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Pressuremeter test parameters of a compacted illitic soil

under thermal cycling

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ABSTRACT: The incorporation of heat exchangers into geostructures leads to changes

in the temperature of the adjacent soil. Important issues arise about the effect of

temperature variations on hydro-mechanical soil behaviour. The objective of this paper

is to improve the understanding and the quantification of the impact of temperature

variation on the bearing capacity of thermo-active piles. Currently, the design of the

deep foundations is based on the results of in situ penetrometer or pressuremeter tests

while there is no published data on the effect of temperature on in situ tests results

allowing to specifically assess thermo-active piles behaviour. Herein, an experimental

device is developed to carry out mini-pressuremeter tests under controlled laboratory

conditions. Mini-pressuremeter tests are performed on an illitic soil in a thermo-

regulated metric scale container subjected to a range of temperature from 1 to 40 °C.

The results showed a slight decrease in the pressuremeter modulus (E_p) and a significant

decrease in the creep pressure (p_l) and the limit pressure (p_l) as the temperature

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increased. The results showed also the reversibility of this effect during a heating-cooling cycle through the investigated temperature range, while the effect of a cooling-heating cycle was only partially reversible. In the case of several thermal cycles, the effect of the first cycle proved to be decisive on the soil parameters.

Keywords: Laboratory tests; Pressuremeter; Temperature; Compacted soils; Illite.

1 Introduction

In the last years, developments in the field of shallow geothermal energy systems using geostructures such as thermal piles, diaphragm walls, tunnel linings, basement slabs are being achieved (Brandl, 2006; Fromentin et al., 1999; Laloui et al., 2003). These thermo-active elements in contact with the subsurface soil are used to inject heat in the summer and/or extract heat in the winter. Under service conditions, the temperature of the heat exchanger fluid within the pipes is generally between 4 and 30 °C (Péron et al., 2011) and the circulation of the fluid results in cyclic changes in the surrounding ground temperature.

In the case of energy piles, on one hand, the thermal expansion or contraction causes additional thermal stresses in the pile which can influence the skin friction and the end bearing and consequently the load transfer (Arairo et al., 2015; Bourne-Webb et al., 2013). On the other hand, the cyclic thermal variation imposed to the soil may affect its hydro-mechanical properties. Thus, the behaviour of an energy pile is directly related to the coupling of thermal, hydraulic and mechanical phenomena. Under thermal conditions, the soil properties such as thermal conductivity, heat capacity, density, water content and hydraulic conductivity could be modified. These modifications may affect the bearing capacity as well as the friction behaviour of the soil. To study the temperature effect on ground structures, the complex interactions between temperature

variations and induced stresses and strains, which may affect long-term structure performance, must be considered (McCartney et al., 2010a). Despite the growing number of energy piles (Brandl, 2006), only few recommendations are published (e.g. GSHP Association, 2012), there is still no standard for the design of this type of piles.

Different studies were conducted on the behaviour of energy piles and their thermomechanical response in situ (Akrouch et al., 2014; Bourne-Webb et al., 2009; Laloui et al., 2003) and in laboratory (McCartney et al., 2010b; Tang et al., 2013). All these studies focused on the pile behavior and the effect of temperature variations on the soil parameters has been completely neglected although, a cyclic thermal variation may have an important impact on the shear strength, yielding and critical-state soil properties. Especially, the contraction of the elastic domain (yield locus) with the increasing temperature has been shown in the literature for various clay materials. Several relevant studies addressing the so-called thermal softening are cited in Table 1. However, to take into account these evolutions, it is necessary: (1) to connect the surrounding soil parameters to the soil-pile mechanical interaction behavior, (2) to study the effect of lowest and highest temperatures and (3) to consider the influence of cyclic evolution of the temperature. In this paper, the impact of cyclic thermal variations in the range from 1 to 40 °C on the parameters which are directly linked to the soil bearing and to the soil-pile friction is quantified.

Calculation of the bearing capacity of deep foundations are currently based on pressuremeter or/and penetrometer test results (AFNOR, 2000, 1996; Standard ASTM, 1999, 1995). Eslami et al. (2015) studied the effect of temperature variation on penetration test parameters in two compacted soils but there is still no available data on the effect of temperature on the pressuremeter parameters. Thus, herein and for the first time, a mini-pressuremeter test adapted to laboratory conditions is used to quantify the

temperature effect on the mechanical behavior of the soil. An experimental method is developed to carry out mini-pressuremeter tests on an homogeneous soil compacted at a controlled density and water content and subjected to monotonous or cyclic thermal variations.

In the following sections, first the material preparation and the experimental device are described, then, the obtained results are presented and the effect of temperature variations on the pressuremeter parameters is discussed.

2 Material and methods

In this section, the studied material is presented and experimental device and the minipressuremeter tests are detailed.

2.1 Basic characterisation of the material

The tested material is an illitic soil named Arginotech® from the east of Germany. It contains 77% illite, 10% kaolinite, 12% calcite and traces of quartz and feldspar. Illite is a non-swelling clay found in various proportions in soils due to the illitization of several minerals (Lynch, 1997). Mini-penetration tests, conducted on this soil by Eslami et al. (2015) showed the sensitivity of this material to temperature variations.

