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## Satellite radar altimetry water elevations performance over a 200 m wide river: Evaluation over the Garonne River

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1 **Satellite radar altimetry water elevations performance over a 200 m wide river: evaluation**  
2 **over the Garonne River**

3  
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25 **Abstract**

26 For at least 20 years, nadir altimetry satellite missions have been successfully used to first  
27 monitor the surface elevation of oceans and, shortly after, of large rivers and lakes . For the last 5-  
28 10 years, few studies have demonstrated the possibility to also observe smaller water bodies than  
29 previously thought feasible (river smaller than 500 m wide and lake below 10 km<sup>2</sup>). The present  
30 study aims at quantifying the nadir altimetry performance over a medium river (200 m or lower  
31 wide) with a pluvio-nival regime in a temperate climate (the Garonne River, France). Three  
32 altimetry missions have been considered: ENVISAT (from 2002 to 2010), Jason-2 (from 2008 to  
33 2014) and SARAL (from 2013 to 2014).

34 Compared to nearby in situ gages, ENVISAT and Jason-2 observations over the lower Garonne  
35 River mainstream (110 km upstream of the estuary) have the smallest errors, with water elevation  
36 anomalies root mean square errors (RMSE) around 50 cm and 20 cm, respectively. The few  
37 ENVISAT upstream measurements have RMSE ranging from 80 cm to 160 cm. Over the estuary,  
38 ENVISAT and SARAL water elevation anomalies RMSE are around 30 cm and 10 cm, respectively.  
39 The most recent altimetry mission, SARAL, does not provide river elevation measurements for  
40 most satellite overflights of the river mainstream. The altimeter remains “locked” on the top of  
41 surrounding hilly areas and does not observe the steep-sided river valley, which could be 50 m to  
42 100 m lower. This phenomenon is also observed, for fewer dates, on Jason-2 and ENVISAT  
43 measurements. In these cases, the measurement is not “erroneous”, it just does not correspond to  
44 water elevation of the river that is covered by the satellite. ENVISAT is less prone to get ‘locked’ on  
45 the top of the topography due to some differences in the instrument measurement parameters,  
46 trading lower accuracy for more useful measurements. Such problems are specific to continental  
47 surfaces (or near the coasts), but are not observed over the open oceans, which are flatter.

48 To overcome this issue, an experimental instrument operating mode, called the  
49 DIODE/DEM tracking mode, has been developed by CNES (Centre National d’Etudes Spatiales)  
50 and has been tested during few Jason-2 cycles and during the first SARAL/AltiKA cycle. This

51 tracking mode “forces” the instrument to observe a target of interest, i.e. water bodies. The example  
52 of the Garonne River shows, for one SARAL ground track, the benefit of the DIODE/DEM tracking  
53 mode for a steep-sided river reach, which is not detected using the nominal instrument operating  
54 mode. Yet, this mode relies on ancillary datasets (a priori global DEM and global land/water mask),  
55 which are critical to obtain river valley observation. The ultimately computed elevations along the  
56 satellite tracks, loaded on board, should have an absolute vertical accuracy around 10 m (or better).  
57 This case also shows, when the instrument is correctly observing the river valley, that the altimeter  
58 can detect water bodies narrower than 100 m (like an artificial canal).

59

60 In agreement with recent studies, this work shows that altimeter missions can provide useful  
61 water elevation measurements over a 200 m wide river with RMSE as low as 50 cm and 20 cm, for  
62 ENVISAT and Jason-2 respectively. The seasonal cycle can be observed with the temporal sampling  
63 of these missions (35 days and 10 days, respectively), but short term events, like flood events, are  
64 most of the time not observed. It also illustrates that altimeter capability to observe a river is highly  
65 dependent of the surrounding topography, the observation configuration, previous measurements  
66 and the instrument design. Therefore, it is not possible to generalize at global scale the minimum  
67 river width that could be seen by altimeters.

68 This study analyzes, for the first time, the potential of the experimental DIODE/DEM  
69 tracking mode to observe steep-sided narrow river valleys, which are frequently missed with  
70 nominal tracking mode. For such case, using the DIODE/DEM mode could provide water elevation  
71 measurements, as long as the on board DEM is accurate enough. This mode should provide many  
72 more valid measurements over steep-sided rivers than currently observed.

73

74 **Keywords:** satellite altimetry; Garonne River; DIODE/DEM mode; ENVISAT; Jason-2; SARAL

75

76 **1. Introduction**

77 Continental waters play a key role in the Earth water cycle and are subject to complex  
78 interactions at the interface between the atmosphere and ocean. These waters directly impact human  
79 societies through food consumption, agriculture, and industrial activities. Continental waters need to  
80 be monitored, especially during flood or drought events. They are also directly impacted by human  
81 activities, through pumping, river embankment, dams, reservoirs, and hydraulic infrastructure. The  
82 monitoring of the spatial distribution and temporal variability of surface waters still remains  
83 challenging: there could be around 117 million of water bodies with an area above 0.002 km<sup>2</sup> on the  
84 continental surfaces according to a recent study (Verpoorter et al., 2014). The biggest lakes and  
85 rivers are of course important to the study of global hydrological process and water/carbon cycles.  
86 But smaller lakes and rivers begin to draw attention, as they might also play a non-negligible role  
87 because of their numbers (e.g. Downing et al., 2010). In situ monitoring of water bodies at global or  
88 even regional scales is very heterogeneous, as it depends on local gage networks,. Moreover, in situ  
89 measurements are considered sensitive information and are not always freely available to the  
90 research community. In this context, satellite measurements are a viable complementary source of  
91 information and especially those from nadir altimeters, even if they will not replace in situ  
92 measurements.

93 Initially designed to monitor the dynamic topography of the ocean, satellite radar altimeters  
94 have proven their abilities to observe continental surfaces water bodies elevations, allowing long-  
95 term observations of water level variations of lakes (e.g., Birkett, 1995; Ponchaut and Cazenave,  
96 1998; Medina et al., 2008; Crétaux et al., 2011; 2015), rivers (e.g., Birkett, 1998; Frappart et al.,  
97 2006a; Santos da Silva et al., 2010; Biancamaria et al., 2011; Michailovsky et al., 2012) and  
98 floodplains (e.g., Frappart et al., 2005; 2006b; 2012; Lee et al., 2009; Santos da Silva et al., 2010).  
99 As altimetry measurements demonstrated their abilities to provide reliable water stages over large  
100 water bodies, they started to be used with or to substitute missing in situ data, especially in large  
101 remote river basins. Major hydrological applications are currently the followings: calibration of  
102 hydrodynamics models (e.g., Wilson et al., 2007; Getirana, 2010; Yamazaki et al., 2012; Paiva et al.,

103 2013), estimate of discharge using either rating curves (e.g., Kouraev et al., 2004; Papa et al., 2012),  
104 routing models (e.g., León et al., 2006; Hossain et al., 2014; Michailovsky and Bauer-Gottwein,  
105 2014), or coupling with measurements of river velocities from multi-spectral images (e.g.,  
106 Tarpanelli et al., 2015), estimate of surface water storage in large river basins (e.g., Frappart et al.,  
107 2012; 2015a), lakes and reservoirs storage dynamics (e.g. Crétaux et al., 2016) and low water maps  
108 of the groundwater table (Pfeffer et al., 2014). Nadir altimetry is quite a mature technology, as the  
109 very first scientific altimeters flew more than thirty years ago. Continuity of measurements in time  
110 is guaranteed by incoming follow-on missions like Jason-3 (launched 17 January 2016), Sentinel-  
111 3A (launched 16 February 2016), Sentinel-3B (currently planned for 2017), Sentinel-3C and -3D  
112 (which should be launched around 2021; ESA, 2016) and Jason-Continuity of Service/Sentinel-6 (in  
113 2020 for Jason-CS A and 2026 for Jason-CS B; Scharroo et al., 2015).

114 Radar altimetry, however, has several limitations for monitoring land hydrology. The main  
115 restrictions are its low spatial (one measurement every 175 to 400 m along track, for an instrument  
116 footprint with several kilometers radius and an intertrack distance at the equator between 80 km to  
117 315 km, depending on the mission) and temporal resolutions (repeat cycle of 10 days for  
118 Topex/Poseidon and Jason-1/2/3 missions and 35 days for ERS-1&2, ENVISAT and SARAL, 27  
119 days for Sentinel-3 missions). Because of these limitations, the use of altimetry data has been  
120 limited to large (tens of km wide or wider) water bodies. Moreover, altimeters miss most water level  
121 extrema during extreme flow periods or fail to study rapid hydrological events such as flash floods.  
122 However, the altimeter performance depends not only on river size but also on the surrounding  
123 topography (better performance over flat areas), on other surrounding water bodies and, to some  
124 extent, on vegetation that will affect the reflected electromagnetic wave (Frappart et al., 2005;  
125 Frappart et al. 2006a; Frappart et al., 2006b; Santos da Silva et al., 2010; Crétaux et al., 2011; Ricko  
126 et al., 2012).

127 Improvements in nadir altimetry sensors performance, in the quality of the corrections  
128 applied to the altimetry range and in measurements post-processing have allowed measurements of

129 water stage variations over small-to-medium rivers and small lakes. Small rivers are 40 to 200 m  
130 wide, while medium rivers have widths between 200 and 800 m. Small rivers discharge from 10 to  
131  $100 \text{ m}^3 \cdot \text{s}^{-1}$  while medium range between 100 and  $1000 \text{ m}^3 \cdot \text{s}^{-1}$  (Meybeck et al., 1996). For lakes,  
132 small lakes have areas  $<0.01 \text{ km}^2$  (Verpoorter et al., 2014). On the basis of a global inventory of  
133 lakes from optical satellite images, Verpoorter et al. (2014) showed that small lakes are the most  
134 numerous (90 million, from a total of 117 million lakes worldwide), but cover a total area (0.27% of  
135 non-glaciated land surfaces) much smaller than bigger lakes (which cover a total of 3.5%). Despite  
136 their importance for land hydrology and water resources management, a large number of rivers and  
137 lakes are poorly gaged (Alsdorf and Lettenmeier, 2003). Few studies have demonstrated the  
138 possibility to accurately monitor water levels of small water bodies (e.g., Santos da Silva et al.,  
139 2010; Michailovsky et al., 2012; Baup et al., 2014; Frappart et al., 2015a; Sulistioadi et al., 2015).  
140 The present study aims at doing this benchmarking for a medium river: the Garonne River in  
141 France. Section 2 presents the study domain, in situ gages and radar altimetry missions used.  
142 Section 3 compares in situ and altimetry water elevations along the river main course and its  
143 estuary, discuss the sources of errors and investigate potential solution for future altimetry missions  
144 to improve measurements. Conclusions and perspectives are provided in section 4.

145

## 146 **2. Study domain and Methodology**

147

### 148 **2.1. Garonne River basin presentation and available data**

149

150 The Garonne River (Figure 1) is located in Southwest France and drains an area of 56,000  
151  $\text{km}^2$ . Its mean annual discharge near its outlet, at Tonneins where the river width is around 200 m  
152 (Figure 1), is around  $600 \text{ m}^3 \cdot \text{s}^{-1}$ . At the global scale, according to the Global Runoff Data Center  
153 (GRDC) discharge database, it is the 120<sup>th</sup> largest river in the world by its annual discharge and the  
154 3<sup>rd</sup> in mainland France. It is therefore a medium river according to Meybeck et al. (1996) (section  
155 1).

156 The Garonne River has a pluvio-nival regime, with a low flow period between July and

157 October and high flow period between December and April. The river source is located in the  
158 Pyrénées Mountains (South of the basin, Figure 1) and outflows to the Atlantic Ocean via the  
159 Gironde estuary. The Garonne supports an agricultural activity that uses 70% of the total water  
160 uptake (mainly from surface waters) during low flow period (Sauquet et al., 2009; Martin et al.,  
161 2016). For more details on the Garonne basin, see Martin et al. (2016).

