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PALAEO

Phytoliths as paleoenvironmental indicators, West Side Middle Awash Valley, Ethiopia

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Abstract

Phytolith assemblage analysis offers the potential to refine our knowledge of paleoecosystems where grasses and sedges predominate. In this work, Holocene and Pleistocene sediments from an arid tropical region in Ethiopia have been analyzed for their phytolith content, presented as detailed counts and diagrams according to the Twiss classification. The aim is to test the usefulness of phytolith assemblages to indicate paleoenvironments at Middle Awash, where fossil pollen grains are poorly preserved in sediments that yielded abundant archaeological remains. The vegetation in the Middle Awash subdesertic valley is currently a shrub steppe dominated by C4 grasses adapted to arid conditions, with a narrow riparian forest limited to the Awash River. Our results show that modern surface samples, Holocene and Pleistocene sediments contain well-preserved and different phytolith assemblages, and therefore that no translocation processes from modern soil to geological strata seem to occur. Fossil records and modern assemblages are interpreted using phytolith ratios to estimate the density of the tree cover, the aridity and the proportion of C3 versus C4 grasses, as applied to phytolith assemblages from North America and West Africa. The phytolith assemblages from modern soil samples correctly reflect the proportion of trees and shrubs versus grasses, different in the riparian vegetation and the shrub steppe. Modern phytolith assemblages appear to be a mixed signature of local and regional vegetation. Phytolith analysis of the Holocene sample suggests a grassland, where the grass community is constituted by the Chloridoideae subfamily, adapted to warm and dry conditions and where C3-Pooideae cover the highlands. Phytolith analysis of the Pleistocene sample evidences grassland formation with scattered woody elements, where C4-Panicoideae grasses, adapted to warm and humid conditions dominate the grass cover. However, these conclusions need to be confirmed by more complete study on phytolith assemblages from modern vegetation from Ethiopia. © 1999 Elsevier Science B.V. All rights reserved.

Keywords: phytoliths; paleoenvironment; Middle Awash; Ethiopia

1. Introduction

Phytoliths are microscopic opal-A particles (Jones and Segnit, 1969) that precipitate in cells and/or

between cells of living plant tissues. They occur in many plant families (Piperno, 1988), but they are especially abundant, diverse and distinctive in the Poaceae (Twiss et al., 1969; Twiss, 1987, 1992). Due to this distinctiveness, phytolith analysis allows for separation between forest and grassland, C3 and C4 grasslands and, among the C4 grasslands, those

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dominated by the subfamily Chloridoideae and those by Panicoideae (Twiss, 1987; Fredlund and Tieszen, 1994, 1997; Alexandre et al., 1997). Fossil phytolith assemblages from soils and lake sediments have been used to reconstruct paleovegetation patterns, especially forest/grassland ecotones (Alexandre et al., 1997). Further, trends in grassland dynamics have been investigated through the evolution of the Poaceae associations (Kurman, 1985; Fredlund et al., 1985; Bartolome et al., 1986; Fisher et al., 1995; Fredlund and Tieszen, 1997).

This paper presents a preliminary study on phytolith assemblages, applied to paleoenvironmental reconstruction from the West Side of the Middle Awash Valley (Ethiopia), an area which is of particular interest due to its richness in archaeological and anthropological sites. This work was designated as a complementary study to fossil pollen grain analysis, which has shown that grasses were an abundant component of the vegetation in the Ethiopian Rift at least since the Pliocene (Bonnefille, 1995). Because pollen grains are poorly preserved in Middle Awash fossil sediments and because grass pollen is indistinguishable at the generic level, phytolith assemblage analysis have been tried as a new and complementary paleoecological indicator. Here we report on a pilot investigation designated to test whether samples from Holocene and Pleistocene deposits yielded phytoliths, and how they differ from the modern ones.

2. Environmental setting of the Middle Awash

The West Side of the Middle Awash hominid site (10°30'N, 40°25'E) is located in the southern Afar Rift, straddling the Awash River (Fig. 1). At present, the Middle Awash basin structural context leads to climatic differences between the study area and the northwestern Ethiopian escarpment.

The Ethiopian plateau, between 1800 and 2500 m of altitude, is characterized by annual rainfall higher than 1400 mm/yr, and by a mean annual temperature of about 16°C (F.A.O., 1965). It is the domain of afro-montane forest with scattered *Juniperus procera* forest on the drier slopes of the escarpment up to 3300 m of elevation (Friis, 1986). Above the forest altitudinal zone, montane grasslands include grasses

belonging to the Pooideae (e.g. *Poa*, *Agrostis*, *Avena*) and Panicoideae (e.g. *Andropogon*, *Pennisetum*) subfamilies, Cyperaceae and scattered shrubs and trees (Sermolli, 1957).

On the rift escarpment of the northwestern Ethiopian plateau, between 1500 and 800 m, mean annual rainfall is mostly between 500 and 800 mm/yr and the mean annual temperature about 26°C. Evergreen and semi-evergreen bushland and thicket occur on steep slopes, with trees and succulents. Grasses belong to subfamilies Panicoideae (*Chrysopogon*, *Cenchrus*, *Brachiaria*), Arundinoideae (*Aristida*) and Chloridoideae (*Chloris*, *Schmidtia*) (Gillett, 1941; Gilliland, 1952; Sermolli, 1957; White, 1983).

In the Awash valley, below 500 m, rainfall ranges from 250 to 500 mm/yr, and the mean annual temperature is about 30°C. The climate is characterized by two rainy seasons, respectively, in March–April and July–August–September (Delliquadri, 1958). With a mean annual evapotranspiration potential above 1055 mm, the region is considered arid. *Acacia*–*Commiphora* scattered bushland, thicket and the shrub steppe occur here. Grass cover of the steppe is dominated by C4 grasses belonging to the Chloridoideae subfamily (e.g. *Chloris* and *Cynodon*) and, in the small temporary pools, by the Panicoideae subfamily (e.g. *Chrysopogon*, *Cenchrus*, *Andropogon*) (Gillett, 1941; Gilliland, 1952; Sermolli, 1957; White, 1983).

