Evaluation and Aggregation Properties of Thermal Infra-Red-Based Evapotranspiration Algorithms from 100 m to the km Scale over a Semi-Arid Irrigated Agricultural Area

Malik Bahir, Gilles Boulet, Albert Olioso, Vincent Rivalland, Belen Gallego-Elvira, Maria Mira, Julio-Cesar Rodriguez, Lionel Jarlan, Olivier Merlin

▶ **To cite this version:**

Malik Bahir, Gilles Boulet, Albert Olioso, Vincent Rivalland, Belen Gallego-Elvira, et al.. Evaluation and Aggregation Properties of Thermal Infra-Red-Based Evapotranspiration Algorithms from 100 m to the km Scale over a Semi-Arid Irrigated Agricultural Area. Remote Sensing, 2017, 9 (11), pp.1178. 10.3390/rs9111178 . hal-01651305

HAL Id: hal-01651305
https://hal.science/hal-01651305

Submitted on 28 Nov 2017

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

remote sensing

MDPI

Article

Evaluation and Aggregation Properties of Thermal Infra-Red-Based Evapotranspiration Algorithms from 100 m to the km Scale over a Semi-Arid Irrigated Agricultural Area

Malik Bahir [1,2], Gilles Boulet [1,*], Albert Olioso [2], Vincent Rivalland [1], Belen Gallego-Elvira [3], Maria Mira [2,4], Julio-Cesar Rodriguez [5], Lionel Jarlan [1] and Olivier Merlin [1]

1 CESBIO, Université de Toulouse, CNES/CNRS/IRD/UPS, 31401 Toulouse, France; malikbahir@gmail.com (M.B.); vincent.rivalland@cesbio.cnes.fr (V.R.); lionel.jarlan@ird.fr (L.J.); olivier.merlin@cesbio.cnes.fr (O.M.)

2 EMMAH, INRA, Université d'Avignon et des Pays de Vaucluse, 84000 Avignon, France; albert.olioso@inra.fr (A.O.); maria.mira@uab.cat (M.M.)

3 NERC Centre for Ecology & Hydrology, Wallingford OX10 8BB , Oxfordshire, UK; belgal@ceh.ac.uk

4 Grumets Research Group, Department of Geography, Universitat Autònoma de Barcelona (UAB), 08193 Bellaterra, Catalonia, Spain

5 Departamento de Agricultura y Ganadería, Universidad de Sonora, 83000 Hermosillo, Sonora, Mexico; jcrod@guayacan.uson.mx (J.C.R.)

* Correspondence: gilles.boulet@ird.fr; Tel.: +33-561-558543

Received: 23 August 2017; Accepted: 10 November 2017; Published: 17 November 2017

Abstract: Evapotranspiration (ET) estimates are particularly needed for monitoring the available water of arid lands. Remote sensing data offer the ideal spatial and temporal coverage needed by irrigation water management institutions to deal with increasing pressure on available water. Low spatial resolution (LR) products present strong advantages. They cover larger zones and are acquired more frequently than high spatial resolution (HR) products. Current sensors such as Moderate-Resolution Imaging Spectroradiometer (MODIS) offer a long record history. However, validation of ET products at LR remains a difficult task. In this context, the objective of this study is to evaluate scaling properties of ET fluxes obtained at high and low resolution by two commonly used Energy Balance models, the Surface Energy Balance System (SEBS) and the Two-Source Energy Balance model (TSEB). Both are forced by local meteorological observations and remote sensing data in Visible, Near Infra-Red and Thermal Infra-Red spectral domains. Remotely sensed data stem from Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) and MODIS sensors, respectively, resampled at 100 m and 1000 m resolutions. The study zone is a square area of 4 by 4 km^2 located in a semi-arid irrigated agricultural zone in the northwest of Mexico. Wheat is the dominant crop, followed by maize and vegetables. The HR ASTER dataset includes seven dates between the 30 December 2007 and 13 May 2008 and the LR MODIS products were retrieved for the same overpasses. ET retrievals from HR ASTER products provided reference ET maps at LR once linearly aggregated at the km scale. The quality of this retrieval was assessed using eddy covariance data at seven locations within the 4 by 4 km^2 square. To investigate the impact of input aggregation, we first compared to the reference dataset all fluxes obtained by running TSEB and SEBS models using ASTER reflectances and radiances previously aggregated at the km scale. Second, we compared to the same reference dataset all fluxes obtained with SEBS and TSEB models using MODIS data. LR fluxes obtained by both models driven by aggregated ASTER input data compared well with the reference simulations and illustrated the relatively good accuracy achieved using aggregated inputs (relative bias of about 3.5% for SEBS and decreased to less than 1% for TSEB). Results also showed that MODIS ET estimates compared well with the reference simulation (relative bias was down to about 2% for

SEBS and 3% for TSEB). Discrepancies were mainly related to fraction cover mapping for TSEB and to surface roughness length mapping for SEBS. This was consistent with the sensitivity analysis of those parameters previously published. To improve accuracy from LR estimates obtained using the 1 km surface temperature product provided by MODIS, we tested three statistical and one deterministic aggregation rules for the most sensible input parameter, the surface roughness length. The harmonic and geometric averages appeared to be the most accurate.

Keywords: energy balance; Thermal Infra-Red; scaling

1. Introduction

Monitoring water uptakes from groundwater reservoirs is important for water resources managers as well as water allocation and water rights regulation authorities. This is especially true in semi-arid lands where many aquifers face an average depletion of 0.5 to 1.5 m/year due to unsustainable pumping for irrigation [1]. EvapoTranspiration (ET) is the largest water loss of most agricultural areas. In many cases, it represents the only loss of water on the long term if one assumes that irrigation excess mostly contributes to deep drainage and that the corresponding water flux eventually reaches the groundwater table and participates to recharge. Groundwater uptake can therefore be scaled to total cumulative ET on an annual basis at least. Up to now, regional scale evaluation of ET relies on distributed information obtained either through hydrological modeling, Remote Sensing (RS) or a combination of both [2].

In order to estimate ET from the sole RS and meteorological data, methods based on Thermal Infra-Red (TIR) information are therefore increasingly used in the context of irrigation monitoring [3,4]. Due to its tight coupling with surface energy balance and water stress, surface temperature information provides a way to estimate ET independently from any knowledge on the various components of the water balance.

Two kinds of methods are implemented to retrieve water stress and actual ET rates from a combination of TIR as well as Visible (VIS) and Near Infra-Red (NIR) RS data: those generally referred to as "contextual" methods, and those known as "residual" or "single pixel" methods [5].

"Contextual" methods estimate the water status of a given pixel by scaling its value between non-evaporating and fully-evaporating conditions according to maximum and minimum surface temperature values observed on the image. It assumes that cold and hot pixels correspond to wet and dry conditions, respectively. Since those extremes depend as well on the surface conditions, particularly the amount of vegetation, extremes are derived for each class of either the Normalized Differential Vegetation Index (NDVI) [6], albedo [7] or both [8]. These methods have operational applications for irrigated lands with large fields and high-resolution (HR) TIR data (hectometric resolution data from sensors on board LANDSAT platforms or from the ASTER sensor). Their applicability drops in the context of dryland agriculture and low-resolution (LR) data (e.g., kilometric resolution data from MODIS sensors) [9–11]. Those methods are therefore more adapted to HR TIR data, which are available at a weekly interval only, and usually at lower frequencies due to cloud conditions. In order to use the routinely available LR data acquired daily over the globe with a revisit consistent with typical timescales of the hydrological cycle (droughts of few days to several weeks), data sharpening techniques [12–17] can be used to infer surface temperature at a higher spatial resolution, but they all rely on an a priori link between vegetation cover and surface water status, whose validity can be challenged.

"Residual" or "single pixel" methods, on the other hand, derive an evaporation rate in actual conditions for each pixel independently from the others. Latent heat is computed as the residual term of the surface energy balance equation. The various terms of that energy balance (net radiation, soil heat flux, sensible heat flux) are derived for a single source representing the soil vegetation composite (single source models, e.g., SEBS [18,19]) or for two sources with two separate energy

budgets for the soil and the vegetation (e.g., TSEB [20,21]). The latter is often considered as more realistic and provides acceptable estimates of water stress levels [22]. Those methods are more easily applicable to derive operational products because they can be fed with a combination of LR sensors and reanalysis data. There is a growing number of LR ET products based on TIR data ([3,23]) acquired by LR sun-synchronous satellite platforms such as MODIS or geostationary satellites. Contrarily to "contextual" methods, which are rather insensitive to systematic errors on Land Surface Temperature (LST) estimates, the "residual" methods add the uncertainties related to each flux (net radiation, soil heat flux, sensible heat flux) computed from an error prone estimate of the LST. Because they use aerodynamic equations for the computation of the sensible heat flux, they are also sensitive to the quality of the available information on surface roughness and meteorological data such as air temperature and wind speed on a pixel basis. Their validity remains therefore questionable.

The Simple Energy Balance System (SEBS) is a single-source model that uses the Monin–Obukhov Similarity Theory (MOST) to estimate the sensible heat flux (H) as the transfer of heat from the surface to the height of atmospheric parameters measurements (air temperature and relative humidity, wind speed, incoming shortwave radiation). Its specificity lies in the relationship between the roughness length for heat transfer and the roughness length for a momentum transfer, which is derived from the crop height parameters through the kB^{-1} parameter.

The Two-Source Energy Balance (TSEB) is a two-source model, which achieves two energy budgets separately, one for the soil and one for the vegetation. It partitions the available energy between the two compartments according to the vegetation fraction cover. The two-source approach has the advantage of not requiring the inclusion of the kB^{-1} parameter used in the calculation of turbulent fluxes to relate the aerodynamic and surface temperatures [18].

Performance assessment of HR retrievals of latent heat fluxes is usually achieved through intercomparison with flux measurements obtained from local measurements (lysimeters, Bowen Ratio ...), nowaday mainly from eddy covariance (EC) systems [22]. For LR data, the amount of EC stations to be installed, ideally for each field within a pixel, requires a multiplication of EC towers, which is largely intractable. Therefore, a limited amount of validation exercises of ET products at LR have been carried out. There are mostly three ways of validating LR ET products at scales representative of LR products, i.e., the km scale and above: (1) using one flux tower for an homogeneous landscape (generally natural vegetation or forests), (2) by scintillometry [24–27], or (3) through water balance modelling at the catchment scale [28,29]. Regarding the first category, many studies have evaluated ET products on FLUXNET biomes with large extension, almost never in intensive agricultural areas [30–32]. Concerning the second category, scintillometer applications have been limited by the number of dates or models involved, therefore lacking any generalization of their outcomes. The third category is only applicable for large catchments with well-known water budget components and limited surface water–groundwater interactions.

LR ET products need to be better evaluated. In the absence of observed latent heat flux data at the relevant scale, one can rely on comparing aggregated HR ET retrievals at the Landsat or ASTER overpass times (airborne sensors have been used also) and those retrieved at LR. These methods have been often used for the evaluation of other biophysical products such as Leaf Area Index (LAI), fraction of absorbed photosynthetically active radiation or albedo [33,34]. It can also be useful to combine HR and LR ET products in order to maximize revisit frequency while taking advantage of the field-scale outputs from the HR products. Multiscale evaluation has rarely been carried out. The method proposed here to evaluate LR ET retrievals consists in validating HR retrievals using EC data for the main land use classes, and comparing LR and HR retrievals aggregated to the LR pixel.

Parameter estimation at LR is also an issue. In particular, deriving all parameters necessary to run ET retrieval algorithms at the km scale proves to be difficult [35]. This is particularly the case for the roughness length parameterization. The validity at LR of the methods initially used to derive surface parameters from HR RS data is questionable. In particular, estimating an effective vegetation height over various fields within the kilometer square is difficult and cannot be simply derived from

an average reflectance value obtained at low resolution. On the other hand, HR data in the VIS and NIR wavelengths are routinely available and lead to more robust methods to estimate land use and crop cover types as well as their corresponding biophysical parameters such as crop height. There is therefore a need to apply scaling procedures to some surface variables such as crop height or minimum surface resistance, which cannot be easily derived directly from existing LR RS data.

Moreover, since HR and LR products are not available at the same frequency, producing continuous ET though temporal interpolation will likely result in combining products at both scales. To do so, one must ensure the consistency in scaling products from HR to LR. Those products should at least be unbiased, i.e., aggregated HR fluxes to LR should match the LR ET products generated from LR data on the same date. Scaling properties are frequently studied by using LR RS data generated from HR data to get rid of discrepancies in the input variables from different sensors on different platforms. To analyse scaling properties and the subsequent biases, radiances are linearly aggregated to generate surface temperatures and reflectances at LR. McCabe and Wood [36] for one date only and, more recently Ershadi, McCabe et al. [37], for several dates within a season, have used the single-source model SEBS to assess those biases, and found scale discrepancies on instantaneous LE retrieval of up to 15%. They attributed this difference to a change in the roughness length parameters of the land surface due to the aggregation.

