

# Perspectives of seismic imaging using FWI with reciprocity misfit functional

Florian Faucher

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## Perspectives of seismic imaging using FWI with reciprocity misfit functional

### Florian Faucher<sup>1</sup>,

Giovanni Alessandrini<sup>2</sup>, Hélène Barucq<sup>1</sup>, Maarten V. de Hoop<sup>3</sup>, Romina Gaburro<sup>4</sup> and Eva Sincich<sup>2</sup>.

### Journées Ondes Sud-Ouest, CEA Cesta, France Marth 12<sup>th</sup>-14<sup>th</sup>, 2019





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### Overview



- Introduction
- Time-Harmonic Inverse Problem, FWI
- Reconstruction procedure using dual-sensors data
- Numerical experiments
  - Comparison of misfit functions
  - ullet Changing the numerical acquisition with  $\mathcal{J}_{\mathcal{C}}$
- Conclusion

Intro

### Introduction

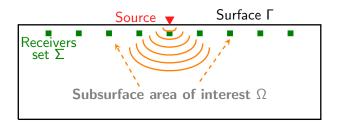


Experiments

### Seismic inverse problem



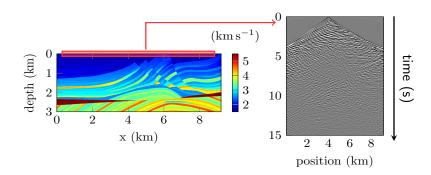
Reconstruction of subsurface Earth properties from seismic campaign: collection of wave propagation data at the surface.



### Seismic data

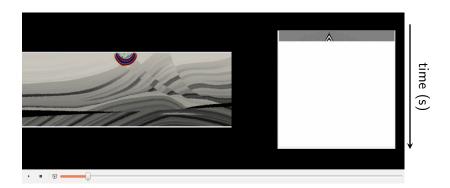


We work with back-scattered partial data from one-side illumination on large domain.



Intro

We work with back-scattered partial data from one-side illumination on large domain.

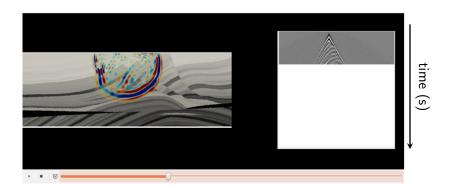


### Seismic data

Intro



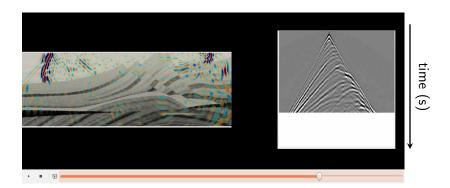
We work with back-scattered partial data from one-side illumination on large domain.



### Seismic data

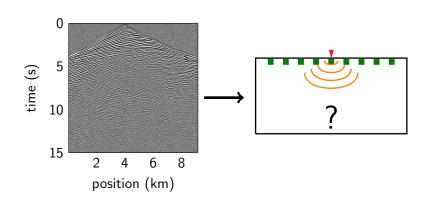


We work with back-scattered partial data from one-side illumination on large domain.



Intro

**Inverse problem**: from seismic traces to subsurface?



nonlinear, ill-posed inverse problem.

2 Time-Harmonic Inverse Problem, FWI

### Time-harmonic wave equation



We consider propagation in acoustic media, given by the Euler's equations, heterogeneous medium parameters  $\kappa$  and  $\rho$ :

$$\begin{cases} -i\omega \rho(\mathbf{x})\mathbf{v}(\mathbf{x}) = -\nabla p(\mathbf{x}), \\ -i\omega p(\mathbf{x}) = -\kappa(\mathbf{x})\nabla \cdot \mathbf{v}(\mathbf{x}) + f(\mathbf{x}). \end{cases}$$

p: scalar pressure field,  $\kappa$ : bulk modulus,

 $\mathbf{v}$ : vectorial velocity field,  $\rho$ : density,

f: source term.

 $\omega$ : angular frequency.



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p: scalar pressure field,  $\kappa$ : bulk modulus,

 $\mathbf{v}$ : vectorial velocity field,  $\rho$ : density,

 $\omega$ : angular frequency. f: source term.

The system reduces to the Helmholtz equation when  $\rho$  is constant,

$$(-\omega^2 c(\mathbf{x})^{-2} - \Delta) p(\mathbf{x}) = 0,$$

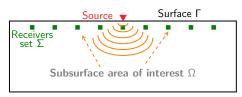
with 
$$c(\mathbf{x}) = \sqrt{\kappa(\mathbf{x})\rho(\mathbf{x})^{-1}}$$
.

