed sedimentary fluxes analysis could confirm or infirm a pre-
451 Messinian connection between the Ebro Basin and the Mediterranean Sea.

452 The segmentation and nature of the substratum triggered the formation of knickpoints due to differential
453 subsidence between the shallow VB, the transitional MB and the deep LPB, and controlled fluvial processes
454 and erosion during the MSC sea-level drop. These knickpoints induced step-by-step major regressive-
455 retrogressive fluvial erosional processes which shift through time during sea-level drop. This work highlights
456 the general major relationship between kinematic, subsidence evolution, base-level fall and depositional
457 response particularly well expressed during the outstanding Messinian Salinity Crisis event.

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Acknowledgments

This work was co-funded by a grant from the French government under the program "Investissements d'Avenir" the "Laboratoire d'Excellence" LabexMER (ANR-10-LABX-19), and ISBLUE (ANR-17-EURE-0015) and by a grant from the Regional Council of Brittany. It was further supported by CNRS and IFREMER, with additional support from the French Actions-Marges program (JL Rubino & P. Unternehr) and the GRI Méditerranée (Groupement Recherche et Industrie TOTAL-UPMC). The Data base was built thanks to the SIGEOF spanish site, to academic cruises from France including PROGRES, AMED-SARDINIA, SEEPGOL, VALSIS) and Spain. Additional industrial seismic lines were provided by Schlumberger. The biostratigraphic studies and re-evaluations were performed by biostratigraphers from TOTAL. The authors acknowledge the fruitful and constructive reviews by Antonio Pedrera, Agnes Maillard and two anonymous reviewers, as well as those from the editor Liviu Matenco that greatly improved the manuscript.

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746 Figures

748 **Fig.1.** A Present-day NW Mediterranean physiographic areas **and** previously published major
749 physiographic features and seismic units assigned to the Messinian event (compiled from Stampfli and
750 Hoëcker, 1989; Frey-Martinez *et al.*, 2004; Maillard *et al.*, 2006a, b; Del Olmo, 2011; Garcia *et al.*, 2011; Urgeles
751 *et al.*, 2011; Bache *et al.*, 2009, 2015). Major geodynamic boundaries are represented in the Valencia-Menorca
752 Basin (VB-MB): C.F.Z. = Central Fracture Zone; N.B.F.Z. = North Balearic Fracture Zone) (Maillard and
753 Mauffret, 1999; Pellen *et al.*, 2016) and in the Liguro-Provence Basin (LPB): domain 1 = Unthinned
754 Continental Crust; domain 2a = Thinned Continental Crust.; domain 2b = Highly thinned Continental Crust;
755 domain 3 = Exhumed lower crust; domain 4 = Atypical oceanic crust) (Moulin *et al.*, 2015; Leroux *et al.*, 2015).

756 **Fig.2.** Map showing our new interpretation of the sedimentary units and main morphological
757 observations associated to the Messinian event in the VB-VM. The Messinian drainage network (in red) could
758 be followed from the shelf to the VB and up to the MB. Position of vertical sections a to f shown in black (see
759 Fig. 3).

760 **Fig.3.** Seismic sections a-f showing the morphological evolution of the trans-VB Messinian valley system,
761 its outlet and the related Messinian detrital and evaporitic deposits. The red marker symbolizes MES and S20
762 rough surface, the violet markers symbolize the smoother S21 surface at the outlet of the valley system.
763 Transect01 is a composite dip section converted in depth (see methods and supplementary figures for more
764 details) from the VB to the deep LPB. Note in particular the progressive downward shifts of top of units 12a
765 and 12c). Position of the profiles is given in Fig. 2.

766 **Fig. 4.** Detailed seismic section highlighting the development of a Miocene ante-MSC platform along
767 the Ebro-Tortosa area (section g) which mark Km2 knickpoint. The development of SU12a is associated to a
768 regressive surface of marine erosion (RSME) (section g) and to a major change in the origin of sediment supply
769 (section h). MES morphological region (section g) is deduced from Urgeles *et al.* (2011).

770 **Fig. 5.** Detailed seismic section highlighting the erosional truncations at the transition between the VB
771 and the MB domain (section i) and development of the progradational seismic units SU12c in the deep and
772 distal MB area (section j). We interpret this SU12c unit as the LST related to the MSC, it corresponds to the
773 most distal progradating detritic unit observed in the area.

774 **Fig. 6.** Correlation of seismic units and Well Benicarlo C1, for the Miocene to Pliocene deposits.
775 (Borehole position is shown on Figure 1 & 2). The right part of the figure shows the vertical records of the
776 concerned species, the horizontal lines indicate which level has been considered as corresponding to the first
777 or last appearance datum (FAD or LAD) of some of them. Biozones and related ages refer to Gradstein *et al.*
778 (2012). Caliper (blue curve) and Gamma Ray (red curve) are represented.

779 **Fig.7.** Simplified schematic section illustrating the role of multiple versus single knickpoints in
780 retrogressive erosion response during a relative sea level fall. **a, b, c** correspond to the rejuvenation phase of
781 the retrogressive erosion linked to the relative sea level successive falls (step 1, 2, 3).

782 **Fig.8.** Schematic and balanced cross section of the VB-MB and LPB (modified from Bache, 2008)
783 illustrating the impact of sills and sea-level drop on the fluvial valley system development and paleo-depositional
784 profile process during the Messinian event. During sea-level drop, erosion occurs both as regressive erosion of
785 the fluvial system but also as retrogressive erosion from more distal knickpoints (see text for further
786 explanation).

787 **Fig.9.** Detailed seismic section highlighting the magmatic pulse during the Lower Unit deposition and
788 Pliocene-Quaternary time at the Menorca Basin and Liguro-Provence Basin transition (North Balearic Fracture
789 Zone area) (Location and units color code as in Figs. 4 & 5).

792 SUPPLEMENTARY MATERIAL

793
794 **Supplementary material 1. A)** Time-depth velocity model used for the time to depth conversion of the
795 Transect01 using velocity information from boreholes (Benicarlo C1, Bocarte 1, Castellon G1, Tarragona D2
796 and Fornax 1) and ESP refraction data in the VB and GOL.

797 **Supplementary material 2.** Backstripping reconstruction along the regional Transect01 profile (in
798 depth). Input parameters for the VB domain include palaeobathymetry, age, and porosity of sedimentary
799 packages based on Urgeles *et al.* (2011). Rabineau *et al.* (2014) presented a new method to quantify post-rift
800 subsidence by direct use of sedimentary geometries. An estimated total subsidence value up to 960 m/Ma for

801 the deep LPB domain was estimated for Pliocene and Quaternary times ([Rabineau et al., 2014](#)) and applied to
802 the 2-D backstripping. MES, S20 and base of Halite surfaces were reported on Fig. 3. Q1, Q2, P1 and P2 are
803 respectively dated at 0.9, 1.8, 2.49 and 3.2 Ma following [Urgeles et al. \(2011\)](#) and [Leroux et al. \(2015b\)](#) studies.

