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Aurélie Dessier, Paco Bustamante, Tiphaine Chauvelon, Martin Huret, Marc Pagano, et al.. The spring mesozooplankton variability and its relationship with hydrobiological structure over year-to-year changes (2003–2013) in the southern Bay of Biscay (Northeast Atlantic). *Progress in Oceanography*, 2018, 166, pp.76-87. 10.1016/j.pocean.2018.04.011 . hal-01926540

HAL Id: hal-01926540

<https://hal.science/hal-01926540v1>

Submitted on 19 Nov 2018

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1 **The spring mesozooplankton variability and its relationship with hydrobiological**
2 **structure over year-to-year changes (2003–2013) in the southern Bay of Biscay**
3 **(Northeast Atlantic)**

4

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24 **Abstract**

25 Mesozooplankton can be considered the most important secondary producers in
26 marine food webs because they hold an intermediate position between the phytoplankton
27 assemblage and the upper trophic levels. They also are a robust indicator of climatic and
28 hydrological conditions. We conducted an analysis of the interannual variability of the spring
29 mesozooplankton assemblage, as sampled by the PELGAS fisheries survey in the southern
30 part of the Bay of Biscay (Northeast Atlantic Ocean) between 2003 and 2013. We examined
31 hydrology and trophic drivers to explain the variability. Our results revealed that the
32 subsurface temperature, the subsurface salinity, the biomasses of subsurface pico-, nano-, and
33 microphytoplankton, and the copepod assemblage exhibited a recurrent spatial pattern that
34 was driven mainly by freshwater and nutrient inputs from the main rivers. The
35 mesozooplankton assemblage was dominated by copepods (82%), composed of coastal,
36 neritic, and oceanic copepod genera that paralleled the various hydrological fronts converging
37 in the southern Bay of Biscay. The copepod community displayed high temporal-variability;
38 there were three periods of abundant adult copepods throughout the southern Bay of Biscay.
39 The copepod community was structured primarily around the drive for resource control,
40 especially by the microphytoplankton biomass (24.3% of the total variability), and to a lesser
41 extent by hydrological features (13.7% of the total variability).

42 **Introduction**

43 Mesozooplankton (200–2,000 μm) play a pivotal role in marine ecosystems by transferring
44 energy from primary producers to the upper trophic levels. They are very sensitive to
45 hydroclimatic features, so climate-mediated changes in zooplankton abundance and
46 composition may affect upper trophic levels and fisheries (Beaugrand, 2003). Zooplankton
47 monitoring therefore represents a powerful tool to detect, understand, and anticipate how
48 global changes induce modifications in pelagic ecosystems (Beaugrand et al., 2003;
49 Beaugrand, 2004; Perry et al., 2004; Richardson, 2008; Hinder et al., 2014). Although
50 extensive and long-term zooplankton surveys have been undertaken (ICES Working Group on
51 Zooplankton Ecology Zooplankton Status Report: <http://wgze.net/zooplankton-status-report>;
52 O'Brien et al., 2013), they still are in relatively short supply compared to fish time-series from
53 commercial catches (Batchelder et al., 2012).

54 In this study, we investigated the assemblage structure of mesozooplankton (at the
55 level of the species, genus, or family) in the Bay of Biscay, an important Northeast Atlantic
56 fisheries area and a relatively stable ecosystem that primarily is structured around bottom-up
57 forces (e.g., Lassalle et al., 2014). Our investigation covers a large spatio-temporal scale: from
58 the Spanish coast at 46°N to the French coast at 3°35'W, from the spring of 2003 through the
59 spring of 2013. Previous mesozooplankton studies in the Bay of Biscay have been conducted
60 in the spring as well because it is the key period for plankton blooms and the reproductive
61 period for anchovies and sardines (Huret et al., this issue). Although some of these studies
62 considered large spatial and/or temporal scales, they had more limited scopes; some relied on
63 the LOPC (Laser Optical Plankton Counter) technique and were limited to size-structure
64 analyses (Sourisseau and Carlotti, 2006), while others used the CPR (Continuous Plankton
65 Recorder) method and were restricted to the “ultra subsurface” (0–7 m) (Beaugrand et al.
66 2000a, 2000b). Studies with high taxonomic resolution focused on a single field survey

67 (Albaina and Irigoien, 2007; Irigoien et al., 2011) or were conducted on limited spatial scales
68 despite their large temporal scales (Albaina and Irigoien, 2007; Stenseth et al., 2006; Valdés
69 and Moral, 1998; Valdés et al., 2007; Gonzalez-Gil et al., 2015; Cabal et al., 2008; Bode et
70 al., 2012 and 2013). Other studies focused on one or two species on a high spatio-temporal
71 scale, such as Bonnet et al. (2005) on *Calanus helgolandicus* and Lindley and Daykin (2005)
72 on *Temora stylifera* and *Centropages chierchiae*. Finally, only two studies were quite similar
73 to ours, conducting large spatio-temporal surveys of the Bay of Biscay, albeit with a lower
74 taxonomic resolution (groups were defined by image analysis) (Irigoien et al. 2009;
75 Vandromme et al. 2014). Irigoien et al. (2009) identified permanent features in the spatial
76 distribution of spring zooplankton between 1998 and 2006, with a higher abundance of large
77 organisms over the shelf break and offshore areas. Vandromme et al. (2014) used a
78 combination of LOPC data and WP2 sampling processed with the Zooscan and employed it
79 over the same network of stations used in our study between 2005 and 2012. They found a
80 negative relationship between the zooplankton biomass and normalized biomass size spectra
81 slopes, thus suggesting a clear association between zooplankton size distribution,
82 productivity, and transfer efficiency.

83 The novelty of the present study lies in the temporal and spatial coverage of the
84 mesozooplankton community structure, with a focus on the main mesozooplankton groups
85 and the key genera or species of copepods. These aspects have not been investigated
86 previously in the Bay of Biscay. From a large data set of annual spring PELGAS (‘PELagique
87 GAScogne’) surveys (Doray et al., this issue), our study aimed to elucidate which drivers
88 control the spring mesozooplankton community and whether any spatial and/or temporal
89 changes in this community occurred between 2003 and 2013. We had three objectives: 1) to
90 summarize the habitat variabilities in terms of temperature, salinity, and chlorophyll *a* data; 2)
91 to analyze the temporal patterns of the mesozooplankton and copepod spring communities

92 over the studied decade; and 3) to elucidate which local hydrological and trophic variables
93 may account for the variability of mesozooplankton communities over space and time.

94

95 **Materials and Methods**

96 **Sampling area**

97 The Bay of Biscay is characterized by a morphostructural dissymmetry opposing the large
98 northeastern edge of a narrow meridional shelf (Fig. 1). Our study focused on the southern
99 part of the Armorican shelves, which are largely under the influence of the Gironde and the
100 Adour river plumes.

101 **Data collection**

102 Samples were collected during the annual PELGAS surveys (Doray et al., this volume-b),
103 which have been conducted every spring since 2000 over the whole continental shelf of the
104 Bay of Biscay. In this study, we selected a subset of data with a homogenous
105 mesozooplankton sampling protocol, i.e. data related to the southern part of the Bay of Biscay
106 from 2000–2013 (Fig. 1). PELGAS surveys routinely collect several parameters for all system
107 components, including hydrobiology, pelagic fish abundance and distribution, marine
108 mammals, and seabird observations. We used data that had been collected at night from a total
109 of 118 visited stations (Table 1). Each station was separated from the next by approximately
110 44 km in the along-shore direction and 10–25 km in the cross-shelf direction. PELGAS
111 surveys were not carried out on exactly the same calendar dates over the series (especially at
112 the beginning of the studied decade), although they always did take place during springtime
113 over the course of about 15 days (Table 1).

114

115 *Hydrology and phytoplankton biomass*

116 Our station network for hydrology data consisted of 26 to 34 stations per year (Table
117 1, Fig. 1). At each station, the salinity and temperature were measured over the water column
118 with a CTD (Seabird 19+v2) probe. Two water column integrated indices were calculated: the
119 equivalent freshwater depth (m) and the deficit of potential energy or DEP ($\text{kg}\cdot\text{m}^{-1}\cdot\text{s}^{-2}$). As
120 detailed in Huret et al. (2013), the equivalent freshwater depth represents the local depth of
121 freshwater in the absence of mixing. As such, equivalent freshwater depth is a good indicator
122 of the plume influence in a given year and location. The DEP is an index of stratification that
123 is calculated from the vertical distribution of density values.

124 Chlorophyll *a*-based biomass (*chl_a*, see Introduction of Boyer et al. 2009) was used as
125 a proxy of autotrophic biomass (mostly phytoplankton). Water samples (200–500 mL) were
126 collected with Niskin bottles at the subsurface. Once on board, the samples were passed
127 through successive filtrations on three membranes with different porosities (Whatman GF/F
128 with pore sizes of 0.7, 3, and 20 μm and a diameter of 25 mm) to fractionate the samples into
129 three size classes: pico- (< 3 μm), nano- (3–20 μm), and microphytoplankton (>20 μm)
130 including large taxa such as diatoms. Each sample was stored at -20°C for subsequent
131 analysis. Laboratory extraction was performed with a Turner TD-700 fluorometer according
132 to the protocols of Aminot and K  rouel (2005) and of Lorenzen (1967). The *chl_a*-based
133 biomass was expressed in $\mu\text{g}\cdot\text{L}^{-1}$. An index of water-column integrated *chl_a* biomass ($\mu\text{g}\cdot\text{m}^{-2}$)
134 was calculated using the fluorescence data from the vertical profiles (WETStar fluorometer,
135 WET Labs, USA) and corrected by *chl_a*-based biomass data.

136

137 *Mesozooplankton*

138 The mesozooplankton samples were collected by vertical trawls using WP2 nets (0.25
139 m^2 opening, 200 μm mesh size), from a depth of 100 m (or the bottom depth for inshore
140 stations). For each annual survey, ten or 12 stations were selected from four transects; this

141 sampling included coastal, continental shelf, and slope stations (Fig. 1B). After collection, the
142 net samples were preserved in 4% formaldehyde. The organisms were identified by species,
143 genus, family, or more general categories/forms (Table 2; identification protocol of Rose
144 (1933) and Tregoubof and Rose (1957)), and were enumerated using a Leica M3Z stereo
145 microscope (65× to 100× magnification). The abundance was expressed in individuals m⁻³
146 (ind.m⁻³).

147

148 **Data analysis**

149 *Ordinary kriging*

150 To describe the spatio-temporal patterns of the different studied compartments, an
151 ordinary kriging procedure (ArcMap 10.2 using the Geostatistical Analyst toolbox) was
152 applied to the data for the subsurface temperature, salinity, and size-fractionated chl a -based
153 biomass. For each parameter, annual data were compiled from 2003–2013 and interpolated to
154 provide an overview of the spring spatial patterns over the time series. For each parameter,
155 annual kriged maps are presented in Supplemental Figures 1–4. To identify significant
156 temporal variations (i.e., across years) for each parameter (temperature, salinity, and pico-,
157 nano-, and microphytoplankton biomasses), a non-parametric multiple pairwise comparison
158 (Statistica®, Tulsa, OK, USA) of mean ranks was applied to all years from 2003–2013; p-
159 values < 0.05 were considered significant. All other statistical analyses were performed with
160 the R-Cran project free software (R Core Team, 2014). Furthermore, to consider these five
161 water mass variables together throughout the decade, confidence ellipses (surrounding the
162 barycentre of coordinate stations) were determined on a station factor map produced by a
163 principal component analysis (PCA, ‘FactoMineR’ package (Husson et al., 2012)).

164

165 *Mesozooplankton community analysis*

166 The time-variability of the mesozooplankton community was analyzed by considering
167 the mean yearly relative abundance of classical characteristic groups of mesozooplankton
168 (meroplankton, gelatinous plankton, copepods, and other holoplankton). The abundance and
169 taxonomic composition of the dominant group (copepods) was analyzed according to the
170 variability in space and over time. As done in previous studies, a temporal analysis was
171 conducted via a multiple comparison of the mean ranks for all years based on copepod
172 abundance. The spatio-temporal variation in the composition of the adult copepod community
173 was also described with a non-metric multi-dimensional scaling (NMDS), performed using
174 Primer 6 (Plymouth Routines In Multivariate Ecological Research) software. A matrix of
175 sampling points (n=118 samples; see Table 1) was created for the abundance data of copepods
176 species or genera (n= 23). In this matrix, the columns corresponded to the copepod taxa
177 abundances, and the lines to the sampling points. The data were transformed with $\log x + 1$
178 before the Bray-Curtis metric was applied to estimate the similarity between stations. The
179 similarity matrix was then ordinated by NMDS. A SIMPER (SIMilarity PERcentage) analysis
180 was performed to identify the genera or species contributing most to similarity (or
181 dissimilarity) within (or between) three groups of sampling years defined from the time series
182 analysis.