The Awash River rises on the northwestern plateau of central Ethiopia at 3000 m of altitude, and crosses these different climatic and vegetation zones in its descent. Its riparian vegetation, restricted to a narrow riverine forest zone in the Middle Awash Valley, is mostly dominated by shrubs, trees such as *Terminalia*, *Ficus*, *Tamarix*, *Salvadora* and climbers (Bonnefille et al., 1987). The herbaceous layer is mostly constituted by C4 grasses, such as Chloridoideae (*Chloris*, *Lintonia*, *Cynodon*, *Sporobolus*) and Panicoideae (*Pennisetum*, *Sorghum*) and also by C3 grasses belonging to the Arundinoideae subfamily, such as *Phragmites* (Sermolli, 1957). Table 1 gives the list of the grass subfamilies and dominant genera for the western Middle Awash Valley reported from the literature. Attribution to the C3 or C4 photosynthetic pathway is compiled from Smith and Brown (1973), Brown (1977) and Watson et al. (1985).

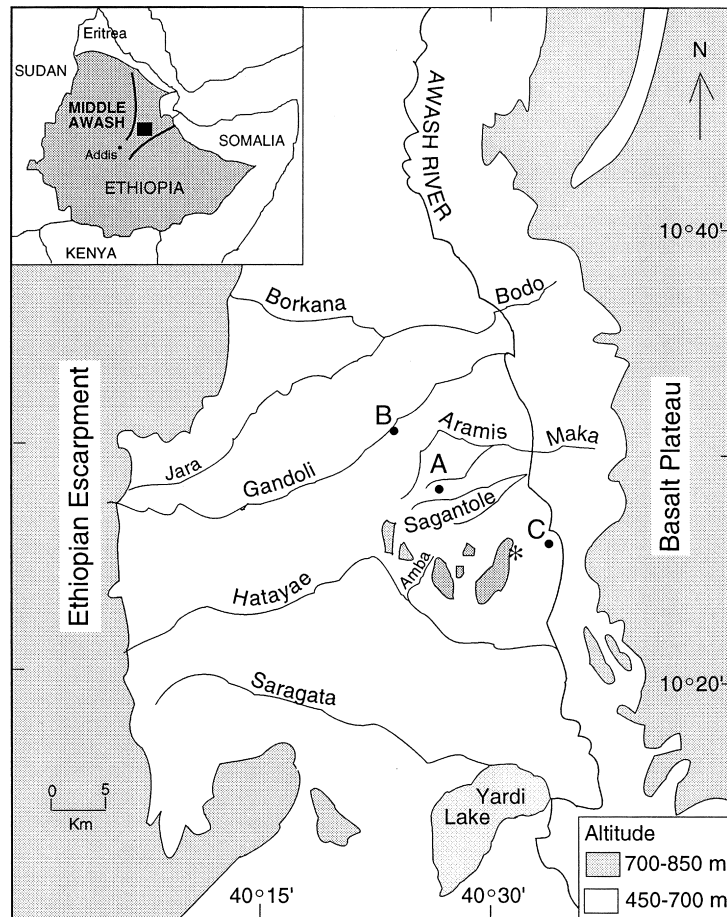


Fig. 1. Location of the samples in the Middle Awash Valley (Ethiopia). A = surface soil sample MA94-103; B = stream mud sample MA94-106; C = surface soil sample MA94-101; * = Aduma fossil site.

3. Field and laboratory methods

3.1. Field sampling

Sediment samples for pollen and phytolith analysis were collected in the Middle Awash valley during the 1994 field season of the Middle Awash Project (White et al., 1994). Five samples were analyzed for this preliminary study, one from Holocene strata, one from Pleistocene, and three from modern surface soil (Table 2). The sampling method for the modern soil consists of collecting at more than twenty spots the upper one cm of the litter deposited at the soil surface. The individual sub-samples collected at random over an area of about hundred square meters

are mixed together. This sampling method provides data on the average modern pollen rain over a few years that correctly reflect the different vegetation types encountered in the lower Awash valley (Bonnefille et al., 1987), and more specifically in the Ethiopian plateau (Bonnefille et al., 1993). The first modern sample (MA94-103, 10°28'N, 40°27'E) (A, Fig. 1), collected under *Acacia* trees near the course of the Aramis stream, is from the subdesertic steppe. The second modern sample (MA94-106, 10°31'N, 40°26'E) (B, Fig. 1), consists of the mud collected after one flood of the Ganduli River, a tributary of the Awash River. The third modern sample (MA94-101, 10°27'N, 40°30'E) (C, Fig. 1), was collected in a swampy area surrounded by trees at the margin of

Table 1

Dominant genera and subfamilies among the modern grasses of the Awash valley, Ethiopia after Sermolli (1957), Gillett (1941), Gilliland (1952); C3 and C4 attribution after Smith and Brown (1973), Brown (1977) and Watson et al. (1985)

Subfamily	Genus	C3/C4
Pooideae	<i>Poa</i>	C3
	<i>Agrostis</i>	C3
	<i>Avena</i>	C3
Chloridoideae	<i>Chloris</i>	C4
	<i>Cynodon</i>	C4
	<i>Lintonia</i>	C4
	<i>Schmidtia</i>	C4
	<i>Sporobolus</i>	C4
Panicoideae	<i>Andropogon</i>	C4
	<i>Brachiaria</i>	C4
	<i>Cenchrus</i>	C4
	<i>Chrysopogon</i>	C4
	<i>Pennisetum</i>	C4
	<i>Sorghum</i>	C4
Arundinoideae	<i>Phragmites</i>	C3
	<i>Aristida</i>	C4

the riparian forest, on the west bank of the Awash river, near Aduma.