804 **Supplementary material 3.** Summary of key seismic surfaces and units mapped across VB, MB, and LPB
805 areas in this study. Unit age estimation are from [Bache et al. \(2015\)](#) and [Gorini et al. \(2015\)](#).

806 **Supplementary material 4.** Short video illustrating the evolution of sea-level and paleogeographies in
807 the VB, MB and LPB between 6 and 5.20 Ma (using Placa 4D©ifremer software).

808

FIGURE 01

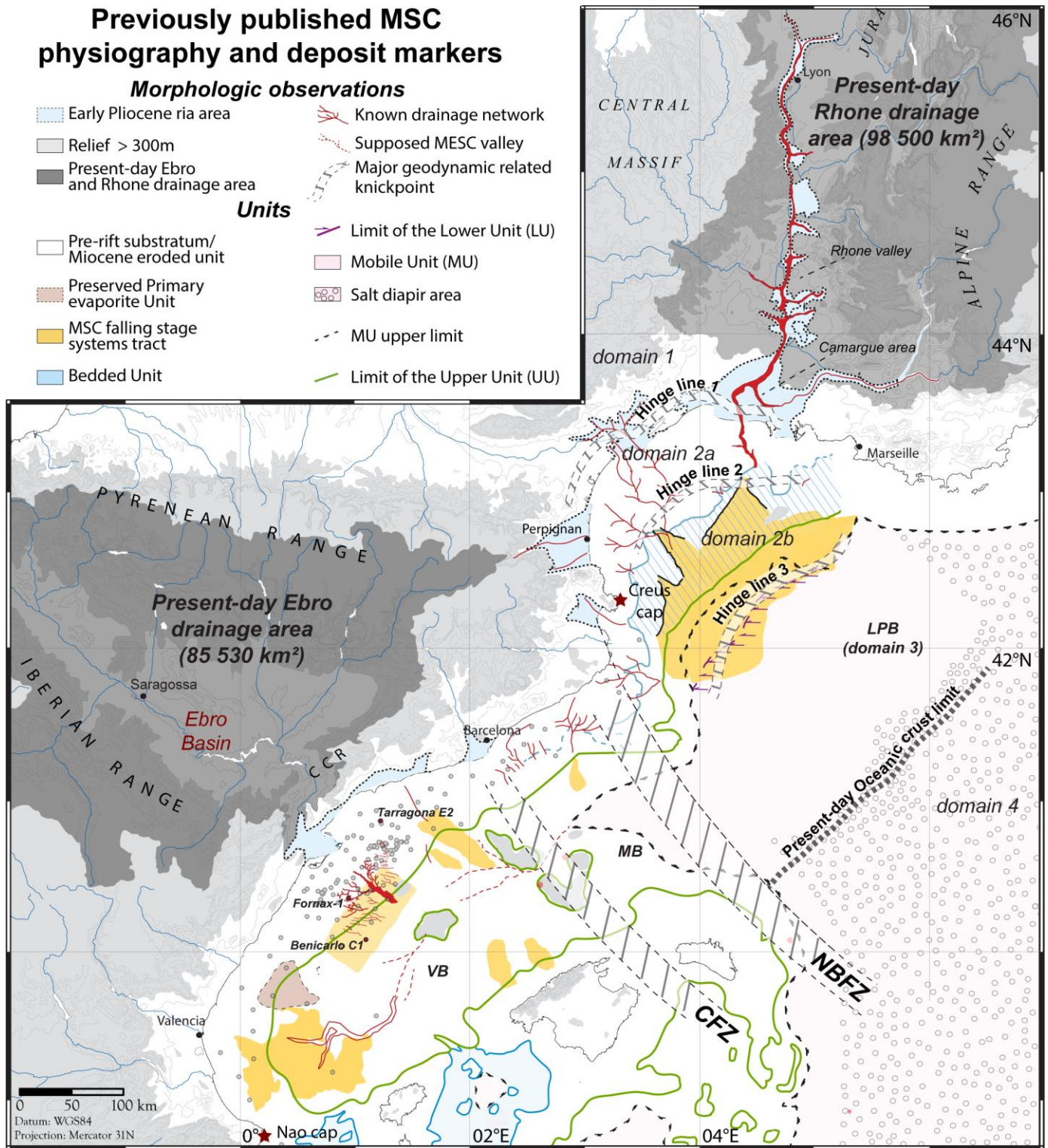


FIGURE 02

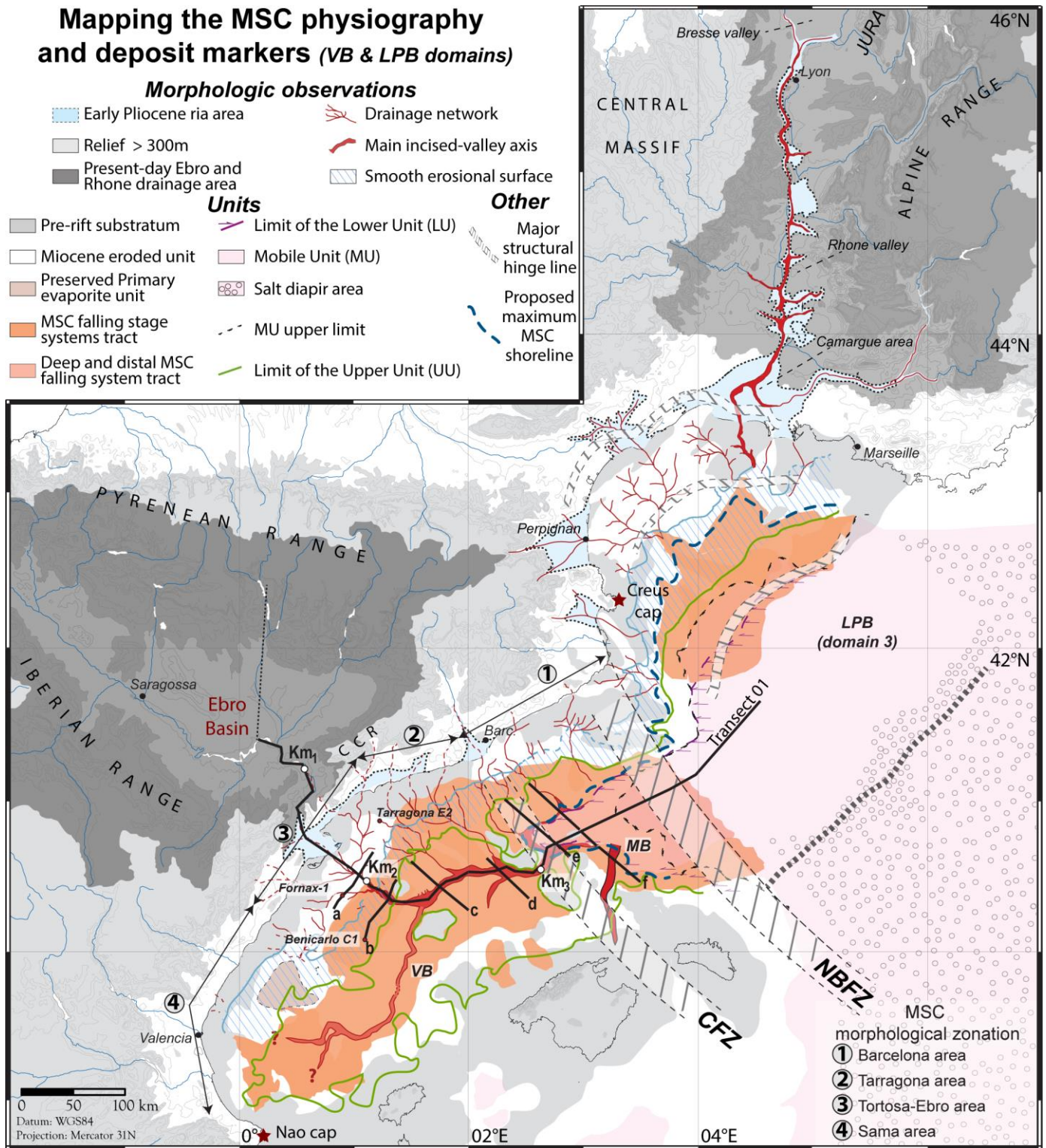
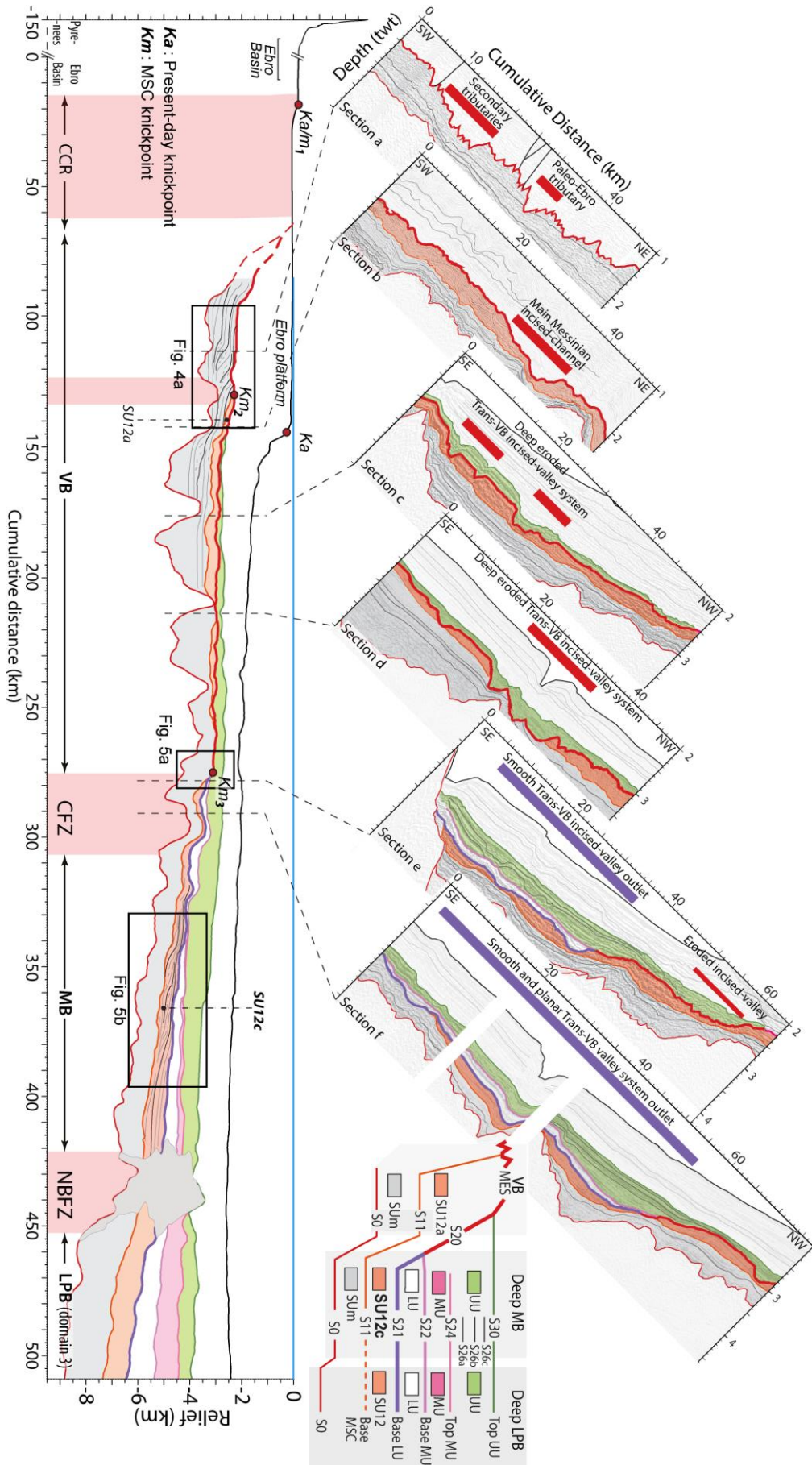
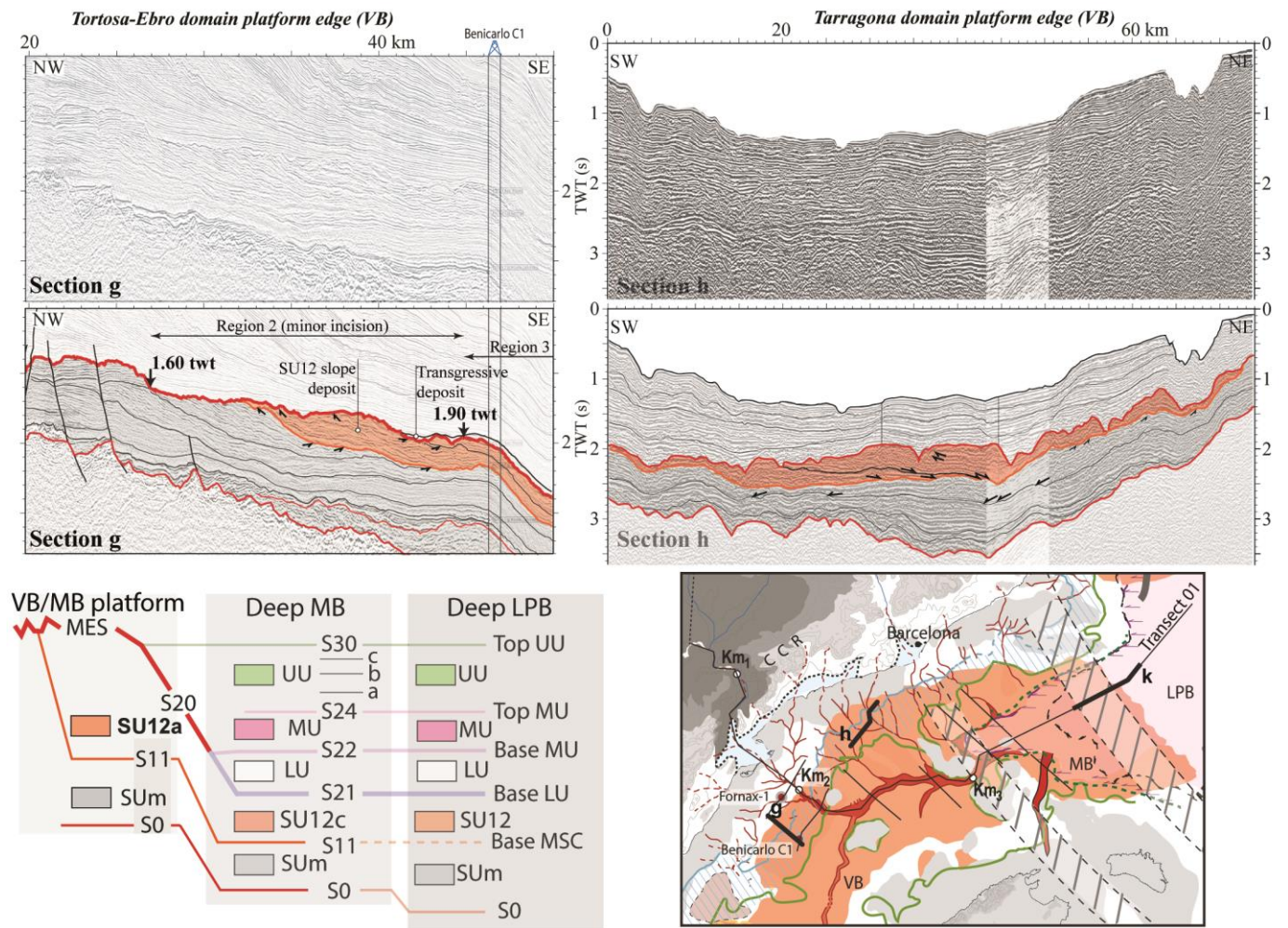