Similar work, within a theoretical frame (aggregation between contrasted patches rather than real RS data), were carried out by Kustas and Norman [38] using the dual source model TSEB. They concluded that, for those very contrasted conditions, there was a bias up to 50 W/m^2 on the instantaneous retrieval between the aggregated latent heat fluxes and the latent heat flux computed from surface average parameters at the km scale. On the other hand, the same authors have performed similar work on real data for a mixture of riparian trees and stressed shrubs [39] and a realistic contribution (fraction cover f_c) of both patches to the LR pixel (with a contribution of the most extreme conditions of less than 20%). They concluded that divergence was significantly less than 50 W/m^2, which is the typical precision of most Energy Balance (EB) models [40] for retrieving instantaneous latent heat flux. They also found that simple averaging of displacement height and an averaging of roughness length based on a lognormal to the power -2 reduced the bias down to less than 10 W/m^2. Both studies pointed out the need to use proper efficient scaling relationships in parameter estimation at LR. There has been abundant literature on the subject from back in the 1990s, mostly based on theoretical works deriving deterministic [41] or statistical [42] relationships between local and regional parameters for land–atmosphere exchange modelling. For roughness length, Wassenaar et al. [43] conclude that a geometric averaging of the roughness length is performing best; this is also the scaling proposed by Taylor [44].

Previous studies with SEBS [45,46] agree on the important role of roughness parameters in the estimation of ET. They provide at least three main findings: (1) that, for SEBS, the main uncertainty is linked to roughness estimation [47]; (2) that the original formulation of the roughness length for heat transfer tends to overestimate ET with the original use of the soil component of the kB^{-1} parameter [48]; and (3) that the limited accuracy of height estimate (and subsequently the roughness length for momentum transfer) from RS data is rarely taken into account, with few empirical relationships that tend to diverge for high NDVI values [49]. The latter becomes more significant when LR RS data are used to derive surface roughness lengths. However, none of those aggregation rules (apart from [43]) have been evaluated with actual data acquired at LR on very heterogeneous land surfaces.

TSEB appears to provide good ET estimates and to be less sensitive to roughness length parameters [22]. One expects flux conservation between HR and LR. The relative contribution of the soil and the vegetation temperatures through the vegetation fraction cover parameter can be the issue at LR and the separated soil and vegetation fluxes conservation would be affected [50].

Up to now, most studies on scaling properties have focused on one of the most common models only (SEBS, TSEB or SEBAL). The widely used SEBS and TSEB "single pixel" models have not been compared on the same landscape [22,51]. Furthermore, intercomparison has been carried out

usually for a limited amount of HR images [46,52,53]. In addition, few studies were dedicated to the comparison of separate satellite sensors.

In this paper, we address two objectives:

- To investigate ET flux scaling properties from HR to LR using data from the same sensor (i.e., ASTER), as well as data stemming from different sensors onboard the same platform (i.e., ASTER, MODIS).
- To develop and evaluate new and existing scaling relationships based on easily obtainable RS quantities to relate local HR and LR roughness lengths.

We use two of the most common 'single-pixel' models, TSEB and SEBS, to estimate ET fluxes at the satellite overpass time and aggregate the values to a daily time scale using the algorithm presented in Delogu et al. [54]. The results are analyzed for seven dates throughout an entire growing season in a semi-arid irrigated area in northwest Mexico, for which both models have been evaluated using EC measurements in [22].

2. Data

2.1. Study Area

The study area is located in the Yaqui region (27.4°N, 19.9°W), Sonora, in the northwest Mexico. The region covers 225,000 ha between the Cortez Sea (northeast) and the Sierra Madre mountains (southwest), and is the largest agricultural area in the country (Figure 1). More than half of the fields are cultivated with winter wheat, the dominant crop, and the rest with a mix of broccoli, beans, chili, potatoes, peas, safflower, orange trees and maize. When considering the type of crops and the spatial patterns, the study area is representative of the region and supports the idea that the results presented here can be extended to the whole agricultural region. The climate is considered as semi-arid with an annual potential ET of 2233 mm (mean value between years 1971 and 2000) and a low annual precipitation rate of 290 mm essentially spread between June and September. Only 42.8 mm of rain falls between January and June. About 90% of the water use concern irrigation, the water being withdrawn from the Alvaro Obregon Reservoir with a capacity of 3 km^3. In this hydrological context, an accurate method to estimate water losses by ET is essential for managing the water resources at the regional scale. Between December 2007 and May 2008 an international field measurement campaign was set up on a 4 $\times$ 4 km^2 zone located in the South of Obregon city and an important HR and LR RS dataset covering the same period and area was collected.

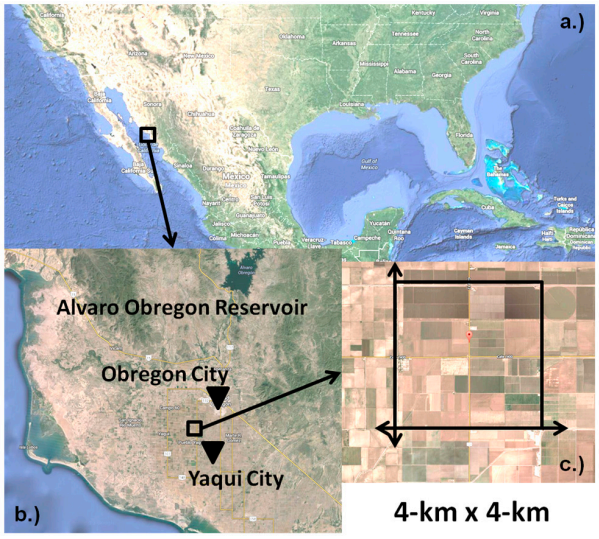

Figure 1. Study zone location in North America (**a**); the Sonora state (**b**); and the irrigated perimeter (**c**).

2.2. In Situ Measurements

The meteorological dataset was acquired at a height of 10 m from an automatic weather station located at the center of the 4 by 4 km^2 square (Figure 1c). Half hourly global solar and atmospheric incoming radiations, air temperature and relative humidity as well as wind speed were recorded. Eddy covariance data was acquired at 2 to 3 m at seven sites located on different crop plots to represent the diversity of crop types and phenological stages as well as contrasted simultaneous soil water conditions related to irrigation patterns (Figure 2). At each site, net radiation, soil heat flux (at 0.05 m depth), surface temperature and soil moisture (at 0.05 and 0.3 m depth) were measured each 10 s before being averaged over 30 min periods. The latent and the sensible heat fluxes were acquired at a frequency of 10 Hz, processed using the FLUXNET guidelines [55] and converted to 30 min flux average. The devices used for all the automatic measurements at the different EC and meteorological stations as well as the fluxes quality analysis were described in Chirouze et al. [22].

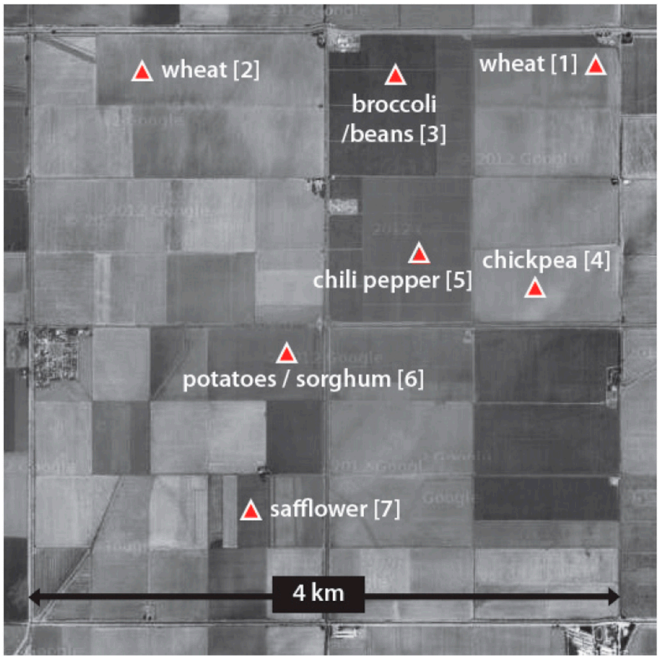

Figure 2. Satellite view of the study zone with respective positions of the eddy covariance flux towers and their associated crop type (Imagerie© 2012 Cnes/Spot Image. DigitalGlobe. Données cartographiques© 2012 Google (City, US State abbrev., Country ; Inst. Nat. Estat. y Geog, INEGI).

2.3. Remote Sensing Data

The ASTER (Advanced Spaceborne Thermal Emission and Reflection Radiometer) sensor on board the Terra satellite provided a dataset made of four products and eight dates (30 December 2007, 23 February 2008, 10 March 2008, 11 April 2008, 27 April 2008, 6 May 2008, 13 May 2008 and 29 May 2008) covering the 2007–2008 cereals growing period. The first three products consisted of surface reflectances in different spectral ranges and the last one was a top of the canopy radiative surface temperature product at 90 m resolution (Table 1). All products are atmospherically corrected and include LST retrieval includes emissivity correction. The local overpass time was between 10:30 a.m. and 11:00 a.m. and the swath of each product was always the same (60 by 60 km^2).

The MODIS (MODerate resolution Imaging Spectroradiometer) sensor also onboard the Terra satellite offered a daily global coverage at LR. The MODIS dataset was a blend of four products at 1 km resolution directly usable in both energy balance models (Table 1). They were extracted for the same dates as the ASTER dataset. ASTER and MODIS overpass times are approximately coincident since they both are onboard the same satellite platform.

Table 1. ASTER and MODIS products used in this study.

Sensor	Products	Frequency	Resolution [m]	Bands/Product/Subdatasets
ASTER	AST07XT—VNIR	16 days	15	Band1 Band2 Band3N
	AST07XT—SWIR		30	Band4 Band5 Band6 Band7 Band8 Band9
	AST05		90	Surface emissivity [-]
	AST08		90	Surface temperature [°K]
MODIS	MOD11A1	Daily	~1000	Surface temperature [°K] Emissivity Band31 Emissivity Band32 QC_Day (Quality control)
	MOD13A2	16 days	~1000	NDVI VI Quality QA (Quality control)
	MOD15A2	8 days	~1000	Lai_1 km FparLai_QC (Quality control)
	MCD43B3	16 days	~1000	Black Sky Albedo SW White Sky Albedo SW

2.4. Remote Sensing Data Preprocessing

The quality information contained in each product was taken into account and led to the removal of the last acquisition date reducing the ASTER dataset to seven dates. Concerning MODIS, all synthesis products (NDVI, albedo) were linearly interpolated between the available dates.

In order to provide easier comparison, ASTER images and MODIS images were resampled at 100 m and 1 km resolution, respectively [16]. All ASTER products were resampled using a bicubic interpolation and reprojected to the UTM12 system on a ~100 m resolution (100.49 m × 101.46 m) grid that fitted the study area. MODIS products were resampled at 1 km resolution on a grid based on the largest zone covered by the ASTER images using a bicubic method and reprojected in the UTM12 projection system. This projected the data on a squared regular grid instead of the irregular parallelepipoidal pixel shape of the original sinusoidal projection of the MODIS products. The grid defined by the sinusoidal projection of the MODIS products is centered on the intersection between the Equator and the Greenwich meridian where the sinusoidal grid produces square-like pixels, whereas our study area is located far from both latitudinal and longitudinal references. This resulted in a native resolution of about 900 by 1800 m^2. Therefore, in what follows, the aggregation performance for MODIS will be assessed at a more representative 2000 m spatial resolution instead of the usual 1 km resolution, in agreement with previous work on surface temperature disaggregation using the same dataset and on the same area [16].

3. Energy Balance Models Parameterization and ET Calculation

3.1. Available Energy

In both TSEB and SEBS, the available energy is considered as the difference between the total net radiation (R_n) and the ground heat flux (G). The net radiation is estimated with the same equation [22]:

$$R_n = (1 - \alpha)R_{sw} + \varepsilon R_{lw} - \varepsilon\sigma {T_{surf}}^4 \quad (1)$$

where R_{sw} and R_{lw} are, respectively, the shortwave and longwave surface incoming radiation, α the albedo, ε the emissivity and T_{surf} the radiative temperature of the surface. R_{sw} is taken from the meteorological station and $R_{lw} = 1.24(e_a/T_a)\sigma\ {T_a}^4$, where e_a and T_a are the actual vapour pressure and the temperature of the air measured by the meteorological station. The ground heat flux (G) is

estimated according to the SEBAL model formulation [4], which proved to be the more accurate in this context [22]:

$$G = \left[T_{surf}\ (0.0038 + 0.0074\ \alpha)\ \left(1 - 0.98\ NDVI^4\right)\right]\ R_n. \quad (2)$$

3.2. SEBS Model

The SEBS model [19] derives the latent heat (LE) flux as the residual term of the energy balance equation, $LE = R_n - G - H$, where H is the sensible heat flux.

