

Quantitative Convergence and Stability of Seismic Inverse Problems.

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Quantitative Convergence and Stability of Seismic Inverse Problems.

<u>Florian Faucher</u>¹,

Hélène Barucq 1 , Henri Calandra 2 and Guy Chavent 1 .

Reconstruction Methods for Inverse Problems Indam Workshop, Roma, Italy, May 28^{th} – June 1^{st} , 2018.

Istituto Nazionale di Alta Matematica

 2 Total E& P

Seismic inverse problem

Reconstruction of subsurface Earth properties from seismic campaign: collection of wave propagation data at the surface.

- \blacktriangleright Reflection (back-scattered) partial data,
- \triangleright only from the surface of the (large) domain,

nonlinear, ill-posed inverse problem.

1 Time-Harmonic Inverse Problem, FWI

2 Quantitative stability and convergence of FWI

- **•** Finite Curvature/Limited Deflection problem
- Numerical convergence estimates
- Numerical stability estimates
- 3 Numerical experiments

The forward problem wave equation depends on the medium:

Time-harmonic wave equation

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 \blacktriangleright acoustic isotropic (c)

$$
(-\omega^2 c^{-2} - \Delta)p = 0,
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lastic isotropic (λ, μ, ρ) $-\rho\omega^2\mathbf{u}-\nabla\left(\lambda\nabla\cdot\mathbf{u}\right)-\nabla\cdot\left(\mu\left[\nabla\mathbf{u}+(\nabla\mathbf{u})^T\right]\right)=0.$

Time-harmonic wave equation

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- lastic isotropic (λ, μ, ρ) $-\rho\omega^2\mathbf{u}-\nabla\left(\lambda\nabla\cdot\mathbf{u}\right)-\nabla\cdot\left(\mu\left[\nabla\mathbf{u}+(\nabla\mathbf{u})^T\right]\right)=0.$
- **F** anisotropy (stiffness tensor, ρ) $-\omega^2 \rho \mathbf{u} - \nabla \cdot (\mathbf{C} \epsilon(\mathbf{u})) = 0.$

Viscous behavior are considered with complex coefficients.

Full Waveform Inversion (FWI)

FWI provides a quantitative reconstruction of the subsurface by solving a minimization problem,

$$
\min_{m\in\mathcal{M}}\quad \mathcal{J}(m)=\frac{1}{2}\|F(m)-d\|^2.
$$

- \blacktriangleright d are the observed data.
- \blacktriangleright $F(m)$ represents the simulation using an initial model m:

$$
F \; : \; m \;\;\rightarrow \;\;\big\{p(\mathbf{x}_1),p(\mathbf{x}_2),\ldots,p(\mathbf{x}_{n_{rev}})\big\}.
$$

P. Lailly

The seismic inverse problem as a sequence of before stack migrations Conference on Inverse Scattering: Theory and Application, SIAM, 1983

A. Tarantola

Inversion of seismic reflection data in the acoustic approximation Geophysics, 1984

A. Tarantola

Inversion of travel times and seismic waveforms Seismic tomography, 1987
Florian Faucher – Converge

Convergence of time harmonic FWI $-$ May 28– June 1st, 2018 6/28

Numerical methods

- \triangleright Forward problem resolution with Discontinuous Galerkin methods,
- \triangleright parallel computation, HPC, large-scale optimization.

Multi-frequency algorithm, stability and convergence.

Plan

- Finite Curvature/Limited Deflection problem
- Numerical convergence estimates
- Numerical stability estimates

Helmholtz inverse problem from back-scattered partial data, quantitative reconstruction with iterative optimization.

$$
\min_{m \in \mathcal{M}} \quad \mathcal{J}(m) = \frac{1}{2} \|\mathcal{F}(m) - d\|^2.
$$

Convergence radius

 \blacktriangleright initial model needs to be within the radius of convergence.

Stability

 \blacktriangleright ensures reconstruction accuracy.

Helmholtz inverse problem from back-scattered **partial** data, quantitative reconstruction with iterative optimization.

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Numerical estimates to guide the procedure

- \blacktriangleright frequency;
- \blacktriangleright parametrization;
- \blacktriangleright geometry;
- \blacktriangleright forward problem, ...

1/ Convergence radius

 \triangleright initial model needs to be within the radius of convergence.

Least squares minimization problem

$$
\min_{m \in \mathcal{M}} \quad \mathcal{J}(m) = \frac{1}{2} \|\mathcal{F}(m) - d\|^2.
$$

M. V. de Hoop, L. Qiu, O. Scherzer

An analysis of a multi-level projected steepest descent iteration for nonlinear inverse problems in Banach spaces subject to stability constraints Numerische Mathematik 2015

G. Chavent

Nonlinear least squares for inverse problems: theoretical foundations and step-by-step guide for applications. Springer 2010

G. Chavent and K. Kunisch

On weakly nonlinear inverse problems.

