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ARTICLE TYPE

DEM investigations of internal erosion: grain transport in the light of micromechanics.

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Abstract

Internal erosion by suffusion can change dramatically the constitutive behavior of granular materials by modifying the fabric of granular materials. In this study, the effect of an internal fluid flow on granular materials is investigated at the material point scale using the numerical coupling between a Discrete Element Method (DEM) and a Pore-scale Finite Volume (PFV) coupling scheme. The influence of the stress state and the hydraulic loading (direction and intensity) on the occurrence of grain transport in dense widely graded granular samples is thus investigated and interpreted in terms of micromechanics. In particular, it is shown that, grain transport is increased when the macroscopic flow direction is aligned with the privileged force chain orientation. The stress induced microstructure modifications are shown to influence the transport distances by controlling the number of rattlers.

KEYWORDS:

Suffusion ; DEM ; Fluid grain interactions ; Grain Transport ; Force chain ; Granular material

1 | INTRODUCTION

Internal erosion is the process by which a saturated soil can lose a part of its solid fraction under the action of an internal fluid flow. If this phenomenon occurs for instance in oil reservoir, the resulting sand production may clog the reservoir and reduce oil production (1). If internal erosion occurs in earthen dikes or dams or in natural deposits of loose rocks in mountain rivers, the modifications of the internal structure of the soil can generate mechanical instability that will lead to unexpected failures in case of sudden changes in the loading conditions (2, 3, 4).

Among the four types of internal erosion identified today (3), suffusion, that consists in the selective erosion of the finest particles of a soil, affects directly the internal fabric of granular materials. Its modeling at the continuum scale is still a complex issue as it requires the identification of internal state variables able to account for the microstructure state. To overcome the lack of reliable constitutive models, efforts have been historically devoted to anticipate and prevent the occurrence of grain transport. Consequently, many suffusion susceptibility criteria have been proposed to ensure soil self-filtration. These criteria are usually constructed either on the particle size distribution (PSD) or constriction size distribution (CSD) curves. As a result they often lack of microscale justifications. Moreover, they may prove to be too conservative as they totally dismiss the consequences of suffusion on the mechanical behavior of granular materials. The notion of internal stability derived from these criteria is thus completely different from the notion of mechanical stability.

Thanks to ever increasing computation power together with recent efforts to develop numerical schemes accounting for fluid/-grain interactions locally (5, 6, 7, 8, 9), the full three-dimensional modeling of suffusion at the scale of representative elementary volumes (REV) is now accessible. The results of such numerical simulations can be used to analyze locally the elementary

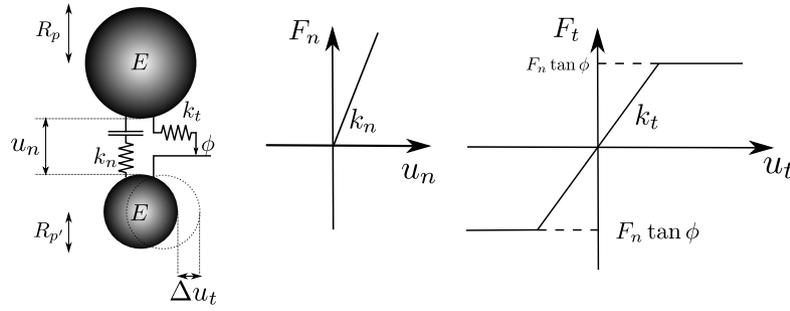


FIGURE 1 Elasto-frictional contact law used in DEM simulations.

mechanisms responsible for suffusion and their consequences on the mechanical behavior of granular materials. Such a micromechanical approach will provide a physical meaning to additional terms introduced in phenomenological models or drive the definition of physically sound micromechanically-based models.

In this paper, REV of granular materials are considered through a discrete element modeling approach and fluid/grains interactions are accounted for using a pore network model developed by Chareyre et al. (6). This numerical scheme offers a local scale modeling at a relatively cheap computational cost in order to enable fully resolved simulations in three dimensions. This contrasts with most of the approaches found in the literature that either consider 2D problems and introduce a hydraulic radius to recover the connectivity of the fluid phase (8, 10, 11) or use in 3D a non resolved CFD approach (12, 13, 14, 15, 16, 17, 18). The objective of this paper is to investigate the microscale impact on an internal fluid flow on the microstructure of widely graded granular materials for different stress states and hydraulic gradients. In particular, the occurrence of grain transport is analyzed with respect to micromechanical features of the granular assemblies.

This paper is organized as follows. In a first section, the numerical method used to tackle the 3D problem of suffusion at the material point scale is reviewed. A particular attention is devoted to the definition of the flow boundary conditions and to an erosion criterion. Then, the relative occurrence of transport and clogging is analyzed for different stress states and hydraulic gradients and related to micromechanical features of granular materials in the second section.

2 | MATERIAL POINT SCALE SIMULATIONS OF SUFFUSION

Down to the microscale, soil can be seen as a collection of individual particles interacting together at contact points. However, at the engineering scale, this discrete nature cannot be accounted for and is hidden behind a macroscopic continuum description. Between the macroscopic and microscopic scales, an intermediate scale exist for which a continuum media can be defined by averaging discrete quantities. This scale is known as the Representative Elementary Volume (REV) scale in the sense of Hill (19) and a quantitative definition with respect to suffusion can be found in Wautier et al. (20). In the following sections, the modeling of suffusion at this scale is detailed.

2.1 | DEM/PFV numerical modeling

At the REV scale, individual soil particles are accounted for using a discrete element method (DEM). In DEM, only two ingredients are needed to account for the evolution of a mechanical system: the geometry of the particles and the interaction forces between them. In this study only non-cohesive granular materials are considered in the form of poly-disperse assemblies of spheres. Particle interactions are modeled with the classical elasto-frictional contact law proposed by Cundall and Strack(21) and illustrated in Figure 1 .

In this law, a contact is created as two particles overlap. The two contacting bodies are then assumed to be linked by two fictitious linear springs and a slider. Based on the direction of the vector joining the particle centers, a normal force F_n is defined as proportional to the overlapping distance u_n between the two particles where the contact normal stiffness k_n is proportional to the material's Young modulus E and to the harmonic average of the two particle radii R_p and R_p' ($k_n = E \frac{2R_p R_p'}{R_p + R_p'}$). In addition to the normal force, a tangential force F_t is introduced. This tangential force is proportional to the relative tangential displacement u_t

TABLE 1 Mechanical parameters used in the elasto-frictional contact law implemented in YADE.

Parameters	Value
Density (ρ_p)	3,000 kg.m ⁻³
Young Modulus (E)	356 MPa
Stiffness ratio (α)	0.42
Inter-particle friction angle (ϕ)	35°
Particle-wall friction angle	0°
Number of particles	5,000

between the two particles where the horizontal contact stiffness is a fraction α of its normal counterpart ($k_t = \alpha k_n$). Compared to the normal displacement, the tangential displacement u_t cannot be defined directly on the current geometry. It is indeed defined in an incremental form with Δu_t being the relative tangential displacement of the contact point between two successive geometrical configurations. The linear behavior of the tangential force F_t is limited by a Coulomb friction condition based on an internal friction angle ϕ , which defines the largest accessible ratio F_t/F_n (see Figure 1). The input parameters used in this elasto-frictional contact law are reported in Table 1. They are chosen equal to those of Hadda et al. (22) and Wautier et al. (20). After computing all inter-particle contact forces, the induced particle displacements are integrated based on Newton's second law of motion using the DEM open source code YADE (23).