The particle-size distribution of the soil was determined using a laser Malvern Mastersizer 2000® device (Figure 1a). Almost 85% of the particles of the illitic material are smaller than 0.002 mm (clay particles), and the remaining 15% are smaller than 0.02 mm (silt particles). Its liquid limit (LL) is 65%, the plastic limit (PL) is 34%, and the plasticity index (PI) reaches 31 (AFNOR, 1993). The obtained standard Proctor curve of the material (AFNOR, 1999) allows defining an optimum water content of 31.3% and a maximum dry density of 1.43 Mg/m³ (Figure 1b). The material is classified

as "A3" in the French standard for soil classification (GTR, 2000) and as a fat clay, MH, according to the Unified Soil Classification System (Standard ASTM, 2006).

2.2 Experimental device and methods

To perform mini-pressuremeter tests under controlled temperatures, a homogeneous soil with a relatively large volume is required as well as an experimental device to impose thermal variations to the soil. In this section, the developed methodology to obtain a meter-scale sample with homogeneous properties (temperature, water content and dry densities) is described.

2.2.1 The container

A cylindrical stainless steel container of 800 mm in height and 600 mm in diameter is made. A stainless steel tube is welded to the outside of the container. This tube is connected to a Vulcatherm® thermoregulator to allow the circulation of an ethylene glycol-water solution at the target temperature in the tube (-20 to 90 °C). As a result, the soil, compacted in the container, is heated or cooled from the outer lateral surface of the container. The evolution of the temperature within the sample is centripetal. Insulating sleeves are placed on the tube to reduce heat exchange with the surrounding atmosphere. The top of the massif is insulated with a plastic film to preserve the initial water content. Finally, the whole device is placed in a box made of extruded polystyrene plates of 40 mm thick to reinforce the insulation (Figure 2).

2.2.2 Sample preparation and compaction

First of all, the required amount of water is added to material ($w_{initial} \cong 17\%$) in a large capacity fixed-speed (38 rpm) mixer MIX120® to reach the target water content: $w_{OPN} = 31.3\%$ (Figure 3a). The wet material is stored in 8 plastic drums for a minimum

period of 5 days to ensure a good homogenization. A pneumatic compactor is used to compact the material in the container in several layers at ambient temperature (approximately 20 °C). The compactor applied dynamic forces on a metallic plate of 4 mm in thick and 600 mm in diameter to allow a homogeneous compaction of the soil. The volume of each layer is controlled by three rulers fixed to the inside wall of the container (Figure 3b).

Two compaction processes were tested: (1) a compaction in 8 layers of 100 mm in thickness and (2) a compaction in 11 layers of 70 mm in thickness. The two processes were tested and the resulting dry density evolution according to the depth was checked thanks to the extraction of several soil cores. Results of these tests, presented in section 3.1, led to the choice of a compaction in 11 layers of 70 mm in thickness for all test series.

2.2.3 Instrumentation of the sample

Seven thermal sensors PT100 are positioned within the soil at various depths and various distances from the wall of the container as shown in Figure 4. The sensors are plugged to a data logger to monitor the temperature evolution inside the compacted soil during the heating/cooling phases.

2.3 Mini-pressuremeter tests

Only six mini-pressuremeter tests are performed in each massif to prevent edge effects and cross-influence between the tests. The test points are positioned on a concentric circle twice smaller than the container. The tests are carried out at 150 mm from the edge of the container and the distance between the test points is also150 mm (Figure 5). All test points have the same position so that the results can be compared. The pressuremeter tests are conducted with an APAGEO® mini-pressuremeter probe of 380

mm in height and 28 mm in diameter (Figure 6). Before each pressuremeter test, a core with a diameter equal to that of the mini-pressuremeter probe is extracted using a core sampler. The 630 mm core is divided into small pieces to measure the water content and the density of the material as a function of the depth. The center of the probe is positioned in the middle of the height of the compacted soil (385 mm). The probe is connected to a pressure-volume controller GDS®. The pressure controlled test consists in applying increasing pressure with equal increments of 25 kPa for at least one minute. The equilibrium injected volume is measured for each increment and the volume versus pressure curve is plotted. The test is stopped when the injected volume reached 140,000 mm³ (the volume variation limit of the probe). Immediately after the pressuremeter test at a given temperature, the borehole is filled up with the same material, at the same water content, to avoid influencing the following tests. In this study, the effect of temperature variations on the resistance of the mini-pressuremeter membrane is measured by placing the probe in a climatic chamber at a given temperature during the calibration test of the membrane resistance. Figure 7 shows the developed experimental device elements for the mini-pressuremeter tests.