Florian Faucher SIAM Journal on Applied Mathematics 1996 Florian Faucher – May 28– June 1st, 2018 10/28

1/ Convergence radius

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Least squares minimization problem

$$
\min_{m \in \mathcal{M}} \quad \mathcal{J}(m) = \frac{1}{2} \|\mathcal{F}(m) - d\|^2.
$$

Finite Curvature/Limited Deflection guarantees **uniqueness** of the solution and **unimodality**: no local minimum.

M. V. de Hoop, L. Qiu, O. Scherzer

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SIAM Journal on Applied Mathematics 1996 Florian Faucher – Convergence of time harmonic FWI – May 28– June 1st, 2018 10/28

 $\forall m_0, \Delta_m \in \mathcal{M}, \qquad P \; : \; t \in [0,1] \rightarrow \mathcal{F}\big(m_0 + (2t-1)\Delta_m\big).$

Florian Faucher $-$ Convergence of time harmonic FWI $-$ May 28– June 1st, 2018 11/28

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Florian Faucher $-$ Convergence of time harmonic FWI $-$ May 28- June 1st, 2018 12/28

G. Chavent

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For a given m_0 , we estimate the maximal distance Δ_m that still verifies the Limited Deflection property $\Theta(P) \leq \frac{\pi}{2}$ $\frac{1}{2}$.

Larger size are required when missing a priori info.

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 \triangleright Δ_m is estimated in the direction δ_k ,

$$
\|\Delta_m^{\delta_k}\| \leq \frac{\pi}{4} \frac{\|\mathcal{F}'(m_0)(\delta_k)\|}{\|\mathcal{F}''(m_0)(\delta_k,\delta_k)\|}.
$$

Larger size are required when missing a priori info.

Numerical estimate of the size $\Delta_m^{\delta_k}$

- \bullet with direction δ_k (the geometry of the unknown),
- **2** with the frequency ω ,
- ³ with the parametrization.

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Numerical estimate of the size $\Delta_m^{\delta_k}$ with

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- **2** with the frequency ω .

(c) direction δ_3

$$
(-\omega^2 c^{-2} - \Delta)p = 0
$$

 \blacktriangleright $\big| -\omega^2 = (s + 2i\pi f)^2$

$$
\blacktriangleright s = 0
$$
: Fourier domain $\omega^2 = (2\pi f)^2$,

$$
f = 0
$$
: Laplace domain $-\omega^2 = s^2$.

C. Shin and Y. H. Cha

Waveform inversion in the laplace domain ; Waveform inversion in the laplace fourier domain Geophysical Journal International 2008–2009

W. Ha, S. Pyun, J. Yoo and C. Shin

Acoustic full waveform inversion of synthetic land and marine data in the laplace domain ; Geophysical Prospecting 2010

P. V. Petrov and G. A. Newman

Three-dimensional inverse modelling of damped elastic wave propagation in the fourier domain.

Geophysical Journal International 2014

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 \blacktriangleright ensures reconstruction accuracy.

The stability indicates how minimizing the data recovers the model

$$
||m_1-m_2||\leq \mathcal{C}||\mathcal{F}(m_1)-\mathcal{F}(m_2)||.
$$

The stability constant C depends on the frequency and the number of unknowns.

Ē

G. Alessandrini

Stable determination of conductivity by boundary measurement Applicable Analysis 1988

E. Beretta, M. V. de Hoop, F. and O. Scherzer

Inverse boundary value problem for the Helmholtz equation: quantitative conditional Lipschitz stability estimates.

SIAM J. Math. Analysis 2016

G. Alessandrini, M.V. de Hoop, R. Gaburro and E. Sincich

Lipschitz stability for a piecewise linear Schrödinger potential from local Cauchy data arXiv:1702.04222, 2017

The stability indicates how minimizing the data recovers the model

 $\|m_1 - m_2\| \leq C \|\mathcal{F}(m_1) - \mathcal{F}(m_2)\|.$

Stability & Convergence estimates

From our estimates, we see different conditions for stability and convergence.

Quantitative estimates provide an initial, comprehensive, relation to guide the iterative procedure.

0./ No initial information on the subsurface

Elastic Marmousi model

Multi-parameters inversion 17×3.5 km, true models.

$$
-\rho\omega^2\mathbf{u}-\nabla\left(\lambda\nabla\cdot\mathbf{u}\right)-\nabla\cdot\left(\mu\left[\nabla\mathbf{u}+(\nabla\mathbf{u})^T\right]\right)=0.
$$

(c) Density

Elastic Marmousi model

Multi-parameters inversion 17×3.5 km, starting models.

(a) P-wave speed (b) S-wave speed

(c) Density

Elastic Marmousi reconstruction 17×3.5 km

Frequency from 1 to 10Hz, unknown density, multi-level algorithm.