SEBS computes turbulent exchanges between the source at the aerodynamic level and the reference level z, where atmospheric forcing data is available, usually just a few meters above the crop height. Heat transfers are based on the Monin-Obukhov Similarity Theory (MOST) in the Atmospheric Boundary Layer (ABL). Bulk ABL similarity functions proposed by Brutsaert 1999 [56] describe the wind and temperature profiles in the turbulent environment (Equations (3) and (4) (see [56,57]):

$$H = \rho\ c_p \frac{T_{aero} - T_a}{r_a} = \rho\ c_p \frac{T_{aero} - T_a}{k u_*}\left[ln\left(\frac{z - d_0}{z_{0h}}\right) - \Psi_h\left(\frac{z - d_0}{L}\right) + \Psi_h\left(\frac{z_{0h}}{L}\right)\right], \quad (3)$$

$$u = \frac{u_*}{k}\left[ln\left(\frac{z - d_0}{z_{0m}}\right) - \Psi_m\left(\frac{z - d_0}{L}\right) + \Psi_m\left(\frac{z_{0m}}{L}\right)\right], \quad (4)$$

where r_a is the aerodynamic resistance to heat transfer at the surface–atmosphere interface, u is the wind velocity at level z, $k = 0.4$ is the von Karman's constant, d_0 is the displacement height ($d_0 \approx h_c \times 2/3$, h_c being the crop height), z_{0h} is the roughness length for heat transfer and z_{0m} the roughness height for momentum transfer ($z_{0m} \approx h_c \times 0.123$). T_a and T_{aero} are the temperature of the air at reference and aerodynamic levels, respectively, ρ is the density of the air and c_p is the specific heat capacity of air. Ψ_m and Ψ_h are the stability correction functions for momentum and sensible heat transfer and L is the Monin–Obukhov length.

The main originality of SEBS lies in the derivation of an equation that relates the roughness length for heat transfer to the roughness length for momentum transfer according to the soil and vegetation individual drag coefficients:

$$z_{0h} = \frac{z_{0m}}{exp(kB^{-1})}, \quad (5)$$

$$kB^{-1} = A_1\ f_c^2 + A_2\ f_c\ f_s + kB_s^{-1}\ f_s^2, \quad (6)$$

where f_c is the vegetation fraction cover (in a pixel), and f_s the bare soil fraction cover with $f_c + f_s = 1$. A_1 describes the vegetation aerodynamic properties, ${kB_s}^{-1}$ the bare soil properties, and A_2 represents the interactions between vegetation and bare soil. All these terms are estimated as in Su et al. [19].

The second originality of SEBS is that the model provides bounding relationships to constrain the latent heat flux between two hypothetical extreme surface wetness conditions (dry and wet, which correspond to non-evaporative and potential conditions respectively):

$$H_{dry} = R_n - G;\ LE_{dry} = 0, \quad (7)$$

$$H_{wet} = R_n - G - LE_{wet} = \frac{\left[R_n - G - \frac{\rho\ c_p\ (e_s - e_a)}{r_{ew}\ \Upsilon}\right]}{\left(1 + \frac{\Delta}{\Upsilon}\right)}, \quad (8)$$

where LE_{dry} and LE_{wet} are the latent heat fluxes in dry and wet conditions, e_s the saturation vapor pressure temperature T_a, γ the psychrometric constant and Δ the slope of the saturation vapor pressure at temperature T_a.

The relative evaporation (Λ_r) is then computed to estimate LE as $LE = \Lambda_r\ LE_{wet}$:

$$\Lambda_r = 1 - \frac{H - H_{wet}}{H_{dry} - H_{wet}}, \quad (9)$$

which corresponds to the residual calculation of LE bounded by the dry and wet conditions' values. This potentially induces thresholds, which can affect the nonlinearity of the scaling properties between HR and LR (a threshold locally encountered at HR is possibly not present for average conditions at BR).

3.3. TSEB Model

TSEB [20] solves two different energy budgets for the soil and the vegetation components. The net radiation flux, estimated as described in Equation (1), is partitioned according to the vegetation and bare soil cover fractions following:

$$R_{n,s} = H_s + LE_s + G = R_n\, f_s, \tag{10}$$

$$R_{n,c} = H_c + LE_c = R_n\, f_c, \tag{11}$$

$$H_s = \rho\, c_p\, \frac{T_s - T_a}{r_a + r_s} \text{ and } H_c = \rho\, c_p\, \frac{T_c - T_a}{r_a}. \tag{12}$$

H_c and LE_c are the sensible and latent heat fluxes at the vegetation/atmosphere interface and H_s and LE_s the same heat fluxes at the bare soil/atmosphere interface. T_c and T_s are, respectively, the vegetation and bare soil temperature, r_s [21] is an additional resistance to describe the resistance to heat transfer in the ABL at the bare soil/atmosphere interface, and r_a the atmospheric resistance to heat transfer at the surface–atmosphere interface is expressed according to MOST.

TSEB computes the soil and the vegetation temperatures from their respective energy balance equation as well as the link between the total radiative surface temperature and the component temperatures:

$$T_{surf} = \left[f_c\, T_c^4 + f_s\, T_s^4\right]^{1/4}. \tag{13}$$

TSEB first assumes that the vegetation is unstressed and transpires at a potential rate defined by the Priestley–Taylor formulation:

$$LE_c = 1.3\, f_g\, \frac{\Delta}{\Delta + \Upsilon}\, R_{n,c}, \tag{14}$$

where f_g is the relative fraction of the vegetation that is green. Since LAI is estimated through the NDVI, which is an indicator of the green vegetation development, and f_g is here fixed to 1.

The canopy energy budget (Equation (11)) provides a first estimate of T_c by introducing LE_c in potential conditions (Equation (14)) in Equation (11). When T_c is known, T_s can be computed from Equation (13) and then H_s from its formulation detailed in Equation (12). Finally, LE_s is calculated as the residual term of the soil energy balance (Equation (10)). If the resulting LE_s is positive, the solution is reached and the hypothesis of an unstressed vegetation is considered as valid even if neither the soil or the vegetation evaporate at a potential rate. If the resulting LE_s is negative, the unstressed vegetation assumption is challenged, so LE_c value is decreased until LE_s is equal to 0. Then, new H_s and T_s values are computed from Equations (10) and (12) and new T_c, H_c and LE_c values from Equation (13), Equation (12) and Equation (11), respectively. If LE_c is positive, a new solution is reached. In addition, if LE_c is negative, the vegetation part is also considered as fully stressed and dry and LE_c is set to 0. H_c can then be computed from Equation (11) and a new T_c value from Equation (12). T_s can then be estimated according to Equation (13) and new H_s and G values from Equation (12) and Equation (10), respectively.

3.4. Daily ET Fluxes Generation

The latent heat fluxes computed by SEBS and TSEB are obtained at the time of satellite overpass. In order to compute daily evapotranspiration ET_d, we used an extrapolation algorithm based on an empirical diurnal shape of the evaporative fraction during a day in clear sky conditions (see [54]).

If one obtains an instantaneous value of the ET flux and owns a daytime dataset that allows computing a cumulative available energy, the corresponding daily ET can be estimated as:

$$ET_d = EF \times AE_d = \frac{ET_i}{AE_i} \times AE_d, \quad (15)$$

where ET_d is the daily cumulative ET value, ET_i the instantaneous evapotranspiration flux, which corresponds to the latent heat flux $ET_i = \frac{LE}{\lambda}$ (with λ being the latent heat of vaporisation), $EF = \frac{ET_i}{AE_i}$ is the evaporative fraction (considered as constant during daytime) defined as the ratio between the instantaneous ET value and the instantaneous available energy value, and AE_d is the daily cumulative available energy computed as the daily cumulative net Radiation flux (R_{nd}), considering that the cumulative soil heat flux at the daily scale can be neglected. R_{nd} is estimated assuming that the R_{nd}/R_{ni} ratio has a sinusoidal evolution along the year as stated in Gomez et al. [58]:

$$AE_d = R_{nd} = \left[a_1 + a_2 sin\left(\frac{2\pi(JD + a_3)}{365}\right)\right] R_{ni}, \quad (16)$$

where R_{ni} is the instantaneous net radiation flux, JD is the day of the year and a_1, a_2 and a_3 are obtained for each hour of the day using mean local measurement to calibrate the relationship (not shown).

4. Methods

4.1. General Design of the Experiments

In order to evaluate the impact of spatial resolution on flux estimates by the two models, SEBS and TSEB, we designed several numerical experiments based on the MODIS and the ASTER data. Figure 3 describes the flowchart of the four flux processing chains with each Energy Balance (EB) model (aggregated ASTER, LR ASTER, MODIS). The sole difference between those chains is the type and scale of the RS input data. The way one obtains those inputs is described in Sections 4.2–4.4 at and eventually the way one aggregates those inputs in the case of SEBS (Section 4.5).

- In a preliminary step, HR maps of ET were computed with both models from ASTER products (dataset named 'HR-ASTER' including spectral surface reflectances, spectral surface emissivities, radiative surface temperature and surface fluxes). ET maps were evaluated against the eddy correlation measurements performed in seven crop fields. These maps were aggregated at the kilometric resolution to be used in the following steps as a reference dataset to evaluate ET maps obtained at low resolution. This aggregation was done considering that surface fluxes can be averaged using a simple arithmetic mean. The reference dataset at low resolution was named 'agg-ASTER'.
- In the second step, LR maps of ET were produced from the high-resolution ASTER products aggregated at the kilometric resolution (equivalent to MODIS resolution). LR RS products were used for all inputs of both models. This dataset was named 'LR-ASTER'. ET maps at LR were evaluated against 'agg-ASTER' ET maps. Since both input datasets at LR and HR came from the same sensor, the biases between the two estimations of ET were only related to how the nonlinear relationships in the model translates the variability of inputs at HR into an average LR outputs that can be significantly different than the one generated using LR inputs.
- In a third step, LR maps of ET were produced from the MODIS products (surface temperature, emissivities, albedo, NDVI) at 2 km resolution, as it would be done in a standard application of SEBS and TSEB using MODIS data. These dataset (named 'MODIS') was evaluated against 'agg-ASTER' ET maps. In this case, differences between ET maps were related to a combination of the differences in products, input parameter derivations and heterogeneity/nonlinearity issues.
- In a fourth step, we analysed the possibility to derive SEBS and TSEB input parameters at low resolution by aggregating parameters estimated from high-resolution data. This scenario

considered the possible use of high-resolution images in the solar domain from Earth observation satellites for deriving model inputs at the kilometric resolution. Decametric information are now increasingly available, in particular thanks to Sentinel2 satellites. We expected that this scenario would provide a deeper analysis of the potential source of biases in deriving ET and to develop more adequate aggregation rules for the relevant inputs.

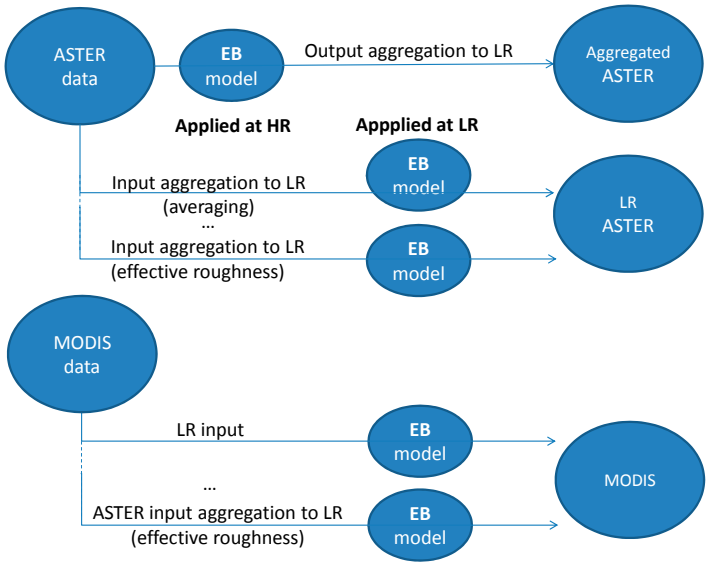

Figure 3. Model application rationale (energy balance model is either SEBS or TSEB).

Several aggregation rules were tested by comparing the three LR ET maps generated by each model ('MODIS', 'LR-ASTER' with input parameters derived from LR-inputs and 'LR-ASTER' with effective values of the input parameters, derived from HR data) with the reference 'agg-ASTER' ET map for the same model.

The aggregation rules tested with ASTER input data were also evaluated with the MODIS inputs' dataset. It means that aggregated ("effective") surface roughness lengths values derived from ASTER data were used with the other MODIS inputs to produce ET maps. The results of this comparison will be useful to many MODIS data users, which, in general, benefit from a HR VIS/NIR dataset extracted from Landsat, for example. It can lead them to derive effective LR surface roughness lengths from HR data in order to improve the parameterization of the EB models at LR and then the ET output fluxes.

