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# Experimental mutual solubilities of CO<sub>2</sub> and H<sub>2</sub>O in pure water and NaCl solutions

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#### Abstract

Solubility measurements in conditions relevant to CO<sub>2</sub> geological storage are sparse, especially solubility data of water in the CO<sub>2</sub>-rich phase. New data are acquired by in situ Raman spectroscopy in NaCl aqueous. An effect of NaCl concentration is observed on calibration curves in the H<sub>2</sub>O-rich phase: the slope of the fitted curve  $I_{CO2}/I_{H2O} = f(s_{CO2})$ is equal to 4.97, 4.25, and 3.93 at 0, 1, and 2 M NaCl, respectively. On the contrary, the Raman signal in the CO<sub>2</sub>-rich phase is not affected by salinity. The slope of the fitted curve  $A_{CO2}/A_{H2O} = f(y_{CO2}/y_{H2O})$  is equal to 1.72 whatever the NaCl concentration.

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Keywords: CO<sub>2</sub> geological storage; solubility; Raman spectroscopy.

#### 1. Introduction

Carbon dioxide is ubiquitous on Earth, as a gas composing almost 400 ppm of the atmosphere, through solubilization as dissolved species responsible for acidification of the oceans, or trapped in minerals as a major component of sedimentary rocks.  $CO_2$  plays also an important role as a fluid circulating in the crust, originating from hydrocarbon oxidation, carbonate dissolution or magma degassing. Consequently,  $CO_2$  is present under various conditions of pressure and temperature ranging from very low values at the surface of the Earth to pressures reaching

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several kilobars and temperatures of several hundreds of degrees in deep Earth. The determination of the thermodynamic properties of  $CO_2$  is of prime interest to understand and predict its behavior through the geological and atmospheric cycles or for carbon sequestration. The geological storage of  $CO_2$  would be efficient if it is in a critical state, i.e. at temperature higher than 31.1°C and pressure higher than 73.8 bar. Supercritical  $CO_2$  behaves like a gas by filling all the available volume but has a liquid-like density, increasing storage capacity [1]. Considering a hydrostatic pressure, these conditions are reached above 800 m. At this depth, assuming a thermal gradient in the crust of around 25°C.km<sup>-1</sup>, temperature is around 35°C. These P-T conditions have to be considered as minima for an efficient  $CO_2$  storage so most of the storage reservoirs should be deeper, implying higher temperature and pressure (35 to 150°C and 80 to 250 bar or more) [2]. Numerous studies were published under various pressure-temperature-composition (*PTX*) ranges to cover  $CO_2$  storage conditions, in particular in the H<sub>2</sub>O-CO<sub>2</sub> system. In this system, one of the key parameter to define thermodynamic equilibria is  $CO_2$  solubility. If  $CO_2$  solubility in water is rather well documented at temperature higher than 100 °C, measurements at lower temperature and mutual solubility of water in the  $CO_2$ -rich phase are sparse [3,4].

The usual experimental and analytical techniques are based on the sampling but this operation may induce a perturbation of the system (pressure and temperature decrease, fluid displacement), which can be of a great influence under conditions where the system is close to instability. Moreover, it is not suitable for the analysis of the vapor phase and real-time monitoring of the system. Raman spectroscopy was successfully used to quantify dissolved  $CO_2$  in aqueous media [5–10]. In a previous study, Caumon et al. [10] presented measurements of  $CO_2$  solubility in pure water in the water-rich phase and H<sub>2</sub>O solubility in the  $CO_2$ -rich phase by in situ Raman spectroscopy at 65 °C and 100 °C using the IMAGES experimental device [11]. In this paper, we present extended data of mutual solubilities of  $CO_2$  and H<sub>2</sub>O at 100 °C in NaCl aqueous solutions. The results are compared to the previously published data at 100 °C in pure water using the same experimental device.

#### Nomenclature

A <sub>CO2</sub>	area of the Raman peaks of gaseous or supercritical CO <sub>2</sub>
A <sub>H2O</sub>	area of the Raman peak of gaseous water
I <sub>CO2</sub>	intensity (peak height) of the Raman peak of dissolved CO <sub>2</sub>
I <sub>H2O</sub>	intensity (peak height) of the Raman bending vibration band of liquid water
SCO2	solubility of CO <sub>2</sub> in liquid water in mol.kg <sup>-1</sup> <sub>H2O</sub>
YCO2	molar fraction of $CO_2$ in the $CO_2$ -rich (vapour) phase
Ун20	molar fraction of $H_2O$ in the CO <sub>2</sub> -rich (vapour) phase
-	

#### 2. Material and methods

#### 2.1. Experimental device

The IMAGES experimental device was described in Caumon et al. [10]. Briefly, it consisted of an autoclave of 2L made of 316 stainless steel. A Teflon bowl was placed in the autoclave to prevent corrosion and laser reflections on the metallic walls. Temperature was set to  $100.0 \pm 0.1$  °C. Pressure and temperature were continuously recorded. Pressure variations were of  $\pm 0.3$  bar at equilibrium. Two Raman probes (immersion optics from Kaiser Optical Systems, model IO-18L-VIS) were crimped in the cover and connected via an optical fiber to a RXN1 Raman spectrometer (Kaiser Optical Systems, Inc.). One probe was in the H<sub>2</sub>O-rich phase, and the other one in the CO<sub>2</sub>-rich phase (Fig. 1).