3 Results

In the framework of this study, a preliminary series and four test series (six minipressuremeter tests for each series) were carried out (Table 2). The preliminary series aimed to check the efficiency and repeatability of the compaction method to obtain a homogeneous massif (water content and dry density). It also enabled to assess the reproducibility of the mini-pressuremeter tests performed in the soil and to estimate the related uncertainties of the test parameters. The four following test series quantified the effect of a monotonic or a cyclic thermal variation on the pressuremeter parameters. The

performed tests in each sample were numbered as follows Xx; with X, the massif temperature when performing test (i.e. 1, 20 or 40 °C) and x, a letter indexing the test that increments for each new test at the same temperature.

In the following, the preliminary series will be presented. Then we focus on the evolution over one thermal cycle by studying the first and the second series. Finally we conclude on the impact of several thermal cycles by analysing the results of the third and the fourth test series.

3.1 Preliminary series

The preliminary test series, carried out at 20 °C, aimed to verify the homogeneity of compacted soil (Figure 8), as described in section 2.2.2. Figure 8a shows the variation of the dry density versus depth for a soil compacted in 8 layers of 100 mm. In each layer, ρ_d is 7% higher than the target at the top of the layer, and 6% lower than the target value at the bottom of the layer. This density gradient was measured within each layer and the soil is heterogeneous. Figure 8b shows the dry density profile when the soil is compacted in 11 layers of 70 mm (4th test series). As it was expected, the gradient of ρ_d decreased significantly, the values were closer to the target value (in average \pm 2.5%). Thus, having fixed the accuracy of the dry density to 5%, the compaction in 11 layers of 70 mm in thickness was chosen for the preparation of the soil studied in the following test series.

Five mini-pressuremeter tests were performed for the preliminary series under the same conditions (Figure 9). Each pressuremeter curve consists in 3 domains: (i) the probe inflates to achieve the contact with the wall of the hole, (ii) the volume increases linearly with increasing pressure allowing the calculation of the pressuremeter modulus E_p (the soil pseudo-elastic reaction against the probe pressure) and (iii) large

displacements take place and the soil becomes plastic. The creep pressure (p_f) is in the boundary between the second and the third domain of the curve. The pressuremeter parameters (E_p, p_f, p_l) were determined from the pressuremeter curves (Table 3). The results showed a good repeatability of the tests and allowed the estimation of the uncertainty of each parameter according to the Student's t-distribution (because of the limited number of measurements) for a 90% confidence interval: E_p : $\pm 10\%$, p_f : $\pm 5\%$, p_l : $\pm 5\%$.

3.2 Effect of one thermal cycle

To quantify the impact of temperature variations on pressuremeter parameters, two test series were carried out. In this part, each test series is explained then results are compared.

3.2.1 *First series: heating-cooling cycle* (20-40-20 °C)

For the first test series, starting from 20 °C, the soil was heated to 40 °C and then cooled to 20 °C at the end. Figure 10 shows the temperature variation as recorded by the sensor S5 located at 150 mm from the edge of the massif and at the middle of its height (Figure 4). The thermal equilibrium within the limits of ± 1 °C was reached after 65h.

The average water content of the soil close to the test area (from 200 to 580 mm in depth) was measured (Table 4). At 40 °C, w was slightly lower than at 20 °C, this might be due to: (1) the evaporation of water during the core drilling and the sample weighting or/and (2) an heterogeneity of the soil. According to the location of the successive tests (Figure 5), the first hypothesis is most likely. The average dry densities were almost identical, confirming the homogeneity of this parameter over the massif.

Two mini-pressuremeter tests were performed at each temperature step: 20a, 20b, then 40a, 40b and finally 20c, 20d. The test 20a failed due to a membrane leakage. The results of the five other tests are presented in Figure 11.

The evaluation of the pressuremeter parameters (E_p , p_l and p_f), allowed the quantitative comparison of the tests at different temperatures (Table 4). The values of p_f and p_l decreased significantly when the massif was heated to 40 °C. Thus, the soil was yielded under lower pressures at 40 °C than at 20 °C. In other words, the results showed a softening of the soil with heating. The pressuremeter parameters for 20b and 20c-d, tests after decreasing the temperature from 40 to 20 °C, were close, which showed a good reversibility of the effect of the temperature variation in the range of 20-40 °C. The softening due to the heating and the reversibility of the observed effects on the creep pressure and the pressure limit were also noticed on the pressuremeter modulus (E_p) to a lower extent.

3.2.2 *Second series: cooling-heating cycle* (20-1-20 °C)

For the second series, the soil was compacted and then successively brought from 20 °C to 1 °C and then again to 20 °C. The average water content at the test depth varied from 30.8% to 32.6%. This variation shows the homogeneity of the water content within the massif. The average dry densities vary from 1.21 to 1.43 Mg/m³. Figure 12 shows the temperature evolution in the massif as measured by sensor S5 (Figure 4). After temperature equilibrium, two pressuremeter tests were conducted at the end of each temperature step. The results of the six tests are shown in Figure 13.

