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Velocity-Density Systematics of Fe-5wt%Si: Constraints on Si Content in the Earth's Inner Core

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Key Points: 9

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10 11	•	Isothermal compression of Fe-5wt.%Si at high temperatures up to 1.1 Mbar and 2100 K $$
12	•	Measurement of Vp of Fe-5wt.%Si under quasihydrostatic conditions to 1.1 Mbar
13	•	Si cannot be the sole light element in the Earth's inner core

• Si cannot be the sole light element in the Earth's inner core

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14 Abstract

¹⁵ The elasticity of hcp-Fe-5wt%Si has been investigated by synchrotron X-ray

¹⁶ diffraction up to 110 GPa and 2100 K, and by picosecond acoustics measure-

¹⁷ ments at ambient temperature up to 115 GPa. The established Pressure-Volume-

¹⁸ Temperature (PVT) equation of state shows that the density of the Earths inner core

¹⁹ can be matched by an Fe-Si alloy with 5wt% Si for all reasonable core temperatures, but

that its compressional and shear velocities remain too high with respect to seismolog-

ical observations. On the other hand, Fe-Si alloys whose velocities are expected to get close to seismological observations are too dense at relevant temperatures. Thus, based

on these combined velocity-density measurements, silicon is not likely to be the sole light

element in the inner core.

25 1 Introduction

Iron and iron alloys at extreme conditions have garnered significant interest due 26 to their relevance to Earths deep interior. While a first 1D reference seismological model 27 providing density (ρ) , compressional (Vp) and shear (Vs) sound velocities as a function 28 of depth into the Earth (PREM Preliminary Reference Earth Model) was established 29 in the 1980s (Dziewonski and Anderson, 1981), there is still considerable debate over the 30 chemical composition of the Earths core. On cosmochemical grounds, iron is considered 31 to be the most abundant element in the Earth's core (e.g., Allégre et al., 1995; McDonough 32 and Sun, 1995). However, early on in the study of Fe at extreme conditions, it was ob-33 served that Fe is too dense at the pressure (P) and temperature (T) conditions of the 34 core to be the sole element present (Birch, 1952). Earths solid inner core exists at pres-35 sures of 330-360 GPa, and at temperatures of 5000-7000K based on the melting curve 36 of Fe (Anzellini et al., 2013; Boehler, 1993; Nguyen and Holmes, 2004). As a consequence 37 of the density difference between Fe and PREM, there needs to be some quantity of el-38 ements lighter than Fe alloyed to it in order to compensate for this density deficit. Among 39 the potential light element candidates. Si has been favored by many recent studies but 40 without reaching a firm consensus. Regardless of the nature of accretionary materials 41 and redox path, all recent core differentiation models based on metal-silicate partition-42 ing support the presence of Si in the core (Fischer et al., 2015; Siebert et al., 2013). The 43 presence of Si is also advocated on the basis of isotopic arguments (Fitoussi et al., 2009). 44 The possible presence and the quantity of Si in the Earths core has important implica-45 tions for geodynamic processes and the bulk redox state of the Earths interior (Hirose 46 et al., 2017; Wade and Wood, 2005; Wood et al., 1990). One possible way to constrain 47 the Si content of the inner core is the comparison between seismological data and exper-48 imental measurements, or calculations, of ρ , Vp and Vs of candidate materials at per-49 tinent PT conditions (Sakairi et al., 2018; Sakamaki et al., 2016). Based off the measure-50 ment of Vp vs. ρ at extreme conditions, estimates have varied from 1-2 wt% Si (Anto-51 nangeli et al., 2010; Badro et al., 2007) to $\sim 8 \text{ wt\%}$ Si (Fischer et al., 2014; Mao et al., 52 2012), with the most recent measurements putting an upper limit of 5wt% Si (Antonan-53 geli et al., 2018; Sakairi et al., 2018). 54

Thanks to the adaptation of Picosecond Acoustics (PA) to the Diamond Anvil Cell 55 (DAC), it is possible to make direct measurements of the acoustic travel time of Fe-alloys, 56 and metals in general, at very high pressures (Decremps et al., 2014,0). Additionally, PA 57 has fewer limitations on sample dimensions than conventional synchrotron-based tech-58 niques, allowing the measurement of Vp under quasihydrostatic conditions to Mbar pres-59 sures. We thus used PA to probe acoustic echoes and the compressional sound velocity 60 of an Fe-Si allov with 5wt. % Si up to 115 GPa. We complemented these measurements 61 with synchrotron x-ray diffraction measurements in laser-heated diamond anvil cells up 62 to 110 GPa and 2100 K, deriving a P-V-T equation of state. Our results provide tight 63 constraints on the Si abundance in the Earths inner core. 64