FIGURE 03



817

FIGURE 04



818

819

FIGURE 05

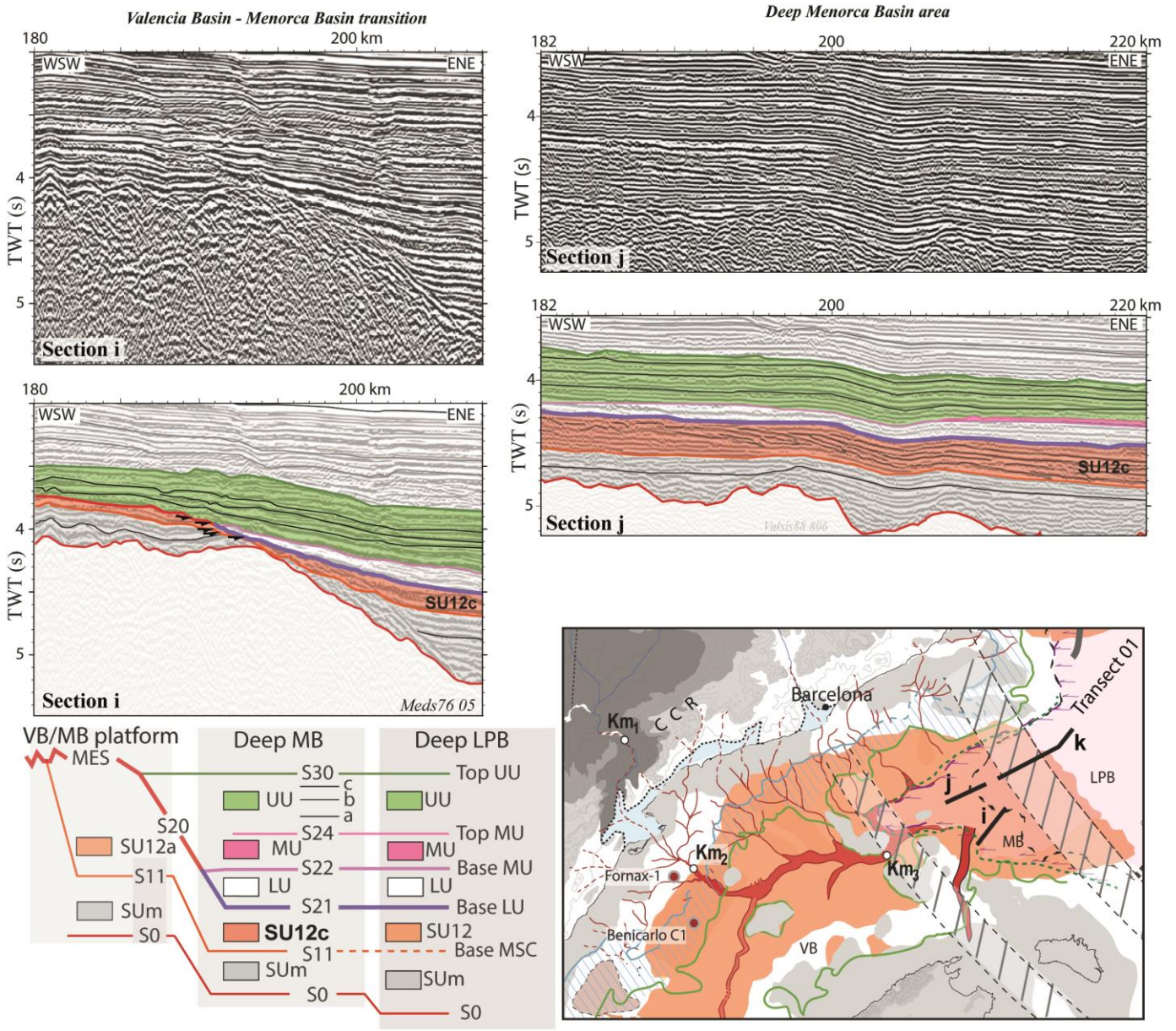
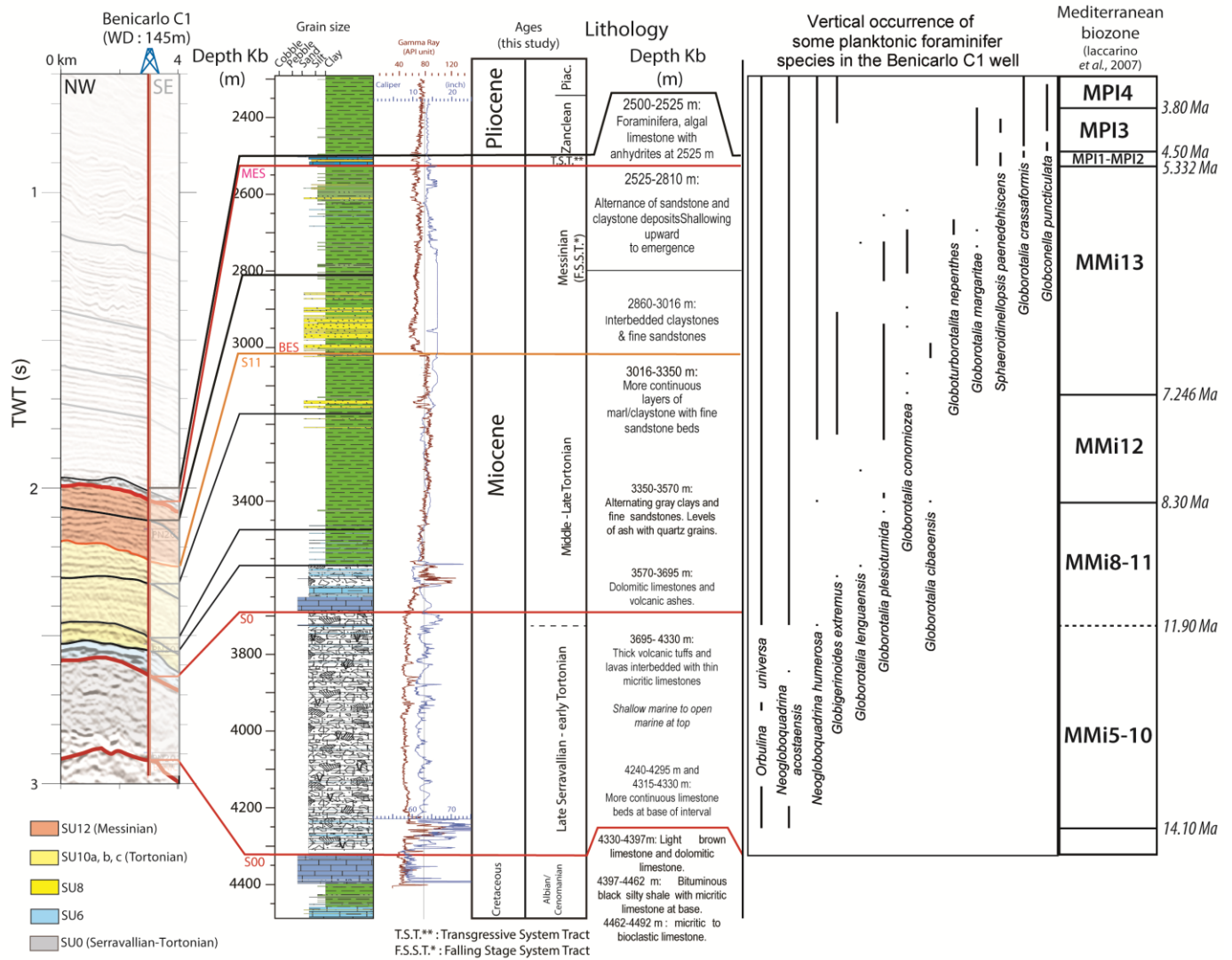