183 Finally, to quantify the relative influences of the food-resource (*chl_a*) and abiotic parameters
184 (temperature, salinity) that underlie the patterns observed in the copepod community
185 structure, a variation partitioning analysis (using “vegan” and “ade4” packages in the R-Cran
186 project free software) was applied (Volis et al., 2011). We considered biotic variables
187 (subsurface and integrated water column values of *chl_a* for pico-, nano-, and
188 microphytoplankton fractions) and abiotic variables (subsurface values of temperature and
189 salinity, and three water column integrated indices; the DEP, the equivalent fresh water, and
190 the mixed-layer depth). First the variables to include in the analysis (ie variables that

191 influenced the most the copepod distribution) were selected through forward selection (pack
192 for package). After the forward selection, two matrices (H matrix for abiotic parameter and C
193 matrix for biotic parameters) were built using only the most influential variables. Finally,
194 using these two matrices, the variation partitioning was performed aiming at identifying the
195 part of technology copepod distribution variation explained by each variable. By this way, our
196 aim is to evaluate if copepod distribution is better explained by hydrological parameters than
197 by food resources. Indeed, in ecological matrices each variable has a pure variation and a part
198 of variation linked to the other variables. This covariation between variables can be evaluated
199 through the variation partitioning.

200 Variation partitioning evaluates diverse components of variation: 1) the pure effect of
201 each individual matrix, 2) the redundancy of the two explicative matrices, and 3) the residual
202 effects that are unexplained by the chosen variables. Redundancy analysis (RDA), partial
203 redundancy analysis (pRDA), and Monte Carlo permutation (999 permutations) tests were
204 used to test the significance of the canonical axis.

205

206 **Results**

207 **Subsurface salinity and temperature**

208 The absence of a spatial pattern in the decadal average map (Fig. 2) means that no
209 spring-time mesoscale process had a strong recurrent impact on the surface temperature in the
210 southern Bay of Biscay. However, a coast-to-offshore positive gradient can be observed in
211 certain years (e.g., in 2006, 2010, and 2013; see Fig. S1), revealing the intermittent
212 occurrence of an upwelling process along the southern French coast. The Gironde plume
213 brought characteristically lower temperatures in 2005 and 2006, and a latitudinal gradient also
214 appeared in some years (e.g., in 2010, 2012, and 2013), albeit without any consistent pattern
215 across years. These anomalies may be explained by the changing meteorological conditions

216 (e.g., wind, irradiance) during the 15-day survey conduction. During the studied decade
217 (2003–2013), the subsurface temperature varied from 10.8°C to 19.8°C, highlighting the large
218 interannual variability in the surface warming and the stratification process during this rapidly
219 changing season. The non-parametric multiple pairwise comparison allowed for a separation
220 of the years into three groups: 2003 was the warmest, 2007 and 2011 had intermediate
221 temperatures, and the remaining years were the coldest (see Fig. S1).

222 The spatial pattern of salinity was driven by the plumes of the Gironde and the Adour
223 rivers (Fig.2). Despite some interannual variability in the extension of the plumes (maxima in
224 2007, 2009, and 2013; minima in 2011 and 2012), the interannual variability of the subsurface
225 salinity was not significantly different between years (multiple pairwise comparison, $p=0.05$)
226 (see Fig. S2). Subsurface salinities varied from 26.07 psu to 36.4 psu (see Fig. S2).

227

228 **Chlorophyll *a*-based biomasses**

229 The average picture of the *chl a* biomasses (Fig.2) revealed a homogeneous
230 concentration over the shelf, with lower values off-shelf, for the two smallest size classes (< 3
231 μm and $3\text{--}20 \mu\text{m}$). For the largest size class ($> 20 \mu\text{m}$), the highest values were found over a
232 coastal strip only. Generally, biomasses were higher in the area influenced by the Gironde
233 plume as compared to the most other southern locations.

234 Substantial differences appeared when analyzing the annual maps of *chl a* biomasses
235 (see Figs. S3, S4, and S5). For the smallest size class (Fig. S3), significant interannual
236 differences allowed the years to be divided into five groups: 2003 exhibited the lowest values
237 ($0.20 \pm 0.38 \mu\text{g.L}^{-1}$) and 2009–2010 exhibited the highest ($0.66 \pm 0.41 \mu\text{g.L}^{-1}$). For the $3\text{--}20$
238 μm size class, the interannual variability revealed four groups of years, with mean biomasses
239 ranging from $0.17 \pm 0.26 \mu\text{g.L}^{-1}$ to $0.45 \pm 0.37 \mu\text{g.L}^{-1}$. For the $> 20 \mu\text{m}$ size class, no
240 significant interannual differences were found between years.

241 The non-parametric multiple pairwise comparison of the five environmental variables
242 (subsurface salinity and temperature, subsurface pico-, nano-, and microphytoplankton
243 biomasses) over the decade could not extract significant differences between years.

244 The average contribution of each size class over the studied decade (Fig. 3) was 44%
245 for picophytoplankton ($< 3 \mu\text{m}$), 23% for nanophytoplankton ($3\text{--}20 \mu\text{m}$), and 33% for
246 microphytoplankton ($> 20 \mu\text{m}$) (Fig. 3). The relative contributions varied from 23% (2003) to
247 67% (2009) for the picophytoplankton, from 19% (2008) to 42% (2003) for the
248 nanophytoplankton, and from 10% (2007) to 42% (2008) for the microphytoplankton
249 (including larger taxa such as diatoms).

250

251 **Mesozooplankton**

252 *Total abundance and the contribution of the main groups*

253 The mean spring abundance of mesozooplanktonic organisms over the study period
254 varied from $1,321 \pm 391 \text{ ind.m}^{-3}$ (2008) to $4,986 \pm 355 \text{ ind.m}^{-3}$ (2005) (Fig. 4A). Between
255 2003 and 2006, abundances were higher ($4,454$, $3,233$, $4,985$, and $4,191 \text{ ind.m}^{-3}$ for each year,
256 respectively) than the mean decadal value ($2,995 \pm 1,195 \text{ ind.m}^{-3}$), while the mean
257 abundances were lower than the mean decadal value between 2007 and 2013, with the
258 exceptions of 2010 and 2012 (with $3,349$ and $3,061 \text{ ind.m}^{-3}$, respectively). The spring
259 abundance of mesozooplanktonic organisms revealed three temporal phases: from 2003–2006
260 (higher than the mean decadal value), from 2007–2009 (lower than the mean decadal value),
261 and from 2010–2013 (close to the mean decadal value, except 2011, which was below the
262 mean decadal value). The lowest contribution of copepods was registered in 2006 (Fig. 4B).
263 The gelatinous organisms encountered in 2006 were mainly cnidaria and siphonophora as
264 well as major meroplankton organisms (Bivalvia larvae and Cirripedia). On average, more
265 than 80% of the spring mesozooplankton community comprised copepods, 8% meroplankton,

266 8% gelatinous plankton, and 4% other holoplanktonic organisms (Fig. 4B). Depending on the
267 year, the relative abundance of copepods varied from 42% (2006) to 87% (both 2011 and
268 2009), those of meroplankton from 1% (2009) to 22% (2006), those of gelatinous organisms
269 from 3% (2011) to 24% (2006), and those of the other holoplanktonic organisms from 2%
270 (2009) to 11% (2010). Meroplanktonic organisms were found mostly along the coast,
271 gelatinous organisms were scattered, and the other holoplanktonic organisms were widespread
272 among the sampling stations (data not shown). Based on these observations, the following
273 results focus on the main group: the copepods.

274

275 *Copepod patterns*

276 Over the 11 springtime surveys, copepodites (16% of the total copepod abundance)
277 and nauplii (1.4%) of miscellaneous copepod species were recorded.

278 Twenty-five adult copepod taxa (at the family, genus, or species level) were recorded;
279 orders included Harpacticoida, Poecilostomatoida, Cyclopoida, and Calanoida, as well as
280 copepodites and copepod nauplii (Table 2). The spatial distribution of these 25 taxa is shown
281 in Fig. 5. Irrespective of the year, the copepod abundance followed a longitudinal gradient:
282 coastal waters had the highest values ($2,848 \pm 2,507 \text{ ind.m}^{-3}$), shelf waters had intermediate
283 values ($2,335 \pm 3,787 \text{ ind.m}^{-3}$), and the continental slope had the lowest values ($1,304 \pm 1,546$
284 ind.m^{-3}), except for in 2010 ($1,467 \text{ ind.m}^{-3}$). One outlier of copepod abundance ($21,127 \text{ ind.m}^{-3}$)
285 was recorded in 2005 at a single station on the northern coast. From a temporal viewpoint
286 and based on the annual abundances at each station, only 2009 could be distinguished from
287 the other years (Fig. S6). Moreover, eight of these 25 adult taxa (*Oithona* spp., *Acartia* spp.,
288 *Temora longicornis*, *Oncaea* spp., Calanidae, *Calanus helgolandicus*, Paracalanidae, and
289 *Euterpina acutifrons*) alone represented more than 1.5% of the copepod abundance (Table 2).
290 Their spring spatial abundances in the Bay of Biscay are plotted in Fig. 5. *Oithona* spp. was

291 present preferentially from the continental shelf to the slope; *Acartia* spp., *Oncaea* spp., and
292 *Temora longicornis* were most common at the coastal stations; and the genera *Acartia*,
293 *Oithona*, *Temora*, and *Oncaea* were the most abundant within the copepod community (66%
294 between 2003 and 2013; see details in Table 2).

295

296 *Spatio-temporal variations of adult copepods*

297 The NMDS ordination of the adult copepod abundance data (stress value of 0.23,
298 indicating a good ordination) discriminated the temporal variations better than the spatial
299 variations. The plots of the three time-sampling groups previously discerned by the temporal
300 analysis of abundance (see above) showed that the adult copepod community in the 2003–
301 2006 period was clearly distinct from the 2007–2009 period (Fig. 6). The last period (2010–
302 2013) was intermediate, suggesting a return toward the initial situation (2003–2006) (Fig. 6).
303 The SIMPER analysis (Tables 3 and 4) confirmed this trend, with the same four taxa (*Oithona*
304 spp., *Oncaea* spp., *Acartia* spp., and *Temora longicornis*) explaining > 65% of the similarity
305 within the 2003–2006 and 2010–2013 periods, whereas in 2007–2009, Paracalanidae (15%
306 contribution for this group) replaced *Temora longicornis* (7.8%) in these top-four taxa.
307 Overall, the within-group similarity was higher for the 2003–2006 (67.0%) and the 2010–
308 2013 (59.9%) periods than for the 2007–2009 period (52.3%). The level of dissimilarity
309 between the groups was highest between 2003–2006 and 2007–2009 (52.0%), and lowest
310 between 2003–2006 and 2010–2013 (45.5%).

311

312 *Hydrological versus trophic control on the spring copepod community*

313 The forward selection identified four significant variables that significantly drive the copepod
314 distribution: microphytoplankton (> 20 µm) biomass at the subsurface, equivalent freshwater
315 height, DEP, and subsurface temperature (Table 5). The variation partitioning highlighted that

316 49.2% of the variation was explained by both matrices H and C with the most variance
317 (24.3%) attributable to chl a -based microphytoplankton biomass and the next-most variance
318 (13.7%) attributable to hydrological parameters (Figure 7, Supplemental table 1). The part of
319 the variation due by the interaction between the 2 matrices (ie. between hydrological
320 parameters and subsurface microphytoplankton) was 11.2 %. Partial RDA (without
321 interaction between variables) concurred that the distribution of the copepod community was
322 influenced significantly by the microphytoplankton subsurface biomass, equivalent freshwater
323 height, and DEP (Supplemental table 1). The contribution of the subsurface temperature was
324 not significant in the variation partitioning analysis.