65 2 Methods

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2.1 Sample Preparation

All experiments were performed on an Fe-Si alloy synthesized by melt spinning at 67 ICMPE laboratory in Thiais, France (melt temperature T = 1823 K, wheel velocity v 68 = 20 m/s, (Morard et al., 2011)). Scanning electron microscopy measurements have shown 69 the alloy to be chemically homogeneous and contain 5.2(2) wt% Si, and so the alloy will 70 be referred to hereafter as Fe-5Si for brevity. Fe-5Si was also measured by grazing-incidence 71 x-ray diffraction (XRD) at ambient conditions to check phase purity and density, show-72 ing minimal texture and excellent polycrystallinity. Experiments were performed using 73 Le Toullec-type membrane-driven DACs with diamonds of culet size ranging from 250 74 μm down to 100 μm and equipped with Re gaskets. All experiments at ambient tem-75 perature were performed with Ne as the pressure-transmitting medium (PTM) to en-76 sure quasihydrostatic conditions up to Mbar conditions. Samples were scraped from a 77 large, thin ribbon, and sample dimensions were chosen such that there was no bridging 78 between sample and diamonds, nor contact between gasket and sample. For PA exper-79 iments, the thickness of the sample was determined by the measured initial travel time 80 and derived Vp from literature elastic constants (see 2.3). In PA Run 1 the sample was 81 determined to be $3.73(6) \ \mu m$ thick, and for Run 2 the sample was determined to be 1.84(3)82 μm thick. For high-T XRD the sample was 1-3 μm thick, and for ambient-T XRD the 83 sample was 5-7 μ m thick, estimated by optical microscopy (during sample loading) and 84 observations of diffraction intensity relative to PTM. 85

2.2 X-ray Diffraction

For room temperature measurements, Fe-5Si was loaded alongside Mo (Pressure 87 calibration: $K_0 = 260(1)$ GPa, K' = 4.19(5), using the reference volume of Mo - $V_0 =$ 88 31.17 Å³, Litasov et al., 2013; Ross and Hume-Rothery, 1963) or Pt ($K_0 = 277.3$ GPa, 89 $K' = 5.12, V_0 = 60.38 \text{ Å}^3$ Dorogokupets and Oganov, 2007) as a pressure calibrant, with 90 Ne as the PTM. For experiments at high temperatures, the samples were sandwiched 91 between KCl disks which served as the PTM, and to thermally insulate the sample from 92 the diamonds. KCl also served as the pressure calibrant. The samples pressure at high 93 temperatures was determined by diffraction of KCl, with the method for estimating tem-94 perature following Campbell et al. (2009) and the KCl pressure calibration of Dewaele 95 et al. (2012). By virtue of its very low thermal pressure, uncertainties on pressure due 96 to thermal gradients across KCl are below reported error bars. After loading the DAC, 97 the assembly was left open to dry in a vacuum oven at 130° C for at least 1 hr after which 98 the DAC was closed. This practice ensures that moisture content is minimal within the 99 sample chamber. Angle-dispersive XRD measurements were performed at beamline ID27 100 at the ESRF (Mezouar et al., 2005). The radiation was monochromatic ($\lambda = 0.3738$ nm) 101 focused to approximately 3 x 3 μ m² (horizontal x vertical FWHM). Diffraction patterns 102 were collected on a MarCCD camera, with collection times of 30-60 seconds. Samples 103 were heated on both sides by two continuous Nd:YAG fibre lasers (TEM00), each one 104 delivering up to 200 W. Hot spots were approximately 20 μ m in diameter, much larger 105 than the FWHM of the focused X-ray beam. All temperatures were measured by the spec-106 troradiometric method, using a Planck fit of the observed blackbody radiation from the 107 center of the heating spot, as described by Mezouar et al. (2017). While absolute errors 108 in temperature are on the order of 150 K, the measured temperature was seen to vary 109 by less than 30 K during pattern integration (averaged over 3-5 measurements per diffrac-110 tion pattern). Diffraction images were calibrated against a CeO_2 standard, and then ra-111 dially integrated using Dioptas image processing software (Prescher and Prakapenka, 2015). 112 The integrated diffraction pattern was then analysed by use of Le Bail fits in the soft-113 ware Jana2006 (Václav et al., 2014). 114