4.2. Estimating Surface Parameters at HR from ASTER Products

Reflectances for each band were combined to estimate the land surface properties (Figure 3) required by both models. Albedo (α) was computed according to Liang [59]:

$$\alpha^{ASTER} = 0.484\,\rho_1 + 0.335\,\rho_2 - 0.324\,\rho_5 + 0.551\,\rho_6 + 0.305\,\rho_8 - 0.367\,\rho_9 - 0.0015, \quad (17)$$

where ρ_i referred to the spectral reflectance from band i.

Broadband emissivity (ε) was computed from spectral emissivities ε_i following Ogawa et al. [60]):

$$\varepsilon^{ASTER} = 0.121\,\varepsilon_{11} + 0.194\,\varepsilon_{12} + 0.323\,\varepsilon_{14} + 0.242. \quad (18)$$

NDVI was calculated from the NIR (ρ_{3N}) and red-VIS (ρ_2) bands as:

$$NDVI^{ASTER} = \frac{\rho_{3N} - \rho_2}{\rho_{3N} + \rho_2}. \quad (19)$$

LAI and fraction cover (f_c) were estimated from *NDVI* values following [61]:

$$LAI^{ASTER} = -\frac{1}{1.14} ln \frac{NDVI_\infty - NDVI}{NDVI_{\infty x} - NDVI_0}, \quad (20)$$

with $NDVI_\infty$ = 0.94 and $NDVI_0$ = 0.14, and:

$$fc^{ASTER} = \frac{NDVI - NDVI_0}{NDVI_\infty - NDVI_0}, \quad (21)$$

To ensure that all distributed parameters were derived from the sole remotely sensed information, crop heights (h_c) were also estimated from NDVI values. Knowing the repartition of the different land use classes on the 4 × 4 km^2 square, it was clear that the majority of the area was covered by cereals. Hence, for each pixel (i), crop height was estimated as a linear regression between its temporal maximum and minimum NDVI values (respectively, $NDVI_{max}(i)$ and $NDVI_{min}(i)$) and realistic maximum and minimum cereal crop heights' values. The period covered by the ASTER dataset contained dates from the whole winter cereal growing period from seeding to harvesting. Therefore, the minimum height was that of a ploughed bare soil of 0.05 m equivalent height and a realistic maximum was set to 1.3 m, which led to:

$$h_c(i) = \frac{1.3 - 0.05}{NDVI_{max}(i) - NDVI_{min}(i)} \times NDVI(i). \quad (22)$$

All ASTER data obtained from the original ASTER products constituted the hereafter called 'HR-ASTER' input dataset (~100 m resolution). Vegetation height translates directly into roughness length z_{om} using the simple rule of the thumb: $z_{om}(i) = 0.123 \times h_c(i)$.

4.3. Estimating Surface Parameters at LR from ASTER Product

To compute the fluxes at LR using the Energy Balance (EB) models and low resolution ASTER data, ASTER products were linearly aggregated to ~1000 m resolution, so that they approached the original MODIS product resolution (Figure 3). Special care was taken for the aggregation of surface temperature: radiative fluxes were linearly aggregated ($\varepsilon\sigma {T_{su}}^4$, with σ = 5.67 × 10^{-8} Wm^{-2}K^{-1} the Stefan–Boltzmann constant) before estimating the surface temperatures at LR by inverting the Stefan–Boltzman equation.

In order to account for geolocation errors between the two sensors, ASTER NDVI and surface temperature data at low resolution were compared to the equivalent MODIS products by seeking the maximum correlation considering different spatial shifts from −1000 m to +1000 m. A maximum of correlation was reached for a unique spatial shift, on each direction, for each date and each product. This spatial shift was applied to the product when the maximum of correlation was obtained with coefficient exceeding 0.8. For dates with lower correlation coefficients, an average of the spatial shift obtained on the other dates was applied to each direction and each product.

The spatial shifts were applied before resampling. Final surface temperature products were then computed using resampled emissivity and radiances and led to the ASTER input dataset (Figure 3).

Eventually, input parameters for the two models (albedo, emissivity, LAI, fraction cover, vegetation height) were derived using the same procedures as used at HR. This new dataset was called 'LR-ASTER'.

4.4. Estimating Surface Parameters at LR from MODIS Product

Because MODIS *LAI* presented abnormally large values when compared to ASTER, it was estimated from the MODIS NDVI product using the same relationships as for the ASTER datasets.

This also applied to fractional cover f_c. Albedo (blue sky value) was estimated from black-sky and white-sky MODIS values according to [62,63]:

$$\alpha^{MODIS} = (1 - S)\ \alpha_{Black\ Sky} + S\ \alpha_{White\ Sky}, \quad (23)$$

where $\alpha_{BlackSky}$ (directional hemispherical reflectance) was the direct component and was a function of the solar zenith angle and $\alpha_{WhiteSky}$ (bihemispherical reflectance) was the diffuse component. Black-sky and white-sky albedos corresponded to the extreme cases of purely direct or diffuse illumination. S is the fraction of diffuse solar radiation, which varied according to the atmospheric content in aerosol and water vapor and the solar zenith angle. S was fixed to 0.15 for all dates. The use of a constant value for S can introduce discrepancies in albedo and then in the available energy fluxes estimates. It could be estimated from MODIS aerosol product (MOD04-L2), but, in our area in dry season, it was usually low and had only a small impact on blue-sky albedo values. Broadband emissivity was computed from the emissivity products in band 31 and band 32 (ε_{b31} *and* ε_{b32}) with the relationship detailed in Liang [64]:

$$\varepsilon^{MODIS} = 0.4587\ \varepsilon_{b31} + 0.5414\ \varepsilon_{b32}. \quad (24)$$

Crop height values at MODIS scale were computed similarly as for ASTER values using MODIS NDVI product (Equation (24)).

4.5. Aggregation Rules for the Input Parameters

The possibility to monitor ET with a daily time step relies on the use of LR data in the thermal infrared such as those provided by MODIS sensors (with a spatial resolution in the order of 1 km). Other required information can be derived at the same resolution from MODIS data, but other sensors with a higher spatial resolution, in particular on board of Earth Observation satellites such as Sentinel-2, Landsat 8, FORMOSAT or others, can be very useful to introduce the representation of the spatial variability within the MODIS pixels. This offers the possibility of deriving input parameters at LR from their estimation at HR. When combining several platforms or using both Sentinel-2 satellites, time revisit can be down to just a few days.

The spatial aggregation of input parameters may follow different rules depending on the linearity of the equations. Concerning surface energy balance, we expected that roughness length was a significant parameter in this issue, as its behaviour is nonlinear either in the calculation of the turbulent fluxes or in its estimation from remote sensing data [43,47,65]. We also know from the literature review in Section 1 that SEBS is known for its sensitivity to the parameterization of surface roughness lengths. It was therefore important to assess the impact of aggregation rules for this parameter. In the case of other parameters such as albedo, fraction cover or emissivities, an almost linear behaviour was expected.

First, three simple aggregation methods were used to generate LR roughness length from HR data: simple linear average, geometric mean (Equation (25)) and harmonic mean (Equation (26)) [66]:

$$log\left(z_{om_{eff}}\right) = \sum_{i=1}^{n} \alpha_i\ log(z_{om_i}), \quad (25)$$

and

$$\frac{1}{z_{om_{eff}}} = \sum_{i=1}^{n} \frac{\alpha_i}{z_{om_i}}, \quad (26)$$

where α_i is the relative pixel area of each individual roughness (or HR roughness). All three aggregation rules were implemented to derive effective input parameters at LR.

We also wanted to test more deterministic methods such as those proposed by [43,47], which were based on the inversion of the momentum flux equations in the boundary layer at LR over heterogeneous areas. Here, we considered the inversion of the sensible heat flux, which made it possible to estimate effective parameters by considering the aggregation of the aerodynamic conductance. If one assumes that surface temperatures were close to air temperature T_a, the aerodynamic resistance in neutral

conditions is a good approximation of the aerodynamic conditions with true stability corrections. In these conditions, the sensible heat flux expression reduces to:

$$H = \rho\, c_p\, g_{a0} \left(T_{surf} - T_a\right), \tag{27}$$

$$g_{a0} = \frac{1}{r_{a0}} = \frac{k^2 u}{ln\left(\frac{z-d}{z_{om}}\right)\left(ln\left(\frac{z-d}{z_{om}}\right) + kB^{-1}\right)}, \tag{28}$$

where g_{a0} is the aerodynamic conductance to heat transfer at the surface–atmosphere interface in neutral conditions (inverse of the aerodynamic resistance r_{a0} in the same conditions). If one assumed that, in those conditions, surface temperatures were similar at HR and LR, and our purpose was to find effective roughness values reducing as much as possible the difference between the HR and LR aerodynamic conductances g_{a0}. The solutions were reached numerically by finding the effective (LR) surface roughness lengths that solves Equation (30):

$$\sum_{i=1}^{n} \alpha_i g_{a0i} \approx g_{a0eff}, \tag{29}$$

or:

$$\sum_{i=1}^{n}\left[ln\left(\frac{z-\alpha_i d_{0i}}{\alpha_i z_{0mi}}\right) + \alpha_i kB^{-1}{}_i\right] - \left(ln\left(\frac{z-d_{0eff}}{z_{0meff}}\right) - kB^{-1}{}_{eff} = 0, \tag{30}$$

where g_{a0i} is the aerodynamic conductance of HR pixel i, g_{a0eff} the effective value at LR, and $kB^{-1}{}_{eff}$ the effective kB^{-1}, which allows deriving the effective roughness z_{0meff} at LR (Equation (5)). Once computed, the effective surface roughness lengths were used as input into the models at low resolution together with the 'LR-ASTER' input dataset. The new ET maps were compared to the 'agg-ASTER' ET maps.

For each model, the fluxes conservation at LR was quantified by looking at the difference between 'LR-ASTER' and 'agg-ASTER' ET maps when using either the LR roughness estimated at LR or effective parameter values produced from HR data using a simple linear averaging as well as Equations (25), (26) or (30).

5. Results

5.1. Preliminary Step: Evaluation of Flux Estimations at High Resolution from ASTER Data

In a previous study [22], ET maps were obtained on the same area using SEBS and TSEB driven by a combination of the ASTER temperature product and the Vis-NIR Formosat satellite reflectances, as well as in situ crop height measurements. The resulting fluxes were evaluated against the eddy correlation (EC) measurements over seven crop fields whose size is large enough to respect fetch requirements for all dates. This gave an indication of the expected performances for considering the HR fluxes maps as reliable: absolute relative biases were about 18% for SEBS and 23% for TSEB.

In the present study, for a better consistency, ASTER products were used to derive all of the input parameters of SEBS and TSEB, including crop height. The parameters of the f_c(NDVI) and LAI(NDVI) relationships (Equations (20) and (21), respectively) were drawn empirically from the in situ measurements performed in [61]. A comparison to the same EC measurements showed an absolute relative bias lower than 1% for SEBS and about 20% for TSEB (see Figure 4). We thus considered the HR ASTER fluxes as realistic. Moreover, we checked that the available energy fluxes compared well to the measured values (relative bias of 5% and 9% for R_n and G, respectively).

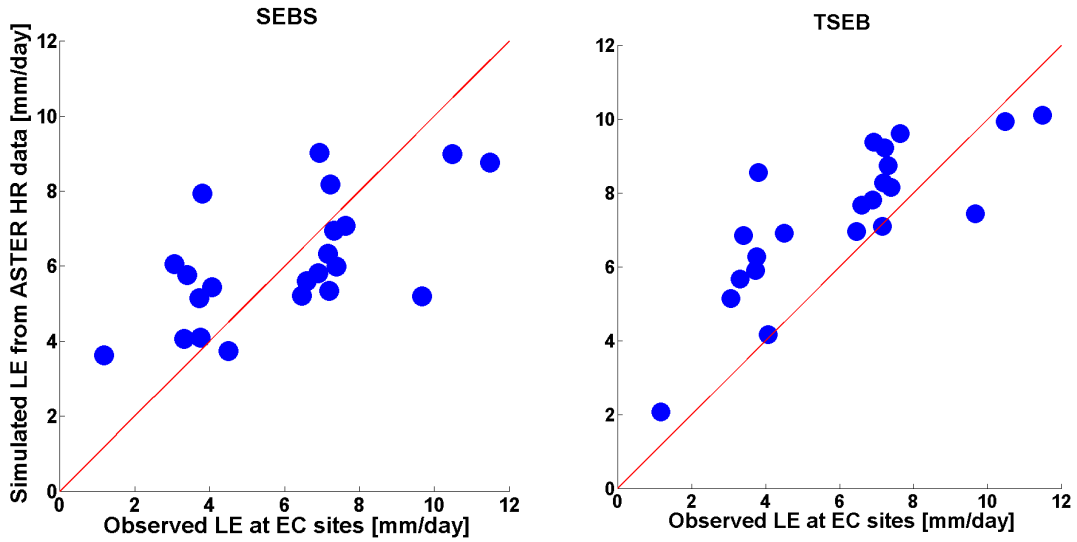

Figure 4. Validation of SEBS (**a**) and TSEB (**b**) HR latent heat fluxes obtained using our simple approach to derive crop heights.