In order to provide a full microscopic modeling of suffusion, the local fluid impact has to be simulated also at the local scale. From a conceptual point of view, accounting for fluid/grain interaction consists in the simulation of a laminar flow in a pore space which geometry is explicitly known with no-slip boundary conditions on the fluid/grain interfaces. Under the additional assumption of fluid incompressibility, the local fluid problem consists in solving Stokes and mass conservation equations

$$\begin{cases} \mu \nabla^2 \mathbf{u} - \nabla p + \mathbf{f} = \mathbf{0} \\ \nabla \cdot \mathbf{u} = 0 \end{cases}, \quad (1)$$

where \mathbf{u} is the fluid velocity, p the fluid pressure field, μ the fluid dynamic viscosity and \mathbf{f} a body force taken here equal to zero as gravity forces are dismissed.

Given the fluid velocity field, the action of the fluid on the different particles of the material is deduced by integrating the shear and normal stresses on grain interfaces. Eventually, these additional forces can be used in DEM to update the grain positions and the pore space geometry.

Different numerical schemes exist in the literature in order to solve for the fluid phase. The four main approaches are summarized below:

- A first approach is to simulate the fluid flow with resolved Computational Fluid Dynamics (CFD) techniques. In this approach, the pore space is triangulated in order to solve Equation (1) using Finite Element Methods (FEM). As tens to hundreds FEM elements per pore are required to get accurate results, the fluid solving step is extremely expensive and published studies are extremely scarce (6). Moreover, frequent remeshing is needed to account for microstructure evolutions.
- As a result, much of the CFD approaches found in the litterature are indeed unresolved CFD approaches (12, 13, 14, 15, 16, 17, 18). Unresolved CFD is a hybrid approach as the fluid problem is not solved at the pore scale but on fluid cells much larger than the particles. As a result, Equation (1) has to be changed to account for the local porosity of the granular media. Compared to resolved CFD, unresolved CFD is computationally less expensive and requires no remeshing but in the meantime the microscale description of the problem is lost and ad hoc macroscopic assumptions are needed to account for the fluid flow (Darcy's law and Kozeny-Carman expression for the permeability for instance).
- A third option to account for the fluid at the local scale is to use Lattice Boltzmann Methods (LBM) (24, 25, 5, 26, 27, 11). These methods rely on a statistical physics approach and Equation (1) is replaced by Boltzmann discrete equation. The fluid behavior is described through a collision-propagation scheme (24). As for resolved CFD approaches, this numerical scheme requires a discretization of the pore space but LBM relies on a fixed lattice that is not changed as particles move. Even if this numerical technique can be easily parallelized, its implementation in 3D is still very expensive. In addition, it should be noted that, incompressible flow represents a high difficulty for the conventional LBM, which is entirely based on density fluctuations.

- A fourth class of approach lies in the use of pore network models (28, 29, 30, 31, 32). Indeed, these methods (initially developed to predict the permeability of porous media) are based on a discretization of the void space into connected pores. The complexity of the fluid flow is then reduced to fluid exchange laws between adjacent pores. As a result, the accuracy of these approaches strongly relies on adequate definitions of how fluids are exchanged between pores in terms of the local pore geometry. A second difficulty consists in computing the fluid forces on the surrounding particles as the fluid state is defined as piecewise constant within each pore. However, provided that these difficulties can be overcome, these approaches provide a microscale fully coupled fluid/grain modeling at a relatively cheap computational cost.

In this paper, the efficient Pore-scale Finite Volume method introduced by Chareyre et al. (6) is used to account for internal fluid flows in granular materials at the local pore scale and in three dimensions.

In the PFV method, a regular Delaunay triangulation is used to simplify the geometry of the void space by defining a connected network of pores. Based on this partition of the void space, the mass conservation equation gives for an incompressible fluid (and for no-slip boundary conditions on the fluid/grain interfaces) a relation between the time derivative of the fluid volume V_i^f contained in pore Ω_i and fluid fluxes q_{ij} from pores i to j as

$$\frac{dV_i^f}{dt} = - \sum_{j=j_1}^{j_4} q_{ij}. \quad (2)$$

As Stokes equation (1) predicts a linear relationship between the fluid velocity and the pressure gradient, the fluid exchange between two adjacent pores is sought in a Darcy (or Poiseuille) form. By assuming the fluid pressure to be constant in each pore, fluid fluxes are computed as

$$q_{ij} = g_{ij} \frac{p_i - p_j}{\ell_{ij}} \quad (3)$$

where g_{ij} is the hydraulic conductivity of constriction linking pore i to pore j and ℓ_{ij} is the inter-pore distance.

In the PFV model, Equation (1) is thus reduced to the two Equations (2) and (3) and the unknowns of the problem are the pore pressures p_i and the fluid fluxes q_{ij} where $(i, j) \in \llbracket 1, N_{\text{pores}} \rrbracket$, N_{pores} being the number of pores. As a result, the fluid flow problem is reduced to a linear algebra problem and basically to the inversion of a sparse matrix for which several efficient algorithms exist.

In Equation (3), the definitions of the geometrical parameters g_{ij} and ℓ_{ij} are not given. The accuracy of the PFV model relies indeed on astute definitions of these two parameters based on the local geometry of the constrictions joining two adjacent pores. A detailed definition of g_{ij} and ℓ_{ij} as well as a validation of the PFV scheme against fully resolved CFD simulations can be found in Chareyre et al. (6).

2.2 | Representative elementary volume definition and stress boundary conditions

In this study, non periodic cubic volumes of granular materials are considered. Such REV's are composed of spherical particles following the widely graded particle size distribution (PSD) illustrated in Figure 2. This PSD is such that the ratio between the largest and smallest particles is 10 and such that the coefficient of uniformity $Cu = d_{60}/d_{20} = 1.98$ where d_{20} and d_{60} are the particle diameters corresponding to respectively 20 % and 60 % of the mass passing fractions.

After generating a cloud of 5,000 non-overlapping spheres surrounded by six bounding planes defining a cube, the particles are inflated and allowed to rearrange according to the radius expansion technique. This process is stopped when the confining pressure applied to the bounding planes reaches 20 kPa and the normalized unbalanced force (F_{unb}) of the system decreases below $1.5 \cdot 10^{-2}$. By definition F_{unb} is equal to the mean summary force on the N_p particles divided by the mean contact force magnitude on the N_c contacts:

$$F_{\text{unb}} = \frac{\frac{1}{N_p} \sum_{p=1}^{N_p} \|\mathbf{F}_p\|}{\frac{1}{N_c} \sum_{c=1}^{N_c} \|\mathbf{F}_c\|} \quad (4)$$

where $\mathbf{F}_p = \sum_{c_p} \mathbf{F}_{c_p}$ is the sum of all contact forces acting on particle p . F_{unb} has no units and provides an intrinsic measure of how close to equilibrium a mechanical system is. During this process, the inter-particle friction angle is maintained to 35° in order to prepare a dense granular material with a void index $e = 0.55$. The resulting sample shown in Figure 2 is then rescaled such

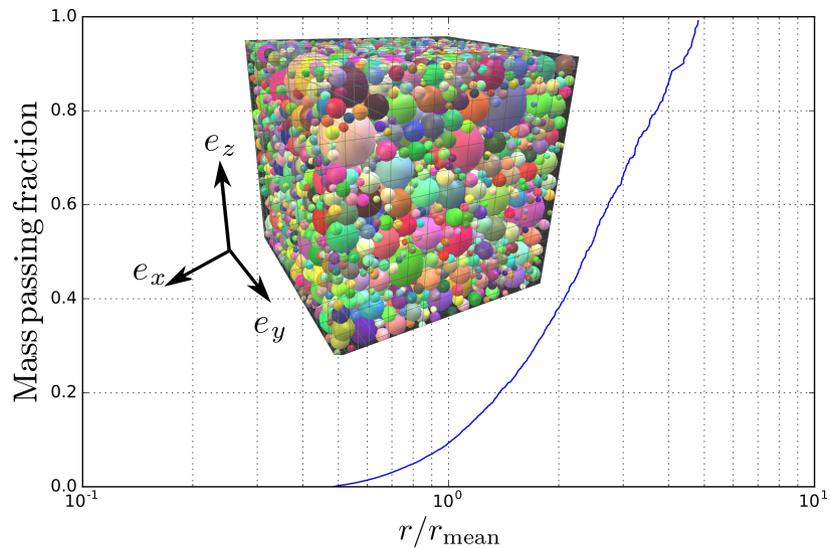


FIGURE 2 Cumulative distribution of the mass passing fraction with respect to the particle sizes. Particle radii are normalized by the mean radius value $r_{\text{mean}} = 2.10^{-4}$ m. The generated sample in the isotropic state is shown.

that $r_{\text{mean}} = 2.10^{-4}$ m. It should be highlighted here that despite the large aspect ratio between the largest and smallest particles, the used PSD is not identified as internally unstable according to Kezdi's filter rule (33) or Kenney and Lau criterion (34).