The mini-pressuremeter curves presented three domains as with the others mini-pressuremeter curves but, on the contrary to the heating tests (cf. 3.2.1), the cooling to 1 °C leads to a hardening of the soil. Thus, the pseudo-elastic domain was longer in the

case of the tests at 1 °C and the parameters p_f , p_l and E_p increased (Table 5). When the temperature raised back to 20 °C (20g and 20h), the obtained parameters decreased back but remained slightly higher than those given by tests 20e and 20f (before cooling). Thus, only a partial reversibility of the impact of the temperature variation was observed: the material appears to be definitely affected by the cooling phase to 1 °C.

3.2.3 First and second series comparison

The first two series of heating test (20-40-20 °C) and cooling test (20-1-20 °C) were used to study the effect of one temperature cycle. The preparation of a massif of 0.22 m^3 made it difficult to obtain exactly the same initial conditions in these two test series. Thus, the obtained values for the first tests at 20 °C of both test series (20b, 20e and 20f) were different. To compare these tests, the average values obtained at the end of the first step at 20 °C, $\overline{E_p}(20)$, $\overline{p_f}(20)$ and $\overline{p_l}(20)$, were considered as reference values and the other values were normalized with respect to the reference values according to the equations (1), (2) and (3).

$$E_{n.p} = \frac{E_p(T)}{\overline{E_p}(20)} \tag{1}$$

$$p_{n.f} = \frac{p_f(T)}{p_f(20)} \tag{2}$$

$$p_{n,l} = \frac{p_l(T)}{\overline{p_l(20)}} \tag{3}$$

Figure 14 shows the evolution of the normalized parameters at each step of each test series. The initial values at 20 °C, used as references, were close to 1.0 and the studied parameters decreased as temperature increased. The temperature effect was more pronounced for $p_{n,f}$ and $p_{n,l}$ than E_{n-p} . Cooling cycle evidenced the partial reversibility of the hardening. The values of $p_{n,f}$ increased by 30% as the soil was cooled

from 20 to 1 °C while it decreased by 7% as temperature reduced to 20 °C (Table 6). In the same way, $p_{n,l}$ increased by 21% with cooling then decreased by 9% with returning to 20 °C.

Heating cycle evidenced the reversibility of the thermal softening. Heating the soil from 20 to 40 °C decreased $p_{n,f}$ by 15%, while returning back to 20 °C increased it by 13%. In the same manner, $p_{n,l}$ decreased by 14% with heating and increased by 20% with returning to 20 °C.

3.3 *Effect of several temperature cycles*

The third and fourth test series, carried out on different massifs consisted in tests at 20, 1 and 40 °C. The initial temperature was 20 °C and the massif was brought to 1 and 40 °C successively completing three cooling-heating cycles, each heating or cooling step lasting at least 70 hours (Figure 15 and Figure 16). The thermal equilibrium was reached for each step before performing the mini-pressuremeter test.

Table 7 and Table 8 present all the obtained parameters and properties. The average water contents and the average dry densities variations according to the thermal solicitation were lower than 1%: the massifs were homogeneous.

The mini-pressuremeter tests were conducted at the end of the selected steps as shown in the Figure 15 and Figure 16. The pressuremeter curves (Figure 17 and Figure 18) were consistent with the three pressuremeter domains. In agreement with the previous series, a softening took place after heating the soil to 40 °C and a hardening was observed after cooling it to 1 °C, whatever the order of testing.

The pressuremeter parameters were calculated at each step of each series and presented in Figure 19. Over the third series, the p_f value increased slightly with cooling the massif from 20 to 1 °C, then it remained almost constant with heating/cooling

cycles (1-40-1 °C). The p_l values increased with heating and decreased with cooling, these results were consistent with those of the previous test series. Over the fourth series, the first pressuremeter test at 20 °C was inoperable and the corresponding values in Figure 19 were estimated assuming similarity between the 3^{rd} and the 4^{th} series.

The parameters evolution of the third and the fourth series were comparable. After the first thermal cycle, the variation of p_f with temperature falls within the parameter uncertainty domain and this parameter reached a stable value (Figure 19a, b). The evolution of p_l was different, the temperature impact remained measurable while being totally reversible. Its change rate could be estimated to 1.5 kPa/°C in the studied temperature range (1-40 °C) (Figure 19c, d).

As with the previous test series, the variation of E_p with temperature was lower than the other pressuremeter parameters (Figure 19e, f). The E_p evolution for the third massif was erratic and inconsistent with the previous series and seems not to be relevant. On the contrary, a slight E_p decrease, consistent with the previous tests, was measured when the fourth massif was heated.

4 Discussion

For the studied material at this initial state and under these test conditions, the creep pressure and the limit pressure decreased with increasing temperature which evidenced the thermal softening of the material. These results are consistent with the literature results obtained during oedometric or triaxial tests (Table 1).

The effect of first temperature cycle was more significant than the following cycles. The same phenomenon exists in the other domains of the soil mechanics. For example, during the investigation of the effect of cyclic wetting and drying on the characteristics of clays, Al-Homoud et al. (1995), Basma et al. (1996) and Nowamooz et

al. (2013) showed that the first cycle caused the most reduction in swelling potential and observed that the amount of shrinkage for the first cycle was the largest and steadily decreased in later cycles. They explained it by more important modification in the clay structure after the first cycle than the subsequent cycles.