### **Dual-sensors** devices



The inverse problem aims the recovery of the subsurface medium parameters from surface measurements of pressure **and** normal (vertical) velocity:

$$\mathcal{F}: m = (\kappa, \rho) \rightarrow \{\mathcal{F}_p ; \mathcal{F}_v\} = \{p(\mathbf{x}_1), p(\mathbf{x}_2), \dots, p(\mathbf{x}_{n_{rcv}}); \\ v_n(\mathbf{x}_1), v_n(\mathbf{x}_2), \dots, v_n(\mathbf{x}_{n_{rcv}})\}.$$





D. Carlson, N. D. Whitmore et al.

Increased resolution of seismic data from a dual-sensor streamer cable – Imaging of primaries and multiples using a dual-sensor towed streamer SEG. 2007 - 2010



CGG & Lundun Norway (2017-2018)

 $Top Se is \ acquisition \ (\verb|www.cgg.com/en/What-We-Do/Offshore/Products-and-Solutions/Top Se is) \\$ 



FWI provides a **quantitative reconstruction** of the subsurface parameters by solving a minimization problem,

$$\min_{m \in \mathcal{M}} \quad \mathcal{J}(m) = \frac{1}{2} \|\mathcal{F}(m) - d\|^2.$$

- ▶ d are the observed data,
- $\triangleright$   $\mathcal{F}(m)$  represents the simulation using an initial model m:



P. Lailly

The seismic inverse problem as a sequence of before stack migrations Conference on Inverse Scattering: Theory and Application, SIAM, 1983



A. Tarantola

Inversion of seismic reflection data in the acoustic approximation Geophysics, 1984



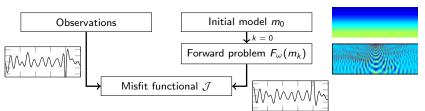
A Tarantola

Inversion of travel times and seismic waveforms Seismic tomography, 1987

### FWI, iterative minimization

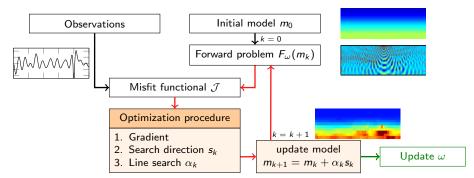
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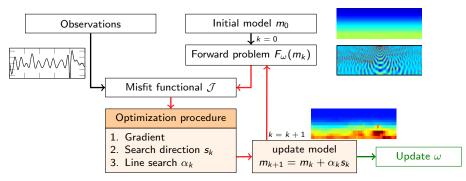
### FWI, iterative minimization





### FWI, iterative minimization

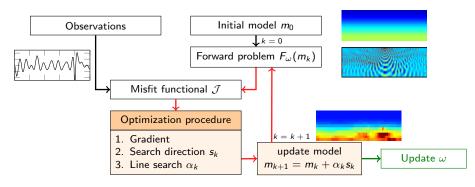




### Numerical methods

- ▶ Adjoint-method for the gradient computation, L-BFGS method,
- forward problem resolution with Discontinuous Galerkin methods,
- parallel computation, HPC, large-scale optimization,
- Rk: the code also works for elastic anisotropy and viscous media.

## Inria-



- $ightharpoonup > 10^5$ : unknowns per physical parameter,
- $> 10^6$ : matrix size for discretization,
- we also study stability and convergence of the algorithm . . .

Plan



Reconstruction procedure using dual-sensors data



The appropriate misfit functional to minimize with pressure and vertical velocity measurements.

► Compare the pressure and velocity fields separately:

$$\mathcal{J}_{L2} = \sum_{\text{source}} \frac{1}{2} \|\mathcal{F}_{p}^{(s)} - d_{p}^{(s)}\|^{2} + \frac{1}{2} \|\mathcal{F}_{v}^{(s)} - d_{v}^{(s)}\|^{2}.$$



The appropriate misfit functional to minimize with pressure and vertical velocity measurements.