325

326 **Discussion**

327 *Hydrological and biological features of the mesozooplankton habitat*

328 Throughout the 2003–2013 decade, surface water warming occurred continually in the
329 Bay of Biscay, from the south to the north and from the coast to the open sea. The water mass
330 circulation in the Bay of Biscay is driven by a combination of large-scale and local forces (Le
331 Boyer et al., 2013); the warming of water masses in spring is modulated by slope currents,
332 shelf residual circulation, and the Iberian Poleward Current from Galicia to the Bay of Biscay
333 on a larger scale, and wind and irradiance conditions on a local scale (Koutsikopoulos and Le
334 Cann, 1996; Puillat et al., 2004; Rubio et al., 2013). From our study, the spatial pattern of
335 variability in the south of the Bay appeared to be independent of the interannual variability of
336 the hydrobiological parameters (e.g., the subsurface temperature). From 2003–2013, the
337 survey observed a mix of warm (in 2003, 2007, and 2011), intermediate, and cold years. This
338 variation in temperature regimes is consistent with results previously reported in this area
339 (e.g., Huret et al., 2013).

340 The size structure of the phytoplankton community contribute to the structure of
341 pelagic food webs, so our study considered three size classes (pico-, nano-, and
342 microphytoplankton). Indeed, mesozooplanktonic organisms are major grazers of nano- and
343 microphytoplankton (Marquis et al., 2007). Calvo-Díaz et al. (2008) recorded the monthly
344 contributions (in Iberian Peninsula waters from 2003–2006) of each size fraction of
345 phytoplankton, and showed that the spring (i.e., April) contribution of picophytoplankton was
346 about 20%, corresponding to the minimum recorded over an annual cycle. By contrast, the
347 contributions of nanophytoplankton and microphytoplankton were around 50% and 30%,
348 respectively. In our study, the mean balance between pico- and nanophytoplankton varied
349 widely, most likely because our spatial coverage involved more oceanic features beyond
350 200 m of bathymetry. It is important to note that the survey dates were not exactly the same
351 throughout the decade and hence added further variability. For example, compared to the later
352 surveys, the 2003 survey captured significantly different hydrology, plankton concentration,
353 and size structure.

354 To summarize, taking into account both the hydrological and the phytoplanktonic
355 components of the mesozooplankton habitat, the Bay of Biscay appeared to support a
356 consistent spatial structure across years, due mainly to the continental supply of freshwater
357 and nutrients, although a degree of interannual variability occurred in the range of the
358 observed values.

359

360 *The spring mesozooplankton community*

361 Among the prominent groups of the mesozooplankton community, meroplanktonic
362 organisms were restricted to the coastal area in our study, as previously reported by Ayata et
363 al. (2011). By contrast, gelatinous organisms were dispersed throughout the Bay of Biscay.
364 No clear spatio-temporal development of gelatinous plankton was observed, although this

365 group (mainly cnidaria and siphonophora) was particularly abundant in 2006. Their
366 proliferation may be linked to climate change, eutrophication, and/or habitat modifications
367 (e.g., Lo et al., 2008). However, the WP2 net used in the study probably was not the most
368 suitable device for monitoring gelatinous plankton.

369 In the present study, copepods represented the most dominant holoplankton taxa of the
370 mesozooplankton community, as highlighted by various studies in the same area (e.g.,
371 Albaina and Irigoien, 2007; Irigoien et al., 2009; Valdés et al., 2007; Villate et al., 2014,
372 1997). Previous observations of the spatial distribution of the spring size structure of
373 mesozooplankton have reported a negative coastal-to-offshore gradient (e.g., Vandromme et
374 al., 2014), as also observed in the present study for copepod taxa. Here, the sample
375 represented three dominant genera and one dominant species (> 10% of the copepod
376 abundance) *Acartia* spp., *Oncaea* spp., *Oithona* spp., and *Temora longicornis*—consistent
377 with previous investigations (e.g., Albaina and Irigoien, 2007; Irigoien et al., 2011).
378 According to the literature regarding the copepod community in the Bay of Biscay, *Temora*
379 *longicornis* and *Pseudocalanus elongatus* are dominant neritic species that occur
380 preferentially on the continental shelf, whereas *Acartia* spp. dominates stations under coastal
381 influence, and *Calanus helgolandicus* and *Eucalanidae* dominate under oceanic influence.
382 More specifically, *C. helgolandicus* is found in the southern part of the Bay of Biscay and in
383 the surface layers in the north during the spring (Bonnet et al. 2005).

384

385 *Temporal trends in the mesozooplankton and copepod communities*

386 Climatic indicators increasingly are being explored to identify major changes affecting
387 plankton (e.g., Wouters et al., 2015) and fish communities (Guénette and Gascuel, 2012). On a
388 similar spatial scale as in the present study, previous research detected oscillations mainly in
389 the vicinity of the Gironde river plume, which represents a major source of nutrient inputs

390 into the Bay of Biscay for both plankton (David et al., 2005) and fish (Pasquaud et al., 2012).
391 Overall temperatures in the Gironde estuary increased significantly around 1987 and again
392 around 2001 (Chaalali et al., 2013).

393 During the decade studied here, the year 2005 brought additional abrupt changes in
394 mesozooplankton abundance and diversity in the Gironde estuary—a relatively closed
395 system—and in the Arcachon basin (Chaalali, personal communication). It is possible that
396 this abrupt change also occurred in the Bay of Biscay, albeit with a time lag that is linked to
397 the Bay’s resilience properties. As described above, three temporal phases of adult copepod
398 dominance were observed in the present study: 2003–2006, 2007–2009, and 2010–2013, with
399 abundance values that were higher, lower, and higher, respectively, than the spring mean
400 decadal abundance (including 2003–2013). This temporal oscillation also correlated with
401 taxonomic composition changes (see Tables 3 and 4). However, we observed that 2010–2013
402 witnessed a reversal to the initial situation (2003–2006) with regard to both the abundance
403 and taxonomic composition of the adult copepod community. The only anomaly detected was
404 the high percentage of gelatinous organisms in the spring of 2006, which correlated with the
405 low percentage of copepods observed that year. These gelatinous organisms comprised
406 cnidaria and siphonophora primarily. It is reasonable to consider predation as a possible
407 cause, as this would result in top-down control on copepods and spatial occupation of a very
408 close ecological niche between carnivorous copepods, cnidarian, and siphonophora
409 organisms.

410

411 *Factors controlling the spring copepod community*

412 The structure and functioning of planktonic food webs depend largely on the
413 hydrodynamics of the water column. Our study found that four variables of the pelagic habitat
414 (subsurface temperature, stratification with the deficit of potential energy, equivalent

415 freshwater height, and subsurface microphytoplankton biomass) accounted for more than 49%
416 of the variability in the main mesozooplanktonic community. Zarauz et al. (2008) suggested
417 that the mesozooplankton biomass is driven mainly by hydrogeographic data (e.g., latitude,
418 longitude, surface temperature, salinity, stratification index, and water depth), and to a lesser
419 extent by potential trophic resources (e.g., the nano-microplankton biomass). By contrast, the
420 present study showed that for the same study site (same scale and season), the spring copepod
421 community (described taxonomically) was influenced more by trophic variables (e.g.,
422 microphytoplankton biomass; 24.3% of the total variability) than by hydrographic variables
423 (13.7% of the total variability) over the 11-year time period. These differences can be
424 explained by the fact that Zarauz et al. (2008) studied mesozooplankton at the biomass level
425 only, whereas our study took into account the abundance of taxa (species, gender, or family)
426 to describe the community. These methods do not provide the same ecological information.
427 Moreover, the temporal dataset was not exactly the same: Zarauz et al. (2008) studied three
428 spring seasons (2004–2006) while the present study included eleven spring seasons (2003–
429 2013). Finally, the two studies used different trophic variables for the statistical analysis:
430 nano-microplankton biomass in Zarauz et al. (2008) vs. phytoplankton biomasses by size
431 classes in the present study.

432 At the scale and with the spring timing of our study, the trophic link (i.e., the
433 subsurface microphytoplankton biomass) appeared to be the major driver of the copepod
434 community, while hydrographic variables appeared to play less important roles. There are two
435 potential reasons for this: i) the trophic dimension probably represents a time integration of
436 hydrological conditions (i.e., more integrated than the temperature and salinity variables in the
437 present context), and ii) the planktonic food web probably has an ecological succession of
438 physical and geographical parameters that control the initial phytoplanktonic blooms in late
439 winter. Nutrients then become limited in spring, ultimately influencing the production and life

440 cycle of mesozooplankton, which thus are maintained by trophic control during the PELGAS
441 spring surveys. However, copepods have other potential trophic sources, such as ciliates and
442 heterotrophic flagellates, that can influence the mesozooplankton distribution but were not
443 considered in our study.

444 According to the present observations, both the thermal stratification and equivalent
445 freshwater height significantly affected the copepod community, whereas subsurface salinity
446 did not. This probably is because thermal stratification and equivalent freshwater height are
447 more integrated spatially and temporally. However, it is surprising the water column
448 integrated chl a index did not have a prominent effect in the non-parametric multiple pairwise
449 comparison of mean ranks. Thus, an approach using functional traits appears to be more
450 efficient for understanding the springtime copepod community, as copepods are major
451 predators of microphytoplankton (Marquis et al., 2007). Moreover, although fewer stations
452 were sampled in our study than in Irigoien et al. (2011), we explained a greater percentage of
453 the community's variability, probably because our data set spanned a longer period of time
454 and therefore captured more variability. However, half of the variation in the copepod
455 community over the studied decade remained unexplained. We concur with Irigoien et al.
456 (2011) that other biological functional traits of the present community—including the effect
457 of copepod density on feeding activities—merit consideration.

458

459 *Concluding remarks*

460 This study of annual springtime PELGAS survey data is the only analysis of the southern
461 Bay of Biscay's mesozooplankton community between 2003 and 2013. We claim the
462 following contributions:

- 463 • We demonstrated that the spatial structure found in the Bay of Biscay is attributable to
464 continental outflow, although some interannual variability occurred in the range of
465 the values observed.
- 466 • We generated interpolation maps that represent a valuable foundation for future
467 studies in the Bay of Biscay.
- 468 • We identified three time-oscillation phases (in terms of both abundance and taxonomic
469 composition) in the mesozooplankton community, with a major change recorded in
470 2006.
- 471 • We demonstrated, for the first time in this area, the influence of trophic variables (i.e.,
472 microphytoplankton biomass) from 2003–2013.

473

474 **Acknowledgments**

475 We thank the entire PELGAS team – researchers, technicians, students, captains and crews of
476 R/V “Thalassa” – who have contributed to the collection of the PELGAS samples since 2003.
477 We would like to express our thanks to two teams’ laboratory, LIENSs and EMH and
478 especially, M. Jacques, P. Petitgas, M. Doray, and P. Bourriau. We are grateful to Proof-
479 reading and Dr. Sophie Domingues-Montanari for English corrections. We also grateful to
480 Céline Lavergne for its statistical advices. The IUF (Institut Universitaire de France) is
481 acknowledged for its support to PB as a Senior Member. This research was supported through
482 a PhD grant for A. Dessier from the Conseil Régional de Poitou-Charentes and by the
483 European project REPRODUCE (EraNet-Marifish, FP7). This work was supported by the
484 “Plateau Microscopie” of the LIENSs laboratory.

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654

655 **List of Figures**

656 **Figure 1:** Map of the Bay of Biscay showing the location of the sampling stations of A) sub-
657 surface water environmental parameters (temperature, salinity, size-fractionated biomass of
658 chlorophyll *a*), and B) the mesozooplankton community. Isobaths 100 m (dotted line), 200 m
659 (solid line) and 500 m (dashed line) are drawn.

660 **Figure 2:** Results of the annual (2003-2013) spatial interpolation of sub-surface temperature
661 (°C), salinity (psu), picophytoplankton ($chl a < 3 \mu m$) biomass, nanophytoplankton ($3 \mu m$
662 $< chl a < 20 \mu m$), and microphytoplankton ($chl a > 20 \mu m$) biomass ($\mu g \cdot L^{-1}$) in the southern Bay
663 of Biscay.

664 **Figure 3:** Interannual variation of the mean relative biomass of surface size-fractionated
665 chlorophyll *a* between 2003 and 2013 in the southern Bay of Biscay; in black for the
666 picophytoplankton biomass ($< 3 \mu m$), in light grey for the nanophytoplankton biomass ($3-$
667 $20 \mu m$) and in dark grey ($> 20 \mu m$) for the microphytoplankton biomass.