At the REV scale, a Cauchy stress tensor σ can be defined from the forces applied on the bounding walls as long as the system remains quasi-static. According to soil mechanics convention, compressions are counted positive. The stress ratio η is then introduced as the ratio between the deviatoric stress q and the mean pressure p

$$\begin{cases} \eta = \frac{q}{p} \\ q = \sqrt{\frac{3}{2} \sigma_{\text{dev}} : \sigma_{\text{dev}}} \\ \sigma_{\text{dev}} = \sigma - p \mathbf{1} \\ p = \frac{1}{3} \text{Tr}(\sigma) \end{cases} \quad (5)$$

where $\mathbf{1}$ stands for the identity tensor and ":" stands for the double dot contraction product.

From the bounding walls displacements, a macroscopic homogeneous strain tensor ϵ is also defined. The volumetric strain is then obtained as

$$\epsilon_v = \text{Tr}(\epsilon) \quad (6)$$

where Tr is the trace operator.

From the obtained equilibrium state, a drained triaxial loading is imposed to the sample in the form of a double-step procedure. First, the confining pressure (σ_0) is increased from 20 kPa to 100 kPa by allowing the bounding walls to move. Once a new equilibrium state is reached ($F_{\text{unb}} < 10^{-4}$, dimensionless), a vertical compression strain rate $\dot{\epsilon}_{zz} = 0.01 \text{ s}^{-1}$ is applied up to 3 % of deformation while keeping the same lateral confining pressure σ_0 . This strain rate is chosen similar to the one used in previous numerical studies (22, 35) and is supposed to be sufficiently small so that the loading can be considered as quasi-static. In Figure 3, the stress ratio η and volumetric strain ϵ_v responses are shown. A typical contractive/dilative behavior of a dense soil is observed.

While the stress ratio increases, two samples are saved for $\eta \in \{0.7, 0.9\}$ (corresponding to diamonds in Figure 3). As a result, two granular material REV's with slightly different stress and microstructure states are obtained. This ends the sample preparation procedure.

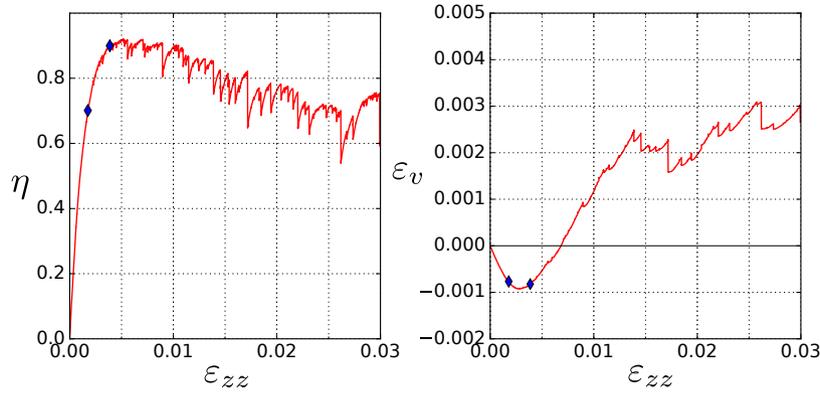


FIGURE 3 Drained triaxial test response. The diamond points correspond to the mechanical states considered for the stability analysis.

TABLE 2 The six DEM/PFV simulations carried out.

Stress state	Hydraulic gradient		Flow direction	
	$I = 0.1$	$I = 1$	Vertical (e_z)	Horizontal (e_x)
$\eta = 0.7$	x		x	x
$\eta = 0.9$	x	x	x	x

2.3 | Flow boundary value problem and bounce back erosion criterion

By adjusting the displacements of the six bounding planes around the considered granular samples, the stress or strain state of the material is controlled at the REV scale. Concerning the fluid boundary conditions, hydraulic gradients are imposed by controlling the water pressure or flux on the six bounding planes. In the following, two flow directions are considered:

- parallel to the principal loading direction (along z axis) and referred to as *vertical flow case* hereafter;
- perpendicular to the principal loading direction (along x axis) and referred to as *horizontal flow case* hereafter.

Figure 4 summarizes the boundary value problem solved for the two flow cases. On upstream faces, the fluid pressure is set to p_0 while on downstream faces null fluid pressure and null displacement are imposed. On the remaining faces of the sample, zero fluxes conditions are imposed ($q_x = q_y = 0$ for the vertical case and $q_z = q_y = 0$ for the horizontal one). The upstream pressure p_0 is prescribed such that the hydraulic gradient $I = \frac{p_0}{\rho g \ell}$ is set either to 0.1 or 1 with $\rho = 1000 \text{ kg.m}^{-3}$ the fluid density, $g = 9.81 \text{ m.s}^{-2}$ the earth gravity and $\ell \in \{\ell_x, \ell_z\}$ the sample length in the flow direction (since gravity forces are not considered, I is simply a dimensionless rescaling of p_0). It should be emphasized that the additional fluid pressure p_0 applied only on the upstream wall generates an effective stress gradient between the upstream and downstream walls of the sample. As a result, in the direction of the pressure gradient, the downstream effective stress is p_0 higher than the upstream one. To ensure the controllability of the system in YADE, zero displacement conditions $u_x = 0$ or $u_z = 0$ are imposed to the downstream walls. Since the effective stress is now varying at the scale of the sample, the problem of interest is no longer a homogenization problem and the concept of REV vanishes. However, as the fluid pressure drop p_0 is small compared to the effective stress σ_0 and if we neglect the flow induced microstructure heterogeneities at the scale of the sample, the concept of REV is restored approximatively. Under this hypothesis, the fluid can be seen as an external perturbation that makes the microstructure evolve without any change in the stress state.

In total, six DEM/PFV simulations are defined by considering different stress/microstructure states and different hydraulic loading conditions. They are summarized in Table 2 .

In this paper, non-periodic boundary conditions are used in order to avoid collective motion of the particles under the action of the fluid. As a result fluid forces can be applied to all particles. However, non periodic boundary conditions generate strong boundary effects when grain transport is considered. Indeed, transported particles are either i) extracted out from the core of the