The geological samples were taken from outcrops at the Aduma archaeological excavation (Fig. 1). The Pleistocene sample MA94-097 was collected from the second shell bed which contains Middle Stone Age archaeological artefacts in the 'Ardu' beds, exposed at the A4 archaeological trench (A. Brooks, pers. commun.). The sample MA94-098 comes from the black montmorillonite layer, lying on the top of

the beds at the same locality. Its age is most probably Holocene.

3.2. Chemical treatment

Approximately 10 g of sediment sample was used for each phytolith extraction. The phytoliths were extracted by dissolution of carbonates using acetic acid buffered at pH 3, oxidation of the organic matter, using 30% hydrogen peroxide at 90°C until reaction subsides, and mechanical removal of clays. Densimetric separation of phytoliths from residue is achieved by separation in a zinc bromide solution with a density of 2.3 (Kelly, 1990). This light fraction, including opal phytoliths, diatoms and volcanic ashes, was mounted on microscope slides, using glycerin as a mounting medium. Microscopic observations were done at $\times 600$ magnification.

4. Phytolith classification and counting

The extracted phytoliths were counted and classified according to the classification of Twiss (Twiss et al., 1969; Twiss, 1992), improved and completed by phytolith shape descriptions of Mulholland (1986, 1989), Mulholland and Rapp (1992) and Fredlund and Tieszen (1994). The different morphotypes identified in the Middle Awash samples are illustrated in Plate I. Short-cell phytoliths from grass epidermis are related to three Poaceae subfamilies (Metcalfe, 1960; Brown, 1984; Twiss, 1987; Mulholland, 1989). The rondel and rectangle morphotypes fol-

Table 2

List of Middle Awash samples treated for phytolith analysis

Samples	Sites	Coordinates or stratigraphic position	Nature of samples
<i>Modern</i>			
MA94-103	Aramis site, subdesertic steppe	10°28'N, 40°27'E, altitude 500 m	surface litter soil
MA94-106	Ganduli river bank	10°31'N, 40°26'E, altitude 600 m	stream mud
MA94-101	Aduma site, Awash riparian forest	10°27'N, 40°30'E, altitude 500 m	surface litter soil
<i>Holocene</i>			
MA94-098	Aduma, section A4 (archaeological trench)	top geological layer above Ardu beds	black clay (vertisol) with plant remains
<i>Pleistocene</i>			
MA94-097	Aduma, section A4 (archaeological trench)	Middle Stone Age archaeological level	second shell bed with fish bones

low the definition of Mulholland (1989) and belong to the pooid type (Twiss et al., 1969). They occur dominantly in the subfamily Pooideae, which are C3 grasses abundant in cool, moist climates and at high elevations where available soil moisture is high during the growing season (Tieszen et al., 1979; Livingstone and Clayton, 1980). According to the same authors, the saddle morphotype, which belongs to the chloridoid type, is produced in high proportion by the Chloridoideae, represented by C4 grasses adapted to warm and dry climatic or edaphic conditions. The cross and dumbbell morphotypes belong to the panicoid type and occur dominantly in the subfamily Panicoideae, mainly represented by warm-season C4 grasses adapted to high available soil moisture (Tieszen et al., 1979; Twiss, 1980). It should be noted that pooid phytoliths have also been described in the intercostal zones of the epidermis of Chloridoideae and Panicoideae, but are produced in small amount (Brown, 1984; Mulholland, 1989).

Among the other grass subfamilies, Bambusoideae and Arundinoideae produce the same general morphotypes as discussed previously; no morphotypes are distinctive to the subfamily level (Metcalf, 1960; Brown, 1984; Mulholland, 1986, 1989; Ollendorf et al., 1988). Morphotypes elongate smooth, elongate sinuous, point and fan-shaped are phytoliths from epidermis long-cells and bulliform-cells of grasses. They are without taxonomic significance in the adopted classification (Twiss et al., 1969).

This classification of phytolith shapes has been criticized because multiplicity and redundancy of many grass morphotypes prevent the attribution of phytoliths to species or genus (Rovner, 1971; Brown, 1984; Mulholland, 1989). Because several shapes of phytolith occur in many grass taxa, a single phytolith morphotype cannot belong to a single grass species. But on the other hand, a phytolith assemblage can indicate the dominant subfamily constituting the grass association (Fredlund and Tieszen, 1994). The phytolith assemblages are also helpful to distinguish C3 from C4 grasslands and, among the C4 grasslands those dominated by Chloridoideae from those dominated by Panicoideae (Fredlund and Tieszen, 1994, 1997). Dry and wet C4 grasslands of West Africa were also well discriminated by phytolith assemblages using Twiss classification (Alexandre et al., 1997).

Non-Poaceae families also produce characteristic morphotypes. The cone-shape morphotype is attributed to the Cyperaceae (sedge) (Le Cohu, 1973; Ollendorf, 1987; Piperno, 1988) and the circular crenate to the Palmae (Piperno, 1988). The circular rugose morphotype, produced by the sclerenchym of ligneous dicotyledon (Geis, 1973; Scurfield et al., 1974; Welle, 1976; Laroche, 1976; Bozarth, 1992) is attributed to woody elements (Alexandre et al., 1997). The corkcell-like morphotype may not have any taxonomic significance because different families produce it (Ellis, 1979). These morphotypes have been identified in the phytolith assemblages of the Middle Awash samples (Plate I).

SEM photos of East African grass epidermis (Palmer and Tucker, 1981, 1983; Palmer et al., 1985; Palmer and Gerbeth-Jones, 1986) were consulted to check the applicability of Twiss classification to the grasses from the Middle Awash region. We found that East African Pooideae are rich in rectangles and rondels, Chloridoideae are characterized by abundance of saddles and Panicoideae by the abundance of crosses and dumbbells, which fits the usual pattern (Twiss et al., 1969). However, there are some exceptions. *Sporobolus*, frequent on riverbanks and largely represented in mud flats surrounding East African lakes (Menaut, 1983) as well as *Phragmites*, abundant nearby the Awash River (Sermolli, 1957), do not have a phytolith production that corresponds to the above classification. Actually, *Sporobolus*, which is a C4-Chloridoideae grass, produces the pooid type, and *Phragmites*, which is a C3-Arundinoideae species, produces the chloridoid type. These exceptions will be taken into account while discussing the results.