668 **Figure 4:** Interannual variation of the mesozooplankton abundance between 2003 and 2013:
669 A) mean decadal and annual abundances ($ind \cdot m^{-3} \pm SD$) of the entire mesozooplankton
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671 organisms belonging to copepods, gelatinous, other holoplankton and meroplankton groups
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673 **Figure 5:** Annual spatial distribution of abundance ($ind \cdot m^{-3}$) for major families, genera or
674 species contributing for more than 1.5% in abundance to the copepods community between
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676 from each station throughout the decade.

677 **Figure 6:** Non-metric Multi-Dimensional Scaling (NMDS) on adult's copepods dataset in two
678 panels, above: the plot of the sampling points and below, plot of the copepod taxa (only
679 keeping the ones having correlation coefficient > 0.25). Spatial location of each station was

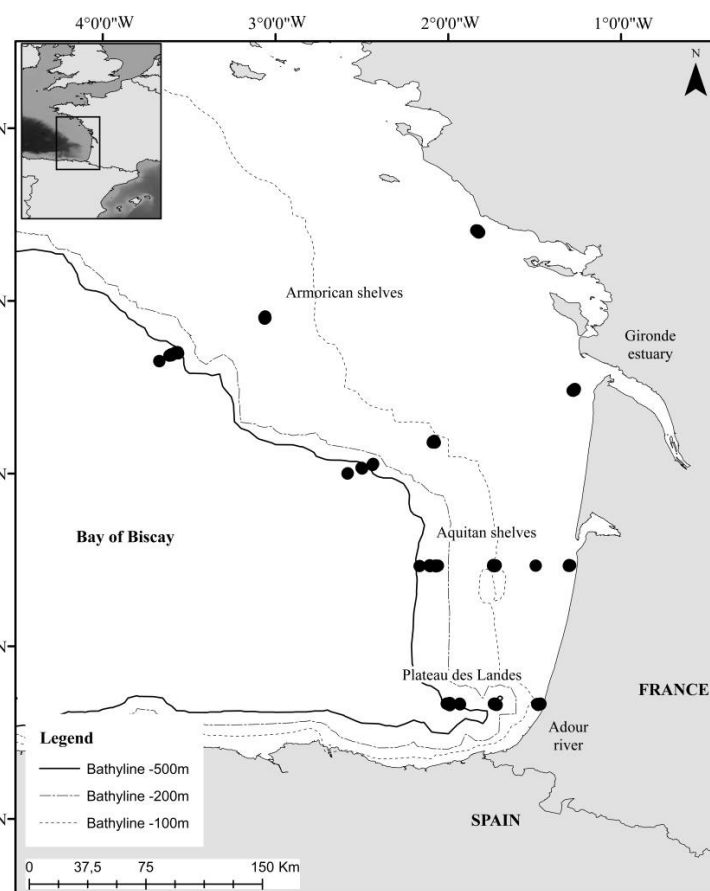
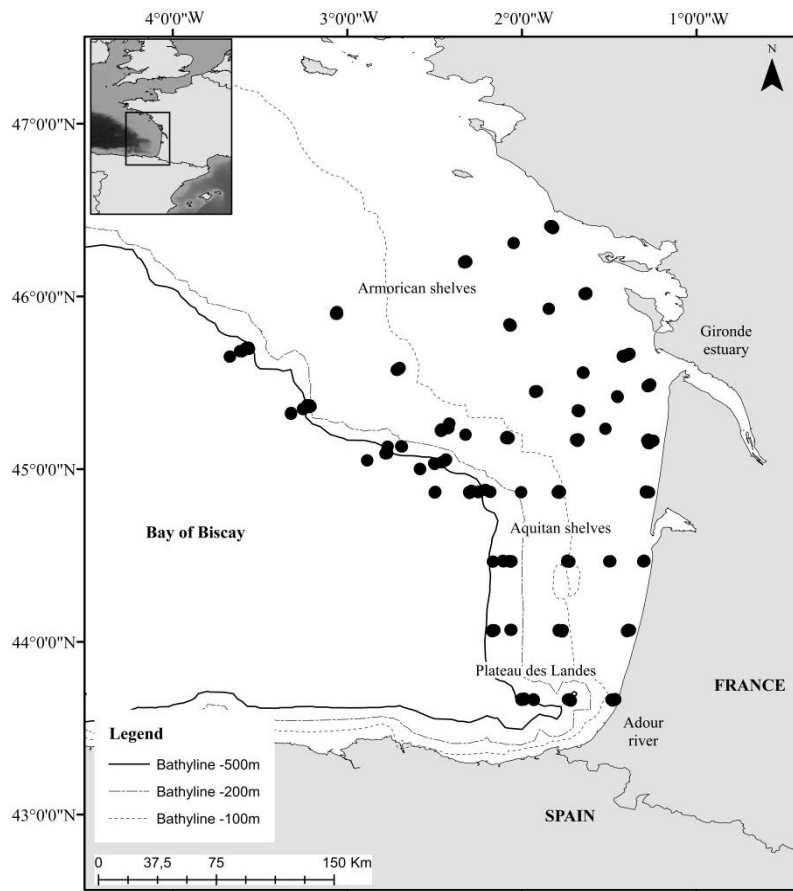
680 reported by their situation in the Bay of Biscay: coastal (C), continental shelf (Sh) and
681 continental slope (Sl). The choice of the three temporal groups was based on observations
682 following Fig. 4A.

683

684 Figure 7: Venn diagram based on a variation partitioning presenting the explained variability
685 of copepods community with two matrices, H and C. *H* was built with subsurface
686 temperature, deficit of potential energy and equivalent freshwater height and C with
687 microphytoplankton (> 20 μm) biomass at the subsurface (the chl*a* based-biomass > 20 μm).

688 The external square represents the whole variation of the copepod community. Each circle
689 represents the explanatory tables and values are the part of the variation explained by each
690 explanatory table. The fraction “Int” is the intersection of the amount of variation explained
691 by both types of explanatory variables. Statistically significant pure fraction of variation of
692 copepod community is given in supplemental table 1.

