

Numerical modelling of the surge-derived pyroclastic ow of 25 June 1997 at Soufriere Hills Volcano, Montserrat

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▶ To cite this version:

Valentin Gueugneau, Karim Kelfoun, Timothy H. Druitt. Numerical modelling of the surge-derived pyroclastic ow of 25 June 1997 at Soufriere Hills Volcano, Montserrat. European Geosciences Union - General Assembly, Apr 2018, Vienne, Austria. hal-02077439

HAL Id: hal-02077439 https://hal.science/hal-02077439

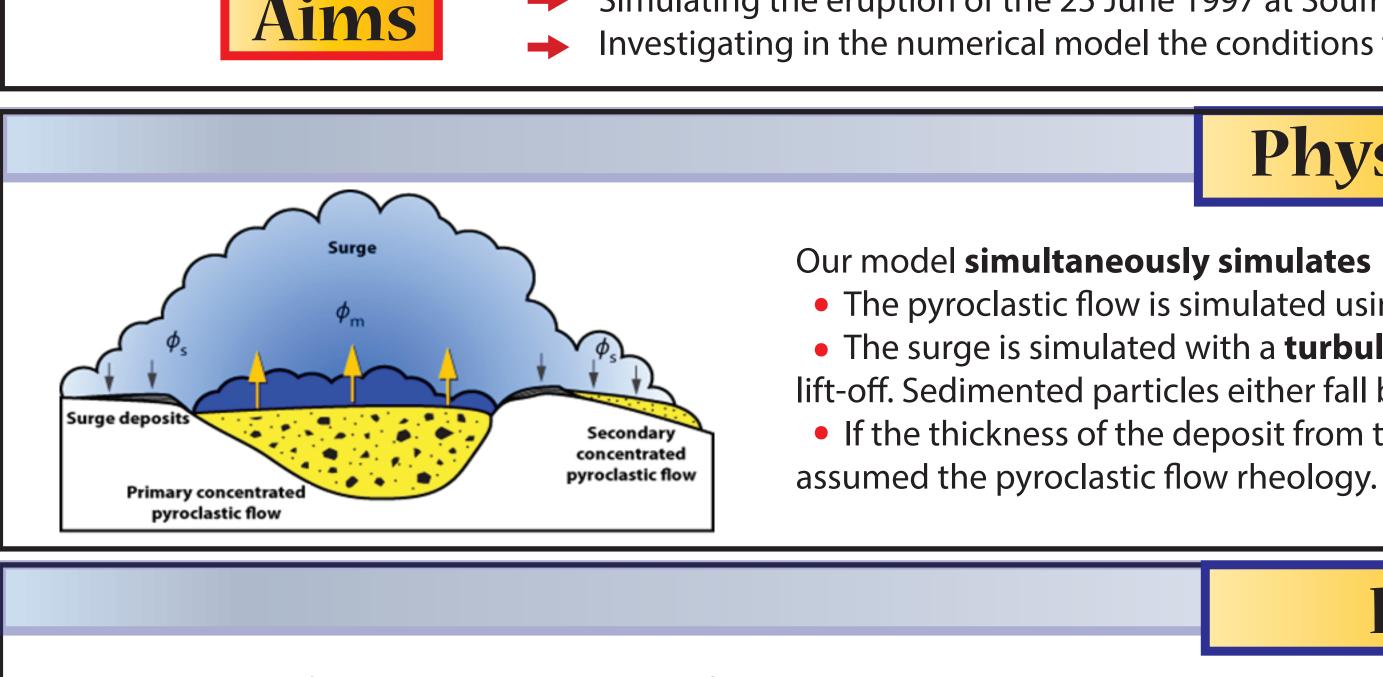
Submitted on 22 Mar 2019

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Numerical modelling of the surge-derived pyroclastic flow of 25 June **1997 at Soufriere Hills Volcano, Montserrat.**

Pyroclastic density currents (PDCs) are volcanic flows usually composed of 2 distinct parts: the pyroclastic flow and the overriding ash-cloud surge. Deposits from ash-cloud surges associated with dome-collapse block-and-ash flows (BAFs) can be immediately remobilized to form surge-derived pyroclastic flows (SDPF). SDPFs, such as those generated on 25 June 1997 at Soufrière Hills, Montserrat, are hazardous because they can follow drainages different from those exploited by the BAF. Better understaning of their mechanism of formation is essential for risk mitigation. SDPFs have never been reproduced numerically in the past. Simulating the eruption of the 25 June 1997 at Soufrière Hills, Montserrat, because the SDPFs have been well described. Aims Investigating in the numerical model the conditions that favor the generation of SDPFs during dome-collapse events.