5. Results

Phytoliths were recovered in Pleistocene, Holocene and all three of the modern samples that were processed. The number of phytoliths counted ranged from 750 up to more than 1800 per slide in order to discuss at least 200 short-cell morphotypes for each sample. Phytoliths with unrecognizable shape were counted and listed in the unidentified category. The percentages of pooid, chloridoid and panicoid phytoliths are calculated on the sum of the

PLATE I

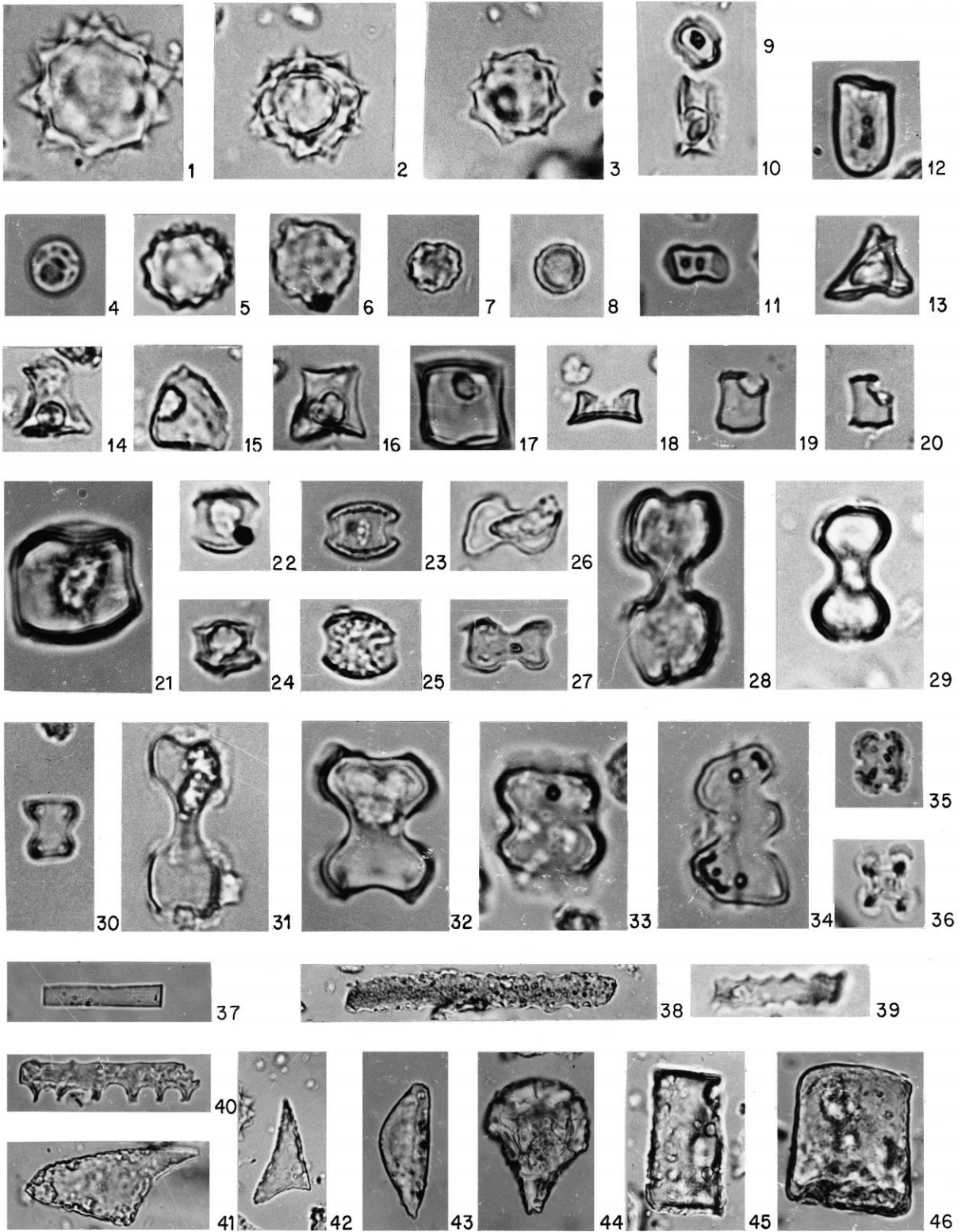


Table 3

Detailed counts of Poaceae and other plant phytoliths, recovered from modern and fossil samples from the western Middle Awash valley (Ethiopia)

Morphotypes	Surface soil			Holocene	Pleistocene	
	MA94-103	MA94-106	MA94-101	MA94-098	MA94-097	
Poaceae short-cell phytoliths						
rectangle	22	45	62	72	56	
rondel	126	43	78	67	95	
saddle	86	36	91	28	24	
dumbbell	66	126	132	39	66	
cross	23	41	29	5	13	
Σ short-cells	323	291	392	211	254	
Poaceae long-cell phytoliths						
elongate smooth	57	62	42	1332	75	
elongate sinuous	90	55	52	70	36	
fan shaped	95	91	106	97	90	
point shaped	69	56	69	30	108	
Σ long-cells	311	264	269	1529	309	
Ligneous	circular rugose	54	101	435	33	100
Palmeae	circular crenate	11	5	0	2	1
Cyperaceae	cone shape	0	12	13	3	5
All plants	corkcell like	156	83	198	78	64
	Σ non-Poaceae	221	201	646	116	170
	unidentified	17	14	27	19	23
	total	872	770	1334	1875	756

short-cell morphotypes alone (Σ *short cell*), to emphasize their variability in each sample. Percentages of Poaceae long-cell morphotypes and non-Poaceae morphotypes are presented in percent of the total

sum of phytoliths (unidentified category excluded). The detailed counts are provided in Table 3. The relative percentages of the modern and fossil phytolith assemblages are presented in a pollen-like diagram in Fig. 2.