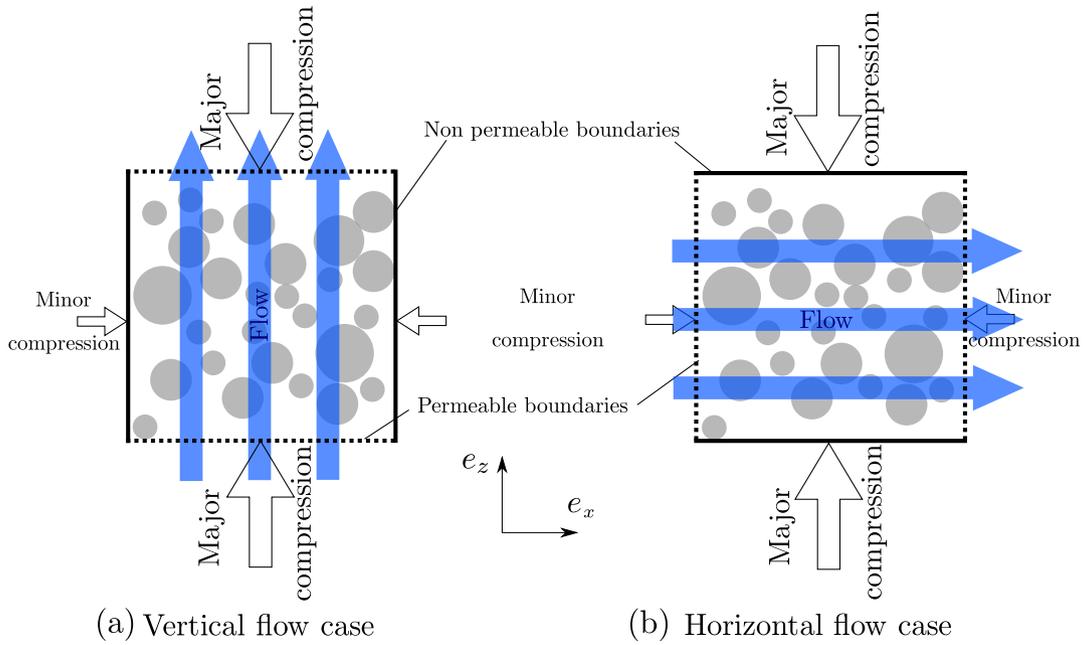


FIGURE 4 Boundary value problems solved for two flow directions. In the vertical case (a), zero flux boundary conditions are applied in directions x and y while on the top face a pressure p_0 is imposed while keeping $p = 0$ on the bottom face. In the horizontal case (b), zero flux boundary conditions are applied in directions z and y while on the left face a pressure p_0 is imposed while keeping $p = 0$ on the right face.

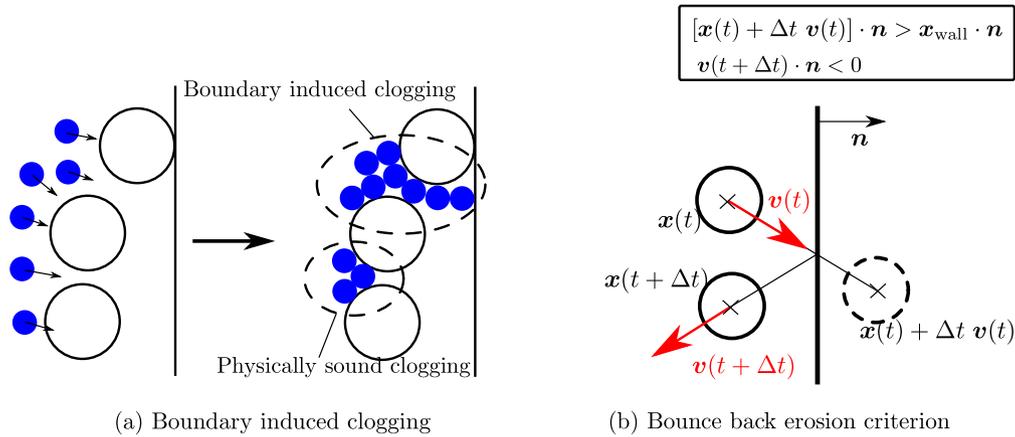


FIGURE 5 Illustration of the non-physical clogging induced by the downstream wall.

sample as they reach the downstream wall or ii) filtrated by the primary fabric as they are blocked by a too narrow constriction. In the first case, the downstream wall induces artificial clogging as illustrated in Figure 5 .

To avoid such boundary effects, it is thus important to detect and delete particles reaching the downstream wall. A naive and computationally expensive approach is to look for particles with a single contact with the downstream wall. The main problem with this criterion is that it has to be run at every time step as extracted particles often bounce several times and the contact with the downstream wall never lasts for long. As a result, a prediction/correction algorithm was preferred to this first approach. This algorithm is illustrated in Figure 5 and consists in the following two steps:

- At time t , rattlers are identified. By assuming that their velocity $\mathbf{v}(t)$ remains constant over the interval $[t, t + \Delta t]$, rattlers that will virtually go through the downstream wall during this interval are identified as candidates to particle removal. This is the prediction step.
- At time $t + \Delta t$, the scalar product between the updated velocity $\mathbf{v}(t + \Delta t)$ and the outward normal to the downstream wall \mathbf{n} is computed for rattlers candidates to particle removal. $\mathbf{v}(t + \Delta t) \cdot \mathbf{n} < 0$ means that the considered particle bounced back on the downstream wall during the time interval $[t, t + \Delta t]$ and will be removed from the simulation. This is the correction step.

This approach reduces the computational cost needed to detect extracted particles as loops over the number of particles are required only from time to time and not at every DEM time step. In practice, Δt is set to be equal to 1,000 DEM time steps. To improve computational efficiency, when extracted particles are identified, they are not deleted immediately. Indeed, as the PFV scheme relies on a particle based tessellation of the pore space (see Section 2.1), remeshing is needed each time a particle is deleted. As a result, extracted particles are kept in memory and are only deleted after several runs of the prediction/correction algorithm (every 5,000 iterations in practice which is small enough to avoid that particle that bounced back collide with another particle).

3 | INTERNAL FLUID FLOW IMPACT ON GRANULAR MICROSTRUCTURES

For cohesionless granular materials, internal fluid flows can modify the constitutive mechanical behavior at the continuum scale through two kinds of mechanisms:

- a first change lies in the transport of rattlers, i.e. particles with no contacts with their neighbors in the absence of gravity. Under the action of an internal fluid flow, these particles move through the pore space and are either eroded or clogged depending on their size with respect to those of pores and constrictions.
- a second flow impact consists in introducing additional forces on the primary skeleton of the granular material (10, 36, 37). This results in stress redistribution and possibly to grain detachment if the existing contact network is forced to rearrange.

In the following, the first aspect is emphasized while the second one has been isolated by artificially removing rattlers and was discussed in a previous paper (37).

3.1 | Numerical results for grain transport

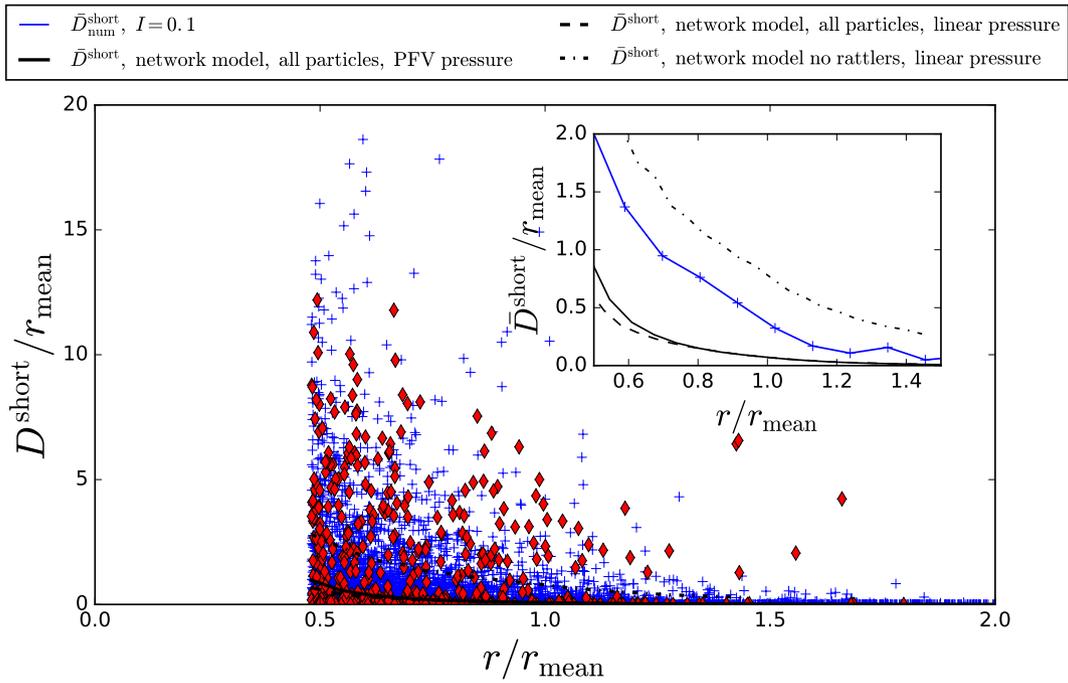
By tracking the positions of all particles during the fluid flow simulations, travel distances can be estimated for all the particles. The effective travel distances may be computed by integrating the norm of the particle velocities over time, but if transient evolutions are dismissed, the shortest displacements of all particles can be defined by simply comparing the initial and final positions of all particles. For a given particle p , moving an initial position $\mathbf{x}_p(t_0)$ to a final one $\mathbf{x}_p(t_f)$ in DEM/PFV simulations, the associated shortest travel distance reads

$$D_{\text{num}}^{\text{short}}(p) = \|\mathbf{x}_p(t_f) - \mathbf{x}_p(t_0)\|. \quad (7)$$