The temperature effect on the pressuremeter parameters could be due to the combination of the change in pore water viscosity and the expansion/contraction of the soil components which leads sometimes to the reorganization of the soil microstructure. Indeed, the decrease in pore water viscosity by heating can facilitate sliding between grains, thereby reducing the measured parameters. The heating of the massif from 20 to 40 °C led to the decrease in the studied parameters. It could be mainly due to the decrease in pore water viscosity and thermal softening and consequently it was reversible with cooling the massif to 20 °C. On the contrary, the cooling of the massif from 20 to 1 °C led to the increase in the pressuremeter parameters and since the effect of cooling was partially reversible, it could be concluded that there was also the reorganization of the soil microstructure during the cooling which might be irreversible. This reorganization took place during the first cycle and the effect of the subsequent cycles was less significant. These hypotheses should be confirmed by an investigation on the modification of the microstructure of soil with heating and cooling.

5 Conclusions

The objective of this paper was to quantify the effect of cyclic temperature variations on soil parameters derived from the pressuremeter test used in pile design. As a result and for the first time, a device was developed to measure the mini-pressuremeter parameters in laboratory conditions and at different temperatures. The evolution of the pressuremeter modulus (E_p) , the creep pressure (p_f) and the limit pressure (p_l) was

measured under imposed thermal conditions on a clayey compacted soil which has been selected for its high sensitivity to temperature variations.

The results of the mini-pressuremeter tests on the compacted illitic soil showed a decrease in the creep pressure and the limit pressure as temperature increased (thermal softening), while the variation of the pressuremeter modulus was less significant, this could be due to the well-known higher variability of this parameter. The results tended to show the reversibility of the temperature effects on the measured parameters for one heating cycle in the tested temperature range (20-40 °C), whereas the effect of the first cooling phase to 1 °C was only partially reversible. For the following cycles, the effect of the temperature variation was lower. After more than one cycle, p_f seemed to reach an equilibrium value independent of the temperature. The p_l kept its temperature dependence beyond the 1st cycle but its evolution (1.5 kPa/°C) was reversible in the studied temperature range (1-40 °C).

The limit pressure is the key parameter of calculating ultimate load of piles, so its variation with temperature should be taken into account to estimate the bearing capacity of geothermal piles particularly when they are built in temperature sensitive soils.

The trends demonstrated in this laboratory experiments study should be confirmed during full-scale in situ pressuremeter tests which need to be conducted in the proximity of an instrumented thermo-active pile.

6 ACKNOWLEDGEMENTS

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Table 1. Temperature effect on elastic domain (yield locus).

Authors	Studied soil	Measuring methods	Parameters	Effect of heating
Hueckel & Baldi, (1990)	Saturated silty clays	Drained triaxial shear test	$T = 20, 90 ^{\circ}C$	Contraction
Tanaka et al. (1997)	Saturated illite	Undrained triaxial shear test	T = 28, 65, 100 °C	Contraction
Graham et al. (2001)	Saturated illite	Isotropic consolidation (Triaxial)	T = 28, 65, 100 °C	Contraction
Cekerevac & Laloui (2004)	Saturated kaolin	Drained triaxial shear test	$ \rho_{d} = 1.35 \text{ Mg/m}^3 $ T = 22, 60, 90 °C	Contraction
Marques et al., (2004)	Saturated clay	Undrained triaxial shear test and oedometric test	$\rho \cong 1.5 \text{ Mg/m}^3$ T = 10, 20, 50 °C	Contraction
Uchaipichat & Khalili (2009)	Unsaturated Bourke silt	Suction and temperature controlled compression shear tests	$\rho_d = 1.53 \text{ Mg/m}^3$ T = 25, 40, 60 °C	Contraction

Table 2. Details of mini-pressuremeter tests series.

Series	Temperature (°C)	W _{target} (%)	$ ho_{d.target} \ ({ m Mg/m^3})$	Objective		
Preliminary series	20	31.3	1.29	To verify the efficiency of the compaction protocol and massif homogeneity. To test the possibility to perform 6 tests in a massif.		
1 st series	20-40-20	31.3	1.29	To quantify the effect of heating: 2 tests at each temperature (20a-b, 40a-b, 20c-d)		
2 nd series	20-1-20	31.3	1.29	To quantify the effect of cooling: 2 tests at each temperature (20e-f, 1a-b, 20g-h)		
3 rd series	20-1-40-1-40-1-40-1	31.3	1.29	To quantify the effect of temperature cycle: test during the first 4 steps et des the last 2 steps (20i, 1c, 40c, 1d, 40d, 1e)		
4 th series	20-1-40-1-40-1-40	31.3	1.29	To quantify the effect of temperature cycle: 1 test at each temperature (1f, 40e, 1g, 40f, 1h, 40g)		

Table 3. Pressuremeter parameters for preliminary tests series.