Dissolution features that could indicate selective destruction (Fredlund and Tieszen, 1997) have not been observed in the modern and in the Pleistocene samples. However, the elongate morphotypes of the Holocene sample presented particularly corroded surfaces.

5.1. Modern phytolith assemblages

In the modern assemblages, there are about 2% of unidentified taxa. The Cyperaceae (cone-shape) are not abundant (less than 2%), even in the locally wet environments such as the riparian forest of Awash. They are absent from the banks of the Ganduli, which are not locally wet for more than a few days at a time. A low representation of Cyperaceae is expected due to the fact that this morphotype is

PLATE I

Non-Poaceae phytoliths (1–8) and Poaceae phytoliths (9–46). Phytoliths 1–36, $\times 1000$; phytoliths 37–46, $\times 500$.

- 1–3. Circular crenate (Palmae type).
4. Corkcell like (all plants).
- 5–8. Circular rugose (ligneous dicotyledons type).
- 9–17. Pooid type.
- 9–10. Rondel with circular base and circular flat top
- 11–12. Rondel with oblong base and top.
- 13–16. Rondel with circular base and ridge top.
17. Rectangle.
- 18–25. Chloridoid type: saddle.
- 26–36. Panicoid type.
- 26–34. Dumbbell.
- 35–36. Cross.
- 37–40. Elongate.
- 41–43. Point-shaped.
- 44–46. Fan-shaped.

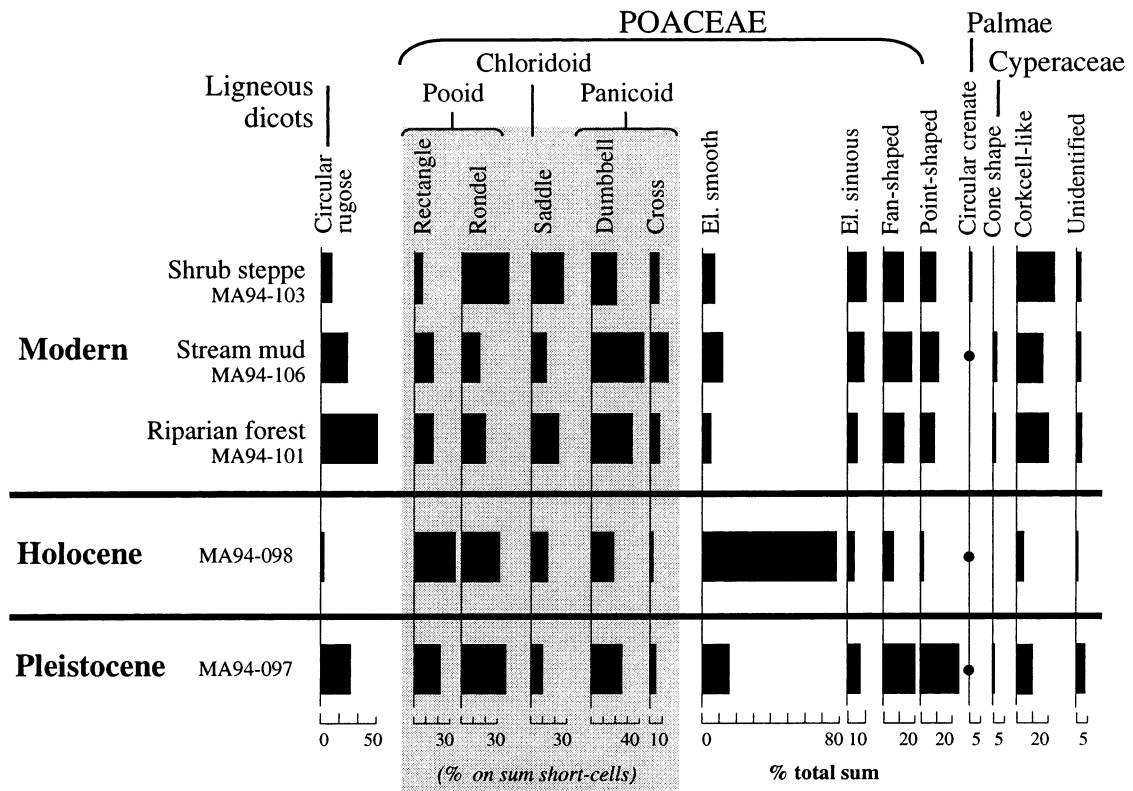


Fig. 2. Phytolith percentage diagram of modern and fossil samples from the Middle Awash anthropological site, Ethiopia.

highly subjected to dissolution, fragmentation and weathering processes (Alexandre et al., 1997). The Palmae (circular crenate) morphotypes account for less than 2%. Palm trees are not present in the area, but in the Lower Awash Valley. The corkcell-like morphotype is more abundant in the steppe (18%) than in the Awash riparian forest sample (15%) and the Ganduli mud (11%).

The steppe phytolith assemblage is characterized by a low proportion of the ligneous dicotyledon morphotype (6%) and a high content of Poaceae phytoliths (more than 65% of the total sum). Among the Poaceae phytoliths, rondels represent 39%, rectangles 7%, dumbbells plus crosses 27% and saddles 27% of the short-cell sum (Fig. 2). The long-cell phytoliths represent 35% of the total phytolith assemblage. The stream mud of Ganduli presents a phytolith assemblage where ligneous dicotyledon phytoliths are more abundant than in the steppe (13%). Poaceae morphotypes account for 73%.