2.3 Picosecond Acoustics

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These experiments are performed in pump-probe configuration, where laser pulses 116 generated by a Maitai Ti:Sapphire laser ($\lambda = 800$ nm, pulse duration = 100 fs) are sep-117 arated into two beams which are focused at the two opposing faces of the metallic sam-118 ple. The majority of the intensity ($\sim 80\%$) of the laser is directed towards the pump side 119 (30-100 mW depending on experimental conditions), where the beam generates a small 120 thermal stress at the surface of the sample. The relaxation of this thermal stress gen-121 erates an elastic wave which propagates through the sample. The probe beam, compris-122 123 ing the remainder of the lasing intensity, is analysed by means of an interferometer, in order to detect the change of phase of reflectivity at the sample surface. This quantity 124 changes abruptly upon arrival of the acoustic wave, and therefore provides an accurate 125 determination of the acoustic travel time across the material. Further details of the setup 126 are provided elsewhere (Decremps et al., 2015). At ambient conditions, the value of Vp 127 was derived from single crystal elastic constants (Machová and Kadečková, 1977) using 128 the Hashin-Strikman average (Hashin and Shtrikman, 1962) and combined with the mea-129 sured acoustic travel time to determine initial thickness. The acoustic travel time used 130 to determine the thickness of the sample was an average of several measurements across 131 the sample surface, with the travel time of each location being derived from the time dif-132 ference between the first and second acoustic echo. This procedure was performed be-133 fore and after each experiment for each sample. The thickness at pressure was assumed 134 to scale as $(V/V_0)^{1/3}$ where V₀ is the V₀ of bcc Fe-5Si, measured to be 23.34(4)³ and 135 V is determined from the fitted EoS for hcp-Fe-5Si. Thickness and travel times as a func-136 tion of pressure were then combined to determine Vp up to Mbar pressures. Ruby flu-137 orescence (Mao et al., 1986; Sokolova et al., 2013) and the Raman shift of the center of 138 the diamond culet (Akahama and Kawamura, 2006) were used to assess pressure. De-139 termination by ruby spectra and diamond edge Raman were within 2 GPa of each other 140 at all pressures where both were measured. Ruby fluorescence was measured before and 141 after every travel time measurement, and the reported pressure is an average of the two 142 values. For Run 1, the difference in pressures determined from ruby fluorescence before 143 and after measurement was less than 0.5 GPa, while for Run 2 it was less than 2 GPa. 144 All reported pressures are derived from ruby fluorescence measurements with the cal-145 ibration of Sokolova et al. (2013). 146

147 **3 Results**

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3.1 X-ray Diffraction

Two runs were performed in Ne at 300 K, one of which used Pt as pressure cali-149 brant to 41 GPa, and another run measured to 1.1 Mbar with Mo as the pressure cal-150 ibrant. Diffraction of the pressure calibrant was collected independently from that of the 151 sample by translating the cell a few microns from the sample position. As sample reflec-152 tions were also observable at the calibrant position, the presented volumes are unweighted 153 averages of all measured volumes at a given pressure step, and the reported pressure are 154 averages of Mo/Pt measurements taken at the same pressure step, drift between pres-155 sure measurements was typically less than 0.2 GPa. While measurement error of Mo vol-156 ume and the statistical errors in the calibrant EoS are small (<0.5%), the absolute er-157 ror at 300 K due to the intrinsic uncertainty in pressure calibration and pressure gra-158 dients is $\sim 2-3\%$. The bcc-hcp transition started at about 14 GPa and all bcc reflections 159 were absent by 21 GPa. 160

In another set of experiments, Fe-5Si was compressed in a laser-heated membrane DAC along two high-temperature isotherms, at about 1450 K and at about 2100 K. Temperatures varied by less than 100 K along each isotherm ($1\sigma = 30$ K at 1450 K and 1σ = 50 K at 2100 K). Temperatures were corrected downwards by around 3% following standard methods (Campbell et al., 2009) to account for axial T gradients (however this had a negligible effect on the fitted equation of state parameters). No phase other than
 hcp-Fe-5Si was observed at the HP-HT conditions of the present study, consistent with
 Tateno et al. (2015). Integrated diffraction patterns in Ne and at high temperatures are
 discussed in Supplementary Text S1 and shown in Supplementary Figures S1 and S2.

The ambient temperature component of the P-V-T thermal model employed in the present study consisted of either a 3rd Order Birch-Murnaghan (Equation 1, Birch, 1947) or Vinet (Equation 2, Vinet et al., 1989) EoS:

$$P_{300K}(V) = \frac{3}{2}K_0\left[\left(\frac{V_0}{V}\right)^{7/3} - \left(\frac{V_0}{V}\right)^{5/3}\right]\left\{1 + \frac{3}{4}\left(K' - 4\right)\left[\left(\frac{V_0}{V}\right)^{2/3} - 1\right]\right\}$$
(1)

$$P_{300K}(V) = 3K_0\left(\frac{1-\eta}{\eta^2}\right) exp\left[\frac{3}{2}\left(K'-1\right)\left(1-\eta\right)\right]$$
(2)

¹⁷³ Where V₀, K₀ and K' are the unit cell volume (Å³), bulk modulus (GPa) and $\frac{dK}{dP}$ ¹⁷⁴ at ambient conditions. In Equation 2 specifically, $\eta = \left(\frac{V}{V_0}\right)^{(1/3)}$

¹⁷⁵ The thermal parametrization is shown in Equation 3.

$$P(V,T) = P_{300K}(V) + P_{vib}(V,T) |_{300}^{T} + P_{el+anh}(V,T) |_{300}^{T}$$
(3)

In Equation 3,
$$P_{vib}(V,T)$$
 is given by:

$$P_{vib}\left(V,T\right) = \frac{9NR\gamma_{vib}}{V} \left[\frac{\theta_D}{8} + T\left(\frac{T}{\theta_D}\right)^3 \int_{300}^{\theta_D/T} \frac{x^3}{\exp\left(x\right) - 1} dx\right]$$
(4)