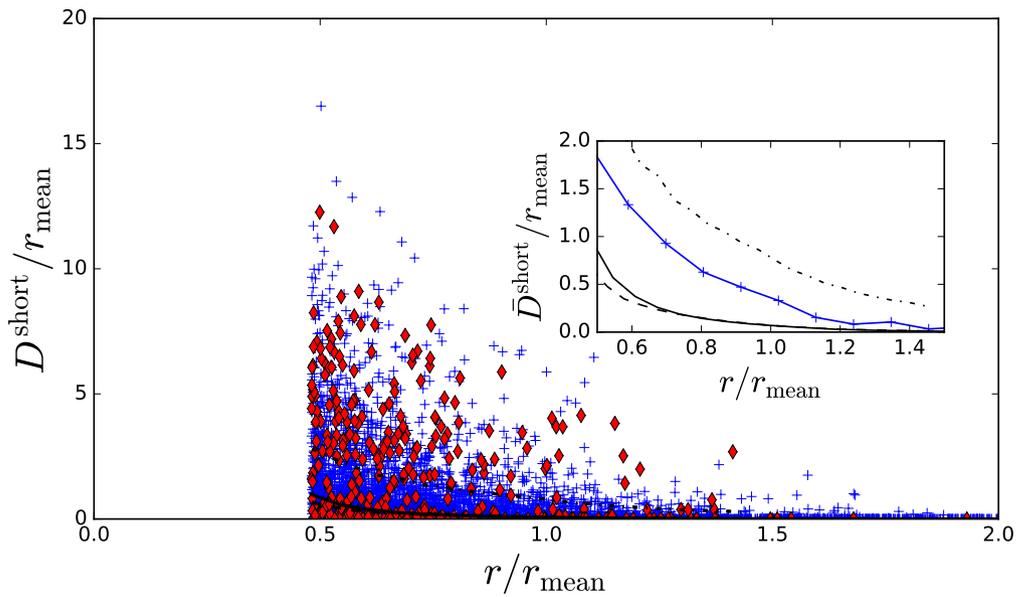
Because of the finite dimension of the sample, $D_{\text{num}}^{\text{short}}(p)$ does not always correspond to a maximum travel distance. Indeed, transported rattlers are either blocked by a too narrow constriction or reach the downstream side of the sample before being removed from the sample. In the latter case, $D_{\text{num}}^{\text{short}}(p)$ corresponds to the initial distance between the particle and the downstream wall. If the sample would have been larger, these particles would have traveled farther away.

In Figures 6 a and 6 b $D_{\text{num}}^{\text{short}}(p)$ are plotted for all particles p for the sample corresponding to $\eta = 0.9$ and subjected to a hydraulic gradient $I = 0.1$ applied in vertical or horizontal directions, respectively.

In Figures 6 a and 6 b the clouds of points corresponding to extracted (red diamonds) and non extracted (blue crosses) particles superimpose, which shows that the samples are sufficiently large to compute mean travel distances representative of material properties and not influenced by sample boundaries. Compared to the sample dimensions $(\ell_x, \ell_y, \ell_z) = (43.7r_{\text{mean}}, 43.7r_{\text{mean}}, 43.5r_{\text{mean}})$ some particles are transported over quite large distances up to 20 times the mean radius, but the average travel distances $\bar{D}_{\text{num}}^{\text{short}}$ are limited ($\bar{D}_{\text{num}}^{\text{short}} < 2r_{\text{mean}}$). Indeed, too few pores and constrictions large enough to allow for transport exist, and clogging dominates over erosion. This is consistent with the fact that the used PSD in Figure 2 is not identified as internally unstable according to Kezdi's filter rule (33) or Kenney and Lau criterion (34).



(a) Vertical flow case



(b) Horizontal flow case

FIGURE 6 Shortest travel distances of all the particles of the initially unstable sample corresponding to $\eta = 0.9$ subjected to a hydraulic gradient $I = 0.1$ in the vertical (a) or horizontal (b) directions. Extracted particles are identified by red diamonds. The mean transport distances $\bar{D}_{\text{num}}^{\text{short}}$ are shown in blue solid lines. In black solid lines the numerical predictions corresponding the transport model of Section 3.2 are shown (pressure map computed using the PFV scheme). Dashed and dot-dashed lines correspond to the predictions computed from the linear approximation of the macroscopic pressure gradient respectively on the pore network computed with all particles and on the pore network computed with only non free particles.

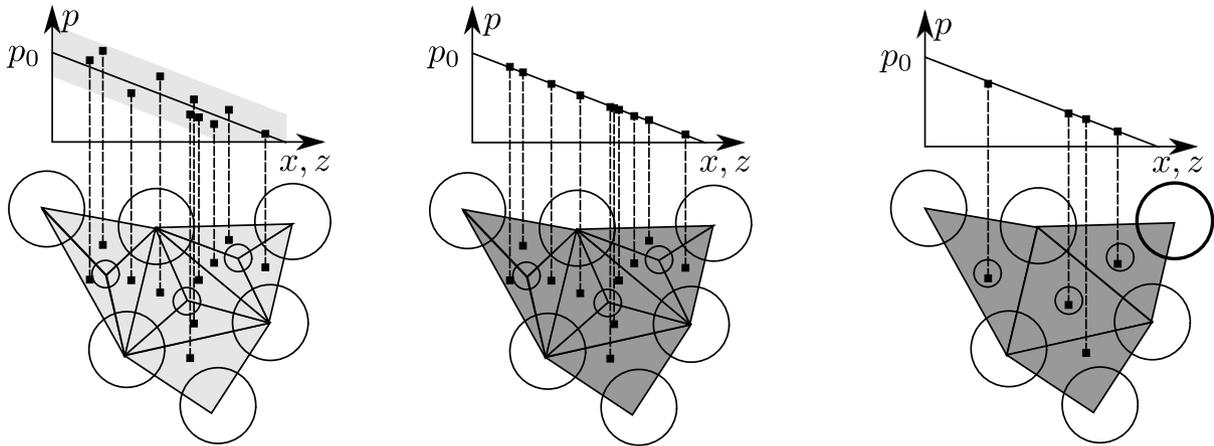
3.2 | Pore network model predictions

In this subsection, the mean travel distances $\bar{D}_{\text{num}}^{\text{short}}$ computed numerically from DEM/PFV simulations are compared with pore network model predictions \bar{D}^{short} . To this respect, three predictions of \bar{D}^{short} based on the three weighted pore networks illustrated in Figure 7 are considered.

- The first pore network relies on a triangulation of the pore space based on all particles and a pressure map computed using the PFV scheme (Figure 7 (a)).
- The second pore network relies on a triangulation of the pore space based on all particles and a pressure map computed using the linear approximation of the macroscopic pressure gradient (Figure 7 (b)).
- The third pore network relies on a triangulation of the pore space based only on non free particles and a pressure map computed using the linear approximation of the macroscopic pressure gradient (Figure 7 (c)).