Among the short-cell Poaceae phytoliths, rectangles and rondels account for 16% and 15%, saddles for 12%, dumbbells plus crosses morphotypes for 57%. Long-cell phytoliths represent 35% of the total assemblage. The Awash riparian forest is characterized by a phytolith assemblage with higher amount of ligneous dicotyledon morphotypes (33%) and relative low amount of Poaceae morphotypes (50%) compared to the shrub steppe and the mud. Among the short-cell Poaceae phytoliths, rectangles account for 16%, rondels for 20%, saddles for 23% and crosses plus dumbbells for 41%. Long-cell phytoliths represent 20% of the total phytolith sum.

5.2. Fossil phytolith assemblages

The Holocene assemblage is characterized by a low content of ligneous dicotyledons morphotype (less than 2%) and by 93% of Poaceae phytoliths, dominated by the elongate smooth morpho-

type (71%). Among the short-cell phytoliths, rectangles account for 34%, rondels for 32%, saddles for 13%, and dumbbells plus crosses for 21%. The corkcell morphotypes represents 4% of the total assemblage. Cyperaceae and Palmae are present.

The Pleistocene phytolith assemblage is characterized by 75% of Poaceae morphotypes and 13% of dicotyledon morphotypes. Among the short-cell Poaceae phytoliths, rectangles account for 22%, rondels for about 38%, saddles for 9% and dumbbells plus crosses morphotypes for 31%. Elongated, fan and point-shaped have high value of 41%. Undetermined phytoliths represent 3% of the total assemblage. The amount of corkcell-like morphotype in this sample is about 8%, and Cyperaceae and Palmae morphotypes account for less than 1%.

Modern and fossil phytolith assemblages present noticeable differences in their composition, which do provide evidence that there is no translocation of phytoliths from the surface soil to the geological layers. Consequently, phytolith assemblages can be considered as potential tools for reconstruction of paleoenvironments at Middle Awash.

6. Interpretation and discussion

In order to interpret phytolith assemblages, three indices have been defined by authors whose studies were done on grasses of American Great Plains (Twiss, 1992) and West African regions (Diester-Haas et al., 1973; Alexandre et al., 1997). These indices are applied to modern phytolith assemblages in an attempt to calibrate the phytolith response of the local vegetation against the climatic and/or edaphic conditions prevailing nowadays in the sub-desertic Middle Awash Valley. By comparison with the results obtained for the modern samples, an interpretation of the fossil assemblages can be made.

6.1. Modern phytolith assemblages and physiognomic aspect of the vegetation

6.1.1. Density of woody elements

The first index, defined as the D/P index, is the ratio of the ligneous dicotyledons morphotype (circular rugose) versus the Poaceae morphotypes (i.e. sum of pooid, chloridoid, panicoid, point and fan shaped) (Fig. 3). It has been used to estimate the density of woody elements in different types

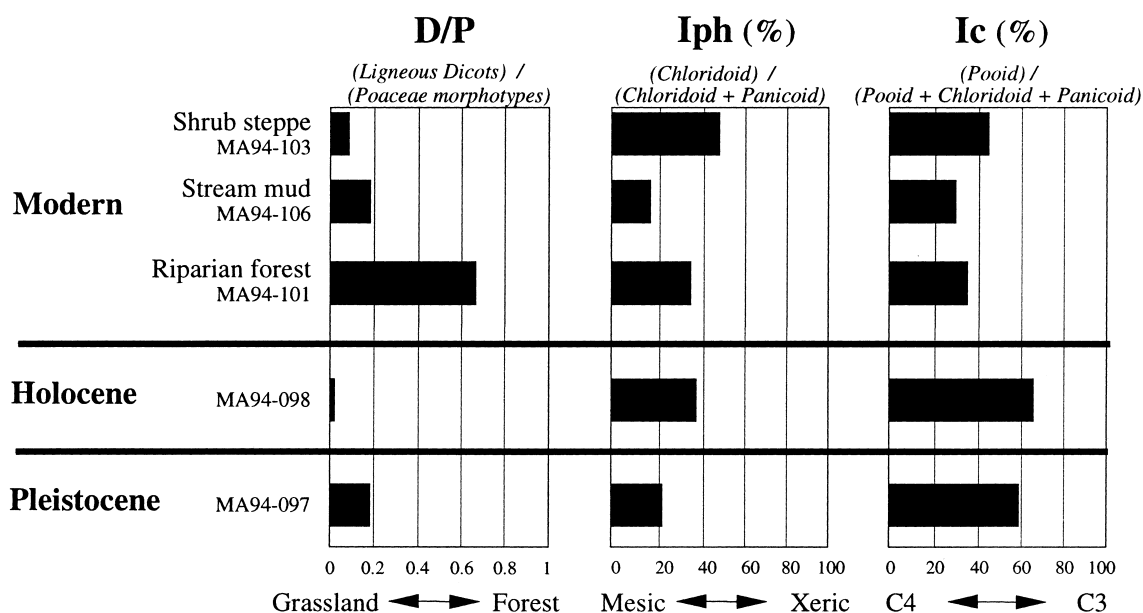


Fig. 3. Three indices [D/P = Tree cover density index (Alexandre et al., 1997); I_{ph} (%) = aridity index (Diester-Haas et al., 1973); and I_c (%) = climatic index (Twiss, 1987)] applied on modern and fossil samples from the Middle Awash anthropological site, Ethiopia.

of grass savannas between 15°N and 4°S in West Africa (Alexandre et al., 1997). In modern samples, a value of 7 was found for the semi-evergreen forest in Congo, whereas values lower than 1 characterized savannas with abundant grass cover in Senegal. In fossil Holocene sediments from Lake Sinnda (Congo), remarkably high D/P values (greater than 150) were found associated with high proportion of arboreal pollen, evidencing equatorial dense humid forest. In the present study at Middle Awash, D/P of the shrub steppe, the Ganduli stream and the margin of the riparian forest are, respectively, 0.1, 0.2 and 0.7 (Fig. 3). These values correspond well with the field observations. Indeed, trees and shrubs are scarce in the shrub steppe and along the Ganduli stream, and more abundant near the Awash riparian forest. Compared to those from West Africa, the D/P values clearly indicate that grasses dominate the vegetation cover at Middle Awash (Fig. 3). The D/P value for the riparian forest is still much lower than the one obtained for rainforest proper in Congo, where the tree cover is dense and stratified. We assume that the D/P would be a reliable index to estimate the density of trees and shrubs in different vegetation types in Africa. However, this conclusion would need to be confirmed by additional studies on phytolith assemblages from modern vegetation in Ethiopia.