FIGURE 06



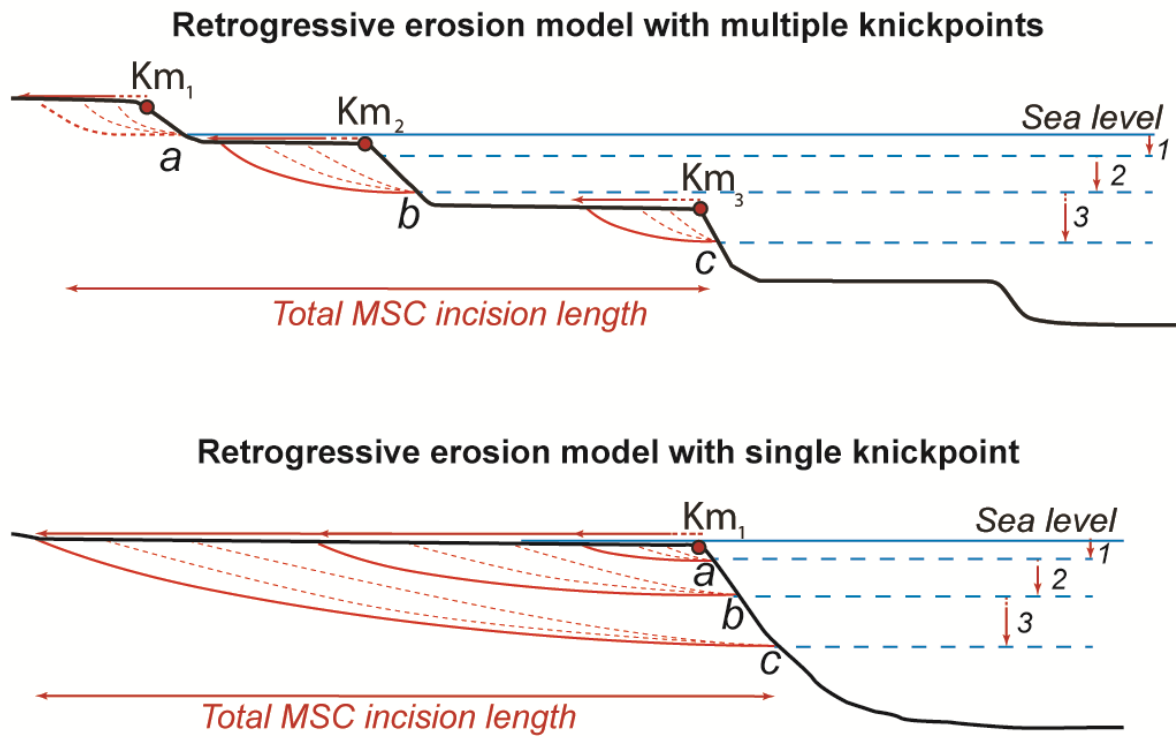
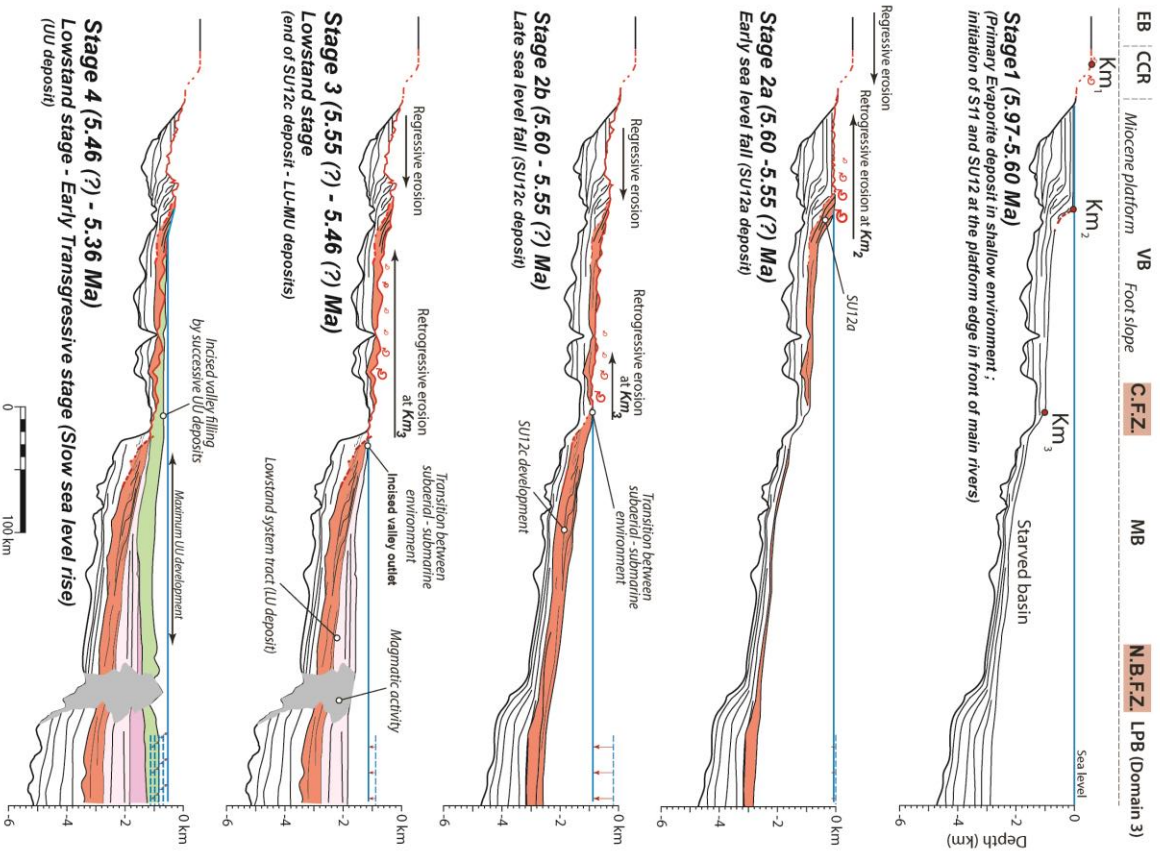