162 Water level and discharge gages on most rivers in France are operated by regional public  
163 agencies (DREAL – Directions Régionales de l'Environnement, de l'Aménagement et du Logement)  
164 and all their measurements are collected by the Service Central de l'Hydrométéorologie et d'Appui à  
165 la Prévision des Inondations (SCHAPI) within the national 'Banque Hydro' database  
166 (<http://www.hydro.eaufrance.fr>). Four gages from this database have been used in this study  
167 (Verdun-sur-Garonne, Lamagistère, Tonneins and Marmande, see Figure 1), as they are on the  
168 Garonne mainstream and provide validated water level measurements. Data are available with a  
169 non-uniform temporal resolution that depends on the water elevation stage (the median time step for  
170 all gages in Figure 1 is below 1 hour). All gages have records starting before 01 January 1 2002 (the  
171 first year of the oldest altimetry mission considered in this study, see section 2.2.3). Water elevation  
172 time series used in this study end 31 December 2010 at Verdun-sur-Garonne, 02 April 2014 at  
173 Lamagistère, 01 February 2015 at Tonneins and 28 March 2014 at Marmande. The river width is  
174 around 130 m at Verdun-sur-Garonne, 150 m at Lamagistère and 200 m at Tonneins and Marmande.  
175 Also, the 15 km reach of the Garonne just upstream of Lamagistère (from the upstream confluence,  
176 at Malause, to Lamagistère) has multiple man made hydraulic infrastructures along the river. There  
177 are five weirs within the reach and a "run-of-the-river" dam at Malause, which induce river slope  
178 breaks. Thus, water elevation variations within this reach and, in particular, at the location of ENV-  
179 773 virtual station (see Figure 1 for its location and section 2.2.4 for definition of virtual station),  
180 might not be comparable to water elevation variations at Lamagistère gage. There is no similar  
181 infrastructure (and no such slope break) near other in situ gages that might impact comparison with  
182 altimeter measurements.



183           Moreover, a Digital Elevation Model (DEM) at 25 m horizontal resolution, with few meters  
184 absolute vertical accuracy, is available over the entire mainland France to the research community  
185 and provided by the Institut National de l'Information Géographique et Forestière (IGN,  
186 <http://professionnels.ign.fr>), a French government agency which is “the official reference for  
187 geographic and forest information in France” (from <http://www.ign.fr/institut/en>).  
188 Time series for three tide gages are also available on the Gironde estuary, seaward of the Garonne  
189 (Figure 1). They are operated by the Service Hydrographique et Océanique de la Marine (SHOM –  
190 <http://refmar.shom.fr>) for the Royan gage (from 19 September 2008 to 31 August 2014) and by the  
191 Grand Port Maritime de Bordeaux for Port-Bloc (from 1 January 2006 to 12 October 2014) and  
192 Richard (from 1 January 2006 to 31 December 2014) gages, with water level measurements every  
193 minute.

194           All elevations (from gages time series and DEM) are referenced to the “Nouveau  
195 Géoréféntiel Français” (NGF-IGN69), the official French vertical reference for the main territory.  
196 Because all these data are available, the Garonne River basin is particularly well suited to evaluate  
197 the capability of nadir altimeters to observe a medium river between 100 m and 200 m wide.

198

## 199 **2.2. Satellite altimetry missions used**

### 200 **2.2.1 Principle of altimetry measurement**

201           The purpose of radar altimeters is to provide the height of the ground surface above a reference  
202 ellipsoid. To do so, the altimeter emits a radar pulse and records the radar echo using a pulse  
203 compression technique. This record, also known as a waveform, contains the value of the returned  
204 power as a function of the distance between the radar and the reflectors. In this study, the term  
205 “range” is equivalent to the distance from the instrument. For technical reasons, the altimeter does  
206 not record all the power backscattered by all targets between the instrument and the lowest ground  
207 elevation within the instrument footprint (all the possible ranges). It only samples a small subset of  
208 these ranges, called the range window or tracking window, whose size is typically between 30 m

209 and 50 or 64 m depending on the instrument, but can reach 1024 m for Envisat in the 20 MHz  
 210 mode. For more details, see Benveniste et al. (2001), Desjonquères et al. (2010) and Steunou et al.  
 211 (2015). A special function of the altimeter is to keep the range window tracking the ground surface  
 212 (see section 2.2.2 for more information related to the on-board tracking system).

213 The two-way travel-time from the satellite to the surface is the measurement that needs to be  
 214 estimated as accurately as possible from the waveform. It corresponds to an instant known, in the  
 215 waveform, as the middle of the leading edge over the ocean. Over other types of surface, this time is  
 216 more complex to retrieve and depends of the retracking algorithm used. It is accurately determined  
 217 by the mission ground segment using retracking algorithms and is used to compute the distance  
 218 between the satellite and the Earth surface, the altimeter range ( $R$ ). Then, the satellite altitude ( $H$ )  
 219 referenced to an ellipsoid is computed from orbit modeling, with an accuracy better than 2 cm (e.g.  
 220 Cerri et al., 2010; Couhert et al., 2015; Dettmering et al., 2015).

221 Taking into account propagation corrections caused by delays from the interactions of  
 222 electromagnetic waves in the atmosphere, and geophysical corrections, the height of the reflecting  
 223 surface ( $h$ ) with reference to an ellipsoid can be estimated as:

$$224 \quad h = H - \left( R + \sum(\Delta R_{propagation} + \Delta R_{geophysical}) \right) \quad (1)$$

225 where  $H$  is the satellite centre of mass height above the ellipsoid, estimated using the precise orbit  
 226 determination (POD) technique,  $R$  is the nadir altimeter range from the satellite center of mass to  
 227 the surface taking into account instrument corrections,  $\sum \Delta R_{propagation}$  is the sum of the  
 228 geophysical and environmental corrections applied to the range, respectively, and  $\sum \Delta R_{geophysical}$  is  
 229 another geophysical correction. Furthermore,  $\sum \Delta R_{propagation}$  is computed as follow:

$$230 \quad \sum \Delta R_{propagation} = \Delta R_{ion} + \Delta R_{dry} + \Delta R_{wet} \quad (2)$$

231 where  $\Delta R_{ion}$  is the atmospheric refraction range delay due to the free electron content associated  
 232 with the dielectric properties of the ionosphere,  $\Delta R_{dry}$  is the atmospheric refraction range delay due  
 233 to the dry gas component of the troposphere,  $\Delta R_{wet}$  is the atmospheric refraction range delay due to  
 234 the water vapor and the cloud liquid water content of the troposphere. Also,  $\sum \Delta R_{geophysical}$

235 corresponds to the following corrections:

$$236 \quad \sum \Delta R_{geophysical} = \Delta R_{solid\ Earth} + \Delta R_{pole} \quad (3)$$

237 where  $\Delta R_{solid\ Earth}$  and  $\Delta R_{pole}$  are the corrections accounting for crustal vertical motions due to the  
238 solid Earth and pole tides, respectively. Over the ocean, other corrections need to be applied to take  
239 into account other physical processes (such as ocean tides, see Chelton et al., 2001, for more  
240 information).

241

### 242 **2.2.2 On board tracking system**

243 As indicated in the previous section, one important function of the altimeter is to modify the  
244 position of its tracking window to make it follow the ground topography, which can rapidly change  
245 over few kilometers on the continents. This is automatically performed on board in “closed-loop”  
246 by the Adaptive Tracking Unit (ATU) from previously received waveforms. Chelton et al. (2001)  
247 and Desjonquères et al. (2010) provide a detailed description of the closed-loop tracking system for  
248 TOPEX and Poseidon-3 altimeters, respectively. The following paragraph provides only a  
249 simplified overview, which is sufficient enough to understand the observations presented in this  
250 study.

251 The principle of the closed loop is that the ATU tries to keep some signal in its tracking  
252 window. On Poseidon-3, this is done by using the so-called “median mode” (Desjonquères, 2010),  
253 which tries constantly to center the signal in the window. If this fails, the level of received signal  
254 decreases dramatically. When the level of received signal becomes lower than a predefined  
255 threshold, the ATU considers that the tracking is lost and switches to a “search” mode in which it  
256 scans a window, with range of a few kilometers range, centered on the estimated satellite altitude.  
257 The scan begins with the smallest range (i.e. closest to the satellite) and the tracking window is  
258 moved until the level of received signal exceeds again another specified threshold (Desjonquères et  
259 al., 2010).

260 This behavior implies that the ground surface observed by nadir altimeters heavily depends

261 on the previously received radar echoes. For example, the geometry of the observations can induce  
262 loss of radar echoes tracking in some circumstances (e.g., if the satellite trajectory crosses a steep-  
263 sided valley perpendicularly) or not, (e.g., if it follows the valley over a long distance). In general,  
264 when the tracking is lost and the ATU is in search mode, the signal received from the top of the hills  
265 is high enough to exceed the threshold making the ATU to stop the search. This can occur before the  
266 tracking window reaches the river that flows in the valley, but of course this depends on the depth  
267 of the valley as well as the size of the tracking window. So the top of the hilly areas often tends to  
268 be observed rather than rivers in the valley. However the exact behavior of the altimeter depends on  
269 the ground reflectivity, the size of the tracking window and the two thresholds mentioned above.  
270 Thus, whenever there are variations of topography, there is no way to control which part of the  
271 scenery will be observed by the radar altimeter in the closed-loop tracking mode.

272 To overcome this challenge, a new tracking mode, the Doris Immediate Orbit on board  
273 DEtermination/Digital Elevation Model (DIODE/DEM) mode, has been implemented on board  
274 Poseidon-3 and AltiKa altimeters. In this mode, tracking range is not estimated in closed-loop, but  
275 in “open-loop”. In this case, the satellite/ground range is not estimated automatically from formerly  
276 measured signal, but using a DEM stored on board and an estimate of the satellite orbital position,  
277 computed on board and in real time by the DIODE navigator function of the DORIS (Doppler  
278 Orbitography and Radio-positioning Integrated by Satellite) instrument (Desjonquères et al., 2010).  
279 The DEM mode was activated during cycles 3, 5, 7, 34, 209 and 220 for Jason-2 and only during a  
280 portion of SARAL cycle 1 (tracks 600 to 800 from 4 April to 10 April 2013), corresponding to  
281 tracks 646 and 773 over the Garonne basin (see Figure 1 for their location). To compute on board  
282 elevations used by the DIODE/DEM tracking mode, CNES used an a priori global DEM and a  
283 global land/water classification (Desjonquères, 2009). If there is water in the land/water  
284 classification within the instrument footprint, then only a priori DEM elevations within the water  
285 mask are used to compute on board elevation (Desjonquères, 2009). However, if there is no water,  
286 then all land elevations are considered. Therefore, for steep-sided regions with no water in the

287 classification (or if the water mask is not correctly geolocalized), the computed on board elevations  
288 can be closer to the top of the hills elevations than the river valley elevations. Furthermore, the  
289 waveform is expected to be centered on the first third of the tracking window. As this window size  
290 is around 50 m for Jason-2 and 30 m for SARAL, it has been estimated that the a priori DEM  
291 accuracy should be around 10 m (even slightly less for SARAL).

292 For Jason-2, the a priori global DEM used is the 1 km Altimetry Corrected Elevation DEM  
293 (ACE; Berry et al., 2000) and the water mask comes from the Generic Mapping Tools (GMT,  
294 <http://gmt.soest.hawaii.edu/>) (Desjonquères, 2009). A comparison between ACE DEM and 25 m  
295 IGN DEM (see section 2.1) tends to show that ACE accuracy over the Garonne valley is better than  
296 10 m. Global-scale ACE uses local Digital Terrain Elevation Data (DTED) and altimetry data from  
297 the ERS-1 (European Remote Sensing-1 satellite) acquired during its geodetic mission. However,  
298 over the Garonne River valley, ACE elevations come only from DTED (exact source is not provided  
299 in the ACE documentation). In addition, the river position in the GMT database is not correctly  
300 geolocalized. Therefore, the on board elevations over the Garonne might be biased toward  
301 elevations on top of the hills. For SARAL, the a priori global DEM corresponds to ACE2 (Berry et  
302 al., 2010) and the land/water mask is derived from Globcover  
303 ([http://due.esrin.esa.int/page\\_globcover.php](http://due.esrin.esa.int/page_globcover.php)). The accuracy of ACE2 also seems to be better than 10  
304 m over the Garonne valley. Globcover correctly geolocalizes the Garonne River but because of its  
305 300 m pixel size and the undersampling to 1 km pixels in the CNES tool, pixels identified as water  
306 do not always correspond to the river surface. These discrepancies can impact the computed  
307 elevations stored on board.

