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► **To cite this version:**

Nikolay I Makarenko, Janna L Maltseva, Eugene Morozov, Roman Tarakanov, Kseniya A. Ivanova. Internal waves in marginally stable abyssal stratified flows. *Nonlinear Processes in Geophysics*, In press, 10.5194/npg-2018-1 . hal-01693709

HAL Id: hal-01693709

<https://hal.science/hal-01693709v1>

Submitted on 26 Jan 2018

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Internal waves in marginally stable abyssal stratified flows

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1 Abstract.

2 The problem on internal waves in a weakly stratified two-layered fluid is studied semi-analytically. We discuss the 2.5-layer
3 fluid flows with exponential stratification of both layers. The long-wave model describing travelling waves is constructed by
4 means of scaling procedure with a small Boussinesq parameter. It is demonstrated that solitary wave regimes can be affected
5 by the Kelvin — Helmholtz instability arising due to interfacial velocity shear in upstream flow.

6 1 Introduction

7 In this paper, we consider an analytical model of internal solitary waves in a two-layer fluid with the density continuously
8 increasing with depth in both layers. This model is a development of non-linear two-layer models previously suggested
9 by Ovsvyannikov (1985), Miyata (1985) and Choi & Camassa (1999), as well as the latest 2.5-layer models considered by
10 Voronovich (2003), Makarenko and Maltseva (2008, 2009a,b). Two-layer approximation is a standard model of sharp pycno-
11 cline in a stratified fluid with constant densities in each layer, but discontinuous at the interface. Correspondingly, the 2.5-layer
12 model takes into account a slight density gradient in stratified layer which is comparable with the density jump at the interface.
13 In all these cases, internal solitary waves can be described in closed form by the solutions resulting from the quadrature

$$14 \left(\frac{d\eta}{dx} \right)^2 = f(\eta) \quad (1)$$

15 for stationary wave elevation $\eta(x)$. The simplest version of non-linearity f appears in a two-layer system, hence, it is the rational
16 function $f(\eta) = P(\eta)/Q(\eta)$ in this case when P is a fourth degree polynomial, and Q depends linearly on η . Equation (1)
17 also appears as a travelling wave equation for nonlinear evolution systems being similar to single-layer dispersive Green –
18 Naghdi model (see Choi & Camassa, 1999). These non-linear dispersive equations can be obtained by means of long-wave
19 perturbation technique as well as by Whitham's variational method. Several authors noted that solitary wave solutions of such
20 approximate models are in good agreement with the numerical solutions of fully nonlinear Euler equations for a perfect two-
21 layer fluid. In this context, Camassa & Tiron (2011) also compared numerical travelling wave-solutions, supported by smooth



1 stratification, with the known explicit solitary-wave solutions in order to optimize a two-layer model of the Euler system with
2 smooth stratification.

3 We apply the method of derivation involving asymptotic analysis of the non-linear Dureuil-Jacotin — Long equation that
4 results from fully nonlinear Euler equations of stratified fluid. Long-wave scaling procedure uses a small Boussinesq parameter
5 which characterizes slightly increasing density in the layers and a small density jump at their interface. This method combines
6 the approaches applied formerly to a pure two-fluid system with perturbation technique discussed for the first time by Long
7 (1965) and developed by Benney and Ko (1978) for a continuous stratification. Parametric range of solitary wave is considered
8 in the framework of the constructed mathematical model. It is demonstrated that these wave regimes can approach the paramet-
9 ric domain of the Kelvin — Helmholtz instability. The stability of solitary travelling-wave solutions of the Euler equations for
10 continuously stratified, near two-layer fluids was studied numerically and analytically by Almgren, Camassa, & Tiron (2012).
11 They demonstrated that the wave-induced shear can locally reach unstable configurations and give rise to local convective
12 instability. This is in good qualitative agreement with the laboratory experiments performed by Grue et al. (2000). It seems that
13 such a marginal stability of long internal waves could explain the formation mechanism of a very long billow trains in abyssal
14 flows observed by Van Haren et al. (2014).

15 2 Basic Equations

16 We consider a 2D motion of inviscid two-layer fluid which is weakly stratified due to gravity in both layers. The fully nonlinear
17 Euler equations describing the flow are

$$18 \quad \rho(u_t + uu_x + vu_y) + p_x = 0, \quad (2)$$

$$19 \quad \rho(v_t + uv_x + vv_y) + p_y = -\rho g, \quad (3)$$

$$20 \quad \rho_t + u\rho_x + v\rho_y = 0, \quad (4)$$

$$21 \quad u_x + v_y = 0, \quad (5)$$

22 where ρ is the fluid density, (u, v) is the fluid velocity, p is the pressure and g is the gravity acceleration. We assume that the
23 flow domain is bounded by the flat bottom $y = -h_1$ and the rigid lid $y = h_2$ (see Fig. 1), with the boundary condition

$$24 \quad v = 0 \Big|_{y=-h_1, y=h_2}. \quad (6)$$

25

26 The layers are separated by the interface $y = \eta(x, t)$ with the equilibrium level at $y = 0$. Non-linear kinematic and dynamic
27 boundary conditions at this interface are

$$28 \quad \eta_t + u\eta_x = v \Big|_{y=\eta}, \quad [p] = 0 \Big|_{y=\eta} \quad (7)$$