Fig. 1: the IMAGES experimental device

#### 2.2. Experimental procedure

NaCl solutions were prepared by weighting of NaCl and Milli-Q water (18.2 M $\Omega$ .cm) to obtain 1.02 ± 0.05 M and 2.00 ± 0.05 M NaCl. The autoclave was filled with 1 L of solution. After temperature equilibrium at 100 °C (~24h), CO<sub>2</sub> was injected step by step up to 200 bar. Equilibrium occurred 4-12 h after injection. The total experiment time was of two weeks.

#### 2.3. Data analysis

Raman spectra were recorded continuously in the CO<sub>2</sub>-rich phase till equilibrium (8 accumulations of 30 s). Raman spectra of the CO<sub>2</sub>-rich phase were recorded at equilibrium (8 accumulations of 15 s). A minimum of 10 spectra were recorded at each step. The spectra were corrected by the spectrometer response (calibration using the HCA accessory, Kaiser Optical Systems, Inc.). In the H<sub>2</sub>O-rich phase, the signal from Teflon superimposed to the one of dissolved CO<sub>2</sub>. The peak intensity ratio were corrected from the apparent " $I_{CO2} / I_{H2O}$ " measured before the injection of CO<sub>2</sub> (blank).

Peak areas and peak intensities were measured using the Labspec6 software following the procedure described in Caumon et al. [10]. The only data processing was the subtraction of a straight baseline. Peak area ratios  $A_{CO2} / A_{H2O}$  were plotted as a function of molar fraction ratio  $y_{CO2} / y_{H2O}$  for the CO<sub>2</sub>-rich phase. Peak intensity ratios  $I_{CO2} / I_{H2O}$  were plotted as a function of CO<sub>2</sub> solubility  $s_{CO2}$  (mol.kg<sup>-1</sup>H<sub>2O</sub>). Solubility and molar fraction were determined using the model of Duan and Sun [12] at the pressure and temperature of the experiment. A linear regression was used to fit the data.

#### 3. Results

Raman data are plotted *vs.* solubility or molar fraction ratio in Fig. 2. In the H<sub>2</sub>O-rich phase (liquid phase), the dispersion of the data points is low: the standard deviation  $\sigma$  of the repeated measurements is lower than 2%. Linear regression yields R<sup>2</sup> better than 0.998. The slope of the lines are 4.97, 4.25 and 3.93 at 0, 1, and 2 M, respectively. It decreases when salinity increases.



Fig. 2: (a) Raman peak intensity ratios as a function of CO<sub>2</sub> solubility and salinity (H<sub>2</sub>O-rich phase); (b) Raman peak area ratios as a function of CO<sub>2</sub>/H<sub>2</sub>O molar fraction ratio and salinity (CO<sub>2</sub>-rich phase). P = 5-200 bar, T= 100°C.

In the CO<sub>2</sub>-rich phase, the dispersion of the data points in Fig. 2 is a little higher ( $1\sigma \le 5\%$ ). This is due to the low amount of gaseous H<sub>2</sub>O in CO<sub>2</sub> (20% at 5 bar, and down to 2% at ~ 150 bar) and, so to a lower signal-to-noise ratio on Raman spectra. Only the data points recorded at pressure lower than ~130-150 bar are plotted. Beyond this point, the molar ratio reverses and the correlation is no more linear [10]. The linear regression yields slopes of 1.72 ± 0.05 whatever the salinity.

#### 4. Discussion

Salinity affects calibration curves in the  $H_2O$ -rich phase. It is probably due to a deformation of the Raman signal of water with salinity. This effect is describe by Wu et al. [13] for the stretching vibration band of water. Raman cross section in liquids depends on the local field factor L. L is sensitive to the refraction index of the medium, and this index changes with salinity. As a consequence, the Raman cross section of water increases with NaCl concentration, but the local field factor can only partially explain the variation of Raman cross section [13]. This effect is first evidenced here for the bending vibration band of liquid water. In the CO<sub>2</sub>-rich phase, salinity doesn't affect calibration. As NaCl is absent from this phase, the only effect is on the molar fraction ratio *vs.* pressure, but the Raman signal is not affected.

The solubilities experimentally measured in this study are plotted in Fig. 3 as a function of pressure. In the H<sub>2</sub>Orich phase, the solubility of CO<sub>2</sub> is coherent with the Duan and Sun model [12] over the full pressure and salinity ranges. In the CO<sub>2</sub>-rich phase, the molar fraction ratio seems to be more dispersed above 100 bar. These points were not included in the fitting procedure (Fig. 2). Applying the calibration, our data points are in between the model of Duan and Sun [12] and Spycher et al. [4] and within the range of literature experimental data [14–17].



Fig. 3: (a) solubility of CO<sub>2</sub> at 100 °C as a function of pressure and NaCl concentration. (b) molar fraction ratio  $y_{CO2}/y_{H2O}$  in the CO<sub>2</sub>-rich phase at 100 °C as a function of NaCl salinity; comparison with the thermodynamic models of Duan and Sun [12] and Spycher et al. [4] and literature experimental data points [14–17].

#### 5. Conclusion

The determination of mutual solubilities of  $H_2O$  and  $CO_2$  is essential for the modelling and understanding of  $CO_2$  geological storage in deep saline aquifer. Here it is demonstrated that the IMAGES experimental device, coupled to a Raman spectrometer, makes possible to produce new solubility data at 100 °C up to 200 bar, in the presence of NaCl. Moreover, it is now possible to measure  $H_2O$  solubility data in the  $CO_2$ -rich phase where data are very sparse.

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