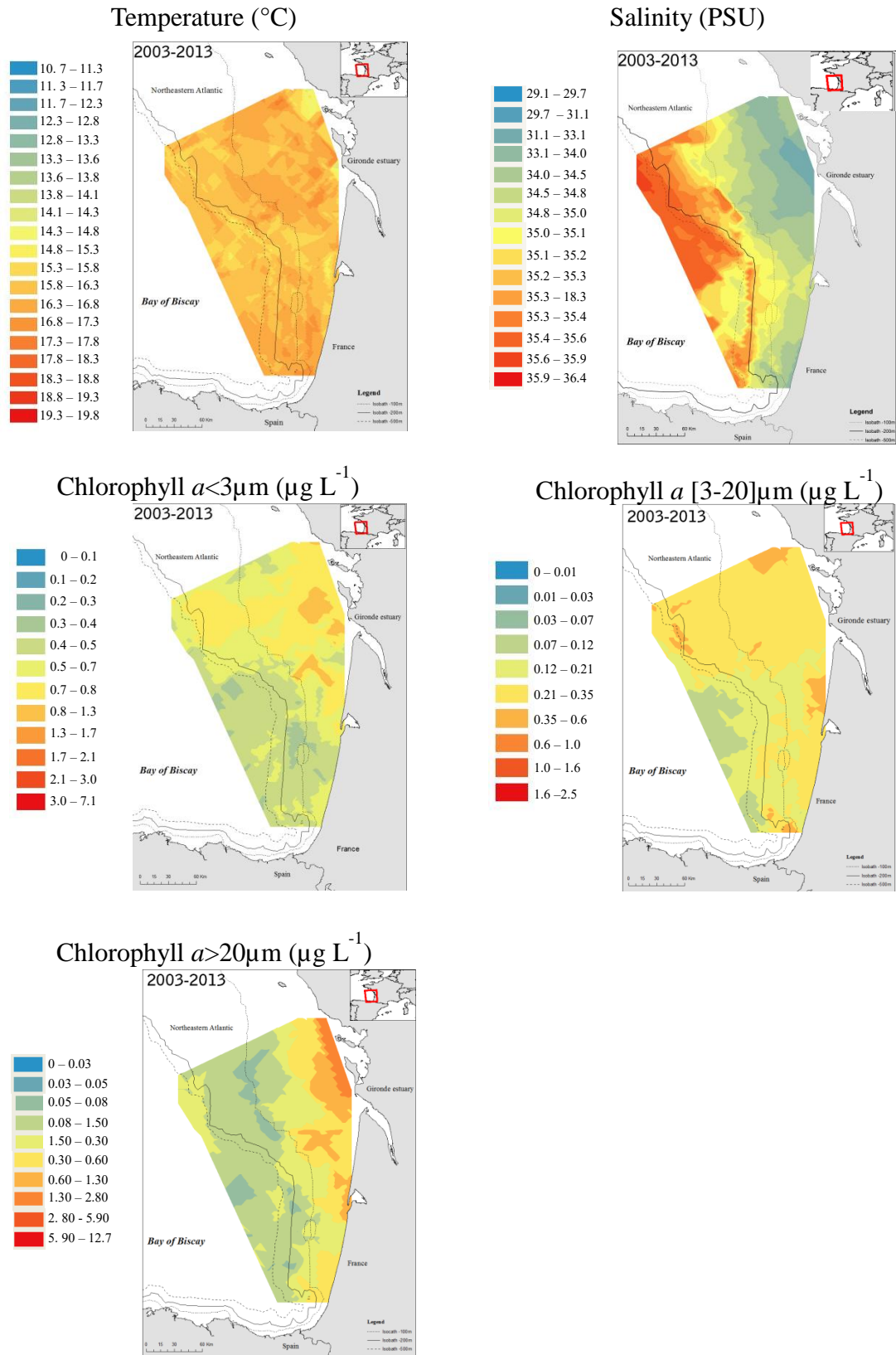
693 **Figure 1**



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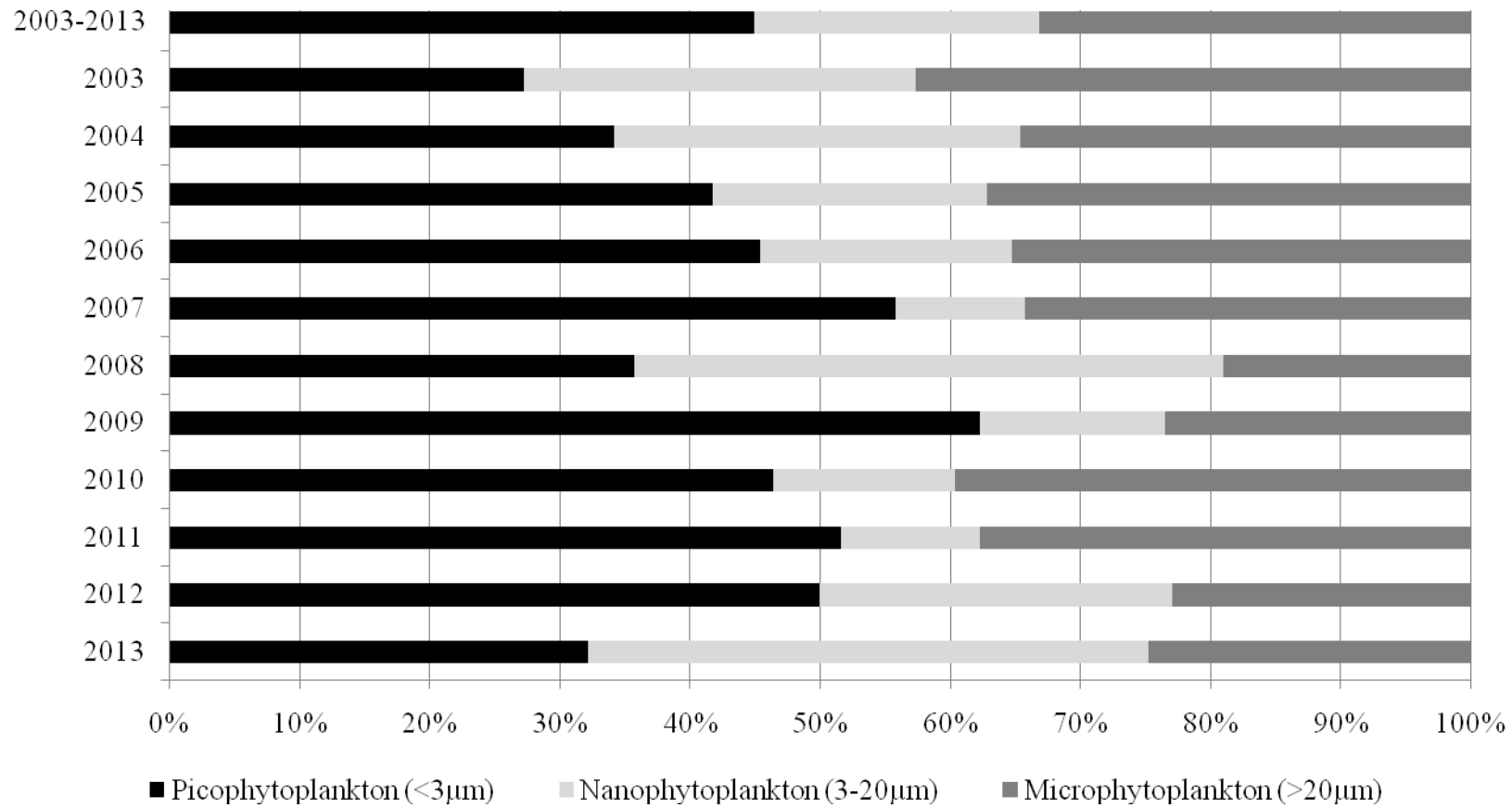
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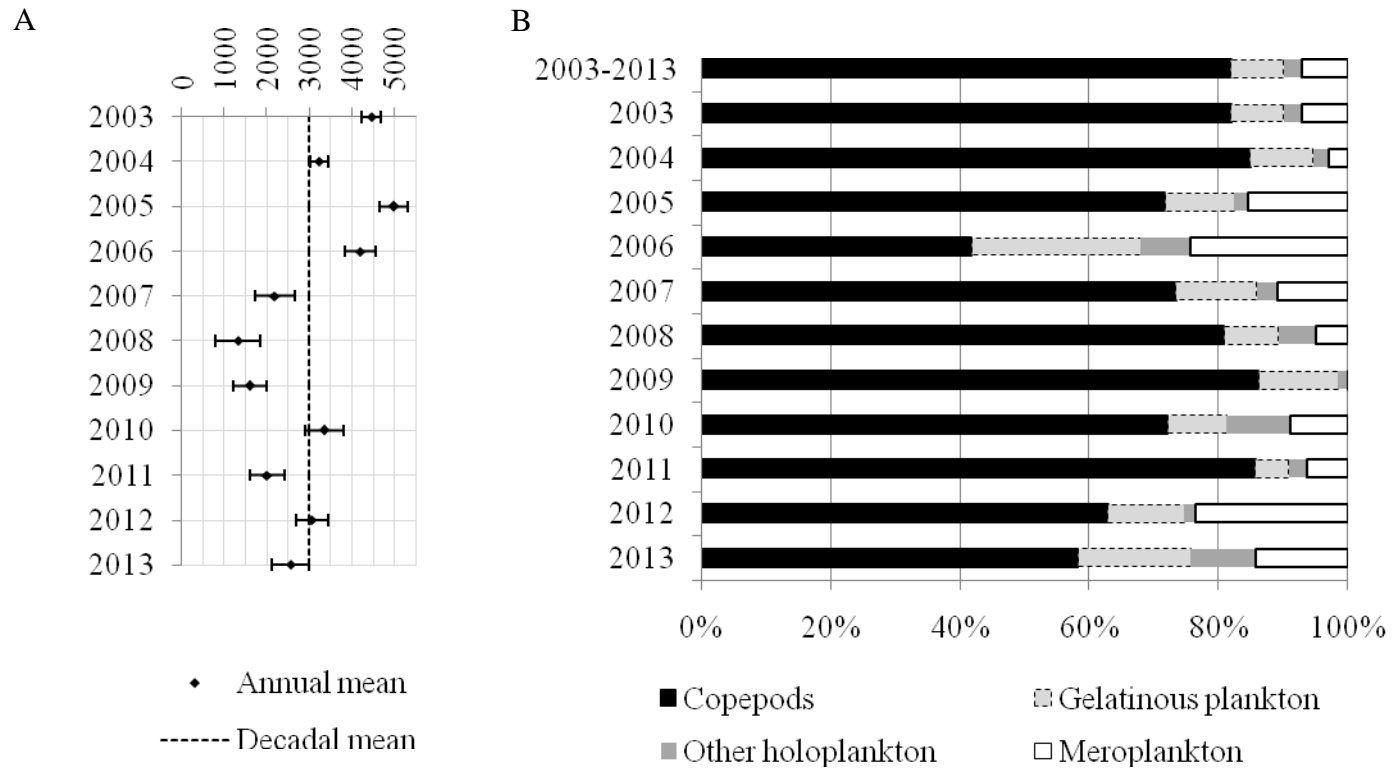
700 **Figure 3**

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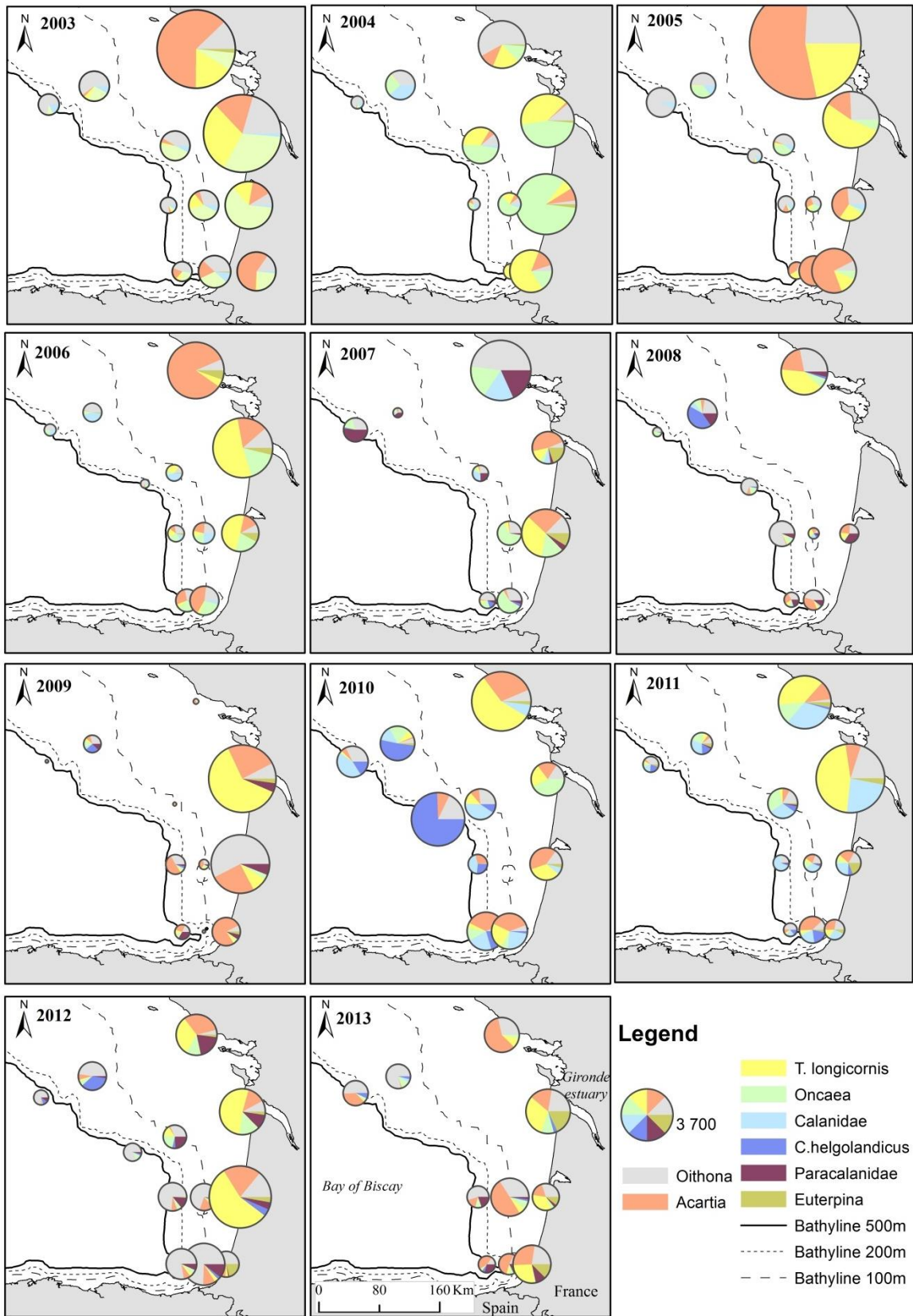
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703 **Figure 4**



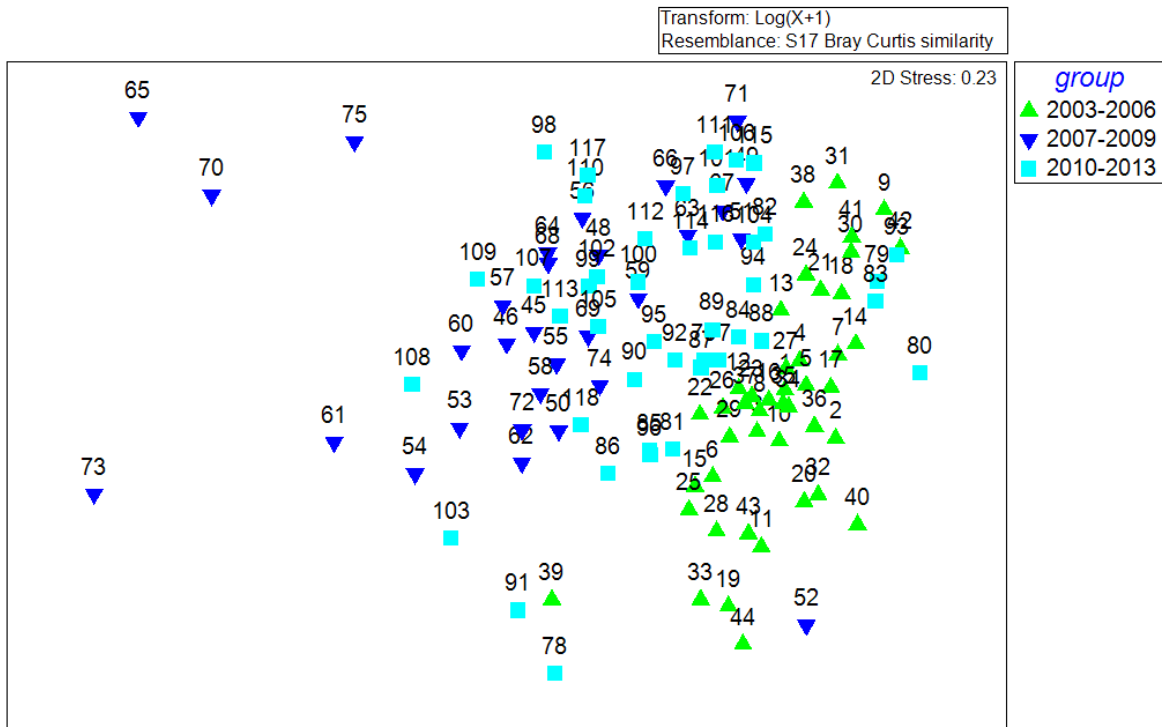
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705 **Figure 5**
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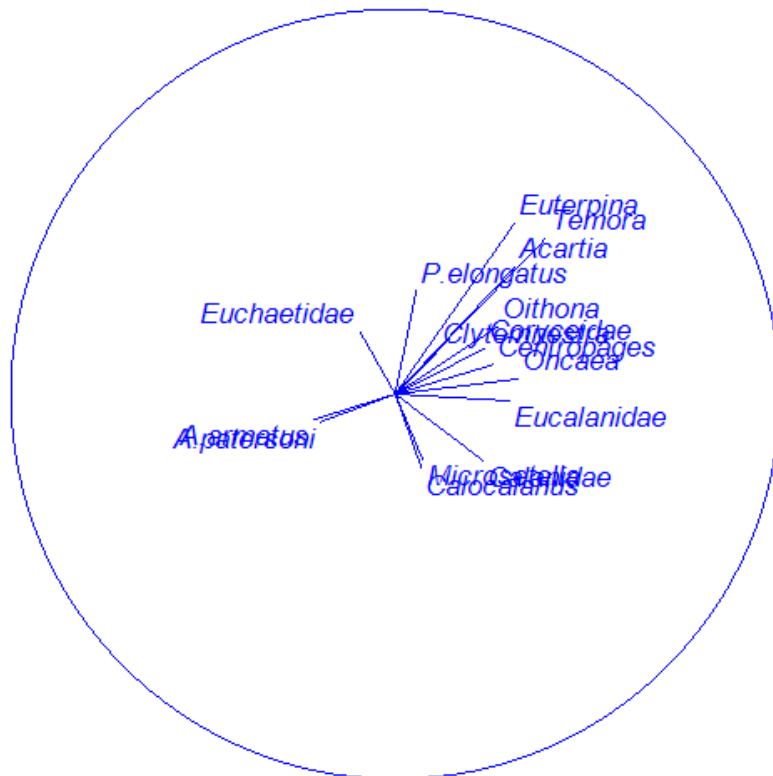