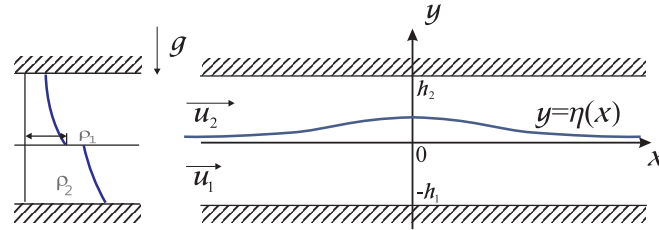


Figure 1. Scheme of the flow

1 where the square brackets denote the discontinuity jump at the interface between the layers. Non-disturbed parallel flow has
 2 no vertical velocity and elevation (i.e. $v = 0$, $\eta = 0$) but the horizontal velocity $u = u_0(y)$ may be piece-wise constant,

$$3 \quad u_0(y) = \begin{cases} u_1 & (-h_1 < y < 0), \\ u_2 & (0 < y < h_2). \end{cases} \quad (8)$$

4 In this stationary case, the fluid density $\rho = \rho_0(y)$ and pressure $p = p_0(y)$ should be coupled by the hydrostatic equation
 5 $dp_0/dy = g\rho_0$. We consider the density profile depending exponentially on height,

$$6 \quad \rho_0(y) = \begin{cases} \rho_1 \exp(-N_1^2 y/g) & (-h_1 < y < 0), \\ \rho_2 \exp(-N_2^2 y/g) & (0 < y < h_2), \end{cases} \quad (9)$$

7 where $N_j = \text{const}$ is the Brunt — Väisälä frequency in the j -th layer, and constant densities ρ_1 and ρ_2 are related as $\rho_2 < \rho_1$.
 8 The special case $N_j = 0$ ($j = 1, 2$) gives a familiar two-fluid system with piece-wise constant density $\rho = \rho_j$ in the j -th layer.

9 Further we consider a steady non-uniform flow, hence we have $\eta_t = 0$ and $u_t = v_t = \rho_t = 0$ in Eqs. (2) – (4). We introduce
 10 the stream function ψ by standard formulae $u = \psi_y$, $v = -\psi_x$, hence the mass conservation implies the dependence $\rho = \rho(\psi)$,
 11 and pressure p can be found from the Bernoulli equation

$$12 \quad \frac{1}{2} |\nabla\psi|^2 + \frac{1}{\rho(\psi)} p + gy = b(\psi). \quad (10)$$

13 Seeking for a solitary-wave solutions, we require that the upstream velocity of the fluid (u, v) tends to $(u_j, 0)$ as $x \rightarrow -\infty$. In
 14 this case, boundary conditions (6) transform to the conditions for the stream function as

$$15 \quad \psi = -u_1 h_1 \Big|_{y=-h_1}, \quad \psi = 0 \Big|_{y=\eta}, \quad \psi = u_2 h_2 \Big|_{y=h_2}. \quad (11)$$

16 It is known (Yih, 1980) that system (2) – (5) can be reduced in a stationary case to the non-linear Dureuil-Jacotin — Long
 17 (DJL) equation for the stream function

$$18 \quad \rho(\psi) \nabla^2 \psi + \rho'(\psi) \left(gy + \frac{1}{2} |\nabla\psi|^2 \right) = H'(\psi). \quad (12)$$

19 Here, the function $H(\psi) = \rho(\psi)b(\psi)$ involves the Bernoulli function $b(\psi)$ and the density function $\rho(\psi)$, so that H is specified
 20 by the upstream condition. More exactly, the density function is determined by the relation $\rho(\psi) = \rho_0(\psi/u_j)$ in the j -th layer,



1 and the Bernoulli function $b(\psi)$ is defined by the formula

$$2 \quad b = \begin{cases} \frac{1}{2} u_1^2 + g \frac{\psi}{u_1} + \frac{g^2}{N_1^2} \left(1 - e^{-\frac{N_1^2 \psi}{g u_1}} \right) & (-h_1 < y < \eta(x)), \\ \frac{1}{2} u_2^2 + g \frac{\psi}{u_2} + \frac{g^2}{N_2^2} \left(1 - e^{-\frac{N_2^2 \psi}{g u_2}} \right) & (\eta(x) < y < h_2). \end{cases}$$

3 As a consequence, we can rewrite the DJL equation (12) as follows:

$$4 \quad \nabla^2 \psi = \frac{N_j^2}{g u_j} \left\{ g \left(y - \frac{\psi}{u_j} \right) + \frac{1}{2} (|\nabla \psi|^2 - u_j^2) \right\}, \quad (13)$$

5 where $j = 1$ is related to the lower layer, and $j = 2$ to the upper layer. Further, in accordance with relations (7) and (10), the
 6 continuity of pressure p provides non-linear boundary condition for stream function ψ

$$7 \quad [\rho(\psi)(|\nabla \psi|^2 + 2gy - 2b(\psi))] = 0|_{y=\eta}. \quad (14)$$

8 Using the explicit form of functions $\rho(\psi)$ and $b(\psi)$, condition (14) can be also rewritten in detail as follows:

$$9 \quad 2g(\rho_1 - \rho_2)\eta = \rho_2 (|\nabla \psi|^2 - u_2^2)|_{y=\eta(x)+0} - \rho_1 (|\nabla \psi|^2 - u_1^2)|_{y=\eta(x)-0}.$$

10 We reformulate this boundary condition in view of conservation of the total horizontal momentum in a steady two-layer flow,
 11 which has integral formulation

$$12 \quad \int_{-h_1}^{h_2} (p + \rho u^2) dy = C$$

13 where constant C is determined by the upstream condition. Excluding pressure p from here using the Bernoulli equation (10)
 14 leads to the integral relation

$$15 \quad \rho_1 \int_{-h_1}^{\eta(x)} e^{-\frac{N_1^2 \psi}{g u_1}} \Psi_1 dy + \rho_2 \int_{\eta(x)}^{h_2} e^{-\frac{N_2^2 \psi}{g u_2}} \Psi_2 dy = C \quad (15)$$