308 Birkett and Beckley (2010) evaluated both the closed-loop and the DIODE/DEM modes (for  
309 cycles 3, 5, 7 and 34) for Jason-2 over 28 lakes and reservoirs around the world with areas spanning  
310 from 380,000 km<sup>2</sup> (Caspian Sea, Kazakhstan) to 150 km<sup>2</sup> (Windsor lake, Bahamas). They  
311 concluded that both modes on Jason-2 are able to monitor water bodies with area around 150 km<sup>2</sup>  
312 and width around 0.8 km. This monitoring capability is an improvement compared to

313 Topex/Poseidon and Jason-1, with the DIODE/DEM having the fastest acquisition time for many  
314 targets. However, for few targets (Chajih lake, 900 km<sup>2</sup>, Windsor lake, 150 km<sup>2</sup>, Brokopondo  
315 reservoir 1,500 km<sup>2</sup> and Powell and Diefenbaker reservoir systems, 500 km<sup>2</sup> and 550 km<sup>2</sup>,  
316 respectively), they noted some loss of data in the DIODE/DEM mode for cycles 3, 5 and 7, whereas  
317 the closed-loop mode was performing well during other cycles. These targets are better (at least  
318 partially) observed during cycle 34, after the on board DEM elevations have been updated and some  
319 altimeter parameters have been tuned in cycle 16. They attributed errors to “inadequate resolution  
320 and/or data in the DEM” and concluded that the on board DEM might not be “optimized for all  
321 regions”. For all investigated lakes and reservoirs, they observed some cases where the closed-loop  
322 successfully observed the water body, contrarily to the DIODE/DEM. For all other cases, both  
323 tracking modes provided similar results (both failed or succeeded to observe water bodies). But  
324 there was no case where the DIODE/DEM mode observed the target and the closed-loop did not.  
325 However, Birkett and Beckley (2010) considered in their study targets that were larger than the  
326 Garonne River.

327

### 328 **2.2.3 Altimetry datasets**

329 This study uses altimetry missions only after 2002, namely the ENVironmental SATellite  
330 (ENVISAT) mission from the European Space Agency (ESA), Jason-2 mission from National  
331 Aeronautics and Space Administration (NASA) and Centre National d'Etudes Spatiales (CNES),  
332 and Satellite for ARGos and ALtika (SARAL)/Altimeter in Ka-Band (AlitKa) mission jointly  
333 developed by the Indian Space Research Organization (ISRO) and CNES. Table 1 sums up the main  
334 characteristics and technical details of these three altimetry missions, which are described in more  
335 detail below.

336 ENVISAT mission was launched by ESA on 01 March 2002. It carried 10 instruments  
337 including the advanced radar altimeter (RA-2). It was based on the heritage of sensors on board the  
338 European Remote Sensing (ERS-1 and 2) satellites. Altimeter RA-2 was a nadir-looking pulse-

339 limited radar altimeter operating at two frequencies: Ku- (13.575 GHz), as ERS-1 and 2, and S- (3.2  
340 GHz) bands. Its goal was to collect radar altimetry over ocean, land and ice caps (Zelli, 1999).  
341 Altimeter RA-2 could change the range resolution to set the range detection window to three sizes:  
342 1024 m, 256 m and 64 m (they are also commonly designated as 20MHz, 80MHz and 320MHz  
343 modes, respectively, corresponding to the bandwidths used to achieve the corresponding window  
344 width; Benveniste et al., 2001; ESA, 2007). Changing the tracking window is particularly useful  
345 over continents to adapt tracking to ground topography changes (regions with rapidly varying  
346 altitude will be better tracked with a relatively wider window, whereas flat regions will be more  
347 precisely observed with a relatively narrower window). The range window is sampled using a fixed  
348 number of bins (or gates). A bin corresponds to a continuous interval of ranges that will be  
349 aggregated to a unique range value, during the analog-to-digital conversion step. Changes in range  
350 window size are done automatically on board, based on the received signal and reference data  
351 stored in the on board memory (ESA, 2007). ENVISAT orbited at an average altitude of 790 km,  
352 with an inclination of 98.54°, on a sun-synchronous orbit with a 35-day repeat cycle. It provided  
353 observations of the Earth surface (oceans and lands) from 82.4° latitude North to 82.4° latitude  
354 South. This orbit was formerly used by ERS-1 and 2 missions, with an equatorial ground-track  
355 spacing of about 80 km. ENVISAT remained on this nominal orbit until October 2010 and its  
356 mission ended 08 April 2012. This study used ENVISAT data from cycles 6 (which started 14 May  
357 2002) to 94 (which ended 21 October 2010).

358 Jason-2 mission was launched on 20 June 2008 as a cooperation between CNES,  
359 EUMETSAT, NASA and NOAA. Its payload is mostly composed of the Poseidon-3 radar altimeter  
360 from CNES, the Advanced Microwave Radiometer (AMR) from JPL/NASA, and a triple system for  
361 precise orbit determination: the DORIS instrument from CNES, a GNSS receiver and a Laser  
362 Retroreflector Array (LRA) from NASA. Jason-2 orbits at an altitude of 1336 km, with an inclination  
363 of 66°, on a 10-day repeat cycle, providing observations of the Earth surface (oceans and lands)  
364 from 66° latitude North to 66° latitude South, with an equatorial ground-track spacing of about 315

365 km. This orbit was formerly used by Topex/Poseidon, and Jason-1. Poseidon-3 radar altimeter is a  
366 two-frequency altimeter, operating at Ku- (13.575 GHz) and C- (5.3 GHz) bands (Desjonquieres et  
367 al., 2010). The tracker range window is the same as previous Poseidon instruments and has a useful  
368 size around 50m (sampled over 104 range bins). Jason-2 measurements are used in this study from  
369 cycle 1 (which started 12 July 2008) to cycle 227 (which ended 09 September 2014) h. Jason-2 raw  
370 data are processed by SSALTO (Segment Sol multimissions d'ALTimétrie, d'Orbitographie).

371 SARAL is a CNES-ISRO joint-mission that was launched on 25 February 2013. Its payload  
372 is composed of the AltiKa radar altimeter and bi-frequency radiometer, and a double system for  
373 precise orbit determination (Steunou et al., 2015): DORIS instrument and a Laser Retroreflector  
374 Array (LRA). SARAL flights on the same nominal orbit than ENVISAT (see above). AltiKa radar  
375 altimeter is a mono-frequency altimeter and the first one to operate at Ka-band (35.75 GHz). The  
376 bandwidth of the signal has been increased (480 MHz of useful bandwidth; Sengenés and Steunou,  
377 2011) with respect to Jason-2 and ENVISAT, improving the range resolution. As the number of  
378 useful bins remains the same as for the Jason altimeters (104 bins), the tracker window size is  
379 around 30 m. For electromagnetic wave in Ka-band, ionospheric delay becomes negligible. In this  
380 study, SARAL/AltiKa measurements from cycle 1 (which started 14 March 2013) to cycle 17  
381 (which ended 30 October 2014) have been used.

382 Altimetry data processed in this study come from the Geophysical Data Records (GDRs) –  
383 GDR T patch 2 for SARAL, GDR v2.1 for ENVISAT, GDR D for Jason-2, delivered by  
384 CNES/ESA/NASA processing centers. Since this study has been performed, a new Jason-2 GDR  
385 version has been released in May 2015 (GDR E). Differences between GDR D and E are expected  
386 to have a low impact on the results presented here, as the foreseen improvement is a better  
387 agreement of the geographically correlated radial orbit drift rate (1 mm/year to less than 0.5  
388 mm/year over roughly 6 years) with respect to JPL (RLSE14A) GPS-only reduced-dynamic orbits  
389 for Jason-2 (for more details, see  
390 [http://www.aviso.altimetry.fr/fileadmin/documents/data/tools/New\\_GDR\\_E\\_orbit\\_20150521.pdf](http://www.aviso.altimetry.fr/fileadmin/documents/data/tools/New_GDR_E_orbit_20150521.pdf)).



391 Similarly, a new SARAL GRD E is now available. But SARAL GDR T patch 2 and SARAL GDR  
392 E must provide similar results for this study. These data are provided in a consistent NetCDF  
393 (Network Common Data Form) format with coherent geophysical corrections for all missions by  
394 Centre de Topographie de l'Océan et de l'Hydrosphère (CTOH –<http://ctoh.legos.obs-mip.fr/>). They  
395 are sampled along the altimeter track at 18, 20 and 40 Hz for ENVISAT, Jason-2 and SARAL  
396 respectively. As explained in the previous section, a so-called retracker algorithm is needed to  
397 estimate the satellite/ground range  $R$  (Eq. 1) and the surface backscattering coefficient from the  
398 received electromagnetic signal. Previous studies showed that Ice-1 retracking algorithm (Wingham  
399 et al., 1986) is one of the most suitable for hydrological studies, in terms of accuracy of water levels  
400 and availability of the data (e.g. Frappart et al., 2006a; Santos da Silva et al., 2010), among the  
401 commonly available retracked data present in the GDRs. However, Santos da Silva et al. (2010)  
402 found, on the Amazon basin, that Ice-2 retracking algorithm (Legrésy and Rémy, 1997) could  
403 provide similar results to Ice-1. Sulistioadi et al. (2015) showed, for some 250 m wide Indonesian  
404 river reaches, that Sea Ice retracking algorithm (Laxon, 1994) can provide, in some cases, slightly  
405 more accurate water levels than Ice-1. In the following, when not explicitly indicated, ranges used  
406 to derive altimeter heights and backscattering coefficients are those derived from the Ice-1  
407 retracking algorithm. ENVISAT GDRs directly provide ranges estimated using Ice-1 and Sea Ice  
408 algorithms. Therefore, in section 3.1, a comparison is presented between Ice-1 and Sea Ice derived  
409 water heights for ENVISAT measurements. Ice-2 retracker has not been considered as it provides  
410 similar results to Ice-1 (Santos da Silva et al., 2010). Similarly, the retracker used over ocean  
411 (Brown, 1977) has not been used, as it has been widely shown that it provides the worst results  
412 compared to the three other retracking algorithms for rivers (e. g. Frappart et al., 2006a; Santos da  
413 Silva et al., 2010; Sulistioadi et al., 2015).

414 Over the ocean, wet troposphere corrections are computed from the on board radiometer  
415 measurements, for each altimetry mission. However, measurements over continents from these  
416 radiometers cannot be used to estimate those corrections, as the ground emissivity (contrarily to

417 water) is much more important than the emissivity from the atmosphere. In this case, the  
418 propagation corrections applied to the range are derived from the Era Interim model outputs by the  
419 European Centre Medium-Range Weather Forecasts (ECMWF) for the dry/wet troposphere range  
420 delays. The range correction accounting for ionosphere delays is estimated using the Global  
421 Ionospheric Maps (GIM).

422

#### 423 **2.2.4 Time series of altimetry-based water levels**

424 Time variations of river levels from radar altimetry measurements are computed at virtual  
425 stations. A virtual station is defined as the intersection between an orbit ground track and a water  
426 body (*i.e.*, lake, reservoir river channel, floodplain or wetland). At these specific locations,  
427 variations from one cycle to the next of height  $h$ , derived from altimeter measurements (see Eq. 1),  
428 can be associated to changes in water level.

429 In this study, we used the Multi-mission Altimetry Processing Software (MAPS) that allows a  
430 refined selection of the valid altimetry data to build virtual stations (Frappart et al., 2015b). Data  
431 processing is composed of three main steps: a coarse delineation of the virtual stations using Google  
432 Earth, a refined selection of the valid altimetry data, and a computation of the water level time-  
433 series. For virtual stations on the Garonne mainstream (Figure 1), the length of the selection is not  
434 constant and varies from 700 m to 2 km. The altimetry-based water level is computed for each cycle  
435 using the median of the selected altimetry heights, along with their respective deviation (*i.e.*, mean  
436 absolute deviations). This process is repeated each cycle to construct the water level time series at  
437 the virtual stations.

438

### 439 **3. Results**

#### 440 **3.1. Multi-satellite water elevation on the Garonne River mainstream**

441 The altimetry-based time series of water elevation at virtual stations shown on Figure 1 have  
442 been compared to the closest in situ station available in the Banque Hydro database (see section

443 2.1), with the exception of virtual stations JA2-070a and JA2-070b. They have not been used in this  
444 study, as the confluence with the Lot River (one of the main Garonne River tributaries) is located  
445 between the closest in situ gage (Tonneins) and the virtual stations. Therefore water elevation at the  
446 gage might not be representative of water elevation at the virtual stations.

447 For other virtual stations, only common dates have been used for the comparison between  
448 altimetry and in situ time series. In situ measurements recorded the same day as the altimetry  
449 measurement are linearly interpolated at the altimetry measurement observation time. If there is no  
450 in situ measurement the same day as the altimetry observation, this observation is not considered. In  
451 situ time series have been referenced to UTC (Coordinated Universal Time) to match the altimeter  
452 time reference. Elevation anomaly time series have been computed for both the altimeters and in  
453 situ gages. The anomalies are computed by removing, from the elevation time series, its temporal  
454 mean over the same common dates between in situ and altimetry time series (e.g. Biancamaria et  
455 al., 2011).