5.2. Flux Estimation at Low Resolution from ASTER LR Data and MODIS Data

ET was estimated using SEBS and TSEB by considering input data at LR derived from ASTER products or MODIS products in a similar way. ET maps at LR were compared to the reference maps obtained by aggregating ET obtained at high resolution from ASTER products ('agg-ASTER').

5.2.1. Available Energy

Available energy was computed with the same method in both models. Very good performances were obtained at LR as well from LR ASTER data as from MODIS data: when compared to 'Agg-ASTER' fluxes, relative biases were smaller than 0.001 (ASTER and MODIS) for R_n and 0.04 (ASTER) and 0.002 (MODIS) for G. The relative bias is the mean relative difference between both estimates normalized to the mean value of the reference value ('Agg-ASTER' in our case). The corresponding Root Mean Square Error (RMSE) values were 0.03 mm·day^{-1} for R_n (ASTER and MODIS) and 0.09 mm·day^{-1} (ASTER) and 0.11 mm·day^{-1} (MODIS) for G. These good performances implied that, for both models, evaluating ET retrievals corresponded to evaluate their capacity to partition the available energy between sensible and latent heat fluxes.

5.2.2. Estimation of ET from LR ASTER with SEBS

The comparison of 'LR-ASTER' ET maps obtained using SEBS to the reference ET maps 'agg-ASTER' showed fairly good performances over the entire dataset, with a relative bias of −0.044 and an RMSE value of 1.0 mm·day^{-1}.

Those performances decreased for individual dates along the season, showing larger biases and RMSE values. More precisely, 'LR ASTER' ET values were smaller than 'agg-ASTER' ET during vegetation growth and higher when the vegetation was senescent (Figure 5). The relative biases were −0.10 on the 30 December 2007, −0.11 on the 23 February 2008, −0.20 on the 10 March 2008, −0.12 on the 11 April 2008, 0.045 on the 27 April 2008, 0.18 on the 6 May 2008 and 0.10 on the 13 May 2008. Those differences globally offset each other and explained the good average performance at the seasonal scale.

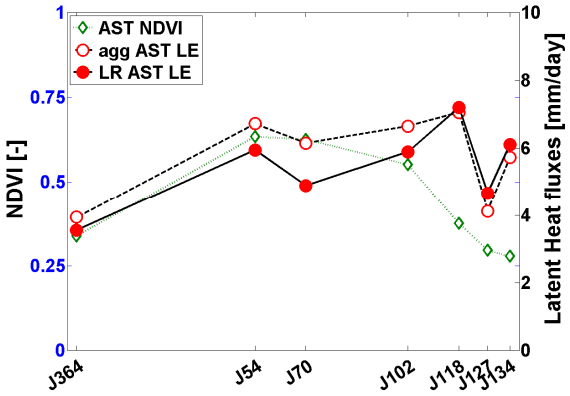

Figure 5. 'agg-ASTER' (agg AST) and 'LR-ASTER' (LR AST) SEBS latent heat flux, along with ASTER NDVI (AST NDVI) time series.

5.2.3. Estimation of ET from LR ASTER with TSEB

The averaged results showed that TSEB fluxes at low resolution were more conservative across the different scales than SEBS, with a relative bias of 0.009 and an RMSE of 0.44 mm·day^{-1}, so that daily ET can be derived with a satisfying precision from LR inputs (Figure 6).

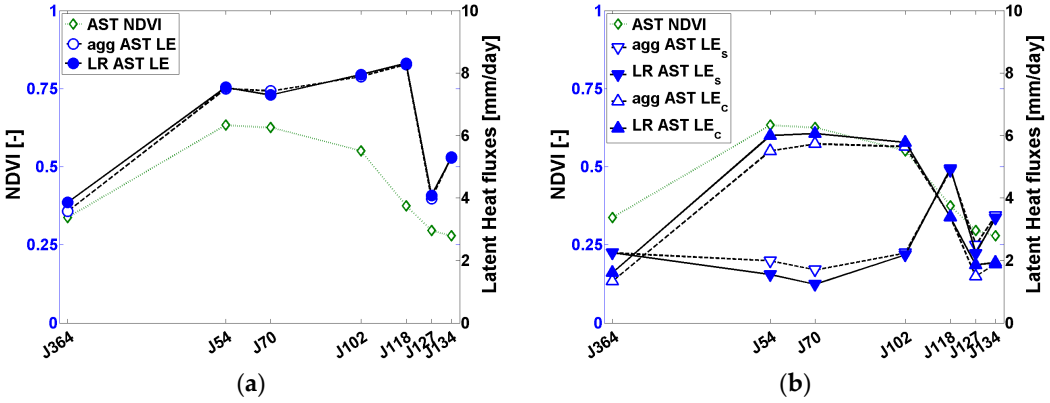

Figure 6. (**a**) 'agg-ASTER' and 'LR-ASTER' TSEB latent heat fluxes time series; (**b**) partition between evaporation LE_s and transpiration LE_v.

However, if the total flux was accurately retrieved using the same parameterizations at high and low resolutions, this did not fully apply to the partition between soil evaporation and plant transpiration, which both presented higher biases (Figure 6). The transpiration obtained from LR inputs was higher than the aggregated HR fluxes (relative bias = 0.065 and RMSE = 0.46 mm·day^{-1}) and the soil evaporation was lower (relative bias = −0.066 and RMSE = 0.38 mm·day^{-1}). Compensation between evaporation and transpiration bias resulted in a very low bias for ET.

5.2.4. Estimation of ET from MODIS Data with SEBS

Results averaged for the seven dates over the 4 km by 4 km (i.e., 2 by 2 pixels at 2000 m spatial resolution) showed performances very close to those obtained with 'LR ASTER'. When SEBS was used to estimate evapotranspiration with MODIS inputs, the relative bias was −0.08 and the RMSE was 1.03 mm·day^{-1}, which was quite similar to the results obtained with 'LR-ASTER' (in particular when considering RMSE). Again, the low biases at the seasonal scale did not reflect the date to date performances and, as for 'LR-ASTER', more contrasted results were obtained for the different vegetation development stages as shown on Figure 7. The resulting relative bias reached high values such as −0.32 on the 10 March 2008, −0.26 on the 11 April 2008 and 0.12 on the 13 May 2008.

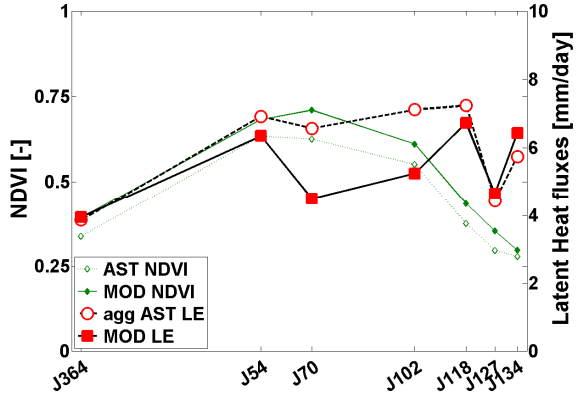

Figure 7. 'agg-ASTER' and MODIS SEBS latent heat flux time series.

5.2.5. Estimation of ET from MODIS Data with TSEB

The averaged results for the 4 × 4 km^2 showed fairly good performances when TSEB was used to estimate ET with MODIS inputs: relative bias on daily ET was 0.03, RMSE 0.96 mm·day^{-1}. However, these results were significantly less favourable than those obtained with 'LR-ASTER'. Model performances regarding soil evaporation and transpiration separately were of similar magnitude as for the total flux when considering RMSE: 0.81 and 1.04 mm·day^{-1} for daily evaporation and transpiration, respectively (Figure 8). As for the results obtained with LR ASTER, the transpiration calculated from MODIS input was higher than the aggregated HR fluxes (relative bias was 0.21) and the evaporation was lower (relative bias was −0.19). Again, compensation between evaporation and transpiration bias resulted in a lower bias for ET. As for the total flux (ET), transpiration and evaporation obtained from MODIS products presented higher discrepancies than from 'LR-ASTER' when compared to 'agg-ASTER'.

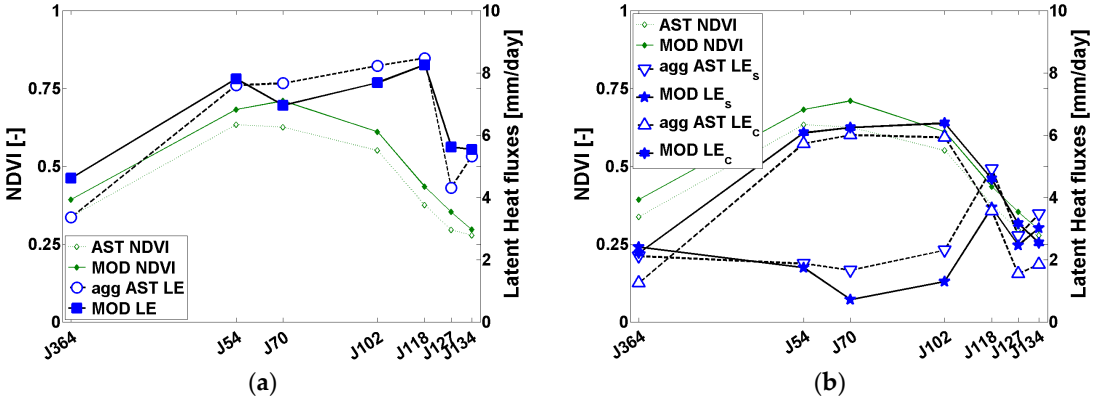

Figure 8. (**a**) 'agg ASTER' and MODIS TSEB latent heat fluxes time series; (**b**) partition between evaporation LE_s and transpiration LE_c.

5.3. Test of Effective Roughness Length Parameterization at Low Resolution

Since TSEB was fairly conservative across scales for total ET, we focused on reducing scale discrepancies for SEBS, using roughness length as a key parameter to explain those differences.

Evaluation

The four aggregation rules of roughness length (linear, geometric i.e., Equation (25), harmonic i.e., Equation (26) and conductance average i.e., Equation (30)) were implemented to derive input roughness lengths for SEBS at LR. The resulting ET calculated by SEBS was compared to the 'agg-ASTER' ET

fluxes. Evaluation of the comparisons is presented in Table 2 when considering LR ASTER dataset and in Table 3 when considering the MODIS dataset. In the two cases, high-resolution roughness lengths were derived from the roughness length maps in the HR ASTER dataset.

For all dates except the 30 December 2007 and the 13 May 2008, it was possible to solve Equation (30) and to derive effective roughness values that reduced the difference between the HR and the LR conductances to zero. For the first and the last dates, no solution to Equation (30) was found using the numerical solver and only a minimum value for the first term of Equation (30) was found (effective roughnesses corresponding to these minimum were used). Compared to the results obtained when roughness lengths were derived directly from LR remote sensing data (Section 5.2), significant reductions in relative biases and/or RMSE values were obtained for most dates when considering the first three types of aggregation methods (either with LR ASTER dataset, Table 2, or MODIS dataset, Table 3). On average, the best results were obtained with the 'harmonic' aggregation rule, with the lowest RMSE (0.52 mm·day^{-1} with LR ASTER and 0.78 mm·day^{-1} with MODIS) and very low relative biases (0.03 and −0.02). The linear aggregation rule resulted only in slight reductions in RMSE and in biases (which, on average, were even increased). The geometric aggregation method gave RMSE on average only slightly higher and relative biases higher than the harmonic aggregation method. For all dates but the 27 April 2008 (low relative bias of 0.04), the relative biases were lower than the previous values obtained in Section 5.2.2. For each date, the relative biases were reduced. This was especially true for the dates that presented biases exceeding 10%.

Table 2. Relative biases for the 'agg ASTER' vs. 'LR ASTER' ET fluxes produced with SEBS at low resolution for different methods to estimate the LR surface roughness length parameter. z_{om}—RS refers to the "remotely sensed" roughness length derived from LR NDVI values, "agg" uses aggregated values from HR to LR using the four methods ("lin": linear, "geo": geometric, "har": harmonic, "grad": deterministic averaging).

Date	z_{om}—RS	z_{om}—agg lin	z_{om}—agg geo	z_{om}—agg har	z_{om}—grad
			Relative bias [–]		
30/12/2007 (J364)	−0.10	−0.10	−0.06	**0.05**	−0.15
23/02/2008 (J54)	−0.11	−0.09	−0.05	**0.00**	−0.07
10/03/2008 (J70)	−0.20	−0.15	−0.09	**0.00**	−0.17
11/04/2008 (J102)	−0.11	−0.09	−0.07	**−0.01**	−0.08
27/04/2008 (J118)	0.02	0.00	**0.01**	0.05	0.00
06/05/2008 (J127)	0.13	−0.10	−0.08	0.10	**−0.02**
13/05/2008 (J134)	0.07	−0.07	−0.06	0.06	−0.10
Seasonal	−0.05	−0.08	−0.05	**0.03**	−0.08
			RMSE [mm/day]		
30/12/2007 (J364)	0.74	0.62	0.44	0.36	1.09
23/02/2008 (J54)	1.49	1.08	0.88	0.70	1.34
10/03/2008 (J70)	1.48	1.11	0.76	0.57	1.61
11/04/2008 (J102)	1.01	0.75	0.56	0.34	1.03
27/04/2008 (J118)	0.48	0.33	0.28	0.42	0.45
06/05/2008 (J127)	1.12	0.60	0.55	0.65	1.26
13/05/2008 (J134)	0.79	0.54	0.55	0.47	1.27
Seasonal	1.08	0.77	0.60	0.52	1.20

Table 3. Relative biases for the 'agg ASTER' vs. MODIS produced with SEBS at low resolution for different methods to estimate the LR surface roughness length parameter (the same symbols as Table 2).