16 where the integrand functions Ψ_j are

$$17 \quad \Psi_j = \psi_y^2 - \psi_x^2 + u_j^2 + 2g \left(\frac{\psi}{u_j} - y \right) - \frac{2g^2}{N_j^2} \left(e^{-\frac{N_j^2 \psi}{g u_j}} - 1 \right),$$

18 and constant C depends on the parameters of the upstream flow as follows:

$$19 \quad C = 2\rho_1 g \left[\left(e^{-\frac{N_1^2 h_1}{g}} - 1 \right) \left(\frac{u_1^2}{N_1^2} + \frac{g^2}{N_1^4} \right) - \frac{g h_1}{N_1^2} \right] + \\ + 2\rho_2 g \left[\left(1 - e^{-\frac{N_2^2 h_2}{g}} \right) \left(\frac{u_2^2}{N_2^2} + \frac{g^2}{N_2^4} \right) - \frac{g h_2}{N_2^2} \right].$$

20 It is important here that the integral relation (15) is equivalent to the boundary condition (14) which is rather simple. This
 21 equivalence can be checked immediately by differentiation the relation (15) with respect to the variable x , so the integrals can
 22 be evaluated explicitly due to Eq.(13). Equation (15) will be used later instead of (14) by the construction model differential
 23 equation for the function $\eta(x)$ describing strongly nonlinear waves.



1 3 Non-Dimensional Formulation

2 Now we introduce scaled independent variables \bar{x} , \bar{y} and scaled unknown functions $\bar{\eta}$, $\bar{\psi}$ in order to reformulate the basic
 3 equations in the dimensionless form. Namely, the fixed ratio h_1/π is used as an appropriate length scale for x , y , η , and
 4 normalized volume discharges $u_j h_j/\pi$ serve as the units for the stream function; thus, we have

$$5 \quad (x, y, \eta) = \frac{h_1}{\pi} (\bar{x}, \bar{y}, \bar{\eta}), \quad \psi = \frac{u_j h_j}{\pi} \bar{\psi}$$

6 separately in the lower layer ($j = 1$) or in the upper layer ($j = 2$). The number π is only introduced here due to the specific
 7 form of trigonometric modal functions which are typical for the exponential density (9). Scaling procedure with this density
 8 profile uses the Boussinesq parameters σ_1 , σ_2 and the Atwood number μ defined by the formulae

$$9 \quad \sigma_j = \frac{N_j^2 h_j}{\pi g} \quad (j = 1, 2), \quad \mu = \frac{\rho_1 - \rho_2}{\rho_2}. \quad (16)$$

10 Here, constants σ_j characterize the slope of the density profile in continuously stratified layers, and parameter μ determines
 11 the density jump at interface.

12 Following Turner (1973), we introduce densimetric (or internal) Froude number

$$13 \quad F_j = \frac{u_j}{\sqrt{g_j h_j}} \quad (j = 1, 2)$$

14 which presents scaled fluid velocity u_j in the j -th layer, defined with reduced gravity acceleration $g_j = (\rho_1 - \rho_2)g/\rho_j$. In
 15 addition to the Froude numbers F_j , it is also convenient to use the pair of the Long's numbers λ_j given by the formula

$$16 \quad \lambda_j = \frac{N_j h_j}{\pi u_j} \quad (j = 1, 2).$$

17 The Long's numbers λ_j are coupled with the Boussinesq parameters σ_1 , σ_2 , the Atwood number μ and the Froude numbers F_j
 18 by the relations

$$19 \quad \lambda_1^2 = \frac{\pi \sigma_1 (1 + \mu)}{\mu F_1^2}, \quad \lambda_2^2 = \frac{\pi \sigma_2}{\mu F_2^2}. \quad (17)$$

20 Finally, we introduce the ratio of undisturbed thicknesses of the layers $r = h_1/h_2$. By that notation, we locate the bottom
 21 as $\bar{y} = -\pi$, and relation $\bar{y} = \pi/r$ defines the rigid lid. Thus, we obtain the equations for scaled stream function $\bar{\psi}$ and non-
 22 dimensional wave elevation $\bar{\eta}$ as follows (bar is omitted throughout what follows):

$$23 \quad \nabla^2 \psi + \lambda_1^2 (\psi - y) = \frac{1}{2} \sigma_1 (|\nabla \psi|^2 - 1) \quad (18)$$

24 in the lower layer $-\pi < y < \eta(x)$, and

$$25 \quad \nabla^2 \psi + \lambda_2^2 r^2 (\psi - ry) = \frac{1}{2} \sigma_2 (|\nabla \psi|^2 - r^2) \quad (19)$$

26 in the upper layer $\eta(x) < y < \pi/r$. Kinematic boundary conditions (11) can be rewritten now as follows:

$$27 \quad \psi(x, -\pi) = -\pi, \quad \psi(x, \eta(x)) = 0, \quad \psi(x, \pi/r) = \pi. \quad (20)$$



1 Correspondingly, Eq. (14) providing continuity of pressure at interface $y = \eta(x)$ leads to nonlinear boundary condition

$$2 \quad 2\eta = F_2^2(|\nabla\psi|^2 - r^2)|_{y=\eta+0} - F_1^2(|\nabla\psi|^2 - 1)|_{y=\eta-0}, \quad (21)$$

3 and the dimensionless version of integral relation (15) takes the form

$$4 \quad \int_{-\pi}^{\eta} e^{-\sigma_1\psi} \Psi_1 dy + \int_{\eta}^{\pi/r} e^{-\sigma_2\psi} \Psi_2 dy = C \quad (22)$$