456 Table 2 shows the correlation coefficient, mean bias (mean of the difference) and Root Mean  
457 Square Error (RMSE) between altimetry and in situ time series for both absolute water elevations  
458 referenced to NGF-IGN69 and water elevation anomalies, along the Garonne River mainstream. For  
459 anomaly time series, the Nash-Sutcliffe (NS) coefficient (Nash and Sutcliffe, 1970) is also  
460 computed. The NS coefficient, which ranges between  $-\infty$  and 1, is widely used to assess how an  
461 estimated time series (most of the time from a model) accurately match (i.e., in time and amplitude)  
462 in situ measurements. The closer to 1 the NS coefficient is, the closest to the in situ time series the  
463 altimetry-based water elevations are. NS above 0.5 can be considered satisfactory (Moriassi et al.,  
464 2007). However, a negative NS means that the estimated time series is a worse “predictor” than the  
465 in situ time series mean and should be considered as unacceptable (Moriassi et al., 2007). In this  
466 study, the NS is not computed for absolute water elevation (bias between absolute in situ and  
467 altimetry water elevation induces negative NS), but for their anomalies. Table 2 also provides  
468 distance between the altimetry virtual station and the gage, number of common dates between

469 altimetry and in situ time series, and the amplitude (maximum minus minimum over the common  
470 dates) of the in situ time series.

471 At Lamagistère and Verdun-sur-Garonne, ENVISAT data are not really correlated to the in  
472 situ measurement and the NS coefficients are negative, indicating poor performance of the  
473 altimeters. For virtual stations ENV-102 and ENV-773, which correspond to the worst results for  
474 ENVISAT, the distance to the in situ gage can partially explain the mismatch. However, the small  
475 river width (~150 m) and the surrounding topography affecting the quality of the altimetry signal  
476 are also likely to be an important source of error. Virtual station ENV-646, which is only 1 km  
477 downstream the gage at Lamagistère, has better RMSE (1.80 m for absolute water elevations and  
478 0.80 for water elevation anomalies), correlation coefficient (0.61) and NS coefficient (-0.53)  
479 compared to upstream ENVISAT virtual stations, even if they cannot be considered as satisfactory.

480 Downstream, at Tonneins and Marmande, where the river width is around 200 m, ENVISAT  
481 altimetry time series are of good quality with correlation coefficient around 0.8, NS around 0.7 and  
482 water elevation anomalies RMSE between 0.5 m and 0.6 m for ENVISAT. For ENVISAT and  
483 Jason-2, the mean bias must be mostly explained by the river slope between the gage and the  
484 altimetry virtual station, as they have the same order of magnitude as the slopes computed from  
485 IGN DEM, even if the DEM vertical accuracy (few meters) prevents a quantitative estimate of the  
486 river slope (which also varies in time). The sign of the bias depends of the position of the virtual  
487 station compared to the gage (positive if the virtual station is downstream and negative if it is  
488 upstream). Yet, some part of this bias might also be related to the altimeter measurement error.

489 Results shown in Table 2 have been computed using Ice-1 retracker algorithm and the  
490 median value of altimetry heights for each observation time (see section 2.2.4). Even if some  
491 studies reported better results using Ice-1 retracker (over the Amazon see Frappart et al., 2005;  
492 Frappart et al., 2006a; Santos da Silva et al., 2010), Sulistioadi et al. (2015) found that Ice-1 was not  
493 always providing the best results for some Indonesian rivers, whose widths were around 250 m. In  
494 this study, the Sea Ice retracker was sometimes performing better than Ice-1. As ENVISAT GDRs

495 provide ranges computed using at least these two retracker algorithms, we computed water  
496 elevation from both retrackers and compared them to in-situ measurements for ENVISAT virtual  
497 stations along the Garonne mainstream. These results are shown in Table 3. Within each virtual  
498 station, the median water elevation is computed for each observation time, as it is more robust than  
499 the mean, when there are outliers (Frappart et al., 2005; Frappart et al., 2006a). However, Santos da  
500 Silva et al. (2010) stated that computing both the median and the mean can provide a “qualitative  
501 indicator of the presence of outliers”. That is why both the median and the mean are shown in Table  
502 3. This table seems to confirm the results from Sulistioadi et al. (2015), the best results are not  
503 always obtained with Ice-1. For the three virtual stations with NS coefficient below 0, two have  
504 better results with Sea Ice. For the two virtual stations with NS coefficient above 0.5, one has better  
505 results with Ice-1, the other with Sea Ice. However, for these two virtual stations, difference  
506 between RMSE for these two retrackers is just a few centimeters, which is small compared the  
507 actual value of the RMSE (more than 50 cm). Therefore, both retrackers are well suited for the  
508 Garonne basin. Results using the median are better, most of the time, than results using the mean  
509 (except for Ice-1 and virtual station ENV-315, where the RMSE using the mean is only 2 cm lower  
510 than the RMSE using the median). From these results it seems that both the median Ice-1 and the  
511 median Sea Ice are well suited for the Garonne basin. Differences between these two retrackers are  
512 one order of magnitude lower than the RMSE obtained from comparison to in situ time series.

513 Figure 2 shows enlargements from Figure 1 on ENVISAT and Jason-2 virtual stations at  
514 Lamagistère (a.), at Tonneins (b.) and in the estuary (c.). Especially, it should be recalled that there  
515 are four weirs between ENV-773 virtual station and Lamagistère gage (as explained in section 2.1),  
516 which are 10 km apart. These weirs cause slope breaks and can explain at least a part of the 1.55 m  
517 water elevation anomalies RMSE for this virtual station. Figure 3 shows ENVISAT (red curves for  
518 Ice-1 and cyan curves for Sea Ice) and in situ (blue curves) water elevation anomaly time series at  
519 the satellite measurement times for virtual stations ENV-646 (Fig 3.a) and ENV-188 (Fi. 3.c). On  
520 this figure, the right panels (b. and d.) show all records in the in situ water elevation anomalies time

521 series (blue dots) and the altimetry water elevation anomaly measurements (red dots for Ice-1 and  
522 cyan dots for Sea Ice) during the common time period for these two virtual stations. These right  
523 panels highlight the coarse altimetry time sampling. On Figure 3.d, ENVISAT seems to roughly  
524 sample the water elevation seasonal cycle, but, because of the 35 days repeat orbit, it cannot  
525 observe intra-monthly variability. This variability can be quite important for a medium river like the  
526 Garonne, for which precipitation and snow melting induce few meters water elevation variations  
527 within few days at Tonneins. Figure 3 corresponds to two virtual stations (ENV-646 and ENV-188)  
528 for which Sea Ice retracker is performing better than Ice-1 retracker, yet the two retracker's time  
529 series remain close. Table 3 shows that for two other virtual stations (ENV-773 and ENV-315) Ice-1  
530 performs better than Sea Ice. This result is in agreement with the results obtained by Sulistioadi et  
531 al. (2015): for medium size rivers, Ice-1 is not always the best retracker. However, water elevation  
532 obtained from both retracker's are close enough and both could be used (there is just few centimeters  
533 difference between them for the two virtual stations, which have a correlation coefficient above  
534 0.8). According to these results and as Sea Ice retracker is not provided in Jason-2 GDRs, only  
535 results using Ice-1 retracker will be shown in the following.

536 For Jason-2 virtual station JA2-070, in between Tonneins and Marmande, correlation  
537 coefficient is equal to 0.98, NS around 0.95 and the RMSE of water elevation anomalies is close to  
538 20 cm. Results for virtual stations JA2-035 show slightly lower agreement with a correlation  
539 coefficient of 0.91 and water elevation anomalies RMSE and NS coefficient of 0.36 cm and 0.82,  
540 respectively. The most noticeable feature in this virtual station is the few dates (62) that measure  
541 river water commonly with the Marmande gage time series (Table 2). In comparison, virtual station  
542 JA2-070 has 150 dates with measurements of river water elevation during the same period. For the  
543 other dates, elevations are 50 m higher than valid measurements of river water elevations and have  
544 therefore been removed during the virtual station time series generation before comparison with in  
545 situ data. These dates (around 40 for JA2-070 and 140 for JA2-035) correspond to cases when the  
546 altimeter remains 'locked' on the surrounding hills (see section 2.2.2 for an explanation of this

547 phenomenon). The Garonne valley is roughly 5 km or less wide at these locations and is surrounded  
548 by hilly areas (see Figure 2) that can be 50 m to 100 m higher than the valley (according to the IGN  
549 DEM and knowledge of the region). These two virtual stations also illustrate the importance of the  
550 geometry of observation. The track 070 is almost parallel to the valley over a long distance (almost  
551 30 km). Therefore, distance variations between the ground and the radar are much smoother  
552 compared to the track 035 that crosses the valley almost orthogonally.

553 Figure 4 presents similar plots than Figure 3, but for the Jason-2 virtual station JA2-070,  
554 using Ice-1 retracker only (red curves and red dots). This virtual station clearly shows better results  
555 than ENVISAT virtual station ENV-188 (Figure 3.c and 3.d), when compared to Tonneins in situ  
556 time series. Besides, with a 10 days repeat orbit, Jason-2 observes higher frequency variations, but  
557 still misses all the local maxima and especially the 2009 and 2014 heavy floods, which lasted only  
558 few days.

559 Table 2 also highlights high mean bias for most SARAL virtual stations, only virtual station  
560 SRL-188 have correlation and errors similar to ENVISAT. For the three other ones, the mean error  
561 goes from 44 m to 105 m, with few dates in the time series, indicating that the altimeter is not  
562 observing the river valley but the surrounding hills. This problem is similar to that already observed  
563 for Jason-2 time series. However, for these SARAL virtual stations and contrary to Jason-2 virtual  
564 stations, there is no measurement on the river. ENVISAT is less affected by such effects, thanks to  
565 its three resolutions (see section 2.2.3) and differences in the closed-loop parameters. This  
566 drawback and potential reasons for the differences between the three missions is discussed in more  
567 detail in section 3.2.

568 In the Gironde estuary, at Richard tide gage (see Figure 1 for its location), both ENVISAT  
569 and SARAL tracks 274 compare unfavorably to in situ measurements (Table 4) with correlation  
570 coefficients of 0.28 and 0.09, respectively. As the absolute vertical reference for this tide gage is not  
571 known, mean bias and the absolute elevation RMSE cannot be computed. The RMSE of water  
572 elevation anomalies is around 1.5 m. Differences between altimetry and in situ time series could be

573 related to instrument error, impact of surrounding lands and the fact that water elevation variations  
574 at the tide gage might not be representative of water elevation variations along the satellite ground  
575 track. Figure 5, shows the measured elevation from ENVISAT/RA2 for track 274 during 21 June  
576 2007 (red line) and the IGN DEM elevation (green curve) on the estuary. It shows that over half of  
577 the estuary, the altimeter remains locked over the surrounding topography (which is a common  
578 issue for nadir altimetry due to the closed-loop tracking mode, as explained in section 2.2.2). These  
579 measurements are not taken into account to compute time series for ENV-274, but represent a  
580 source of error that is likely to affect the altimetry signal in the lower estuary.

581 Results are much better for both ENVISAT and SARAL tracks 859 (Table 4) at Port-Bloc  
582 and Royan tide gages (see Figure 1 and 2.c for their locations). The correlation coefficient is above  
583 0.97, water elevation anomalies RMSE are around 30 cm for ENVISAT and 10 cm for SARAL. The  
584 comparison between anomalies time series measured by Port-Bloc tide gage and ENVISAT track  
585 859 is shown on Figure 6.a and 6.b and SARAL track 859 on Figure 6.c and 6.d. Figures 6.b. and  
586 6.d. also highlight the well-known effect of tidal aliasing in the altimeter water elevation time  
587 series, due to the altimetry satellite orbit repeat period which is much higher than the tides period (e.  
588 g. Le Provost, 2001). As stated by Le Provost (2001), the important difference between altimeter  
589 time sampling (10 days or 35 days) and semidiurnal and diurnal tides period (between 12 to 24  
590 hours), leads to alias these tides “into periods of several months to years”. But the issue of tidal  
591 aliasing is beyond the scope of the present paper.