Mean transport distances are then estimated on these weighted pore network without running any numerical simulations using the propagation criterion introduced in (20). This criterion is based on the definition of a flow intensity indicator in every constriction linking adjacent pores. Based on the Hagen-Poiseuille flow velocity profile for a circular tube and the expression of the drag force of a uniform flow acting on a sphere for low Reynolds numbers (Figure 8), the drag force acting on a particle of radius r in a constriction of radius R and subjected to a pressure gradient ∇p is indeed assumed to scale as

$$F \sim r \nabla p R^2. \quad (8)$$



(a) Pore network on all particles, PFV pore pressure map (b) Pore network on all particles, linear pore pressure map (c) Pore network on non free particles, linear pore pressure map

FIGURE 7 Three different pore networks used to estimate mean travel distances. The network can be constructed either from all particles ((a) and (b)) or only from non free particles (c). The pore pressure map can be derived from the use of the PFV scheme (a) or from the linear approximation of the macroscopic pressure gradient ((b) and (c)).

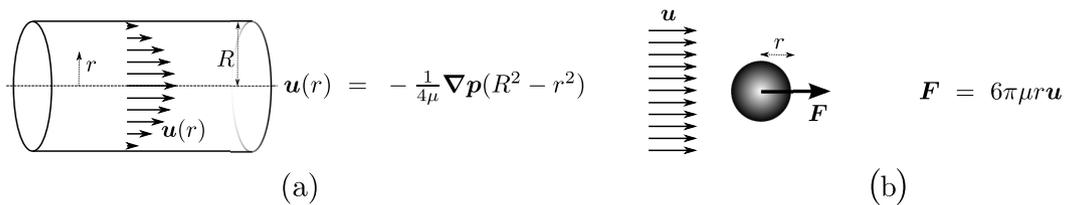


FIGURE 8 Hagen-Poiseuille flow velocity profile in a circular tube (a) and Stokes drag force (b). In both cases viscous flow conditions are assumed.

Given a pore i connected to four neighboring pores $\{j_1, j_2, j_3, j_4\}$, a particle of radius r in the pore i will move to a neighboring pore $k \in \{j_1, j_2, j_3, j_4\}$ if and only if

$$\begin{cases} R_{ik} > r \\ \|\nabla p_{ij}\| R_{ij}^2 \leq \|\nabla p_{ik}\| R_{ik}^2 \quad \forall j \in \{j_1, j_2, j_3, j_4\} \end{cases} \quad (9)$$

where R_{ik} is the constriction radius of the connection between pores i and k . In other words, the particle propagates to the next accessible pore in the direction of maximum drag force. This propagation criterion initially introduced in (20) enables to compute the expected shortest travel distance $D_p^{\text{short}}(r)$ using the propagation criterion constructed from Equations (9) and (8) for any radius r and pore p . The expected mean travel distance $\bar{D}^{\text{short}}(r)$ is then computed by averaging $D_p^{\text{short}}(r)$ over the number of pores.

Mean expected shortest travel distances \bar{D}^{short} computed for the three pore networks of Figure 7 are plotted in Figures 6 a and 6 b. While \bar{D}^{short} computed on the pore network build only on non-rattlers overestimates the real travel distances, \bar{D}^{short} computed on the pore network build on all particles underestimates the real travel distances. Mean expected shortest travel distances \bar{D}^{short} give indeed lower and upper bounds of the transport distances estimated using fully coupled DEM/PFV simulations:

- When the pore network is computed with all particles, pore and constriction sizes are underestimated. For instance, for a free particle in a tetrahedron composed of four particles, four small pores are identified in the corresponding pore network whereas the free particle moves in a single large pore (see Figure 7). As a result, mean travel distances computed from this pore network provide lower bounds for the effective mean travel distances.
- When the pore network is computed with only non free particles, pore and constriction sizes are more relevant but collective motions of the free particles are not accounted for (in particular collective clogging). As a result, mean travel distances computed from this pore network provide upper bounds for the effective mean travel distances.

In Figures 6 a and 6 b, the comparison between the predictions obtained with the pore networks of Figure 7 (a) (pressure map computed using PFV scheme) and Figure 7 (b) (pressure map computed using the linear approximation of the macroscopic pressure gradient) shows that local pore pressure fluctuations tend to increase the travel distances. This result justifies the use of a *resolved* approach to model the fluid flow among the modeling options listed in Section 2.1.

3.3 | Micromechanical insight on particle transport

When comparing Figures 6 a and 6 b, transport and erosion seem to be more important for the vertical than for the horizontal flow direction. This can be quantified in Figure 9 by checking the time evolution of the particle mass fraction removed from the simulation for the six DEM/PFV simulations of Table 2 . In order to present the results in a dimensionless form, a characteristic time scale is defined as $\tau = \ell \sqrt{\frac{2\rho_p}{p_0}}$ with ℓ being the length of the sample in the flow direction ($\ell \in \{\ell_x, \ell_z\}$ depending on the flow case), ρ_p being the particle density and p_0 the fluid pressure drop. τ corresponds indeed to the time needed for a particle initially at rest to move from a distance ℓ under the action of a pressure gradient p_0/ℓ .

Figure 9 , shows that only 1.4-1.9 % of the total mass of the particles is removed from the samples in the end of the simulations. This corresponds to 10-14% of the initial mass of rattlers as for $\eta \in \{0.7, 0.9\}$ the initial rattlers population represents respectively 13.41 % and 13.96 % of the total mass of particles. This result is consistent with the fact that the used particle size distribution is self-filtrating and that clogging dominates over erosion. A larger eroded mass is observed i) for the vertical flow case, ii) for the larger hydraulic gradient $I = 1$ and iii) for the largest stress ratio $\eta = 0.9$. For $t > 3.5\tau$, all extractable particles have reached the downstream size of the sample which shows that the local tortuosity slows down particle extraction by a factor of 3.5.

These observations can indeed be interpreted through micromechanical analysis. It is now well established that stresses in granular materials concentrate on a limited number of particles organized in mesostructures called force chains (38, 39, 40, 41, 42). It has also been shown that, the fraction of particles involved in force chains (the chained particles), is closely related to the macroscopic stress state. From drained triaxial tests, it was shown that, when the axial strain loading starts, force chains are built in the loading directions (20). As long as the material contracts, more and more particles are recruited to build new force chains. For dense granular materials characterized by a contractive/dilative behavior, the number of force chains decreases after the characteristic point (ε_v is minimum) up to the peak stress point (η is maximum). Simultaneously to the decrease in the number of chained particles, more and more rattlers (i.e. particles with no contacts) are released in the pore space. These trends are illustrated in Figure 10 in which the chained particles are identified using the force chain definition proposed by Peters et al. (43).

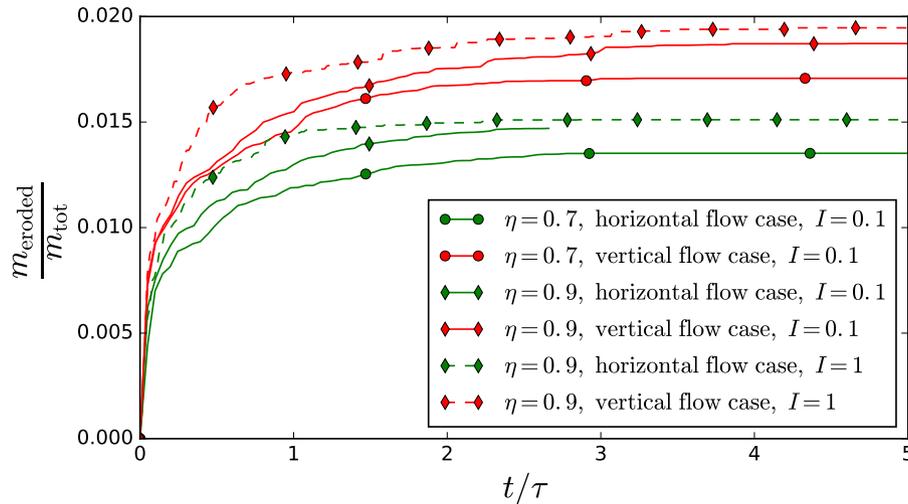


FIGURE 9 Dimensionless time evolution of the eroded mass (corresponding to the mass of deleted particles) for the six DEM/PFV simulations considered.