5 where is denoted

$$6 \quad \Psi_1 = \frac{\mu F_1^2}{2} \left(\psi_y^2 - \psi_x^2 + 1 \right) + \frac{1 + \mu}{\pi} \left(\psi - y - \frac{e^{\sigma_1\psi} - 1}{\sigma_1} \right),$$

$$7 \quad \Psi_2 = \frac{\mu F_2^2}{2r^3} \left(\psi_y^2 - \psi_x^2 + r^2 \right) + \frac{1}{\pi r} \left(\psi - ry - \frac{e^{\sigma_2\psi} - 1}{\sigma_2} \right).$$

9 Constant

$$10 \quad C = \pi\mu \left(F_1^2 + \frac{F_2^2}{r^2} \right) + (1 + \mu) \frac{e^{\sigma_1\pi} - 1 - \sigma_1\pi}{\pi(\lambda_1^2 + \sigma_1^2)} + \frac{1 - \sigma_2\pi - e^{-\sigma_2\pi}}{\pi r^2(\lambda_2^2 + \sigma_2^2)}$$

11 is chosen here so that the horizontal upstream flow given by the solution

$$12 \quad \eta = 0, \quad \psi_0(y) = \begin{cases} y & (-\pi < y < 0), \\ ry & (0 < y < \pi/r) \end{cases} \quad (23)$$

13 satisfies momentum relation (22).

14 The model of fully nonlinear travelling waves in a two-layer irrotational flows, with the interface $y = \eta(x)$ between the fluids
 15 with constant densities ρ_2 in the upper layer and $\rho_1 > \rho_2$ in the lower layer, can be specified as follows. In this limit case, at
 16 least formally, the Boussinesq parameters σ_j and Long's numbers λ_j vanish: $\sigma_1 = \sigma_2 = \lambda_1 = \lambda_2 = 0$. Therefore, we obtain the
 17 Laplace equation

$$18 \quad \nabla^2\psi = 0 \quad (24)$$

19 instead of Eqs. (18)–(19), but all the boundary conditions (20) and (21) still remain unchanged.

20 4 Spectrum of Harmonic Waves

21 In many cases, parametric range of solitary waves can be determined *a priori* as the domain being supercritical with respect to
 22 the spectrum of small-amplitude sinusoidal waves. It is helpful while the critical phase speed can be simply defined from the



1 dispersion relation of infinitesimal waves. In our case, linearizing of Eqs. (18)–(21) for the upstream solution (23) leads to the
 2 dispersion relation

$$3 \quad \Delta(k; F_1, F_2) = 0 \quad (25)$$

4 for stationary harmonic wave-packets

$$5 \quad \eta(x) = a e^{ikx}, \quad \psi = \psi_0(y) + W(y) e^{ikx}.$$

6 Here k is the non-dimensional wave-number, a is the amplitude of interfacial wave, and $W(y)$ is the modal eigenfunction
 7 which describes deformation of streamlines within the fluid layers. For the given Long's numbers λ_1, λ_2 and the Boussinesq
 8 parameters σ_1, σ_2 , we also introduce non-dimensional values

$$9 \quad \varkappa_j = \sqrt{|\lambda_j^2 - k_j^2 - \frac{1}{4} \pi^2 \sigma_j^2|} \quad (j = 1, 2), \quad (26)$$

10 where $k_1 = rk$ and $k_2 = k$ are dimensionless wave-numbers specified for each layer. According to these notations, dispersion
 11 function $\Delta(k; F_1, F_2)$ in (25) has the form

$$12 \quad \Delta = F_1^2 \left(\varkappa_1 \text{Cot}_1 \varkappa_1 + \frac{\pi \sigma_1}{2} \right) + F_2^2 \left(\varkappa_2 \text{Cot}_2 \varkappa_2 - \frac{\pi \sigma_2}{2} \right) - 1$$

13 where functions Cot_j ($j = 1, 2$) are denoted as follows:

$$14 \quad \text{Cot}_j \varkappa_j = \begin{cases} \cot \varkappa_j & (\lambda_j^2 > k_j^2 + \frac{1}{4} \pi^2 \sigma_j^2) \\ \coth \varkappa_j & (\lambda_j^2 < k_j^2 + \frac{1}{4} \pi^2 \sigma_j^2). \end{cases}$$

15 In fact, function Δ takes such a combined form since modal function $W(y)$ depends on y trigonometrically or hyperbolically,
 16 if the radicand term $\lambda_j^2 - k_j^2 - \frac{1}{4} \pi^2 \sigma_j^2$ in (26) is positive or negative. Explicit formulae for these modal eigenfunctions $W(y)$
 17 are given in Appendix A.

18 Spectrum of stationary harmonic waves, defined on the (F_1, F_2) -plane, is formed by the Froude points (F_1, F_2) so that
 19 dispersion function $\Delta(k; F_1, F_2)$, which is even in k , has at least one pair of real roots $\pm k$. Wave modes differ by the number
 20 of these pairs, and this number can change only by passing of the root across the value $k = 0$. Therefore, the modal bounds
 21 should satisfy the equation $\Delta(0; F_1, F_2) = 0$; these bounds are defined by separate branches of the curve

$$22 \quad F_1^2 \left\{ \sqrt{\lambda_1^2 - \left(\frac{\pi \sigma_1}{2}\right)^2} \cot \sqrt{\lambda_1^2 - \left(\frac{\pi \sigma_1}{2}\right)^2} - \frac{\pi \sigma_1}{2} \right\} + \quad (27)$$

$$+ F_2^2 \left\{ \sqrt{\lambda_2^2 - \left(\frac{\pi \sigma_2}{2}\right)^2} \cot \sqrt{\lambda_2^2 - \left(\frac{\pi \sigma_2}{2}\right)^2} + \frac{\pi \sigma_2}{2} \right\} = 1$$

23 where parameters λ_j should be coupled with the Froude numbers F_j using the formulae (17).