592

### 593 **3.2. Tracking issue and possible solution**

594 The challenge of observing the Garonne valley for some altimetry mission (SARAL/AltiKa,  
595 but also Jason-2 and ENVISAT to a lower extent), highlighted in section 3.1, has multiple origins,  
596 as explained in section 2.2.2. The Garonne valley is only 5 km wide at virtual stations ENV-/SRL-  
597 773 and 50 m to 100 m lower than the surrounding hilly areas. Due to the closed-loop tracking  
598 algorithm, nadir altimeter tends to get locked on the top of the hilly areas and miss steep-sided



599 valleys (see section 2.2.2). The portion of SARAL/AltiKa ground track that crosses the Garonne  
600 valley (roughly perpendicularly, see Figure 2.a) is only 7 km long. As the antenna footprint on the  
601 ground is equal to 4 km, considering the 3-dB aperture angle of  $0.6^\circ$  for SARAL that defines the  
602 half-power points of the antenna radiation pattern (Steunou et al., 2015), the instrument still  
603 receives some backscattered energy from the surrounding hilly areas for previous radar echoes  
604 when it is near and over the valley. Therefore, the ATU does not change the position of the tracking  
605 window, which remains locked on the hills. After few kilometers, the hills are not in the antenna  
606 footprint anymore and no more energy is received, resulting in a loss of measurements. By the time  
607 the instrument changes the tracking window position (search mode, see section 2.2.2) and receives  
608 again some energy, acquisition of the Garonne valley has been lost. Figure 7 shows an example of  
609 data loss by AltiKa due to the closed-loop. On this figure, the x-axis represents the latitude along  
610 the SARAL track 773 (see Figure 2.a for its location) and the y-axis corresponds to elevation  
611 referenced to NGF-IGN69. The green curve corresponds to the IGN DEM. The Garonne River  
612 location is indicated by the blue rectangle. SARAL measurement for cycle 2 (in closed-loop) is  
613 shown by the magenta dashed line. During cycle 2, AltiKa remains locked over the hills and loses  
614 tracking over the Garonne valley, as previously explained.

615         Signal-locking over hills is less frequent for Jason-2 virtual station JA2-070, because of a  
616 more favorable observation configuration than SARAL. Jason-2 track flies over the Garonne valley  
617 and follows the river over 30 km before virtual station JA2-070. Therefore, elevation variations  
618 observed by the satellite are smoother than SARAL. Smoother variations along Jason-2 track allow  
619 more time for the closed-loop tracking algorithm to adapt to the hills/valley transition, whereas  
620 ENVISAT and SARAL track 188 is almost perpendicular to the valley (Figure 2.b). Jason-2 virtual  
621 station JA2-035 has a configuration of observation close to SARAL track 188, that is why it also  
622 has few river water elevation measurements.

623         However, ENVISAT better performance compared to SARAL/AltiKa is not due to a different  
624 observation configuration (contrary to Jason-2, as SARAL is on the same orbit), but it must be

625 related to the three window sizes that are chosen automatically on board (64 m, 256 m, 1024 m, see  
626 section 2.2.3). ENVISAT is better suited to observe ground with appreciable slope variations as the  
627 instrument increases the size of its tracking window, which allows measurements in the river valley.  
628 Table 5 shows percentage of data for all ENVISAT cycles acquired with each window size, for all  
629 virtual stations. Measurements for virtual stations close to high relief have more tracking window  
630 size variability (typically the case of ENV-102, ENV-773 and ENV-315). At ENV-188 virtual  
631 station, the river is more distant from high relief, the altimeter is very frequently in 320 MHz mode  
632 (64 m window size) and that is also why it is the only SARAL/AltiKa virtual station observing river  
633 water elevations. However, increasing the tracking window size (with the same number of range  
634 bins) degrades the range resolution of the altimeter. Yet, SARAL better results over the estuary,  
635 compared to ENVISAT, must be linked to its improved range resolution.

636 Nadir altimeters observe a ground surface most of the time but this surface is not always the  
637 most useful for hydrologists. To overcome this issue and force the altimeter to observe the river  
638 valley instead of the surrounding hilly areas, the DIODE/DEM tracking mode has been developed  
639 by CNES (see section 2.2.2). SARAL/AltiKa measurements for virtual station SRL-773 were  
640 performed in DIODE/DEM mode during the first cycle of the mission and they are shown in Figure  
641 7 (red curve). In this mode, AltiKa successfully observes the Garonne valley without data loss and  
642 does not remain locked over the hills, despite the terrain steepness (highest terrain slopes are around  
643 80 m/km at 44.05°N and 100 m/km at 44.06°N).

644 Measurements for track 773 show the potential of the DEM mode to let the altimeter  
645 observe a river within a steep-sided valley. Yet, this mode requires that the a priori DEM stored on  
646 board has better accuracy than the size of the tracking window. For track 646, the on board DEM  
647 value is almost 40 m above the actual Garonne valley elevation. This discrepancy can be related to  
648 the Globcover classification, used in combination with ACE DEM, to compute on board elevations  
649 (see section 2.2.2). Around virtual station SRL-646, there is no water pixel on the Garonne River in  
650 Globcover (contrary to virtual station SRL-773), which biases on board elevation toward the top of

651 the hills elevation. Therefore, AltiKa loses signal over the Garonne even for cycle 1. A similar  
652 discrepancy occurs with Jason-2 cycles in DEM mode. For this altimeter mission, the GMT water  
653 mask used is not correctly geolocalized on the river (section 2.2.2). Thus, elevations computed  
654 along the satellite track and loaded on board are also close to the top of the surrounding hills  
655 elevations and not the river valley. Therefore, these cycles provide similar results to the Jason-2  
656 cycles in closed-loop, which remains locked on the top of the hills for both track 070 and 035. This  
657 example clearly shows DEM tracking mode sensitivity to the databases (a priori DEM and water  
658 mask) used to compute the on board elevations, especially if the tracking window is smaller.

659

### 660 **3.3. Observation of a narrow artificial canal**

661 A frequent question asked about nadir altimetry concerns the minimum water body size that  
662 can be observed with this type of altimeter. It is impossible to answer this question generally. From  
663 previous examples shown in this study, it is clear that the main reasons explaining why a water body  
664 is observed or not at a specific time is more linked with previous waveform history, instrument  
665 settings, ground topography rather than just the water body size. The previous example of the  
666 SARAL track 773 for cycle 1 on the Garonne River is a good example of such situation. All water  
667 bodies within the instrument footprint that backscatter enough energy will be observed in one or  
668 multiple range gates (if they are in the tracking window). In this case, the waveform will have  
669 multiple peaks (with different amplitudes), corresponding to these water bodies. They could also be  
670 observed on multiple consecutive waveforms along the satellite ground track. Retracker, like Ice-1  
671 used in this study, use the whole waveform measured by the altimeter within the tracking window to  
672 estimate one range value. Therefore, Ice-1 provides elevation for only one observed water body (the  
673 first peak), but not for the others.

674 To account for the heterogeneity of the scene observed by the altimeter and all the potential  
675 targets measured by the instrument, it is beneficial to plot the radargram, which corresponds to  
676 waveforms recorded by the altimeter along the satellite track around a virtual station. The radargram

677 for virtual station SRL-773 during cycle 1 (in DIODE/DEM tracking mode) is shown in Figure 8.  
678 This figure shows the history of the returned power along the track over the virtual station. The x-  
679 axis corresponds to time (or along-track latitude) and the y-axis to the range gate number  
680 (equivalent to distance). The intensity of the returned power (normalized to the maximum power  
681 registered during the pass, in decibels) is shown by the color. The parabolic shapes observed on this  
682 figure are characteristic of the signal returned by small size water bodies. The returned signal is  
683 received by the altimeter a few kilometers before and after the satellite crosses the river. The  
684 variation of distance between the river and the radar during this period explains the parabolic shape  
685 (closest approach corresponds to the minimum of the parabola). Two examples of such observations  
686 in Figure 8 are caused by the Garonne River and the narrow artificial canal. Two other points can  
687 also be observed and correspond to very high intensity of the received signal. These points also  
688 correspond to bright targets and the absence of parabola associated to them is caused by the nature  
689 of the retrodiffusion by these targets: they are very specular in contrast to the two points previously  
690 discussed. For more information on diffusive and specular targets responses in a radargram, see, for  
691 example, Tournadre et al. (2006). Figure 2.a shows specifically SARAL track 773 for cycle 1  
692 (dashed red line) and the overflowed water bodies near SRL-773 virtual stations. This specific track  
693 is 1.5 km shifted compared to the nominal ENVISAT/SARAL track, due to some less stringent  
694 requirement on the satellite position during the first cycles. From Figure 2.a, it is clear that the first  
695 specular target and the first diffusive target corresponds to the Southernmost lake/reservoir and the  
696 Garonne River mainstream (southernmost channel on this figure, which is ~150 m wide) of the  
697 image, respectively. The second diffusive target, which is higher than the Garonne mainstream,  
698 corresponds to the artificial canal (northernmost channel, ~70 m wide), which brings cooling water  
699 to the Golfech nuclear power plant. It is less clear what the second specular target is (it could be  
700 another smaller lake/reservoir, a bright man made structure like a road or metallic building roofs or  
701 even another much smaller artificial canal : the “canal du Midi”). Positions of the Garonne River  
702 and the Golfech canal along SARAL track are also indicated on Figure 7 (blue polygons). This

703 example clearly shows that nadir altimeters can observe small targets (river or canal with width  
704 below 100 m), when the tracking window is correctly set.

705

#### 706 **4. Conclusion and perspectives**

707 Nadir altimeters have proven their capability to observe water elevation for major rivers  
708 (like the Amazon, the Congo, etc.). In this study, it has been shown they can also provide  
709 meaningful water elevation for a 200 m wide, steep-sided river: the lower Garonne River in France.  
710 Jason-2 time series measures water elevation with 20 cm RMSE compared to in situ observations at  
711 Tonneins (115 km upstream the estuary), whereas ENVISAT mission had higher RMSE (50 cm)  
712 compared to the same in situ gage. With good reason, Jason-2 10-day repeat orbit is better suited to  
713 observe Garonne River seasonal cycle than the ENVISAT 35-day orbit. Therefore, Jason-2 (and to a  
714 lower extent ENVISAT) repeat period seems appropriate to observe water level variations at the  
715 seasonal cycle, annual, interannual and even decadal time scale (since Jason-2 was launched in  
716 2008). However, Jason-2 time sampling is too coarse for observing daily/hourly high frequency  
717 water level variations for this kind of medium size river. The Garonne River is very sensitive to  
718 short-time intense rain events and quick snowmelt, which induces several meters water elevation  
719 variations in few days and all these rapid events are missed by the satellite altimeters. Upstream  
720 Tonneins, ENVISAT has higher errors and does not seem to measure water elevation as accurately  
721 (Jason-2 does not sample the river upstream).

722 Comparisons between ENVISAT and in situ water elevations showed that Ice-1 retracker  
723 and Sea Ice retracker provide very similar results when the Garonne River is around 200 m-wide,  
724 confirming what was obtained by Sulistioadi et al. (2015). This is due to the peaky shape of the  
725 waveform for small and medium size water bodies. When studying drainage basins with various  
726 river widths, it would be better to use only Ice-1 altimeter heights for consistency between the  
727 altimetry-based time series of water levels. Ice-1 retracking algorithm provides much better results  
728 than Sea Ice over large rivers and wetlands. Ice-1-derived altimeter ranges are available in the GDR

729 for all recent altimetry missions, which is not currently the case for Sea Ice.

730 SARAL/AltiKa mostly fails to observe the river valley and remains locked on the hilly  
731 surrounding areas. Such problem also happens with ENVISAT, but less often. It also happens quite  
732 frequently for the Jason-2 virtual station downstream of Marmande (the other one near Tonneins is  
733 less affected). This issue is related mainly to the closed-loop tracking algorithm which is influenced  
734 by the history of the measurements and the geometry of observation. As a consequence, over the  
735 continents, nadir altimeters tend to be locked over the top of the topography within the instrument  
736 footprint and during previous measurements. This is the case over the Garonne River for  
737 SARAL/AltiKa and for Jason-2 track 035, which crosses the narrow (~5 km) and steep-sided  
738 Garonne River valley almost perpendicularly. However, over 30 km, Jason-2 track 70 is almost  
739 parallel to the river and within the valley, which allows the closed-loop tracking mode to get locked  
740 on the river. ENVISAT provides more measurements on the river, not because of a different  
741 observation geometry (SARAL has the same orbit as ENVISAT), but because of differences in  
742 closed-loop tracking parameters and its three tracking window sizes. Different window sizes help  
743 sample a wider range span. Over the estuary, SARAL/AltiKa provides smaller RMSE (around 10  
744 cm) than ENVISAT (around 30 cm) compared to tide gages.