Test	Temperature (°C)	w (%)	ρ_d (Mg/m ³)	E_p (MPa)	p_f (kPa)	p_l (kPa)
1	20	31.6	1.30	4,02	250	467
2	20	31.6	-	4,26	252	457
3	20	31.8	1.26	4,02	220	409
4	20	31.8	1.33	5,19	250	457
5	20	31.7	1.29	4,97	242	447

Table 4. Pressuremeter parameters for the 1st series: heating cycle (20-40-20 °C).

Test	Temperature (°C)	w (%)	ρ_d (Mg/m ³)	$p_f(kPa)$	p_l (kPa)	E_p (MPa)
20b	20	31.0	1.31	194	355	3.54
40a	40	30.1	1.31	166	321	4.16
40b	40	30.4	1.31	165	288	3.25
20c	20	31.1	1.31	197	379	4.54
20d	20	31.1	1.32	178	352	3.89

Table 5. Pressuremeter parameters for the 2^{nd} series: cooling cycle (20-1-20 °C).

Test	Temperature (°C)	w (%)	ρ_d (Mg/m ³)	$p_f(kPa)$	p_l (kPa)	E_p (MPa)
20e	20	31.8	1.34	132	305	2.47
20f	20	32.0	1.26	133	246	2.28
1a	1	32.2	1.29	181	339	3.40
1b	1	31.8	1.27	163	330	2.68
20g	20	32.1	1.27	157	286	2.36
20h	20	31.6	1.24	162	323	2.97

Table 6. Variation of p_f and p_l for the first and second series of the tests.

Initial temperature (°C)	Final temperature (°C)	$\Delta p_f(\%)$	$\Delta p_l(\%)$
20	1	+30	+21
1	20	-7	-9
20	40	-15	-14
40	20	+13	+20

Table 7. Pressuremeter parameters for the 3^{rd} series: temperature cycling (20-1-40-1-40-1-40-1-40-1).

Test	Temperature (°C)	w (%)	ρ_d (Mg/m ³)	p_f (kPa)	p_l (kPa)	E_p (MPa)
20i	20	31.4	1.25	103	240	5.51
1c	1	31.4	1.26	124	303	5.73
40c	40	32.1	1.25	125	236	7.68
1d	1	31.7	1.25	123	337	4.16
40d	40	31.5	1.28	143	270	6.85
1e	1	31.6	1.24	133	310	5.66

Table 8. Pressuremeter parameters for the 4^{th} series: temperature cycling (20-1-40-1-40-1-40-0°C).

Test	Temperature (°C)	w (%)	ρ_d (Mg/m ³)	$p_f(kPa)$	p_l (kPa)	E_p (MPa)
40e	40	31.2	1.30	174	305	6.24
1g	1	31.5	1.24	170	357	6.97
40f	40	31.2	1.28	170	307	6.18
1h	1	31.0	1.24	160	344	6.16
40g	40	31.1	1.25	153	276	5.76

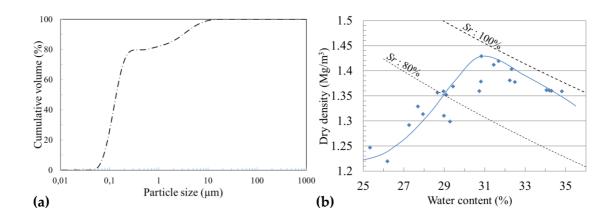


Figure 1 : Characteristic of the studied illitic soil: a) particle size distribution and b) compaction curve, *Sr*: The degree of saturation.

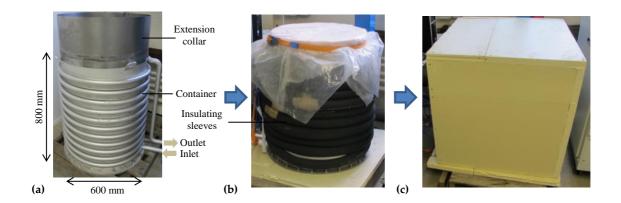


Figure 2: Thermo-regulated metric scale container: (a) cylindrical stainless steel container; (b) massif in the container isolated by insulating sleeves and plastic film; (c) assembly in the box made of extruded polystyrene plates.



Figure 3. Sample preparation: (a) mixing; (b) compaction.

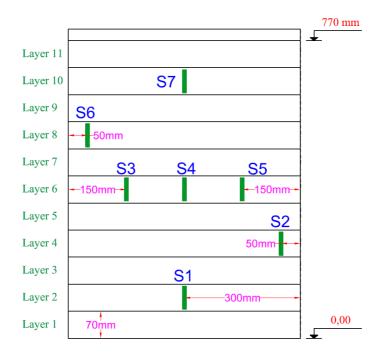


Figure 4. Position of the 7 thermal sensors in the massif.

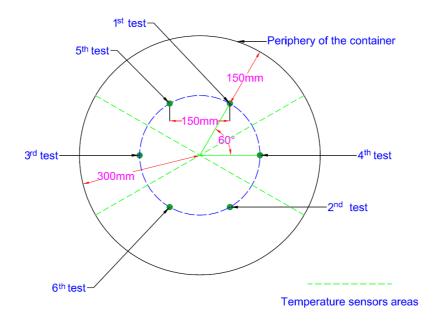


Figure 5. Positions of the test points and thermal sensors in the massif.