24 We emphasize that parameters σ_j characterize the slope of density profile in continuously stratified layers, and μ defines the
 25 density jump at the interface. As usual, all these parameters are small in the case of low stratification. However, the interfacial

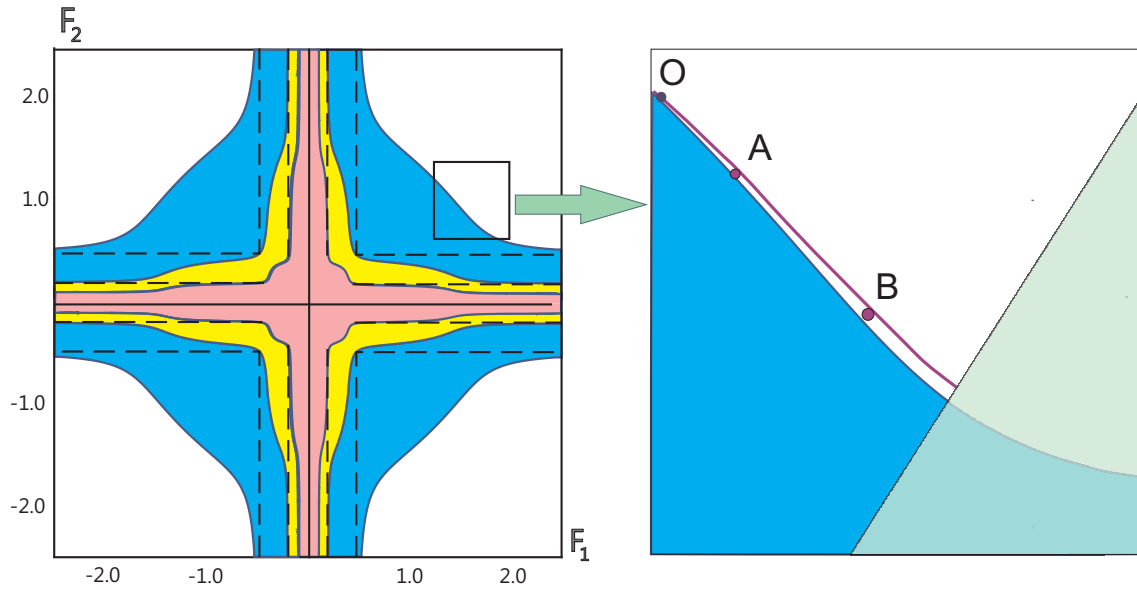


Figure 2. Spectrum of linear waves (colored modes 1-3) (left) and fragment of parametric domain of solitary waves (right).

1 mode dominates over the modes of internal waves in stratified layers when $\sigma_j \ll \mu$ is valid. In this limit case, linearized
 2 boundary conditions (20)–(21), considered together with the linear Laplace equation (24), lead to the standard dispersion
 3 relation of two-layer fluid

$$4 \quad F_1^2 r k \coth r k + F_2^2 k \coth k = 1. \quad (28)$$

5 This relation determines only a single pair of real wave-numbers of the interfacial mode, so the spectral domain of a perfect
 6 2-layer system occupies the unit disk

$$7 \quad F_1^2 + F_2^2 \leq 1. \quad (29)$$

8 The 2.5-layer model starts with the hypotheses that the Boussinesq parameters σ_1 , σ_2 and the Atwood number μ are of the
 9 same order, so we can use a single small parameter σ by setting

$$10 \quad \sigma = \sigma_1 = \sigma_2 = \mu. \quad (30)$$

11 The limit passage $\sigma \rightarrow 0$ is singular because the Long's numbers λ_j involve the ratios σ_j/μ in formulae (17). However,
 12 condition (30) allows us to simplify the spectral portrait, hence modal curve (27) defining the critical wave speeds takes the
 13 form

$$14 \quad \sqrt{\pi} F_1 \cot \frac{\sqrt{\pi}}{F_1} + \sqrt{\pi} F_2 \cot \frac{\sqrt{\pi}}{F_2} = 1. \quad (31)$$



1 Figure 2 demonstrates the parts of the spectrum defined by curve (31) for the dominating modes. The domain covered only
 2 by the first mode is marked with the blue color. Correspondingly, the embedded domain of the second mode is highlighted
 3 with the yellow, and the third mode is marked with the pink color. It is important that this spectrum differs essentially from the
 4 ordinary 2-layer spectrum (29), even the flow is characterized with a pair of the Froude numbers F_1 , F_2 , defined by the same
 5 manner. We specially note that the 2.5-layer spectrum extends infinitely on the spectral plane by involving unbounded Froude
 6 numbers F_j .

7 5 The Non-Linear Long-Wave Model

8 The derivation procedure of non-linear long-wave 2.5-layer model should involve, in accordance with hypothesis (30), the slow
 9 horizontal variable $\xi = \sqrt{\sigma} x$, as it was demonstrated by Benney & Ko (1978) in the case of slight linear stratification. Scaling
 10 with the parameter σ gives the equation

$$11 \quad \sigma\psi_{\xi\xi} + \psi_{yy} + \lambda_1^2(\psi - y) = \frac{1}{2}\sigma(\sigma\psi_{\xi}^2 + \psi_y^2 - 1) \quad (32)$$

12 in the lower layer $-\pi < y < \eta(\xi)$, and

$$13 \quad \sigma\psi_{\xi\xi} + \psi_{yy} + \lambda_2^2 r^2(\psi - y) = \frac{1}{2}\sigma(\sigma\psi_{\xi}^2 + \psi_y^2 - r^2) \quad (33)$$

14 in the upper layer $\eta(\xi) < y < \pi/r$. Kinematic boundary conditions (11) can be rewritten now as follows:

$$15 \quad \psi(\xi, -\pi) = -\pi, \quad \psi(\xi, \eta(\xi)) = 0, \quad \psi(\xi, \pi/r) = \pi. \quad (34)$$