745 To overcome challenges inherent to closed-loop tracking mode that tends to observe top of  
746 the topography instead of steep-sided river, an experimental tracking mode has been developed by  
747 the CNES: the DIODE/DEM mode. For SARAL/AltiKa, this experimental mode has been activated  
748 only during the first cycle. Over the Garonne River, it successfully observed the Garonne valley for  
749 track 773, whereas for other cycles, in closed-loop tracking mode, it failed to observe it. Yet, this  
750 mode requires a priori DEM values (derived from ACE2 for SARAL) and land/water mask (derived  
751 from Globcover for SARAL) to compute DEM on board along the satellite track. This on board  
752 computed DEM must have vertical accuracy better than the tracking window size (which is  
753 typically in the range 30-50 m, see section 2.2.2), otherwise it provides incorrect tracking  
754 commands and misses the river valley (like for SARAL track 646 cycle 1 and Jason-2 track 070 and

755 035 cycles in DEM tracking mode over the Garonne). Therefore, for this tracking mode, it is crucial  
756 to have an a priori, validated, database of the expected elevation for all water bodies the altimeter  
757 will be forced to observe, with vertical accuracy better than the size of the tracking window. The on  
758 board DEM is highly dependent of both input water mask and DEM used, which should be  
759 consistent among themselves. On board DEM can be improved by using satellite imagery (e.g.  
760 Landsat) for more accurate water mask, like the NARWidth database (Allen and Pavelsky, 2015).  
761 For DEM values, it should be assessed if and where current global DEM are compatible with the  
762 used water mask (e.g. elevations around the water mask should be lower than surrounding  
763 topography in steep-sided valleys) and accurate enough. Using new (or soon to be released)  
764 improved DEM, like the new version of Shuttle Radar Topography Mission (SRTM) DEM (released  
765 in September 2014, see <http://www2.jpl.nasa.gov/srtm/>) or the DLR global TanDEM-X DEM (Zink  
766 et al., 2014), should also improve computed on board DEM.

767 Therefore, the DIODE/DEM mode seems promising for future altimetry missions to observe  
768 previously missed steep-sided rivers (or lakes). However, performance, benefits and limits of this  
769 mode for continental hydrology will require more investigation by the scientific community,  
770 especially because two new altimetry satellites have just been launched (Jason-3, 17 January 2016,  
771 and Sentinel-3A, 16 February 2016). These altimeters are equipped with the DIODE/DEM mode  
772 available along with the closed-loop tracking mode.

773

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790

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1054 **Tables**

1055 Table 1. Altimetry missions used in this study and their main characteristics

	Altimetry missions used		
Mission name	ENVISAT	Jason-2	SARAL
Operating agencies	ESA	CNES, EUMETSAT, NASA, NOAA	CNES, ISRO
Mission duration	2002 - 2010 on the nominal orbit (below)	2008 - present	2013 - present
Orbit repeat period	35 days	10 days	35 days
Orbit altitude	790 km	1336 km	790 km
Orbit inclination	98.54°	66°	98.54°
Equatorial ground-track distance	80 km	315 km	80 km
Altimeter name	RA-2	Poseidon-3	AtiKa
Radar frequencies	Ku- and S-bands	Ku- and C-bands	Ka-band
Range window size	64 m, 256 m, 1024 m	50 m	30 m
Number of used range window bins	104	104	104
Along track sampling	18 Hz (~400 m)	20 Hz (~350 m)	40 Hz (~175 m)
DEM tracking mode	No	Yes	Yes
DEM used for DEM mode	-	ACE	ACE2
Cycle(s) in DEM tracking mode	-	3, 5, 7, 34, 209, 220	1

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1057

1058 Table 2. Satellite altimetry virtual stations name, closest in situ gage name, number of common  
 1059 dates in the altimetry and in situ time series, correlation coefficient, mean bias and RMSE between  
 1060 altimetry and in situ water elevation time series, RMSE and Nash-Sutcliffe coefficient between  
 1061 altimetry and in situ water elevation anomaly time series, and amplitude (maximum minus  
 1062 minimum) of the in situ water elevation time series at the common dates, along the Garonne River  
 1063 mainstream (from upstream to downstream virtual stations).

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Alti virtual station (+ length river crossing)	In situ gage (+ river width in m)	Dist alti/gage (km)	Nb dates alti/ in situ	Corr coeff	Mean bias (m)	RMSE (m)	Anom RMSE (m)	Anom NS coeff	In situ amplitude (m)
ENV-102 (250 m)	Verdun-sur-Garonne (130 m)	29.5	48	0.41	27.90	27.92	0.95	-3.10	1.56
ENV-773 (130 m)	Lamagistère (150 m)	10.5	46	0.35	8.66	8.80	1.55	-2.63	3.52
SRL-773 (130 m)	Lamagistère (150 m)	10.5	6	-0.80	105.10	105.13	2.56	-3.88	3.29
ENV-646 (175 m)	Lamagistère (150 m)	1.4	64	0.62	1.61	1.80	0.81	-0.53	2.60
SRL-646 (175 m)	Lamagistère (150 m)	1.4	6	-0.14	103.53	103.55	2.01	-2.00	3.29
ENV-188 (700 m)	Tonneins (200 m)	2.5	70	0.85	-0.23	0.58	0.53	0.73	3.83
SRL-188 (700 m)	Tonneins (200 m)	2.5	11	0.90	-1.77	1.90	0.69	0.80	5.3
JA2-070 (185 m)	Tonneins (200 m)	5.3	165	0.98	-2.56	2.57	0.22	0.95	4.41
JA2-070 (185 m)	Marmande (200 m)	9.2	150	0.98	4.04	4.04	0.20	0.96	4.20
ENV-315 (2 km)	Marmande (200 m)	14.2	21	0.81	-1.27	1.41	0.59	0.66	3.43
SRL-315 (2 km)	Marmande (200 m)	14.2	4	-0.84	44.15	44.19	1.97	-1.36	3.48
JA2-035 (160 m)	Marmande (200 m)	30	62	0.91	-6.95	6.96	0.36	0.82	3.74

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1069

1070 Table 3. Correlation coefficient, RMSE (in meter) and Nash-Sutcliffe coefficient between water  
 1071 elevation anomaly time series for ENVISAT altimetry and the closest in situ gage (see Table 2),  
 1072 along the Garonne River mainstream (from upstream to downstream virtual stations). The  
 1073 comparison is done for two retracker algorithms (Ice-1 and Sea Ice) and using rather the median or  
 1074 the mean of retracked points within the virtual station per observation time. In bold number  
 1075 correspond to the better result obtain for each virtual station and each statistic.

		ENVISAT virtual stations				
		ENV-102	ENV-773	ENV-646	ENV-188	ENV-315
Corr coeff	Median Ice-1	0.42	<b>0.35</b>	0.63	0.85	0.82
	Mean Ice-1	0.41	0.35	0.62	0.85	<b>0.84</b>
	Median Sea Ice	<b>0.50</b>	0.33	<b>0.65</b>	<b>0.88</b>	0.79
	Mean Sea Ice	0.50	0.33	0.65	0.87	0.80
Anom RMSE (m)	Median Ice-1	0.95	<b>1.55</b>	0.81	0.53	0.59
	Mean Ice-1	0.94	1.55	0.80	0.54	<b>0.57</b>
	Median Sea Ice	<b>0.89</b>	1.80	<b>0.70</b>	<b>0.51</b>	0.63
	Mean Sea Ice	0.89	1.80	0.69	0.51	0.61
Anom NS coeff	Median Ice-1	-3.10	<b>-2.63</b>	-0.53	0.73	0.66
	Mean Ice-1	-3.01	-2.63	-0.48	0.72	<b>0.69</b>
	Median Sea Ice	<b>-2.60</b>	-3.90	-0.12	<b>0.76</b>	0.62
	Mean Sea Ice	-2.60	-3.90	<b>-0.09</b>	0.75	0.65

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1077

1078 Table 4. Satellite altimetry virtual stations name, closest tide gage name, number of common dates  
 1079 in the altimetry and in situ time series, correlation coefficient, mean bias and RMSE between

1080 altimetry and in situ water elevation time series, RMSE and Nash-Sutcliffe coefficient between  
 1081 altimetry and in situ water elevation anomaly time series, and amplitude (maximum minus  
 1082 minimum) of the in situ water elevation time series at the common dates, along the Gironde estuary.

1083

Alti virtual station	Tide gage	Dist alti/gage (km)	Nb dates alti/ in situ	Corr coeff	Mean bias (m)	RMSE (m)	Anom RMSE (m)	Anom NS coeff	In situ amplitude (m)
ENV-274	Richard	6.5	35	0.28	-	-	1.51	-0.91	5.26
SRL-274	Richard	6.5	15	0.09	-	-	1.57	-0.21	4.54
ENV-859	Royan	5.0	17	0.98	1.01	1.06	0.31	0.94	3.76
SRL-859	Royan	5.0	15	0.99	0.32	0.34	0.09	0.99	3.27
ENV-859	Port-Bloc	3.7	39	0.97	0.73	0.78	0.26	0.94	3.55
SRL-859	Port-Bloc	3.7	15	0.99	0.14	0.20	0.13	0.98	3.26

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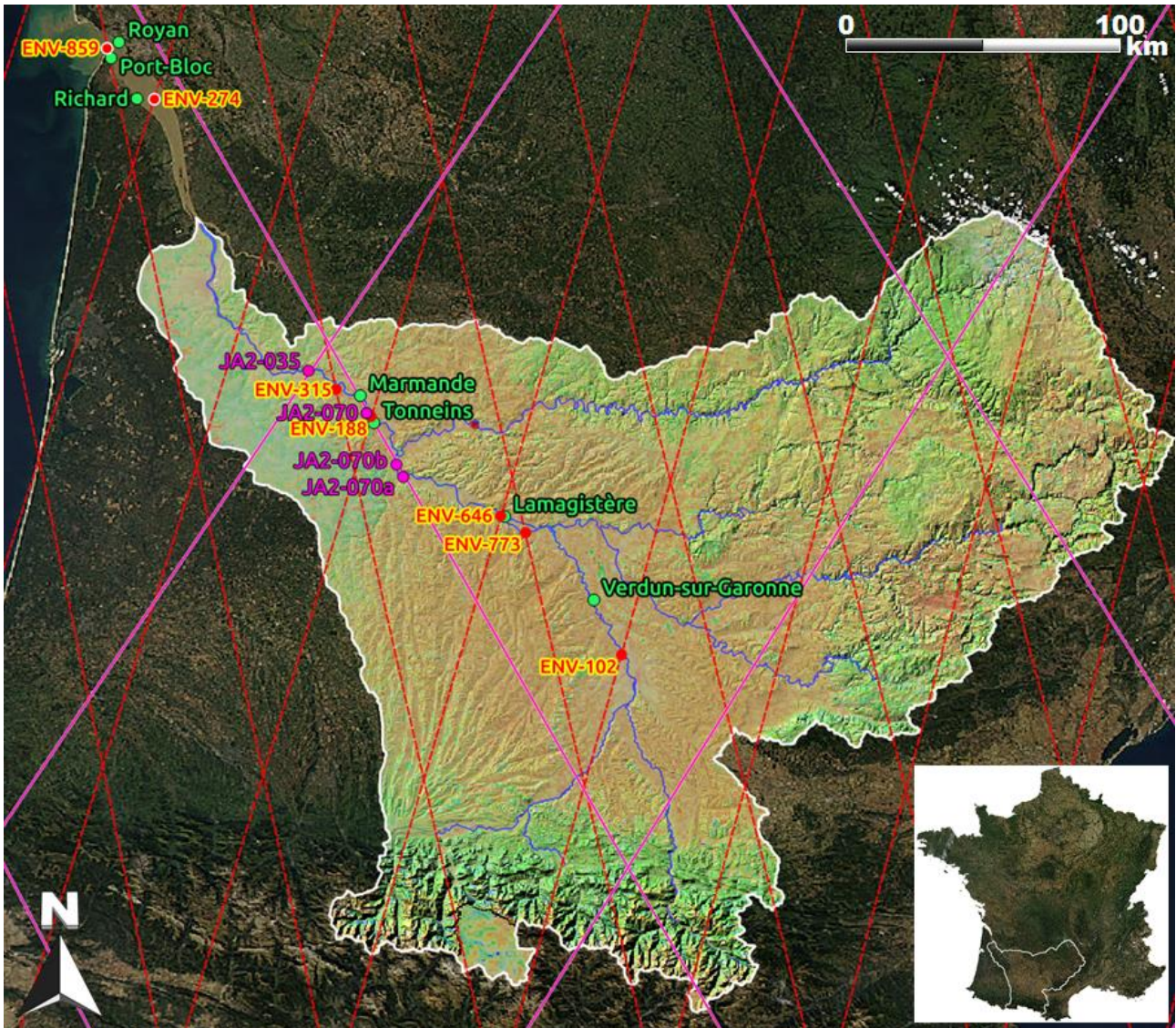
Table 5. Percentage of data, for all ENVISAT cycles, acquired in each tracking mode, for all virtual stations. 320 MHz, 80 MHz and 20 MHz tracking modes correspond to a tracking window size of 64 m, 256 m and 1024 m, respectively.