Date	z_{om}—RS	z_{om}—Agg Lin	z_{om}—Agg Geo	z_{om}—Agg Harmonic	z_{om}—Grad
			Relative bias [-]		
30/12/2007 (J364)	0.02	0.12	0.18	0.18	0.07
23/02/2008 (J54)	−0.08	0.02	0.05	−0.02	−0.04
10/03/2008 (J70)	−0.32	−0.12	−0.08	−0.16	−0.19
11/04/2008 (J102)	−0.26	−0.21	−0.18	−0.14	−0.17
27/04/2008 (J118)	−0.07	−0.13	−0.11	−0.01	−0.03
06/05/2008 (J127)	0.05	−0.24	−0.11	0.11	0.09
13/05/2008 (J134)	0.12	−0.03	0.03	0.04	−0.02
Seasonal	−0.10	−0.09	**−0.04**	**−0.02**	**−0.06**
			RMSE [mm/day]		
30/12/2007 (J364)	0.35	0.65	0.73	0.70	0.35
23/02/2008 (J54)	0.65	0.47	0.53	0.48	0.56
10/03/2008 (J70)	2.13	0.84	0.59	1.21	1.36
11/04/2008 (J102)	2.04	1.68	1.42	1.20	1.40
27/04/2008 (J118)	0.62	1.08	0.89	0.28	0.30
06/05/2008 (J127)	0.69	1.39	1.04	0.62	0.53
13/05/2008 (J134)	0.74	0.40	0.27	0.44	0.68
Seasonal	1.23	1.03	0.86	0.78	0.85

The fourth aggregation method relied on deriving the effective surface roughness lengths from the aggregation of aerodynamic conductance in neutral conditions (Section 4.5). Using the LR ASTER dataset, it was possible to solve Equation (30) for all dates, except for 30 December 2007, and 13 May 2008, and to derive effective roughness values that reduced the difference between the HR and the LR conductances to zero. For the first and the last dates, no solution to Equation (30) was found using the numerical solver and only a minimum value for the first term of Equation (30) was found (effective roughness lengths corresponding to these minimums were used). RMSE were not improved for most dates and on average, RMSE and relative bias were higher than those obtained directly from LR ASTER data (Table 2). The results obtained when using the MODIS dataset (Table 3) presented lower RMSE and lower relative biases than those obtained from the LR ASTER dataset. The performances were always as good as the geometrical aggregation performances and even better at the beginning of the growing season and the beginning of the senescent period.

When considering a comparison of the two datasets for each aggregation methods, it appeared that the use of MODIS data usually resulted in similar or lower performances than with LR ASTER for the direct use of LR data and for the first three aggregation rules. Conversely, in most cases, the fourth aggregation method provided significantly lower RMSE and bias when using MODIS data rather than LR ASTER data.

6. Discussion

6.1. Discussion on Flux Estimation at Low Resolution

The results obtained using SEBS and TSEB with low resolution input data showed that

- TSEB had significantly higher or similar scaling performances (flux conservation across scales) as SEBS, respectively, with LR ASTER data and MODIS data;
- SEBS had similar scaling performances with LR ASTER and MODIS data;
- TSEB scaling performances were significantly lower with MODIS data than with LR ASTER data (for ET mapping as well as for ET partitioning in E and T);
- With TSEB, relative biases between the soil and the vegetation offset each other when one considers the whole season.

Altogether, TSEB was more conservative through scaling than SEBS (see Figure 6 vs. Figure 5). For TSEB, the issue of roughness length aggregation is expressed differently than in SEBS [67]. Whereas SEBS's sensitivity to roughness is explained by the fact that the roughness length for heat exchange (Equation (5)) is proportional to the mechanical roughness length z_{om}, this is not the case for TSEB. In the latter, however, the aerodynamic temperature is tightly linked to the component (soil and vegetation) energy budgets and therefore very sensitive to the vegetation cover fraction f_c. In our case, this translated into different evaporation/transpiration partitions across scales but did not much affect the total flux, which remained quite conservative across scales.

Both models used NDVI to compute input parameters, in particular LAI, vegetation height and then aerodynamic roughness and the fraction cover (Equations (20)–(22)). Significant differences existed between the NDVI for the two sensors ASTER and MODIS (see Figure 7). These differences were related in particular to differences in NDVI construction and in sensors' spectral bands. ASTER NDVI were derived from reflectance products (instantaneous acquisition) while MODIS NDVI were obtained from a 16-day composite product. Red and Near InfraRed spectral bands were much larger for ASTER than for MODIS. Higher reflectances in the red band and lower reflectances in the NIR bands were expected for ASTER sensor resulting in lower NDVI. This induced differences in fraction cover (Equation (21)) and vegetation height (and therefore the roughness length, Equation (22)), somewhat enhanced by the fact that common maximum and minimum NDVI values are used for both sensors for LAI retrieval.

Even though NDVI differences between both sensors were significant, the impact on scaling between 'LR-ASTER' and 'MODIS' was not significantly different for SEBS. As explained in Section 3.2, surface roughness length influences also the additional resistance to ET (the kB^{-1} term in Equation (6)). Roughness lengths values are derived from NDVI products coming from reflectances measured by two different sensors at different spatial resolutions. Both sensors will differ in their representation of the subscale heterogeneity and its associated vegetation pattern. It is also clear that the model sensitivity to the surface roughness lengths induces the high ET flux differences in the middle of the season when the heterogeneity is the most pronounced.

For TSEB however, it was obvious that the use of MODIS inputs reduced the scaling performances obtained with LR ASTER inputs when looking at evaporation and transpiration separately. The partition was indeed very sensitive to fraction cover and thus to discrepancies in reconstructed NDVI values across scales. This illustrated the sensitivity of TSEB to the partition of the available energy between the two sources through the f_c parameter.

As a result, there were higher transpiration and lower soil evaporation for MODIS compared to LR ASTER. However, these differences of opposite signs led to a fairly good estimate of the surface ET flux. For instance, evapotranspiration fluxes were highly biased on 13 May 2008 (relative bias = -0.13 and 0.37, RMSE = 0.49 and 0.84 mm·day^{-1}, respectively, for soil evaporation and transpiration fluxes), but ET was well estimated with MODIS data (see Figure 8).

On the 30 December 2007 and the 5 May 2008 most of the pixels were covered with a mix of green vegetation and bare soil. These two dates were representative of two phenological stages during the growing season (early growth and senescence) when the heterogeneity reached its maximum. The 30 December 2007 corresponded to a sowing period. Hence, depending on the vegetation type and the time-lapse since the sowing date, each plot presented different NDVI values at HR. The 5 May 2008 corresponded to the beginning of the senescence and was close to harvesting. On this day, the area

was covered by a mix of more or less senescent vegetation and bare soil for plots already harvested with a wide range of NDVI values at HR. This information about the heterogeneity of the crop height parameter was somewhat taken into account by the LR ASTER inputs, albeit imperfectly, because it came from the HR ASTER inputs whose spatial resolution allowed a faithful representation of this heterogeneity. The MODIS NDVI values and their low spatial resolution were only able to represent roughly the main vegetation status present in a LR pixel (Figure 9):

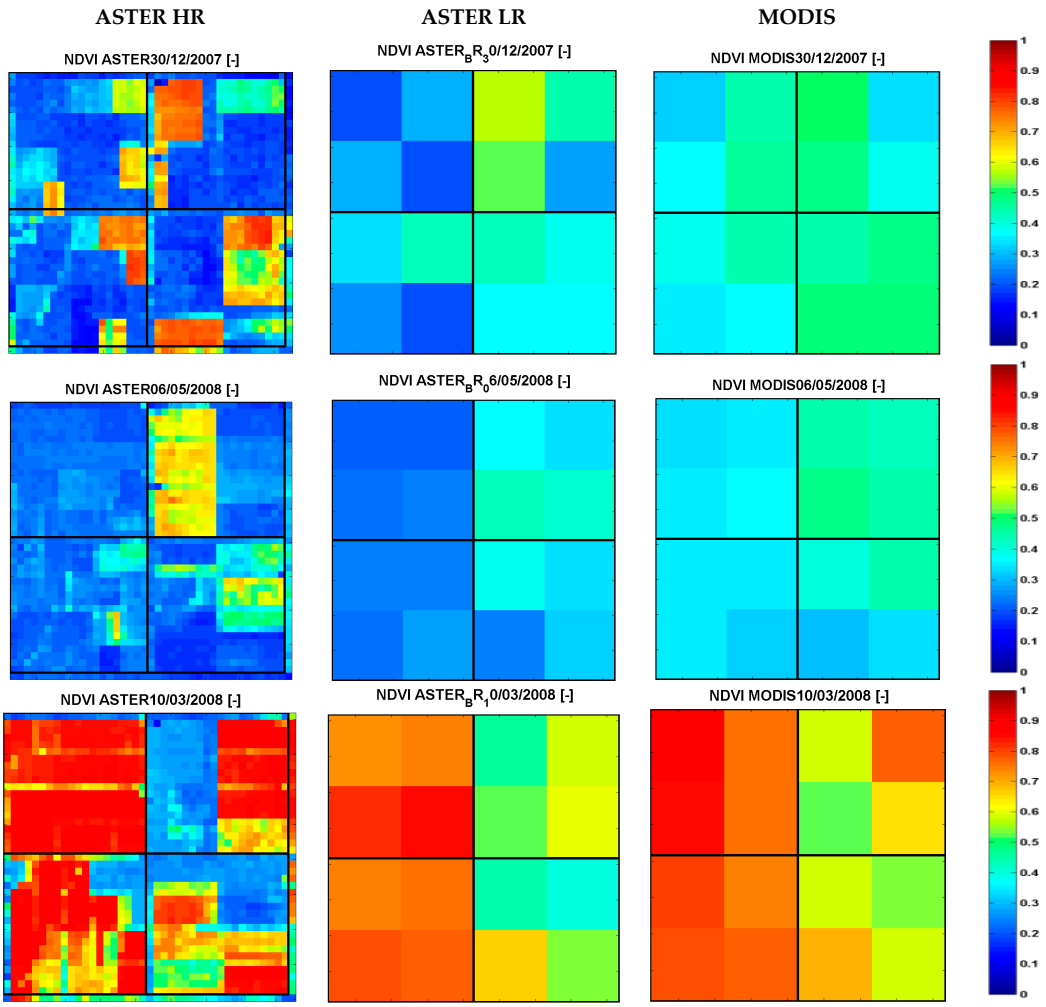

Figure 9. Discrepancies between LR and HR NDVI for three dates within the growing season.

These differences between LR ASTER NDVI and MODIS NDVI directly implied discrepancies in the respective f_c values and, hence, biases in the available energy partition (R_{nc} relative biases = -0.36 and -0.52; RMSE = 1.00 and 1.42 mm·day^{-1}, respectively, for the 30 December 2007 and the 6 May 2008) that resulted in biases in transpiration between 'agg-ASTER' and MODIS instantaneous fluxes (LE_c relative biases = 0.77 and 1.04; RMSE = 1.05 and 1.64 mm·day^{-1}, respectively, the 30 December 2007 and the 6 May 2008). Biases on soil evaporation were smaller and did not offset the bias in transpiration on that particular day, but this did not affect the seasonal average performance of the TSEB model.

Convergence of f_c and roughness probability distributions between LR ASTER and MODIS using adapted maximum and minimum values of NDVI for MODIS could probably decrease this difference. Since the scaling analysis reies solely on the 'LR-ASTER' vs. 'agg-ASTER' intercomparison, the 'MODIS' vs. 'agg-ASTER' intercomparison is meant only to provide an illustration of a "real world" application of the models to moderate resolution data; therefore, the intermediate exercise consisting in

matching MODIS and 'agg-ASTER' NDVI distribution while keeping MODIS LST was not considered in this study.

6.2. Discussion on the Efficiency of Aggregation Rules

Results showed that if HR NDVI data were available, the linear aggregation rule and the deterministic method were less accurate than the two other averaging methods tested in this study. Both the geometric and harmonic aggregation methods reduce all daily biases and RMSE values with an advantage for the harmonic aggregation rule, which also reduces considerably the average bias. This is consistent with [43] who evaluated the geometric average as efficient to estimate effective surface roughness to reduce the error on fluxes estimation. The deterministic rule performed well but did not outperform the harmonic mean, which is a clear indication that aggregation of the roughness length is the main but not the sole element to account for in scaling from HR to LR.