6.1.2. Adaptation to aridity

The second index, first used to appreciate humidity–aridity tendencies in marine cores (Diester-Haas et al., 1973), is the ratio of chloridoid versus total chloridoid plus panicoid phytoliths (Twiss, 1992). It expresses the percentage of Chloridoideae among the C4 grasses. High I_{ph} values suggest grasslands dominated by Chloridoideae, i.e. xerophytic grasses, and hence prevalence of dry climatic conditions. Conversely, low I_{ph} values indicate associations in which Panicoideae, i.e. mesophytic grasses dominate, suggesting humid climatic and/or high available soil moisture. In West Africa an application of this index under the name I_{ph} (Phytolith Index), shows that a boundary value of 30% distinguishes tall-grass savannas of the humid Sudanian zone ($I_{ph} < 30\%$), from the short-grass savannas of the arid Sahelian region in Senegal ($I_{ph} > 30\%$) (Alexandre et al., 1997).

I_{ph} values of the assemblages collected in the shrub steppe, in the stream mud and at the margin of the Awash riparian forest are, respectively, 49, 18 and 36% (Fig. 3). The highest I_{ph} values are for the steppe and the riparian forest, while the lowest is for the mud. This is consistent with the fact that Chloridoideae dominate in the subdesertic steppe and at the margin of the riparian forest, while Panicoideae are abundant near the streams and temporary pools. However, the low value registered in the Ganduli sample may also be explained by panicoid phytoliths input from the tall-grasses found upstream the Ganduli river at the foothills of the Ethiopian escarpment (Fig. 1). The I_{ph} value for the steppe (49%) is higher than the minima value of 30% that characterizes dry grasslands in Senegal (Alexandre et al., 1997). It does correspond well with more arid conditions prevailing in East Africa. The I_{ph} value obtained for the riparian forest is also greater than 30%. We postulate that it does not reflect the local edaphic conditions of the riparian forest itself, but fits the general climate of the region. However, *Phragmites*, a C3-Arundinoideae grass, which is present in the swampy areas along the riversides, and which produces the chloridoid type phytolith, may have contributed to the phytolith assemblage, thus resulting in an overestimation of the aridity.

Although I_{ph} values of the Middle Awash modern samples fit the general conclusions based on samples from West Africa (Alexandre et al., 1997), a more extensive study on modern phytolith assemblages under different local conditions is necessary to fully assess the climatic interpretation of the phytolith index in the arid East African Rift.

6.1.3. Climatic index

Twiss (1987, 1992) defined a climatic index (I_c) to estimate the relative proportion of C3 grasses present in the American Great Plains. It is the ratio of pooid versus pooid plus chloridoid plus panicoid types. High I_c values suggest an abundance of Pooideae, and hence an abundance of C3 grasses. Abundance of C3-Pooideae characterizes high latitudes (Twiss, 1992, fig. 6.2) and high altitudes in the tropics (Tieszen et al., 1979; Livingstone and Clayton, 1980). Consequently high I_c values suggest cool climate. Indeed, modern phytolith assemblages from the North American Great Plains show I_c values

greater than 70% for northern C3-Pooideae-dominated grasslands, whereas southern C4-dominated grasslands show values of about 30% (calculated from [Fredlund and Tieszen, 1994](#)).

In the three modern samples collected in the Middle Awash Valley, the pooid type occurred at frequency of more than 30%. The surprisingly high I_c values would indicate the presence of Pooideae grasses. But their occurrence at Middle Awash is not consistent either with the tropical arid conditions ([Twiss, 1992](#)), or with their location in the lowlands or with the description of the vegetation ([Sermolli, 1957](#)). Inflorescence materials of non-Pooideae grasses may also produce the pooid morphotype in small amounts ([Mulholland, 1989](#)). However, statistical analysis carried out on phytolith assemblages from North American great Plains has shown a good correspondence between proportion of pooid morphotypes and representation of the Pooideae subfamily ([Fredlund and Tieszen, 1994, 1997](#)). The relatively high I_c value of the riparian forest (36%) could be explained by the presence of *Sporobolus*, which produces the pooid type though it belongs to the Chloridoideae subfamily. However, it is more difficult to explain the high I_c values for the steppe (46%) and for the mud (30%) where *Sporobolus* is absent from the local vegetation. If no other grasses at Middle Awash Valley produce pooid phytoliths ([Palmer and Tucker, 1981, 1983; Palmer et al., 1985; Palmer and Gerbeth-Jones, 1986](#)), we must then consider that non-local source areas contribute to the phytolith assemblages of the modern samples at Middle Awash.