16 We find that stream function ψ is expanded in a power series with respect to σ as

$$17 \quad \psi = \psi^{(0)}(\xi, y) + \sigma\psi^{(1)}(\xi, y) + \dots \quad (35)$$

18 where the leading-order term $\psi^{(0)}$ defines the hydrostatic mode, and the coefficient $\psi^{(1)}$ provides the correction due to non-
 19 linear dispersion. All these coefficients $\psi^{(k)}$ can be uniquely determined from equations (32) and (33) (with fixed Long's
 20 numbers λ_1 and λ_2) under kinematic boundary condition (34). Thus, we obtain

$$21 \quad \psi^{(0)} = y - \eta \frac{\sin \alpha_1(y)}{\sin \alpha_1(\eta)} \quad (-\pi < y < \eta),$$

22 and

$$23 \quad \psi^{(0)} = ry - r\eta \frac{\sin \alpha_2(y)}{\sin \alpha_2(\eta)} \quad (\eta < y < \pi/r),$$

24 where

$$25 \quad \alpha_1(y) = \lambda_1(\pi + y), \quad \alpha_2(y) = \lambda_2(\pi - ry). \quad (36)$$

26 The final form of dispersive term $\psi^{(1)}$ is much more complicated, therefore this coefficient is given in Appendix A.



1 Now we substitute power expansion (35) for function ψ into the scaled version of integral relation (22) and truncate the
 2 terms with the powers higher than the first power of σ . By that, equation (22) reduces to the first-order ordinary differential
 3 equation for the wave elevation $\eta(x)$ and is written as

$$4 \left(\frac{d\eta}{dx} \right)^2 = \eta^2 \frac{D(\eta; F_1, F_2)}{Q(\eta; F_1, F_2)}. \quad (37)$$

5 Here function D is given by the formula

$$6 D(\eta; F_1, F_2) =$$

$$7 = \sqrt{\pi} F_1 \cot \alpha_1(\eta) + \sqrt{\pi} F_2 \cot \alpha_2(\eta) + \frac{1}{3} (1-r)\eta - 1$$

8 where α_1 and α_2 should be taken as

$$9 \alpha_1(\eta) = \frac{\pi + \eta}{\sqrt{\pi} F_1}, \quad \alpha_2(\eta) = \frac{\pi - r\eta}{\sqrt{\pi} F_2} \quad (38)$$

10 since we have at the leading order in σ the relations $\lambda_j = 1/\sqrt{\pi} F_j$ ($j = 1, 2$) obtained under condition (30). Denominator Q in
 11 (37) has a complicated form, therefore this function is given in Appendix C. Solitary-wave solutions of Eq. (37) are given in
 12 the implicit form by the formula

$$13 x = \pm \int_a^\eta \sqrt{\frac{Q(s; F_1, F_2)}{D(s; F_1, F_2)}} \frac{ds}{s} \quad (39)$$

14 where parameter a determines non-dimensional amplitude of the wave.

15 Small-amplitude waves can be modelled by simplified weakly nonlinear version of the Eq.(37) having the form

$$16 \left(\frac{d\eta}{dx} \right)^2 = \eta^2 \frac{D_0 + D_1 \eta + D_2 \eta^2}{Q(0; F_1, F_2)} \quad (40)$$

17 where the coefficients $D_0 = D(0; F_1, F_2)$ and $D_1 = D'_\eta(0; F_1, F_2)$ are

$$18 D_0 = \sqrt{\pi} F_1 \cot \frac{\sqrt{\pi}}{F_1} + \sqrt{\pi} F_2 \cot \frac{\sqrt{\pi}}{F_2} - 1,$$

$$19 D_1 = -\cot^2 \frac{\sqrt{\pi}}{F_1} + r \cot^2 \frac{\sqrt{\pi}}{F_2} + \frac{2}{3} (r - 1),$$

21 and the explicit form of coefficient D_2 is not important here. This model takes into account the balance of quadratic and cubic
 22 nonlinearities in the weakly-nonlinear KdV–mKdV – Gardner model (Kakutani and Yamasaki 1978; Gear and Grimshaw,
 23 1983; Helfrich and Melville 2006; Grimshaw et al 2002).

24 Solitary wave regimes are obtained depending on the multiplicity of the roots $a_j(F_1, F_2, r)$ ($j = 1, 2$) of the numerator on
 25 the right-hand side of (40). Profile of solitary wave is given by the formula

$$26 \eta(x) = a \frac{1 - \tanh^2 kx}{1 - \theta^2 \tanh^2 kx}, \quad k = \frac{a\sqrt{3/q_*}}{2\theta},$$



1 with $q_* = Q(0)$, $a = a_1$ and $\theta^2 = a_1/a_2 < 1$, and the bore (internal front) corresponds to the double root $a = a_1 = a_2$, it has
2 the following profile
3
$$\eta(x) = \frac{a}{2} \left(1 + \tanh kx \right), \quad k = \frac{a\sqrt{3/q_*}}{2}.$$

4 Parametric range of strongly nonlinear solitary waves described by Eq. (37) is formed by the domain in (F_1, F_2) -plane
5 where the radical function Q/D in (39) is ensured to be non-negative. It is easy to check that $Q(0; F_1, F_2) > 0$, hence function
6 $Q(s; F_1, F_2)$ is positive in the vicinity of point $s = 0$. Therefore, function D plays the determining role here. Depending on
7 F_1 and F_2 , this function can change the sign even by small s , where the leading-order coefficient D_0 from formula (40)
8 dominates. As a consequence, the map of solitary-wave regimes is formed by the Froude numbers (F_1, F_2) such that inequality
9 $D_0(F_1, F_2) > 0$ holds. Indeed, this inequality defines the range of non-linear waves, which are supercritical with respect to the
10 phase speed of linear harmonic wave-packets (see Fig. 2).