Figure 07

Valencia-Menorca domains



Liguro-Provence domain

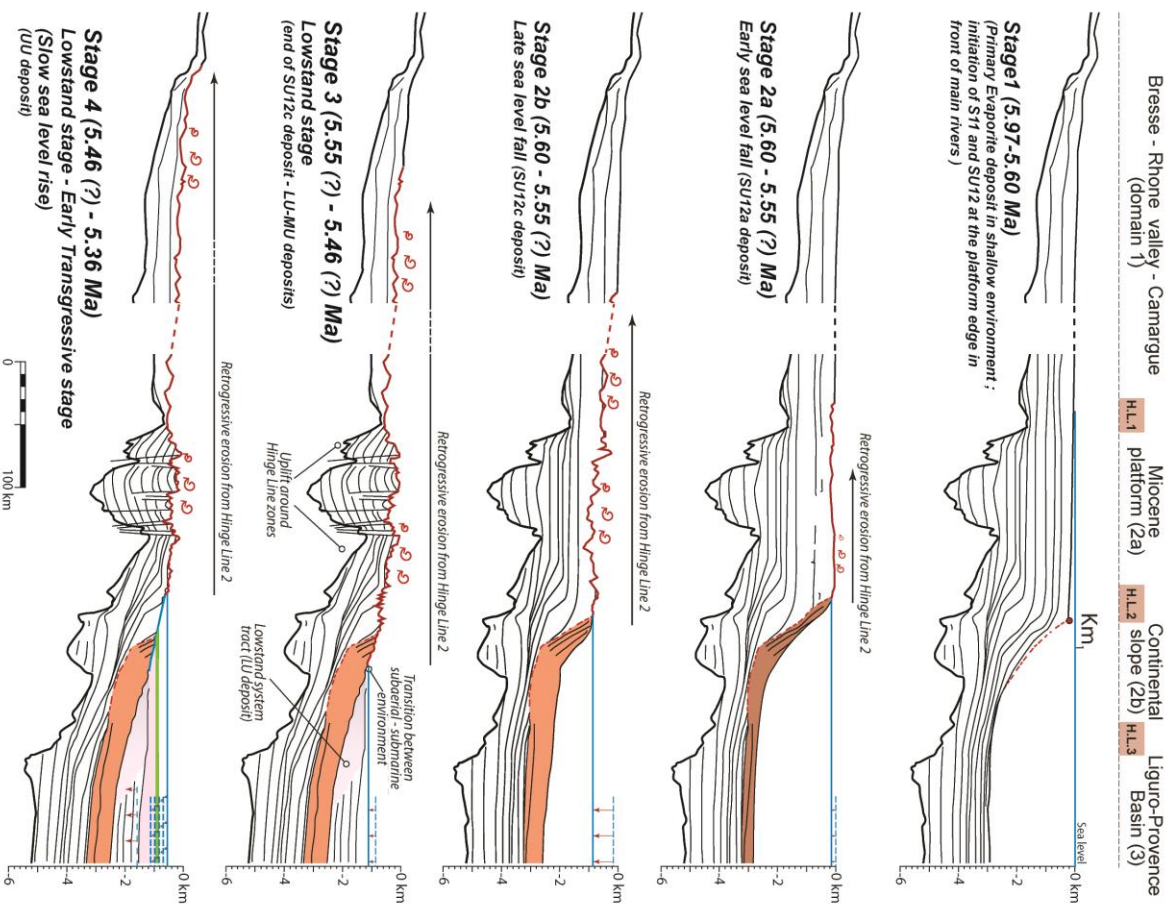
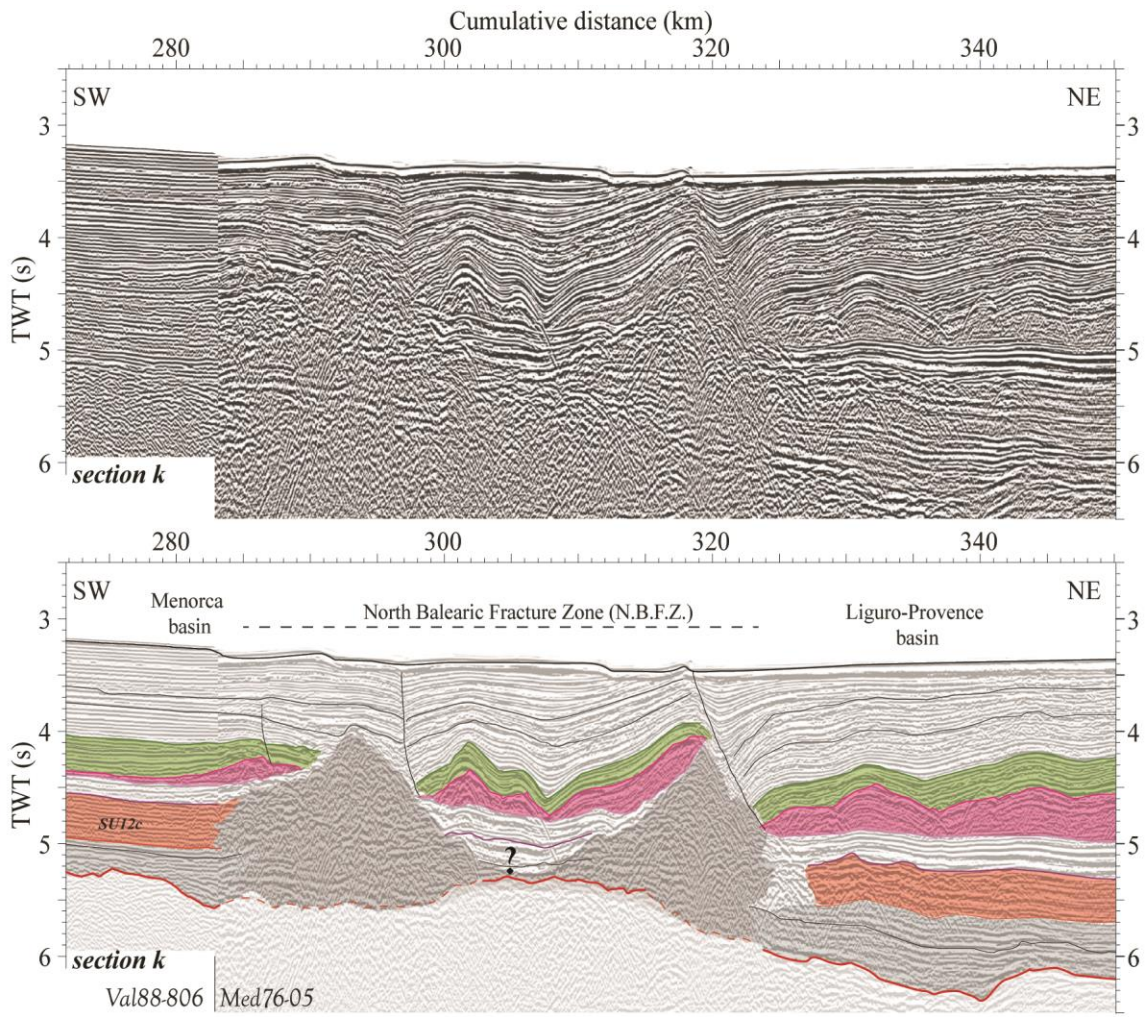