As illustrated in Figure 10 , for the two samples considered in this study, force chains are aligned with the vertical direction (the major compression direction). This load induced anisotropy explains why more particles are extracted from the sample in the vertical flow case. The flow is indeed parallel to the force chains which facilitates grain transport.

The two samples considered in this paper are saved in the dilation regime (see Figure 3). As a result, the number of chained particles decreases when η increases from 0.7 to 0.9 and conversely the number of rattlers for $\eta = 0.7$ is smaller than for $\eta = 0.9$ (see Figure 10). In addition, the macroscopic dilation imposes locally an increase in pore and constriction sizes. All these observations justify why erosion is a little more pronounced in the sample corresponding to $\eta = 0.9$ than in $\eta = 0.7$ despite the fact that it is *more stressed* (see Figure 9). It should also be underlined that the very large proportion of rattlers found in the sample (64% of the particles) is linked to the fact that i) the used PSD is very wide ($r_{\max}/r_{\min} = 10$) and ii) the volume fraction of the small particles is quite large ($Cu = 1.98$). As a result the primary fabric of the considered material relies on few particles and many smaller particles fill the interstitial voids. As illustrated by the small travel distances observed in Figures 6 a and 6 b the rattlers remain trapped because of too small constrictions and collective clogging.

In Figure 9 , when the hydraulic gradient is increased, the fluid forces applied on the particles are larger which explains why $\frac{m_{\text{eroded}}}{m_{\text{tot}}}$ increases faster for $I = 1$ than for $I = 0.1$. If particles are locally transported always in the direction of the maximum flow intensity (as assumed previously in the pore network transport models), the linearity of the PFV scheme (see Section 2.1) would result in similar transport and erosion whatever the intensity of the hydraulic gradient. In Figure 9 , this is obviously not the case as more erosion occurs when hydraulic gradients are more intense. This can result from the fact that fluid forces induce stress redistribution in the primary network of granular materials. The primary fabric of the material deforms which slightly change the constriction sizes. If the existing contact network is not resilient enough, fluid forces can even lead to force chain collapse and production of new rattlers.

As no initially jammed particles (with coordination number $Z_c > 3$) become rattlers at some point in the six DEM/PFV simulations, the flow induced deformation of the primary fabric provides a rational basis to interpret the results in Figure 9 .

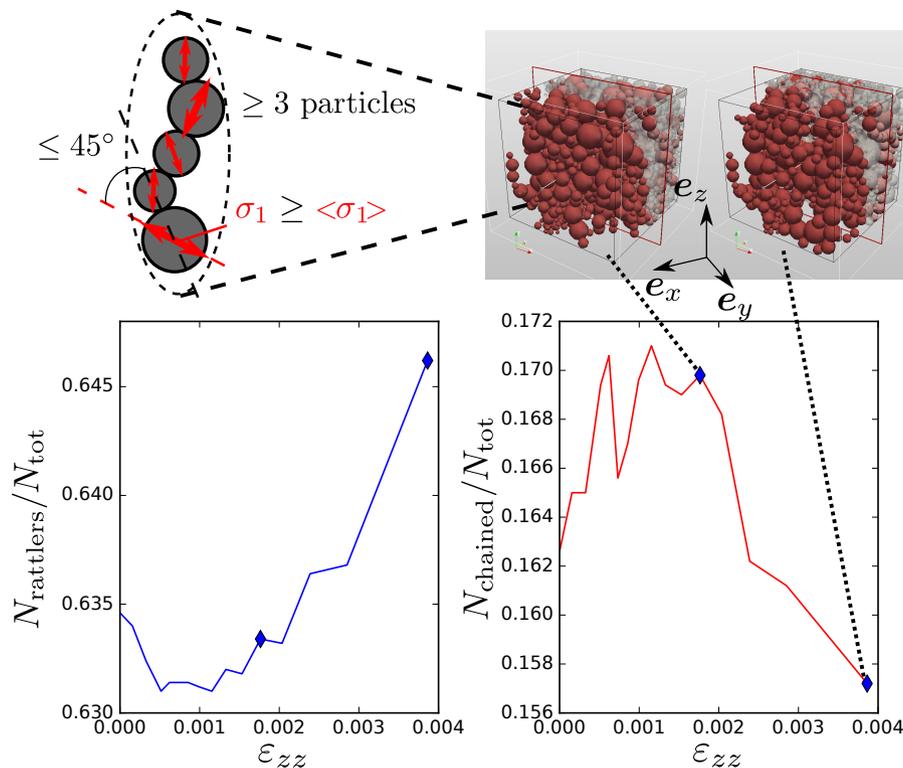


FIGURE 10 Strain evolution of the proportions of rattlers (bottom left) and chained particles (top right) during the first increments of the drained triaxial loading. The two states corresponding to $\eta \in \{0.7, 0.9\}$ are shown with diamond points. For these two states, visualization of the force chains are provided (top right) and the force chain definition of Peters et al. (43) is illustrated (top left).

4 | CONCLUSION AND OUTLOOK

Thanks to fully coupled DEM/PFV simulations, the occurrence of grain transport has been investigated at the REV scale in widely graded granular materials. By considering several hydraulic loadings (two directions and two intensities) and two stress states, the influence of micromechanical features on the travel distances has been highlighted. The stress induced microstructure modifications are shown to influence grain transport as

- for anisotropic stress states, force chains are built in privileged orientations. As a result, transport distances and erosion increase if the macroscopic flow direction is aligned with the principal force chains orientation;
- while subjected to drained triaxial tests, dense granular materials exhibit a contractive/dilative behavior. As they dilate, contact networks weaken and more rattlers are released in the pore space. This feature combined with larger pore sizes results in larger transport distances in dense granular materials subjected to large deviatoric stress.

By comparing the results obtained from simple transport models to fully coupled DEM/PFV simulations, the influence of collective grain motion and local pressure fluctuations has been highlighted. By over or under estimating these collective effects, the proposed models provide lower and upper bounds for grain transport. In addition, by modifying the local pressure maps used in the transport models, local pressure fluctuations around macroscopic pressure gradients have been shown to increase transport distances. This impact underlines the importance of using fully resolved fluid grain coupling schemes while modeling suffusion at the representative elementary volume scale.

In this study, mechanical loading is applied prior to hydraulic loading and no post-erosion mechanical loading is considered. While considering the problem of suffusion in dams or dikes, it will be of primary interest indeed to analyze the fluid flow effect on the mechanical stability of the structure undergoing internal erosion. A proper analysis can be achieved thanks to the combined use of the second-order work criterion introduced by Hill (44) and directional analyses (45). Such a work is currently in progress.