11 6 Waves in Marginally Stable Abyssal Flows

12 Large-amplitude internal waves are generated in abyssal flows due to the interaction of internal tides with irregular bottom
13 topography near underwater ridges (Morozov 1995, Morozov et al. 2010). These waves play a significant role in the energy
14 transformation and mass transport in the oceanic stratified flows while they intensify mixing of the abyssal waters. Note that
15 internal Froude numbers F_1 and F_2 characterize the magnitude of the velocity jump at the interface in upstream flow. The
16 shear $u_1 \neq u_2$ between the layers can initiate the development of the Kelvin — Helmholtz instability which provides non-
17 stationary formation of billow trains (Thorpe 1985; Drazin 2002). Constant two-layer flow is linearly stable under long-wave
18 perturbations if the inequality

$$19 \quad |u_1 - u_2| < \sqrt{\frac{g(\rho_1 - \rho_2)(\rho_1 h_2 + \rho_2 h_1)}{\rho_1 \rho_2}}$$

20 holds, and this flow is unstable in the opposite case. Exactly the same bound for a *variable* difference $|u_1 - u_2|$ and *variable*
21 layer thicknesses h_1, h_2 follows from the *non-linear* stability criteria predicted by the shallow water theory (Ovsyannikov et
22 al. 1985; Gavriluk, Makarenko, Sukhinin 2017). As a consequence, we have the stability domain

$$23 \quad |\sqrt{r}F_1 - F_2| < \sqrt{1+r}$$

24 shown for $r = 3$ in Fig. 2 (right panel) as unshaded domain in the fragment of quarter-plane (F_1, F_2) .

25 Figures 3 and 4 demonstrate fragments of quasi-steady shear flow recorded with a 350 m mooring station located at a depth
26 of 4720 m at the entrance of the Romanche Fracture Zone in the equatorial Atlantic (Van Haren et al. 2014). Trains of internal
27 waves modulated by tide propagate here along a sharp interface corresponding to the 0.85°C isotherm which separates the
28 lower layer of cold Antarctic Bottom Water (AABW) from the overlying warmer layer. By that, moored CTD/LADCP data
29 indicate permanently marginal stability of the flow having the Richardson number $0.25 < Ri < 1$. Tidal amplification of the
30 shear triggers the formation of small-scale overturns which create long trains of the Kelvin — Helmholtz billows. Bold curves

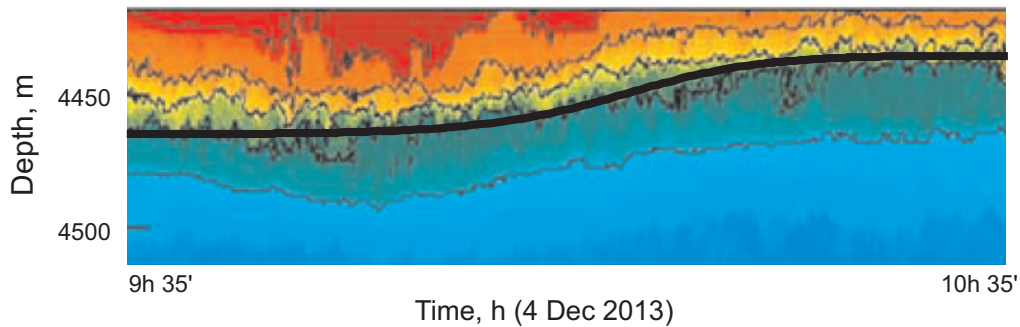


Figure 3. Internal front in abyssal stratified flow

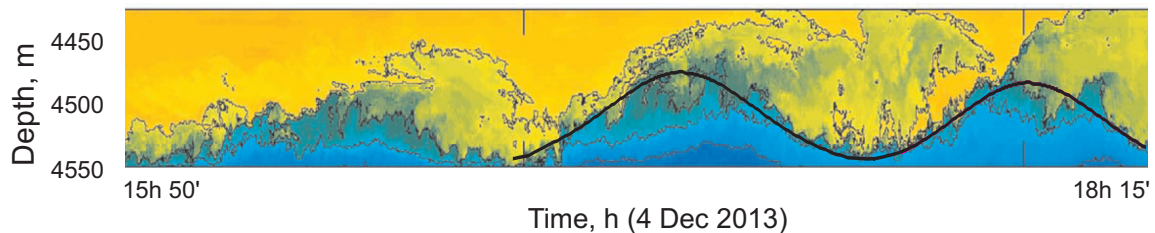


Figure 4. Interfacial solitary waves affected by the Kelvin — Helmholtz instability

1 in Fig. 3 and Fig. 4 show overlapped profiles of internal waves calculated by solution (39). Internal front shown in Fig. 3
2 corresponds to the Froude point A with coordinates $(F_1, F_2) = (1.719, 0.891)$ in Fig. 2 (right panel). Point A belongs to the
3 bore diagram which is tangential to the spectrum boundary at the point O. Figure 4 demonstrates a series of solitary waves
4 with intense overturns, which are distributed uniformly along with gently sloping wave top (Froude point B with coordinates
5 $(F_1, F_2) = (1.634, 0.959)$). It is interesting that similar overturning near the middle part of broad solitary wave was observed
6 in laboratory experiments (Grue et al. 2000).

7 7 Conclusions

8 In this paper we have considered the problem on permanent internal waves at the interface between exponentially stratified
9 fluid layers. An ordinary differential equation describing large amplitude solitary waves has been obtained using the long-wave
10 scaling procedure. Parametric range of solitary waves is characterized, including regimes of broad plateau-shape solitary waves
11 and internal fronts. It is demonstrated that these solitary wave regimes can be affected by the Kelvin — Helmholtz instability
12 generated due to the velocity shear at the interface.