Various phytolith dispersal processes can be considered. Fires are considered to be important regional vectors of phytolith transportation and deposition ([Twiss et al., 1969; Twiss, 1983, 1987](#)), as well as aeolian processes, herbivory, fluvial and colluvial erosion ([Clark, 1988; Fredlund and Tieszen, 1994](#)). In North American Great Plains, extra-local and regional source areas have been estimated to be responsible for 30–70% of the annual phytolith budget for grassland soils ([Fredlund and Tieszen, 1994](#)). At Middle Awash, the geomorphic setting and the distribution of rainfall over the year, as well as the main winds must be considered as potential processes for phytolith dispersion and deposition in the valley. The main winds blow eastward from the Ethiopian highlands and runoff is important in this arid region

where rainfall on the escarpment occurs at high rates but over a short period. Although regional phytolith input from the Ethiopian plateau is not so obviously evidenced by the D/P and I_{ph} indices, the presence of *Palmae* phytolith, though Palm trees are absent in the local vegetation, also suggests that extra-local sources supply modern phytolith assemblages. It is also interesting to note that the pooid type was not found in modern phytolith assemblages from West Africa, where highlands as potential source areas of this phytolith type do not exist ([Alexandre et al., 1997](#)). In this study, I_c values can be more likely explained by high-elevation contributions by wind or runoff. We assume that phytolith assemblages include C3-Pooideae signal from highland vegetation, mixed with local grass composition of the subdesertic steppe.

Finally, our investigation of the modern phytolith assemblages shows that the two local environments, the subdesertic steppe and the riparian forest are well discriminated by the D/P index, which values are in good agreement with the local shrub and tree cover densities. The two other indices are, however, not fitting exactly the local vegetation. The I_{ph} values are in agreement with the aridity of the local grass associations, taking into account potential bias resulting from the presence of *Phragmites*. The significance of the climatic index appears more likely interpreted by a mixture of phytoliths from regional and local source areas. Additional investigation of phytolith assemblages collected in a soil surface developed under highland C3-Pooideae grasslands would allow a better interpretation of modern phytolith assemblages in the Middle Awash Valley. Nevertheless, extra-local contributions may influence D/P and I_{ph} indices but do not control them.

6.2. Fossil assemblages

The interpretation of the fossil assemblages must take into account two facts: (1) Phytolith index I_{ph} can overestimate the aridity of a lacustrine or riparian fossil sample if *Phragmites* was abundant in the local vegetation. (2) Climatic index I_c for this area evidences the vegetation signature on both regional and local scale.

The phytolith assemblage related to the Holocene sample is clearly characterized by a very high pro-

portion (70%) of corroded elongate smooth morphotypes, which has no equivalence in the modern samples. The same phenomena has been observed in four samples of Pleistocene loess sediment from the Central Great Plains of North America (Twiss, 1987). Corroded particles may indicate chemical and/or mechanical erosive action by various agents such as running water or wind. The very low D/P value for the Holocene sample implies that local vegetation was a grassland. The high I_{ph} value (39%) indicates a high contribution of phytoliths from C4-Chloridoideae grasses, although non-Chloridoideae species such as *Phragmites* may have contributed to the assemblage. The high I_c value (66%) implies a significant contribution of phytoliths from C3-Pooideae, although the presence of non-Pooideae species such as *Sporobolus* may have contributed to the assemblage. Since Chloridoideae and Pooideae cannot be found under the same ecological conditions in East Africa (Livingstone and Clayton, 1980), phytoliths of the Holocene sample must have been brought in from different areas. This phytolith assemblage seems to record local grassland rich in Chloridoideae grasses, adapted to a dry environment, and C3-Pooideae grasses would have covered the Ethiopian Plateau. Aberrant 'co-existence' of Chloridoideae and Pooideae as well as abundance of corroded particles in the Holocene assemblage imply potential transport and long-distance phytolith input.

The phytolith assemblage related to the Pleistocene sample, contemporaneous with Middle Stone Age archaeological artefacts, shows a D/P value of 0.2. Following the same arguments as those presented before, the low D/P value indicates an open formation where trees and shrubs are scarce, but more abundant than for the Holocene sample. The low I_{ph} value (23%) indicates the dominance of Panicoideae grasses, and hence suggests humid local and/or regional conditions. The associated chloridoid morphotype, in lower proportion than in modern samples, may be related to local development of Chloridoideae grasses or to *Phragmites*. High I_c value (59%) suggests that Pooideae grasses or *Sporobolus* supplied the phytolith assemblage of the Pleistocene sample. It is unlikely to consider that Pooideae grasses could have developed at 500 m of altitude during the Pleistocene, however, they may

have occurred in the highlands, and contributed as regional source of pooid phytoliths.

Finally, the combination of remarkably low I_{ph} and high I_c values for the Pleistocene sample may be interpreted as attesting two kinds of environments: (1) riparian vegetation developed under suitable edaphic conditions and constituted of warm-season C4 grasses which requires high available soil moisture, such as Panicoideae, *Phragmites* and *Sporobolus*; and (2) grassland, adapted to humid climatic conditions in the lowlands, constituted of C4 grasses among which Panicoideae were dominant. C3-Pooideae grasses would have occurred in the highlands. In both hypotheses, the D/P index is consistent and suggests scarce trees and shrubs.

7. Conclusion

Modern soils, Holocene and Pleistocene sediments from the Middle Awash paleoanthropological site contain well-preserved phytolith assemblages. Their different composition evidence that translocation from surface soil to geological layers is not important in the study area. Therefore, phytolith assemblages are valid tools for investigating past environments.

This study demonstrates that modern phytolith assemblages enable the differentiation of subdesertic steppe from wooded riparian forest, and characterize the composition of the C4-grass associations. Moreover, they record local as well as regional vegetation.

The phytolith assemblage of the Holocene sample records C4-Chloridoideae grassland close to the modern subdesertic shrub steppe, where C3-Pooideae grasses would similarly have covered the highlands. The phytolith assemblage of the Pleistocene sediment indicates an open grassland with more trees and shrubs, developed under humid edaphic conditions or under climate more humid than today. Remarkably, C3-Pooideae grasses are more abundant in the fossil than in the modern samples. The most probable hypothesis is that this might evidence stronger wind or runoff under wetter climatic conditions in the past. Further studies on modern phytolith assemblages are strongly needed to improve the calibration of indices against the modern vegetation and hence improve the validity of phytolith assemblages to reconstruct

paleoenvironments notably for paleoanthropological sites, abundant in the Ethiopian Rift Valley.

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