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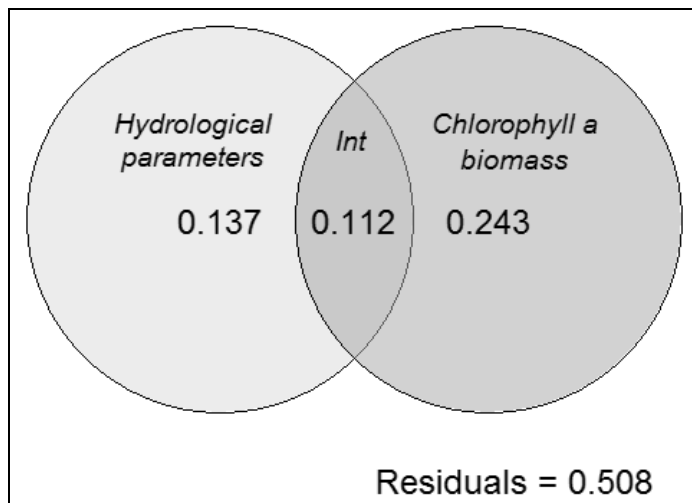
708 **Figure 6**
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715 **Figure 7**
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719 **Table 1:** Number of stations used in kriging (N.Station.Kriging) and multivariate (N.Station.VarPart) analysis per year start and the end dates of
 720 the surveys.
 721

722

723	Year	First Station	Last Station	N.Station.Kriging	N.Station.VarPart
724	2003	30 May	09 June	34	11
725	2004	29 April	10 May	26	10
726	2005	05 May	16 May	30	12
727	2006	02 May	13 May	30	11
728	2007	27 April	08 May	31	10
729	2008	27 April	09 May	31	9
730	2009	26 April	09 May	29	12
731	2010	26 April	09 May	29	10
732	2011	26 April	09 May	28	11
733	2012	27 April	13 May	31	12
734	2013	28 April	16 May	30	10

735

736 **Table 2:** Representative list of taxa and species of mesozooplankton found in the Bay of Biscay with distinction between major groups:
 737 copepods, gelatinous organisms (G), other holoplankton (H) and meroplankton (M) organisms. For adult copepods, order affiliation is
 738 represented by Calanoida (Ca), Poecilostomatoida (P), Cyclopoida (Cy) and Harpacticoida (Ha) following by their relative percentage of
 739 abundance between 2003 and 2013 on copepods community.
 740

Copepods	Taxonomic level	Order(%)	Other taxonomic groups	Taxonomic level	Group
<i>Acartia</i> spp.	Genus	Ca (19.3)	Appendicularia	Class	G
<i>Oithona</i> spp.	Genus	Cy (18.5)	Siphonophorae	Order	G
<i>Temora longicornis</i>	Species	Ca (17.4)	Indetermined Cnidaria	Phylum	G
Copepodites	Development stage	-(16)	Chaetognata	Phylum	G
<i>Oncaea</i> spp.	Genus	P (10.8)	Salpida	Order	G
Calanidae	Family	Ca (4.5)	Doliolida	Order	G
<i>Calanus helgolandicus</i>	Species	Ca (2.8)	Annelida larvae	Phylum	M
Paracalanidae	Family	Ca (2.2)	Cirripedia	Infra-class	M
<i>Euterpina acutifrons</i>	Genus	H (1.6)	Decapoda	Order	M
<i>Centropages</i> spp.	Genus	Ca (1.5)	Scaphopoda	Class	M
Copepods nauplii	Development stage	-(1.4)	Gastropoda larvae	Class	M

741

742 **Table 2:** continued

Copepods	Taxonomic level	Order(%)	Other taxonomic groups	Taxonomic level	Group
Coryceidae	Family	P (1.3)	Bivalvia larvae	Class	M
Eucalanidae	Family	Ca (1.3)	Enteropneusta	Class	M
<i>Metridia</i> spp.	Genus	Ca (0.4)	Anthozoa larva	Class	M
<i>Pseudocalanus elongatus</i>	Species	Ca (0.5)	Ectoprota	Phylum	M
<i>Microsetella</i> spp.	Genus	H (0.1)	Cyphonauta larvae	Development stage	M
<i>Candacia</i> spp.	Genus	Ca (0.1)	Bryozoa	Phylum	M
<i>Pleuromamma</i> spp.	Genus	Ca (0.1)	Echinodermata larvae	Phylum	M
<i>Calocalanus</i> spp.	Genus	Ca (0.1)	Hydrozoa	Class	M
<i>Clytemnestra</i> spp.	Genus	H (<0.05)	Scyphozoa	Class	M
<i>Aegistus</i> spp.	Genus	H (<0.05)	Ostracoda	Class	H
<i>Aetideus armatus</i>	Species	Ca (<0.05)	Isopoda	Order	H
Euchaetidae	Family	Ca (<0.01)	Mysida	Order	H
Rhincalanidae	Family	Ca (<0.01)	Amphipoda	Order	H
<i>Anomalocera patersoni</i>	Species	Ca (<0.01)	Euphausiacea	Order	H
			Cladocera	Infraorder	H
			Chordata	Phylum	H
			Cumacea	Order	H
			Foraminifera	Phylum	H

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744 **Table 3:** Results of the similarity percentage (SIMPER) analysis based on spring adult copepods matrix in the Bay of Biscay between 2003 and
 745 2013.
 746

Group [2003-2006]		Group [2007-2009]		Group [2010-2013]	
Average similarity: 66.83		Average similarity: 52.32		Average similarity: 59.89	
Genus/Species/Family	Contrib%	Genus/Species/Family	Contrib%	Genus/Species/Family	Contrib%
<i>Oithona</i> spp.	22.4	<i>Oithona</i> spp.	24.37	<i>Oithona</i> spp.	24.02
<i>Oncaea</i> spp.	17.8	<i>Acartia</i> spp.	17.96	<i>Acartia</i> spp.	19.21
<i>Acartia</i> spp.	13.04	<i>Oncaea</i> spp.	17.1	<i>Temora longicornis</i>	13.24
<i>Temora longicornis</i>	12.2	Paracalanidae	15.16	<i>Oncaea</i> spp.	10.96
Eucalanidae	11.94	<i>Temora longicornis</i>	7.77	<i>Calanus helgolandicus</i>	7.95
Calanidae	10.54	Coryceidae	3.76	Calanidae	6.97
<i>Centropages</i> spp.	5.31	<i>Calanus helgolandicus</i>	3.04	Paracalanidae	4.84
		Calanidae	2.9	<i>Euterpina acutifrons</i>	4.11

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749 **Table 4:** Results of the dissimilarity percentage (SIMPER) analysis based on spring adult copepods matrix in in the Bay of Biscay between 2003
750 and 2013.
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Groups [2003-2006] and [2007-2009]		Groups [2003-2006] and [2010-2013]		Groups [2007-2009]and [2010-2013]	
Average dissimilarity: 52.01		Average dissimilarity: 45.51		Average dissimilarity: 46.88	
Genus/Species/Family	Contrib%	Genus/Species/Family	Contrib%	Genus/Species/Family	Contrib%
Eucalanidae	11.04	Eucalanidae	11.29	<i>Temora longicornis</i>	10.09
Paracalanidae	10.49	<i>Calanus helgolandicus</i>	9.44	Calanidae	9.78
<i>Temora longicornis</i>	9.62	Calanidae	8.55	<i>Calanus helgolandicus</i>	9.05
Calanidae	8.59	<i>Temora longicornis</i>	8.46	Paracalanidae	8.79
<i>Acartia</i> spp.	8.2	<i>Acartia</i> spp.	7.85	<i>Acartia</i> spp.	8.39
<i>Oncaea</i> spp.	7.2	<i>Oncaea</i> spp.	7.67	<i>Euterpina acutifrons</i>	7.48
<i>Centropages</i> spp.	6.87	Paracalanidae	7.15	<i>Oncaea</i> spp.	7.38
Coryceidae	6.29	<i>Euterpina acutifrons</i>	7.1	<i>Oithon</i> aspp.	6.31
<i>Oithona</i> spp.	5.77	<i>Centropages</i> spp.	6.92	<i>Centropages</i> spp.	5.99
<i>Euterpina acutifrons</i>	4.98	Coryceidae	6.49	Coryceidae	5.97
<i>Metridia</i> spp.	4.54	<i>Metridia</i> spp.	4.79	<i>Metridia</i> spp.	4.71
<i>Calanus helgolandicus</i>	4.4	<i>Oithona</i> spp.	3.74	<i>Pseudocalanus elongatus</i>	3.46
<i>Candacia armata</i>	2.65	<i>Microsetella</i> spp.	3.06	<i>Candacia armata</i>	3.04

754 **Table 5:** Selection of the most influential variables explaining the most the patterns of copepod community using a forward selection. The eigen
 755 values sums (a measure of explained variance, R^2), “ F ” statistic and p-values were reported.

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Variables order	Eigenvalues sums	F	
Sub-Surface Chla>20µm (Microphytoplankton) Based biomass	0.35	65.28	**
Equivalent fresh water height	0.42	15.47	**
Deficit of potential energy	0.45	7.50	*
Sub-surface temperature	0.49	8.84	**

Codes:0 ‘***’;0.001 ‘**’;0.01 ‘*’

760

763 **List of supplemental figures**

764 **Supplemental figure S1:** Annual interpolated maps of results of the annual spatial
765 interpolation of sub-surface temperature measures ($^{\circ}\text{C}$), performed for each year between
766 2003 and 2013 in the southern Bay of Biscay. Thanks to a non-parametric multiple pairwise
767 comparison of mean ranks applied for all years, significant differences of sub-surface
768 temperature according to year are indicated by letters.

769 **Supplemental figure S2:** Annual interpolated maps of results of the spatial interpolation of
770 sub-surface salinity (PSU_{psu}) performed for each year between 2003 and 2013, in the
771 southern Bay of Biscay.

772 **Supplemental figure S3:** Annual interpolated maps of results of the spatial interpolation of
773 sub-surface picophytoplankton ($\text{chl}a < 3\mu\text{m}$) biomass ($\mu\text{g}\cdot\text{L}^{-1}$) performed for each year
774 between 2003 and 2013, in the southern Bay of Biscay. Thanks to a non-parametric multiple
775 pairwise comparison of mean ranks applied for all years, significant differences of sub-surface
776 picophytoplankton biomass according to year are indicated by letters.

777 **Supplemental figure S4:** Annual interpolated maps of results of the spatial interpolation of
778 sub-surface nanophytoplankton ($[3-20]\mu\text{m}$) biomass ($\mu\text{g}\cdot\text{L}^{-1}$) performed for each year
779 between 2003 and 2013, in the southern Bay of Biscay. Thanks to a non-parametric multiple
780 pairwise comparison of mean ranks applied for all years, significant differences of sub-surface
781 nanophytoplankton biomass according to year are indicated by letters.

782 **Supplemental figure S5:** Annual interpolated maps of Results of the spatial interpolation of
783 sub-surface microphytoplankton ($\text{chl}a > 20\mu\text{m}$) biomass ($\mu\text{g}\cdot\text{L}^{-1}$) performed for each year
784 between 2003 and 2013, in the southern Bay of Biscay. Thanks to a non-parametric multiple
785 pairwise comparison of mean ranks applied for all years, significant differences of sub-surface
786 microphytoplankton biomass according to year are indicated by letters.