Compare the pressure and velocity fields separately:

$$\mathcal{J}_{L2} = \sum_{source} \frac{1}{2} \|\mathcal{F}_{p}^{(s)} - d_{p}^{(s)}\|^{2} + \frac{1}{2} \|\mathcal{F}_{v}^{(s)} - d_{v}^{(s)}\|^{2}.$$

Compare the fields multiplication for all combinations:

$$\mathcal{J}_{\mathcal{G}} = \frac{1}{2} \sum_{s_1} \sum_{s_2} \|d_v^{(s_1)T} \mathcal{F}_p^{(s_2)} - d_p^{(s_1)T} \mathcal{F}_v^{(s_2)}\|^2.$$



G. Alessandrini, M.V. de Hoop, F. F., R. Gaburro and E. Sincich

Inverse problem for the Helmholtz equation with Cauchy data: reconstruction with conditional well-posedness driven iterative regularization preprint

$$\mathcal{J}_{\mathcal{G}} = rac{1}{2} \sum_{s_1} \sum_{s_2} \| d_{v}^{(s_1)T} \mathcal{F}_{p}^{(s_2)} - d_{p}^{(s_1)T} \mathcal{F}_{v}^{(s_2)} \|^2.$$

From Euler's equation,  $v_n(\mathbf{x}_i) = -\mathrm{i}(\omega \rho)^{-1} \partial_n p(\mathbf{x}_i)$ .



$$\mathcal{J}_{\mathcal{G}} = \frac{1}{2} \sum_{s_1} \sum_{s_2} \|d_v^{(s_1)T} \mathcal{F}_p^{(s_2)} - d_p^{(s_1)T} \mathcal{F}_v^{(s_2)}\|^2.$$

From Euler's equation,  $v_n(\mathbf{x}_i) = -\mathrm{i}(\omega \rho)^{-1} \partial_n p(\mathbf{x}_i)$ .

**Cauchy data**: the cost function follows **Green's identity**.



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From Euler's equation,  $v_n(\mathbf{x}_i) = -\mathrm{i}(\omega \rho)^{-1} \partial_n p(\mathbf{x}_i)$ .

- ▶ Cauchy data: the cost function follows Green's identity.
- ▶ Reciprocity gap functional in inverse scattering.



D. Colton and H. Haddar

An application of the reciprocity gap functional to inverse scattering theory Inverse Problems 21 (1) (2005), 383398.



G. Alessandrini, M.V. de Hoop, R. Gaburro and E. Sincich

Lipschitz stability for a piecewise linear Schrödinger potential from local Cauchy data arXiv:1702.04222, 2017



G. Alessandrini, M.V. de Hoop, F. F., R. Gaburro and E. Sincich

Inverse problem for the Helmholtz equation with Cauchy data: reconstruction with conditional well-posedness driven iterative regularization preprint

Florian Faucher

### Stability results



Lipschitz type stability is obtained for the Helmholtz equation with partial Cauchy data.

$$||c_1 - c_2|| \le C (\mathcal{J}_{\mathcal{G}}(c_1, c_2))^{1/2}$$

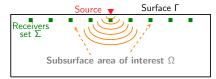
### Stability results



Lipschitz type stability is obtained for the Helmholtz equation with partial Cauchy data.

$$\|c_1-c_2\|\leq \mathcal{C}ig(\mathcal{J}_{\mathcal{G}}(c_1,c_2)ig)^{1/2}$$

Using back-scattered data from one side in a domain with free surface and absorbing conditions,



for piecewise linear parameters.



G. Alessandrini, M.V. de Hoop, R. Gaburro and E. Sincich

Lipschitz stability for a piecewise linear Schrödinger potential from local Cauchy data arXiv:1702.04222, 2017



G. Alessandrini, M.V. de Hoop, F. F., R. Gaburro and E. Sincich

### Additional possibilities



It allows the non-collocation of numerical and observational sources:

$$\mathcal{J}_{\mathcal{G}} = \frac{1}{2} \sum_{s_1} \sum_{s_2} \|d_v^{(s_1)T} \mathcal{F}_p^{(s_2)} - d_p^{(s_1)T} \mathcal{F}_v^{(s_2)}\|^2.$$

- $\triangleright$   $s_1$  is fixed by the observational setup,
- s<sub>2</sub> is chosen for the numerical comparisons.

### Plan

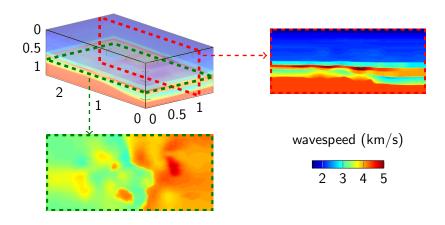


- Mumerical experiments
  - Comparison of misfit functions
  - ullet Changing the numerical acquisition with  $\mathcal{J}_{\mathcal{G}}$

### Experiment setup



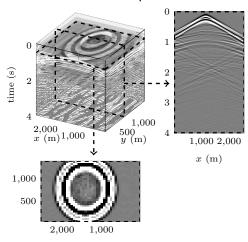
3D velocity model  $2.5 \times 1.5 \times 1.2$ km using dual-sensors data.