FIGURE 08

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FIGURE 09

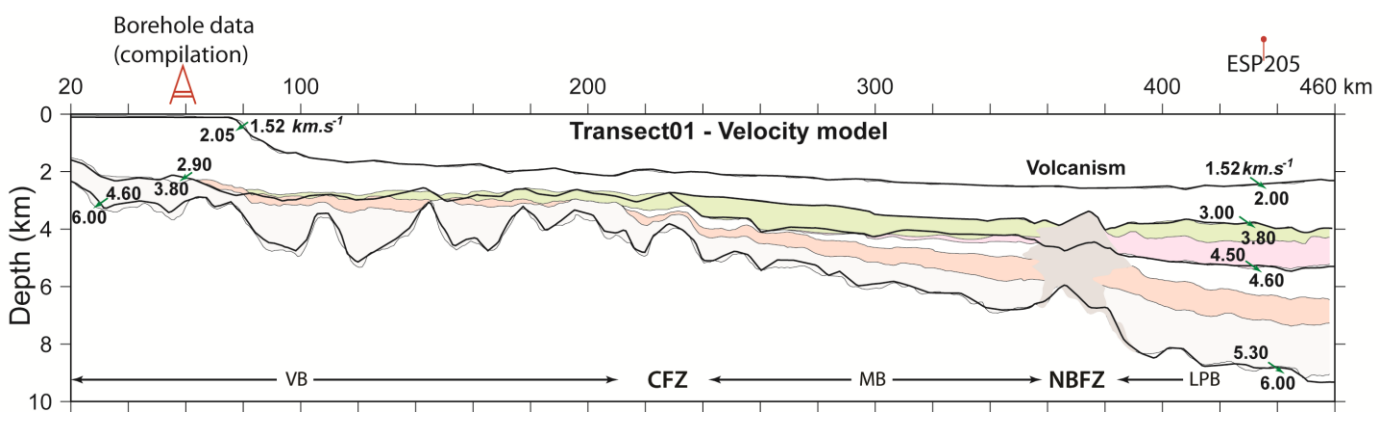


832

833

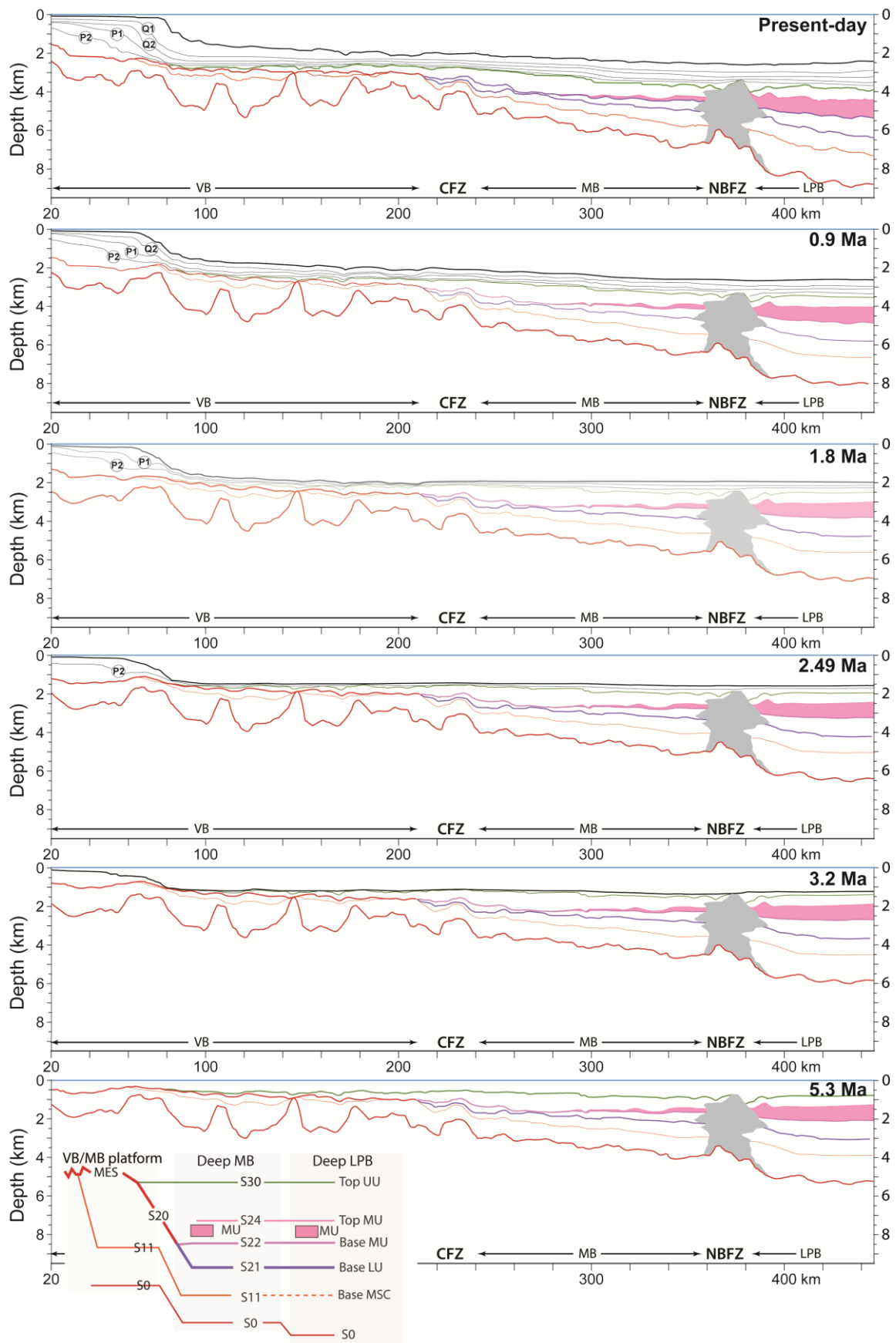
834

SUPPLEMENTARY MAT 01



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SUPPLEMENTARY MAT 02



Seismic discontinuities	Seismic unit	Seismic facies	Spatial location and observations				Age known/supposed	Interpretations	
			VB/MB Platform	VB footslope	MB footslope	MB deep basin LPB basin (dom. 3)			
Messinian Erosional Surface (MES)		Rough erosive discontinuity	Rough erosive discontinuity until ~1.6 twt	Lateral transition to S20 and S30 at ~ 2.4-2.7 twt			Diachron surface: 5.96-5.32 Ma	Fluvial erosional surface linked to sea level fall	
Ravinement Surface (RS) (Hachured blue layer)		Flat discontinuity	- MES/S30 reshaped surface - Conform with SU12 and UU units	—	—	—	5.46 (?) - 5.32 Ma	Marine abrasion surface (flattened by the waves)	
S30		Erosive discontinuity	—	Upper lateral correlation of MES. Locally conform; high amplitude reflection			5.46 (?) - 5.32 (?) Ma	Marine erosional surface linked to sea level oscillation	
	Upper Unit (UU) UUb, c, d	Continuous, sub-parallel reflections of medium-high amplitude	—	UUd development Filling of the incised valley	UUa,b,c,d development with erosional features	UUa, b, c, d max.development Sub-parallel horizons	UU undifferentiated unit Sub-parallel horizons	5.46 (?) - 5.32 Ma	Mixed evaporitic-detrital unit
S26a, b, c		Erosive discontinuity	—	Upper lateral correlation of MES. Locally conform; high amplitude reflection			5.46 (?) - 5.32 Ma	Marine erosional surface linked to sea level oscillation	
	UUa	Similar as UUb, c, d	—	—	- At the Top of MU - UUa development conform to MU top unit		5.46 (?) - 5.32 Ma	Mixed evaporitic-detrital unit	
S24		Mobile Unit (MU) top surface	—	—	—	Onlap on S21 (Lower Unit top surface)			