Deposit map from the 25 June 1997 Soufière Hills dome collaspe (Montserrat) 62°12'W Block-and-ash flow deposit Ash-cloud surge deposit Surge-derived pyroclastic flow deposit MVO Farrell's First major St George's Hi Gages Dom

+ The simulation highlights a mechanism of surge derived-pyroclastic flow formation, highly dependent of the topography :

62°11′W

62°10'W

62°12′W

First, the major bend in the Mosquito Ghaut valley allows the surge detachment and a strong sedimentation on the Farrell's plain.

Second, the accumulation of surge deposits on steep slopes generates destabilizations related to each rheology. These small destabilizations in deposits transform into small surge-derived pyroclastic flows.

Third, due to the valley shape, the SDPFs amalgamate into a single flow in the Belham valley. This SDPF travels 3 km in the opposite direction of the main BAF and becomes a significant hazard in its own right.

References

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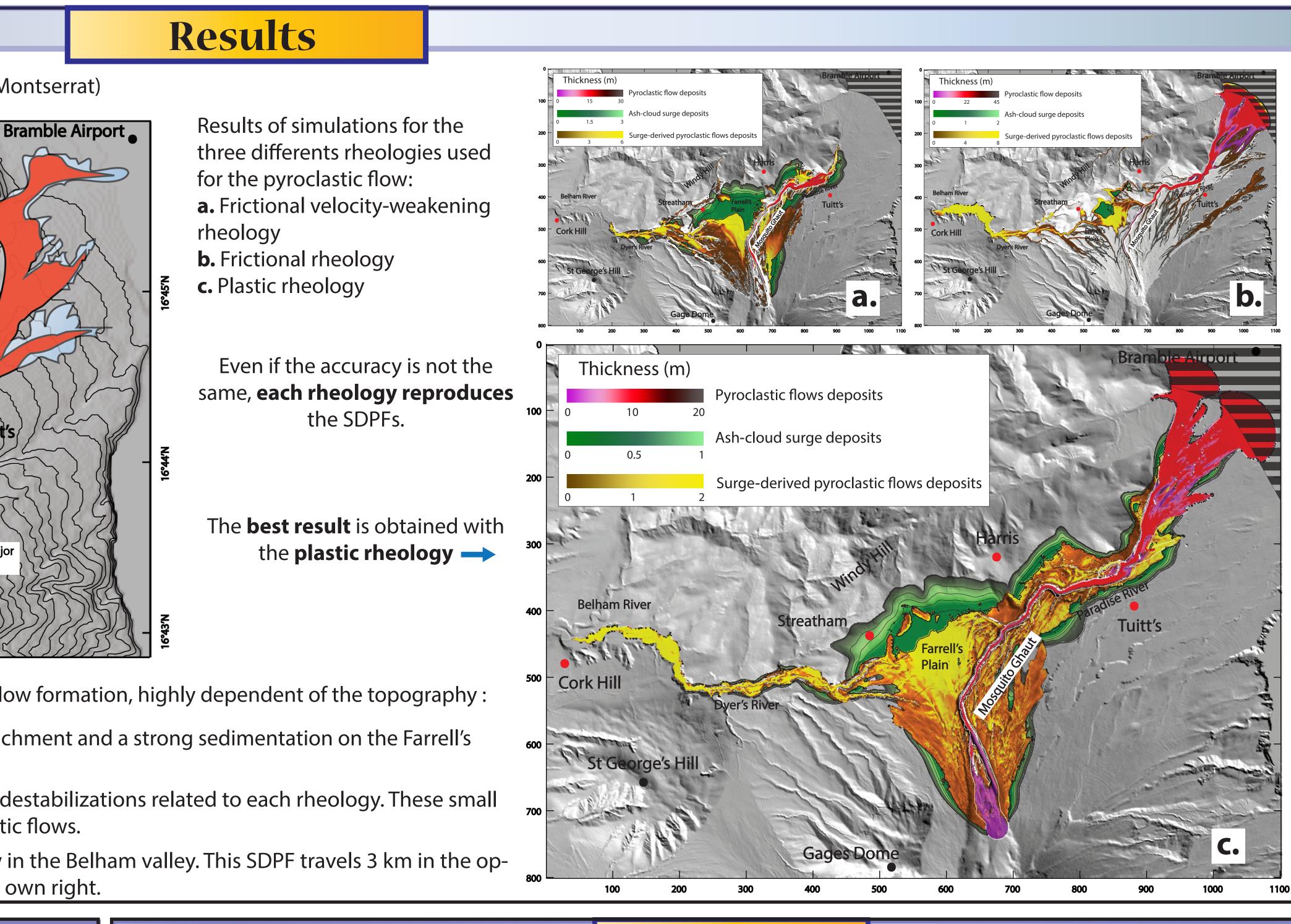
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Introduction