### Experiment setup



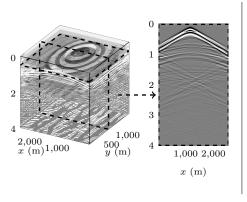
We work with time-domain data acquisition.



### Experiment setup

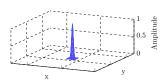


We work with time-domain data (pressure and velocity).



Acquisition for the measures

- ▶ 160 sources.
- ▶ 100 m depth,
- point source,



For the reconstruction, we apply a Fourier transform of the time data.



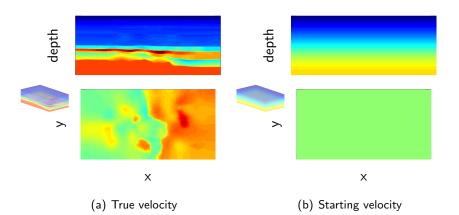
We respect the observational acquisition setup and perform the minimization of  $\mathcal{J}_{L2}$  or  $\mathcal{J}_{G}$ , frequency from 3 to 15Hz.

$$\mathcal{J}_{L2} = \sum_{\text{source}} \frac{1}{2} \|\mathcal{F}_{p}^{(s)} - d_{p}^{(s)}\|^{2} + \frac{1}{2} \|\mathcal{F}_{v}^{(s)} - d_{v}^{(s)}\|^{2}.$$

$$\mathcal{J}_{\mathcal{G}} = \frac{1}{2} \sum_{\text{source source}} \| d_{v}^{(s_{1})T} \mathcal{F}_{p}^{(s_{2})} - d_{p}^{(s_{1})T} \mathcal{F}_{v}^{(s_{2})} \|^{2}.$$

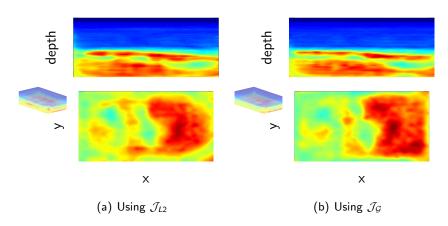


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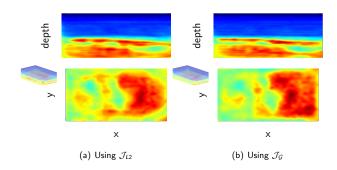


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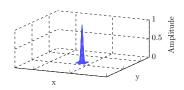
But the major advantage of  $\mathcal{J}_{\mathcal{G}}$  is the possibility to consider alternative acquisition setup.

# Experiment with different obs. and sim. acquisition nua

$$\min \mathcal{J}_{\mathcal{G}} = \frac{1}{2} \sum_{s_1} \sum_{s_2} \| \textit{d}_{\textit{v}}^{(s_1)T} \mathcal{F}_{\textit{p}}^{(s_2)} - \textit{d}_{\textit{p}}^{(s_1)T} \mathcal{F}_{\textit{v}}^{(s_2)} \|^2.$$

Acquisition for the measures  $s_1$ 

- ▶ 160 sources,
- ▶ 100 m depth,
- point source,

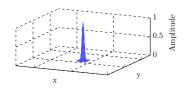


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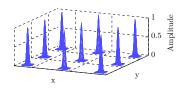
Acquisition for the measures  $s_1$ 

- 160 sources.
- 100 m depth,
- point source,



Arbitrary numerical acquisition  $s_2$ 

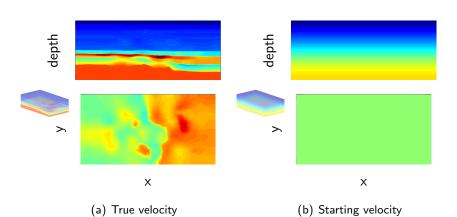
- ► 5 sources.
- ▶ 80m depth.
- multi-point sources,



- No need to known observational source wavelet.
- Differentiation impossible with least squares types misfit.

# Experiment with different obs. and sim. acquisition

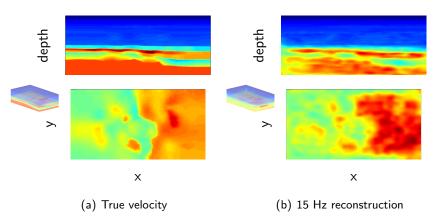
Data from frequency between 3 to 15 Hz, domain size  $2.5 \times 1.5 \times 1.2$ km, Simulation using 5 sources only.