		ENVISAT virtual stations						
		ENV-102 (river)	ENV-315 (river)	ENV-773 (river)	ENV-646 (river)	ENV-188 (river)	ENV-274 (estuary)	ENV-859 (estuary)
ENVISAT tracking modes	320 MHz	28	77	44	76	80	96	100
	80 MHz	71	22	33	23	19	3	0
	20 MHz	0	1	23	1	1	1	0

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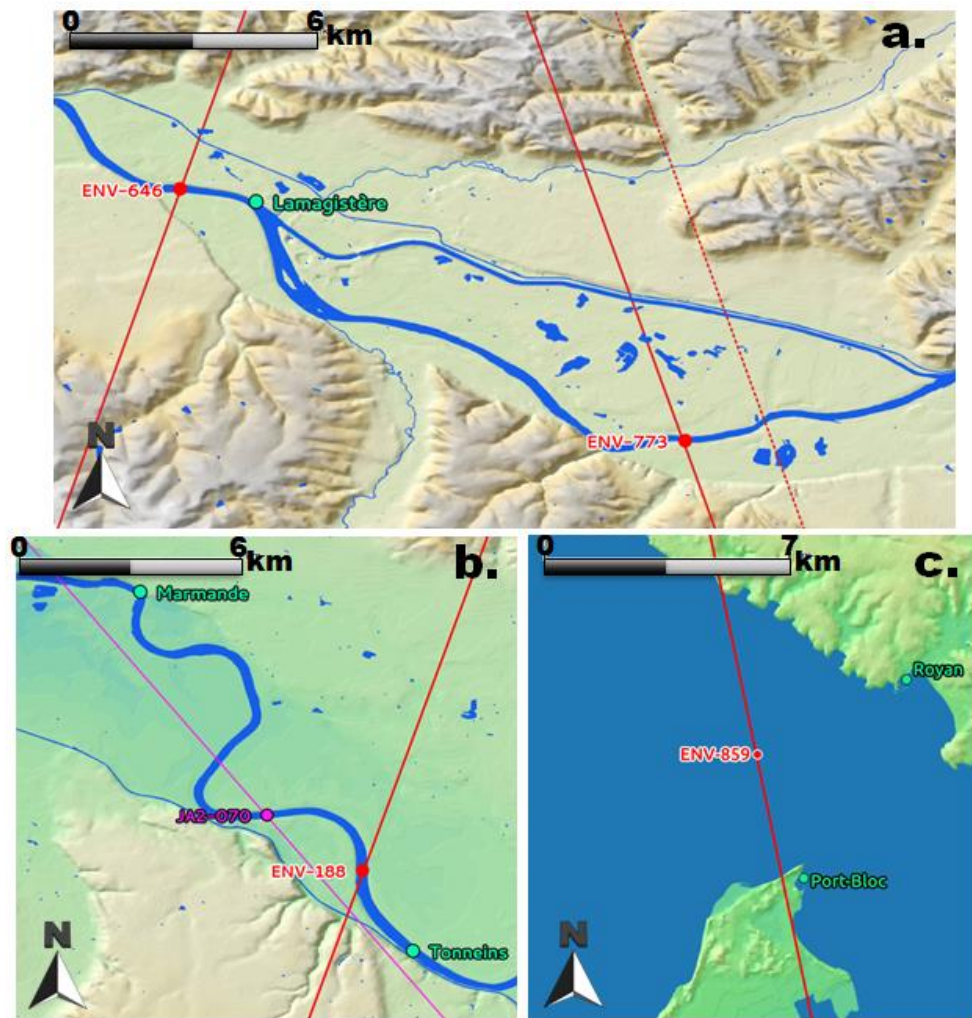
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## Figures



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Figure 1. Garonne basin (white boundaries), with in situ gages (green dots), ENVISAT and SARAL virtual stations (red dots) and Jason-2 virtual stations (magenta dots) used in this study. Dotted red lines correspond to ENVISAT and SARAL ground tracks. Magenta lines correspond to Jason-2 ground tracks. Background colors correspond to a MODIS image from 01 October 2011, whose color has been artificially lightened using a shaded relief, computed from a DEM provided by IGN, over the Garonne basin. Location of the Garonne basin in France main territory is shown in the bottom right hand corner map (white boundaries)



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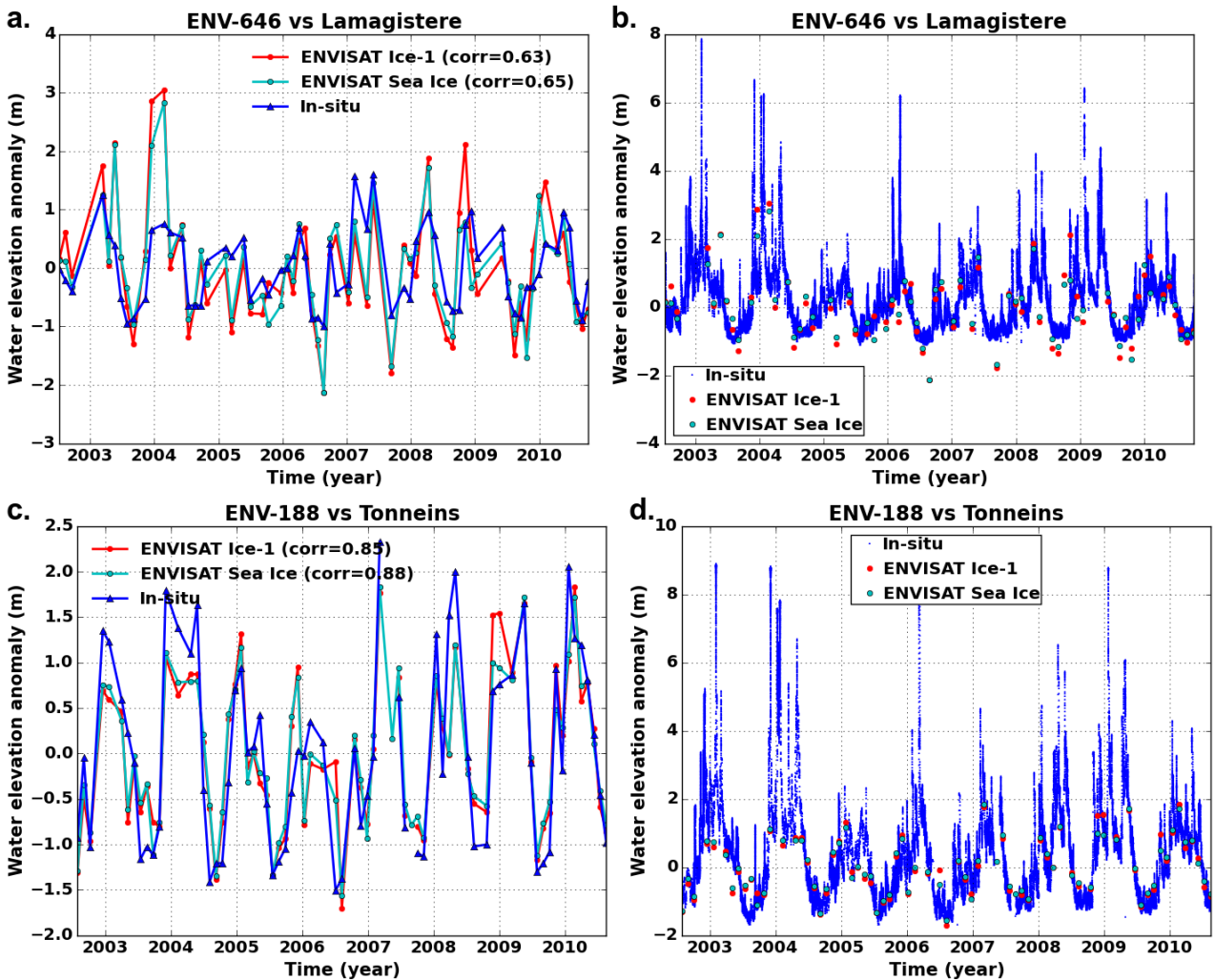
1098 *Figure 2. ENVISAT satellite nominal ground tracks (red line), ENVISAT virtual stations (red dots),*

1099 *Jason-2 satellite nominal ground track (magenta), Jason-2 virtual stations (magenta dot) and in*

1100 *situ gages (green dots) over the Garonne mainstream at Lamagistère (a.), Tonneins (b.) and over*

1101 *the estuary (c.). Dashed red line on panel (a.) corresponds to SARAL ground track 773 for cycle 1*



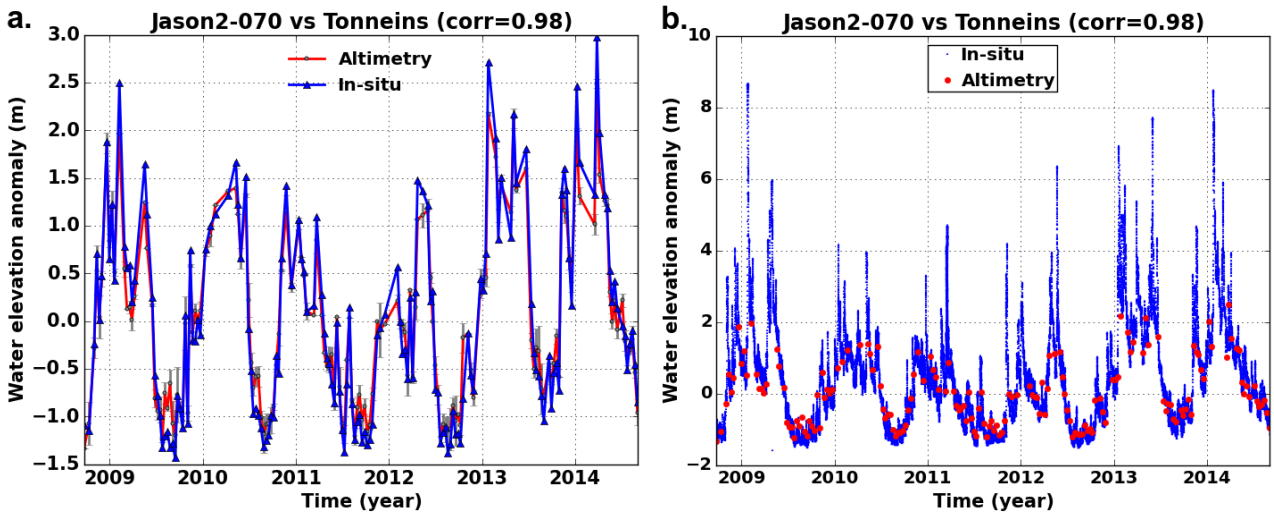


1102  
 1103 *Figure 3. ENVISAT water elevation anomaly computed with Ice-1 retracker (red line on panels a.*  
 1104 *and c. and red dots on panels b. and d.) and with Sea Ice retracker (cyan line on panels a. and c.*  
 1105 *and cyan dots on panels b. and d.), in situ water elevation anomaly at ENVISAT measurement times*  
 1106 *(blue line on panels a. and c.) and all records in the in situ water elevation anomaly time series*  
 1107 *between 2003 and 2010 (blue dots on panels b. and d.) for virtual station 646 and Lamagistère gage*  
 1108 *(panels a. and b.) and virtual station 188 and Tonneins gage (panels c. and d.). In these panels,*  
 1109 *water elevation anomaly time series correspond to water elevations minus the time series temporal*  
 1110 *mean over the same common dates between in situ and altimetry time series. The term ‘corr’*  
 1111 *corresponds to the altimetry time series and the corresponding in situ time series correlation*  
 1112 *coefficient. For more statistics concerning these virtual stations, see Table 2*