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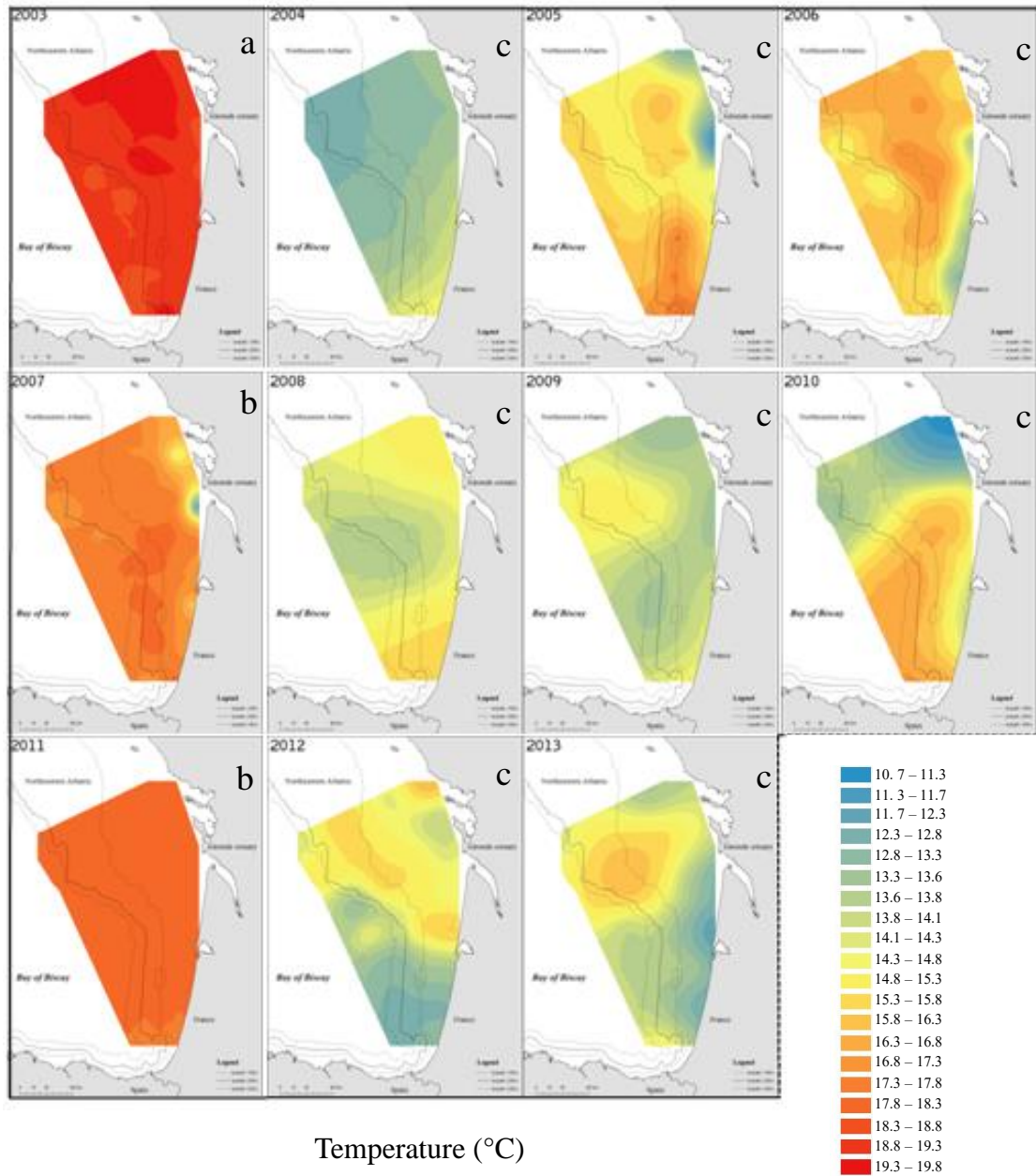
788 **Supplemental figure S6:** Spatial distribution of abundance (ind.m⁻³) of copepods nauplii,
789 copepodites and Harpacticoida, Poecilostomatoida, Cyclopoida, Calanoida (orders) between
790 2003 and 2013 in the southern Bay of Biscay. The size of the pies is proportional to the total
791 abundance of from each station throughout the decade. Significant differences in abundance
792 according to year are indicated by letters.

793 **Supplemental figure 1:**

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798 **Supplemental figure 2:**

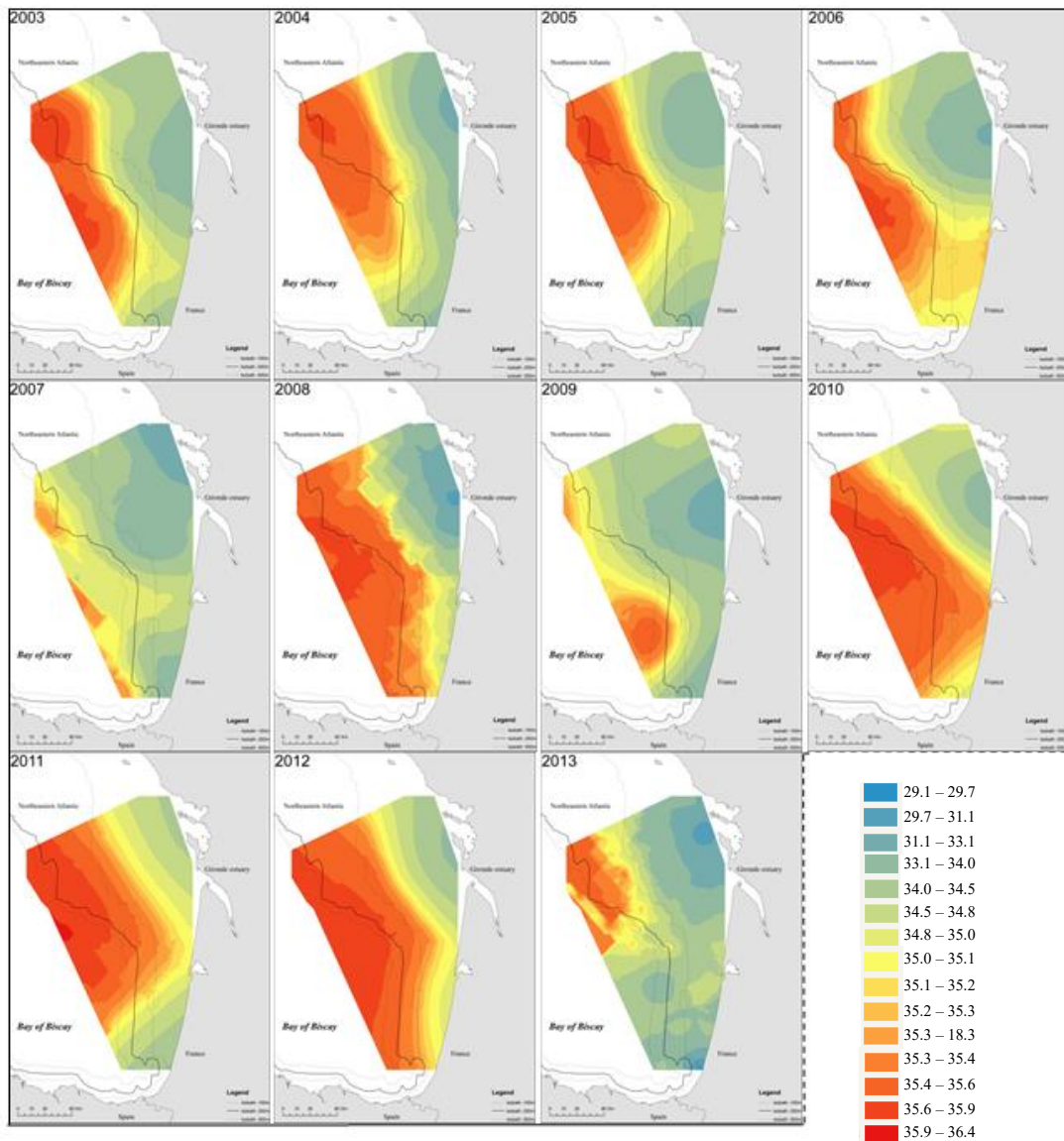
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Salinity (PSU)

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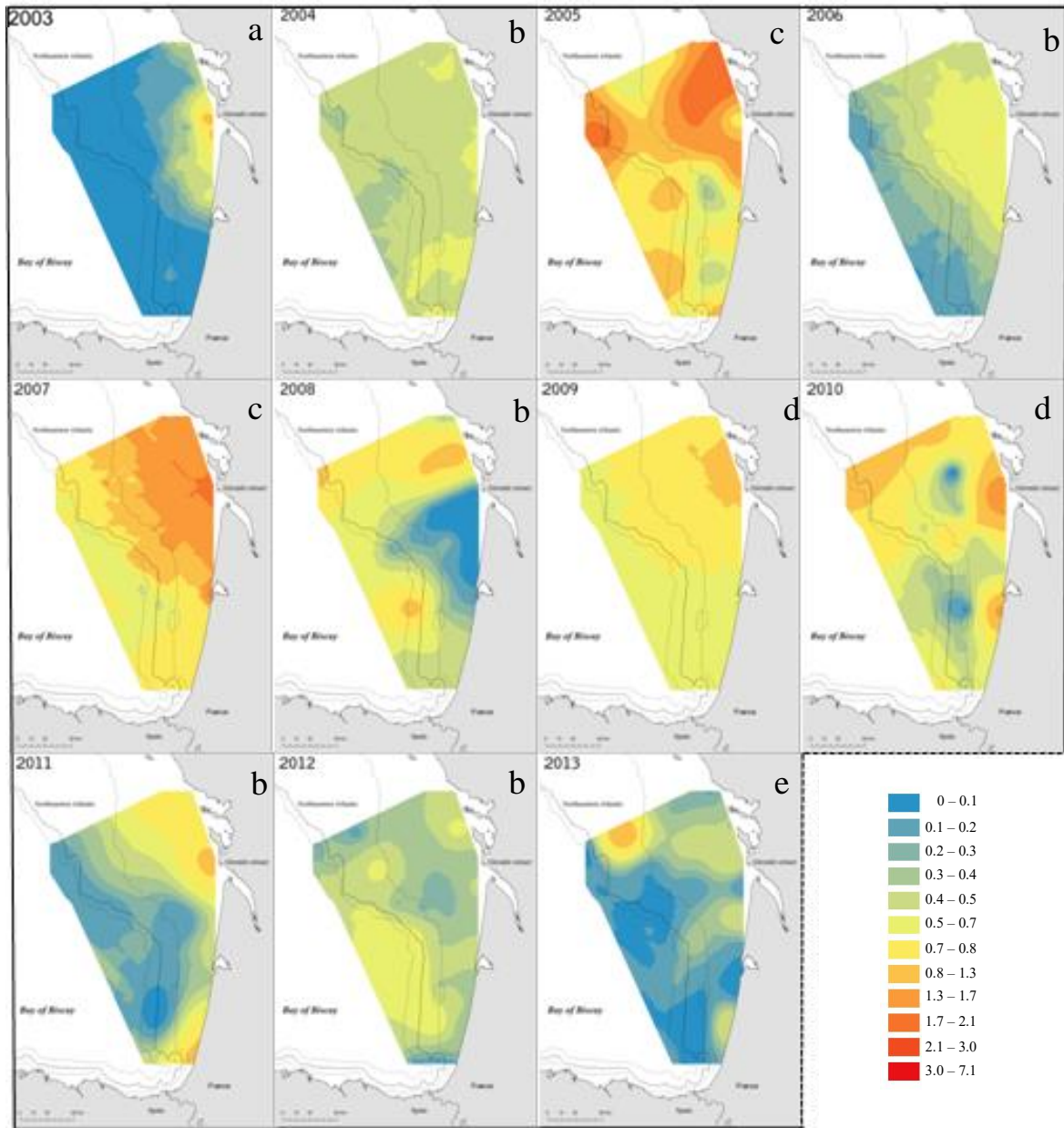
805 **Supplemental figure 3:**

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Chlorophyll *a* < 3 μm ($\mu\text{g L}^{-1}$)