References

- [1] Papamichos E, Vardoulakis I, Tronvoll J, Skjaerstein A. Volumetric sand production model and experiment. *International journal for numerical and analytical methods in geomechanics*. 2001;25(8):789–808.
- [2] Seed H Bolton, Duncan James Michael. The failure of Teton dam. *Engineering Geology*. 1987;24(1-4):173–205.
- [3] Bonelli Stéphane. *Erosion in geomechanics applied to dams and levees*. John Wiley & Sons; 2013.
- [4] King Jonathan, Loveday Ian, Schuster Robert L. The 1985 Bairaman landslide dam and resulting debris flow, Papua New Guinea. *Quarterly Journal of Engineering Geology and Hydrogeology*. 1989;22(4):257–270.
- [5] Yu Zhao, Fan Liang-Shih. Lattice Boltzmann method for simulating particle–fluid interactions. *Particuology*. 2010;8(6):539–543.
- [6] Chareyre Bruno, Cortis Andrea, Catalano Emanuele, Barthélemy Eric. Pore-scale modeling of viscous flow and induced forces in dense sphere packings. *Transport in porous media*. 2012;94(2):595–615.
- [7] Zhao Jidong, Shan Tong. Numerical modeling of fluid-particle interaction in granular media. *Theoretical and Applied Mechanics Letters*. 2013;3(2).
- [8] Han Yanhui, Cundall Peter A. LBM-DEM modeling of fluid-solid interaction in porous media. *International Journal for Numerical and Analytical Methods in Geomechanics*. 2013;37(10):1391–1407.
- [9] Robinson Martin, Ramaioli Marco, Luding Stefan. Fluid-particle flow simulations using two-way-coupled mesoscale SPH-DEM and validation. *International journal of multiphase flow*. 2014;59:121–134.
- [10] Mahabadi Nariman, Jang Jaewon. The impact of fluid flow on force chains in granular media. *Applied Physics Letters*. 2017;110(4):041907.
- [11] Ngoma Jeff, Philippe Pierre, Bonelli Stéphane, Radjai Farhang, Delenne Jean-Yves. Two-dimensional numerical simulation of chimney fluidization in a granular medium using a combination of discrete element and lattice Boltzmann methods. *Physical Review E*. 2018;97(5):052902.
- [12] Zeghal M, El Shamy U. A continuum-discrete hydromechanical analysis of granular deposit liquefaction. *International Journal for Numerical and Analytical Methods in Geomechanics*. 2004;28(14):1361–1383.
- [13] El Shamy Usama, Zeghal Mourad. Coupled continuum-discrete model for saturated granular soils. *Journal of engineering mechanics*. 2005;131(4):413–426.
- [14] Chen Feng, Drumm Eric C, Guiochon Georges. Coupled discrete element and finite volume solution of two classical soil mechanics problems. *Computers and Geotechnics*. 2011;38(5):638–647.
- [15] Zhao Jidong, Shan Tong. Coupled CFD-DEM simulation of fluid–particle interaction in geomechanics. *Powder technology*. 2013;239:248–258.
- [16] Shan Tong, Zhao Jidong. A coupled CFD-DEM analysis of granular flow impacting on a water reservoir. *Acta Mechanica*. 2014;225(8):2449–2470.
- [17] Abdoulaye Hama N, Ouahbi Tariq, Taibi Said, Souli Hanène, Fleureau Jean-Marie, Pantet Anne. Analysis of mechanical behaviour and internal stability of granular materials using discrete element method. *International Journal for Numerical and Analytical Methods in Geomechanics*. 2016;40(12):1712–1729.
- [18] Kawano Kenichi, Shire Tom, O’Sullivan Catherine. Coupled DEM-CFD Analysis of the Initiation of Internal Instability in a Gap-Graded Granular Embankment Filter. In: :10005EDP Sciences; 2017.
- [19] Hill Rodney. Elastic properties of reinforced solids: some theoretical principles. *Journal of the Mechanics and Physics of Solids*. 1963;11(5):357–372.
- [20] Wautier Antoine, Bonelli Stéphane, Nicot François. Scale separation between grain detachment and grain transport in granular media subjected to an internal flow. *Granular Matter*. 2017;19(2):22.
- [21] Cundall Peter A, Strack Otto DL. A discrete numerical model for granular assemblies. *Geotechnique*. 1979;29(1):47–65.
- [22] Hadda Nejib, Nicot François, Bourrier Franck, Sibille Luc, Radjai Farhang, Darve Félix. Micromechanical analysis of second order work in granular media. *Granular matter*. 2013;15(2):221–235.
- [23] al. V.. *Yade Documentation 2nd ed*. The Yade Project; 2015. <http://yade-dem.org/doc/>.
- [24] Lallemand Pierre, Luo Li-Shi. Theory of the lattice Boltzmann method: Dispersion, dissipation, isotropy, Galilean invariance, and stability. *Physical Review E*. 2000;61(6):6546.
- [25] Mansouri Mouloud, Delenne J-Y, El Youssoufi Moulay Saïd, Séridi Ahcene. A 3D DEM-LBM approach for the assessment of the quick condition for sands. *Comptes Rendus Mécanique*. 2009;337(9-10):675–681.
- [26] Cuellar Pablo, Philippe Pierre, Bonelli Stéphane, et al. Micromechanical analysis of the surface erosion of a cohesive soil by means of a coupled LBM-DEM model. In: ; 2015.
- [27] Jaeger Robin, Mendoza Miller, Herrmann Hans J. Mechanism behind erosive bursts in porous media. *Physical review letters*. 2017;119(12):124501.

- [28] Bryant Steven L, King Peter R, Mellor David W. Network model evaluation of permeability and spatial correlation in a real random sphere packing. *Transport in Porous Media*. 1993;11(1):53–70.
- [29] Thompson Karsten E, Fogler H Scott. Modeling flow in disordered packed beds from pore-scale fluid mechanics. *AIChE Journal*. 1997;43(6):1377–1389.
- [30] Bakke Stig, Øren Pål-Eric, others . 3D pore-scale modelling of sandstones and flow simulations in the pore networks. *Spe Journal*. 1997;2(02):136–149.
- [31] Hilpert Markus, Glantz Roland, Miller Cass T. Calibration of a pore-network model by a pore-morphological analysis. *Transport in porous media*. 2003;51(3):267–285.
- [32] Abichou Tarek, Benson Craig H, Edil Tuncer B. Network model for hydraulic conductivity of sand-bentonite mixtures. *Canadian Geotechnical Journal*. 2004;41(4):698–712.
- [33] Kézdi A. *Soil physics (selected topics) developments in geotechnical engineering*. : ; 1979.
- [34] Kenney TC, Lau D. Internal stability of granular filters. *Canadian Geotechnical Journal*. 1985;22(2):215–225.
- [35] Nicot François, Sibille Luc, Darve Félix. Failure in rate-independent granular materials as a bifurcation toward a dynamic regime. *International Journal of Plasticity*. 2012;29:136–154.
- [36] Beguin Rémi, Philippe Pierre, Faure Yves-Henri. Pore-scale flow measurements at the interface between a sandy layer and a model porous medium: Application to statistical modeling of contact erosion. *Journal of Hydraulic Engineering*. 2012;139(1):1–11.
- [37] Wautier Antoine, Bonelli Stéphane, Nicot François. Flow impact on granular force chains and induced instability. under review2018.
- [38] Drescher A, De Jong G De Josselin. Photoelastic verification of a mechanical model for the flow of a granular material. *Journal of the Mechanics and Physics of Solids*. 1972;20(5):337–340.
- [39] Liu CH, Nagel Sydney R, Schecter DA, Coppersmith SN, Majumdar Satya, others . Force fluctuations in bead packs. *Science*. 1995;269(5223):513.
- [40] Radjai Farhang, Wolf Dietrich E, Jean Michel, Moreau Jean-Jacques. Bimodal character of stress transmission in granular packings. *Physical review letters*. 1998;80(1):61.
- [41] Bardenhagen SG, Brackbill JU, Sulsky D. Numerical study of stress distribution in sheared granular material in two dimensions. *Physical Review E*. 2000;62(3):3882.
- [42] Cambou Bernard, Jean Michel, Radjai Fahrang. *Micromechanics of granular materials*. John Wiley & Sons; 2013.
- [43] Peters J F, Muthuswamy M, Wibowo J, Tordesillas A. Characterization of force chains in granular material. *Physical review E*. 2005;72(4):041307.
- [44] Hill R. A general theory of uniqueness and stability in elastic-plastic solids. *Journal of the Mechanics and Physics of Solids*. 1958;6(3):236–249.
- [45] Gudehus G. A comparison of some constitutive laws for soils under radially symmetric loading and unloading. *Canadian Geotechnical Journal*. 1979;20:502–516.

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