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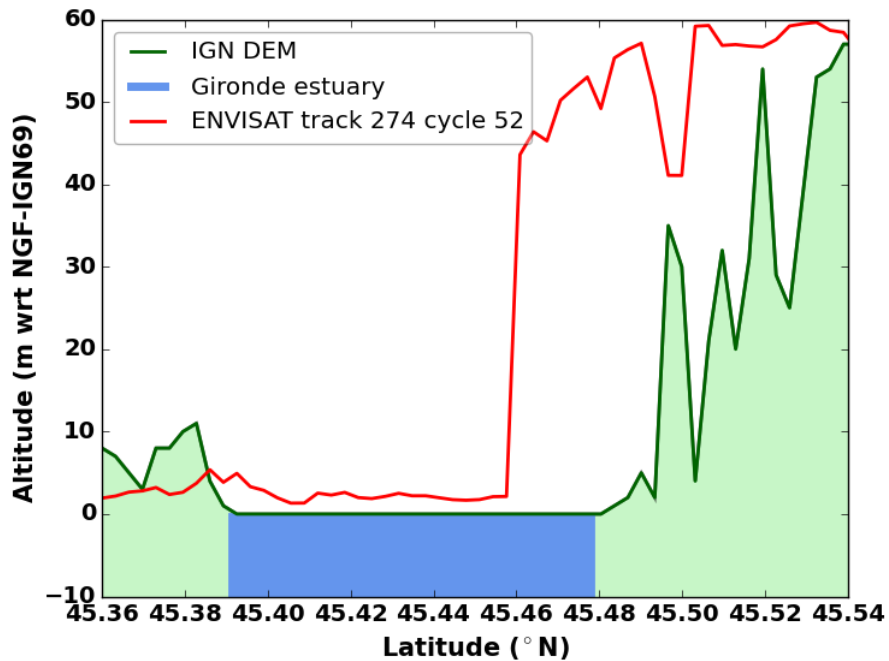
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1117 *Figure 4. Jason-2 virtual station 070 water elevation anomaly (red line on panel a. and red dots on*  
 1118 *panel b.) and in situ water elevation anomaly measured by the Tonneins gage at Jason-2*  
 1119 *measurement times (blue line on panel a.) and all records in the in situ water elevation anomaly*  
 1120 *time series between 2003 and 2010 (blue dots on panel b.)*

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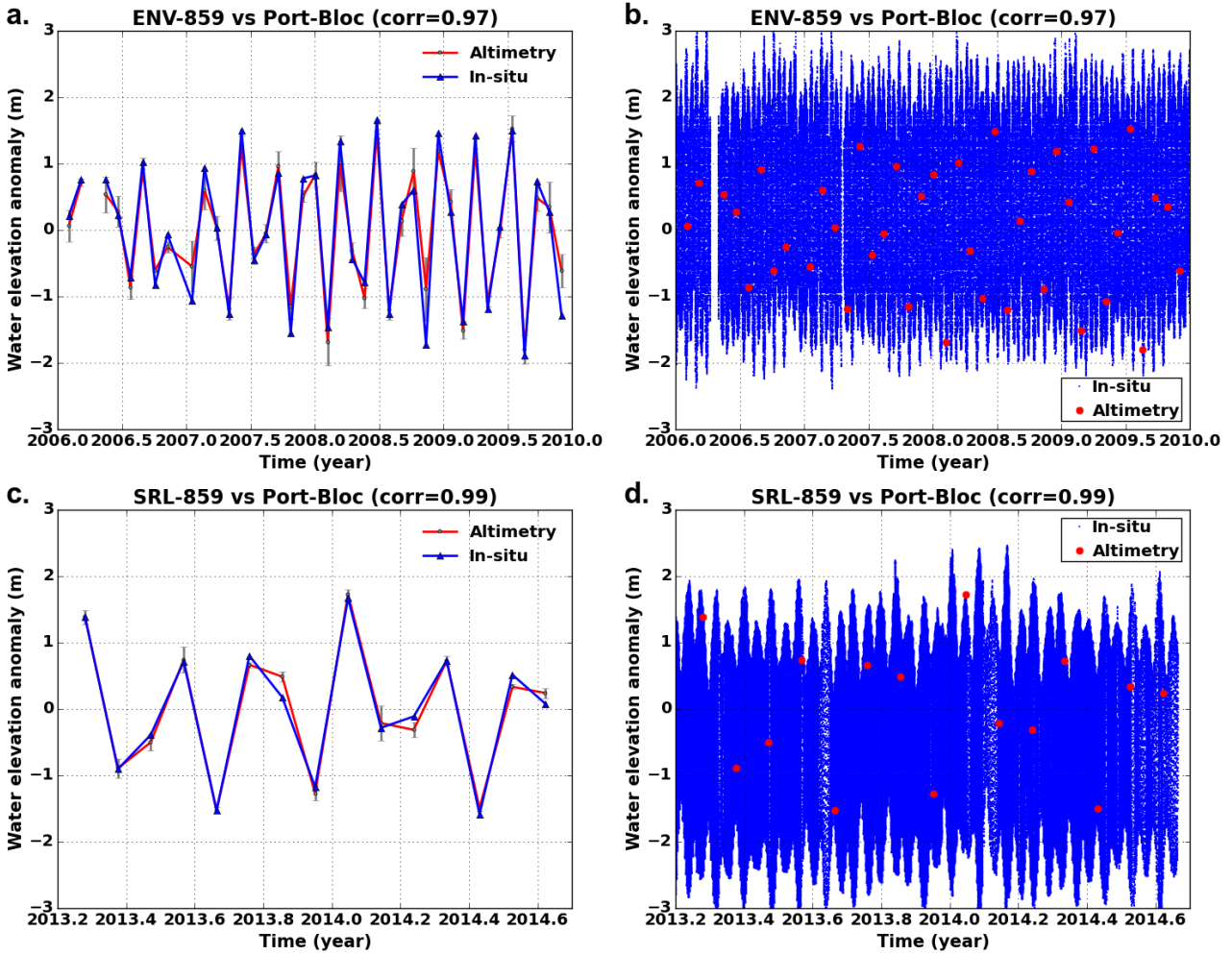


1122

1123 *Figure 5. Comparison between IGN DEM (green line) along the ENVISAT 274 track (see Figure 1)*  
 1124 *over the Gironde estuary (its position is shown by the light blue polygon, its elevation corresponds*  
 1125 *to the DEM and not to the actual water level) and elevation measured by ENVISAT for cycle 59 (21*

1126 June 2007, red curve)

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1129 Figure 6. Comparison between virtual station 859 water elevation anomaly (red line on panels a.

1130 and c. and red dots on panels b. and d.), water elevation anomaly register by the Port-Bloc tide

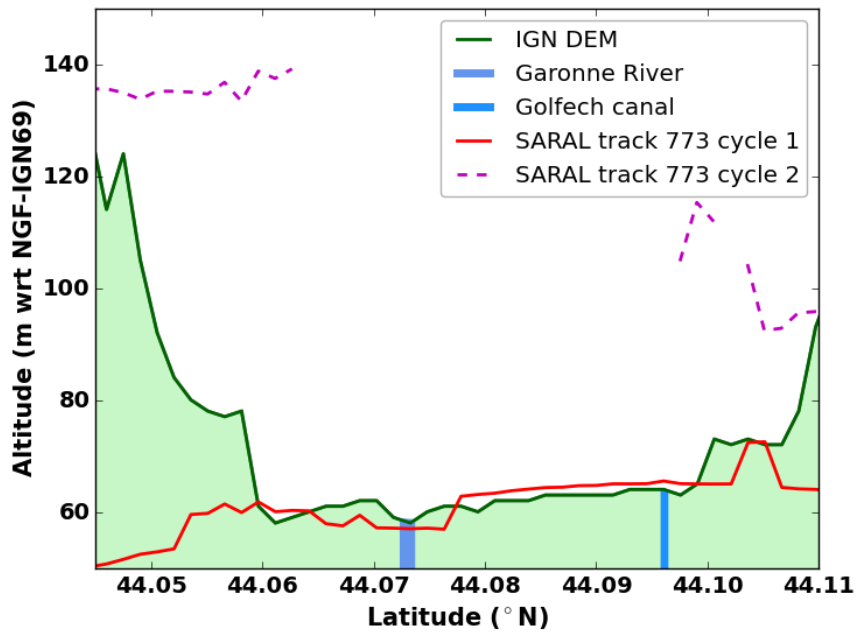
1131 gage (blue line on panels a. and c.) at altimeter measurement times and all tide gage measurements

1132 (blue dots on panels b. and d.) for ENVISAT between 2006 and 2010 (panels a. and b.) and for

1133 Saral between 2013 and 2014 (panels c. and d.)

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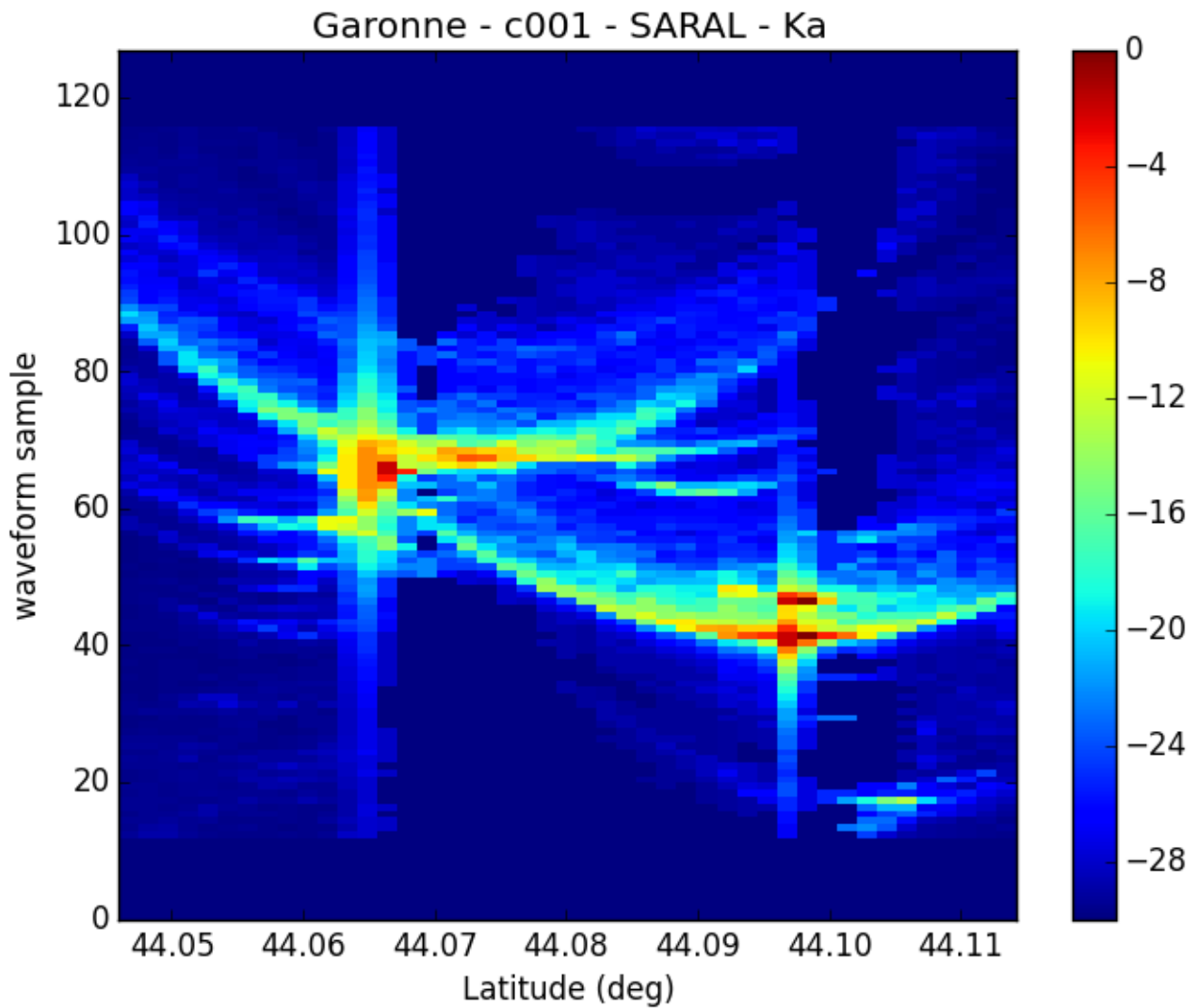
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1137 *Figure 7. Comparison between IGN DEM (green line) along the SARAL 773 track (see Figure 1)*  
 1138 *over the Garonne River (its position is shown by the large blue polygon near 44.07°N, its elevation*  
 1139 *corresponds to the DEM and not to the actual river water level) and the artificial canal (tight light*  
 1140 *blue polygon near 44.09°N), elevations measured by SARAL/AltiKa for cycle 1 (10 April 2013, red*  
 1141 *curve) and for cycle 2 (15 May 2013, dashed magenta curve)*

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1144 *Figure 8. Radargram for SARAL/AltiKa cycle 1 pass over virtual station SRL-773. This is a*  
 1145 *representation of the successive waveforms measured on this pass. Each column represents one*  
 1146 *waveform (y-axis correspond to range gates number). Colors represent returned power received in*  
 1147 *each bin. A logarithmic scale (decibels) is used in order to be able to represent the large dynamic of*  
 1148 *received power.*