Experiments 0000

# Experiment with different obs. and sim. acquisition

Frequency from 3 to 15 Hz,  $2.5 \times 1.5 \times 1.2$  km, Simulation using 5 sources only. -33% computational time.



# Plan

Conclusion

### Conclusion



Seismic inverse problem using pressure and vertical velocity data:

- appropriate cost function to minimize,
- allow minimal information on the acquisition setup,
- other applications,
- perspective: design the most efficient numerical setup,
- Rk: possible for elastic media with measures of traction.

#### Conclusion

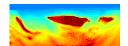


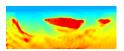
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Quantitative reconstruction algorithm toolbox for time-harmonic wave,

- Discontinuous Galerkin discretization in HPC framework.
- large scale optimization scheme using back-scattered data.
- acoustic, elastic, anisotropy, 2D, 3D, attenuation.





P- and S-wavespeed reconstructions

### Conclusion



Seismic inverse problem using pressure and vertical velocity data:

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- ▶ allow minimal information on the acquisition setup,
- other applications,
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Quantitative reconstruction algorithm toolbox for time-harmonic wave,

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#### APPENDIX

## Stability of the Helmholtz Inverse Problem



$$||c_1^{-2} - c_2^{-2}|| \le \mathcal{C}(||F(c_1^{-2}) - F(c_2^{-2})||)$$



G. Alessandrini

Stable determination of conductivity by boundary measurement Applicable Analysis 1988

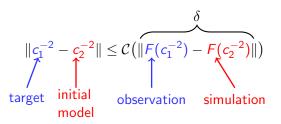


N. Mandache

Exponential instability in an inverse problem for Schrödinger equation Inverse Problems 2001

## Stability of the Helmholtz Inverse Problem





- Stability associate data and model correspondence
- Reconstruction is based on the iterative minimization of the difference between observation and simulation using an initial model.



#### G. Alessandrini

Stable determination of conductivity by boundary measurement Applicable Analysis 1988

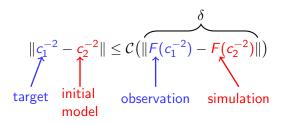


#### N. Mandache

Exponential instability in an inverse problem for Schrödinger equation Inverse Problems 2001

## Stability of the Helmholtz Inverse Problem





- Stability associate data and model correspondence
- $ightharpoonup \mathcal{C}(\delta) \leq \mathcal{C}(\log(1+\delta^{-1}))^{-\alpha}$



#### G. Alessandrini

Stable determination of conductivity by boundary measurement Applicable Analysis 1988



#### N. Mandache

Exponential instability in an inverse problem for Schrödinger equation Inverse Problems 2001

## Conditional Lipschitz stability: assumptions



- c(x) is bounded  $B_1 \le c^{-2}(x) \le B_2$  in  $\Omega$
- ightharpoonup c(x) has a piecewise constant representation of size N

$$c(x)^{-2} = \sum_{k=1}^{N} c_k \chi_k(x)$$

Ω has Lipschitz boundary

$$\|c_1^{-2} - c_2^{-2}\|_{L^2(\Omega)} \le \mathcal{C} \|F(c_1^{-2}) - F(c_2^{-2})\|$$
 (1)



G. Alessandrini and S. Vessella

Lipschitz stability for the inverse conductivity problem Advances in Applied Mathematics 2005



E. Beretta, M. V. de Hoop, F. and O. Scherzer

Inverse boundary value problem for the Helmholtz equation: quantitative conditional Lipschitz stability estimates.

SIAM Journal of Mathematical Analysis 2016

### Formulation



The stability constant is bounded

$$\frac{1}{4\omega^2} e^{K_1 N^{1/5}} \le \mathcal{C} \le \frac{1}{\omega^2} e^{(K(1+\omega^2 B_2)N^{4/7})}$$
 (2)

ightharpoonup depends on the partitioning N and the frequency  $\omega$ 



E. Beretta, M. V. de Hoop, F. and O. Scherzer

Inverse boundary value problem for the Helmholtz equation: quantitative conditional Lipschitz stability estimates  $2016\,$ 

## Conditional Lipschitz stability for Cauchy data



In the case of partial Cauchy data  $(p \text{ and } \partial_{\nu} p)$ , we have that, we can obtain a Lipschitz type stability:

$$\|c_1^{-2} - c_2^{-2}\| \le \mathcal{C} \left( \mathcal{J}_{\mathcal{G}}(c_1^{-2}, c_2^{-2}) \right)^{1/2}$$

Where  $c_1^{-2}$  and  $c_2^{-2}$  are piecewise linear.