7. Conclusions

EC measurements acquired in an extensive experiment in the Yaqui valley in Mexico over a 4×4 km^2 agricultural area were used to validate HR (100 m) ET maps computed with the SEBS and the TSEB models forced by ASTER data, local meteorological data and in situ measured crop heights in [22]. In order to bypass local crop height evaluation, we implemented a simple approach to derive crop heights (thus, indirectly surface roughness lengths) based on a priori maximum and minimum values scaled by NDVI levels. A common relationship valid for cereal covers in this area dominated by cereals or crops with similar height was derived. The performances of the SEBS and the TSEB models, when forced by remotely sensed vegetation height, were found to be similar to those obtained when in situ crop height measurements were used.

The satisfying intercomparison between measured ET and simulated ET at EC flux stations within the test area led to consider the HR ET maps as reliable datasets to study scaling properties between ET estimation at HR and LR (1000 m) in order to take advantage of the abundant availability of data at LR. A reference dataset was thus built by aggregating linearly ET fluxes from 100 m to 1000 m resolution maps for each model. Scaling properties were thus assessed by comparing ET fluxes computed from LR data only, and these reference maps deduced from the sole HR data. LR input data was first produced from HR ASTER data in order to investigate scaling properties with the same sensor, and then from the MODIS sensor in order to investigate the impact of the radiometric differences on the resulting LR ET product from one sensor to the other. We confirmed that the discrepancies between the reference maps deduced from SEBS HR outputs and all maps obtained by running SEBS with LR input data was mostly due to improper scaling of roughness length (i.e., vegetation height), which was consistent with the high model sensitivity to this parameter recognized in many previous studies [36,37]. TSEB was found to be more conservative between scales than SEBS, discrepancies on total ET flux being much smaller than for SEBS in particular when using LR ASTER dataset. However, the flux partitioning between evaporation and transpiration was very sensitive to small differences between fraction cover patterns at high and low resolutions.

Since surface properties such as vegetation height and fraction cover could be derived from time series of HR images because they can provide information for each field (Land Use type, LAI, and fraction cover), we investigated the use of aggregation rules to derive more spatially resolved estimates of those quantities to be used by either model at LR.

Four aggregation rules were tested, linear, harmonic and geometric averaging, as well as a method that numerically found roughness lengths that minimize the difference between aggregated aerodynamic conductances in neutral conditions and the effective aerodynamic conductance at LR. The harmonic mean appeared to be the most accurate both for LR ASTER dataset and MODIS dataset. This exercise showed how the proper scaling of input parameters can improve the consistency of ET estimates across scales.

Those results must, however, be scaled to the large differences between both models as well as between the observations and the model outputs at High Resolution. By providing the best aggregation rule, one ensures that at least the aggregation error does not add up to the other error sources.

This study can be considered as representative of typical situations for semi-arid environment and irrigated crops. It will be interesting to evaluate these results in other climatic conditions, over an area presenting different irrigation practices or without irrigation practices and, hence, submitted to different stress conditions.

It would be of interest to evaluate the performances of these models on a study area with even more contrasting roughness values. The parameterization of the surface roughness lengths has been a real issue and its small range over the study area made the latter relatively more reachable. In an environment presenting crop heights ranging from bare soil to high trees as well as intermediate values (e.g., vegetables and cereals), the exercise would be more difficult.

Radar measurements could help in this way through their ability to go through different vegetation layers to estimate properly the crop height and the vegetation fraction cover parameters.

In the future, the model developed by [47] will be tested to estimate an effective surface roughness length through HR land use, surface temperature, LAI products, and local meteorological observations. This model includes an additional resistance representing the effects of the molecular diffusion through the viscous sub-layers on the sensible heat flux.

Acknowledgments: Financial support by the French Space Agency (Centre National d'Etudes Spatiales, CNES)/ "Terre Océan Surfaces Continentales Atmosphère" (TOSCA) program for the Thermal InfraRed SpaTial System(THIRSTY) phase 0 study through projects "EVAluation de l'EVApotranspiration issue de l'Infra Rouge Thermique" (EVA2IRT) and "EVApotranspiration from SPAce V3" (EVASPA3), as well as the IRD (Institut de Recherche pour le Développement), the ITSON (Instituto Tecnologico de SONora), and the University of Sonora in the setting up of the "Yaqui experiment" are gratefully acknowledged. Maria Mira was endorsed by a postdoctoral contract within the VALi+d program from Generalitat Valenciana (Spain), followed by a CNES postdoctoral contract and a "Juan de la Cierva" postdoctoral contract from the Spanish Ministry of Economy and Competitiveness. Belen Gallego-Elvira was endorsed by a postdoctoral fellowship from the Ramon Areces Foundation.

Author Contributions: Malik Bahir conducted the study and wrote the paper, Gilles Boulet and Albert Olioso designed the protocol, helped with analysing the results and participated in writing the paper, Vincent Rivalland helped with the RS data processing, Belen Gallego-Elvira and Maria Mira provided guidance for the RS products, Lionel Jarlan and Julio-Cesar Rodriguez were responsible for the acquisition of the data collected during the Yaqui Valley experiment, and Olivier Merlin provided guidance on MODIS data resampling.

Conflicts of Interest: The authors declare no conflict of interest.

References

1. Leduc, C.; Pulido-Bosch, A.; Remini, B. Anthropization of groundwater resources in the Mediterranean region: Processes and challenges. *Hydrogeol. J.* **2017**, *25*, 1529–1547. [CrossRef]
2. Saadi, S.; Simonneaux, V.; Boulet, G.; Raimbault, B.; Mougenot, B.; Fanise, P.; Ayari, H.; Lili-Chabaane, Z. Monitoring irrigation consumption using high resolution NDVI image time series: Calibration and validation in the Kairouan Plain (Tunisia). *Remote Sens.* **2015**, *7*, 13005–13028. [CrossRef]
3. Anderson, M.C.; Kustas, W.P.; Norman, J.M.; Hain, C.R.; Mecikalski, J.R.; Schultz, L.; Gonzalez-Dugo, M.P.; Cammalleri, C.; d'Urso, G.; Pimstein, A.; et al. Mapping daily evapotranspiration at field to continental scales using geostationary and polar orbiting satellite imagery. *Hydrol. Earth Syst. Sci.* **2011**, *15*, 223–239. [CrossRef]
4. Bastiaanssen, W.G.M.; Menenti, M.; Feddes, R.A.; Holtslag, A.A.M. A remote sensing surface energy balance algorithm for land (SEBAL). 1. Formulation. *J. Hydrol.* **1998**, *212–213*, 198–212. [CrossRef]
5. Jiang, L.; Islam, S. A methodology for estimation of surface evapotranspiration over large areas using remote sensing observations. *Geophys. Res. Lett.* **1999**, *26*, 2773–2776. [CrossRef]
6. Moran, M.S.; Clarke, T.R.; Inoue, Y.; Vidal, A. Estimating crop water deficit using the relation between surface-air temperature and spectral vegetation index. *Remote Sens. Environ.* **1994**, *49*, 246–263. [CrossRef]
7. Roerink, G.J.; Su, Z.; Menenti, M. S-SEBI: A simple remote sensing algorithm to estimate the surface energy balance. *Phys. Chem. Earth Part B Hydrol. Oceans Atmos.* **2000**, *25*, 147–157. [CrossRef]

8. Merlin, O.; Chirouze, J.; Olioso, A.; Jarlan, L.; Chehbouni, G.; Boulet, G. An image-based four-source surface energy balance model to estimate crop evapotranspiration from solar reflectance/thermal emission data (SEB-4S). *Agric. For. Meteorol.* **2014**, *184*, 188–203. [CrossRef]
9. Tang, R.L.; Li, Z.L.; Chen, K.S.; Jia, Y.Y.; Li, C.R.; Sun, X.M. Spatial-scale effect on the SEBAL model for evapotranspiration estimation using remote sensing data. *Agric. For. Meteorol.* **2013**, *174*, 28–42. [CrossRef]
10. Galleguillos, M.; Jacob, F.; Prévot, L.; French, A.; Lagacherie, P. Comparison of two temperature differencing methods to estimate daily evapotranspiration over a Mediterranean vineyard watershed from ASTER data. *Remote Sens. Environ.* **2011**, *115*, 1326–1340. [CrossRef]
11. Long, D.; Singh, V.P. Assessing the impact of end-member selection on the accuracy of satellite-based spatial variability models for actual evapotranspiration estimation. *Water Resour. Res.* **2013**, *49*, 2601–2618. [CrossRef]
12. Agam, N.; Kustas, W.P.; Anderson, M.C.; Li, F.Q.; Neale, C.M.U. A vegetation index based technique for spatial sharpening of thermal imagery. *Remote Sens. Environ.* **2007**, *107*, 545–558. [CrossRef]
13. Cammalleri, C.; Anderson, M.C.; Gao, F.; Hain, C.R.; Kustas, W.P. Mapping daily evapotranspiration at field scales over rainfed and irrigated agricultural areas using remote sensing data fusion. *Agric. For. Meteorol.* **2014**, *186*, 1–11. [CrossRef]
14. Chen, X.H.; Li, W.T.; Chen, J.; Rao, Y.H.; Yamaguchi, Y. A combination of TsHARP and thin plate spline interpolation for spatial sharpening of thermal imagery. *Remote Sens.* **2014**, *6*, 2845–2863. [CrossRef]
15. Mechri, R.; Ottle, C.; Pannekoucke, O.; Kallel, A. Genetic particle filter application to land surface temperature downscaling. *J. Geophys. Res. Atmos.* **2014**, *119*, 2131–2146. [CrossRef]
16. Merlin, O.; Duchemin, B.; Hagolle, O.; Jacob, F.; Coudert, B.; Chehbouni, G.; Dedieu, G.; Garatuza, J.; Kerr, Y. Disaggregation of modis surface temperature over an agricultural area using a time series of formosat-2 images. *Remote Sens. Environ.* **2010**, *114*, 2500–2512. [CrossRef]
17. Zhan, W.F.; Chen, Y.H.; Zhou, J.; Wang, J.F.; Liu, W.Y.; Voogt, J.; Zhu, X.L.; Quan, J.L.; Li, J. Disaggregation of remotely sensed land surface temperature: Literature survey, taxonomy, issues, and caveats. *Remote Sens. Environ.* **2013**, *131*, 119–139. [CrossRef]
18. Boulet, G.; Olioso, A.; Ceschia, E.; Marloie, O.; Coudert, B.; Rivalland, V.; Chirouze, J.; Chehbouni, G. An empirical expression to relate aerodynamic and surface temperatures for use within single-source energy balance models. *Agric. For. Meteorol.* **2012**, *161*, 148–155. [CrossRef]
19. Su, Z. The Surface Energy Balance System (SEBS) for estimation of turbulent heat fluxes. *Hydrol. Earth Syst. Sci.* **2002**, *6*, 85–99. [CrossRef]
20. Boulet, G.; Mougenot, B.; Lhomme, J.P.; Fanise, P.; Lili-Chabaane, Z.; Olioso, A.; Bahir, M.; Rivalland, V.; Jarlan, L.; Merlin, O.; et al. The sparse model for the prediction of water stress and evapotranspiration components from thermal infra-red data and its evaluation over irrigated and rainfed wheat. *Hydrol. Earth Syst. Sci.* **2015**, *19*, 4653–4672. [CrossRef]
21. Norman, J.M.; Kustas, W.P.; Humes, K.S. Source approach for estimating soil and vegetation energy fluxes in observations of directional radiometric surface temperature. *Agric. For. Meteorol.* **1995**, *77*, 263–293. [CrossRef]
22. Chirouze, J.; Boulet, G.; Jarlan, L.; Fieuzal, R.; Rodriguez, J.C.; Ezzahar, J.; Er-Raki, S.; Bigeard, G.; Merlin, O.; Garatuza-Payan, J.; et al. Intercomparison of four remote-sensing-based energy balance methods to retrieve surface evapotranspiration and water stress of irrigated fields in semi-arid climate. *Hydrol. Earth Syst. Sci.* **2014**, *18*, 1165–1188. [CrossRef]
23. Su, Z.; Dorigo, W.; Fernández-Prieto, D.; Van Helvoirt, M.; Hungershoefer, K.; de Jeu, R.; Parinussa, R.; Timmermans, J.; Roebeling, R.; Schröder, M.; et al. Earth observation Water Cycle Multi-Mission Observation Strategy (WACMOS). *Hydrol. Earth Syst. Sci. Discuss.* **2010**, *7*, 7899–7956. [CrossRef]
24. Jia, L.; Su, Z.B.; van den Hurk, B.; Menenti, M.; Moene, A.; De Bruin, H.A.R.; Yrisarry, J.J.B.; Ibanez, M.; Cuesta, A. Estimation of sensible heat flux using the Surface Energy Balance System (SEBS) and ATSR measurements. *Phys. Chem. Earth* **2003**, *28*, 75–88. [CrossRef]
25. Jia, Z.; Liu, S.; Xu, Z.; Chen, Y.; Zhu, M. Validation of remotely sensed evapotranspiration over the Hai River Basin, China. *J. Geophys. Res. Atmos.* **2012**, *117*. [CrossRef]
26. Kleissl, J.; Hong, S.H.; Hendrickx, J.M.H. New Mexico scintillometer network supporting remote sensing and hydrologic and meteorological models. *Bull. Am. Meteorol. Soc.* **2009**, *90*, 207–218. [CrossRef]