(a) True P-wave speed (b) True S-wave speed

(c) Initial P-wave speed (d) Initial S-wave speed

Elastic Marmousi reconstruction 17×3.5 km

Frequency from 1 to 10Hz, unknown density, multi-level algorithm.

(a) True P-wave speed (b) True S-wave speed

Multi-parameters inversion 31×7 km, true models.

(a) P-wave speed (b) S-wave speed

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Multi-parameters inversion 31×7 km, starting models.

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Frequency from 1 to 10Hz, unknown density, multi-level algorithm.

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(a) True P-wave speed (b) True S-wave speed

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Reminder from the convergence analysis.

\triangleright Salt domes reduce the size of the radius of convergence.

Solution

 \triangleright Start with low or complex frequency.

Using complex frequencies, unknown density, multi-level algorithm.

(a) True P-wave speed (b) True S-wave speed

(c) Initial P-wave speed (d) Initial S-wave speed

Using complex frequencies, unknown density, multi-level algorithm.

(a) True P-wave speed (b) True S-wave speed

 (c) (0Hz,10) P-wave speed (d) (0Hz,10) S-wave speed

Using complex frequencies, unknown density, multi-level algorithm.

(a) True P-wave speed (b) True S-wave speed

(c) 10Hz P-wave speed (d) 10Hz S-wave speed

Conclusion

Comprehensive FWI

- \triangleright Quantitative stability and convergence analysis:
	- \blacktriangleright depending on the frequency and geometry,
	- \blacktriangleright depending on the methods.
- \triangleright We can also quantify the noise robustness.

Perspectives

- \triangleright Analytical relation to develop between the two components,
- \triangleright obtain conditional progression in frequency from stability and convergence estimates

Quantitative reconstruction method for inverse problem

- \triangleright Discontinuous Galerkin discretization in HPC framework,
- \blacktriangleright acoustic, elastic, anisotropy, viscosity; 2D, 3D, dual-sensors data.

Applications and ongoing investigations:

Quantitative reconstruction method for inverse problem

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- \triangleright acoustic, elastic, anisotropy, viscosity; 2D, 3D, dual-sensors data.

Applications and ongoing investigations:

- \triangleright Inverse scattering using obstacles positions,
- visco-elastic reconstruction: five unknowns,
	- \blacktriangleright (λ , μ , ρ , Q_{λ} , Q_{μ}),
	- In unknown Q does not prevent λ recovery,
	- \blacktriangleright procedure to recover attenuation parameter?

\blacktriangleright helioseismology; anisotropy; model parametrization.

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Thank you

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Appendix

Quantitative convergence for MBTT

Numerical estimate of the size $\Delta_m^{\delta_k}$ the problem components

- **1** the frequency ω ,
- **2** with direction δ_k (the geometry of the unknown),
- **3** with the parametrization.

The estimates allow a comparison of methods, in particular we want to compare the dependency of the optimization on the low frequencies.

Quantitative convergence for MBTT

Numerical estimate of the size $\Delta_m^{\delta_k}$ the problem components

- **1** the frequency ω ,
- **2** with direction δ_k (the geometry of the unknown),
- **3** with the parametrization.

the MBTT method decomposes the model with a smooth part (propagator) and the reflectors:

$$
c^{-2}=p+r=p+D\mathcal{F}^*(p)s.
$$

F. Clément and G. Chavent

Waveform inversion through MBTT formulation – 1992

F. Clément, G. Chavent and S. Gómez

Migration-based traveltime waveform inversion of 2-d simple structures: A synthetic example Geophysics 2001

G. Chavent, K. Gadylshin and V. Tcheverda

Reflection fwi in mbtt formulation EAGE 2015

Quantitative convergence for MBTT

Numerical estimate of the size $\Delta_m^{\delta_k}$ the problem components

- **1** the frequency ω ,
- **2** with direction δ_k (the geometry of the unknown),
- **3** with the parametrization.

Comparison of convergence with parametrization,

$$
c^{-2} = p + D\mathcal{F}^*(p)\mathbf{s}.
$$

Quantitative convergence estimates with frequency

Frequency progression for iterative algorithm, $-\omega^2 = (s + 2\mathrm{i}\pi f)^2$.

Other application

- \triangleright Single frequency gives high radius than range of frequencies,
- \triangleright can be apply for other minimization problem,

Finite curvature indicates robustness to noise

 \blacktriangleright Low and complex frequencies more affected.

Stability

 \blacktriangleright

 \blacktriangleright Indicates the accuracy of the reconstruction.

Acoustic Marmousi reconstruction 9.2×3 km

Starting available frequency is 4Hz.

Acoustic Marmousi reconstruction 9.2×3 km

Starting available frequency is 4Hz.

Using Fourier frequencies only: from 4 to 10Hz

Acoustic Marmousi reconstruction 9.2×3 km

Starting available frequency is 4Hz.

(a) Using Fourier frequencies only (b) Using Complex frequencies from (4Hz, 7)