¹⁷⁷ Where γ_{vib} is the vibrational Grüneisen parameter, θ_D is the Debye temperature, ¹⁷⁸ N is the number of atoms per formula unit (N = 2 for hcp-structured Fe-alloys). R is ¹⁷⁹ the ideal gas constant, V is unit cell volume (in units of cm^3/mol) and T is the temper-¹⁸⁰ ature (in K). The volume dependence of the vibrational Grüneisen parameter and De-¹⁸¹ bye temperature are given by Equations 5 and 6 respectively.

$$\left(\frac{\gamma_{vib}}{\gamma_{vib,0}}\right) = \left(\frac{V}{V_0}\right)^q \tag{5}$$

$$\theta_D = \theta_{D,0} \exp\left[\left(\gamma_{vib,0} - \gamma_{vib}\right)/q\right] \tag{6}$$

In Equations 5 and 6 q characterizes the volume dependence of the vibrational contributions to thermal pressure. In the fitting process, $\theta_{D,0}$ was fixed to 422 K.

¹⁸⁴ Due to the similar T^2 dependence of P_{el} and P_{anh} , a single electronic pressure term ¹⁸⁵ was used for fitting in the present thermal model:

$$P_{el}(V,T) = \frac{\gamma_e}{V} \beta_0 \left(\frac{V}{V_0}\right)^k T^2 \tag{7}$$

¹⁸⁶ Where γ_e is the electronic Grüneisen parameter (here fixed to 2 after (Fei et al., ¹⁸⁷ 2016)), β_0 is the electronic heat capacity and k (fixed to 1.34) is an exponent which ¹⁸⁸ characterizes the volume dependence of the electronic contribution to thermal pressure.

All parameters of the presented P-V-T EoS except k, γ_e and θ_D were refined si-189 multaneously by an unweighted least-squares fit to the entire dataset. When the Debye 190 temperature is fixed in literature, it is generally fixed to anywhere between 417-422 K 191 for Fe and Fe-Si alloys (e.g. Dewaele et al., 2006; Fischer et al., 2012,1), however such 192 differences are negligible with regards to the resultant EoS. Additionally, most of the stud-193 ies which have constrained k either experimentally or by means of *ab initio* calculations 194 (Boness et al., 1986; Dewaele et al., 2006; Fei et al., 2016) generally report a k value of 195 ~ 1.34 . Varying k on the order of ~ 0.1 in the fitting process produces deviations in ex-196 trapolated densities at inner core conditions of less than 0.2%, far below the error bar 197 of the extrapolation ($\sim 1.5\%$). It is stressed that while electronic contributions to ther-198 mal pressure are relatively small (up to ~ 5 GPa at 2100 K), it was not possible to fit 199 the present dataset to a purely vibrational model of thermal pressure, as such a model 200 could not simultaneously reproduce the compressional behaviour of both the 1450 K and 201 2100 K isotherms. The fitted EoS parameters of this dataset are shown in Table 1. All 202 datapoints were within 2 GPa of the fitted EoS, irrespective of whether a 3BM or Vinet 203 EoS was used for the ambient temperature compression curve. 204

At ambient temperature, the present dataset is similar to recent measurements of 205 Fe10Ni5Si (Morrison et al., 2018) and those of hcp-Fe in He (Dewaele et al., 2006). Dis-206 crepancies between the presented ambient temperature EoS parameters are primarily due 207 to small systematic differences **observed** at low pressures (below ~ 60 GPa), as Fe, Fe10Ni5Si 208 and Fe5Si all have nearly identical volumes by 1 Mbar. Shown in Figure 1 as raw dat-209 apoints, the presented alloy exhibits a slightly higher volume than hcp-Fe, consistent with 210 literature on Fe-Si alloys (Fischer et al., 2014; Lin et al., 2003a; Tateno et al., 2015). While 211 the Earth's core is likely to be composed of an Fe-Ni alloy, the addition of Ni likely does 212 not significantly change the presented results and conclusions, due to the weak effect that 213 Ni has on Vp relative to Fe and Si (Antonangeli et al., 2010; Liu et al., 2016; Martorell 214 et al., 2013; Wakamatsu et al., 2018) and its weak effect on density (Morrison et al., 2018). 215 However it is stressed that at present there are few constraints on the thermoelastic be-216 haviour of Fe-Ni or Fe-Ni-Si alloys at simultaneous high P-T conditions. 217