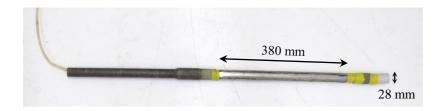


Figure 6. Mini-pressuremeter probe APAGEO®.

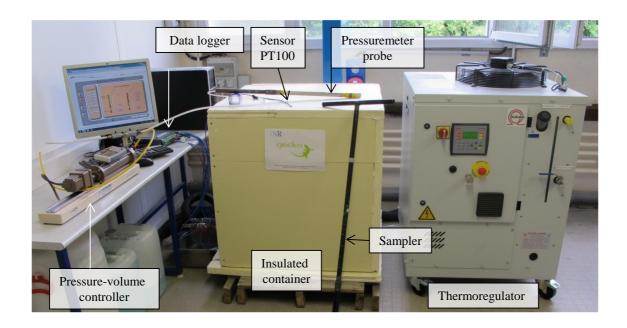


Figure 7. Elements of the developed experimental device.

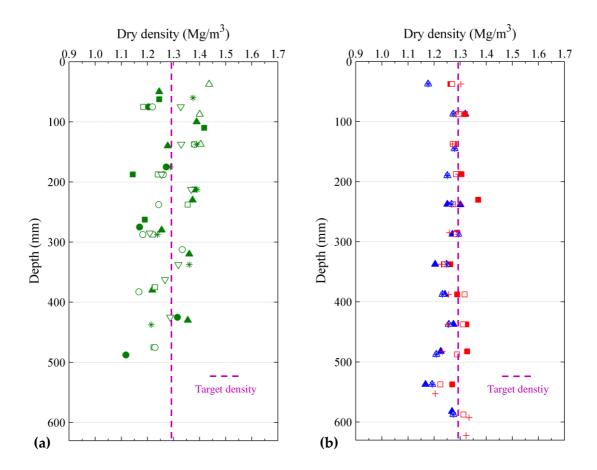


Figure 8. Dry density Profiles: (a) compaction in 8 layers (preliminary series); (b) compaction in 11 layers (4th series).

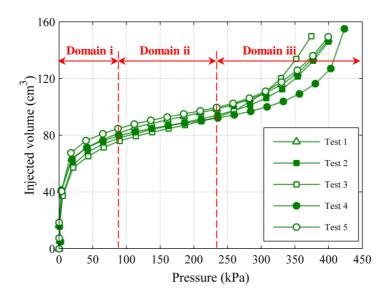


Figure 9. Pressuremeter curves: preliminary tests series.

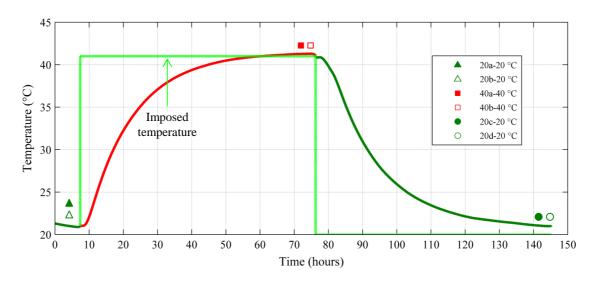


Figure 10. Temperature variation in the 1^{st} massif at 150 mm from the edge of the container (S5) and the chronology of the tests: heating-cooling cycle (20-40-20 °C).

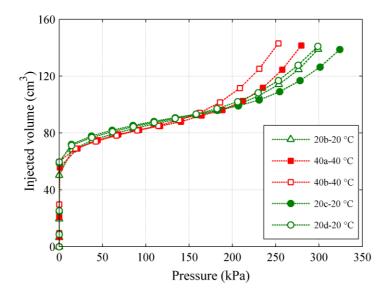


Figure 11. Pressuremeter curves for the 1st series: heating-cooling cycle (20-40-20 °C).

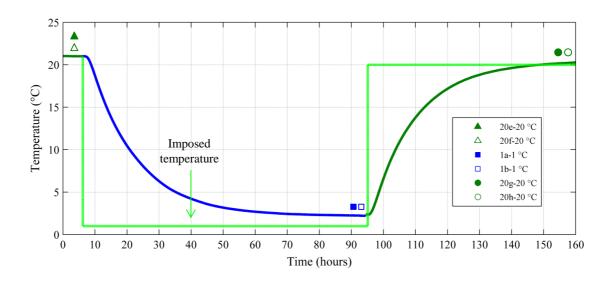


Figure 12. Temperature variation in the 2nd massif at 150 mm from the edge of the container (S5) and the chronology of the tests cooling-heating cycle (20-1-20 °C).