	Mobile Unit (MU)	Transparent facies	—	—	—	Full development in LPB Above LU	5.60 (?) - 5.46 (?) Ma	Halite	
S22		Conform, continuous high amplitude surface	—	—	—	- MU basal surface -onlap on S21 surface			
	Lower Unit (LU)	Bedded, continuous reflectors of low frequency	—	—	—	- Initiation at Trans-VB outlet at Km3 area - Full development in LPB domain	5.60 (?) - 5.46 (?) Ma	Evaporitic or Mixed evaporitic and detrital	
S21		Conform surface ; Lateral evolution of S20	—	—	—	- Slightly erosive at Km3 area - Conform with SU12c and LU			
S20		Erosive discontinuity; margin lateral correlation to MES; Basinward correlation to S21	—	- Erosive surface along trans-VB valley and paleo-relief. - Slightly erosive from either side of the incised valley.		—	—		
	SU12c (Menorca basin)	Prograding clinoform; sub-continuous with moderate amplitude reflectors	—	—	—	- Below LU unit; conformable with LU toward LPB domain; - ONO-ESE prograding trend.	5.97-5.46 (?) Ma	products of messinian erosion	
	SU12a (margin)	sub-continuous with moderate amplitude reflectors; cutt and fill geometry	- Below MES and S20 - Above S11; - Bottomset reflectors	- Downstream extension limited by basement high; - Lateral transition with SU12c - SU12a truncated by SU12b		—	—	5.97-5.46 (?) Ma	Earlier marine products of messinian erosion
	SU12b (mass transport deposit) (Maillard et al., 2006)	Chaotic reflectors	—	- Below MES and S20 - Limited by the trans-VB incised valley		—	—	5.97-5.46 (?) Ma	products of messinian erosion/ link to isostatic rebound?
S11		Erosive discontinuity	- Below SU12 units;	- Below SU12 units; - Slightly erosive near paleo-relief; - Conformable surface in the MB / LPB				5.97-5.46 (?) Ma	Basal marine erosive surface
	Primary Evaporite (Del Olmo, 2011) (not show in this study)	Two continuous reflectors with medium-high amplitude	- Below MES and S20 - Only observed along the Sama area	—	—	—	—	5.97-5.60 Ma	Evaporitic deposit in shallow water depth