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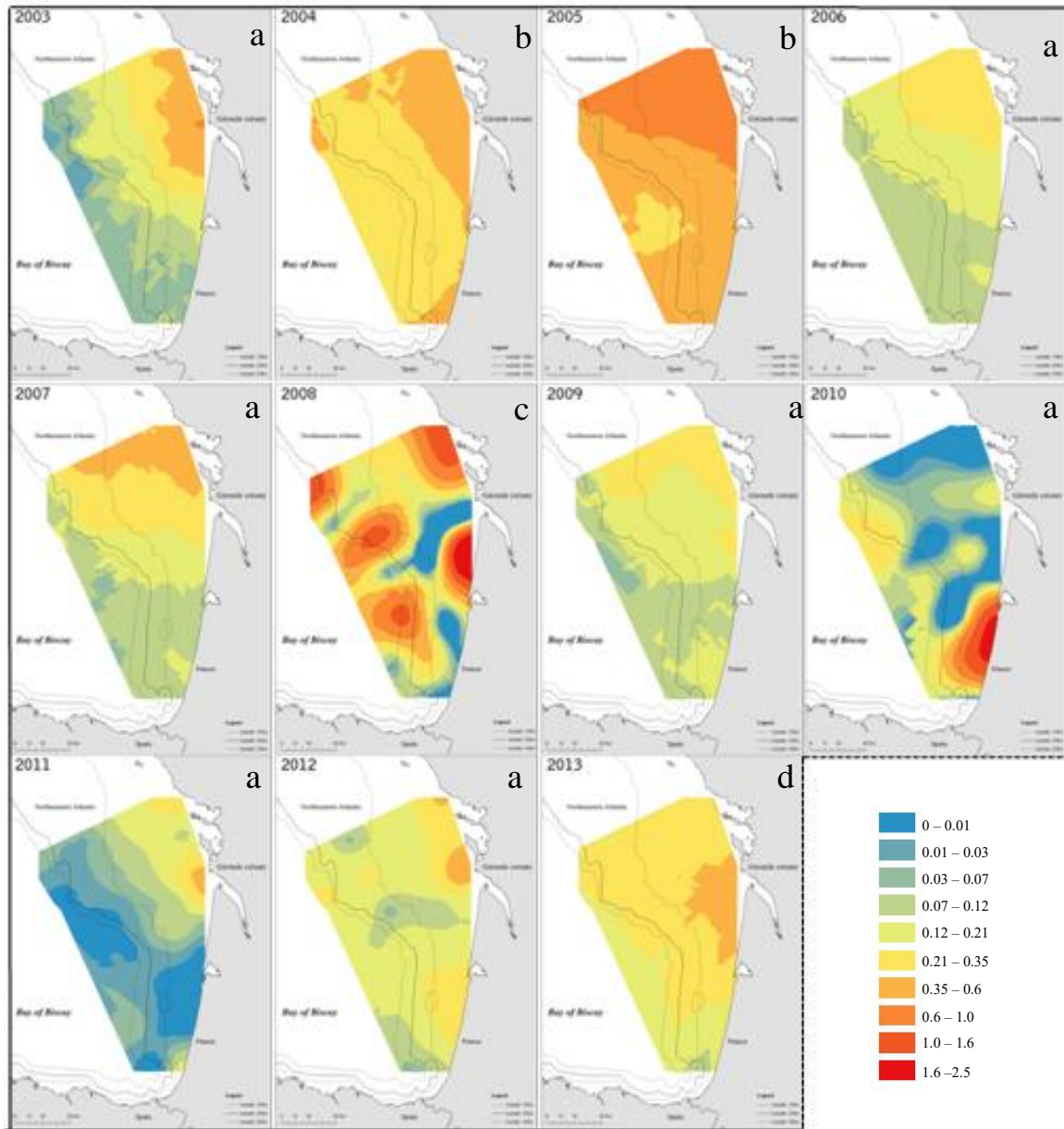
812 **Supplemental figure 4:**

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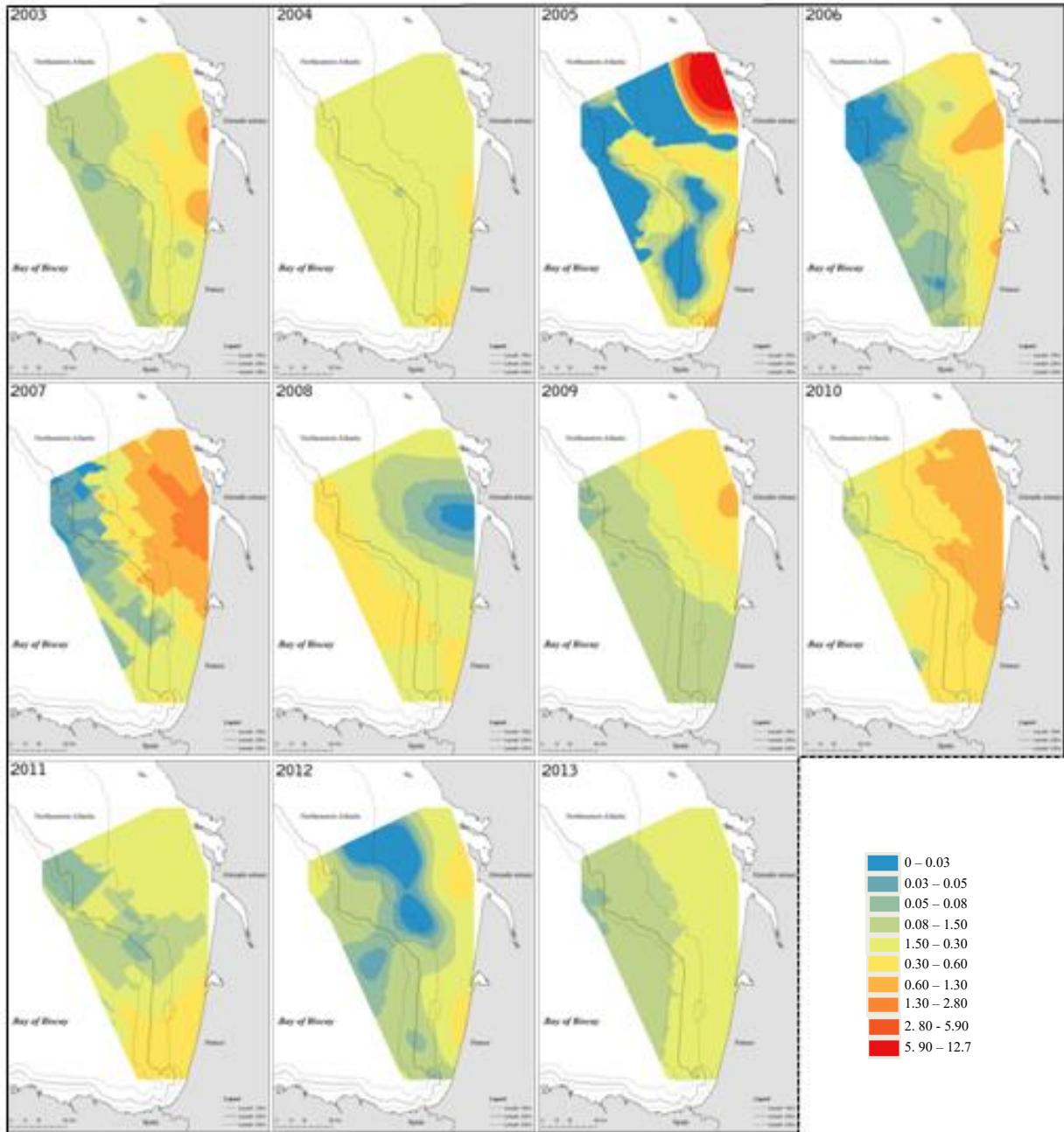


Chlorophyll *a* [3-20] μm ($\mu\text{g L}^{-1}$)

817 **Supplemental figure 5:**

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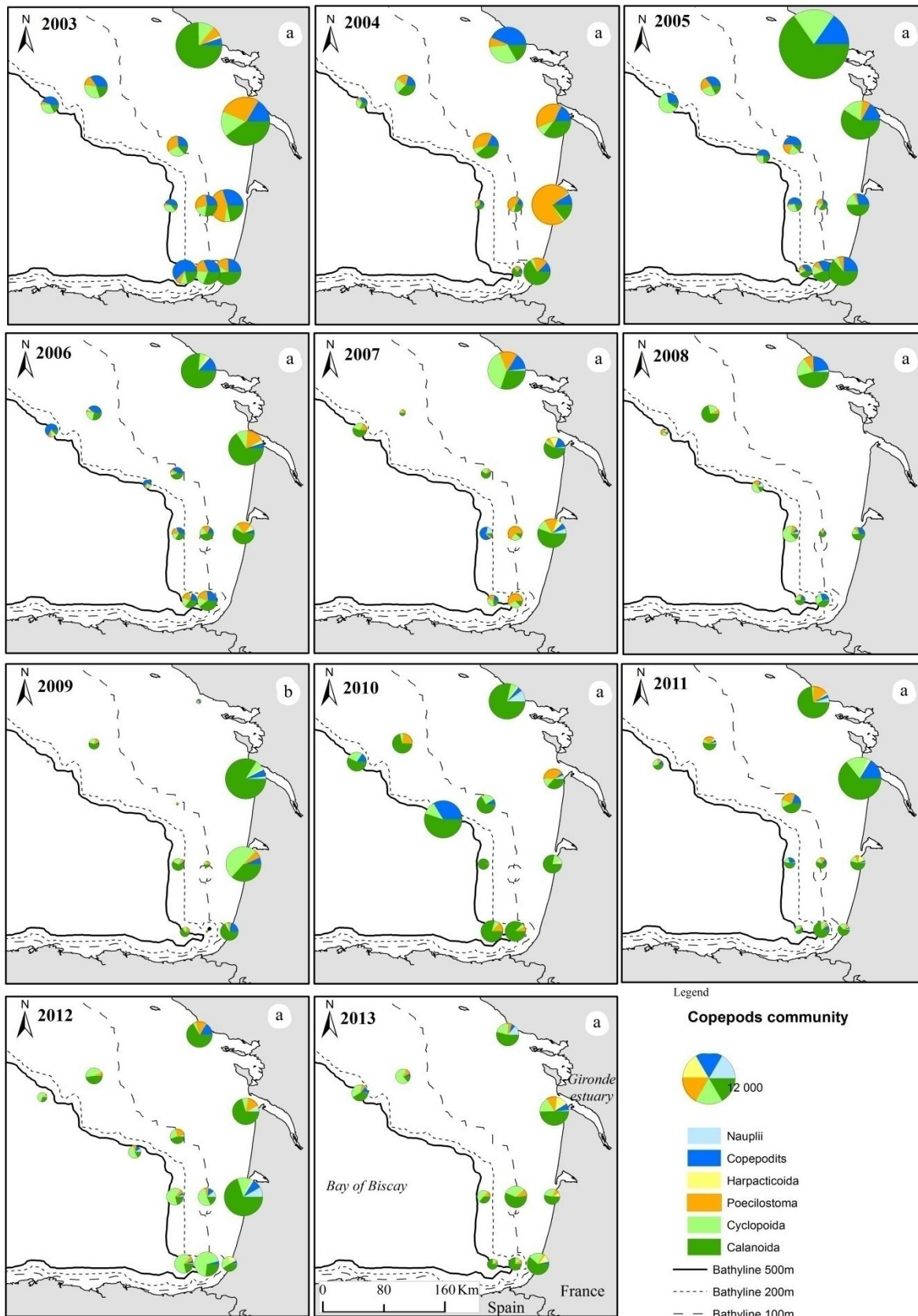
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Chlorophyll *a*_{>20 μm} ($\mu\text{g L}^{-1}$)

820 **Supplemental figure 6:**

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824 **List of supplemental figure :**

825 **Supplemental table 1:** Identification of copepod community drivers by variation partitioning
 826 analysis based on ordination analyses. Matrix H contained environmental variables (sub-
 827 surface temperature values, deficit of potential energy and equivalent freshwater height),
 828 matrix C contained the chlorophyll *a* >20 µm biomass (microphytoplankton biomass).
 829 Multivariate analysis, covariables, component of variation, eigenvalues and statistical p-
 830 values are reported. Explanatory matrix / component of variation; H/h – hydrological
 831 parameters, C/c – chlorophyll *a* based biomass.

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Analysis	Constraining variable	Covariate	Component of the variation	Eigenvalues sums	<i>p-value</i>
ACP	-	-	-		
RDA	H		h + hc	0.24	***
RDA	C		c + hc	0.35	***
RDA	H+C		h + c + hc	0.49	***
RDAp	H	C	h	0.13	***
RDAp	C	H	c	0.24	**

Codes for p-values: 0 '***';0.001 '**';0.01 '*'

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