It is immediately noticeable in Fig. 2 that the P-V-T EoS of Fe-5Si measured here 218 is similar to that of reported thermal parametrizations of hcp-Fe (Dewaele et al., 2006; 219 Fei et al., 2016), considering the small differences in 300 K EoS in each study. Unsur-220 prisingly, the fitted P-V-T EoS parameters of Fe-5Si are in good agreement with those 221 of Fei et al. (2016), which employs the same type of thermal model. It is remarkable that 222 the present XRD dataset composed purely of static compression data is capable of pro-223 ducing a P-V-T EoS which is directly comparable to those incorporating extensive parametriza-224 tion using shock compression data, ab initio calculations and/or NRIXS (e.g. Dewaele 225 et al., 2006; Fei et al., 2016). 226

As the direct measurement of thermal EoS are at the cutting edge of experimen-227 tal capabilities, it has been common in recent past to use *ab initio* parametrizations to 228 account for P_{el} (e.g. Dewaele et al., 2006). Inputs from calculations have been used to 229 constrain P_{el} and fit a purely vibrational model (Yamazaki et al., 2012), or to construct 230 thermal models using purely ambient temperature experimental data (e.g. Lin et al., 2003a; 231 Morrison et al., 2018). Choice of parametrization can change electronic thermal 232 233 pressures at core conditions by nearly 50%, and P_{el} itself is comparable in magnitude to P_{vib} at such conditions. Rescaling $P_{el} + P_{anh}$ of Dewaele et al. (2006) to the 234 formalism used in the present study, $\beta_0 \approx 1.7$. As the fitted β_0 of the present work is 235 $\sim 3.2(9)$ and that of Fei et al. (2016) is 3.91, it is likely for a dilute Fe-alloy (of realistic 236 composition for the Earth's inner core) that P_{el} is larger than that reported by Dewaele 237 et al. (2006). 238



Figure 1. Ambient temperature P-V relations of Fe-5Si (present work), pure Fe (Dewaele et al., 2006) and Fe10Ni5Si (Morrison et al., 2018). It is observed that Fe5Si and Fe10Ni5Si exhibit a small systematic difference up to 60 GPa, with their volumes becoming indistinguishable at higher pressures.



Figure 2. P-V-T dataset measured in the present study. Shown as lines are the results of the 3BM EoS fit described here, and the P-V-T relations reported by literature on hcp-Fe (Dewaele et al., 2006; Fei et al., 2016).

	Fe5Si	Fe5Si	Fe F16	Fe Y12	Fe9Si F14
Formalism	3BM	Vinet	3BM	3BM	3BM
V_0	22.524(62)	22.587(67)	22.428	22.15(5)	23.92(18)
K_0	172.4(6.0)	163.3(6.9)	172.7(1.4)	202(7)	129.1(1.4)
<i>K'</i>	4.64(14)	5.13(16)	4.79(0.05)	4.5(2)	5.29(8)
θ_D	422	422	422	1173(62)	420
γ_0	1.72(13)	1.73(13)	1.74	3.2(2)	1.14(14)
q	0.65(23)	0.67(23)	0.78	0.8(3)	1
β_0	3.20(85)	3.22(85)	3.91	ab initio	-
k	1.34	1.34	1.34	ab initio	-

Table 1. V_0 has units of Å³, K₀ has units of GPa, θ_D has units of K, β_0 has units of cm^3 $mol^{-1}JK^{-2}10^{-6}$, the other parameters are dimensionless. F16 denotes (Fei et al., 2016), Y12 denotes (Yamazaki et al., 2012), and F14 denotes (Fischer et al., 2014). Parameters in bold font are those which have been fixed during the fitting process. The two thermal models of Fe5Si are based on different choices of ambient temperature EoS formalism.



Figure 3. Travel Time measurement by Picosecond Acoustics at 47 GPa. The sharp peak associated with the 1st acoustic echo is clearly visible, followed by a background signal related to the generation of surface waves and Brillouin scattering in the Ne PTM (filtered in the red curve). The difference between the travel time of filtered and unfiltered data changes by a maximum of 2 ps, or an error of 0.1-0.5% of Vp depending on sample thickness. Inset is the sample chamber of Run 2 at ~100 GPa.

3.2 Picosecond Acoustics

Travel times were measured as a function of pressure (Figure 3 and Figure 4) in 240 two independent runs performed on samples of Fe-5Si alloy of different initial thickness. 241 Errors in Vp were about 2% up to 60 GPa, and about 3% by 115 GPa. Shown in Fig-242 ure 3 are background-subtracted time domain signals at 47 GPa before and after data 243 treatment to extract Brillouin oscillations. Larger errors at high pressures are due to an 244 increase in diffuse background caused by the progressive depolarization of the pump and 245 probe beams resulting from increased defect scattering, stress gradients across the di-246 amond anvil, and the cupping of the diamond culet- classical issues encountered by op-247 tical measurements at Mbar conditions (Merkel et al., 1999). Uncertainties of sample thick-248 ness, especially at high pressure, has only a small effect on the travel time compared to 249 the change in velocity. As a matter of fact, by 1 Mbar, the thickness has changed by $\sim 10\%$ 250 relative to ambient pressure, while the acoustic travel time is typically 50% its original 251 value. Sample tilt within the sample chamber has a negligible effect on measured travel 252 times due to the instrumental configuration and focusing strategy employed. In Run 1, 253 it was seen that there were some residual stresses in the sample which induced local vari-254 ation in measured travel times of about 2-3% in the bcc phase, and so for Run 2, the sam-255 ple was additionally annealed under vacuum at ~ 400 K for 12 h. This procedure effec-256 tively reduced scatter in measured travel times to less than 1% in the bcc phase. Nor-257 malized travel times for bcc and hcp Fe5Si are shown in Figure 4. 258