1 *Acknowledgements.* This work was supported by the Russian Foundation for Basic Research (grants No 15-01-03942, 17-08-00085, 18-01-
 2 00648) and Interdisciplinary Program of SB RAS. Roman Tarakanov was supported by the Russian Science Foundation (grant No 16-17-
 3 10149).

4 Appendix A Modal functions of linearized problem

5 Eigenfunction $W(y)$ considered in the strip $-\pi < y < 0$ has the form

$$6 \quad W = a_1 e^{\frac{1}{2}\sigma_1 y} \begin{cases} e^{\varkappa_1 \pi} \frac{\sinh \varkappa_1 (\pi + y)}{\varkappa_1} & (\lambda_1^2 < k_1^2 + \frac{\pi^2}{4} \sigma_1^2) \\ \pi + y & (\lambda_1^2 = k_1^2 + \frac{\pi^2}{4} \sigma_1^2) \\ \frac{\sin \varkappa_1 (\pi + y)}{\varkappa_1} & (\lambda_1^2 > k_1^2 + \frac{\pi^2}{4} \sigma_1^2). \end{cases}$$

7 Similarly, eigenfunction $W(y)$ defined in upper layer which corresponds to the strip $0 < y < \pi/r$ has the form

$$8 \quad W = a_2 e^{\frac{1}{2}\sigma_2 y} \begin{cases} e^{\varkappa_2 \pi} \frac{\sinh \varkappa_2 (\pi - ry)}{\varkappa_2} & (\lambda_2^2 < k_2^2 + \frac{\pi^2}{4} \sigma_2^2) \\ \pi - ry & (\lambda_2^2 = k_2^2 + \frac{\pi^2}{4} \sigma_2^2) \\ \frac{\sin s_1 (\pi - ry)}{\varkappa_2} & (\lambda_2^2 > k_2^2 + \frac{\pi^2}{4} \sigma_2^2). \end{cases}$$

9 The dimensionless wave-numbers \varkappa_j ($j = 1, 2$) are introduced in formula (26), and the factors a_j are the amplitude parameters.

10 Appendix B Dispersive term of the long-wave expansion

11 The coefficient $\psi^{(1)}(\xi, y)$ which gives the correction due to the dispersion in power expansion (35) has the form

$$12 \quad \psi^{(1)} = \frac{\eta(\eta - y)}{2} \frac{\sin \alpha_1(y)}{\sin \alpha_1(\eta)} + \frac{\sin \alpha_1(y)}{2\lambda_1} \left(\frac{\eta}{\sin \alpha_1(\eta)} \right)_{\xi\xi} \times$$

$$13 \quad \times \left\{ (\pi + \eta) \cot \alpha_1(\eta) - (\pi + y) \cot \alpha_1(y) \right\} + \frac{\eta^2}{6} \times$$

$$14 \quad \times \left\{ \frac{\sin \lambda_1(y - \eta) - \sin \alpha_1(y)}{\sin^3 \alpha_1(\eta)} + \frac{1 + \sin^2 \alpha_1(y)}{\sin^2 \alpha_1(\eta)} - \frac{\sin \alpha_1(y)}{\sin \alpha_1(\eta)} \right\},$$

17 in the lower layer $-\pi < y < \eta$. Similarly, we have

$$18 \quad \psi^{(1)} = \frac{r^2 \eta(\eta - y)}{2} \frac{\sin \alpha_2(y)}{\sin \alpha_2(\eta)} + \frac{\sin \alpha_2(y)}{2\lambda_2} \left(\frac{\eta}{\sin \alpha_2(\eta)} \right)_{\xi\xi} \times$$

$$19 \quad \times \left\{ (y - \pi/r) \cot \alpha_2(y) - (\eta - \pi/r) \cot \alpha_2(\eta) \right\} + \frac{r^2 \eta^2}{6} \times$$

$$20 \quad \times \left\{ \frac{\sin \lambda_2 r(\eta - y) - \sin \alpha_2(y)}{\sin^3 \alpha_2(\eta)} + \frac{1 + \sin^2 \alpha_2(y)}{\sin^2 \alpha_2(\eta)} - \frac{\sin \alpha_2(y)}{\sin \alpha_2(\eta)} \right\}$$

23 in the upper layer $\eta < y < \pi/r$. Here the functions α_j are given by the formula (36).



1 Appendix C Denominator of the non-linear long-wave equation

2 Denominator Q in (37) has the form

$$\begin{aligned} &3 \quad 2Q(\eta; F_1, F_2) = \\ &4 \\ &5 \quad \left(\pi F_1^2 - 2\sqrt{\pi} F_1 \eta \cot \alpha_1(\eta) + \eta^2 \cot^2 \alpha_1(\eta) \right) \times \\ &6 \\ &7 \quad \times \left(\frac{\eta + \pi}{\sin^2 \alpha_1(\eta)} - \sqrt{\pi} F_1 \cot \alpha_1(\eta) \right) + \\ &8 \\ &9 \quad + \left(\frac{\pi F_2^2}{r^2} - 2\frac{\sqrt{\pi} F_2}{r} \eta \cot \alpha_2(\eta) + \eta^2 \cot^2 \alpha_2(\eta) \right) \times \\ &10 \\ &11 \quad \times \left(\frac{\pi - r\eta}{\sin^2 \alpha_2(\eta)} - \sqrt{\pi} F_2 \cot \alpha_2(\eta) \right) \end{aligned}$$

12 where functions $\alpha_j(\eta)$ are given by formula (38) which is an approximate version of (36).



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