Physical model

Our model **simultaneously simulates** the two parts of the PDCs: the pyroclastic flow, and the overriding ash-cloud surge: • The pyroclastic flow is simulated using three main rheologies: the plastic rheology, the frictional velocity-weaking rheology, and frictional rheology. • The surge is simulated with a turbulent gas rheology. It gains mass by particle entrainment from the underriding flow, and loses mass by sedimentation, and lift-off. Sedimented particles either fall back into the pyroclastic flow, or form a deposit on the land surface. • If the thickness of the deposit from the surge exceeds the critical threshold for the material rheology, the deposit immediately remobilises to form a SDPF with



Conclusion

First numerical reproduction of surge-derived pyroclastic flows. Validation of the physical model. Highlights the mechanism of formation of SDPFs : a cycle of accumulation-remobilisation of surge deposits on a steep slope, channelised into one valley, forming a surge-derived pyroclastic flow. Highly dependent on the topography. The plastic rheology for the pyroclastic flow seems to be the more accurate.

Numerical modelling

$$\begin{aligned} \frac{\partial h_{d}}{\partial t} + \frac{\partial}{\partial x} (h_{d}u_{x}) + \frac{\partial}{\partial y} (h_{d}u_{y}) &= \frac{\phi_{s} - \phi_{m}}{\rho_{d}} \\ \frac{\partial}{\partial t} (h_{d}u_{x}) + \frac{\partial}{\partial x} (h_{d}u_{x}^{2}) + \frac{\partial}{\partial y} (h_{d}u_{x}u_{y}) \\ &= gh_{d} \sin \alpha_{x} - \frac{1}{2} \frac{\partial}{\partial x} (gh_{d}^{2} \cos \alpha_{x}) + \frac{T_{x}}{\rho_{d}} + \frac{\phi_{s}}{\rho_{d}} v_{x} - \frac{\phi_{m}}{\rho_{d}} u_{x} \\ \frac{\partial}{\partial t} (h_{d}u_{y}) + \frac{\partial}{\partial x} (h_{d}u_{x}u_{y}) + \frac{\partial}{\partial y} (h_{d}u_{y}^{2}) \\ &= gh_{d} \sin \alpha_{y} - \frac{1}{2} \frac{\partial}{\partial y} (gh_{d}^{2} \cos \alpha_{y}) + \frac{T_{y}}{\rho_{d}} + \frac{\phi_{s}}{\rho_{d}} v_{y} - \frac{\phi_{m}}{\rho_{d}} u_{y} \end{aligned}$$

$$\frac{\partial h_{d}}{\partial t} + \frac{\partial}{\partial x}(h_{d}u_{x}) + \frac{\partial}{\partial y}(h_{d}u_{y}) = \frac{\phi_{s} - \phi_{m}}{\rho_{d}}$$

$$\frac{\partial}{\partial t}(h_{d}u_{x}) + \frac{\partial}{\partial x}(h_{d}u_{x}^{2}) + \frac{\partial}{\partial y}(h_{d}u_{x}u_{y})$$

$$= gh_{d}\sin\alpha_{x} - \frac{1}{2}\frac{\partial}{\partial x}(gh_{d}^{2}\cos\alpha_{x}) + \frac{T_{x}}{\rho_{d}} + \frac{\phi_{s}}{\rho_{d}}v_{x} - \frac{\phi_{m}}{\rho_{d}}u_{x}$$

$$\frac{\partial}{\partial t}(h_{d}u_{y}) + \frac{\partial}{\partial x}(h_{d}u_{x}u_{y}) + \frac{\partial}{\partial y}(h_{d}u_{y}^{2})$$

$$= gh_{d}\sin\alpha_{y} - \frac{1}{2}\frac{\partial}{\partial x}(gh_{d}^{2}\cos\alpha_{y}) + \frac{T_{y}}{\rho_{d}} + \frac{\phi_{s}}{\rho_{d}}v_{y} - \frac{\phi_{m}}{\rho_{d}}u_{y}$$