In Run 2, the initial travel time and travel time of the recovered samples are within error bar of each other, indicating negligible plastic deformation of the sample when compressed in Ne PTM up to 1.1 Mbar. Provided the volume decrease at the bcc-hcp transition is accounted for, the measured Vp varies by less than 0.5% depending on the choice of Fe or Fe-Si EoS, well within error at all pressures. The bcc-hcp transition is observed by PA to occur over a pressure range in close agreement with XRD measurements.



Figure 4. Normalized travel times of Fe-5Si from \sim 5-115 GPa. Observed scatter in the bccphase disappears above the bcc-hcp transition. The onset of the bcc-hcp transition measured by PA is in good agreement with that observed by XRD.

Following the transition region determined by both XRD and PA, a sharp rise in 265 Vp is observed. While bcc peaks are absent from XRD by ~ 23 GPa, nonlinear varia-266 tions in acoustic velocity have been observed up to ~ 30 GPa where the variation becomes 267 regular. This is attributed to the slight development of preferred orientations in hcp-Fe-268 5Si at the end of the phase transition - XRD highlights a moderate change in the inten-269 sity ratio of the (002)/(100) peaks between 20 to 30 GPa, while at higher pressures this 270 ratio does not vary any more. Due to the difference in measurement geometry between 271 PA and XRD, the intensity reduction of the (002) peak observed by XRD reflects a larger 272 contribution of the c-axis to the measured travel time and hence a small increase in Vp 273 (Antonangeli et al., 2006). As linearity in the Vp - ρ relation for hcp-Fe-5Si was observed 274 by ~ 30 GPa, the following discussion will be limited to data measured at and above this 275 pressure. 276

277 4 Discussion

4.1 Density of Fe-Si Alloys at the P-T Conditions of Earth's Inner Core



Figure 5. Density of Fe (Fei et al., 2016), Fe-5Si, Fe-10Ni-5Si (Morrison et al., 2018), Fe-3.3Si and Fe-6.7Si (Martorell et al., 2016) at the pressures of the IC. Fe-5Si is within error bar of PREM for all reasonable core temperatures (5000-7000 K). Shaded region denotes expected error bar of PREM (~2%, Masters and Gubbins, 2003). Extrapolations of Fe-5Si to core temperatures and pressures have an error bar of ~1.5%.

Extrapolated to IC pressures and temperatures, the density of Fe-5Si is within er-279 ror of PREM ($\sim 2\%$, Masters and Gubbins, 2003) for all reasonable core temperatures, 280 with the best match for T = 5500 K (Fig. 5). As the compressional behaviour of Fe-5Si 281 and Fe are similar over the wide range of P-T conditions measured in this study, the dom-282 inant mechanism for density reduction even at core conditions is simply the difference 283 in atomic mass between Fe and Si. Indeed, the density reduction between hcp-Fe and 284 hcp-Fe-5Si is similar both at 300 K and at temperatures exceeding 5000 K. While the 285 recent measurements of Fe-10Ni-5Si exhibit less compressible behaviour at ambient tem-286 perature, extrapolations of that EoS to core conditions results in densities very similar to those presented here. It is noted however that updating the Fe-10Ni-5Si thermal model 288 with P_{el} determined here, there is a ~0.1 g cm⁻³ (0.8%) decrease in density of the al-289 loy at 5500 K. This highlights the importance of accurate, high quality volume measure-290 ments at simultaneous high temperatures and high pressures, in order to constrain such 291 effects at the conditions of the Earth's inner core. 292

Our results are well compatible with the most recent ab initio calculations on Fe-Si alloys (Li et al., 2018; Martorell et al., 2016), the 5500 K isotherm extrapolates to the midpoint between calculations of Fe-3.3Si and Fe-6.7Si at the same temperatures. The calculated thermodynamic grüneisen parameter of this alloy is between 1.5 and 1.6 at the ICB for temperatures of between 4500 K and 6500 K, consistent with previous results on hcp-Fe (Dewaele et al., 2006; Fei et al., 2016; Vočadlo et al., 2003). This combined with the observation that Si alloying does not strongly affect the melting curve

- of iron (e.g. Fischer et al., 2013; Morard et al., 2017) indicates that Si-alloying would not
- ³⁰¹ likely have a strong effect on the thermal profile of the inner core.