27. Tang, R.L.; Li, Z.L.; Jia, Y.Y.; Li, C.R.; Sun, X.M.; Kustas, W.P.; Anderson, M.C. An intercomparison of three remote sensing-based energy balance models using large aperture scintillometer measurements over a wheat-corn production region. *Remote Sens. Environ.* **2011**, *115*, 3187–3202. [CrossRef]
28. Corbari, C.; Mancini, M.; Su, Z.; Li, J. Evapotranspiration estimate from water balance closure using satellite data for the Upper Yangtze River basin. *Hydrol. Res.* **2014**, *45*, 603–614. [CrossRef]
29. Velpuri, N.M.; Senay, G.B.; Singh, R.K.; Bohms, S.; Verdin, J.P. A comprehensive evaluation of two MODIS evapotranspiration products over the conterminous United States: Using point and gridded FLUXNET and water balance ET. *Remote Sens. Environ.* **2013**, *139*, 35–49. [CrossRef]
30. Su, H.; Wood, E.F.; McCabe, M.F.; Su, Z. Evaluation of remotely sensed evapotranspiration over the CEOP eop-1 reference sites. *J. Meteorol. Soc. Jpn.* **2007**, *85A*, 439–459. [CrossRef]
31. Su, H.B.; McCabe, M.F.; Wood, E.F.; Su, Z.; Prueger, J.H. Modeling evapotranspiration during SMACEX: Comparing two approaches for local- and regional-scale prediction. *J. Hydrometeorol.* **2005**, *6*, 910–922. [CrossRef]
32. Verstraeten, W.W.; Veroustraete, F.; Feyen, J. Estimating evapotranspiration of European forests from NOAA-imagery at satellite overpass time: Towards an operational processing chain for integrated optical and thermal sensor data products. *Remote Sens. Environ.* **2005**, *96*, 256–276. [CrossRef]
33. Baret, F.; Weiss, M.; Allard, D.; Garrigue, S.; Leroy, M.; Jeanjean, H.; Fernandes, R.; Myneni, R.B.; Privette, J.; Morisette, J.; et al. Valeri: A network of sites and a methodology for the validation of medium spatial resolution land satellite products. *Remote Sens. Environ.* **2013**, *76*, 36–39.
34. Mira, M.; Weiss, M.; Baret, F.; Courault, D.; Hagolle, O.; Gallego-Elvira, B.; Olioso, A. The MODIS (collection V006) BRDF/albedo product MCD43D: Temporal course evaluated over agricultural landscape. *Remote Sens. Environ.* **2015**, *170*, 216–228. [CrossRef]
35. Etchanchu, J.; Rivalland, V.; Gascoin, S.; Cros, J.; Brut, A.; Boulet, G. Effects of multi-temporal high-resolution remote sensing products on simulated hydrometeorological variables in a cultivated area (southwestern France). *Hydrol. Earth Syst. Sci. Discuss.* **2017**, *2017*, 1–23. [CrossRef]
36. McCabe, M.F.; Wood, E.F. Scale influences on the remote estimation of evapotranspiration using multiple satellite sensors. *Remote Sens. Environ.* **2006**, *105*, 271–285. [CrossRef]
37. Ershadi, A.; McCabe, M.F.; Evans, J.P.; Walker, J.P. Effects of spatial aggregation on the multi-scale estimation of evapotranspiration. *Remote Sens. Environ.* **2013**, *131*, 51–62. [CrossRef]
38. Kustas, W.P.; Norman, J.M. Evaluating the effects of subpixel heterogeneity on pixel average fluxes. *Remote Sens. Environ.* **2000**, *74*, 327–342. [CrossRef]
39. Kustas, W.P.; Norman, J.M.; Shmugge, T.J.; Anderson, M.C. Mapping surface energy fluxes with radiometric temperature. In *Thermal Remote Sensing in Land Surface Processes*; CRC Press: Boca Raton, FL, USA, 2004; pp. 205–253.
40. Kalma, J.D.; McVicar, T.R.; McCabe, M.F. Estimating land surface evaporation: A review of methods using remotely sensed surface temperature data. *Surv. Geophys.* **2008**, *29*, 421–469. [CrossRef]
41. Chehbouni, A.; Njoku, E.G.; Lhomme, J.P.; Kerr, Y.H. Approaches for averaging surface parameters and fluxes over heterogeneous terrain. *J. Clim.* **1995**, *8*, 1386–1393. [CrossRef]
42. Kim, C.P.; Stricker, J.N.M.; Feddes, R.A. Impact of soil heterogeneity on the water budget of the unsaturated zone. *Water Resour. Res.* **1997**, *33*, 991–999. [CrossRef]
43. Wassenaar, T.; Olioso, A.; Hasager, C.; Jacob, F.; Chehbouni, A. Estimation of evapotranspiration on heterogeneous pixels. In Proceedings of the First International Symposium on Recent Advances in Quantitative Remote Sensing, Valencia, Spain, 16–20 September 2002; Sobrino, J.A., Ed.; Publicacions de la Universitat de València: Valencia, Spain, 2002; pp. 458–465.
44. Taylor, P.A. Comments and further analysis on effective roughness lengths for use in numerical 3-dimensional models. *Bound.-Layer Meteorol.* **1987**, *39*, 403–418. [CrossRef]
45. Byun, K.; Liaqat, U.W.; Choi, M. Dual-model approaches for evapotranspiration analyses over homo- and heterogeneous land surface conditions. *Agric. For. Meteorol.* **2014**, *197*, 169–187. [CrossRef]
46. Choi, M.; Kustas, W.P.; Anderson, M.C.; Allen, R.G.; Li, F.Q.; Kjaersgaard, J.H. An intercomparison of three remote sensing-based surface energy balance algorithms over a corn and soybean production region (Iowa, US) during SMACEX. *Agric. For. Meteorol.* **2009**, *149*, 2082–2097. [CrossRef]
47. Hasager, C.B.; Jensen, N.O.; Olioso, A. Land cover, surface temperature and leaf area index maps from satellites used for the aggregation of momentum and temperature roughnesses. In Proceedings of

the First International Symposium on Recent Advances in Quantitative Remote Sensing, Valencia, Spain, 16–20 September 2002; pp. 466–473.
48. Michel, D.; Jiménez, C.; Miralles, D.G.; Jung, M.; Hirschi, M.; Ershadi, A.; Martens, B.; McCabe, M.F.; Fisher, J.B.; Mu, Q.; et al. The WACMOS-ET project—Part 1: Tower-scale evaluation of four remote-sensing-based evapotranspiration algorithms. *Hydrol. Earth Syst. Sci.* **2016**, *20*, 803–822. [CrossRef]
49. Jacob, F.; Olioso, A.; Gu, X.F.; Su, Z.B.; Seguin, B. Mapping surface fluxes using airborne visible, near infrared, thermal infrared remote sensing data and a spatialized surface energy balance model. *Agronomie* **2002**, *22*, 669–680. [CrossRef]
50. Anderson, M.C.; Norman, J.M.; Kustas, W.P.; Li, F.; Prueger, J.H.; Mecikalski, J.R. Effects of vegetation clumping on two-source model estimates of surface energy fluxes from an agricultural landscape during SMACEX. *J. Hydrometeorol.* **2005**, *6*, 892–909. [CrossRef]
51. Timmermans, W.J.; van der Kwast, J.; Gieske, A.S.M.; Su, Z.; Olioso, A.; Jia, L.; Elbers, J. Intercomparison of energy flux models using ASTER imagery at the SPARC 2004 site (Barrax, Spain). In Proceedings of the SPARC Final Workshop, Enschede, The Netherlands, 4–5 July 2005; p. 8.
52. French, A.N.; Jacob, F.; Anderson, M.C.; Kustas, W.P.; Timmermans, W.; Gieske, A.; Su, Z.; Su, H.; McCabe, M.F.; Li, F.; et al. Surface energy fluxes with the Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) at the Iowa 2002 SMACEX site (USA). *Remote Sens. Environ.* **2005**, *99*, 55–65. [CrossRef]
53. Van der Kwast, J.; Timmermans, W.; Gieske, A.; Su, Z.; Olioso, A.; Jia, L.; Elbers, J.; Karssenberg, D.; de Jong, S. Evaluation of the Surface Energy Balance System (SEBS) applied to ASTER imagery with flux-measurements at the SPARC 2004 site (Barrax, Spain). *Hydrol. Earth Syst. Sci.* **2009**, *13*, 1337–1347. [CrossRef]
54. Delogu, E.; Boulet, G.; Olioso, A.; Coudert, B.; Chirouze, J.; Ceschia, E.; Le Dantec, V.; Marloie, O.; Chehbouni, G.; Lagouarde, J.P. Reconstruction of temporal variations of evapotranspiration using instantaneous estimates at the time of satellite overpass. *Hydrol. Earth Syst. Sci.* **2012**, *16*, 2995–3010. [CrossRef]
55. Baldocchi, D.; Falge, E.; Gu, L.; Olson, R.; Hollinger, D.; Running, S.; Anthoni, P.; Bernhofer, C.; Davis, K.; Evans, R.; et al. Fluxnet: A new tool to study the temporal and spatial variability of ecosystem–scale carbon dioxide, water vapor, and energy flux densities. *Bull. Am. Meteorol. Soc.* **2001**, *82*, 2415–2434. [CrossRef]
56. Brutsaert, W. Aspects of bulk atmospheric boundary layer similarity under free-convective conditions. *Rev. Geophys.* **1999**, *37*, 439–451. [CrossRef]
57. Su, Z.; Schmugge, T.; Kustas, W.P.; Massman, W.J. An evaluation of two models for estimation of the roughness height for heat transfer between the land surface and the atmosphere. *J. Appl. Meteorol.* **2001**, *40*, 1933–1951. [CrossRef]
58. Gomez, M.; Olioso, A.; Sobrino, J.A.; Jacob, F. Retrieval of evapotranspiration over the Alpilles/ReSeDA experimental site using airborne POLDER sensor and a thermal camera. *Remote Sens. Environ.* **2005**, *96*, 399–408. [CrossRef]
59. Liang, S.L. Narrowband to broadband conversions of land surface albedo I algorithms. *Remote Sens. Environ.* **2001**, *76*, 213–238. [CrossRef]
60. Ogawa, K.; Schmugge, T.; Rokugawa, S. Estimating broadband emissivity of arid regions and its seasonal variations using thermal infrared remote sensing. *IEEE Trans. Geosci. Remote Sens.* **2008**, *46*, 334–343. [CrossRef]
61. Fieuzal, R.; Duchemin, B.; Jarlan, L.; Zribi, M.; Baup, F.; Merlin, O.; Hagolle, O.; Garatuza-Payan, J. Combined use of optical and radar satellite data for the monitoring of irrigation and soil moisture of wheat crops. *Hydrol. Earth Syst. Sci.* **2011**, *15*, 1117–1129. [CrossRef]
62. Lewis, P.; Barnsley, M.J. Influence of the sky radiance distribution on various formulations of the earth surface albedo. In Proceedings of the 6th International Symposium on Physical Measurements and Signatures in Remote Sensing, Val d'Isere, France, 17–21 January 1994; pp. 707–716.
63. Lucht, W.; Schaaf, C.B.; Strahler, A.H. An algorithm for the retrieval of albedo from space using semiempirical BRDF models. *IEEE Trans. Geosci. Remote Sens.* **2000**, *38*, 977–998. [CrossRef]
64. Liang, S. An optimization algorithm for separating land surface temperature and emissivity from multispectral thermal infrared imagery. *IEEE Trans. Geosci. Remote Sens.* **2001**, *39*, 264–274. [CrossRef]
65. Bouguerzaz, F.A.; Olioso, A.; Raffy, M. Modelling radiative and energy balance on heterogeneous areas from remotely-sensed radiances. *Can. J. Remote Sens.* **1999**, *25*, 412–424. [CrossRef]

66. Boulet, G.; Kalma, J.D.; Braud, I.; Vauclin, M. An assessment of effective land surface parameterisation in regional-scale water balance studies. *J. Hydrol.* **1999**, *217*, 225–238. [CrossRef]
67. Kustas, W.; Anderson, M. Advances in thermal infrared remote sensing for land surface modeling. *Agric. For. Meteorol.* **2009**, *149*, 2071–2081. [CrossRef]

© 2017 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (http://creativecommons.org/licenses/by/4.0/).