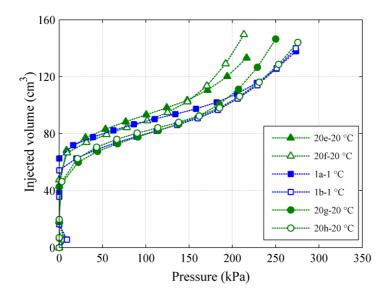


Figure 13. Pressuremeter curves for the 2nd series: cooling-heating cycle (20-1-20 °C).

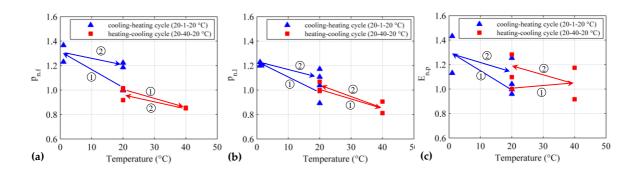


Figure 14. Variation of reduced pressuremeter parameters with temperature: (a) $p_{n,f}$; (b) $p_{n,l}$; (c) $E_{n,M}$.

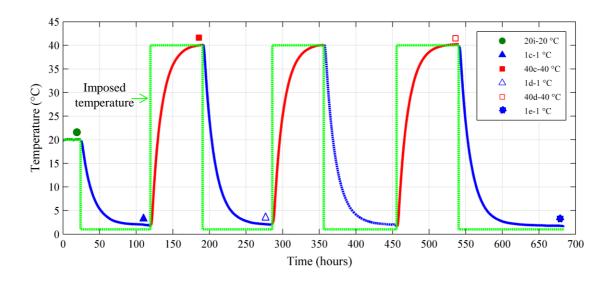


Figure 15. Temperature variation in the massif at 150 mm from the edge of the container and the chronology of the 3rd series.

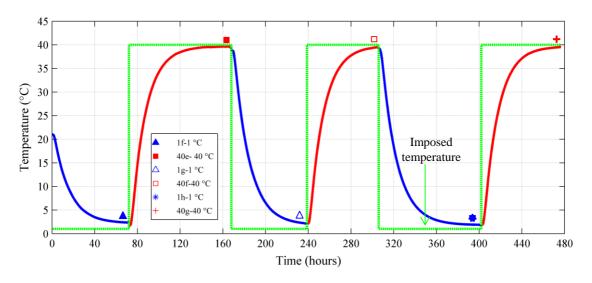


Figure 16. Temperature variation in the massif at 150 mm from the edge of the container and the chronology of the 4^{th} series.

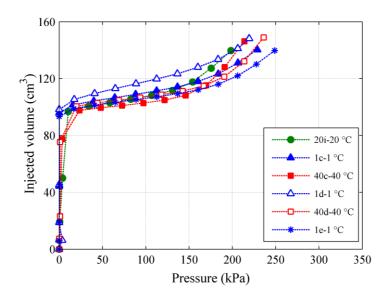


Figure 17. Pressuremeter curves of the 3rd series tests.

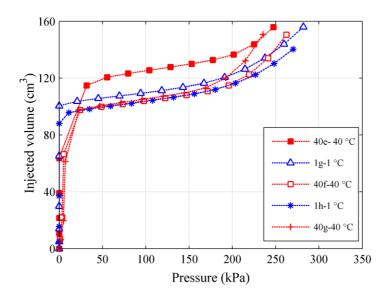


Figure 18. Pressuremeter curves of the 4th series tests.

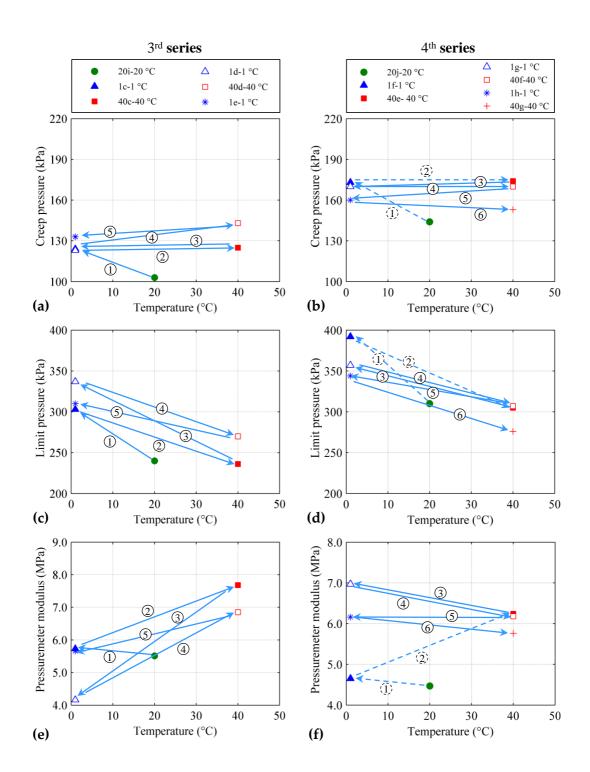


Figure 19. Variation of creep pressure (p_f) (a-b), limit pressure (p_l) (c-d) and pressuremeter modulus (E_p) (e-f) with temperature for the 3^{rd} and the 4^{th} test series. Dotted lines correspond to estimated parameters.