4.2 Sound Velocities of Fe-Si Alloys at Core Conditions



Figure 6. Compressional Sound Velocity vs. Density for Fe and Fe-Si alloys in the hcp structure. Inset: the slope of linear fits to Velocity vs. Density $(dVp/d\rho)$ vs. Si content for the here-discussed Fe-xSi datasets (Antonangeli et al., 2018; Mao et al., 2012; Martorell et al., 2016; Sakairi et al., 2018; Tsuchiya and Fujibuchi, 2009). Downward arrows indicate the maximum estimated magnitude of anharmonic effects at 5500 K and 6500 K (after Martorell et al., 2016).

³⁰³ Vp measurements show a clear linear trend for the entire density range studied, ³⁰⁴ with the fitted parameters $Vp(km/s) = 1.135(20) * \rho(g/cm^3) - 2.27(19)$ as shown in ³⁰⁵ Figure 6. dVp/d ρ of Fe-5Si is reduced with respect to hcp-Fe Antonangeli and Ohtani ³⁰⁶ (2015), although the effect is not as large as that reported by previous IXS measurements ³⁰⁷ on samples with higher Si content (Antonangeli et al., 2018; Mao et al., 2012; Sakairi ³⁰⁸ et al., 2018), as shown in the inset of Figure 6.

Our measurements extrapolate at inner core densities somewhat higher than re-309 cent measurements by IXS on more Si-rich samples (Antonangeli et al., 2018; Mao et al., 310 2012; Sakairi et al., 2018), but are in very good agreement with the result of athermal 311 ab initio calculations on Fe-3.3Si and Fe-6.7Si (Martorell et al., 2016). We also observe 312 agreement between Vp - ρ relations presented here and those reported in Tsuchiya 313 and Fujibuchi (2009). We stress however, that the agreement between our 314 experiments and calculations worsen when looking at velocities versus pres-315 sure, in particular with respect to Tsuchiya and Fujibuchi (2009). In that study, 316 at 40 GPa a linear interpolation of Fe3.3Si and Fe6.7Si exhibits a bulk modulus which 317 is 28% higher than one derived from the present EoS, and the reported density is 4.5%318 higher. 319

Fig. 6 shows that IXS results are generally parallel to each other (Antonangeli et al., 320 2018; Mao et al., 2012; Sakairi et al., 2018), but in disagreement with the present work. 321 It is evident based on the combined results of (Antonangeli et al., 2018; Mao et al., 2012; 322 Sakairi et al., 2018; Tsuchiya and Fujibuchi, 2009), and the results of the present study 323 that $dVp/d\rho$ decreases with increasing Si content (shown inset in Fig. 6). While a lin-324 ear decrease in $dVp/d\rho$ with Si content can rationalize a significant amount of the dif-325 ference between PA and IXS, there are also systematic differences due to the different 326 measurement geometries of the two techniques. PA measures acoustic travel times along 327 the compression axis of the DAC, and as the sample is expected to develop texture, PA 328 will preferentially sample the c-axis of the alloy. By contrast, IXS measures phonon dis-329 persions perpendicular to the DAC compression axis, and as such preferentially samples 330 the basal plane of the Fe-alloy upon development of texture. In this way, textural effects 331 bias PA and IXS measurements in opposite directions. Additionally, IXS measurements 332 require larger sample volumes, and so are often measured in a solid PTM (Sakairi et al.. 333 2018; Sakamaki et al., 2016) or no PTM at all (Antonangeli et al., 2018). We note how-334 ever, that the difference in extrapolations of PA and IXS measurements here observed 335 for the Fe-Si system is much more significant than for the case of hcp-Fe (Antonangeli 336 and Ohtani, 2015, references therein). 337



Figure 7. Vp vs Density measured for Fe-Si alloys in the hcp structure by different experimental techniques (Antonangeli et al., 2018; Lin et al., 2003b; Mao et al., 2012; Sakairi et al., 2018).

It is important to minimize texture by performing the experiments under quasi-338 hydrostatic conditions with noble gas media. While there is still some development of 330 texture under compression in Ne, it is significantly weaker than the texture observed to 340 occur in IXS experiments at comparable pressure conditions (Supplemental Figure S1.S3, 341 Antonangeli et al., 2018; Sakairi et al., 2018). Figure 7 shows the individual datapoints 342 of various studies in the hcp-Fe-Si system. Ultimately, the effect of preferred orientation 343 is a relatively small effect, but is shown to systematically bias extrapolations (upward 344 in the case of NRIXS) and downwards (in the case of IXS) in Vp at core densities, which 345 strongly hinders compositional modelling at core conditions. It can also be observed in 346

Figure 7 that the present work has significantly improved data coverage to typical IXS or NRIXS experiments, and is measured over an extremely wide density range, under quasihydrostatic conditions, allowing for more robust extrapolations to core densities.