$$\frac{\partial h_{s}}{\partial t} + \frac{\partial}{\partial x}(h_{s}v_{x}) + \frac{\partial}{\partial y}(h_{s}v_{y}) = \frac{\phi_{m}}{\rho_{m}}$$

$$\frac{\partial \rho_{s}h_{s}}{\partial t} + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{y}) = \phi_{m} - \phi_{s}$$

$$\frac{\partial}{\partial t}(\rho_{s}h_{s}v_{x}) + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}^{2}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{x}v_{y}) = (\rho_{s} - \rho_{a})gh_{s}\sin\beta_{x}$$

$$-\frac{1}{2\partial x}((\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{x}) + \mathbf{R}_{x} + \phi_{m}u_{x} - \phi_{s}v_{x}$$

$$\frac{\partial}{\partial t}(\rho_{s}h_{s}v_{y}) + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}v_{y}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{y}^{2}) = (\rho_{s} - \rho_{a})gh_{s}\sin\beta_{y}$$

$$-\frac{1}{2\partial y}((\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y}) + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$
we flow are coupled through two empirical mass flux laws:
From the flow to the surge : $\phi_{m} = a_{3}.u^{3}$
From the surge to the flow : $\phi_{s} = (\rho_{s} - \rho_{a}).w_{sed}$
The resistive stresses are adpatable to each rheology :

$$\frac{\mathbf{u}}{\|\mathbf{u}\|} - c_{1}\rho_{d}\|\mathbf{u}\|\mathbf{u} \quad \text{or} \quad \mathbf{T} = h_{d}\rho_{d}\left(g\cos\alpha + \frac{\mathbf{u}^{2}}{r}\right)\tan\phi_{bed}\frac{\mathbf{u}}{\|\mathbf{u}\|} + c_{1}\rho_{d}\|\mathbf{u}\|\mathbf{u}$$

$$\frac{\partial h_{s}}{\partial t} + \frac{\partial}{\partial x}(h_{s}v_{x}) + \frac{\partial}{\partial y}(h_{s}v_{y}) = \frac{\phi_{m}}{\rho_{m}}$$

$$\frac{\partial \rho_{s}h_{s}}{\partial t} + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{y}) = \phi_{m} - \phi_{s}$$

$$\frac{\partial}{\partial t}(\rho_{s}h_{s}v_{x}) + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}^{2}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{x}v_{y}) = (\rho_{s} - \rho_{a})gh_{s}\sin\beta_{x}$$

$$-\frac{1}{2\partial x}((\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{x}) + \mathbf{R}_{x} + \phi_{m}u_{x} - \phi_{s}v_{x}$$

$$\frac{\partial}{\partial t}(\rho_{s}h_{s}v_{y}) + \frac{\partial}{\partial x}(\rho_{s}h_{s}v_{x}v_{y}) + \frac{\partial}{\partial y}(\rho_{s}h_{s}v_{y}^{2}) = (\rho_{s} - \rho_{a})gh_{s}\sin\beta_{y}$$

$$-\frac{1}{2\partial y}((\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y}) + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$
The two flows are coupled through two empirical mass flux laws:

$$-\frac{1}{2\partial y}((\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y}) + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$
The two flows are coupled through two empirical mass flux laws:

$$-\frac{1}{2\partial y}(\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$
The two flows are coupled through two empirical mass flux laws:

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$
The two flows are coupled through two empirical mass flux laws:

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{a})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{y} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

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$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{s})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{s} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{s})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{s} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{s})gh_{s}^{2}\cos\beta_{y} + \mathbf{R}_{s} + \phi_{m}u_{y} - \phi_{s}v_{y}$$

$$-\frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} - \rho_{s})gh_{s}^{2}\cos\beta_{y} + \frac{1}{2}\frac{\partial}{\partial y}(\rho_{s} -$$

plastic $T_0 = cst$

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The simulation of the pyroclastic flows (both primary and secondary) is based on the resolution of mass and momentum balance equations (Kelfoun, 2017; Kelfoun et al., 2017):

The surge is also based on the resolution of mass and momentum balance equations, and we add a density balance equation:

> frictional frictional velocity-weakening $\tan \varphi_{bed} = cst$ $\tan \varphi_{bed} = \frac{1}{\sqrt{1}}$ $(1 + \|\mathbf{u}\| / U_{\mu})$

Contact

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