Having experimentally determined both the P-V-T EoS and Vp of Fe5Si, it is pos-350 sible to combine these quantities to derive shear velocities. Figure 8 shows Vs plotted 351 against density. Considering the model-dependence and large error bars of the present 352 shear velocities $(\sim 8-10\%)$, it is not unambiguously possible to determine whether these 353 vary linearly or sublinearly with density (solid black lines). However, whether the dataset 354 is fit with a linear relation or power-law relation as suggested by (Mao et al., 2012), Vs 355 of Fe5Si is within error of *ab initio* calculations and **seems** higher than hcp-Fe (Anto-356 nangeli and Ohtani, 2015) at inner core densities. Recent IXS measurements on Fe9Si 357 report a sublinear density dependence of Vs (Antonangeli et al., 2018), indicating that 358 this may be a systematic effect of Si alloying. Shear moduli calculated by Tsuchiya 359 and Fujibuchi (2009) are very similar to the present study at 40 and 120 GPa, 360 and the discrepancies observed in Figure 8 arise primarily from the aforemen-361 tioned differences in densities. 362



Figure 8. Derived Vs from the present study and selected literature (Antonangeli et al., 2018; Antonangeli and Ohtani, 2015; Li et al., 2018; Martorell et al., 2016; Tsuchiya and Fujibuchi, 2009). It is observed that the derived Vs of Fe5Si are higher than hcp-Fe at core densities irrespectively of the linear or sub-linear extrapolation (solid black lines). Downward arrows indicate the maximum estimated magnitude of anharmonic effects at 5500 K and 6500 K (after Martorell et al., 2016).

To meaningfully compare obtained Vp and Vs with PREM, high temperature effects have to be accounted for. At a constant density of 13 g/cm³, the T corrections on Vp after experiments by Sakamaki et al. (2016) (on hcp-Fe) and Sakairi et al. (2018) (on Fe-6Si) yield Vp reductions of ~0.09 m s⁻¹ K⁻¹, with almost no difference between Fe and Fe-6Si. Alternatively, by converting the constant pressure simulations of Martorell et al. (2016) to a constant density of 13 g/cm³, it is possible to estimate the mag-

nitude of anharmonic temperature effects to be $\sim 0.11 \text{ m s}^{-1} \text{ K}^{-1}$ for Fe and Fe3.3Si, and $\sim 0.05 \text{ m s}^{-1} \text{ K}^{-1}$ for Fe-6.7Si.

Regardless of the magnitude of anharmonic corrections for Vp at high tempera-371 ture (Li et al., 2018; Martorell et al., 2016; Sakairi et al., 2018; Sakamaki et al., 2016). 372 even when assuming the largest proposed effects (Martorell et al., 2016), PREM 373 Vp is expected to be matched by an Fe-Si alloy **only when** containing <2wt% Si for 374 T = 6500 K and <1wt% Si for T = 5500 K. The largest anharmonic effects reported in 375 recent literature come from *ab initio* calculations (Martorell et al., 2016), but a more re-376 cent ab initio study by the same group using larger simulation cells (Li et al., 2018) 377 supports significantly reduced anharmonic effects compared to Martorell et al. (2016), 378 such that the magnitude of anharmonic effects are more in line with those observed by 379 IXS (Sakairi et al., 2018; Sakamaki et al., 2016). As such, constraints imposed by com-380 pressional sound velocity (at maximum < 2wt% Si) are incompatible with constraints im-381 posed by density (an Si alloy containing 5wt% Si has a density matching PREM at re-382 alistic core conditions). As a result, Si likely cannot be the sole light element in the Earths 383 core. 384

This mismatch between mineral-physics measurements and seismolog-385 ical observables is even worse when considering Vs. Assuming a power-law ex-386 trapolation of Vs to core densities (representing a lower bound to the extrapolated Vs) 387 and the largest possible magnitude of anharmonic effects ($\sim 0.34 \text{ m s}^{-1} \text{ K}^{-1}$ (Martorell 388 et al., 2016)), the Vs of an Fe-Si alloy of any composition cannot match PREM for tem-389 peratures below 6500 K. Considering that such high core temperatures are likely to be 390 unrealistic for a variety of geophysical and mineralogical reasons, this emphasizes the un-391 suitability of Si as the primary light element in the Earth's inner core. This study high-392 lights the important point that density or velocity information alone can only be used 393 to exclude possible core compositions, and must be coupled together in order to develop 394 accurate models of the Earths interior. 395

5 Conclusions

The combination of PVT X-ray diffraction measurements and 300 K picosecond 307 acoustics Vp measurements under quasi-hydrostatic conditions of Fe-5wt% Si show that 398 while an alloy with 5wt% Si can potentially match the PREM density of the Earths in-399 ner core, this quantity of silicon is incompatible with seismological observations of Vp 400 and Vs, even when accounting for anharmonic effects at high temperature. Thus, our 401 paper supports the conclusion suggested by Martorell et al. (2016) that Si can-402 not be considered as the only light element in the Earth's inner core in the absence of 403 other mechanisms, such as partial melting (Vočadlo, 2007) which significantly reduces 404 density as well as sound velocities of Fe-Si alloys at core conditions. 405

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415 S3.

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